https://doi.org/10.5194/egusphere-2026-381
Preprint. Discussion started: 11 March 2026

Evaluation of representation of seasonally frozen ground characteristics in Land Surface Models: JSBACH and CLM

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Abstract. Land surface models (LSMs) differ in simulating winter soil conditions due to complex freeze-thaw processes and snow-soil interactions, leading to uncertainties in spring and summer soil moisture and runoff. This study evaluates standalone simulations from two LSMs, JSBACH and CLM, driven by ERA5 data (1986-2022), compares them with ERA5-Land to study model differences across cold regions, and then examines all three models against the reference RIHMI-WDC observational dataset at 26 sites to better understand how SFG characteristics (timing, duration, and freeze depth) are represented in these models. The research aims to identify biases in simulated seasonally frozen ground (SFG) characteristics, investigate their causes, and assess how snow cover errors propagate into frozen ground biases using site-level evaluation over Russia. The importance of snow parameterization is highlighted in this study through an improvement to the snow density scheme in JSBACH, which reduced its cold bias in soil temperature by up to 10-20 ^{o}C, and this improved version was used for the comparative analysis. Among the models assessed in this study, JSBACH reproduces frozen ground extent most realistically, closely matching reference estimates of SFG and permafrost (PEFT) extent, but it simulates reduced snow depth (mean bias = -14.2 cm), leading to weaker insulation, enhanced soil cooling (mean bias = -3.7^{o}C), and deeper seasonal freezing, whereas CLM simulates soil temperatures comparatively close to observations ($+0.1^{o}$C) under colder air temperatures (-3.4^{o}C) and excessive snow (12.2 cm), indicating overestimated snow insulation. Site-level freeze-thaw evaluation reveals systematic biases across models, including premature autumn freezing and delayed spring thaw, leading to longer frozen ground duration (16 to 19 days) associated with contrasting snow insulation effects in CLM and JSBACH. Soil freezing in JSBACH responded too strongly to surface thermal forcing, whereas ERA5-Land and CLM showed an overestimated relationship between SFG and snow characteristics. This discrepancy indicates that LSMs differ in how control of soil freeze-thaw is partitioned between air temperature and snow processes. The study highlights that improving model performance requires better snow representation and a detailed assessment aimed at enhancing the parameterization of soil thermal and hydraulic properties.

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1 Introduction

Frozen ground is a significant component of the global cryosphere, which includes both seasonally frozen ground (SFG) and permafrost (PEFT) (Peng et al., 2020b). SFG covers nearly fifty percent of the exposed land area in the Northern Hemisphere (including the active layer overlaying PEFT), where freeze-thaw processes significantly influence local and global climate, atmospheric circulation, hydrology, terrestrial ecosystems, and biogeochemistry (Boli et al., 2014; Luo et al., 2016; Schmidt et al., 2011; Li et al., 2002; Jin et al., 2009; Schuur et al., 2009). The coexistence of ice and liquid water in this area modifies the hydraulic and thermal properties of the soil, consequently affecting distribution of water and energy within the land surface system (Jan et al., 2022). Frozen soil can increase overland flow and flood risk due to reduced hydraulic conductivity, and earlier thawing of frozen ground can impact ecosystems by accelerating spring greening through reduced albedo and enhanced soil moisture infiltration (Stuurop et al., 2022; Hua et al., 2025). Previous studies have primarily looked at freeze-thaw processes in PEFT regions, with most research limited to local or regional case studies (Chang et al., 2018; Zhang et al., 1997; Peng et al., 2023). A limited number of studies have shown the dynamics and broader implications of freeze-thaw cycles in SFG on region scale (Luo et al., 2020; Zhao et al., 2022), with only a very few addressing the hemispheric scale (Chen et al., 2022), leaving significant gaps in our understanding of SFG's large-scale behavior and environmental impacts.

Land surface models (LSMs) are essential tool for exploring the physical mechanisms underlying the frozen soil processes. However, many studies have reported persistent challenges in accurately simulating soil temperature, snow cover, and freeze-thaw transitions under low-temperature conditions (Li et al., 2002; Zheng et al., 2017; Luo et al., 2020; Risto et al., 2022). Considerable discrepancies are often observed among different models even when they are driven by identical meteorological forcing, leading to substantial uncertainty in simulated surface water and energy budgets (Zheng et al., 2017; Slater et al., 2007). These inconsistencies mainly come from the highly nonlinear interactions between soil temperature, soil moisture, and ice content, as well as from differing model parameterizations of heat and mass transfer in frozen soils. A crucial prerequisite for evaluating how well models simulate freeze-thaw conditions is the assessment of the spatial distribution of frozen ground on large scales. However, despite its importance, available hemispheric-scale maps for the Northern Hemisphere are still limited. According to Zhang et al. (2000), SFG is estimated to cover approximately 48 million km^2, whereas PEFT occupies about 12.21 to 16.98 million km^2. More recent research, however, suggests that the extent of PEFT may range between 13.60 and 18.97 million km^2 (Ran et al., 2022). Mapping and comparing these distributions across models is a preliminary yet crucial step in identifying and addressing the uncertainties related to the status of frozen ground.

Apart from spatial distribution, a comprehensive understanding of SFG also requires investigating its key characteristics, such as the timing of freeze onset, thaw or end date, duration of the frozen period, and maximum freezing depth, along with their primary controlling factors (Chen et al., 2022; Li et al., 2021). Numerous studies have shown that air temperature and snow cover are the dominant factors modulating SFG characteristics, as they particularly influence the depth and duration of soil freezing (Wang and Chen, 2022; Zhang et al., 2021; Osokin et al., 2000; Zhang, 2005; Lawrence and Slater, 2010; Peng et al., 2016). Observational evidence consistently shows that rising air temperatures lead to delayed freezing, earlier thawing, a shorter frozen period, and shallower freezing depth (Li et al., 2021, 2012; Wang et al., 2015; Frauenfeld et al., 2011). Meanwhile, snow

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cover plays a dual and complex role through its insulating effect, it buffers the soil from air temperature fluctuations, and its influence depends strongly on snow characteristics such as timing, duration, thickness, density, and structure. In northern latitudes, shallow snow and shorter snow cover duration tend to enhance soil freezing by allowing greater heat loss from the ground to the atmosphere while under thick and persistent snowpacks, the insulating effect can dominate to such an extent that the influence of snow cover on soil freezing depth surpasses that of air temperature (Zhang, 2005; Lawrence and Slater, 2010). Given this complexity, LSMs must realistically capture the interactions between snow insulation, soil moisture, and energy exchange to simulate SFG behavior reliably. However, there is a notable lack of studies over Russia that systematically compare LSMs driven by the same forcing dataset to examine SFG characteristics. Offline or standalone model configurations provide a strategic framework for such analyses, as it allows for the detailed examination of soil freeze-thaw dynamics without atmospheric feedbacks (Matthes et al., 2025).

This study investigates two LSMs: JSBACH and CLM, each utilizing different formulations for soil and snow processes, with both models driven by ERA5 reanalysis data (Hersbach et al., 2020). Additionally, the ERA5-Land offline model (ERA5L; Muñoz-Sabater et al. (2021)), which is also driven by ERA5 data, is included to provide context for the range of uncertainties in offline simulations, rather than serving as a baseline. Model performance is evaluated using in situ observational data across Russia, focusing on how well frozen soil characteristics are represented. By examining the simulated winter soil and snow characteristics and their coupling relationships, the research seeks to identify the respective strengths and weaknesses of each model. Specifically, the paper aims to address the following questions:

1. How well do JSBACH and CLM simulate SFG characteristics compared to observations?

2. What factors lead to biases in the timing of SFG?

3. How do errors in simulated snow cover characteristics affect biases in SFG characteristics?

The paper is structured as follows: Section 2 discusses the data and methods, which include descriptions of LSMs, model setup, in situ observational data, definition of SFG and snow cover characteristics, and the evaluation metrics. In Sect. 3, results and discussion are presented, which examine model improvements, large-scale spatial performance, site-level validation, and the causes of SFG bias. Finally, Sect. 4 provides the conclusions of the study.

2 Data and Methods

2.1 Land surface model description

2.1.1 JSBACH

The Jena Scheme for Biosphere-Atmosphere Coupling in Hamburg version 4 (JSBACHv4, Schneck et al., 2022), developed at the Max Planck Institute for Meteorology, is a part of ICON-Land, which is the land surface component of the ICOsahedral Nonhydrostatic (ICON) model (current ICON release 2025.04; https://gitlab.dkrz.de/icon/icon-model). It is the successor of

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the earlier JSBACHv3 (Reick et al., 2021) and provides the lower boundary for the atmosphere in coupled simulations and can also be applied in standalone mode as a comprehensive terrestrial ecosystem model. In ICON-Land, JSBACHv4 processes are organized modularly, allowing clear separation and flexible coupling of individual components (Jungclaus et al., 2022). JSBACH simulates key processes of the land surface system including the surface energy and water balance, soil hydrology and thermodynamics, vegetation phenology, photosynthesis, and the coupling of carbon and nitrogen cycles (Ekici et al., 2014; Brovkin et al., 2009; Raddatz et al., 2007). To account for heterogeneity within grid cells, JSBACH uses a tiling approach in which different land-cover types are simulated separately and then aggregated to the grid scale (Reick et al., 2021).

Recent model versions feature an enhanced representation of soil thermal and hydrological processes. Soil thermal properties such as heat capacity and conductivity are calculated dynamically as a function of the relative amounts of liquid water, ice, and organic matter in each layer (Ekici et al., 2014; Schneck et al., 2022). The standard discretization resolves five soil layers (node depths: 0.065, 0.254, 0.913, 2.902, and 5.7 m), enabling a realistic simulation of the active layer and deeper subsurface thermal buffering. Moisture phase transitions, incorporating the concept of supercooled water, are represented explicitly, and the presence of soil ice reduces hydraulic conductivity, thereby limiting vertical water transport, restricting the effective rooting depth to the thawed active layer, and improving the realism of plant water stress responses in boreal and PEFT regions. The current ICON release (2025.04) also incorporates enhancements to the soil hydrology scheme, such as vertical variations in organic matter content, which directly affect hydraulic and thermal properties, refinements in percolation and drainage through the inclusion of soil ice impedance, and a simple wetland parameterization (De Vrese et al., 2023). A more detailed description of the model parameterizations can be found in Reick et al. (2021, 2013); De Vrese et al. (2023); Schneck et al. (2022).

Snow is represented using a multilayer scheme with up to five layers, where the upper four layers are fixed at 5 cm and the bottom layer grows dynamically. The layers are hydrologically inactive, so meltwater transfer and refreezing within the snowpack are not simulated, and snowmelt infiltration into the soil is handled separately by the hydrology module (Ekici et al., 2014). Snow densification over time alters the thermal conductivity of the snowpack and its insulating effect on the underlying soil. Earlier formulations used a time-dependent approach with a constant decay factor, which often caused rapid densification and shallow snow depths. After updating the snow aging parameterization to depend on snow temperature (Heise et al., 2006), as described in Sect. 3.1, it has resulted in better representation of snow depth evolution and an improved simulation of the soil thermal regime beneath the snowpack. This update is now officially included in ICON-Land, and further details on the improvements are provided in Sect. 3.1.

