



1 Glycan purification reveals persistence of algal fucoidan in 2 millennia-old marine sediments

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10 **Abstract:**

11 Marine sediments are an important sink for the carbon fixed by photosynthetic algae in the sunlit ocean. Glycans
12 represent a substantial fraction to the organic carbon buried globally in marine sediments. However, assigning the
13 origin and contribution of specific glycans remains challenging with commonly used methods that are designed to
14 measure the total glycan carbon content. Glycan-specific detection using monoclonal antibodies previously
15 revealed diatom- and brown algae-derived fucoidans in marine sediments of various ages. These complex anionic
16 glycans may contribute to long-term carbon sequestration, while their quantitative significance remains unknown.
17 Here, we combined anion exchange chromatography with glycan-specific monoclonal antibodies to isolate and
18 detect fucoidans from marine sediments. Applying this method to a sediment core from the Bransfield Strait in the
19 Southern Ocean revealed that fucoidan was buried and persisted for the extent of the Holocene (11.8 ka). Our
20 results highlight that glycan purification from sediments can help to resolve the importance of different algal
21 glycans to the sequestration of carbon in the ocean.

22 **1 Introduction:**

23 Marine sediment is the largest long-term reservoir of organic carbon, sequestering carbon over geological
24 timescales (Burdige, 2007; Li et al., 2023). Processes like the biological carbon pump transport organic carbon
25 fixed in the sunlit ocean to the sediments in form of aggregates, faecal pellets, and other particles (Eppley and
26 Peterson, 1979; Cavan et al., 2015; Le Moigne, 2019). The particles reaching the seafloor are compositionally
27 diverse, originating from multiple sources and undergoing various transformations in the water column (Duarte
28 and Cebrián, 1996; Aluwihare et al., 2002; Dittmar et al., 2021; Hedges et al., 2001; Chin et al., 1998; Engel et al.,
29 2004). From this mixture of organic molecules, specific biomolecules have been targeted to reconstruct past
30 climate, ecological dynamics, and carbon cycling processes (Lee et al., 2016; Salmeán et al., 2022; Vorrath et al.,
31 2023a).

32 Glycans (polymeric carbohydrates) of unidentified origin contribute 5-30% of organic carbon in marine sediments
33 (Burdige, 2007; Miyajima et al., 2001; Vidal-Melgosa et al., 2022). Previously used methods were designed to
34 answer questions about total glycan carbon concentrations and bulk monosaccharide compositions (Burdige et al.,
35 2000; Taylor et al., 1999; Vidal-Melgosa et al., 2022). While certain glycans are enriched in unique monomers
36 specific to certain algal phyla, most monosaccharides are shared across different phyla as building blocks for
37 diverse glycans (Arnosti et al., 2021). Therefore, quantifying monosaccharides or ratios of monosaccharides as



attempted in earlier studies (Cowie and Hedges, 1984; Klok et al., 1984; Miyajima et al., 2001), may not resolve for specific glycan types or reveal the source of the glycans. To gain more molecular resolution, recent studies have used structure-specific monoclonal antibody-based approaches (Hellige et al., 2026; Salmeán et al., 2022; Vidal-Melgosa et al., 2022). These studies semi-quantitatively revealed that brown algal- and diatom-derived fucoidans persist in sediment for various periods of time, indicating the potential roles of these complex anionic glycans towards long-term carbon sequestration. However, the quantification of these glycans is necessary to understand their exact contributions to the various components of the biological carbon pump and sedimentary organic matter.

Other than the limitations of the available analytical methods, glycan quantification in the sediments is further complicated by the presence of interfering salts, humic compounds and other contaminants. Alginate, an anionic brown algal glycan, was recently quantified in marine sediments using alginate lyases and other colorimetric assays (Nakazato et al., 2026). This was only possible after a multi-step extraction and clean-up protocol that involved the removal of phenolic compounds, other glycans like pectin, and other pigments. The quantification of the purified alginate highlights the potential of brown algal glycans in carbon sequestration (Nakazato et al., 2026). Similarly, coccolith-associated polysaccharides were extracted and isolated from coccolith fossils in the sediment using anion exchange chromatography (Lee et al., 2016). The purification step enabled quantification of uronic acid content in the specific polysaccharides. It was observed that the uronic acid content in the coccolith-associated polysaccharides correlates with the internal saturation state of the coccolith vesicle indicative of past environmental pCO_2 (Lee et al., 2016). These studies highlight that the extraction and purification of specific glycans from complex mixtures enable their quantification and use as proxies for carbon sequestration and/or environmental conditions involved in their synthesis and deposition. We considered that similar approaches may apply to fucoidans. Our aim was to extract and isolate fucoidans from millennia-old marine sediments from the diatom dominated Bransfield Strait.

2 Materials and Methods

2.1 Study site and sample procurement

Sediment samples from the Bransfield Strait were provided as subsamples of sediment sections obtained from a multi corer (PS97/72-2) and piston corer (PS97/72-1) during the *RV Polarstern* expedition PS97, both taken at the same station (Latitude: -62.0065, Longitude: -56.0647, see Fig. 1). A total of ten sections, corresponding to ten depths ranging from 2 to 868 cm (see Table 1) were analysed. Sediments from this region are characterized as silt-bearing diatomaceous clay, with up to 40% comprising diatom assemblages i.e., 31.49×10^6 valves per gram (Cárdenas et al., 2019; Lamy, 2016). Resting spores of *Chaetoceros* sp. were the dominant contributors (Cárdenas et al., 2019). Previous radiocarbon measurements indicated that the sediment core spanned the extent of the Holocene i.e., deepest depth (868 cm) is 11.8 ka old (Vorrath et al., 2023b, a). The sediment sections were freeze-dried and then homogenized thoroughly with a mortar and pestle before all further analysis. For negative control, combusted sand was used (450 °C for 4 h).

