



1 **Flux and Radiocarbon Evidence of Urban Carbon Emission Reductions under**
2 **Climate Mitigation Policies**

3
4 Dan Liang,^{1,2,3,†} Hongxing Zhang,^{4,†} Zhenchuan Niu,^{1,5,6,7,*} Sen Wang,^{2,8} Weijian
5 Zhou,^{1,6} Xuxiang Li,⁵ Jingjing Jia,⁹ Yifan Tao,⁹ Guowei Wang,^{1,3} Xue Feng,^{1,3} Wanyu
6 Liu,^{1,10} Zhao Jin,¹ Yongming Han,¹ Xuefeng Lu,^{1,6} Jan B. C. Pettersson¹¹ and Xiangrui
7 Kong¹¹

8 ¹ State Key Laboratory of Loess Science, Institute of Earth Environment, Chinese Academy of
9 Sciences, Xi'an 710061, China

10 ² College of Urban and Environmental Sciences, Northwest University, Xi'an 710127, China

11 ³ Xi'an Institute for Innovative Earth Environment Research, Xi'an, China

12 ⁴ State Key Laboratory of Regional and Urban Ecology, Research Center for Eco-Environmental
13 Sciences, Chinese Academy of Sciences, Beijing 100085, China

14 ⁵ Institute of Global Environmental Change, Xi'an Jiaotong University, Xi'an, 710049, China

15 ⁶ Shaanxi Provincial Key Laboratory of Accelerator Mass Spectrometry Technology and Application,
16 Xi'an AMS Center, Xi'an 710061, China

17 ⁷ National Observation and Research Station of Regional Ecological Environment Change and
18 Comprehensive Management in the Guanzhong Plain, Shaanxi, China

19 ⁸ Shaanxi Xi'an Urban Ecosystem National Observation and Research Station, National Forestry and
20 Grassland Administration, Xi'an 710127, China

21 ⁹ State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry, Institute
22 of Atmospheric Physics, Chinese Academy of Sciences, Beijing, 100029, China

23 ¹⁰ Interdisciplinary Research Center of Earth Science Frontier, Beijing Normal University, Beijing,
24 100875, China

25 ¹¹ Department of Chemistry and Molecular Biology, University of Gothenburg, Gothenburg 41390,
26 Sweden

27 *Correspondence: Zhenchuan Niu (niuzc@ieecas.cn)

28 † Dan Liang and Hongxing Zhang contributed equally to this work.



29 **Abstract.** Urban areas account for the majority of global CO₂ emissions, yet direct
30 observational evidence of emission reductions under climate mitigation policies
31 remains limited. Here we combine eddy covariance measurements with tree-ring
32 radiocarbon (¹⁴C) to quantify changes in urban CO₂ emissions and attribute their
33 sources in two megacities, Beijing and Xi'an, from 2020 to 2024. We find that all
34 urban sites remained net carbon sources but exhibited consistent interannual declines
35 in CO₂ fluxes, with the strongest reductions during heating season and traffic peak
36 hours. Compared to 2021, the CO₂ fluxes in Beijing and Xi'an in 2024 reduced by
37 21.6% (1.1 kg C m⁻² yr⁻¹) and 41.2% (0.7 kg C m⁻² yr⁻¹), respectively. Independent
38 constraints from tree-ring ¹⁴C reveal a concurrent decrease in fossil fuel-derived CO₂
39 (CO_{2ff}), demonstrating that the observed emission reductions are primarily driven by
40 declining fossil fuel combustion. Analysis of energy consumption data indicates that
41 these reductions are associated with substitution of coal and petroleum with electricity
42 and natural gas, as well as rapid adoption of new energy vehicles. Our results provide
43 direct observational evidence that urban carbon emissions can decline under climate
44 mitigation policies and highlight the central role of energy system transformation in
45 achieving sustained emission reductions.

46



47 **1 Introduction**

48 The rapid rise in atmospheric carbon dioxide (CO₂) concentrations in recent decades,
49 driven by large-scale fossil fuel use and land-use changes, has disrupted Earth's
50 energy balance and accelerated global warming (Solomon et al., 2009). Because more
51 than 70% of global CO₂ emissions originate from urban areas, cities are central to
52 climate mitigation efforts (Canadell et al., 2009). Among major emitters, China plays
53 a pivotal role due to its scale and rapid energy transition (McGlade and Ekins, 2015).
54 In 2020, China launched the “dual-carbon” strategy, committing to peak carbon
55 emissions before 2030 and achieve carbon neutrality by 2060 (MFA, PRC, 2020). The
56 policy framework provides a unique opportunity to examine how urban emission
57 patterns evolve under ambitious mitigation targets and what factors drive these
58 changes.

59 Previous studies in Chinese cities have explored pathways for reducing urban
60 carbon emission, highlighting measures such as stricter environmental regulations and
61 optimization of urban energy structures based on energy statistics and model-based
62 analysis (Hao et al., 2022; Huang et al., 2022; Wang et al., 2024b; Duan and Zhu,
63 2025). However, assessing the real-world effectiveness of these policies requires
64 direct observational evidences, particularly to disentangle the contributions of fossil
65 fuel and biogenic sources to urban CO₂ emissions.

66 The eddy covariance (EC) technique is a widely used *in-situ* approach for
67 quantifying CO₂ exchange between the surface and atmosphere (Baldocchi, 2003; Yu
68 et al., 2006), and has been extensively applied across diverse ecosystems, including
69 urban environments (Hirata et al., 2008; Yu et al., 2008; Wen et al., 2010; Ward et al.,
70 2013; Xiao et al., 2014; Ueyama and Ando, 2016; Wei et al., 2021; Xueref-Remy et
71 al., 2023; Wang et al., 2024a; Zhu et al., 2024). EC measurements provide direct



insights into the magnitude, temporal variability, and interannual trends of urban CO₂ fluxes, revealing whether cities act as net carbon sources or sinks (Ward et al., 2013; Xueref-Remy et al., 2023). However, EC alone cannot distinguish the relative contributions of fossil fuel and biological sources to these fluxes. Therefore, an independent tracer is required to quantify fossil fuel CO₂ (CO_{2ff}) and attribute observed change in urban carbon emissions.

Radiocarbon (¹⁴C) is a robust tracer for quantifying CO_{2ff}. Given that the formation time of fossil fuels far exceeds the half-life of ¹⁴C (5730 ± 40 years), fossil-derived CO₂ is effectively depleted in ¹⁴C (Godwin et al., 1962), in contrast to biogenic CO₂ that reflects contemporary atmospheric levels (Turnbull et al., 2016). Trees assimilate atmospheric CO₂ through photosynthesis and preserve its Δ¹⁴C signature in annual growth rings (Rakowski, 2011). Tree-ring Δ¹⁴C time series therefore enable reconstruction of atmospheric CO_{2ff} concentrations and provide an independent constraint on fossil fuel contributions to urban carbon emissions.

Beijing and Xi'an, two megacities with populations exceeding ten million in northern and western China, respectively, are undergoing rapid energy system transformations under the “dual-carbon” strategy. These cities provide representative cases for evaluating how urban CO₂ emissions respond to large-scale mitigation policies and for identifying the underlying driving mechanisms. In this study, we combined EC and tree-ring ¹⁴C measurements to investigate urban carbon emissions in Beijing and Xi'an during 2020–2024. Specifically, we aimed to (i) quantify the temporal dynamics of CO₂ emissions under the “dual-carbon” strategy and identify key periods of emission reduction; (ii) constrain fossil fuel contributions using CO_{2ff} derived from tree-ring ¹⁴C; and (iii) identify the dominant drivers of emission changes by integrating energy consumption data. Our results provide observation-based



97 evidence for evaluating urban carbon mitigation strategies and offer insights
98 applicable to other rapidly developing cities worldwide.

99

100 **2 Methods**

101 **2.1 Site description**

102 Beijing, the capital of China and the core city of the Beijing-Tianjin-Hebei region, is
103 located in northern North China Plain, bordered by the Yanshan Mountains to the
104 north and the Taihang Mountains to the west. It has a warm temperate, semi-humid
105 continental monsoon climate with four distinct seasons. The mean annual temperature
106 ranges from 11 to 13 °C, and the mean annual precipitation is approximately 500–
107 600 mm based on long-term meteorological records (<https://www.beijing.gov.cn/>).