2.1.2 CLM

The Community Land Model (CLM) is an advanced land surface modeling system developed as the terrestrial component of the Community Earth System Model (CESM; http://www.cesm.ucar.edu/models/cesm2/, Danabasoglu et al. (2020)) within the Community Terrestrial Systems Model framework (CTSM; https://github.com/ESCOMP/CTSM). Beyond CESM, CLM has also been integrated into other Earth System Modeling frameworks, including the Norwegian Earth System Model (NorESM), Euro-Mediterranean Center on Climate Change Earth System Model (CMCC-ESM2), etc. This study uses version

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Table 1. Key structural and parametric differences in soil and snow process representations between JSBACH and CLM

Model features	**JSBACH**	**CLM**
Total soil column depth	∼9.8 m	∼49.5 m (bedrock below 8.6 m)
Number of soil layers (default)	5 soil layers	20 soil layers + 5 bedrock layers
Number of snow layers (default)	5 (Hydrologically inactive)	10
Number of PFTs	11	15
Concept of supercooled water	Yes (Niu and Yang, 2006)	Yes (Niu and Yang, 2006)
Soil heat capacity scheme	De Vries (1953)	De Vries (1953)
Soil thermal conductivity	Johansen (1977)	Farouki (1981)
Bottom Boundary condition	Zero flux	Zero flux
Snow density formulation	Temperature dependent	Multi-process: compaction, destructive+constructive+melt metamorphism (Anderson, 1976)
Fresh snow density parameterization	50 kg/m^3 (constant)	Not constant (temperature dependent)
Snow thermal conductivity	Calonne et al. (2011)	Jordan (1991)

CTSM5.1.dev128_cm3_v1 (https://github.com/CMCC-Foundation/CTSM; hereafter referred to as CLM), which is designed to simulate complex land-atmosphere interactions and can be used in both regional and global modeling configurations, offering flexibility for diverse scientific applications. The soil thermal and hydrological schemes are tightly coupled, with phase change processes explicitly accounting for latent heat effects during freezing and thawing, and soil ice content dynamically affecting both thermal conductivity and hydraulic properties. This coupling is particularly important for simulating PEFT dynamics and seasonal freeze-thaw cycles in cold regions. The model uses a comprehensive representation of terrestrial hydrological processes spanning the full water cycle, from canopy interception through snowpack dynamics to deep groundwater interactions. A notable distinction of CLM compared to JSBACH is its explicit treatment of hydrologically active snow layers, allowing simulation of liquid water percolation, refreezing, and storage within the snowpack itself (Lawrence et al., 2019), processes that directly affect snow thermal properties and meltwater timing. The model uses a sophisticated subgrid hierarchy where each grid cell is divided into land units (vegetated, urban, glacier, lake), which are further subdivided into soil columns and plant functional types (PFTs). The default soil column configuration comprises 25 vertical layers, including 20 hydrologically and biogeochemically active soil layers overlying 5 bedrock layers, extending to a total depth of 49.5 m with bedrock starting below 8.6 m (Lawrence et al., 2008b), although this vertical discretization can be adjusted for specific applications. A detailed description of the CLM model physics and parameterizations is provided by Lawrence et al. (2019), with additional technical information available in the official model documentation (https://www.cesm.ucar.edu/models/clm/). Table 1 shows a side-by-side comparison of the key structural and parametric formulations in JSBACH and CLM.

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2.2 Standalone model setup

To study frozen soil processes, JSBACH and CLM were implemented in standalone mode to enable detailed evaluation of their frozen soil representations without feedbacks from coupled atmospheric dynamics. The models were forced with ERA5 reanalysis (Hersbach et al., 2020) data at a 3-hourly temporal resolution, providing globally consistent atmospheric forcing at a spatial resolution of 31 km, which ensures reliable surface meteorological conditions for multi-decadal simulations. For CLM, the BGC-CROP biogeochemistry crop model was activated (not used for JSBACH), and it was spun up in two phases (Koven et al., 2013; Lawrence et al., 2019): an initial preindustrial run to equilibrate carbon and nitrogen pools, followed by a historical simulation from 1850 to 1940 (forced with default Global Soil Wetness Project (GSWP3) data) to establish realistic component balances and interannual variability, with the final spin-up state serving as the initial condition for the transient simulation from 1941 to 2022, forced with 3 hourly ERA5 data. JSBACH was run continuously from 1941 to 2022 without the two phase spin-up used for CLM. The analysis for both models focused on 1986-2022 to match the availability of observational datasets. The models were run on their native grids using a 30-minute time step, with JSBACH on the ICON R02B06 grid (40 km resolution) and CLM on an approximately 0.5^o grid (0.47^o × 0.63^o latitude-longitude resolution), producing daily output for analysis.

2.3 In situ observations

To assess the performance of the models and to better understand the magnitude of differences between them, we first evaluated the simulated frozen soil characteristics over the entire mid- to high-latitude region and then conducted a detailed site-level analysis using 26 meteorological stations from the All-Russian Research Institute of Hydrometeorological Information-World Data Centre (RIHMI-WDC; http://meteo.ru/data/), indicated by red dots in Fig. 1, as this dataset provides uniform soil temperature records at multiple depths along with other key variables required for this study. The stations are located in regions of Russia (mainly Siberia) that experience a complete annual freeze-thaw cycle and provide continuous records from 1986, with the time series carefully screened to avoid large data gaps. If a station remains frozen to the last layer depth throughout the entire observation period, it is classified as PEFT and excluded from the analysis. Stations with only shallow or short-duration freezing were excluded because this study focuses on areas where freezing extends below the top soil, strongly influencing seasonal soil thermal and hydrological regimes and allowing effective evaluation of the models' ability to represent freeze-thaw dynamics. This ensures that the selected sites are representative of zones where freeze-thaw dynamics are most pronounced and have significant implications for land-atmosphere interactions. The following variables were used for the site-level analysis: soil temperature at multiple depths (20 cm, 40 cm, 80 cm, 120 cm, 160 cm, 240 cm, and 320 cm), 2 m air temperature, and snow depth. To facilitate a direct model-to-observation comparison, the gridded model output was spatially interpolated to the geographic coordinates of each station. Details of the 26 selected stations, including their three letter codes, geographic coordinates, and elevations, are provided in Table A1, where the sites are arranged in increasing order of latitude.

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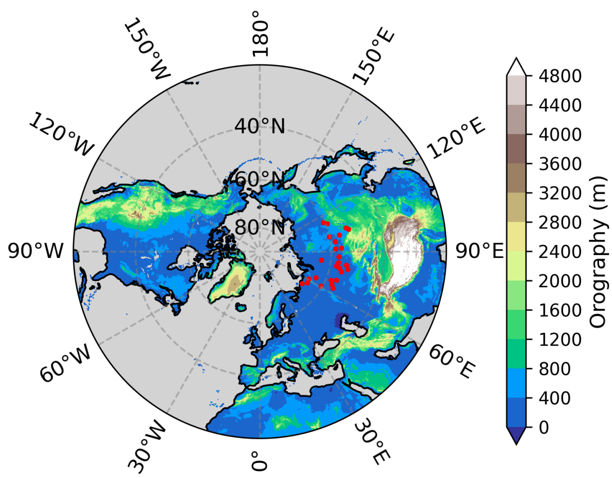

Figure 1. Orography map with red dots showing 26 RIHMI-WDC observational sites used for site-level analysis.

2.4 Definition of SFG and snow cover characteristics

To quantitatively evaluate model performance in representing cold region processes during the annual freezing-thawing cycle (September 1 of a given year to August 31 of the following year (Luo et al., 2020)), several parameters were determined using simulated and observed data. These are defined as follows:

1. SFG Onset and End date: Onset (freeze) occurs on the first day after September 1 when the daily mean soil temperature at 20 cm depth falls below 0 ^{o}C and remains below this threshold for at least 15 consecutive days. End (thaw) occurs when, following the onset, the daily mean soil temperature at 20 cm depth rises above 0 ^{o}C and remains above this threshold for at least 15 consecutive days.

2. SFG Duration (SFGD): The number of days between the SFG onset and end dates, representing the complete period during which the soil at the 20 cm depth remains SFG.

3. Snow cover Onset and End date: Onset (snow accumulation) occurs on the first day when snow depth exceeds 5 cm (Akyurek et al., 2023) and is maintained for at least the following 10 consecutive days (Bender et al., 2020). End (snow disappearance) occurs when, following the onset, snow depth falls below 5 cm and remains below this threshold for at least 10 consecutive days.

4. Snow Cover Duration (SCD): This is the total period of persistent snow cover, calculated as the number of days between the snow cover onset date and end date, when snow depth falls below the 5 cm threshold at the end of the season.

5. Maximum Freeze Depth (MFD): The MFD for each year is estimated using the approach of Frauenfeld et al. (2004), based on monthly soil temperature profiles. It is defined as the maximum depth of the 0 ^{o}C isotherm within the annual

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freeze-thaw period. This parameter provides an index of the intensity of seasonal freezing. It should be noted that MFD differs from active layer thickness (ALT) as MFD represents the maximum depth of soil freezing in SFG, while ALT refers to the annually thawed layer in PEFT regions. In addition to the MFD, the annual air freezing index (FI) and thawing index (TI) provide an indication of the intensity of the seasonal freeze-thaw cycle that governs SFG. FI is calculated as the cumulative sum of daily mean 2 m air temperatures below 0 ^{o}C during the freezing period, considered from 1 July of a given year to 30 June of the following year. TI is the cumulative sum of daily mean 2 m air temperatures above 0 ^{o}C during the thawing period, defined from 1 January to 31 December of the same year (Zhang et al., 2005; Steurer and Crandell, 1995).

2.5 Evaluation metrics

The evaluation of model performance against in situ observations was conducted using the following standard statistical metrics: the Root Mean Square Error (RMSE), Mean Bias Error (MBE), and Pearson Correlation Coefficient (R). The RMSE quantifies the average magnitude of the differences (errors) between simulated and observed values, the MBE indicates the systematic tendency of the model to over- or under-estimate observations, and R measures the strength and direction of the linear relationship between simulated and observed values. These metrics are defined as:

$$RMSE = \sqrt{\frac{1}{N}\sum_{i=1}^{N}(S_i - O_i)^2} \tag{1}$$

$$MBE = \frac{1}{N}\sum_{i=1}^{N}(S_i - O_i) \tag{2}$$

$$R = \frac{\sum_{i=1}^{N}(S_i - \bar{S})(O_i - \bar{O})}{\sqrt{\sum_{i=1}^{N}(S_i - \bar{S})^2 \sum_{i=1}^{N}(O_i - \bar{O})^2}} \tag{3}$$

where S_i and O_i denote the simulated and observed values at time step i, $\bar{S}$ and $\bar{O}$ are their respective means, and N is the total number of data points.

In addition to the above metrics, the timing and duration of SFG and snow cover characteristics were evaluated using the median bias and interquartile range (IQR), which describe the central tendency and spread of model deviations from observations. These metrics were selected because they are less sensitive to outliers and provide a more representative summary of model performance across stations and years with varying distributions. The difference between simulated and observed values was calculated for each station and year for each SFG characteristic as:

$$\Delta_{s,y} = SFG_{s,y}^{sim} - SFG_{s,y}^{obs} \tag{4}$$

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where $s = 1, 2, \ldots, 26$ denotes the stations and $y = 1986, 1987, \ldots, 2022$ denotes the years. These differences were then pooled across all stations to form a single dataset of deviations.

The overall median bias was computed as the median of all pooled differences:

$$\text{Median Bias} = \text{median}(\Delta_{s,y}) \tag{5}$$

and the IQR, representing the spread of model deviations, was defined as:

$$IQR = Q_{75}(\Delta_{s,y}) - Q_{25}(\Delta_{s,y}) \tag{6}$$

where Q_{75} and Q_{25} are the 75th and 25th percentiles of the pooled distribution, respectively. This approach provides an overall assessment of model performance across all stations and years, capturing both the systematic bias and the variability in simulating SFG characteristics.

