

1 Synoptic drivers of the August 2024 record-breaking rainfall in the Chadian
2 Sahara: Dynamics, thermodynamics, and socio-economic consequences

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37 Abstract

38 This study examines the synoptic atmospheric mechanisms behind the extreme rainfall
39 event of August 2024 in northern Chad and their devastating socio-economic impacts. Based on
40 the CHIRPS, TAMSAT, and ERA5 datasets (1983–2024), Theil-Sen and Mann-Kendall analyses
41 reveal significant increases, over several decades, in precipitation and integrated water vapor
42 transport. Crucially, a 2003 climatological transition threshold shifted the regional baseline from
43 persistent late-20th-century deficits into a moister "new normal," ultimately culminating in the
44 record-breaking August 2024 anomaly.

45 During this month, unprecedented accumulations reached 100–400 mm over the Tibesti, Borkou,
46 and Ennedi massifs, representing parametric return periods of 124.5 years in CHIRPS and 451.4
47 years in TAMSAT, with standardized anomalies exceeding $+8\sigma$ (locally $>+12\sigma$). Our analysis
48 of lower-tropospheric convergence, specific humidity, vertical velocity (ω), and moist static
49 energy (MSE) reveals a major shift from the typical West African monsoon regime, validated via
50 a 1000-iteration grid-point Monte Carlo permutation test ($p \leq 0.05$). In August 2024, the
51 Intertropical Front (ITF) defined by the 15°C dew point isodrosotherm at 925 hPa shifted
52 abnormally northward, reaching 20°N – 22°N , which allowed moist air to penetrate deep into the
53 Saharan zone. This shift was driven by strengthened convergence at 850 hPa, anomalous
54 cyclonic vorticity, and a significant increase in low-level humidity, where specific humidity and
55 Total Column Water Vapor (TCWV) reached their absolute historical records (97.67th empirical
56 percentile). Furthermore, negative ω anomalies throughout the troposphere indicate a northward
57 extension of the monsoon's upward branch, creating a deep vertical ascending chimney. Strong
58 positive MSE anomalies and equivalent potential temperature exceeding the strict 2σ statistical
59 threshold over desert regions further highlight a thermodynamic enrichment of the atmospheric
60 column. Together, these signals point to a highly effective dynamic-thermodynamic coupling that
61 fueled intense convective systems. The synchronization between these atmospheric conditions
62 and the synoptic forcing of African Easterly Waves (AEWs) between August 10 and 14
63 generated extreme local rainfall anomalies, redefining the hydrological balance between aquifer
64 recharge and increased risks of flash flooding.

65 This hydroclimatic shift caused devastating impacts: nationwide, floods affected 2 million people
66 and caused 576 fatalities. Regionally, impacts were severe and heterogeneous, affecting 67,613
67 individuals in Borkou with 23,174 houses destroyed, and claiming 60 lives in Tibesti. Beyond
68 livestock, crop, and gold-mining infrastructure losses, the floods unearthed thousands of
69 conflict-era landmines, creating a grave secondary hazard. These findings underscore hyper-arid
70 zones' vulnerability to extreme hydrological events, highlighting the urgent need to transition
71 from reactive disaster relief to proactive climate adaptation.

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73 Keywords: Extreme rainfall; Chadian Sahara; flooding; Intertropical Front; Moist Static Energy;
74 west African monsoon; Socio-economic impacts

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1. Introduction

Extreme precipitation episodes constitute one of the most critical manifestations of contemporary climate variability, posing a growing challenge to the resilience of human societies (Seneviratne et al., 2021). Their impact is particularly devastating in arid and semi-arid regions, where structural water scarcity renders natural and socio-economic systems extremely vulnerable to hydrometeorological anomalies (Saha et al., 2020). Under global warming, the intensification of hydrological cycle has led to a documented increase in the frequency and magnitude of these events, often exceeding local adaptive capacities (Allan and Soden, 2008; Trenberth et al., 2015; Ehtasham et al., 2024; Adeyeri, 2025). Specifically, surpassing critical warming thresholds could increase intense precipitation in Central Africa by up to 25%, exacerbating flooding risks and population exposure (Zakariahou et al., 2024). This vulnerability is especially concerning for the Lake Chad Basin, where projections indicate a resurgence of rainfall extremes alongside an extension of dry spells (Adeyeri et al., 2019).

The Sahara and its southern margins operate under a climatic regime traditionally governed by tropospheric subsidence and erratic rainfall. In this environment, intense rain events arise from exceptional atmospheric configurations linked to the dynamics of the Saharan Heat Low (SHL) and regional moisture convergence (Selami et al., 2021). This dynamics is part of the rainfall recovery observed since the end of the 20th century, driven by oceanic warming and a strengthened monsoon moisture budget (Sindikubwabo et al., 2018; Biasutti, 2019). However, paleoclimatic records serve as a reminder that this system is prone to non-linear and abrupt responses; the history of "Lake Mega-Chad" demonstrates that transitions between aridity and "African Humid Periods" can be sudden (Simon et al., 2015; Pausata et al., 2017; Pausata et al., 2020).

Northern Chad, a confluence point between the central Sahara and the Sahel, represents a highly sensitive climatic transition zone where the balance between water resources and agro-pastoral activities is precarious (Raimond et al., 2014; Mahamat Nour et al., 2017). Rainfall distribution is strictly governed by the latitudinal position of the Intertropical Front (ITF), the surface expression of the Intertropical Convergence Zone (ITCZ) marking the boundary between the dry harmattan and the moist monsoon (Nicholson, 2013). An anomalous northward migration of the ITF allows for the injection of Atlantic moisture into the heart of the Sahara, triggering deep convection (Sultan and Janicot, 2003; Parker et al., 2005). This dynamics is further modulated by intra-seasonal oscillations (ISO), which influence water vapor transport and moist static energy (MSE) fluxes (Siwe et al., 2025; Wamba et al., 2023) and is often exacerbated by synoptic perturbations, such as African Easterly Waves (AEW) (Kiladis et al., 2006; Lafore et al., 2017).

Beyond dynamics, thermodynamic and radiative factors are decisive. MSE offers a diagnostic framework integrating temperature, humidity, and the vertical structure of the atmosphere (Neelin and Held, 1987; Holloway and Neelin, 2009). High MSE values indicate a warmer and moister boundary layer, which directly enhances the convective available potential energy while

116 simultaneously reducing convective inhibition, facilitating the organization of mesoscale
117 convective systems responsible for extreme accumulations (Romps, 2015). Recent studies
118 demonstrate that these events result from a synergy between thermodynamic instability, radiative
119 flux anomalies, and dynamic forcings such as moist enthalpy advection and potential vorticity
120 anomalies (Kenfack et al., 2024, 2025; Tchana et al., 2025). Simultaneously, vertical velocity
121 characterizes the strength of updraft, where negative ω anomalies represent the classic signature
122 of organized convection in West Africa (Taylor et al., 2017).

123 At the global scale, 2024 was characterized by unprecedented warming and a surge in
124 hydrometeorological extremes (Zhang et al., 2025). In August 2024, exceptionally intense
125 rainfall hit the Saharan regions of Chad and the Lake Chad Basin, leading to catastrophic
126 socio-economic impacts, including unusual flooding, infrastructure collapse, agricultural losses,
127 and widespread population displacement (OCHA, 2024). This crisis highlighted the acute
128 vulnerability of local socio-ecosystems, which are historically adapted to a hyper-arid regime
129 with low mean annual rainfall. In such environments, even a minor absolute increase in
130 precipitation represents a severe hydrometeorological hazard, triggering disproportionate societal
131 disruptions due to the lack of infrastructure designed for torrential rain. While the 2024 anomaly
132 aligns with a broader Sahelian trend associated with an anomalously northward ITCZ (Zhang et
133 al., 2025; Nicholson, 2013), the specific atmospheric mechanisms driving rainfall in the true
134 Saharan desert remain insufficiently documented. Most existing literature focuses on the
135 Sahelian and Sudanian belts, leaving a gap in our understanding of the joint dynamic and
136 thermodynamic processes and their socio-economic consequences within the hyper-arid Sahara
137 and the Lake Chad Basin.

138 This study aims to fill this gap by characterizing the forcings behind the August 2024
139 extreme event through an integrated analysis of moisture convergence, remote moisture
140 advection, vertical velocity, and MSE. By linking these physical mechanisms to field
141 observations of socio-economic impacts, this work provides critical insights for climate
142 adaptation and risk management in the Sahara and Lake Chad Basin (Biasutti, 2019; Vizy and
143 Cook, 2022). The article is structured as follows: Section 2 details the study area, data and
144 methodology; Section 3.1 characterizes the extreme rainfall event and its spatial-temporal
145 evolution, while the rest of Section 3 explores the underlying dynamic and thermodynamic
146 physical mechanisms; Section 4 details the associated socio-economic impacts and local
147 vulnerabilities; and Section 5 concludes the study.

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149 2. Study Area, Data and Methods

150 2.1. Study Area

151 This study focuses on northern Chad, extending 16°-24°N, 13°-25°E, an area
152 administratively comprising the regions of Tibesti, Borkou, Ennedi West, and Ennedi East
153 (Figure 1). The landscape is dominated by vast sandy plains and longitudinal dune systems. The
154 regional climate is hyper-arid, characterized by high diurnal and seasonal thermal ranges and a
155 mean annual precipitation below 50 mm (UNESCO, 1977; Nicholson, 2011). Despite its harsh

environment, the Saharan sector of Chad accounts for approximately 13% of the national population (approx. 2.73 million inhabitants; INSEED-Chad). Economic subsistence is primarily linked to oasis-based agriculture, specifically date palm cultivation, and nomadic pastoralism centered on camel herding. The primary urban centers include Faya-Largeau (pop. 100,000), Fada (pop. 50,000), Bardaï (pop 30,000) and Amdjarass (pop. 2,000). These cities, like all inhabited localities, are established in oases to take advantage of the meager water resources available, given the hyper-arid environment.