108 Xi'an, the core city of Northwest China, is situated in the Guanzhong Basin at the
109 southern margin of the middle reaches of the Yellow River, bordered by the Qinling
110 Mountains to the south and the Loess Plateau to the north. It has a warm temperate,
111 semi-humid continental monsoon climate with pronounced seasonal variability. The
112 mean annual temperature ranges from 13.5 to 16.6°C, and the mean annual
113 precipitation is 506–1007 mm based on climatological statistics from 1991 to 2024
114 (<https://www.xa.gov.cn/>).

115 One CO₂ flux observation site in Beijing is located on a 325 m meteorological
116 tower at the Institute of Atmospheric Physics (IAP), Chinese Academy of Sciences
117 (39.97 °N, 116.37 °E; Figure 1). This site lies in Haidian District within the North
118 Fourth Ring Road, approximately 8 km from Tian'anmen Square, representing a
119 highly urbanized core area with dense population. The surrounding land cover
120 consists of research and education facilities, residential areas, major transportation
121 corridors, and urban green spaces, providing a representative setting for urban carbon



122 flux measurements. Tree-ring samples was collected at Madian Park, located 1.8 km
123 from the IAP site. Vegetation within the flux footprint is dominated by typical urban
124 greening species in northern China, including deciduous species (*e.g.*, *Populus*
125 *tomentosa*, *Sophora japonica*, *Ginkgo biloba*, *Forsythia suspensa*) and evergreen
126 species (*e.g.*, *Pinus tabuliformis*, *Platycladus orientalis*, *Pinus bungeana*).

127 The second CO₂ flux observation site in Beijing is located on the rooftop of an
128 office building at the Research Center for Eco-Environmental Sciences (RCEES),
129 Chinese Academy of Sciences (40.01°N, 116.35°E; Figure 1). The site lies within the
130 North Fifth Ring Road and is adjacent to Tsinghua University, Peking University, and
131 other academic and research institutions, representing a high-density urban
132 environment. The surrounding land cover consists primarily of research and
133 educational buildings, residential areas, urban roads, and artificial green spaces.
134 Tree-ring samples were collected at Dongsheng Bajia Country Park, located 1.1 km
135 from the RCEES. Vegetation within the flux footprint is dominated by common urban
136 landscape species, including deciduous types (*e.g.*, *Sophora japonica*, *Ginkgo biloba*,
137 *Koelreuteria paniculata*) and evergreen species (*e.g.*, *Pinus tabuliformis*, *Pinus*
138 *bungeana*, *Sabina chinensis*).

139 One CO₂ flux observation site in Xi'an is situated at the Tang City Wall Relics Park
140 (TCWRP; 34.23°N, 108.89°E; Figure 1), within the urban core area inside the South
141 Third Ring Road. Population density exceeds 17,000 persons km⁻² within this area,
142 indicating a highly concentrated urban environment (Seventh National Population
143 Census, 2020). The surrounding land cover consists of residential areas, commercial
144 buildings, major traffic roads, and vegetation, providing a representative setting for
145 urban carbon flux measurements. Tree-ring samples were collected within the
146 TCWRP, approximately 0.1 km from the flux observation site. Vegetation within the



147 flux footprint includes typical urban greening species, with deciduous trees (*e.g.*,
148 *Amygdalus persica*, *Ginkgo biloba*, *Jasminum nudiflorum*) and evergreen species (*e.g.*,
149 *Pinus bungeana*, *Magnolia grandiflora*, *Osmanthus fragrans*, *Picea* spp., *Cedrus*
150 spp.).

151 The second CO₂ flux observation site in Xi'an is located at Guanzhong (GZ;
152 34.07°N, 108.35°E; Figure 1), within the National Observation and Research Station
153 of Regional Ecological Environment Change and Comprehensive Management in the
154 Guanzhong Plain, approximately 50 km from downtown Xi'an. The surrounding land
155 cover consists of villages, vegetation and rural roads, representing a typical rural
156 landscape. Vegetation is dominated by deciduous species (*e.g.*, *Ginkgo biloba* and
157 *Davidia involucrate*) and evergreen species (*e.g.*, *Abies* spp., *Taxus chinensis*, and
158 *Pinus tabulaeformis*), along with bamboo (*Fargesia qinlingensis*) and farmland. The
159 area is characterized by sparse settlements and extensive vegetation cover, providing a
160 representative setting for CO₂ exchange in rural ecosystem.

161 No significant land cover changes were observed in Beijing and Xi'an during
162 2021–2024, as confirmed by analysis of 30 m resolution land cover datasets (Yang et
163 al., 2024) in conjunction with Territorial Space Master Plan (2021–2035) in both
164 cities.

165

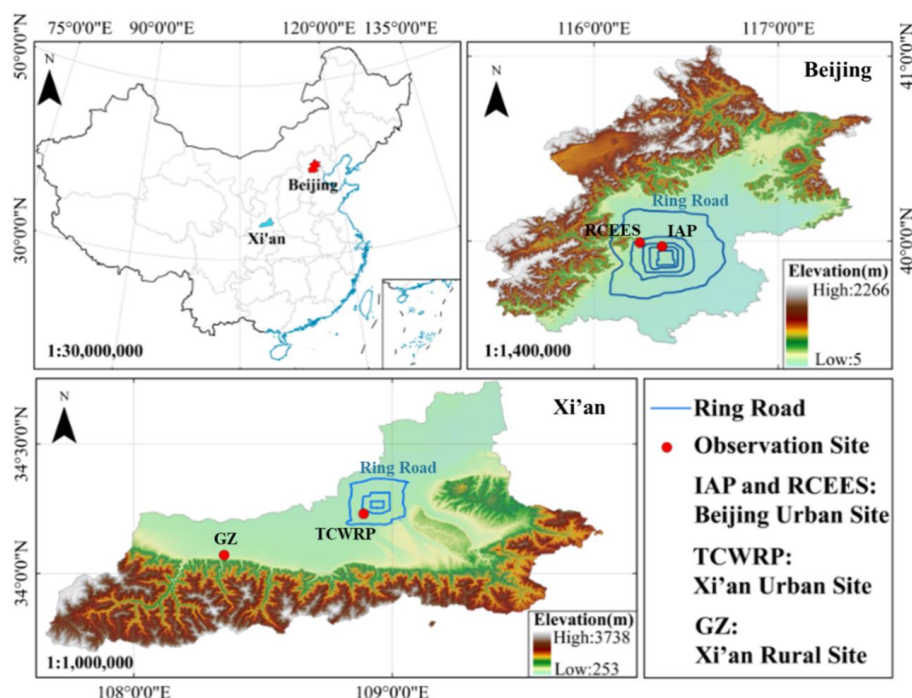


Figure 1. Locations of the observation sites in Beijing and Xi'an, China. The upper-left map shows the positions of Beijing and Xi'an within China. Red dots indicate the observation sites, and blue lines represent the ring roads in each city. IAP, RCEES, TCWRP, and GZ denote the Institute of Atmospheric Physics, the Research Center for Eco-Environmental Sciences, the Tang City Wall Relics Park, and the Guanzhong site, respectively.

166

167 2.2 Measurements

168 CO₂ flux measurements were conducted at the Beijing urban sites (from 1 January
169 2020 to 30 December 2024), the Xi'an urban site (from 21 January 2021 to 30
170 December 2024), and the Xi'an rural site (from 16 February 2022 to 30 December
171 2024). The EC system consisted of an open-path CO₂/H₂O infrared gas analyzer
172 (urban site: LI-7500A; rural site: LI-7500DS; LI-COR Inc., USA) coupled with a
173 three-dimensional ultrasonic anemometer (WindMaster, GILL Inc., UK), installed
174 facing the prevailing wind direction. The analyzers were calibrated every six months
175 using standard CO₂ gases (0 ppm and 395.46 ppm) and a dew-point generator (LI-610,



176 LI-COR Inc., USA).

177 At the IAP site, the flux instrument was installed at a height of 140 m on a 325 m
178 iron tower. At the RCEES site, the system was mounted on a 5 m tripod atop a 38 m
179 building, while at the TCWRP site it was mounted at 15 m on a monitoring tower. At
180 the rural GZ site, the instrument was mounted at 20 m on a similar tower. In all cases,
181 the EC system was positioned at more than 1.5 times the mean canopy height of the
182 surrounding vegetation and operated at a sampling frequency of 10 Hz. Signal
183 strength was routinely monitored, and maintenance was performed when it fell below
184 0.8.

