



1 **Differentiable Hybrid Hydrological Model for Short-Term**
2 **Flood Forecasting with Future Meteorological Information**

3 Leijing Li^{1,2,3}, Liutianjiao Hong¹, Jianzhu Li¹, Peng Shi³, Shuaihang Wang³, Jiyang Tian⁴

4 ¹State Key Laboratory of Hydraulic Engineering Simulation and Safety, Tianjin University, Tianjin

5 300072, China

6 ²Guizhou Water Conservation Science and Research Institute, Guiyang 550002, Guizhou, China

7 ³Guizhou Provincial Hydrology and Water Resources Bureau, Guiyang 550002, Guizhou, China

8 ⁴ China Institute of Water Resources and Hydropower Research, Beijing 100038, China

9 *Corresponding author: Liutianjiao Hong, PhD

10 E-mail: hongliutianjiao@tju.edu.cn

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12 **Abstract:** Short-term flood forecasting is essential for flood control operation and risk warning in
13 small- and medium-sized basins, yet its accuracy is constrained by rainfall uncertainty, hydrological
14 model structural limitations, and watershed physical response characteristics. To address these
15 challenges, this study developed a differentiable hybrid hydrological forecasting framework that
16 coupled future meteorological information with physically representations of runoff generation,
17 baseflow recession, flow concentration, and channel routing. A physics-informed neural architecture
18 search (PINAS) routing scheme was further proposed and systematically compared with
19 differentiable Muskingum routing and convolutional unit-hydrograph routing, while a purely data-
20 driven LSTM model was used as the benchmark. The framework was evaluated in three
21 representative watersheds in Hebei Province, China, namely the Shahe basin (2210 km²), Jumahe
22 basin (1760 km²), and Liulin basin (57.4 km²), using rainfall–runoff observations from the flood



23 seasons of 2000–2023 and radar-echo data from 2018–2023. Multi-lead flood forecasting
24 experiments were conducted for lead times of 0.5–2.0 h. The results showed that the LSTM
25 benchmark achieves high numerical accuracy in some high-peak flood events, but lacks explicit
26 physical constraints. In contrast, the model with PINAS routing exhibited the most stable overall
27 performance among the hybrid hydrological models, achieving a more balanced representation of
28 flood-peak propagation, hydrograph smoothing, and recession preservation. The data-dependency
29 analysis indicated that increasing the amount of training data improved forecasting stability, but
30 model skill does not increase strictly monotonically with sample size, sample representativeness
31 and flood-type coverage were also critical. SHAP interpretability analysis further revealed that
32 forecasts in the Shahe and Jumahe basins were mainly controlled by radar echoes and their
33 extrapolated features, whereas the small Liulin Basin was more strongly influenced by measured
34 areal rainfall and gauge rainfall. These findings demonstrated that integrating physical constraints,
35 differentiable runoff–routing structures, and future meteorological information can improve the
36 stability, physical consistency, and hydrological interpretability of short-term flood forecasting in
37 small- and medium-sized basins.

38 **Keywords:** Flood forecasting; Hybrid hydrological model; Neural architecture search; Radar echo
39 extrapolation; SHAP interpretability

40 **1 Introduction**

41 Floods are one of the most destructive natural hazards worldwide, and the impacts are further
42 amplified by rapid hydrometeorological variability and the pronounced spatiotemporal nonlinearity
43 of rainfall–runoff transformation (Matanó et al., 2024). Short-term flood forecasting is useful for



44 emergency response, reservoir operation, flood management and decision making. However, errors
45 in calculation of runoff generation and channel routing may lead to large inaccuracies in forecasted
46 peak discharge and arrival time, making accurate and reliable flood forecasting highly challenging.
47 (Xiang et al., 2025; Tanhapour et al., 2025).

48 Two main types of flood forecasting techniques are data-driven model (DDM) and process-
49 driven model (PDM). Although PDM can characterize runoff generation and routing with strong
50 physical interpretability, the forecasting capability is limited by structural deficiencies and
51 parameter uncertainty, particularly when calibration primarily relies on outlet measured discharge
52 in basins with limited data (Gan et al., 2018; Cibirin et al., 2014). With strong non-linear
53 representational capabilities, DDM, especially deep learning techniques, show significant potential
54 in rainfall-runoff forecasting. However, purely data driven models may show reduced generalization
55 and lack clear physical consistency and interpretability with extreme data (Amini et al., 2024). These
56 shortcomings have motivated rapid development of hybrid hydrological modeling, which aims to
57 combine the explanatory structure of physical mechanism with the learning-based components (Xu
58 et al., 2024). Differentiable modeling explicitly constructs hydrological process equations and
59 embeds them into an automatic differentiation framework, enabling end-to-end training of the
60 models (Gao et al., 2024; Huynh et al., 2025; Jamaat et al., 2025; Feng et al., 2022). In contrast to
61 Physical Informed Neural Networks (PINNs), which impose physical constraints through loss
62 functions, differentiable models retain a clear mechanistic framework and introduce learning
63 components that are difficult to parameterize (Herat et al., 2025). Recently, several studies have
64 developed distributed hybrid hydrological models for flood simulation. Zhong et al. (2024)
65 developed a distributed physics-informed deep learning model under a differentiable parameter



66 learning framework, highlighting how regionalized deep learning can map watershed attributes to
67 parameters and improve forecasting in data-scarce contexts. Wang et al. (2024) advanced physics-
68 encoded deep learning for distributed modeling and proposed that routing remains difficult to embed
69 in neural architectures in a physically consistent manner. Song et al. (2025) further discussed
70 national-scale differentiable modeling and showed that routing choices can substantially affect
71 computational cost and scalability in high-resolution settings. These studies showed that, building
72 reliable distributed hybrid hydrologic models requires not only a physically meaningful runoff-
73 generation backbone, but also a routing design that is computationally feasible and physically
74 defensible.

75 For distributed hybrid hydrologic models, channel routing is becoming more widely
76 acknowledged as a bottleneck (Mendoza et al., 2023; Collins et al., 2025). Although Muskingum
77 routing is still the default option, it necessitates recursive calculations from upstream to downstream,
78 which significantly raise computational costs during iterative training (Ayvaz et al., 2017). When
79 the routing is trained in a differentiable manner, the data-driven routing operator performs
80 exceptionally well and show promise in learning complex spatiotemporal relationships (He et al.,
81 2025a). However, the learned routing can turn into a black-box operation that may violate
82 hydrologic constraints, fail to maintain physically meaningful wave dynamics, strictly conserve
83 water mass, and degrade robustness and out-of-sample generalization, which are particularly
84 important for flood forecasting and model transfer across basins (He et al., 2025b). Such routing
85 deficiencies can most obviously show up as distorted rising limbs and biased peak timing in event-
86 based flood nowcasting for small-to-medium size watersheds with short lead times, decreasing the
87 operational effectiveness of hybrid models. As a result, creating a routing approach that is both



88 physically consistent and learnable continues to be a major research issue.

89 The selection of a routing architecture that achieves the best balance between expressiveness,
90 stability, and physical plausibility remains a fundamental methodological challenge even after the
91 idea of physics informed learnable routing has been adopted (Feng et al., 2022). The components of
92 differentiable models are usually determined by manual design, which can be challenging to transfer
93 across basins, temporal resolutions, and event regimes (Sawadekar et al., 2025). Manually selected
94 designs may quickly become either under-parameterized or over-parameterized in short-term lead
95 time (Astagneau et al., 2024). Neural architecture search (NAS) has recently been used in the field
96 of hydrological forecasting to automate model structure design, such as water level prediction (Ng
97 et al., 2025) and flood forecasting (Yao et al., 2023; Li et al., 2024), thereby reducing reliance on
98 manual expertise and iterative parameter tuning. Similar studies have also been carried out in related
99 hydro-meteorological forecasting, such as NAS-assisted rainfall forecast post-processing (Wen et
100 al., 2024). In the field of hydrological machine learning, architecture search has been recognized as
101 a more reproducible and comparable technique for model construction compared to empirically
102 selecting network architectures (Chadalawada et al., 2020). Furthermore, the routing design must
103 satisfy conservation and non-negativity constraints, merely using general network architecture
104 search techniques may produce correct but physically incomprehensible solutions. Therefore,
105 incorporating NAS into routing design while limiting searches within the physics-informed routing
106 operator space to yield convergence structures with better forecasting capabilities and greater
107 conformity to hydrological physical plausibility.

108 Apart from the structure of the hydrological model, the main limiting factor in short-term event
109 forecasting is typically the uncertainty in meteorological parameters. (Cafiero et al., 2025). Effective



110 forecasting requires some depiction of future precipitation, as measured rainfall and radar data can
111 only provide information about the past and present (Li et al., 2024). Because meteorological radar
112 can provide high-resolution and rapidly updated observations of convective systems, it plays an
113 important role in precipitation nowcasting. To generate future reflectivity fields and accompany
114 precipitation nowcasting, radar echo extrapolation techniques have been developed (Ravuri et al.,
115 2021). These techniques range from conventional motion-field advection to deep-learning sequence
116 models (Zheng et al., 2024; Song and Han, 2025). However, the practical application value of radar
117 in hydrological forecasting is still unclear. This is because nowcasting accuracy usually decreases
118 quickly with lead time, and hydrological reaction filters meteorological uncertainty in a scale-
119 dependent way. Therefore, current research focuses on determining the extent to which various data
120 derived from radar echo analysis can improve the accuracy of flood forecasts.

121 Motivated by the challenges in hydrological forecasting, A hybrid multiscale flood forecasting
122 framework is proposed, which focuses on the runoff generation and concentration, employing a
123 modular routing representation for forecasting of multi-period hydrograph, and attempt to
124 incorporate future meteorological information based on radar echo extrapolation to improve forecast
125 accuracy. The Physics-Informed Neural Architecture Search (PINAS) routing has been proposed,
126 which is guided by physics-consistency constraints and learned in a physically motivated
127 architecture space, and compared it with learnable convolution unit hydrograph routing and
128 differentiable Muskingum routing. In summary, the objective of this study is to investigate the
129 following scientific questions:

130 (1) How does routing design affect hydrological processes throughout entire watershed areas
131 during short lead times?