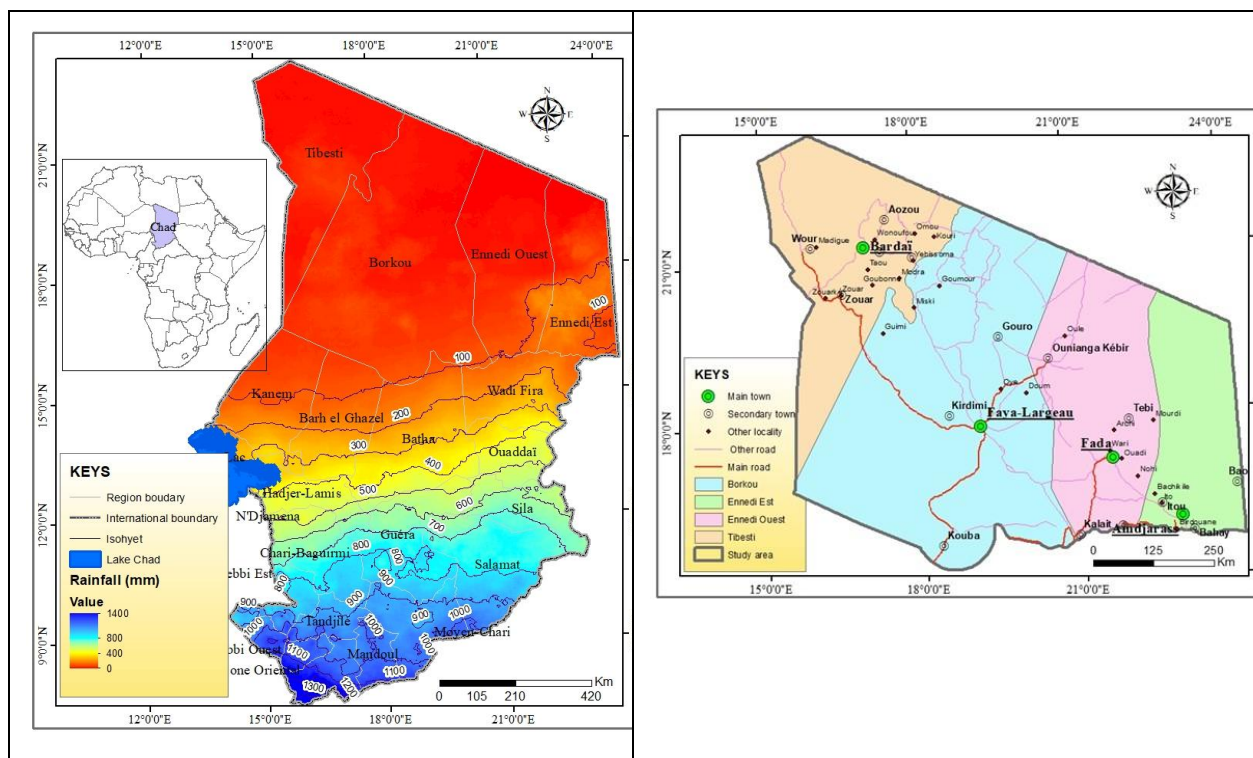


Figure 1. Study area in Chad. Caption A shows that this predominantly desert area is bordered to the south by the 100 mm isohyet. Caption B presents the four administrative provinces whose socioeconomic losses from the August 2024 floods were analyzed. It is worthnoting that despite the desert environment, numerous settlements are located there, generally in oases.

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2.2. Data sources

2.2.1. Satellite Precipitation Estimates

To analyze the long-term interannual variability, monotonic trends, and the specific magnitude of the August 2024 rainfall event, monthly precipitation datasets spanning the 1983-2024 period were gathered from two distinct high-resolution satellite-based products: the Tropical Applications of Meteorology using SATellite Data v.3.1 (TAMSAT; Maidment et al., 2017) and the Climate Hazards Group InfraRed Precipitation with Station data v.2.0 (CHIRPS; Funk et al., 2015).

177 A critical constraint in the hyper-arid Borkou-Ennedi-Tibesti (BET) region of northern Chad is
178 the severe scarcity or total absence of operational, continuous in-situ rain gauges. Consequently,
179 hydro-climatic monitoring relies exclusively on satellite remote sensing. CHIRPS blends global
180 satellite infrared imagery ($0.05^\circ \times 0.05^\circ$ resolution) with available station observations to
181 generate long-term gridded time series, which is highly effective for drought and trend
182 monitoring (Funk et al., 2015). The dataset was accessed from the Climate Hazards Group
183 (<https://www.chc.ucsb.edu/data>). Concurrently, TAMSAT estimations are derived from Meteosat
184 Thermal Infrared (TIR) imagery, using a Cold Cloud Duration (CCD) threshold approach
185 calibrated against regional ground-based gauges (Maidment et al., 2017). TAMSAT features a
186 high spatial resolution of 0.0375° (~ 4 km), which is optimized for capturing convective rainfall
187 distributions across complex Saharan topography (Wainwright et al., 2021). Relying on both
188 products is necessary to address structural uncertainties inherent to satellite retrieval algorithms
189 over hyper-arid terrains where direct gauge validation is unavailable.

190

191 2.2.2. Atmospheric Reanalysis (ERA5)

192 Atmospheric dynamics, vertical coupling, and moisture transport patterns were analyzed
193 using the ERA5 reanalysis, the fifth-generation global atmospheric climate dataset developed by
194 the European Centre for Medium-Range Weather Forecasts (ECMWF) (Hersbach et al., 2020).
195 ERA5 employs a multi-decadal advanced data assimilation scheme, ensuring a physically
196 consistent co-dependence between dynamic and thermodynamic variables across data-sparse
197 regions such as the Sahara desert (Tarek et al., 2020; Ssenyunzi et al., 2020). Previous climate
198 assessments have confirmed that ERA5 accurately reproduces monsoon circulation, moisture
199 influxes, and synoptic-scale waves over the Sahelian and Saharan belts (Johannsen et al., 2019;
200 Cook and Vizzy, 2021).

201 For this study, data were extracted at a horizontal resolution of $0.25^\circ \times 0.25^\circ$ on standard
202 pressure levels spanning from the lower troposphere to the upper troposphere (850 hPa to 200
203 hPa). The monthly variables retrieved include the horizontal components of wind (u , v), vertical
204 velocity (ω), geopotential height (z), temperature (T), and specific humidity (q). It should be
205 noted that, to avoid mathematical edge effects and preserve physical consistency with the
206 ECMWF modeling framework, the lower-tropospheric wind divergence (d) fields were extracted
207 directly from the native outputs of the ECMWF spectral model archives rather than computed via
208 post-processing finite-difference schemes. The analysis focuses on the month of August from
209 1983 to 2024. While the total 42-year period is utilized for interannual trend estimations, the
210 1983-2023 window ($N=41$ years) is strictly designated as the standard climatological baseline
211 mean from which the exceptional structural anomalies of August 2024 are evaluated.

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213 2.2.3. Data on flood losses

214 Due to Chad's fragile socioeconomic and security context, systematic disaster recording
215 is primarily facilitated by international humanitarian organizations and United Nations agencies
216 in collaboration with the Chadian government. For the 2024 flood event, impact data were

217 obtained from the United Nations Office for the Coordination of Humanitarian Affairs (OCHA),
218 which synthesizes inter-agency field reports. These data are validated by the Ministry of
219 Territorial Administration and Decentralization, alongside the Ministry of Social Action,
220 Solidarity, and Humanitarian Affairs.

221 The dataset includes standardized metrics on:

- 222 • Human casualties (mortality and injury rates);
- 223 • Displaced populations and affected households;
- 224 • Agricultural losses (hectares of devastated cropland);
- 225 • Infrastructure damage (destruction of homes and commercial facilities);
- 226 • Livestock mortality.

227 While national summaries typically emphasize the southern and central provinces due to their
228 higher absolute population densities (e.g., Mayo-Kebbi Est, Lac, and Tandjilé), this study
229 specifically isolates data for the northern Saharan provinces. Although the absolute number of
230 victims in these arid regions is lower compared to the south, the relative impact, defined as the
231 ratio of affected individuals to the total provincial population, is among the highest recorded in
232 2024. This relative approach is essential for assessing the exceptional nature of climate extremes
233 in sparsely populated hyper-arid zones. Data was accessed via the OCHA Chad humanitarian
234 portal (<https://www.unocha.org/chad>).

235 2.3. Methodological Framework and Statistical Metrics

236 To diagnose the physical mechanisms driving the exceptional rainfall of August 2024, a
237 suite of diagnostic variables was employed. The methodological framework distinguishes
238 between thermodynamic precursors (the energy and moisture required for convection) and
239 dynamic forcing (the atmospheric mechanisms triggering vertical motion).

240 2.3.1. Climatological Anomalies and Spatial Significance

241 To highlight the physical anomalies driving the extreme event of August 2024, monthly
242 climatological anomalies at each grid point were computed as the direct departure of the
243 observed August 2024 values from the long-term historical mean, which was established over a
244 strict 41-year reference baseline period spanning from 1983 to 2023.

245 To assess whether these spatial anomalies represent a statistically significant departure from the
246 background noise or are merely a product of random internal climate variability, a
247 non-parametric Monte Carlo permutation test was systematically applied to the anomaly fields.
248 This approach is highly suitable for atmospheric dynamics over hyper-arid zones because it does
249 not rely on parametric assumptions of normal probability distributions (Livezey and Chen,
250 1983). For each variable, the historical time series (1983 - 2024) was randomly shuffled 1000
251 times to empirically reconstruct a local null distribution of climatological anomalies at each grid
252 point. The true observed August 2024 anomaly was then positioned within this empirical

distribution. Anomalies falling outside the 2.5th and 97.5th percentiles of the permuted distribution ($p \leq 0.05$, two-tailed test) were diagnosed as statistically significant at the 95% confidence level. These structurally robust grid points are explicitly highlighted via stippling on the corresponding spatial maps.

2.3.2. Long-Term Trend Analysis

The long-term monotonic trajectories of regional precipitation and column-integrated moisture transport over the 42 years (1983-2024) were quantified using the non-parametric Theil-Sen estimator, while trend significance was evaluated using the Mann-Kendall test. The Sen's slope represents the median of all slopes calculated between all possible pairs of data points in the time series, ensuring structural immunity to extreme outliers such as the August 2024 anomaly itself (Sen, 1968). To contextualize the physical magnitude of these trends, the relative change (Δ_{rel} , expressed in %) was calculated by multiplying the Sen's slope by the total length of the time series ($\Delta_t = 42$ years), normalized by the long-term climatological mean ($\bar{X}$) obtained from the 1983 - 2023 baseline:

$$\Delta_{rel} = \left(\frac{\text{Sen's Slope} \times \Delta t}{\bar{X}} \right) \times 100 \quad (1)$$

2.3.3. Extreme hydro-climatic quantified diagnostics

To quantify the exact recurrence interval and rarity of the August 2024 precipitation, a parametric framework was implemented. Because precipitation over Saharan environments is highly intermittent, strictly non-negative, and positively skewed, standard Gaussian frameworks are physically invalid. Instead, a continuous two-parameter Gamma probability density function ($f(x)$) was fitted to the multi-decadal historical August series:

$$f(x, \alpha, \beta) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-\frac{x}{\beta}} \text{ for } x > 0 \quad (2)$$

Where $\alpha > 0$ is the shape parameter, $\beta > 0$ is the scale parameter, and $\Gamma(\alpha)$ represents the standard Gamma function. These parameters were optimized via maximum likelihood estimation (MLE), providing asymptotically unbiased and numerically robust estimators for meteorological series (Husak et al., 2007). Once the cumulative distribution function ($F(x)$) was solved, the return period (T , in years) for the August 2024 rainfall magnitude (x_{2024}) was derived as the inverse of its exceedance probability:

$$T = \frac{1}{1 - F(x_{2024})} \quad (3)$$

To evaluate the severity of the low-tropospheric moisture anomalies without assuming a parametric background distribution, the historical specific humidity and total column water vapor

values were ranked against the 42-year record using the non-parametric Weibull plotting position formula:

$$P = \frac{m}{N_{total}+1} \quad (4)$$

where “m” represents the ascending rank of the observed value (with m = 42 for the absolute historical maximum) and $N_{total} = 42$ is the total number of years in the 1983 - 2024 series. The corresponding percentile rank is expressed as $PR = P \times 100$, allowing an objective determination of the rarity of the thermodynamic preconditioning. Following Makkonen (2006), this specific formulation represents the only mathematically consistent plotting rule for extreme value analysis, preventing the overestimation of return periods often induced by alternative empirical plotting positions and providing a realistic assessment of weather-related risks in hyper-arid environments.

2.4. Atmospheric Circulation and Thermodynamic Indicators

2.4.1. Intertropical Front (ITF) Position

The ITF serves as a critical diagnostic for monitoring the northward penetration of the West African Monsoon (WAM) into the Saharan heat low. In this study, the latitudinal position of the ITF is identified using the 15°C isodrosotherm (dew point temperature $T_d=15^\circ\text{C}$) at the 925 hPa level, following the criteria established by Nicholson (2013). This boundary marks the interface between the moist south-westerly monsoon flow and the dry north-easterly Harmattan winds.

