



Green Above, Dry Below: Satellite Multi-Indices Reveal Contrasting Drought Trajectories Across the Tekeze Basin (2000–2025)

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Research Highlights

- Seven satellite-derived drought indices integrated over Tekeze Basin (2000-2025)
- Vegetation indices show positive trends (71-86%) while soil moisture declines
- Northern Tekeze sub-basin identified as a persistent multi-dimensional drought hotspot
- TCI-VCI correlation reaches $r = 0.97$; TVDI-VHI reaches $r = -0.97$ in peak season
- Multi-index framework reveals decoupling masked by single-index monitoring



31 Abstract

32 Drought is one of the most pervasive and damaging climate-related hazards in the Horn of Africa,
33 with profound impacts on agricultural productivity, water resources, ecosystem functioning, and
34 socioeconomic development. The Tekeze Basin of northern Ethiopia is particularly vulnerable due
35 to its high rainfall variability, recurrent dry spells, and dependence on rain-fed agriculture. Although
36 numerous drought studies have been conducted in Ethiopia, few have simultaneously integrated
37 vegetation, thermal, soil moisture, precipitation, and surface water indicators to provide a
38 comprehensive assessment of drought dynamics. This study presents a multi-dimensional
39 spatiotemporal analysis of drought conditions in the Tekeze Basin (2000–2025) using seven satellite-
40 derived drought indices: the Vegetation Condition Index (VCI), Temperature Condition Index (TCI),
41 Vegetation Health Index (VHI), Temperature Vegetation Dryness Index (TVDI), Soil Moisture
42 Condition Index (SMCI), Precipitation Condition Index (PCI), and Normalized Difference Water
43 Index (NDWI). The indices were derived from MODIS, soil moisture, and CHIRPS precipitation
44 datasets. Long-term pixel-wise temporal trends were assessed using the non-parametric Mann-
45 Kendall (MK) test and Sen's slope estimator across the growing season (April–September), while
46 inter-index relationships were examined through monthly Pearson correlation analysis. Results
47 reveal that VCI, TCI, VHI, and NDWI exhibited predominantly positive trends over the study period
48 (71.5–86.4% positive spatial coverage), indicating general improvement in vegetation and water
49 conditions across most of the basin. In contrast, TVDI and SMCI showed dominant negative trends
50 (82.2% and 55.3% negative coverage, respectively), indicating increasing land-surface dryness and
51 declining soil moisture, particularly in the basin interior and lowlands. PCI displayed spatially
52 heterogeneous patterns (54.7% negative coverage), signaling progressive precipitation deficits in the
53 northern sub-basin. Monthly correlation analysis revealed strong positive relationships between TCI
54 and VCI ($r = 0.48\text{--}0.97$) and strong negative relationships between TVDI and VHI ($r = -0.54$ to
55 -0.97) across all months. SMCI and PCI showed strengthening positive intercorrelations during the
56 peak rainy season (June–September). The contrasting trajectories observed between vegetation-
57 based indicators and soil moisture-related indices suggest a potential decoupling between vegetation
58 response and underlying hydrological conditions. These findings demonstrate the value of multi-
59 index remote sensing approaches for capturing the complex and multidimensional nature of drought
60 and provide important scientific evidence for drought early warning, water resource management,
61 and climate adaptation planning in northern Ethiopia.

62 Keywords: Drought Monitoring; Spatiotemporal Analysis; Growing Season; Remote sensing;
63 Tekeze Basin;

64 1. Introduction

65 Drought is a complex and slowly developing natural hazard that affects more people globally than
66 any other climate-related disaster and causes substantial economic, environmental, and social losses
67 (Wilhite, 2000; Sheffield and Wood, 2011). Unlike sudden-onset hazards such as floods and
68 earthquakes, drought evolves gradually and manifests through multiple interconnected processes,
69 including precipitation deficits, soil moisture depletion, vegetation stress, reduced streamflow, and
70 declining water availability. These impacts are particularly severe in semi-arid and arid regions
71 where livelihoods depend heavily on rain-fed agriculture and natural resources (Dai, 2013; Masih et
72 al., 2014).