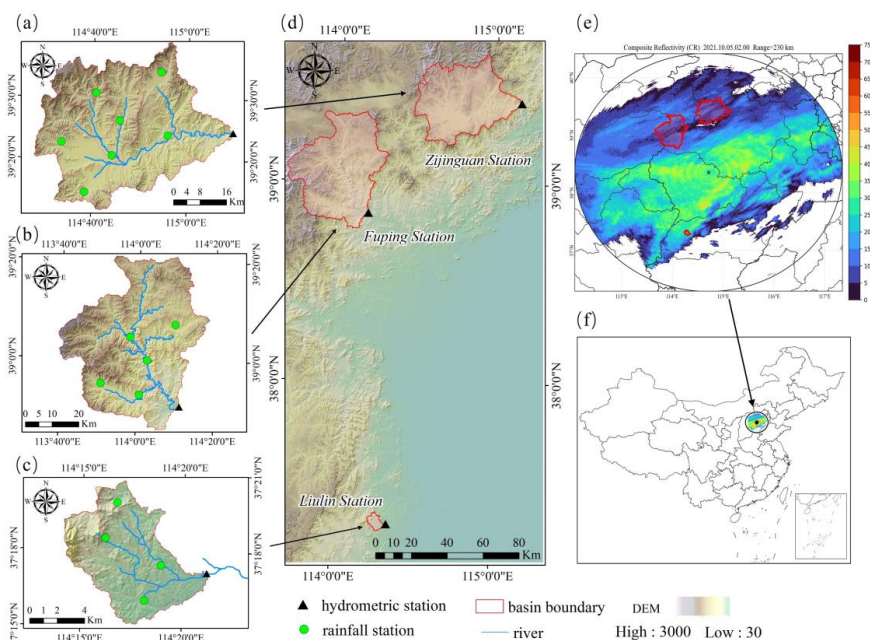
132 (2) Is it possible for PINAS to improve hydrographs simulation while maintaining physical
133 consistency and transferability across events and basins?

134 (3) In the short-term forecasting of flood events, which characteristics will have a decisive
135 impact? Will the radar echo data bring about positive contributions?

136 **2 Materials and Methods**

137 **2.1 Study area and data**

138 Three representative basins in Hebei Province, China, namely the Jumahe, Shahe, and Liulin
139 basins, have been selected as case study areas. The Shahe basin covers an area of 2210 km², a main-
140 channel length of 124 km, and an average basin width of 34.5 km. Six meteorological stations are
141 located within the basin, and the Fuping hydrological station is situated at the basin outlet. The
142 Jumahe basin covers an area of 1760 km², with a main channel length of 81.5 km and an average
143 basin width of 25.4 km. It contains eight meteorological stations, and discharge is measured at the
144 Zijinguan hydrological station at the outlet. The Liulin basin with a drainage area of 57.4 km². Five
145 meteorological stations are distributed within the basin, and the Liulin hydrological station serves
146 as the outlet control station.



147

148 Figure 1. Overview of the study area. (a) Jumahe basin, (b) Shahe basin, (c) Liulin basin, (d)

149 Location of the research area, (e) Radar scanning range, (f) The position of the radar.

150 The data were obtained from the Hebei Provincial Hydrology Bureau and the Hebei Provincial

151 Meteorological Bureau. The rainfall-runoff records during the flood seasons from 2000 to 2023 and

152 radar echo data during the flood seasons from 2018 to 2023 for the three study basins have been

153 collected. The rainfall-runoff records were temporally interpolated to a 30 min resolution. Typical

154 flood events were identified based on the criterion that each event should exhibit a distinct and

155 complete rising and recession process. The radar-echo data were temporally matched with the flood

156 events and used as optional meteorological inputs to the forecasting models. In total, 38 typical

157 flood events were identified in the Shahe Basin, with 4250 h of available radar-echo maps, 25 typical

158 flood events were identified in the Jumahe Basin, with 1512 h of radar-echo maps, and 45 typical

159 flood events were identified in the Liulin Basin, with 498 h of radar-echo maps.



160 **2.2 Flood forecasting model**

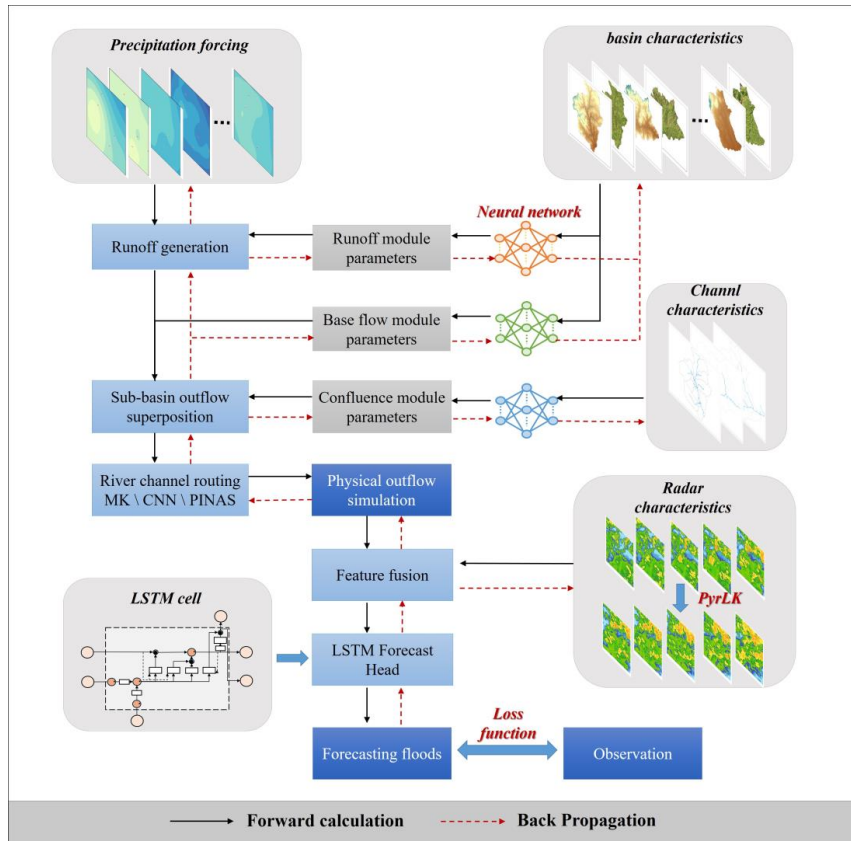
161 **2.2.1 Model framework**

162 The differentiable hybrid hydrological model suitable for short-term flood forecasting in small
163 and medium-sized basins were developed. The model was constructed based on key hydrological
164 processes and integrates runoff generation, baseflow recession, lateral inflow aggregation, channel
165 routing, and multi-lead state correction into an end-to-end trainable framework. Different to purely
166 data-driven models, the proposed framework first generated initial flow with physical meaning
167 through a differentiable hydrological core, and then nonlinearly corrected the physical prior by
168 combining historical state memory, future meteorological characteristics, and lead-time-specific
169 information.

170 The model consisted of two main components. The first part is the differentiable hydrological
171 core, which produced the physical prior discharge Q_{phys} . This core was guided by hydrological
172 mechanisms, in which watershed topographic characteristics were represented in a differentiable
173 form, runoff generation was described using physical formulations, and baseflow dynamics were
174 represented through a differentiable recession process. Within this framework, channel routing was
175 treated as a key learnable component. Three routing schemes, namely Muskingum routing,
176 convolutional unit-hydrograph (CNN Unit) routing, and Physics-Informed Neural Architecture
177 Search (PINAS) routing, were structurally substituted and compared to evaluate the influence of
178 different routing formulations on flood hydrograph representation. The second part was the multi-
179 lead forecasting head, which integrated the physical-prior discharge, historical hydrological memory,
180 and future meteorological characteristics to generate the final forecast discharge. This design
181 enables the model to retain the hydrological process constraints while adapting to the nonlinear



182 response differences between different basins and flood events through data-driven corrections.



184 Figure 2. The diagram of the hybrid hydrological model

185 2.2.2 Differentiable hybrid hydrological core

186 The differentiable hybrid hydrological core was designed to simulate the main process of
187 rainfall converting into downstream river flow. This module consisted of four components, runoff
188 generation, baseflow recession, lateral inflow aggregation, and channel routing. All components
189 were formulated in a differentiable manner, allowing the model to be trained end-to-end through
190 backpropagation while maintaining hydrologically meaningful process constraints.

191 2.2.2.1 Runoff generation

192 The runoff generation adopted differentiable SCS-CN expression, which linked rainfall loss



193 with effective runoff. The curve number parameters were used to describe the combined influence
194 of land surface conditions, previous wetness status, and rainfall-runoff threshold behavior. To make
195 the model better adapted to different flood generation conditions, the runoff generation parameters
196 were treated as learnable variables and restricted within physically reasonable ranges through
197 bounded differentiable mappings. This prevents nonphysical parameter values and numerical
198 instability during training process. In this module, rainfall inputs were first transformed into
199 effective rainfall after accounting for hydrological losses, and the resulting effective rainfall was
200 then converted into fast lateral inflow. The differentiable SCS-CN expression retained the threshold
201 and nonlinear characteristics of the rainfall-runoff conversion, thereby improving the hydrological
202 rationality of the model under limited sample conditions.

