

Nitrous oxide dynamics across nitrogen and pH gradients in headwater streams

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Abstract

Headwater streams in agricultural landscapes can contribute substantially to nitrous oxide (N₂O) emissions, yet the environmental controls on stream N₂O dynamics remain poorly resolved, particularly in systems with low pH. We investigated 72 Danish headwater streams spanning broad gradients in pH (5.0 - 8.8), land use, and soil type to identify the main drivers of N₂O variability. Nitrate (NO₃⁻) was the strongest predictor of N₂O saturation, and its positive association with N₂O intensified under acidic conditions according to linear mixed models. In the same model -aAmmonium, dissolved organic carbon, and stream depth also showed significant but weaker positive relationships with N₂O. Spatial differences among streams explained considerably more variation than seasonal or regional patterns, underscoring the dominance of local factors. Streams with pH < 6 consistently exhibited higher N₂O saturation, and generalized additive modelling indicated a marked decline in N₂O levels beginning near pH 6. Despite generally high N₂O saturation, approximately 9 % of observations displayed undersaturation, which occurred mainly in streams with low NO₃⁻ concentrations and across all seasons. Our results indicate that acidic, weakly buffered catchments may enhance in-stream N₂O accumulation even at moderate nitrogen levels. These findings highlight the need to consider pH-related controls when assessing N₂O dynamics in freshwater networks and when designing mitigation strategies for agricultural landscapes.

1. Introduction

Nitrous oxide (N₂O) is a potent greenhouse gas and ozone-depleting substance, with agriculture as its main anthropogenic source (Ravishankara et al., 2009; Ipcc, 2019). While terrestrial emissions are well studied, emissions from aquatic systems, especially, ditches, streams and rivers, remain less constrained (Maavara et al., 2019; Seitzinger and Kroeze, 1998; Silverthorn et al., 2025). Recent evidence suggests riverine networks may contribute significantly to global N₂O budgets (Yao et al., 2020). In agricultural areas, a considerable share of applied nitrogen (N) fertilizers bypasses crop uptake and is transported into groundwater and surface waters, eventually reaching coastal waters. Along the hydrological continuum, dissolved N undergoes various biogeochemical transformations that can generate N₂O, either in soils and groundwater before entering streams, or directly within the stream channels.

Nitrous oxide is produced through a variety of biotic and abiotic transformations of N compounds in soils, sediments, and waters (Quick et al., 2019). Nitrification, the microbial oxidation of ammonia (NH₃) or ammonium (NH₄⁺) to nitrate (NO₃⁻), can release N₂O as a by-product during its initial step (Butterbach-Bahl et al., 2013). In addition, nitrifier denitrification, a process carried out by ammonia-oxidizing bacteria under low-oxygen conditions, can also contribute to N₂O emissions (Wrage et al., 2001). Heterotrophic denitrification is the sequential reduction of NO₃⁻ to gaseous dinitrogen (N₂), with N₂O formed as an intermediate step just before the final reduction to N₂ (Knowles, 1982). Although nitrification and denitrification are typically regarded as the main pathways of N₂O production, abiotic processes, as well as their interaction, may also contribute substantially (Zhu-Barker et al., 2015; Grabb et al., 2017).

Soil acidification can strongly influence denitrification, particularly the terminal step in which N_2O is reduced to N_2 (Šimek and Cooper, 2002; Firestone et al., 1980). This reduction is catalyzed by nitrous oxide reductase (*nosZ*), and under acidic conditions it is often impaired, resulting in elevated $\text{N}_2\text{O}/(\text{N}_2 + \text{N}_2\text{O})$ ratios (Firestone et al., 1980). Although this effect is well documented in soils, the influence of low pH on N_2O levels in aquatic systems is less investigated (Audet et al., 2020) or limited to near-neutral pH (Clough et al., 2011; Baulch et al., 2012). To address this gap, we investigated the interplay between land use, hydromorphology, and stream water chemistry across a diverse set of Danish headwater streams. Denmark is particularly well suited for such an investigation because of its extensive acidic soils (podzols). Although these soils are generally less favorable for farming, they remain widely cultivated and heavily fertilized. This combination of naturally low pH and intensive agricultural practices creates conditions that are especially suitable for investigating the links between stream water acidity, nutrient inputs, and processes such as N_2O production. We hypothesize that streams with low pH exhibit disproportionately large N_2O saturation, particularly in agricultural catchments with large N inputs. In addition, we explore how different environmental factors beyond N and pH contribute to spatial and temporal variability in N_2O saturation, and which sources of variance, such as seasonal and temporal, are most influential.

2. Methods

2.1 Site and catchment characteristics

Sampling was conducted in 72 distinct headwater streams located in Denmark, selected to obtain a wide range in pH, soil type, and land cover (Fig. 1+2). Many of these streams have been deepened and channelized, especially those located in agricultural landscapes. Due to regional geological variation, stream pH exhibits a pronounced east-west gradient (Thodsen et al., 2024). Eastern Denmark, including Zealand, Funen, and the easternmost part of Jutland, is underlain by calcareous glacial tills deposited during the Weichselian glaciation (Madsen, 1987). These sediments, originating from an ice advance from the Baltic Sea, are rich in chalk, which imparts a higher buffering capacity and contributes to elevated pH levels in surface waters (Fig. 1). In contrast, western Jylland is dominated by older, more leached glacial and meltwater deposits with lower carbonate content, resulting in reduced buffering capacity and generally lower stream pH. The catchments associated with each sampling location were delineated using a national digital elevation model (DEM) with a 10 m spatial resolution, obtained from the Danish Agency for Data Supply and Infrastructure. Delineation was carried out in ArcGIS Pro (version 3.5; ESRI Inc., Redlands, CA, USA) using the Fill, Flow Direction, and Watershed tools. Stream slope was quantified as the change in water-surface elevation over a 100 m reach, measured with a Leica GPS at the sampling site and a point 100 m upstream; where canopy impaired satellite reception, slope was derived from the DEM. Soil pH and soil type (JB number ~~at 30–60 cm depth~~) for each catchment were extracted for five depths (0-15, 15-30, 30–60, 60-100, and 100-200 cm depth) from Adhikari (2013). Land use within each catchment was quantified by calculating the percentage cover of major land-use categories using the polygon-based national land-use map (scale 1:25,000) from the Danish Area Information System

(Nielsen et al., 2000). Daily precipitation data were obtained from the national 10-km grid, whereas air pressure data were from the corresponding 20-km grid covering the sampling sites, both provided by the Danish Meteorological Institute.

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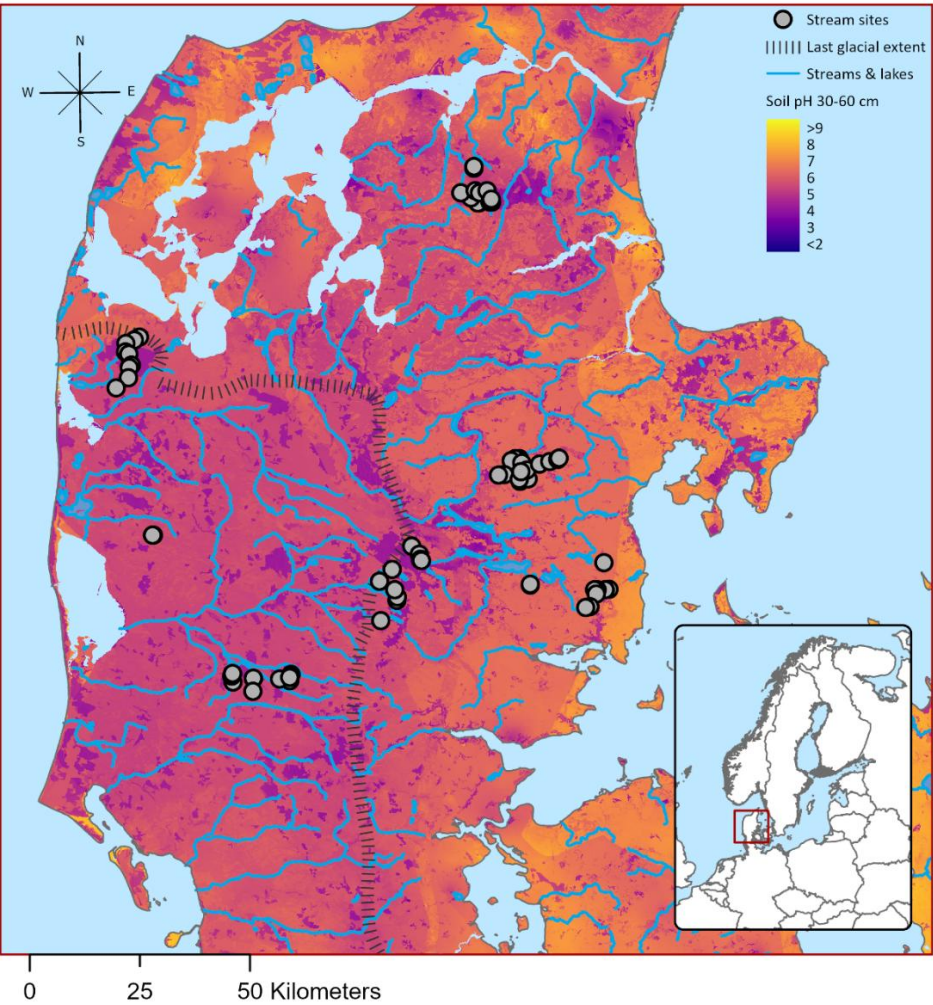


Figure 1. Map of the 72 stream study sites, the pH of soil in a depth of 30-60 cm below surface, and the extent of the ice advance during the Weichselian glaciation (land area south-west of this line was ice free).





Figure 2. Photographs of Danish headwater streams with stream identity number and sampling date.

2.2 Sampling

Water samples were taken from all streams seasonally during summer (19–26 August 2022), autumn (3–14 November 2022), winter (17–30 January 2023), and spring (19 April – 4 May 2023) (Fig. 2). Of these, ten streams (Stream 1–10) were sampled monthly from April 2022 to April 2023. During the summer sampling campaign, 12 of the 72 streams were dry, with three remaining dry until the winter sampling period. Water samples were collected 20–30 cm from the stream bank using a 250 mL polypropylene bottle and stored in a transportable cooler. At each site, stream water velocity, width, and depth were measured. Stream water velocity and depth were measured across the stream with three points (right, middle and left). Stream velocity was measured using a vane wheel anemometer, positioned $0.4 \times \text{depth}$ from the streambed (Höntzsch GmbH & Co. KG, Germany). The discharge was calculated from average stream dimensions and average velocity. In-stream ∇ vegetation cover

was visually estimated and recorded as a percentage, while the presence of iron ochre deposits in the stream was assessed on a scale from 0 to 3 (low to high abundance). To maintain consistency, the same observer conducted all assessments along a 2-m stream reach at each sampling location. Stream water temperature, dissolved oxygen, electrical conductivity, and pH were measured in situ using a multi-parameter probe (YSI Professional Plus, Xylem Analytics, USA). Samples for dissolved N₂O and carbon dioxide (CO₂) analysis were collected by drawing 50 mL of stream water into a syringe, followed by the addition of 10 mL of ambient air. The syringe was sealed using a stopcock and shaken vigorously for 1 minute by holding the piston, while keeping the syringe in the shade to minimize changes in temperature. The 10 mL headspace gas was then transferred into a 5.9 mL pre-evacuated glass vial. Samples were stored in the dark until analysis. At each sampling location In addition, 10 mL ambient air collected above the water surface was injected into a 5.9 mL vial ambient air sample was collected above the water surface at each sampling location. Data on precipitation and atmospheric pressure were obtained from the Danish meteorological institute.

