



1 **Measurement report: Molecular-level insights into the seasonal**
2 **formation of organosulfates in a mountain atmosphere**

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Abstract: Organosulfates (OSs) are important components of secondary organic aerosol, yet their molecular composition and formation processes in mountain atmospheres remain poorly constrained. In this study, PM_{2.5} samples collected during summer and winter at a high-altitude background site at Wuyi Mountain, southeastern China, were analyzed using high-performance liquid chromatography coupled with orbitrap mass spectrometry (HPLC-Orbitrap MS) to examine how seasonal changes in precursor availability and atmospheric processing influence OS molecular composition. Clear seasonal differences were observed in the molecular composition and structural characteristics. Summer OSs were enriched in more highly oxygenated species (with $\overline{\text{OS}}_{\text{c}} > -1$ accounting for 72% of the total OS signal abundance), whereas winter OSs were characterized by larger (C₅–C₂₃), less oxygenated (O₄–O₇), and more unsaturated molecules, reflecting more diverse precursor influences. Among 35 structurally identified OSs, isoprene- and monoterpene-derived species were enriched in summer, while aliphatic-like OSs and monoterpene-derived nitrooxy-OSs were more abundant in winter, indicating seasonal variations in precursor availability and the relative importance of atmospheric processing. Mechanistic analysis suggested that summer OS formation was more closely associated with sulfate-rich and highly acidic conditions favorable for multiphase processing, whereas wintertime OS formation reflected greater influences from elevated atmospheric oxidants and nitrogen-related chemistry. Regional transport further contributed to wintertime OS composition by increasing anthropogenic precursor inputs and promoting atmospheric aging. These findings demonstrate that seasonal OS variability in mountain environments reflects the interplay between precursor availability and atmospheric processing, providing field-based molecular evidence for the coupling of biogenic and anthropogenic influences in secondary organic aerosol formation and evolution.

Keywords: Organosulfates, Molecular composition, Seasonal variation, Multiphase chemistry, Mountain atmosphere



1. Introduction

Organosulfates (OSs) are an important class of sulfur-containing components in secondary organic aerosol (SOA), accounting for 3–30% of organic matter mass in ambient fine particles (Mazzoleni et al. 2010; Tolocka et al. 2012; Hettiyadura et al. 2018). They may also contribute up to 12% of particulate total sulfur, playing a significant role in the global biogeochemical sulfur cycle (Lukács et al. 2009). OSs are primarily formed through the oxidation of volatile organic compounds (VOCs), followed by sulfate esterification and related aqueous-phase or multiphase reactions involving SO₂ or H₂SO₄. Due to the presence of sulfate functional groups (–OSO₃[–]), OSs generally exhibit high hygroscopicity and thermal instability, influencing aerosol water solubility and subsequent chemical transformations (Brüggemann et al. 2020). OSs may also modify particle surface properties and cloud condensation nuclei activity, thereby influencing aerosol-cloud interactions and climate-relevant processes (Lin et al. 2013; Hansen et al. 2015; Estillore et al. 2016). Despite important insights from laboratory and field studies, the processes governing the abundance and molecular composition of OSs in the ambient atmosphere remain incompletely understood.

OSs can be broadly classified into biogenic and anthropogenic types based on their precursor sources. Biogenic OSs are mainly derived from the oxidation of plant-emitted isoprene and terpenes (Stone et al. 2012; He et al. 2014), whereas anthropogenic OSs have been linked to aromatic compounds, alkanes and unsaturated fatty acids emitted from fossil fuel combustion, biomass burning, and cooking activities (Riva et al. 2015; Riva et al. 2016; Blair et al. 2017). A variety of formation processes have been proposed for different OS classes. For example, isoprene-derived OSs are closely linked to the acid-catalyzed ring-opening reactions of epoxide intermediates on sulfate-containing particles (Lin et al. 2013; Surratt et al. 2008). Chamber studies have also shown that monoterpenes (e.g., β-pinene) can be oxidized by O₃ to form alcohol intermediates that subsequently undergo sulfate esterification in the particle phase (Iinuma et al. 2007). Additional OS formation may occur through heterogeneous reactions of SO₂ with monoterpene-derived SOA, unsaturated fatty acids, or organic peroxides (Passananti et al. 2016; Yao et al. 2019; Zhu et al. 2019). For nitrooxy-OSs, nighttime NO₃ radical-initiated oxidation of VOCs has been identified as an important source of organic nitrate and peroxy intermediates that may subsequently undergo particle-phase processing (Wang et al. 2020; Yang et al. 2023). The relative importance of these processes depends



78 on precursor composition and environmental conditions, including sulfate availability,
79 aerosol acidity, aerosol liquid water content (ALWC), and atmospheric oxidants (e.g.,
80 O₃) (Duporté et al. 2020; Aoki et al. 2020; Duporté et al. 2016). However, only a limited
81 number of laboratory-proposed processes have been evaluated through field
82 observations, and a substantial fraction of ambient OSs remains structurally unresolved
83 (Chen et al. 2021). Molecular-level investigation that link OS composition with
84 precursor sources and atmospheric conditions are therefore needed to better understand
85 their formation in the ambient atmosphere.

