



1 **Measurement Report: Quantitative Analysis of Aerosol Acidity and**
2 **Its Driving Factors in Guanzhong Area, Northwest China**

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16 **Abstract.** Aerosol acidity significantly affects atmospheric chemistry and human
17 health, yet its driving factors remain controversial. This study systematically
18 examined a year-long characteristics of water-soluble inorganic ions in PM_{2.5} in the
19 semi-arid Guanzhong Plain, Northwest China, and quantitatively analyzed its acidity
20 and driving factors. The annual mean pH of PM_{2.5} was 3.8 ± 1.0 (winter > spring >
21 summer > autumn). As pollution increased, aerosol pH shifted from the acidic range
22 (2-5) on clean days to a near-neutral range (3-6) on polluted days. Sensitivity tests and
23 driver analysis revealed that atmospheric temperature (22.3%-33.8%), NH_x (gas
24 NH₃+NH₄⁺, 11.4%-44.9%), and SO₄²⁻ (8.5%-10.8%) were common key factors
25 influencing pH across all seasons. Among these, temperature played a dominant role
26 in seasonal variations, while NH_x was the primary contributor during autumn and
27 winter. Notably, Ca²⁺ emerged as a unique driver specific to spring, the relative
28 standard deviation (RSD) was 8.8%. Relative humidity (RH) exhibited a distinctive
29 non-linear regulatory effect, i.e., aerosol pH initially decreases and subsequently
30 increases with rising RH (inflection point occurred at 60%–85%). This phenomenon
31 is primarily attributed to an abrupt change in liquid water content triggered by the
32 deliquescence of hygroscopic components. Our results enhance the understanding of
33 the contributions of various factors to aerosol pH and offers a new perspective for
34 developing PM_{2.5} control strategies in the semi-arid and ammonia-rich region.

35 **Keywords:** Aerosol acidity, aerosol liquid water, quantitative analysis, driving factors,
36 Guanzhong Plain



37 **1. Introduction**

38 Atmospheric aerosols play a vital role in the Earth's atmospheric system,
39 significantly impacting regional and global air quality, climate change, and human
40 health (Liu et al., 2025; Zhang et al., 2017). Among these properties, aerosol acidity,
41 typically characterized by pH, is one of the most critical parameters. Research
42 indicates that acidity can directly or indirectly regulate the gas-particle partitioning of
43 semi-volatile species and the rates of chemical reactions (Guo et al., 2016; Paglione et
44 al., 2021). It drives pH-dependent aqueous-phase reactions and impacts other
45 processes related to particle formation and growth. Additionally, acidity can enhance
46 the formation of secondary organic aerosol (SOA) and influence its transformation
47 pathways through acid-catalyzed reactions (Rengarajan et al., 2011; Shi et al., 2019).
48 Moreover, aerosol pH affects the dissolution of trace metals and the concentration of
49 their toxic forms, with stronger acidity promoting metal dissolution via acid
50 dissociation (Ding et al., 2019; Shi et al., 2011). High aerosol acidity levels have been
51 linked to increased risks of respiratory illnesses and specific cancers. These
52 acidity-driven atmospheric processes collectively contribute to the generation of
53 intricate air pollution, the deposition of pollutants through dry and wet mechanisms,
54 and their overall effects on human health and the climate system (Mao et al., 2009;
55 Pye et al., 2020). Therefore, the exploration of particle pH is crucial for enhancing air
56 quality management strategies and policy development, which is highly important for
57 alleviating the health and environmental consequences of air pollution. The direct
58 measurement of atmospheric aerosol acidity presents challenges (Xu et al., 2025),
59 primarily due to the difficulties in quantifying the aqueous-phase concentrations of
60 semi-volatile compounds and H^+ in ambient particulate matter. Thus, thermodynamic
61 modellings, such as E-AIM (Clegg et al., 1998) and ISORROPIA (Nenes et al., 1998),
62 were frequently used as a proxy method for assessing aerosol acidity in recent
63 decades.

64 Aerosol acidity displays notable spatiotemporal variability, playing a crucial role
65 in atmospheric environmental research. Globally, aerosols in Europe and North



66 America generally exhibit higher acidity levels compared to those in China, with pH
67 values 1–2 units lower (Zhang et al., 2021). In North America, coastal cities in
68 Canada frequently exhibit greater aerosol acidity than inland regions, primarily due to
69 elevated humidity that enhances the aqueous-phase oxidation of SO₂ (Brook et al.,
70 1997). In China, aerosol acidity typically rises from north to south. Specifically,
71 during winter in the heavily polluted North China Plain, pH values typically range
72 from 4.1 to 4.9 (Ding et al., 2019; Liu et al., 2017; Shi et al., 2019; Tan et al., 2018).
73 In the Guanzhong region, aerosol pH can reach as high as 5 (Wang et al., 2016).
74 Conversely, coastal cities in southeastern China, such as Guangzhou and Xiamen,
75 display significantly lower mean pH values of 2.5 and 3.5, respectively (Jia et al.,
76 2020; Xu et al., 2025). Furthermore, vertical differences in aerosol acidity are also
77 noteworthy. As altitude increases, aerosol acidity generally intensifies, with higher
78 acidity observed at mountain summits compared to surface levels (Feng et al., 2023;
79 Guo et al., 2016). Temporally, aerosol acidity is typically higher in summer than in
80 winter, while patterns in spring and autumn fluctuate based on regional climatic and
81 emission characteristics (Tan et al., 2018; Xu et al., 2025). In recent years, aerosol
82 acidity has gained prominence in academic research and discussion due to its critical
83 role in influencing aerosol physicochemical properties (Cui et al., 2025; Liu et al.,
84 2021; Paglione et al., 2021; Tao et al., 2025). However, its levels are governed by
85 numerous, complex, and region-dependent factors, leading to significant spatial
86 heterogeneity. Therefore, investigating the dynamic changes in aerosol acidity
87 throughout different seasons and understanding the underlying mechanisms are
88 crucial for enhancing our comprehension of haze formation and developing precise air
89 pollution control strategies.

90 Through quantitative analysis of aerosol pH driving factors and elucidation of
91 PM_{2.5} chemical composition and meteorological impacts across seasons, we can better
92 understand spatiotemporal variation pattern (Jia et al., 2020; Xu et al., 2025). Previous
93 studies showed that in Beijing, SO₄²⁻, NH_x, and T significantly affect PM_{2.5} pH, while
94 Ca²⁺ and RH serve as dominant factors in spring and summer, respectively (Ding et al.,
95 2019). In contrast, in southeastern Chinese cities such as Guangzhou and Xiamen,



96 meteorological parameters like T and RH exert an even greater influence on pH
97 variation than chemical components (Jia et al., 2020; Xu et al., 2025). These findings
98 underscore the complexity of the factors influencing aerosol acidity. However,
99 existing studies have largely focused on compositional analyses in specific cities,
100 leaving our understanding of the trends in aerosol pH values and their primary drivers
101 and relative contributions still insufficient. The discrepancies in conclusions from
102 different research further highlight the geographical specificity of this issue.

103 Guanzhong Plain is a vital economic and agricultural center in northwestern
104 China, the rapid industrial and urban development has resulted in severe air pollution
105 complexities. The mechanisms of air pollution formation and its driving factors in
106 the region are significantly shaped by its inland geographical characteristics and
107 socio-economic conditions. The unique semi-enclosed topography, with the Loess
108 Plateau to the north and the Qinling Mountains to the south, restricts pollutant
109 dispersion, leading to the accumulation of various pollutants and enabling complex
110 chemical reactions (Huang et al., 2014; Liu et al., 2025; Shen et al., 2024). The area is
111 known for its intensive agricultural practices, contributing to high levels of ammonia
112 emissions that influence aerosol acid-base equilibrium. Additionally, long-distance
113 transportation from the nearby Loess Plateau introduces alkaline mineral dust and
114 crustal cations (e.g., Ca^{2+} , Mg^{2+}) that can neutralize acidic components effectively.
115 Furthermore, distinct seasonal climates (e.g., stable conditions in winter, high
116 temperatures in summer) further influence the transformation and removal processes
117 of pollutants. The combination of these factors makes the Guanzhong Plain a unique
118 location for a comprehensive study on aerosol pH trends and key influencing factors,
119 which would differ from those in southeastern coastal cities of China, where marine
120 air masses and hot, humid meteorological conditions predominantly influence air
121 quality.