185 Flux footprint characteristics were determined using the two-dimensional
186 parameterized model of Kljun et al. (2025). For the three urban sites, the dominant
187 land cover types within the 80% flux footprint were buildings, urban roads, and
188 vegetation, with corresponding fetch distances of approximately 4200 m (IAP), 314 m
189 (RCEES), and 200 m (TCWRP) (Figure S1a-c). At the GZ site, the 80% footprint was
190 dominated by villages, rural roads, and vegetation (Figure S1d).

191

192 **2.3 Data processing and gap-filling**

193 Raw EC data were occasionally lost or corrupted due to adverse weather conditions,
194 power outages, or instrument malfunctions. Pre-processing was therefore applied to
195 ensure data quality and reliability (Amar et al., 2023). The 10 Hz raw data were
196 processed using EddyPro software (LI-COR Inc., USA), including spike removal,
197 time-lag correction, coordinate rotation (double rotation), and ultrasonic virtual
198 temperature correction. The processed high-frequency data were used to compute
199 covariance fluxes, which were then aggregated into 30-minute intervals. The
200 Webb-Pearman-Leuning (WPL) density correction was applied to account for thermal



201 expansion and water vapor effects, along with frequency response correction for
202 spectral attenuation, yielding CO₂ fluxes at 30-minute resolution.

203 Quality flags (0, 1, 2) were assigned according to the data quality control method of
204 Foken et al. (2004) Half-hourly CO₂ flux data were excluded if they met any of the
205 following criteria: (1) incomplete datasets (flagged as -9999); (2) quality flag = 2; (3)
206 anomalous negative nighttime flux; (4) fluxes during rainfall events; and (5) stable
207 nighttime conditions with friction velocity (u^*) ($u^* < 0.33 \text{ m s}^{-1}$ at the IAP site; $u^* <$
208 0.16 m s^{-1} at the RCEES site; $u^* < 0.10 \text{ m s}^{-1}$ at the TCWRP site; $u^* < 0.38 \text{ m s}^{-1}$ at
209 the GZ site) (Amiro et al., 2006). Data gaps caused by sensor malfunction, power
210 failure, or quality filtering were filled using the mean diurnal variation (MDV)
211 method (Falge et al., 2001) with a 10-day moving window. All CO₂ fluxes in this
212 study were reported on a carbon mass basis (e.g., $\text{g C m}^{-2} \text{ d}^{-1}$ or $\text{mg C m}^{-2} \text{ h}^{-1}$).

213

214 **2.4 Tree-ring sample collection, preparation, and ¹⁴C measurement**

215 Tree-ring samples were collected in Beijing and Xi'an in January 2026. The sampled
216 tree species were *Pinus tabulaeformis* and *Pinus bungeana*. Core samples were
217 extracted from multiple directions using a 10 mm diameter increment borer, with 2–3
218 cores collected per tree. The cores were wrapped in clean aluminum foil, and
219 sampling information (time, location, and direction) was recorded.

220 The samples were air-dried and sanded to clearly distinguish annual rings based on
221 differences between earlywood and latewood under a microscope. Tree rings
222 corresponding to the years 2021–2024 were manually separated and subsequently
223 subjected to organic solvent extraction, bleaching, and alkalization to isolate
224 α -cellulose.

225 Approximately 3 mg of α -cellulose was mixed with excess wire-shaped copper



oxide and sealed in a 6 mm quartz tube under vacuum ($< 10^{-5}$ torr). The sealed tubes were combusted in a muffle furnace at 850 °C for 3 h. The resulting CO₂ was purified in a vacuum line using a liquid nitrogen cold trap (−196 °C) and an ethanol-liquid nitrogen trap (−90 °C). The purified CO₂ was then reduced to graphite using the Zn/Fe reduction method (Slota et al., 1987). The graphite was pressed into aluminum targets and analyzed using an accelerator mass spectrometer (0.2 MV MICADAS) at the Xi'an Accelerator Mass Spectrometry Center. The measurement precision was better than 2‰.

The ¹⁴C content of a sample is typically expressed as $\Delta^{14}\text{C}$, defined as (Stuiver and Polach, 1977):

$$\Delta^{14}\text{C} = \left[\frac{(^{14}\text{C}/^{12}\text{C})_{\text{SN}}}{(^{14}\text{C}/^{12}\text{C})_{\text{abs}}} - 1 \right] \times 1000\text{‰} \quad (1)$$

where $(^{14}\text{C}/^{12}\text{C})_{\text{SN}}$ is the isotopic ratio in the sample after correction for $\delta^{13}\text{C}$ -based fractionation and radioactive decay, and $(^{14}\text{C}/^{12}\text{C})_{\text{abs}}$ is corresponding ratio of the absolute international radiocarbon standard.

2.5 Fossil fuel CO₂ (CO_{2ff}) calculation

CO_{2ff} is typically calculated using the three-source partitioning method, in which ambient atmospheric CO₂ (CO_{2obs}) is divided into three components: background atmospheric CO₂ (CO_{2bg}), CO_{2ff}, and biological CO₂ (CO_{2bio}). Here, Δ_{obs} , Δ_{bg} , Δ_{ff} and Δ_{bio} represent the ¹⁴C composition of each component, respectively.

$$\text{CO}_{2\text{obs}} = \text{CO}_{2\text{bg}} + \text{CO}_{2\text{ff}} + \text{CO}_{2\text{bio}} \quad (2)$$

$$\Delta_{\text{obs}} \text{CO}_{2\text{obs}} = \Delta_{\text{bg}} \text{CO}_{2\text{bg}} + \Delta_{\text{ff}} \text{CO}_{2\text{ff}} + \Delta_{\text{bio}} \text{CO}_{2\text{bio}} \quad (3)$$

The CO_{2bg} concentrations were used as a proxy for CO_{2obs} concentrations, which were not directly available from the tree-ring records.

Combining equations (2) and (3), the formula for calculating CO_{2ff} is as follows:



$$\text{CO}_{2ff} = \frac{\text{CO}_{2bg}(\Delta_{obs} - \Delta_{bg})}{\Delta_{ff} - \Delta_{obs}} - \frac{\text{CO}_{2other}(\Delta_{bio} - \Delta_{obs})}{\Delta_{ff} - \Delta_{obs}} \quad (4)$$

Equation (4) can be simplified by omitting the error term on the right-hand side:

$$\text{CO}_{2ff} = \frac{\text{CO}_{2bg}(\Delta_{obs} - \Delta_{bg})}{\Delta_{ff} - \Delta_{obs}} \quad (5)$$

Fossil fuels contain no ^{14}C , i.e., $\Delta_{ff} = -1000\%$, since the half-life of ^{14}C is much shorter than the formation time of fossil fuels (Levin et al., 2003). The CO_{2bg} and Δ_{bg} required for calculating CO_{2ff} were obtained from the annual mean CO_2 concentration data at the Waliguan Observatory (<https://gml.noaa.gov/data/data.php?site=WLG>) and the Northern Hemisphere Zone 1 (NH-1) atmospheric background $\Delta^{14}\text{C}$ data (Hua et al., 2022), respectively. Missing data $\Delta^{14}\text{C}$ were supplemented using an annual decrease rate of 4%.

261

2.6 Bottom-up statistical aggregation of energy consumption data in Beijing and

Xi'an City

To assess the influence of anthropogenic energy consumption on CO_2 fluxes, energy consumption data for Beijing and Xi'an were compiled from official statistical sources.

For Beijing, data for 2020–2024 were obtained from the energy balance sheets reported in the *Beijing Statistical Yearbook* (2021–2025). These datasets include seven terminal energy consumption sectors: (1) agriculture, forestry, animal husbandry, and fishery; (2) industry; (3) construction; (4) transportation, storage, and postal services; (5) wholesale, retail, hospitality, and catering; (6) residential consumption; and (7) others.