86 High-altitude mountain sites provide particularly useful natural environments for
87 investigating OS formation, which are characterized by abundant biogenic VOC
88 emissions, persistently high relative humidity (RH), and frequent influence from
89 regional transport of anthropogenic pollutants (Zhao et al. 2020; Wang et al. 2021a; Cao
90 et al. 2022). These conditions can promote photochemical aging, gas-particle
91 partitioning, and aqueous-phase or heterogeneous reactions (Hallquist et al. 2009;
92 McNeill 2015). Elevated ALWC associated with high RH may enhance the uptake and
93 diffusion of gaseous and dissolved precursors, thereby facilitating sulfate esterification
94 and related OS formation. Summer observations in the Qinling Mountains, for example,
95 indicated that elevated RH favored the formation of monoterpene- and long-chain
96 alkane-derived OSs and that aqueous-phase reactions contributed to multigenerational
97 oxidation (Li et al. 2024). At the same time, the coexistence of biogenic precursors and
98 regionally transported anthropogenic pollutants may increase the diversity of OS
99 precursors and atmospheric processing. Consequently, OS molecular composition in
100 mountain environments may reflect the combined influences of precursor emissions,
101 atmospheric oxidation, multiphase chemistry, and regional transport. However, how the
102 relative importance of these factors varies seasonally and influences OS formation
103 remains poorly understood.

104 In this study, PM_{2.5}-bound OSs collected during summer and winter at a high-
105 altitude background site on Wuyi Mountain were characterized using high-performance
106 liquid chromatography coupled with orbitrap mass spectrometry (HPLC-Orbitrap MS).
107 The objectives were to (1) characterize the molecular composition, structural
108 characteristics, and seasonal variability of OSs; (2) identify major OS classes and
109 determine their relative contributions in summer and winter; and (3) investigate the
110 possible formation processes and regional transport influences of different OS groups



111 by integrating molecular observations with sulfate, aerosol acidity, ALWC, oxidants,
112 and backward-trajectory analysis.

113 **2. Materials and methods**

114 **2.1 Sample Collection**

115 Ambient PM_{2.5} samples were collected from 3rd June to 26th July 2016 and from
116 27th December 2016 to 24th February 2017, representing summer and winter,
117 respectively (Table S1 in the Supporting Information). The sampling site is located on
118 the Wuyi Mountain (27.64°N, 117.99°E) in Fujian province. The samples were
119 collected on pre-baked (450 °C, 3 h) quartz fiber filters (20.3 × 25.4 cm; Whatman,
120 QM-A, Clifton, US) for 46 h using a high-volume PM_{2.5} air sampler (Tisch TE-6070V-
121 2.5-HVS, Cleveland, US) at a flow rate of 1.05 m³ min⁻¹. After collection, the samples
122 were wrapped in pre-baked (450 °C, 8 h) aluminum foil and stored at -20 °C until
123 analysis.

124 **2.2 Chemical Analysis**

125 A portion of the filter sample was ultrasonically extracted with 3 mL of methanol
126 for 20 min, repeated three times. The extracts were then filtered with a 0.22 µm PTFE
127 pore syringe filter and concentrated to near dryness under a gentle stream of nitrogen.
128 The residue was re dissolved in 70 µL of methanol, transferred into a 1 mL centrifuge
129 tube, and centrifuged at 12000 rpm for 20 min. The supernatant was subsequently
130 transferred into LC vials and analyzed within 24 h. The extracts were analyzed using a
131 Dionex Ultimate 3000 UHPLC (Thermo Fisher Scientific, Waltham, US) coupled to an
132 Orbitrap Exploris 120 mass spectrometer (Thermo Fisher Scientific, Waltham, US). A
133 HSS T3 column (2.1 × 100 mm, 1.8 µm particle size) was used to separate the organics
134 in the sample extracts. Details of the elution program are provided in the Supporting
135 Information. The injection volume is 5 µL. The electrospray ionization source was
136 operated in negative mode (ESI-) with a sheath gas flow of 35.0 L min⁻¹, capillary
137 temperature of 320 °C, spray voltage of -3.2 kV, and mass resolution of 140,000. Full-
138 scan data were acquired over the *m/z* range of 90–500.

139 **2.3 Data Processing**

140 The MS data were processed using Xcalibur 4.0 (Thermo Scientific) and the open-
141 source software MZmine (<http://mzmine.github.io/>) to extract *m/z* values, molecular
142 formulas, retention times, and peak areas of detected organic compounds. The



143 molecular formula assignment was mainly based on detected m/z values with a mass
144 tolerance of ± 2 ppm, with elemental constraints set to $0.3 \leq H/C \leq 3.0$, $O/C \leq 3.0$, N/C
145 ≤ 0.5 , $S/C \leq 0.2$. The double bond equivalent (DBE) of the assigned formulas
146 (C_cH_hO_oN_nS_s) was calculated as:

147
$$DBE = (2C + 2 - H + N) / 2$$

148 where C, H, and N represent the number of carbon, hydrogen, and nitrogen atoms,
149 respectively.

