

Middle atmosphere chemical and dynamical effects in the CCMI-2022 stratospheric aerosol injection scenario

Andrin Jörimann^{1,2}, Timofei Sukhodolov¹, Simone Tilmes³, David Plummer⁴, Shingo Watanabe^{5,6}, Hideharu Akiyoshi⁷, Gabriel Chiodo⁸, Daniele Visioni⁹, Sandro Vattioni^{1,10}, Eugene Rozanov¹, Ewa Monika Bednarz^{11,12}, Béatrice Josse¹³, Yousuke Yamashita⁷, and Thomas Peter¹⁰

¹Physikalisch-Meteorologisches Observatorium Davos / World Radiation Center, Davos, Switzerland

²Institute for Particle Physics and Astrophysics, ETH Zürich, Zürich, Switzerland

³Atmospheric Chemistry, Observations and Modeling Laboratory, National Center for Atmospheric Research, Boulder, CO, USA

⁴Climate Research Division, Environment and Climate Change Canada, Montreal, QC, Canada

⁵Japan Agency for Marine-Earth Science and Technology (JAMSTEC), Yokohama, Japan

⁶Advanced Institute for Marine Ecosystem Change (WPI-AIMEC), Tohoku University, Sendai, Japan

⁷National Institute for Environmental Studies, Tsukuba, Japan

⁸Institute of Geosciences, Spanish National Research Council, Madrid, Spain

⁹Department of Earth and Atmospheric Sciences, Cornell University, Ithaca, NY, USA

¹⁰Institute for Atmospheric and Climate Sciences, ETH Zürich, Zürich, Switzerland

¹¹Cooperative Institute for Research in Environmental Sciences (CIRES), University of Colorado Boulder, Boulder, CO, USA

¹²NOAA Chemical Sciences Laboratory (NOAA CSL), Boulder, CO, USA

¹³Météo-France, CNRS, Univ. Toulouse, CNRM, Toulouse, France

Correspondence: Andrin Jörimann (andrin.joerimann@pmodwrc.ch)

Abstract. Stratospheric aerosol injection (SAI) could slow surface warming ~~and help prevent some irreversible tipping points in the climate system. However,~~ however, potential side effects include changes in stratospheric ozone and ~~warming due to infrared absorption by the aerosol, which can alter surface ultraviolet radiation, hydrology, and weather through stratosphere-troposphere coupling changes in regional surface temperatures and precipitation resulting from tropical lower stratospheric warming.~~ Previous multi-model studies have reported ~~large model discrepancies~~ substantial discrepancies among models regarding these effects. Here we present results from ~~an experiment within~~ the Chemistry-Climate Model Initiative Phase 2 (CCMI-2022), designed to constrain inter-model uncertainties by applying a common, transient stratospheric aerosol forcing to five chemistry-climate models ~~using an SAI scenario offsetting all that offsets~~ surface warming after 2025 in a moderate greenhouse gas emission scenario. ~~All Simulations were analyzed between 2025-2099, and all~~ models show a global total column ozone decrease in the first three decades of ~~no more than at most~~ ~10 DU relative to a no-SAI case. ~~Despite sizable differences in stratospheric heating, the models yield a qualitatively similar pattern of ozone redistribution. Changes in Tropical lower stratospheric heating differs by up to 4 K between models, but despite that, the models agree very well on the region of influence of~~ key processes, ~~such as the ClO_x activation and NO_x like chlorine activation, nitrogen oxide passivation, and the strengthening of the deep branch of the Brewer-Dobson circulation, are widely robust across all models, though their relative importance and contribution to ozone changes.~~ Therefore, the sign of the ozone anomalies due to SAI is consistent in all stratospheric regions except the lower polar stratosphere, although the contribution of different processes varies consider-

ably. In three of the models, we separate ~~chemical and dynamical~~ pure chemical from dynamical (heating and nonlinear) contributions, and find ~~significant nonlinearities from feedbacks between chemistry and dynamics, highlighting where model development and sensitivity experiments are most needed to advance the understanding of the middle atmosphere in future mitigation scenarios involving~~ that towards the end of the century, dynamical effects dominate ozone anomalies, except in the lower polar stratosphere, where heterogeneous chemistry plays a major role. Our findings highlight the need for sensitivity experiments on the absorptive heating and resulting dynamical effects under SAI.

1 Introduction

Human-made climate change is progressing ~~at a faster rate~~ faster than ever before (Allan et al., 2021; Pörtner et al., 2022; Forster et al., 2024). Global warming is likely to exceed the 1.5 K warming threshold set by the Paris agreement in the next years (Liu and Raftery, 2021), increasing the likelihood that some irreversible tipping points of the climate system will be triggered (McKay et al., 2022; Möller et al., 2024). While the only permanent and sustainable solution is reaching net zero emissions, temporary amelioration of the surface warming could provide an opportunity to mitigate some of the adverse effects of climate change. Budyko (1977) and Crutzen (2006) proposed researching solar radiation modification (SRM) as a possible means of combating the rise in global surface temperatures. Stratospheric aerosol injection (SAI) is among the most discussed SRM concepts. The idea is to inject a sulfate aerosol precursor gas, which would artificially enhance the natural stratospheric aerosol layer – similar to a large explosive volcanic eruption. The resulting aerosol layer would scatter solar shortwave (SW) radiation, reflecting some of it back to space and thus ~~result in~~ producing a net cooling effect on the Earth system. Even if it were only possible to limit the rise in global mean surface temperature to 1.5 K relative to pre-industrial, rather than offsetting the entire anthropogenic effect, many of the impacts of climate change would likely be mitigated (Tollefson, 2018).

SAI carries potential side effects, some of which could be detrimental to the environment. Any perturbation to the natural stratospheric aerosol layer influences stratospheric temperatures, chemistry, and circulation (~~Robock, 2000; Marshall et al., 2022~~) (Robock, 2000; Marshall et al., 2022; World Meteorological Organization, 2022). Dynamical changes in the stratosphere can also propagate down into the troposphere, affecting regional surface climate beyond the intended global cooling effect (~~Hegerl and Solomon~~) (Hegerl and Solomon, 2009; Banerjee et al., 2021; Wunderlin et al., 2024; Tilmes et al., 2025b). When considering SAI with sulfate aerosol particles, stratospheric temperature responds according to their radiative properties. Besides reflecting SW solar radiation, liquid sulfate droplets also absorb longwave (LW) radiation, largely from the Earth, effectively trapping terrestrial energy and heating up their surroundings. This absorptive heating is expected to lead to changes in regional precipitation and increase warming in the polar regions (Wunderlin et al., 2024). Likewise, altered meridional stratospheric temperature gradients after both volcanic eruptions and potential SAI deployment have been linked to changes in atmospheric dynamics (Marshall et al., 2022; Bednarz et al., 2023a). Another serious societal concern is ozone depletion, facilitated by ~~the increased aerosol surface area. Heterogeneous catalytic ozone destruction~~ heterogeneous reactions that occur ~~primarily in the lower stratosphere on surfaces. These heterogeneous reactions~~ are limited by the available ~~aerosol~~ surface area and the presence of ~~ozone-depleting substances (ODS, World Meteorological Organization, 2022). These ozone-depleting substances (ODS),~~

50 therefore they are primarily important in the lower stratosphere (World Meteorological Organization, 2022). The ODS contain halogens, most notably chlorine and bromine, whose concentrations have been elevated through human activity for several decades. Even though the use of chlorofluorocarbons (CFC) and hydrochlorofluorocarbons (HCFC) has been strongly regulated, some emissions still remain (Western et al., 2023; An et al., 2025) and due to the long lifetimes of those compounds, concentrations will remain elevated in the stratosphere for decades (Bourguet et al., 2025). Because of ~~this~~ the large ozone
55 destruction potential that could be activated by SAI, several studies have focused on the extent of ozone loss caused by sulfur-based SAI (e.g., Pitari et al., 2014; Weisenstein et al., 2022).

Following Decision XXXI/2 of the Montreal Protocol on Substances that Deplete the Ozone Layer, which requested ‘information and research related to solar radiation management and its potential effect on the stratospheric ozone layer’ (UNEP, 2019), the Chemistry Climate Model Initiative in its second phase (CCMI-2022) proposed a new ~~SAI-scenario-with-increased~~
60 ~~specified-surface-area-densities-model experiment, that simulates the effects of SAI between 2025 and 2100, by prescribing stratospheric aerosols properties, including surface area density and radiative properties from prescribed stratospheric aerosol amounts and spatial distributions for the time period 2025 to 2100,~~ following an interactive SAI model simulation (Plummer et al., 2021; Tilmes et al., 2025a).

To simulate a climate intervention scenario with stratospheric aerosol injection in its full complexity, a chemistry-climate
65 model (CCM) with interactive aerosol ~~module-microphysics~~ is required. In recent years, the number of such CCMs has increased, enabling highly complex multi-model comparisons (e.g., Weisenstein et al., 2022; Vioni et al., 2023; Bednarz et al., 2023b). ~~However, these studies also show~~ The sophistication of the different models is still variable, however, and these studies have shown that different aerosol modules can produce different aerosol burdens, ~~SADs~~ surface area densities (SAD), and spatial distributions, leading to large uncertainties in the SAI effects (Pitari et al., 2014; Tilmes et al., 2021; Lee et al., 2026)
70 . Recently, Tilmes et al. (2025b) showed that even within the same model, substantial differences in radiative efficiency and size-dependent sedimentation can arise when using a modal versus a sectional interactive aerosol scheme. The CCMI-2022 experiment circumvents such uncertainties arising from the representation of microphysical processes, by specifying a common stratospheric aerosol forcing in all participating models (Plummer et al., 2021), while also avoiding inter-model spread arising from differing equilibrium climate sensitivities (ECS), which necessitate different SAI injection amounts to achieve equivalent
75 surface cooling. Here, we analyze the five models that contributed to this inter-comparison, describing the effects of SAI on the middle atmosphere, focusing on points of agreement and differences, and assess the remaining challenges.

2 Methods

2.1 Experimental setup

The CCMI-2022 scenario aims to identify the most important dynamic and/or chemical uncertainty factors in modeling the
80 radiative and chemical effects of a prescribed artificially enhanced stratospheric aerosol layer. All models use the same prescribed aerosol layer in order to achieve a consistent forcing representation. Even models with interactive stratospheric sulfur cycle and microphysical module for the growth and development of sulfate aerosols replace this interaction with the prescribed

aerosol. This study also does not aim to generate realistic SAI effects in the troposphere, as the sea surface temperatures are fixed as the climatological average of the years 2020-2029. We do not consider any tropospheric signals in this analysis, ~~except for temperature, where the SAI effect must be close to zero to confirm the validity of the surface climate prescription in each model.~~

The base experiment uses two model configurations: senD2-fix, a reference simulation with repeating 2025 aerosol climatology (replacing refD2 described below); and senD2-sai, an SAI simulation with a continuously increasing stratospheric aerosol burden. The difference between the two isolates the effects of SAI on the middle atmosphere. The resulting anomalies arise from chemical and dynamical changes that can also influence each other. The chemical impacts are directly driven by heterogeneous chemical reactions on the prescribed enhanced ~~surface area density (SAD)~~ SAD, while the dynamics respond ~~primarily~~ to the changes in radiative heating rates from both aerosol and chemical changes. The higher heating rates are a consequence of the increase in IR-absorbing stratospheric aerosol volume, and are in addition moderated by changes in ultraviolet (UV)-absorbing ozone. In the the updated CCMI-2022 experiment presented here, we have added a third model simulation called senD2-chem, which simulates only the chemical effects of SAI, in order to better disentangle the separate roles of chemical and dynamical influences. The senD2-chem simulation is calculated by imposing only the ~~transient~~ changing SAD for the chemistry modules in the models, while all radiative properties of the aerosols are kept at 2025 conditions.

Table 1 summarizes the specified forcings for the three simulations senD2-fix, senD2-sai, and senD2-chem. ~~Greenhouse;~~ transient greenhouse gas (GHG) forcings are adopted from the Shared Socioeconomic Pathway 2 (SSP2-4.5; "middle of the road" climate change scenario). ~~The,~~ while transient ozone-depleting substance (ODS) concentrations are taken from the 2018 report Scientific Assessment of Ozone Depletion (World Meteorological Organization, 2019). The more recent 2022 report was not yet available when some of these simulations were first started. The original description of the CCMI-2022 experiment is given in the Stratosphere-troposphere Processes And their Role in Climate (SPARC) Newsletter 57 (Plummer et al., 2021), pp. 22-30, with more additional description ~~and preliminary WACCM results of the experimental setup and single-model results~~ including comparisons to the interactive aerosol simulation are shown in Tilmes et al. (2025a). The current study includes updates to the reference scenario and the additional simulation senD2-chem, which was not originally included in CCMI-2022. Subsequently, we document the protocol for these simulations in more detail.

2.1.1 refD2 (CCMI baseline future projection)

For the purposes of this analysis, the refD2 simulation can be regarded as a legacy case. We describe refD2 briefly, because some of its output is used as a forcing for the other simulations. refD2 consists of a historical and a scenario-driven part. It starts in 1960 and uses historical data on GHGs, ozone and aerosol precursors, and open burning data up to ~~2014.~~ 2014 (Plummer et al., 2021). After that, these inputs follow the "middle of the road" climate change scenario SSP2-4.5 to the year 2100, with sea surface temperatures (SST) and sea-ice cover (SIC) ~~either~~ calculated using a fully coupled ocean model ~~if available—or;~~ except for one model, where these boundary conditions are specified from another model with a coupled ocean (detailed in Section 2.3.3). The refD2 simulations are not used for analysis here, since senD2-fix with fixed SSTs and SIC provides a better reference for isolating the impact on the stratosphere. However, the 2020-2029 climatological SSTs and SIC

Table 1. Specified forcings for the three simulations in the updated CCMI-2022 experiment. Adapted from Plummer et al. (2021).

forcings	senD2-fix	senD2-sai	senD2-chem
GHGs ¹ , ozone/aerosol precursors, open burning	SSP2-4.5 ("middle of the road" Shared Socioeconomic Pathway)		
ODSs ²	World Meteorological Organization (2019)		
SSTs ³ , sea-ice	repeating climatology (2020-2029) from refD2 <u>separate simulation (refD2;</u> see Section 2.1.1)		
QBO ⁴	nudged (SOCOL, CMAM, NIES) or internally generated (WACCM, MIROC)		
stratospheric aerosol (derived from one simulation with 4-point injection scheme for all simulations, see Section 2.2)	repeating 2025 climatology from simulation (chemical & optical)	transient simulation forcing with SAI (chemical & optical)	transient simulation forcing with SAI (chemical only) + repeating 2025 climatology (optical properties)

(1) greenhouse gases
(2) ozone-depleting substances
(3) sea surface temperatures
(4) quasi-biennial oscillation

from refD2 of each model served as input for the corresponding senD2-fix and senD2-sai simulations (see Table 1). A full description of refD2 with all its boundary conditions is given by Plummer et al. (2021).

2.1.2 senD2-fix (reference simulation)

120 ~~The For the~~ senD2-fix simulation ~~uses the climatological SSTs and SIC for the period 2020-2029 from~~, we use repeating
decadal mean SST and SIC climatologies from the respective refD2 simulations of each model and ~~the background~~-climatolog-
ical stratospheric aerosol fields ~~for 2025 from WACCM as fixed inputs from the year 2025 of the prescribed aerosol data (prior~~
to SAI to represent a background state) as fixed input. This background climatology has been derived separately and adapted
to each model's individual radiation scheme (see Section 2.2). Other boundary conditions are identical to refD2 and can be
125 found in Table 1. senD2-fix undergoes a five-year spin-up phase and begins its run on January 1, 2020 with SSP2-4.5 GHGs and
~~WMO~~ World Meteorological Organization ODSs throughout the 21st century Table 1. The senD2-fix simulation is not included
in the original CCMI-2022 experiment description. ~~It is intended to provide reference simulations that optimally isolate the~~
~~models' responses to the SAI forcing. To eliminate uncertainties from the ocean, SSTs and SIC are prescribed using a repeating~~
~~decadal (2020-2029) climatology from each model's respective refD2 simulations. For the stratospheric aerosol, a repeating~~
130 ~~year 2025 climatology is prescribed to simulate time-invariant conditions. This climatology has been derived separately and~~
~~adapted to each model's individual radiation scheme (see Section 2.2. Other boundary conditions are identical to refD2 and~~
~~can be found in~~, but was designed to provide a more ideal reference simulation.

2.1.3 senD2-sai

The senD2-sai simulation is identical to senD2-fix with the exception of the stratospheric aerosol forcing. To simulate the effects of SAI, an enhanced and evolving aerosol layer above the tropopause is introduced into the models. This aerosol layer comes from a single model, namely [WACCMCESM2\(WACCM6\)](#), and is uniformly prescribed in all models. The injection that creates this aerosol layer maintains the global mean surface temperature of the year 2025, which means that the injection rate increases over time. This means that for the first year (2025), the stratospheric aerosol forcing in senD2-fix and senD2-sai is identical, and thereafter the forcing in senD2-sai increasingly diverges from senD2-fix. ~~The injection strategy uses four points at $\pm 15^\circ$ and 30° latitude and the source gas for the aerosol particles is sulfur dioxide (see). The model forcing includes optically and chemically relevant properties of the enhanced stratospheric aerosol layer.~~ A detailed description of the forcing, including its simulation, derivation, and processing, can be found in Section 2.2.

2.1.4 senD2-chem

In addition to senD2-fix, we introduce senD2-chem as an additional simulation to extend the analysis. To ensure comparability, there is only one difference between these simulations and the senD2-fix and senD2-sai simulations, namely the way in which stratospheric aerosols are implemented. The goal of senD2-chem is to separate the effects of SAI on stratospheric chemistry from its radiative effects. This is achieved by maintaining the radiative properties of the stratospheric aerosol at the 2025 levels (of the forcing derived in WACCM, (Section 2.2), whereas the full time-dependent surface area density for use in the stratospheric chemistry module is prescribed to be identical to senD2-sai.

There is one complication, however, with the ~~senD2-chem implementation that experimental setup that potentially~~ introduces an inconsistency between ~~the some~~ models. When the stratospheric aerosol is ~~considered to be fully~~ radiatively inactive, the photolysis rates do not take scattering of UV radiation on aerosol particles into account, which can impact photochemistry in the middle atmosphere. Neglecting the aerosol effect on photolysis ~~would therefore not just separate in senD2-chem would therefore only isolate~~ the "chemical-only heterogeneous chemistry" effect ~~of the aerosol with respect to senD2-fix. However, the full chemistry effect, in senD2-chem, but more specifically, the "heterogeneous chemistry" effect. The complete chemical effect, in turn, is separated by still accounting principle, should also account~~ for the photolysis effect, while keeping the radiative aerosol properties at background (2025) levels. ~~This "full-chemistry" treatment is only possible in models that calculate photolysis rates online and thus can accommodate changes in photolysis due to the aerosol.~~ In practice, ~~not every model can easily choose between simulating senD2-chem in one way or the other~~ the implementation is challenging and hence differs across the model ensemble. Therefore, ~~one model out of the three models that performed senD2-chem simulations, two models (CMAM) does and (SOCOL do~~ not adjust photolysis rates to the increasing aerosol load, while the other ~~two do so (-, NIES) ; - does so.~~ This means that in CMAM and SOCOL the senD2-sai simulations also exclude the aerosol effect on photochemistry, potentially leading to a minor bias in the ozone tracer.

Table 2. Summary of the key modules for radiation, transport and (stratospheric) chemistry used in the prescribed version of each model.

model	long-wave radiation module	short-wave radiation module	transport scheme	chemistry module
WACCM	RRTMG (16 bands)	RRTMG (14 bands)	flux-form semi-Lagrangian (Lin and Rood, 1996)	TSMLT (Emmons et al., 2020)
SOCOL	RRTMG (16 bands)	RRTMG (14 bands)	flux-form semi-Lagrangian (Lin and Rood, 1996)	MEZON (Rozanov et al., 1999), (Egorova et al., 2003)
CMAM	Morcrette (1989) (6 bands)	Fouquart and Bonnel (1980) (4 bands)	spectral advection	Jonsson et al. (2004)
MIROC	mstrnX (23 bands) (Sekiguchi and Nakajima, 2008)	mstrnX (14 bands) (Sekiguchi and Nakajima, 2008)	flux-form semi-Lagrangian (Lin and Rood, 1996)	CHASER (Sudo et al., 2002), (Watanabe et al., 2011)
NIES	mstrnX (11 bands) (Sekiguchi and Nakajima, 2008)	mstrnX (21 bands) (Sekiguchi and Nakajima, 2008)	flux-form semi-Lagrangian (Lin and Rood, 1996)	Akiyoshi (2000) Sessler et al. (1996)

165 **2.2 Stratospheric aerosol forcing**

The aerosol layer applied in this work was generated similarly to the Stratospheric Aerosol Geoengineering Large Ensemble (GLENS) project documented by Tilmes et al. (2018). This utilized a feedback control algorithm (MacMartin et al., 2017) to adjust SO₂ emissions at four locations~~several kilometers above the tropopause~~ (15 and 30° N and S, 180° W; see Fig. 2 in Richter et al., 2018) in CESM1(WACCM) with the aim of maintaining global mean surface temperature (T0), the inter-
170 hemispheric temperature gradient (T1), and the equator-to-pole temperature gradient (T2) in a "high anthropogenic emission" scenario (RCP8.5). Different experiments were performed that injected at altitudes around 6 km and 1 km about the tropical tropopause (Tilmes et al., 2021).