2.3 Water chemistry

The samples for analysis of dissolved organic carbon (DOC), NO₃⁻, NH₄⁺, and soluble reactive phosphorus (SRP) were filtered within 24 hours in the laboratory using Whatman GF/C filters (pore size 0.45 µm), which were rinsed with 300 ml demineralized water before use. Both filtered and unfiltered samples were stored cold (3–5 °C) and dark until analysis. The samples for NO₃⁻ and sulphate (SO₄²⁻) determination were analysed according to Danish standard Ds/En Iso 10304 (2009) by ion chromatography (Dionex ICS-1500 IC-system) with an anion Micro Membrane Suppressor (AMMS III 4 mm) as basic eluent. The system was equipped with a guard column (IonPac AG22) and a separator column (IonPac AS22). The eluent was a mixture of 4.5 mM Na₂CO₃ and 1.4 mM NaHCO₃. All samples for ion chromatography were filtered through a double-layered 0.22 µm glass fibre filter (SNY2225, Frisenette ApS, Knebel, Denmark). Ammonium (NH₄⁺-N), total P (TP) and SRP were measured colorimetrically on a spectrophotometer (Shimadzu 1700, Shimadzu Corp., Kyoto, Japan) according to the Danish/European standard methods Ds/En Iso 11732 (2005) for NH₄⁺ and Ds/En Iso 6878 (2004) for TP and SRP. Total organic carbon, DOC and TN were both measured on a TOC-L analyser equipped with a TNM-L module (Shimadzu, Kyoto, Japan) at a temperature of 720 °C following Ds/En Iso 1484 (1997) and Ds/En 12260 (2003), respectively.

2.4 Dissolved ~~nitrous oxide~~gas analysis

The ~~headspace~~ concentrations of N₂O and CO₂ in headspace samples and ambient air samples were determined using a dual-inlet Agilent 7890 GC system interfaced with a CTC CombiPal autosampler (Agilent, Denmark) configured and calibrated with standard gases according to Petersen et al. (2012), with detection limits of 0.15 ppm for N₂O. The gas chromatograph was equipped with an electron capture detector (ECD) for N₂O analysis and a thermal conductivity detector (TCD) for CO₂ measurements. The headspace concentration of N₂O and CO₂ (ppm) were converted to moles using the ideal gas law, accounting for the headspace volume, laboratory temperature, and site-specific atmospheric pressure. The aqueous concentrations of N₂O and CO₂ remaining in the water phase were calculated from the headspace gas concentrations according

140 to Henry's law and using Henry's constant corrected for water temperature and atmospheric pressure at the sampling time
(Weiss and Price, 1980). ~~The total dissolved concentration originally present in the stream water was obtained as the sum of
the gas transferred to the headspace and the remaining gas dissolved in the water. The expected headspace contribution from
air was based on the ambient air sample, which was subtracted from the measured headspace values to correct for the ambient
air introduced during the equilibration step.~~ Stream water was considered undersaturated with respect to N₂O when dissolved
145 concentrations were lower than ~~the measured air concentration the theoretical equilibrium concentration calculated from
Henry's law the global annual mean atmospheric concentration of 0.336 ppm.~~ The saturation levels were classified as
undersaturated (<95 %), ~atmospheric equilibrium (95 %–105 %), and oversaturated (>105 %) following Aho et al. (2023).
~~These threshold intervals were applied to explicitly incorporate the measurement uncertainty associated with the gas
concentration determinations and the subsequent saturation calculations. Analytical precision of the GC measurements was~~
150 ~~quantified to be 2.6 % CV (SD = 0.0075 µg N L⁻¹, n = 7, range= 0.3-2.0 µg N L⁻¹).~~

2.5 Sediment analysis

~~Dry matter content was determined following Ds/En 15934 (2012) and the loss on ignition, following Ds/En 15934 (2012).~~
155 ~~For more details see appendix.~~

2.6 Data analysis and statistical analysis

All statistical analyses were performed using R version 4.4.3 (R Core Team, 2025). The N₂O saturation was log transformed
to ensure normality. Linear mixed-effects models (LMM) were used to identify environmental drivers of N₂O saturation using
the lmer() function from the lmer4 R package (Bates et al., 2015). To account for the hierarchical data structure and repeated
160 measurements within the same stream, we included a random intercept for stream ~~identity number~~. Prior to model development
the correlations between predictor variables were checked, to ensure that parameters with a correlation coefficient above 0.5
were not included (Fig. A1). In addition, we calculated the variance inflation factor (VIF) to further evaluate potential
collinearity among predictors (Table A1). All continuous predictor variables were standardized (z-transformed: centered and
scaled to unit variance) to ensure comparability of effect estimates and facilitate interpretation of interaction terms. Model
165 selection followed a backward elimination procedure, where non-significant predictors were sequentially removed. Model
assumptions were evaluated through residual diagnostics, including visual inspection of Q-Q plots and residual vs. fitted plots
(Zuur et al., 2009). Linear mixed-effects models were also applied to assess temporal and spatial differences in N₂O saturation,
while accounting for repeated measures within streams. Pairwise comparisons between regions and seasons were conducted
using estimated marginal means via the emmeans R package (Lenth, 2025). P-values were adjusted for multiple testing using
170 the Bonferroni method to control the family-wise error rate (Haynes, 2013). The relationship between log-transformed N₂O

saturation and stream pH was modeled using a generalized additive model (GAM) with a smooth pH term fitted by Restricted Maximum Likelihood (REML). Derivatives of the smooth term were used to identify pH ranges where N₂O changed significantly with pH. The first pH at which the derivative became significantly negative was defined as the onset of decline, and the pH with the most negative derivative as the point of steepest decline. Robustness of the onset threshold was evaluated using bootstrap resampling (n = 500).

3 Results

3.1 Spatiotemporal variation

During the study period (April 2022 to March 2023) the mean (± SD) annual precipitation was 741±102 mm and mean air temperature was 8.9±0.4 °C (Table 1). Mean stream water N₂O concentration between all sampling sites were 2.4±2.8 µg N L⁻¹ ranging from 0.1 to 22.2 µg N L⁻¹. In the ten streams monitored monthly, N₂O saturation showed substantial variability both temporally and ~~across-between-individual~~ streams. The stream with the lowest water pH (Stream 6, mean pH=5.8±0.2) generally exhibited higher N₂O saturation (Fig. 3, Fig. A2). For most streams N₂O saturation peaked during late winter and early spring coinciding with elevated NO₃⁻ concentrations (Fig. 3, Fig. A3). In contrast, NH₄⁺ levels remained low and relatively stable throughout the study period (Fig. A4). Consistent with patterns observed for streams with monthly monitoring, mean N₂O saturation measurements across all distinct stream systems were significantly higher during winter, compared to summer ~~N₂O saturation~~ (Table A2+A3). Summer N₂O saturation was also significantly lower compared to spring and autumn (Table A3). Stream-level differences explained the largest share of variance, with an intra-class correlation coefficient (ICC) of 0.67 according to our LLM (between-stream variance = 0.119; residual = 0.059). The remaining 33 % of the variance was attributable to within-stream variability. Regional variation accounted for a smaller proportion (ICC = 0.15; variance = 0.030; residual = 0.173), while seasonal differences were negligible (ICC = 0.016; variance = 0.0032; residual = 0.195).

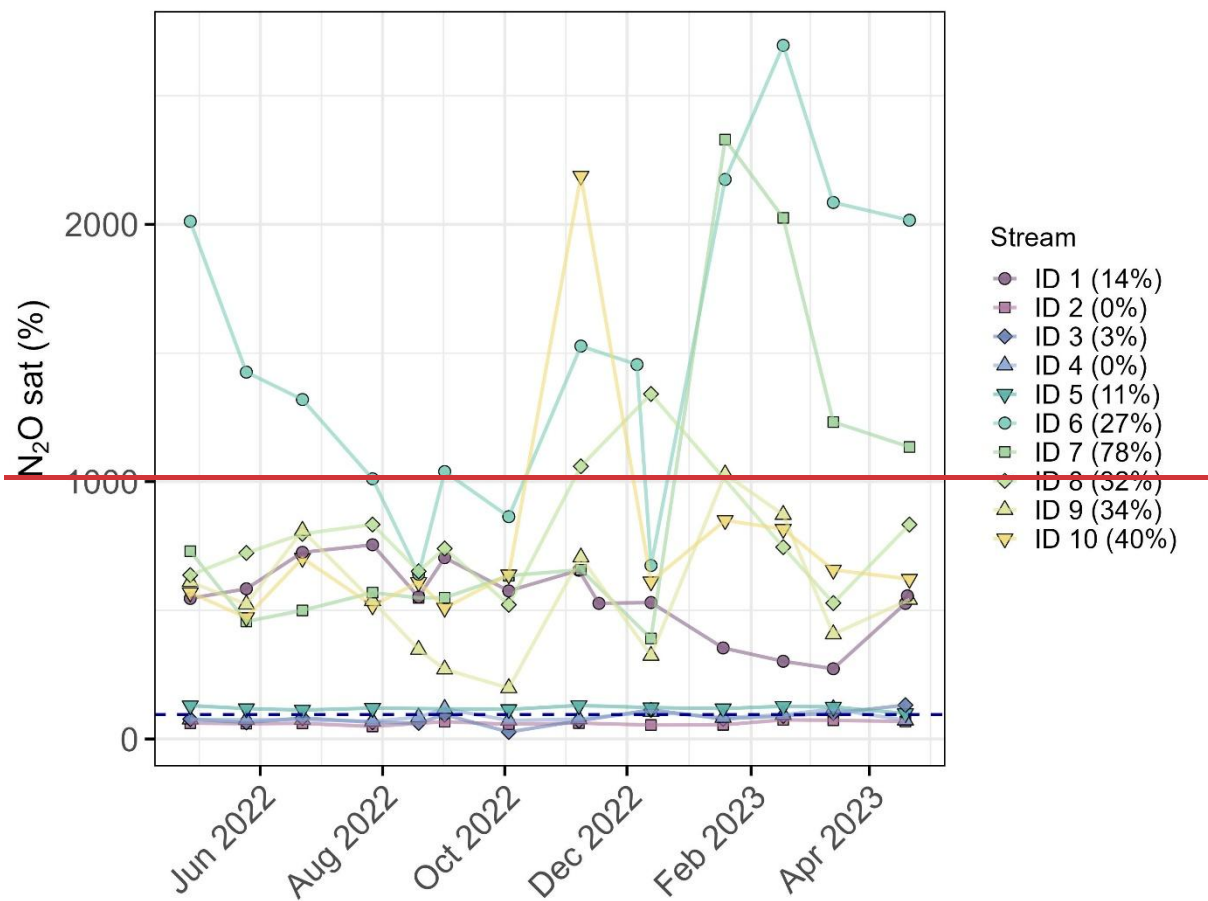
Table 1. Mean, standard deviation (SD), min and max of key parameters for Danish headwater streams (sample size=378), where N₂O=nitrous oxide, CO₂=carbondioxide, ~~NO₃⁻=nitrate~~, TN= total nitrogen, NO₃⁻=nitrate, NH₄⁺=ammonium, TOC=total organic carbon, DOC =dissolved organic carbon, TP=total phosphorus, SRP =soluble reactive phosphorus, SO₄²⁻=sulphate, and O₂= dissolved oxygen.