3 Results and Discussions

3.1 JSBACH snow aging parameterization

JSBACH simulation output exhibited unrealistic cold soil temperatures relative to CLM and ERA5L over cold regions of the Northern Hemisphere, with the underlying cause of this bias and the corresponding solution documented in a dedicated technical report archived on Zenodo (see code availability). A detailed investigation revealed that one major cause of this cold bias was the old snow densification scheme (hereafter referred to as the old setup), which employed a constant decay factor (Verseghy, 1991). This formulation caused overly rapid snow densification and consequently shallower snowpacks. The reduced snow depth critically diminished the snowpack's insulating capacity, allowing the soil to be more strongly influenced by extreme cold air temperatures and leading to excessive soil cooling during winter. This systematic behavior pointed to a deficiency in the model's representation of snow processes.

To address this issue, a revised snow aging formulation (hereafter referred to as the new setup) was implemented by introducing a temperature-dependent decay factor (Heise et al., 2006). This modification allows snow densification to vary dynamically with snow temperature, improving the representation of snow aging and the temporal evolution of snow density and thickness. Figure B1 illustrates the difference in the seasonal climatological mean (1986-2022) of snow depth and soil temperature at 20 cm depth between the new and old setups. The new parameterization increases snow depth by up to 15-20 cm in cold regions, thereby enhancing the simulated snow insulation effect resulting from increased snowpack and raising soil temperatures by approximately 10-20 ^{o}C. In contrast, some coastal regions exhibit a reduction in snow depth with almost no corresponding change in soil temperature, likely due to the strong dependence on snow temperature, which may lead to reduced snow accumulation under humid, cold coastal conditions. Overall, this change has improved JSBACH's performance by making its simulated snow and soil thermal states more similar to those of CLM and ERA5L. The updated snow parameterization has since been incorporated into the ICON model repository, and all subsequent analyses presented in this study are based on the new setup.

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3.2 Model performances

3.2.1 Frozen ground distribution

Accurately simulating the spatial extent of frozen ground provides a fundamental test of a LSM's ability to represent the cold region soil thermal state. Figure 2 shows the simulated frozen ground distribution, classified into PEFT and SFG, by ERA5L, JSBACH, and CLM. The resulting distributions were evaluated against reference values reported by Ran et al. (2022) for a PEFT extent of $13.60-18.97\times10^6 km^2$ and by Zhang (2003) for an SFG extent of $48.12\times10^6 km^2$.

Comparison of simulated frozen ground extent reveals notable inter-model differences relative to the reference distribution. ERA5L simulates a PEFT area of $7.26\times10^6 km^2$, representing only 38-53% of the reference range and indicating a substantial underestimation of year-round soil freezing conditions. In contrast, JSBACH simulates $13.76\times10^6 km^2$ of PEFT, which is close to the lower bound of the reference range, suggesting a more realistic representation of year-round freezing conditions, while CLM demonstrates intermediate performance, simulating $10.63\times10^6 km^2$ of PEFT, which is below the reference range. For SFG, both ERA5L ($47.35\times10^6 km^2$) and JSBACH ($49.76\times10^6 km^2$) produce estimates that are in close agreement with the reference value, with ERA5L underestimating and JSBACH overestimating the extent by about $\sim1-1.5\times10^6 km^2$. CLM simulates an SFG extent of $49.4\times10^6 km^2$, which is similar to that simulated by JSBACH. Spatial analysis reveals that CLM tends to classify the Tibetan Plateau as SFG rather than PEFT, which is inconsistent with known high-altitude PEFT distributions in this region (Cao et al., 2023; Ran et al., 2021). This misclassification leads to reduced PEFT extent and improving the representation of PEFT processes over the Tibetan Plateau would likely bring CLM's simulated extents closer to reference estimates while enhancing the reliability of its frozen ground simulation. Overall, among the three models, JSBACH provides the most realistic representation of frozen ground distribution, capturing both PEFT and SFG extents in closest agreement with Ran et al. (2022). While CLM shows potential, its performance is limited by the need for improved soil thermal parameterization, whereas ERA5L is mainly constrained by its shallow soil structure, restricting its capability to represent deep PEFT dynamics (Cao et al., 2020, 2022).

3.2.2 Spatial patterns

This section examines spatial variations in simulated soil moisture (SM), soil temperature (ST), and snow depth (SD) to evaluate inter-model consistency and identify key differences in the representation of variables governing frozen ground dynamics. A comparative assessment of JSBACH and CLM, alongside ERA5L, provides a systematic evaluation of the relative performance of each model, highlighting their respective strengths, limitations, and sources of bias in simulating cold region soil and snow processes.

Figure 3, 4, and 5 presents the spatial climatology (1986-2022) of SM, ST at 20 cm depth, and SD, respectively, averaged for the four seasons over mid- to high-latitude regions. Analysis of total SM, defined here as the combined liquid and ice content in the top 20 cm, reveals distinct hydrological behaviors among the models. ERA5L and JSBACH exhibit comparable magnitudes of total SM in the upper 20 cm and show minimal seasonal variability, whereas CLM simulate more pronounced seasonal

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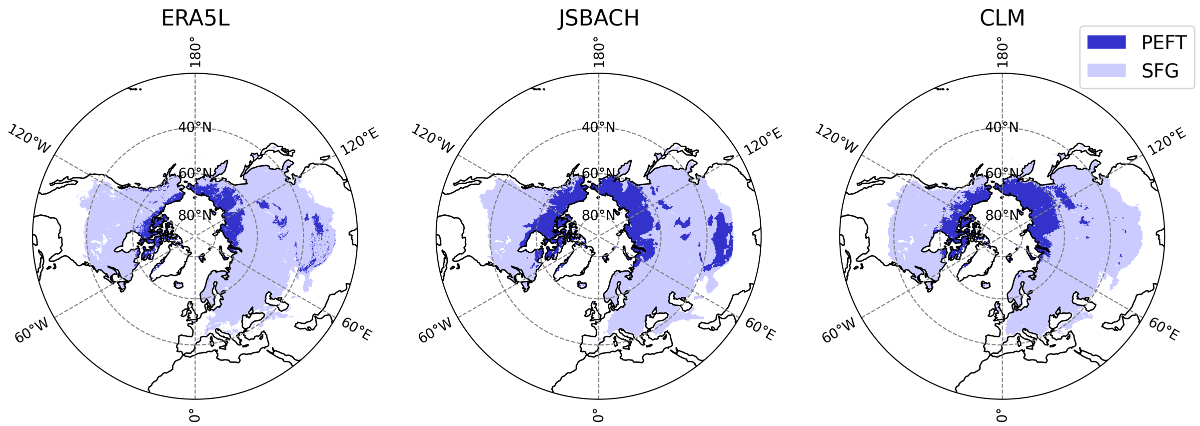

Figure 2. Distribution of frozen ground classified into SFG and PEFT regions based on modeled soil temperatures from ERA5L, JSBACH, and CLM. Grid cells are classified as SFG when frozen for at least 15 days per year, and as PEFT when frozen year-round in any soil layer for more than 50% of the years from 1986 to 2022.

cycles in total SM. JSBACH and CLM separately represent soil liquid and ice contents, which are not provided individually in ERA5L. In cold regions, CLM shows a larger fraction of total SM in the frozen phase compared to JSBACH (Fig. B2), indicating stronger near-surface soil freezing. Comparison of the top layer liquid SM with ESA-CCI satellite observations (Fig. B3) shows that CLM generally underestimates liquid SM across boreal and arctic regions. Although summer thaw increases liquid water content, autumn SM remains lower than observed, suggesting persistently drier near-surface conditions before freeze-up, which may further promote rapid freezing in autumn.

The spatial patterns of ST at 20 cm depth further highlight model disparities. CLM exhibits relatively warm ST over the Tibetan Plateau, corresponding to its misclassification of PEFT as SFG in the same region, although its cold soil temperature extent in other regions during winter appears generally reasonable. These warmer conditions may be attributed to the soil thermal conductivity scheme in CLM, which has been reported to perform less effectively over soils containing broken pieces of stones and gravels, such as those prevalent on the Tibetan Plateau (Pan et al., 2024; Yang et al., 2021, 2022). JSBACH captures the overall extent of cold STs well, reproducing frozen conditions across high-latitude regions during winter and the subsequent thaw in spring and summer. However, without direct comparison with observations, it remains uncertain whether JSBACH systematically over- or under-estimates ST in specific regions. ERA5L simulates less cold ST than both JSBACH and CLM across high-latitude regions, particularly during winter and spring. The comparatively warmer soil conditions are consistent with its underestimation of PEFT extent and likely originates from its soil or snow parameterizations within the ERA5L framework. Previous studies have also reported similar limitations of ERA5L in representing near-surface thermal conditions during cold seasons (Cao et al., 2020, 2022).

Figure 5 illustrates model performance in simulating SD across the mid- to high-latitude Northern Hemisphere. JSBACH simulates shallower snowpacks than the other two models across most regions, even after the improvements introduced in Sect. 3.1.

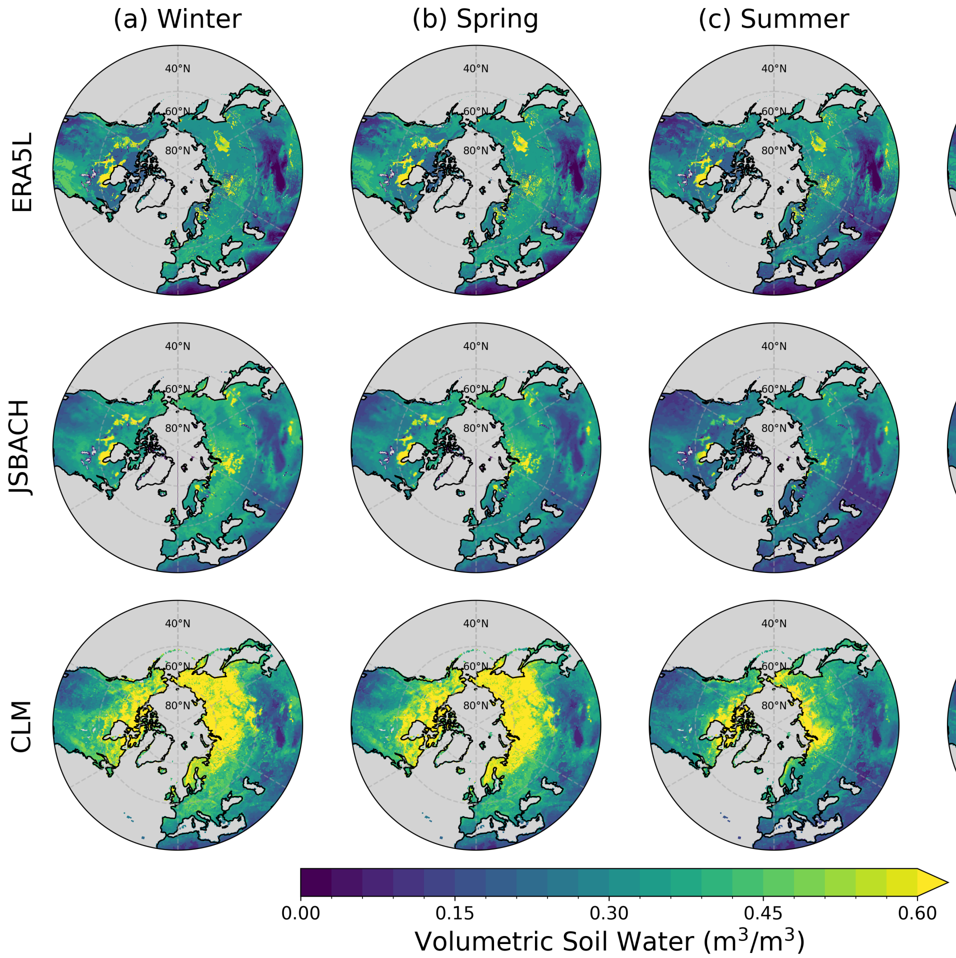

Figure 3. Seasonal climatology (1986-2022) of total soil moisture (liquid + ice) at 20 cm depth across Northern Hemisphere mid- to high-latitude regions (30^o-90^o N), as simulated by ERA5L (top row), JSBACH (middle row), and CLM (bottom row).

