



1 **Nocturnal production of N₂O₅ and ClNO₂ in Delhi: driving factors and impacts**

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30 **ABSTRACT.**

31 Nitryl chloride (ClNO_2) is an important Cl^\bullet precursor, originating from the heterogeneous reactions of
32 dinitrogen pentoxide (N_2O_5) on chloride-containing particles. This N_2O_5 - ClNO_2 chemical process
33 plays critical roles in chloride activation, nitrate formation, and thus air pollution. Here we present
34 field measurements made in the early springtime of 2023 in Delhi and compare with a previous study
35 conducted during the winter of 2019. We found elevated nocturnal levels of N_2O_5 and ClNO_2 ,
36 averaging 13 and 80 ppt, respectively, which are approximately doubled compared to observations in
37 2019. This change is primarily driven by the reduced nighttime NO levels, from 124 ± 25 ppb in 2019
38 to 44 ± 9 ppb in 2023. In addition, the chloride concentration (nighttime average $4.7 \mu\text{g}/\text{m}^3$) in Delhi
39 is among the highest reported globally, driving efficient conversion of N_2O_5 to ClNO_2 . Decreased NO
40 and elevated ClNO_2 levels lead to higher NO_3^\bullet and Cl^\bullet production that promote the oxidation of
41 organics. Consistently, we observed increased fractions of gaseous nitrogen- and chlorine-containing
42 organic products and a higher oxidation state of the organic aerosols. Our findings highlight the need
43 for increased attention to atmospheric secondary pollution and stringent chlorine emissions control
44 with the reduction of NO_x in Delhi.

45



events under different NO regimes remains unclear due to limited field observations in Delhi.

In this work, we present a recent field measurement of N_2O_5 , ClNO_2 and other relevant inorganic and organic species in both gas and particle-phases during the early springtime of 2023 in Delhi. This timing captures a broader range of NO concentrations (below 0.1 to hundreds of ppb) and chemical conditions, which overlaps with but extends beyond previous studies that primarily focused on the most polluted winter months. We characterize the changes in N_2O_5 - ClNO_2 chemistry relative to the previous observation in winter 2019 (Haslett et al., 2023) and examine the underlying driving factors and impacts. The field derived kinetic parameter, quantified as the product of N_2O_5 uptake coefficient ($\gamma_{\text{N}_2\text{O}_5}$) and ClNO_2 yield (f_{ClNO_2}), is also estimated and compared to those reported from other atmospheric environments. We find enhanced N_2O_5 production driven by decreased NO concentrations, and the formation of ClNO_2 is dependent on the chloride level in Delhi. The product of $\gamma_{\text{N}_2\text{O}_5} \times f_{\text{ClNO}_2}$ (average ~ 0.2) is at the upper end of the values observed around the world, indicating efficient N_2O_5 conversion to ClNO_2 in Delhi. We further investigate the impacts of associated increases in $\text{NO}_3\cdot$ and $\text{Cl}\cdot$ production rates on the atmospheric oxidation of organics, utilizing concurrent measurements of a wide range of molecular-level organic oxidation products.

2 METHODS

2.1 Field observation overview

The field campaign was performed at the Indian Institute of Technology Delhi campus (IIT Delhi, 28.54°N, 77.19°E; ~ 230 m above the sea level) from 15 February to 14 March in 2023. The valid dataset collected during the early spring period (2/23-3/14) was used for analysis in this study. IIT Delhi is a representative continental urban site situated at the north-western of Indo-Gangetic Plain, which is frequently affected by intense anthropogenic emissions from biomass and waste burning activities, traffic exhaust, power plants, and industries (Kumar et al., 2022; Bryant et al., 2023; Kashyap et al., 2019; Rai et al., 2020), and potential biogenic emissions from the surrounding vegetations (Fig. S4). All the instruments were installed in an air-conditioned laboratory on the 8th floor, the roof level of the building, at a sampling height of approximately 30 m above the ground. Details of the measurement design are found in the supplementary information (Text. S1).

FIGAERO-I-CIMS. N_2O_5 , reactive chlorines, and oxygenated organics were identified as iodide clusters ($\text{M}\cdot\text{I}$) by a Filter Inlet for Gases and AEROSols Iodide-Chemical Ionization-Time of Flight Mass Spectrometer (FIGAERO-I-HR-ToF-CIMS, hereafter as CIMS) (Lopez-Hilfiker et al., 2014). Basic principles and configurations were similar to those described in previous field studies (Le Breton et al., 2018a; Le Breton et al., 2018b). Each CIMS operation cycle consisted of background (flushing inlet with ultrahigh purity N_2) (~ 2 min), ambient gas (~ 16 min), a second background (~ 2 min), and $\text{PM}_{2.5}$ desorption measurements (~ 50 min) (Fig. S5). The raw signals in counts per second (cps) were normalized by the sum of the reagent ions ($\text{I}^- + \text{I}\cdot\text{H}_2\text{O}^-$) and multiplied by 10^6 . During the field



campaign, ~0.3 lpm perfluoropentanoic acid in N₂ was injected into the CIMS (identified as C₅HF₉O₂I⁻) every 1~2 days for mass calibration. The sensitivities of acetic acid, which was generated from a permeation source (permeation rate of 184.5 ng/min at 40°C), was measured periodically on 2/27, 2/28, 3/5, 3/10, 3/12, and 3/14 to track the instrument stability, which was 0.27 ± 0.05 cps/ppt with a variation of 17.7%.

Calibrations of gaseous species were conducted after the campaign under RH conditions (IH₂O⁻ to I⁻ ratio 0.42 ± 0.02) similar to those observed during the field measurement in Delhi (IH₂O⁻ to I⁻ ratio 0.40 ± 0.08), with sensitivities of 7.5, 3.6, 0.013, 13.6, and 0.29 cps/ppt for N₂O₅, ClNO₂, HCl, chloroacetic acid, and acetic acid, respectively. The sensitivity of acetic acid measured post-campaign (0.29 cps/ppt) agreed well with the on-site calibration result (0.27 cps/ppt), and therefore no additional scaling was applied for the on-site sensitivities of the detected species. Sensitivities for Cl₂ (3.4 cps/ppt), NCl₃ (3.1 cps/ppt), and NHCl₂ (0.05 cps/ppt) were determined from our previous study (Chen et al., 2025b). The detection limits were determined as three times the standard deviation of the hourly averaged background signals, which were 0.6 ppt for N₂O₅, 1.6 ppt for ClNO₂, 0.6 ppt for Cl₂, 22.6 ppt for NHCl₂, 0.3 ppt for NCl₃, 209.1 ppt for HCl, and 2.5 ppt for chloroacetic acid. For the remaining detected gaseous compounds, the maximum value of the calibrated sensitivities (chloroacetic acid, 13.6 cps/ppt) was used to derive conservative concentration estimates. For particulate compounds, the calibrated sensitivity of levoglucosan (2.2×10^5 ions/ng) was applied to obtain lower-limit estimates, as the detection of levoglucosan has recently been demonstrated to proceed at the collision limit (Aggarwal et al., 2025).

EDXRF. Energy Dispersive X-ray Fluorescence (EDXRF; SPECTRO-XEPOS, AMETEK Inc.) was used to determine the concentrations of K and Cl in PM_{2.5} samples. ED-XRF is a sensitive and non-destructive technique which requires negligible sample pre-treatment and has been widely used for measuring elements in ambient particles (Ogrizek et al., 2022; Unga et al., 2025). A total of 83 PM_{2.5} samples were collected at 3 lpm on PTFE filters (25 mm in diameter) via a PM_{2.5} cyclone every 12 hours from 2/22 to 2/28 and every 4 hours from 3/1 to 3/14. We increased the frequency of filter switches by the end of the campaign to capture the variations of Cl element concentrations within a night, i.e., before (19:00~23:00) and post (23:00~3:00) midnight. The mid-point of the measurement period, i.e., 1:00 indicates measurement from 23:00~3:00, 5:00 for 3:00~7:00, 9:00 for 7:00~11:00, 13:00 for 11:00~15:00, 17:00 for 15:00~19:00, and 21:00 for 19:00~23:00, was used as the representative timestamp for each filter sample.

The EDXRF instrument detects the total concentrations of elements in the collected filter samples, while the CIMS detects compounds that evaporate at temperatures up to 200 °C. Persistent signals of HCl·I⁻ were observed from PM_{2.5} desorption during the CIMS measurement (**Fig. S5**), likely originating from the thermal decomposition of non-refractory chloride (e.g. NH₄Cl). The trend of chloride measured by CIMS (pCl⁻-CIMS) largely followed that of the XRF-measured chloride (Cl-



XRF) ($r = 0.70$) (**Fig. S6c**). However, the campaign-averaged pCl-CIMS ($0.02 \mu\text{g}/\text{m}^3$) was approximately two orders of magnitude lower than Cl-XRF ($2.3 \mu\text{g}/\text{m}^3$), likely due to the use of levoglucosan sensitivity for quantification and negligible decomposition and desorption of alkali chlorides (e.g., NaCl and KCl) in CIMS measurements. Nevertheless, pCl-CIMS was used qualitatively to complement the temporal trend of Cl-XRF to hourly time-resolution.

Supporting measurements at IIT Delhi include trace gases (NO, NO₂, O₃, SO₂, and CO) and meteorological parameters (T, RH, wind speed, and wind direction). Hourly PM_{2.5} concentrations and solar radiation data were obtained from the nearest national monitoring station, the R.K. Puram station operated by Delhi Pollution Control Committee (DPCC), which is located approximately 3 km northwest of IIT Delhi. Intercomparison of the trace gases, T, and RH measured at IIT Delhi and R.K. Puram showed overall good agreement (**Fig. S6a**); therefore, data from R.K. Puram were used to supplement the analysis when measurements were not available at IIT Delhi. The aerosol surface area (Sa) was estimated from the linear correlation between the measured Sa at IIT Delhi and PM_{2.5} in a previous study (Gani et al., 2020) (**Fig. S6b**, $r = 0.78$). The solar radiation was converted to photolysis frequency of J_{NO_2} utilizing the method from Trebs et al (Trebs et al., 2009). The estimated J_{NO_2} was then used to derive real-time photolysis frequency of NO₃[•], O₃ and reactive chlorines (Chen et al., 2025b). Additional details about the field observations are shown in **Text S1**.

Comparison with the measurements in 2019 (Haslett et al., 2023) was used to investigate the evolution of N₂O₅-ClNO₂ chemistry under different chemical regimes in Delhi. It is noted that throughout the manuscript, “in 2023” and “in 2019” refer to the corresponding field campaigns conducted from 23 February to 14 March in 2023 (this study) and from 11 January to 5 February in 2019 (the previous campaign), respectively.

2.2 Kinetic calculations

N₂O₅-ClNO₂ chemistry. Key kinetic parameters were derived to investigate factors controlling the variations of N₂O₅ and ClNO₂ and compare with previous studies. The field-constrained uptake parameter, defined as the product ($\gamma_{\text{N}_2\text{O}_5} \times f_{\text{ClNO}_2}$) of the N₂O₅ uptake coefficient ($\gamma_{\text{N}_2\text{O}_5}$) and ClNO₂ yield (f_{ClNO_2}), as is shown in Eq. 1, was calculated hourly using neighboring measurements during the nighttime following Eq. 3b in Mielke et al. (Mielke et al., 2013) (Eq. 1), assuming the local heterogeneous uptake of N₂O₅ was the sole source of the measured ClNO₂ and that ClNO₂ experienced negligible nocturnal losses.

$$[\gamma_{\text{N}_2\text{O}_5} \times f_{\text{ClNO}_2}]_{t_0} = \frac{[\text{ClNO}_2]_{t_1}}{\int_{t_0}^{t_1} \frac{1}{4} \bar{c} S_a [\text{N}_2\text{O}_5]_{t_0} dt} \quad (\text{Eq. 1})$$

where \bar{c} is the average molecular velocity of N₂O₅ ($\sqrt{\frac{8kT}{\pi M_w}}$, k is Boltzmann’s constant, 1.38×10^{-23} J/K).



181 $^{23} \text{ m}^2 \text{ kg s}^{-2} \text{ K}^{-1}$, T is temperature in Kelvin, M_w is the molecular weight of N_2O_5). The theoretical
 182 values of the uptake parameter ($\gamma_{\text{N}_2\text{O}_5} \times f_{\text{ClNO}_2}$) range from 0 to 0.1, as the maximum values of $\gamma_{\text{N}_2\text{O}_5}$
 183 and f_{ClNO_2} are generally below 0.1 and 1, respectively (Tham et al., 2018). However, nocturnal
 184 transport, i.e., changes in air masses, can invalid this method and yield either negative (corresponding
 185 to decreases in ClNO_2 , occurrence frequency of $\sim 38\%$) or unusually high values exceeding 0.1 (sharp
 186 increases in ClNO_2 , occurrence frequency of $\sim 17\%$). The negative values were not considered for
 187 further analysis and the extremely high values were discussed separately (**Section 3.2, Fig. 4**).