~~In~~ For the CCMI-2022 experiment, using the newer version CESM2-WACCM6 (see Section 2.3.1), the same four latitudinal locations for emission are chosen, ~~but only about with altitudes around~~ 1 km above the tropopause (yellow squares in Fig.
175 1), ~~and~~. In addition, the requirement of preserving the inter-hemispheric temperature gradient (T1) is dropped~~in favor of~~ obtaining a, therefore the same injection amount is used for both Northern Hemisphere (NH) and Southern Hemisphere (SH) injections. This has been done to produce a more symmetric stratospheric aerosol distribution and aerosol optical depth in the two hemispheres, which facilitates multi-model comparisons (Tilmes et al., 2025a). Further, in contrast to GLENS using RCP-8.5, CCMI-2022 employs the feedback-controlled SAI to maintain global mean surface temperature at 2020-2030
180 climatological levels in the "middle-of-the-road" scenario SSP2-4.5.

The stratospheric aerosol properties determined by CESM2-WACCM6 in the feedback-control run are then used as input to the other models. The input variables as function of latitude, altitude, and time are: (i) the SAD for the heterogeneous chemistry, and (ii) the optical aerosol properties for the radiative calculations, i.e. extinction coefficients, single scattering

albedos, and asymmetry factors, each as function of wavelength. However, different chemistry-climate models have different
185 treatments of the stratospheric aerosol in terms of different model grids (lat-alt) and different spectral wavelength bands in the
radiative codes. This makes it necessary to prepare the data individually for each model (Table 2). Therefore, individual model
forcings were derived from the original CESM2-WACCM6 stratospheric aerosol layer using the REtrieval Method for optical
and physical Aerosol Properties in the stratosphere (REMAP) (Jörmann et al., 2025). The exact procedure for CCMI-2022
is detailed in Section 3.5 of Jörmann et al. (2025). From the CESM2-WACCM6 data we produced the stratospheric aerosol
190 forcing for each model on a zonal average latitude-altitude field ranging from 87.5° S to 87.5° N (in 5° steps) and from 0.5 to
39.5 km (in 0.5 km steps). The forcing provides monthly data from January 2020 to December 2100 (though we use the period
2025-2099 in this analysis). The data are archived in the ETH research collection (Jörmann, 2023).

Figure 1 shows the enhanced senD2-sai zonal mean ~~winter (DJF) and summer (JJA)~~ stratospheric aerosol surface area
densities averaged over two decades (2060-2079) for the boreal winter months December, January, February (DJF) and the
195 boreal summer months June, July, August (JJA). This looks nearly identical across the five models, with small differences near
the tropopause, whose height differs between models, as shown by the lines in Fig. 1. As only stratospheric SAD is prescribed,
a stratosphere mask selects only values above the tropopause of each individual model, effectively truncating any non-zero grid
cells below. In particular at higher latitudes the models cut out different parts of the forcing. Therefore, small inconsistencies in
the modeled stratospheric aerosol (in the prescribed SAD and optical quantities) exist in the lowest levels of the stratosphere.
200 However, the effect of the tropopause truncation on the latitude-weighted quantities is minor, as shown in Fig. 2a for the
stratospheric SAD column.

2.3 Chemistry-Climate Models

We use chemistry-climate model data to assess the impacts of stratospheric aerosol injection on the middle atmosphere. The
interplay of dynamical and chemical processes is vital for a good representation of nonlinear effects, notably affecting ozone
205 production and destruction. This is congruent with the mission statement of the Chemistry-Climate Model Initiative (CCMI,
2022). To date, five models performed the CCMI-2022 experiment: CESM2-WACCM6, SOCOLv4, CMAM, MIROC-ES2H,
and CCSRNIES-MIROC3.2. Versions of these models have all been used to produce realistic stratospheric ozone distributions
and are well established in studies on ozone recovery (e.g. Dhomse et al., 2018; Friedel et al., 2023; Benito-Barca et al., 2025).
Each model's radiation modules, transport schemes, as well as chemistry modules are compiled in Table 2. The six heterogeneous

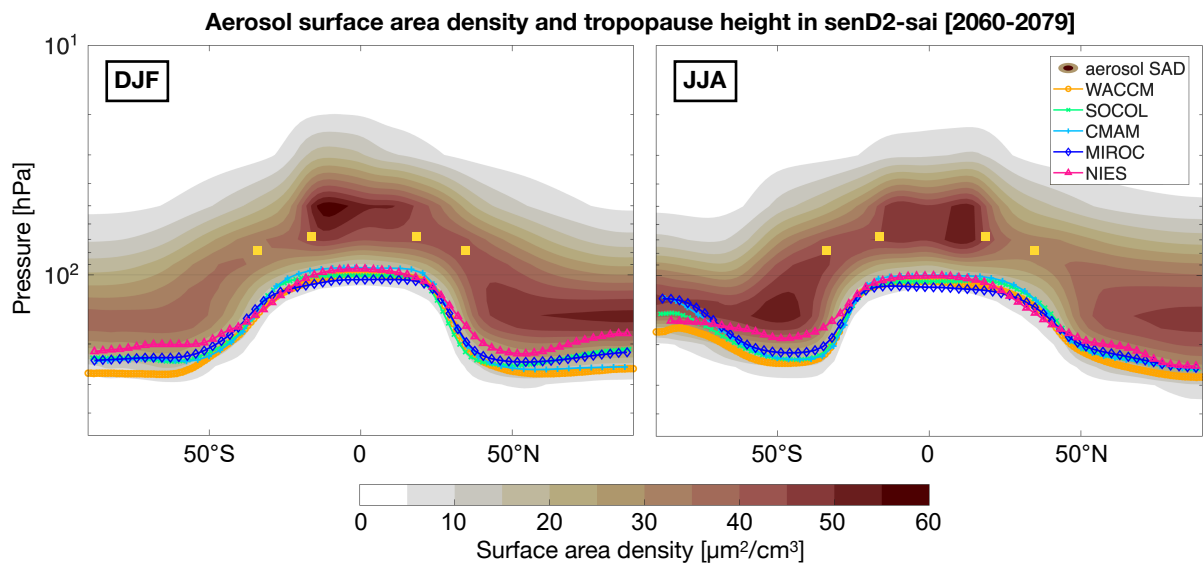
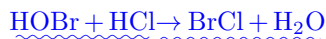
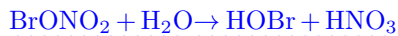
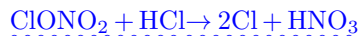
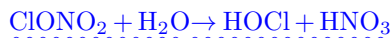


Figure 1. Artificially enhanced stratospheric aerosol zonal mean surface area density (SAD) output of a senD2-sai simulation ([near-identical to prescribed SAD input](#)). Contours: SAD from WACCM averaged over 20 years (2060-2079) for boreal winter (DJF, left) and summer months (JJA, right). Lines: Average tropopause pressure from all models (color-coded) over the same time period. Yellow squares: approximate positions of SO₂ emissions used to generate the depicted SAD.

210 [chemical reactions on sulfate aerosol surfaces present in all five models comprise:](#)



Two models (MIROC and NIES; see descriptions below) include more reactions on sulfate aerosol. For a complete list of reactions of all models, including heterogeneous reactions on nitric acid trihydrate (NAT) and ice aerosols, see Table D1 in the [Appendix](#).

In the analysis, as well as in tables and figures we shorten the model names to "WACCM", "SOCOL", "CMAM", "MIROC",

and "NIES". We use ensembles with three members (realizations) for senD2-sai and one realization for senD2-fix and for senD2-chem per model.

2.3.1 CESM2-WACCM6 ("WACCM")

225 The Whole Atmosphere Community Climate Model Version 6 (WACCM) is the atmospheric component of the Community Earth System Model Verion 2(CESM2) (Danabasoglu et al., 2020) and has a horizontal resolution of $0.95^{\circ} \times 1.25^{\circ}$ and 70 vertical levels with a top at around 140 km. WACCM includes comprehensive chemistry in the troposphere, stratosphere, mesosphere, and lower thermosphere (TSMLT) (Gettelman et al., 2019; Emmons et al., 2020), including 231 solution species, 583 chemical reactions broken down into 150 photolysis reactions, 403 gas-phase reactions, 13 tropospheric, and 17 stratospheric
230 heterogeneous reactions. The photolytic calculations are based on both inline chemical modules and a lookup table approach (Madronich and Flocke, 1998). WACCM includes a prognostic representation of tropospheric and stratospheric aerosols using Modal Aerosol Microphysics version 4 (MAM4) (Liu et al., 2016)~~and sulfur emissions from volcanoes and other sources (Mills et al., 2017)~~. The QBO is internally generated. The model is coupled to the Community Land Model version 5 (CLM5) (Lawrence et al., 2019). The ~~configurations used here run~~ configuration used here runs with a prescribed stratospheric aerosol
235 distribution and prescribed SSTs and sea ice. WACCM did not perform any senD2-chem simulations and ran the senD2-fix simulation until the end of 2083 (instead of 2099) due to resource limitations.

2.3.2 SOCOLv4 ("SOCOL")

This is the ECHAM6-based version of the CCM Solar Climate Ozone Links (SOCOL) of the Swiss Federal Institute of Technology Zürich (ETH Zurich) and Physical Meteorological Observatory Davos (PMOD, Sukhodolov et al., 2021). The
240 dynamical core of SOCOLv4 is ECHAM6 (Stevens et al., 2013), which is interactively coupled to the chemistry module MEZON (Rozanov et al., 1999; Egorova et al., 2003), the Ocean Model of the Max Planck Institute for Meteorology Hamburg (MPIOM, Jungclaus et al., 2013), and the prognostic stratospheric aerosol microphysical model (AER Weisenstein et al., 2007; Sheng et al., 2015). It has a resolution of approximately $1.9^{\circ} \times 1.9^{\circ}$ (T63), 47 vertical levels from the surface to 0.01 hPa, a dynamical time step of 7.5 min and 99 chemical species undergoing 304 reactions, including the major catalytic cycles of
245 ozone destruction, as well as 16 heterogeneous reactions on surfaces of polar stratospheric clouds and aqueous sulfuric acid aerosols. Photolysis rates are calculated from a pre-computed look-up table and do not account for changes in aerosols. The QBO is nudged to the observed equatorial wind profiles. Different ensemble members were produced by changing the CO₂ concentration by ± 0.5 %. For the purpose of the present work the ocean module MPIOM and the aerosol module AER are deactivated. Stratospheric chemistry in the prescribed model version has been tuned to decrease the bias in the total column
250 ozone values with respect to the model version with an interactive aerosol scheme. This was done by halving the heterogeneous reaction rate coefficients on sulfuric acid aerosol surfaces. Potential reasons for the oversensitivity of halogen activation on prescribed aerosol surfaces might include the usage of monthly and zonal mean data, which can result in ~~too large~~ reaction rates under in-situ low ~~temperature~~ temperatures to be too large.

2.3.3 CMAM

255 The Canadian Middle Atmosphere Model (CMAM) has been jointly developed by the University of Toronto, York University, and Environment and Climate Change Canada (ECCC). It is based on a vertically extended version of the 3rd generation Canadian Centre for Climate Modelling and Analysis (CCCma) Atmospheric General Circulation Model as described in Scinocca et al. (2008). The simulations used here were run at T47 spectral resolution ($3.8^\circ \times 3.8^\circ$ on the linear transform grid used for the calculation of physics) on 80 vertical levels up to the model lid at 0.0008 hPa. The atmospheric chemistry mechanism
260 includes the HO_x , NO_x , ClO_x and BrO_x catalytic cycles that are important for controlling ozone in the stratosphere, ~~along with CH_4 , N_2O , six long-lived halocarbons and two very short-lived bromine-containing halocarbons.~~ Photolysis rates are calculated from a pre-computed look-up table and do not account for changes in aerosols in senD2-chem. The representation of polar stratospheric clouds includes supercooled ternary solution and ice polar stratospheric clouds (PSCs), which are calculated diagnostically based on local thermodynamic conditions. SSTs and sea-ice for the ~~refD2~~, senD2-fix~~and~~, senD2-sai, and
265 senD2-chem simulations were taken from CCCma simulations submitted to CMIP6 using the 5th generation Canadian Earth System Model (CanESM5) (Swart et al., 2019).

2.3.4 MIROC-ES2H ("MIROC")

The Model for Interdisciplinary Research on Climate - Earth System Model version 2H (MIROC-ES2H) was mainly developed by the Japan Agency for Marine-Earth Science and Technology (JAMSTEC) (Kawamiya et al., 2020). The present version
270 of MIROC-ES2H is built on MIROC6 (Tatebe et al., 2019) and its atmospheric model has a spectral horizontal resolution of T85 (1.4° latitude $\times$ 1.4° longitude), which is loosely coupled with a lower horizontal resolution T42 (2.8° latitude $\times$ 2.8° longitude) chemistry climate model, Chemical Atmospheric General Circulation Model for Study of Atmospheric Environment and Radiative forcing (CHASER) (Sudo et al., 2002). These atmospheric models share 81 vertical levels from the surface to ~ 0.004 hPa, with a vertical resolution of 0.7 km in the lower stratosphere, ~ 1.2 km in the upper stratosphere, and ~ 3 km in the
275 lower mesosphere. The model spontaneously generates ~~the a~~ QBO in the equatorial lower stratosphere. The configurations of atmospheric chemistry and aerosol modules are outlined in Watanabe et al. (2011). In the present simulations, the microphysics model for sulfate aerosols and volcanic ash (Sekiya et al., 2016) was deactivated in order to prescribe stratospheric aerosols. Photolysis rates are calculated online and capture the effect of UV scattering by an enhanced stratospheric aerosol~~in~~. This model did not perform the senD2-chem simulation. The three senD2-sai ensemble members have different initial conditions,
280 which are based on the CMIP6 SSP2-4.5 ensemble simulations.

2.3.5 CCSRNIES-MIROC3.2 ("NIES")

The CCSRNIES-MIROC3.2 CCM was constructed on version 3.2 of the MIROC3.2 atmospheric general circulation model (K1 Model Developers, 2004), incorporating the gas-phase stratospheric chemistry module that was developed at NIES (Akiyoshi, 2000) and heterogeneous reaction module developed by Sessler et al. (1996). The spatial resolution is a T42 spectral truncation
285 ($2.8^\circ \times 2.8^\circ$) in the horizontal direction, and the model has 34 vertical levels of hybrid sigma–pressure vertical coordinates from

the surface to 0.01 hPa. The stratospheric chemistry module includes 42 photolysis reactions, 142 gas-phase chemical reactions and 13 heterogeneous reactions for sulfuric aerosols, supercooled ternary solution, nitric acid trihydrate and ice. Photolysis rates are calculated online and capture the effect of UV scattering by an enhanced stratospheric aerosol ~~in senD2-chem~~.

3 Results

290 3.1 Heating and Total Column Ozone Response

We start the analysis with the global mean stratospheric aerosol surface area density (SAD) integrated over all vertical levels above the tropopause to retrieve the average column loading shown in Fig. 2a. This serves to verify that the aerosol forcing remains consistent in the model output. By design, there is no temporal trend in senD2-fix while the stratospheric aerosol loading increases steadily in the senD2-sai simulation, with very good agreement between all models. The magnitude of the artificial aerosol forcing is comparable to the G6-1.5K-SAI experiment within the Geoengineering Model Intercomparison Project (GeoMIP Vioni et al., 2024), albeit slightly larger, because the temperature control starts earlier here. An anomalously strong annual cycle is present over the South Pole in CMAM (not shown), where the only output available is total SAD – as opposed to SAD from the sulfuric aerosol partition only. The surplus in CMAM thus comes from the addition of diagnostically calculated polar stratospheric clouds in the South Pole region and does not affect the analysis.

300 Next, we discuss the ~~main~~-effects of the simulated SAI ~~in the tropical annual mean~~ on the lower tropical stratospheric temperature. The direct consequence of an increase in stratospheric ~~aerosol sulfate~~ is heating, as the droplets strongly absorb outgoing terrestrial radiation and, in addition, the much smaller amount of incoming solar infrared radiation. All models experience strong ~~heating in the Junge layer~~ annual mean lower stratospheric heating, which is shown for the region between 115 and 70 hPa and 30° S and 30° N in Fig. 2b. The temperature increase varies between slightly more than 0.5 K per decade in WACCM and CMAM and about 1 K per decade in MIROC. Here, the models that include the same radiation scheme (WACCM and SOCOL: RRTMG; and MIROC and NIES: mstrnX) align closely in their temperature response, with WACCM and SOCOL showing about 1-2 degrees less warming than MIROC and NIES by 2080. This indicates that differences in the radiation scheme may cause significant differences.

Figure 2c shows the SAI signal (defined as the difference in the SAI simulation compared to no SAI simulation over the same time) in global mean total column ozone (TCO) in all five models (solid lines) and the chemistry-only signal (dashed lines). All models indicate a decrease in global mean TCO during at least the first two decades compared to the respective no-SAI senD2-fix simulation. In the later part of the century, models disagree on the sign of the response, with MIROC and NIES showing a relative ozone increase due to SAI, WACCM and CMAM showing near-zero changes, and SOCOL continuing to show a negative ozone response throughout the length of the simulation. As mentioned in the SOCOL model description, prescribing sulfate aerosols in SOCOL resulted in its excessive chemical sensitivity to aerosol increase at coldest temperature locations, i.e. the lowermost stratosphere. This is further amplified in the polar areas due to feedbacks between chlorine activation, temperature, and vortex strength effects. Although SOCOL ozone anomalies outside of the lowermost stratosphere are in a good agreement with other models with no apparent biases (Fig. 3), this effect dominates the global mean TCO sign response

for the given aerosol forcing. Regionally, this is the most apparent in the springtime South and North Pole TCO Fig. A1, where
320 SOCOL shows the strongest ozone depletion. Note, however, that the Antarctic ozone depletion due to SAI was particularly
strong also in the SOCOL G6 experiment with interactive aerosol microphysics, showing continuous TCO decline there even
at the end of the 21st century (Wunderlin et al., 2024). In addition to causing changes in global mean total column ozone, the
~~distribution of stratospheric ozone concentrations~~ zonal distribution of TCO values will remain perturbed while SAI is active, as
shown in Fig. A2. Models generally agree that the SH ozone response tends to be more negative by 2070-2083 than in the NH,
325 where TCO is mostly positive. The positive TCO response in the NH is the strongest in MIROC and NIES, which contributes
to their global mean TCO anomalies being the most positive and similarly distant with respect to CMAM as SOCOL, but with
opposite sign Fig. 2c.

While the injection rate for generating the SAI scenario increases almost linearly with time (see Fig. 2b of Tilmes et al.,
2025a), the increase in the aerosol SAD column weakens over time, ~~which is to be expected when the aerosol particles become~~
330 ~~large enough that gravitational settling becomes an important sink (Heekendorn et al., 2009).~~ The stronger increase in SAD
in the first 10-20 years is the result of increasing particle size with increasing injections (Tilmes et al., 2021). In contrast,
stratospheric temperatures continue to rise unabated, ~~likely due to the initially decreasing and then increasing behavior of~~
~~ozone and the associated changes in heating due to UV absorption~~ with continuously increasing mass. In the following sections
we explore the response of the ozone distribution and the role of different processes.