For Xi'an, energy consumption data for 2021–2023 were derived from the *Shaanxi Provincial Statistical Yearbooks* and *Xi'an Statistical Yearbooks* (2022–2024). The



275 Xi'an Statistical Yearbook provides detailed information on transportation activity
276 (e.g., subway network length and passenger turnover), but includes direct energy
277 consumption data only for the industrial sector (coal, oil, natural gas, and electricity).

278 To construct a comprehensive municipal-level energy inventory for Xi'an, a
279 bottom-up scaling approach was applied following Shan et al. (2017). Provincial-level
280 energy balance data for Shaanxi were disaggregated to the city level using
281 sector-specific conversion factors based on the relative contributions of Xi'an and
282 Shaanxi Province. The sectoral classification is consistent with that used for Beijing.
283 Industrial energy consumption was directly obtained from Xi'an Statistical Yearbooks,
284 while the remaining sectors were estimated using conversion factors (Table S1).

285 The municipal energy consumption for each sector was calculated as:

$$286 \quad Index_i = \frac{P_{city}}{P_{province}} \times 100\% \quad (6)$$

287 where P_{city} and $P_{province}$ represent the relevant activity or energy indicators for
288 Xi'an and Shaanxi Province, respectively.

289

290 **3 Results and discussion**

291 **3.1 Interannual reduction of CO₂ fluxes**

292 Figure 2 shows the daily CO₂ fluxes at the observation sites in Beijing and Xi'an. All
293 urban sites exhibited a pronounced "U-shaped" seasonal pattern, with higher CO₂
294 fluxes in winter and lower values in summer. This seasonal contrast reflects enhanced
295 emissions during the heating period, combined with reduced photosynthetic uptake
296 under low temperature and light conditions, whereas stronger photosynthesis in
297 summer leads to lower net CO₂ fluxes (Delgado-Balbuena et al., 2013; Park et al.,
298 2019).

299 Despite differences in magnitude, all urban sites acted as net carbon sources and



300 displayed clear interannual declines in CO₂ fluxes over the study period (Figure 2a-c).
301 For example, annual net carbon exchange decreased from 2021 to 2024 at both
302 Beijing sites (from 7.2 to 6.1 kg C m⁻² yr⁻¹ at IAP, and from 3.0 to 1.9 kg C m⁻² yr⁻¹
303 at RCEES), with a similar downward trend observed at the Xi'an urban site (from 1.7
304 to 1.0 kg C m⁻² yr⁻¹). Notably, anomalously low CO₂ fluxes occurred during periods
305 affected by COVID-19 restrictions, with reduced values observed in Beijing in 2020
306 and in Xi'an in 2022 during the strict lockdown period (Figure 2a-c), highlighting the
307 strong influence of anthropogenic activity on urban carbon emissions. In addition,
308 lower CO₂ flux values were recorded during the heavy snow period in Beijing,
309 December 2023 (two snowfall events, 14.7 mm snow water equivalent;
310 <https://data.cma.cn/>), likely driven by reduced anthropogenic emissions associated
311 with the snowfall.

312 CO₂ fluxes during the Spring Festival further reveal the impact of human activity
313 on urban carbon dynamics. In 2020, festival CO₂ fluxes in Beijing did not differ
314 significantly from non-holiday winter levels, indicating that pandemic-related
315 restrictions suppressed mobility throughout both periods and masked the typical
316 holiday-related emission reduction. In contrast, from 2021 to 2024, CO₂ fluxes during
317 the Spring Festival were consistently lower than non-holiday winter values at both
318 Beijing sites, with an average reduction of 23.6 ± 18.6% (Figure 2a, b). A similar
319 pattern was observed in Xi'an, where Spring Festival CO₂ fluxes decreased by 28.9 ±
320 4.7% relative to non-holiday periods (Figure 2c). These consistent reductions
321 demonstrate the substantial impact of holiday-associated population migration and
322 reduced anthropogenic activity on urban carbon emissions (Huang et al., 2012; Liu et
323 al., 2024).

324

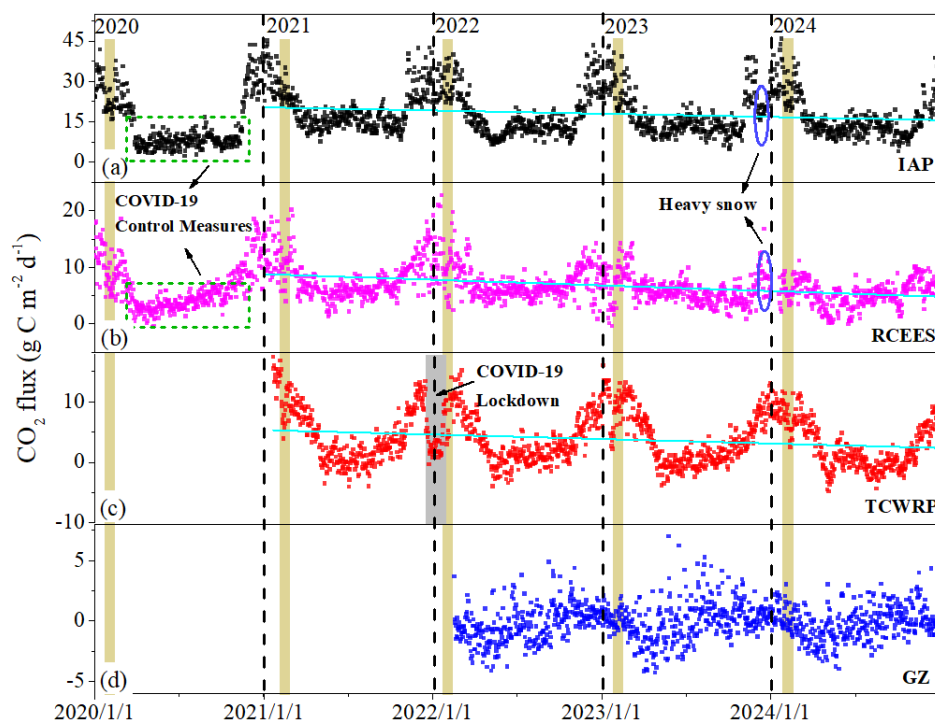


Figure 2. Daily CO₂ fluxes at the observation sites in Beijing and Xi'an. Panels (a) and (b) show the two urban sites in Beijing (IAP and RCEES), while panels (c) and (d) show the urban (TCWRP) and rural (GZ) sites in Xi'an. Yellow shading indicates the Spring Festival period, and gray shading denotes the COVID-19 lockdown period. Green dashed boxes and blue ellipses indicate periods affected by COVID-19 control measures and heavy snow events, respectively. Cyan lines represent linear trends in CO₂ fluxes at the urban sites.

325

326 In contrast, the rural site in Xi'an (GZ) functioned as a weak carbon sink or
 327 near-neutral system during the observation period, with substantially lower net carbon
 328 exchange than the urban sites (Figure 2d). This pattern reflects reduced anthropogenic
 329 emissions and greater vegetation influence, with continuous CO₂ uptake by
 330 surrounding ecosystems (Havu et al., 2024). The seasonal variability of CO₂ flux was
 331 less pronounced than at urban sites, and net carbon uptake occurred throughout most
 332 of the year. The strongest sink was observed in spring (March–May), corresponding to
 333 the early growing season when increasing temperature enhances photosynthesis while



334 ecosystem respiration remains relatively low (Davidson et al., 2005).

335 In contrast to urban sites, CO₂ fluxes at the rural site showed little difference
336 between the Spring Festival and non-holiday periods (Figure 2d), indicating minimal
337 sensitivity to short-term changes in human activity. This suggests that carbon
338 dynamics at the rural site are primarily controlled by natural processes rather than
339 anthropogenic emissions. These findings highlight the contrasting roles of urban and
340 rural environments and underscore the importance of enhancing rural carbon sinks to
341 support low-carbon urban development.

342

343 **3.2 Further identification of reduction periods of CO₂ fluxes**

344 We further examined the temporal patterns of urban carbon emission reductions by
345 analyzing diurnal variations in CO₂ fluxes (Figure 3). All sites exhibited a clear
346 day-night cycle, reflecting CO₂ release during nighttime and uptake during daytime.
347 At the urban sites in Beijing and Xi'an, CO₂ fluxes showed a pronounced bimodal
348 pattern, with peaks corresponding to morning and evening traffic rush hours (Figure
349 3a-c), consistent with observations in other cities (Bergeron and Strachan, 2011; Liu
350 et al., 2012).