188 The production rate ($P(\text{NO}_3)$) and reactivity ($R(\text{NO}_3)$) of NO_3 were calculated to evaluate N_2O_5
 189 budgets, which are shown in Eq. 2 and Eq. 3, respectively. The ClNO_2 production efficiency ($\varepsilon(\text{ClNO}_2)$,
 190 Eq. 4), indicating the fraction of the total generated NO_3 overnight that is ultimately transformed into
 191 ClNO_2 , was calculated for comparison with previous studies (Eger et al., 2019; Xia et al., 2025). The
 192 integration period was set from 20:00 to 04:00 to minimize influences of potential air mass mixing
 193 during the day-night and night-day transition periods.

194
$$P(\text{NO}_3) = k_{\text{NO}_2+\text{O}_3}[\text{NO}_2][\text{O}_3] \quad (\text{Eq. 2})$$

195
$$R(\text{NO}_3) = k_{\text{NO}+\text{NO}_3}[\text{NO}] + J_{\text{NO}_3} + k_{\text{hete}}[\text{NO}_2]K_{\text{eq}} + \sum k_i[\text{VOC}_i] \quad (\text{Eq. 3})$$

196
$$\varepsilon(\text{ClNO}_2) = \frac{[\text{ClNO}_2]_{\text{max}}}{\int_{20:00}^{04:00} P(\text{NO}_3)} \quad (\text{Eq. 4})$$

197 where $k_{\text{NO}_2+\text{O}_3} = 1.2 \times 10^{-13} \times \exp(-2450/T) \text{ cm}^3/\text{molecule/s}$ and $k_{\text{NO}+\text{NO}_3} = 1.8 \times$
 198 $10^{-11} \times \exp(110/T) \text{ cm}^3/\text{molecule/s}$ (Burkholder et al., 2020). J_{NO_3} is the photolysis frequency of
 199 NO_3 and is retrieved from the Tropospheric Ultraviolet and Visible (TUV) radiation model and scaled
 200 with the field constrained J_{NO_2} as discussed previously, k_{hete} is the first-order heterogeneous loss
 201 rate coefficient of N_2O_5 , which is calculated as $\frac{\bar{c}\gamma_{\text{N}_2\text{O}_5}}{4} S_a$. Here $\gamma_{\text{N}_2\text{O}_5}$ was parametrized as a function
 202 ($\gamma_{\text{N}_2\text{O}_5} = RH \times 5.2 \times 10^{-4}$) of ambient RH up to 57% RH and then being a constant (0.03) according
 203 to the experimental measurements of N_2O_5 uptake on aqueous organic aerosols (Evans and Jacob,
 204 2005; Thornton et al., 2003). This parameterization method was supported by the consistency between
 205 the simulated nocturnal ClNO_2 variations using the RH-parameterized $\gamma_{\text{N}_2\text{O}_5}$ and the observed ClNO_2
 206 levels, as detailed in **Text S2**. K_{eq} is the equilibrium constant between NO_3 , NO_2 , and N_2O_5 , where
 207 $K_{\text{eq}} = 5.5 \times 10^{-27} \times \exp(10724/T)$ (Wängberg et al., 1997). $\sum k_i[\text{VOC}_i]$ is the sum of
 208 NO_3 reactivity with different VOCs and we applied the same value of 0.081 s^{-1} in 2023 as in
 209 2019 (Haslett et al., 2023) due to the lack of concurrent VOCs measurements in 2023. This assumption
 210 is not expected to significantly affect the estimated total NO_3 reactivity in 2023, which was still
 211 dominated by the reactivity with NO (15.6 s^{-1} , **Fig. S7a**).

212 **Production rates of $\text{Cl}\cdot$ and $\text{OH}\cdot$ radicals.** The $\text{Cl}\cdot$ production rate ($P(\text{Cl}\cdot)$) is calculated according



to Eq. 5, where J is the photolysis frequency of each photolabile chlorines. For the 2023 campaign, we calculated $P(\text{Cl}\cdot)$ both from ClNO_2 photolysis and from all the detected photolabile reactive chlorines. A direct comparison between 2019 and 2023 was conducted using $P(\text{Cl}\cdot)$ from ClNO_2 photolysis. In 2019, only ClNO_2 photolysis was considered as a source of $\text{Cl}\cdot$, since mixing ratios of other $\text{Cl}\cdot$ precursors were not reported in the previous study (Haslett et al., 2023). Additionally, we estimated a lower-limit $\text{OH}\cdot$ production rate ($P(\text{OH}\cdot)$) in Delhi from the photolysis of O_3 in the presence of water vapor (Eq. 6) (Dunlea and Ravishankara, 2004).

$$P(\text{Cl}) = J_{\text{ClNO}_2}[\text{ClNO}_2] + 2 \times J_{\text{Cl}_2}[\text{Cl}_2] + J_{\text{NCl}_3}[\text{NCl}_3] + J_{\text{NHCl}_2}[\text{NHCl}_2] \quad (\text{Eq. 5})$$

$$P(\text{OH}) = \frac{2 \times J_{\text{O}^1\text{D}}[\text{O}_3] + k_{\text{H}_2\text{O}}[\text{H}_2\text{O}]}{k_{\text{H}_2\text{O}}[\text{H}_2\text{O}] + k_{\text{N}_2}[\text{N}_2] + k_{\text{O}_2}[\text{O}_2]} \quad (\text{Eq. 6})$$

where $J_{\text{O}^1\text{D}}$ is the photolysis frequency of O_3 , $k_{\text{H}_2\text{O}} = 1.6 \times 10^{-10} \times \exp(60/T)$, $k_{\text{N}_2} = 2.2 \times 10^{-11} \times \exp(110/T)$, and $k_{\text{O}_2} = 3.3 \times 10^{-11} \times \exp(55/T)$ (Burkholder et al., 2020).

3 RESULTS AND DISCUSSION

3.1 Characteristics of N_2O_5 and ClNO_2 variations

We observed frequent nocturnal enrichment of N_2O_5 and ClNO_2 in 2023 (Figs. 1a-b). Appreciable N_2O_5 and ClNO_2 up to 566 and 1340 ppt were measured during the nighttime, with mixing ratios commonly exceeding 10 and 100 ppt, respectively (Fig. S8). As shown in Table S1, the maximum N_2O_5 mixing ratio in Delhi is a third to half of those observed in urban areas of the United States (Thornton et al., 2010; Osthoff et al., 2008) and England (Bannan et al., 2015), and approximately an order of magnitude lower than those measured in aged air plumes from industrial regions in China (Chen et al., 2023; Wang et al., 2016; Ye et al., 2021). By contrast, the peak ClNO_2 levels in Delhi are much higher than those reported in urban regions of the North America (Thornton et al., 2010; McNamara et al., 2020; Wang et al., 2023a) and Europe (Bannan et al., 2015; Phillips et al., 2012; Priestley et al., 2018), and similar to or 1–2 times lower than those observed in the North China Plain (Tham et al., 2016b; Xia et al., 2021; Peng et al., 2021; Chen et al., 2025a), which implies efficient conversion of N_2O_5 to ClNO_2 in Delhi.

The measured N_2O_5 generally presents transient, spike-like patterns in 2023 (Fig. S8), peaking around the early evening (18:00–20:00) when both NO and particulate Cl are relatively low, and tapering off as the night proceeds (Figs. 1a, c-d). In comparison, the ClNO_2 mixing ratio was gradually built up after midnight with the increase of particulate Cl, reaching the highest level during the early morning hours (06:00–08:00, Figs. 1b, d). Overall, the nocturnal N_2O_5 mixing ratio shows a significant negative dependency ($r = -0.64$, $p < 0.01$) on the NO concentration (Fig. 2a), and ClNO_2 exhibits a positive correlation with the particulate Cl level (Fig. S9a), indicating the critical roles of NO and chloride in controlling N_2O_5 and ClNO_2 formation in Delhi.



246 We further examined N_2O_5 and ClNO_2 variations every night and classified them into the
 247 enhanced and non-enhanced cases (**Figs. S10-S11**). Among the enhanced cases, elevated N_2O_5 mixing
 248 ratios exceeding 10 ppt were characterized by both relatively low NO (median 3 ppb) and particulate
 249 Cl (median $0.7 \mu\text{g}/\text{m}^3$) concentrations, and appreciable ClNO_2 levels larger than 100 ppt were observed
 250 under moderate NO (median 7 ppb) and high particulate Cl (median $3.3 \mu\text{g}/\text{m}^3$) conditions (**Fig. S10**).
 251 No notable N_2O_5 or ClNO_2 mixing ratios were measured among the non-enhanced cases when NO
 252 levels were relatively high (median 31 ppb) (**Fig. S11**). The occurrence frequency of the enhanced
 253 cases (11 out of total 19 nights) is higher than the non-enhanced (8 nights) cases.

254 The nocturnal abundances of N_2O_5 and ClNO_2 increase in 2023 compared to 2019. Overall, the
 255 nighttime N_2O_5 and ClNO_2 mixing ratios in 2023, with an average of 13.1 ppt and 80.1 ppt, respectively,
 256 are significantly (Mann-Whitney test, $p < 0.01$) higher than those measured in 2019 (4.5 and 36.3 ppt,
 257 respectively) (**Figs. 1a-b**). The nocturnal mean ClNO_2 production efficiency ($\epsilon(\text{ClNO}_2)$) also increases
 258 from 3.0% in 2019 to 4.3% in 2023, which is comparable to the values reported in marine and coastal
 259 areas (Eger et al., 2019; Xia et al., 2025). In addition, the diurnal profile of N_2O_5 in 2023 was inverted
 260 compared to 2019, with higher levels in the early evening rather than during the daytime (**Fig. 1a**).
 261 Consistently, distinct diurnal variations of N_2O_5 were also observed under relatively low- and high-
 262 NO conditions in 2023 (**Fig. S12**). The emerging early evening peak of N_2O_5 leads to a new N_2O_5 -
 263 driven ClNO_2 enhancement pattern in 2023, characterized by a strong positive correlation ($r = 0.93$)
 264 between N_2O_5 and ClNO_2 under conditions of elevated N_2O_5 (10–400 ppt) and limited particulate Cl
 265 concentrations ($0.1\sim 3.4 \mu\text{g}/\text{m}^3$) (**Figs. 1e-f**). This pattern resembles those observed in other continental
 266 air masses (Thornton et al., 2010; Bannan et al., 2015; Zhou et al., 2018), while the ClNO_2 to N_2O_5 ratios
 267 are slightly higher in Delhi, ranging from 0.3 to 3.2, with an average of 1.2. Notably, despite the
 268 substantial decline in particulate Cl from 2019 ($20.8 \mu\text{g}/\text{m}^3$) to 2023 ($4.7 \mu\text{g}/\text{m}^3$), the higher ClNO_2
 269 levels in 2023 indicate that the promoting effect of elevated N_2O_5 potentially outcompetes the limiting
 270 impact of reduced chloride. Additionally, chloride is likely still sufficient to maintain a near unity
 271 ClNO_2 yield in Delhi (**Text S2.2 and Fig. S13a**).

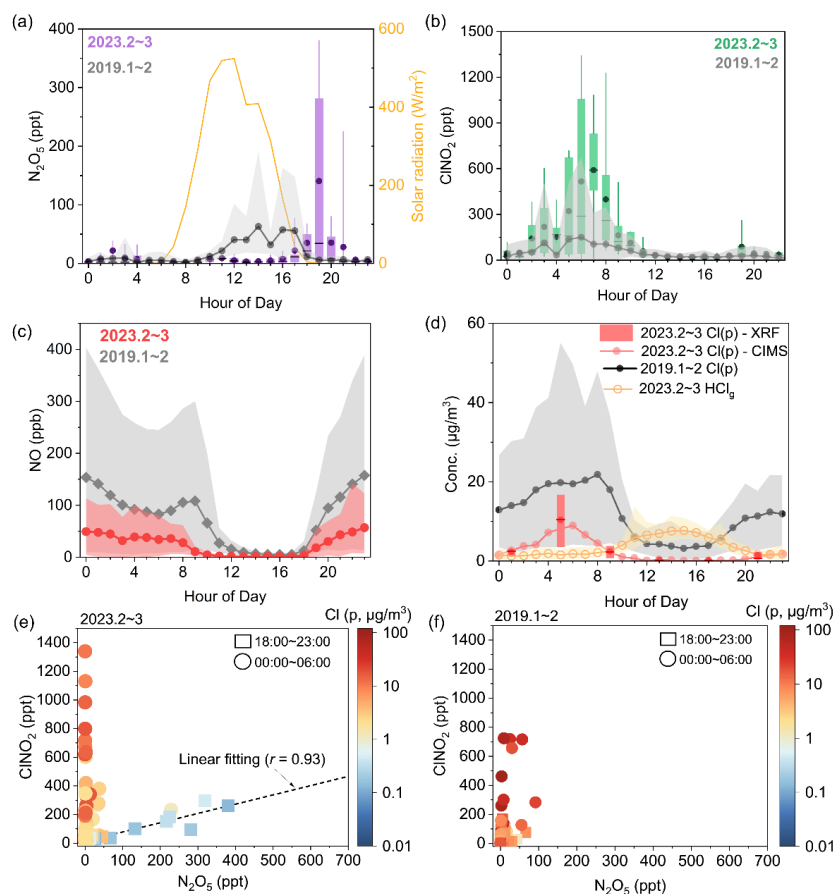


Fig. 1 Field characterization of N_2O_5 and ClNO_2 in Delhi. Campaign-averaged diurnal variations of (a) N_2O_5 , (b) ClNO_2 , (c) NO, and (d) particulate Cl and gaseous HCl. Scatter plots of N_2O_5 versus ClNO_2 mixing ratios in (e) 2023 and (f) 2019, color-coded by the particulate Cl concentrations. The observation (Haslett et al., 2023) in 2019 is included for comparison, where the gaseous HCl measurement is not available. The right axis in (a) indicates solar radiation during the campaign in 2023. The exact sunrise (07:07~07:15) time in 2019 is about 30 min later than that in 2023 (06:55~06:32), while the sunset time in 2019 (17:24~18:03) is about 30 min earlier according to the records from timeanddate website. Gaseous HCl and particulate Cl measured by CIMS are also shown in (c). The upper and lower edge of the box and whisker in (a), (b) and (d) represent the 25th and 75th, and 10th and 90th percentiles, respectively. The dots and squares denote mean values and the lines inside the boxes indicate median levels. The shaded area in (a)-(d) indicates the 10th and 90th percentile range.