The reduced SD likely strengthens soil-atmosphere coupling, allowing greater surface heat loss and contributing to colder STs, which can be confirmed through the site-level analysis (Sect. 3.3). In contrast, CLM produces SD patterns broadly similar to ERA5L, though with higher values in several regions, potentially enhancing insulation and reducing soil freezing intensity. For validation purposes, ERA5 provides more reliable SD estimates than ERA5L because it assimilates both ground-based and satellite snow observations (Cao et al., 2022; Kouki et al., 2023; Sarpong and Nazemi, 2025). A comparison between ERA5L and ERA5 (Fig. B4) further indicates that ERA5L tends to overestimate SD during winter and spring.

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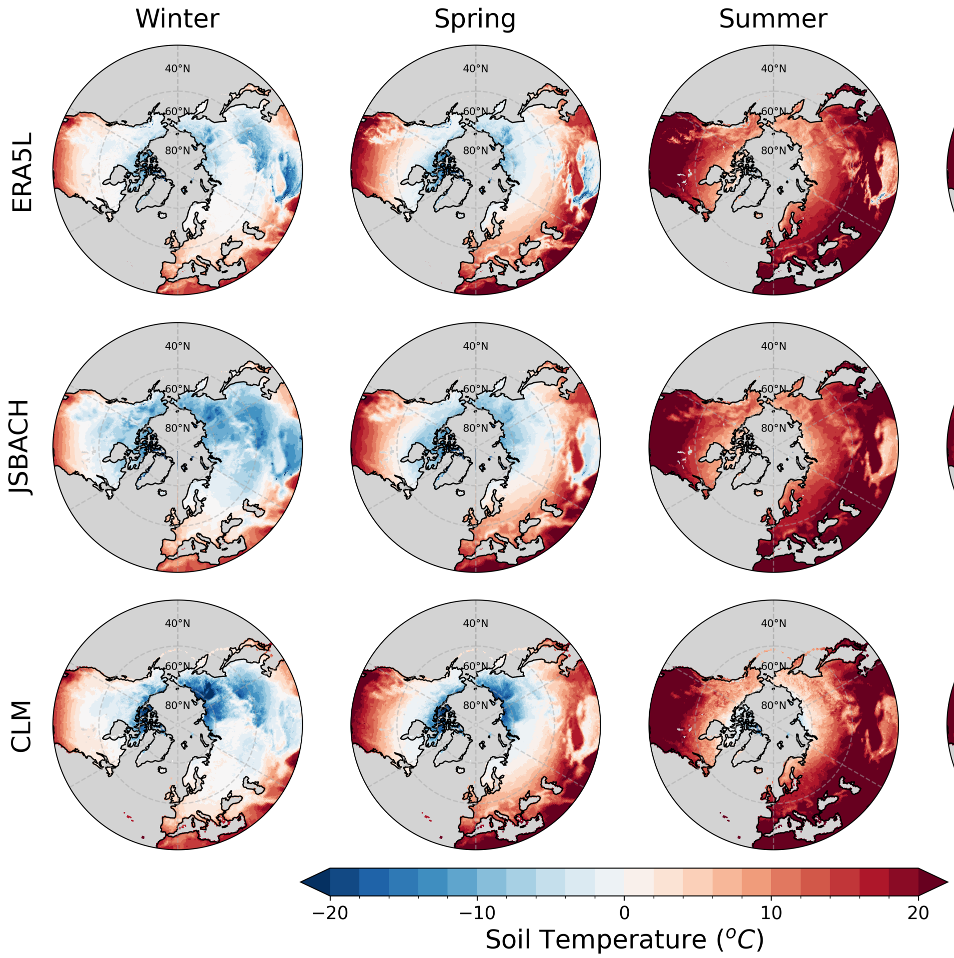

Figure 4. Same as Fig. 3 but for soil temperature at 20 cm depth.

3.3 Site-level Evaluation

3.3.1 Comparison of soil temperature, snow depth and air temperature

This section complements the spatial analysis by providing a direct observation-based evaluation of model performance through comparison with in situ data from 26 stations across Russia (shown in Fig. 1). Figure 6 shows box plots of the monthly means of 2 m air temperature (AT), SD, and ST at 20 cm depth during the winter season, along with the corresponding statistical metrics (RMSE, MBE, and R) computed across all stations. Such comparative analyses remain valuable for identifying consistent model deficiencies and systematic performance characteristics.

The comparison of AT indicates that JSBACH demonstrates the lowest error metrics (RMSE = 1 ^oC), followed by ERA5L (RMSE = 1.3 ^oC) and CLM (RMSE = 3.6 ^oC). Although all models were forced with same atmospheric fields from ERA5,

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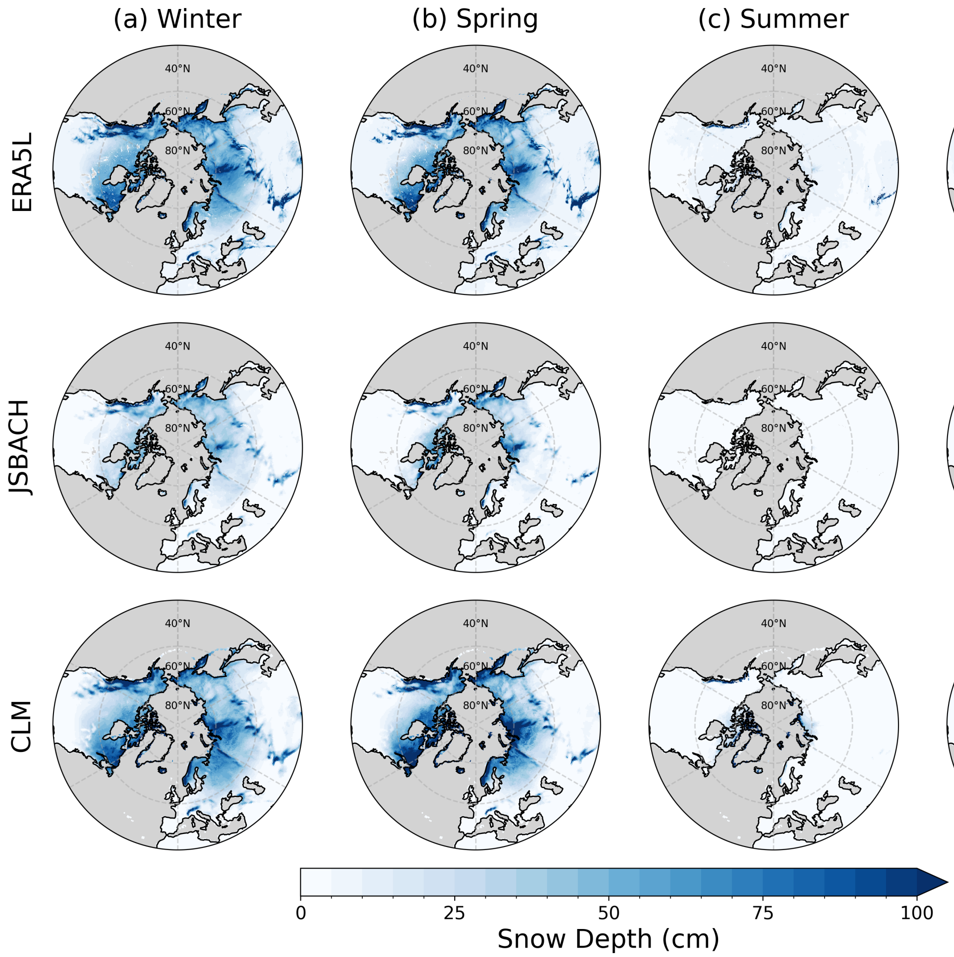

Figure 5. Same as Fig. 3 but for snow depth.

variations in AT arise due to model specific adjustments or corrections applied within their respective structures. Both ERA5L and CLM temperatures were adjusted using lapse rate correction, and in CLM the 2 m air temperature includes surface roughness length and displacement height. JSBACH exhibits low AT errors but shows a cold soil bias (MBE = -3.7 ^oC, RMSE = 4.9 ^oC), likely linked to insufficient snow cover (MBE = -14.2 cm, RMSE = 18.9 cm), that reduces snow insulation and increases heat loss from the soil. In contrast, CLM exhibits a cold air temperature bias (MBE = -3.4 ^oC) and excessive snow accumulation (MBE = +12.2 cm, RMSE = 18.5 cm), while showing a smaller soil temperature bias (MBE = +0.1 ^oC, RMSE = 2.9 ^oC) relative to the other two models. The large RMSE and MBE values associated with SD across all models highlight the persistent challenge of accurately representing snow accumulation and melt processes at the observation sites. These discrepancies in simulated snow cover directly influence the soil thermal regime through snow-soil thermal coupling, which is examined in the subsequent sections. The inter-model differences mainly arise from how each model represents key physical

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processes such as soil heat conduction, snow insulation, and surface energy exchange, rather than from differences in spatial resolution alone. This study recognizes the potential spatial representativeness issues that arise when comparing gridded model outputs with point based in situ observations. Differences in spatial resolution, grid-cell heterogeneity, and boundary condition datasets describing soil properties likely contribute to some of the simulated biases. In addition, since all simulations are driven by the same ERA5 meteorological forcing and frozen soil dynamics in high latitude regions are known to be sensitive to forcing uncertainties (Guo et al., 2012; Lawrence et al., 2012), potential biases in ERA5, particularly related to snow in mid to high latitude regions (Kouki et al., 2023; Bian et al., 2019; Guo and Yang, 2022), may introduce a common baseline error. Although such scale mismatches and forcing uncertainties are unavoidable limitations in global model configurations, the consistent differences observed across sites still provide meaningful information about the systematic strengths and weaknesses of each model and reflect contrasts in the internal snow and soil process representations of the models.

3.3.2 Snow insulation effect

The systematic bias in SD presented in previous section has important implications for surface energy exchange and soil thermal regulation, as the insulating capacity of the snowpack strongly influences winter STs. JSBACH's underestimation of SD and the overestimation by CLM and ERA5L suggest that differences in simulated STs may arise not only from atmospheric forcing or soil parameterizations but also from variations in the snowpack's insulation capacity. This relationship was also evident in Sect. 3.1, where increasing SD produced notable changes in ST. To assess the snow insulation effect, we adopted the approach used in earlier model evaluations by Wang et al. (2016).
Figure 7 (upper panel) shows the relationship between the temperature difference between soil and air (TDSA) and SD using monthly mean values during the cold season. Observations indicate that TDSA increases with snow accumulation and stabilizes around 35-45 cm, beyond which additional SD exerts little influence on the temperature offset. ERA5L reproduces the observed relationship between SD and thermal insulation, whereas CLM tends to overestimate the insulating effect. A lack of data under thin snow conditions (0-5 cm) is also evident in CLM, while JSBACH underestimates the insulation effect, exhibiting a weaker TDSA-SD relationship associated with its generally thinner snow cover.