351 A clear interannual decline in CO₂ fluxes was observed during the period from
352 November to April (Figure 3a-c), corresponding to the heating season. In particular,
353 hourly CO₂ fluxes at the IAP, RCEES, and TCWRP sites were consistently lower in
354 2024 compared to 2021 (Figure S2a-c). This seasonal concentration of emission
355 reductions highlights the dominant role of heating-related energy use and its transition
356 toward cleaner energy sources in driving the observed declines.

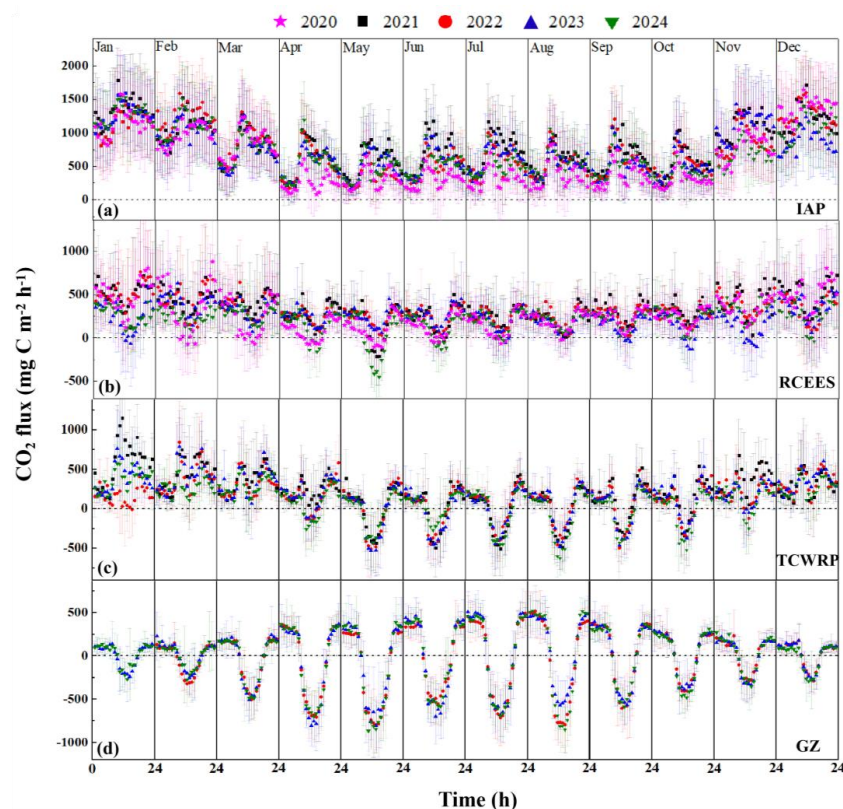


Figure 3. Monthly diurnal variations of CO₂ fluxes at the observation sites in Beijing and Xi'an during 2020–2024. Panels (a) and (b) show the urban sites in Beijing (IAP and RCEES), while panels (c) and (d) show the urban (TCWRP) and rural (GZ) sites in Xi'an.

357

358 To identify periods of carbon mitigation, CO₂ fluxes were analyzed separately for
359 heating (November 15–March 15) and non-heating seasons (March 16–November 14).



360 The corresponding diurnal variations are shown in Figure 4. In Beijing, CO₂ fluxes
361 during the main period of anthropogenic activity (07:00–23:00) reached their lowest
362 levels in 2020, reflecting reduced human mobility during COVID-19 restrictions
363 (Zhao et al., 2025).

364 During the heating seasons from 2021 to 2024, CO₂ fluxes exhibited a clear
365 interannual decline across all urban sites in both Beijing and Xi'an (Figure 4a, c, e).
366 This decrease was observed throughout the daytime period and was also evident
367 during nighttime hours at the Beijing sites. Notably, the largest reductions occurred
368 during traffic peak hours, indicating a strong contribution from changes in
369 transportation-related emissions.

370 During the non-heating season, reductions in CO₂ fluxes were also evident,
371 particularly during peak traffic periods (Figure 4b, d, f). These consistent declines
372 across both seasons suggest that emission reductions are not limited to heating-related
373 sources but also reflect broader changes in urban energy use and transportation
374 systems. This pattern is consistent with ongoing energy structure adjustments and the
375 decarbonization of urban transport.

376 However, interannual variations in summer CO₂ fluxes differed between Beijing
377 and Xi'an, despite the consistent seasonal patterns across all urban sites (Figure 3a-c).
378 In Beijing, summer CO₂ fluxes showed a gradual decline from 2021 to 2024 at both
379 urban sites, whereas no consistent interannual trend was observed at the urban site in
380 Xi'an. This contrast suggests that emission reductions in Beijing extend beyond the
381 heating season, while in Xi'an summer fluxes are more strongly influenced by local
382 variability and non-anthropogenic factors.

383 At the rural site in Xi'an, CO₂ fluxes exhibited a typical “U-shaped” diurnal cycle,
384 with nighttime CO₂ release driven by ecosystem respiration and daytime uptake



dominated by photosynthesis (Figure 3d) (Zhan et al., 2024). During the growing season (April–September), strong daytime carbon uptake and nighttime release were consistently observed, with no significant interannual differences between 2022 and 2024 (Figure S2d). These patterns indicate that CO₂ fluxes at the rural site are primarily controlled by biogenic processes, with vegetation photosynthesis and ecosystem respiration governing the observed variability.

Although the observation sites in Beijing and Xi'an differ in location, instrument height, spatial footprint, and surface characteristics, a consistent interannual decline in CO₂ fluxes was observed across all urban sites. This decline was evident during both the heating season and peak traffic periods, indicating a robust and systematic reduction in urban carbon emissions.

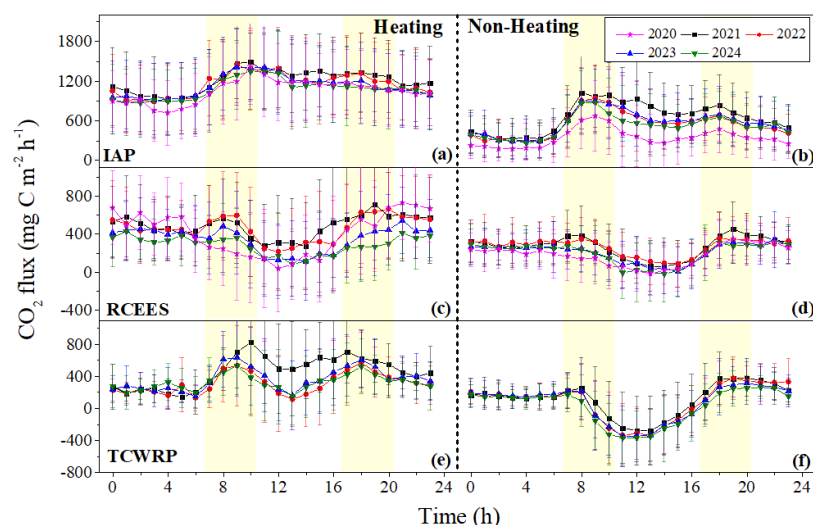


Figure 4. Diurnal variations of CO₂ fluxes at urban sites in Beijing and Xi'an during heating and non-heating seasons. IAP and RCEES represent the Beijing sites, and TCWRP represents the Xi'an site. Yellow shading indicates peak traffic

396

397 3.3 Reduction of CO_{2ff} concentrations traced by tree-ring ¹⁴C

398 The interannual decline in CO₂ fluxes observed at the urban sites raises the question



399 of whether these changes are driven by fossil fuel emissions or biogenic processes. To
400 address this, we analyzed fossil fuel-derived CO₂ (CO_{2ff}) using tree-ring ¹⁴C. CO_{2ff}
401 concentrations at all urban sites in Beijing and Xi'an exhibited a clear and consistent
402 decline from 2021 to 2024 (Figure 5). This decrease was observed across all sites,
403 with the largest absolute reduction at the TCWRP site and the lowest overall

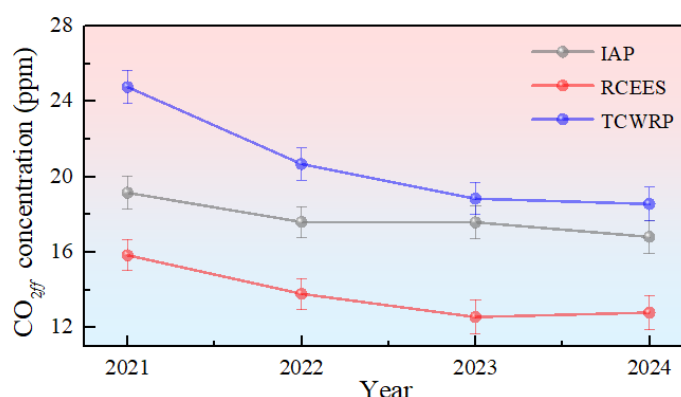


Figure 5. Interannual variations of fossil fuel CO₂ (CO_{2ff}) concentrations based on tree-ring ¹⁴C at the urban sites in Beijing and Xi'an from 2021 to 2024.