122 In this study, we conducted a year-long observation of water-soluble inorganic
123 compositions of $\text{PM}_{2.5}$ in the Guanzhong region, and utilized an ISORROPIA-II model
124 to access aerosol pH and liquid water content. It comprehensively examined the
125 variations of these parameters throughout different seasons and under diverse



126 pollution levels. Additionally, sensitivity tests and quantitative assessments were
127 conducted to identify key factors influencing aerosol pH and to elucidate their relative
128 contributions across seasons. This investigation establishes a robust scientific
129 foundation for the implementation of precise ammonia emission reduction tactics,
130 aiming to effectively decrease atmospheric fine particulate matter levels.

131 **2. Materials and Methods**

132 **2.1 Sample collection**

133 PM_{2.5} samples were collected at the National Observation and Research Station
134 of Regional Ecological Environment Change and Comprehensive Management in the
135 Guangdong Plain, Shaanxi Province, China (34°3'39"N, 108°20'37"E; 503 m a.s.l.).
136 This station is positioned at the northern foothills of the Qinling Mountains (Fig. 1)
137 and functions as a regional background site. It is located approximately 60 km
138 southwest of Xi'an city, with no significant industrial, traffic, or other substantial
139 anthropogenic pollution sources. The sampling campaign spanned from January 1 to
140 December 21, 2022, and a total of 295 valid samples were obtained during the study.
141 Daily sampling was conducted from 10:00 to 09:00 the next day, lasting for 23 hours.
142 Samples were collected using a medium-flow sampler (HC-1010, China Qingdao
143 Company, China) equipped with quartz fiber filters (Whatman QMA, USA) with
144 pre-baked (450 °C, 5 h) quartz fiber filters (Whatman, QMA, USA) at an airflow rate
145 of 100 L min⁻¹. After sampling, each filter sample was promptly enveloped in
146 aluminum foil, sealed, and frozen at -20°C for subsequent laboratory analysis. To
147 ensure quality control, field blank samples were collected both before and after
148 sampling. This involved placing pre-treated blank filters into the sampler, allowing
149 them to stand for 10 minutes without activating the pump, and then sealing and
150 storing them using the same procedure as the actual samples. The chemical analysis
151 data for all samples were adjusted by deducting the average concentration measured
152 from the two blank samples to eliminate potential systematic errors.

153 Daily average concentrations of PM₁₀, SO₂, NO₂, CO, and O₃-8h (defined as the
154 maximum 8-hour average O₃ concentration) were obtained from the local government.



155 Meteorological parameters, including ambient temperature (T), relative humidity
156 (RH), wind speed (WS), and atmospheric pressure (AP), were acquired from the
157 National Observation and Research Station for Regional Ecological Environment
158 Change and Comprehensive Management in the Guanzhong Plain, Shaanxi Province,
159 China.

160 **2.2 Analysis of water-soluble inorganic ions and ammonium partitioning ratio**

161 Water-soluble inorganic ions (WSIIs), including SO_4^{2-} , NO_3^- , NH_4^+ , Cl^- , Ca^{2+} , K^+ ,
162 Mg^{2+} , and Na^+ , were analyzed using ion chromatography (Metrohm-940, Switzerland).
163 To enhance measurement accuracy, a fixed amount of lithium bromide was added to
164 the eluent. For quality assurance, one sample out of every ten was randomly selected
165 for re-analysis. All WSII concentrations were corrected using field blanks, which
166 exhibited concentrations less than 10% of those in ambient samples. The limits of
167 detection (LODs) for SO_4^{2-} , NO_3^- , NH_4^+ , Cl^- , Ca^{2+} , K^+ , Mg^{2+} , and Na^+ were 0.027,
168 0.026, 0.0010, 0.0087, 0.0012, 0.0011, 0.0008, and 0.0005 $\mu\text{g m}^{-3}$, respectively (Liu
169 et al., 2025). The mass concentrations of all detected WSII were significantly higher
170 than their corresponding LODs.

171 As detailed in the preceding section, this study employed ammonium partitioning
172 ratio (NHR) to quantify the transformation of gaseous ammonia into particulate
173 ammonium (Xie et al., 2020; Zhang et al., 2021). This ratio is defined as the molar
174 concentration of particulate ammonium ($n(\text{NH}_4^+)$, $\mu\text{mol m}^{-3}$) relative to the total molar
175 concentration of inorganic ammonium ($n(\text{NH}_3) + n(\text{NH}_4^+)$, $\mu\text{mol m}^{-3}$), calculated
176 using the following equation:

$$177 \quad \text{NHR} = \frac{n(\text{NH}_4^+)}{n(\text{NH}_4^+) + n(\text{NH}_3)} \quad (1)$$

178 **2.3 Calculation of aerosol pH**

179 In this study, the ISORROPIA-II model was utilized to estimate aerosol
180 thermodynamic parameters. The model was operated in "forward" mode, based on the
181 assumption that aerosols were in a "metastable" state (Fountoukis et al., 2009; Nenes
182 et al., 1998), with inputs including concentrations of SO_4^{2-} , NH_x , NO_3^- , Cl^- , Na^+ , Ca^{2+} ,
183 K^+ , Mg^{2+} , RH and T. Aerosol pH was calculated using the following equation:



184
$$\text{pH} = -\lg \frac{1000\gamma_{\text{H}^+}\text{H}_{\text{air}}^+}{\text{ALWC}} \quad (2)$$

185 where γ_{H^+} is the hydronium ion activity coefficient (assumed = 1), H_{air}^+ ($\mu\text{g m}^{-3}$)
186 indicates the concentration of hydronium ions per unit volume of air, and ALWC (μg
187 m^{-3}) signifies the aerosol liquid water content. It should be noted that the
188 model-calculated pH values are more reliable when the RH falls within the range of
189 30% to 95% (Guo et al., 2016)

190 Neglecting the influence of gaseous species, such as NH_3 , HNO_3 , and HCl , on
191 aerosol acidity often results in an underestimation of acidity. The effect of these gases
192 on aerosol pH can only be dismissed when their atmospheric concentrations are
193 exceedingly low (Guo et al., 2015; Hennigan et al., 2015). In China, atmospheric
194 ammonia concentrations are relatively high, and the study region is characterized by
195 elevated ammonia levels; thus, NH_3 cannot be overlooked. The specific estimation
196 method for NH_3 is outlined in Section 3.2. However, previous research has shown that
197 the concentrations of gaseous HNO_3 and HCl in urban Xi'an remained below $1 \mu\text{g m}^{-3}$
198 throughout all seasons. Since the Qinling site serves as a background location,
199 concentrations there are anticipated to be even lower. Given these negligibly low
200 levels, we utilized the measured concentrations of particulate NO_3^- and Cl^- as inputs
201 for TNO_3 and TCl , respectively (Nah et al., 2023).

202 **2.4 Multiple linear regression model**

203 In ammonia-rich regions, aerosol pH, particularly during winter, is primarily
204 influenced by NH_3 levels (Tao and Murphy, 2019). However, the frequent absence of
205 long-term and continuous NH_3 observational data impedes the analysis of long-term
206 trends in $\text{PM}_{2.5}$ chemical properties. To fill this data gap and explore variations in
207 aerosol pH in the research area, this study aimed to create a model for estimating the
208 NHR based on easily accessible routine monitoring data to estimate NH_3
209 concentrations. Given the established relationships between NHR and factors such as
210 NO_3^- , SO_4^{2-} , and T, we utilized multiple linear regression analysis with SPSS. This
211 method identified species significantly correlated with NHR from the existing dataset
212 and established quantitative relationships among them. Model parameters were
213 estimated using the ordinary least squares method. Following model development, its



214 validity and accuracy were rigorously assessed through a series of statistical methods,
215 including the correlation coefficient, F-test, and t-test (Wei et al., 2023).

216 **2.5 Sensitivity tests and quantitative analysis**

217 This study assessed the impact of various factors on aerosol pH through
218 sensitivity tests. The variables analyzed included meteorological parameters (T, RH)
219 and chemical components (SO_4^{2-} , NH_x , NO_3^- , Cl^- , Na^+ , Ca^{2+} , K^+ , Mg^{2+}) (Tao and
220 Murphy, 2019) The specific sensitivity analysis adhered to the methodology outlined
221 by Ding et al. (2019): Within the ISORROPIA-II model, the measured value of a
222 target variable i was used sequentially as input alongside the average values of all
223 other parameters. By observing the resulting trend in the output pH, the individual
224 effect of variable i on aerosol pH could be assessed.

225 Assuming the aerosol pH estimated under scenario 1 (pH_1) differed from that
226 under scenario 2 (pH_2), the pH difference ($\Delta\text{pH} = \text{pH}_2 - \text{pH}_1$) was thus caused by the
227 variations in the factors listed above. The quantitative analysis method adhered to the
228 approach outlined by Zhou et al. (2022). To quantify the contribution of a single
229 variable j within the ISORROPIA-II model, the value of factor j was altered from its
230 Scenario 1 value to its Scenario 2 value, while all other parameters remained constant.
231 The resulting change in pH was recorded as pH_j , which represents the independent
232 contribution of that specific factor to the total ΔpH . The remaining change in pH, not
233 accounted for by this procedure, is attributed to $\Delta\text{pH}_{\text{others}}$, reflecting the combined
234 effects of synergistic variations among the factors.

