



1 A two-stage Bias-Correction and Super-Resolution Framework 2 for Post-Processing Climate Model Outputs

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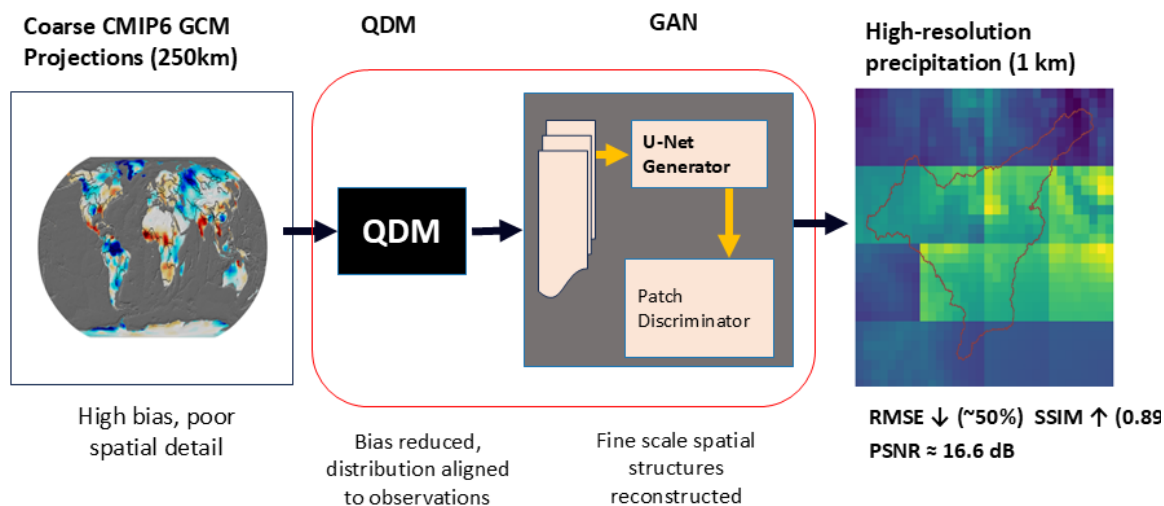
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24 Abstract

25 General circulation models (GCMs) underpin climate change assessments, yet their coarse spatial resolution and
26 systematic biases constrain their direct use in regional applications. Post-processing approaches such as bias correction
27 and statistical downscaling are therefore widely applied, yet these steps are often implemented independently, leading
28 to inconsistencies between corrected statistics and spatial structure. This study presents a reproducible two-stage
29 framework that integrates Quantile Delta Mapping (QDM) for bias correction with a deep learning-based super-
30 resolution method to improve the statistical fidelity and spatial detail of climate model outputs. The framework is
31 evaluated using precipitation, runoff, and evapotranspiration from three Coupled Model Intercomparison Project Phase
32 6 (CMIP6) GCMs (Canadian Earth System Model (CanESM5), Hadley Centre Global Environment Model (HadGEM3-
33 GC31-LL), and Max Planck Institute Earth System Model (MPI-ESM1-2-HR)), over the Hadejia-Jama'are River Basin
34 in northern Nigeria. We first demonstrate that raw model outputs exhibit substantial biases, with domain-averaged root
35 mean square errors (RMSE) of 48.6-57.3 mm month⁻¹ for precipitation, 0.93-7.51 mm month⁻¹ for runoff, and 33.8-58.7
36 mm month⁻¹ for evapotranspiration. QDM substantially reduces systematic errors, lowering precipitation RMSE to 23.8-



37 27.8 mm month⁻¹, runoff RMSE to 0.24-1.85 mm month⁻¹, and evapotranspiration RMSE to 3.7-4.3 mm month⁻¹, while
38 preserving projected distributional changes, as confirmed by Kolmogorov-Smirnov ($D \leq 0.072$) and Wasserstein (\leq
39 11.86) metrics. In the second stage, a conditional Generative Adversarial Network (GAN) super-resolves the bias-
40 corrected precipitation fields from ~250 km to 14 km (regridded to 1 km), with performance assessed using
41 complementary pixel-wise, structural, and distributional metrics, and outperforming bilinear and bicubic interpolation
42 in terms of spatial coherence and structural similarity. The proposed QDM-GAN framework is fully documented and
43 reproducible, with openly available code and data sources, and is intended as a modular post-processing tool that can
44 support downstream modeling applications requiring bias-corrected, high-resolution climate inputs.
45



46

47 **Key words:**

48 Generative Adversarial Networks (GANs), Precipitation, CMIP6, GCMs, RCMs, sub-km spatial resolution.

49 **1. Introduction**

50 General circulation models (GCMs) and regional climate models (RCMs) form the backbone of contemporary climate
51 simulations, providing physically consistent representations of large-scale atmospheric and land-surface processes (Flato
52 et al., 2013). However, the spatial resolution of most GCMs, typically on the order of 100-250 km, remains insufficient
53 for many regional and local-scale applications (Yu et al., 2018, Thao et al., 2022; Wang 2023). At these resolutions,
54 sub-grid heterogeneity associated with topography, land cover, and convective processes is poorly resolved, and
55 systematic biases relative to observations persist (Maraun et al., 2017; Thao et al., 2022). When GCM outputs are used
56 directly as forcings for impact or Earth-system component models, these limitations can propagate into downstream



57 analyses, motivating the widespread use of post-processing techniques (Labeurthre et al., 2024). At the same time,
58 evaluating whether post-processed outputs preserve both statistical fidelity and realistic spatial structure remains an
59 important challenge in climate model assessment.

60 Downscaling approaches aimed at bridging the scale gap between global simulations and local climate information can
61 be broadly categorized into dynamical and statistical methods (Calel and Stainforth, 2017; Hudson, 2021). Dynamical
62 downscaling, implemented through RCMs, explicitly resolves regional processes and has achieved grid spacings as fine
63 as a few kilometers in recent studies (Afrizal and Surussavadee, 2018; Ajjur, 2024; Merks et al., 2020). Despite these
64 advances, RCM simulations remain computationally expensive and inherit uncertainties from both the driving GCM and
65 the regional model configuration, including sensitivity to boundary conditions and parameterizations (Ghimire et al.,
66 2018; Goodwin and Cael, 2021; Wang and Tian, 2022). These uncertainties can be particularly consequential for
67 hydrological variables such as runoff and evapotranspiration (Teng et al., 2015; Wang and Tian, 2022).

68 Statistical approaches to downscaling offer computational efficiency and flexibility and therefore remain central to many
69 climate modeling workflows, particularly where ensemble processing or long time series are required (Hudson, 2021).
70 Bias correction is a key component of statistical post-processing, aiming to reduce systematic discrepancies between
71 modeled and observed climatology (Cannon et al., 2015; Maraun et al. 2017). Quantile-based bias-correction methods
72 are among the most widely used and have been shown to effectively adjust mean state and variability (Maraun, 2016).
73 However, classical quantile mapping approaches typically assume stationary biases and may distort projected climate
74 change signals, particularly for extremes (Maraun et al., 2017; Rajulapati and Papalexiou, 2023).