404 concentrations at the RCEES site. The consistent downward trends in CO_{2ff} provide
405 direct evidence for a reduction in fossil fuel-derived CO₂ in urban environments.

406

407 Further analysis reveals a strong positive correlation between interannual variations
408 in CO₂ fluxes and CO_{2ff} concentrations across the three urban sites. In particular, the
409 IAP site shows an exceptionally high correlation ($R^2 = 0.99$, $p < 0.01$), indicating that
410 variations in total CO₂ flux closely track changes in fossil fuel emissions. This strong
411 coupling likely reflects the large footprint of the IAP site, which integrates emissions
412 over a regional scale and is less influenced by local variability. Taken together, the
413 concurrent decline in CO₂ fluxes and CO_{2ff} concentrations demonstrates that reduced
414 fossil fuel combustion is the primary driver of the observed decrease in urban carbon



emissions.

3.4 Analysis of driving factors for the interannual decline in urban CO₂ fluxes using energy consumption data

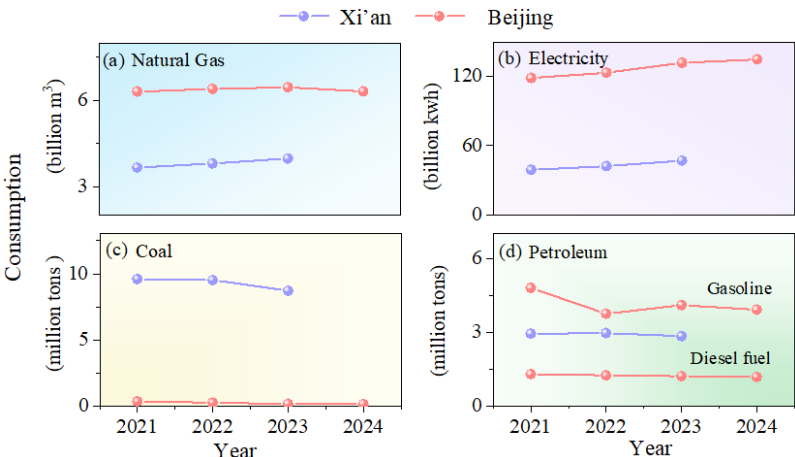


Figure 6. Trends in energy consumption in Beijing (2021–2024) and Xi'an (2021–2023), including (a) natural gas, (b) electricity, (c) coal, and (d) petroleum.

The interannual decline in CO₂ fluxes during the heating season is closely linked to structural changes in urban energy use, as indicated by energy consumption data from both Beijing and Xi'an (Figure 6, S3, S4). Both cities show a clear transition toward cleaner energy, characterized by increasing electricity use and declining coal consumption. This shift, together with a moderate increase in natural gas use, reduces CO₂ emissions associated with heating, as natural gas and electricity have lower carbon intensity than coal (United States Energy Information Administration, 2016). These results suggest that energy structure transformation plays a key role in driving emission reductions during the heating season.

Reductions in CO₂ fluxes during traffic peak periods are associated with changes in transportation systems. Both cities exhibit declining consumption of petroleum-based



433 fuels and increasing adoption of low-carbon mobility options. In particular, the rapid
434 expansion of metro networks and the large-scale deployment of new energy vehicles
435 (NEVs) indicate a transition toward electrified and public transport systems.
436 Particularly in terms of NEVs, Beijing's fleet size reached over 1 million units in
437 2024 (a 97.2% increase compared to 2021), and Xi'an's exceeded 400,000 units (a
438 268.7% rise compared to 2021). These changes are consistent with the observed
439 reductions in CO₂ fluxes during peak traffic hours, highlighting the importance of
440 transport decarbonization in urban emission reductions.

441 Together with the observed decline in fossil fuel-derived CO₂ (CO_{2ff}), these results
442 demonstrate that reduced fossil fuel combustion is the primary driver of decreasing
443 urban carbon emissions. More broadly, the findings indicate that sustained emission
444 reductions in urban environments are driven by structural transitions in energy and
445 transportation systems, rather than by short-term variability.

446

447 **4 Conclusions**

448 In summary, this study quantified the temporal evolution and driving mechanisms of
449 urban CO₂ fluxes in Beijing and Xi'an from 2020 to 2024 using eddy covariance
450 measurements combined with tree-ring ¹⁴C. Urban sites in both cities acted as
451 persistent carbon sources but exhibited clear interannual declines in CO₂ fluxes from
452 2021 to 2024. The most pronounced reductions occurred during the heating season
453 and traffic peak hours, highlighting the dominant influence of anthropogenic activities.
454 In contrast, the rural site in Xi'an functioned as a weak carbon sink with no significant
455 interannual trend. Independent constraints from tree-ring ¹⁴C reveal a concurrent
456 decline in CO_{2ff}, demonstrating that reduced fossil fuel combustion is the primary
457 driver of the observed emission decreases. Analysis of energy consumption data



458 further indicates that these reductions are linked to structural changes in urban energy
459 systems, including the substitution of coal and petroleum with electricity and natural
460 gas, as well as the rapid expansion of low-carbon transportation. Overall, our results
461 provide direct observational evidence that urban carbon emissions can decline under
462 sustained climate mitigation policies. They highlight that energy system
463 transformation and transport decarbonization are central to achieving urban emission
464 reductions, while enhanced rural carbon sinks may provide additional support for
465 low-carbon development.

466

467 **Data availability.** The data that support the findings of this study are available in the
468 Figshare repository (<https://figshare.com/s/a019e730251e29032259>).

469

470 **Supplement.** The supplement related to this article is available online.

471

472 **Author contributions.** Z.N. conceived the study. Z.N., D.L., H.Z., X.L., J.J., Y.T.,
473 G.W. and X.F. collected data. Z.N. and D.L. performed sampling and ^{14}C sample
474 preparation. X.L. performed ^{14}C measurements. Z.N., X.K., S.W., W.Z., W.L., Z.J.,
475 Y.H. and J.P. provided valuable discussions. D.L. drafted the paper, which was revised
476 by Z.N., X.K. and J.P.

477

478 **Competing interests.** The contact author has declared that none of the authors has
479 any competing interests.