235 **3. Results and discussion**

236 **3.1 Seasonal variations in $\text{PM}_{2.5}$ and its chemical composition**

237 **3.1.1 Seasonal variations in $\text{PM}_{2.5}$ and air pollutants**

238 The annual variation for $\text{PM}_{2.5}$ mass concentrations, gaseous pollutants (SO_2 , CO ,
239 NO_2 , and O_3 -8h), and meteorological parameters (T, RH, WS, and AP) are shown in
240 Fig. 2, with seasonal statistics summarized in Table 1. The sampling period in 2022
241 was categorized into four seasons based on local heating patterns and human activity
242 trends: winter (January 1 to March 12, and November 15 to December 21), spring
243 (March 16 to May 31), summer (June 1 to August 30), and autumn (September 1 to



244 November 14). During the observation period, the PM_{2.5} mass concentration varied
245 from 9.8 to 221.7 $\mu\text{g m}^{-3}$, with an annual mean of $65.3 \pm 61.8 \mu\text{g m}^{-3}$. This
246 concentration is lower than those reported for other typical areas in the Guanzhong
247 Plain (Liu et al., 2025; Zhang et al., 2019), aligning with expectations of minimal
248 local emission influence at this site. In comparison to the annual mean of 105.6 μg
249 m^{-3} recorded at a Qinling background station in 2012 (Niu et al., 2016), the
250 concentration observed in this study is approximately 36% lower, suggesting a
251 significant reduction in air pollution in recent years due to effective air quality control
252 measures. PM_{2.5} concentrations exhibited marked seasonal variations, showing the
253 seasonal order of winter ($93.5 \pm 41.7 \mu\text{g m}^{-3}$) > spring ($61.4 \pm 40.8 \mu\text{g m}^{-3}$) > autumn
254 ($43.1 \pm 25.1 \mu\text{g m}^{-3}$) > summer ($33.2 \pm 13.7 \mu\text{g m}^{-3}$). The majority of days exceeding
255 the daily PM_{2.5} limitation (Grade II, $75 \mu\text{g m}^{-3}$) from the ambient air quality standards
256 of China occurred in winter, with about 68% of daily average concentrations
257 surpassing the limitation (Table S1). This trend underscores the significant influence
258 of heating activities and adverse meteorological conditions on air quality.

259 The concentrations of gaseous pollutants also demonstrated significant seasonal
260 variations. CO and NO₂, primary pollutants mainly from fossil fuel combustion, had
261 higher levels in autumn and winter and lower levels in spring and summer,
262 particularly elevated during the heating period compared to the non-heating period.
263 This pattern is linked to increased emissions during heating demands in autumn and
264 winter, along with unfavorable meteorological conditions hindering pollutant
265 dispersion (Liu et al., 2025). Due to its strong dependence on solar radiation intensity
266 and temperature, O₃-8h value peaked in July during summer ($156.8 \mu\text{g m}^{-3}$) and
267 subsequently decreasing as sunlight weakened with varying intensity. This anti-phase
268 variation underscores the seasonal differences in the dominant formation mechanisms
269 for different pollutants.

270 **3.1.2 Water-soluble inorganic ions in PM_{2.5}**

271 The concentrations of WSIIIs (SO₄²⁻, NO₃⁻, NH₄⁺, Cl⁻, Ca²⁺, K⁺, Mg²⁺, Na⁺) in
272 PM_{2.5} are shown in Fig. 1 and Table 1. The annual average concentration of total
273 measured WSIIIs was $24.5 \pm 20.5 \mu\text{g m}^{-3}$, representing approximately 38% of the



274 PM_{2.5} mass. Significant seasonal variations were noted, with concentrations following
275 the order: winter ($39.5 \pm 22.0 \mu\text{g m}^{-3}$) > autumn ($19.8 \pm 15.9 \mu\text{g m}^{-3}$) > spring ($18.6 \pm$
276 $12.9 \mu\text{g m}^{-3}$) > summer ($8.5 \pm 4.2 \mu\text{g m}^{-3}$). Among the WSIs, secondary inorganic
277 aerosols (i.e., SO_4^{2-} , NO_3^- , and NH_4^+ (SNA)) were the predominant constituents,
278 collectively contributing 89.0% to the total WSIs. Notably, NO_3^- exhibited the
279 highest concentration ($8.8 \mu\text{g m}^{-3}$, 35.9%), followed by SO_4^{2-} ($5.1 \mu\text{g m}^{-3}$, 20.7%) and
280 NH_4^+ ($4.0 \mu\text{g m}^{-3}$, 16.3%), indicating nitrate as the most abundant ionic species. This
281 might be due to driven by enhanced atmospheric oxidation and rising ammonia
282 emissions via pathways like OH radical and N_2O_5 reactions. The ternary plot in Fig.
283 S1 illustrates the molar concentration ratios of SO_4^{2-} , NO_3^- , and NH_4^+ across seasons.
284 The proportion of NO_3^- within SNA was markedly greater in winter compared to
285 summer, while the situation was reversed for SO_4^{2-} . These seasonal fluctuations
286 primarily stem from the thermal instability and volatility of NH_4NO_3 at elevated T
287 (Xu et al., 2025), the fractional contributions of NO_3^- reached minimum in summer.
288 Simultaneously, heightened photochemical processes during summer promote the
289 secondary production of SO_4^{2-} . Moreover, most data points cluster nearer to the NH_4^+
290 vertex. This directly indicates a relative abundance of atmospheric NH_3 during the
291 observation period, adequate to neutralize the majority of sulfuric and nitric acids for
292 ammonium salt formation. This constitutes direct chemical proof elucidating the
293 relatively elevated aerosol pH levels in the ammonia-rich Guanzhong Plain region.
294 The contributions of the remaining ions (Cl^- , K^+ , Mg^{2+} , Ca^{2+} , Na^+) were relatively low,
295 collectively constituting 12.6% of total WSIs. Cl^- and K^+ , commonly utilized as
296 tracers for anthropogenic combustion, exhibited highest concentration in winter, likely
297 linked to seasonal biomass burning in western and northern China and subsequent
298 regional transport processes (M. Liu et al., 2025; Sun et al., 2013; Xie et al., 2020).
299 These chemical characteristics, particularly the relative predominance of nitrate,
300 reflect the evolution of PM_{2.5} chemical properties influenced by China's air pollution
301 control policies. Since the implementation of the Air Pollution Prevention and Control
302 Action Plan in 2013, stringent industrial emission controls and coal reduction
303 measures have yielded significant results, leading to a substantial decrease in SO_2



304 concentrations. Su et al. (2024) reported that SO₂ concentrations in Xi'an decreased by
305 approximately 79% from 2006 to 2021. This reduction has facilitated a shift in aerosol
306 pollution type from sulfate-dominated to nitrate-dominated. This transformation in
307 pollution type has emerged as a new characteristic of regional air pollution across
308 China (Krotkov et al., 2016; Mao et al., 2022; Z. Wang et al., 2022).

309 Linear regression analysis (Fig. S2) revealed a significant correlation between
310 the equivalent concentrations of cations and anions throughout all seasons. The annual
311 average of the regression slope (AE/CE) was 0.92, indicating that PM_{2.5} was generally
312 weakly alkaline during the study period. This finding is consistent with other studies
313 conducted in northern Chinese cities (Cao et al., 2013; Shi et al., 2017), yet it
314 contrasts with historical records indicating near-neutral to slightly acidic aerosols in
315 Xi'an in 2005 and Beijing in 2014 (Tan et al., 2018; Shen et al., 2007). The alteration
316 in acid-base characteristics is a result of reduced sulfur emissions due to China's air
317 pollution control efforts, impacting the chemical composition of PM_{2.5}. This
318 transformation corresponds with findings in Southwestern China from 2012 to 2016
319 (Chen et al., 2017, 2019; Huang et al., 2022). Seasonally, the AE/CE ratios for all
320 seasons in 2022 remained below 1. The lowest ratios occurred in spring (0.88) and
321 summer (0.89), which may be associated with the frequent dust events during these
322 seasons that introduce more alkaline dust particles. It is noteworthy that the
323 measurement methods used in this study did not include CO₃²⁻ and HCO₃⁻, which are
324 significant contributors to the anion deficit.