75
76 Quantile Delta Mapping (QDM) was introduced to address this limitation by preserving relative or absolute changes in
77 quantiles between historical and future simulations while correcting baseline biases (Cannon et al., 2015; Pierce et al.,
78 2015) QDM has been applied successfully across multiple hydroclimatic variables and regions and is widely used in
79 climate impact studies (Tong et al., 2021; Ugolotti et al., 2023). Nevertheless, QDM operates independently at each grid
80 point and does not address the loss of spatial detail inherent in coarse-resolution climate model outputs (Gergel et al.,
81 2024).

82 In parallel, recent developments in machine learning have enabled new approaches for statistical downscaling through
83 image super-resolution techniques. Convolutional neural networks and generative adversarial networks (GANs) have
84 demonstrated skill in reconstructing fine-scale spatial structure from coarse inputs by learning mappings from paired
85 low- and high-resolution data (Stengel et al., 2020; Wang and Tian, 2022). GAN-based approaches, in particular, have
86 been shown to recover sharp gradients, terrain-induced variability, and localized extremes that are smoothed by
87 conventional interpolation methods (Oyama et al., 2023). However, many machine-learning-based downscaling studies
88 focus primarily on spatial enhancement and apply limited or no bias correction, which can result in high-resolution fields
89 that retain systematic statistical errors (Besombes et al., 2021; Rampal et al., 2024).



90 Despite progress in both bias correction and machine-learning-based downscaling, relatively few studies have integrated
91 these components within a unified and reproducible evaluation workflow. As a result, bias correction and super-
92 resolution are often treated as separate steps, with limited assessment of how corrections applied in one stage influence
93 statistical fidelity and spatial realism in the other (Gergel et al., 2024; Rajulapati and Papalexiou, 2023; Ugolotti et al.,
94 2023). This fragmentation complicates reproducibility and comparison across studies, particularly in the context of Earth
95 system model post-processing.

96 This paper addresses this gap by developing and evaluating a two-stage post-processing framework that combines
97 Quantile Delta Mapping with a conditional GAN-based super-resolution model. The framework is applied to multiple
98 CMIP6 GCMs and assessed using complementary statistical, distributional, and structural metrics following established
99 practices in bias-correction and super-resolution literature (Cannon et al., 2015; Stengel et al., 2020). Every processing
100 step is documented and supported by open-access code and data to facilitate transparency and reproducibility.
101 Consequently, this study provides a transferable workflow for generating and evaluating bias-corrected, high-resolution
102 climate model outputs intended for downstream hydrological and environmental applications.

103 **2. Data and Methods**

104 **2.1 Study Area Description**

105 The two-stage framework is demonstrated over Hadejia-Jama'are River Basin (HJRB) in northern Nigeria,
106 geographically bounded by 10°-13° N and 7°-11.5° E (Figure 1). The basin serves as a critical water resource hub for
107 agricultural production, domestic water supply, and ecological stability. Despite its importance, the basin suffers from
108 notable data scarcity, which complicates both hydrological modeling and long-term resource planning. The region's
109 climate is characterized by high evapotranspiration, marked wet and dry seasons, and pronounced rainfall variability,
110 rendering local communities and ecosystems highly susceptible to climate the climate change impacts.

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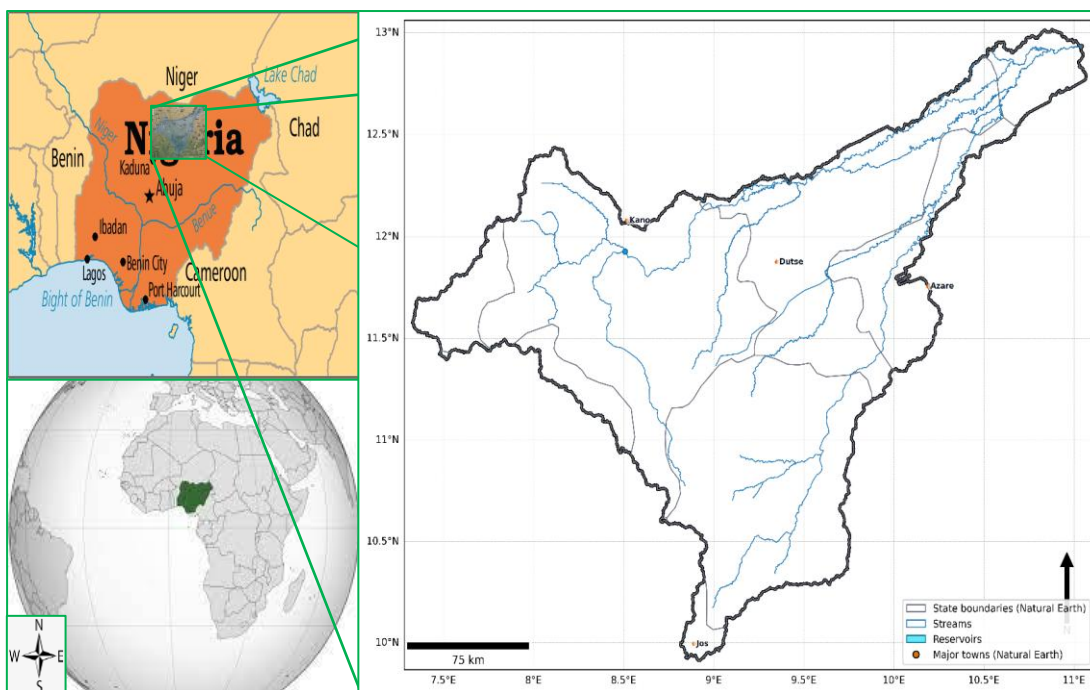


Figure 1: Study area location

113 2.2 Data Sources

114 Observed satellite data

115 Precipitation:

116 Precipitation estimates for the study area were obtained from the Climate Hazards Group InfraRed Precipitation with
117 Stations (CHIRPS) product (Funk et al., 2015) which merges satellite infrared data with in-situ station observations to
118 capture high-resolution rainfall variability. It provides global coverage at a spatial resolution of 0.05° (~5 km) from
119 1981 to near-real-time, updated monthly or more frequently (Horn et al., 2018). CHIRPS integrates infrared Cold
120 Cloud Duration (CCD) measurements from geostationary satellites with quality-controlled rain-gauge observations,
121 thereby reducing biases often found in purely satellite-based products (Bandara et al., 2021). Its relatively fine spatial
122 resolution and long-term temporal record make CHIRPS well-suited for hydrological applications in data-scarce
123 regions like the Hadejia-Jama'are River Basin. CHIRPS has been validated and found to be effective in monitoring
124 extreme precipitation events, contributing to a better understanding of rainfall variability in various regions (Duan et
125 al., 2019). Studies have shown that CHIRPS performs comparably to other satellite-based precipitation products,
126 making it a reliable choice for hydrological modeling (Bandara et al., 2021; Bichet and Diedhiou, 2018). Furthermore,



127 CHIRPS has been utilized in various hydrological evaluations, demonstrating its utility in regions with limited ground-
128 based rainfall data (Belay et al., 2019; Pandey et al., 2021). The integration of CHIRPS data into hydrological models
129 enhances the accuracy of simulations, particularly in areas where traditional rain gauge networks are sparse which is
130 typical of HJRB.