480

481 **Acknowledgements.** This research was jointly supported by the National Natural

482



483 Science Foundation of China (Nos. 42173082 and 42330114), and Natural Science

484 Basic Research Program of Shaanxi (No. 2024JC-JCQN-34).

485



486 **References**

- 487 Amar, G., Mamtimin, A., Wang, Y., Wang, Y., Gao, J., Yang, F., Song, M., Aihaiti, A., Wen, C.,
488 and Liu, J.: Factors controlling and variations of CO₂ fluxes during the growing season
489 in Gurbantungut Desert, Ecol. Indic., 154, 110708,
490 <https://doi.org/10.1016/j.ecolind.2023.110708>, 2023.
- 491 Amiro, B. D., Barr, A. G., Black, T. A., Iwashita, H., Kljun, N., McCaughey, J. H.,
492 Morgenstern, K., Murayama, S., Nesic, Z., Orchansky, A. L., and Saigusa, N.: Carbon,
493 energy and water fluxes at mature and disturbed forest sites, Saskatchewan, Canada,
494 Agri. Forest Meteorol., 136, 237–251, <https://doi.org/10.1016/j.agrformet.2004.11.012>,
495 2006.
- 496 Baldocchi, D. D.: Assessing the eddy covariance technique for evaluating carbon dioxide
497 exchange rates of ecosystems: past, present and future, Global Change Biol., 9, 479–492,
498 <https://doi.org/10.1046/j.1365-2486.2003.00629.x>, 2003.
- 499 Bergeron, O. and Strachan, I. B.: CO₂ sources and sinks in urban and suburban areas of a
500 northern mid-latitude city, Atmos. Environ., 45, 1564–1573,
501 <https://doi.org/10.1016/j.atmosenv.2010.12.043>, 2011.
- 502 Canadell, J. G., Ciais, P., Dhakal, S., Le Quéré C., Patwardhan, A. and Raupach, M. R.: The
503 Global Carbon Cycle 2- the Human Perturbation of the Carbon Cycle, Paris, France. No.
504 10. UNESCO-SCOPE-UNEP Policy Briefs, [http://unesdoc.unesco.org/](http://unesdoc.unesco.org/images/0018/001861/186137e.pdf)
505 [images/0018/001861/186137e.pdf](http://unesdoc.unesco.org/images/0018/001861/186137e.pdf) last access June 2018, 2009.
- 506 Davidson, E. A., Richardson, A. D., Savage, K. E. and Hollinger, D. Y.: A distinct seasonal
507 pattern of the ratio of soil respiration to total ecosystem respiration in a
508 spruce-dominated forest, Global Change Biol., 12, 230–239,
509 <https://doi.org/10.1111/j.1365-2486.2005.01062.x>, 2005.
- 510 Delgado-Balbuena, J., Arredondo, J. T., Loescher, H. W., Huber-Sannwald, E.,
511 Chavez-Aguilar, G., Luna-Luna, M., and Barretero-Hernandez, R.: Differences in plant
512 cover and species composition of semiarid grassland communities of central Mexico and



- 513 its effects on net ecosystem exchange, *Biogeosciences*, 10, 4673–4690,
514 <https://doi.org/10.5194/bg-10-4673-2013>, 2013.
- 515 Duan, Z. and Zhu, X.: The carbon paradox in urban China: High-quality development under
516 the dual-control policy, *Sustain. Cities and Soc.*, 130, 106585.
517 <https://doi.org/10.1016/j.scs.2025.106585>, 2025.
- 518 Falge, E., Baldocchi, D., Olson, R., Anthoni, P., Aubinet, M., Bernhofer, C., Burba, G.,
519 Ceulemans, R., Clement, R., Dolman, H., Granier, A., Gross, P., Grünwald, T., Hollinger,
520 D., Jensen, N., Katul, G., Keronen, P., Kowalski, A., Lai, C. T., Law, B. E., Meyers, T.,
521 Moncrieff, J., Moors, E., Munger, J. W., Pilegaard, K., Rannik, Ü., Rebmann, C., Suyker,
522 A., Tenhunen, J., Tu, K., Verma, S., Vesala, T., Wilson, K., and Wofsy, S.: Gap filling
523 strategies for defensible annual sums of net ecosystem exchange, *Agr. Forest Meteorol.*,
524 107, 43–69, [https://doi.org/10.1016/S0168-1923\(00\)00225-2](https://doi.org/10.1016/S0168-1923(00)00225-2), 2001.
- 525 Foken, T., Göckede, M., Mauder, M., Mahrt, L., Amiro, B. and Munger, W.: Post-Field Data
526 Quality Control. In: Lee, X., Massman, W., Law, B. (eds) *Handbook of*
527 *Micrometeorology, Atmospheric and Oceanographic Sciences Library*, 29, 181–208,
528 https://doi.org/10.1007/1-4020-2265-4_9, 2004.
- 529 Godwin, H.: Half-life of radiocarbon, *Nature*, 195, 984, 1962.
- 530 Hao, J., Gao, F., Fang, X., Nong, X., Zhang, Y. and Hong, F.: Multi-factor decomposition and
531 multi-scenario prediction decoupling analysis of China’s carbon emission under dual
532 carbon goal, *Sci. Total Environ.*, 841, 156788,
533 <https://doi.org/10.1016/j.scitotenv.2022.156788>, 2022.
- 534 Havu, M., Kulmala, L., Lee, H. S., Saranko, O., Soininen, J., Ahongshangbam, J., and Järvi,
535 L.: CO₂ uptake of urban vegetation in a warming Nordic city, *Urban For. Urban Gree.*,
536 94, 128261, <https://doi.org/10.1016/j.ufug.2024.128261>, 2024.
- 537 Hirata, R., Saigusa, N., Yamamoto, S., Ohtani, Y., Ide, R., Asanuma, J., Gamo, M., Hirano, T.,
538 Kondo, H., Kosugi, Y., Li, S., Nakai, Y., Takagi, K., Tani, M., and Wang, H.: Spatial
539 distribution of carbon balance in forest ecosystems across East Asia, *Agr. Forest*



- 540 Meteorol., 148, 761–775, <https://doi.org/10.1016/j.agrformet.2007.11.016>, 2008.
- 541 Hua, Q., Turnbull, J. C., Santos, G. M., Rakowski, A. Z., Ancapichún, S., Pol-Holz, R. D.,
542 Hammer, S., Lehman, S. J., Levin, I., Miller, J. B., Palmer, J. G., and Turney, C. S. M.:
543 Atmospheric radiocarbon for the period 1950–2019, Radiocarbon, 64, 723–745,
544 <https://doi.org/10.1017/RDC.2021.95>, 2022.
- 545 Huang, K., Zhuang, G., Lin, Y., Wang, Q., Fu, J. S., Zhang, R., Li, J., Deng, C., and Fu, Q.:
546 Impact of anthropogenic emission on air quality over a megacity-revealed from an
547 intensive atmospheric campaign during the Chinese Spring Festival, Atmos. Chem.
548 Phys., 12, 11631–11645, <https://doi.org/10.5194/acp-12-11631-2012>, 2012.
- 549 Huang, R., Zhang, S. and Wang, P.: Key areas and pathways for carbon emissions reduction in
550 Beijing for the “Dual Carbon” targets, Energ. Policy, 164, 112873,
551 <https://doi.org/10.1016/j.enpol.2022.112873>, 2022.
- 552 Kljun, N., Calanca, P., Rotach, M. W. and Schmid, H. P.: A simple two-dimensional
553 parameterisation for Flux Footprint Prediction (FFP), Geosci. Model Dev., 8, 3695–3713,
554 <https://doi.org/10.5194/gmd-8-3695-2015>, 2015.
- 555 Levin, I., Kromer, B., Schmidt, M. and Sartorius, H.: A novel approach for independent
556 budgeting of fossil fuel CO₂ over Europe by ¹⁴CO₂ observations, Geophys. Res. Lett., 30,
557 2194, <https://doi.org/10.1029/2003GL018477>, 2003.
- 558 Liu, H., Feng, J., Järvi, L. and Vesala, T.: Four-year (2006–2009) eddy covariance
559 measurements of CO₂ flux over an urban area in Beijing, Atmos. Chem. Phys., 12,
560 7881–7892, <https://doi.org/10.5194/acp-12-7881-2012>, 2012.
- 561 Liu, Z., Zeng, N., Han, P., Cai, Q., Yao, B., Wang, P., Quan, W., Tang, W., Liang, Z., Liu, Y.,
562 and Cory, M.: Detection of Chinese Spring Festival in Beijing using in-situ CO₂
563 observations and atmospheric inversion, Atmos. Environ., 325, 120446,
564 <https://doi.org/10.1016/j.atmosenv.2024.120446>, 2024.
- 565 McGlade, C. and Ekins, P.: The geographical distribution of fossil fuels unused when limiting
566 global warming to 2 °C, Nature, 517, 187–190, <https://doi.org/10.1038/nature14016>,
567 2015.



