



1 **Technical note: Cavity ring-down spectroscopy is a precise and accurate method for**
2 **measuring $^{15}\text{N-NO}_3^-$ in soil extracts from ^{15}N tracer experiments**

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8 **Abstract.** Cavity ring-down spectroscopy (CRDS) provides a low-cost alternative to isotope
9 ratio mass spectrometry (IRMS) for the analysis of $^{15}\text{N-NO}_3^-$ in nitrogen cycle assays, but its
10 performance has not been tested systematically for highly enriched samples. Here we report
11 results from 271 parallel analyses of $^{15}\text{N-NO}_3^-$ in soil extracts after chemical conversion to N_2O ,
12 including low (0.36-1.45 at%) and high labelled samples (0.46-8.49 at%). We demonstrate that
13 CRDS measurements deliver comparable accuracy and precision to continuous-flow Gas
14 Bench-IRMS analysis (slope: 0.982 and 0.972, RMSE: 0.0018 and 0.254 at% for low- and
15 high-labelled samples, respectively), thereby offering a robust and affordable alternative for
16 classic $^{15}\text{N-NO}_3^-$ analysis.

17 **Keywords:** Gross nitrification rate, soil nitrogen cycle, ^{15}N tracing, isotope pool dilution, $^{15}\text{N-}$
18 nitrate

19 Accurate determination of soil nitrogen (N) transformation rates has become increasingly
20 critical due to the global intensification of the N cycle, which stems from mineral fertilizer
21 application, land-use changes, and climate-driven acceleration of N turnover in natural
22 ecosystems (Fowler et al., 2013; Gruber & Galloway, 2008). Within this framework,
23 nitrification represents a pivotal process governing terrestrial N retention and loss, as it directly
24 drives both the emission of gaseous N species to the atmosphere and the leaching of mobile
25 nitrate (NO_3^-) into aquatic systems (Niu et al., 2016). Quantification of gross soil N
26 transformation rates routinely relies on stable isotope assays utilizing ^{15}N . These approaches
27 encompass isotope pool dilution techniques, which measure the dilution of an added ^{15}N label
28 within a specific inorganic pool, and tracer experiments, which track the downstream
29 transformation of a labelled species into another pool (Jia et al., 2022). In both methodologies,
30 analysing the $^{15}\text{N}:^{14}\text{N}$ ratio of NO_3^- is fundamental to resolving specific transformation
31 pathways (Jia et al., 2022). Conventionally, this isotope analysis requires the conversion of



32 NO_3^- to nitrous oxide (N_2O) via chemical reduction (Altabet et al., 2019; Lachouani et al.,
33 2010) or microbiological denitrification (Casciotti et al., 2002; Sigman et al., 2001). The
34 product N_2O is subsequently separated by gas chromatography, cryogenically preconcentrated,
35 and analysed via isotope ratio mass spectrometry (IRMS).

36 Despite its widespread application, conventional IRMS analysis is constrained by significant
37 operational and financial bottlenecks. These systems are expensive and complex to maintain.
38 They require specialized experts to manage the continuous use of high-purity helium, manual
39 liquid nitrogen refills, and intricate parts like gas benches and cryogenic traps. In contrast,
40 cavity ring-down spectroscopy (CRDS) is emerging as an accessible alternative for the stable
41 isotope analysis of ^{15}N . The CRDS method offers substantially lower capital and running costs,
42 utilizing small volumes of N_2O free air rather than expensive carrier gases. The simplified
43 architecture of this instrument minimizes maintenance requirements, eliminating the need for
44 cryogenic liquids and complex moving parts. While a recent study has demonstrated the
45 efficiency of CRDS for evaluating NO_3^- stable isotopes at natural abundance levels (McKay et
46 al., 2026), its performance has not yet been systematically cross-evaluated against IRMS using
47 ^{15}N -labelled tracer samples. Furthermore, neither analytical platform is natively optimized to
48 accommodate highly enriched isotope signatures, which can introduce analytical challenges
49 such as memory effects or detector saturation. Therefore, a direct comparison between these
50 two independent methods is essential. This will help establish clear operating limits and ensure
51 accurate data at high enrichment levels typical for gross N transformation rate determination.