To complement the TDSA-SD analysis, the probability distribution functions (PDFs) of TDSA were examined under two snow regimes, enabling a more detailed evaluation of temperature difference associated with varying snow conditions (Fig. 7, lower panel). This approach allows comparison of the modal values and spread of the TDSA distributions, with the dataset divided into shallow ($\leq$20 cm) and thick ($\geq$45 cm) snow categories to assess how SD influences the insulating effect (Wang et al., 2016). The observed TDSA distributions show modal values of approximately 9 ^{o}C for shallow snow and 15 ^{o}C for thick snow, with the modal TDSA shifting by about 6 ^{o}C toward higher values under deeper snow, confirming the stronger insulating effect of thick snowpacks. ERA5L closely replicates the observed distributions for both regimes, showing a comparable spread and slightly different modal value for shallow snow depth. In contrast, CLM exhibits a narrow distribution that is sharply peaked at a higher TDSA of around 11 ^{o}C for shallow snow, exceeding the observed modal value, with a similar overestimation evident

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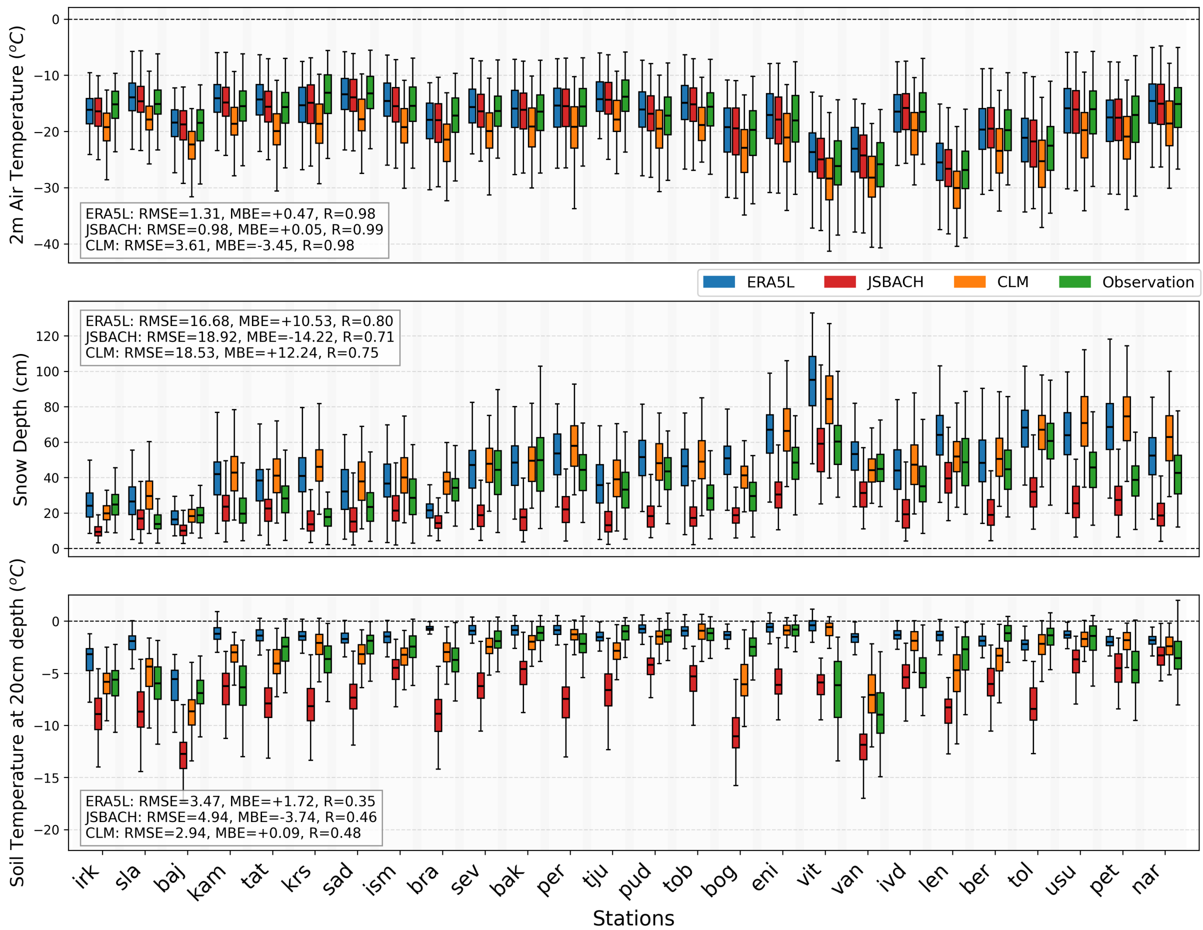

Figure 6. Comparison of simulated and observed monthly mean AT, SD, and ST at 20 cm depth during the cold season (DJF) across 26 stations. Box plots show the distribution of monthly values, and the corresponding statistical metrics: RMSE, MBE, and R are presented for ERA5L (blue), JSBACH (red), and CLM (orange) relative to observations (green).

under deeper snow conditions. JSBACH, on the other hand, shows lower modal values ($\sim 7\ ^{o}$C), consistent with its weaker insulation performance under thin snow cover and its limited ability to capture extreme events. The stepwise shift observed in JSBACH for thick snow likely results from the limited number of deep snow cases in the model.

3.3.3 SFG and snow cover characteristics

Following the assessment of snow insulation effects, it is important to evaluate how well the models reproduce the observed intensity of frozen ground conditions. Figure 8 compares the simulated and observed Maximum Freezing Depth (MFD) across all stations, illustrating the intensity of SFG for ERA5L, JSBACH, and CLM. ERA5L is structurally constrained by its shallow soil column($\sim$2 m), which limits its ability to capture the observed deep freezing (up to 3.2 m). In ERA5L, eight out of the

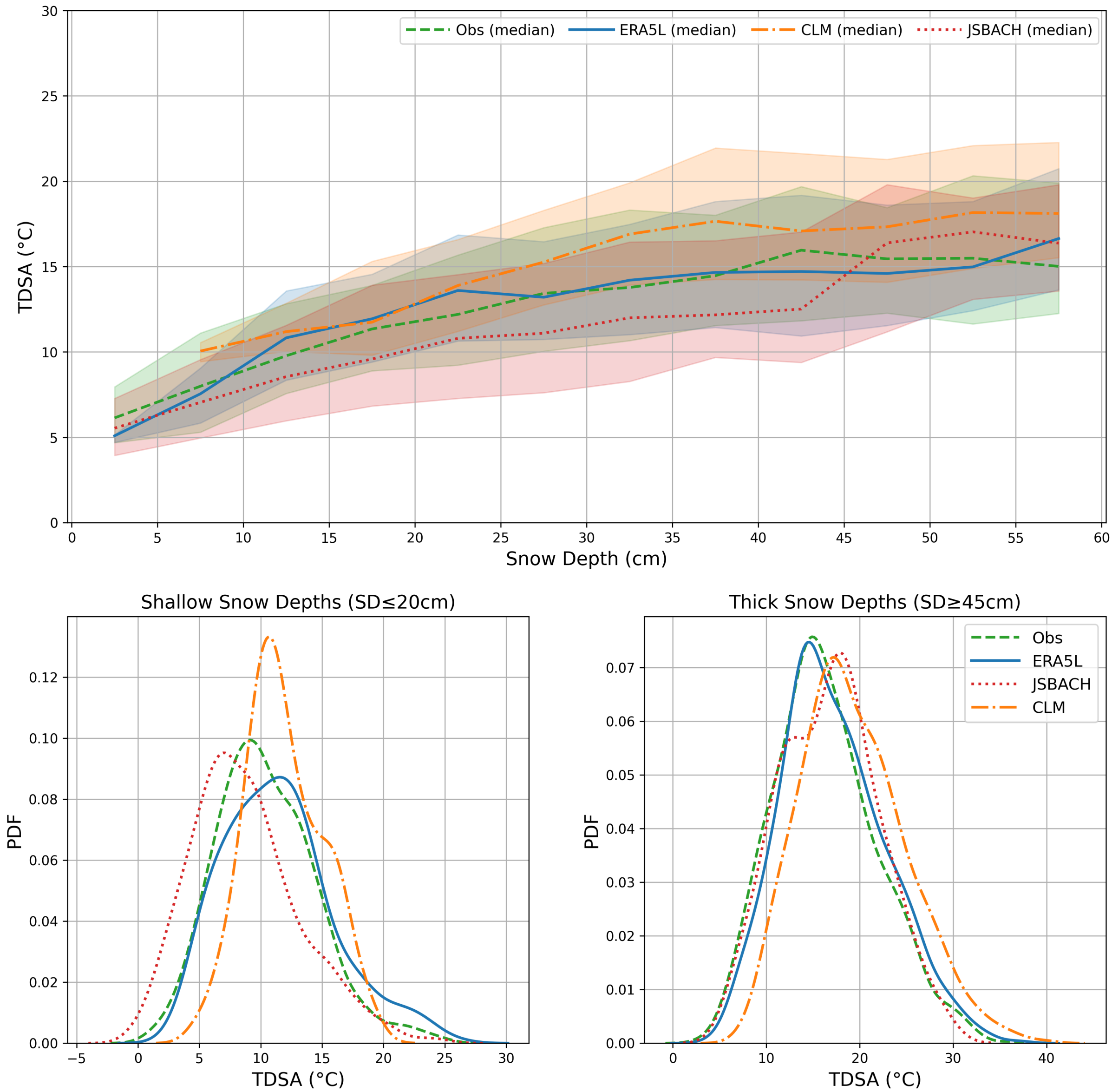

Figure 7. Relationship between wintertime monthly mean SD and soil-air temperature difference (TDSA) across observational and model datasets, shown as median TDSA (dotted lines, 5 cm SD bins) and IQR (shaded) for observations (Obs; green), ERA5L (blue), JSBACH (red), and CLM (orange) in upper panel. Lower panel shows probability density functions (PDFs) of TDSA for shallow (SD $\leq$ 20 cm) and thick (SD $\geq$ 45 cm) snow conditions.

26 stations reached the model's maximum soil depth for several years, which are not shown in the figure. JSBACH generally overestimates MFD (MBE = 1.7 m, RMSE = 2.2 m), with a low correlation coefficient (R = 0.36), suggesting that insufficient snow insulation allows deeper ground freezing. For four stations, freezing reached the maximum soil depth in some years.

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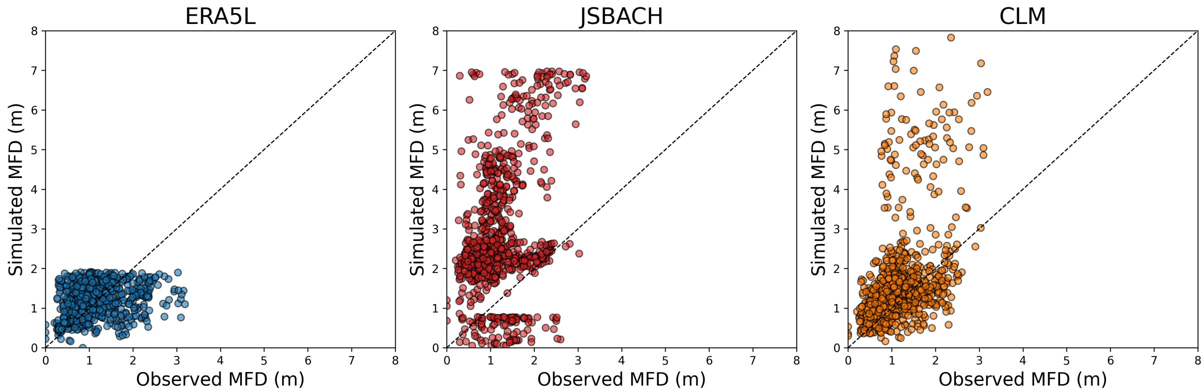

Figure 8. Scatter plot of observed versus simulated Maximum Freezing Depth (MFD) across all stations, showing annual MFD values for ERA5L (blue), JSBACH (red), and CLM (orange). The 1 : 1 line indicates perfect agreement between simulations and observations.