150 Potential OSs and nitrooxy-OSs were first screened based on elemental
151 composition and oxygen content, and structural assignment was supported by
152 diagnostic sulfate-related fragment ions, including HSO₄⁻ at m/z 96.9599, SO₄⁻ at m/z
153 95.9522, HSO₃⁻ at m/z 80.9651, and SO₃⁻ at m/z 79.9573, as well as characteristic
154 neutral losses such as H₂SO₄. For nitrooxy-OSs, the additional presence of nitrogen-
155 containing fragment ions (e.g., NO₂⁻), together with sulfate-related fragments, provided
156 supporting evidence for nitrogen-containing functional groups. Compound identities
157 were assigned based on diagnostic spectral features and comparison with previously
158 reported compounds in the literature. Because authentic standards were unavailable for
159 most species, the reported abundances are based on relative signal intensities and are
160 used primarily for comparative analysis across seasons and OS classes, assuming
161 comparable ionization efficiencies among structurally related compounds.

162 2.4 Additional Measurements

163 Water-soluble inorganic ions were quantified by an ion chromatograph (IC,
164 DIONEX, ICS2500/ICS2000); NO₂ was monitored by a NO_x analyzer (Thermo Fisher
165 Scientific, model 42i). Meteorological parameters (RH and temperature) were provided
166 by a weather station (Met one Instrument Inc.). In addition, the ALWC and aerosol pH
167 were estimated using the thermodynamic model (ISORROPIA-II) with measured
168 inorganic aerosol composition (NH₄⁺, SO₄²⁻, NO₃⁻, and Cl⁻), ambient temperature, and
169 RH as input parameters (Fountoukis et al. 2007). The model was operated in metastable
170 mode, assuming fully deliquescent particles without solid phases.

171 2.5 Backward-Trajectory and Cluster Analysis

172 Air-mass transport pathways were investigated using the National Oceanic and
173 Atmospheric Administration (NOAA) Hybrid Single-Particle Lagrangian Integrated
174 Trajectory model (http://ready.arl.noaa.gov/HYSPLIT_traj.php). 72-h backward
175 trajectories arriving at the Wuyi Mountain sampling site were calculated using the
176 archived meteorological data obtained from the National Center for Environmental



177 Prediction's Global Data Assimilation System (GDAS). The resulting trajectories were
178 grouped using HYSPLIT trajectory cluster analysis, which classifies trajectories
179 according to similarities in their spatial transport pathways. For each season, three
180 trajectory clusters were retained, and the percentage assigned to each cluster represents
181 the fraction of all calculated trajectories belonging to that cluster.

182 **3. Results and discussion**

183 **3.1 Molecular Overview of Organic Sulfates**

184 Fig. 1a summarizes the number of molecular formulas and signal abundance of
185 organic compounds detected in PM_{2.5} samples collected at Wuyi Mountain. A total of
186 2,283 and 3,253 molecular formulas were identified in summer and winter, respectively.
187 Sulfur-containing compounds, mainly CHOS and CHONS species, accounted for more
188 than 30% of all the detected formulas in both seasons, highlighting the importance of
189 sulfur-containing organic compounds in mountain aerosol. Based on molecular formula
190 counts, approximately 90% of CHOS and 30% of CHONS formulas satisfied the
191 oxygen criterion required to accommodate the assignment of $-\text{OSO}_3\text{H}$ and $-\text{ONO}_2$
192 groups ($\text{O}/(4\text{S} + 3\text{N}) \geq 1$) (Fig. 1b, Fig. S1), suggesting that they were potential OSs or
193 nitrooxy-OSs (Wang et al. 2016; Cai et al. 2020). In terms of signal abundance, potential
194 OSs accounted for comparable fractions of total CHOS signals in summer and winter
195 ($> 92\%$), whereas nitrooxy-OSs contributed a larger fraction of CHONS signals in
196 winter ($\sim 72\%$) than in summer ($\sim 61\%$). The total signal abundances of potential OSs
197 and nitrooxy-OSs in winter were 1.4 and 2.1 times those in summer, respectively. These
198 results indicate a greater wintertime contribution of sulfur- and nitrogen-containing
199 organic compounds, particularly species associated with nitrogen-involved atmospheric
200 processing.

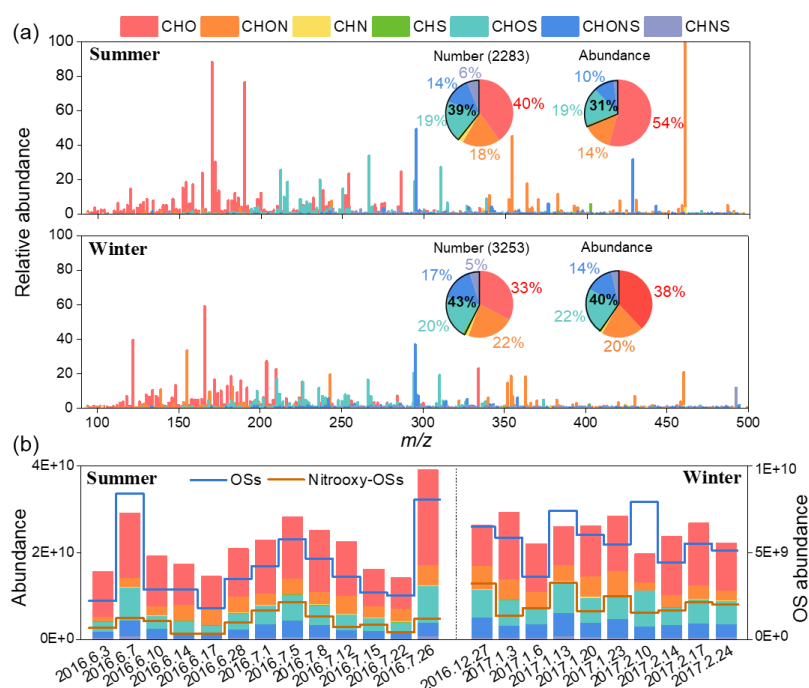


Figure 1. The number and abundance of detected organic compounds in summer and winter at Wuyi Mountain, with relative contributions of different elemental classes based on molecular number and signal abundance (a). Temporal variation in the abundance of the different elemental classes, potential OSs and nitrooxy-OSs (b).