- 568 MFA, PRC (Ministry of Foreign Affairs of the People's Republic of China). Statement by H.E.
569 Xi Jinping President of the People's Republic of China At the General Debate of the 75th
570 Session of The United Nations General Assembly, [https://www.ndrc.gov.cn/](https://www.ndrc.gov.cn/wsdwhfz/202111/t20211111_1303691_ext.html)
571 [wsdwhfz/202111/t20211111_1303691_ext.html](https://www.ndrc.gov.cn/wsdwhfz/202111/t20211111_1303691_ext.html), 2020.
- 572 Park, T., Chen, C., Macias-Fauria, M., Tømmervik, H., Choi, S., Winkler, A., Bhatt, U. S.,
573 Walker, D. A., Piao, S., Brovkin, V., Nemani, R. R., and Myneni, R. B.: Changes in
574 timing of seasonal peak photosynthetic activity in northern ecosystems, *Global Change*
575 *Biol.*, 25, 2382–2395, <https://doi.org/10.1111/gcb.14638>, 2019.
- 576 Rakowski, A. Z.: Radiocarbon method in monitoring of fossil fuel emission, *Geochronometria*,
577 38, 314–324, <https://doi.org/10.2478/s13386-011-0044-3>, 2011.
- 578 Shan, Y., Guan, D., Liu, J., Mi, Z., Liu, Z., Liu, J., Schroeder, H., Cai, B., Chen, Y., Shao, S.,
579 and Zhang, Q.: Methodology and applications of city level CO₂ emission accounts in
580 China, *J. Clean. Prod.*, 161, 1215–1225, <https://doi.org/10.1016/j.jclepro.2017.06.075>,
581 2017.
- 582 Slota, P., Jull, A.T., Linick, T. and Toolin, L.: Preparation of small samples for ¹⁴C accelerator
583 targets by catalytic reduction of CO, *Radiocarbon*, 29, 303–306,
584 <https://doi.org/10.1017/S0033822200056988>, 1987.
- 585 Solomon, S., Plattner, G., Knutti, R. and Friedlingstein, P.: Irreversible climate change due to
586 carbon dioxide emissions, *Proc. Natl. Acad. Sci. U. S. A.*, 106, 1704–1709,
587 <https://doi.org/10.1073/pnas.0812721106>, 2009.
- 588 Stuiver, M. and Polach, H. A.: Discussion reporting of ¹⁴C data, *Radiocarbon*, 19, 355–363,
589 <https://doi.org/10.1017/S0033822200003672>, 1977.
- 590 Turnbull, J. C., Keller, E. D., Norris, M. W. and Wiltshire, R. M.: Independent evaluation of
591 point source fossil fuel CO₂ emissions to better than 10%, *Proc. Natl. Acad. Sci. U. S. A.*,
592 113, 10287–10291, <https://doi.org/10.1073/pnas.1602824113>, 2016.
- 593 Ueyama, M. and Ando, T.: Diurnal, weekly, seasonal, and spatial variabilities in carbon
594 dioxide flux in different urban landscapes in Sakai, Japan, *Atmos. Chem. Phys.*, 16,
595 14727–14740, <https://doi.org/10.5194/acp-16-14727-2016>, 2016.



- 596 United States Energy Information Administration (EIA), Carbon Dioxide Emission
597 Coefficients: Carbon Dioxide Emissions Coefficients by Fuel, Department of Energy,
598 Washington D.C, 2016.
- 599 Wang, J., Yu, G., Han, L., Yao, Y., Sun, M. and Yan, Z.: Ecosystem carbon exchange across
600 China's coastal wetlands: Spatial patterns, mechanisms, and magnitudes, *Agr. Forest*
601 *Meteorol.*, 345, 109859, <https://doi.org/10.1016/j.agrformet.2023.109859>, 2024a.
- 602 Wang, P., Huang, G., Li, Y., Liu, Y. and Li, Y.: An ecological input-output CGE model for
603 unveiling CO₂ emission metabolism under China's dual carbon goals, *Appl. Energ.*, 365,
604 123277, <https://doi.org/10.1016/j.apenergy.2024.123277>, 2024b.
- 605 Ward, H. C., Evans, J. G. and Grimmond, C. S. B.: Multi-season eddy covariance
606 observations of energy, water and carbon fluxes over a suburban area in Swindon, UK,
607 *Atmos. Chem. Phys.*, 13, 4645–4666, <https://doi.org/10.5194/acp-13-4645-2013>, 2013.
- 608 Wei, D., Qi, Y., Ma, Y., Wang, X., Ma, W., Gao, T., Huang, L., Zhao, H., Zhang, J., and Wang,
609 X.: Plant uptake outpaces permafrost and plant carbon loss on the Tibetan Plateau, *Proc.*
610 *Natl. Acad. Sci. U. S. A.*, 118, e2015283118, <https://doi.org/10.1073/pnas.2015283118>,
611 2021.
- 612 Wen, X., Wang, H., Wang, J., Yu, G. and Sun, X.: Ecosystem carbon exchanges of a
613 subtropical evergreen coniferous plantation subjected to seasonal drought, 2003–2007,
614 *Biogeosciences*, 7, 357–369, <https://doi.org/10.5194/bg-7-357-2010>, 2010.
- 615 Xiao, W., Liu, S., Li, H., Xiao, Q., Wang, W., Hu, Z., Hu, C., Gao, Y., Shen, J., Zhao, X.,
616 Zhang, M., and Lee, X.: A flux-gradient system for simultaneous measurement of the
617 CH₄, CO₂, and H₂O fluxes at a lake–air interface, *Environ. Sci. Technol.*, 48, 14490–
618 14498, <https://doi.org/10.1021/es5033713>, 2014.
- 619 Xueref-Remy, I., Milne, M., Zoghbi, N., Lelandais, L., Riandet, A., Armengaud, A., Gille, G.,
620 Lanzi, L., Oppo, S., Brégonzio-Rozier, L., Blanc, P., Yohia, C., Piazzola, J., and
621 Delmotte, M.: Analysis of atmospheric CO₂ variability in the Marseille city area and the
622 north-west Mediterranean basin at different time scales, *Atmos. Environ.*: X, 17, 100208,
623 <https://doi.org/10.1016/j.aeaoa.2023.100208>, 2023.



- 624 Yang, J. and Huang, X.: The 30 m Annual Land Cover Datasets and Its Dynamics in China
625 from 1985 to 2024 [Dataset], Zenodo., <https://doi.org/10.5281/zenodo.12779975>, 2024.
- 626 Yu, G., Wen, X., Sun, X., Tanner, B. D., Lee, X. and Chen, J.: Overview of ChinaFLUX and
627 evaluation of its eddy covariance measurement, *Agr. Forest Meteorol.*, 137, 125–137,
628 <https://doi.org/10.1016/j.agrformet.2006.02.011>, 2006.
- 629 Yu, G., Zhang, L., Sun, X., Fu, Y., Wen, X., Wang, Q., Li, S., Ren, C., Song, X., Liu, Y., Han,
630 S., and Yan, J.: Environmental controls over carbon exchange of three forest ecosystems
631 in eastern China, *Global Change Biol.*, 14, 2555–2571,
632 <https://doi.org/10.1111/j.1365-2486.2008.01663.x>, 2008.
- 633 Zhan, Y., Deng, Y., Wang, J., Liu, Q. and Wang, W.: Environmental response characteristics of
634 the carbon and water fluxes above complex urban surfaces of a subtropical megacity in
635 China, *Phys. Chem. Earth*, 135, 103681, <https://doi.org/10.1016/j.pce.2024.103681>,
636 2024.
- 637 Zhao, L., Zhang, H., Sun, X., Huang, B., Lu, F., Wang, X., and Ouyang, Z.: Dynamics of
638 urban carbon emissions: A comparative analysis of pre- and post-lockdown through eddy
639 covariance measurements, *J. Environ. Sci.*, 158, 735–747,
640 <https://doi.org/10.1016/j.jes.2025.01.023>, 2025.
- 641 Zhu, N., Wang, J., Luo, D., Wang, X., Shen, C. and Wu, N.: Using eddy covariance
642 observations to determine the carbon sequestration characteristics of subalpine forests in
643 the Qinghai-Tibet Plateau, *Biogeosciences*, 21, 3509–3522,
644 <https://doi.org/10.5194/bg-21-3509-2024>, 2024.
- 645