131

132 **Evapotranspiration data:**

133 The Evapotranspiration (ET) datasets were obtained from MODIS/061/MOD16A2GF product, accessed through the
134 NASA Earthdata portal (Running et al., 2021). This dataset, referred to as MODIS ET, employs a Penman-Monteith-
135 type algorithm that integrates various meteorological inputs, such as air temperature and solar radiation, along with
136 MODIS land cover and leaf area index (LAI) products). The dataset is as an 8-day composite with a nominal spatial
137 resolution of 500 meters, upscaled to a 1-kilometer resolution and aggregated to monthly time steps in this work. The
138 physically based approach of the MODIS ET product captures dynamic vegetation conditions, accounting for both
139 energy-driven and water-driven controls on ET, which is crucial for accurate modeling of water balance (Kim et al.,
140 2012). However, it is important to note that discrepancies have been observed between MODIS ET estimates and ground
141 measurements, often attributed to biases in the input data from both MODIS and global meteorological datasets (Aguilar
142 et al., 2018; Kim et al., 2012). The MODIS ET dataset has been widely adopted in global and regional water balance
143 studies, providing a consistent and regularly updated resource since its inception in 2000. Its application extends to
144 various fields, including hydrology, ecology, and environmental science, making it a vital tool for researchers aiming to
145 understand evapotranspiration dynamics on both regional and global scales (Liu et al., 2020).

146

147 **Surface Runoff:**

148 Surface runoff estimates for this study were sourced from the Global Land Data Assimilation System (GLDAS), a
149 NASA-led initiative that integrates satellite-based and ground-based observations into sophisticated land surface models
150 (Qing et al., 2021). The datasets were accessed through the NASA Goddard Earth Sciences Data and Information
151 Services Center (GES DISC) (Beaudoin and Rodell, 2024). GLDAS outputs are available at various spatial resolutions,
152 typically 0.25° or 0.10°, and temporal resolutions ranging from 3-hourly to monthly. For this research, a monthly
153 aggregated surface runoff (Qs) dataset was utilized, covering the period from 2000 to the present, which aligns with the
154 available MODIS ET data. A physically consistent approach was employed by GLDAS to produce land-surface fluxes
155 through the assimilation of satellite-derived precipitation, snow cover, and soil moisture products. This integration was
156 designed to mitigate biases that may arise from purely modeled runoff estimates, thus enhancing the reliability of the
157 data (Gruber et al., 2019).

158 Table 1 summarized the features of each of the satellite-observed datasets. Each of these offers distinct advantages for
159 hydrological modeling, especially in regions with limited ground-based observations. CHIRPS supplies high-resolution



160 rainfall estimates critical for driving surface-water dynamics, while MODIS ET captures seasonally evolving water use
 161 by vegetation, and GLDAS runoff helps constrain near-surface hydrological fluxes under changing precipitation
 162 regimes. Together, they enable a more robust quantification of water balance components across the HJRB, supporting
 163 both validation of bias-corrected GCM outputs and subsequent downscaling efforts.

164

165 Table1: Summary of satellite-observed variables

166

Variable	Data Source	Temporal Coverage	Spatial Resolution	Temporal Resolution	References
Precipitation (mm)	CHIRPS (Climate Hazards Group)	1981-Present, near real-time	0.05° (~5 km)	Daily, Pentadal, Monthly	(Center, 2025) and (Funk et al., 2015)
Evapotranspiration (mm)	MODIS (MOD16A2GF)	2000-Present	500 m to 1 km	8-day composites (aggregable)	(Running et al., 2021)
Surface Runoff (mm)	GLDAS (e.g., Noah LSM)	2000-Present (varies by version)	0.25° or 0.10° (~28-11 km)	3-hourly, Daily, or Monthly	(Beaudoing and Rodell, 2024)

167 **CMIP6 GCs Model Datasets**

168 Three advanced Earth System Models from the Coupled Model Intercomparison Project Phase 6 (CMIP6) (MPI-ESM1-
 169 2-HR, HadGEM3-GC31-LL, and CanESM5) were selected to represent both historical (pre-2015) and future scenarios
 170 under distinct Shared Socioeconomic Pathways (SSPs). The data were downloaded from Earth System Grid Federation
 171 (ESGF) under Analytics and Informatics Management Systems (AIMS) at Lawrence Livermore National Laboratory
 172 data portal (ESGF, 2025) which archives and distributes CMIP6 outputs in compliance with standardized metadata and
 173 file formats, specifically NetCDF.

174 The selection of these CMIP6 models was justified by a recent study conducted by Wada et al. (Wada et al., 2023) which
 175 examined the performance of CMIP6 models (100 ensembles) using an advanced envelop-based selection approach in
 176 the same study area. The study emphasized that MPI-ESM1-2-HR, HadGEM3-GC31-LL, and CanESM5 exhibit lower
 177 biases and better seasonal cycle representation, making them suitable for climate impact assessments in the area.
 178 Furthermore, the selected models have been shown to effectively simulate extreme climate events, which is key for
 179 understanding potential future impacts. For instance, HadGEM3-GC31-LL has been recognized for its robust
 180 performance in simulating extreme precipitation and temperature events (Nishant et al., 2022).

181 Monthly mean fields for precipitation (pr), surface runoff (mrros), and evapotranspiration (evspsbl) were utilized,
 182 employing the r1i1p1f1 realization for each model to ensure consistent forcing protocols and initial conditions across
 183 all scenarios. Table 2 presented a summary of the model and detail of the variables used which are publicly accessible
 184 via the CMIP6 license and data usage policies.

185



186 Table 2: Summary of CMIP6 Datasets used

Models	Historical (years)	Future (2015-2100) SSPs	Variables	Native Resolution
MPI-ESM1-2-HR (from Planck Institute for Meteorology, Germany)	1850-2014	SSP126, SSP245, SSP585	pr, mrros, evspsbl	~1.0° atmosphere (approx)
HadGEM3-GC31-LL (from Met Office Hadley Centre, United Kindom)	1850-2014	SSP126, SSP245, SSP585	pr, mrros, evspsbl	~1.25° × 1.875°
CanESM5 (from Canadian Centre for Climate Modelling and Analysis, Canada)	1850-2014	SSP126, SSP245, SSP585	pr, mrros, evspsbl	~2.8° (atmosphere)

187

188 In this study, the historical data (2000-2014) were used as a baseline, while future projections (2015-2100) under
 189 *SSP126*, *SSP245*, and *SSP585* are used to explore a range of emission and socioeconomic pathways.