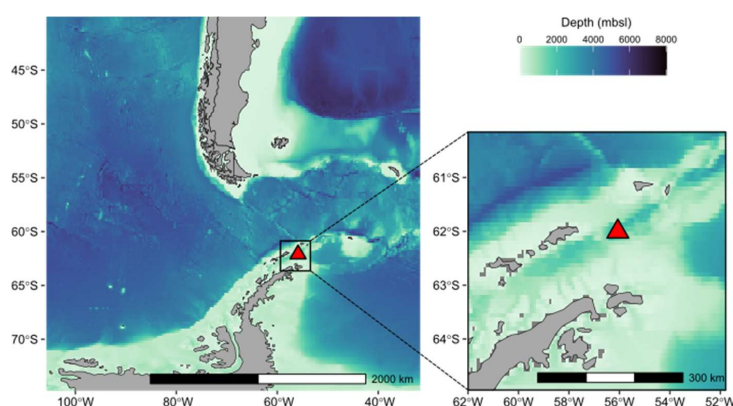


Figure 1: Location of sampling station in the Bransfield Strait. The background colours indicate bathymetry (mbsl i.e., meters below sea level) obtained from the NOAA using the ‘getNOAA.bathy()’ function in the ‘marmap’ package in R (version 4.2.1). The larger map on the left is shown at a spatial resolution of 10 arc-minutes, while the inset is at a spatial resolution of 5 arc-minutes.

2.2 Total organic carbon

Total organic carbon was measured as described earlier (Müller et al., 2025). Depending on sample availability, approximately 0.1 g of the homogenized sediment samples was decalcified by treatment with 12.5% hydrochloric acid and dried on a hot plate. Samples of known weight was then combusted at 1800° C in a stream of oxygen. Total organic carbon was quantified by measuring the released CO₂ with a LecoCS744 elemental analyser. The total organic carbon values are given as carbon weight (mg) per gram dry sediment.

2.3 Sequential glycan extraction

The glycans were extracted using a modified sequential extraction protocol (Hellige et al., 2026; Vidal-Melgosa et al., 2022). For every 10 mg of freeze-dried and homogenized sediment, 200 µL of the first solvent i.e., MilliQ water was added, and the sediment slurry was vortexed thoroughly. The slurry was then sonicated for 1 h in a sonication bath and centrifuged at 4000x g for 20 minutes. The supernatant was considered the MilliQ extract. The protocol was repeated for the pellet with same volumes of second solvent i.e., EDTA (0.3 M, pH 7.5), and the final solvent i.e., NaOH (4 M) + NaBH₄ (0.1% w/v). All extracts were stored at -20°C until further processing.

2.4 Extracted carbohydrate concentrations

Extracted carbohydrate concentrations were obtained by the phenol sulfuric assay (DuBois et al., 1956). 200 µL of the sequential extracts were mixed with 200 µL of 5% (v/v) Phenol and 1 mL of concentrated sulfuric acid (> 95%) was slowly added and mixed. The tubes were incubated at room temperature for 10 minutes and at 30 °C for 20 minutes to ensure completion of the reaction. After incubation, 200 µL was taken wells of a 96 well-plate. The absorbance was measured at 490 nm using a plate reader. Samples were referenced to increasing concentrations of D-glucose standards for quantification.



98 **2.5 Anion exchange chromatography**

99 MilliQ and EDTA extracts were pooled together and filtered through a 0.22 µm syringe filter. Two HiTrap ANX
100 FF (high sub) columns (column volume = 1mL) were connected in series to an Äkta start purification system. The
101 top column served as a guard column to trap any particles, pigments, and gas bubbles. Once the columns were
102 washed of any preservatives, they were equilibrated first with 10 column volumes (CV) of an equilibration buffer
103 (20 mM Tris-HCl, pH 7.5). Depending on the sample availabilities, the filtered samples were applied onto the
104 anion exchange column i.e. 8 mL for 11 cm and 464 cm, and 15 mL for the rest. 10 CV of the equilibration buffer
105 was then passed again through the column to wash any unbound molecules. The flow-through and the wash
106 fractions were collected together. A salt gradient spanning 20 CV starting with the equilibration buffer and ending
107 at the elution buffer (20 mM Tris-HCl, 5 M NaCl, pH 7.5) was applied for the elution of the bound molecules.
108 Additionally, 10 CV of the elution buffer was passed to elute any tightly bound molecules. All the elution steps
109 were collected as 1 mL fractions. In between the chromatography runs, the columns were washed with 0.1 M
110 NaOH. After every 4 runs, the columns were extensively cleaned with 1 M NaOH. A blank chromatography run
111 was also performed and the resultant fractions were used for background subtraction during all subsequent
112 analysis.

113 **2.6 Enzyme-linked immunosorbent assay (ELISA)**

114 ELISA protocol was followed as previously described (Cornuault et al., 2014; Vidal-Melgosa et al., 2021). The
115 monoclonal antibody BAM1 (SeaProbes) was used as a primary antibody to detect fucoidans (Torode et al., 2015).
116 Anti-rat antibody conjugated to horseradish peroxidase enzyme (IgG; Sigma-Aldrich) was used as the secondary
117 antibody. After the addition of the substrate (Tetramethylbenzidine, Slow kinetic form, Sigma-Aldrich), the plates
118 were incubated in dark until few samples had a strong colour development. The reaction was stopped with same
119 development time for all samples by adding 1 M HCl. The absorbance was measured at 450 nm.

120 **2.7 Deoxyhexose quantification assay (L-Cysteine assay)**

121 The deoxyhexose concentrations were estimated using a modified deoxyhexose quantification assay (Dische and
122 Shettles, 1948). 100 µL of sample were hydrolysed with 234 µL of concentrated sulfuric acid (>95%) at 99°C for
123 10 minutes on a heat block. The samples were then cooled rapidly on ice for 10 minutes. Subsequently, 10 µL of
124 3% (w/v) L-cysteine hydrochloride solution was added to the cool samples. After incubation in the dark for 1.5
125 hours, 200 µL of the solution was transferred into 96-well plates and the absorbance at 396 nm and 427 nm was
126 measured. Sample measurements were referenced to L-fucose standards in MilliQ water that were run in parallel.
127 The difference in absorbance at 396 nm and 427 nm in the standard curve is proportional to the concentration of
128 deoxyhexoses in the samples.