335 3.2 Spatial ozone distribution changes

While ~~global TCO gives an idea~~ annual and seasonal global TCO anomalies provide a view of the integrated ozone response,
Fig. 3 provides a much more detailed view of the spatial anomalies in ozone mixing ratios (senD2-sai - senD2-fix; "full
SAI"). The same figure but with ozone number concentration anomalies is provided in the Appendix (Fig. A3). As expected
with a steadily increasing aerosol injection rate, the anomalies increase over time, and regional changes become statistically
340 significant (t-test at 99 %). In all models the strongest signal occurs in the tropics, where a strong positive anomaly between 15
hPa and 5 hPa is sandwiched between two negative anomalies centered at about 40 hPa and 3 hPa. Directly above the tropical
tropopause, some but not all models exhibit a small positive ozone anomaly. In the last of these 14-year averages, the middle
stratospheric positive anomaly centered around 10 hPa and the negative anomaly above extend well into the mid-latitudes
and all the way to the polar regions in some cases. The positive anomaly gets weaker farther away from the tropics, however,
345 the negative upper stratospheric anomaly reaches a maximum at different ~~extra-tropic~~ extra-tropical latitudes – depending on
the model. This meridional trend is only observed in the mixing ratio anomalies, though, due to the climatological ozone
maximum in the tropics (see Fig. A3 for comparison). Finally, all models show significant ozone depletion over Antarctica (see
also Fig. A1; more pronounced in Antarctic springtime, but clearly also visible in the ~~the annual average~~ annual average Fig.
A2), whereas over the Arctic there are only weak ~~anomalies with low statistical significance and little agreement among the~~
350 ~~different models in the multi-year mean~~ negative TCO anomalies in the first few decades, which then become neutral and even
consistently positive in CMAM, MIROC, and NIES in the second half of the century Fig. A2. Even for boreal springtime, Fig.
A1 records no ~~clear~~ strong, lasting negative TCO anomalies in the Arctic ~~outside of variability~~ after 2050, except for SOCOL,

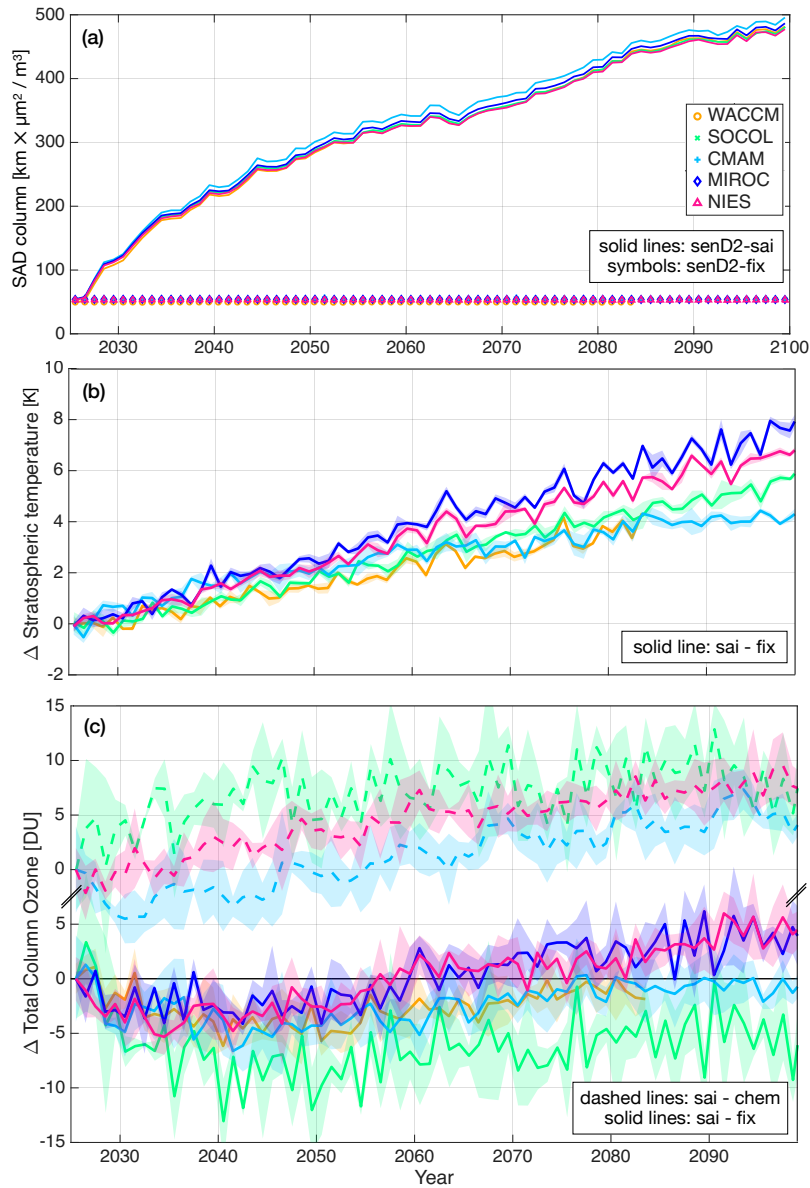


Figure 2. Stratospheric aerosol loading and resulting temperature and ozone changes for senD2-sai in the five models (color-coded). (a) Global and annually averaged column-integrated aerosol surface area density. Symbols: fixed aerosol forcing used in senD2-fix. Solid lines: aerosol forcing in senD2-sai required to maintain the original surface climate. (b) Stratospheric temperature anomaly due to SAI, senD2-sai - senD2-fix, averaged from 30°S to 30°N and from 115 to 70 hPa. (c) Global mean total ozone anomalies relative to 2025 due to full SAI (senD2-sai - senD2-fix; solid lines) and due to the dynamic signal alone (senD2-sai - senD2-chem; dashed lines). [Data for WACCM senD2-fix is only available until 2083.](#) Note the interrupted ordinate. Shaded areas in (b) and (c): confidence interval of one standard deviation.

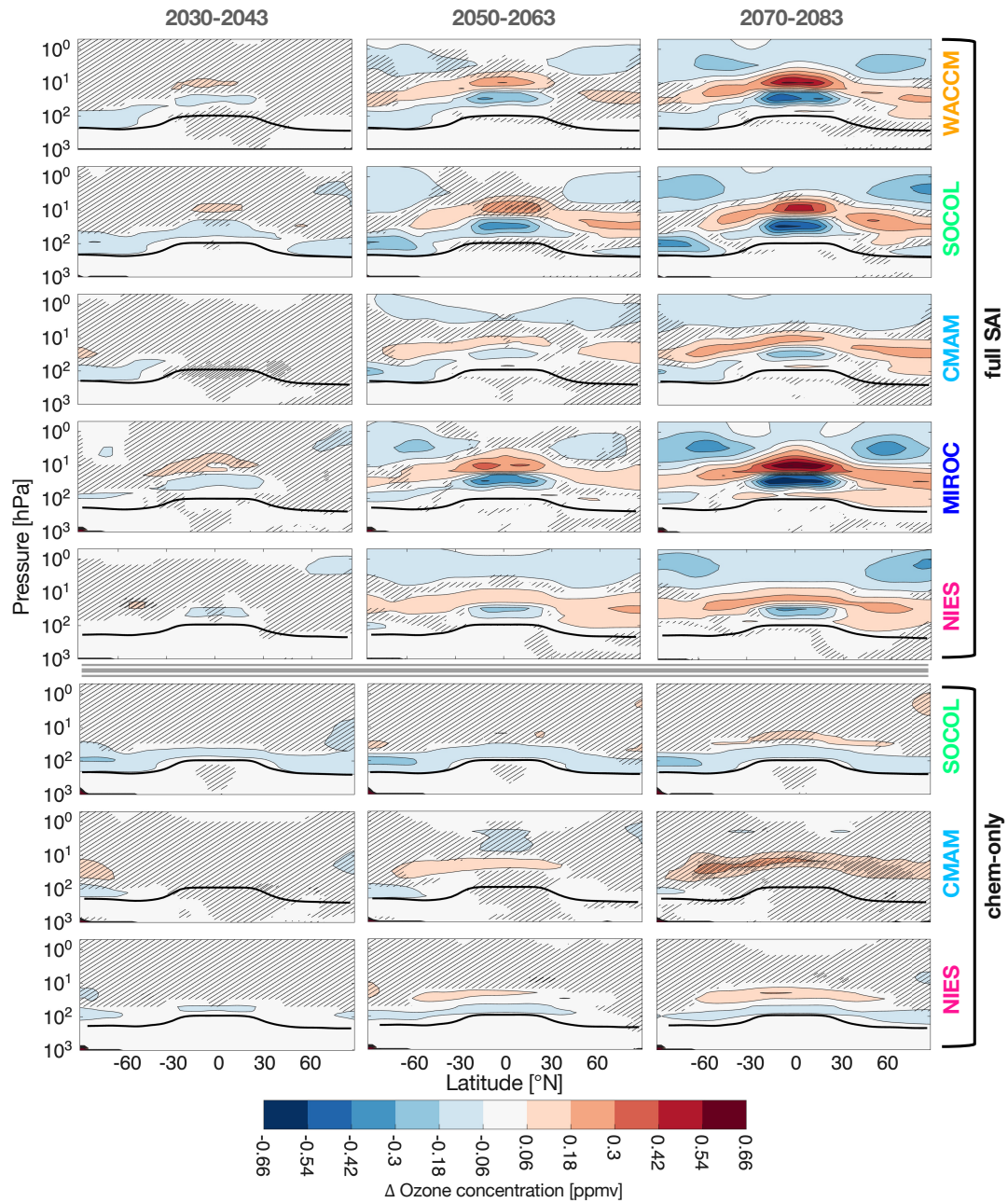


Figure 3. Zonal mean ozone mixing ratio anomalies in the full SAI scenario (senD2-sai - senD2-fix; top five rows) and for three of the models in the SAI chemistry-only scenario (senD2-chem - senD2-fix; bottom three rows). Anomalies are averaged over the 14-year time periods specified on top of each column. Black lines: mean senD2-sai tropopauses. Hatched areas: not significant at a 1 % level in a two-sample t-test.

where strong ozone depletion of some 20-40 DU ~~lasts~~persists until the end of the century ~~with no pronounced trend after the initial drop in the first 15 years, suggesting a balance between the competing effects of increasing SAI and decreasing CFC concentrations. Similar behavior is shown by WACCM in the Antarctic springtime~~ (Fig. A1).

The purely chemistry-driven signal – although differences in the implementation of photolysis exist – is diagnosed from senD2-chem - senD2-fix ("chem-only", see Section 2.1.4). Its anomalies (bottom three rows of Fig. 3) are both simpler and of smaller magnitude than the full SAI signal, revealing the dominant role of the circulation in shaping the SAI impacts on the stratosphere. In the lower stratosphere, there is a negative ozone anomaly in two models, pole-to-pole in a layer about 5 km thick directly above the tropopause. Above the negative anomaly there are hints of a weaker positive ozone anomaly confined to the low and mid-latitudes. In CMAM the chem-only ozone mixing ratio signal remains statistically insignificant throughout the simulation, possibly because of high variability in the single senD2-chem realization. However, in the ozone number concentration anomalies in Fig. A3 the polar negative anomalies are more visibly pronounced, which shows a consensus between the models over chemical ozone depletion in the lower stratosphere over the South Pole.

3.3 Dynamical effects

When using sulfur-based aerosol particles for SAI at low latitudes, an unavoidable side effect is the aerosol-induced heating of the tropical lower stratosphere. This direct aerosol heating affects atmospheric circulation and transport of species, including ozone, and any changes in dynamics further feed back into and modulate the temperature response and rates of chemical reactions. We later attempt to disentangle these different processes with our additional senD2-chem experiment, but start the analysis of dynamical effects with the impacts of SAI (i.e. senD2-sai - senD2-fix).

The strongest temperature increase due to SAI is found in the tropical lower stratosphere and is roughly limited to 30° N and S. Despite strongly enhanced aerosol mass concentrations at the poles the models consistently produce a cooling over the Antarctic and in some cases over the Arctic as well. Due to limited insolation and relatively weak infrared emission from the surface, the absorptive properties of the particles are not the dominant control over temperature there. Instead, polar temperature changes are more strongly indicative of changes in dynamical heating (due to changes in downwelling and/or wave forcing), or changes in polar ozone itself. Positive extra-tropical zonal wind anomalies in Fig. 4 indicate a substantial strengthening of the westerlies that form the polar vortex (reported in Visioni et al., 2020; Tilmes et al., 2022), especially in the Southern ~~hemisphere~~Hemisphere. The polar vortex intensification – as previously mentioned – is most pronounced in SOCOL, which coincides with the largest cooling in the polar lower stratosphere and thus strongest increase in meridional temperature gradient out of all models. This polar lower stratospheric cooling can be attributed to the isolation of the polar air within the strengthened polar vortex, as well as enhanced polar ozone loss under increased SAD from SAI, both of which feed back into the zonal wind anomalies. There are indications of a weakening of the SH subtropical jet in a very confined region near the tropopause at 30° S from significant negative zonal wind anomalies, and in some cases in the NH as well. These are more pronounced in seasonal zonal wind anomalies (not shown here). For this experiment it can be expected that anomalies in the subtropical jets remain weaker than in some other SAI modeling experiments, because here the stratospheric aerosol is the only difference between senD2-sai and senD2-fix; the surface climate is similar in the two simulations. In studies that

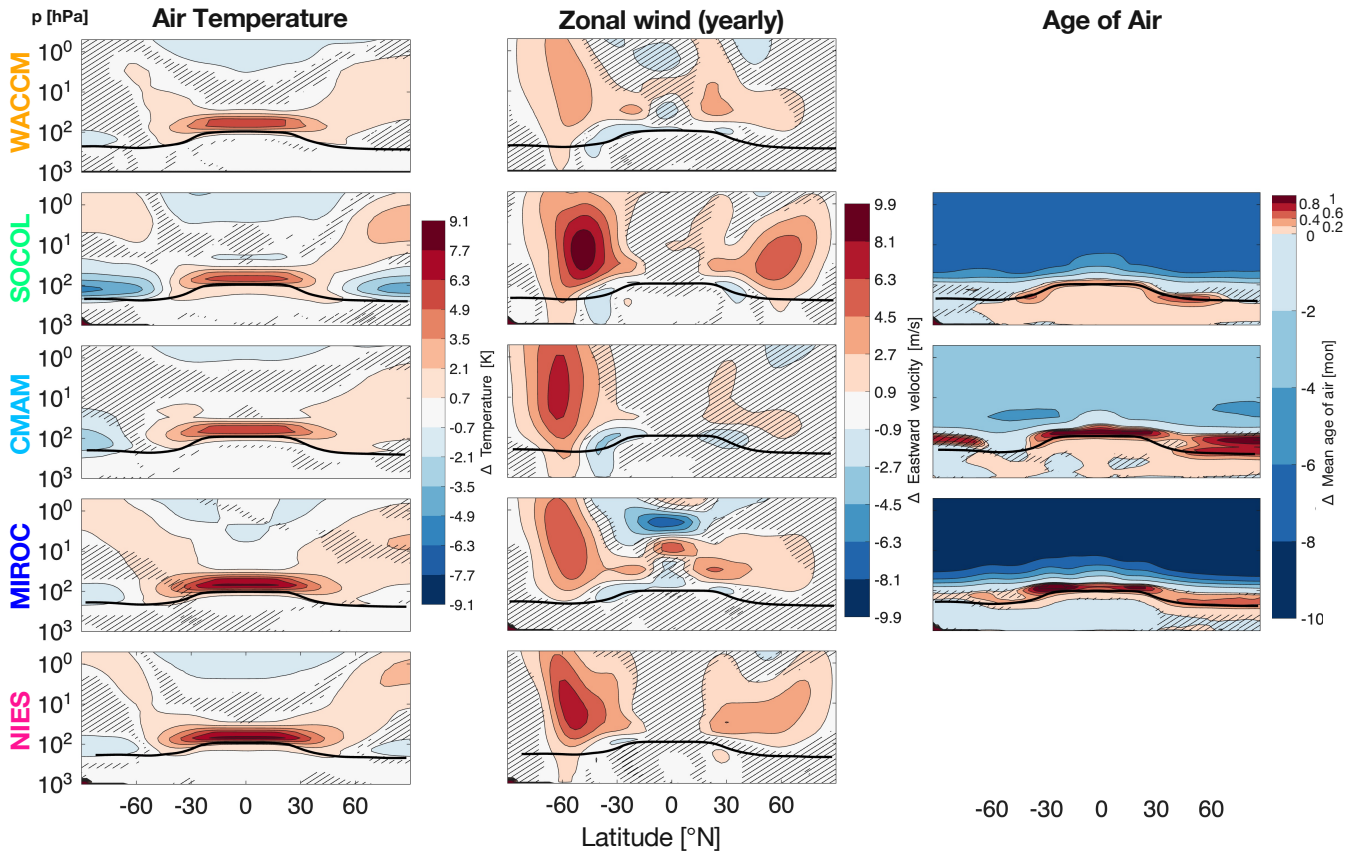


Figure 4. Zonal mean air temperature, zonal wind, and age of air anomalies due to SAI (senD2-sai - senD2-fix) for the 2060-2079 period. Black lines indicate the senD2-sai tropopause. Hatched areas are not significant at a 1 % level in a two-sample t-test.

compare a surface warming scenario with an SAI scenario, larger anomalies due to differences in the wave breaking intensity that contributes to the subtropical jets are expected (cf. Fig. 2e in Wunderlin et al., 2024). Three of the five models included an ideal-idealized age of air tracer, and the resulting changes in mean age of air (in months) are shown in Fig. 4. All three models show positive anomalies in the tropical lower stratosphere in a narrow region near the tropopause, suggesting a slow-down of the shallow branch of the **BDC** Brewer-Dobson circulation (BDC), and slowed tropical upwelling in the upper troposphere and lower stratosphere. Conversely, the air parcels in the entire middle and upper stratosphere up to the mesosphere experiences shorter residence times due to the aerosol heating-driven intensification of the deep branch of the BDC above the aerosol layer. The magnitude of this effect was found to be closely related to the magnitude of lower stratospheric heating within one model (Bednarz et al., 2023a) and here we see some indication of such a correlation existing also across the models here (with the largest heating rates observed in MIROC and the lowest in CMAM (see Fig. 2b). These substantial inter-model differences in zonal wind and age of air anomalies also suggest that the effect of an altered stratospheric aerosol burden on stratospheric dynamics, specifically transport barriers, may be majorly model-specific, which has previously been shown to be

[the case for the Mt. Pinatubo eruption \(Perny et al., 2025\)](#). An additional signal that is only present in WACCM and MIROC, the two models that have an internally-generated QBO, is a potential modulation of the QBO ([not investigated further here](#)), as evidenced by changes in the zonal winds in the tropical middle stratosphere. A slow-down of the QBO or even a complete stall are potential effects with sulfur-based SAI that have been found under specific injection strategies with strong aerosol heating (Franke et al., 2021; Jones et al., 2022), [however injections outside the equator minimize this effect \(Richter et al., 2018\)](#).

3.4 Tropopause and cold point anomalies

~~The heating that is inherently tied to sulfur-based SAI is the source of various secondary side effects that change components of the atmospheric system.~~ Besides the strong effects on stratospheric circulation, the tropopause height and cold point temperature are both affected [by SAI](#). The tropopause pressure anomalies (Fig. 5) are directly related to the temperature anomalies right above the tropopause. With strong heating the stratosphere effectively expands, lowering the tropopause height and thus resulting in positive tropopause pressure anomalies, ~~observed~~ [simulated](#) in all models in the tropics and well into the extra-
tropics. In comparison to other SAI scenarios the tropopause pressure anomalies only arise from stratospheric heating – not from tropospheric cooling, which does not exist in this experimental setup (compare e.g. Fig. 3 in Lee et al., 2023). The inter-model differences are very well constrained in the tropics with the individual models diverging more and more towards the poles, even showing different directions in change over the Arctic. This ~~well~~ reflects the level of ambiguity in temperature anomalies in those same regions [well](#). Additionally to changing the altitudinal temperature profile, the aerosol heating raises the tropopause minimum temperatures, i.e. the tropopause cold point in the tropics. Fig. 5 shows minimum temperatures over 14-year time windows in an early and late window of the simulations. All models exhibit a warming trend in the tropopause cold point temperature in the senD2-sai simulation compared to senD2-fix. The warming across all models stays within 1 K for the first time window, then reaches a median between 2-4 K in the second window, where the scenario is more progressed. The tropopause cold point change has implications for the moistening of the stratosphere, as the tropopause acts as the bottleneck for water vapor to pass through into the stratosphere (see below).