190 2.3 Methodology

191 2.3.1 Quantile Delta Mapping (QDM) for Bias Correction

192 Biases in Global Climate Model (GCM) outputs were corrected using Quantile Delta Mapping (QDM), which extends
 193 the standard quantile mapping (QM) approach by aligning the quantile structure of simulated data with observations
 194 while preserving the model-projected distributional changes (Cannon et al., 2015; Pierce et al., 2015). Unlike classical
 195 QM, which can suppress or overcorrect projected changes, QDM ensures that relative or absolute shifts in quantiles are
 196 retained.

197 Let $X_m(t)$ denote the model simulation at time t , and $X_o(t)$ represent the corresponding observed value. The percentile
 198 of a model value under the future distribution is computed as:

$$199 \quad p(t) = F_{mod,fut}(X_{m,fut}(t)) \quad (1)$$

200 where $F_{mod,fut}$ is the future model cumulative distribution function (CDF). The corresponding baseline mapped value is
 201 obtained as:

$$202 \quad \tilde{X}(t) = F_{obs,ref}^{-1}(p(t)), \quad (2)$$

203 with $F_{obs,ref}^{-1}$ denoting the inverse CDF of the observed reference distribution. QDM then applies a quantiles specific
 204 adjustment (Eq. 3, 4).

$$205 \quad \text{Additive form (e.g., temperature): } X^*(t) = \tilde{X}(t) + [F_{mod,fut}^{-1}(p) - F_{mod,ref}^{-1}(p)] \quad (3)$$

$$206 \quad \text{Multiplicative form (e.g., precipitation, runoff): } X^*(t) = \tilde{X}(t) \times \frac{F_{mod,fut}^{-1}(p)}{F_{mod,ref}^{-1}(p)} \quad (4)$$



207 The QDM procedure was applied independently to precipitation (pr), evapotranspiration (evspsbl), and runoff (mrros).
208 Both observed and model datasets were interpolated to a common 128×128 grid. Following (Cannon et al., 2015) and
209 (Tong et al., 2021), the implementation comprised three steps: (i) estimation of empirical CDFs for observed and
210 modeled data, (ii) derivation of quantile-wise correction factors, and (iii) application of these factors to future CMIP6
211 simulations under SSP126, SSP245, and SSP585.

212 Validation of the QDM-corrected simulations against historical observations involved:

- 213 a) Quantitative assessment using Root Mean Squared Error (RMSE), Nash-Sutcliffe Efficiency (NSE), and coefficient
214 of determination (R^2).
- 215 b) Qualitative assessment through time-series plots, scatter diagrams, and quantile-quantile (Q-Q) plots.

216 This ensured both statistical improvements and the preservation of distributional integrity in corrected simulations.

217 2.3.2 GAN-Based Downscaling

218 To further enhance spatial resolution, Generative Adversarial Networks (GANs) (Goodfellow et al., 2014) were
219 employed. In this framework, a **Generator (G)** learns to transform coarse-resolution GCM fields into high-resolution
220 climate fields, while a **Discriminator (D)** learns to distinguish between real high-resolution data and generated outputs.
221 A conditional GAN (cGAN) structure was implemented (Isola et al., 2017; Mirza and Osindero, 2014), conditioning
222 both G and D on low-resolution model input.

223 Loss Functions:

224 The generator was trained using a composite objective (Eq. 3, 4 and 5)

$$225 \quad L_G = \alpha L_{MSE} + \beta L_{adv}, \quad (3)$$

226 where L_{MSE} penalizes pixel-wise errors, and L_{adv} enforces realism. The discriminator optimized a binary cross-
227 entropy objective:

$$228 \quad L_D = -E[\log D(y_{true})] - E[\log (1 - D(G(x_{coarse})))], \quad (4)$$

229 where $D(y_{true})$ denotes the probability assigned to real high-resolution data. A perceptual loss (Ledig et al., 2017) was
230 also integrated:

$$231 \quad L_G = L_{GAN} + \lambda_{L1} \|\hat{y} - y\|_1 + \lambda_{perc} \|\Phi(\hat{y}) - \Phi(y)\|_1, \quad (5)$$

232 where $\Phi(\cdot)$ denotes VGG-16 feature maps, $\lambda_{L1} = 100$, and $\lambda_{perc} = 0.1$.



233 **Data and Preprocessing:**

234 Monthly precipitation from bias-corrected CMIP6 models ($\sim 250\text{km}$ resolution) and 14 km reanalysis data were paired
235 for 180 months. Each pair was stored as a 128×128 array, normalized to $[-1,1]$ following Radford et al. (2016). Data
236 were split into 80% training and 20% validation sets, with augmentation (rotations, flips) applied only to training
237 samples.

238 **Network Architecture and Training:**

- 239 • Generator: A U-Net architecture with four encoder-decoder levels and skip connections.
- 240 • Discriminator: 64×64 PatchGAN with spectral normalization.
- 241 • Optimization: Adam optimizer ($\beta_1 = 0.5, \beta_2 = 0.999$); learning rates initialized at 2×10^{-4} (generator) and
242 4×10^{-4} (discriminator), annealed to 10^{-6} .

243 Training was performed on an NVIDIA GeForce GTX 1050 GPU with a batch size of 4, using early stopping based on
244 minimum validation loss.

245 **Assessment Framework and Evaluation Metrics:**

246 Downscaling performance was assessed using three complementary and widely accepted metrics in super-resolution
247 (SR): Mean Squared Error (MSE), Peak Signal-to-Noise Ratio (PSNR), and Structural Similarity Index (SSIM). MSE
248 was computed on the rescaled rainfall fields to assess amplitude errors in the physical variable after normalization;
249 PSNR (a monotonic transform of MSE) provided a distortion measure in decibels that is standard in SR benchmarks;
250 and SSIM captured luminance/contrast/structure consistency to better reflect perceived and structural quality than pixel-
251 wise errors alone. The joint use of MSE/PSNR and SSIM follows common SR practice and addresses the known
252 distortion-perception trade-off (Arabbœv et al., 2024).

253 Results were additionally benchmarked against bicubic interpolation, the conventional “no-learning” baseline in SR
254 (including remote-sensing SR) because it is deterministic, fast, and grounded in cubic convolution theory. Demonstrating
255 improvements over bicubic therefore isolates the added value of learned (adversarial) models in reconstructing high-
256 frequency detail beyond interpolation’s inherent smoothing (Keys, 1981; Wang et al., 2022).

257 Finally, inclusion of an adversarially trained model (GAN) enabled assessment of the extent to which perceptual
258 optimization improves spatial realism relative to conventional interpolation and pixel-wise optimization approaches,
259 showing that adversarial/perceptual losses can restore sharper, more realistic textures relative to pure MSE optimization,
260 while acknowledging the perception-distortion trade-off in evaluation (Ledig et al., 2017).



261 **Computational Environment and Reproducibility**

262 All data preprocessing, bias correction, model training, and evaluation were conducted in a Windows 10 (64-bit)
263 environment using Python (v3.9). The computational platform was a Dell G3 3579 workstation equipped with an Intel®
264 Core™ i5-8300H CPU (2.30 GHz; 4 cores, 8 threads) and 24 GB RAM. Deep learning experiments were executed on
265 an NVIDIA GeForce GTX 1050 GPU (4 GB VRAM) using NVIDIA driver version 560.94 and CUDA version 12.6.