	Mean	SD	Min	Max
Temperature (<u>°C</u> , water)	9.3	3.6	-0.2	19.5
Flow velocity (m s ⁻¹)	0.24	0.22	0.005	1.49
N ₂ O (µg N L ⁻¹)	2.4	2.8	0.1	22.2
N ₂ O (sat, %)	608	722	27	5991
CO ₂ (µg C L ⁻¹)	2282	1981	69	12981
TN (mg L ⁻¹)	3.3	2.8	0.02	14.9
NO ₃ ⁻ (mg N L ⁻¹)	2.6	2.6	0.02	13.7

NH₄⁺ (mg N L⁻¹)	0.08	0.30	0.0001	4.42	
TOC (mg L⁻¹)	10.3	9.9	0.7	60.4	
DOC (mg L⁻¹)	9.2	9.3	1.1	54.3	
TP (mg L⁻¹)	0.27	1.54	0.01	28.63	
SRP (mg L⁻¹)	0.04	0.07	0.001	0.94	
SO₄²⁻ (mg L⁻¹)	40	161	2	2810	200
O₂ (mg L⁻¹)	9.6	2.7	0.5	18.7	
pH	7.1	0.7	5.0	8.8	
Stream depth (cm)	20.6	13.7	0.2	68.3	
Stream width (cm)	149	84	25	660	
Stream slope (‰)	7.0	9.5	0.0	68.5	
Intensive agriculture (%)	48	31	0	94	205
Forest (%)	26	27	0	88	

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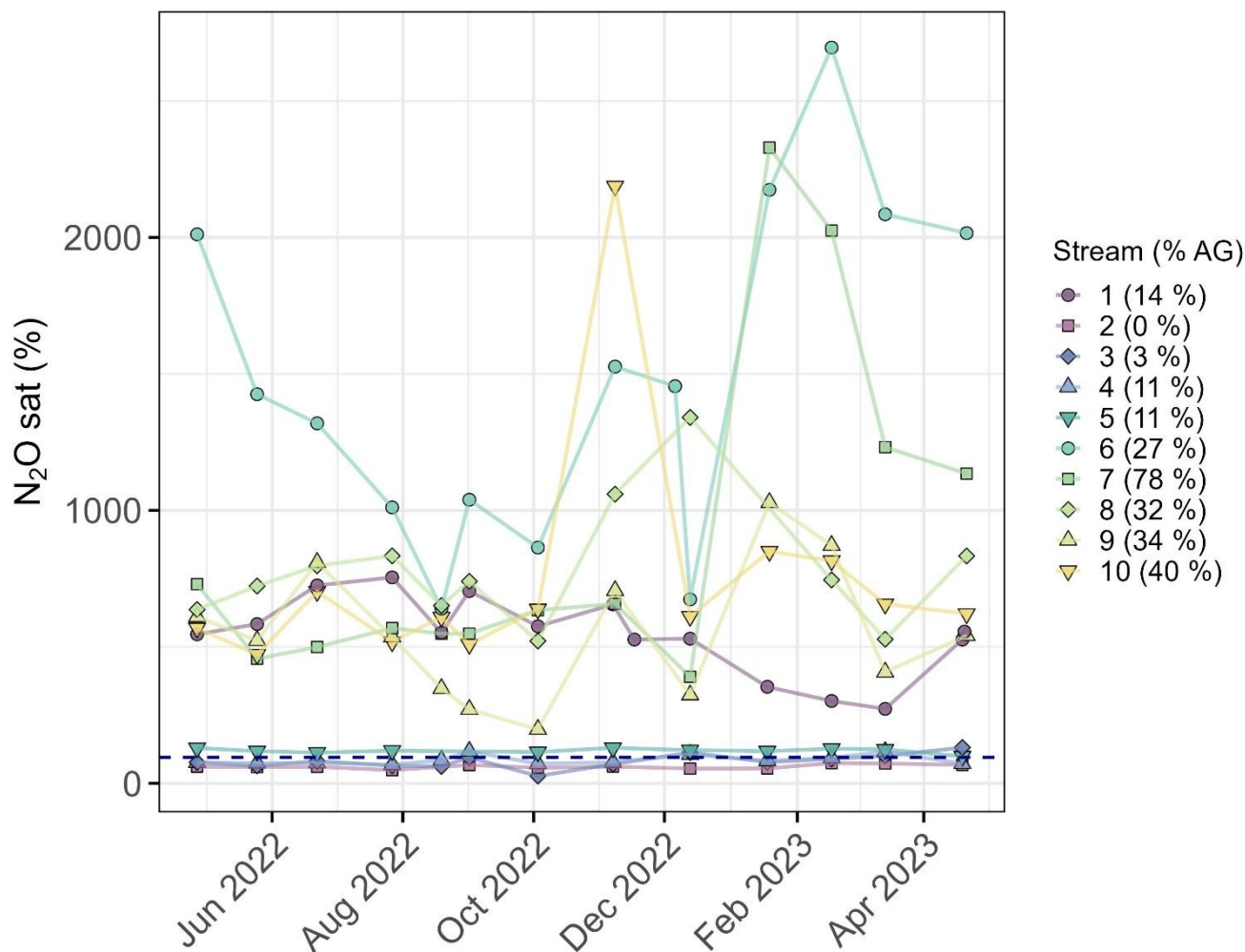
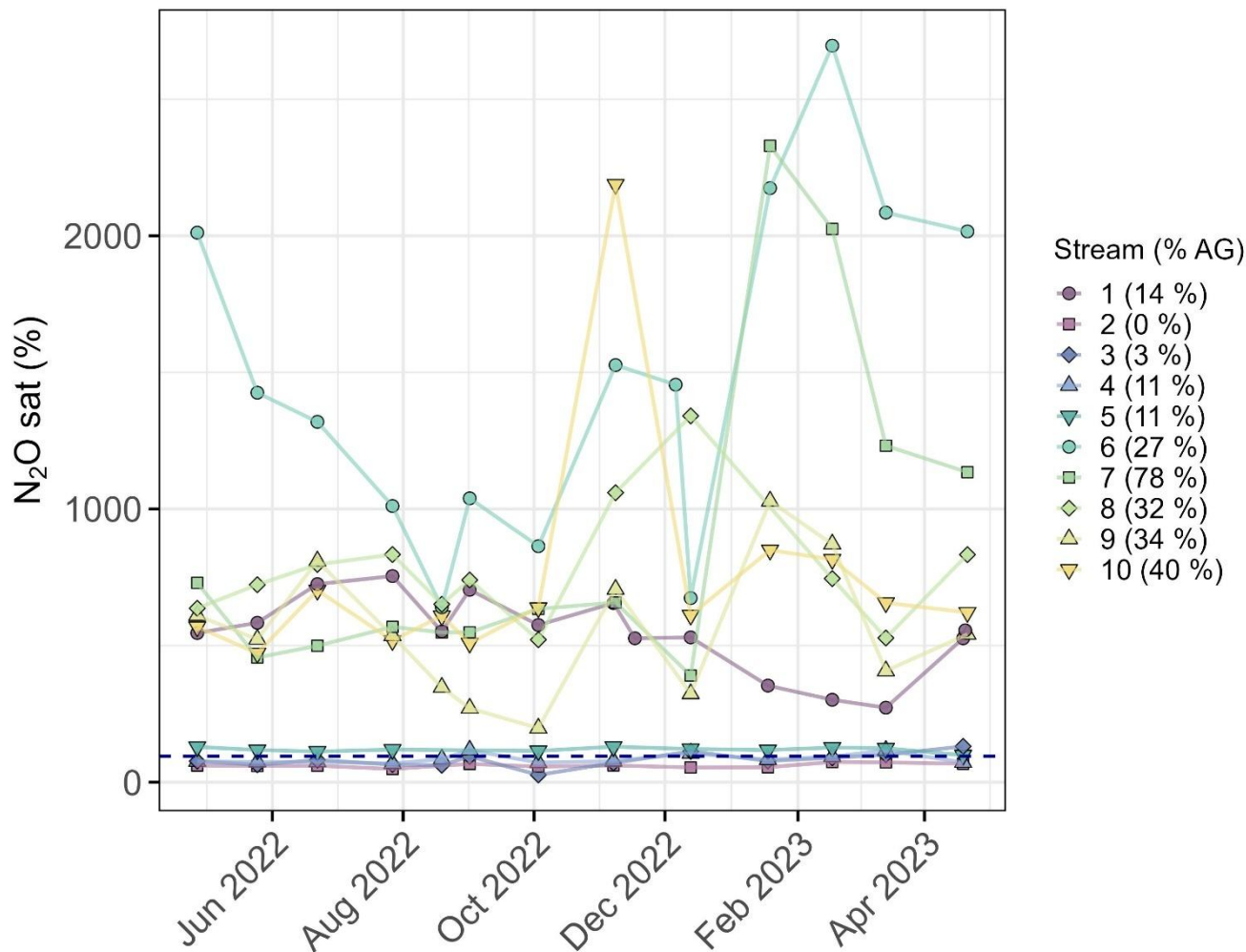


Figure 3. Nitrous oxide (N₂O) saturation (%) in stream water from ten streams (ID 1–10) monitored monthly. Stream-specific percentages of intensive agriculture within the catchment are given in the legend. The dark blue dashed line indicates the threshold below which N₂O concentrations are considered undersaturated relative to atmospheric equilibrium.

3.3 Drivers of nitrous oxide dynamics

The final model retained NO₃⁻, pH, NH₄⁺, DOC, land use, and stream depth as significant fixed effects (Table 2, Fig. A5). All retained predictors were positively correlated with N₂O concentrations, except pH. A significant negative interaction with pH (Estimate = -0.10, $p < 0.001$) was included, indicating that N₂O increased with NO₃⁻ and the effect of NO₃⁻ was amplified under acidic conditions (Table 2). The model captured notable variation between stream locations (random intercept variance = 0.041±0.020), although a considerable amount of unexplained variation remained (residual variance = 0.041±0.020) (Fig.

230 A6). Fixed effects alone explained approximately 47 % of the variation in N₂O (marginal R²), while the full model (including stream-level random effects) accounted for about 73 % (conditional R²) (Table A4). Initially SRP appeared as a significant predictor of N₂O saturation, but this relationship was driven by a single high observation and became non-significant upon its exclusion. ~~CO₂ was excluded due to a strong correlation with pH which might have introduced multicollinearity into the statistical model.~~ Variables such as water temperature, oxygen in stream water, precipitation, sulphate and flow velocity were
235 excluded from the final model due to a lack of statistical significance. CO₂ was excluded due to a strong correlation with pH which might have introduced multicollinearity into the statistical model (Fig. A1). Nevertheless, the effect of the interaction between pH and CO₂ on N₂O saturation was investigated separately to shed more light on the nature of this relationship (Fig. A5 F). A LLM analysis revealed a significant interaction between pH and CO₂ in relation to N₂O saturation (CO₂ × pH: Estimate = 0.066 ± 0.019, *p* < 0.001). While N₂O saturation decreased with increasing pH overall, the positive relationship
240 between CO₂ and N₂O became stronger at higher pH values. This was further investigated by another LLM showing that CO₂ was not significantly related to N₂O saturation at pH < 6 (β = 0.05 ± 0.05, *p* = 0.29), whereas at pH > 6 N₂O saturation increased strongly with CO₂ (β = 0.21 ± 0.02, *p* < 0.001).