285 **3.2 Drivers of enhanced N₂O₅-ClNO₂ chemistry**

286 Lower NO concentrations in 2023 relative to 2019 reduce NO₃[•] losses to reactions with NO,
 287 resulting in elevated nocturnal N₂O₅ levels. The reaction with NO dominated NO₃[•] losses both in 2019
 288 and 2023 (~99%, **Fig. S7a**), contributing to the low abundances and short-lived occurrences of N₂O₅
 289 in Delhi (Haslett et al., 2023). The strong anti-correlation between nocturnal N₂O₅ and NO persists in
 290 2023, where the N₂O₅ mixing ratios overlap with those in 2019 when NO concentrations are within
 291 the same range (approximately 2–100 ppb) (**Fig. 2a**). However, in 2023, the frequency of relatively
 292 low-NO conditions (<2 ppb) rose from 0 to 8% and extreme NO levels (>100 ppb) nearly disappeared,
 293 indicating that NO titration in suppressing N₂O₅ production has diminished (**Fig. 2a**). With the decline
 294 in nighttime NO concentrations from 124 ± 25 ppb in 2019 to 44 ± 9 ppb in 2023 (**Fig. 1c**),
 295 the absolute NO₃[•] reactivity towards NO dropped by ~66% from 83.7 s⁻¹ to 28.4 s⁻¹, extending
 296 nighttime NO₃[•] lifetimes approximately two folds in 2023 (0.04 s). The reduced NO₃[•] losses to NO
 297 increased the availability of NO₃[•] for reacting with NO₂, thereby enhancing N₂O₅ production.
 298 Additionally, no significant correlation between particulate Cl and ClNO₂ was observed in 2019, likely
 299 because the chloride increases were accompanied by a substantial elevation in NO, which suppressed
 300 N₂O₅ formation (**Fig. S9b**).

301 Higher NO₃[•] production rates (P(NO₃)) enhance N₂O₅ formation. Due to the increases in
 302 nocturnal O₃ levels (from 3.5 to 6.8 ppb) and the reaction rate constant between NO₂ and O₃ (from
 303 2.36×10^{-17} to 2.93×10^{-17} cm³ molec⁻¹ s⁻¹), the average nocturnal P(NO₃) in 2023 almost doubled
 304 (0.75 ppb h⁻¹) compared to 2019 (0.41 ppb h⁻¹) (**Fig. 2b**). The observed P(NO₃) in Delhi 2023 exceeds
 305 those reported in the United States (Wang et al., 2023b; Noxon et al., 1980) and Europe (Ljungström and
 306 Hallquist, 1996; Wang et al., 2023b), and 1–2 times lower than those observed in China (Wang et al.,
 307 2024; Yan et al., 2021; Chen et al., 2023). Consistently, the N₂O₅ mixing ratios in 2023 peak after sunset
 308 around 18:00–19:00 (**Fig. 1a**), coinciding with the period of elevated P(NO₃) (**Fig. 2b**) and relatively
 309 low NO levels (**Fig. 1c**). Other factors, such as the decreased aerosol surface area (2637 ± 152 and
 310 1411 ± 162 μm²/cm³ in 2019 and 2023, respectively) results in reduced first-order heterogeneous loss
 311 rate coefficient of N₂O₅ (k_{hete}) from 0.74 ± 0.09 s⁻¹ in 2019 to 0.14 ± 0.02 s⁻¹ in 2023, which also
 312 contributes to the elevated nocturnal N₂O₅ mixing ratios observed in 2023. Overall, the enhanced
 313 NO₃[•] production and reduced losses to NO led to elevated NO₃[•] and N₂O₅ levels in 2023, aligning
 314 with the higher estimated steady-state NO₃[•] mixing ratios (0.25 ppt vs. 0.03 ppt in 2019) (**Fig. S7b**).
 315 It should be emphasized that the 2019 and 2023 measurements were conducted in different seasons,
 316 i.e., January to early February (winter) in 2019 and late February to mid-March (early spring) in 2023.
 317 The observed increase in O₃ and decrease in NO and aerosol surface areas during the 2023 campaign
 318 are influenced by the inter-month variability (e.g., change in boundary layer height and solar radiation)
 319 of these pollutants (**Figs. S14a-b**).

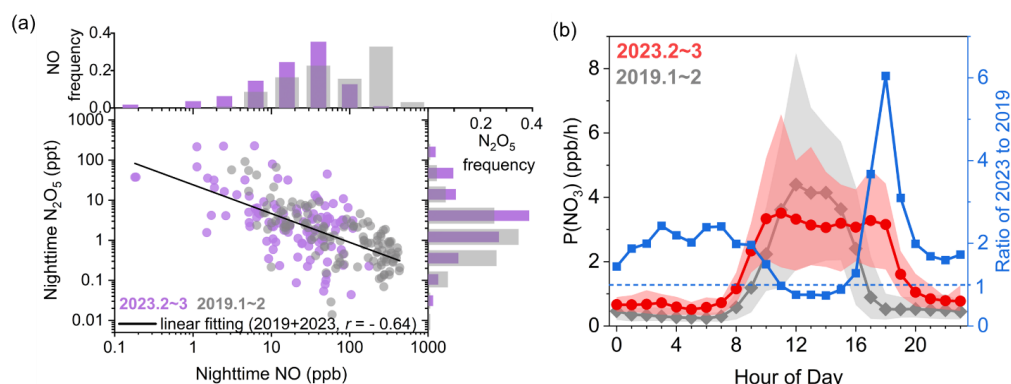


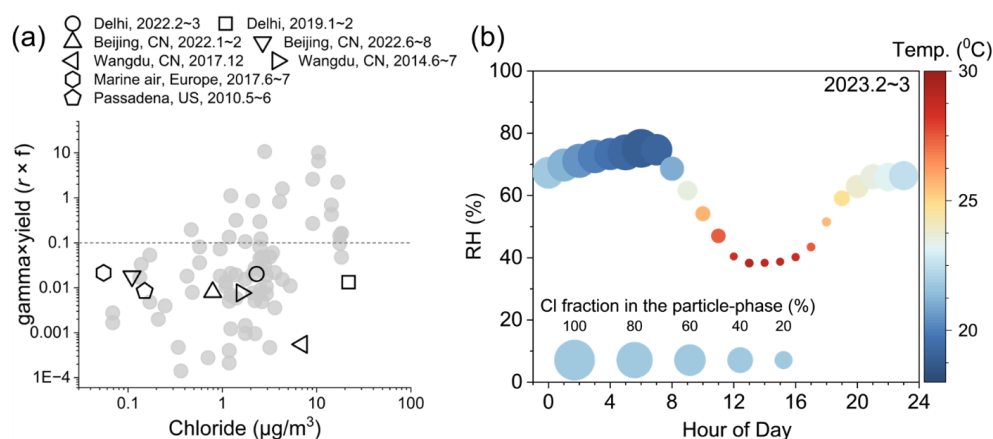
Fig. 2 Drivers of nocturnal N₂O₅ enhancement. (a) Scatter plot of nighttime NO versus N₂O₅. The top and right subfigures in (a) show the occurrence frequencies of NO and N₂O₅ in different concentration bins, respectively. Nighttime is defined as the period from 20:00 to 04:00, excluding the transition period between day and night when the air mass was unstable. (b) Campaign-averaged diurnal variations of P(NO₃) in 2023 and 2019, respectively. The right axis in (b) displays the 2023-to-2019 ratio of P(NO₃). The dots and squares denote mean values and the shaded area indicates the 10th and 90th percentile range.

High particulate Cl concentrations drive significant N₂O₅ uptake and ClNO₂ production in Delhi. The field-derived uptake parameter ($\gamma \times f$) in Delhi is similar with those observed in marine air (~ 0.02 on average)¹¹ and several times to an order of magnitude higher than those reported in other continental regions ($6 \times 10^{-4} \sim 8 \times 10^{-3}$) (Xia et al., 2021; Tham et al., 2018; Xia et al., 2020; Mielke et al., 2013). As is shown in **Fig. 3a**, this parameter increases with the particulate Cl level in Delhi, reflecting the combined effects of chloride in promoting N₂O₅ uptake and ClNO₂ yield observed in prior laboratory studies (Bertram and Thornton, 2009; Jahl et al., 2021). By contrast, despite comparable high chloride was observed in Wangdu (Xia et al., 2021) (upper triangle in **Fig. 3b**), both the uptake parameter ($\sim 6 \times 10^{-4}$ on average) and the nocturnal ClNO₂ level (typically < 50 ppt) are significantly lower than those measured in Delhi. This is likely due to the lower ambient water content ($\sim 0.4\%$) compared to Delhi ($\sim 1.5\%$), which limits aerosol liquid water content and thereby suppressed N₂O₅ hydrolysis on particles (Xia et al., 2021; Tham et al., 2018).

The high chloride in Delhi is largely in the form of semi-volatile NH₄Cl, originating from gas to particle partitioning. We observed a concurrent late-evening enrichment of non-refractory chloride measured by CIMS and Cl element measured by XRF (**Fig. 1d**), indicating the semi-volatile nature of particulate Cl (i.e., NH₄Cl) in Delhi. Additionally, direct measurements of both HCl and particulate Cl enabled us to examine the gas–particle partitioning behavior of Cl. As shown in **Fig. 3b**, the particle-phase Cl fraction amounted to up to 89% on average at 6:00 whereas it remained below 20% in the afternoon (12:00–18:00). This diurnal variation is attributed to lower temperatures (15–25 °C) and



347 higher RH (56-92%) in the early morning hours, which favored the co-condensation of HCl and water
 348 vapor into particles as NH_4Cl under the NH_3 -rich conditions commonly found in Delhi (Acharja et al.,
 349 2023; Gunthe et al., 2021). A recent study suggests that increasing RH promotes the migration of
 350 chloride from the particle bulk phase to the gas-particle interface (Fauré et al., 2024). This phenomenon
 351 helps explain the observed post-midnight accumulation of ClNO_2 and consistently low N_2O_5 levels in
 352 Delhi, which are attributed to enhanced heterogeneous conversion of N_2O_5 to ClNO_2 on aerosol
 353 surfaces. The observed thermodynamically-driven partitioning behavior of Cl agrees with previous
 354 model simulations (Gunthe et al., 2021; Chen et al., 2022b).



355 **Fig. 3 Drivers of nocturnal ClNO_2 enhancement in Delhi.** (a) Dependence of the field-observed
 356 uptake parameter ($\gamma \times f$) on particulate Cl concentrations in 2023. The campaign-average value for
 357 2023 (excluding unusually high values of $\gamma \times f > 0.1$) is shown as an open circle. The values derived
 358 from previous observations in Delhi 2019 (Haslett et al., 2023), Beijing (Chen et al., 2025b), Wangdu
 359 (Xia et al., 2021; Tham et al., 2018), Pasadena (Mielke et al., 2013), and marine air in Europe (Eger et
 360 al., 2019) are shown as different markers. The reference line of $\gamma \times f$ equals to 0.1 is also plotted. (b)
 361 Campaign-averaged diurnal variation of RH in 2023, color-coded by ambient temperature. The size of
 362 the datapoints is proportional to the observed Cl fraction in the particle-phase, which is calculated as
 363 $\text{Cl}_p/(\text{Cl}_p + \text{HCl}_g)$.
 364

365 ClNO_2 levels in Delhi were influenced by the transport of ClNO_2 -laden biomass burning plumes
 366 from upwind regions, and vertical intrusion associated with the breakup of the residual layer. We
 367 observed some cases with unusually high $\gamma \times f$ values exceeding 0.1 (Tham et al., 2018), implying
 368 additional ClNO_2 sources beyond the local N_2O_5 uptake. Most of the high values were found during
 369 episodes of easterly ($50\text{--}130^{\circ}$) winds with elevated particulate Cl concentrations (Fig. 4a), when the
 370 air masses passed over biomass-burning hotspots (Fig. S15). This suggests substantial ClNO_2
 371 production on high chloride containing biomass burning aerosols, as observed in previous laboratory
 372 experiments (Ahern et al., 2018; Jahl et al., 2021). A typical case is shown in Fig. S16, where ClNO_2



373 abruptly increases concurrently with the sharp rise in particulate Cl, biomass-burning tracers (K and
 374 HCN, **Table S3**), and easterly wind speeds. Overall, the nighttime Cl strongly correlates with K ($r =$
 375 0.82) and HCN ($r = 0.73$) (**Fig. 4b-c and Fig. S17**), indicating significant Cl emissions from biomass
 376 burning in Delhi. The ambient nighttime Cl to K ratio (1.6 ± 0.8) in Delhi is about three-fold higher
 377 (0.4 ± 0.2) than those measured from biomass/biofuel emission sources (**Table S3**), indicating
 378 additional sources of Cl potentially from plastic-contained garbage burning, industrial processes, and
 379 coal combustion (Gunthe et al., 2021). We consistently observed a moderate correlation between Cl
 380 and the coal combustion tracer SO_2 ($r = 0.51$). In contrast, the poor correlation of total Cl with NO
 381 ($r = -0.30$) and CO ($r = -0.08$) indicated limited contributions from vehicle exhaust, consistent with
 382 their low Cl emission factors (**Table S3**). Additionally, a continued rise in ClNO_2 was observed after
 383 sunrise (6:00–8:00; **Fig. 1b and Fig. S16**), likely due to the downward mixing of ClNO_2 -rich air
 384 masses aloft, as reported in other locations (Tham et al., 2016a; Haslett et al., 2023). We estimated that
 385 108–805 ppt of ClNO_2 resided in the residual layer (**Text S3**), which is in the range (100–3000 ppt) of
 386 previous measurements in polluted continental regions (Wang et al., 2017b; Wagner et al., 2012; Young
 387 et al., 2012; Riedel et al., 2013). N_2O_5 and ClNO_2 production from NO_x , O_3 , and chloride within the
 388 residual layer can also contribute to the ground ClNO_2 levels (Wang et al., 2017b; Tham et al., 2016b),
 389 but it is difficult to quantify in this study due to the lack of vertically-resolved measurements.