325 The linear relationship between the observed and predicted NH₄⁺, based on local
326 ammonia availability, is depicted in Fig. S3. When assuming that NO₃⁻ and SO₄²⁻
327 exist in the forms of (NH₄)₂SO₄ and NH₄NO₃, the slope (k) for each season
328 approaches 1, indicating complete conversion to (NH₄)₂SO₄. NH₄⁺ first reacts with
329 SO₄²⁻ to form (NH₄)₂SO₄. After SO₄²⁻ is completely neutralized, any excess NH₄⁺ then
330 reacts with NO₃⁻ to form NH₄NO₃. The k < 1 signifies NH₄⁺ excess post-reaction with
331 SO₄²⁻ and NO₃⁻. In practice, the observed excess NH₄⁺ may interact with other anions,
332 such as Cl⁻ from biomass burning, potentially forming NH₄Cl.

333 **3.2 Establishment and validation of NH₃ concentration calculation methods**



334 NH_3 emissions significantly impact atmospheric chemical processes and act as a
335 crucial precursor for the formation of secondary particulate matter (Liu et al., 2019;
336 Saraswati et al., 2019). The multiphase buffer theory clarifies that the $\text{NH}_3/\text{NH}_4^+$
337 conjugate pair core buffering system regulates aerosol acidity (Zheng et al., 2020).
338 Guanzhong Plain as a typical ammonia-rich region, elevated NH_3 emissions not only
339 suggest an increased potential for secondary aerosol formation but also indicate that
340 aerosol acidity is subject to robust internal buffering regulation (Zhang et al., 2021;
341 Zheng et al., 2023). However, the lack of online monitoring data for gaseous NH_3 at
342 the sampling site limits the accuracy of pH estimation using thermodynamic models.
343 Therefore, it is important to establish a reliable method for estimating NH_3 . In this
344 study, we adapted the methodology of Wei et al. (2023), with the researchers using
345 multiple linear regression analysis, to estimate the NHR relying on our previous
346 short-term scale datasets (Wu et al., 2020). This is because the direct estimation of
347 NH_3 is prone to interference from meteorological factors and the gas-particle
348 distribution of acidic precursors, resulting in unstable outcomes. Consequently, this
349 study opts to estimate NH_3 through NHR and assesses the method's accuracy (Meng
350 et al., 2018; Wei et al., 2023).

351 Table 2 and S2 summarize the total results of multiple linear regression analysis
352 in each season. The F-test and t-test results reveal significant correlations for all
353 seasons ($R > 0.82$, $P < 0.01$), confirming the efficacy of the established statistical
354 regression equations. Subsequently, a Pearson correlation analysis was performed to
355 elucidate the statistical associations between NHR and diverse environmental factors.
356 As shown in Fig. 3 and Table S3, NHR exhibited significant correlations ($P < 0.01$)
357 with all selected environmental factors in both winter and spring. Among these, the
358 highest correlation coefficient was observed between NO_3^- and SO_4^{2-} ($R = 0.96$, $P <$
359 0.01). Across all four seasons, NHR showed good correlations with either NO_3^- or
360 SO_4^{2-} , but with distinct seasonal variations. In winter and spring, NHR displayed the
361 highest Pearson correlation coefficients with NO_3^- ($R = 0.86$, $P < 0.01$ in winter; $R =$
362 0.74 , $P < 0.01$ in spring). Conversely, in summer and autumn, NHR correlated most
363 significantly with SO_4^{2-} ($R = 0.75$, $P < 0.01$ in summer; $R = 0.79$, $P < 0.01$ in autumn).



364 These seasonal distinctions primarily arise from fluctuations in the thermodynamic
365 stability of different ammonium salts and the seasonal alterations in their formation
366 pathways (Tao et al., 2025; Zhang et al., 2021).

367 We also accessed the reliability of MRL model results by comparing the
368 estimated NHR with the observed NHR in 2016 and 2017. The validation results (Fig.
369 4) indicated significant correlations across all seasons (with the intercept constrained
370 to zero, all $R^2 > 0.94$, $P < 0.01$). Thus, the NH_3 concentrations in this study could be
371 subsequently determined by inversely calculating the estimated NHR. Finally, these
372 NH_3 values, along with the measured NH_4^+ data, were incorporated into the
373 ISORROPIA-II model to simulate aerosol pH (refer to Section 3.3 for further details).

374 **3.3 Seasonal variation in aerosol pH and pH under different pollution levels**

375 Figure 5 shows the monthly averages of aerosol pH, aerosol liquid water content
376 (ALWC), and H_{air}^+ , as calculated by the ISORROPIA-II model during the 2022
377 sampling period. The annual average aerosol pH at the Qinling observation site was
378 3.8 ± 1.0 . The seasonal pH averages ranked as follows: winter (4.9 ± 0.4) > spring
379 (3.4 ± 0.7) > autumn (3.2 ± 0.7) > summer (2.8 ± 0.7). The average ALWC
380 concentration was $60.3 \mu\text{g m}^{-3}$, exhibiting the pattern: winter ($80.4 \pm 112.9 \mu\text{g m}^{-3}$) >
381 autumn ($76.2 \pm 121.0 \mu\text{g m}^{-3}$) > spring ($55.7 \pm 132.2 \mu\text{g m}^{-3}$) > summer (10.0 ± 25.1
382 $\mu\text{g m}^{-3}$). Overall, aerosol pH was highest in winter and lowest in summer, with spring
383 and autumn values intermediate. This trend closely aligns with the seasonal evolution
384 of $\text{PM}_{2.5}$ chemical composition and meteorological conditions, reflecting the seasonal
385 pH variation pattern observed in Beijing (Ding et al., 2019). Winter aerosol pH was
386 similar to those reported in Xi'an previously and slightly higher than those recorded
387 during the COVID-19 period (Liu et al., 2025). A comparison of aerosol acidity levels
388 across various regions in China is detailed in Table S4, revealing substantial regional
389 variability. In contrast to southeastern coastal cities such as Shanghai and Xiamen,
390 typical inland cities like Xi'an and Beijing generally display higher aerosol pH (Fu,
391 2022; Liu et al., 2025; Xu et al., 2025).

392 To further explore the seasonal characteristics in $\text{PM}_{2.5}$ acidity/alkalinity under
393 different pollution levels, this study classified $\text{PM}_{2.5}$ mass concentrations into two



394 categories based on ambient air quality standards of China: clean days ($< 75 \mu\text{g m}^{-3}$)
395 and polluted days ($> 75 \mu\text{g m}^{-3}$). A systematic comparison of aerosol pH and ALWC
396 variations corresponding to these pollution levels across different seasons was
397 performed (Table S5). Under clean conditions, aerosol pH ranged from 2 to 5,
398 whereas under polluted and heavily polluted conditions, it was primarily concentrated
399 between 3 and 6. As air quality declined and $\text{PM}_{2.5}$ concentrations increased from
400 clean to polluted days, the concentrations of aerosol components and pH values
401 exhibited an upward trend across all seasons. However, the behaviors of RH, ALWC,
402 and H_{air}^+ displayed seasonal discrepancies. In spring and summer, RH, ALWC, and
403 H_{air}^+ decreased as $\text{PM}_{2.5}$ concentrations increased. Conversely, during autumn and
404 winter, these parameters increased alongside rising $\text{PM}_{2.5}$ concentrations. These
405 findings reveal fundamental differences in the dominant chemical mechanisms of
406 $\text{PM}_{2.5}$ pollution formation across seasons.

407 During spring and summer, pollution formation is mainly controlled by
408 gas-phase photochemical oxidation (Chen et al., 2025; Li et al., 2021). The enhanced
409 solar radiation and high temperatures during these seasons collectively create a highly
410 oxidizing atmospheric environment. While promoting the transformation of
411 precursors, these conditions also drive the decomposition of semi-volatile NH_4NO_3 ,
412 leading to the release of gaseous HNO_3 and NH_3 (Guo et al., 2018; Tao et al., 2025).
413 Consequently, despite an increase in $\text{PM}_{2.5}$ concentrations, both ALWC and H_{air}^+
414 concentrations exhibit a decreasing or stabilizing trend due to the reduction of
415 hygroscopic nitrate, while RH remains relatively low because of higher T. In contrast,
416 autumn and winter are characterized by pollution processes driven by heterogeneous
417 and aqueous-phase chemical reactions. The cold temperatures and stagnant weather
418 conditions lead to high RH, promoting the aqueous-phase oxidation of pollutants and
419 inhibiting the evaporation of NH_4NO_3 (Wang et al., 2016; Zhang et al., 2021). This
420 results in the significant formation of highly hygroscopic NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$,
421 which triggers an increase in ALWC. This increase further intensifies heterogeneous
422 reactions and concentrates acidic products, leading to a concurrent rise in H_{air}^+
423 concentration. Consequently, a complex pollution scenario emerges, characterized by



424 a simultaneous increase in $PM_{2.5}$ ALWC, and H_{air}^+ concentrations.