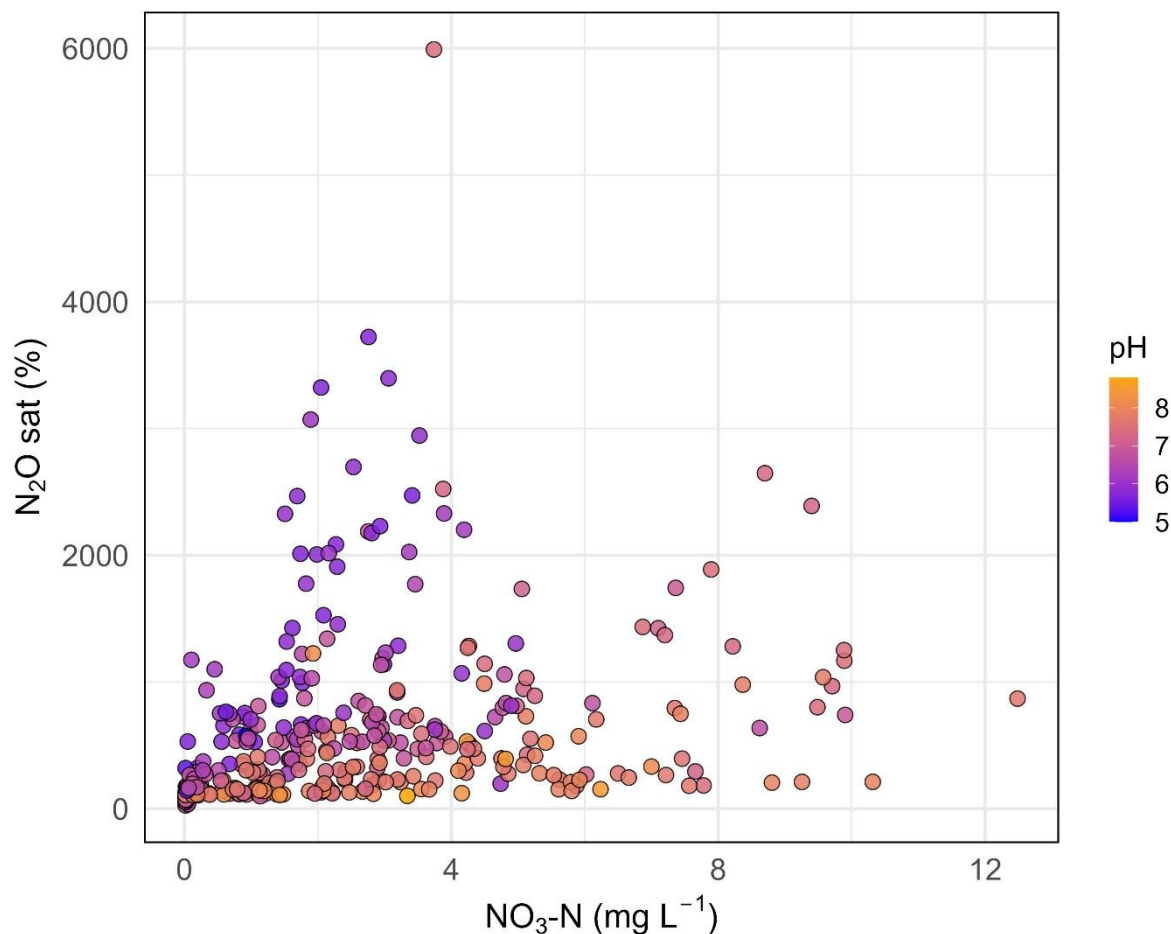
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Table 2. Results from linear mixed-effects model analysis with individual streams as random effects with pH, nitrate (NO_3^-), dissolved organic carbon (DOC), ammonium (NH_4^+), stream depth and interaction between pH and NO_3^- . SE= standard error and df= degrees of freedom. All continuous predictor variables were standardized (z-transformed: centered and scaled to unit variance).

	Estimate	SE	df DF	t value	p value
(Intercept)	2.587	0.027	69	93.9	< 0.05
NO_3^-	0.226	0.021	300	10.7	< 0.001
pH	-0.181	0.022	231	-8.03	< 0.001
NH_4^+	0.036	0.011	310	3.10	< 0.01
DOC	0.048	0.014	362	3.38	< 0.001
Stream depth	0.069	0.016	361	4.37	< 0.001

250 3.4 The effect of pH

The N₂O saturation increased with NO₃⁻ concentrations, but the magnitude of this response varied markedly with pH (Fig. 4). The highest N₂O saturations were observed at low pH, particularly below pH 6, where several observations exceeded 2,000 % saturation even at moderate NO₃⁻ levels. Derivative analysis of the GAM indicated that N₂O saturation began to decline at pH ≈ 5.95, with the steepest decrease occurring at pH ≈ 6.37, defining a transition zone over which N₂O saturation dropped most sharply (Fig. A7). Bootstrap resampling confirmed the robustness of the start-of-decline estimate, with a median of 5.96 and a 95 % confidence interval of 5.83–6.19. The N₂O saturation was significantly higher for streams with pH < 6 (mean: 1275±927 %~~μg L⁻¹~~) compared to streams with pH > 6 (mean: 485±576 %~~μg L⁻¹~~) according to LLM analysis taking stream ~~identity-number~~ into account ($\beta = 0.4 \pm 0.1$, $p < 0.001$).



260 **Figure 4. Nitrous oxide (N₂O) saturation versus nitrate (NO₃⁻-N) concentration. The color gradient indicates the pH in the stream water.**

3.5 Streams as net nitrous oxide sinks

Streams acted as net N₂O sinks in approximately 9 % of observations (n=35), as indicated by water column undersaturation relative to atmospheric equilibrium. Mean N₂O undersaturation was 69 %, with a range of ranging from 27 to 94 %. Nitrous
265 oxide undersaturation was observed primarily in streams with low NO₃⁻ concentrations, and there was no seasonal tendency. Prolonged periods of undersaturation were observed in three semi-natural streams whose catchments were characterized by a low proportion of agricultural land use. One stream (Stream 2) was undersaturated during all sampling campaigns, whereas Streams 3 and 4 were undersaturated in approximately 75 % of the campaigns. In addition, a single occurrence of undersaturation was observed in a stream receiving water from an agriculturally dominated catchment (stream 25).

270 4 Discussion

4.1 In-stream nitrous oxide dynamics and variation

Our observed N₂O concentrations (mean 2.4±2.8 µg N L⁻¹ or 608±722 %) were somewhat higher than levels from past studies of stream and rivers: median 1.0 µg N L⁻¹ in Sweden (Audet et al., 2020), mean 0.8 ug N L⁻¹ or 236 % in Michigan in United States (Beaulieu et al., 2008), 114±20 % in New Zealand (Clough et al., 2011) and 184±1,081 % in United States (Aho et al.,
275 2023). Stream water N₂O concentrations are the result of complex interactions among potential transport pathways (e.g. in-stream production, external input), within-stream physical factors (e.g. turbulence, slope) and biogeochemical conditions (e.g. pH, redox conditions). Here, we nonetheless assessed whether statistically robust relationships exist between N₂O saturation and selected in-stream environmental variables associated with N₂O production and consumption. Our finding that NO₃⁻ was a significant predictor of N₂O saturation is consistent with earlier studies (Baulch et al., 2012; Clough et al., 2011; Beaulieu et
280 al., 2009). These findings may imply that heterotrophic denitrification contributes substantially to N₂O dynamics, whether through enhanced process rates or altered product ratios under elevated NO₃⁻ levels, though further evidence is needed to confirm this. Metagenomic analysis of the sediment from two agricultural streams from our stream network (Peterse et al., in prep.) confirmed the presence of denitrification genes (*nirS*, *norB*, *nosZ*). Our results further indicate that the effect of NO₃⁻ on N₂O levels is stronger under acidic conditions. This implies that low pH in streams may promote elevated N₂O
285 concentrations even under low N availability, consistent with Audet et al. (2020), who observed similar N₂O concentrations in forested (pH<6) and agricultural streams despite much higher N concentrations in the latter.

Hydrological characteristics such as water residence time, slope, stream depth, flow velocity, and physical structure also influence N₂O concentration by affecting biogeochemical processes and gas exchange (Marzadri et al., 2021; Tonina et al., 2021; Mulholland et al., 2008; Mwanake et al., 2023). Our results show that deeper streams tend to exhibit higher N₂O
290 saturation (Fig. A7). This pattern may be partially explained by the negative association between stream depth and lower pH

levels. Physical constraints on gas exchange may also contribute, as deeper water bodies typically have lower surface area-to-volume ratios, potentially limiting the evasion of N₂O to the atmosphere, leading to higher N₂O concentrations in the water column. Several potentially important controls on N₂O concentrations were not accounted for in our study design, particularly those related to hydrological flow paths. Groundwater, often oversaturated with N₂O (Jurado et al., 2017; Von Der Heide et al., 2008), can significantly influence stream concentrations, particularly in headwater systems where it constitutes a major portion of baseflow (Bisson et al., 2025). Drain systems can deliver concentrated pulses of NO₃⁻ and N₂O directly into streams. Reay et al. (2003) reported N₂O concentrations at drain outlets up to five times higher than those measured 100 meters downstream, suggesting rapid degassing upon discharge. Our study showed that seasonal patterns were weak, with slightly lower N₂O saturation in summer. This may reflect reduced substrate availability, enhanced N assimilation, and/or increased denitrification to N₂ under warmer conditions. Overall, our findings align with previous studies reporting inconsistent seasonality (Audet et al., 2020; Beaulieu et al., 2008; Hama-Aziz et al., 2017; Aho et al., 2022; Cole and Caraco, 2001; Rosamond et al., 2012), reinforcing the importance of local over temporal drivers.

4.2 The direct effect of pH

In our study, N₂O saturation was negatively correlated with pH, and this effect was most pronounced at pH values below 6. Whether this relationship reflects a direct influence of pH or an indirect one mediated by other environmental variables remains unresolved, and it will most likely depend on the dominant N₂O production pathway. Only a small number of studies have investigated N₂O concentrations in streams with low pH (Audet et al., 2020). In contrast, the effect of pH on N₂O is well documented in soils, where numerous studies have demonstrated that low pH can impair N₂O reduction and shift end-product ratios toward N₂O rather than N₂ (Šimek and Cooper, 2002; Firestone et al., 1980; Liu et al., 2014; Qiu et al., 2024). This impairment is linked to reduced functionality of *nosZ*, the enzyme responsible for reducing N₂O to N₂ (Firestone et al., 1980). Metagenomic data from the sediment from two of our study streams (Peterse et al., ~~submitted in prep.~~) revealed that *nosZ* was only slightly less abundant in an acidic stream compared with a neutral stream. However, previous studies have demonstrated that gene presence alone is a poor indicator of N₂O reduction capacity under low pH conditions, as acidic environments primarily impair *nosZ* enzyme functionality rather than gene abundance (Liu et al., 2014; Olaya-Abril et al., 2021). The limited number of studies in streams means that the generality of the pH–N₂O relationship across diverse landscapes is still poorly understood. One reason this effect has not been widely reported previously may be that earlier studies examined streams with relatively narrow and high pH ranges such as 7.2–9.0 (Clough et al., 2011), or 7.55–8.43 (Baulch et al., 2012). In other cases, pH was measured, but results were not shown (Beaulieu et al., 2009; Hinshaw and Dahlgren, 2013) or pH was not measured (Baulch et al., 2011a; Stow et al., 2005). More research is needed to determine whether similar pH–N₂O interactions operate under different climatic, hydrological, and biogeochemical conditions.

4.3 Indirect effects of pH

A range of environmental conditions and interactions influenced by pH may shift N transformation pathways toward processes that yield higher N₂O production. A range of environmental conditions and interactions influenced by pH may shape patterns of N₂O dynamics. In particular, dissolved CO₂ was strongly correlated with pH and showed a positive association with N₂O saturation overall, however, this association strengthened progressively at higher pH but was not significant at pH < 6. Overall, these results indicate that observed CO₂–N₂O relationships vary systematically across the pH gradient. Spatially, lower pH in streams draining western Jutland is consistent with sandy, low-alkalinity catchments with limited buffering capacity. Under such conditions, shifts in inorganic carbon speciation favor dissolved CO₂ (Wetzel, 2001). In addition, low buffering capacity further increases chemical pH sensitivity to dissolved CO₂. This strong physicochemical coupling complicates interpretation of CO₂–N₂O relationships in acidic systems. At higher, circumneutral pH, the significant CO₂–N₂O association may be explained by dissolved CO₂ acting as an integrated indicator of ecosystem metabolism and/or a direct control on N₂O production. Elevated CO₂ concentrations under circumneutral conditions may reflect increased respiratory activity and organic matter mineralization, processes that are often associated with enhanced nitrogen turnover and greater substrate availability for nitrification and denitrification (Hotchkiss et al., 2015; Mulholland et al., 2008). Increased respiratory oxygen demand could also promote localized low-oxygen conditions in sediments and biofilms, which are conducive to N₂O production. However, as oxygen availability was not measured directly, these interpretations remain inferential. Experimental evidence from recent sediment incubations indicates that elevated CO₂ can coincide with increased ammonia-oxidizing bacterial abundance and enhanced N₂O production, particularly under low DOC:NO₃⁻ ratios (Mwanake et al., 2025). While these experimental findings are consistent with the field-based patterns observed here, their applicability to heterogeneous stream environments is uncertain. In conclusion, because CO₂ and pH are inherently linked through carbonate system chemistry, their individual influences on N₂O dynamics cannot be disentangled in this study. As a result, neither the direction of causality nor the relative importance of biological versus physicochemical controls can be resolved. Under acidic conditions, nitrification intermediates such as hydroxylamine (NH₂OH) and nitric oxide (NO) can undergo chemical conversion to N₂O via non-enzymatic pathways (Zhu-Barker et al., 2015). A significant, but weak, positive correlation between NH₄⁺ and N₂O indicate that nitrification, nitrifier denitrification, coupled nitrification-denitrification or nitrification coupled to abiotic processes may also contribute to N₂O production. The NH₄⁺ concentrations in streams were generally low, but it is possible that NH₄⁺ was released through the mineralization of organic matter in sediments, providing a substrate for nitrification (Arango and Tank, 2008). Supporting this, sediment data from ten streams with monthly monitoring showed that NH₄⁺ concentrations were consistently higher in sediments than in the overlying water (Peterse et al., in prep.). Microbial community profiling via 16S rRNA gene sequencing revealed nitrifiers in ten streams, with acidophilic Nitrosotaleaceae dominating acidic sites, indicating that pH might alter community structure (Peterse et al., in prep.). Another potential N₂O source is chemodenitrification (Wankel et al., 2017), where oxidized N species are reduced by Fe²⁺, which may be important in iron-rich, acidic soils such as western Jutland, but its contribution could not be confirmed because Fe²⁺ and abiotic processes were not measured directly.