Fig. 2a shows the molecular distributions of OSs and nitrooxy-OSs as functions of m/z , carbon number, and oxygen number. In both seasons, OS molecular mass generally increased with carbon number. However, clear seasonal differences were observed in molecular compositional space. Summer OSs were mainly distributed within the low-to-medium carbon-number range (C_5 – C_{12}), whereas winter OSs extended to substantially higher carbon numbers, reaching C_{23} . As for oxygen-number distribution, summer OSs covered a wider range (O_4 – O_{11}) and included a larger proportion of highly oxygenated species, whereas winter OSs were mainly concentrated between O_4 and O_7 . The combination of higher carbon numbers and lower oxygen contents indicates that winter OSs occupied a distinct region of molecular compositional space characterized by larger and less oxygenated molecules. Nitrooxy-OSs exhibited comparatively narrower molecular distribution in both seasons, with carbon numbers mainly ranging from C_5 to C_{15} and oxygen numbers from O_7 to O_{11} . Compared with summer, winter nitrooxy-OSs displayed higher signal abundance and a more concentrated



220 compositional distribution, suggesting enhanced formation and/or accumulation under
221 winter conditions.

222 The Van Krevelen (VK) diagrams provide further information on the structural
223 characteristics of OSs and nitrooxy-OSs (Fig. 2b). OSs occupied a broad region in H/C
224 and (O–3S)/C space, indicating substantial compositional variability and the
225 coexistence of aliphatic and unsaturated structures (Cai et al. 2020; Lin et al. 2012). In
226 both seasons, OS signal abundance was dominated by compounds with H/C ratios
227 greater than 1.0 and DBE values lower than 5, which accounted for 98% and 95% of
228 the total OS signals in summer and winter, respectively. These features are consistent
229 with sulfate esterification of relatively saturated or weakly unsaturated oxidation
230 products, including those associated with biogenic precursors such as isoprene and
231 monoterpenes (Thoma et al. 2025). Compared with summer, winter OSs showed a
232 modest shift toward lower H/C and higher DBE values, indicating a greater contribution
233 from more unsaturated molecular features (Mazzoleni et al. 2012). Nitrooxy-OSs
234 occupied a narrower VK space and exhibited substantially higher signal abundances in
235 winter, further supporting their formation under nitrogen-involved oxidation conditions.
236 Moreover, OSs spanned a broad range of carbon oxidation state ($\overline{\text{OS}}_c$) from -2 to 2,
237 reflecting substantial variability in molecular oxidation levels (Fig. S2). Highly
238 oxidized OSs ($\overline{\text{OS}}_c > -1$) accounted for 80% of the molecular number and 72% of the
239 signal abundance in summer, compared with 64% and 56%, respectively, in winter. This
240 result is consistent with a higher degree of molecular oxidation and active sulfate
241 esterification of biogenic oxidation products (Kroll et al. 2011). Nitrooxy-OSs exhibited
242 moderate-to-high carbon oxidation states, with $\overline{\text{OS}}_c$ values mainly between -1 and 2,
243 and showed higher molecular counts and signal abundances in winter than in summer.

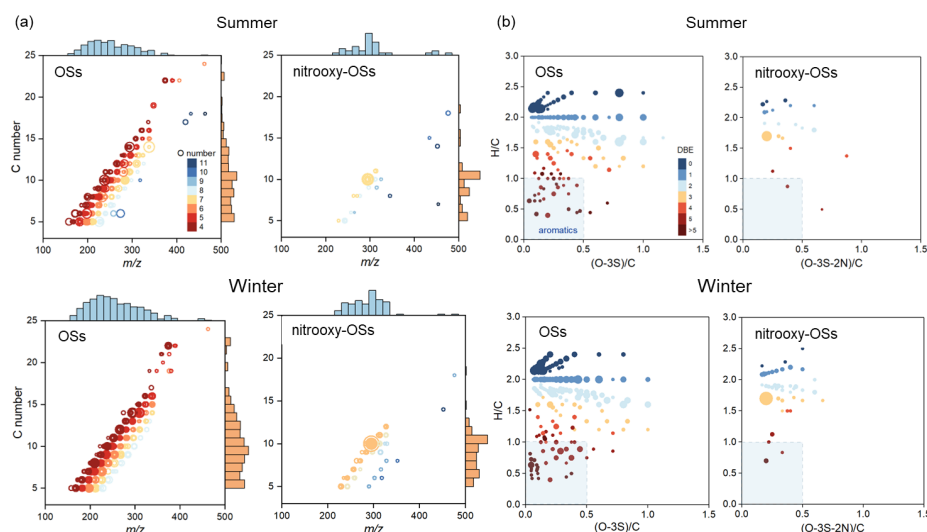


Figure 2. Distributions of molecular mass, carbon number, oxygen atom number (a) and Van Krevelen diagrams for OSs and nitrooxy-OSs in summer and winter.