129 **2.8 Desalting by silica purification**

130 Protocol for desalting was followed as previously described (Schultz-Johansen et al., 2025). 100 µL of the anion
131 exchange fractions were diluted with 1900 µL of absolute ethanol, so that the final concentration of ethanol was
132 95% (v/v). The diluted sample was then loaded onto a NEB Monarch DNA Cleanup spin column. The samples
133 were spun at 15,000 x g for 5 minutes and the flow-through was discarded. Once all the sample was loaded onto
134 the column, the column was washed twice with 95% ethanol. The samples bound to the silica column were then
135 eluted with MilliQ into a clean collection tube.



136 **2.9 Monosaccharide quantification by high performance anion exchange chromatography with pulsed**
137 **amperometric detection (HPAEC-PAD)**

138 Prior to monosaccharide quantification, the samples were either desalted using silica purification (selected
139 fractions) or diluted 10-15 times (MilliQ extracts and EDTA extracts) so that the concentration of salt was low
140 enough for the HPAEC-PAD system. The samples were then acid hydrolysed in 1 M HCl (final conc.) for 24 h at
141 100 °C. Acid hydrolysed samples were dried in a centrifugal vacuum concentrator for 3 hours. A set of standard
142 mixtures containing known concentrations of monosaccharides in 1 M HCl were dried simultaneously with the
143 samples. The extracts were filtered through a 0.22 µm nylon-spin filter. All other fractions and standards were
144 spun at 21000 g for 10 minutes. 100 µl of the samples and 200 µl of the standard were transferred into HPLC vials
145 with an inlet. The vials were closed with slitted silicone sealing discs and placed in the auto-sampler of the Dionex
146 ICS-5000+ system. The columns used in the system were CarboPac PA10 analytical column (2 x 250 mm) and
147 CarboPac PA10 guard column (2 x 50 mm). After loading the samples onto the columns, the neutral and amino
148 sugars were separated using an isocratic phase of 18 mM NaOH, while the acidic monosaccharides were eluted
149 using a gradient reaching up to 200 mM NaCH₃COO. The peaks of the eluted sugars were identified using
150 retention times of the known monosaccharide standards. The peak areas were converted to concentration of the
151 monosaccharides (µg/L) based on the standard curves.

152 **3 Results**

153 **3.1 Contribution of carbohydrates to the buried organic carbon**

154 To assess the contribution of carbohydrates to buried organic carbon, we first quantified the total organic carbon
155 in the sediment by elemental analysis after combustion. Organic carbon concentrations ranged between 6.7 mgC.
156 g_{dw}⁻¹ (mg carbon per gram dry weight of sediment) at the lowest depth (868 cm) and 10.9 mgC. g_{dw}⁻¹ at the
157 uppermost depth (2 cm) (see Table 1).

158 Carbohydrates extracted sequentially with MilliQ water, EDTA and NaOH were quantified using phenol-sulfuric
159 acid assay. The extraction method was previously found suitable to extract different algae- and plant-derived types
160 of complex glycans from marine sediments (Vidal-Melgosa et al., 2022, Hellige et al., 2026). For example, EDTA
161 chelates cations cross-linked to anionic glycans such as pectin and other uronic acid-containing glycans, while
162 NaOH is required to extract crystalline, insoluble glycans such as cellulose. Because this approach extracts both,
163 free monomers and the ones polymerized in form of glycans, they are collectively described as carbohydrates at
164 this step of the analysis. While the NaOH extracts had the highest carbohydrate concentrations among the three
165 extracts (Fig. S1), they were omitted from the next steps as they are not amicable for the anion exchange
166 chromatography either due to the high pH or due to high salt concentrations after neutralization. The combined
167 carbohydrate carbon concentrations in the MilliQ and EDTA extracts ranged between 0.23 mgC. g_{dw}⁻¹ and 1.047
168 mgC. g_{dw}⁻¹ in 192 cm and 4 cm sediment depths, respectively (see Table 1). Carbohydrates in the MilliQ and
169 EDTA extracts constituted 2.45 – 10% of the total organic carbon across the sediment depths.

170 As fucose is a deoxyhexose sugar and a key constituent of fucoidan, deoxyhexoses were quantified in the MilliQ
171 and EDTA extracts using the L-cysteine assay. Deoxyhexose concentrations ranged between 0.009 mgC. g_{dw}⁻¹ and
172 0.044 mgC. g_{dw}⁻¹ in 868 cm depth and 6 cm depth, respectively (see Table 1). Therefore, on average the



173 deoxyhexoses contributed 5% of the carbohydrates extracted using MilliQ and EDTA and 0.22% of the total
174 organic carbon.

175 **Table 1: Contribution of carbohydrate carbon to the sedimentary organic carbon pool.**

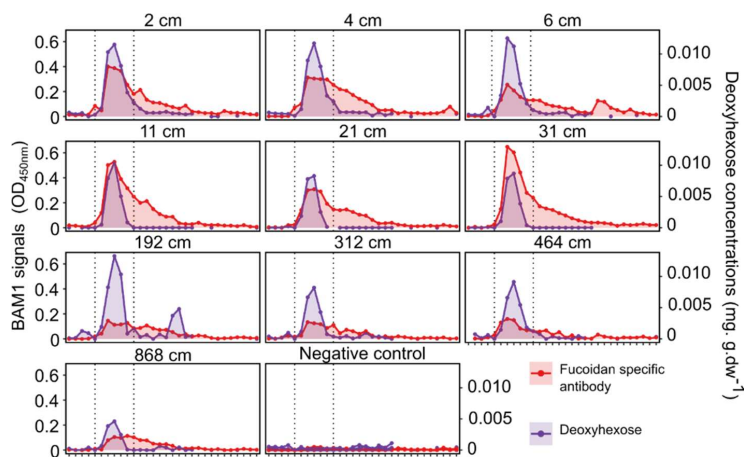
Depth	Age	Total organic carbon	Extracted carbohydrate concentration	Deoxyhexose concentration – before AEX	Deoxyhexose concentration – after AEX
[cm]	[ka BP]	[mgC. g _{dw} ⁻¹]	[mgC. g _{dw} ⁻¹]	[mgC. g _{dw} ⁻¹]	[mgC. g _{dw} ⁻¹]
2	-0.05	10.9	0.555	0.024	0.017
4	-0.042	10.2	1.047	0.030	0.016
6	-0.033	9.58	0.545	0.044	0.016
11	-0.015	9.95	0.404	0.027	0.011
21	0.024	7.85	0.240	0.018	0.009
31	0.074	9	0.349	0.018	0.010
192	2.935	9.36	0.230	0.008	0.016
312	4.162	9.37	0.292	0.015	0.009
464	5.766	8.58	0.336	0.016	0.011
868	11.738	6.66	0.335	0.009	0.005

176 Note: Age models of 2 cm – 31 cm depth intervals were established previously by excess ²¹⁰Pb dating (Vorrath et
177 al., 2020b, a), and the rest were determined previously by ¹⁴C dating (Vorrath et al., 2023b, a).