203 2.2.2.2 Baseflow calculation

204 The baseflow component was represented by a differentiable recession formulation to describe
205 slow-response runoff. The baseflow state was updated recursively over time, and its recession rate
206 was controlled by a recession coefficient. To ensure the stability of the numerical and hydrological
207 values, the recession coefficient was limited within a physically reasonable range, so that the
208 baseflow dynamics followed an exponential decay behavior. The baseflow component was then
209 combined with the routed fast-flow component at the basin outlet, forming part of the physical-prior
210 discharge. This design preserved the influence of antecedent hydrological conditions on current
211 discharge in short-term flood forecasting, and was particularly important for maintaining recession
212 flows and representing post-peak hydrograph behavior.

213 2.2.2.3 Lateral inflow aggregation and river-network topology

214 For spatial representation, the basins were divided into multiple subbasin units, and the



215 generated rapid runoff was allocated as lateral inflow according to the area fraction of each subbasin.
216 The lateral inflow volume of each subbasin was further mapped to the corresponding river section
217 as the local input for the calculation of river network paths. The river network was represented as a
218 directed topological structure composed of reach nodes. The inflow volume of each river segment
219 includes the local lateral inflow volume and the outflow volume of its upstream river segment. Flow
220 propagation was then performed reach by reach from upstream to downstream according to a
221 predefined topological order. This structure ensured physically consistent flow accumulation and
222 propagation along the river network.

223 **2.2.3 Channel routing schemes**

224 Channel routing was a key component linking subbasin runoff generation to outlet discharge
225 forecasting. To assess the impact of different routing structures on the performance of flood
226 hydrographs, three differentiable channel routing schemes were embedded into a unified
227 hydrological core: Muskingum routing, convolutional unit-hydrograph routing, and Physics-
228 Informed Neural Architecture Search routing. The formulas for the channel routing were shown in
229 Table 1.

230 **2.2.3.1 Muskingum routing**

231 The Muskingum method represented flood-wave propagation through a storage–discharge
232 relationship. For each river reach r , two parameters K_r and X_r were respectively defined, where K_r
233 denoted the characteristic travel-time scale and X_r controlled the relative weighting of inflow and
234 outflow in channel storage. After standard finite-difference discretization, the routed outflow was
235 expressed as a recursive update, in which the current outflow was represented as a linear
236 combination of the current inflow, previous inflow, and previous outflow. To enable end-to-end



237 differentiable training while avoiding numerical oscillations, K_r and X_r were treated as learnable
238 parameters but constrained within physically plausible ranges through differentiable mappings. In
239 addition, the routing stability condition $K_r(1-X_r) \geq \Delta t/2$ was imposed.

240 2.2.3.2 Convolutional unit-hydrograph routing

241 In the convolutional unit-hydrograph routing scheme, channel routing was formulated as a
242 linear time-invariant impulse-response process. Each reach r was assigned a discrete causal unit-
243 hydrograph kernel $w_{r,j}$. Non-negativity and volume-conservation constraints are imposed on the
244 kernel to ensure physical interpretability, preventing negative discharge generation and preserving
245 runoff volume in the convolutional sense while allowing flood peak smoothing and propagation
246 delay. The routed outflow $Q_r(t)$ was obtained by the discrete convolution between the reach inflow
247 $I_r(t)$ and the unit-hydrograph kernel. Because both the convolution operation and the normalized
248 kernel parameterization were differentiable, this routing operator was jointly trained with the
249 upstream rainfall–runoff module through end-to-end backpropagation.

250 2.2.3.3 Physics-Informed Neural Architecture Search routing

251 To enhance the adaptability of the routing structure to different basin scales and flood wave
252 propagation characteristics, a routing scheme based on physical information called Physical
253 Informed Neural Architecture Search (PINAS) was proposed. This method combined neural
254 architecture search (NAS) with physics-informed constrained to automatically identify appropriate
255 combinations and scales of routing operators under hydrological constraints. The channel routing
256 was represented as a routing stack composed of multiple differentiable operators, and the
257 appropriate operator combinations and scale parameters were selected from the predefined search
258 space to form a routing structure that was adaptable to different river network configurations.



259 The search space included causal convolution, dilated convolution, multi-scale convolution,
 260 kernel mixing, gated connections, pure-delay mapping, and identity mapping. During structural
 261 computation, the input sequence was transformed layer by layer, while differentiable mixing
 262 between kernel output and residual gated connections was allowed. To maintain physical
 263 consistency, smoothness regularization was introduced. The regularization terms from different
 264 layers were aggregated into a routing regularization term, and a volume-consistency residual was
 265 calculated. During the training process, physical constraints were incorporated into the loss function
 266 in a weighted manner. Candidate architectures were first screened using limited training and
 267 validation batches, and the selected architectures were then evaluated more thoroughly. A
 268 complexity penalty was also included in the scoring function to avoid unnecessarily complicated
 269 routing structures. Compared with Muskingum and CNN-Unit, PINAS learned routing operator
 270 combinations that were adaptive to different reaches and flood events under physical constraints,
 271 thereby enhancing the representation of complex flood-wave propagation processes.

272 Table 1. The formula for the channel routing

Type	Formula	Description
Unit-hydrograph kernel	$w_{r,j} = \frac{\text{softplus}(w_{r,j})}{\sum_{i=0}^{k-1} \text{softplus}(w_{r,i})}$	Non-negative and mass-conserving parameterization of the convolutional unit kernel.
Unit-hydrograph routing	$Q_r(t) = \sum_{j=0}^{k-1} w_{r,j} I_r(t-j)$	Channel routing based on the discrete convolution.
Muskingum denominator	$D_r = K_r(1 - X_r) + \frac{\Delta t}{2}$	Denominator used to calculate the Muskingum routing coefficients.
Muskingum routing coefficients	$C_{0,r} = \frac{-K_r X_r + \Delta t / 2}{D_r}$ $C_{1,r} = \frac{K_r X_r + \Delta t / 2}{D_r}$ $C_{2,r} = \frac{K_r(1 - X_r) - \Delta t / 2}{D_r}$	Stable recurrence coefficients of the differentiable Muskingum routing scheme.



Muskingum routing recurrence	$Q_r(t + \Delta t) = C_{0,r}I_r(t + \Delta t) + C_{1,r}I_r(t) + C_{2,r}Q_r(t)$	Recursive update of routed outflow in the Muskingum scheme.
PINAS routing stack	$Q_r(t) = (f^{(B)} \circ f^{(B-1)} \circ \dots \circ f^{(1)})(I_r(t))$	Differentiable routing stack composed of multiple operator layers.
Operator selection	$f^{(b)}(x) = \sum_{o \in \mathcal{O}} z_{b,o} o(x), \quad z_{b,o} \in \{0,1\}, \quad \sum_{o \in \mathcal{O}} z_{b,o} = 1$	Differentiable relaxation of operator selection
Routing regularization	$R_{route} = \frac{1}{N_r} \sum_{r=1}^{N_r} \sum_{b=1}^B \text{reg}(\theta_{r,b})$	Physics-inspired regularization imposed on routing parameters.

273 Note: r indexes river reaches and s indexes sub-basins. t denotes the discrete time index, and Δt is
 274 the hydrological model time step. $I_r(t)$, $L_r(t)$, and $Q_r(t)$ denote reach inflow, lateral inflow, and reach
 275 outflow, respectively. $U(r)$ is the set of immediate upstream reaches of reach r . Sub-basin runoff
 276 $Q_s(t)$ is mapped to the channel network through the assignment function $m(s)$. In the convolutional
 277 unit-hydrograph routing scheme, $w_{(r,j)}$ denotes the normalized kernel weight. In Muskingum routing,
 278 Kr is the propagation time-scale parameter, Xr is the weighting coefficient, and $C_{0,r}$, $C_{1,r}$, and $C_{2,r}$
 279 are the recurrence coefficients. In PINAS, B is the number of differentiable routing layers, $f^{(b)}(\cdot)$
 280 denotes the operator layer at depth b , $\theta_{r,b}$ denotes the routing parameters of reach r at layer b , and
 281 R_{route} is the routing regularization term.

282 2.2.4 Multi-lead forecasting head and state-correction mechanism

283 The differentiable hydrological core generated a physically meaningful prior discharge,
 284 denoted as Q_{phys} . However, in short-term flood forecasting, uncertainties associated with future
 285 rainfall evolution, initial hydrological states, and model structural deficiencies may lead to
 286 deviations between the physical-prior discharge and the actual discharge. Therefore, a multi-lead
 287 forecasting head was constructed after the hydrological core to dynamically correct the physical-
 288 prior discharge through a state-correction mechanism, thereby producing the final multi-lead flood
 289 forecasts.



290 Let t denote the current forecasting time, and $t+k$ denote the target time corresponding to the k
291 forecast lead, where $k=1,2,\dots,K$. In this study, the rainfall data were represented at a 30 min temporal
292 resolution, and $K=4$, corresponding to lead times of 0.5, 1.0, 1.5, and 2.0 h. For each lead time, the
293 hydrological core first generated the corresponding physical-prior discharge Q_{t+k}^{phys} . To perform
294 state-dependent correction of this physical prior, the model further constructed an integrated context
295 vector C_t , which consisted of three types of information: the historical memory state h_t^{mem} , the
296 encoded future meteorological forcing f_t^{fut} , and the hydrological state features s_t^{hyd} . Specifically,
297 h_t^{mem} was encoded from historical hydrometeorological sequences and historical discharge records,
298 and was used to represent antecedent wetness conditions, rainfall–runoff response memory, and
299 hydrograph persistence. The term f_t^{fut} was derived from future meteorological forcing
300 information, mainly including extrapolated radar echoes and their basin-scale statistical descriptors,
301 and was used to describe the intensity, spatial position, and movement tendency of short-term future
302 rainfall systems. The vector s_t^{hyd} represents hydrological state information, including the current
303 discharge state, discharge variation rate, antecedent precipitation index, rainfall statistics, and
304 physical-prior discharge.