3.2 Molecular Identification and Seasonal Variation

Based on HPLC-Orbitrap MS fragmentation patterns, 35 compounds with diagnostic sulfate ester features were structurally assigned (Fig. 3a, Table S1). These compounds were grouped into five classes according to molecular characteristics and likely precursor associations: OSs containing two or three carbon atoms (C_2/C_3 OS), isoprene-derived OSs (OSi), monoterpene-derived OSs (OSm), monoterpene-derived nitrooxy-OSs (NOSm), and aliphatic-like OSs (OSali). OSi was characterized by ions at m/z 198.9913 ($C_4H_7O_7S^-$) and m/z 197.0120 ($C_5H_9O_7S^-$), which are commonly attributed to products formed via reactive uptake and subsequent sulfate esterification of isoprene oxidation intermediates (Cai et al. 2020). OSm was mainly represented by m/z 249.0791 ($C_{10}H_{17}O_5S^-$), a species linked to monoterpene oxidation products such as those derived from limonene. The NOSm group was represented by m/z 294.0646 ($C_{10}H_{16}NO_7S^-$), which has been associated with monoterpene oxidation under NO_x -rich conditions or nighttime NO_3 -radical chemistry, followed by incorporation of nitrate- and sulfate-containing functionalities (Wang et al. 2020; Wang et al. 2021b). For the C_2/C_3 OS group, the most abundant ions included m/z 152.9860 ($C_3H_5O_5S^-$), m/z 154.9651 ($C_2H_3O_6S^-$), and m/z 168.9808 ($C_3H_5O_6S^-$), corresponding to hydroxyacetone sulfate (HAS), glycolic acid sulfate (GAS), and lactic acid sulfate (LAS), respectively. These low-carbon-number OSs are generally considered products of highly oxidized biogenic and anthropogenic precursors, including carboxylic acids and hydroxyl-



ketone, and their formation is often associated with aqueous-phase or multiphase reactions (Wang et al. 2023). OSali species were characterized by relatively high H/C ratios and low degrees of unsaturation, consistent with aliphatic-like molecular structures. They may originate from the oxidation of saturated or weakly functionalized organic precursors followed by sulfate esterification, although their exact sources remain uncertain. The molecular composition of OSali also differed between seasons: a species with m/z 209.0850 ($C_8H_{17}O_4S^-$) was more prominent in summer, whereas a species with m/z 153.0224 ($C_4H_9O_4S^-$) dominated in winter.

The relative contributions of the five OS groups varied substantially between seasons (Fig. 3b). OSi and OSm showed higher relative abundances in summer than in winter, consistent with enhanced emissions of isoprene and monoterpene under warmer and more irradiated conditions. The sulfur oxidation ratio (SOR, calculation shown in the Text S3) was higher in summer (0.83 ± 0.08) than in winter (0.42 ± 0.23 ; Fig. 3c, Fig. S3), indicating more efficient conversion of SO_2 to sulfate during summer despite lower anthropogenic SO_2 emissions (Duan et al. 2026). The resulting increase in sulfate levels likely enhanced sulfate-related transformations and multiphase processing. C_2/C_3 OS were strongly correlated with OSi in both seasons ($R^2=0.90$, Fig. S4), suggesting related precursor sources and/or similar atmospheric processing. However, the ratios of C_2/C_3 OS to OSi were approximately twice as higher in winter as in summer, indicating a greater relative contribution of C_2/C_3 OS under winter conditions. In contrast, NOSm exhibited higher abundances in winter and contributed substantially to total OS signals in several winter samples (Fig. S5). This enhancement coincided with higher NO_2 and particulate nitrate (NO_3^-) concentrations, as well as the elevated nitrogen oxidation rates in winter (NOR, average 0.54 ± 0.10) than in summer (0.01 ± 0.00), indicating a more nitrogen-rich secondary aerosol environment. Lower winter temperatures may further favor the partitioning and persistence of nitrate-containing species in the particle phase. Together, these observations suggest that NO_x -related oxidation and subsequent multiphase processing contributed more strongly to the formation and accumulation of nitrooxy-OSs during winter.

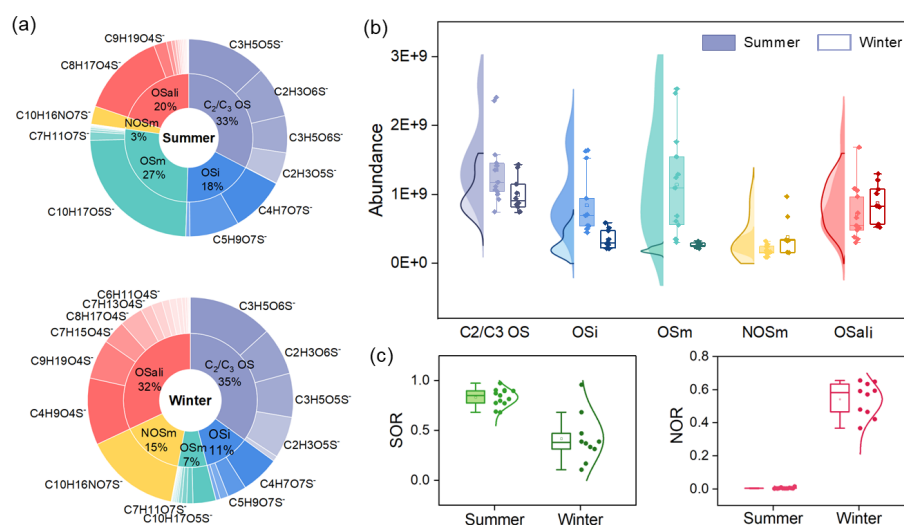


Figure 3. Seasonal variations in the chemical composition (a) and abundance (b) of five group OSs, together with SOR and NOR.