CLM performs comparatively better (MBE = 0.6 m, RMSE = 1.3 m, R = 0.43), with most data points clustering near the 1 : 1 reference line and within the expected range, except at a few stations where freezing extended beyond 4 m.

To get a full picture of model behavior, it is also essential to examine the timing and duration of SFG and snow cover. Figure 9 compares the simulated and observed timing of seasonal transitions in frozen ground and snow cover by showing the biases in onset, end date, and duration across all stations, with panels (a-c) representing SFG characteristics and (d-f) representing snow cover characteristics. The SFG onset timing indicates that all models simulate an earlier freeze onset at 20 cm depth relative to observations across most stations, with median biases ranging from -5 to -11.5 days. ERA5L shows a bias of -5 days (IQR: -19 to $+1$ days), CLM -7 days (IQR: -19 to 0 days), and JSBACH -11.5 days (IQR: -24 to -4 days). This early freezing suggests that soil cooling begins prematurely, reflecting challenges in representing autumn land-atmosphere interactions related to soil thermal characteristics and snow evolution processes.

Early freezing onset and delayed thaw together extend the frozen period in all models (median bias: 19 to 26 days) as shown in Fig. 9c. The SFG end date or thaw date (Fig. 9b) is delayed across all models, with median biases ranging from $+9$ to $+16$ days. JSBACH exhibits the smallest bias ($+9$ days), whereas CLM and ERA5L show larger positive biases ($+12$ to $+16$ days) and greater variability. The timing and duration of snow cover (Fig. 9d-f) further explain these SFG biases. JSBACH simulates delayed snow onset (median bias = $+10$ days) and earlier snow disappearance (median bias = -15 days), resulting in a shorter snow cover duration (median bias = -27 days). This outcome is consistent with its underestimated SD and weaker insulation, which enhance soil heat loss and allow for deeper soil freezing. We speculate that snow density still needs adjustment, as it strongly influences snow depth, consequently, snow insulation. Cold soil conditions and low snow accumulation are interlinked, with snowfall occurring but snow depth remaining low, which delays reaching the 5 cm threshold and reduces insulation, allowing the soil to stay cold, and further research is needed to identify all the contributing factors. In contrast, ERA5L and CLM simulate earlier snow onset (median bias = -3 to -7 days) and delayed melt (median bias: $+9$ to $+15$

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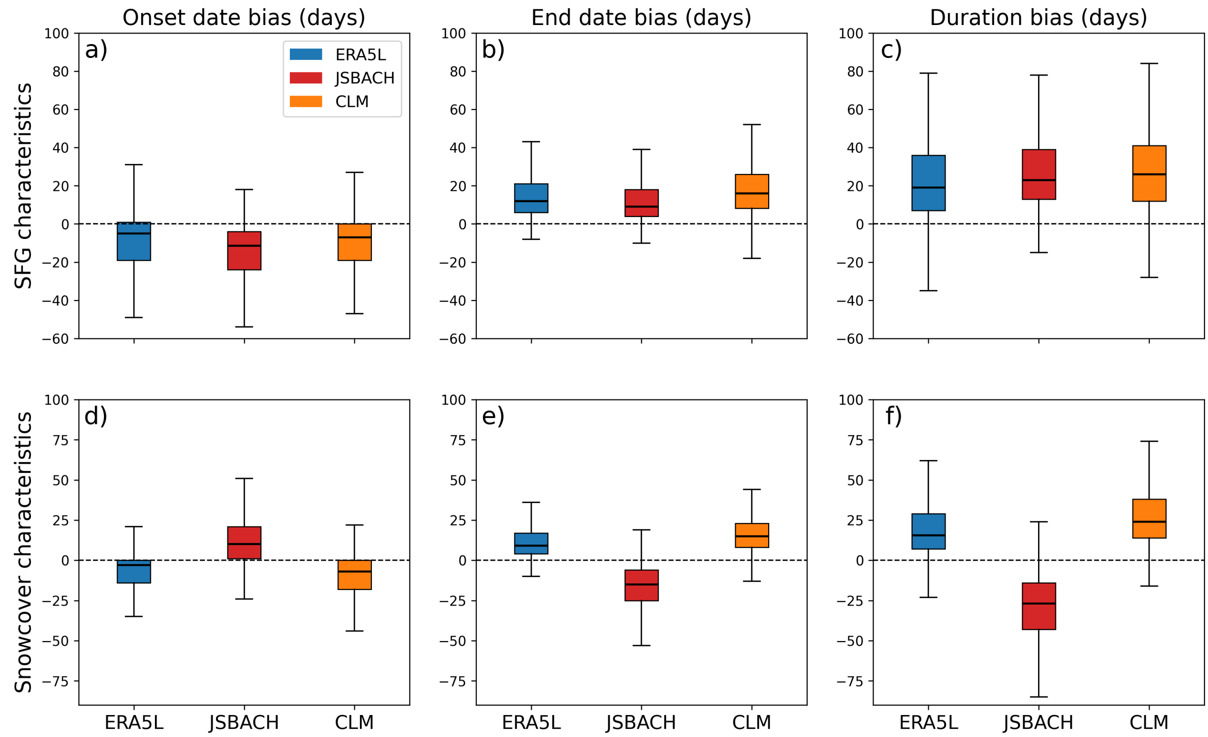

Figure 9. Comparison of simulated and observed timing of SFG and snow cover characteristics. The figure shows biases in onset date (first column), end date (second column), and duration (third column) across all stations. Panels (a-c) represent biases in SFG characteristics, while panels (d-f) show biases in snow cover characteristics. Positive (negative) values indicate later (earlier) occurrence relative to observations.

days), resulting in an extended snow cover duration (+15.5 to +24 days). The prolonged frozen soil period in these models likely arises from excessive snow accumulation and insulation, which inhibit soil-atmosphere heat exchange and delay thaw, whereas JSBACH's weaker snow dependence and stronger AT control lead to smaller variability in thaw timing.

3.3.4 Factors affecting seasonal freeze-thaw variability in cold region soils

In the previous sections, model biases in the timing, duration, and intensity of SFG have been examined. To gain a comprehensive understanding of the underlying mechanisms, it is important to investigate how these SFG characteristics respond to environmental drivers. Previous research has predominantly focused on the direct effects of AT or SD on MFD (Peng et al., 2017; Frauenfeld et al., 2011; Wang and Chen, 2022) or SFG duration (Luo et al., 2020; Li et al., 2021), often neglecting the onset timings of freeze and thaw and their relationship with snow cover characteristics. While MFD reflects the cumulative effect of continuous soil freezing, the processes controlling its formation and disappearance remain insufficiently understood (Wang and Chen, 2024). This section analyzes the relationships between SFG characteristics and snow cover properties, in-

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cluding timing, duration, and depth, along with seasonal ATs and cumulative thermal indices, specifically annual freezing and thawing degree days, offering an integrated perspective on the factors influencing seasonal freeze-thaw variability.

Figure 10 shows statistically significant correlation matrices between SFG and snow cover characteristics, seasonal ATs, winter SD, as well as thermal indices for observations and models. This correlation analysis reveals distinct differences between observed relationships and their representation in the models. In observations (Fig. 10a), both SFG onset and snow cover onset exhibit weak correlations with all seasonal variables ($R \leq 0.30$), indicating that the freeze and snow accumulation events are not primarily controlled by mean seasonal conditions. Instead, these transitions are more strongly influenced by short-term weather variability, microclimatic conditions, and local soil properties. The SFG end date shows a close relationship with snow cover end date ($R = 0.63$), spring AT ($R = -0.57$), and the TI ($R = -0.50$), indicating that snow persistence and accumulated melt energy largely determine the timing of soil thaw. The SFG duration exhibits the strongest association with the autumn AT ($R = -0.52$; (Smith et al., 2004)) and SFG end date ($R = 0.52$), followed by FI ($R = 0.51$), with a moderate dependence on snow cover duration also ($R = 0.37$). These relationships suggest that colder autumns and higher cumulative freezing energy extend the frozen period. The weaker correlation with SFG onset ($R = 0.10$) compared to the stronger link with SFG end date ($R = 0.52$) further demonstrates that the length of the frozen season is more sensitive to processes delaying thaw in spring than to those initiating freezing in autumn. Altogether, the observations indicate that snow cover governs thaw timing, and together with the intensity of the annual FI, determines the duration of seasonal ground freezing.

ERA5L (Fig.10b) and CLM (Fig. 10c) reproduce several of the observed dependencies but often amplify the strength of correlations, suggesting an excessive sensitivity of the simulated soil thermal regime to AT and snow controls. For instance, both models exhibit very strong correlations between the SFG end date and snow cover end date ($R > 0.9$), and between snow cover end date and winter SD (R $\approx$ 0.7), which substantially exceed the observed relationships. Moreover, both models do not account for the high heterogeneity in local conditions present at observation sites, as land-surface properties are spatially smoothed, which likely contributes to the strong dependence of snow cover onset on seasonal variables, particularly autumn air temperature (AT). In addition, precipitation phase partitioning in the models is governed by grid-scale temperature thresholds, making the onset of snow accumulation primarily temperature driven. Consequently, local factors that modify the near-surface temperature regime, such as vegetation dynamics, soil thermal properties, microtopography, and surface energy exchanges, cannot be represented as specifically as in station observations. Collectively, these amplified correlations indicate that excessive snow accumulation in ERA5L and CLM produces a tighter-than-observed coupling between snowpack evolution and subsurface freezing, thereby reducing the influence of other environmental controls on SFG variability. JSBACH (Fig. 10d) on the other hand, shows different behavior, with SFG onset exhibiting relatively strong correlations with autumn AT ($R = 0.71$) and the FI ($R = -0.60$). This indicates that soil freezing in JSBACH is overly responsive to seasonal temperature conditions, lacking the thermal buffering typically provided by snow insulation. The model's limited snow accumulation and relatively shallow snowpack fail to adequately decouple the soil from atmospheric temperature variations, resulting in a more direct soil response to AT. The SFG end date is primarily controlled by the spring AT ($R = -0.78$) and TI ($R = -0.74$), with a moderate correlation to snow cover end ($R = 0.54$). This indicates that soil thaw process is primarily driven by thermal energy input rather than by snow disappearance, in contrast to other models, likely due to the stronger influence of atmospheric conditions.