305 Due to rainfall uncertainty, watershed response lag, and routing phase errors differ among
306 forecast lead times, a lead-aware mechanism was introduced. The forecasting head formulations
307 was shown in Table 2. For the k lead time, the model integrated the context vector C_t , the lead-time
308 embedding e_k , the future meteorological forcing feature x_{t+k}^{fut} , and the physical-prior discharge
309 Q_{t+k}^{phys} to construct the lead aware feature representation $z_{t,k}$. Here, e_k was a learnable lead-time
310 embedding that represented lead-dependent error growth, information decay, and response delay.
311 The variable x_{t+k}^{fut} denoted the future meteorological input at the k lead time, mainly derived from



312 radar-echo extrapolation and its statistical features. The lead-aware feature $z_{t,k}$ was then used as the
 313 direct input for residual correction. After obtaining $z_{t,k}$, the model estimated the discharge correction
 314 term ΔQ_{t+k} through a nonlinear residual correction function $f_{\theta}(\cdot)$. The residual term was designed to
 315 compensate for structural biases in the hydrological core arising from localized intense rainfall,
 316 multi-peak flood processes, radar extrapolation errors, and complex nonlinear routing dynamics.
 317 Finally, the residual correction ΔQ_{t+k} was added to the physical-prior discharge Q_{t+k}^{phys} , and a non-
 318 negative mapping function $g^+(\cdot)$ was applied to obtain the final forecast discharge Q_{t+k} , ensuring
 319 that the output satisfies the basic physical constraint of non-negative discharge. Accordingly, the
 320 multi-lead forecasting head could be interpreted as a physical-prior-driven state-correction
 321 mechanism. The core idea is to learn a residual correction ΔQ_{t+k} conditioned on the physical prior
 322 discharge Q_{t+k}^{phys} , historical memory, future meteorological information, and lead-time-specific
 323 characteristics. This structure preserved the fundamental physical constraints associated with runoff
 324 generation, flow concentration, and channel routing, while enhancing the adaptability to hydrograph
 325 shape, peak magnitude, and recession behavior across different forecast lead times.

326 Table 2. Forecasting-head formulations

Type	Formula	Description
Integrated context vector	$C_t = [h_t^{mem}; f_t^{fut}; s_t^{hyd}]$	Integrated context vector combining historical memory, future forcing, and hydrological state information.
Lead-aware feature	$z_{t,k} = \phi(C_t, e_k, x_{t+k}^{fut}, Q_{t+k}^{phys})$	Lead-aware feature representation for the k-th forecast horizon.
Residual correction	$\Delta Q_{t+k} = f_{\theta}(z_{t,k})$	Data-driven correction term for the physical-prior discharge.
Final forecast	$Q_{t+k} = g^+(Q_{t+k}^{phys} + \Delta Q_{t+k})$	Final discharge forecast constrained by a non-negative mapping.

327 Note: t denotes the current forecasting time, and k is the forecast lead step. C_t is the integrated
 328 context vector, where $[;]$ denotes vector concatenation. h_t^{mem} is the historical memory state,



329 f_t^{fut} is the encoded future meteorological forcing, and s_t^{hyd} is the hydrological state feature
 330 vector. e_k is the learnable lead-time embedding, $\varphi(\cdot)$ is the nonlinear feature-fusion function, $f_{\theta}(\cdot)$ is
 331 the residual correction function with parameters θ , and $g^+(\cdot)$ is a non-negative mapping function
 332 used to ensure physically feasible discharge predictions.

333 2.3 Radar echo extrapolation

334 The future meteorological information was obtained through short-term extrapolation of the
 335 radar echo sequence. The Pyramidal Lucas–Kanade (PyrLK) optical-flow method was used to
 336 estimate a pixel-wise two-dimensional motion field $v(x)=[u(x),v(x)]$. The formulations were shown
 337 in Table 3. Under a Lagrangian persistence assumption, the most recent radar echo was advected
 338 forward to produce multi-step extrapolated fields. To accommodate large storm displacements,
 339 PyrLK adopted a coarse-to-fine image pyramid, where large-scale motion was first estimated at
 340 coarser levels and then iteratively refined at finer levels. In practice, outliers were rejected by a
 341 combination of forward–backward consistency checking and displacement-magnitude thresholding,
 342 followed by spatial interpolation and mild smoothing to obtain a continuous motion field.
 343 Extrapolation was implemented through backward warping with bilinear interpolation for non-
 344 integer source locations. The procedure was applied directly to dBZ echoes, and the extrapolated
 345 echoes were used as future forcing inputs for the multi-lead-time flood forecasting model.

346 Table 3. Lucas-Kanade optical-flow formulations

Type	Formula	Description
Brightness conservation	$R(\mathbf{x}, t) \approx R(\mathbf{x} + \mathbf{v}(\mathbf{x})\delta, t + \delta), \delta = 6\text{min}$	Approximate conservation of radar-echo intensity between two consecutive frames.
Optical-flow constraint	$R_x u + R_y v + R_t = 0$	Linearized optical-flow constraint equation.



Local least-squares estimation	$\min_{\Delta \mathbf{d}} \sum_{\mathbf{p} \in \mathcal{W}} (\nabla R(\mathbf{p}, t) \cdot \Delta \mathbf{d} + R_t(\mathbf{p}, t))^2$	Estimation of local displacement increments in the Lucas-Kanade window.
Pyramid displacement update	$\mathbf{d}^{(\ell)} = 2\mathbf{d}^{(\ell+1)} + \Delta \mathbf{d}^{(\ell)}$	Coarse-to-fine displacement refinement in the image pyramid.
Frozen-advection extrapolation	$\hat{R}(\mathbf{x}, t + k\delta) = R(\mathbf{x} - k\mathbf{v}(\mathbf{x})\delta, t), \quad k = 1, 2, \dots$	Backward-mapping extrapolation under a frozen advection field.

347 Note: $R(x, t)$ denote radar reflectivity in dBZ at location $x=(x, y)$ and time t , $v(x)=[u(x), v(x)]$ is the
 348 two dimensional motion field, R_x, R_y, R_t denote spatial and temporal gradients, W denote the local
 349 Lucas–Kanade window used for least-squares estimation, ℓ denote pyramid levels in the coarse-to-
 350 fine scheme, $\hat{R}(\mathbf{x}, t + k\delta)$ is the extrapolated reflectivity, where k is the lead step.

351 2.4 Control experiment

352 To evaluate the effectiveness of the proposed hybrid hydrological model, the Long Short-Term
 353 Memory model (LSTM) (Kratzert et al., 2018) was implemented as a control experiment. This
 354 model directly learned the nonlinear input-output relationship from historical information sequences
 355 and future meteorological forcing. Therefore, its forecasting results were used to represent the
 356 baseline performance achievable by a black-box deep-learning model under the same input
 357 conditions, providing a reference for assessing the added value of physical constraints in the hybrid
 358 hydrological framework.

359 2.5 Loss function

360 A hybrid loss function was constructed by combining data-fitting terms with physics-
 361 consistency constraints to improve multi-lead flood hydrograph forecasting, peak-flow forecasting
 362 accuracy, and physical consistency. To enhance model sensitivity to flood peaks, a weighted error
 363 term was introduced in the logarithmic discharge domain. Physical consistency was enforced
 364 through a physics fitting loss, a residual consistency loss, and a relative volume error constraint



365 during end-to-end training. In addition, the regularization residuals produced by the PINAS routing
 366 module were incorporated into the objective function with corresponding weights, thereby
 367 constraining the learned routing structure while maintaining training stability. The Loss-function
 368 formulations were shown in Table 4.

369 Table 4. Loss-function formulations

Type	Formula	Description
Total objective	$\mathcal{L} = \mathcal{L}_{\text{data}} + \lambda_{\text{vol}} \mathcal{L}_{\text{vol}} + s(e) \lambda_{\text{route}} R_{\text{route}}$	The training objective combines a data-fitting term with physically motivated constraints.
Data-fitting loss	$\mathcal{L}_{\text{data}} = \frac{1}{BK} \sum_{b=1}^B \sum_{k=1}^K \sum_{\tau \in \mathcal{Q}} W_{bk} \rho_{\tau}(y_{bk} - \hat{q}_{bk}^{(\tau)})$	Quantile-based data-fitting loss.
Volume-consistency loss	$\mathcal{L}_{\text{vol}} = \frac{1}{B} \sum_{b=1}^B \frac{\left \sum_k \hat{q}_{bk}^+ - \sum_k y_{bk}^+ \right }{\left \sum_k y_{bk}^+ \right + \varepsilon}, \quad x^+ = \max(x, 0)$	Constrains the cumulative forecast discharge volume to match the measured volume.
Ramp-up schedule	$s(e) = \min\left(1, \frac{e}{e_0}\right)$	Epoch-dependent ramp-up factor used to gradually introduce physics-based constraints.

370 Note: B is the batch size, K is the lead times, y_{bk} is the observation, W_{bk} denotes a composite
 371 weighting matrix, λ_{vol} and λ_{route} are hyperparameters controlling the strength of the volume-
 372 consistency and routing-regularization terms, $s(e)$ is a ramp-up schedule used to gradually introduce
 373 physics constraints during training.

374 2.6 Shapley Additive Explanation

375 The Shapley Additive Explanation (SHAP) was used to interpret the forecasting results. SHAP
 376 is based on the concept of Shapley values from cooperative game theory, where a single prediction
 377 is decomposed into the marginal contributions of each input feature to the model output. Therefore,
 378 it provides local explanations at the sample level and can be aggregated over multiple flood events



379 to quantify the importance of global features. Given the temporal input structure, including gauge
380 rainfall, areal rainfall, antecedent precipitation index, historical radar echoes, and future radar-echo
381 features, individual feature values of different time steps were used as basic explanatory units to
382 calculate the Shapley values. To enhance hydrological interpretability, grouped SHAP analysis was
383 further conducted by aggregating Shapley values according to variable category and temporal scale.
384 Specifically, the contributions of the same hydrological variable over the full historical window
385 were summed to obtain variable-level importance. This enabled diagnosis of the relative
386 contribution of different hydrometeorological factors to flood forecasting.