390 Since the above discussions are built on two short-term campaigns, we further assess the
 391 prevalence of N_2O_5 and ClNO_2 production in Delhi based on the long-term distributions of the
 392 identified driving factors, i.e., NO, $\text{P}(\text{NO}_3)$, and particulate Cl (**Fig. S2**). The occurrence distribution
 393 of nighttime NO and $\text{P}(\text{NO}_3)$ during the 2023 campaign largely overlapped with the conditions
 394 commonly occurred in Delhi (**Figs. S2a-b**), indicating comparable N_2O_5 production potential
 395 throughout the years in Delhi. In addition, the frequency distribution of Cl_p in 2023 campaign is
 396 comparable to that observed from January to April of 2022 and much higher than those reported from
 397 May to September (**Fig. S2c**). The seasonal variation of chloride reflects the temperature-driven
 398 evaporation of semi-volatile NH_4Cl as discussed previously. Given the critical role of Cl_p in promoting
 399 N_2O_5 uptake and driving ClNO_2 production in Delhi, these results suggest that elevated nighttime N_2O_5
 400 and ClNO_2 levels are most likely observed during the warm and cold seasons, respectively, in Delhi.

401

402

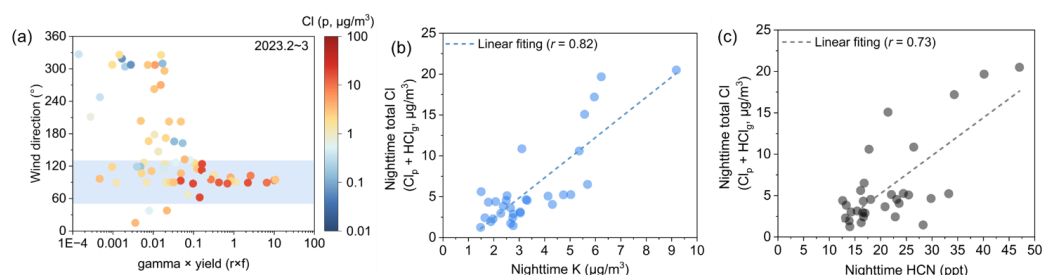


Fig. 4 The key source of Cl emission in Delhi. (a) Scatter plot of uptake parameter ($\gamma \times f$) and wind direction color-coded by particulate Cl concentrations. The shaded area in (a) indicates the region with wind direction ranging from 50° to 130° . Correlation of nocturnal total Cl (sum of Cl_p and HCl_g) with (b) particulate K and (c) gaseous HCN. The reported HCN mixing ratio can be considered a lower-limit estimate (see **Methods 2.1**).

3.3 Impacts on atmospheric radical production and VOCs oxidation

Enhanced N_2O_5 -ClNO₂ chemistry at lower NO levels emphasizes the important roles of $\text{Cl}\cdot$ and $\text{NO}_3\cdot$ in initiating the oxidation of VOCs in Delhi. The average daytime Cl production rate ($P(\text{Cl})$) from ClNO₂ photolysis in 2023 (0.035 ppb/h) is about approximately twice that observed in 2019 (0.015 ppb/h). The maximum daytime-averaged $P(\text{Cl})$ (0.24 ppb/h) from the photolysis of all the detected $\text{Cl}\cdot$ precursors (Cl_2 , ClNO₂, NHCl_2 , and NCl_3) is up to an order of magnitude higher than values reported for the urban inland areas of Europe and North America (0.02–0.07 ppb/h) (Priestley et al., 2018; McNamara et al., 2020; Faxon et al., 2015), and remains several times lower than those observed during the pollution episodes in China (0.5–1.1 ppb/h) (Chen et al., 2023; Tham et al., 2016b; Liu et al., 2017). ClNO₂ photolysis dominates the production of $\text{Cl}\cdot$ around the sunrise hours (7:00–8:00) when the $P(\text{Cl})_{\text{total}}$ to $P(\text{OH})$ ratio (42 on average) was also the highest (**Fig. 5a**), suggesting pronounced Cl-initiated oxidation of VOCs under these conditions in Delhi.

Consistently, we observed specific gaseous organic oxidation products indicative of the prevalence of chlorine chemistry (**Fig. 5b**). We identified organochlorinated species, i.e., $\text{C}_2\text{H}_3\text{O}_2\text{Cl}$ and $\text{C}_3\text{H}_5\text{O}_2\text{Cl}$ (presumably chloroacetic acid and chloropropionic acid), which were previously measured in chamber experiments from $\text{Cl}\cdot$ reacting with ethylbenzene (Jahn et al., 2024) and isoprene (Wang et al., 2022) and used as markers for the occurrence of chlorine chemistry in field studies (Priestley et al., 2018; Li et al., 2025). The daytime concentration fraction of $\text{C}_2\text{H}_3\text{O}_2\text{Cl}$ in the total measured gaseous organics increases from ~0.9% in 2019 to ~1.8% in 2023. Additionally, the daytime average peak mixing ratio of $\text{C}_2\text{H}_3\text{O}_2\text{Cl}$ (~100 ppt) in 2023 is about 5–30 times higher than those measured in Hong Kong (12–20 ppt) (Li et al., 2025) and Manchester (~3 ppt) (Priestley et al., 2018), indicating significant $\text{C}_2\text{H}_3\text{ClO}_2$ production in Delhi. We also identified a series of hydroxynitrates (HN) homologues, i.e., $\text{C}_8\text{H}_{17}\text{O}_4\text{N}$, $\text{C}_{10}\text{H}_{21}\text{O}_4\text{N}$, and $\text{C}_{12}\text{H}_{25}\text{O}_4\text{N}$, which were potentially early-generation products from the chlorine-initiated oxidation of $\text{C}_{8,10,12}$ -alkanes under high NO_x conditions



433 (Wang and Hildebrandt Ruiz, 2018). Here their sharp increase in the early morning hours (7:00~8:00)
 434 when $\text{Cl}\cdot$ was likely the dominant atmospheric oxidant (**Figs. 5a-b**), together with the positive
 435 correlations between these HN and the product of $\text{P}(\text{Cl})$ and NO ($r_{7:00\sim8:00}=0.59$, **Fig. 5e**), underscores
 436 the $\text{Cl}\cdot$ initiated alkane oxidation in the presence of NO as an important source of HN in Delhi. It is
 437 noted that these HN can also be formed via OH-initiated oxidation processes (Lim and Ziemann, 2009),
 438 but would then likely peak later in the day when OH concentrations are higher. Here, a quantitative
 439 comparison between the $\text{Cl}\cdot$ - and OH-initiated reaction pathways is missing in current study due to the
 440 lack of VOCs precursor measurements and warrants further investigation.

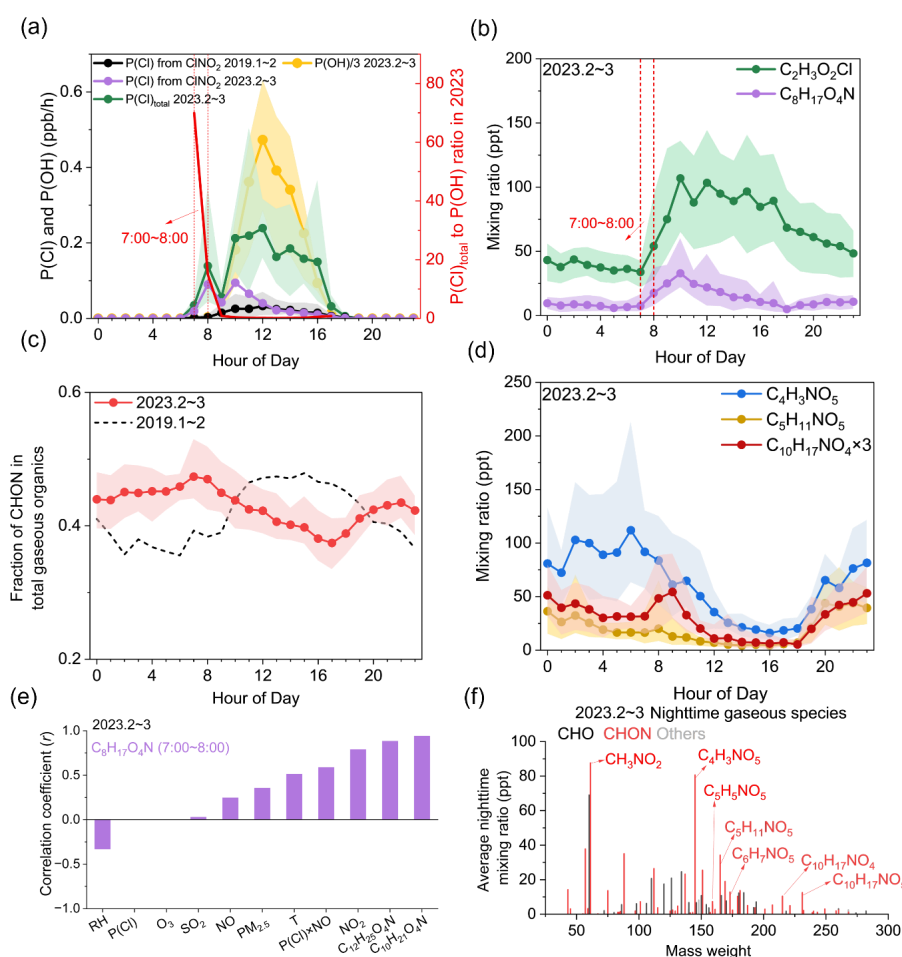
441 The NO_3 -initiated nocturnal oxidation of VOCs was strengthened during the 2023 campaign. The
 442 fraction of nitrogen-containing organics (CHON) in the total measured gaseous oxygenated organic
 443 compounds in 2023 (~46%) was similar to that observed in 2019 (42%). However, the diurnal pattern
 444 of the CHON fraction was inversed, exhibiting a moderate increase (~9%) during the nighttime (**Fig.**
 445 **5c**), which resembles those found in other locations (Ye et al., 2021; Huang et al., 2019), indicating
 446 enhanced nocturnal NO_3 -driven organic oxidation in 2023. We further identified the nighttime gaseous
 447 organic species with the criteria of average nighttime (20:00~04:00) to daytime (7:00~16:00) mixing
 448 ratio larger than 1. Most of the nighttime species are nitrogen-containing compounds, primarily
 449 including $\text{C}_{4,5,6}\text{H}_{3,5,7}\text{O}_5\text{N}$ and $\text{C}_{10}\text{H}_{17}\text{O}_{4,5}\text{N}$ (**Fig. 5f**), which are typical first-generation oxidation
 450 products from NO_3 -induced oxidation of heterocyclics (i.e., furan, methylfuran, and dimethylfuran)
 451 (Chen et al., 2022a; Joo et al., 2019; Jiang et al., 2020) and monoterpenes (Jenks et al., 2023; Ayres et
 452 al., 2015), respectively. These oxidation products exhibited strong intra-group correlations ($r_{20:00\sim4:00}$
 453 $= 0.71$ to 0.77), whereas the correlations between the $\text{C}_{4,5,6}\text{H}_{3,5,7}\text{O}_5\text{N}$ and $\text{C}_{10}\text{H}_{17}\text{O}_{4,5}\text{N}$ groups were
 454 weak ($r_{20:00\sim4:00} < 0.2$), confirming their production from distinct precursors. Moreover, their low
 455 correlations with CO and NO ($r_{20:00\sim4:00} < 0.4$) largely exclude direct combustion emissions as the
 456 dominant source. Notably, observations of furan-derived $\text{NO}_3\cdot$ oxidation products have currently been
 457 limited to laboratory studies (Chen et al., 2022a; Joo et al., 2019; Jiang et al., 2020). Our findings offer
 458 direct field evidence of NO_3 -initiated furan oxidation in the atmosphere.