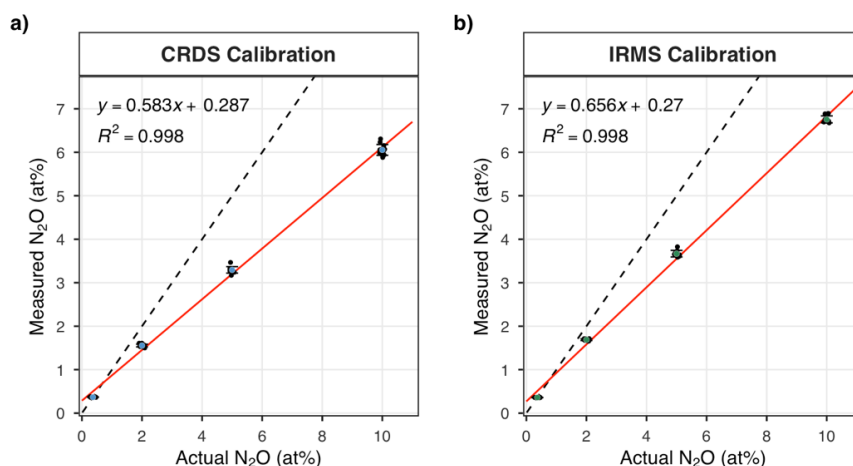
52 Here, we present a method of intercomparison study comprising 271 soil extracts analyzed for
53 the $^{15}\text{N}:^{14}\text{N}$ ratios in NO_3^- . The sample material originated from a ^{15}N tracer experiment
54 designed to determine mineral N transformation rates in permafrost sediments from disturbance
55 sites in the West Alaskan Yedoma region. The experimental design included two labelling
56 treatments, where either NH_4^+ or NO_3^- pool was enriched with ^{15}N . The NH_4^+ treatment
57 resulted in a set of samples with lower ^{15}N enrichment in NO_3^- as all the $^{15}\text{N}\text{-NO}_3^-$ present in
58 the sample came from the conversion of $^{15}\text{N}\text{-NH}_4^+$ by nitrification. We amended soil with
59 ammonium at either 10 or 25 at% $^{15}\text{N}\text{-NH}_4^+$, targeting a ^{15}N abundance of 1–10 at% in the soil
60 NH_4^+ pool. This resulted in NO_3^- extracts with an at% ranging between 0.36 to 1.45 at% ^{15}N .
61 The second set of samples with higher ^{15}N enrichment originated from the NO_3^- treatment,
62 where we targeted an initial enrichment of 1–10 at% in the soil nitrate pool and subsequent
63 dilution of the label according to the principle of isotope pool dilution experiment as NO_3^- is
64 produced from native non-labeled NH_4^+ (Di et al., 2000). This was achieved by amending the



65 soil samples with nitrate at either 10 or 98 at% $^{15}\text{N}\text{-NO}_3^-$ and resulted in NO_3^- extracts ranging
 66 0.46–8.49 at%. Samples were incubated in the dark at 5°C for 4, 24, 48, or 72 hours followed
 67 by extraction with 1M KCl at a 1:3 (v/v) soil-to-solution ratio. Samples were shaken for 1 hour
 68 at 21°C and 175 rpm and subsequently filtered using paper filters (Whatman®-Cytiva, Global
 69 Life Sciences, UK).

70 Nitrate in the extracts was converted to N_2O using the Ti (III) Cl_3 method (Altabet et al., 2019),
 71 with a detailed description given in the supplementary methods. Prior to the reaction, NO_3^-
 72 concentrations in the KCl extracts (0–6000 $\mu\text{mol NO}_3^-\text{-N L}^{-1}$) were equalized to a target
 73 concentration of 14.3 $\mu\text{mol NO}_3^-\text{-N L}^{-1}$ by diluting with 1M KCl. Samples falling below the
 74 target concentration were excluded. Titanium reduction reactions were conducted in 12 mL
 75 glass exetainer vials (Labco®, Lampeter, UK). Within each reaction batch of 70–80 number of
 76 samples, we included standards (STDs, also in 1M KCl) with a target concentration and four
 77 different at% levels (natural abundance, 2, 5, and 10 at%) as well as analytical blanks (1M
 78 KCl). For isotope analysis, 1 mL of headspace gas was transferred from each reaction vial into
 79 a separate secondary exetainer vial, evacuated and prefilled with 25 mL N_2O free synthetic air
 80 (5.0 purity). To prevent cross-contamination, a separate syringe was used for each treatment.
 81 The secondary exetainer was analyzed via CRDS, while the remainder of the gas in the original
 82 reaction exetainer was analyzed by GB-precon-IRMS.