178 3.2 Chromatography assisted purification of fucoidans from Bransfield Strait sediments

179 We aimed to quantify the fucoidan concentrations in sediments from the Bransfield Strait. As we had limited
180 quantities of sediment samples from the Bransfield Strait, we used sediments from the Wadden Sea to test and
181 establish methodologies used in this study (see Fig. S2, Fig. S3). The established method, i.e. anion exchange
182 chromatography, was then used to isolate fucoidans from the bulk glycans extracted (MilliQ and EDTA extracts)
183 from Bransfield Strait. Epitope detection chromatography, a method described in previous studies uses specific
184 monoclonal antibodies for the detection of glycan epitopes after anion exchange purification (Cornuault et al.,
185 2014; Vidal-Melgosa et al., 2021). Similarly, we tested the chromatographic fractions for fucoidans using the
186 specific monoclonal antibody BAM1. Additionally, the deoxyhexose concentrations (as an estimate for L-fucose)
187 in the fractions were measured as a proxy for fucoidans. The concentrations of deoxyhexoses and the antibody
188 signals both peaked in the same elution window (see Fig. 2). Summed deoxyhexose concentrations in this peak
189 ranged between 0.005 mgC. g_{dw}⁻¹ to 0.017 mgC. g_{dw}⁻¹ in sediment depths 868 cm and 2 cm, respectively (see Table
190 1). Except for depth 192 cm, 35-70% of the deoxyhexose in the MilliQ and EDTA extracts were recovered using
191 anion exchange chromatography (see Fig. S4). At depth 192 cm, a 200% recovery was observed, indicating
192 possible matrix effects causing underestimation of the crude extracts or overestimation in the purified samples.

193



194

195 **Figure 2: Chromatographic purification of fucoidans from marine sediments confirmed by different**
196 **analyses.** Chromatographic fractions (x-axis) were probed for fucoidans using the monoclonal antibody BAM1,
197 and deoxyhexose concentrations were measured with the L-cysteine assay. The monoclonal antibody signals
198 (OD_{450nm}) in each fraction are shown in red, while the deoxyhexose concentrations are shown in purple. The region
199 where the relative antibody signals and the deoxyhexose concentrations peak and overlap are indicated by the
200 vertical dotted lines in each facet. This region signifies the fucoidans isolated from the corresponding depth.

201 3.3 Monosaccharide composition before and after anion exchange chromatography

202 Monosaccharide composition was quantified in the pooled MilliQ and EDTA extracts and also in the pooled
203 fractions from the elution window using HPAEC-PAD. Galactose was the primary monosaccharide in both the
204 crude extracted carbohydrates (average relative abundance of 22.9%) and in the isolated fucoidans (average
205 relative abundance 28%). Fucose, galactose, glucose, xylose, galacturonic acid, and glucuronic acid have higher
206 relative abundances in the isolated fucoidans in comparison to the crude extracts (see Fig. 3). Arabinose, rhamnose,
207 and glucosamine show lower relative abundances in the isolated fucoidans. There are notable increases in the
208 relative abundances of xylose and glucose in sediment depths 464 cm and 868 cm, respectively. The differences
209 in monosaccharide compositions before and after anion exchange chromatography indicates the selection of certain
210 monosaccharides and omission of others.

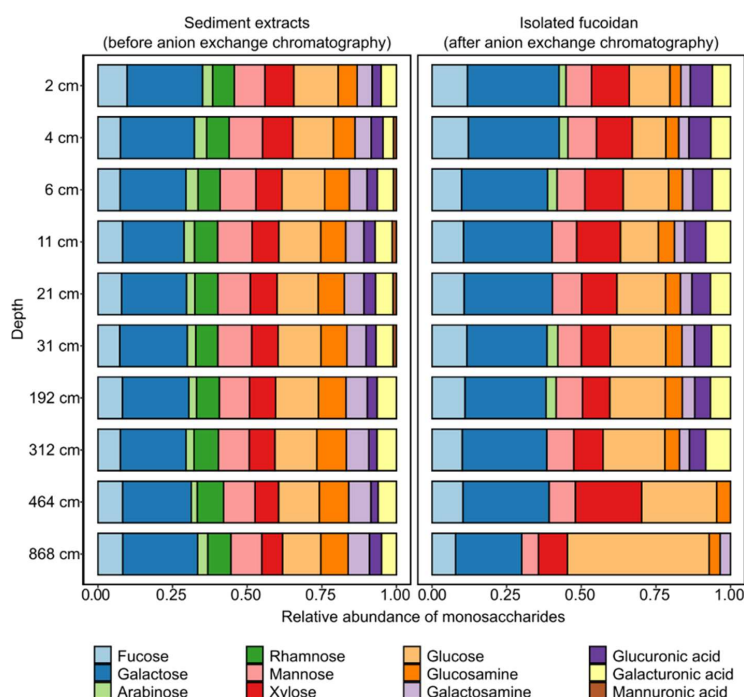


Figure 3: Enrichment of fucoidan monosaccharides after anion exchange chromatography. Relative abundances of monosaccharides before (left) and after (right) anion exchange chromatography are compared. Certain monosaccharides often associated with the building blocks of fucoidan like fucose, galactose, glucose, xylose, and the uronic acids are enriched after anion exchange chromatography. Monosaccharides like rhamnose and arabinose, however, have a lower abundance indicating selection of certain glycans.