459 Another prominent nighttime compound potentially associated with $\text{NO}_3\cdot$ induced oxidation is
 460 $\text{C}_5\text{H}_{11}\text{NO}_5$ (**Fig. 5f**), which exhibited the largest nocturnal enhancement among all the observed species,
 461 increasing by a factor of 3.6. $\text{C}_5\text{H}_{11}\text{NO}_5$ correlates well with $\text{C}_{10}\text{H}_{17}\text{O}_5\text{N}$ ($r_{20:00\sim4:00} = 0.71$) and shows
 462 a marked increase between 18:00 and 19:00 (**Fig. 5d**), coinciding with the peak of the estimated
 463 $\text{NO}_3\cdot$ levels (~2.3 ppt) (**Fig. S7b**). It is thus likely a co-product from the reaction of $\text{NO}_3\cdot$ with
 464 monoterpenes. However, this compound has not been reported from any NO_3 -initiated oxidation
 465 processes (Jenks et al., 2023; Ayres et al., 2015; Xu et al., 2025). Further investigation is required to
 466 elucidate the formation mechanisms of $\text{C}_5\text{H}_{11}\text{NO}_5$ in Delhi's atmosphere.

467 In addition to the specific organic products discussed above, the overall oxidation state of the
 468 measured organic aerosols (OA) is elevated in 2023 compared to 2019. The bulk oxygen-to-carbon



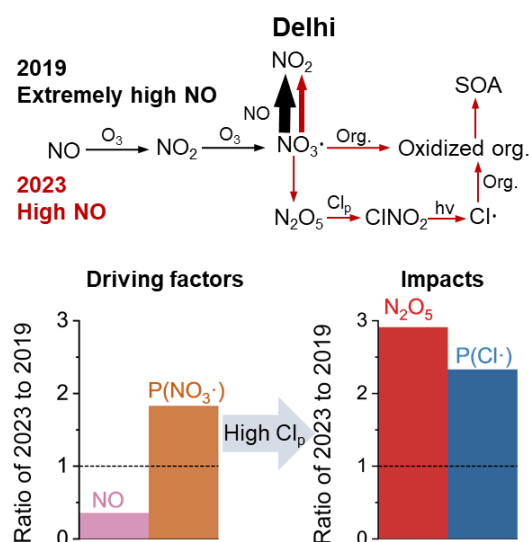
469 (O/C) ratio of OA (0.76 on average) was higher than the value (0.66) reported in the 2019 campaign
 470 (Huang et al., 2024), indicating an increased contribution from more oxidized or aged secondary
 471 organic aerosols (SOA) (Jimenez et al., 2009). A clear negative dependence of the O/C ratio on NO
 472 was observed at night when NO exceeded about 10 ppb (Figs. S18a-b), agreeing with the suppression
 473 effect of NO on organic oxidation by increasing the competitive $\text{RO}_2 + \text{NO}$ channel under the radical-
 474 limited regime commonly found in Delhi (Nelson et al., 2021; Kenagy et al., 2024). Taken together, the
 475 characteristics of the bulk organics provide additional field evidence for the enhanced atmospheric
 476 organics oxidation in 2023.



477
 478 **Fig. 5 Enhanced daytime $\text{Cl}\cdot$ and nocturnal $\text{NO}_3\cdot$ initiated oxidation of organics in Delhi.** (a) $\text{P}(\text{Cl})$
 479 in 2023 and 2019, and $\text{P}(\text{OH})$ from the photolysis of O_3 in 2023. The $\text{P}(\text{Cl})_{\text{total}}$ 2023 includes
 480 $\text{Cl}\cdot$ production from the sum of ClNO_2 , Cl_2 , NHCl_2 , and NCl_3 photolysis. The $\text{P}(\text{Cl})_{\text{total}}$ to $\text{P}(\text{OH})$ ratio
 481 in 2023 is shown in the right axis. The value of $\text{P}(\text{OH})$ is divided by a factor of 3 to place it on the



482 same scale as $P(\text{Cl})$. Average diurnal variations of tracer compounds indicating (b) $\text{Cl}\cdot$ and (d)
 483 $\text{NO}_3\cdot$ initiated oxidation, respectively. (c) Comparison of the diurnal variations of CHON compounds
 484 fraction in total detected gaseous organics in 2019 and 2023. (e) Correlation coefficients, using
 485 observations between 7:00 and 8:00 of $\text{C}_8\text{H}_{17}\text{NO}_4$ with $P(\text{Cl})_{\text{total}}$, meteorological parameters, trace
 486 gases, and related homologue species. (f) Molecular composition and average nighttime mixing ratios
 487 for the nighttime gaseous species. The shaded area in (a-d) indicates the 10th and 90th percentiles and
 488 the dots denote mean values.



489
 490 **Fig. 6 Amplified N_2O_5 and ClNO_2 production and impacts on secondary organic aerosol**
 491 **formation in Delhi.** A schematic comparing N_2O_5 - ClNO_2 chemistry between the 2019 and 2023
 492 campaigns. Red arrows indicate reactions that are largely enhanced during the 2023 campaign.

493 4 Conclusions

494 This study provides an extensive characterization of N_2O_5 - ClNO_2 chemistry in Delhi (**Fig. 6**). We
 495 observed a clear dependency on available nocturnal NO that high NO suppresses N_2O_5 production and
 496 reduces the formation of ClNO_2 while decreased NO promotes nocturnal N_2O_5 - ClNO_2 chemistry. The
 497 former pattern is consistent with the previously reported significant NO suppression in 2019 (Haslett
 498 et al., 2023). With moderate NO in 2023, substantial ClNO_2 is produced after midnight hours in Delhi,
 499 which is driven by the elevated RH and chloride levels that promote efficient conversion of N_2O_5 to
 500 ClNO_2 . The abundance of chloride strongly correlates with the biomass burning tracers, indicating
 501 emissions from biomass burning as an important source of Cl in Delhi. The increased ClNO_2 levels at
 502 lower nocturnal NO conditions implies that the atmospheric oxidation of organics is amplified both
 503 during the nighttime (via $\text{NO}_3\cdot$ pathways) and daytime (via $\text{Cl}\cdot$ pathways) in Delhi.



504 The differing behaviors of N_2O_5 and ClNO_2 depending on the levels of nocturnal NO emphasize
 505 the need of continuous monitoring of air pollutants in Delhi, which is essential to understand the
 506 evolution of emission sources and atmospheric chemical processes. The complexity of the air pollution
 507 situation in Delhi has been further scrutinized where we noted that as chloride levels declined, nitrate
 508 has become the leading inorganic component since late 2019 (**Figs. S3a-b**) and plays an increasingly
 509 important role in exacerbating Delhi's $\text{PM}_{2.5}$ pollution (**Fig. S3c**). Thus, future studies should pay
 510 attention to elucidating the dominant production pathways of nitrate in Delhi, especially the
 511 contribution from N_2O_5 heterogeneous hydrolysis (Wang et al., 2017a; Yan et al., 2023).

512 The present study suggests that NO_x emission control may intensify atmospheric oxidation of
 513 organics and secondary organic aerosol (SOA) formation in Delhi, whereas concurrent chlorine
 514 emission regulations may help mitigate the associated negative impacts. Despite recent reductions, the
 515 chloride concentrations in Delhi remain among the highest reported worldwide (**Fig. S1a**). Previous
 516 studies have revealed the critical role chloride has in exacerbating haze formation in Delhi by
 517 enhancing aerosol liquid water content (ALWC) (Chen et al., 2022b; Gunthe et al., 2021), while our
 518 study implies that, with the decrease in NO concentrations, chloride can also be readily activated into
 519 photolabile gases, contributing to the formation of secondary pollutant, e.g., SOA, through generating
 520 radicals and advancing atmospheric organics oxidation. Chloride is mostly of anthropogenic origins in
 521 Delhi (Sharma et al., 2019; Gunthe et al., 2021) and we identified biomass burning as one of the major
 522 sources of chlorine during our measurements. It is expected that regulations on biomass burning along
 523 with transition to clean residential fuels (e.g., natural gas and electricity) would benefit $\text{PM}_{2.5}$
 524 abatement in Delhi by not only cutting primary particle emissions, but also suppressing secondary
 525 pollutant formation. To support such strategies, an updated chlorine emission inventory combined with
 526 modelling studies, is necessary to quantitatively evaluate the contribution of chlorine chemistry to
 527 $\text{PM}_{2.5}$ formation under different NO_x reduction scenarios and inform targeted emission control
 528 strategies. Finally, the interplay between nitrogen and halogen chemistry in India warrants continued
 529 investigation in the context of ongoing industrialization, vehicle electrification, and a changing climate.

530 **Supporting Information**

531 Text S1 to S3, Table S1 to S3 and Figs. S1 to S20

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 538 University of Gothenburg are gratefully acknowledged for their contributions to the EDXRF analyses.



539 **Data Availability**

540 The dataset used for this study is available at <https://cloud.tsinghua.edu.cn/d/cb099983b459474fa1df/>.
 541 Measurement data from the Sri Aurobindo Marg monitoring station in Delhi is available at
 542 [https://www.dpccairdata.com/dpccairdata/display/AallAdvanceSearchMet.php?stName=U3JpQXVy](https://www.dpccairdata.com/dpccairdata/display/AallAdvanceSearchMet.php?stName=U3JpQXVyYmluZG9NYXJn)
 543 [YmluZG9NYXJn](https://www.dpccairdata.com/dpccairdata/display/AallAdvanceSearchMet.php?stName=U3JpQXVyYmluZG9NYXJn). The Tropospheric Ultraviolet and Visible Radiation (TUV) Model is accessible at
 544 https://www.acom.ucar.edu/Models/TUV/Interactive_TUV. The exact sunrise and sunset time during
 545 the 2023 and 2019 campaign are retrieved from the website
 546 <https://www.timeanddate.com/sun/india/delhi>.

547 **Author Contributions**

548 M.H., Y.C., C.W., and J.J. designed the study. E.T., J.P., C.W., R.P., and M.H. organized and performed
 549 the field observations in Delhi. R.P., H.M., G.H., and G.T. provided the measurement space, helped
 550 setting up the instruments, and facilitated the measurements. Y.C., E.T. and J.P. calibrated and
 551 quantified the data. Y.C. analyzed and visualized the data. Y.C. wrote the original draft with inputs
 552 from J.J., M.H., and C.W. All co-authors contributed to scientific discussions and reviewing the
 553 manuscript.

554 **Author Information**

555 The authors declare no competing financial interests.

556 **References**

- 557 Acharja, P., Ghude, S. D., Sinha, B., Barth, M., Govardhan, G., Kulkarni, R., Sinha, V., Kumar, R., Ali,
 558 K., Gultepe, I., Petit, J.-E., and Rajeevan, M. N.: Thermodynamical framework for effective mitigation
 559 of high aerosol loading in the Indo-Gangetic Plain during winter, *Scientific Reports*, 13, 13667,
 560 <https://doi.org/10.1038/s41598-023-40657-w>, 2023.
 561 Aggarwal, S., Bansal, P., Wang, Y., Jorga, S., Macgregor, G., Rohner, U., Bannan, T., Salter, M., Zieger,
 562 P., Mohr, C., and Lopez-Hilfiker, F.: Identifying key parameters that affect sensitivity of flow tube
 563 chemical ionization mass spectrometers, *Atmos. Meas. Tech.*, 18, 4227-4247,
 564 <https://doi.org/10.5194/amt-18-4227-2025>, 2025.
 565 Ahern, A. T., Goldberger, L., Jahl, L., Thornton, J., and Sullivan, R. C.: Production of N₂O₅ and ClNO₂
 566 through Nocturnal Processing of Biomass-Burning Aerosol, *Environ Sci Technol*, 52, 550-559,
 567 <https://doi.org/10.1021/acs.est.7b04386>, 2018.
 568 Ayres, B. R., Allen, H. M., Draper, D. C., Brown, S. S., Wild, R. J., Jimenez, J. L., Day, D. A.,
 569 Campuzano-Jost, P., Hu, W., de Gouw, J., Koss, A., Cohen, R. C., Duffey, K. C., Romer, P., Baumann,
 570 K., Edgerton, E., Takahama, S., Thornton, J. A., Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Wennberg,
 571 P. O., Nguyen, T. B., Teng, A., Goldstein, A. H., Olson, K., and Fry, J. L.: Organic nitrate aerosol
 572 formation via NO₃ + biogenic volatile organic compounds in the southeastern United States, *Atmos.*
 573 *Chem. Phys.*, 15, 13377-13392, <https://doi.org/10.5194/acp-15-13377-2015>, 2015.
 574 Bannan, T. J., Booth, A. M., Bacak, A., Muller, J. B. A., Leather, K. E., Le Breton, M., Jones, B., Young,
 575 D., Coe, H., Allan, J., Visser, S., Slowik, J. G., Furger, M., Prévôt, A. S. H., Lee, J., Dunmore, R. E.,