4.4 Nitrous oxide undersaturation

355 Freshwater ecosystems have been recognized as potential sinks for N₂O for some time (Kroeze et al., 2007), yet this role remains largely ignored in both global and national greenhouse gas inventories (Aho et al., 2023; Webb et al., 2019). Kroeze et al. (2007) proposed that inland waters can ~~be function as net~~ N₂O ~~sinks-undersaturated~~ when N₂O consumption processes exceed in situ production and atmospheric invasion. In our study undersaturated was observed for three streams located in areas with low impact from agriculture, and thus low stream water N concentrations. This finding is consistent with previous
360 studies demonstrating that low N availability is a key factor influencing N₂O undersaturation in streams (Kroeze et al., 2007; Baulch et al., 2011b; Aho et al., 2023). Samples collected from streams across the United States, as part of the National Ecological Observatory Network, showed that 30 % of the samples (678 of 2,288) were undersaturated. In our study undersaturation was observed across all seasons, with no statistically significant seasonal variation, implying that temporal factors such as temperature or hydrological conditions may play a lesser role than N status. Besides NO₃⁻, N₂O undersaturation
365 has also been linked to low oxygen and high DOC, (Baulch et al., 2011b; Borges et al., 2019).

5. Conclusions

Our study indicates that streams draining acidic catchments with low buffering capacity can exhibit elevated N₂O saturation even when NO₃⁻ concentrations are only moderate. Consequently, landscapes with acidic soils and limited buffering capacity could contribute more to regional N₂O emissions than currently assumed in large-scale inventories, although the mechanisms
370 behind the pH sensitivity require further investigation. Globally, acidic soils cover nearly one-third of the ice-free land surface (Soil Atlas of Europe, 2005), and in many regions, they coincide with areas of intensive agriculture (Guo et al., 2010). Implementing management strategies that reduce N inputs or surpluses in such systems could help mitigate in-stream N₂O formation. However, more research is needed to determine whether similar patterns occur in other regions of the world, particularly in tropical and subtropical landscapes. Despite generally high N₂O saturation, a small fraction of the observed
375 streams showed undersaturation, occurring mainly in low-NO₃⁻ streams with no seasonal variation, demonstrating that headwater streams can function as both sources and sinks of N₂O.

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Author contribution

JA ~~and MC designed the study~~~~conceived the original idea~~, JA and MC planned the field campaigns; MC and IP performed the measurements; MC analyzed the data; MC wrote the manuscript draft; JA, AV, IP and NW reviewed and edited the manuscript. JA, NW, and AV supervised the project.

385 Data availability statement

All data used in this study are available from Zenodo at <https://zenodo.org/records/18457449>.

Competing interests

The authors declare that they have no competing interests.

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Appendix A

2. Methods

2.5 Sediment analysis

545 Dry matter content was determined following Ds/En 15934 (2012). Approximately 5 g of homogenized soil was weighed into pre-weighed ceramic crucibles and dried at 105°C for 70 hours. Samples were weighed in triplicate per site. After drying, samples were transferred to a desiccator and stored until further analysis. The desiccator remained open for 20 minutes post-drying to prevent condensation. The loss on ignition, following Ds/En 15934 (2012), was determined by annealing the dry matter at 550 °C ± 25 °C for at least 2 hours in a muffle furnace. The crucibles were then cooled briefly on a metal plate, followed by cooling in a desiccator. The desiccator lid remained open for at least one hour post transfer to avoid condensation. 550 Once samples reached room temperature, they were weighed to determine organic matter content. The carbon to nitrogen (C:N) ratio of stream sediment samples was determined using a Thermo Scientific™ FlashSmart™ Elemental Analyzer. Approximately 25–30 mg of oven-dried sediment was sealed in tin capsules and combusted at 950°C in a quartz reactor packed with quartz wool, copper oxide, and electrolytic copper. Nitrogen and carbon were subsequently separated chromatographically and detected via thermal conductivity.

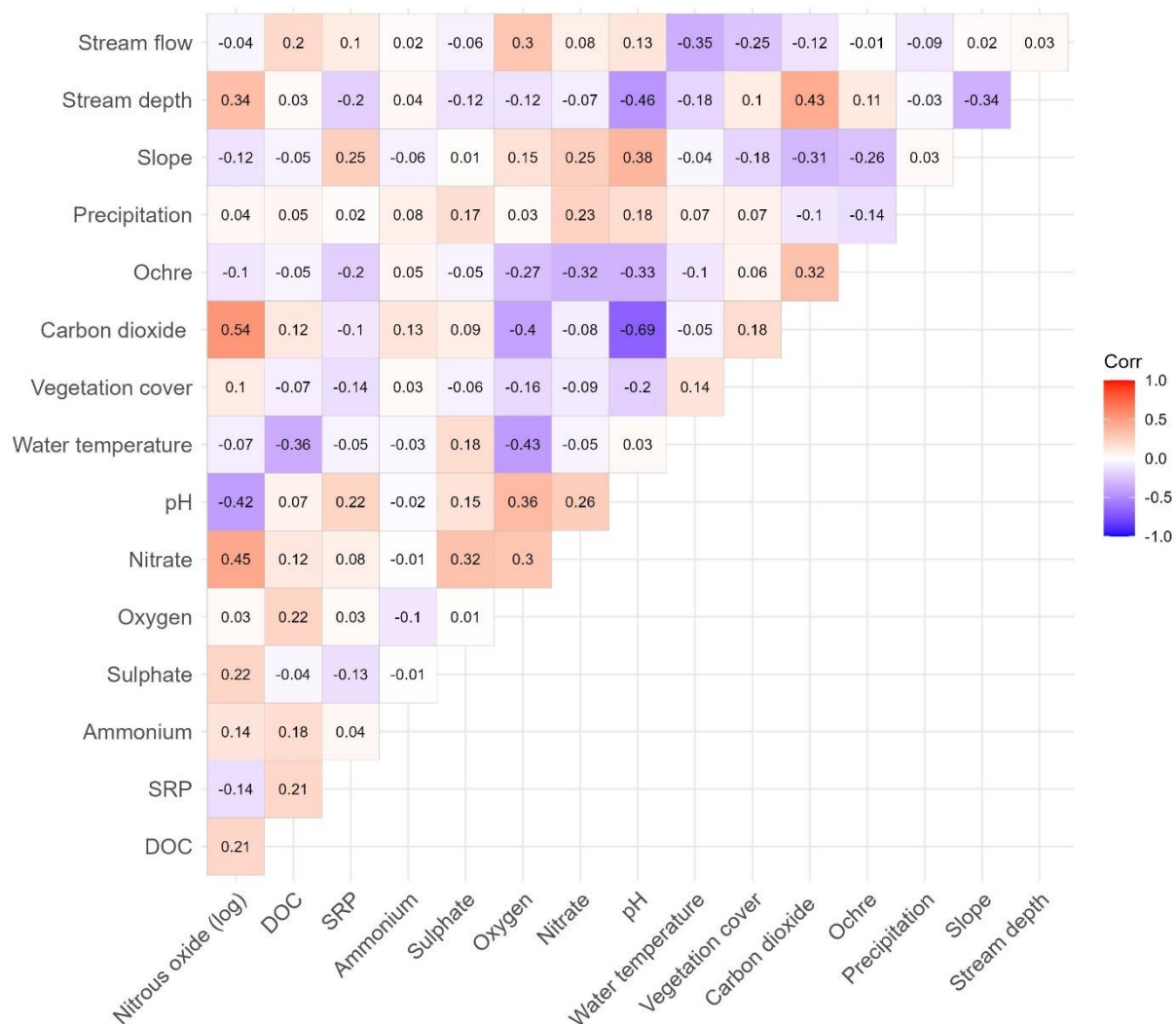
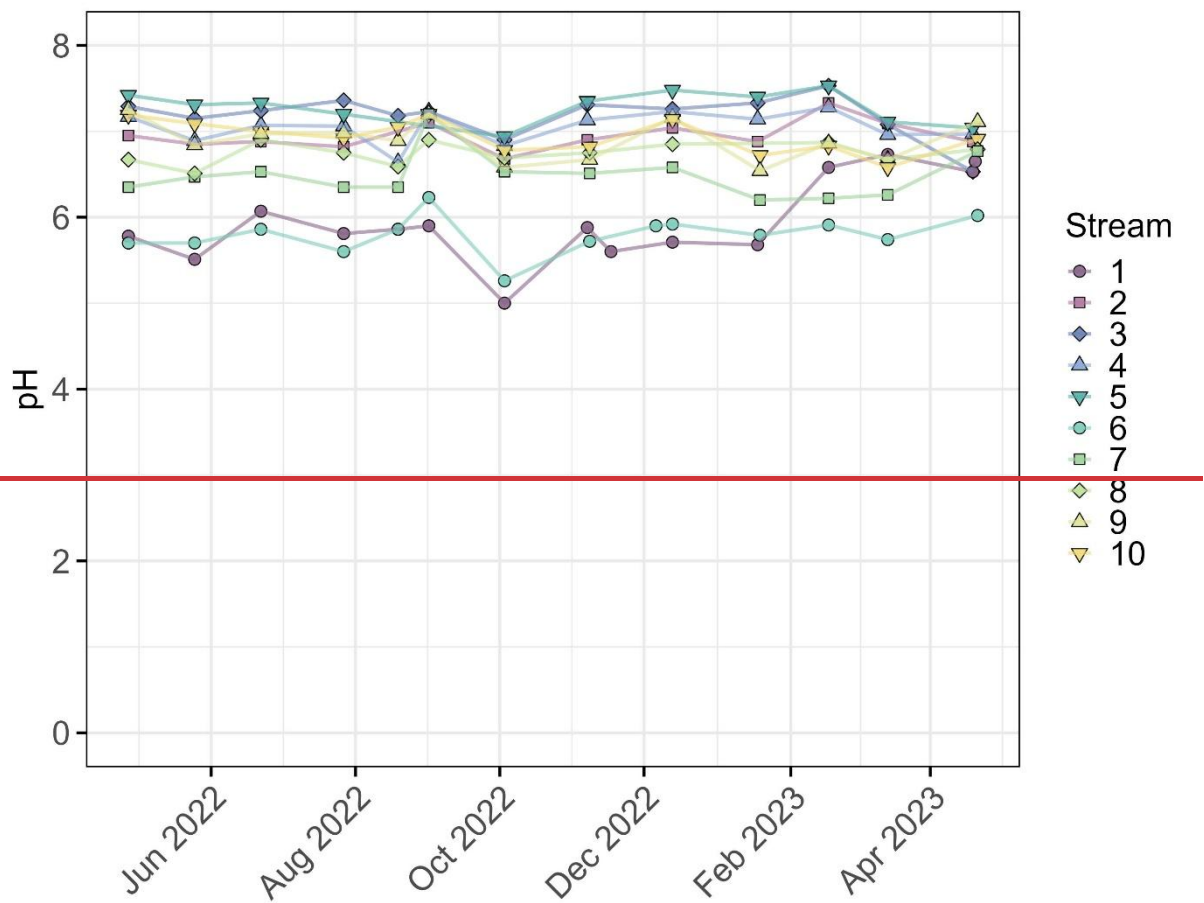


Figure A1. Correlation plots based on Pearson's correlation coefficients, where unit of nitrous oxide is percentage saturation (log transformed), stream flow is m s^{-1} , stream depth is cm, slope is permille, precipitation is mm, ochre is on scale from 0-3, vegetation cover in percentage, water temperature in Celsius degrees, carbon dioxide is $\mu\text{g L}^{-1}$, and units of nitrate, oxygen, sulphate, ammonium, SRP (soluble reactive phosphorus) and DOC (dissolved organic carbon) are mg L^{-1} .