3.5 Effects of aerosol heating on key chemical species

With the changes to the temperature profile come important changes in the concentration of a number of chemical species. Fig. 6 shows water vapor (H_2O), methane (CH_4), and nitrous oxide (N_2O) mixing ratio percentage anomalies as a result of SAI. These species are all greenhouse gases and they can be a source of active agents participating in ozone depletion. All the anomalies above the tropopause in this figure are positive as a result of temperature and circulation changes. The strongest ~~percent~~ [percentage](#) increase in H_2O is found just above the tropical tropopause layer (TTL) in SOCOL, which features the largest tropopause cold point warming, freezing out less water and allowing air with a higher humidity to pass into the stratosphere. The ~~percent~~ [percentage](#) (as well as absolute) anomalies assume their highest values just above the TTL in all models and decrease with height and towards the poles. As a result the signal is less stratified than for CH_4 and N_2O , indicating a shorter lifetime due to ~~photolysis to hydrogen monoxide~~ [the reaction with excited, monatomic oxygen \$\text{O}\(^1\text{D}\)\$ to form hydroxyl radical \(OH\)](#). The strongest increase in CH_4 and N_2O is recorded in MIROC, which produced the highest heating rates [and the largest](#)

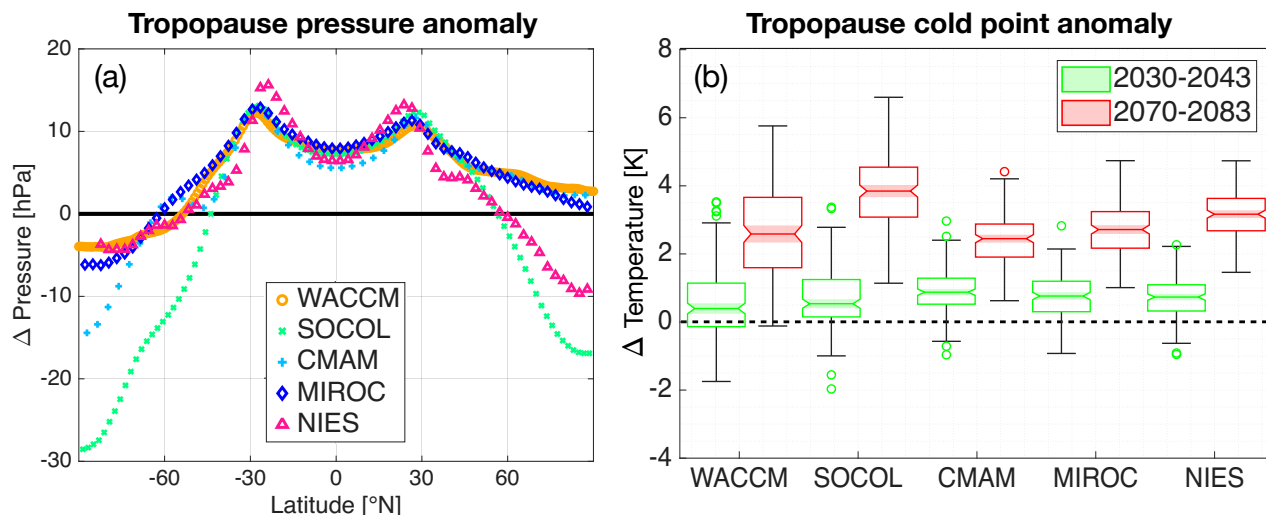


Figure 5. (a) Tropopause pressure anomaly as function of latitude caused by SAI (senD2-sai - senD2-fix) averaged over 2060-2079/2070-2083. (b) Anomaly in tropopause cold point temperature between 30° S and 30° N due to early-stage (green) and advanced (red) SAI (senD2-sai - senD2-fix). Whisker plot shows the spread of monthly mean data (minimum, first quartile, median, third quartile, maximum, and outliers).

increase in tropical upwelling. The other models show more modest positive anomalies, with CMAM even lacking a significant widespread signal in CH₄ anomalies. The strength of the CH₄ and N₂O anomalies are is very well anti-correlated with that of the age of air anomalies in the three models that provide them Fig. 2, but this is not the case for H₂O, which confirms that an additional process acts on H₂O, while the other two species are more tracer-like. The CH₄ and N₂O anomalies are largely stratified and increase with greater height, because these two gases are relatively long-lived and thus able to reach the upper stratosphere, where their concentrations in senD2-sai increase due to the stronger circulation. In the absolute CH₄ and N₂O anomalies there are maxima between 10 and 1 hPa (not shown here), while the percentage anomalies continue increasing with height due to smaller and smaller background concentrations. Fig. 6 demonstrates that the TTL temperature and circulation changes in the middle atmosphere exert meaningful control on the abundance and distribution of chemical species that play a role as source gases of ozone-depleting agents that we will consider next.

3.6 Partitioning of ozone-depleting agents in active and passive compounds

Chemistry in the middle atmosphere is characterized by chemical families that produce or destroy ozone in catalytic cycles and interact with each other by forming reservoir compounds. Changes to those interactions and changes to the concentrations of involved species result in new dynamic equilibria and a new ratio of active (i.e. capable of participating in ozone-depleting reactions) and passive (reservoir) species. The three major chemical families that are affected by sulfur-based SAI are the nitrogen oxide (NO_x), hydrogen-oxide-odd hydrogen (HO_x) and chlorine oxide (ClO_x) families. We assume bromine oxides (BrO_x) to play a minor role and do not include them in the analysis, although some bromine family reactions are included

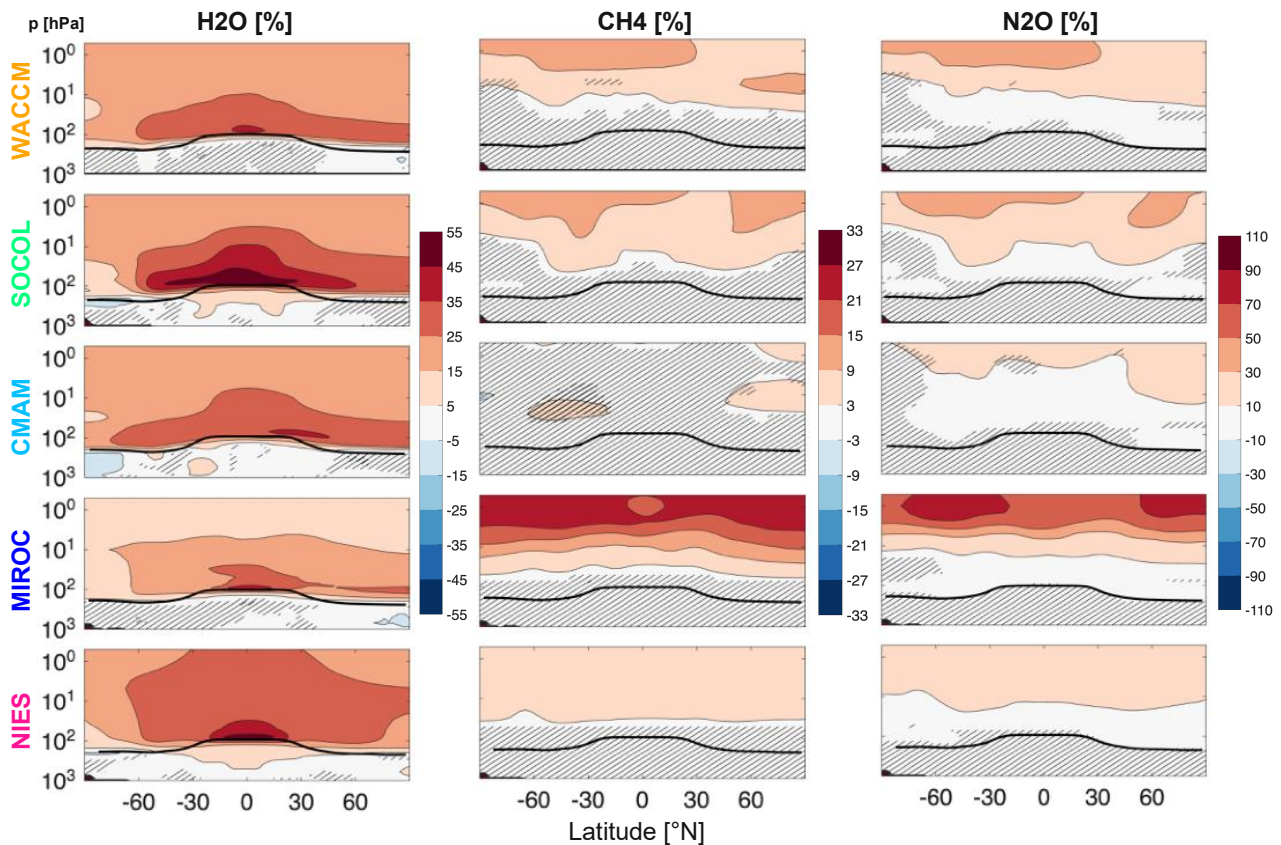


Figure 6. Relative zonal mean anomalies (in %) of water vapor, methane, and nitrous oxide mixing ratio caused by SAI (senD2-sai - senD2-fix) for the period 2060-2079. Black lines: senD2-sai tropopause. Hatched areas: not significant at a 1 % level in a two-sample t-test.

in the models. For NO_x the models are in qualitative agreement over a bipolar structure with negative anomalies in the lower and middle stratosphere – i.e. in the region where NO_x -catalyzed ozone depletion is most important – and positive anomalies above 10-5 hPa. In fact, the reversal from negative to positive signal is in very good agreement across the models, which is reflected in Fig. 9a, where only a small hatched region in the tropics around 5 hPa indicates that not all models agree on the sign of ozone change. The main disagreement comes from CMAM, which has the strongest NO_x loss that extends up slightly higher than the other models. The negative anomaly can be explained by the SAI-induced acceleration of N_2O_5 hydrolysis on aerosol surfaces that consumes NO_x and is further facilitated by enhanced water vapor availability. The positive anomaly higher up can be attributed to the dynamically-induced N_2O increases, which is a source gas for NO_x . Even better agreement is shown in the HO_x response to SAI, where all models show the same spatial structure stay within a factor of 1.5 of each other in absolute concentration change (see Fig. B1). The positive HO_x anomaly stems from [the](#) water vapor increase. The ClO_x trends are more complex, since the heterogeneous chemistry of this family is more sensitive to temperature. With aerosol surface area density levels many times that of the background state, chlorine activation on aerosol surfaces is expected to intensify in all

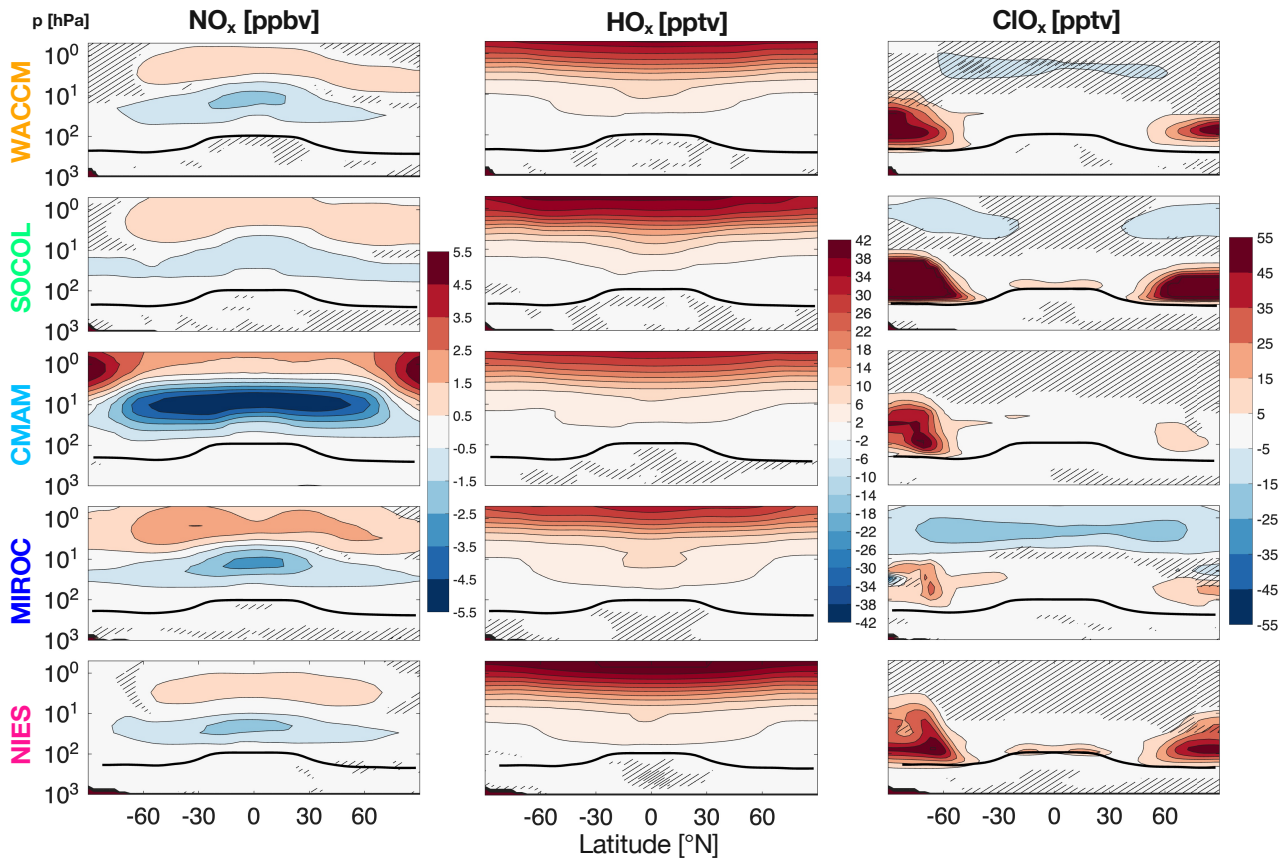


Figure 7. Zonal mean mixing ratio anomalies of active nitrogen oxide ($\text{NO} + \text{NO}_2 + 2 \times \text{N}_2\text{O}_5$), active hydrogen oxide ($\text{OH} + \text{HO}_2$), and active chlorine (oxide) ($\text{Cl} + \text{ClO} + 2 \times \text{Cl}_2\text{O}_2$) caused by to SAI (senD2-sai - senD2-fix) for the period 2060-2079. Black lines: senD2-sai tropopause. Hatched areas: not significant at a 1 % level in a two-sample t-test.

places, where the aerosol intersects the regions that reach low enough temperatures. Anthropogenic chlorinated compounds (e.g. CFCs) are still in the atmosphere after 2060, albeit at lower levels than in present day. Therefore, the models agree that SAI promotes the available chlorine to be converted into active species (ClO_x) in the lower stratosphere. The main hotspots for chlorine activation are the poles, where winter and spring conditions enable heterogeneous chlorine activation reaction chains, and in some cases the tropical tropopause. The degree of chlorine repartitioning into active species is particularly sensitive to the model, as demonstrated in the differences in magnitude of the anomalies in Fig. 7. In the upper half of the stratosphere there is also no clear consensus between the models with significant negative anomalies in only two models. We note here that absolute anomalies do not take into account the background state of each model, which may already contain different climatologies. We include percentage anomalies of the same species in the appendix (Fig. B1), which are especially highly variable in ClO_x , but also show substantial variance across models for HO_x and NO_x . Figure B1 shows that the upper stratospheric ClO_x percentage anomaly is very small compared to the lower stratosphere.

Since CFC concentrations and aerosol SAD are both prescribed in this experiment, differences in chlorine activation can only come from the model-specific treatment of heterogeneous reaction or ambient conditions (e.g. temperature). We provide a comprehensive table of heterogeneous reactions and which surfaces they are active on in each model in Table D1, which indeed shows that apart from the six main reactions on sulfate aerosols that all models include (listed in Section 2.3), only some models include more heterogeneous reactions. While these additional reactions may be able to explain, why NIES shows the strongest tropical lower stratospheric chlorine activation, it fails to explain why MIROC, with the same reactions as NIES, shows lower chlorine mixing ratio anomalies overall, even being close to zero in the tropical lower stratosphere despite MIROC having a lower tropopause cold point temperature than NIES (Fig. 5b). The number of reactions included also does not generally correlate with the intensity of chlorine activation in the other models. Therefore, it is more likely that differences in the implementation of heterogeneous reactions (e.g. reaction rates), or differences in ambient conditions are the cause of this variability.

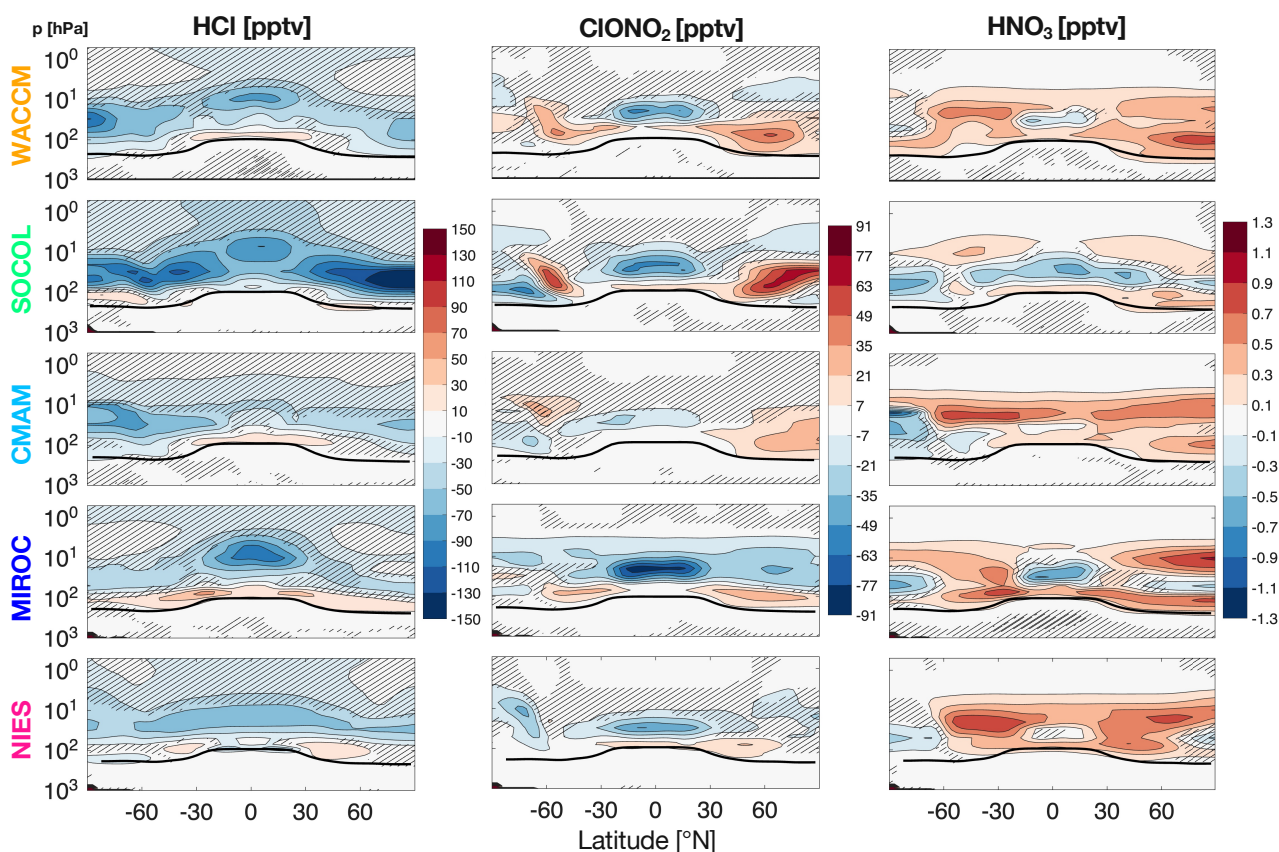


Figure 8. Zonal mean mixing ratio anomalies of hydrochloric acid, chlorine nitrate, and nitric acid mixing ratio $[(\text{senD2-sai} - \text{senD2-fix}) / \text{senD2-fix}]$ caused by SAI for the period 2060-2079. Black lines: senD2-sai tropopause. Hatched areas: not significant at a 1 % level in a two-sample t-test.

To better understand the activation/deactivation effect of SAI we also provide the anomalies of some important reservoir gases that represent the passive partition. Given the right conditions, ClO_x and NO_x can be converted into hydrochloric acid (HCl), chlorine nitrate (ClONO_2), and ~~nitrous~~nitric acid (HNO_3) – or, in reverse, the reservoir gases can be activated. We show mixing ratio anomalies of these three reservoir gases due to SAI in Fig. 8. In the lower and middle tropical stratosphere there is a strong dynamical control on all reservoir gases, which creates a similar dipole structure of positive anomalies around and just above the tropopause and negative anomalies between some 70 and 30 hPa. The increase in the lower stratosphere is especially well recorded in the percentage anomalies of the reservoir gases (supplied in the Appendix Fig. B2) and comes from the slowed shallow branch of the BDC, which draws tropospheric air up into that region less quickly and exports it out of that region less efficiently than without SAI (cf. Fig. 4). This causes a relative increase in concentrations, as tropospheric air has lower concentrations of all three reservoir gases, and the effect is mainly limited to the tropics, with some leakage into the mid-latitudes. The negative tropical anomalies above partly stem from the strengthened deep branch. This pattern is only broken in the HCl anomalies of SOCOL and NIES, which are entirely negative in the tropical lower and middle stratosphere. This comes from the chemical conversion of HCl into ClO_x on sulfate aerosol particles, as evidenced by the ClO_x enhancement there in the same two models (Fig. 7). Heterogeneous chemical chlorine activation also leads to negative HCl anomalies across models from some kilometers up from the tropopause into the middle stratosphere in the extratropics. Further up, in the middle stratosphere and towards the mesosphere, HCl does not show significant anomalies, ~~as there is~~because temperatures are high enough to evaporate any aerosol that may be transported up from below, leaving no surface area for heterogeneous chlorine activation to take place on. Chlorine nitrate mostly lacks significant anomalies that are consistent across the models, suggesting a complex change in the new chemical equilibrium due to multiple processes – or, possibly, differences in the heterogeneous chemical reactions that are included in each model and influence ClONO_2 . In HNO_3 some negative anomalies exist in the lower southern polar stratosphere (and sparsely in the North, although models disagree) and extend into the troposphere and are likely the result of denitrification. Apart from these regions and the dynamical negative anomaly in the tropics, however, HNO_3 anomalies are widely positive with a peak around 20-30 hPa, due to N_2O_5 hydrolysis (correlating with NO_x depletion).