- 576 Hopkins, J. R., Hamilton, J. F., Lewis, A. C., Whalley, L. K., Sharp, T., Stone, D., Heard, D. E., Fleming,
 577 Z. L., Leigh, R., Shallcross, D. E., and Percival, C. J.: The first UK measurements of nitryl chloride
 578 using a chemical ionization mass spectrometer in central London in the summer of 2012, and an
 579 investigation of the role of Cl atom oxidation, *Journal of Geophysical Research: Atmospheres*, 120,
 580 5638-5657, <https://doi.org/10.1002/2014jd022629>, 2015.
- 581 Bertram, T. H., and Thornton, J. A.: Toward a general parameterization of N₂O₅ reactivity on aqueous
 582 particles: the competing effects of particle liquid water, nitrate and chloride, *Atmos. Chem. Phys.*, 9,
 583 8351-8363, <https://doi.org/10.5194/acp-9-8351-2009>, 2009.
- 584 Brown, S. S., and Stutz, J.: Nighttime radical observations and chemistry, *Chemical Society Reviews*,
 585 41, 6405-6447, <https://doi.org/10.1039/C2CS35181A>, 2012.
- 586 Bryant, D. J., Nelson, B. S., Swift, S. J., Budisulistiorini, S. H., Drysdale, W. S., Vaughan, A. R.,
 587 Newland, M. J., Hopkins, J. R., Cash, J. M., Langford, B., Nemitz, E., Acton, W. J. F., Hewitt, C. N.,
 588 Mandal, T., Gurjar, B. R., Shivani, Gadi, R., Lee, J. D., Rickard, A. R., and Hamilton, J. F.: Biogenic
 589 and anthropogenic sources of isoprene and monoterpenes and their secondary organic aerosol in Delhi,
 590 India, *Atmos. Chem. Phys.*, 23, 61-83, <https://doi.org/10.5194/acp-23-61-2023>, 2023.
- 591 Burkholder, J., Sander, S., Abbatt, J., Barker, J., Cappa, C., Crounse, J., Dibble, T., Huie, R., Kolb, C.,
 592 and Kurylo, M.: Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies
 593 Evaluation Number 19, JPL Publication 19-5, Jet Propulsion Laboratory, 2020.
- 594 Chen, K., Mayorga, R., Raeofy, N., Lum, M., Woods, M., Bahreini, R., Zhang, H., and Lin, Y.-H.:
 595 Effects of Nitrate Radical Levels and Pre-Existing Particles on Secondary Brown Carbon Formation
 596 from Nighttime Oxidation of Furan, *ACS Earth and Space Chemistry*, 6, 2709-2721,
 597 <https://doi.org/10.1021/acsearthspacechem.2c00244>, 2022a.
- 598 Chen, X., Xia, M., Wang, W., Yun, H., Yue, D., and Wang, T.: Fast near-surface ClNO₂ production and
 599 its impact on O₃ formation during a heavy pollution event in South China, *Sci Total Environ*, 858,
 600 159998, <https://doi.org/10.1016/j.scitotenv.2022.159998>, 2023.
- 601 Chen, X., Jiang, Y., Zong, Z., Wang, Y., Sun, W., Wang, Y., Xia, M., Guan, L., Liu, P., Zhang, C., Chen,
 602 J., Mu, Y., and Wang, T.: Atmospheric Reactive Halogens Reshaped by the Clean Energy Policy and
 603 Agricultural Activity in a Rural Area of the North China Plain, *Environmental Science & Technology*,
 604 <https://doi.org/10.1021/acs.est.4c13986>, 2025a.
- 605 Chen, Y., Wang, Y., Nenes, A., Wild, O., Song, S., Hu, D., Liu, D., He, J., Hildebrandt Ruiz, L., Apte,
 606 J. S., Gunthe, S. S., and Liu, P.: Ammonium Chloride Associated Aerosol Liquid Water Enhances Haze
 607 in Delhi, India, *Environ Sci Technol*, <https://doi.org/10.1021/acs.est.2c00650>, 2022b.
- 608 Chen, Y., Xia, M., Zhang, J., Tsiligiannis, E., Wu, C., Yan, C., Cai, R., Zheng, G., Li, Y., Guo, J., An,
 609 Z., Li, Y., Zhao, X., Qu, Q., Hua, C., Wang, Z., Wang, S., Liu, Y., Cao, L., He, K., Kulmala, M.,
 610 Hallquist, M., Wang, T., Worsnop, D., and Jiang, J.: Chloramine chemistry as a missing link in
 611 atmospheric chlorine cycling, *Science Advances*, 11, eadv4298,
 612 <https://doi.org/10.1126/sciadv.adv4298>, 2025b.
- 613 Dentener, F. J., and Crutzen, P. J.: Reaction of N₂O₅ on tropospheric aerosols: Impact on the global
 614 distributions of NO_x, O₃, and OH, *Journal of Geophysical Research: Atmospheres*, 98, 7149-7163,
 615 <https://doi.org/https://doi.org/10.1029/92JD02979>, 1993.
- 616 Dunlea, E. J., and Ravishankara, A.: Measurement of the rate coefficient for the reaction of O(¹D) with



- 617 H₂O and re-evaluation of the atmospheric OH production rate, *Physical Chemistry Chemical Physics*,
 618 6, 3333-3340, 2004.
- 619 Eger, P. G., Friedrich, N., Schuladen, J., Shenolikar, J., Fischer, H., Tadic, I., Harder, H., Martinez, M.,
 620 Rohloff, R., Tauer, S., Drewnick, F., Fachinger, F., Brooks, J., Darbyshire, E., Sciare, J., Pikridas, M.,
 621 Lelieveld, J., and Crowley, J. N.: Shipborne measurements of ClNO₂ in the Mediterranean Sea and
 622 around the Arabian Peninsula during summer, *Atmospheric Chemistry and Physics*, 19, 12121-12140,
 623 <https://doi.org/10.5194/acp-19-12121-2019>, 2019.
- 624 Evans, M. J., and Jacob, D. J.: Impact of new laboratory studies of N₂O₅ hydrolysis on global model
 625 budgets of tropospheric nitrogen oxides, ozone, and OH, *Geophysical Research Letters*, 32,
 626 [https://doi.org/https://doi.org/10.1029/2005GL022469](https://doi.org/10.1029/2005GL022469), 2005.
- 627 Fauré, N., Chen, J., Artiglia, L., Ammann, M., Bartels-Rausch, T., Kanji, Z. A., Wang, S., Pettersson,
 628 J. B. C., Thomson, E. S., Gladich, I., and Kong, X.: Formation of Sodium Chloride on the Surface of
 629 Sulfate-Rich Gobi Desert Salt in Response to Water Adsorption, *ACS ES&T Air*, 1, 1373-1382,
 630 <https://doi.org/10.1021/acsestair.4c00092>, 2024.
- 631 Faxon, C., Bean, J., and Ruiz, L.: Inland Concentrations of Cl₂ and ClNO₂ in Southeast Texas Suggest
 632 Chlorine Chemistry Significantly Contributes to Atmospheric Reactivity, *Atmosphere*, 6, 1487-1506,
 633 <https://doi.org/10.3390/atmos6101487>, 2015.
- 634 Finlayson-Pitts, B. J., Ezell, M. J., and Pitts, J. N.: Formation of chemically active chlorine compounds
 635 by reactions of atmospheric NaCl particles with gaseous N₂O₅ and ClONO₂, *Nature*, 337, 241-244,
 636 <https://doi.org/10.1038/337241a0>, 1989.
- 637 Gani, S., Bhandari, S., Patel, K., Seraj, S., Soni, P., Arub, Z., Habib, G., Hildebrandt Ruiz, L., and Apte,
 638 J. S.: Particle number concentrations and size distribution in a polluted megacity: the Delhi Aerosol
 639 Supersite study, *Atmos. Chem. Phys.*, 20, 8533-8549, <https://doi.org/10.5194/acp-20-8533-2020>, 2020.
- 640 Gunthe, S. S., Liu, P., Panda, U., Raj, S. S., Sharma, A., Darbyshire, E., Reyes-Villegas, E., Allan, J.,
 641 Chen, Y., Wang, X., Song, S., Pöhlker, M. L., Shi, L., Wang, Y., Kommula, S. M., Liu, T., Ravikrishna,
 642 R., McFiggans, G., Mickley, L. J., Martin, S. T., Pöschl, U., Andreae, M. O., and Coe, H.: Enhanced
 643 aerosol particle growth sustained by high continental chlorine emission in India, *Nature Geoscience*,
 644 14, 77-84, <https://doi.org/10.1038/s41561-020-00677-x>, 2021.
- 645 Haslett, S. L., Bell, D. M., Kumar, V., Slowik, J. G., Wang, D. S., Mishra, S., Rastogi, N., Singh, A.,
 646 Ganguly, D., Thornton, J., Zheng, F., Li, Y., Nie, W., Liu, Y., Ma, W., Yan, C., Kulmala, M.,
 647 Daellenbach, K. R., Hadden, D., Baltensperger, U., Prevot, A. S. H., Tripathi, S. N., and Mohr, C.:
 648 Nighttime NO emissions strongly suppress chlorine and nitrate radical formation during the winter in
 649 Delhi, *Atmos. Chem. Phys.*, 23, 9023-9036, <https://doi.org/10.5194/acp-23-9023-2023>, 2023.
- 650 Huang, W., Saathoff, H., Shen, X., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical Characterization
 651 of Highly Functionalized Organonitrates Contributing to Night-Time Organic Aerosol Mass Loadings
 652 and Particle Growth, *Environmental Science & Technology*, 53, 1165-1174,
 653 <https://doi.org/10.1021/acs.est.8b05826>, 2019.
- 654 Huang, W., Wu, C., Gao, L., Gramlich, Y., Haslett, S. L., Thornton, J., Lopez-Hilfiker, F. D., Lee, B.
 655 H., Song, J., Saathoff, H., Shen, X., Ramisetty, R., Tripathi, S. N., Ganguly, D., Jiang, F., Vallon, M.,
 656 Schobesberger, S., Yli-Juuti, T., and Mohr, C.: Variation in chemical composition and volatility of
 657 oxygenated organic aerosol in different rural, urban, and mountain environments, *Atmos. Chem. Phys.*,



- 24, 2607-2624, <https://doi.org/10.5194/acp-24-2607-2024>, 2024.
- Jahl, L. G., Bowers, B. B., Jahn, L. G., Thornton, J. A., and Sullivan, R. C.: Response of the Reaction Probability of N₂O₅ with Authentic Biomass-Burning Aerosol to High Relative Humidity, *ACS Earth and Space Chemistry*, 5, 2587-2598, <https://doi.org/10.1021/acsearthspacechem.1c00227>, 2021.
- Jahn, L. G., McPherson, K. N., and Hildebrandt Ruiz, L.: Effects of Relative Humidity and Photoaging on the Formation, Composition, and Aging of Ethylbenzene SOA: Insights from Chamber Experiments on Chlorine Radical-Initiated Oxidation of Ethylbenzene, *ACS Earth and Space Chemistry*, 8, 675-688, <https://doi.org/10.1021/acsearthspacechem.3c00279>, 2024.
- Jenks, O. J., DeVault, M. P., Ziola, A. C., Morris, M. A., Schueneman, M. K., Stark, H., Jimenez, J. L., Ziemann, P. J., and de Gouw, J. A.: Investigation of Gas-Phase Products from the NO₃ Radical Oxidation of Δ-3-Carene, *ACS Earth and Space Chemistry*, 7, 1097-1106, <https://doi.org/10.1021/acsearthspacechem.3c00020>, 2023.
- Jiang, J., Carter, W. P. L., Cocker, D. R., III, and Barsanti, K. C.: Development and Evaluation of a Detailed Mechanism for Gas-Phase Atmospheric Reactions of Furans, *ACS Earth and Space Chemistry*, 4, 1254-1268, <https://doi.org/10.1021/acsearthspacechem.0c00058>, 2020.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., E., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326, 1525-1529, <https://doi.org/doi:10.1126/science.1180353>, 2009.
- Joo, T., Rivera-Rios, J. C., Takeuchi, M., Alvarado, M. J., and Ng, N. L.: Secondary Organic Aerosol Formation from Reaction of 3-Methylfuran with Nitrate Radicals, *ACS Earth and Space Chemistry*, 3, 922-934, <https://doi.org/10.1021/acsearthspacechem.9b00068>, 2019.
- Kashyap, P., Kumar, A., Kumar, R. P., and Kumar, K.: Biogenic and anthropogenic isoprene emissions in the subtropical urban atmosphere of Delhi, *Atmospheric Pollution Research*, 10, 1691-1698, <https://doi.org/https://doi.org/10.1016/j.apr.2019.07.004>, 2019.
- Kenagy, H. S., Heald, C. L., Tahsini, N., Goss, M. B., and Kroll, J. H.: Can we achieve atmospheric chemical environments in the laboratory? An integrated model-measurement approach to chamber SOA studies, *Science Advances*, 10, eado1482, <https://doi.org/10.1126/sciadv.ado1482>, 2024.
- Kercher, J. P., Riedel, T. P., and Thornton, J. A.: Chlorine activation by N₂O₅: simultaneous, in situ detection of ClNO₂ and N₂O₅ by chemical ionization mass spectrometry, *Atmos. Meas. Tech.*, 2, 193-204, <https://doi.org/10.5194/amt-2-193-2009>, 2009.
- Kumar, V., Giannoukos, S., Haslett, S. L., Tong, Y., Singh, A., Bertrand, A., Lee, C. P., Wang, D. S., Bhattu, D., Stefanelli, G., Dave, J. S., Puthussery, J. V., Qi, L., Vats, P., Rai, P., Casotto, R., Satish, R., Mishra, S., Pospisilova, V., Mohr, C., Bell, D. M., Ganguly, D., Verma, V., Rastogi, N., Baltensperger,