4 Discussion

In this study, we estimated the contribution of fucoidans towards carbon sequestration in millennia-old marine sediments. We were able to overcome limitations of existing analytical methods via separation of fucoidans from the sediment matrix using anion exchange chromatography. Our approach was built on previous applications of anion exchange chromatography for quantification and characterization of anionic glycans from diverse biological and environmental sources (Cornuault et al., 2014; Huang et al., 2021; Sichert et al., 2021; Vidal-Melgosa et al., 2021; Buck-Wiese et al., 2023; Krull et al., 2026). Fucoidan-specific monoclonal antibody BAM1 (Torode et al., 2015) also provided support for the isolation of fucoidan from the sediment matrix. While other glycans could have coeluted with the fucoidans, the monosaccharide constituents of the sediment fucoidan corroborated with previously reported fucoidans (Huang et al., 2021; Ponce and Stortz, 2020; Vidal-Melgosa et al., 2021). Future enzyme-based methods could provide additional structure-specific quantification and confirmation of fucoidans (Becker et al., 2017; Buck-Wiese et al., 2023; Nakazato et al., 2026). However, this is limited by the availability of appropriate enzymes, as marine sediments may comprise structurally diverse and poorly characterized fucoidans from multiple sources. The presence of diverse glycans and polyphenols in complex sediment matrices can also obscure the enzymatic quantification of glycans (Nakazato et al., 2026). The necessity of clean-up methods further



highlights the value of anion exchange chromatography, as used in this study. Therefore, by enabling the isolation of fucoidans, anion exchange chromatography provides a basis for accurately evaluating the roles of these glycans in long-term carbon sequestration.

Fucoidans isolated from the Bransfield Strait likely have a diatomaceous origin. While fucoidans are primarily exuded by both brown algae and diatoms (Buck-Wiese et al., 2023; Huang et al., 2021; Vidal-Melgosa et al., 2021), this region is diatom-dominated with a high proportion of *Chaetoceros* resting spores in the sediment (Cárdenas et al., 2019; Lamy, 2016). However, previous studies also have reported the presence of brown algal detritus (Fischer and Wiencke, 1992; Reichardt, 1987). Similar monosaccharide compositions, with high percentages of galactose have been observed in fucoidans isolated from both diatoms and brown algae (Huang et al., 2021; Ponce and Stortz, 2020; Vidal-Melgosa et al., 2021). Hence, unambiguous assignment of the origin of fucoidan would require a higher structural resolution. Recently, oligosaccharide fingerprints obtained by MALDI-MS analysis of fucoidan hydrolysates have been shown to distinguish the different biological origins (Reyes-Weiss et al., 2024). In the future, oligosaccharide fingerprints of fucoidans from the sediment can be compared to fingerprints of various micro- and macroalgal fucoidans, to better understand their origins. Overall, our study demonstrates that isolation of glycans from marine sedimentary records using anion exchange chromatography provides the foundation for resolving their origins and assessing the roles of various algae in carbon sequestration.

Fucoidan sequesters carbon for climate-relevant timescales. Assuming a constant rate of fucoidan deposition, we find that at least 30% of the fucoidan at the uppermost depth ($0.017 \text{ mgC} \cdot \text{g}_{\text{dw}}^{-1}$) was sequestered for the entire extent of the Holocene (i.e., $0.005 \text{ mgC} \cdot \text{g}_{\text{dw}}^{-1}$). The binding of fucoidans to the anion exchange column also suggest that they remain sufficiently intact in millennia-old sediments, because low molecular weight molecules are washed away (Buck-Wiese et al., 2023). The preservation of fucoidan carbon has also been semi-quantitatively observed in centuries- to millennia-old sediments (Vidal-Melgosa et al., 2022; Salmeán et al., 2022). These studies were unable to discern if the preservation was attributed to the recalcitrance of fucoidan or the harsh depositional environments like hyper-salinity and anoxia (in the Red Sea), and periodic anoxia (in the Koljö Fjord). However, our findings indicate that fucoidan resists bacterial degradation, as it is transported to the sediment in a 2000 m deep oxic water column. Other than the direct contribution to carbon sequestration, fucoidan also transports other molecules to the sediment by forming aggregates (Vidal-Melgosa et al., 2021). The presence of fucoidan in sinking particles could protect labile organic carbon from bacterial degradation, analogous to their proposed roles in the extracellular matrix of diatoms and brown algae (Bligh et al., 2022) and similar to the shielding effects of mineral associations (Keil et al., 1994). Therefore, fucoidans are important components of the biological carbon pump and may contribute to long-term carbon sequestration through different mechanisms.

5 Outlook

The cultivation of macroalgae is increasingly being considered as a potential climate mitigation strategy. Evidence of algae detritus in marine sediments or the increased accumulation of organic carbon in sediment near algae farms point to carbon sequestration potential of brown algae (Fischer and Wiencke, 1992; Reichardt, 1987; Krause-Jensen and Duarte, 2016; Duarte et al., 2025). However, little is known about the far-field carbon sequestration efficiency of algae farms, due to the lack of stable long-term proxies that can be specifically attributed to the brown algal input. Our study provides a framework for the quantification of fucoidans in sediments, thereby enabling



270 their use as proxies to assess carbon sequestration. This approach therefore may facilitate the long-term monitoring
271 of brown algal carbon sequestration beyond the immediate vicinity of cultivation sites.

272 **Data availability**

273 Data reported in this paper will be submitted to the online data archive PANGAEA, and will be made accessible
274 to the public at the time of publication.

275 **Author contributions**

276 AAM, and JHH designed research with contributions from all authors. SVM and JHH supervised the project. AAM
277 performed all laboratory methods. LR developed and validated the modified L-cysteine assay (manuscript in prep).
278 JM collected and provided sediment samples from the Bransfield Strait. AAM, KWB, and JHH led the
279 interpretation and discussion of data and manuscript preparation, with inputs from all authors.

280 **Competing interests**

281 The authors have declared no competing interest.

282 **Acknowledgements**

283 The authors thank technicians at the MARUM - Centre for Marine Environmental Sciences, University of Bremen,
284 namely Brit Kockisch for total organic carbon measurement; and the Max Plank Institute for Marine Microbiology
285 Bremen, Germany namely Katharina Föll for HPAEC-PAD measurement. The authors also thank student assistant
286 Kumaresh Jawahar Natarajan for sample processing. We further acknowledge the captain and crew of RV
287 *Polarstern* expedition PS97 (GPF grant number AWI_PS97_01).