3.7 Multi-model mean and key regions

In this section we synthesize the key findings from across all models and point out the inter-model agreements and differences. Fig. 9 shows the ozone anomalies averaged over all five participating models for the same time period (2060-2079, inclusive) used before. We show the anomalies in ozone concentration due to SAI ("full") in panel (a). Panel (b) shows the anomalies from only the chemical effects of SAI ("chem") and panel (c) shows the difference between "full" and "chem", which equates to the linearly isolated dynamical effects ("dyn") of SAI. This does not represent a clean isolation of dynamical effects, but rather an approximation, where nonlinearities and differences in treatment of photolysis due to aerosol UV scattering in senD2-chem are still included. Since panels (b) and (c) require the senD2-chem simulation, they only include the three models for which these simulations are available: SOCOL, CMAM, NIES. The "full" anomaly for these three models only – with respect to which "dyn" is derived – is given in Fig. C1 in the appendix. Shading in the entire figure indicates that at least one of the models disagrees about the sign of the anomaly. We define some key regions in the latitude-vertical domain of the panels to distinguish

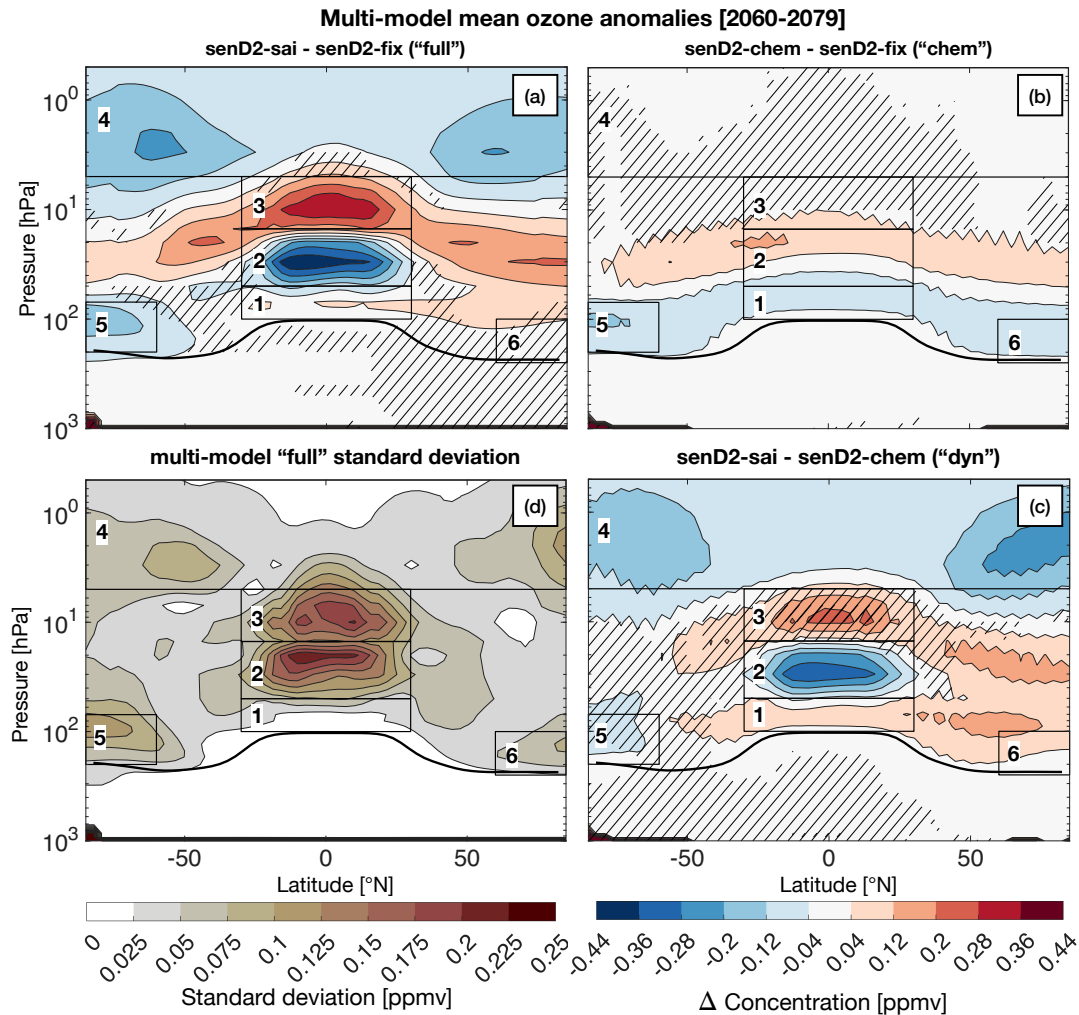


Figure 9. Multi-model mean zonal mean ozone mixing ratio anomalies due to full SAI (senD2-sai - senD2-fix; panel a), the standard deviation of full SAI in panel (d), only chemical effects of SAI ("chem"; panel b), isolated dynamical effects ("dyn"; panel c). Panel (a) is the multi-model mean of all five models, panels (b) and (c) show multi-model means of the three models that simulated the senD2-chem simulation. The "full" model mean of the same models that were used in panels (b) and (c) is not shown here, but added in Fig. C1. The standard deviation in panel (d) pertains to panel (a). The black rectangles show regions with a dominant process shaping the ozone anomaly signal. The black line indicates the tropopause. Regions where the sign of all models do not agree are hatched.

between different dominant processes. We give a short summary for each of these regions to describe the multi-model ozone anomalies:

- 1 In the [tropical](#) lowermost stratosphere the SAI effect on ozone is smallest out of the chosen regions, because the negative chemical effect and the positive dynamical effect almost cancel each other out. While the deep BDC branch increases,

the shallow branch is weakened (see Sect. 3.3). This leads to less transport of ozone-poor air from the troposphere to the tropical lower stratosphere, less efficient export to mid-latitudes and thus an increase in ozone concentration. This effect is partly counteracted by tropical chlorine activation, which is apparent in "chem", but the anomaly is still consistently positive across all models (see also the ozone number concentration anomaly in Fig. A3).

- 525 2 In the tropical lower stratosphere the models agree well in "full", "chem", and "dyn". The strong negative anomaly in "full" can be attributed almost entirely to dynamics. The chemical contribution in region 2 changes from negative (bottom) to positive (top) in this region, but is weak overall. The dynamical signal arises from the strong aerosol absorption heating, which increasingly draws comparatively ozone-poor air up from below. The largest variability here is observed towards the top of the box, which indicates that the negative anomaly varies one or two model levels in extent between
- 530 the models and the reversal to positive comes earlier (lower down) in some than in others.
- 3 Region 3 receives its positive anomaly in all models from the increased BDC deep branch strength, like region 2. Air comparatively rich in ozone from the main formation region below is transported faster in the tropical pipe and brings higher ozone concentrations to region 3. However, there is also a significant chemical contribution of NO_x suppression through N_2O_5 hydrolysis, which curbs catalytic ozone destruction. From the "chem" anomalies this contribution seems
- 535 minor, but this does not take feedback between dynamics and chemistry into account. With strong stratospheric moistening (see Fig. 6; not present in senD2-chem by design) N_2O_5 hydrolysis is strengthened significantly (see HNO_3 in Fig. 8) and may become the dominant process.
- 4 Similar to region 3 the warming of the TTL cold point and acceleration of the BDC increases the concentration of key chemical species, in this case the HO_x and NO_x precursor gases H_2O and N_2O . This leads to significant ozone depletion
- 540 in the upper stratosphere and into the mesosphere. The "chem" anomalies do not record this effect, since it only emerges when dynamical effects drive chemical changes.
- 5 The South polar lower stratosphere experiences a decrease in Antarctic lower stratospheric ozone and ozone column (equivalent to a delay in recovery) consistent across models, even late in the second half of the ~~21st~~^{21st} century. This can be attributed to heterogeneous chlorine activation on the increased aerosol surface, apparent in "chem". Even though
- 545 "dyn" shows almost no significant anomaly in region 5, there is also interplay between chemistry and dynamics in this region. Figure 4 demonstrates the dynamical isolation of the (South) polar regions with the strong zonal wind strengthening and temperature decrease. ~~This provides improved conditions for~~ The colder conditions promote aerosol formation and growth, which further ~~benefits~~enhances the chemical ozone destruction.
- 6 The North polar lower stratosphere does not experience as much isolation as the South ~~pole~~^{Pole} and conditions for ozone depletion do not form as consistently, hence dynamical changes (including inter-annual variability) can play a comparatively larger role. The models therefore do not agree on the sign of the anomaly. SOCOL and NIES both calculate a stronger chemical effect with resulting negative effects for ozone, while the other models tend to show positive anomalies by the second half of the century, likely indicative of the enhancement in ozone transport to Arctic
- 550

stratosphere under strengthened deep branch of BDC (although some differences could also be caused by the large inter-annual variability characterizing the Arctic stratosphere and hence difficulty in isolating the forced response to SAI in these simulations (see also TCO in Fig. 10). The significant positive anomaly in "chem" makes it clear, however, that when positive anomalies appear in "full", it is the dynamical effect that has to counter and overwhelm chlorine activation.

Figure 9 resolves the intricate latitude-pressure pattern of ozone anomalies with height, where positive and negative anomalies are stacked above each other. This illustrates that SAI causes a major redistribution of ozone in the middle atmosphere (which has implications for stratospheric conditions), even in places, where the total column ozone does not change. Nevertheless, changes in TCO are a good indicator for changes in radiation received at the surface. In Fig. 10 we show the multi-model mean temporal evolution of TCO in Hovmöller plots; again including "full" (panel a) from all five models including standard deviation in panel (d) and "chem" and "dyn" for SOCOL, CMAM and NIES. A multi-model mean of these three models and their standard deviation is provided in Fig. C2.

In the "full" multi-model mean the subtropics generally show little inter-model agreement in the sign of the anomaly (hatched regions). The tropics part reveals that the vertically stacked anomalies from Fig. 9a almost cancel out. Over time, a growing negative trend emerges, reaching up to around 8 DU of ozone depletion with some consistency across models. The negative effect of "chem" – as the ozone decrease in the lower stratosphere from halogen activation dominates over the ozone increase above it from NO_x deactivation – remains roughly constant with time in low-latitudes, as SAD increases but background halogen levels available for activation decrease with time. It is countered by "dyn" (which accounts for direct transport effects and thermodynamical influence on chemistry) at first, although models partially disagree in sign. By 2060, however, "dyn" also starts to add a significant negative contribution in the tropics, strengthening the chemical effect and leading to a combined negative anomaly of ~ 5 DU in "full" by 2080.

In the South, SAI causes ozone depletion throughout the entire simulated period that the models agree well on, despite inter-annual variability in polar conditions (polar vortex strength and temperatures). While there is still a positive signal in "dyn" in the Southern mid-latitudes from SAI-induced changes in both the shallow and deep branch of the BDC that balances the negative "chem" contribution, it wanes towards the South ~~pole~~^{Pole}. In contrast, the positive "dyn" anomaly frequently extends into polar latitudes in the North and grows considerably in strength over time, because the North pole is dynamically less isolated. This causes a transition from significant negative to significant positive TCO anomalies in the "full" multi-model mean.

The transition occurs in the 2050s and marks a change in regime, where the ~~the~~ positive TCO anomaly of altered dynamics dominates the negative chemical anomaly that is still persistent in "chem" even by the end of the century. Figure 10a shows that in addition to vertical ozone redistribution, competing effects also cause a hemispheric imbalance in the CCMI-2022 SAI scenario – even though the global mean TCO analysis (see Fig. 2c) largely suggests small global ozone changes in the second half of the 21st century for many models.

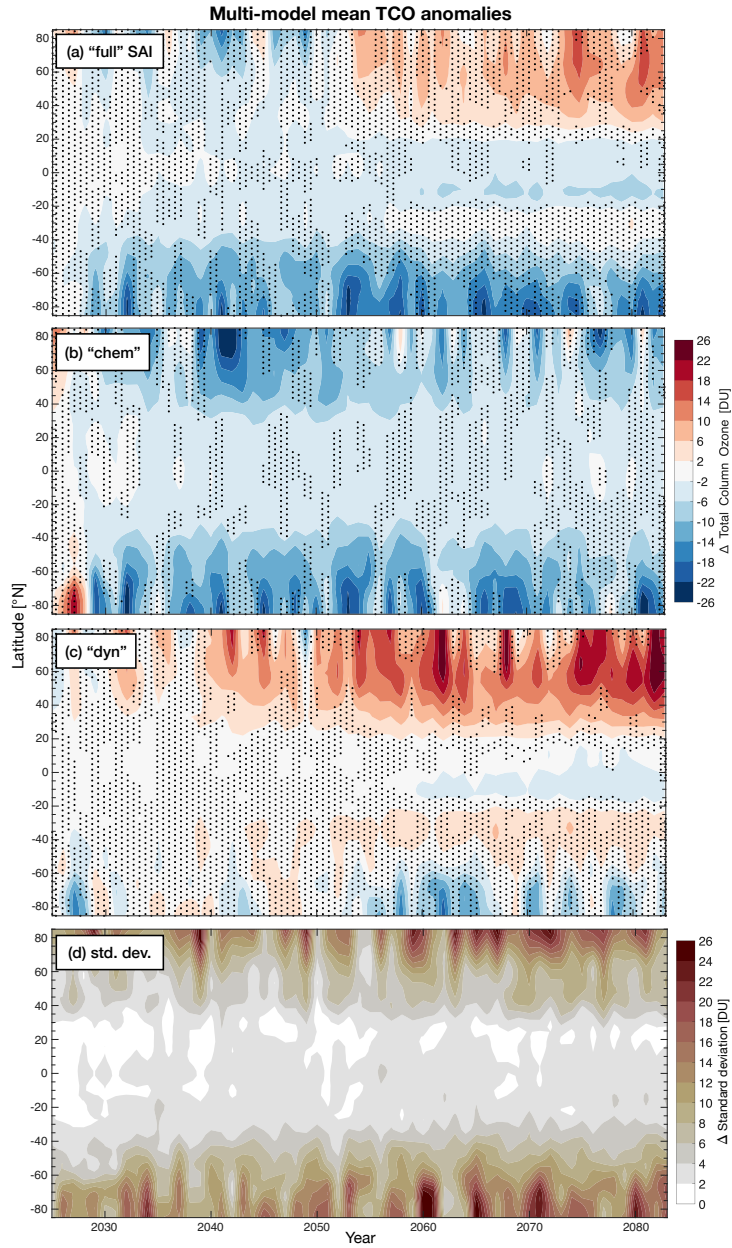


Figure 10. Multi-model mean zonal mean total column ozone anomalies due to full SAI (senD2-sai - senD2-fix; panel a), only chemical effects of SAI ("chem"; panel b), isolated dynamical effects ("dyn"; panel c) and the standard deviation of full SAI in panel (d). Panel (a) is the multi-model mean of all five models, panels (b) and (c), and (d) show multi-model means of the three models that simulated the senD2-chem simulation. The "full" some-model mean of the models that were used in panels (c) and (d) is not shown here, but added in Fig. C2. The standard deviation in panel (d) pertains to panel (a). Regions where the sign of all models do not agree are stippled.

4.1 Agreement and differences between models

In this analysis, we examined the chemical and dynamical response of the middle atmosphere to a specific SAI scenario with uniformly prescribed stratospheric aerosol properties in five models. This comparison primarily serves to highlight where the models agree well and to understand the characteristic features of individual models and where they diverge. Since this experiment eliminated differences from model-dependent aerosol microphysics modules, the remaining differences we show in this study indicate which other model components require further testing and development to improve multi-model projections under SAI.

Overall, the models qualitatively agree very well qualitatively on ozone concentration anomalies in different regions, which points to a good process understanding. In many regions feedbacks between dynamics and chemistry representation of the processes involved. The separation of processes using the additional senD2-chem simulation demonstrated that in many regions heating-derived effects, i.e., dynamical changes (transport response to temperature changes) and dynamically-induced changes in chemistry – which we also call nonlinear feedback effects – define the SAI response, and even though they are largely consistently captured by the five model analyzed, some uncertainties still exist. Since the SAI aerosol was prescribed in this experiment, the remaining uncertainties are indicative of the processes that need to be improved in the models to further constrain the inter-model uncertainty. Even though these effects, which contribute to ozone anomalies, appear regionally consistent across the five models analyzed, important uncertainties remain in their magnitudes. For one, the importance and persistence of chlorine activation – a key contributing factor to the ozone anomalies in the (polar) lower stratosphere – strongly varies across models, yet it is a key defining part of the signal in the (even when disregarding the high sensitivity in SOCOL), with SOCOL and NIES even showing opposite trends above the tropical tropopause compared to the other three models, in the form of an increase in ClO_x and decrease in HCl. High variability is also recorded in the polar lower stratosphere, especially at the poles, where ozone holes from anthropogenic ozone-depleting substances are still a major issue. Similarly, where springtime total column ozone anomalies averaged over 60-90° latitude range from around -90 to -10 DU and -30 to +30 DU over the South and North Pole, respectively, by the year 2100.

More variability between models, despite prescribing the same aerosol optical (radiative) properties in all models, the temperature is found in the rate of change in the center of the aerosol layer lower tropical stratospheric temperature, which ranges from just over 0.5 (CMAM) to just over 1 K (MIROC) per decade in Fig. 2b. The differences are correlated with the use of different radiation schemes. This results in a high variance in the transport of chemical species, especially in the tropical pipe. We argue that these differences should be studied in coordinated idealized simulations to further increase the robustness of projections of SAI scenarios. Dedicated experiments could isolate the inter-model spread due to heterogeneous chemistry and the radiative transfer scheme. Besides that, we find that despite the It is exemplified in MIROC that shows the strongest acceleration of the deep branch of the Brewer-Dobson circulation producing the highest CH_4 and N_2O increases in the upper stratosphere. For water vapor, in turn, SOCOL shows the highest positive percentage anomaly (over 50 %) in the lower stratosphere due to its large tropopause cold point warming, while NIES has a strong positive H_2O anomaly of some 30 % that extends up highest into

the upper stratosphere. CMAM projects particularly strong NO_x passivation in the middle stratosphere and WACCM shows the smallest zonal wind anomalies, possibly due to its rather low equator-to-pole temperature gradient in the lower stratosphere. Due to these various features unique to specific models, the multi-model mean presented in this study should be interpreted with the experimental setup and the model spread in mind. Even still, the spatial pattern of ozone anomalies and thus the extent in latitude and altitude of the dominant processes controlling ozone in each region, is in remarkably good agreement across all models.

There are various ways to investigate key differences between the models using more idealized simulations. To further isolate differences in heterogeneous chemistry, model experiments could be constrained by nudging the wind and temperature fields, in addition to fixed aerosol fields, thereby removing all dynamical differences among the models. Differences could arise from handling the production of nitric acid trihydrate and ice clouds, in both tropics and high latitudes, or differences in the heterogeneous reaction rates. In addition, radiative transfer schemes could be tested through offline calculations, while impacts of photolysis rates could also be tested, but only in models that can handle both prescribed and interactive photolysis schemes. A follow-up "dynamics-only" senD2-dyn simulation is also planned for a more elaborate analysis of the chemistry-transport-dynamical feedbacks.