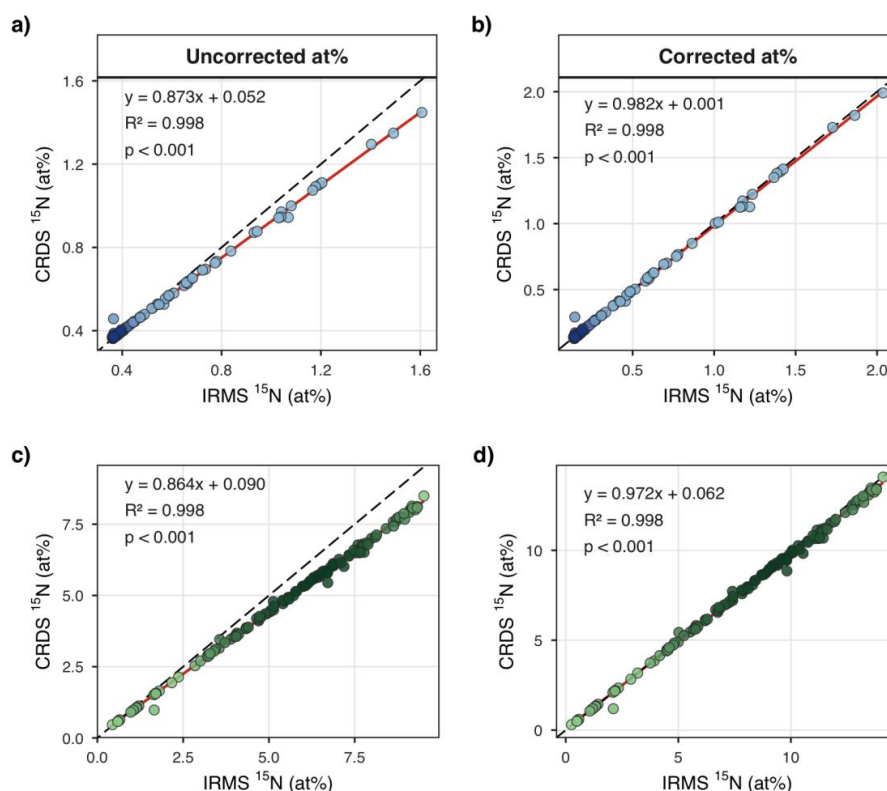
83 CRDS measurements were performed using a Picarro G5131-i isotopic N_2O analyzer coupled
 84 with a custom-built autosampler (Open Autosampler SAM, Frank Krijnen, University of
 85 Saskatchewan, Wozniak, 2025). To introduce samples to the CRDS analyzer, the transfer line
 86 and autosampler needle were flushed with synthetic air for 3 min. The autosampler needle was
 87 then lowered into a vial, and the sample was pulled into the analyzer inlet through a moving
 88 needle for 8 min. The instrument regulates the cavity pressure through a metered valve placed
 89 between cavity and outlet vacuum pump. Placing the needle in a vial therefore means that
 90 sample air is pulled into the cavity until the pressure in the vial drops to cavity pressure, at
 91 which point the instrument measures the air volume in the cavity until the needle is retracted
 92 to ambient air. Isotope ratios were calculated by averaging the measured $\delta^{15}\text{N}$ ratios during the
 93 last 6 minutes of the measurement. Complementary GB-Precon-IRMS analyses were
 94 conducted using a Thermo Delta V Advantage IRMS equipped with Gas Bench II and a PreCon
 95 unit. Finally, raw isotope ratios were converted to at% and calibrated against standards using
 96 batch-specific linear regressions (Figure 1). This calibration also served to correct the
 97 significant blank caused by N contamination present in the sample matrix (1M KCl, Figure S1).



98

99 **Figure 1.** Calibrations for N₂O atom percent (at%) analysis by **a)** CRDS (Picarro G5131-i) and
 100 **b)** IRMS (GB-Precon-IRMS) based on standards with natural abundance, 2 at%, 5 at%, and 10
 101 at% of ¹⁵N-N₂O. Solid lines show the linear regression fit, and dashed lines show the 1:1
 102 relationship (perfect agreement between measured and actual values). Small grey circles are
 103 individual measurements of each standard; the filled circle (black) is the mean, with error bars
 104 representing standard deviation. Regression equations and R² values are shown in each panel.

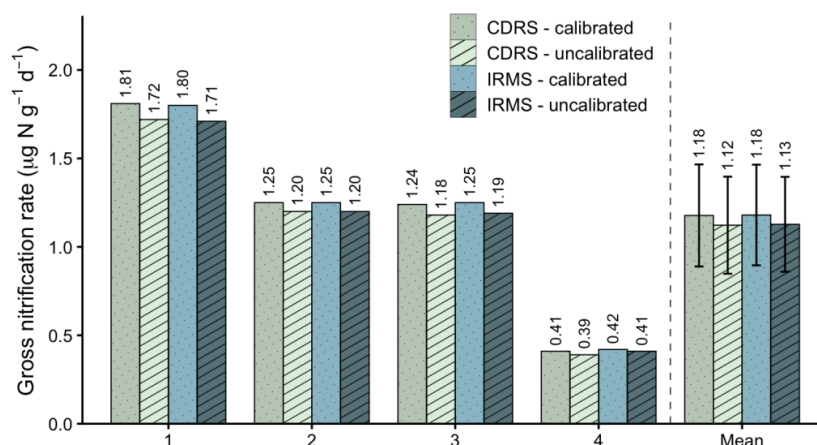
105 We found strong agreement between IRMS and CRDS measurements across all samples (R²=
 106 0.998 for each treatment/both treatments) (Figure 2). For uncorrected at% data, the precision
 107 (root mean square error; RMSE) was 0.036 at% for the low labelled samples from the ¹⁵NH₄⁺
 108 treatment and 0.745 at% for the highly labelled samples from the ¹⁵NO₃⁻ treatment, with an
 109 accuracy (slope) of 0.873 ± 0.003 and 0.864 ± 0.004 , respectively (**Figure 2a, c**). Applying the
 110 at% correction narrowed the data scatter, improving precision to 0.0018 at% (NH₄⁺) and 0.254
 111 at% (NO₃⁻), while the slopes shifted to 0.982 ± 0.004 and 0.972 ± 0.004 (**Figure 2b, d**).
 112 Because neither instrument is optimized for highly labelled samples (>3 at%), this excellent
 113 agreement across a wide enrichment range highlights the robustness of both analytical methods.