288 **Financial support**

289 This work has been funded by EU Consolidator ERC-grant, C-Quest, “Discover molecular pathways for glyco-
290 carbon sequestration” (101044738)

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292 **References**

- 293 Aluwihare, L. I., Repeta, D. J., and Chen, R. F.: Chemical composition and cycling of dissolved organic matter in
294 the Mid-Atlantic Bight, Deep Sea Res. Part II Top. Stud. Oceanogr., 49, 4421–4437,
295 [https://doi.org/10.1016/S0967-0645\(02\)00124-8](https://doi.org/10.1016/S0967-0645(02)00124-8), 2002.
- 296 Arnosti, C., Wietz, M., Brinkhoff, T., Hehemann, J.-H., Probandt, D., Zeugner, L., and Amann, R.: The
297 Biogeochemistry of Marine Polysaccharides: Sources, Inventories, and Bacterial Drivers of the Carbohydrate
298 Cycle, Annu Rev Mar Sci, 13, 81–108, <https://doi.org/10.1146/annurev-marine-032020>, 2021.
- 299 Becker, S., Scheffel, A., Polz, M. F., and Hehemann, J. H.: Accurate quantification of laminarin in marine organic
300 matter with enzymes from marine microbes, Appl. Environ. Microbiol., 83, [https://doi.org/10.1128/AEM.03389-](https://doi.org/10.1128/AEM.03389-16)
301 16, 2017.
- 302 Bligh, M., Nguyen, N., Buck-Wiese, H., Vidal-Melgosa, S., and Hehemann, J.-H.: Structures and functions of
303 algal glycans shape their capacity to sequester carbon in the ocean, Curr. Opin. Chem. Biol., 71, 102204,
304 <https://doi.org/10.1016/j.cbpa.2022.102204>, 2022.
- 305 Buck-Wiese, H., Andskog, M. A., Nguyen, N. P., Bligh, M., Asmala, E., Vidal-Melgosa, S., Liebeke, M.,
306 Gustafsson, C., and Hehemann, J.-H.: Fucoid brown algae inject fucoidan carbon into the ocean, Proc. Natl. Acad.
307 Sci., 120, e2210561119, <https://doi.org/10.1073/pnas.2210561119>, 2023.
- 308 Burdige, D. J.: Preservation of Organic Matter in Marine Sediments: Controls, Mechanisms, and an Imbalance in
309 Sediment Organic Carbon Budgets?, Chem. Rev., 107, 467–485, <https://doi.org/10.1021/cr050347q>, 2007.
- 310 Burdige, D. J., Skoog, A., and Gardner, K.: Dissolved and particulate carbohydrates in contrasting marine
311 sediments, Geochim. Cosmochim. Acta, 64, 1029–1041, [https://doi.org/10.1016/S0016-7037\(99\)00361-0](https://doi.org/10.1016/S0016-7037(99)00361-0), 2000.
- 312 Cárdenas, P., Lange, C. B., Vernet, M., Esper, O., Srain, B., Vorrath, M.-E., Ehrhardt, S., Müller, J., Kuhn, G.,
313 Arz, H. W., Lembke-Jene, L., and Lamy, F.: Biogeochemical proxies and diatoms in surface sediments across the
314 Drake Passage reflect oceanic domains and frontal systems in the region, Prog. Oceanogr., 174, 72–88,
315 <https://doi.org/10.1016/j.pocean.2018.10.004>, 2019.
- 316 Cavan, E. L., Le Moigne, F. a. C., Poulton, A. J., Tarling, G. A., Ward, P., Daniels, C. J., Fragoso, G. M., and
317 Sanders, R. J.: Attenuation of particulate organic carbon flux in the Scotia Sea, Southern Ocean, is controlled by
318 zooplankton fecal pellets, Geophys. Res. Lett., 42, 821–830, <https://doi.org/10.1002/2014GL062744>, 2015.
- 319 Chin, W.-C., Orellana, M. V., and Verdugo, P.: Spontaneous assembly of marine dissolved organic matter into
320 polymer gels, Nature, 391, 568–572, 1998.
- 321 Cornuault, V., Manfield, I. W., Ralet, M.-C., and Knox, J. P.: Epitope detection chromatography: a method to
322 dissect the structural heterogeneity and inter-connections of plant cell-wall matrix glycans, Plant J., 78, 715–722,
323 <https://doi.org/10.1111/tj.12504>, 2014.
- 324 Cowie, G. L. and Hedges, J. I.: Carbohydrate sources in a coastal marine environment, Geochim. Cosmochim.
325 Acta, 48, 2075–2087, [https://doi.org/10.1016/0016-7037\(84\)90388-0](https://doi.org/10.1016/0016-7037(84)90388-0), 1984.
- 326 Dische, Zacharias. and Shettles, L. B.: A specific color reaction of methylpentoses and a spectrophotometric
327 micromethod for their determination, J. Biol. Chem., 175, 595–603, [https://doi.org/10.1016/S0021-](https://doi.org/10.1016/S0021-9258(18)57178-7)
328 9258(18)57178-7, 1948.
- 329 Dittmar, T., Lennartz, S. T., Buck-Wiese, H., Hansell, D. A., Santinelli, C., Vanni, C., Blasius, B., and Hehemann,
330 J. H.: Enigmatic persistence of dissolved organic matter in the ocean, Nat. Rev. Earth Environ., 2, 570–583,
331 <https://doi.org/10.1038/s43017-021-00183-7>, 2021.
- 332 Duarte, C. M. and Cebrián, J.: The fate of marine autotrophic production, Limnol. Oceanogr., 41, 1758–1766,
333 <https://doi.org/10.4319/lo.1996.41.8.1758>, 1996.
- 334 Duarte, C. M., Delgado-Huertas, A., Marti, E., Gasser, B., Martin, I. S., Cousteau, A., Neumeyer, F., Reilly-
335 Cayten, M., Boyce, J., Kuwae, T., Hori, M., Miyajima, T., Price, N. N., Arnold, S., Ricart, A. M., Davis, S.,
336 Surugau, N., Abdul, A.-J., Wu, J., Xiao, X., Chung, I. K., Choi, C. G., Sondak, C. F. A., Albasri, H., Krause-