73



74 The Horn of Africa is recognized as one of the world's most drought-prone regions, experiencing
75 frequent drought episodes driven by climate variability, changing atmospheric circulation patterns,
76 and increasing climate extremes (Funk et al., 2015; Haile et al., 2020). Ethiopia has experienced
77 numerous severe drought events over recent decades, resulting in widespread crop failures, livestock
78 losses, food insecurity, and socioeconomic disruption (Bewket and Conway, 2007; Viste et al.,
79 2013). The country's rainfall regime is strongly influenced by the Inter-Tropical Convergence Zone
80 (ITCZ), Indian Ocean sea-surface temperature variability, and complex topographic conditions,
81 which contribute to pronounced spatial and temporal rainfall variability (Diro et al., 2011; Seleshi
82 and Zanke, 2004). Among Ethiopia's major river basins, the Tekeze Basin is particularly vulnerable
83 because of its semi-arid climate, recurrent rainfall deficits, land degradation, and strong dependence
84 on seasonal precipitation for agricultural production and water resources (Nyssen et al., 2004;
85 Gebremicael et al., 2013; Tesfaye et al., 2017).

86 Reliable drought monitoring is essential for effective drought preparedness, early warning, and
87 resource management. Traditional drought indicators such as the Standardized Precipitation Index
88 (SPI) and Palmer Drought Severity Index (PDSI) have been widely applied for drought assessment
89 and operational monitoring (Palmer, 1965; McKee et al., 1993). Although these indices effectively
90 characterize meteorological drought conditions, they do not directly capture vegetation stress, soil
91 moisture dynamics, land surface thermal conditions, or changes in surface water availability that
92 collectively determine drought impacts on ecosystems and agriculture. Consequently, increasing
93 attention has been directed toward remote sensing approaches that provide spatially continuous
94 observations of drought-related variables over large geographic areas (AghaKouchak et al., 2015;
95 Zhang et al., 2017).

96 Advances in satellite remote sensing have enabled the development of numerous drought indicators
97 representing different components of the land-atmosphere system. Vegetation-based indices such as
98 the Vegetation Condition Index (VCI) and Vegetation Health Index (VHI) quantify vegetation stress
99 and productivity (Kogan, 1990; Kogan, 1997). The Temperature Condition Index (TCI) reflects
100 thermal stress associated with drought conditions (Kogan, 1995), whereas the Temperature
101 Vegetation Dryness Index (TVDI) provides information on land-surface dryness through vegetation-
102 temperature relationships (Sandholt et al., 2002). Soil Moisture Condition Index (SMCI)
103 characterizes moisture availability within the root zone (Du et al., 2013), while the Precipitation
104 Condition Index (PCI) represents meteorological drought conditions based on rainfall variability
105 (Rhee et al., 2010). The Normalized Difference Water Index (NDWI) provides additional
106 information on vegetation water content and surface moisture conditions (Gao, 1996). Because each
107 index captures a different aspect of drought, no single indicator can adequately represent the full
108 complexity of drought processes.

109 Recent studies have increasingly demonstrated the advantages of integrating multiple drought
110 indicators to improve drought characterization and monitoring accuracy (Yao et al., 2020; Ayana et
111 al., 2022; Woldeesenbet et al., 2023). Multi-index frameworks are particularly valuable in
112 environmentally heterogeneous regions where meteorological, agricultural, and hydrological
113 drought conditions may evolve differently across space and time. Despite growing interest in multi-
114 dimensional drought assessment, comprehensive long-term studies integrating vegetation,
115 temperature, soil moisture, precipitation, and water-related indicators remain limited in Ethiopia.
116 Existing drought investigations in northern Ethiopia have primarily focused on individual drought
117 indicators, short analysis periods, or localized study areas, leaving significant uncertainty regarding
118 the interactions among different drought dimensions at the basin scale.



119

120 The Tekeze Basin represents a critical case study for advancing drought monitoring in Ethiopia
121 because of its high climatic variability, ecological sensitivity, and socioeconomic importance. The
122 basin encompasses diverse topographic and climatic conditions that may produce substantial spatial
123 heterogeneity in drought responses. Furthermore, the agricultural growing season extending from
124 April to September, including the Belg and Kiremt rainfall periods, is crucial for crop production,
125 water availability, and rural livelihoods throughout the basin (Segele and Lamb, 2005; Conway et
126 al., 2009). Understanding drought dynamics during this period is therefore essential for
127 strengthening drought risk management and adaptation planning.

128 Against this background, this study presents a comprehensive multi-index assessment of drought
129 dynamics in the Tekeze Basin during the growing season from 2000 to 2025. By integrating seven
130 satellite-derived drought indices representing vegetation condition, thermal stress, soil moisture
131 status, precipitation variability, and surface water conditions, the study seeks to provide a more
132 complete understanding of drought evolution than can be achieved through individual indicators
133 alone. Particular attention is given to identifying potential divergences among drought dimensions,
134 including situations in which vegetation conditions improve while underlying soil moisture and
135 hydrological conditions deteriorate.