- 699 U., Tripathi, S. N., Prévôt, A. S. H., and Slowik, J. G.: Highly time-resolved chemical speciation and
 700 source apportionment of organic aerosol components in Delhi, India, using extractive electrospray
 701 ionization mass spectrometry, *Atmos. Chem. Phys.*, 22, 7739-7761, [https://doi.org/10.5194/acp-22-](https://doi.org/10.5194/acp-22-7739-2022)
 702 [7739-2022](https://doi.org/10.5194/acp-22-7739-2022), 2022.
- 703 Le Breton, M., Hallquist, Å. M., Pathak, R. K., Simpson, D., Wang, Y., Johansson, J., Zheng, J., Yang,
 704 Y., Shang, D., Wang, H., Liu, Q., Chan, C., Wang, T., Bannan, T. J., Priestley, M., Percival, C. J.,
 705 Shallcross, D. E., Lu, K., Guo, S., Hu, M., and Hallquist, M.: Chlorine oxidation of VOCs at a semi-
 706 rural site in Beijing: significant chlorine liberation from ClNO₂ and subsequent gas- and particle-phase
 707 Cl-VOC production, *Atmospheric Chemistry and Physics*, 18, 13013-13030,
 708 <https://doi.org/10.5194/acp-18-13013-2018>, 2018a.
- 709 Le Breton, M., Wang, Y., Hallquist, Å. M., Pathak, R. K., Zheng, J., Yang, Y., Shang, D., Glasius, M.,
 710 Bannan, T. J., Liu, Q., Chan, C. K., Percival, C. J., Zhu, W., Lou, S., Topping, D., Wang, Y., Yu, J., Lu,
 711 K., Guo, S., Hu, M., and Hallquist, M.: Online gas- and particle-phase measurements of organosulfates,
 712 organosulfonates and nitrooxy organosulfates in Beijing utilizing a FIGAERO ToF-CIMS,
 713 *Atmospheric Chemistry and Physics*, 18, 10355-10371, <https://doi.org/10.5194/acp-18-10355-2018>,
 714 2018b.
- 715 Li, M., Xia, M., Lin, C., Jiang, Y., Sun, W., Wang, Y., Zhang, Y., He, M., and Wang, T.: Mechanistic
 716 insights into chloroacetic acid production from atmospheric multiphase volatile organic compound-
 717 chlorine chemistry, *Atmos. Chem. Phys.*, 25, 3753-3764, <https://doi.org/10.5194/acp-25-3753-2025>,
 718 2025.
- 719 Lim, Y. B., and Ziemann, P. J.: Chemistry of Secondary Organic Aerosol Formation from OH Radical-
 720 Initiated Reactions of Linear, Branched, and Cyclic Alkanes in the Presence of NO_x, *Aerosol Science*
 721 *and Technology*, 43, 604-619, <https://doi.org/10.1080/02786820902802567>, 2009.
- 722 Liu, X., Qu, H., Huey, L. G., Wang, Y., Sjostedt, S., Zeng, L., Lu, K., Wu, Y., Hu, M., Shao, M., Zhu,
 723 T., and Zhang, Y.: High Levels of Daytime Molecular Chlorine and Nitryl Chloride at a Rural Site on
 724 the North China Plain, *Environ Sci Technol*, 51, 9588-9595, <https://doi.org/10.1021/acs.est.7b03039>,
 725 2017.
- 726 Ljungström, E., and Hallquist, M.: Nitrate radical formation rates in scandinavia, *Atmospheric*
 727 *Environment*, 30, 2925-2932, [https://doi.org/https://doi.org/10.1016/1352-2310\(96\)00006-4](https://doi.org/https://doi.org/10.1016/1352-2310(96)00006-4), 1996.
- 728 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A.,
 729 Hallquist, M., Worsnop, D., and Thornton, J. A.: A novel method for online analysis of gas and particle
 730 composition: description and evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO),
 731 *Atmospheric Measurement Techniques*, 7, 983-1001, <https://doi.org/10.5194/amt-7-983-2014>, 2014.
- 732 Ma, W., Chen, X., Xia, M., Liu, Y., Wang, Y., Zhang, Y., Zheng, F., Zhan, J., Hua, C., Wang, Z., Wang,
 733 W., Fu, P., Kulmala, M., and Liu, Y.: Reactive Chlorine Species Advancing the Atmospheric Oxidation
 734 Capacities of Inland Urban Environments, *Environmental Science & Technology*, 57, 14638-14647,
 735 <https://doi.org/10.1021/acs.est.3c05169>, 2023.
- 736 McNamara, S. M., AR, W. R., Wang, S., Thanekar, S., Boone, E. J., Kolesar, K. R., Peterson, P. K.,
 737 Simpson, W. R., Fuentes, J. D., Shepson, P. B., and Pratt, K. A.: Springtime Nitrogen Oxide-Influenced
 738 Chlorine Chemistry in the Coastal Arctic, *Environ Sci Technol*, 53, 8057-8067,
 739 <https://doi.org/10.1021/acs.est.9b01797>, 2019.



- McNamara, S. M., Kolesar, K. R., Wang, S., Kirpes, R. M., May, N. W., Gunsch, M. J., Cook, R. D.,
 Fuentes, J. D., Hornbrook, R. S., Apel, E. C., China, S., Laskin, A., and Pratt, K. A.: Observation of
 Road Salt Aerosol Driving Inland Wintertime Atmospheric Chlorine Chemistry, *ACS Cent Sci*, 6, 684-
 694, <https://doi.org/10.1021/acscentsci.9b00994>, 2020.
- Mielke, L. H., Stutz, J., Tsai, C., Hurlock, S. C., Roberts, J. M., Veres, P. R., Froyd, K. D., Hayes, P.
 L., Cubison, M. J., Jimenez, J. L., Washenfelder, R. A., Young, C. J., Gilman, J. B., Gouw, J. A., Flynn,
 J. H., Grossberg, N., Lefer, B. L., Liu, J., Weber, R. J., and Osthoff, H. D.: Heterogeneous formation
 of nitryl chloride and its role as a nocturnal NO_x reservoir species during CalNex - LA 2010, *Journal
 of Geophysical Research: Atmospheres*, 118, 10638-10652, <https://doi.org/10.1002/jgrd.50783>, 2013.
- Nelson, B. S., Stewart, G. J., Drysdale, W. S., Newland, M. J., Vaughan, A. R., Dunmore, R. E.,
 Edwards, P. M., Lewis, A. C., Hamilton, J. F., Acton, W. J., Hewitt, C. N., Crilley, L. R., Alam, M. S.,
 Şahin, Ü. A., Beddows, D. C. S., Bloss, W. J., Slater, E., Whalley, L. K., Heard, D. E., Cash, J. M.,
 Langford, B., Nemitz, E., Sommariva, R., Cox, S., Shivani, Gadi, R., Gurjar, B. R., Hopkins, J. R.,
 Rickard, A. R., and Lee, J. D.: In situ ozone production is highly sensitive to volatile organic
 compounds in Delhi, India, *Atmos. Chem. Phys.*, 21, 13609-13630, <https://doi.org/10.5194/acp-21-13609-2021>, 2021.
- Noxon, J. F., Norton, R. B., and Marovich, E.: NO₃ in the troposphere, *Geophysical Research Letters*,
 7, 125-128, <https://doi.org/10.1029/GL007i002p00125>, 1980.
- Ogrizek, M., Kroflič, A., and Šala, M.: Critical review on the development of analytical techniques for
 the elemental analysis of airborne particulate matter, *Trends in Environmental Analytical Chemistry*,
 33, e00155, <https://doi.org/10.1016/j.teac.2022.e00155>, 2022.
- Osthoff, H. D., Roberts, J. M., Ravishankara, A. R., Williams, E. J., Lerner, B. M., Sommariva, R.,
 Bates, T. S., Coffman, D., Quinn, P. K., Dibb, J. E., Stark, H., Burkholder, J. B., Talukdar, R. K.,
 Meagher, J., Fehsenfeld, F. C., and Brown, S. S.: High levels of nitryl chloride in the polluted
 subtropical marine boundary layer, *Nature Geoscience*, 1, 324-328, <https://doi.org/10.1038/ngeo177>,
 2008.
- Peng, X., Wang, W., Xia, M., Chen, H., Ravishankara, A. R., Li, Q., Saiz-Lopez, A., Liu, P., Zhang, F.,
 Zhang, C., Xue, L., Wang, X., George, C., Wang, J., Mu, Y., Chen, J., and Wang, T.: An unexpected
 large continental source of reactive bromine and chlorine with significant impact on wintertime air
 quality, *Natl Sci Rev*, 8, nwaa304, <https://doi.org/10.1093/nsr/nwaa304>, 2021.
- Phillips, G. J., Tang, M. J., Thieser, J., Brickwedde, B., Schuster, G., Bohn, B., Lelieveld, J., and
 Crowley, J. N.: Significant concentrations of nitryl chloride observed in rural continental Europe
 associated with the influence of sea salt chloride and anthropogenic emissions, *Geophysical Research
 Letters*, 39, n/a-n/a, <https://doi.org/10.1029/2012gl051912>, 2012.
- Priestley, M., le Breton, M., Bannan, T. J., Worrall, S. D., Bacak, A., Smedley, A. R. D., Reyes-Villegas,
 E., Mehra, A., Allan, J., Webb, A. R., Shallcross, D. E., Coe, H., and Percival, C. J.: Observations of
 organic and inorganic chlorinated compounds and their contribution to chlorine radical concentrations
 in an urban environment in northern Europe during the wintertime, *Atmospheric Chemistry and
 Physics*, 18, 13481-13493, <https://doi.org/10.5194/acp-18-13481-2018>, 2018.
- Rai, P., Furger, M., El Haddad, I., Kumar, V., Wang, L., Singh, A., Dixit, K., Bhattu, D., Petit, J.-E.,
 Ganguly, D., Rastogi, N., Baltensperger, U., Tripathi, S. N., Slowik, J. G., and Prévôt, A. S. H.: Real-



781 time measurement and source apportionment of elements in Delhi's atmosphere, *Science of The Total*
 782 *Environment*, 742, 140332, <https://doi.org/10.1016/j.scitotenv.2020.140332>, 2020.

783 Riedel, T. P., Bertram, T. H., Crisp, T. A., Williams, E. J., Lerner, B. M., Vlasenko, A., Li, S. M.,
 784 Gilman, J., de Gouw, J., Bon, D. M., Wagner, N. L., Brown, S. S., and Thornton, J. A.: Nitryl chloride
 785 and molecular chlorine in the coastal marine boundary layer, *Environ Sci Technol*, 46, 10463-10470,
 786 <https://doi.org/10.1021/es204632r>, 2012.

787 Riedel, T. P., Wagner, N. L., Dubé, W. P., Middlebrook, A. M., Young, C. J., Öztürk, F., Bahreini, R.,
 788 VandenBoer, T. C., Wolfe, D. E., Williams, E. J., Roberts, J. M., Brown, S. S., and Thornton, J. A.:
 789 Chlorine activation within urban or power plant plumes: Vertically resolved ClNO₂ and Cl₂
 790 measurements from a tall tower in a polluted continental setting, *Journal of Geophysical Research:*
 791 *Atmospheres*, 118, 8702-8715, <https://doi.org/10.1002/jgrd.50637>, 2013.

792 Sharma, G., Sinha, B., Pallavi, Hakkim, H., Chandra, B. P., Kumar, A., and Sinha, V.: Gridded
 793 Emissions of CO, NO_x, SO₂, CO₂, NH₃, HCl, CH₄, PM_{2.5}, PM₁₀, BC, and NMVOC from Open
 794 Municipal Waste Burning in India, *Environmental Science & Technology*, 53, 4765-4774,
 795 <https://doi.org/10.1021/acs.est.8b07076>, 2019.

796 Simpson, W. R., Brown, S. S., Saiz-Lopez, A., Thornton, J. A., and Glasow, R.: Tropospheric halogen
 797 chemistry: sources, cycling, and impacts, *Chem Rev*, 115, 4035-4062,
 798 <https://doi.org/10.1021/cr5006638>, 2015.

799 Tham, Y. J., Wang, Z., Li, Q., Yun, H., Wang, W., Wang, X., Xue, L., Lu, K., Ma, N., Bohn, B., Li, X.,
 800 Kecorius, S., Größ, J., Shao, M., Wiedensohler, A., Zhang, Y., and Wang, T.: Significant concentrations
 801 of nitryl chloride sustained in the morning: investigations of the causes and impacts on ozone
 802 production in a polluted region of northern China, *Atmos. Chem. Phys.*, 16, 14959-14977,
 803 <https://doi.org/10.5194/acp-16-14959-2016>, 2016a.