4.2 Limitations of aerosol forcing implementation and future recommendations

Despite the best efforts of the CCMI-2022 experiment to eliminate differences in the aerosol forcing, some aspects of implementation in the models may have accounted for minor inconsistencies. ~~We shortly discuss these and offer suggestions to further improve uniform implementation.~~

4.3 Limitations of aerosol forcing implementation and future recommendations

Prescribing stratospheric aerosol derived from one model in another model poses a distinct challenge: the tropopause height is not necessarily the same. Consequently, if a stratospheric mask is applied to the aerosol data, some of the aerosol load may be truncated. This problem has not been addressed in this study, but in the future one could interpolate the stratospheric aerosol input onto the model-specific part of the stratosphere from the tropopause to an isotherm defining the aerosol evaporation threshold at every timestep to include the entire forcing. This method would also allow for the implementation of a full latitude-longitude-pressure field in any new experiments with a prescribed aerosol forcing. Accounting for zonal variations in the aerosol is secondary to meridional variations due to fast zonal mixing time scales, but could introduce a bias in heterogeneous chemistry. Processes like chlorine activation are limited by aerosol surface, the formation of which is sensitive to temperature and thus often appears regionally and bound to orographic features. Imposing a zonal mean aerosol forcing may not be sufficient to accurately reflect the occurrence and impact of heterogeneous chemistry. However, this would still not solve the problem of the inconsistency between the in-situ model transport barriers and the shape of the monthly mean aerosol forcing, which can result in different amounts of aerosol being taken into account within the wintertime vortex, important for ozone chemistry.

This can be solved by reshaping the forcing along a potential vorticity coordinate system (i.e., using equivalent latitude instead of the geographic latitude-longitude grid).

Code and data availability. Parts of the model data are available for download from the Centre for Environmental Data Analysis (CEDA) Archive (<https://data.ceda.ac.uk/badc/ccmi/data/post-cmpi6/ccmi-2022>) (Center for Environmental Data Analysis (CEDA), 2025). The full post-processed data used for this analysis can be downloaded from Zenodo (<https://doi.org/10.5281/zenodo.18331210>) (Jörmann, 2026). The REMAP code and its products are freely available for download from the ETH research collection (see code and data availability statement in Jörmann et al., 2025).

Appendix A: Aerosol-forcing-detailsSupplementary ozone anomalies

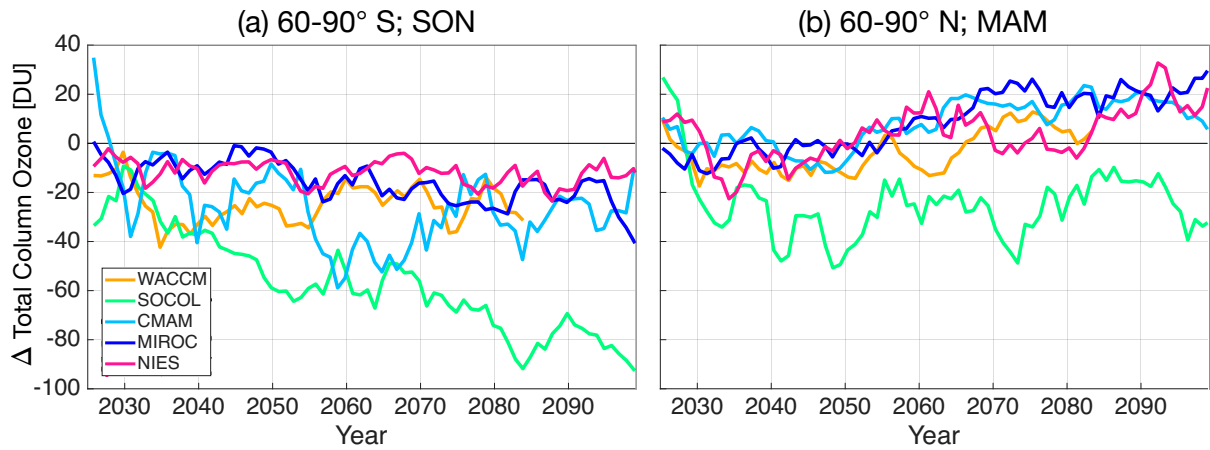


Figure A1. Mean total column ozone anomalies due to full SAI (senD2-sai - senD2-fix) from 60 to 90° S in austral spring (SON; panel a) and from 60° to 90° N in boreal spring (MAM; panel b). All timeseries have been treated with a moving mean filter of 5 years width.

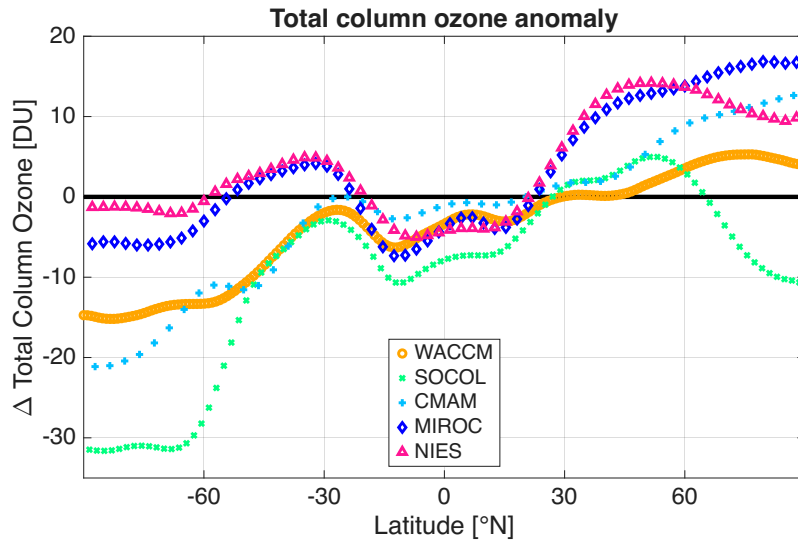


Figure A2. Zonal-mean Mean total column ozone number density anomalies in the due to full SAI scenario (senD2-sai - senD2-fix; top five rows) and for three from 2070-2083 as function of the models in the SAI chemistry-only scenario (senD2-chem - senD2-fix; bottom three rows) latitude. Anomalies are averaged over the 14-year time periods specified on top of each column. Black lines: mean senD2-sai tropopause. Hatched areas: not significant at a 1-level in a two-sample t-test.

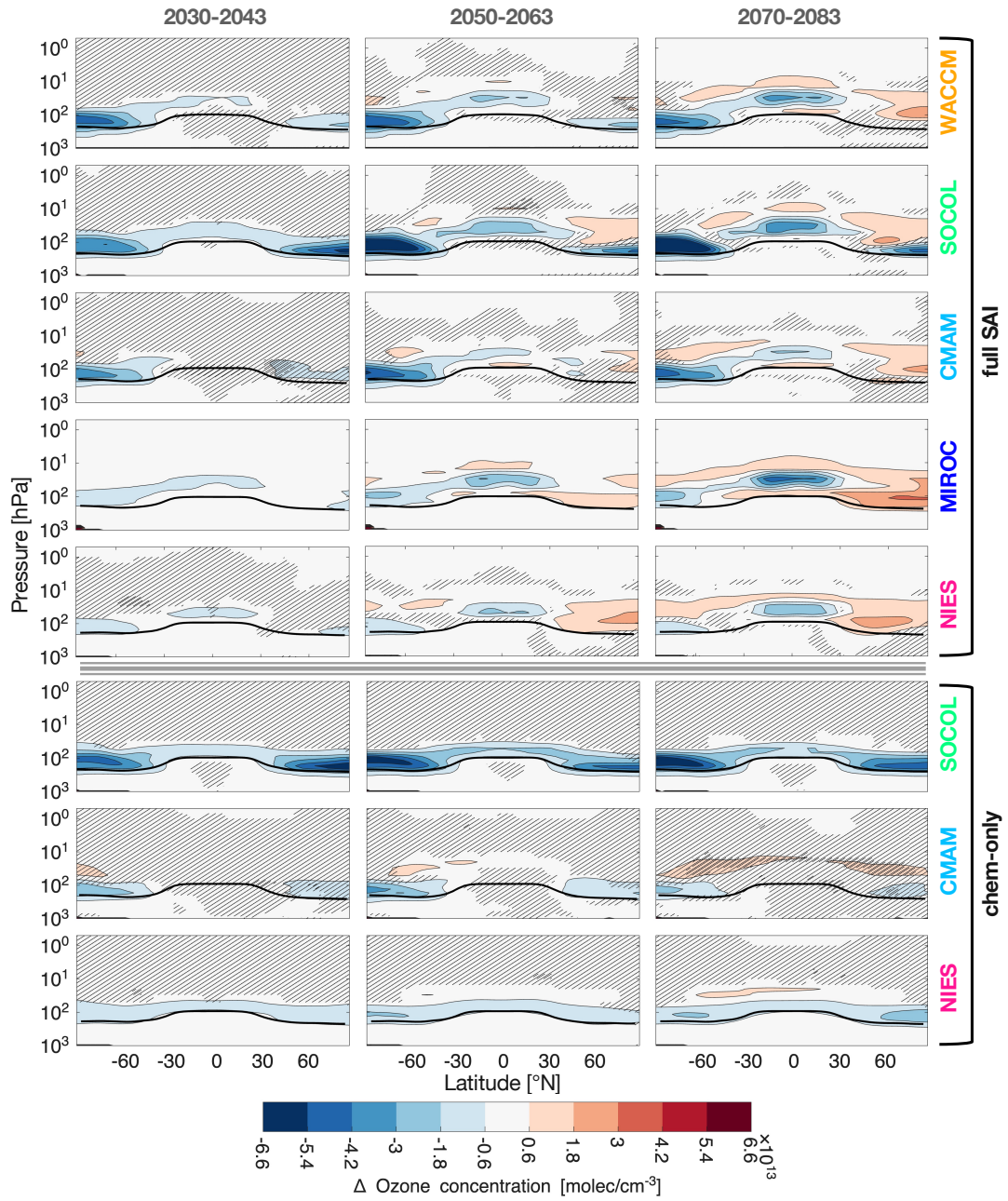


Figure A3. Zonal mean ozone number density anomalies in the full SAI scenario (senD2-sai - senD2-fix; top five rows) and for three of the models in the SAI chemistry-only scenario (senD2-chem - senD2-fix; bottom three rows). Anomalies are averaged over the 14-year time periods specified on top of each column. Black lines: mean senD2-sai tropopause. Hatched areas: not significant at a 1 % level in a two-sample t-test.

Appendix B: Active families and reservoir gases percentage anomalies

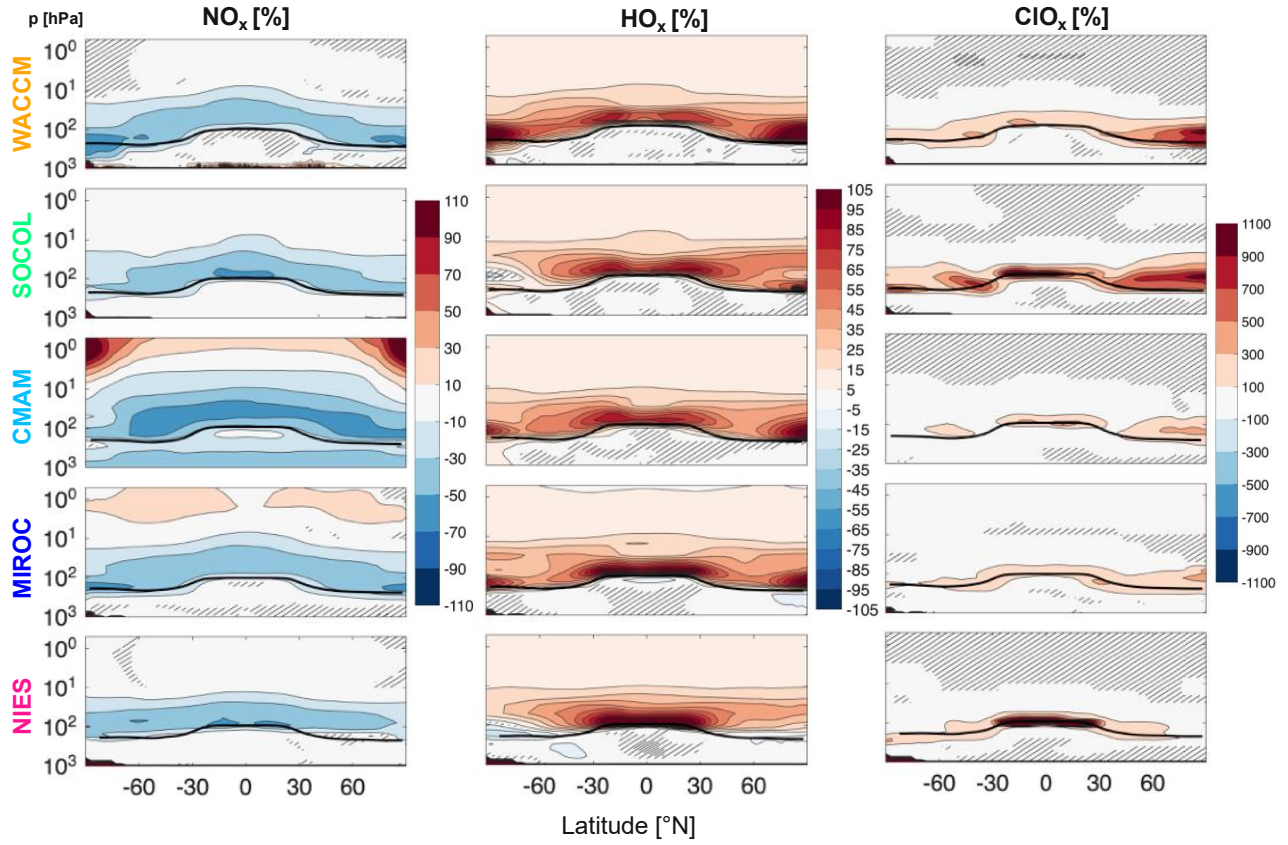


Figure B1. Zonal mean active nitrogen oxide ($\text{NO} + \text{NO}_2 + 2 \times \text{N}_2\text{O}_5$), active hydrogen oxide ($\text{OH} + \text{HO}_2$), and active chlorine (oxide) ($\text{Cl} + \text{ClO} + 2 \times \text{Cl}_2\text{O}_2$) percentage anomalies due to SAI $[(\text{senD2-sai} - \text{senD2-fix}) / \text{senD2-fix}]$ for the 2060-2079 period. Black lines indicate the senD2-sai tropopause. Hatched areas are not significant at a 1 % level in a two-sample t-test.

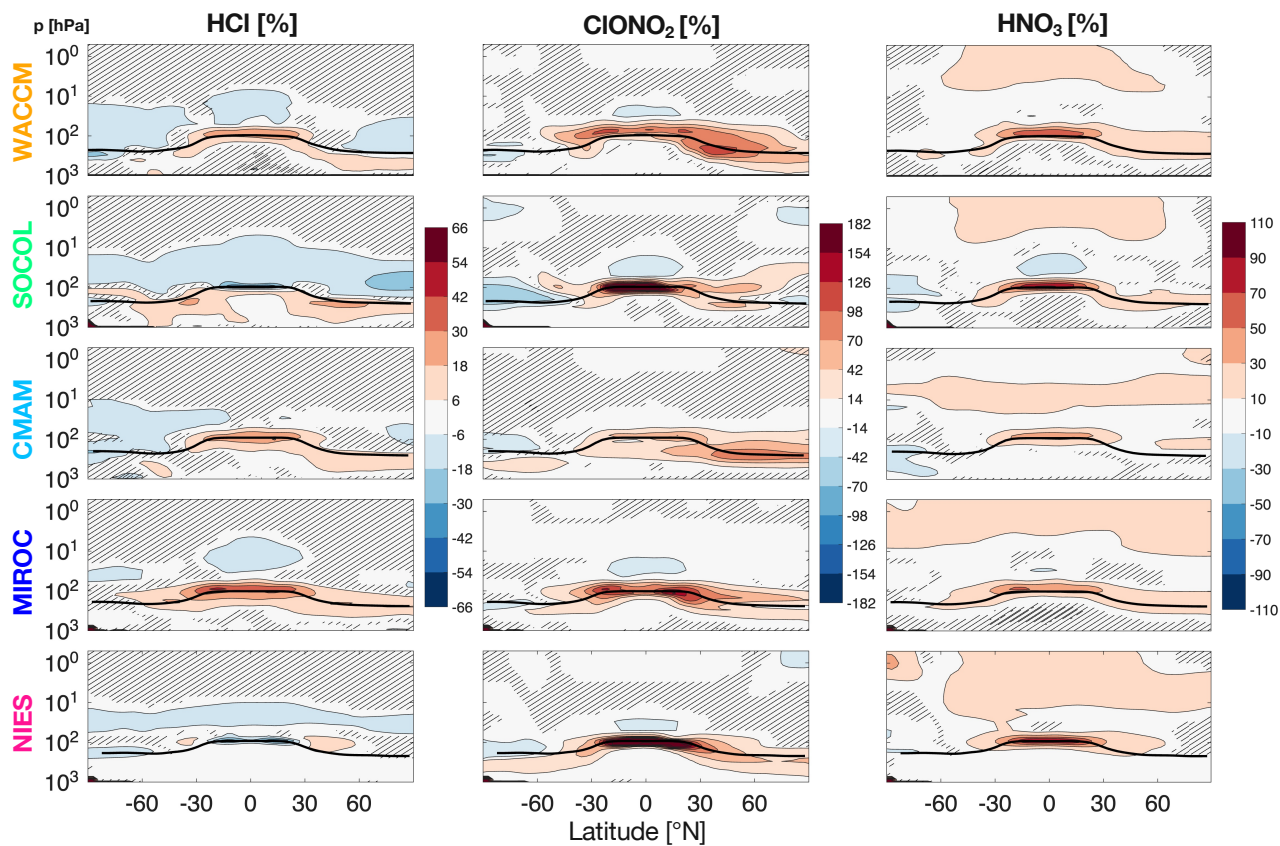


Figure B2. Relative zonal mean anomalies (in %) of hydrochloric acid, chlorine nitrate, and nitric acid mixing ratio $[(\text{senD2-sai} - \text{senD2-fix}) / \text{senD2-fix}]$ caused by SAI for the 2060-2079 period. Black lines: senD2-sai tropopause. Hatched areas: not significant at a 1 % level in a two-sample t-test.

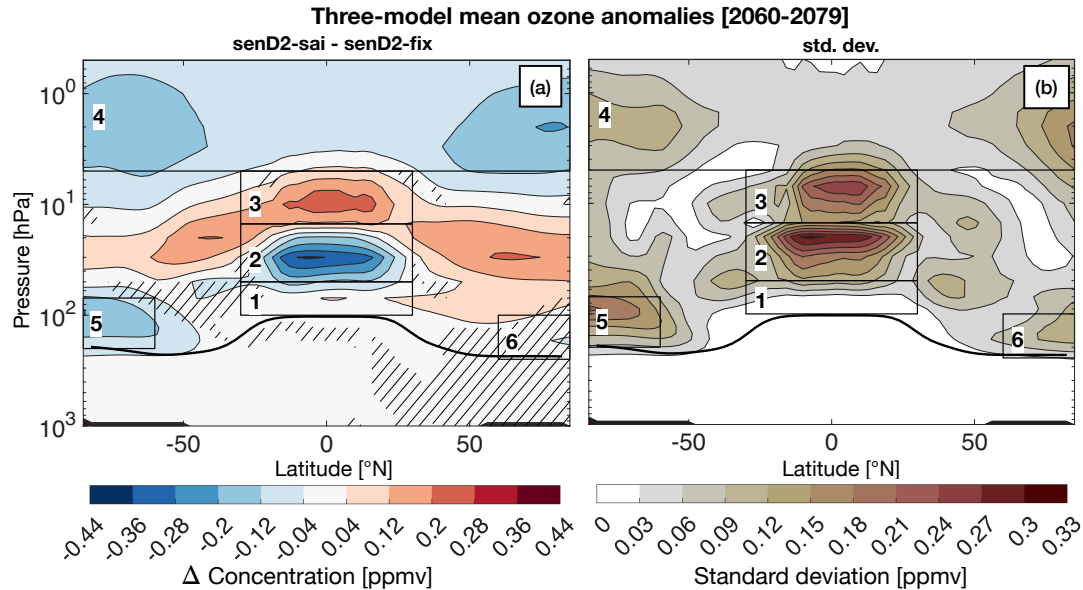


Figure C1. Three-model mean zonal mean ozone anomalies due to full SAI (senD2-sai - senD2-fix; panel a) and standard deviation in panel (b). Panel (a) is the multi-model mean of the three models that produced the chemistry-only simulations: SOCOL, CMAM, and NIES. These are the same models that are used to make the plots in panels (b) and (c) of Fig. 9. The black rectangles show regions with a dominant process shaping the ozone anomaly signal. The black line indicates the tropopause. Regions where the sign of all models do not agree are hatched.