136 The specific objectives of this study are to: (1) characterize the spatial and temporal patterns of
137 drought across the Tekeze Basin during the growing season using seven satellite-derived drought
138 indices; (2) quantify long-term drought trends through pixel-wise Mann–Kendall trend analysis and
139 Sen's slope estimation; and (3) evaluate interrelationships among vegetation, thermal, soil moisture,
140 precipitation, and water-related drought indicators using monthly correlation analysis. The findings
141 are expected to contribute to improved drought monitoring frameworks, support operational drought
142 early warning systems, and provide scientific evidence for climate-resilient water and agricultural
143 management in northern Ethiopia.

144 2. Data and Methods

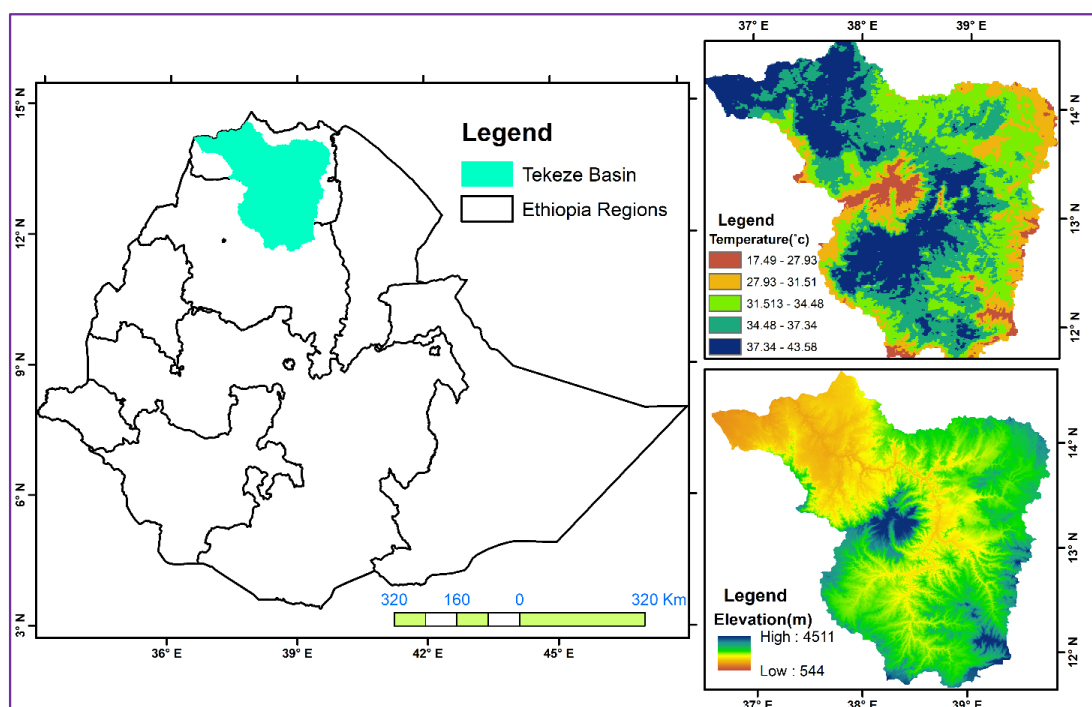
145 2.1 Study Area

146 The Tekeze Basin (approximately 36°–40°E, 11.5°–15°N) is located in northern and northwestern
147 Ethiopia and covers an area of approximately 82,000 km² (Figure 1). The basin drains northward
148 into Sudan, where the Tekeze River merges with the Setit and subsequently joins the Atbara River
149 before discharging into the Nile. The climate of the Tekeze Basin is predominantly semi-arid to arid,
150 with mean annual rainfall ranging from less than 300 mm in the northern and eastern lowlands to
151 over 1,200 mm in the western highland escarpments. Rainfall is highly seasonal, concentrated in two
152 distinct wet seasons: the short Belg rains (March–May, contributing 10–25% of annual rainfall) and
153 the main Kiremt rainy season (June–September), contributing 60–80% of annual rainfall (Seleshi
154 and Zanke, 2004; Segele and Lamb, 2005; Diro et al., 2011; Nicholson, 2018). Mean annual
155 temperatures range from approximately 12°C in the western highlands to more than 34°C in the Afar
156 lowlands. Land cover in the basin includes rainfed subsistence agriculture (predominantly in the
157 highlands and mid-elevations), semi-arid shrubland and open woodland (intermediate zones),
158 seasonal and perennial riverine vegetation along major tributaries, and bare or sparsely vegetated
159 surfaces in the Afar lowlands (Nyssen et al., 2004; Gebre et al., 2015). Ongoing watershed
160 rehabilitation and conservation programs, including area check dams and hillside terracing, have



161 progressively restored vegetation in parts of Tigray and northern Amhara, and the effects of these
162 interventions are detectable in satellite-derived vegetation trends (Descheemaeker et al., 2022;
163 Tesfaye et al., 2017).