804 Tham, Y. J., Wang, Z., Li, Q., Yun, H., Wang, W., Wang, X., Xue, L., Lu, K., Ma, N., Bohn, B., Li, X.,
 805 Kecorius, S., Größ, J., Shao, M., Wiedensohler, A., Zhang, Y., and Wang, T.: Significant concentrations
 806 of nitryl chloride sustained in the morning: investigations of the causes and impacts on ozone
 807 production in a polluted region of northern China, *Atmospheric Chemistry and Physics*, 16, 14959-
 808 14977, <https://doi.org/10.5194/acp-16-14959-2016>, 2016b.

809 Tham, Y. J., Wang, Z., Li, Q., Wang, W., Wang, X., Lu, K., Ma, N., Yan, C., Kecorius, S., Wiedensohler,
 810 A., Zhang, Y., and Wang, T.: Heterogeneous N₂O₅ uptake coefficient and production yield of ClNO₂
 811 in polluted northern China: roles of aerosol water content and chemical composition, *Atmos. Chem.*
 812 *Phys.*, 18, 13155-13171, <https://doi.org/10.5194/acp-18-13155-2018>, 2018.

813 Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: N₂O₅ hydrolysis on sub-micron organic aerosols:
 814 the effect of relative humidity, particle phase, and particle size, *Physical Chemistry Chemical Physics*,
 815 5, 4593-4603, <https://doi.org/10.1039/B307498F>, 2003.

816 Thornton, J. A., Kercher, J. P., Riedel, T. P., Wagner, N. L., Cozic, J., Holloway, J. S., Dube, W. P.,
 817 Wolfe, G. M., Quinn, P. K., Middlebrook, A. M., Alexander, B., and Brown, S. S.: A large atomic
 818 chlorine source inferred from mid-continental reactive nitrogen chemistry, *Nature*, 464, 271-274,
 819 <https://doi.org/10.1038/nature08905>, 2010.

820 Trebs, I., Bohn, B., Ammann, C., Rummel, U., Blumthaler, M., Königstedt, R., Meixner, F. X., Fan,
 821 S., and Andreae, M. O.: Relationship between the NO₂ photolysis frequency and the solar global



- 822 irradiance, *Atmos. Meas. Tech.*, 2, 725-739, <https://doi.org/10.5194/amt-2-725-2009>, 2009.
- 823 Unga, F., Calzolari, G., Chiari, M., Cuccia, E., Colombi, C., Franciosa, M., Dinoi, A., Merico, E.,
- 824 Pennetta, A., Gómez-Sánchez, N., Mapelli, C., Pareti, S., Perrino, C., Yubero, E., and Contini, D.:
825 Determination of aerosol composition by ED-XRF on Teflon and quartz substrates: potentialities and
826 limits, *Aerosol Research*, 3, 405-415, <https://doi.org/10.5194/ar-3-405-2025>, 2025.
- 827 Wagner, N. L., Riedel, T. P., Roberts, J. M., Thornton, J. A., Angevine, W. M., Williams, E. J., Lerner,
828 B. M., Vlasenko, A., Li, S. M., Dubé, W. P., Coffman, D. J., Bon, D. M., de Gouw, J. A., Kuster, W.
829 C., Gilman, J. B., and Brown, S. S.: The sea breeze/land breeze circulation in Los Angeles and its
830 influence on nitryl chloride production in this region, *Journal of Geophysical Research: Atmospheres*,
831 117, <https://doi.org/https://doi.org/10.1029/2012JD017810>, 2012.
- 832 Wang, C., Liggitto, J., Wentzell, J. J. B., Jorga, S., Folkerson, A., and Abbatt, J. P. D.: Chloramines as
833 an important photochemical source of chlorine atoms in the urban atmosphere, *Proceedings of the*
834 *National Academy of Sciences*, 120, e2220889120, <https://doi.org/doi:10.1073/pnas.2220889120>,
835 2023a.
- 836 Wang, D. S., and Hildebrandt Ruiz, L.: Chlorine-initiated oxidation of n-alkanes under high-NOx
837 conditions: insights into secondary organic aerosol composition and volatility using a FIGAERO–
838 CIMS, *Atmos. Chem. Phys.*, 18, 15535-15553, <https://doi.org/10.5194/acp-18-15535-2018>, 2018.
- 839 Wang, D. S., Masoud, C. G., Modi, M., and Hildebrandt Ruiz, L.: Isoprene–Chlorine Oxidation in the
840 Presence of NOx and Implications for Urban Atmospheric Chemistry, *Environmental Science &*
841 *Technology*, 56, 9251-9264, <https://doi.org/10.1021/acs.est.1c07048>, 2022.
- 842 Wang, H., Lu, K., Chen, X., Zhu, Q., Chen, Q., Guo, S., Jiang, M., Li, X., Shang, D., Tan, Z., Wu, Y.,
843 Wu, Z., Zou, Q., Zheng, Y., Zeng, L., Zhu, T., Hu, M., and Zhang, Y.: High N2O5 Concentrations
844 Observed in Urban Beijing: Implications of a Large Nitrate Formation Pathway, *Environmental*
845 *Science & Technology Letters*, 4, 416-420, <https://doi.org/10.1021/acs.estlett.7b00341>, 2017a.
- 846 Wang, H., Wang, H., Lu, X., Lu, K., Zhang, L., Tham, Y. J., Shi, Z., Aikin, K., Fan, S., Brown, S. S.,
847 and Zhang, Y.: Increased night-time oxidation over China despite widespread decrease across the globe,
848 *Nature Geoscience*, <https://doi.org/10.1038/s41561-022-01122-x>, 2023b.
- 849 Wang, J., Wang, H., Tham, Y. J., Ming, L., Zheng, Z., Fang, G., Sun, C., Ling, Z., Zhao, J., and Fan,
850 S.: Measurement report: Atmospheric nitrate radical chemistry in the South China Sea influenced by
851 the urban outflow of the Pearl River Delta, *Atmos. Chem. Phys.*, 24, 977-992,
852 <https://doi.org/10.5194/acp-24-977-2024>, 2024.
- 853 Wang, T., Brown, S. S., Dubé, W. P., Tham, Y. J., Zha, Q., Xue, L., Poon, S., Wang, Z., Blake, D. R.,
854 Tsui, W., and Parrish, D. D.: Observations of nitryl chloride and modeling its source and effect on
855 ozone in the planetary boundary layer of southern China, *Journal of Geophysical Research:*
856 *Atmospheres*, 121, 2457-2475, <https://doi.org/10.1002/2015jd024566>, 2016.
- 857 Wang, X., Jacob, D. J., Eastham, S. D., Sulprizio, M. P., Zhu, L., Chen, Q., Alexander, B., Sherwen,
858 T., Evans, M. J., Lee, B. H., Haskins, J. D., Lopez-Hilfiker, F. D., Thornton, J. A., Huey, G. L., and
859 Liao, H.: The role of chlorine in global tropospheric chemistry, *Atmospheric Chemistry and Physics*,
860 19, 3981-4003, <https://doi.org/10.5194/acp-19-3981-2019>, 2019.
- 861 Wang, Z., Wang, W., Tham, Y. J., Li, Q., Wang, H., Wen, L., Wang, X., and Wang, T.: Fast
862 heterogeneous N2O5 uptake and ClNO2 production in power plant and industrial plumes observed in



- the nocturnal residual layer over the North China Plain, *Atmospheric Chemistry and Physics*, 17, 12361-12378, <https://doi.org/10.5194/acp-17-12361-2017>, 2017b.
- Wängberg, I., Etzkorn, T., Barnes, I., Platt, U., and Becker, K. H.: Absolute Determination of the Temperature Behavior of the $\text{NO}_2 + \text{NO}_3 + (\text{M}) \leftrightarrow \text{N}_2\text{O}_5 + (\text{M})$ Equilibrium, *The Journal of Physical Chemistry A*, 101, 9694-9698, <https://doi.org/10.1021/jp972203o>, 1997.
- Xia, M., Peng, X., Wang, W., Yu, C., Sun, P., Li, Y., Liu, Y., Xu, Z., Wang, Z., Xu, Z., Nie, W., Ding, A., and Wang, T.: Significant production of ClNO_2 and possible source of Cl_2 from N_2O_5 uptake at a suburban site in eastern China, *Atmospheric Chemistry and Physics*, 20, 6147-6158, <https://doi.org/10.5194/acp-20-6147-2020>, 2020.
- Xia, M., Peng, X., Wang, W., Yu, C., Wang, Z., Tham, Y. J., Chen, J., Chen, H., Mu, Y., Zhang, C., Liu, P., Xue, L., Wang, X., Gao, J., Li, H., and Wang, T.: Winter ClNO_2 formation in the region of fresh anthropogenic emissions: seasonal variability and insights into daytime peaks in northern China, *Atmospheric Chemistry and Physics*, 21, 15985-16000, <https://doi.org/10.5194/acp-21-15985-2021>, 2021.
- Xia, M., Jiang, Y., Dai, J., Liu, Y., Yan, C., Kulmala, M., and Wang, T.: Chlorine Activation in Marine Air: Insights From Chemical Budgets of Molecular Chlorine and Hypochlorous Acid, *Journal of Geophysical Research: Atmospheres*, 130, e2024JD042568, <https://doi.org/https://doi.org/10.1029/2024JD042568>, 2025.
- Xu, T., Takeuchi, M., Rivera-Rios, J. C., and Ng, N. L.: Multigeneration Chemistry in Secondary Organic Aerosol Formation from Nitrate Radical Oxidation of Isoprene, *ACS Earth and Space Chemistry*, 9, 411-423, <https://doi.org/10.1021/acsearthspacechem.4c00417>, 2025.
- Yan, C., Tham, Y. J., Nie, W., Xia, M., Wang, H., Guo, Y., Ma, W., Zhan, J., Hua, C., Li, Y., Deng, C., Li, Y., Zheng, F., Chen, X., Li, Q., Zhang, G., Mahajan, A. S., Cuevas, C. A., Huang, D. D., Wang, Z., Sun, Y., Saiz-Lopez, A., Bianchi, F., Kerminen, V.-M., Worsnop, D. R., Donahue, N. M., Jiang, J., Liu, Y., Ding, A., and Kulmala, M.: Increasing contribution of nighttime nitrogen chemistry to wintertime haze formation in Beijing observed during COVID-19 lockdowns, *Nature Geoscience*, <https://doi.org/10.1038/s41561-023-01285-1>, 2023.
- Yan, Y., Wang, S., Zhu, J., Guo, Y., Tang, G., Liu, B., An, X., Wang, Y., and Zhou, B.: Vertically increased NO_3 radical in the nocturnal boundary layer, *Science of The Total Environment*, 763, 142969, <https://doi.org/https://doi.org/10.1016/j.scitotenv.2020.142969>, 2021.
- Yang, X., Wang, Q., Ma, N., Hu, W., Gao, Y., Huang, Z., Zheng, J., Yuan, B., Yang, N., Tao, J., Hong, J., Cheng, Y., and Su, H.: The impact of chlorine chemistry combined with heterogeneous N_2O_5 reactions on air quality in China, *Atmos. Chem. Phys.*, 22, 3743-3762, <https://doi.org/10.5194/acp-22-3743-2022>, 2022.
- Ye, C., Yuan, B., Lin, Y., Wang, Z., Hu, W., Li, T., Chen, W., Wu, C., Wang, C., Huang, S., Qi, J., Wang, B., Wang, C., Song, W., Wang, X., Zheng, E., Krechmer, J. E., Ye, P., Zhang, Z., Wang, X., Worsnop, D. R., and Shao, M.: Chemical characterization of oxygenated organic compounds in the gas phase and particle phase using iodide CIMS with FIGAERO in urban air, *Atmospheric Chemistry and Physics*, 21, 8455-8478, <https://doi.org/10.5194/acp-21-8455-2021>, 2021.
- Young, C. J., Washenfelder, R. A., Roberts, J. M., Mielke, L. H., Osthoff, H. D., Tsai, C., Pikelnaya, O., Stutz, J., Veres, P. R., Cochran, A. K., VandenBoer, T. C., Flynn, J., Grossberg, N., Haman, C. L.,



904 Lefer, B., Stark, H., Graus, M., de Gouw, J., Gilman, J. B., Kuster, W. C., and Brown, S. S.: Vertically
905 Resolved Measurements of Nighttime Radical Reservoirs in Los Angeles and Their Contribution to
906 the Urban Radical Budget, *Environmental Science & Technology*, 46, 10965-10973,
907 <https://doi.org/10.1021/es302206a>, 2012.
908 Zhou, W., Zhao, J., Ouyang, B., Mehra, A., Xu, W., Wang, Y., Bannan, T. J., Worrall, S. D., Priestley,
909 M., Bacak, A., Chen, Q., Xie, C., Wang, Q., Wang, J., Du, W., Zhang, Y., Ge, X., Ye, P., Lee, J. D., Fu,
910 P., Wang, Z., Worsnop, D., Jones, R., Percival, C. J., Coe, H., and Sun, Y.: Production of N₂O₅ and
911 ClNO₂ in summer in urban Beijing, China, *Atmospheric Chemistry and Physics*, 18, 11581-11597,
912 <https://doi.org/10.5194/acp-18-11581-2018>, 2018.

913