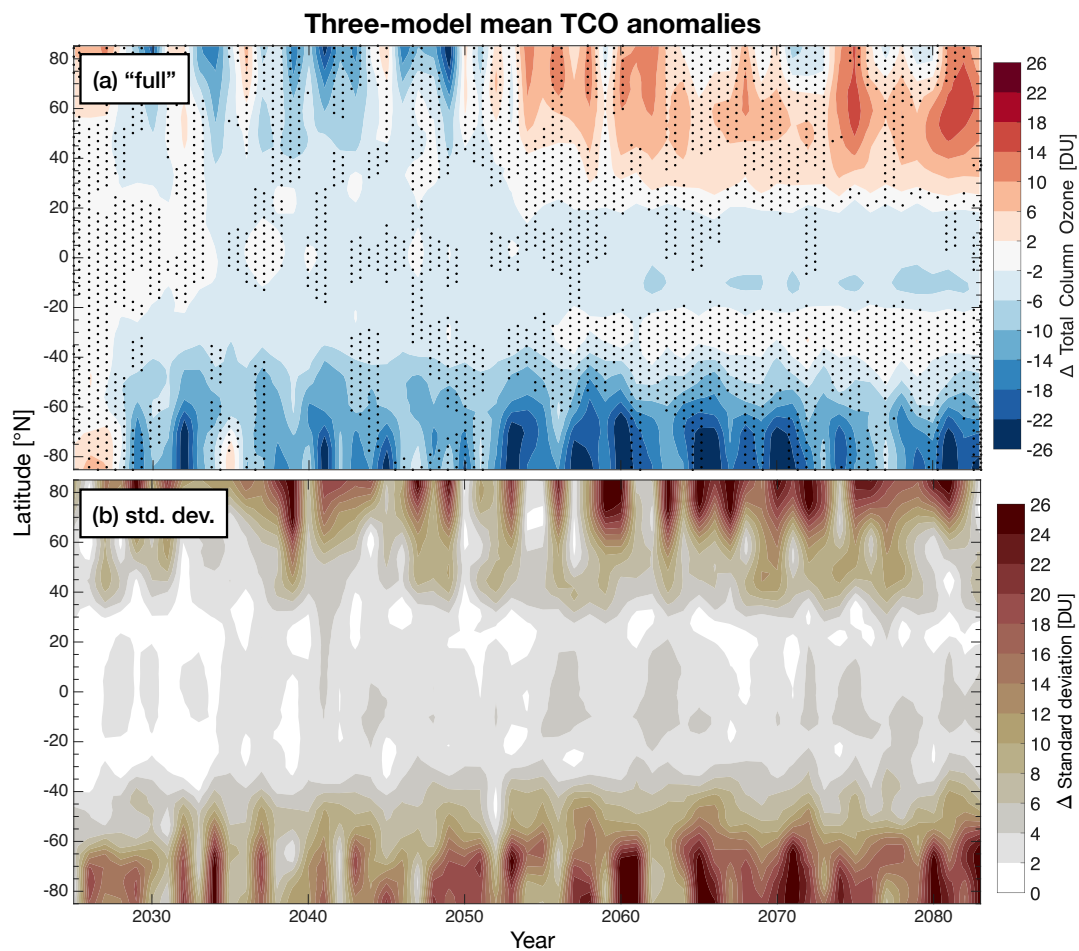


Figure C2. Three-model mean zonal mean total column ozone anomalies due to full SAI (senD2-sai - senD2-fix; panel a) and the standard deviation of full SAI in panel (b). Panel (a) is the multi-model mean of the three models that produced the chemistry-only simulations: SOCOL, CMAM, and NIES. These are the same models that are used to make the plots in panels (b) and (c) of Fig. 10. Regions where the sign of all models do not agree are stippled.

Appendix D: Heterogeneous chemical reactions

Table D1. Heterogeneous chemical reactions on aerosol surfaces included in the models. For each reaction the type of aerosol it is modeled on is listed. Note that CMAM does not include any NAT chemistry. The table is comprehensive

<u>Reaction</u>	<u>WACCM</u>	<u>SOCOL</u>	<u>CMAM</u>	<u>MIROC</u>	<u>NIES</u>
<u>$\text{N}_2\text{O}_5 + \text{H}_2\text{O} \rightarrow \text{HNO}_3$</u>	<u>all</u>	<u>all</u>	<u>STS¹,ICE³</u>	<u>all</u>	<u>all</u>
<u>$\text{N}_2\text{O}_5 + \text{HCl} \rightarrow \text{Cl} + \text{NO}_2 + \text{HNO}_3$</u>				<u>NAT²,ICE</u>	<u>NAT,ICE</u>
<u>$\text{N}_2\text{O}_5 + \text{HBr} \rightarrow \text{Br} + \text{NO}_2 + \text{HNO}_3$</u>				<u>NAT,ICE</u>	<u>NAT,ICE</u>
<u>$\text{ClONO}_2 + \text{H}_2\text{O} \rightarrow \text{HOCl} + \text{HNO}_3$</u>	<u>all</u>	<u>all</u>	<u>STS,ICE</u>	<u>all</u>	<u>all</u>
<u>$\text{ClONO}_2 + \text{HCl} \rightarrow 2\text{Cl} + \text{HNO}_3$</u>	<u>all</u>	<u>all</u>	<u>STS,ICE</u>	<u>all</u>	<u>all</u>
<u>$\text{ClONO}_2 + \text{HBr} \rightarrow \text{BrCl} + \text{HNO}_3$</u>				<u>NAT,ICE</u>	<u>NAT,ICE</u>
<u>$\text{BrONO}_2 + \text{H}_2\text{O} \rightarrow \text{HOBr} + \text{HNO}_3$</u>	<u>all</u>	<u>STS,ICE</u>	<u>STS</u>	<u>all</u>	<u>all</u>
<u>$\text{BrONO}_2 + \text{HCl} \rightarrow \text{BrCl} + \text{HNO}_3$</u>			<u>ICE</u>	<u>NAT,ICE</u>	<u>NAT,ICE</u>
<u>$\text{BrONO}_2 + \text{HBr} \rightarrow 2\text{Br} + \text{HNO}_3$</u>				<u>NAT,ICE</u>	<u>NAT,ICE</u>
<u>$\text{HOCl} + \text{HCl} \rightarrow 2\text{Cl} + \text{H}_2\text{O}$</u>	<u>all</u>	<u>all</u>	<u>STS,ICE</u>	<u>all</u>	<u>all</u>
<u>$\text{HOCl} + \text{HBr} \rightarrow \text{BrCl} + \text{H}_2\text{O}$</u>				<u>all</u>	<u>all</u>
<u>$\text{HOBr} + \text{HCl} \rightarrow \text{BrCl} + \text{H}_2\text{O}$</u>	<u>STS,ICE</u>	<u>STS,ICE</u>	<u>STS,ICE</u>	<u>all</u>	<u>all</u>
<u>$\text{HOBr} + \text{HBr} \rightarrow 2\text{Br} + \text{H}_2\text{O}$</u>				<u>all</u>	<u>all</u>

⁽¹⁾ supercooled ternary solution (liquid) aerosol
⁽²⁾ nitric acid trihydrate (liquid) aerosol
⁽³⁾ solid ice aerosol

Author contributions. AJ, TS, GC, SV, and TP designed the study. AJ, TS, ST, DP, SW, HA, and YY performed model simulations and curated the output data. AJ analyzed the data and compiled the manuscript with inputs from all other authors.

Competing interests. At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

665 *Acknowledgements.* Andrin Jörimann acknowledges the support from the Swiss National Science Foundation (SNSF) project AEON (grant
no. 200020E_219166). We acknowledge the use of the Scientific Colour Maps developed by Fabio Crameri (Crameri, 2023) and the Vi-
sually Distinct Colors Generator (<https://mokole.com/palette.html>) to ensure perceptually uniform and colorblind-friendly visualizations in
this work. Support for Sandro Vattioni was provided by the ETH Research grant no. ETH-1719-2 as well as by the Harvard Geoengi-
neering Research Program. Support for Gabriel Chiodo was provided by the Swiss Science Foundation within the Ambizione grant no.
670 PZ00P2_180043. Gabriel Chiodo, Sandro Vattioni, Timofei Sukhodolov and Simone Tilmes also received funding from the Simons Fun-
dation (grant no. SFI-MPS-SRM-00005217 and grant no. SFI-MPS-SRM-00005208 and SFI-MPS-SRM-00005203). Gabriel Chiodo addi-
tionally received funding from the European Commission within the ERC-StG project 101078127. Shingo Watanabe was supported by the
MEXT-Program for The Advanced Studies of Climate Change Projection (SENTAN) (grant number JPMXD0722681344), and MIROC-
ES2H simulations were conducted using the Earth Simulator at JAMSTEC. Hideharu Akiyoshi was supported by KAKENHI (JP24K00700,
675 JP24H00751 and JP25K07401) of the Ministry of Education, Culture, Sports, Science, and Technology, Japan. Yousuke Yamashita was
supported by KAKENHI (JP24K00700, JP24H00751, JP25K00377) and the Environmental Research and Technology Development Fund
(JPMEERF24S12201) of the Environmental Restoration and Conservation Agency, Japan. NEC SX-AURORA TSUBASA at NIES were
used to perform NIES model simulations. [Simone Tilmes acknowledges support from the NOAA Climate Program Office Earth's Radiation
Budget Awards Number 03-01-07-001 and NA22OAR4310477 and the CESM project, which is supported primarily by the National Science
680 Foundation. Computing and data storage resources, including the Cheyenne supercomputer \(doi:10.5065/D6RX99HX\), were provided by
the Computational and Information Systems Laboratory \(CISL\) at NCAR.](#)

References

- Akiyoshi, H.: Modeling of Chemistry and Chemistry-radiation Coupling Processes for the Middle Atmosphere and a Numerical Experiment on CO₂ Doubling with a 1-D Coupled Model, *Journal of the Meteorological Society of Japan. Ser. II*, 78, 563–584, https://doi.org/10.2151/jmsj1965.78.5_563, 2000.
- Allan, R. P., Hawkins, E., Bellouin, N., and Collins, B.: IPCC, 2021: Summary for Policymakers, *Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change*, <https://doi.org/10.1017/9781009157896.001>, 2021.
- An, M., Yao, B., Western, L. M., Prinn, R. G., Zhao, X., Hu, J., Mühle, J., Reimann, S., Vollmer, M. K., Harth, C. M., O’Doherty, S., Weiss, R. F., Chi, W., Xu, H., Yu, Y., Ganesan, A. L., and Rigby, M.: Persistent emissions of ozone-depleting carbon tetrachloride from China during 2011–2021, *Nature Geoscience*, 18, 593–598, <https://doi.org/10.1038/s41561-025-01721-4>, 2025.
- Banerjee, A., Butler, A. H., Polvani, L. M., Robock, A., Simpson, I. R., and Sun, L.: Robust winter warming over Eurasia under stratospheric sulfate geoengineering—the role of stratospheric dynamics, *Atmospheric Chemistry and Physics*, 21, 6985–6997, 2021.
- Bednarz, E. M., Butler, A. H., Visionsi, D., Zhang, Y., Kravitz, B., and MacMartin, D. G.: Injection strategy – a driver of atmospheric circulation and ozone response to stratospheric aerosol geoengineering, *Atmospheric Chemistry and Physics*, 23, 13 665–13 684, <https://doi.org/10.5194/acp-23-13665-2023>, 2023a.
- Bednarz, E. M., Visionsi, D., Kravitz, B., Jones, A., Haywood, J. M., Richter, J., MacMartin, D. G., and Braesicke, P.: Climate response to off-equatorial stratospheric sulfur injections in three Earth system models – Part 2: Stratospheric and free-tropospheric response, *Atmospheric Chemistry and Physics*, 23, 687–709, <https://doi.org/10.5194/acp-23-687-2023>, 2023b.
- Benito-Barca, S., Abalos, M., Calvo, N., Garny, H., Birner, T., Abraham, N. L., Akiyoshi, H., Dennison, F., Jöckel, P., Josse, B., Keeble, J., Kinnison, D., Marchand, M., Morgenstern, O., Plummer, D., Rozanov, E., Strode, S., Sukhodolov, T., Watanabe, S., and Yamashita, Y.: Recent Lower Stratospheric Ozone Trends in CCMI-2022 Models: Role of Natural Variability and Transport, *Journal of Geophysical Research: Atmospheres*, 130, <https://doi.org/10.1029/2024JD042412>, 2025.
- Bourguet, S., Stone, K., and Lickley, M.: Semi-empirical estimates of stratospheric circulation and the lifetimes of chlorofluorocarbons and carbon tetrachloride, *Communications Earth & Environment*, 6, 531, <https://doi.org/10.1038/s43247-025-02500-0>, 2025.
- Budyko, M. I.: On present-day climatic changes, *Tellus*, 29, 193–204, 1977.
- CCMI: IGAC/SPARC Chemistry-Climate Model Initiative, <https://blogs.reading.ac.uk/ccmi/ccmi-2022>, last accessed: 02.01.2026, 2022.
- Center for Environmental Data Analysis (CEDA): CCMI-2022 Monthly Chemistry-Climate Model Data, <https://catalogue.ceda.ac.uk/uuid/92dddf542adc44b5898f535be4179705>, last accessed: 22.01.2026, 2025.
- Cramer, F.: Scientific colour maps, <https://doi.org/10.5281/zenodo.8409685>, 2023.
- Crutzen, P. J.: Albedo enhancement by stratospheric sulfur injections: a contribution to resolve a policy dilemma?, *Climatic change*, 77, 211, <https://doi.org/10.1007/s10584-006-9101-y>, 2006.
- Danabasoglu, G., Lamarque, J., Bacmeister, J., Bailey, D. A., DuVivier, A. K., Edwards, J., Emmons, L. K., Fasullo, J., Garcia, R., Gettelman, A., Hannay, C., Holland, M. M., Large, W. G., Lauritzen, P. H., Lawrence, D. M., Lenaerts, J. T. M., Lindsay, K., Lipscomb, W. H., Mills, M. J., Neale, R., Oleson, K. W., Otto-Bliesner, B., Phillips, A. S., Sacks, W., Tilmes, S., van Kampenhout, L., Vertenstein, M., Bertini, A., Dennis, J., Deser, C., Fischer, C., Fox-Kemper, B., Kay, J. E., Kinnison, D., Kushner, P. J., Larson, V. E., Long, M. C., Mickelson, S., Moore, J. K., Nienhouse, E., Polvani, L., Rasch, P. J., and Strand, W. G.: The Community Earth System Model Version 2 (CESM2), *Journal of Advances in Modeling Earth Systems*, 12, <https://doi.org/10.1029/2019MS001916>, 2020.

- Dhomse, S. S., Kinnison, D., Chipperfield, M. P., Salawitch, R. J., Cionni, I., Hegglin, M. I., Abraham, N. L., Akiyoshi, H., Archibald, A. T.,
720 Bednarz, E. M., Bekki, S., Braesicke, P., Butchart, N., Dameris, M., Deushi, M., Frith, S., Hardiman, S. C., Hassler, B., Horowitz, L. W.,
Hu, R.-M., Jöckel, P., Josse, B., Kirner, O., Kremser, S., Langematz, U., Lewis, J., Marchand, M., Lin, M., Mancini, E., Marécal, V.,
Michou, M., Morgenstern, O., O'Connor, F. M., Oman, L., Pitari, G., Plummer, D. A., Pyle, J. A., Revell, L. E., Rozanov, E., Schofield,
R., Stenke, A., Stone, K., Sudo, K., Tilmes, S., Visionsi, D., Yamashita, Y., and Zeng, G.: Estimates of ozone return dates from Chemistry-
Climate Model Initiative simulations, *Atmospheric Chemistry and Physics*, 18, 8409–8438, <https://doi.org/10.5194/acp-18-8409-2018>,
725 2018.
- Egorova, T., Rozanov, E., Zubov, V., and Karol, I. L.: Model for Investigating Ozone Trends (MEZON), *Izvestiya, Atmos. Oceanic Phys.*,
39, 277–292, 2003.
- Emmons, L. K., Schwantes, R. H., Orlando, J. J., Tyndall, G., Kinnison, D., Lamarque, J., Marsh, D., Mills, M. J., Tilmes, S., Bardeen,
C., Buchholz, R. R., Conley, A., Gettelman, A., Garcia, R., Simpson, I., Blake, D. R., Meinardi, S., and Pétron, G.: The Chem-
730 istry Mechanism in the Community Earth System Model Version 2 (CESM2), *Journal of Advances in Modeling Earth Systems*, 12,
<https://doi.org/10.1029/2019MS001882>, 2020.
- Forster, P. M., Smith, C., Walsh, T., Lamb, W. F., Lamboll, R., Hall, B., Hauser, M., Ribes, A., Rosen, D., Gillett, N. P., Palmer, M. D., Rogelj,
J., von Schuckmann, K., Trewin, B., Allen, M., Andrew, R., Betts, R. A., Borger, A., Boyer, T., Broersma, J. A., Buontempo, C., Burgess,
S., Cagnazzo, C., Cheng, L., Friedlingstein, P., Gettelman, A., Gütschow, J., Ishii, M., Jenkins, S., Lan, X., Morice, C., Mühle, J., Kadow,
735 C., Kennedy, J., Killick, R. E., Krummel, P. B., Minx, J. C., Myhre, G., Naik, V., Peters, G. P., Pirani, A., Pongratz, J., Schleussner, C.-F.,
Seneviratne, S. I., Szopa, S., Thorne, P., Kovilakam, M. V. M., Majamäki, E., Jalkanen, J.-P., van Marle, M., Hoesly, R. M., Rohde, R.,
Schumacher, D., van der Werf, G., Vose, R., Zickfeld, K., Zhang, X., Masson-Delmotte, V., and Zhai, P.: Indicators of Global Climate
Change 2023: annual update of key indicators of the state of the climate system and human influence, *Earth System Science Data*, 16,
2625–2658, <https://doi.org/10.5194/essd-16-2625-2024>, 2024.
- 740 Fouquart, Y. and Bonnel, B.: Computations of solar heating of the Earth's atmosphere: A new parameterization, *Beitr. Phys. Atmos.*, 53,
35–62, 1980.
- Franke, H., Niemeier, U., and Visionsi, D.: Differences in the quasi-biennial oscillation response to stratospheric aerosol modification depend-
ing on injection strategy and species, *Atmospheric Chemistry and Physics*, 21, 8615–8635, <https://doi.org/10.5194/acp-21-8615-2021>,
2021.
- 745 Friedel, M., Chiodo, G., Sukhodolov, T., Keeble, J., Peter, T., Seeber, S., Stenke, A., Akiyoshi, H., Rozanov, E., Plummer, D., Jöckel, P.,
Zeng, G., Morgenstern, O., and Josse, B.: Weakening of springtime Arctic ozone depletion with climate change, *Atmospheric Chemistry
and Physics*, 23, 10 235–10 254, <https://doi.org/10.5194/acp-23-10235-2023>, 2023.
- Gettelman, A., Mills, M. J., Kinnison, D. E., Garcia, R. R., Smith, A. K., Marsh, D. R., Tilmes, S., Vitt, F., Bardeen, C. G., McInerney, J., Liu,
H.-L., Solomon, S. C., Polvani, L. M., Emmons, L. K., Lamarque, J.-F., Richter, J. H., Glanville, A. S., Bacmeister, J. T., Phillips, A. S.,
750 Neale, R. B., Simpson, I. R., DuVivier, A. K., Hodzic, A., and Randel, W. J.: The Whole Atmosphere Community Climate Model Version 6
(WACCM6), *Journal of Geophysical Research: Atmospheres*, 124, 12 380–12 403, <https://doi.org/10.1029/2019JD030943>,
2019.
- Heckendorn, P., Weisenstein, D., Fueglistaler, S., Luo, B. P., Rozanov, E., Schraner, M., Thomason, L. W., and Peter, T.: The impact of geo-
engineering aerosols on stratospheric temperature and ozone, *Environmental Research Letters*, 4, 045 108, [https://doi.org/10.1088/1748-
9326/4/4/045108](https://doi.org/10.1088/1748-

755 9326/4/4/045108), 2009.
- Hegerl, G. C. and Solomon, S.: Risks of Climate Engineering, *Science*, 325, 955–956, <https://doi.org/10.1126/science.1178530>, 2009.