2.4.2. Integrated Vapor Transport (IVT)

To capture the horizontal moisture streams fueling the convective anomalies over northern Chad, the time series of monthly mean interannual vertically Integrated Vapor Transport (IVT) magnitude was computed. Originally conceptualized to distinguish concentrated filamentary moisture fluxes from broader background fields (Zhu and Newell, 1998), the analysis of IVT magnitudes has become a robust standard for diagnosing synoptic-scale moisture transport and associated hydrometeorological extremes (Gimeno et al., 2014). The zonal (IVT_u) and meridional (IVT_v) mass-weighted components ($\text{Kg.m}^{-1}.\text{s}^{-1}$) were obtained by integrating the product of the specific humidity (q) and horizontal wind components (u, v) across the active tropospheric column (Lavers and Ramos, 2019):

$$IVT_u = \frac{1}{g} \int_{P_{top}}^{P_{surf}} u \cdot q \cdot dp \quad ; \quad IVT_v = \frac{1}{g} \int_{P_{top}}^{P_{surf}} v \cdot q \cdot dp \quad (5)$$

where g is the standard gravitational acceleration (9.80665m.s^{-2}), q is the specific humidity (Kg.Kg^{-1}), u and v are the horizontal wind velocities (m.s^{-1}), and dp is the pressure increment (Pa). Vertical integration bounds were restricted from $p_{\text{surf}} = 1000$ hPa to $p_{\text{top}} = 200$ hPa. This configuration captures the full structural depth of the monsoon circulation, including the coupled lower-tropospheric moisture surges preconditioned by the Saharan heat low, the horizontal shear zones along the African Easterly Jet (AEJ), and upper-level transitions near the Tropical Easterly Jet (TEJ) (Gimeno et al., 2016; Awolaye et al., 2025). The total IVT magnitude (IVT_{mag}) was then derived as:

$$\text{IVT}_{\text{mag}} = \sqrt{\text{IVT}_u^2 + \text{IVT}_v^2} \quad (6)$$

In the context of the African monsoon belt, evaluating IVTmag is critical as vertically coupled moisture fluxes exceeding specific synoptic thresholds have been directly linked to enhanced organized convection, intense spatially focused precipitation events, and regional aerosol scavenging dynamics (Awolaye et al., 2025).

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2.4.3. Low-Tropospheric Convergence and Vertical Velocity

The dynamic mechanisms forcing local vertical ascent were diagnosed using two structural fields: low-tropospheric horizontal wind convergence and tropospheric vertical velocity (ω). Wind convergence ($-\nabla \cdot \mathbf{V}$) was extracted directly at the 850 hPa level from the native ECMWF spectral model divergence archive (d) to prevent numerical truncation errors arising from post-processing finite-difference schemes. Concurrently, the vertical velocity (ω , expressed in Pa.s^{-1}) was examined across multiple pressure levels to assess the deep vertical coupling of the column. Negative values ($\omega < 0$) represent strong updrafts and synoptic-scale convective ascent (Taylor et al., 2017; Tchana et al., 2025).

341

2.4.4. Thermodynamic Columns and Convective Instability

To evaluate the thermodynamic structure and convective potential governing the anomalous August 2024 event, we analyze two complementary conserved thermodynamic tracers: Moist Static Energy (MSE) and equivalent potential temperature (θ_e). Moist Static Energy (MSE) is used to evaluate the convective potential and stability of the atmosphere (Neelin and Held, 1987; Kenfack et al., 2025). It integrates sensible heat, potential energy, and latent heat. It also quantifies the total energy content of an air parcel by integrating sensible heat, potential energy, and latent heat, and is defined as:

$$\text{MSE} = C_p T + gz + L_v q \quad (7)$$

Where:

C_p is the specific heat of dry air at constant pressure ($1004 \text{ J Kg}^{-1} \text{ K}^{-1}$); T is the absolute temperature (K), g is the gravitational acceleration (9.81 m s^{-1}), L_v is the latent heat of

vaporization, z is the geopotential height (m), and q is the specific humidity (Kg Kg^{-1}). By examining MSE anomalies, we can establish if the August 2024 extremes were driven by an unusual amount of moisture (latent heat) or intense localised surface heating (sensible heat).

In parallel, the equivalent potential temperature θ_e (K), which dictates the absolute temperature a water-vapor-saturated air parcel would attain if all its moisture condensed pseudo-adiabatically, was calculated following Bolton's (1980) formulation:

$$\theta_e = \theta \cdot \exp\left(\frac{L_v q}{C_p T_L}\right) \quad (8)$$

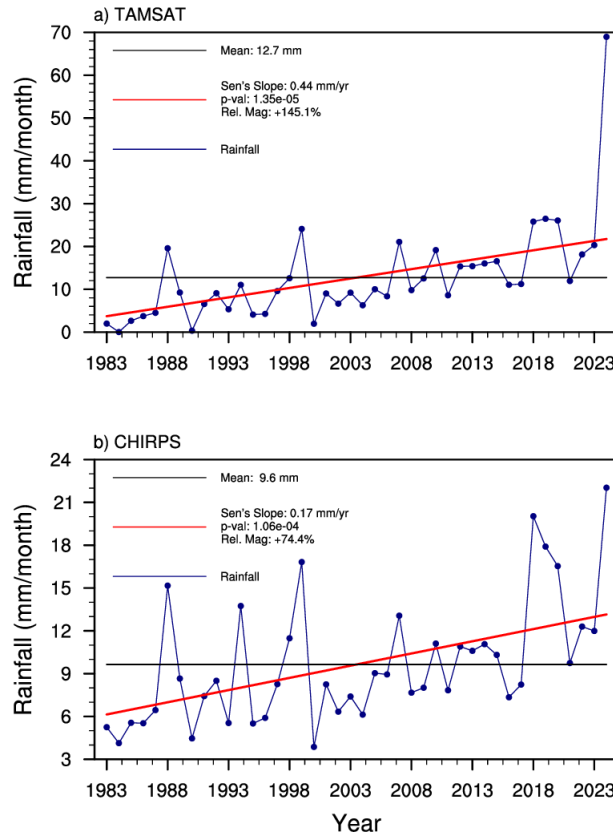
Where θ is the dry potential temperature and T_L is the temperature at the lifted condensation level. In the tropical troposphere, MSE and $C_p \theta_e$ are tightly coupled and approximately linearly related ($\text{MSE} \approx C_p \theta_e$), as both conserve total energy during moist moist-adiabatic processes (Neelin and Held, 1987).

From a diagnostic standpoint, an increase in low-level MSE (and thus θ_e) operates via a critical dual thermodynamic mechanism: it enhances the thermal buoyancy of a lifted air parcel, thereby expanding the convective available potential energy aloft, while simultaneously eroding the convective inhibition barrier at the base of the atmospheric profile, facilitating spontaneous and deep convective triggering.

3. Results

3.1. Multi-Decadal Rainfall Trends, IVT Evolution, and the Hydro-Climatic Transition

The interannual variation in August rainfall over the northern Chadian Sahara ($16^\circ\text{--}24^\circ\text{N}$, $13^\circ\text{--}25^\circ\text{E}$) characterizes a hyper-arid hydro-climatic regime defined by extreme scarcity and inherent high interannual variability. Historically, monthly totals over this region hover around a very low climatological baseline, with a long-term mean of 12.7 mm for TAMSAT and 9.6 mm for CHIRPS. This hyper-arid dynamic is intrinsically governed by the Saharo-Sahelian transition zone, where precipitation depends strictly on the erratic northward excursions of the Intertropical Front (ITF) and the seasonal migration of the West African Monsoon (WAM) (Nicholson et al., 2013; 2018). An analysis of the 42-year time series (1983 - 2024), comparing the TAMSAT (Fig. 2a) and CHIRPS (Fig. 2b) satellite products, reveals a remarkable structural consistency in capturing regional climate variability. Both independent datasets accurately identify historical dry cycles such as the severe drought years of 1984, 1990, and 2000, as well as past wet episodes like 1988, 1999, and the 2018-2020 corridor. Although a systematic magnitude offset exists between the two products due to their distinct retrieval algorithms, their long-term climate trajectories are perfectly synchronized.



389

390 Figure 2: August precipitation trends and patterns (solid blue line) averaged over Northern Chad
 391 (16° - 24°N, 13° - 25°E) from 1983 to 2024. Panels display (a) TAMSAT and (b) CHIRPS
 392 satellite series. The solid red line represents the monotonic linear Sen's slope; the solid black line
 393 marks the long-term climatological mean.

394

395 Both datasets exhibit a robust, continuous, and statistically significant monotonic upward trend
 396 over the entire 1983 - 2024 period, with exceptionally high confidence levels ($p = 1.35 \times 10^{-5}$ for
 397 TAMSAT and $p = 1.06 \times 10^{-4}$ for CHIRPS). Rather than an abrupt structural regime shift or a
 398 statistical break in the trend line itself, the year 2003 marks a critical climatological transition
 399 threshold. This year represents the exact chronological midpoint where the long-term linear
 400 Sen's slope intersects the historical baseline mean. Before 2003, the regional signal was heavily
 401 weighted by the extreme aridity and persistent rainfall deficits of the late 20th-century drought,
 402 with the vast majority of years falling well below the long-term average. After 2003, due to a
 403 sustained humidification trend, a profound transition occurred. The frequency of extreme dry
 404 years drops sharply, replaced by a dense clustering of positive anomalies and closely spaced
 405 rainfall peaks (notably post-2017), establishing a moister "new normal" background state over
 406 the Chadian Sahara.

407 This multi-decadal evolution is quantified by a positive Sen's slope of +0.44 mm/year for
 408 TAMSAT and +0.17 mm/year for CHIRPS, corresponding to substantial relative magnitudes of
 409 +145.1% and +74.4% over the 42-year record, respectively. This long-term wetting trend over

Northern Chad aligns with the wider "greening" and hydrometeorological intensification signals documented across the central Sahel (Brandt et al., 2015).

To establish a clear physical connection between this long-term precipitation trend and large-scale atmospheric drivers, a multi-decadal time series of the Integrated Vapor Transport (IVT) during August was analyzed over the same domain (Fig. 3). As an integrated indicator of lower-tropospheric moisture supply and wind convergence, the IVT exhibits an exceptional, statistically significant co-variability with the satellite rainfall series.

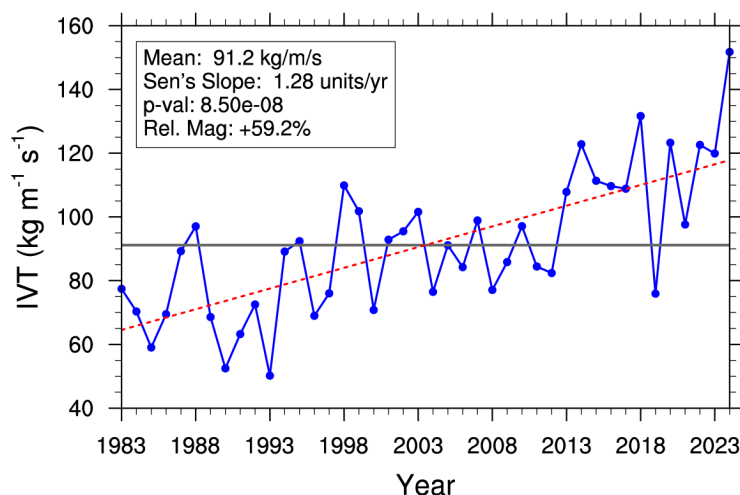


Figure 3: Same as in fig 2, but for Integrated Vapor Transport (IVT, $\text{kg m}^{-1} \text{s}^{-1}$). The dashed red line represents the Sen's slope; the solid black line marks the historical mean ($91.2 \text{ kg m}^{-1} \text{s}^{-1}$).