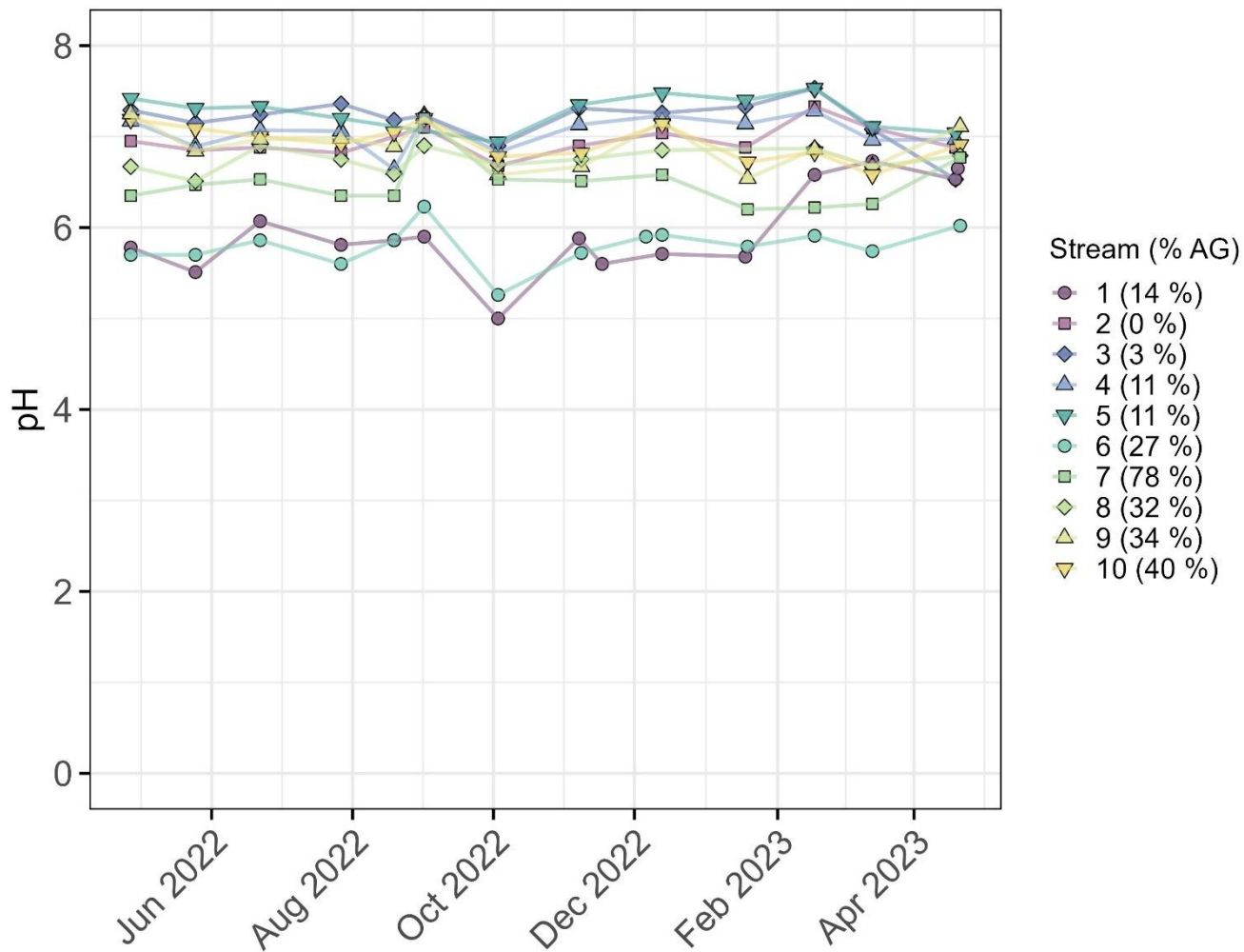
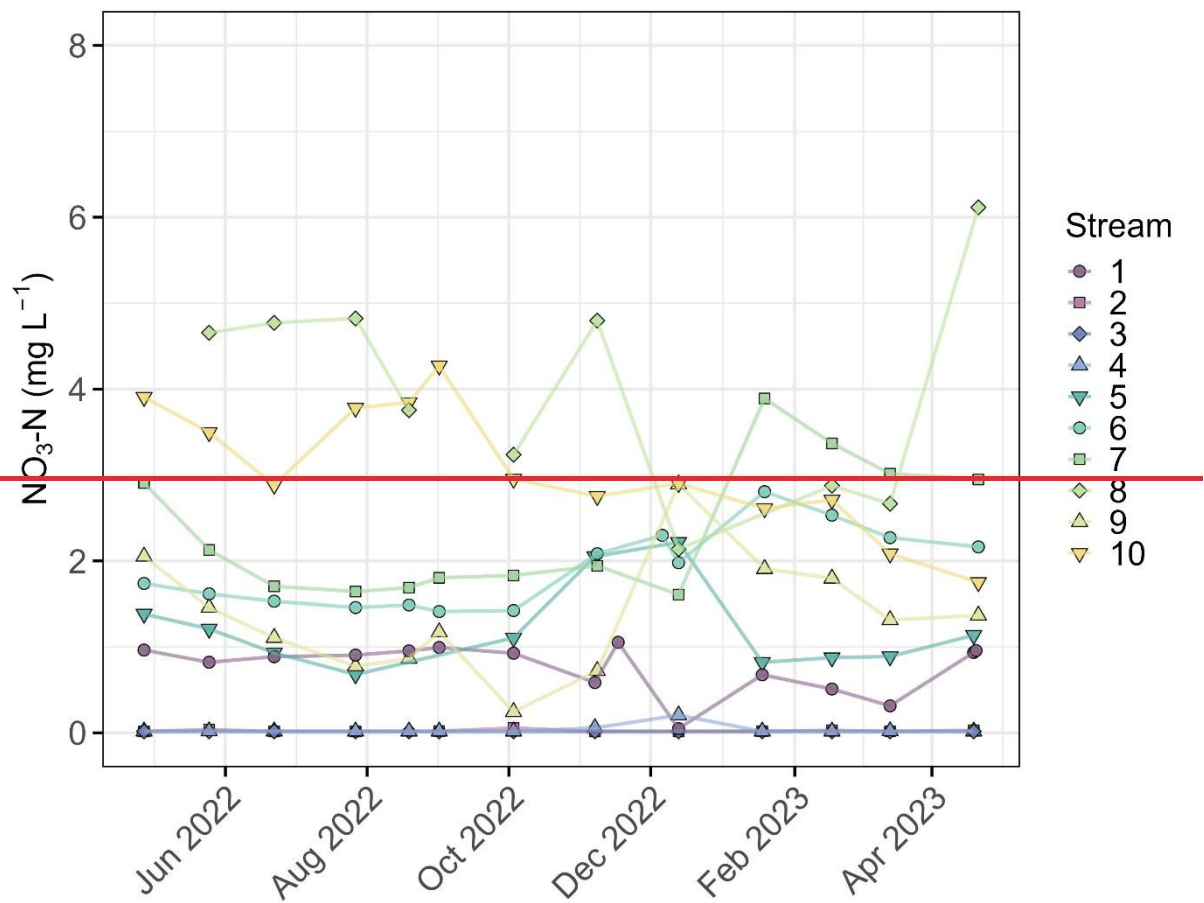


Figure A2. The pH measured in the water column from the ten streams (1-10) with monthly monitoring. Stream-specific percentages of intensive agriculture within the catchment are given in the legend.