387 **3 Experiment Design**

388 **3.1 Hydrometeorological fusion for flood forecasting**

389 A land–atmosphere coupled forcing configuration was established to integrate hydrological
390 information with future meteorological inputs. Three channel-routing modules, namely CNN-Unit,
391 Muskingum, and PINAS, were comparatively evaluated. In addition, a purely deep-learning LSTM
392 model was implemented as the benchmark for flood hydrograph forecasting. The comparison
393 focused on the ability of different configurations to reproduce flood peak magnitude, runoff volume,
394 and the rising and recession limbs of the hydrograph, thereby quantifying the contribution of
395 different routing mechanisms to short-term flood forecasting.

396 **3.2 Data dependency analysis for flood forecasting**

397 To assess the dependence of model performance on data availability and its applicability under
398 limited data conditions, training sample reduction experiments were conducted. By gradually
399 reducing the proportion of training data from 25% to 100%, this study repeatedly trained and



400 evaluated the model on the same test set. This experiment was designed to characterize how
401 forecasting performance changes with training sample size, thereby providing guidance for data
402 requirements and model transfer in operational applications.

403 **3.3 Interpretability analysis of forecasting results**

404 In order to determine the factors that have the greatest impact on the accuracy of flood
405 forecasting, an interpretability analysis based on Shapley values to attribute the causes of the
406 forecast results were employed. The PINAS model was used for interpreting representative flood
407 events, and the direction and degree of each contribution to the forecasting were quantified. To
408 accommodate the temporal structure of the input data, feature contributions were further
409 aggregated along the time dimension, yielding both variable-level importance and temporal
410 contribution patterns. These results were used to diagnose the key influencing factors of flood
411 forecasts at different lead times.

412 **3.4 Model Training and Evaluation Metrics**

413 All models were implemented in PyTorch and trained on the NVIDIA GeForce RTX 4060 Ti
414 GPU. The AdamW optimizer was used with a batch size of 32 and a maximum training epoch of
415 200. Cosine annealing learning rate scheduling was employed, with an initial learning rate of 0.001.
416 The optimal model weights were selected through an early stopping strategy based on validation-
417 set performance. The accuracy of flood forecasting was evaluated using four deterministic metrics.
418 The relative peak error (REP) was used to quantify the bias in flood-peak magnitude, while the
419 relative runoff volume error (REV) was used to measure the bias in total flood volume. Root-mean-
420 square error (RMSE) and Nash–Sutcliffe efficiency (NSE) were further employed to
421 comprehensively assess hydrograph fitting performance. The evaluation metrics were shown in



422 Table 5.

423 Table 5. Evaluation metrics

Type	Formula	Description
NSE	$NSE = 1 - \frac{\sum_{t=1}^N (\hat{Q}(t) - Q(t))^2}{\sum_{t=1}^N (Q(t) - \bar{Q})^2}$	Hydrograph skill relative to the observed.
REP	$REP = \frac{\hat{Q}_p - Q_p}{Q_p} \times 100\%$	Relative error of peak discharge.
REV	$REV = \frac{\hat{V} - V}{V} \times 100\%$	Relative error of runoff volume.
RMSE	$RMSE = \sqrt{\frac{1}{N} \sum_{n=1}^N (\hat{Q}(t) - Q(t))^2}$	Root mean square error of the forecast hydrograph.

424 Note: $Q(t)$ is the measured discharge hydrograph, $\hat{Q}(t)$ is the forecasted discharge hydrograph,
 425 Q_p and \hat{Q}_p denote the measured and forecasted peak discharges, V and \hat{V} denote the
 426 measured and forecasted runoff volume, T_p and \hat{T}_p denote the measured and forecasted peak
 427 time, \bar{Q} represent the average value of the measured discharge.

428 4 Results

429 4.1 Evaluation of Flood Forecasting Accuracy

430 As shown in Figure 3, the forecasting accuracy of all models deteriorated to varying degrees
 431 as the lead time increased. NSE generally decreased and RMSE increased for most events,
 432 indicating that uncertainties associated with future rainfall information, watershed state evolution,
 433 and routing related errors accumulated progressively. The LSTM model performed well in several
 434 high-peak flood events, particularly in controlling peak-discharge and runoff-volume errors, which
 435 suggested that LSTM was able to learn complex nonlinear rainfall–runoff relationships from
 436 historical samples. However, because it lacked explicit hydrological process constraints, LSTM was



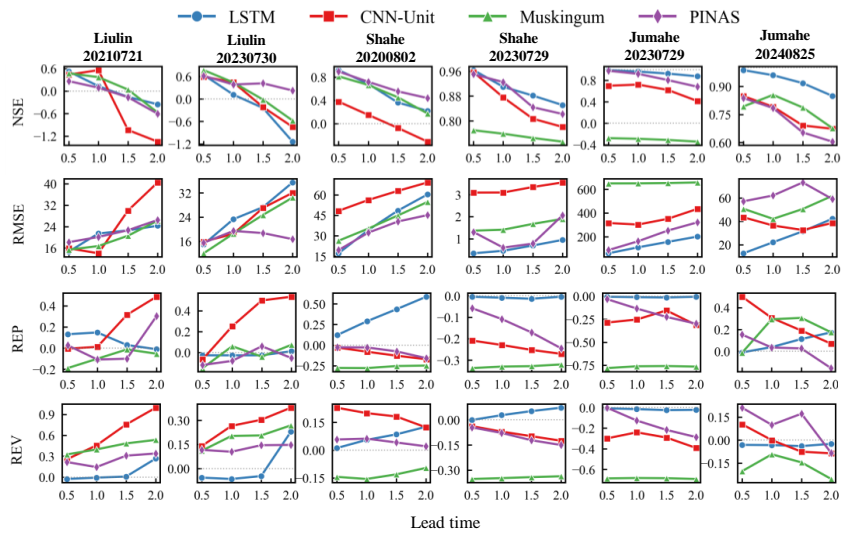
437 still prone to phase displacement and local hydrograph distortion, especially for multi-peak floods
438 in small basins and under longer lead times. Among the three hybrid hydrological models, PINAS
439 exhibited the most stable overall performance. Its NSE was generally higher than those of CNN-
440 Unit and Muskingum across most events and lead times, which indicated that the physically
441 constrained searchable routing structure was more flexible in adapting to different basin scales and
442 flood propagation characteristics. By contrast, CNN-Unit represented channel response through
443 unit-hydrograph convolution and was able to capture smoothing and lag effects to some extent, but
444 it was susceptible to peak amplification, spurious peaks, or phase shifts under strongly nonstationary
445 flood conditions. The Muskingum model had clear physical interpretability. However, its linear
446 storage–discharge assumption limited its ability to represent complex nonlinear flood-routing
447 processes, particularly during high-peak or rapidly varying flood events, where peak attenuation
448 and runoff-volume underestimation were more likely to occur.

449 Model performance also varied substantially among basins and flood types. As shown in Figure
450 4, the 20230729 event in the Shahe Basin exhibited a well defined dominant peak, and all models
451 reproduced the hydrograph reasonably well, which indicated that both the LSTM and the hybrid
452 hydrological models could achieve high accuracy when the rainfall–runoff response was relatively
453 regular. For the 20200802 event in the Shahe Basin, PINAS preserved the hydrograph shape more
454 effectively than CNN-Unit and Muskingum, suggesting that the searchable routing structure helped
455 reduce the accumulation of phase errors. The 20230729 event in the Jumahe Basin was characterized
456 by a high peak and large discharge. LSTM provided the best fit for peak magnitude and runoff
457 volume, while PINAS performed best among the hybrid models. In contrast, Muskingum
458 substantially underestimated the hydrograph, reflecting the limited applicability of traditional linear



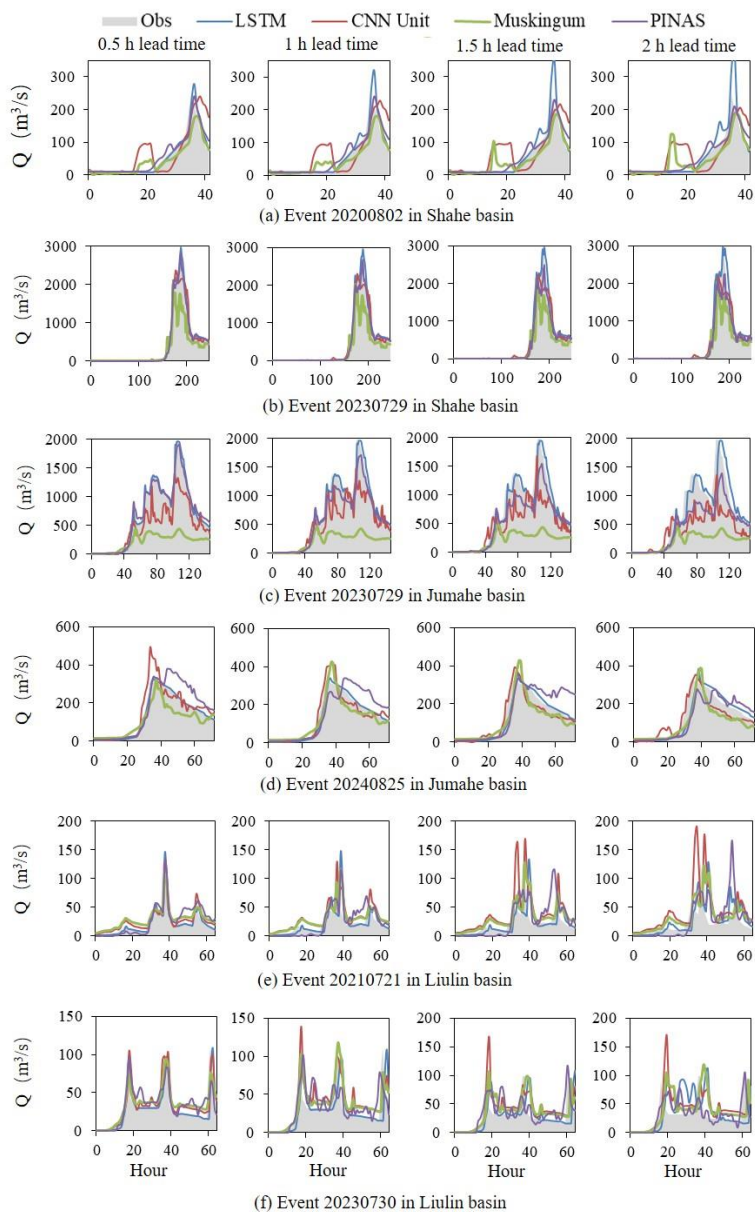
459 routing under strongly nonlinear flood propagation conditions. The two flood events in the Liulin
460 Basin were characterized by rapid small basin response and multi-peak hydrographs, which
461 considerably increased forecasting difficulty for all models. PINAS showed relatively stable
462 performance for the 20230730 event, indicating its advantage over fixed form physically based
463 routing schemes in representing multi-peak responses. Hydrograph comparisons further revealed
464 distinct sources of model error. LSTM showed strong capability in fitting peak magnitude, but its
465 hydrograph evolution was mainly sample-driven and lacked physical constraints on peak timing and
466 recession behavior. CNN-Unit could simulate propagation delay and hydrograph smoothing to some
467 extent, but it was sensitive to nonstationary inputs. Muskingum provided clear physical
468 interpretation but tended to over-smooth complex flood hydrographs. PINAS achieved a better
469 balance among peak propagation, hydrograph smoothing, and recession preservation by searching
470 routing-operator combinations under physical constraints, making it the best-performing scheme
471 among the three hybrid hydrological models.