The IVT time series confirms a robust atmospheric moisture-loading trend, featuring a highly significant Sen's slope of $+1.28 \text{ units/year}$ ($p = 8.50 \times 10^{-8}$) and a relative magnitude increase of $+59.2\%$. More specifically, the IVT reproduces exactly the same physical behavior as the precipitation data:

1. Late 20th-century suppressed state: Persistent negative IVT anomalies dominate the 1980s and 1990s, indicating a severely restricted moisture supply to the region.
2. The 2003 Co-Dependent Transition: The IVT line permanently crosses its historical climatological mean ($91.2 \text{ kg m}^{-1} \text{s}^{-1}$) around 2003, transitioning the regional atmosphere into a consistently heightened moisture-loading baseline.
3. The 2024 Extreme Manifestation: The record-breaking event of August 2024 sits at the absolute tail end of this multi-decadal trend, where the IVT reaches an unprecedented historical peak exceeding $150 \text{ kg m}^{-1} \text{s}^{-1}$.

The record-breaking precipitation anomaly of August 2024, which reached an unprecedented regional average of 69.71 mm/month in TAMSAT and 22.35 mm/month in CHIRPS, represents the absolute extreme manifestation of this long-term hydro-climatic transition. Rather than occurring as an isolated, random synoptic anomaly, the exceptional magnitude of the 2024 event was fundamentally favored by this altered, post-2003 moister background state. The long-term intensification of the atmospheric moisture transport background state effectively maximized the

439 regional susceptibility to extreme monsoonal rainfall. This setup allowed contemporary dynamic
440 and thermodynamic forcings to trigger an unprecedented northward penetration of the WAM,
441 shifting the regional risk paradigm from perpetual hyper-aridity to acute flash flood hazards.
442 This multi-decadal shift toward a structurally moister baseline suggests that the Chadian Sahara
443 has transitioned into a hydroclimatic regime that is fundamentally more susceptible to intense
444 convective forcing. Within this altered background state, the exceptional event of August 2024
445 does not emerge as an isolated, uncoupled anomaly but rather as the absolute tail-end expression
446 of a long-term intensification in regional moisture loading. To fully understand how this
447 large-scale climatological transition manifests on a local scale, it is necessary to move beyond
448 domain-averaged time series and dissect the explicit spatial architecture, structural magnitude,
449 and statistical return periods of the August 2024 anomalies across the hyper-arid core.

450

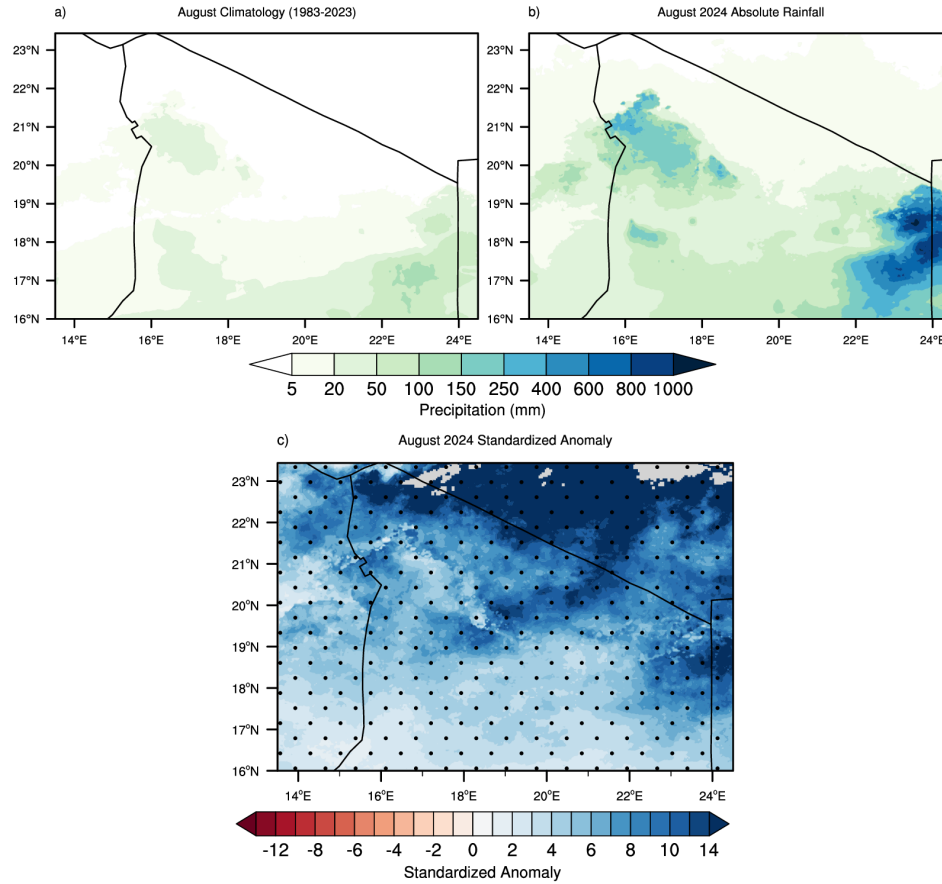
451 3.2. Magnitude and Statistical Significance of the August 2024 Rainfall Anomalies

452

453 In response to the long-term hydroclimatic transition and the regional-scale
454 intensification of Integrated Vapor Transport (IVT), the August 2024 event manifested as a
455 highly intense and spatially heterogeneous deployment, which had been masked until now by
456 domain-averaged metrics. Examining fine-scale spatial structures thus becomes essential to map
457 how this historical anomaly is distributed across the core hyper-arid northern sector of Chad. To
458 evaluate this spatial architecture and the exceptional severity of the August 2024 rainfall within
459 this hyper-arid zone, we examine the distribution of monthly accumulated precipitation alongside
460 probabilistic Z-scores. Figure 4 illustrates the monthly accumulated rainfall and corresponding
461 standardized anomalies over northern Chad, compared against the 41-year historical baseline
462 climatology (1983–2023).

463 Although the preceding regional temporal analysis (Section 3.1) demonstrated a remarkable
464 synchronization between the TAMSAT and CHIRPS datasets, the spatial characterization
465 presented here (Figure 4) relies exclusively on the TAMSAT product. This methodological
466 choice is motivated by TAMSAT’s algorithmic design optimized for tracking African convective
467 systems in arid zones. Furthermore, since the analysis critically depends on the calculation of
468 standardized anomalies (Z-scores), the systematic magnitude offset observed between the two
469 satellite products is mathematically normalized. This ensures that the spatial structure and
470 statistical significance of the 2024 rainfall plume remain robust and independent of the selected
471 product.

472 The baseline climatology (Fig. 4a) highlights the extreme and uniform aridity that typically
473 characterizes the region. Within the central Saharan sectors encompassing the massifs and
474 plateaux of the Tibesti, Borkou, and Ennedi, mean August accumulations are nearly non-existent
475 and remain strictly below 5 mm.



476

477 Figure 4: Spatial distribution, magnitude, and statistical significance of August precipitation over
 478 Northern Chad. Panels display (a) the multi-decadal baseline climatology (1983-2023), (b)
 479 absolute monthly rainfall accumulations for August 2024, and (c) the corresponding standardized
 480 anomalies (Z-score). Stippling in panel (c) denotes statistically significant anomalies ($p \leq 0.05$)
 481 derived from a grid-point-wise Monte Carlo simulation.

482 In contrast to this baseline state, the spatial structure of the August 2024 rainfall (Fig. 4b)
 483 exhibits a remarkable northward thrust of precipitation well beyond the 20°N parallel. This
 484 rainfall activity was organized around two primary, well-defined core nodes:

- 485 ➤ The first developed over the rugged topography of the Ennedi and extended westward
 486 into the Borkou region, generating unprecedented absolute accumulations ranging from
 487 150 to 400 mm.
- 488 ➤ The second manifested further to the northwest, directly impinging on the volcanic
 489 massifs of the Tibesti region, with totals ranging from 100 to 250 mm.

490 The standardized anomaly field (Fig. 4c) provides a mathematical context for the extraordinary
 491 magnitude of this event. Localized departures from the historical mean are exceptional, with
 492 Z-scores systematically exceeding $+8\sigma$, and localized peaks transcending $+12\sigma$ within the
 493 Borkou and Ennedi sectors. Such values underscore the profound statistical deviation of this
 494 event from regional historical observations.

Finally, to ensure that this signal is reliable and does not result from random variability, a grid-point-wise statistical Monte Carlo significance test was applied. The resulting empirically derived significance mask (represented by the stippling in Fig. 4c) completely envelopes the Tibesti, Borkou, and Ennedi regions. The spatial continuity of this mask confirms that the northward advance of the monsoon and the associated hyper-arid precipitation constitute a statistically significant anomaly ($p \leq 0.05$) that is unprecedented since the beginning of the satellite era, indicative of an exceptional displacement of the West African Monsoon (WAM) boundaries during this period.

Table 1: Parametric Return Periods and Non-Parametric Thermodynamic Percentile Ranks for the August 2024 Extreme Event over the Northern Chad study domain (16°N–24°N, 13°E–25°E). Rainfall metrics are based on a Gamma distribution fitted via Maximum Likelihood Estimation (MLE) over the 1983–2024 baseline. Specific humidity (q) and Total Column Water Vapor (TCWV) metrics represent empirical ranks and percentile ranks calculated using the Weibull formula over the same climatological period (N = 42 years).

Diagnostic Category	Parameter / Dataset	Observed Value (August 2024)	Empirical Rank (Out of 42)	Calculated Percentile Rank (%)	Estimated Return Period (T, years)
Pluviometric Metrics	CHIRPS Rainfall	22.35 mm	–	–	124.5
	TAMSAT Rainfall	69.71 mm	–	–	451.4
Thermodynamics Metrics	ERA5 Specific Humidity q at 850 hPa	10.67g.kg ⁻¹	42	97.67%	–
	ERA5 Specific Humidity q at 925 hPa	11.90g.kg ⁻¹	42	97.67%	–
	ERA5 Column Water Vapor TCWV (Total Column)	14.83kg.m ⁻²	42	97.67%	–

While the spatial architecture of Figure 4 illustrates the geographic extent of this event along the Tibesti-Borkou-Ennedi corridor, it raises, above all, the question of its statistical return period