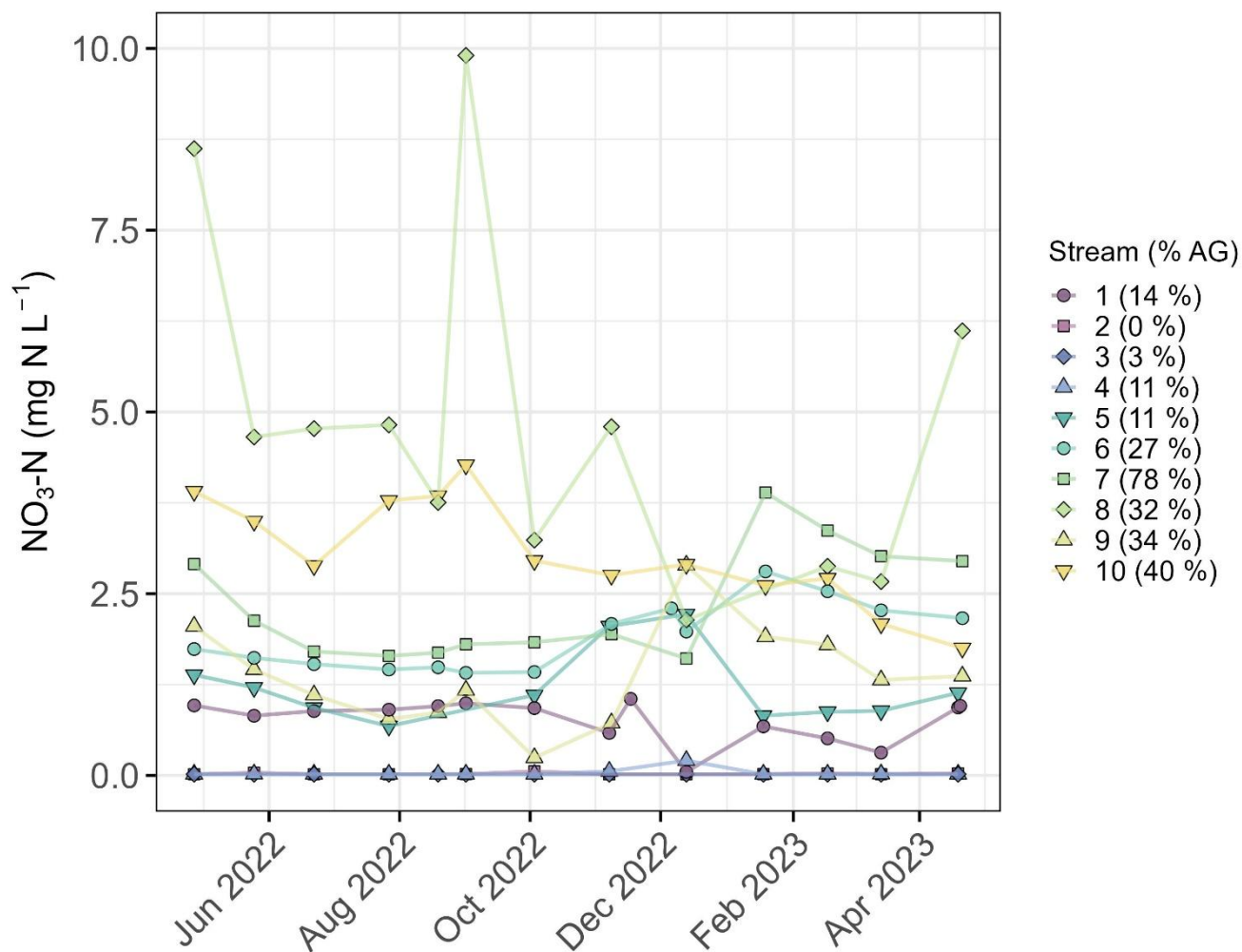
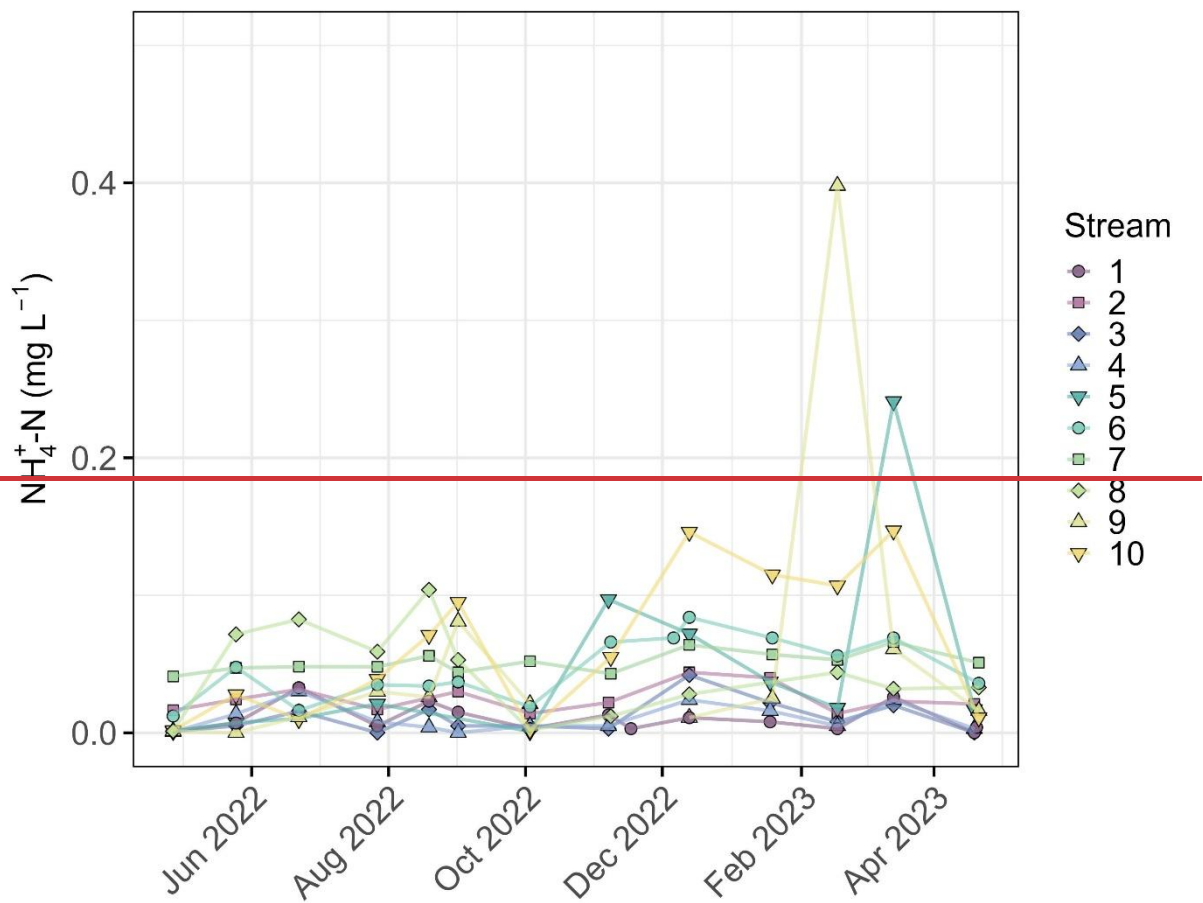


Figure A3. The nitrate (NO₃⁻) concentration of water from the ten streams (1-10) with monthly monitoring. Stream-specific percentages of intensive agriculture within the catchment are given in the legend.



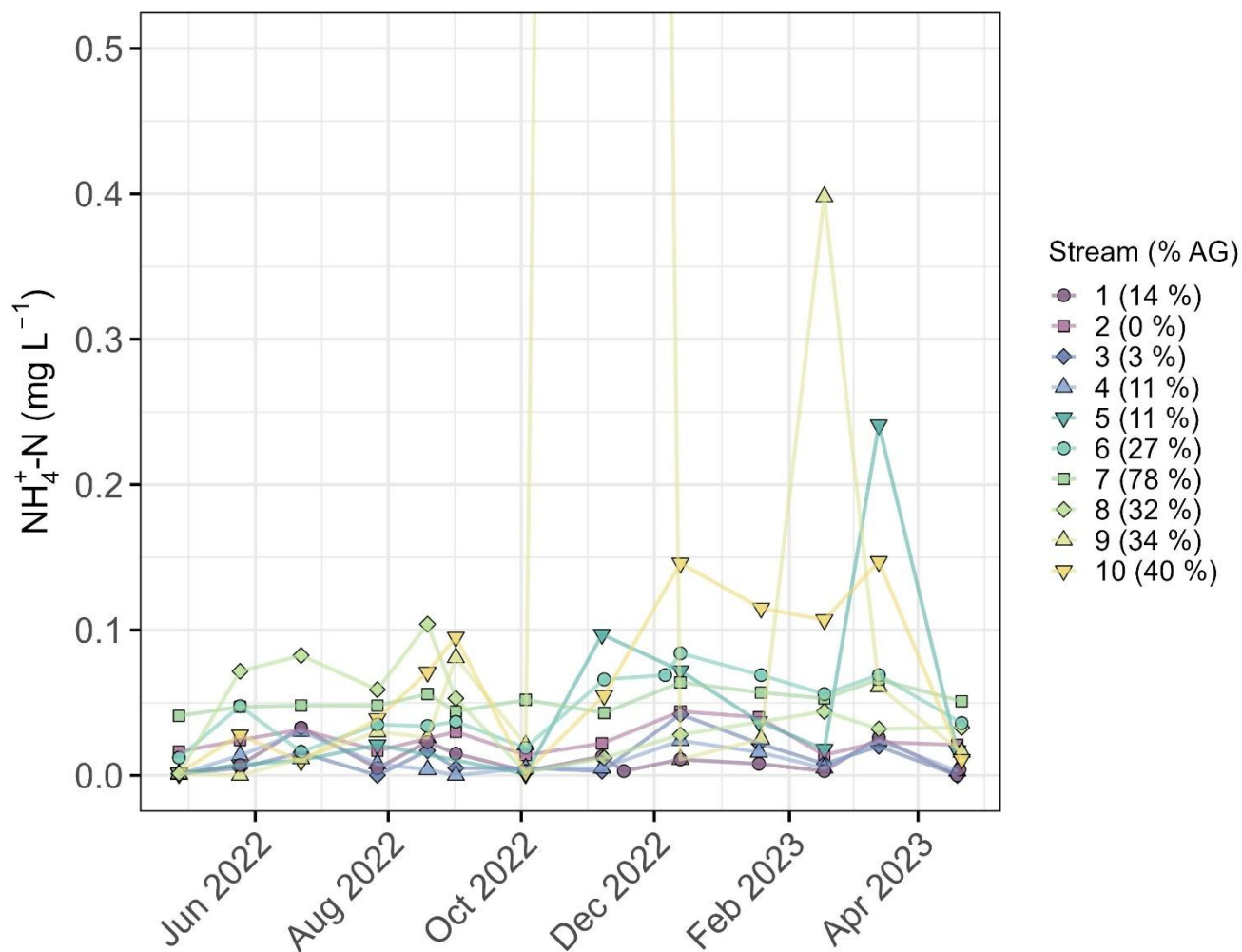


Figure A4. The ammonium (NH_4^+) concentration of water from the ten streams (1-10) with monthly monitoring. A NH_4^+ concentration of 4.4 mg L^{-1} from stream 9 from the 8 November 2022 was omitted from the graph for visual reasons. Stream-specific percentages of intensive agriculture within the catchment are given in the legend.

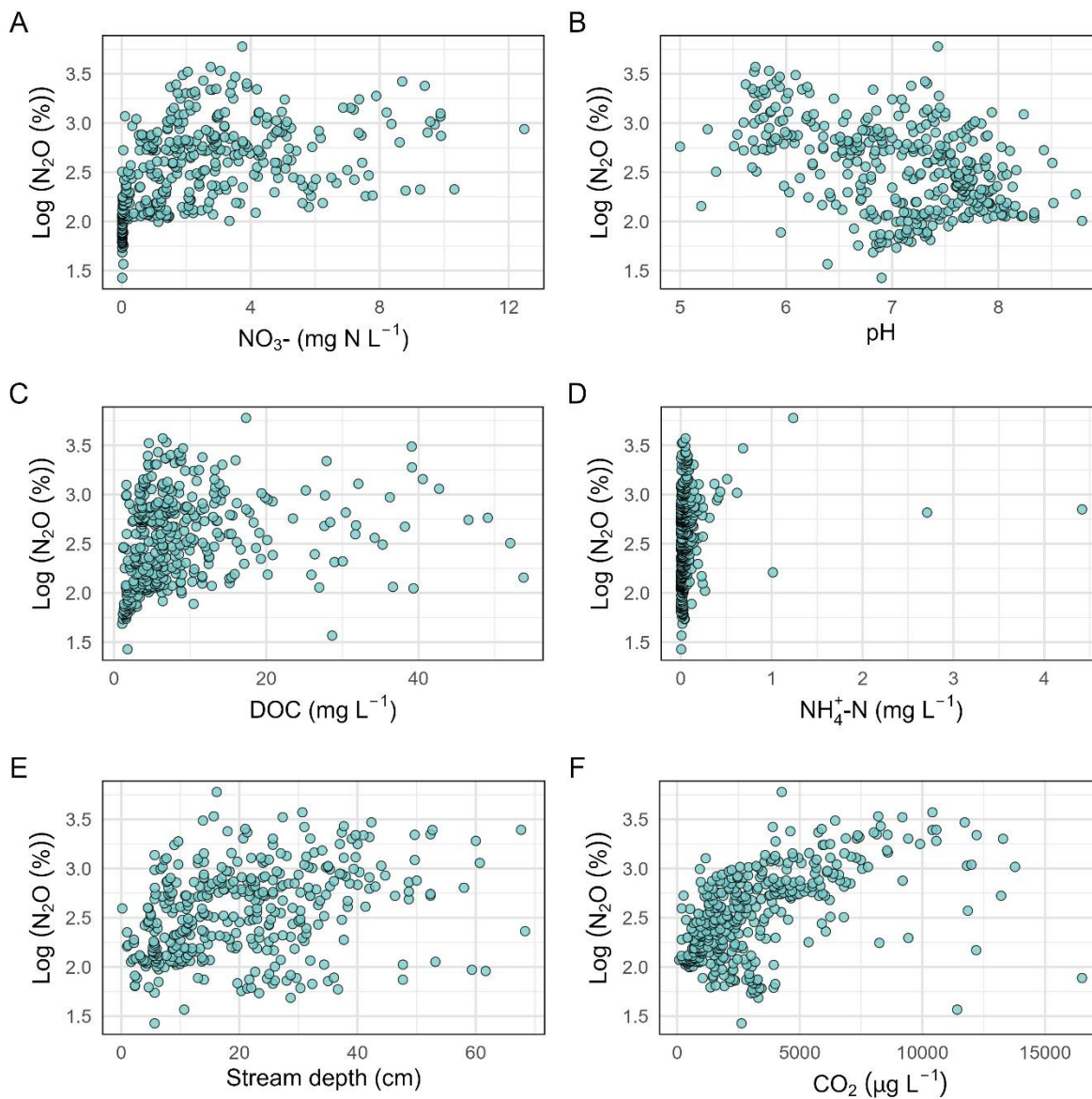


Figure A5. Stream water nitrous oxide (N₂O) saturation plotted against nitrate (NO₃⁻) (A), pH (B), dissolved organic carbon (DOC) (C), ammonium (NH₄⁺) (D), ~~and~~ stream depth (E) and stream water carbon dioxide (CO₂) (F) for Danish headwater streams.