425 **3.4 Sensitivity analysis for aerosol pH**

426 Sensitivity analysis of aerosol pH and ALWC in response to meteorological
427 conditions and chemical composition is a crucial step for understanding and
428 predicting atmospheric environmental changes. This analysis not only clarifies the
429 nonlinear processes involved in secondary aerosol formation but also assesses the
430 impact of variations in anthropogenic emissions on atmospheric chemical systems
431 (Nah et al., 2023; Zhou et al., 2022). In this study, sensitivity analyses were performed
432 for each season, incorporating both particulate-phase compositions and
433 meteorological factors. The effects of SO_4^{2-} , NO_3^- , NH_x (gas NH_3 + NH_4^+), Cl^- , Ca^{2+} ,
434 K^+ , Mg^{2+} , RH, and T on aerosol pH, ALWC, and H_{air}^+ were evaluated, as illustrated in
435 Fig. 6. The primary drivers influencing aerosol pH variation across all seasons were
436 NH_x , SO_4^{2-} , and T, while Ca^{2+} in spring and RH in summer also emerged as important
437 influencing factors. This pattern is consistent with the drivers identified in the Beijing
438 area (Ding et al., 2019). Regarding ALWC, RH emerged as the most significant factor,
439 followed by NO_3^- and SO_4^{2-} .

440 Sulfate as a key component of secondary inorganic aerosols, is one of the most
441 significant acidic sources in atmospheric fine particulate matter. Sulfuric acid,
442 generated through heterogeneous oxidation processes, is a strong acid that directly
443 increases the concentration of H_{air}^+ in the particulate phase and reduces the pH (Shah
444 et al., 2018; Zhang et al., 2021). As illustrated in Fig. 6, when the levels of alkaline
445 substances such as NH_3 are maintained constant, aerosol pH significantly declines
446 with increasing SO_4^{2-} concentration. The concurrent rise in ALWC and H_{air}^+ further
447 substantiates its direct role in driving acidification processes (Ding et al., 2019). The
448 increase in ALWC and H_{air}^+ is particularly pronounced during summer and autumn.
449 This sustained rise continuously depletes alkaline buffering substances in the
450 particulate phase, resulting in heightened hydrogen ion activity and increased acidity
451 (Guo et al., 2018). Conversely, while nitrate similarly elevates H_{air}^+ and ALWC, its
452 overall impact on pH is generally weaker than that of sulfate due to its high volatility
453 (Ding et al., 2019). Except during summer when the pH remained relatively stable,



454 pH values decreased with rising NO_3^- concentrations in all other seasons.

455 Ammonia (NH_3), the primary alkaline gas, plays a crucial role in buffering
456 aerosol pH. In ammonia-rich environments, NH_3 effectively counteracts acidic
457 components, stabilizing pH and resulting in a nonlinear pH response to precursor
458 concentrations (Cui et al., 2025; Karydis et al., 2021; Niu et al., 2016).
459 Simultaneously, the dissolution of NH_3 into the particle phase to form NH_4^+ markedly
460 enhances the hygroscopicity of particles, causing an increase in ALWC as NH_x
461 concentrations rise (Jia et al., 2020; Liu et al., 2019). Studies indicate that NH_3 first
462 reacts with H_2SO_4 and subsequently with HNO_3 to form NH_4NO_3 . After most nitrate
463 is converted to NH_4NO_3 , it becomes difficult to dissolve additional NH_3 into aerosol
464 droplets (Ding et al., 2019; Xu et al., 2025). The results of this study (Fig. 6) show
465 that the response of pH to NH_x is non-linear. When the NH_x mass concentration is
466 within the range of 3–10 $\mu\text{g m}^{-3}$, aerosol pH increases significantly, H_{air}^+ decreases
467 sharply, and ALWC increases with rising NH_x concentration. Nevertheless, when the
468 NH_x concentration surpasses a specific threshold, the trends in pH and H_{air}^+ decelerate
469 and subsequently remain relatively stable. It should be noted that the results may be
470 less accurate when NH_x concentration is below 3 $\mu\text{g m}^{-3}$, as high temperature and
471 intense solar radiation promote the thermal decomposition of semi-volatile
472 ammonium salts. The seasonal variations observed in this study further substantiate
473 the regulatory role of atmospheric ammonia.

474 Ca^{2+} is an important crustal element in mineral dust, displaying distinct regional
475 chemical characteristics and effects on aerosol acidity. According to the
476 ISORROPIA-II model, Ca^{2+} typically combines with SO_4^{2-} to form the less soluble
477 CaSO_4 solid phase, reducing its solubility and diminishing its effectiveness for
478 neutralizing atmospheric acidity (Fountoukis et al., 2007; G. Wang et al., 2016). In the
479 Guanzhong Plain, the regulatory impact of Ca^{2+} on aerosol pH is particularly crucial,
480 especially due to the region's proximity to the Loess Plateau and its semi-arid climate.
481 Dust transport from the northwest could shape the alkaline atmospheric conditions in
482 this area. This phenomenon results in neutral aerosol pH in the free troposphere and
483 contributes to the alkalinity of cloud water at high-altitude sites (Liu et al., 2024; Shen



484 et al., 2024). Elevated concentrations of Ca^{2+} can increase aerosol pH by decreasing
485 both ALWC and H_{air}^+ , with the most pronounced impact observed during the frequent
486 dust events of spring. These events transport significant amounts of alkaline mineral
487 dust, enhancing acid neutralization, consistent with studies in northern China (Shi et
488 al., 2019; Tan et al., 2018). This study revealed that the relative standard deviation
489 (RSD) for Ca^{2+} concentration variation in spring reached 8.8%, notably higher than
490 the corresponding RSD of 7.5% reported for the North China Plain during the same
491 season (Ding et al., 2019). Conversely, in southern coastal urban areas like Xiamen
492 and Guangzhou, which lack robust dust origins, the impact of Ca^{2+} on aerosol pH is
493 typically negligible (Jia et al., 2020; Xu et al., 2025). This finding directly confirms
494 the greater influence of dust sources in the Guanzhong Plain.

495 In addition to the particulate chemical composition, meteorological conditions,
496 specifically RH and T, significantly influence aerosol acidity. This study revealed that
497 during winter, spring, and autumn, aerosol pH displayed a non-linear response to
498 increasing RH, characterized by an initial decrease followed by a subsequent increase.
499 The inflection point typically occurred within the RH range of 60%–85%. Beyond
500 this critical threshold, ALWC entered a phase of exponential growth. This
501 phenomenon is probably driven by the deliquescence of the predominant hygroscopic
502 components within the aerosol, primarily NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$, which have
503 deliquescence relative humidities of approximately 62% and 80% (Tang, 1996).
504 Given that actual atmospheric particles are mixtures of various components, their
505 hygroscopic and phase-state characteristics are more complicated (Su et al., 2022;
506 Zheng et al., 2020). In the initial stage, an increase in RH results in a higher ALWC,
507 which facilitates the heterogeneous uptake and aqueous-phase oxidation of acidic
508 precursors, resulting in the formation of strong acids. During this phase, the
509 acid-forming effect dominates the pH change (Du et al., 2020; Zhang et al., 2021).
510 Once RH exceeds a critical threshold, ALWC exhibits an exponential growth trend. At
511 this point, the intense dilution effect of RH on H^+ begins to surpass the ongoing
512 acid-forming effect (Karydis et al., 2021). Furthermore, this hygroscopic growth is
513 self-reinforcing, i.e. the increased ALWC provides a more extensive aqueous medium



514 for heterogeneous reactions of gaseous precursors, leading to further formation of
515 hygroscopic salts and further promoting ALWC growth, and thus establishing a
516 positive feedback loop (Su et al., 2022). Concurrently, under high RH conditions,
517 thermodynamic factors favor the partitioning of semi-volatile ammonium salts into
518 the particle phase. This neutralization process further depletes H^+ (Tao et al., 2025).
519 These combined effects ultimately lead to the observed recovery in pH. Ambient
520 temperature can influence aerosol pH by affecting gas-particle partitioning (Ding et
521 al., 2019). The evaporation of semi-volatile ammonium salts plays a crucial role in
522 aerosol acidification by transferring H^+ to the particle phase, consequently decreasing
523 pH (Guo et al., 2018). Elevated ambient temperatures intensify this acidification
524 mechanism synergistically. Thermodynamically, elevated temperatures drive the
525 volatilization process. Concurrently, they reduce ALWC to concentrate H^+ and further
526 enhancing acidity.