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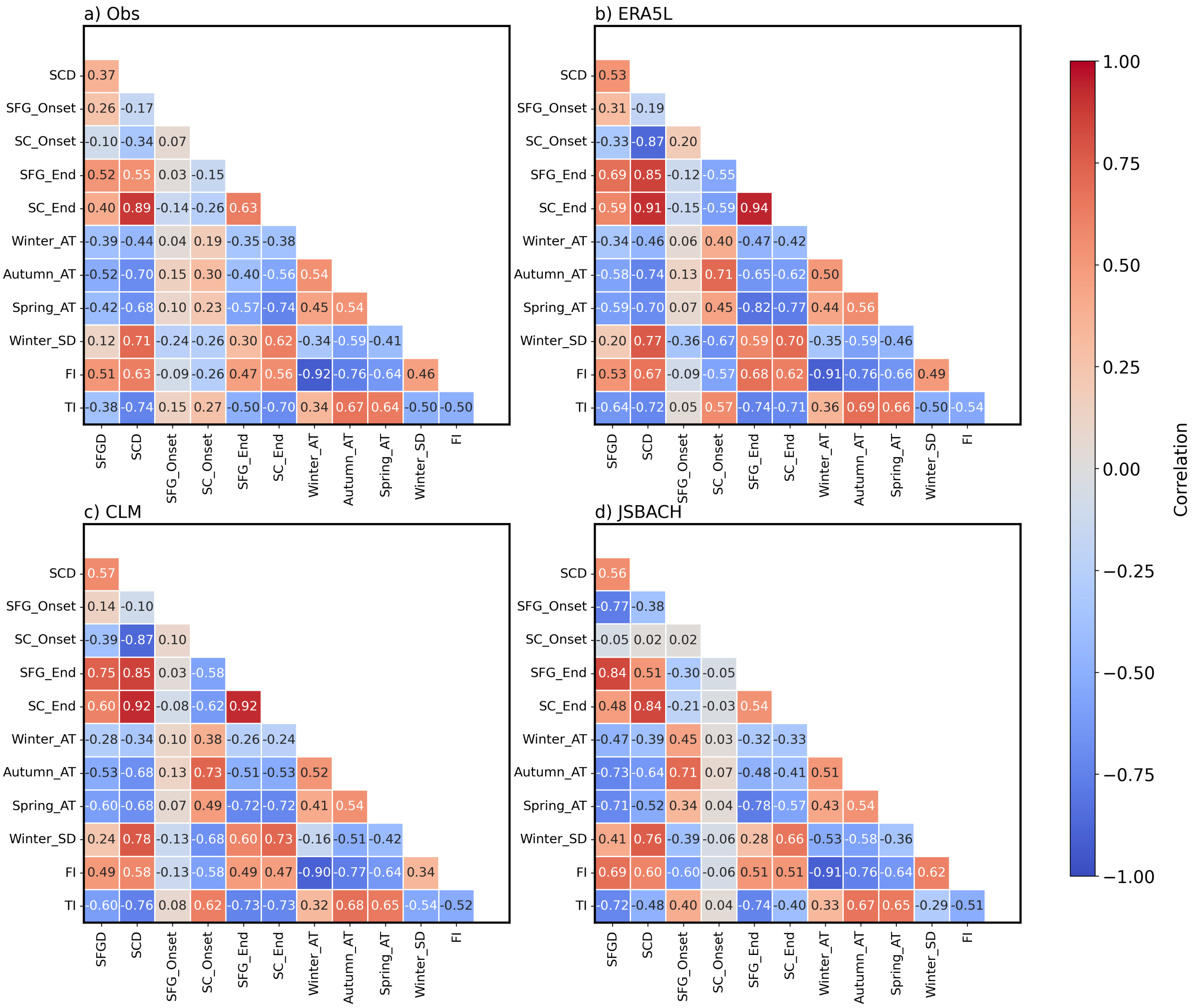

Figure 10. Correlation heat maps showing relationships between SFG characteristics (SFG_Onset, SFG_End, SFGD), snow cover characteristics (SC_Onset, SC_End, SCD), and key environmental variables, including seasonal ATs (Winter_AT, Spring_AT, Autumn_AT), winter SD (Winter_SD), annual FI, and TI. Panels show observations (upper left), ERA5L (upper right), CLM (lower left), and JSBACH (lower right). Positive and negative correlations are shown in red and blue, respectively, with the color intensity indicating the correlation strength.

Together, these results highlight that observed SFG characteristics arise from a balanced interaction between snow dynamics and seasonal thermal/freeze energy, whereas models diverge in how they distribute control between AT and snow insulation. ERA5L and CLM exaggerate the thermal decoupling effect of snow, while JSBACH underrepresents it, leading to contrasting biases in the simulated freeze-thaw behavior.

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4 Conclusions

This study investigated frozen soil characteristics in JSBACH and CLM, two land surface models over mid-to high-latitude region, in standalone configurations driven by 3-hourly ERA5 data (1986-2022) with ERA5L included as an additional dataset for the comparison. JSBACH's baseline cold bias was notably reduced by using a temperature-dependent snow densification scheme, highlighting the importance of snow density parameterization in accurately modeling the soil thermal regime, and all results discussed here use the improved JSBACH version.

The spatial analysis showed that, among the models considered in this study, JSBACH best captured the frozen ground extent, closely matching reference estimates for both SFG and PEFT, although its lower snow accumulation suggested weaker insulation and enhanced soil cooling. CLM showed intermediate skill between JSBACH and ERA5L, with a lack of PEFT over the Tibetan Plateau and excess snow and soil ice, while ERA5L indicated less cold high-latitude soils with excessive snow. These spatial patterns and model behaviors were further evaluated through site-level analysis using observations from 26 Russian SFG stations. At most stations, the models exhibited premature soil freezing and delayed spring thaw, reflecting persistent challenges in capturing autumn land-atmosphere interactions and snow evolution processes. JSBACH's shorter, less effective snow cover and underestimation of SD led to weaker insulation and a stronger direct response of STs to AT. In contrast, ERA5L and CLM produced an extended snow cover period due to earlier snow onset and delayed melt, where excessive snow accumulation suppressed soil-atmosphere heat exchange and prolonged the frozen ground period. Errors in simulating snow timing, depth, and duration propagate into freeze-thaw biases by altering soil insulation and subsurface energy exchange, resulting in contrasting model biases. The results demonstrate that observed SFG characteristics over Siberia are governed by the interplay between snow dynamics and seasonal thermal energy, while the models diverge in how control is partitioned between AT and snow processes. Our findings indicate that both JSBACH and CLM would benefit from improved representation of snow properties, with CLM additionally requiring a better representation of the partitioning between liquid and ice fraction of soil water. Future studies should examine the combined effects of snow processes, vegetation, soil properties, and soil moisture on soil freeze-thaw dynamics to enhance land surface model performance.

Code availability. The scripts used for data processing and analysis are available on Zenodo: https://doi.org/10.5281/zenodo.18110146. Also, a detailed report describing the improvements implemented in JSBACH is available at the same link.

CLM model code: https://github.com/CMCC-Foundation/CTSM

JSBACH model code: https://gitlab.dkrz.de/jsbach/jsbach/-/tree/bcaeb223249befb8c251efff29096fbe848b1b03

Data availability. ERA5-Land daily averaged data (1950–2022) were obtained from the Copernicus Climate Change Service (C3S) Climate Data Store (CDS; Muñoz-Sabater et al. (2021); https://doi.org/10.24381/cds.e9c9c792).

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470 Daily observational data from RIHMI-WDC were retrieved from http://meteo.ru/data/.

Appendix A: Station details

Table A1. Details of the selected meteorological stations with their geographic coordinates, elevation, and station codes, listed in order of increasing latitude.

Station	Latitude (^{o} N)	Longitude (^{o} E)	Elevation (m)	Code
Irkutsk	52.27	104.35	467	irk
Slavgorod	52.97	78.65	125	sla
Bajandaj	53.10	105.53	757	baj
KAMEN'-NA-OBI	53.80	81.28	127	kam
Tatarsk	55.20	75.97	110	tat
KRASNOJARSK, opytnoe pole	56.03	92.75	277	krs
Sadrinsk	56.08	63.63	88	sad
Isim	56.10	69.43	82	ism
Bratsk	56.28	101.75	411	bra
SEVERNOE	56.35	78.35	124	sev
BAKCHAR	57.00	82.07	109	bak
PERVOMAJSKOE	57.07	86.22	114	per
Tjumen'	57.12	65.43	101	tju
PUDINO	57.57	79.43	96	pud
Tobol'sk	58.15	68.25	45	tob
Bogucany	58.38	97.45	131	bog
Enisejsk	58.45	92.10	77	eni
Vitim	59.45	112.58	186	vit
Vanavara	60.33	102.27	259	van
Ivdel'	60.68	60.45	93	ivd
LENSK	60.72	114.88	241	len
Berezovo	63.93	65.05	27	ber
TOL'KA	63.98	82.08	31	tol
Ust'-Usa	65.97	56.92	77	usu
Petrun'	66.43	60.77	61	pet
NAR'JAN-MAR	67.63	53.03	10	nar

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Appendix B: Additional Figures

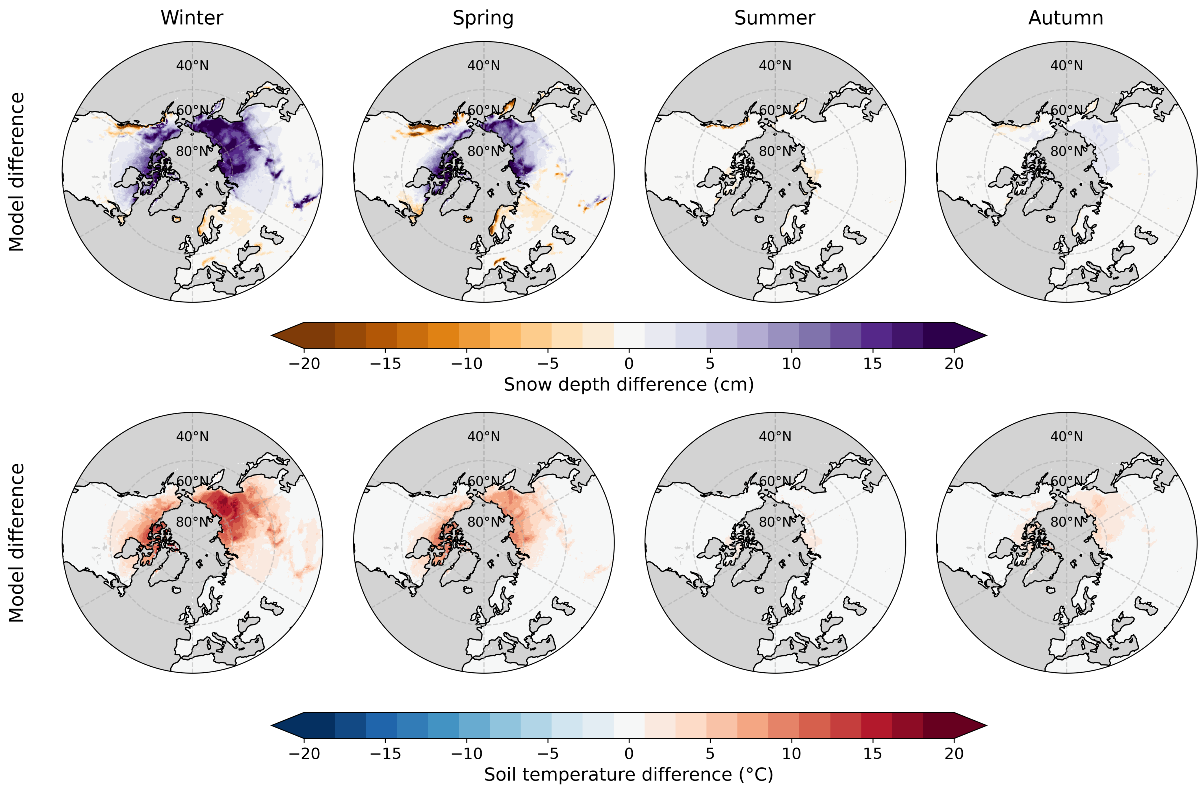

Figure B1. Seasonal climatological mean differences (1986-2022) between the new and old JSBACH setups (model difference) for snow depth (upper panel) and soil temperature at 20 cm depth (lower panel).