266 Bias-correction and data-processing steps were implemented using NumPy (v1.23), pandas (v1.5), xarray (v2023.1),
267 and SciPy (v1.9). Geospatial operations and visualization were performed using rasterio (v1.3), GeoPandas (v0.12),
268 Matplotlib (v3.7), and Cartopy (v0.21). The deep learning architecture was implemented in PyTorch (v2.0) with CUDA
269 acceleration enabled.

270 Training of the conditional GAN was performed for 200 epochs using a batch size of 4. On the GTX 1050 GPU (4 GB
271 VRAM), training required approximately 8-10 hours of wall-clock time per model. Inference for a single monthly
272 precipitation field required less than 3 seconds. These runtimes demonstrate that the framework can be executed on mid-
273 range consumer hardware without requiring high-performance computing infrastructure.

274 To ensure reproducibility, fixed random seeds were applied across NumPy and PyTorch libraries. Deterministic training
275 settings were enabled where supported by the backend framework. Training and validation splits were generated
276 deterministically using predefined seeds to guarantee consistent data partitioning across runs. All hyperparameters,
277 including learning rates, optimizer parameters ($\beta_1 = 0.5$, $\beta_2 = 0.999$), loss-function weights (λ_{L1} and λ_{perc}), batch
278 size, and epoch count, are explicitly defined in configuration files within the project repository.

279 The complete source code, including preprocessing scripts, QDM implementation, neural network architectures,
280 configuration files, training routines, and evaluation utilities, is publicly available at
281 <https://doi.org/10.5281/zenodo.18746067> (Ibrahim et al., 2026). All input datasets (CMIP6 via ESGF, CHIRPS,
282 MODIS, and GLDAS) are publicly accessible through the official portals cited in the Data Availability section.

283 **3. Results and Discussion**

284 The assessment framework adopted in this study evaluates post-processing performance across three complementary
285 dimensions: (i) statistical fidelity, measured using bias, RMSE, Kolmogorov-Smirnov, and Wasserstein metrics; (ii)
286 structural fidelity, evaluated using SSIM and visual reconstruction diagnostics; and (iii) spatial enhancement capability,
287 assessed relative to interpolation baselines. This multi-metric strategy is necessary because no single metric adequately



288 characterizes the performance of climate-model post-processing methods, particularly for hydroclimatic fields
 289 containing sharp gradients and localized extremes.

290 3.1 QDM Results

291 Pre- and Post-Correction Model Performance

292 Evaluation of raw GCM outputs revealed significant discrepancies relative to observed hydroclimatic records,
 293 particularly for precipitation, surface runoff, and evapotranspiration (Table 3). Precipitation exhibited strong correlations
 294 ($R^2 = 0.84 - 0.92$) but large negative biases, with MPI-ESM1-2-HR showing the highest temporal correspondence (
 295 $R^2 = 0.917$) yet underestimating rainfall by $31.68\text{mm}/\text{month}$, and CanESM5 yielding the smallest bias ($-4.57\text{mm}/$
 296 month) but with slightly weaker correlation. Surface runoff was best captured by MPI-ESM1-2-HR ($R^2 = 0.742$, bias
 297 $= +0.45\text{mm}/\text{month}$), while HadGEM3 overestimated runoff substantially ($+4.22\text{mm}/\text{month}$). Evapotranspiration
 298 was consistently overpredicted, with CanESM5 producing the highest bias ($+44.98\text{mm}/\text{month}$). These findings
 299 confirm that raw GCMs, despite moderate-to-high correlations, contained substantial magnitude and distributional errors
 300 that limited their suitability for hydrological applications.

301 Table 3: Monthly domain-averaged evaluation metrics for the raw outputs of HadGEM3, MPI-ESM1-2-HR, and
 302 CanESM5.

Model	Variable	Correlation (R^2)	RMSE (mm)	Bias (mm/month)
HadGEM3	pr	0.855	57.253	-26.764
	mrros	0.467	7.505	+4.222
	evspsbl	0.577	42.749	+28.190
MPI-ESM1-2-HR	pr	0.917	52.442	-31.68
	mrros	0.742	0.930	+0.452
	evspsbl	0.844	34.279	+23.562
CanESM5	pr	0.840	48.569	-4.574
	mrros	0.334	2.265	+0.067
	evspsbl	0.632	58.102	+44.980

303
 304 Following QDM correction, model performance improved markedly (Table 4). Precipitation biases narrowed to
 305 $+5.60 - 10.90\text{mm}/\text{month}$ and RMSE values dropped by nearly half ($23.79 - 27.84\text{mm}$). Surface runoff errors were
 306 reduced to near zero, with RMSE consistently below $2\text{mm}/\text{month}$. Evapotranspiration showed the most dramatic
 307 improvement, with mean biases converging to $\sim +1.7\text{mm}/\text{month}$ and RMSE reduced by over 85% ($3.71 - 4.30\text{mm}$).
 308
 309



310 **Table 4:** Comparison of monthly domain-averaged Bias and RMSE values for raw (pre-correction) and QDM-adjusted
 311 (post-correction) outputs of HadGEM3, MPI-ESM1-2-HR, and CanESM5.

Model	Variable	Bias Pre (mm/month)	Bias Post (mm/month)	RMSE Pre (mm)	RMSE Post (mm)
HadGEM3	pr	-26.76	+5.60	57.25	24.91
	mrros	+4.22	-0.072	7.51	0.24
	evspsbl	+28.19	+1.69	42.75	4.30
MPI-ESM1-2-HR	pr	-31.68	+6.23	52.44	23.79
	mrros	+0.45	+0.068	0.93	0.27
	evspsbl	+23.56	+1.69	34.28	3.86
CanESM5	pr	-4.57	+10.90	48.57	27.84
	mrros	+0.07	+1.14	2.27	1.85
	evspsbl	+44.98	+1.69	58.10	3.71

312
 313 QDM correction markedly reduced systematic errors across precipitation (**pr**), surface runoff (**mrros**), and
 314 evapotranspiration (**evspsbl**). Relative to the raw GCM fields, precipitation RMSE was nearly halved; runoff RMSE
 315 declined to < 2 mm month⁻¹; and evapotranspiration RMSE decreased by > 85%. Distributional agreement also
 316 improved: both the Kolmogorov-Smirnov statistic (D) and the 1-Wasserstein distance were reduced across variables
 317 (see Table 5).

318
 319 **Table 5:** Domain-averaged performance metrics (Bias, RMSE, Kolmogorov-Smirnov D, and Wasserstein distance) after
 320 QDM correction for HadGEM3, MPI-ESM1-2-HR, and CanESM5.