527 **3.5 Quantitative analysis of pH driving factors**

528 To quantitatively analyze aerosol pH, the contributions of various influencing
529 factors to aerosol pH and ALWC were further quantified (Fig. 7 and Fig. S4).
530 According to the multiphase buffer theory (Zheng et al., 2020), T significantly drives
531 the seasonal variation of aerosol pH by influencing pK_a^* . The results presented in Fig.
532 7 confirm this assertion, indicating that T plays a dominant role (22.3%–31.6%) in
533 driving the seasonal variation of aerosol pH. T favored enhanced aerosol acidity in
534 winter and spring, while the opposite trend was observed in summer and autumn.
535 Moreover, NH_x emerged as a key factor influencing aerosol pH during winter and
536 autumn, contributing between 22.9% and 44.9%. When the higher winter NH_x
537 concentration was replaced with the lower concentration observed in spring, while
538 maintaining other winter parameters constant, the simulated ΔpH was negative,
539 indicating increased acidity. In the winter, the widespread use of coal for heating leads
540 to significant emissions of acidic gases (Wang et al., 2016). The decrease in
541 concentrations may exacerbate acidification through two mechanisms. On one hand, it
542 weakens the basic buffering capacity driven by NH_3 , reducing its neutralizing effect
543 on sulfuric and nitric acids. On the other hand, it disrupts the gas-particle equilibrium



544 of NH_4NO_3 , driving the equilibrium towards dissociation and resulting in the
545 production of more gaseous HNO_3 . These two processes together contribute to a
546 decline in system pH. In contrast, the opposite trend is observed in the autumn. The
547 primary influencing factors in spring were T at 31.9% and NVCs at 24.7%, while in
548 summer, they were T at 33.8% and RH at 17.8%. These results suggest that NVCs,
549 including Ca^{2+} , serve as significant drivers of aerosol pH variation in spring, while RH
550 plays a similar role in summer. This observation is consistent with findings from
551 studies conducted in other northern Chinese cities (Ding et al., 2019). Moreover, the
552 synergistic effects of various influencing factors leading to $\Delta\text{pH}_{\text{others}}$ are also
553 significant, particularly in autumn, where they contributed 26.6% to aerosol pH
554 variation.

555 Throughout the observation period, chemical components contributed an average
556 of 47.7% annually to pH, while meteorological factors contributed 35.7%,
557 highlighting the greater impact of chemical components over meteorological factors.
558 This characteristic significantly differs from that observed in southeastern coastal
559 regions of China. For example, studies conducted in Xiamen and Guangzhou have
560 demonstrated that the contribution of meteorological factors frequently surpasses that
561 of chemical components (Jia et al., 2020; Xu et al., 2025). Seasonally, chemical
562 components were more influential in winter (58.7%), spring (42.7%), and autumn
563 (49.2%). Conversely, meteorological factors displayed their most significant impact in
564 summer, with a contribution rate reaching 51.6%, thus serving as the primary driver
565 of pH variation during this season.

566 **4 Conclusions**

567 This study evaluates the aerosol acidity across four seasons in the Guanzhong
568 region of China and quantitatively analyzes the factors contributing to seasonal pH
569 variations. Over the study period, $\text{PM}_{2.5}$ in the study area was generally weakly
570 alkaline, with an annual mean aerosol pH of 3.8 ± 1.0 . This trend corresponds with the
571 seasonal changes in $\text{PM}_{2.5}$ chemical composition and meteorological conditions.
572 Aerosol pH increased alongside rising $\text{PM}_{2.5}$ concentrations, ranging from 2–5



573 (average: 3.4 ± 0.9) on clean days to 3–6 (4.8 ± 0.5) on polluted and heavily polluted
574 days. Sensitivity analysis and quantitative analysis indicated that NH_x , SO_4^{2-} , RH, and
575 T are key drivers driving pH across all seasons. T emerged as the primary driver of
576 seasonal pH fluctuations (contributing 22.3%–31.6%), with NH_x playing a significant
577 role in autumn and winter (contributing 22.9%–44.9%), while Ca^{2+} in spring and RH
578 in summer acted as specific drivers for those seasons. pH showed higher sensitivity to
579 SO_4^{2-} than to NO_3^- . Furthermore, NH_x not only displayed substantial buffering
580 capacity on pH but also, upon conversion to NH_4^+ in the particle phase, enhanced
581 particle hygroscopicity, resulting in increased ALWC with higher NH_x concentrations.
582 Meteorological conditions, beyond chemical components, exerted a non-linear
583 regulatory influence on acidity. During winter, spring, and autumn, aerosol pH
584 initially decreases and subsequently increases with rising RH, exhibiting an inflection
585 point within the RH range of 60%–85%. This behavior might be related to the
586 explosive growth in ALWC that occurs when the predominant hygroscopic
587 components, primarily NH_4NO_3 and $(\text{NH}_4)_2\text{SO}_4$, reach their mixed deliquescence
588 point. Our results discovered the main driving factors of aerosol acidity and alkalinity
589 in the Guanzhong Plain, which would be helpful for deeply understanding the
590 heterogenous formation of atmospheric secondary aerosols in the semi-arid region.

591 **Data availability**

592 The data in this study are available at <https://zenodo.org/records/18455742>
593 (Wang et al., 2026).

594 **Author contributions**

595 Jianjun Li conceived and designed the study, and the revision of the manuscript.
596 Qian Wang conducted the literature search, performed sample and data analysis, and
597 wrote the manuscript. Xiao Guo, Minxia Shen, Yali Liu, Yifan Zhang, Lu Li, Weining
598 Qi, Yue Cao, Shicong Li, Zhuoer Dong and Wenting Dai collected particulate samples
599 and supervised the experiments. All authors provided critical feedback on the
600 manuscript and approved the final version.

601 **Conflict of Interest**



602 The authors declare no conflicts of interest relevant to this study.

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609



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859 **Table 1 Multiple linear regression results between NHR and its drivers (major particulate**
 860 **ions and meteorological factors) in each season.**

	Annual	Winter	Spring	Summer	Autumn
PM _{2.5} (µg m ⁻³)	65.3 ± 61.8	93.5 ± 41.7	61.4 ± 40.8	33.3 ± 13.7	43.1 ± 25.1
PM ₁₀ (µg m ⁻³)	97.5 ± 66.3	146.0 ± 61.1	87.7 ± 66.6	39.3 ± 14.3	94.2 ± 49.1
T (K)	289.3 ± 10.1	278.2 ± 4.2	291.1 ± 5.3	302.0 ± 4.0	290.3 ± 5.1
RH (%)	66.5 ± 20.2	65.5 ± 20.9	62.2 ± 19.8	61.2 ± 19.1	80.0 ± 14.8
WS (m s ⁻¹)	0.7 ± 0.3	0.7 ± 0.2	0.8 ± 0.3	0.9 ± 0.3	0.6 ± 0.2
SO ₂ (µg m ⁻³)	8.9 ± 4.6	10.0 ± 4.5	4.9 ± 2.0	8.6 ± 3.4	11.9 ± 5.0
CO (mg m ⁻³)	0.7 ± 0.4	1.1 ± 0.5	0.5 ± 0.1	0.4 ± 0.1	0.6 ± 0.2
NO ₂ (µg m ⁻³)	25.9 ± 15.1	37.4 ± 13.2	25.0 ± 8.4	9.8 ± 4.5	26.4 ± 14.4
O ₃ (µg m ⁻³)	99.6 ± 50.1	67.8 ± 33.2	109.5 ± 40.2	150.2 ± 43.2	83.4 ± 39.7
WSIIs (µg m ⁻³)	24.5 ± 20.5	39.5 ± 22.0	18.6 ± 12.9	8.5 ± 4.2	19.8 ± 15.9
SNA (µg m ⁻³)	21.8 ± 19.4	35.5 ± 21.1	16.1 ± 12.3	7.0 ± 3.7	18.1 ± 15.5
SO ₄ ²⁻ (µg m ⁻³)	5.1 ± 3.7	7.3 ± 4.7	4.1 ± 1.7	3.4 ± 2.1	4.4 ± 3.1
NO ₃ ⁻ (µg m ⁻³)	8.8 ± 10.8	16.3 ± 11.9	6.1 ± 7.0	1.0 ± 1.2	8.7 ± 10.8
NH ₄ ⁺ (µg m ⁻³)	4.0 ± 4.7	7.3 ± 5.7	2.4 ± 2.6	1.1 ± 0.8	3.8 ± 4.1
Cl ⁻ (µg m ⁻³)	0.5 ± 0.7	1.1 ± 0.8	0.3 ± 0.2	0.1 ± 0.1	0.3 ± 0.3
Ca ²⁺ (µg/m ³)	1.4 ± 1.6	1.8 ± 1.8	1.9 ± 1.6	0.8 ± 0.6	0.6 ± 0.3
K ⁺ (µg m ⁻³)	0.3 ± 0.4	0.6 ± 0.4	0.2 ± 0.1	0.1 ± 0.1	0.2 ± 0.2
Mg ²⁺ (µg m ⁻³)	0.1 ± 0.1	0.2 ± 0.1	0.2 ± 0.1	0.1 ± 0.1	0.1 ± 0.1
Na ⁺ (µg/m ³)	0.7 ± 0.4	0.9 ± 0.5	0.7 ± 0.2	0.6 ± 0.2	0.6 ± 0.1
NH ₃ (µg m ⁻³)	12.8 ± 10.3	22.6 ± 10.0	5.9 ± 4.3	12.1 ± 5.5	4.5 ± 3.9
NH _x (µg m ⁻³)	16.7 ± 14.1	29.7 ± 15.1	8.3 ± 6.8	13.2 ± 6.3	8.4 ± 7.9
NHR	0.2 ± 0.1	0.2 ± 0.1	0.2 ± 0.1	0.1 ± 0.1	0.4 ± 0.1