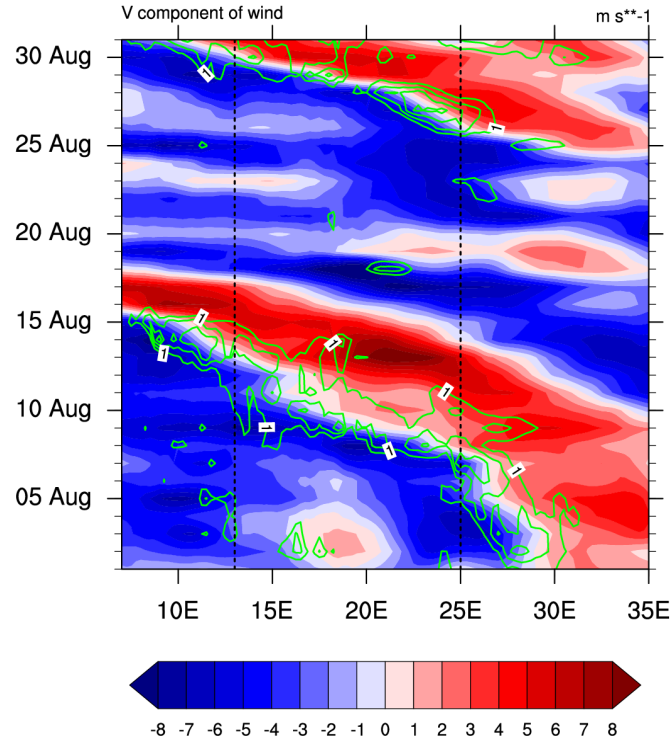
511 and the underlying atmospheric conditions that allowed such anomalies to persist. To place the
 512 extreme nature of this plume within a historical context spanning several decades, we employ a
 513 combined diagnostic framework that links parametric return periods of precipitation to
 514 nonparametric empirical ranks of tropospheric moisture content. To quantify the historical rarity
 515 of this event, a rigorous dual-probabilistic diagnostics framework was applied to the core
 516 convective zone (16°–24°N, 13°–25°E). The results, summarized in Table 1, contrast the
 517 parametric return periods (T) derived from the optimized two-parameter Gamma distribution
 518 against the empirical thermodynamic percentile ranks (PR) computed via the mathematically
 519 consistent Weibull plotting position (Makkonen, 2006). The regionalized August 2024 rainfall
 520 anomaly exhibits a return period of 124.5 years according to CHIRPS (mean: 22.35mm), and
 521 reaches up to 451.4 years according to the TAMSAT satellite product (mean: 69.71mm). The
 522 structural discrepancy between these two satellite estimates reflects well-known algorithmic
 523 sensitivities in hyper-arid, gauge-sparse regions. TAMSAT relies exclusively on Thermal
 524 Infrared (TIR) Cold Cloud Duration (CCD) thresholds calibrated specifically for African
 525 convective regimes (Maidment et al., 2014, 2017). Consequently, it effectively captured the
 526 persistent, deep cloud tops and prolonged convective activity that sustained this specific event,
 527 likely yielding a physically representative depiction of the local convective intensity. Conversely,
 528 while CHIRPS integrates TIR data, its methodology heavily relies on the blending of in-situ rain
 529 gauge observations for structural bias correction (Funk et al., 2015). In the hyper-arid northern
 530 territory of Chad, the severe scarcity of operational, real-time reporting gauges forces the
 531 CHIRPS algorithm to heavily weight its background climatology, typically inducing a
 532 conservative underestimation during unprecedented, high-magnitude anomalies. Despite this
 533 systematic magnitude offset, both independent datasets structurally agree on the centennial to
 534 multi-centennial scale of the anomaly. This out-of-boundary pluviometric response was
 535 systematically fueled by an extraordinary, vertically coherent thermodynamic state throughout
 536 the tropospheric column. The monthly mean specific humidity (q) reached absolute historical
 537 records at both the lower-tropospheric base (11.90g kg⁻¹ at 925 hPa) and the core of the
 538 monsoonal layer (10.67 g kg⁻¹ at 850 hPa). This profound moistening is further corroborated by
 539 the Total Column Water Vapor (TCWV), which attained an unprecedented historical peak of
 540 14.83kg m⁻². For all three hydro-thermodynamic parameters, August 2024 ranked 42nd out of the
 541 42-year climatological record, placing the month at the 97.67th empirical percentile of the
 542 historical distribution (Table 1). This joint analysis confirms that the historic rainfall observed in
 543 the Saharan regions of Chad resulted from a perfect synergy between, on the one hand, a local
 544 moisture reservoir reaching its absolute maximum, thereby preventing the inflow of dry air and
 545 lowering the condensation level due to lifting, and, on the other hand, a dynamically optimized
 546 triggering mechanism at the synoptic scale. However, while these extraordinary thermodynamic
 547 conditions created a highly conducive and moisture-saturated background state, they constitute a
 548 necessary but not sufficient condition for such extreme convective development. To fully
 549 understand how this localized precipitable water was dynamically lifted and organized into
 550 plumes of intense precipitation across the hyper-arid heartland, it is essential to shift our focus

551 from hydrostatic column measurements to the synoptic-scale transient triggers that governed
552 regional atmospheric circulation during August 2024.

553 3.3. Synoptic-Scale Triggering: African Easterly Waves (AEWs) Dynamics

554 To study the synoptic-scale atmospheric mechanisms underlying the exceptional
555 precipitation anomalies observed in August 2024, Figure 5 presents a longitude-time
556 (Hovmöller) diagram of the meridional wind and relative vorticity at 700 hPa, averaged over the
557 Saharan belt (16°N–24°N). During the first half of the month (from August 1 to 16), the regional
558 atmosphere was featured by robust and well-defined activity of the AEWs. This dynamic is
559 visually identifiable by distinct diagonal corridors propagating westward, composed of
560 alternating meridional components of north (blue) and south (red) winds. It is important to note
561 that between August 10 and 14, a high-amplitude wave trough penetrated the specific study area
562 bounded by the meridians 13°E and 25°E (vertical dotted lines). This synoptic event was
563 characterized by a powerful acceleration of the southerly wind exceeding $+4 \text{ m s}^{-1}$, which acted
564 as a dynamic conveyor belt, ensuring significant moisture advection from the monsoon
565 reservoirs in the low latitudes toward the hyper-arid interior. This hydrodynamic driving
566 mechanism was closely linked to a significant core of cyclonic relative vorticity (green contour $\geq$
567 $1 \times 10^{-5} \text{ s}^{-1}$), ensuring the large-scale upwelling and vertical organization of the necessary cells.
568 This intense dynamic forcing on a synoptic scale explains the onset, development, and
569 concentration of the extreme hydrometeorological events observed over the Tibesti, Borkou, and
570 Ennedi mountain ranges during this specific period spanning several days. In contrast, the second
571 half of August 2024 (after August 16) was marked by a notable change in weather conditions.
572 The organized, in-phase, and propagating AEW structures have weakened considerably, as
573 evidenced by the disruption of the diagonal wind bands. This state of suppression gave way to
574 more stationary atmospheric features and localized, isolated vorticity maxima, centered mainly
575 near the eastern boundary (25°E) during the last few days of the month. This evolution indicates
576 that, while the first half of the month was dominated by transient wave-induced convective
577 bursts, the second half shifted toward a more localized convective regime of thermodynamic
578 origin, within the context of stationary synoptic patterns.

579



580

581 Figure 5: Longitude-time (Hovmöller) diagram of 700 hPa meridional wind (shaded, m.s⁻¹) and
 582 relative vorticity (contours, 10⁻⁵ s⁻¹) averaged over the Saharan band (16°N–24°N) during August
 583 2024. Dashed vertical lines delimit the core study domain (13°E–25°E).

584

585 3.2. Atmospheric drivers

586 3.4. Climatological mean and anomalous state of lower-tropospheric (850 hPa) circulation, wind 587 convergence, and Intertropical Front (ITF) positions.

588 The progressive AEW packets analyzed in the previous section are not merely isolated
 589 daily weather events; on the contrary, their recurring occurrence and intensity are likely to
 590 modulate the large-scale monthly atmospheric state. To quantify the net, integrated impact of
 591 these synoptic drivers on the lower troposphere, this section examines the monthly mean
 592 dynamic configuration at 850 hPa. More specifically, Figure 6 displays the spatial distribution of
 593 850 hPa wind convergence, horizontal wind vectors, and the latitudinal position of the
 594 Intertropical Front (ITF), defined by the 15°C dew point isodrosotherm. Three states are
 595 compared: the 1983–2023 August climatology (Fig. 6a), the August 2024 observed state (Fig.
 596 6b), and the resultant absolute anomalies (Fig. 6c).

597 In the climatological mean (Fig. 6a), the solid black contour marking the ITF exhibits a
 598 zonal orientation localized between 15°N and 18°N, while the color shading reveals a
 599 well-defined band of low-level convergence (blue shading; -1 to -3 × 10⁻⁵ s⁻¹) strictly bounded to
 600 the Sahelian belt. South of this boundary, the horizontal wind field is characterized by persistent,
 601 large-scale southwesterly vectors exceeding 5 m s⁻¹. Conversely, inside the red box defining the
 602 northern Chadian study domain (16°N–24°N, 13°E–25°E), the background shading is dominated
 603 by positive divergence values (light yellow to white; +0.5 to +2 × 10⁻⁵ s⁻¹), followed by strong

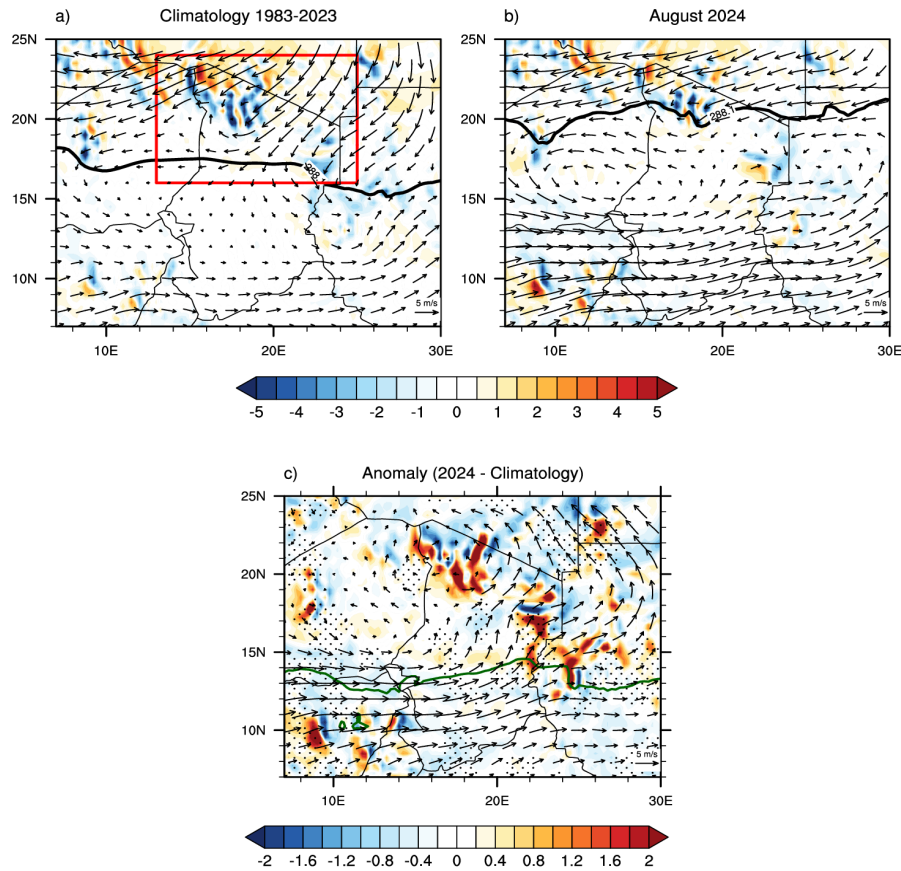
northeasterly wind arrows. Dynamically, this indicates that significant low-level convergence is normally restricted to the humid monsoonal flow over the Sahel. Northern Chad, including the Tibesti, Borkou, and Ennedi massifs, is climatologically dominated by a strongly divergent, subsiding regime associated with the descending branch of the Hadley cell and dry Harmattan winds. These synoptic conditions maintain absolute atmospheric stability and structurally inhibit deep convection, a configuration perfectly aligned with the established canonical structure of the West African Monsoon (WAM) (Sultan & Janicot, 2003; Nicholson, 2013).