114

115 **Figure 2.** Comparison of CRDS and IRMS measurements for ^{15}N - $^{15}\text{NO}_3^-$ at% in soil extracts.
 116 Left column shows uncorrected measurements for samples originating from (a) $^{15}\text{NH}_4^+$ and (c)
 117 $^{15}\text{NO}_3^-$ labelled treatments. The right column shows blank and at% corrected measurements
 118 for (b) $^{15}\text{NH}_4^+$ labelled samples and (d) $^{15}\text{NO}_3^-$ labelled samples. Solid red lines represent the
 119 linear regression fit; dashed lines represent the perfect 1:1 instrumental agreement.

120 To validate even further the efficacy of CRDS for gross N transformation assays, we compared
 121 gross nitrification rates calculated using the Kirkham and Bartholomew (1955) model (Figure
 122 3). This comparison was performed on samples from four randomly selected isotope pool
 123 dilution assays. Again, we find a high degree of agreement between the two methods. The
 124 integrated average rate for 68-hour interval was identical between both CRDS and IRMS (1.18
 125 $\mu\text{g N g}^{-1}\text{d}^{-1}$), which is commonly reported in N cycling. The convergence of these values
 126 demonstrates that CRDS provides a robust, high-throughput alternative to traditional IRMS for
 127 measuring gross N transformation rates in biogeochemical ^{15}N field studies utilizing ^{15}N - NO_3^-
 128 measurements.



129

130 **Figure 3.** Comparison of gross nitrification rates ($\mu\text{g N g}^{-1}\text{d}^{-1}$) across four individual assays
 131 (1–4) and their overall mean. Rates were measured using both Cavity Ring-Down Spectroscopy
 132 (CDRS) and Isotope Ratio Mass Spectrometry (IRMS). Dotted bars represent calibrated rates,
 133 while patterned bars represent uncalibrated rates. Values above each bar indicate the calculated
 134 rate. Error bars on the mean bars represent the standard error (SE, $n = 4$).

135 As demonstrated above, CRDS provides precision and accuracy comparable to conventional
 136 IRMS for determining ^{15}N at% in soil NO_3^- from highly enriched samples. Furthermore, CRDS
 137 operates effectively at lower N_2O concentrations than IRMS. Consequently, future studies
 138 should evaluate the performance of this method across different soil types and NO_3^-
 139 concentrations below the $14.3 \mu\text{mol N L}^{-1}$ targeted here. Given its reduced consumable costs
 140 and lower maintenance requirements, CRDS represents an attractive alternative to IRMS,
 141 potentially facilitating the broader adoption of ^{15}N tracer techniques in N cycle research.

142 **Author contribution.** MEM and LK conceptualized the study, supervised and acquired
 143 funding. WH conducted the experiment and laboratory analysis with the help of PMRM, JG,
 144 MEM and LK. WH, MEM, PMRM, JG, LK, processed the data. WH, MEM and LK drafted
 145 the manuscript after which WH finalize it. All authors reviewed the manuscript.

146 **Code and data availability.** The code and data that support the findings of this study is openly
 147 available at <https://doi.org/10.6084/m9.figshare.33328221>

148 **Competing interests.** The authors declare that they have no conflict of interest.

149 **Acknowledgements.** All measurements were conducted in the Kuopio Isotope Facility, which
 150 is part of the Biocenter Finland. The study was financed by the Thaw-N (no. 349503 and



151 353858), MeMo2 (354501), and N₂O predict (362253) projects and the Atmosphere and
 152 Climate Competence Center (357905) flagship funded by the Research Council of Finland. The
 153 colleagues at the permafrost research section of Alfred Wegener Institute, Potsdam, are
 154 acknowledged for logistical support related to sample acquisition.

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