472 In summary, PINAS showed stronger cross event robustness and long lead adaptability than
473 CNN-Unit and Muskingum, which indicated that incorporating a physically constrained searchable
474 routing structure into a differentiable hydrological framework improved the ability to represent
475 complex watershed response processes in flood forecasting.



476

477 Figure 3. Comparison of accuracy evaluation metrics across models



478

479 Figure 4. Forecast of Flood Process

480 4.2 Data dependency analysis

481 To evaluate the applicability and sample efficiency of the model under conditions of limited

482 data, the PINAS model was trained using 25%, 50%, 75% and 100% of the data to quantify the



483 impact of training sample size on flood forecasting skill. As shown in Figure 5, increasing the
484 training data volume significantly improved the forecasting stability of PINAS, although model
485 performance did not strictly show a monotonic improvement trend across all events and indicators.
486 When using the complete training dataset, the model typically achieved higher NSE and lower
487 RMSE in most flood events, which indicated that sufficient training samples enable the model to
488 more comprehensively learn the relationship between basin runoff generation and path response,
489 thereby improving the overall fitting effect of the hydrological map. In contrast, the results obtained
490 using 25% and 50% of the training data showed greater variability, particularly at longer lead times,
491 where NSE declined more rapidly and RMSE increased more significantly, which suggested that
492 insufficient samples amplified the effects of future meteorological input errors, state evolution
493 uncertainty, and routing-related phase errors on forecast performance.

494 The benefits of increasing training data were mainly reflected in three aspects. First, the ability
495 to identify the dominant flood peak was improved. As shown in Figure 6, for flood events with a
496 pronounced main peak, such as the 20230729 events in the Shahe and Jumahe basins, the
497 hydrographs obtained using 100% of the training data were closer to the measured flood processes,
498 with better preservation of peak magnitude, peak timing, and post-peak recession behavior. Second,
499 systematic bias was reduced. As the training data increased, REP and REV generally became smaller,
500 indicating that more sample coverage helped reduce overestimation or underestimation of flood-
501 peak discharge and total runoff volume. Third, long lead times became more stable. Low-data
502 configurations showed more evident error growth at the 1.5 and 2.0 h lead times, whereas training
503 with the full dataset partially suppressed the accumulation of errors with increasing lead times.
504 Nevertheless, in some events, the 75% or 50% data configurations achieved results that were

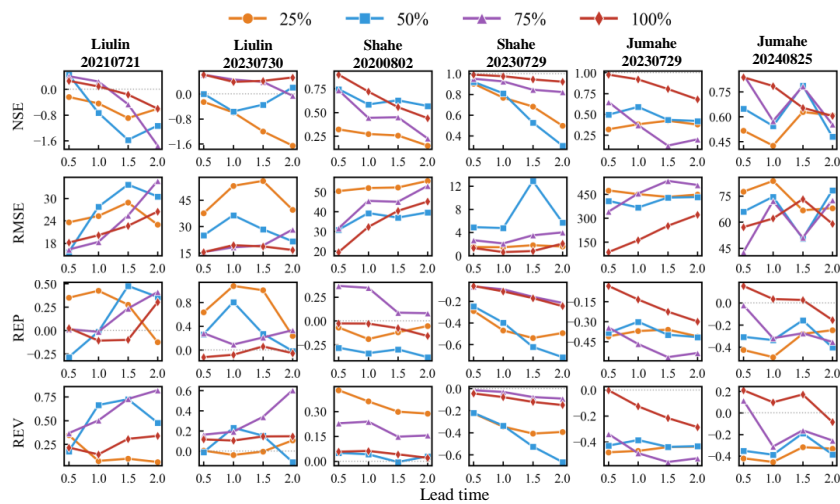


505 comparable to those obtained using 100% of the training data at specific lead times. This behavior
506 indicated that model performance was governed not only by the number of training samples, but
507 also by sample representativeness, flood type coverage, event composition in the training set, and
508 uncertainty in parameter optimization. For flood forecasting with limited event samples, simply
509 increasing the sample size may not directly translate into improved accuracy if the added samples
510 differ substantially from the hydrological response characteristics of the testing events. The
511 sensitivity to training data also varied among basins. The typical large-flood events in the Shahe and
512 Jumahe basins showed a relatively clear positive response to data expansion, especially for high-
513 peak events with a well defined dominant peak, where full-data training substantially improved peak
514 and volume estimation. By contrast, the Liulin Basin was characterized by rapid hydrological
515 response, multi-peak hydrographs, and localized intense rainfall forcing, making model
516 performance more sensitive to the composition of the training data. Even under relatively large
517 training data fractions, some long-lead forecasts still exhibited considerable variability, suggesting
518 that forecasting errors in small basins were controlled more by peak-timing displacement and multi-
519 peak structure mismatch than by sample size alone.

520 Overall, PINAS demonstrated a certain degree of sample efficiency. In some events, acceptable
521 forecasting performance was achieved using 50%–75% of the training data, indicating that the
522 physical constraints and searchable routing structure partly alleviated uncertainty caused by data
523 scarcity. However, from the perspective of cross basin and cross event stability, a larger training
524 dataset remained essential for achieving high accuracy. These results indicated that the performance
525 of PINAS depended not only on the model architecture, but also on whether the training samples
526 adequately represented typical flood types, extreme flood processes, and flash response

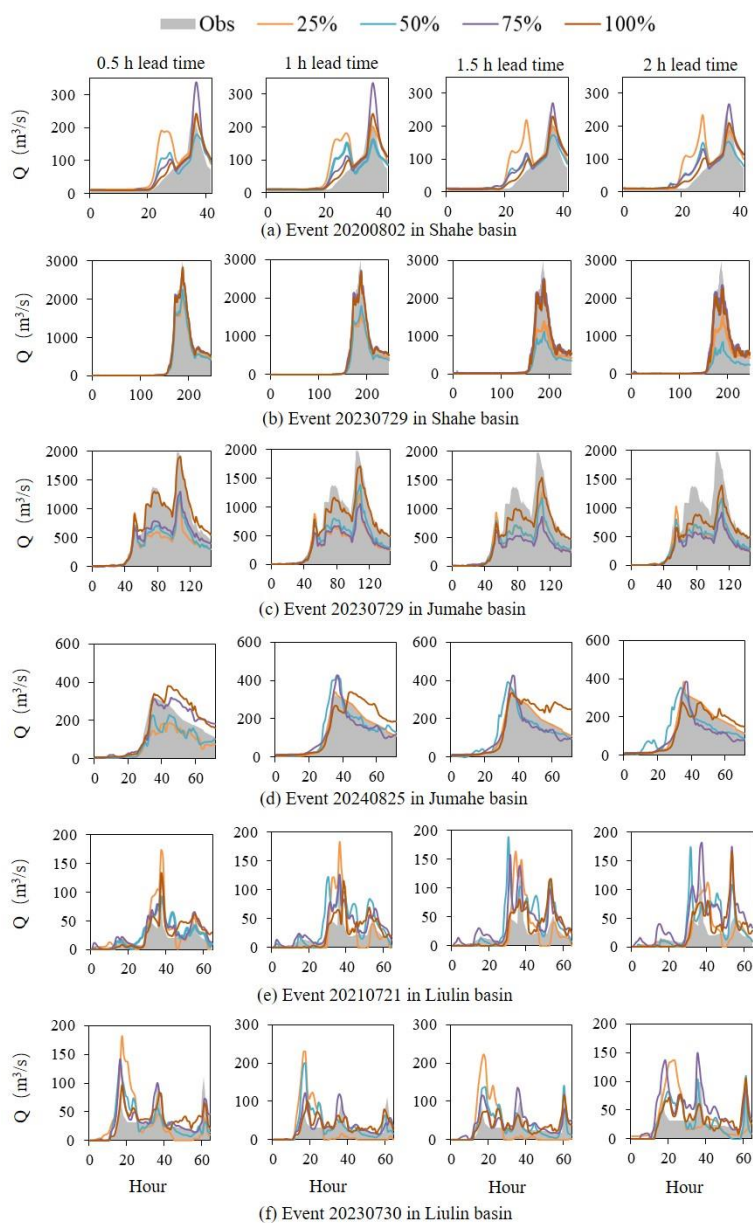


527 characteristics in small basins.