3.3 Possible Formation Processes and Transport Influence

Seasonal variations in atmospheric composition and meteorological conditions provide useful constraints on the processes associated with OS formation (Fig. 4). Difference in sulfate availability, aerosol acidity, oxidant levels, and ALWC indicated distinct atmospheric environments in summer and winter. Summer aerosols were more acidic and sulfate-rich, with an average pH of 2.8 ± 0.7 and SO_4^{2-} concentration of $3.6 \pm 2.8 \mu\text{g m}^{-3}$, compared with 3.7 ± 0.7 and $1.58 \pm 0.8 \mu\text{g m}^{-3}$, respectively, in winter (Fig. S6). Greater aerosol acidity during the warm season has also been reported in previous studies, where temperature-dependent ammonia partitioning and reduced particle neutralization contributed to the seasonal variation in aerosol pH (Ding et al. 2019; Ren et al. 2021). Although O_x ($\text{NO}_2 + \text{O}_3$) concentrations were lower in summer ($75.0 \pm 16.8 \mu\text{g m}^{-3}$), the higher temperatures ($26.3 \pm 2.2 \text{ }^\circ\text{C}$) likely favored photochemical oxidation and the production of oxygenated biogenic intermediates. Most OS groups (except NOSm) closely tracked SO_4^{2-} concentrations and showed inverse relationships with aerosol pH (Fig. 4a), indicating preferential occurrence under sulfate-rich and acidic conditions. C_2/C_3 OSs exhibited a particularly strong positive correlation with SO_4^{2-} ($r = 0.8$, $p < 0.001$; Fig. 4b, Fig. S7), consistent with sulfate-



316 associated aqueous-phase or multiphase processing. Previous studies have proposed
317 that oxygenated precursors containing adjacent carbonyl and hydroxyl groups can
318 facilitate the formation of reactive intermediates (e.g., cyclic sulfate structures), which
319 readily react with sulfur trioxide or sulfuric acid to form low-carbon-number OSs
320 (Wang et al. 2023). OSi also showed positive associations with SO_4^{2-} and acidity ($r >$
321 0.7 , $p < 0.01$), consistent with the uptake and acid-catalyzed conversion of isoprene
322 oxidation products on sulfate-containing particles (Piletic et al. 2013; He et al. 2018).
323 These observations suggest that sulfate availability and aerosol acidity may play an
324 important role in shaping OS distribution by influencing environments (Nguyen et al.
325 2014; Xu et al. 2015). In comparison, OSm and OSali exhibited weaker relationships
326 on SO_4^{2-} and acidity, implying that their variability may additionally depend on
327 precursor-specific processes, including heterogeneous reactions of SO_2 with
328 monoterpene-derived SOA or relatively reduced organic compounds (Yao et al. 2019;
329 Zhu et al. 2019).

330 Summer was also characterized by high RH ($82.5 \pm 7.7\%$), and approximately 80%
331 of the samples had ALWC-to-dry $\text{PM}_{2.5}$ mass ratios greater than 0.3 (Fig. S8),
332 suggesting that particles were predominantly in liquid or near-liquid states (Liu et al.
333 2023). Such conditions can enhance molecular diffusion and the uptake of soluble
334 gaseous precursors, thereby facilitating aqueous-phase and heterogeneous reactions
335 (Liu et al. 2017; Wu et al. 2018; Liu et al. 2019). However, the weak correlations
336 between OS abundance and ALWC indicate that particle-phase water was not the
337 primary factor governing the temporal variability of OSs. Instead, ALWC likely
338 provided a reaction medium whose effect depended on the simultaneous availability of
339 organic precursors, sulfate, and acidity (Ervens et al. 2011). The negative correlations
340 between several OSs and RH ($r < -0.7$, Fig. S9) further indicate that RH mainly reflects
341 meteorological variability (e.g., cloud/fog influence). This pattern suggests that,
342 although elevated RH promotes multiphase processing, this effect may diminish under
343 persistently high-humidity conditions approaching saturation. Similar behavior has
344 been reported in urban regions such as Shanghai (Wang et al. 2021b), where higher RH



345 was associated with increased ALWC and elevated aerosol pH due to enhanced
346 neutralization, thereby suppressing acid-catalyzed OS formation (Riva et al. 2016;
347 Duporté et al. 2016). In addition, enhanced ALWC may also introduce additional
348 physical effects, including increased aqueous partitioning, dilution, and changes in
349 particle phase state, which can reduce the apparent particulate concentrations of OSs
350 (Guo et al. 2015; Shiraiwa et al. 2017).