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862

863 **Table 2 Multiple linear regression results between NHR and its drivers (major particulate**
 864 **ions and meteorological factors) in each season.**

Season	Prediction Equation	R	Sig.
Winter	NHR = 1.204 - 0.003SO ₄ ²⁻ + 0.006NO ₃ ⁻ + 0.023K ⁺ + 0.328Mg ²⁺ - 0.017Ca ²⁺ - 0.004T + 0.001RH	0.88	P < 0.01
Spring	NHR = 0.533 + 0.006SO ₄ ²⁻ + 0.005NO ₃ ⁻ - 0.018K ⁺ - 0.065Mg ²⁺ - 0.017Ca ²⁺ - 0.001T - 0.009RH	0.83	
Summer	NHR = -1.389 + 0.007SO ₄ ²⁻ + 0.006NO ₃ ⁻ + 0.01K ⁺ + 0.021Mg ²⁺ - 0.003Ca ²⁺ + 0.005T - 0.066RH	0.82	
Autumn	NHR = 0.656 + 0.006SO ₄ ²⁻ + 0.003NO ₃ ⁻ - 0.003K ⁺ + 0.064Mg ²⁺ - 0.040Ca ²⁺ - 0.001T + 0.012RH	0.86	

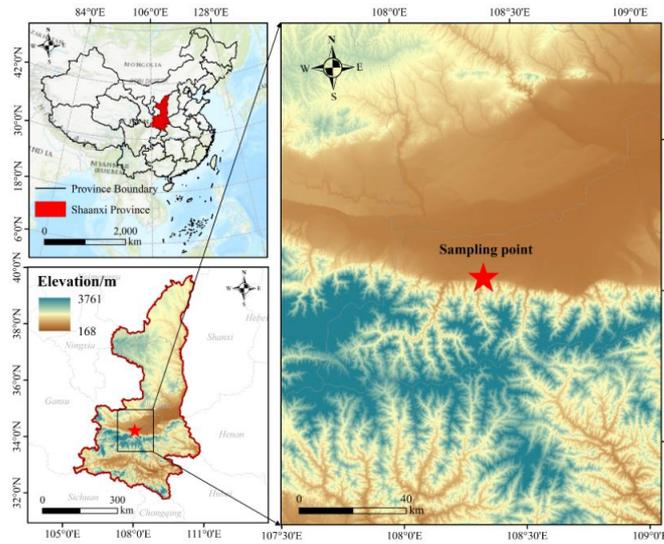
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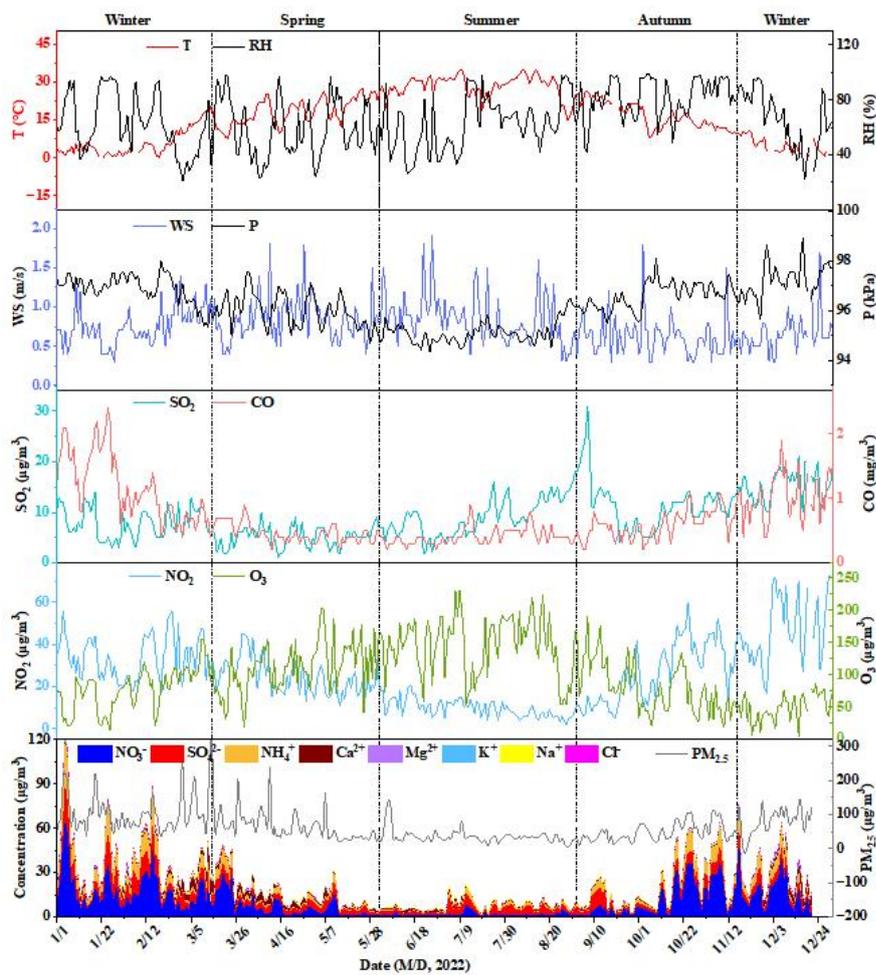
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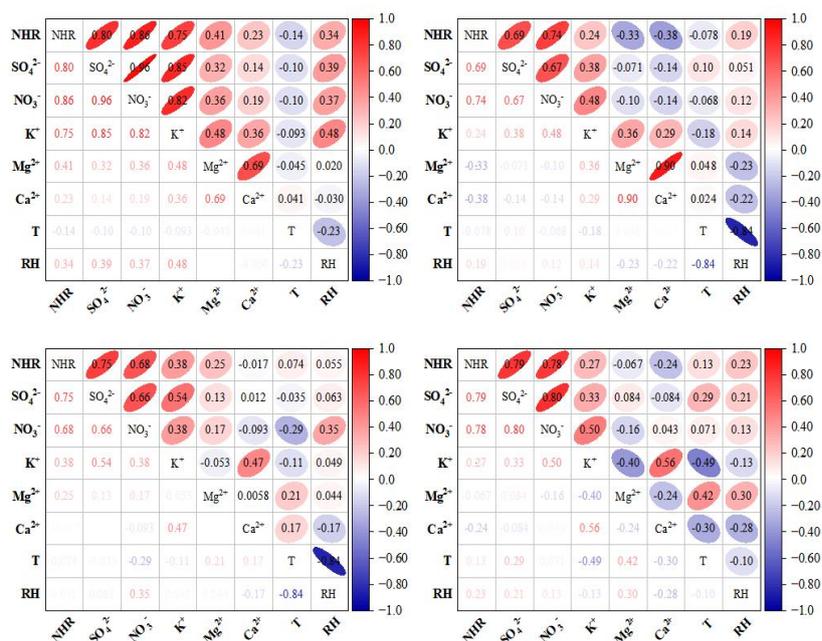
Fig.1 Locations of monitoring stations in Qinling



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874 **Fig. 2** The time series data for PM_{2.5} mass concentrations, criteria air pollutants (SO₂, CO,
875 **NO₂, and O₃-8h**), and meteorological parameters (T, RH, WS, and AP).