528

529 Figure 5. Forecasting accuracy indicators of the model trained with 25% - 100% of the data volume



530

531 Figure 6. Training model to forecast flood process with 25% - 100% of data volume

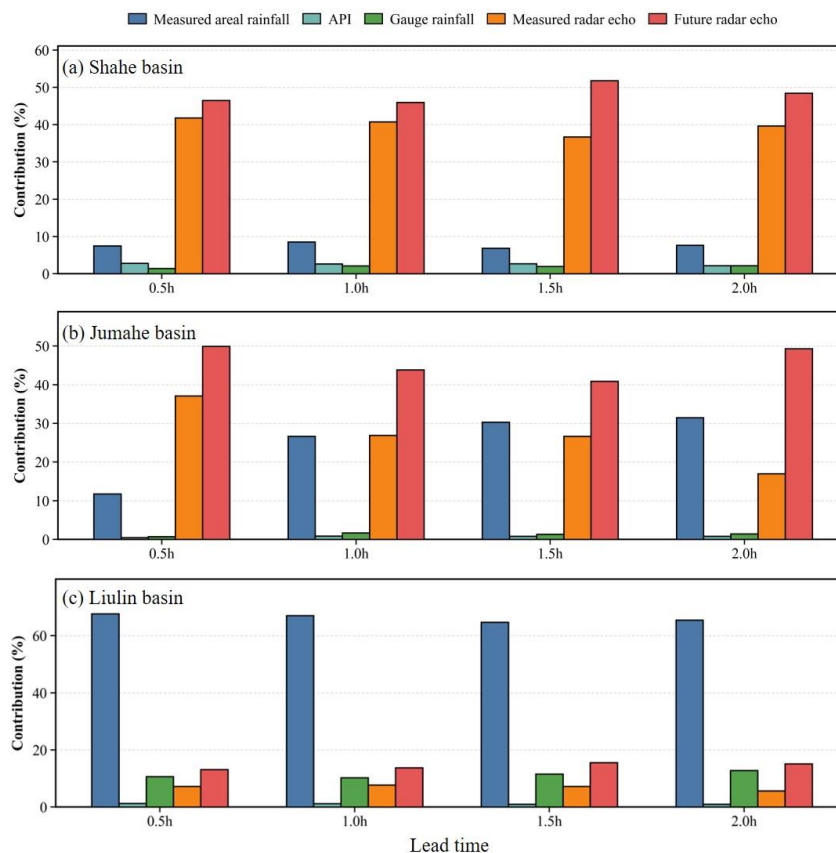
532 4.3 Interpretability analysis

533 To further interpret how the PINAS model utilized hydrometeorological information for flood

534 forecasting, SHAP analysis was conducted to quantify the contributions of different input factors.



535 SHAP values were aggregated by variable category and temporal dimension to identify both global
536 and local contributions of different input information to the model outputs. As shown in Figure 7,
537 the grouped SHAP results revealed clear basin differences in the dominant information sources.
538 Forecasts in the Shahe and Jumahe basins were mainly controlled by radar echo information, with
539 future radar echoes contributing the most and observed radar echoes ranking second. In the Shahe
540 Basin, future radar echoes accounted for approximately 50% of the total contribution across
541 different lead times, while observed radar echoes contributed approximately 40%, which indicated
542 that, for the Shahe Basin, the rainfall spatial structure and short-term evolution represented by radar
543 echoes provided critical information for inferring future flood dynamics. In the Jumahe Basin, as
544 the lead time increased, the contribution of areal rainfall gradually increased, whereas the
545 contribution of radar echo gradually decreases, which suggested that, in the larger basin, flood
546 response was controlled by antecedent accumulated rainfall and the future spatial evolution of
547 rainfall systems. In contrast, the dominant factor in the Liulin Basin was measured areal rainfall,
548 whose contribution remained above 60%, while radar echo variables contributed relatively less. This
549 pattern was consistent with the short concentration time and rapid flood response of the Liulin Basin,
550 indicating that short-term forecasting in this basin relied more strongly on measured rainfall as a
551 direct indicator of watershed saturation and rapid runoff-generation conditions.



552

553 Figure 7 SHAP Importance Grouping

554 The SHAP heatmaps further revealed the temporal roles of different information sources during

555 the forecasting process. As shown in Figure 8, for the Shahe Basin, the maximum value features of

556 both measured and future radar echoes made strong contributions at the current time step and the

557 adjacent future periods, which indicated that the model primarily used radar echo intensity extremes

558 to identify intense rainfall cells that could potentially trigger flood peaks. Meanwhile, features such

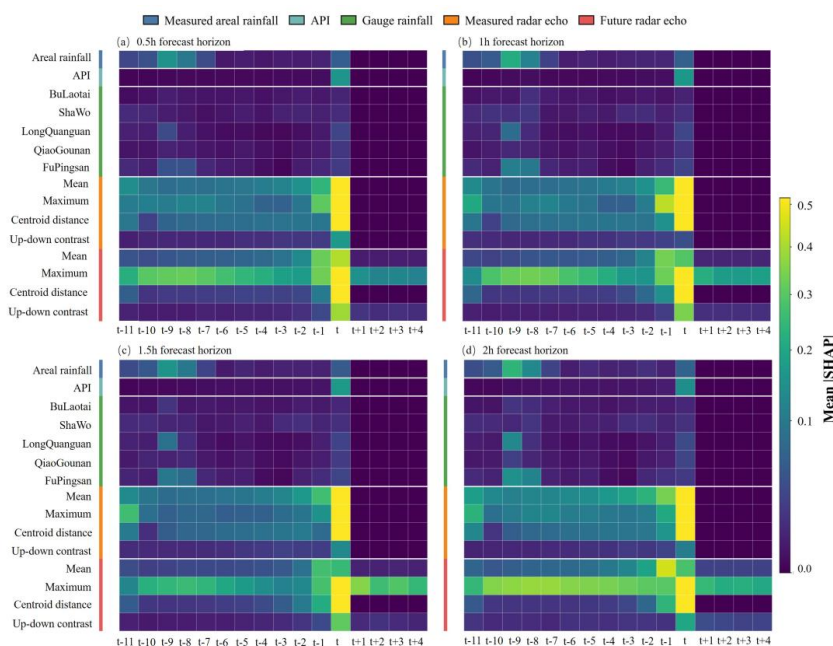
559 as mean radar echo intensity and centroid distance also contributed to the forecasting, suggesting

560 that the model considered not only rainfall intensity, but also the spatial displacement of rainfall

561 systems relative to the basin when inferring flood wave propagation. As shown in Figure 9, for the

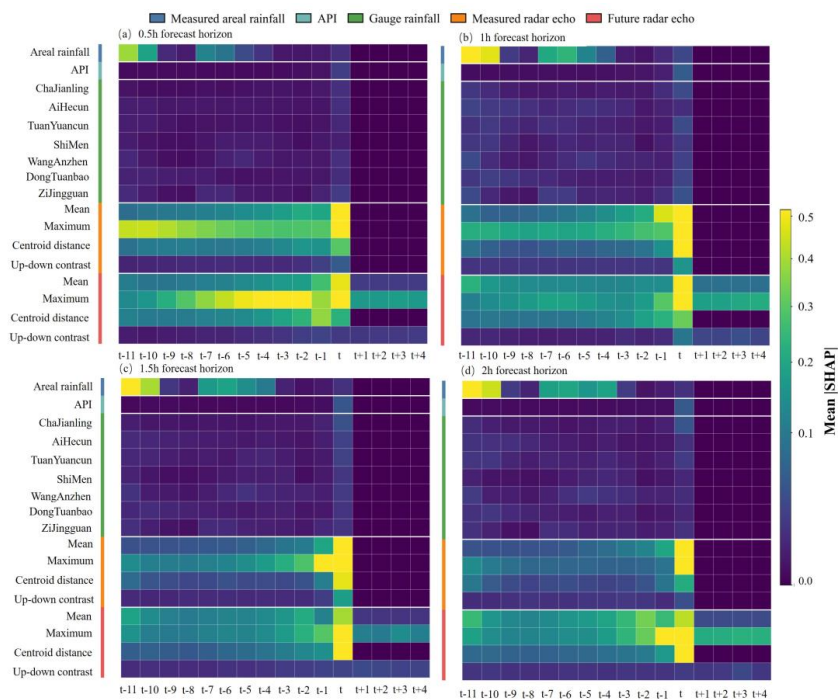


562 Jumahe Basin, areal rainfall showed relatively high contributions during earlier historical periods,
 563 whereas radar echo features contributed more strongly at the current time and over future periods,
 564 which indicated that flood forecasting in this basin depended simultaneously on antecedent rainfall
 565 accumulation and short-term information on rainfall movement. As shown in Fig. 10, for the Liulin
 566 Basin, areal rainfall maintained a substantial contribution throughout the entire historical window,
 567 while gauge rainfall also showed relatively high importance at certain time periods and stations. In
 568 contrast, the contribution of radar-echo information was concentrated mainly around the current
 569 time and adjacent periods, which suggested that flood forecasting in small basins was more strongly
 570 controlled by rainfall accumulation and localized gauge rainfall, whereas the additional value of
 571 radar extrapolation information was relatively limited.