351 Winter samples exhibited higher O_x concentrations ($95.9 \pm 20.9 \mu\text{g m}^{-3}$) despite
352 lower temperatures ($10.0 \pm 3.8^\circ\text{C}$), reflecting an atmospheric environment with elevated
353 NO_2 and O_3 . NOSm showed a positive relationship with O_x , particularly at higher
354 organic aerosol loadings ($r = 0.75$), suggesting that oxidant availability and the amount
355 of particle-phase organic material jointly affected its abundance. The coexistence of
356 NO_2 and O_3 provides conditions favorable for nighttime NO_3 -radical formation, which
357 can initiate monoterpene oxidation and generate organic nitrate and peroxy
358 intermediates. Subsequent partitioning and multiphase reactions of these products may
359 contribute to nitrooxy-OS formation (Jiang et al. 2022; Wang et al. 2018). The absence
360 of clear relationships between NOSm and SO_4^{2-} , aerosol acidity, or RH (Fig. S7) further
361 suggests that its wintertime variability was more closely associated with NO_x -related
362 oxidation conditions than with bulk sulfate or particle-water availability. OSali also
363 exhibited higher abundances in winter showed a different pattern from NOSm. Its
364 positive associations with SO_2 and NH_4^+ ($r \sim 0.6$) suggest an influence from
365 anthropogenic precursor inputs and anthropogenically influenced secondary aerosol
366 processing. Similar increases in the anthropogenic contribution to OSs have been
367 observed during winter in Tianjin (Ding et al. 2022). OSali may form through the
368 oxidation of relatively reduced organic precursors followed by sulfate esterification and
369 particle-phase processing. Lower temperatures, enhanced regional pollution, and
370 increased partitioning of low-volatility products may further favor their wintertime
371 accumulation.

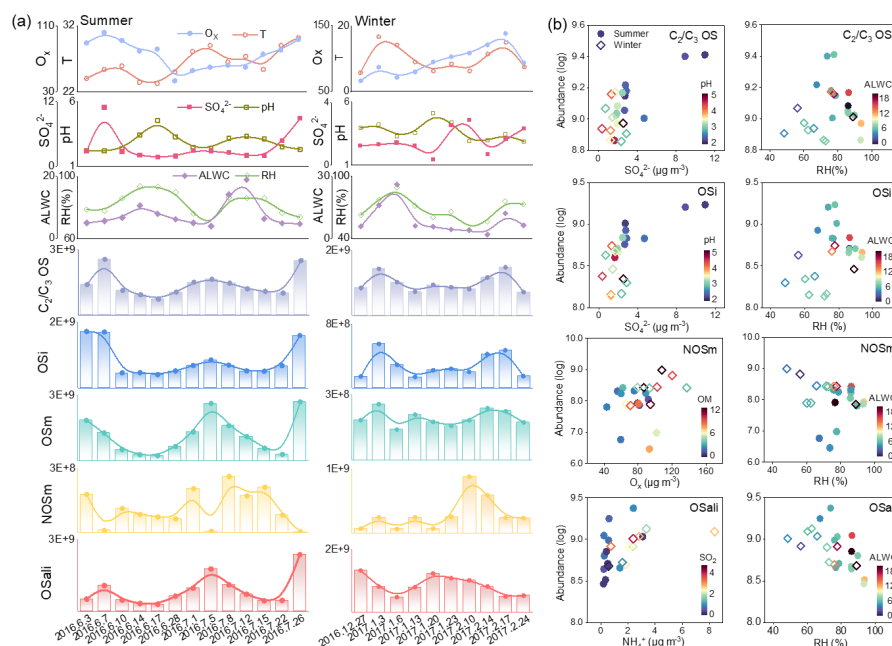


Figure 4. Temporal variations of O₃, T, SO₄²⁻, pH, ALWC, RH and OSs (a), as well as the relationship between them during summer and winter (b).

To assess the influence of regional transport, 72 h-backward trajectories were analyzed (Fig. 5). In summer, 68 % of the trajectories were associated with air masses transported within Fujian Province, while a further 14 % originated over the South China Sea. The predominance of regional air masses from southeastern China, together with strong summertime biogenic emissions, is consistent with the high relative contributions of isoprene- and monoterpene-derived OSs and the importance of in situ photochemical and multiphase processing. Marine air masses may also modify aerosol water content and composition (Hodshire et al. 2019; Moore et al. 2024). However, chloride contributed only marginally to PM_{2.5} mass even for these ocean-influenced air masses, suggesting that the contribution of marine-derived chloride to PM_{2.5} was limited during the sampling periods. Winter air masses showed a greater influence from continental transport. Approximately 57% of trajectories passed over Jiangxi and Hunan provinces, while 27% arrived from northwestern inland regions. These pathways indicate increased exposure to anthropogenically influenced continental air masses



before reaching Wuyi Mountain. During transport, oxidation, gas–particle partitioning, and multiphase aging may alter both precursor composition and the molecular characteristics of particulate OSs. This is consistent with the higher wintertime abundances of NOSm and OSali and their associations with nitrogen-rich and anthropogenically influenced aerosol conditions. These results suggest that OS composition at this mountain site reflects the combined influences of regional transport and in situ atmospheric processing, with regional transport playing a more prominent role in winter by modifying precursor availability and aerosol chemical environments.