In August 2024 (Fig. 6b), the thick black ITF line displays a pronounced northward deformation, looping up to the 20°N–22°N latitudinal band, while the core of negative divergence values (blue patches; -2 to $-4 \times 10^{-5} \text{ s}^{-1}$) undergoes a substantial northward migration into the Saharan interior. Simultaneously, the horizontal wind vectors show a total reorientation within the study domain, where the traditional northeasterlies are replaced by an uncharacteristic influx of southwesterly and southerly arrows penetrating past 20°N. This interior flow was embedded within a broader synoptic configuration, characterized by enhanced cyclonic circulation gyres visible both to the west and east of the study domain. Visually and physically, this confirms a radical breakdown of the typical desert circulation. This deep latitudinal displacement allowed intense southwesterly monsoonal surges to bypass their usual boundaries and directly establish a powerful dynamic convergence engine over the complex terrain of the Tibesti massif, the Borkou depressions, and the Ennedi plateau. This localized convergence provided the forced mechanical ascent necessary to overcome the hyper-arid background stability, shifting a normally stable region into a highly active convective state. This structural configuration closely mirrors the diagnostic frameworks of extreme multi-day rainfall events documented in arid northern Africa, where topographically locked moisture flux convergence drives unprecedented precipitation events (Vondou et al., 2025).

The absolute anomaly field (Fig. 6c) reveals a vast, cohesive cell of enhanced convergence anomalies (blue shading; reaching -0.6 to $-1.6 \times 10^{-5} \text{ s}^{-1}$) spanning the 15°N–22°N domain, overlaid by a prominent cyclonic (counter-clockwise) wind vector circulation centered over the Chad-Niger border. The solid green contour, tracking the +1 K dew point anomaly threshold, extends well north of 15°N and runs spatially in-phase with the maximum convergence anomalies. Physically, this feature indicates that the dynamic lifting engine was perfectly co-located with a highly anomalous tropospheric moisture loading, which effectively suppressed dry entrainment and favored widespread condensation over the Saharan massifs. Mathematically, this cyclonic wind anomaly explicitly proves that horizontal wind convergence was complemented by a substantial increase in low-tropospheric cyclonic vorticity. Such an environment is vital for providing the mesoscale organization and hydrodynamic longevity required to sustain organized and long-lived convective systems in extreme environments (Lavaysse et al., 2009; Flamant et al., 2018). Crucially, the extensive stippling mask confirms that these enhanced convergence anomalies are statistically significant ($p \leq 0.05$) based on the 1000-iteration Monte Carlo permutation test. Because this significant signal is spatially in-phase with the anomalous southwesterly moisture advection, it confirms a total, non-random disruption

644 of the Saharan divergent regime, representing a profound, transient northward expansion of the
 645 Sahelian convective climate boundary into the hyper-arid core of Northern Chad (Taylor et al.,
 646 2017; Biasutti, 2019). In summary, the out-of-boundary rainfall event of August 2024 resulted
 647 from a perfect synergy between the exceptional northward migration of the ITF, reinforced
 648 low-level convergence at 850 hPa, and the establishment of a regional cyclonic circulation that
 649 optimized the vertical lifting of warm, moist air masses over Northern Chad. These findings
 650 strongly align with recent literature linking increased WAM variability to a higher frequency of
 651 hydroclimatic extremes along the Sahara-Sahel transition zone (Janicot et al., 2008; Taylor et al.,
 652 2017; Biasutti, 2019).

653



654

655 **Figure 6: Spatial distribution of lower tropospheric dynamics at 850 hPa in August, showing**
 656 **wind convergence (shading; $\times 10^{-5} \text{ s}^{-1}$), horizontal wind vectors (arrows; m s^{-1}), and tropospheric**
 657 **moisture boundaries.** The panels show: (a) the climatological mean from 1983 to 2023, (b)
 658 conditions observed in August 2024, and (c) absolute anomalies for August 2024 relative to the
 659 climatological reference. In panels (a) and (b), the solid black contour line defines the absolute
 660 position of the intertropical front (ITF) via the dew point isodrosotherm at 15°C (288.15 K). In
 661 panel (c), the solid green line follows the dew point anomaly threshold of +1 K (ΔT_d) associated
 662 with the monsoon intrusion, and the stippling indicate regions where convergence anomalies are
 663 statistically significant at a 95% confidence level ($p \leq 0.05$), based on a Monte Carlo permutation
 664 test with 1,000 iterations. The reference vector located at the bottom right of each figure

665 corresponds to a wind speed of 5 m s^{-1} . The red frame in panel (a) delineates the specific study
666 area in northern Chad (16°N – 24°N , 13°E – 25°E).

667

668 3.5 Spatial Distribution of Low-Level Humidity and Moisture Flux

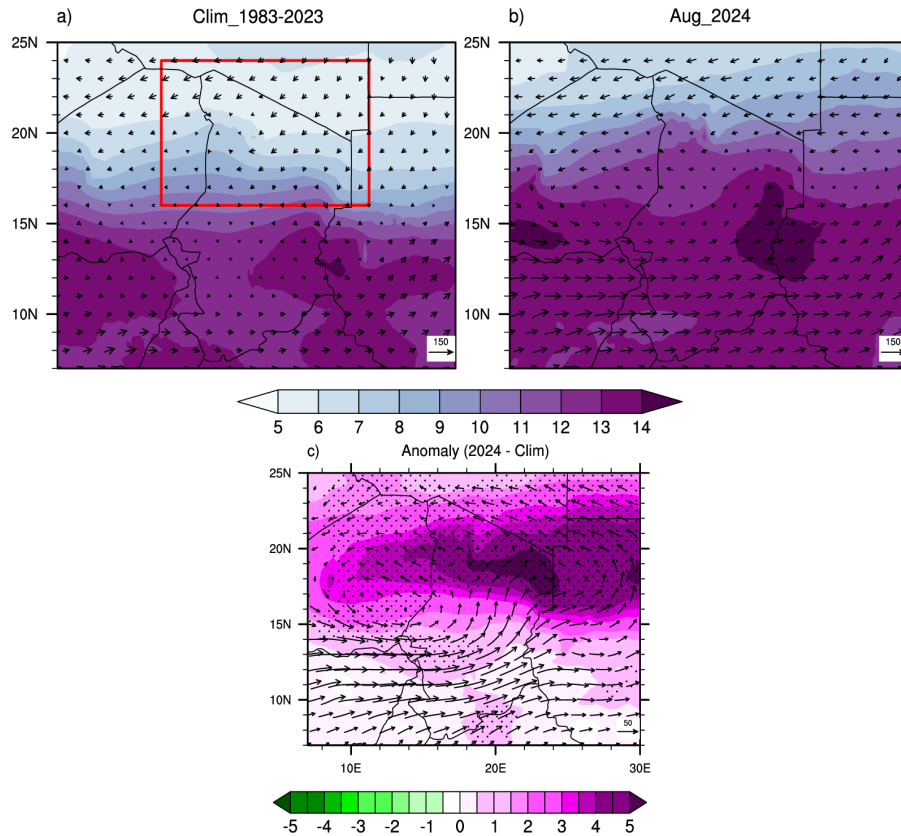
669 While the anomalous 850 hPa wind convergence and northward ITF displacement
670 analyzed in Section 3.4 establish the necessary structural lifting engine over Northern Chad,
671 dynamic forcing alone cannot trigger extreme precipitation without a sufficient and continuous
672 supply of atmospheric water vapor. To assess the thermodynamic environment supporting this
673 localized convergence, it is essential to examine the availability and transport of low-level
674 moisture. Consequently, this section evaluates the spatial distribution of specific humidity and
675 horizontal moisture flux fields at 850 hPa. Figure 7 displays the spatial distribution of 850 hPa
676 specific humidity along with horizontal moisture flux vectors over West and Central Africa.
677 Three configurations are compared: the 1983–2023 August historical climatology (Fig. 7a), the
678 observed conditions in August 2024 (Fig. 7b), and the resulting absolute anomalies (Fig. 7c).

679 The climatological mean (Fig. 7a) highlights a sharp meridional moisture gradient across the
680 region. High specific humidity values (purple shading; exceeding 10 to 14 g kg^{-1}) remain strictly
681 confined south of the 15°N – 18°N latitudinal belt, marking the canonical northernmost boundary
682 of the West African Monsoon (WAM) moisture pool. South of this transition zone, the horizontal
683 moisture flux vectors exhibit a powerful, organized southwesterly routing, driving a continuous
684 supply of monsoonal water vapor from the Atlantic toward the Sahelian corridor. Conversely,
685 inside the red box defining the Northern Chadian study domain (16°N – 24°N , 13°E – 25°E), the
686 lower troposphere is climatologically arid, characterized by low specific humidity values (light
687 blue to white shading; dropping below 6 to 8 g kg^{-1}) and a weak, predominantly zonal moisture
688 transport. This asymmetric structural configuration shows that, under typical summer conditions,
689 the Saharan desert interior is structurally isolated from significant low-level tropical moisture
690 influxes (Nicholson, 2013; Biasutti, 2019).

691 In August 2024 (Fig. 7b), this mean baseline was heavily disrupted. The low-level moisture pool
692 underwent an exceptional northward migration, pushing the 10 to 12 g kg^{-1} specific humidity
693 thresholds deep into the hyper-arid core of Northern Chad, well past 20°N . Concurrently, the
694 horizontal moisture flux vectors reveal a massive, uncharacteristic reorientation. The
695 climatological continental desert regime was entirely replaced by an intensified, cross-equatorial
696 southwesterly moisture conveyor that bypassed its traditional Sahelian limits to directly feed the
697 complex topography of the Tibesti, Borkou, and Ennedi massifs. Physically, this structural
698 change confirms that the 850 hPa layer was subjected to a continuous, large-scale tropical
699 moisture loading, satisfying a critical thermodynamic prerequisite for widespread and persistent
700 high-latitude rainfall anomalies.