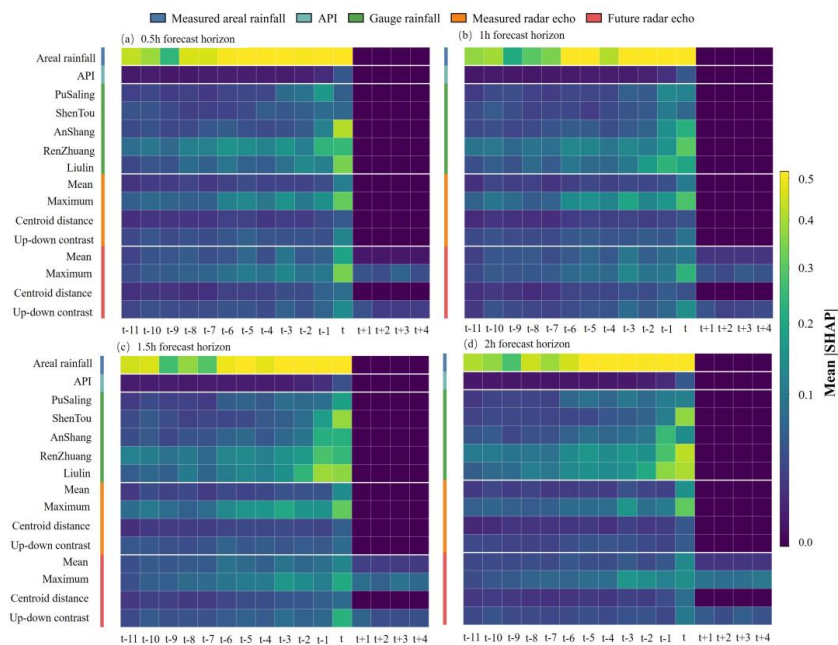
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573 Figure 8. SHAP heatmap of feature contributions for the Shahe Basin model



574

575 Figure 9. SHAP heatmap of feature contributions for the Jumahe Basin model



576

577 Figure 10. SHAP heatmap of feature contributions the Liulin Basin model



578 Overall, the SHAP results indicated that the forecasting produced by the hybrid hydrological
579 model based on PINAS were not dependent on a single dominant input variable. Instead, the model
580 adaptively assigned importance to different information sources based on the scale of the basin and
581 the characteristics of the flood event. For medium-sized basins such as the Shahe and Jumahe basins,
582 radar echoes and the short-term extrapolated features played a central role in flood forecasting,
583 particularly in characterizing the movement of future rainfall systems and the subsequent evolution
584 of flood peaks. In contrast, for the Liulin Basin, measured areal rainfall and gauge rainfall remained
585 the primary information sources, suggesting that in rapidly responding watersheds, measured areal
586 rainfall and local rainfall observations were more crucial for short-term forecasting. These findings
587 provided interpretability evidence for the necessity of hydrometeorological data fusion and
588 demonstrated that the hybrid hydrological model can learn the contribution patterns of hydrological
589 significance features within the framework of physical constraints.

590 **5 Discussion**

591 The differentiable hybrid hydrological forecasting framework coupled with future
592 meteorological information was developed and its short-term flood forecasting capability was
593 evaluated across representative basins and flood events. The results indicated that physically
594 constrained routing, particularly PINAS, can improve the stability of hydrograph forecasting.
595 Nevertheless, several limitations need to be further discussed.

596 The uncertainty of future meteorological forcing remained the main limiting factor for flood
597 forecasting. In this study, radar echoes were extrapolated to provide rainfall information within the
598 next 0-2 hours. However, the optical flow field calculation relied on the continuity of echo



599 movement and cannot fully reflect the rapid deformation of convection. Previous studies have
600 shown that radar based nowcasting was generally effective at short lead times, but forecast errors
601 increased rapidly for intense convective rainfall, and traditional advection based approaches
602 struggle to represent nonlinear precipitation evolution (Ravuri et al., 2021). Therefore, the decline
603 in accuracy at longer lead times may not only be attributed to the errors of the hydrological model,
604 but also to the propagation of radar echo calculation errors to runoff forecasting. Future work should
605 adopt ensemble radar extrapolation methods (Chung et al., 2025), deep generative nowcasting
606 (Franch et al., 2025), or multi-source numerical weather prediction (Wang et al., 2024) to replace
607 deterministic future forcing with probabilistic precipitation inputs, thereby explicitly propagating
608 meteorological uncertainty through the hydrological forecasting chain.

609 The hybrid hydrological model mainly covered aspects such as runoff generation, baseflow
610 recession, basin confluence, and channel routing, but does not explicitly account for spatially
611 heterogeneous soil moisture, dynamic land-surface conditions, reservoir and gate operations, and
612 urbanization effects. These processes can significantly alter the rainfall-runoff relationship in small
613 and medium-sized watersheds, especially during extreme rainstorms or under different antecedent
614 wetness conditions. Recent studies have suggested that differentiable hydrological modelling
615 provides a promising approach for connecting process-based models with deep learning, because it
616 enables end-to-end training while retaining physically meaningful states and parameters (Tsai et al.,
617 2021). Therefore, future extensions should introduce learnable soil-water states, dynamic infiltration
618 parameters, and hillslope concentration parameters, shifting the framework from physically
619 constrained routing toward a distributed forecasting model with interpretable hydrological
620 states (Ettalbi et al., 2025).



621 Data dependency analysis indicated that model performance was more influenced by the
622 representativeness of the samples and the coverage of flood types. Similar findings have been
623 observed in studies on differentiable parameter learning and physics-based deep learning, where a
624 larger training dataset can enhance generalization ability and physical consistency, provided that the
625 samples cover a sufficiently diverse set of hydro-meteorological conditions (Feng et al., 2024;
626 Huynh et al., 2025). This study collected flood events from three different scales and response
627 characteristic basins, but the total number of events was still limited, especially for extreme floods
628 and multi-peak hydrographs. Future work should expand the number of basins and climatic
629 geomorphological settings, while also using physically consistent data augmentation or extreme-
630 event resampling to improve model learning for rare high-peak and flash flood events.

631 The assessment in this study mainly relies on deterministic indicators, which are useful for
632 comparing hydrological fit, peak flow error, and flood volume error, but they are insufficient to
633 comprehensively describe the risk uncertainty in operational flood forecasting. Operating short-term
634 flood forecasting not only requires the hydrological map, but also requires confidence intervals, the
635 probability of exceeding the warning threshold, and the estimation of false alarm and missed alarm
636 risks. Recent studies have shown that quantile regression, mean variance estimation, extreme
637 quantile regression, and generative probability models can provide useful uncertainty information
638 for flood or runoff forecasting (Pasche and Engelke, 2024; Jahangir and Quilty, 2024). Therefore,
639 future work should expand the current framework by incorporating methods such as quantile
640 regression, deep ensembles, or conditional generative models, moving towards probabilistic
641 prediction, so that the model can provide deterministic hydrological maps and reliable flood risk
642 intervals at different lead times. (Xiang et al., 2025; Ou et al., 2025)



643 In general, this study demonstrated that integrating future meteorological information with
644 physically informed channel routing can improve short-term flood forecasting for small and
645 medium-sized basins. However, the model performance was still affected by the uncertainty of
646 future rainfall, the representativeness of the samples, and the simplified process representation.
647 Future research should focus on probabilistic meteorological forcing, distributed physical state
648 representation, and enhanced training for extreme events, in order to support the transition of hybrid
649 intelligent flood forecasting models from experimental research to reliable operational applications.

650 **6 Conclusion**

651 This study developed a differentiable hybrid hydrological forecasting framework coupled with
652 future meteorological information for short-term flood forecasting in small and medium-sized
653 basins. Flood forecasting experiments were conducted in the Shahe, Jumahe, and Liulin basins. The
654 main conclusions are as follows.

655 (1) Forecasting accuracy decreased as lead time increasing, indicating the progressive
656 accumulation of uncertainty from future forcing, hydrological state evolution, and channel routing
657 phase errors. Among the three hybrid hydrological models, PINAS showed the most stable overall
658 performance. Compared with Muskingum and CNN unit, PINAS used a physically constrained
659 searchable routing structure to more flexibly represent basin-scale differences and flood-wave
660 propagation characteristics, achieving a better balance among peak propagation, hydrograph
661 smoothing, and recession representation.

662 (2) The size of the training sample had a significant impact on the performance of PINAS.
663 Increasing the amount of training data usually leads to an improvement in NSE, a reduction in



664 RMSE, and a decrease in the deviations of peak flow and runoff. However, the improvement was
665 not entirely a monotonic trend, indicating that sample representativeness, flood-type coverage, and
666 the inclusion of extreme events are also critical for model generalization.

667 (3) SHAP interpretability analysis revealed significant differences in the main sources of
668 information among different basins. Forecasts in the Shahe and Jumahe basins were mainly
669 influenced by measured data and future radar echoes, indicating that the spatial structure of rainfall
670 and the short-term evolution of storms were of great significance in medium-sized basins. In contrast,
671 forecasts in Liulin Basin were primarily driven by measured areal rainfall and gauge rainfall,
672 reflecting the strong dependence of small basin floods on measured rainfall accumulation and
673 localized rainfall measurements.

674 Overall, integrating the hybrid hydrological model with future meteorological information can
675 enhance the stability and hydrological rationality of short-term flood forecasting for small and
676 medium-sized basins. The proposed framework provides a feasible approach for intelligent flood
677 forecasting models that combine forecasting accuracy, physical consistency, and interpretability.

678 **CRediT authorship contribution statement**

679 **Leijing Li:** Conceptualization, Methodology, Writing–original draft, Writing–review &
680 editing. **Liutianjiao Hong:** Conceptualization, Methodology, Software, Validation, Writing–
681 original draft. **Jianzhu Li:** Validation, Writing – review & editing. **Peng Shi:** Methodology,
682 Visualization. **Shuaihang Wang:** Supervision, Data curation. **Jiyang Tian:** Data curation.

683 **Declaration of Competing Interest**

684 The authors declare that they have no known competing financial interests or personal
685 relationships that could have appeared to influence the work reported in this paper.



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689 **Data availability**

690 Data will be made available on request.

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