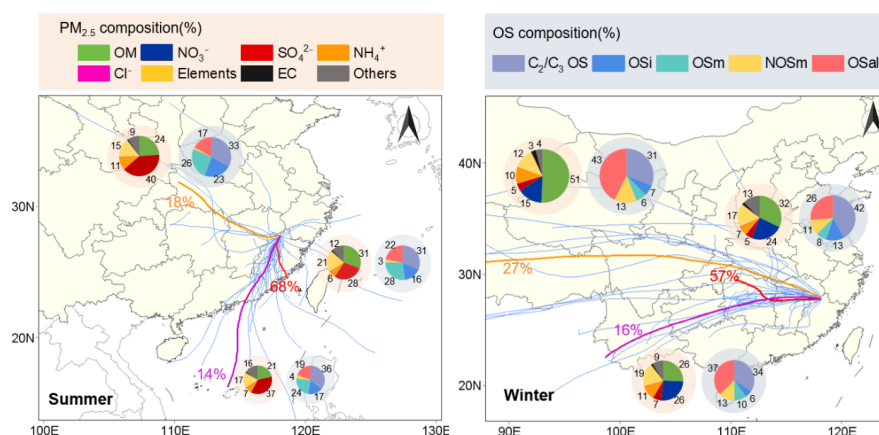


Figure 5. Seasonal 72 h-backward trajectory clusters, the associated PM_{2.5} and OSs composition at the Wuyi Mountain site. Percentages indicate the fraction of trajectories in each cluster, and bold lines represent the corresponding mean trajectories. For each cluster, the left and right pie charts show PM_{2.5} composition (OM, NO₃⁻, SO₄²⁻, NH₄⁺, Cl⁻, elements, EC, and others) and OS composition (C₂/C₃ OS, OSi, OSm, NOSm, and OSali), respectively. Pie chart sizes are proportional to the corresponding PM_{2.5} mass concentration and the total OSs abundance, respectively.

4. Conclusions

This study investigated the molecular composition, seasonal variability, and possible formation processes of PM_{2.5}-bound OSs at a high-altitude background site on Wuyi Mountain using HPLC-Orbitrap MS. Among the detected sulfur-containing compounds, approximately 90% of CHOS and 30% of CHONS formulas met the compositional criteria for potential OSs and nitrooxy-OSs, respectively. Their total signal abundances in winter were 1.4 and 2.1 times those in summer. Nevertheless, summer OSs exhibited a higher overall degree of molecular oxidation, with species



412 having $\overline{OS}_c > -1$ accounting for 72% of the total OS signal abundance compared with 56%
 413 in winter. Winter OSs were characterized by larger (C_5 – C_{23}), less oxygenated (O_4 – O_7),
 414 and more unsaturated molecules. Among the 35 structurally assigned OSs, isoprene-
 415 and monoterpene-derived species were relatively more abundant in summer, whereas
 416 monoterpene-derived nitrooxy-OSs and aliphatic-like OSs contributed more strongly
 417 in winter. The seasonal differences were associated with changes in both precursor
 418 availability and atmospheric processing. In summer, higher sulfate concentrations
 419 ($3.6 \pm 2.8 \mu\text{g m}^{-3}$) and greater aerosol acidity (2.8 ± 0.7) were closely linked to the
 420 abundance of C_2/C_3 OSs and isoprene-derived OSs ($|r| > 0.7$, $p < 0.01$), supporting an
 421 important role for sulfate-related multiphase processing. Although particle water was
 422 abundant, its weak correlations with OS abundance suggest that it primarily provided a
 423 reaction medium rather than directly governing OS variability. In winter, higher O_x
 424 ($95.9 \pm 20.9 \mu\text{g m}^{-3}$) and a more nitrogen-rich aerosol environment were associated
 425 with the enhanced contribution of nitrooxy-OSs, while the increase in aliphatic-like
 426 OSs was more closely related to anthropogenic precursor inputs and secondary aerosol
 427 processing. Backward-trajectory analysis further indicated a stronger influence of
 428 regional transport in winter, which likely altered precursor composition and promoted
 429 atmospheric aging. Our results suggest that seasonal OS composition in mountain
 430 environments reflects shifts in the relative importance of multiple coexisting precursor
 431 sources and atmospheric processes. These findings highlight how interactions among
 432 biogenic and anthropogenic precursor inputs, inorganic aerosol composition,
 433 atmospheric oxidation, multiphase processing, and regional transport regulate
 434 secondary organic aerosol composition and evolution in mountain environments.
 435 Further quantitative and experimental studies, including direct measurements of
 436 precursor VOCs and reactive oxidants, are needed to better constrain the seasonal
 437 contributions of individual OS formation processes.

438

439 **ASSOCIATED CONTENT**

440 **Data availability.** Raw data used in this study can be obtained from
 441 <https://doi.org/10.5281/zenodo.22119962> (Wang et al., 2026). They are also available



442 on request by contacting the corresponding authors.

443

444 ***Supplement.*** The supplement related to this article is available online at:

445

446 ***Competing interests.*** The contact author has declared that none of the authors has any
447 competing interests.

448

449 ***Author contributions.*** YWH and WX designed the study. Data analysis was done by
450 TW and YWH. TW and WX interpreted data, prepared the display items, and wrote the
451 manuscript. All authors commented on and discussed the manuscript.

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