Model	Variable	Mean_Obs	Mean_Corr	Domain_Bias	KS_D	KS_p	Wasserstein	RMSE
MPI-ESM1-2-HR	pr	71.867	78.092	6.226	0.0441	2.63E-05	6.263	23.785
	mrros	0.236	0.304	0.068	0.0601	1.74E-09	0.068	0.266
	evspsbl	12.842	12.842	1.693	0.0108	0.889	0.0178	3.862
HadGEM3-GC31-LL	pr	71.867	77.471	5.604	0.0579	0.0155	5.625	24.91
	mrros	0.236	0.164	-0.072	0.2319	1.56E-34	0.122	0.24
	evspsbl	12.842	12.841	1.692	0.0111	0.9999	0.0174	4.295
CanESM5	pr	71.867	82.761	10.895	0.0717	0.0468	11.862	27.842
	mrros	0.236	1.373	1.137	0.2696	1.35E-23	1.1379	1.845
	evspsbl	12.842	12.84	1.691	0.0111	1	0.0188	3.706

321



322 The corrected simulations reproduced seasonal cycles and interannual variability more faithfully, and Q-Q plots. Figures
323 2 showed closer adherence to the 1:1 line across quantiles. Collectively, these findings underscore the inadequacy of
324 uncorrected GCM outputs for hydrological applications and demonstrate the efficacy of QDM in correcting both mean
325 state and variability, thereby providing a robust basis for subsequent downscaling.

326



327

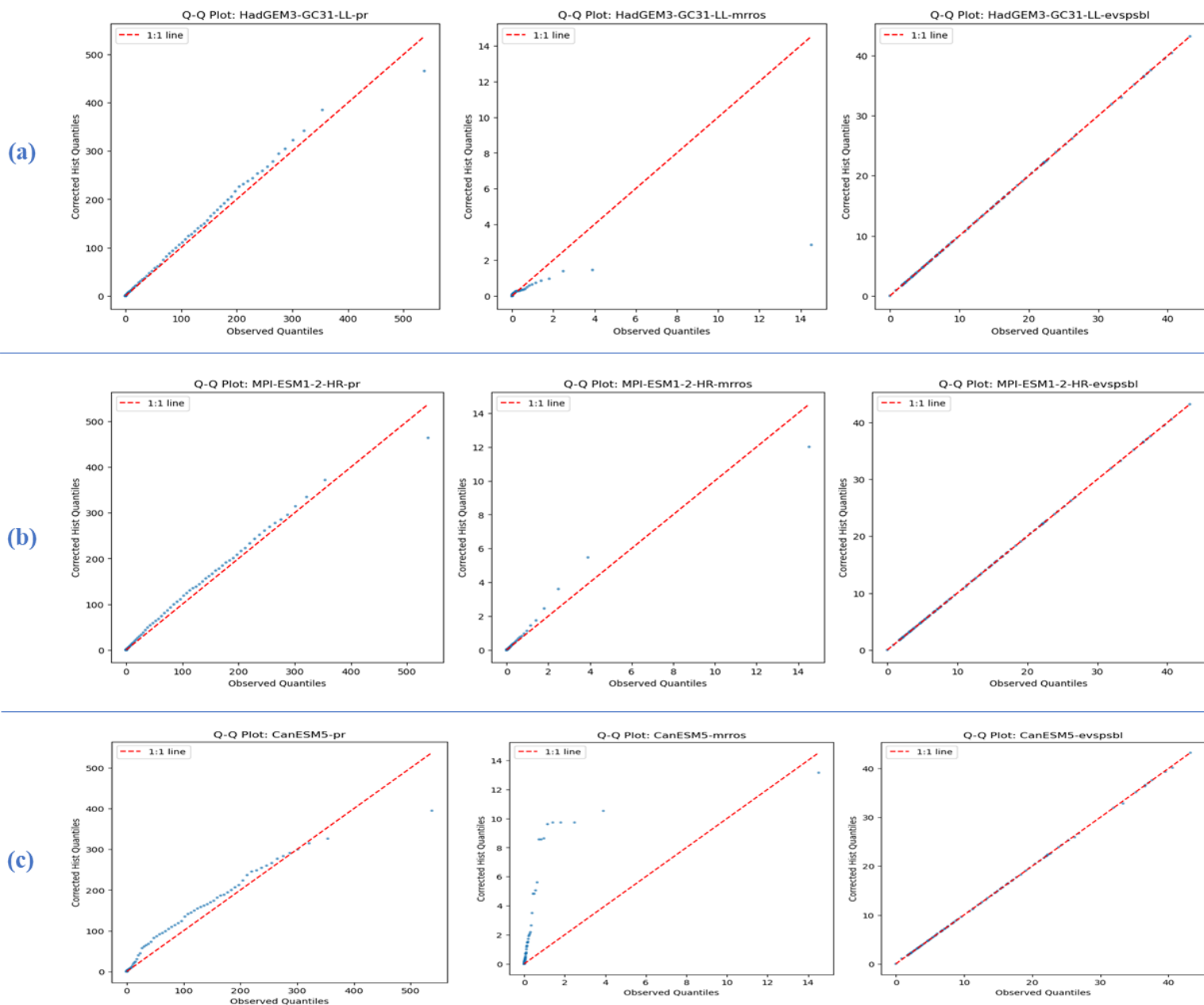


Figure 2: Quantile-quantile (Q-Q) plots comparing observed and QDM-corrected precipitation (mm month^{-1}), surface runoff (mm month^{-1}), and evapotranspiration (mm month^{-1}) for the (a) HadGEM3-GC31-LL, (b) MPI-ESM1-2-HR, and (c) CanESM5 models.



328 **3.2 GAN Downscaling results**

329 **GAN Training and Evaluation Summary**

330 Figure 3 show the training loss curves of the generator and discriminator. Over 200 epochs, the generator’s loss steadily
331 declined from ~88.7 to ~21.6, while discriminator loss remained consistently low (<0.08), indicating stable adversarial
332 dynamics and the absence of mode collapse. Validation loss improved from 0.864 to 0.172, confirming good
333 generalization. These trends suggest effective convergence and balanced adversarial training.