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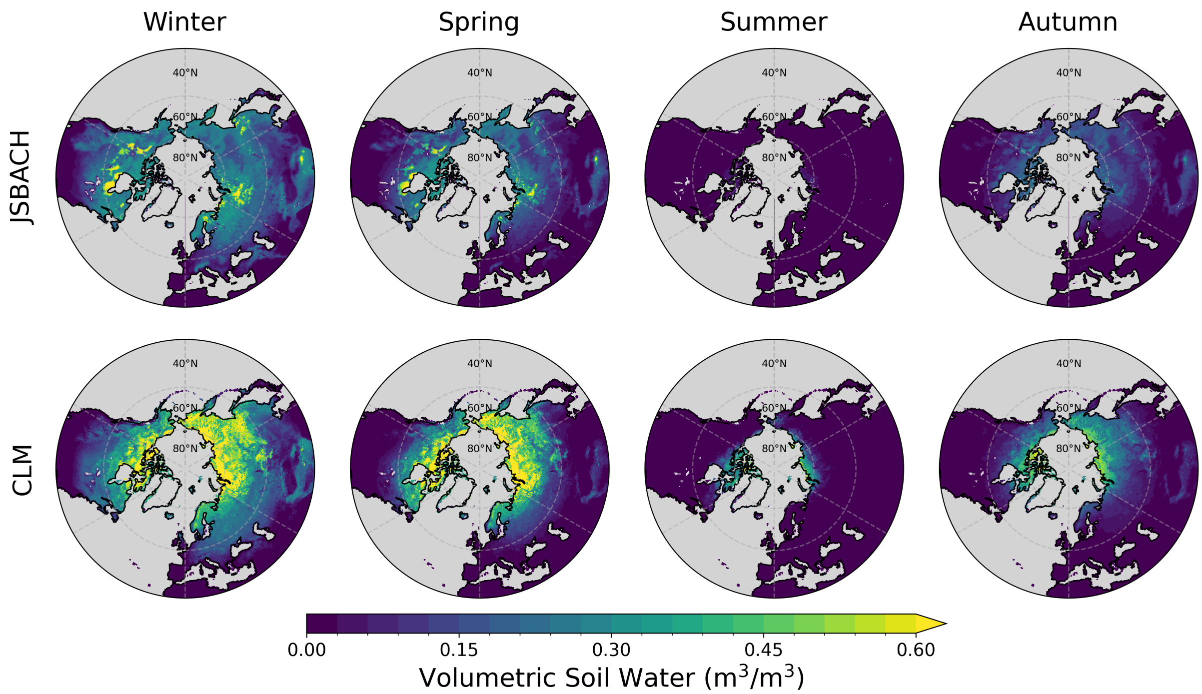

Figure B2. Comparison of the seasonal climatology (1986-2022) of soil ice content (0-20 cm) over mid- to high-latitude regions (30^o-90^o N) of the Northern Hemisphere between JSBACH (top row) and CLM (bottom row).

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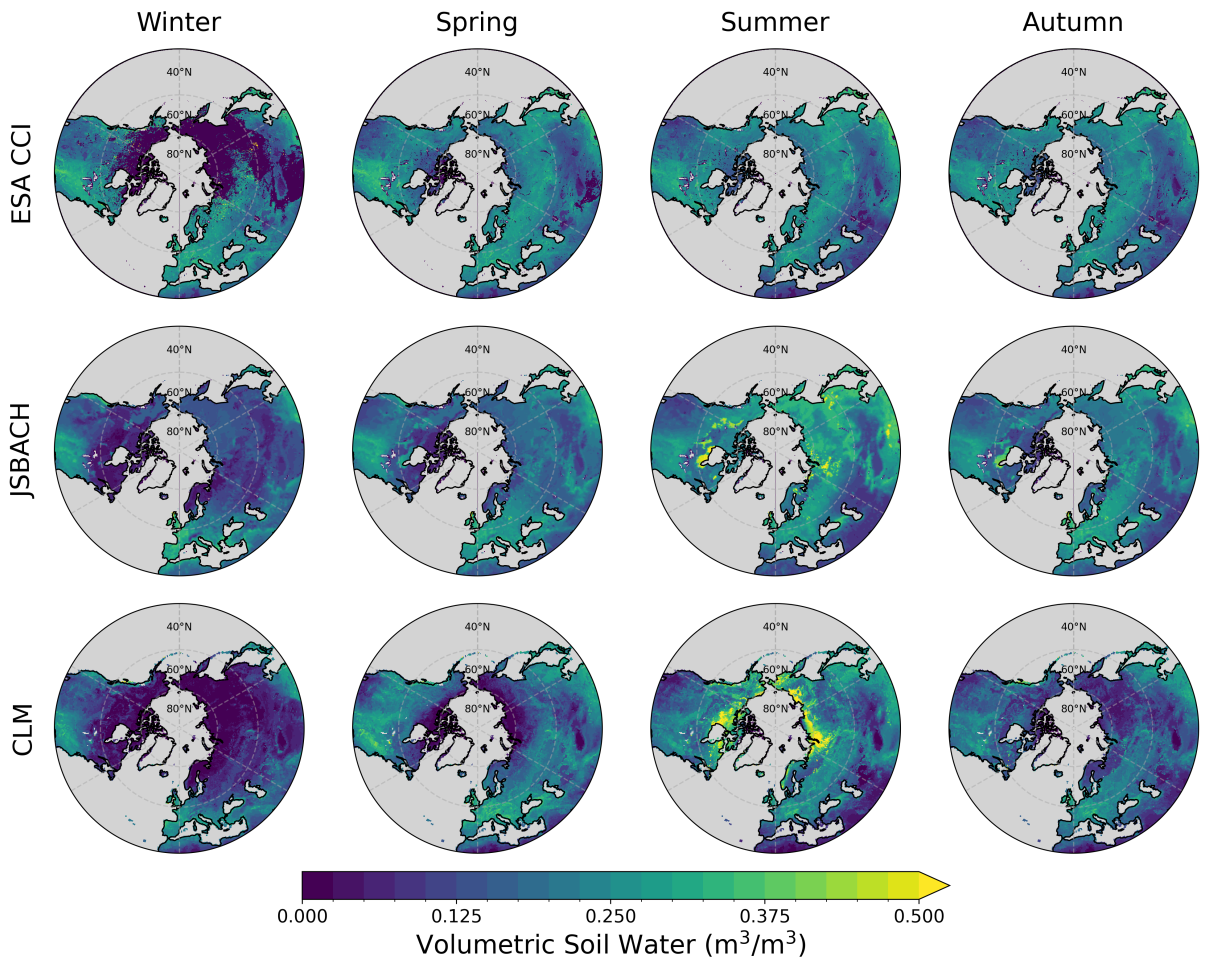

Figure B3. Comparison of the seasonal climatology (1986-2022) of top-layer soil liquid water content over mid- to high-latitude regions (30^o-90^o N) of the Northern Hemisphere, as derived from satellite data (ESA-CCI; top row) and simulated by models: JSBACH (middle row) and CLM (bottom row).

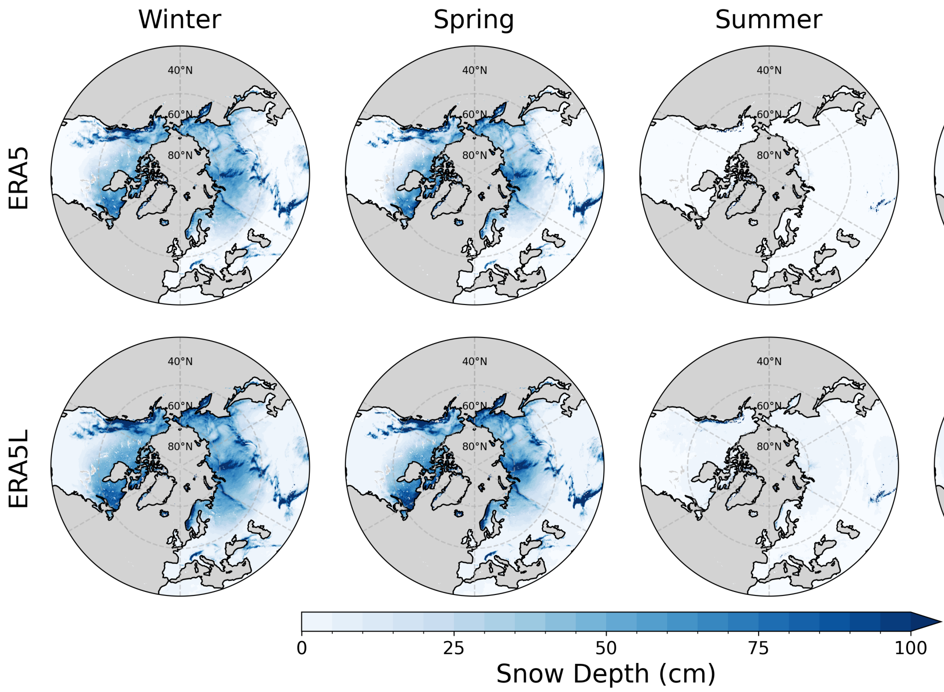

Figure B4. Comparison of the seasonal climatology (1986-2022) of SD over mid- to high-latitude regions (30^o-90^o N) of the Northern Hemisphere between ERA5 (top row) and ERA5L (bottom row).

Author contributions. MP and BA designed the research framework and defined the methodology. MP conducted the JSBACH simulation with the help of TS and KF and CLM simulation with the help of AS, DP, and MB. MP collected the data, carried out the code development and calculations, and generated all figures; ZL helped in verifying the code implementation. AS, KF, and ZL provided methodological and statistical guidance for the data analysis. MP led the interpretation of the results with contributions from all co-authors. The results were discussed collectively by all authors. MP drafted the original manuscript, and all authors contributed to reviewing, editing, and approving the final version.

Competing interests. The authors declare that there are no conflicts of interest, financial or otherwise, related to this study.

Acknowledgements. Mittal Parmar gratefully acknowledges funding from the DWD IDEA S4S project FS-SF (4823IDEAP2). This work made use of computational resources provided by the Deutsches Klimarechenzentrum (DKRZ), granted by its Scientific Steering Committee (WLA) under project ID bb1064, as well as resources from the Center for Scientific Computing (CSC) at Goethe University Frankfurt and the JUNO computing facilities at CMCC, Bologna. The authors thank Danny Risto (Goethe University Frankfurt), Silvio Gualdi and

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Enrico Scoccimarro from CMCC, Bologna, for their valuable insights and support throughout this research. The authors acknowledge the use of AI-assisted tools (e.g., ChatGPT and DeepSeek) for support with code scripting and rephrasing parts of the manuscript. We also acknowledge funding from the European Research Council under the European Union's Horizon 2020 research and innovation program as part of the Q-Arctic project (grant agreement No 951288) to T.S. and from the China Scholarship Council (CSC) sponsorship to Z.L. (No.202006040064).

Financial support. This research has been funded by the IDEA S4S project FS-SF (4823IDEAP2) and Open-access publication of this article was supported by Goethe University Frankfurt.

https://doi.org/10.5194/egusphere-2026-381
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