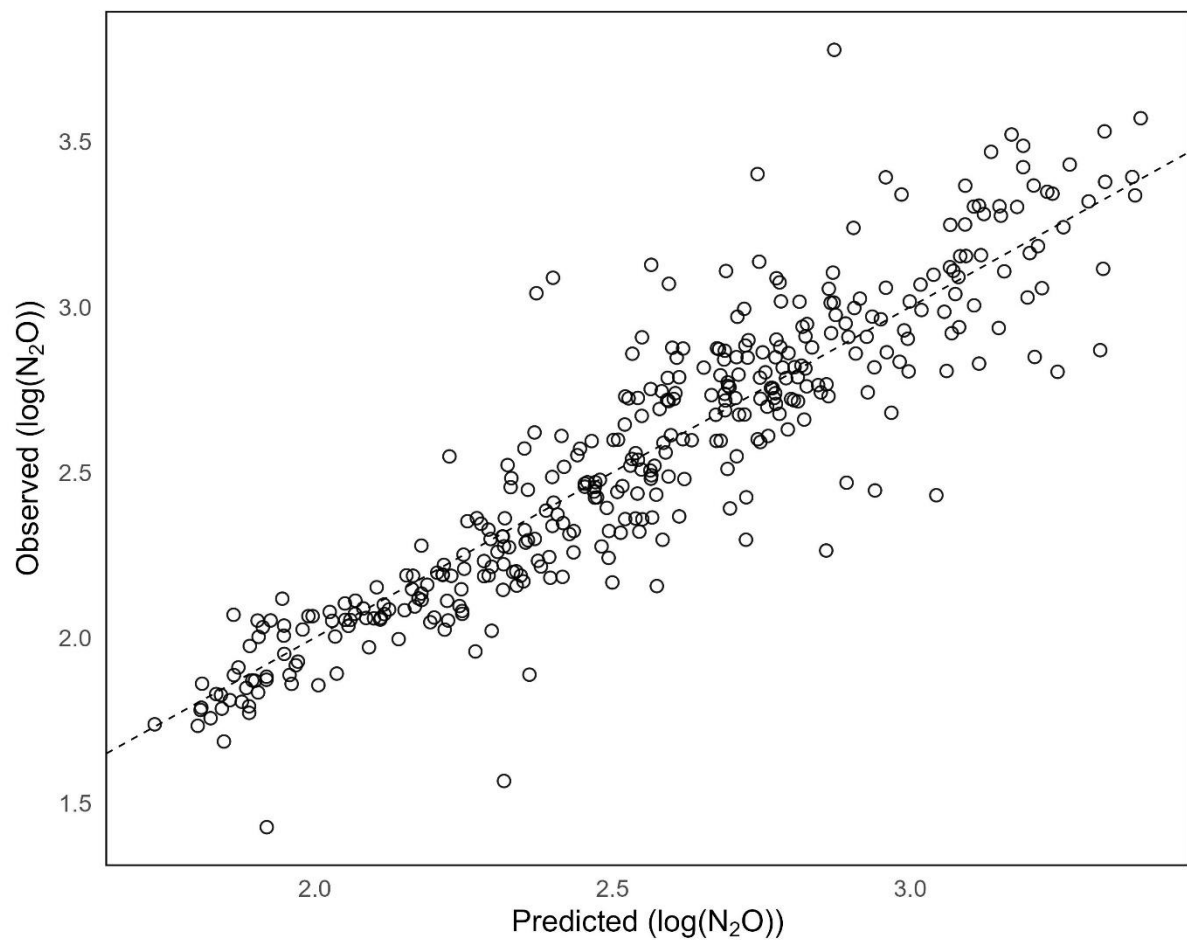


Figure A6. Predicted versus observed values of stream water nitrous oxide (N_2O) saturation.

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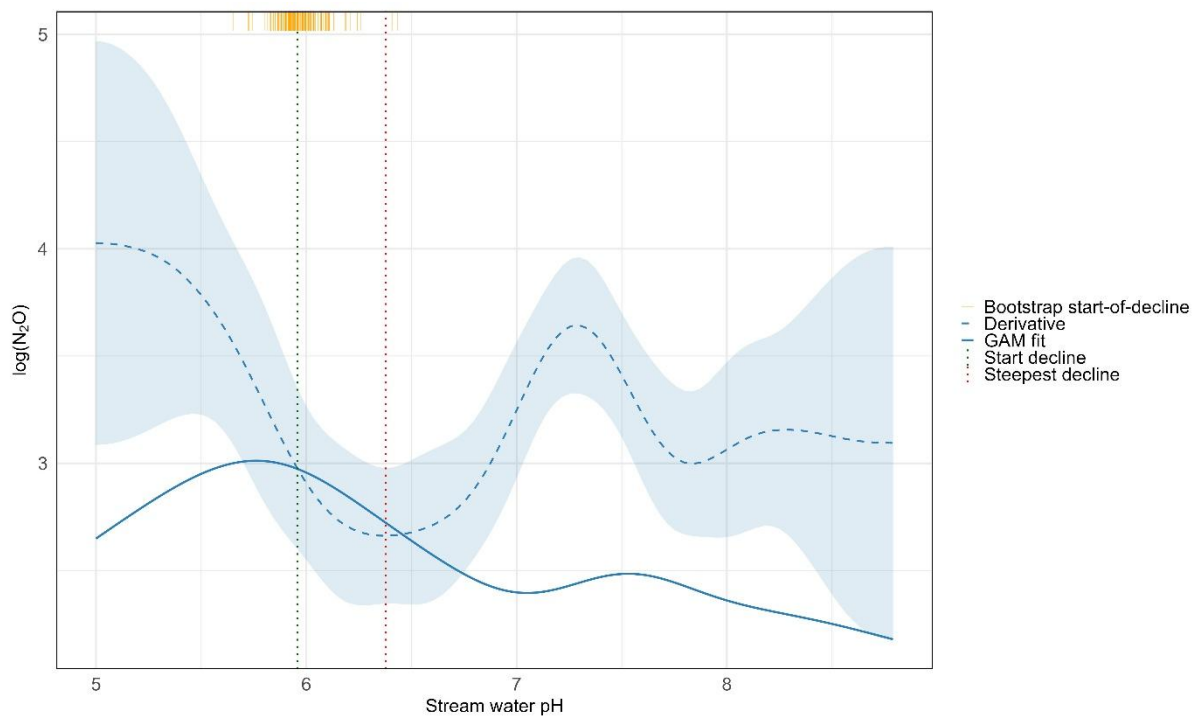


Figure A7. Generalized additive model (GAM) of log(N₂O) saturation vs. stream water pH. Solid blue line: predicted log(N₂O); dashed blue line: derivative with 95 % confidence interval (shaded). Vertical dotted lines: start of decline (green) and steepest decline (red). Orange rug: bootstrap distribution of start-of-decline pH values.

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Table A1. Variance inflation factor (VIF) analysis of linear mixed-effects model (parameters included: nitrate, pH, stream depth, land use and interaction nitrate*pH). All continuous predictor variables were standardized (z-transformed: centered and scaled to unit variance), stream water during summer 2022, autumn 2022, winter 2022 and spring 2023, where N2O =nitrous oxide, NO₃⁻ =nitrate, NH₄⁺ =ammonium, and DOC =dissolved organic carbon.

Variable	VIF
NO ₃ ⁻ _z	1.3
pH_z	1.2
DOC_z	1.1
NH ₄ ⁺ _z	1.0
Depth_z	1.2
NO ₃ ⁻ _z:pH_z	1.3

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Table A2. Characteristics of stream water during summer 2022, autumn 2022, winter 2022 and spring 2023, where N₂O =nitrous oxide, ~~TN= total nitrogen~~, NO₃⁻ =nitrate, ~~TN= total nitrogen~~, NH₄⁺ =ammonium, CO₂= carbon dioxide, TOC =total organic carbon, DOC =dissolved organic carbon, TP=total phosphorus, SRP =soluble reactive phosphorus, SO₄²⁻=sulphate, and O₂= dissolved oxygen.

	Winter	Spring	Summer	Autumn
Sample size	93	110	85	90
Precipitation (mm)	243 (48)	145 (21)	153 (28)	200 (25)
Temperature (water)	4.7 (1.6)	9.4 (2.0)	13.3 (2.2)	10.0 (1.4)
Flow velocity (m s ⁻¹)	0.34 (0.23)	0.27 (0.20)	0.13 (0.15)	0.19 (0.20)
N ₂ O sat (%)	710 (766)	570 (658)	431 (415)	632 (800)
N ₂ O (µg L ⁻¹)	3.1 (3.3)	2.1 (2.4)	1.4 (1.4)	2.5 (3.0)
TN (mg L ⁻¹)	3.5 (2.4)	3.1 (2.4)	2.6 (2.8)	3.2 (2.7)
NO ₃ ⁻ (mg L ⁻¹)	2.7 (2.1)	2.5 (2.3)	2.2 (2.7)	2.6 (2.5)
NH ₄ ⁺ (mg L ⁻¹)	0.11 (0.29)	0.04 (0.07)	0.05 (0.12)	0.11 (0.48)
CO ₂ (µg L ⁻¹)	2042 (1759)	2036 (1794)	2227 (1990)	2899 (2319)
TOC (mg L ⁻¹)	16.6 (13.7)	9.2 (7.8)	5.4 (3.1)	9.5 (7.0)
DOC (mg L ⁻¹)	14.7 (12.9)	8.4 (7.4)	4.7 (2.7)	8.2 (6.3)
TP (mg L ⁻¹)	0.20 (0.49)	0.16 (0.42)	0.44 (3.10)	0.33 (1.02)
SRP (mg L ⁻¹)	0.05 (0.04)	0.03 (0.02)	0.05 (0.03)	0.09 (0.04)
SO ₄ ²⁻ (mg L ⁻¹)	25.1 (14.7)	24.2 (14.0)	32.1 (20.2)	30.6 (19.1)
O ₂ (mg L ⁻¹)	10.9 (2.3)	10.7 (2.3)	8.1 (2.3)	8.2 (2.5)
pH	7.0 (0.7)	7.1 (0.8)	7.1 (0.7)	7.0 (0.7)
Stream depth (cm)	25 (14)	20 (13)	16 (14)	21 (12)

Table A3. Results from linear mixed-effects model analysis with pairwise test of seasonal difference in N2O saturation (%), where SE= standard error and DF=degrees of freedom.

	Estimate	SE	df	t ratio	p value
Autumn - Spring	0.04	0.03	303	1.3	0.57
Autumn - Summer	0.15	0.04	304	4.1	<0.001
Autumn - Winter	-0.04	0.03	304	-1.0	0.73
Spring - Summer	0.11	0.03	304	3.1	<0.05
Spring - Winter	-0.08	0.03	303	-2.4	0.08
Summer - Winter	-0.18	0.04	305	-5.2	<0.001

Table A4. Model performance of linear mixed-effects model (parameters included: nitrate, pH, ammonium, stream depth, land use and interaction nitrate*pH). All continuous predictor variables were standardized (z-transformed: centered and scaled to unit variance).

Model	AIC	AIC weight	AICc	AICc weight	BIC	BIC weight	R2 (cond.)	R2 (marg.)	ICC	RMSE
	1.6	<.001	2.1	<.001	36.9	<.001	0.734	0.465	0.503	0.185

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