334

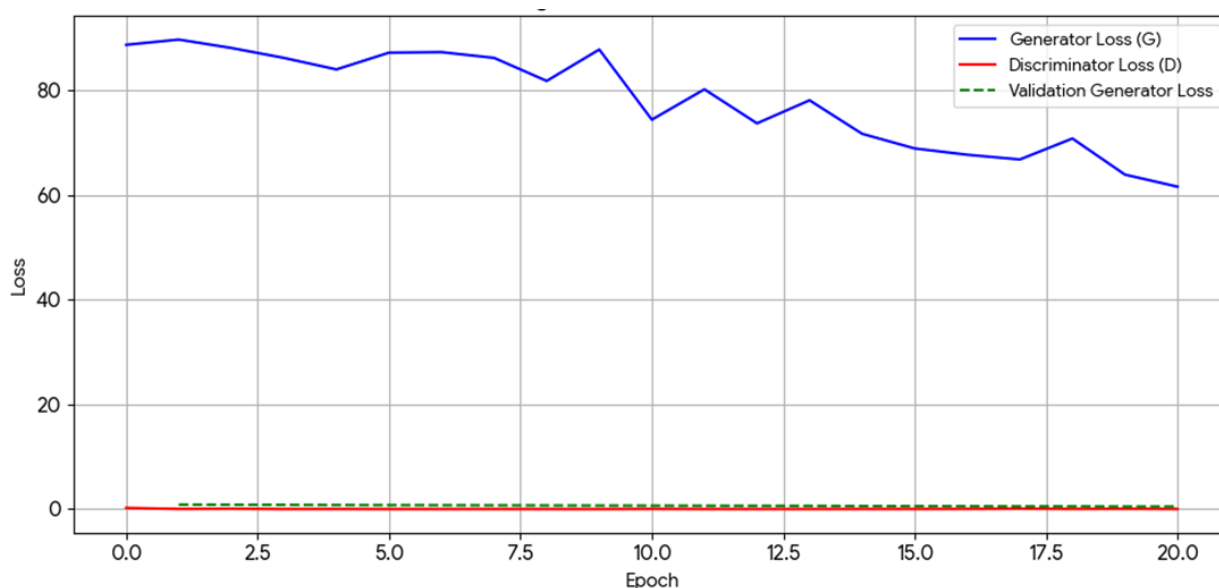


Figure 3 Training and Validation Loss Curve

335

336 Final validation on 180 LR-HR sample pairs yielded strong results: MSE = 0.028, PSNR = 16.6 dB, and SSIM = 0.888.

337 The relatively high SSIM indicates strong structural preservation, while PSNR values, although moderate, are typical of
338 perceptual super-resolution methods that prioritize realism over strict pixel fidelity.

339 As shown in Table 6, the GAN clearly outperformed interpolation baselines, reducing RMSE by ~22% relative to bicubic
340 (9.1 to 7.0 mm/day) and improving SSIM from 0.65 to 0.89.

341

342



343 **Table 6: Summary of validation metrics for bilinear, bicubic, and GAN-based downscaling.**

Method	RMSE (mm/day)	SSIM
Bilinear	9.4	0.62
Bicubic	9.1	0.65
GAN (proposed)	7.0	0.89

344

345 Although PSNR remained moderate (~16.6 dB), this outcome is expected when models optimize perceptual or
 346 adversarial objectives. Theory and experiments describe a perception versus distortion trade-off: methods that favor
 347 spatial realism often score lower on pixelwise distortion metrics such as MSE and PSNR, yet they produce outputs that
 348 and are more physically plausible spatial structure (Blau and Michaeli, 2018; Johnson et al., 2016).

349 The high SSIM (approximately 0.89) supports this interpretation because SSIM emphasizes structural similarity rather
 350 than exact pixel matching. In geophysical images, small phase shifts of narrow rain bands or orographic edges can inflate
 351 MSE and depress PSNR without degrading structural fidelity, which SSIM captures more faithfully (Wang et al., 2004).
 352 Preservation of structural organization is particularly important for hydroclimatic applications because runoff
 353 generation, infiltration, and hydrological connectivity are highly sensitive to the spatial coherence of precipitation fields
 354 rather than pixelwise agreement alone.

355 Adversarial and perceptual super resolution methods have also been shown to recover coherent small-scale features in
 356 Earth system data that interpolation misses, including topography aligned gradients and mesoscale organization. Studies
 357 report sharper, more realistic convective structures than baseline interpolation, consistent with our qualitative gains
 358 (Oyama et al., 2023; Reddy et al., 2023; Stengel et al., 2020). These results justify the claim that the GAN prioritized
 359 perceptual and structural realism and successfully reconstructed orographic gradients and convective patterns despite
 360 only moderate PSNR.

361 Interpolation baselines were included because they represent the operational standard for spatial resampling in many
 362 climate-impact workflows. Improvements beyond interpolation therefore indicate genuine recovery of unresolved
 363 spatial variability rather than simple grid refinement.

364 **Visual and Structural Analysis**

365 Figure 4 compares downscaled precipitation fields produced by bilinear, bicubic, and GAN methods against the 1 km
 366 reference. Interpolation approaches produced over-smoothed outputs, with blurred rainbands and diffuse convective
 367 patterns. These artifacts suppressed extremes and degraded spatial coherence, yielding low SSIM and PSNR (0.28 and
 368 3.5 dB in selected cases).



369 In contrast, the GAN restored sharper gradients, convective cells, and textural variability. Minor textural artifacts, typical
370 of GAN outputs, were observed but did not substantially degrade structural coherence. These results highlight the
371 limitation of PSNR as a sole evaluation metric, as it penalizes perceptual detail. The GAN's higher SSIM confirms its
372 superiority in reproducing spatial coherence and variability.

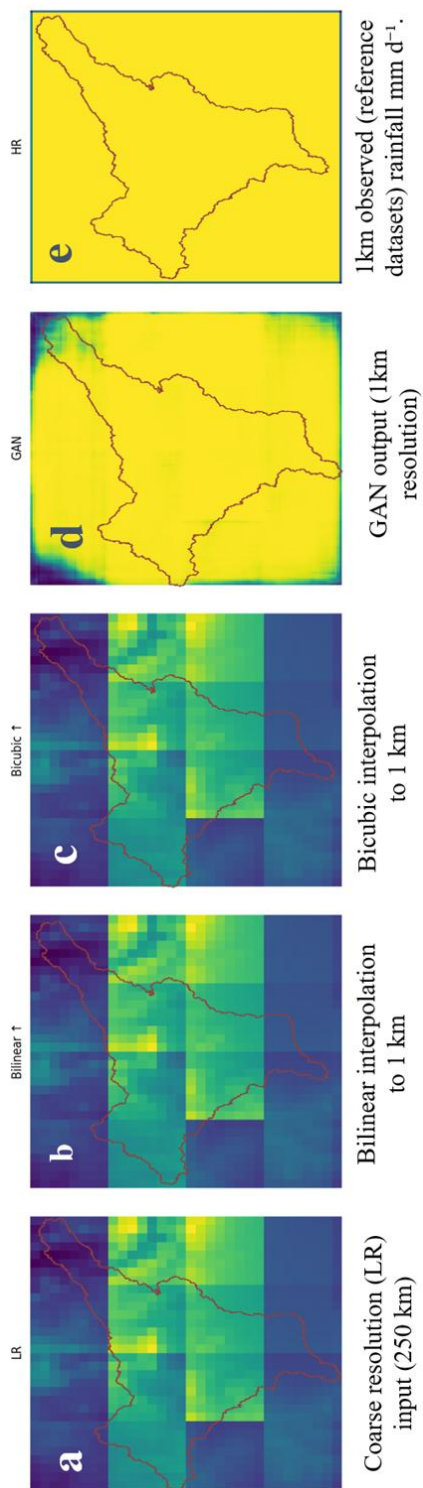


Figure 4. Visual comparison over selected validation patches: (a) input low-resolution GRACE field; (b) bilinear interpolation; (c) bicubic interpolation; (d) GAN-super-resolved (SR) output; and (e) high-resolution (HR) reference precipitation (mm month^{-1}).



374 **Validation Against Observed Data**

375 Further validation against reference datasets reinforced these results. As summarized in table 7, the GAN consistently
376 outperformed bicubic interpolation across metrics, achieving a ~1.4 dB gain in PSNR and a substantial improvement in
377 SSIM (0.62 to 0.888).