701 The absolute anomaly field (Fig. 7c) provides a quantitative diagnostic of this hydroclimatic
702 shift. A vast, spatially cohesive core of positive specific humidity anomalies (green shading;
703 reaching $+2$ to $+4 \text{ g kg}^{-1}$) covers the entire 15°N – 23°N Saharan interior. Crucially, the extensive

704 stippling mask generated by the 1000-iteration Monte Carlo permutation test confirms that these
 705 moist anomalies are statistically significant ($p \leq 0.05$) over virtually the entire northern Chadian
 706 domain. This significant moisture surplus runs perfectly in-phase with robust anomalous
 707 southerly and southwesterly moisture flux arrows. This alignment explicitly proves that the local
 708 humidification of the desert boundary layer was structurally driven by enhanced active advection
 709 from the tropical Atlantic and the monsoonal belt rather than localized evaporation. In summary,
 710 the out-of-boundary convective activation of August 2024 resulted from an optimal
 711 dynamical-thermodynamical coupling. The exceptional low-level moisture advection
 712 documented here (Fig. 7) acted in perfect synergy with the strong 850 hPa wind convergence
 713 anomalies previously analyzed. By shifting the normally dry and stable Saharan boundary layer
 714 into an uncharacteristically humid and moist state, this combined synoptic-scale mechanism
 715 effectively sustained a highly active and long-lived convective regional climate.
 716



718 Figure 7: Spatial distribution of 850 hPa specific humidity and horizontal moisture flux fields in
 719 August over West and Central Africa. The red box outlines the specific study domain over
 720 Northern Chad (16°N–24°N, 13°E–25°E). Panels display: (a) the 1983–2023 climatological
 721 mean, (b) the observed conditions in August 2024, and (c) the absolute anomalies for August
 722 2024 relative to the historical baseline. Shading denotes specific humidity (g kg^{-1}), and vectors
 723 represent horizontal moisture flux ($\text{g kg}^{-1} \text{ m s}^{-1}$). In panel (c), stippling highlights regions where
 724 specific humidity anomalies are statistically significant at the 95% confidence level ($p \leq 0.05$)

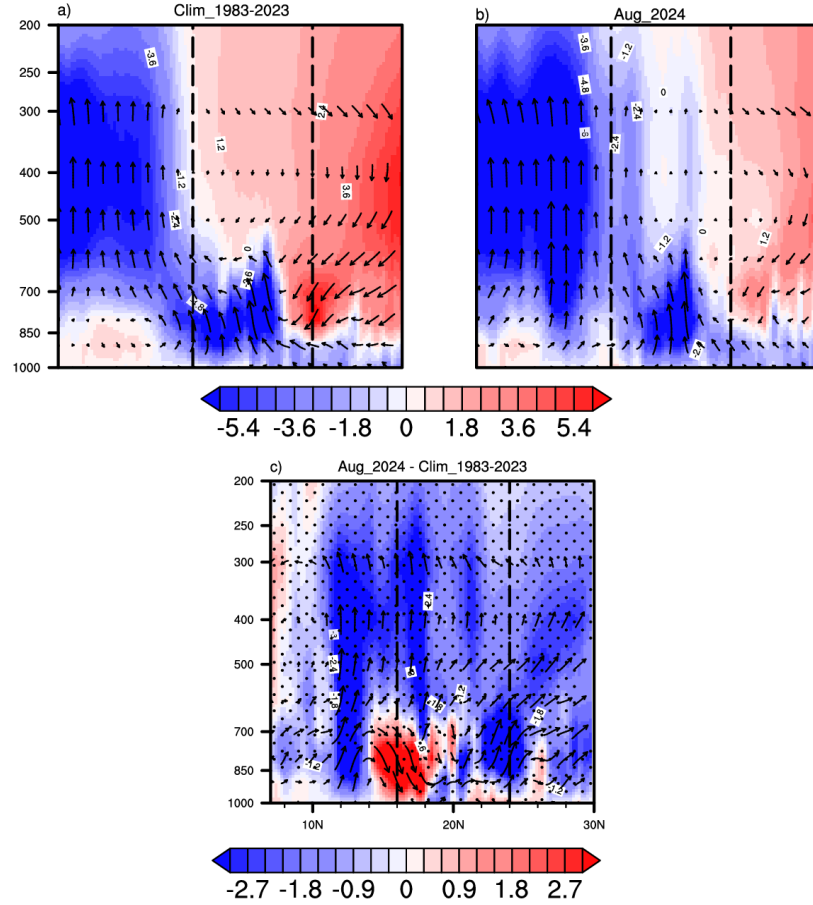
based on a 1000-iteration Monte Carlo permutation test. Reference vectors corresponding to 150 g kg⁻¹ m s⁻¹ (panels a, b) and 50 g kg⁻¹ m s⁻¹ (panel c) are displayed in the bottom-right corners.

3.6. Atmospheric Vertical Structure and Meridional Circulation Cross-Sections

The horizontal accumulation of low-level moisture and enhanced moisture fluxes documented in Section 3.5 over the Saharan interior cannot trigger extreme rainfall anomalies without a corresponding vertical mechanism capable of lifting this moist air mass throughout the troposphere. To understand the dynamic framework supporting these surface thermodynamic shifts, it is essential to examine the meridional-vertical structure of the atmospheric column. Consequently, this section investigates the vertical overturning regional circulation. Figure 8 presents the latitude-pressure cross-sections of vertical velocity (ω , shaded) and superimposed meridional-vertical circulation vectors ($\mathbf{v}$, ω) averaged over the 13°E–25°E longitudinal band for the August historical climatology (Fig. 8a), August 2024 (Fig. 8b), and their absolute anomalies (Fig. 8c). Following the standard meteorological convention, negative ω values (blue shading) denote upward atmospheric motion, whereas positive values (red shading) correspond to subsidence.

In the historical climatology (Fig. 8a), the regional overturning circulation exhibits a classic asymmetric Hadley-type cell configuration. Deep and intense ascending motions ($\omega < 0$; blue shading) are structurally confined between 10°N and 15°N, extending from the lower boundary layer up to the upper troposphere (~300 hPa). This deep vertical convective core represents the mean position of the summer monsoonal ascending branch. North of 18°N, including the northern Chadian study domain (16°N–24°N, highlighted by the vertical dashed lines), the troposphere is heavily dominated by widespread subsidence ($\omega > 0$; red shading), particularly in the middle and upper layers (600–200 hPa). This pervasive mid-tropospheric capping downwelling structurally stabilizes the Saharan atmosphere, explaining the hyper-arid baseline of the desert interior under typical summer conditions.

During August 2024 (Fig. 8b), this mean vertical circulation experienced a profound latitudinal reorganization. The core of deep monsoonal ascent underwent an exceptional northward migration and intensification, with robust negative ω fields (blue shading) spreading continuously into the Saharan interior up to 21°N–22°N. The circulation vectors reveal a powerful, uninterrupted vertical conveyor bridging the lower layers (1000–850 hPa) directly with the upper tropospheric outflow up to 200 hPa. This structural shift implies that the climatological Saharan subsidence capping was completely broken down and replaced by a deep, regional-scale ascending chimney, creating an environment highly favorable for sustained regional moist convection.



759

760 Figure 8: Latitude-pressure cross-sections of meridional-vertical atmospheric circulation and
 761 thermodynamic structure for August averaged over the 13°E - 25°E longitudinal band. Shading
 762 displays the vertical velocity (ω , $\times 10^{-2} \text{ Pa s}^{-1}$), where negative values denote upward motion and
 763 positive values correspond to subsidence. Vectors represent the combined meridional and vertical
 764 wind components (v and ω , with vertical velocity scaled for visualization). Panels show (a) the
 765 1983-2023 climatological mean, (b) August 2024, and (c) the absolute anomalies (August 2024
 766 minus climatology). In panel (c), stippling isolates regions where anomalies are statistically
 767 significant at the 95% confidence level ($p \leq 0.05$) based on a 1000-iteration Monte Carlo
 768 permutation test. Dashed vertical lines delimit the latitudinal domain of the study area (16°N -
 769 24°N).

770 The absolute anomaly field (Fig. 8c) confirms the statistical robustness of this dynamic
 771 displacement, with the 1000-iteration Monte Carlo permutation test isolating large areas of
 772 significant changes ($p \leq 0.05$; stippling mask) within the Saharan interior. A prominent, deeply
 773 rooted negative ω anomaly (blue shading) occupies the 16°N–24°N latitudinal belt, indicating an
 774 unprecedented intensification of upward vertical velocity relative to the climatology. Directly
 775 south of this feature, between 15°N and 18°N and confined to the lower troposphere (1000–700
 776 hPa), a localized positive ω anomaly (red shading) is visible, accompanied by
 777 downward-pointing anomalous circulation vectors. When cross-examined with the absolute
 778 fields (Figs. 8a, 8b), where blue shading (upward motion) persists during both periods, this

779 positive anomaly does not represent physical synoptic-scale subsidence or an invasion of dry
780 desert air. Instead, it indicates a localized, non-statistically significant reduction in the magnitude
781 of the monsoonal updrafts. Because the main convective core of the West African Monsoon
782 underwent a pronounced northward migration beyond 18°N in August 2024, the vertical ascent
783 within the traditional 15°N–18°N latitudinal belt was weaker in absolute terms than its strong
784 historical climatological counterpart, mechanically generating a positive anomaly value.

785 In summary, Figure 8 demonstrates that the extreme Saharan rainfall of August 2024 was driven
786 by an optimal dynamical-thermodynamical synergy. The exceptional low-level moisture
787 advection previously analyzed acted in perfect harmony with a major vertical reorganization of
788 the West African Monsoon circulation, providing the large-scale lifting engine necessary to
789 sustain deep and widespread atmospheric destabilization over Northern Chad.

790

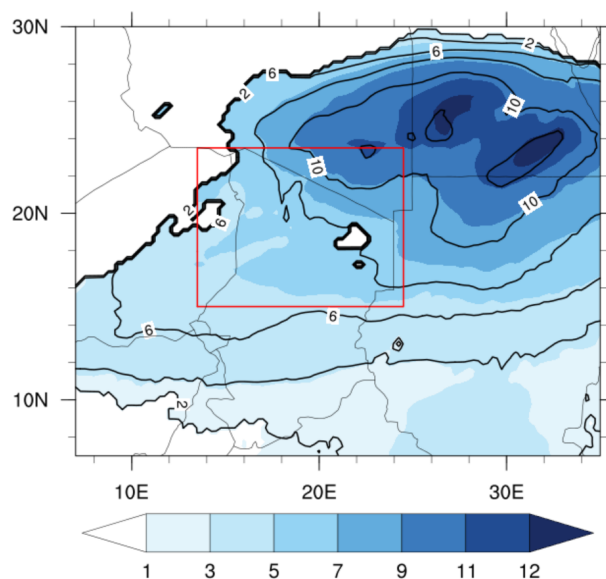
791 3.7 Spatial patterns of moist static energy anomalies

792 While the anomalous meridional-vertical overturning analyzed in Section 3.6 provides the deep
793 dynamical lifting engine necessary to breach the Saharan subsidence capping, the longevity and
794 intensity of such regional updrafts depend heavily on the continuous availability of thermal and
795 latent energy at the surface. To map the horizontal footprint and intensity of the energy reservoir
796 feeding this enhanced vertical circulation, it is essential to examine the spatial distribution of key
797 atmospheric thermodynamic tracers. Consequently, this section investigates the regional patterns
798 of low-level energy accumulation.

799 Figure 9 illustrates the spatial distribution of low-level moist static energy (MSE; shaded, kJ
800 kg⁻¹) anomalies and the corresponding contours of equivalent potential temperature (θ_e ; black
801 contours, K) anomalies over the Sahelian and Saharan sectors. As an integrative thermodynamic
802 state variable combining sensible heat, geopotential, and latent enthalpy contributions, low-level
803 MSE serves as a robust diagnostic tool to link sub-synoptic dynamic forcing (such as wind
804 convergence) to the thermodynamic conditions required to fuel deep convection (Neelin and
805 Held, 1987; Romps, 2015). To guarantee the statistical robustness of the featured structures, only
806 highly significant anomalies exceeding twice the climatological standard deviation ($> 2\sigma$) are
807 represented in this figure. Robust positive MSE anomalies dominate the central and eastern
808 Saharan latitudes, displaying a well-defined zonal structure. The core of these statistically
809 significant thermodynamic anomalies is centered between 18°N and 25°N, directly
810 encompassing the Northern Chad study area designated by the red bounding box. This spatial
811 configuration reflects an exceptional energetic enrichment of the lower troposphere over an arid
812 region typically characterized by a low historical energy baseline. Conversely, the weaker or
813 absent anomalies observed further south imply a significant northward shift of the moist
814 monsoon energy reservoir into the Saharan transition zone. In accordance with the theoretical
815 framework of Romps (2015), the accumulation of high boundary-layer MSE shifts the ascending
816 air parcel's moist adiabat, directly maximizing its thermal buoyancy integral aloft and unlocking
817 the potential for deep convective initiation. The contours of θ_e anomalies exhibit a spatial
818 geometry that perfectly mirrors the MSE fields, with isolines strictly bounding the core of the