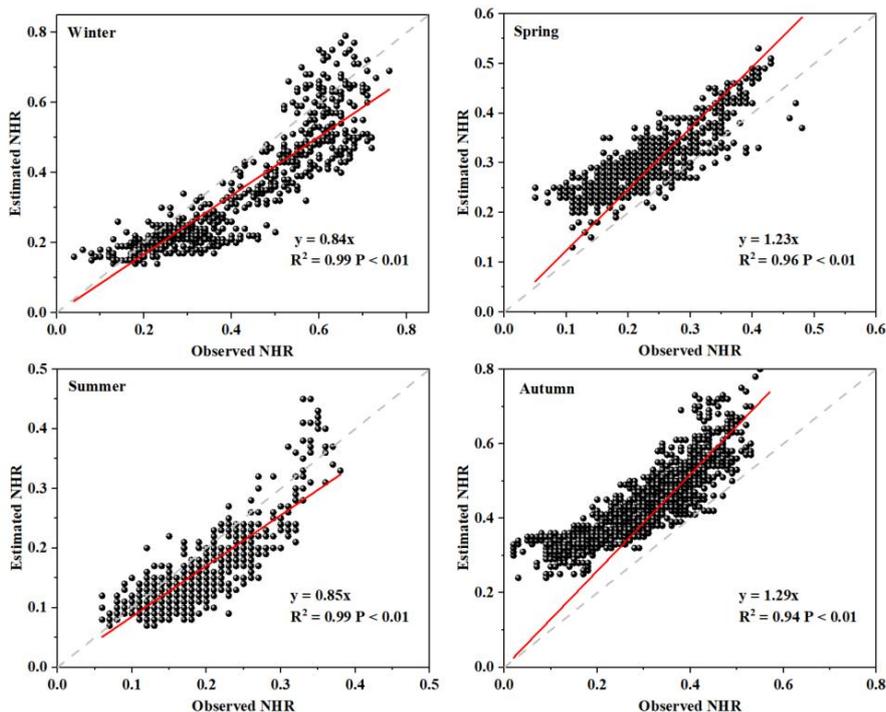
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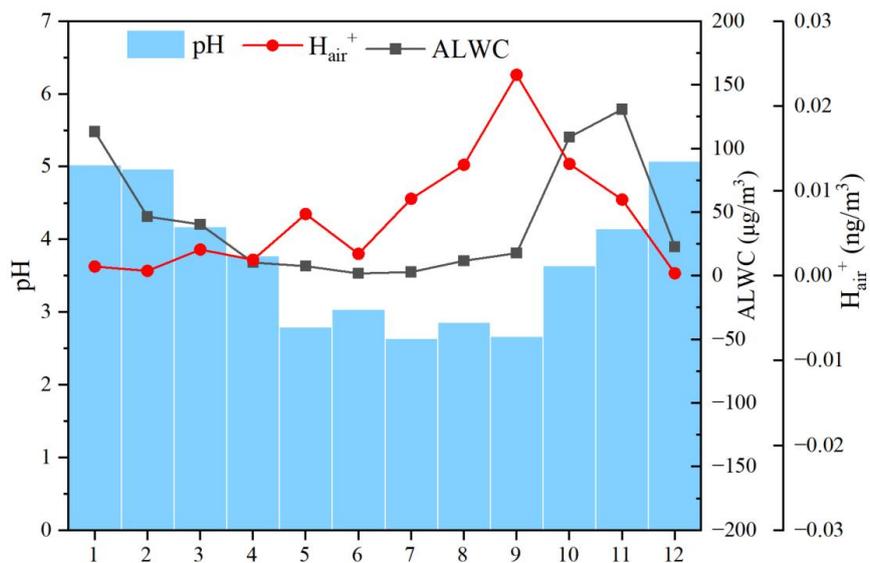
Fig. 3 Pearson correlation among various factors



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Fig. 4 Correlation between estimated and observed daily average values of NHR.



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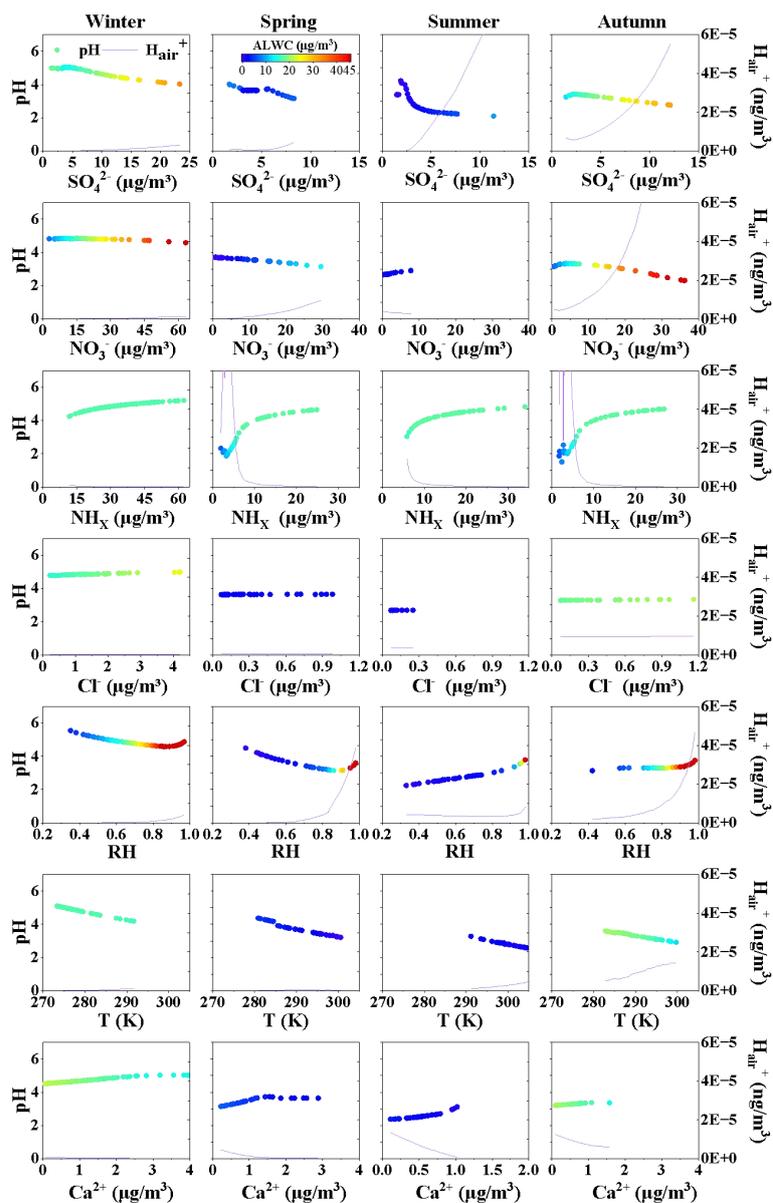
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Fig.5 Monthly average aerosol pH, ALWC and H_{air}^+ changes



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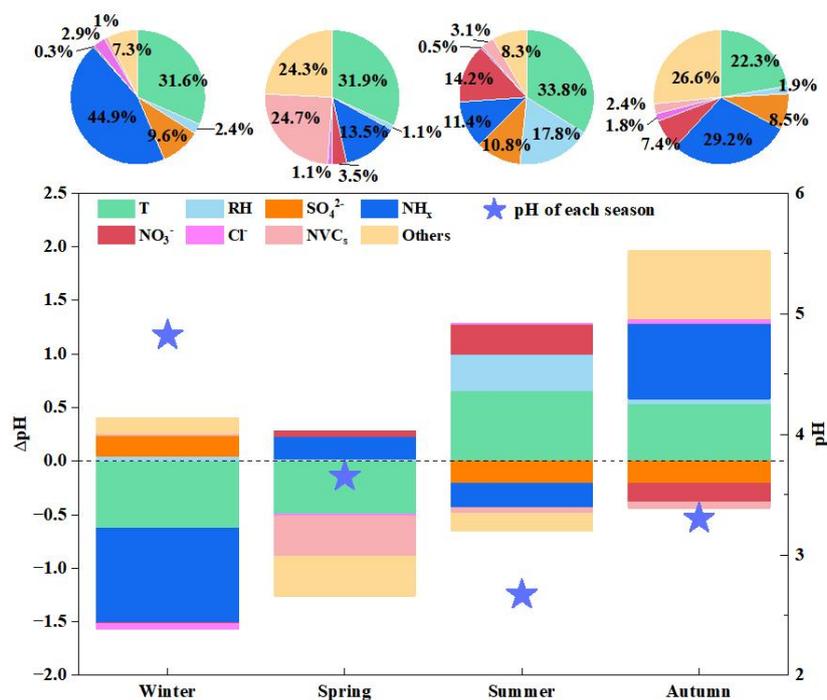
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Fig. 6 Sensitivity analysis of influencing factors



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890 **Fig. 7** The contribution of various influencing factors to the seasonal pH. The asterisk
 891 represents the pH of aerosols in each season. The pie chart shows the relative contributions
 892 of various influencing factors to the seasonal pH (absolute values of pH was taken for pie
 893 chart).