378 Table 7: *Comparative validation metrics for bicubic interpolation and GAN-based downscaling.*

Method	MSE	PSNR (dB)	SSIM
Bicubic interpolation	0.037	15.2	0.62
GAN (proposed)	0.028	16.6	0.888

379

380 The slight reduction in validation PSNR relative to training indicates minimal overfitting, while the high SSIM
381 demonstrates consistent structural preservation.

382 **Limitations and Future Improvements**

383 Two main limitations were observed. First, the moderate PSNR reflects a common tradeoff in adversarial super-
384 resolution, where some pixel-level accuracy is exchanged for sharper and more realistic textures. Second, the model was
385 trained and applied on single images, so temporal consistency was not enforced, even though the source GCM fields are
386 spatiotemporal. As a result, consecutive maps can show small frame-to-frame jitter and less coherent evolution of storms.
387 A third limitation concerns the absence of a formal component-wise ablation analysis. While the present study evaluates
388 the integrated QDM-GAN framework against interpolation baselines, the isolated contributions of QDM-only and GAN-
389 only configurations were not explicitly quantified. Although the complementary roles of bias correction and spatial
390 super-resolution are well established in the literature, controlled experiments separating their marginal effects would
391 provide a more rigorous assessment of additive benefits. This aspect is therefore identified as a priority for future
392 investigation.

393 Future improvements should address these points directly. Spatiotemporal architectures should be adopted to learn
394 dynamics and maintain consistency across time, and temporal discriminators or coherence losses should be introduced
395 so that sequences evolve smoothly. Component-wise benchmarking and sensitivity analyses should be conducted to
396 better quantify the relative contributions of statistical correction and learned spatial enhancement. Pixel losses should
397 be combined with perceptual or feature-matching losses to reduce artifacts while keeping fine details, and the weights
398 on L1 and adversarial terms should be adjusted during training to balance accuracy and realism. Physics-aware
399 constraints and simple mass-balance checks should be included to preserve physical plausibility, and probabilistic
400 outputs should be considered to quantify uncertainty.



401 **Integrated Discussion**

402 The integrated QDM and GAN framework highlights the complementary strengths of statistical bias correction and
403 machine learning based spatial enhancement. QDM corrected climatological biases, and the GAN restored high
404 frequency spatial features that interpolation methods could not reproduce. In quantitative terms, monthly
405 evapotranspiration bias decreased from about +45 mm to approximately +1.7 mm after QDM, and the GAN yielded a
406 22% reduction in precipitation RMSE together with a clear increase in SSIM relative to bicubic interpolation.

407 These outcomes are consistent with studies that combine trend-preserving bias adjustment with deep learning for super-
408 resolution downscaling, which report concurrent gains in mean-state accuracy, distributional fidelity, and spatial
409 structure (Gergel et al., 2024; Oyama et al., 2023). Other work (Besombes et al., 2021; Rampal et al., 2024) also notes
410 that GAN based downscaling can underrepresent extremes unless the training objective and conditioning signals are
411 tuned to emphasize tails, which aligns with the mixed performance observed for extremes in this study.

412 Residual biases in precipitation and runoff, particularly for CanESM5, therefore remain a challenge for the reliable
413 simulation of extremes, consistent with documented CMIP6 biases in precipitation and extremes as well as known
414 CanESM5 regional biases (Abdelmoaty et al., 2021; Li et al., 2022; Swart et al., 2019; Teng et al., 2015).

415 Even with these limitations, the hybrid approach constitutes a substantial improvement over conventional interpolation
416 and raw GCM outputs. The modular design facilitates extension to additional variables and regions and supports
417 uncertainty aware hydrological impact assessments, which makes the method suitable for climate adaptation and water
418 resources planning in data scarce environments.

419 The framework was evaluated across three structurally different CMIP6 GCMs to reduce dependence on any single
420 model configuration and to assess methodological robustness across varying climatological biases. Although the
421 framework improves deterministic reconstruction skill, uncertainty associated with GCM structural differences and
422 adversarial downscaling remains non-negligible and should be propagated in downstream hydrological applications.

423 **4. Conclusion**

424 This study developed and evaluated a two-stage post-processing framework that integrates Quantile Delta Mapping
425 (QDM) for non-stationary bias correction with a conditional Generative Adversarial Network (GAN) for statistical
426 super-resolution of climate model outputs. The objective was to provide a reproducible and modular workflow capable
427 of improving both distributional fidelity and spatial structure in coarse-resolution GCM simulations.

428 Application of QDM substantially reduced systematic biases in precipitation, runoff, and evapotranspiration, improving
429 mean-state alignment and distributional consistency relative to observational datasets. These improvements were
430 confirmed using conventional error metrics as well as Kolmogorov-Smirnov and Wasserstein distance diagnostics.

431 Subsequent application of the GAN enhanced spatial realism by reconstructing finer-scale gradients and localized



432 variability that were not recoverable through interpolation alone. Compared with interpolation baselines, the integrated
433 framework achieved lower RMSE and higher structural similarity for precipitation fields.

434 The results demonstrate that statistical bias correction and learned super-resolution address complementary dimensions
435 of model error: QDM improves climatological consistency, while the GAN enhances spatial heterogeneity. When
436 applied sequentially, these components provide improvements in both statistical and structural performance relative to
437 raw GCM outputs and interpolation-based approaches.

438 The study emphasizes methodological transparency and reproducibility. All processing steps, hyperparameters, and
439 evaluation procedures are documented, and the complete implementation is publicly available to support replication and
440 extension. The framework is designed to be modular and adaptable, and additionally provides a transferable basis for
441 assessing the statistical fidelity and spatial realism of machine learning-based climate downscaling approaches across
442 different variables and regions, although further evaluation, including controlled ablation analyses, sensitivity testing,
443 and incorporation of physical consistency constraints would strengthen future applications. The study therefore
444 contributes not only a post-processing workflow, but also an assessment-oriented framework for evaluating statistical
445 and structural consistency in climate-model refinement.

446 Overall, the proposed QDM-GAN workflow provides a reproducible post-processing tool for generating bias-corrected,
447 high-resolution climate model fields suitable for downstream modeling tasks that require improved statistical
448 consistency and spatial detail.

449 **Code and data availability**

450 The code used in this study is openly available via Zenodo at <https://doi.org/10.5281/zenodo.18746067> (Ibrahim et al.,
451 2026). The repository contains the full reproducible workflow implemented as documented Jupyter notebooks. Input
452 datasets are publicly available and are not redistributed.

453 **Author contributions**

454 Abba Ibrahim: Conceptualization and Writing-original draft. Aimrun Wayayok: Conceptualization, review, and editing.
455 Helmi Zulhaidi Mohd Shafri: Conceptualization, supervision. Noorellimia Mat Toridi: Conceptualization and
456 supervision. Idris Wada Muhammad: Data collection and preprocessing.

457 **Competing interests**

458 The authors declare that they have no competing financial interests or personal relationships that could have appeared
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474

475

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