- 337 Jensen, D., Bruhn, A., Boderskov, T., Hancke, K., Funderud, J., Borrero-Santiago, A. R., Pascal, F., Joanne, P.,
338 Ranivoarivelo, L., Collins, W. T., Clark, J., Gutierrez, J. F., Riquelme, R., Avila, M., Macreadie, P. I., and Masque,
339 P.: Carbon burial in sediments below seaweed farms matches that of Blue Carbon habitats, *Nat. Clim. Change*, 15,
340 180–187, <https://doi.org/10.1038/s41558-024-02238-1>, 2025.
- 341 DuBois, Michel., Gilles, K. A., Hamilton, J. K., Rebers, P. A., and Smith, Fred.: Colorimetric Method for
342 Determination of Sugars and Related Substances, *Anal. Chem.*, 28, 350–356,
343 <https://doi.org/10.1021/ac60111a017>, 1956.
- 344 Engel, A., Silke, T., Riebesell, U., Rochelle-Newall, E., and Zondervan, I.: Polysaccharide aggregation as a
345 potential sink of marine dissolved organic carbon, *Nature*, 428, 929–932, <https://doi.org/10.1038/nature02506>,
346 2004.
- 347 Eppley, R. W. and Peterson, B. J.: Particulate organic matter flux and planktonic new production in the deep ocean,
348 *Nature*, 282, 677–680, <https://doi.org/10.1038/282677a0>, 1979.
- 349 Fischer, G. and Wiencke, C.: Stable carbon isotope composition, depth distribution and fate of macroalgae from
350 the Antarctic Peninsula region, *Polar Biol.*, 12, 341–348, <https://doi.org/10.1007/BF00243105>, 1992.
- 351 Hedges, J. I., Baldock, J. A., Gélinas, Y., Lee, C., Peterson, M., and Wakeham, S. G.: Evidence for non-selective
352 preservation of organic matter in sinking marine particles, *Nature*, 409, 801–804,
353 <https://doi.org/10.1038/35057247>, 2001.
- 354 Hellige, I., Akeerath Mundanatt, A., Massing, J. C., and Hehemann, J.-H.: Fucoidan carbon is stored in coastal
355 vegetated ecosystems, *Biogeosciences*, 23, 387–398, <https://doi.org/10.5194/bg-23-387-2026>, 2026.
- 356 Huang, G., Vidal-Melgosa, S., Sichert, A., Becker, S., Fang, Y., Niggemann, J., Iversen, M. H., Cao, Y., and
357 Hehemann, J.-H.: Secretion of sulfated fucans by diatoms may contribute to marine aggregate formation, *Limnol.*
358 *Oceanogr.*, 66, 3768–3782, <https://doi.org/10.1002/lno.11917>, 2021.
- 359 Keil, R. G., Montluçon, D. B., Prahl, F. G., and Hedges, J. I.: Sorptive preservation of labile organic matter in
360 marine sediments, *Nature*, 370, 549–552, <https://doi.org/10.1038/370549a0>, 1994.
- 361 Klok, J., Cox, H. C., Baas, M., Schuyl, P. J. W., De Leeuw, J. W., and Schenck, P. A.: Carbohydrates in recent
362 marine sediments-I. Origin and significance of deoxy- and O-methyl-monosaccharides, *Org Geochem*, 7, 73–84,
363 [https://doi.org/10.1016/0146-6380\(84\)90138-4](https://doi.org/10.1016/0146-6380(84)90138-4), 1984.
- 364 Krause-Jensen, D. and Duarte, C. M.: Substantial role of macroalgae in marine carbon sequestration, *Nat. Geosci.*,
365 9, 737–742, <https://doi.org/10.1038/NGEO2790>, 2016.
- 366 Krull, J., Sidhu, C., Solanki, V., Bligh, M., Rößler, L., Singh, R. K., Huang, G., Robb, C. S., Teeling, H., Seeberger,
367 P. H., Schweder, T., Crawford, C. J., and Hehemann, J.-H.: Sulfated mannan of diatoms selects host-specific
368 microbiota in the sunlit ocean, *Microbiome*, 14, 112, <https://doi.org/10.1186/s40168-026-02379-9>, 2026.
- 369 Lamy, F.: The expedition PS97 of the research vessel POLARSTERN to the Drake Passage in 2016, *Berichte zur*
370 *Polar- und Meeresforschung = Reports on Polar and Marine Research*, Alfred-Wegener-Institut, Helmholtz-
371 Zentrum für Polar- und Meeresforschung, https://doi.org/10.2312/BZPM_0701_2016, 2016.
- 372 Le Moigne, F. A. C.: Pathways of Organic Carbon Downward Transport by the Oceanic Biological Carbon Pump,
373 *Front. Mar. Sci.*, 6, <https://doi.org/10.3389/fmars.2019.00634>, 2019.
- 374 Lee, R. B. Y., Mavridou, D. A. I., Papadakis, G., McClelland, H. L. O., and Rickaby, R. E. M.: The uronic acid
375 content of coccolith-associated polysaccharides provides insight into coccolithogenesis and past climate, *Nat.*
376 *Commun.*, 7, 13144, <https://doi.org/10.1038/ncomms13144>, 2016.
- 377 Li, Z., Zhang, Y. G., Torres, M., and Mills, B. J. W.: Neogene burial of organic carbon in the global ocean, *Nature*,
378 613, 90–95, <https://doi.org/10.1038/s41586-022-05413-6>, 2023.
- 379 Miyajima, T., Ogawa, H., and Koike, I.: Alkali-extractable polysaccharides in marine sediments: Abundance,
380 molecular size distribution, and monosaccharide composition, *Geochim. Cosmochim. Acta*, 65, 1455–1466,
381 [https://doi.org/10.1016/S0016-7037\(00\)00612-8](https://doi.org/10.1016/S0016-7037(00)00612-8), 2001.