819 energy maxima. The tight spatial co-location between the maximum MSE shading and the
 820 elevated θ_e contours highlights their direct thermodynamic covariation arising from shared
 821 conservation properties during moist adiabatic displacements within the planetary boundary
 822 layer. The northward extension of the 6 K and 10 K anomaly contours deep into the Sahara
 823 Desert confirms that the observed regional warming and moisture enrichment were
 824 co-dependent, fueled by the intense horizontal advection of moist monsoonal air coupled with
 825 anomalous surface diabatic heating. This closely coupled MSE- θ_e configuration marks a
 826 profound reorganization of the regional desert thermodynamic environment relative to its
 827 baseline. Such persistent, large-scale positive energy anomalies exceeding the 2σ threshold are
 828 recognized precursors to extreme rainfall over West Africa when spatially aligned with
 829 lower-tropospheric convergent dynamic structures (Taylor et al., 2017). In a broader context, the
 830 heavy lower-tropospheric energy loading illustrated in Figure 9 is consistent with recent
 831 mechanistic views linking the intensification of the West African monsoon system to a higher
 832 frequency of intense rainfall over North Africa (Biasutti, 2019). This thermodynamic
 833 configuration highlights the role of atmospheric moisture preconditioning and its interaction with
 834 the Sahelian dryline boundary, which are critical precursors for driving high-impact convective
 835 extremes across the region (Vizy and Cook, 2022). The accumulation of high surface MSE over
 836 Northern Chad acted as a powerful thermodynamic primer; once matched with the persistent 850
 837 hPa wind convergence and moisture pump documented in the preceding sections, it sustained the
 838 development and longevity of the severe convective systems observed in August 2024. **This
 839 synergy aligns perfectly with the regional-scale mechanisms described by Akinsanola et al.
 840 (2019), whose moisture budget analyses demonstrate that rainfall intensification across the
 841 central and eastern Sahel is intrinsically driven by a co-forcing between dynamic convergence in
 842 the low-to-mid troposphere and the thermodynamic enrichment of surface moisture.**

843



844

Figure 9. Spatial anomalies of moist static energy (MSE; shaded, kJ kg^{-1}) and equivalent potential temperature (θ_e ; black contours, K). The map illustrates the coupling between regional thermodynamic variability and the spatial organization of MSE over the Sahel-Saharan region. The superposition of these fields highlights the northward expansion of high-energy air masses and the intensification of the monsoon-related energy reservoir.

850

3.8. Socio-economic impacts of the August 2024 extreme event

In 2024, Chad experienced the most catastrophic flooding in its recorded history. Nationwide, the impact was profound: nearly 2 million people were affected, resulting in 576 confirmed fatalities. The floods submerged approximately 432,800 hectares of agricultural land, destroyed 217,700 homes, and led to the loss of 72,100 head of livestock. Furthermore, infrastructure damage was extensive, with 3,058 schools severely impacted. Geographically, the crisis was universal, affecting all 23 provinces. Notably, the Saharan provinces, typically spared due to their hyper-arid climate, suffered disproportionately heavy losses.

In the desert zone, initial warnings were issued in July as flooding began in the south. However, torrential rains struck the northern desert region with unprecedented intensity between August 9th and 14th, 2024. National meteorological records documented 126.5 mm of rainfall over eight days in areas where total annual precipitation rarely reaches 100 mm (Tchadinfos). These levels represent the highest rainfall totals in 60 years. Visual evidence (Fig. 8) corroborates these records, showing water levels reaching knee-depth at the Faya-Largeau market, where commercial and material losses were severe.

866



867

Figure 10: Flooded streets and commercial stalls at the Faya-Largeau market (Source: Tchad Info Société).

870

Within the specific study area, the flooding impacts are categorized by human displacement, mortality, agricultural devastation, and infrastructural loss. While national reports often aggregate these figures, detailed provincial data is typically only prioritized where significant

losses of life occur. Consequently, the situations in Borkou (4 deaths) and Tibesti (60 deaths) are better documented than in the Ennedi provinces, which reported no fatalities. However, OCHA situational reporting as of October 1, 2024, allows for a more granular assessment of the affected populations across all study provinces. This disparity highlights the critical need for developing nations to enhance civil protection services and systematic data collection during hydrometeorological disasters (Saha et al., 2018; Mazhin et al., 2021). Table 1 presents the damage breakdown for each province covered by this study.

881

882 Table 2: Losses due to flooding in the study provinces

883

Region	Number of families affected	Number of people affected	Number of dead recorded	Number of houses/shops destroyed	number of animals killed
Borkou	10,402	67,613	04	23,174 houses	861
Ennedi East	35,980	233, 872			
Ennedi West	1, 842	11,977			
Tibesti	23	150	60	Thousands shops	

Beyond the quantified losses in Table 1, the regional economy, heavily reliant on trade and mining, suffered significant shocks. Thousands of shops were destroyed, and merchandise was lost. In the Tibesti region, gold mining operations suffered immense equipment losses. A particularly grave secondary hazard emerged as the 2024 floods unearthed thousands of landmines remaining from previous conflicts, creating a long-term security and humanitarian risk for the local population.

890

891 Conclusion

This study elucidates the atmospheric mechanisms and long-term hydro-climatic shifts driving the unprecedented, record-breaking rainfall event observed in August 2024 across the hyper-arid Chadian Sahara (16°N–24°N, 13°E–25°E), while documenting its catastrophic socio-economic impacts. Based on a 42-year climatological reference period (1983–2024), the results demonstrate that, far from being an isolated synoptic fluctuation or one unrelated to other phenomena, the historical anomaly observed in August 2024 represents the ultimate manifestation of a persistent and statistically significant trend toward increased humidity over several decades. This long-term transition, fundamentally marked by the year 2003 as a key

900 climatological threshold, is characterized by a continuous trend toward increased atmospheric
901 moisture content, with a +59.2% rise in integrated water vapor transport (IVT). This altered
902 baseline which has been wetter since 2003 has maximized regional vulnerability to extreme
903 monsoon precipitation, structurally shifting the regional hazard paradigm from perpetual
904 hyper-aridity to severe hydrological risks.

905 On a regional scale, the exceptional magnitude of the August 2024 event is evidenced by
906 accumulated monthly rainfall totals that reached up to 250-400 mm over the rugged topographies
907 of the Tibesti, Borkou, and Ennedi massifs, representing the highest rainfall totals in 60 years.
908 These accumulations correspond to extraordinary standardized anomalies systematically
909 exceeding $+8\sigma$ (and locally transcending $+12\sigma$), which are shown to be statistically
910 significant ($p \leq 0.05$) via grid-point Monte Carlo testing. Parametric probability frameworks
911 reveal that this rainfall plume represents a centennial to multi-centennial phenomenon, with
912 return periods estimated between 124.5 years (CHIRPS) and 451.4 years (TAMSAT). This
913 extreme pluviometric response was sustained by an unprecedented, vertically coherent
914 thermodynamic loading, wherein monthly specific humidity at 925 hPa and 850 hPa, alongside
915 Total Column Water Vapor (TCWV), reached their absolute historical records (97.67th empirical
916 percentile), establishing a highly conducive, moisture-saturated column over the Saharan interior.
917 The synthesis of atmospheric diagnostics reveals that this out-of-boundary event was governed
918 by an optimal, highly synchronized dynamic-thermodynamic coupling. Synoptically, the first
919 half of the month was intensely triggered by high-amplitude African Easterly Waves (AEWs)
920 between August 10 and 14, acting as a powerful dynamic conveyor belt for low-latitude moisture
921 advection. On a monthly scale, this transient wave activity drove a radical breakdown of the
922 typical desert circulation, forcing a profound northward migration of the Intertropical Front
923 (ITF), which reached latitudes of 20° - 22° N, supported by a synchronized intensification of the
924 West African Monsoon (WAM).

925 Consequently, the climatological low-level divergence and mid-tropospheric subsidence capping
926 that historically stabilize the desert interior were completely replaced by robust 850 hPa wind
927 convergence anomalies, a sharp increase in lower-tropospheric specific humidity, and an
928 uninterrupted, deep vertical ascending chimney reaching the upper troposphere (~ 200 hPa). This
929 is evidenced by pronounced negative ω anomalies extending from the lower to the upper
930 troposphere, which perfectly coupled in space with robust lower-tropospheric moist static energy
931 (MSE) and equivalent potential temperature (θ_e , derived via Bolton's formulation) anomalies
932 exceeding the 2σ statistical threshold over Northern Chad. This substantial energetic loading
933 indicates a profound thermodynamic enrichment of the atmospheric column in regions typically
934 dominated by dry air and moisture deficits.

935 The accumulation of high boundary-layer energy acted as a powerful thermodynamic primer;
936 once matched with the persistent low-level convergence pump and horizontal moisture
937 advection, it provided the necessary instability to sustain the vertical lifting and longevity of the
938 convective systems, resulting in the exceptional rainfall totals observed.

939 The socio-economic consequences of this hydroclimatic shift were historically devastating. On a
940 national scale, the floods affected nearly 2 million people, claiming 576 lives, submerging
941 432,800 hectares of agricultural land, and destroying 217,700 homes and 3,058 schools. Within
942 the hyper-arid study domain, the impacts were severe and spatially heterogeneous: the Borkou
943 province recorded over 67,000 affected individuals and the destruction of 23,174 structures,
944 while the Tibesti region suffered a heavy human toll with 60 confirmed fatalities. Beyond
945 structural and livestock losses, the severe rainfall caused major economic shocks to local trade

946 and gold mining operations due to massive equipment destruction. Crucially, the floods
947 generated a grave secondary humanitarian hazard by unearthing thousands of conflict-era
948 landmines, introducing long-term security risks for local populations.

949 Finally, the scale of this disaster underscores the profound vulnerability of hyper-arid regions to
950 sudden hydrological hazards. Given that the observed intensification of the regional hydrological
951 cycle suggests an increasing recurrence of such high-impact convective extremes under global
952 climate change, there is an urgent need for regional authorities to transition from reactive disaster
953 management to proactive climate adaptation strategies. Furthermore, these findings provide a
954 robust empirical benchmark for improving the representation of monsoon-desert interfaces in
955 climate models, highlighting the necessity for future research utilizing high-resolution
956 convection-permitting simulations to track the long-term evolution of these saharan
957 hydroclimatic shifts.

958

959 CODE AVAILABILITY

960 Figures shown in this study are plotted using the NCAR Command Language (NCL;
961 <https://doi.org/10.5065/D6WD3XH5>, NCAR Command Language, 2017) and QGIS. Codes can
962 be obtained from the corresponding author.

963

964 DATA AVAILABILITY STATEMENT

965 The ERA5 data were obtained from the Copernicus Climate Change Service (C3S) Climate Data
966 Store (<https://cds.climate.copernicus.eu>). Climate Hazards Group InfraRed Precipitation with
967 Station data (CHIRPS) is available through the link <https://www.chc.ucsb.edu/data/chirps3>
968 The TAMSAT rainfall estimates and derived products are based on Meteosat thermal infra-red
969 (TIR) imagery provided by EUMETSAT (<https://data.tamsat.org.uk/data-download/rainfall/>)

970

971 AUTHOR CONTRIBUTIONS

972 CWT: conceptualization; data analysis; formal analysis; investigation; methodology; writing
973 original draft; project administration; review and editing. FS: supervision; conceptualization;
974 data analysis; investigation; project administration; writing original draft; review and editing.
975 ATT: methodology; resources; validation; writing; review; project administration and editing.

976

977 COMPETING INTERESTS

978 The authors declared that they have no conflict of interests.

979

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986

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