- Jones, A., Haywood, J. M., Scaife, A. A., Boucher, O., Henry, M., Kravitz, B., Lurton, T., Nabat, P., Niemeier, U., Séférian, R., Tilmes, S., and Vioni, D.: The impact of stratospheric aerosol intervention on the North Atlantic and Quasi-Biennial Oscillations in the Geoengineering Model Intercomparison Project (GeoMIP) G6sulfur experiment, *Atmospheric Chemistry and Physics*, 22, 2999–3016, <https://doi.org/10.5194/acp-22-2999-2022>, 2022.
- Jonsson, A. I., deGrandpre, J., Fomichev, V. I., McConnell, J. C., and Beagley, S. R.: Doubled CO₂-induced cooling in the middle atmosphere: Photochemical analysis of the ozone radiative feedback, *Journal of Geophysical Research: Atmospheres*, 109, <https://doi.org/https://doi.org/10.1029/2004JD00509>, 2004.
- Jungclaus, J. H., Fischer, N., Haak, H., Lohmann, K., Marotzke, J., Matei, D., Mikolajewicz, U., Notz, D., and von Storch, J. S.: Characteristics of the ocean simulations in the Max Planck Institute Ocean Model (MPIOM) the ocean component of the MPI-Earth system model, *Journal of Advances in Modeling Earth Systems*, 5, 422–446, <https://doi.org/10.1002/jame.20023>, 2013.
- Jörmann, A.: REMAP-CCMI-2022-sai: Stratospheric aerosol data for use in the CCMI-2022 stratospheric aerosol injection scenario, ETH Research Collection [data set], <https://doi.org/10.3929/ethz-b-000714654>, previously distributed through ftp://iacftp.ethz.ch/pub_read/andrinj/CCMI-2022-sai (last accessed 02.01.2026), 2023.
- Jörmann, A.: CCMI-2022 post-processed model data [data set], <https://doi.org/10.5281/zenodo.18331210>, 2026.
- Jörmann, A., Sukhodolov, T., Luo, B., Chiodo, G., Mann, G., and Peter, T.: RETrieval Method for optical and physical Aerosol Properties in the stratosphere (REMAPv1), *Geoscientific Model Development*, 18, 6023–6041, <https://doi.org/10.5194/gmd-18-6023-2025>, 2025.
- K1 Model Developers: K-1 coupled GCM (MIROC) description (K-1 Technical Report, 34 pp.), Tech. rep., University of Tokyo; National Institute for Environmental Studies (NIES); Frontier Research Center for Global Change (FRCGC), https://ccsr.aori.u-tokyo.ac.jp/~hasumi/miroc_description.pdf, 2004.
- Kawamiya, M., Hajima, T., Tachiiri, K., Watanabe, S., and Yokohata, T.: Two decades of Earth system modeling with an emphasis on Model for Interdisciplinary Research on Climate (MIROC), *Progress in Earth and Planetary Science*, 7, <https://doi.org/10.1186/s40645-020-00369-5>, 2020.
- Lawrence, D. M., Fisher, R. A., Koven, C. D., Oleson, K. W., Swenson, S. C., Bonan, G., Collier, N., Ghimire, B., van Kampenhout, L., Kennedy, D., Kluzek, E., Lawrence, P. J., Li, F., Li, H., Lombardozzi, D., Riley, W. J., Sacks, W. J., Shi, M., Vertenstein, M., Wieder, W. R., Xu, C., Ali, A. A., Badger, A. M., Bisht, G., van den Broeke, M., Brunke, M. A., Burns, S. P., Buzan, J., Clark, M., Craig, A., Dahlin, K., Drewniak, B., Fisher, J. B., Flanner, M., Fox, A. M., Gentine, P., Hoffman, F., Keppel-Aleks, G., Knox, R., Kumar, S., Lenaerts, J., Leung, L. R., Lipscomb, W. H., Lu, Y., Pandey, A., Pelletier, J. D., Perket, J., Randerson, J. T., Ricciuto, D. M., Sanderson, B. M., Slater, A., Subin, Z. M., Tang, J., Thomas, R. Q., Martin, M. V., and Zeng, X.: The Community Land Model Version 5: Description of New Features, Benchmarking, and Impact of Forcing Uncertainty, *Journal of Advances in Modeling Earth Systems*, 11, 4245–4287, <https://doi.org/10.1029/2018MS001583>, 2019.
- Lee, W. R., Vioni, D., Bednarz, E. M., MacMartin, D. G., Kravitz, B., and Tilmes, S.: Quantifying the Efficiency of Stratospheric Aerosol Geoengineering at Different Altitudes, *Geophysical Research Letters*, 50, <https://doi.org/10.1029/2023GL104417>, 2023.
- Lee, W. R., Tilmes, S., and Bednarz, E. M.: Exploring divergent long-term stratospheric aerosol injection scenarios with the G2-SAI and ARISE-hybrid experiments, <https://doi.org/10.5194/egusphere-2026-1004>, 2026.
- Lin, S.-J. and Rood, R. B.: Multidimensional Flux-Form Semi-Lagrangian Transport Schemes, *Monthly Weather Review*, 124, 2046–2070, [https://doi.org/10.1175/1520-0493\(1996\)124<2046:MFFSLT>2.0.CO;2](https://doi.org/10.1175/1520-0493(1996)124<2046:MFFSLT>2.0.CO;2), 1996.
- Liu, P. R. and Raftery, A. E.: Country-based rate of emissions reductions should increase by 80% beyond nationally determined contributions to meet the 2°C target, *Communications Earth & Environment*, 2, 29, <https://doi.org/10.1038/s43247-021-00097-8>, 2021.

- 795 Liu, X., Ma, P.-L., Wang, H., Tilmes, S., Singh, B., Easter, R. C., Ghan, S. J., and Rasch, P. J.: Description and evaluation of a new four-mode version of the Modal Aerosol Module (MAM4) within version 5.3 of the Community Atmosphere Model, *Geoscientific Model Development*, 9, 505–522, <https://doi.org/10.5194/gmd-9-505-2016>, 2016.
- MacMartin, D. G., Kravitz, B., Tilmes, S., Richter, J. H., Mills, M. J., Lamarque, J., Tribbia, J. J., and Vitt, F.: The Climate Response to Stratospheric Aerosol Geoengineering Can Be Tailored Using Multiple Injection Locations, *Journal of Geophysical Research: Atmospheres*, 122, <https://doi.org/10.1002/2017JD026868>, 2017.
- 800 Madronich, S. and Flocke, S.: The role of solar radiation in atmospheric chemistry, pp. 1–26, Springer Verlag (edited by P. Boule), 1998.
- Marshall, L. R., Maters, E. C., Schmidt, A., Timmreck, C., Robock, A., and Toohey, M.: Volcanic effects on climate: recent advances and future avenues, *Bulletin of Volcanology*, 84, <https://doi.org/10.1007/s00445-022-01559-3>, 2022.
- McKay, D. I. A., Staal, A., Abrams, J. F., Winkelman, R., Sakschewski, B., Loriani, S., Fetzer, I., Cornell, S. E., Rockström, J., and Lenton, T. M.: Exceeding 1.5°C global warming could trigger multiple climate tipping points, *Science*, 377, <https://doi.org/10.1126/science.abn7950>, 2022.
- 805 Mills, M. J., Richter, J. H., Tilmes, S., Kravitz, B., MacMartin, D. G., Glanville, A. A., Tribbia, J. J., Lamarque, J., Vitt, F., Schmidt, A., Gettelman, A., Hannay, C., Bacmeister, J. T., and Kinnison, D. E.: Radiative and Chemical Response to Interactive Stratospheric Sulfate Aerosols in Fully Coupled CESM1(WACCM), *Journal of Geophysical Research: Atmospheres*, 122, <https://doi.org/10.1002/2017JD027006>, 2017.
- 810 Morcrette, J. J.: Description of the radiation Scheme in the ECMWF model, Technical Memorandum 165, European Centre for Medium Range Weather Forecasting, <https://www.ecmwf.int/en/elibrary/75744-description-radiation-scheme-ecmwf-model>, 1989.
- Möller, T., Högnér, A. E., Schleussner, C.-F., Bien, S., Kitzmann, N. H., Lamboll, R. D., Rogelj, J., Donges, J. F., Rockström, J., and Wunderling, N.: Achieving net zero greenhouse gas emissions critical to limit climate tipping risks, *Nature Communications*, 15, 6192, <https://doi.org/10.1038/s41467-024-49863-0>, 2024.
- 815 Perny, K., Sukhodolov, T., Kuchar, A., Arsenovic, P., Rosati, B., Brühl, C., Dhomse, S. S., Jörimann, A., Laakso, A., Mann, G., Niemeier, U., Pitari, G., Quaglia, I., Sekiya, T., Sudo, K., Timmreck, C., Tilmes, S., Visionsi, D., and Rieder, H. E.: Assessing the stratospheric temperature response to volcanic sulfate injections by Mt. Pinatubo: insights from the Interactive Stratospheric Aerosol Model Intercomparison Project [preprint], <https://doi.org/10.5194/egusphere-2025-5915>, 2025.
- Pitari, G., Aquila, V., Kravitz, B., Robock, A., Watanabe, S., Cionni, I., Luca, N. D., Genova, G. D., Mancini, E., and Tilmes, S.: Stratospheric ozone response to sulfate geoengineering: Results from the Geoengineering Model Intercomparison Project (GeoMIP), *Journal of Geophysical Research: Atmospheres*, 119, 2629–2653, <https://doi.org/10.1002/2013JD020566>, 2014.
- 820 Plummer, D., Nagashima, T., Tilmes, S., Archibald, A., Chiodo, G., Fadnavis, S., Garny, H., Josse, B., Kim, J., Lamarque, J.-F., et al.: CCMI-2022: A new set of Chemistry-Climate Model Initiative (CCMI) community simulations to update the assessment of models and support upcoming ozone assessment activities, *SPARC Newsletter*, 57, 22–30, https://aparc-climate.org/wp-content/uploads/2025/10/SPARCnewsletter_Jul2021_web.pdf, 2021.
- 825 Pörtner, H. O., Roberts, D. C., Adams, H., Adler, C., Aldunce, P., Ali, E., Begum, R. A., Betts, R., Kerr, R. B., and Biesbroek, R.: Climate change 2022: impacts, adaptation and vulnerability, Contribution of Working Group II to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change, 2022.
- Richter, J. H., Tilmes, S., Glanville, A., Kravitz, B., MacMartin, D. G., Mills, M. J., Simpson, I. R., Vitt, F., Tribbia, J. J., and Lamarque, J.: Stratospheric Response in the First Geoengineering Simulation Meeting Multiple Surface Climate Objectives, *Journal of Geophysical Research: Atmospheres*, 123, 5762–5782, <https://doi.org/10.1029/2018JD028285>, 2018.
- 830 Robock, A.: Volcanic eruptions and climate, *Reviews of geophysics*, 38, 191–219, 2000.

- Rozanov, E. V., Zubov, V. A., Schlesinger, M. E., Yang, F., and Andronova, N. G.: The UIUC three-dimensional stratospheric chemical transport model: Description and evaluation of the simulated source gases and ozone, *J. Geophys. Res.*, 104, 11 755–11 781, <https://doi.org/10.1029/1999JD900138>, 1999.
- Scinocca, J. F., McFarlane, N. A., Lazare, M., Li, J., and Plummer, D.: Technical Note: The CCCma third generation AGCM and its extension into the middle atmosphere, *Atmospheric Chemistry and Physics*, 8, 7055–7074, <https://doi.org/10.5194/acp-8-7055-2008>, 2008.
- Sekiguchi, M. and Nakajima, T.: A k-distribution-based radiation code and its computational optimization for an atmospheric general circulation model, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 109, 2779–2793, <https://doi.org/10.1016/j.jqsrt.2008.07.013>, 2008.
- Sekiya, T., Sudo, K., and Nagai, T.: Evolution of stratospheric sulfate aerosol from the 1991 Pinatubo eruption: Roles of aerosol microphysical processes, *Journal of Geophysical Research: Atmospheres*, 121, 2911–2938, <https://doi.org/10.1002/2015JD024313>, 2016.
- Sessler, J., Good, P., MacKenzie, A. R., and Pyle, J. A.: What role do type I polar stratospheric cloud and aerosol parameterizations play in modelled lower stratospheric chlorine activation and ozone loss?, *Journal of Geophysical Research: Atmospheres*, 101, 28 817–28 835, <https://doi.org/10.1029/96JD02546>, 1996.
- Sheng, J., Weisenstein, D. K., Luo, B., Rozanov, E., Stenke, A., Anet, J., Bingemer, H., and Peter, T.: Global atmospheric sulfur budget under volcanically quiescent conditions: Aerosol-chemistry-climate model predictions and validation, *J. Geophys. Res.- Atmos.*, 120, 256–276, <https://doi.org/10.1002/2014JD021985>, 2015.
- Stevens, B., Giorgetta, M., Esch, M., Mauritsen, T., Crueger, T., Rast, S., Salzmann, M., Schmidt, H., Bader, J., Block, K., Brokopf, R., Fast, I., Kinne, S., Kornblueh, L., Lohmann, U., Pincus, R., Reichler, T., and Roeckner, E.: Atmospheric component of the MPI-M Earth System Model: ECHAM6, *Journal of Advances in Modeling Earth Systems*, 5, 146–172, <https://doi.org/10.1002/jame.20015>, 2013.
- Sudo, K., Takahashi, M., Kurokawa, J.-i., and Akimoto, H.: CHASER: A global chemical model of the troposphere 1. Model description, *Journal of Geophysical Research: Atmospheres*, 107, ACH 7–1–ACH 7–20, <https://doi.org/10.1029/2001JD001113>, 2002.
- Sukhodolov, T., Egorova, T., Stenke, A., Ball, W. T., Brodowsky, C., Chiodo, G., Feinberg, A., Friedel, M., Karagodin-Doyennel, A., and Peter, T.: Atmosphere–ocean–aerosol–chemistry–climate model SOCOLv4. 0: description and evaluation, *Geoscientific Model Development*, 14, 5525–5560, 2021.
- Swart, N. C., Cole, J. N. S., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., Anstey, J., Arora, V., Christian, J. R., Hanna, S., Jiao, Y., Lee, W. G., Majaess, F., Saenko, O. A., Seiler, C., Seinen, C., Shao, A., Sigmond, M., Solheim, L., von Salzen, K., Yang, D., and Winter, B.: The Canadian Earth System Model version 5 (CanESM5.0.3), *Geoscientific Model Development*, 12, 4823–4873, <https://doi.org/10.5194/gmd-12-4823-2019>, 2019.
- Tatebe, H., Ogura, T., Nitta, T., Komuro, Y., Ogochi, K., Takemura, T., Sudo, K., Sekiguchi, M., Abe, M., Saito, F., Chikira, M., Watanabe, S., Mori, M., Hirota, N., Kawatani, Y., Mochizuki, T., Yoshimura, K., Takata, K., O’ishi, R., Yamazaki, D., Suzuki, T., Kurogi, M., Kataoka, T., Watanabe, M., and Kimoto, M.: Description and basic evaluation of simulated mean state, internal variability, and climate sensitivity in MIROC6, *Geoscientific Model Development*, 12, 2727–2765, <https://doi.org/10.5194/gmd-12-2727-2019>, 2019.
- Tilmes, S., Richter, J. H., Kravitz, B., MacMartin, D. G., Mills, M. J., Simpson, I. R., Glanville, A. S., Fasullo, J. T., Phillips, A. S., and Lamarque, J.-F.: CESM1 (WACCM) stratospheric aerosol geoengineering large ensemble project, *Bulletin of the American Meteorological Society*, 99, 2361–2371, 2018.

- Tilmes, S., Richter, J. H., Kravitz, B., MacMartin, D. G., Glanville, A. S., Vioni, D., Kinnison, D. E., and Müller, R.: Sensitivity of Total Column Ozone to Stratospheric Sulfur Injection Strategies, *Geophysical Research Letters*, 48, <https://doi.org/10.1029/2021GL094058>, 2021.
- Tilmes, S., Vioni, D., Jones, A., Haywood, J., Séférian, R., Nabat, P., Boucher, O., Bednarz, E. M., and Niemeier, U.: Stratospheric ozone response to sulfate aerosol and solar dimming climate interventions based on the G6 Geoengineering Model Intercomparison Project (GeoMIP) simulations, *Atmospheric Chemistry and Physics*, 22, 4557–4579, <https://doi.org/10.5194/acp-22-4557-2022>, 2022.
- 875 Tilmes, S., Bednarz, E. M., Jörimann, A., Vioni, D., Kinnison, D. E., Chiodo, G., and Plummer, D.: Stratospheric Aerosol Intervention experiment for the Chemistry–Climate Model Initiative, *Atmospheric Chemistry and Physics*, 25, 6001–6023, <https://doi.org/10.5194/acp-25-6001-2025>, 2025a.
- Tilmes, S., Vioni, D., Quaglia, I., Zhu, Y., Bardeen, C. G., Vitt, F., and Yu, P.: Uncertainties of SAI efficiency and impacts depending on the complexity of the aerosol microphysical model, <https://doi.org/10.5194/egusphere-2025-4274>, 2025b.
- 880 Tollefson, J.: IPCC says limiting global warming to 1.5 [degrees] C will require drastic action, *Nature*, 562, 172–174, 2018.
- UNEP: United Nations Environment Program, Decision XXXI/2: Potential areas of focus for the 2022 quadrennial reports of the Scientific Assessment Panel, the Environmental Effects Assessment Panel and the Technology and Economic Assessment Panel, <https://ozone.unep.org/treaties/montreal-protocol/meetings/thirty-first-meeting-parties/decisions/decision-xxxi2-potential-areas-focus-2022-quadrennial-reports-scientific-assessment-panel>, last accessed: 02.01.2026, 2019.
- 885 Vioni, D., MacMartin, D. G., Kravitz, B., Lee, W., Simpson, I. R., and Richter, J. H.: Reduced Poleward Transport Due to Stratospheric Heating Under Stratospheric Aerosols Geoengineering, *Geophysical Research Letters*, 47, <https://doi.org/10.1029/2020GL089470>, 2020.
- Vioni, D., Bednarz, E. M., Lee, W. R., Kravitz, B., Jones, A., Haywood, J. M., and MacMartin, D. G.: Climate response to off-equatorial stratospheric sulfur injections in three Earth system models – Part 1: Experimental protocols and surface changes, *Atmospheric Chemistry and Physics*, 23, 663–685, <https://doi.org/10.5194/acp-23-663-2023>, 2023.
- 890 Vioni, D., Robock, A., Haywood, J., Henry, M., Tilmes, S., MacMartin, D. G., Kravitz, B., Doherty, S. J., Moore, J., Lennard, C., Watanabe, S., Muri, H., Niemeier, U., Boucher, O., Syed, A., Egbebiyi, T. S., Séférian, R., and Quaglia, I.: G6-1.5K-SAI: a new Geoengineering Model Intercomparison Project (GeoMIP) experiment integrating recent advances in solar radiation modification studies, *Geoscientific Model Development*, 17, 2583–2596, <https://doi.org/10.5194/gmd-17-2583-2024>, 2024.
- Watanabe, S., Hajima, T., Sudo, K., Nagashima, T., Takemura, T., Okajima, H., Nozawa, T., Kawase, H., Abe, M., Yokohata, T., Ise, T., Sato, H., Kato, E., Takata, K., Emori, S., and Kawamiya, M.: MIROC-ESM 2010: model description and basic results of CMIP5-20c3m experiments, *Geoscientific Model Development*, 4, 845–872, <https://doi.org/10.5194/gmd-4-845-2011>, 2011.
- 895 Weisenstein, D. K., Penner, J. E., Herzog, M., and Liu, X.: Global 2-D intercomparison of sectional and modal aerosol modules, *Atmos. Chem. Phys.*, 7, 2339–2355, <https://doi.org/10.5194/acp-7-2339-2007>, 2007.
- Weisenstein, D. K., Vioni, D., Franke, H., Niemeier, U., Vattioni, S., Chiodo, G., Peter, T., and Keith, D. W.: An interactive stratospheric aerosol model intercomparison of solar geoengineering by stratospheric injection of SO₂ or accumulation-mode sulfuric acid aerosols, *Atmospheric Chemistry and Physics*, 22, 2955–2973, 2022.
- 900 Western, L. M., Vollmer, M. K., Krummel, P. B., Adcock, K. E., Crotwell, M., Fraser, P. J., Harth, C. M., Langenfelds, R. L., Montzka, S. A., Mühle, J., O’Doherty, S., Oram, D. E., Reimann, S., Rigby, M., Vimont, I., Weiss, R. F., Young, D., and Laube, J. C.: Global increase of ozone-depleting chlorofluorocarbons from 2010 to 2020, *Nature Geoscience*, 16, 309–313, <https://doi.org/10.1038/s41561-023-01147-w>, 2023.
- 905

World Meteorological Organization: Scientific Assessment of Ozone Depletion: 2018, Global Ozone Research and Monitoring Project–Report No. 58, 588 pp., ISBN 9781732931718, <https://csl.noaa.gov/assessments/ozone/2018/>, 2019.

World Meteorological Organization: Scientific Assessment of Ozone Depletion: 2022, GAW Report No. 278, Geneva, Switzerland, ISBN 978-9914-733-99-0, <https://csl.noaa.gov/assessments/ozone/2022/>, 2022.

- 910 Wunderlin, E., Chiodo, G., Sukhodolov, T., Vattioni, S., Visioni, D., and Tilmes, S.: Side Effects of Sulfur-Based Geoengineering Due To Absorptivity of Sulfate Aerosols, *Geophysical Research Letters*, 51, <https://doi.org/10.1029/2023GL107285>, 2024.