- 382 Müller, D., Liu, B., Geibert, W., Holtappels, M., Sander, L., Miramontes, E., Taubner, H., Henkel, S., Hinrichs,
383 K.-U., Bethke, D., Dohrmann, I., and Kasten, S.: Depositional controls and budget of organic carbon burial in fine-
384 grained sediments of the North Sea – the Helgoland Mud Area as a natural laboratory, *Biogeosciences*, 22, 2541–
385 2567, <https://doi.org/10.5194/bg-22-2541-2025>, 2025.
- 386 Nakazato, S., Kawahara, T., Ichii, R., Akita, S., Inoue, A., Fujita, M., Kurihara, H., Nakaya, M., and Ooki, A.:
387 Technical note: Development of an extraction protocol and colorimetric analysis for alginate in marine sediment,
388 *Biogeosciences*, 23, 2747–2759, <https://doi.org/10.5194/bg-23-2747-2026>, 2026.
- 389 Ponce, N. M. A. and Stortz, C. A.: A Comprehensive and Comparative Analysis of the Fucoidan Compositional
390 Data Across the Phaeophyceae, *Front. Plant Sci.*, 11, <https://doi.org/10.3389/fpls.2020.556312>, 2020.
- 391 Reichardt, W. T.: Burial of Antarctic macroalgal debris in bioturbated deep-sea sediments, *Deep Sea Res. Part*
392 *Oceanogr. Res. Pap.*, 34, 1761–1770, [https://doi.org/10.1016/0198-0149\(87\)90024-0](https://doi.org/10.1016/0198-0149(87)90024-0), 1987.
- 393 Reyes-Weiss, D. S., Bligh, M., Rhein-Knudsen, N., Hehemann, J.-H., Liebeke, M., Westereng, B., and Horn, S.
394 J.: Application of MALDI-MS for characterization of fucoidan hydrolysates and screening of *endo*-fucoidanase
395 activity, *Carbohydr. Polym.*, 340, 122317, <https://doi.org/10.1016/j.carbpol.2024.122317>, 2024.
- 396 Salmeán, A. A., Willats, W. G. T., Ribeiro, S., Andersen, T. J., and Ellegaard, M.: Over 100-Year Preservation
397 and Temporal Fluctuations of Cell Wall Polysaccharides in Marine Sediments, *Front. Plant Sci.*, 13,
398 <https://doi.org/10.3389/fpls.2022.785902>, 2022.
- 399 Schultz-Johansen, M., Parnami, J., Hellige, I., and Hehemann, J.-H.: Silica-based purification and desalting of
400 fucoidan, *chemRxiv [preprint]*, <https://doi.org/10.26434/chemrxiv-2025-fc6bq>, 11 February 2025.
- 401 Sichert, A., Le Gall, S., Klau, L. J., Laillet, B., Rogniaux, H., Achmann, F. L., and Hehemann, J. H.: Ion-exchange
402 purification and structural characterization of five sulfated fucoidans from brown algae, *Glycobiology*, 31, 352–
403 357, <https://doi.org/10.1093/glycob/cwaa064>, 2021.
- 404 Taylor, I. S., Paterson, D. M., and Mehler, A.: The quantitative variability and monosaccharide composition of
405 sediment carbohydrates associated with intertidal diatom assemblages, *Biogeochemistry*, 45, 303–327,
406 <https://doi.org/10.1007/BF00993005>, 1999.
- 407 Torode, T. A., Marcus, S. E., Jam, M., Tonon, T., Blackburn, R. S., Hervé, C., and Knox, J. P.: Monoclonal
408 antibodies directed to fucoidan preparations from brown algae, *PLoS ONE*, 10,
409 <https://doi.org/10.1371/journal.pone.0118366>, 2015.
- 410 Vidal-Melgosa, S., Sichert, A., Francis, T. B., Bartosik, D., Niggemann, J., Wichels, A., Willats, W. G. T., Fuchs,
411 B. M., Teeling, H., Becher, D., Schweder, T., Amann, R., and Hehemann, J. H.: Diatom fucan polysaccharide
412 precipitates carbon during algal blooms, *Nat. Commun.*, 12, <https://doi.org/10.1038/s41467-021-21009-6>, 2021.
- 413 Vidal-Melgosa, S., Lagator, M., Sichert, A., Priest, T., Pätzold, J., and Hehemann, J.-H.: Not digested: algal
414 glycans move carbon dioxide into the deep-sea, <https://doi.org/10.1101/2022.03.04.483023>, 4 March 2022.
- 415 Vorrath, M.-E., Müller, J., Rebolledo, L., Cárdenas, P., Shi, X., Esper, O., Opel, T., Geibert, W., Muñoz, P., Haas,
416 C., Lange, C. B., Lohmann, G., and Mollenhauer, G.: Geochemistry and biomarker of sediment core PS97/072-2,
417 PANGAEA [dataset], <https://doi.org/10.1594/PANGAEA.918736>, 2020a.
- 418 Vorrath, M.-E., Müller, J., Rebolledo, L., Cárdenas, P., Shi, X., Esper, O., Opel, T., Geibert, W., Muñoz, P., Haas,
419 C., Kuhn, G., Lange, C. B., Lohmann, G., and Mollenhauer, G.: Sea ice dynamics in the Bransfield Strait, Antarctic
420 Peninsula, during the past 240 years: a multi-proxy intercomparison study, *Clim. Past*, 16, 2459–2483,
421 <https://doi.org/10.5194/cp-16-2459-2020>, 2020b.
- 422 Vorrath, M.-E., Müller, J., Cárdenas, P., Opel, T., Mieruch, S., Esper, O., Lembke-Jene, L., Etourneau, J., Vieth-
423 Hillebrand, A., Lahajnar, N., Lange, C. B., Leventer, A., Evangelinos, D., Escutia, C., and Mollenhauer, G.:
424 Deglacial and Holocene sea-ice and climate dynamics in the Bransfield Strait, northern Antarctic Peninsula, *Clim.*
425 *Past*, 19, 1061–1079, <https://doi.org/10.5194/cp-19-1061-2023>, 2023a.
- 426 Vorrath, M.-E., Müller, J., Cárdenas, P., Mieruch, S., Esper, O., Opel, T., Lembke-Jene, L., Etourneau, J., Vieht-
427 Hillebrand, A., Lahajnar, N., Lange, C. B., Leventer, A., Evangelinos, D., Escutia, C., and Mollenhauer, G.:



428 Radiocarbon ages from the sediment core PS97/72-1, PANGAEA [dataset],
429 <https://doi.org/10.1594/PANGAEA.952274>, 2023b.

430

431

432

433

434