164 The Tekeze Basin has experienced several catastrophic drought events during the study period,
165 including the 2002–2003, 2008–2009, 2015–2016 (El Niño-driven), and 2021–2023 (compounded
166 Tigray conflict–drought) crises, making it a priority region for long-term drought monitoring and
167 early warning system development.



168
169 **Figure 1.** Location map of the Tekeze Basin in Ethiopia, showing its geographical location, major
170 elevation, temperature, and administrative boundaries.

171 2.2 Datasets

172 **2.2.1 MODIS Vegetation Indices (MOD13Q1 V6.1):** The Moderate Resolution Imaging
173 Spectroradiometer (MODIS)/Terra MOD13Q1 product provides NDVI and EVI. NDVI was used to
174 compute VCI, VHI, and NDWI, and to parameterize the dry edge of the TVDI, LST/NDVI feature
175 space (Didan, 2015).

176 **2.2.2 MODIS Land Surface Temperature (MOD11A2 V6.1):** The MODIS/Terra MOD11A2
177 product provides daytime land surface temperature (LST). LST data were used to compute TCI,
178 VHI, and TVDI (Wan et al., 2015).

179 **2.2.3 NASA-USDA SMAP Enhanced L3 Soil Moisture:** Surface soil moisture data were obtained
180 from the NASA-USDA Global Soil Moisture product derived from the Soil Moisture Active Passive
181 (SMAP) Level-3 Enhanced passive microwave sensor. SMAP data (available continuously from



182 April 2015) were supplemented for the 2000–2015 period using the SMAP-Sentinel fusion product
183 and cross-validated with MODIS-based surface moisture proxies (TVDI) to ensure temporal
184 consistency across the full 2000–2025 period. SMCI was derived from these harmonized soil
185 moisture data.

186 **2.2.4 CHIRPS v2.0 Precipitation:** Monthly precipitation data from the Climate Hazards Group
187 Infrared Precipitation with Station (CHIRPS) were used to compute PCI. CHIRPS provides quasi-
188 global (~50°N–50°S) daily and monthly precipitation from 1981 to near-present, blending cold cloud
189 duration satellite estimates with in-situ gauge observations. CHIRPS has been rigorously validated
190 over Ethiopia and East Africa and has demonstrated superior performance relative to other gridded
191 products in the region (Funk et al., 2015; Dinku et al., 2018; Ayehu et al., 2018).

192 2.3 Drought Assessment Methodology

193 This study integrates four primary remote sensing and climate datasets into a unified Google Earth
194 Engine (GEE) processing platform using Python API in Google Colaboratory. We acquire MODIS
195 MOD13Q1 (version 6) vegetation indices (NDVI, NDWI) at 250 m resolution (16-day composites),
196 MODIS MOD11A2 (version 6) land surface temperature LST) at 1 km resolution (8-day
197 composites), NASA-USDA SMAP enhanced soil moisture at ~9 km resolution, and CHIRPS
198 (version 2) precipitation at ~5.5 km resolution. Within GEE, we compute seven complementary
199 drought indices: the VCI and TCI, which we combine into the VHI; the TVDI derived from the LST-
200 NDVI feature space; the SMCI from SMAP data; PCI from CHIRPS; and NDWI from MODIS
201 reflectance bands. All indices are computed monthly for the six-month growing season (April–
202 September) across the 26-year study period (2000–2025) at 250 m pixel resolution, enabling
203 spatially continuous, temporally consistent characterization of multiple drought dimensions
204 simultaneously. All datasets were reprojected to the WGS84/Geographic coordinate system, and
205 clipped to the Tekeze Basin boundary derived from the HydroSHEDS Level 5 hydrographic dataset
206 (Lehner et al., 2008). These operations were performed within GEE and QGIS. Following index
207 computation in GEE, we apply pixel-wise Mann-Kendall trend testing and Sen's slope estimation to
208 quantify long-term spatiotemporal drought trends across the entire basin, generating trend maps that
209 show both the direction (MK Tau) and magnitude (Sen's slope) of change for each index. In parallel,
210 we use R statistical software (ggplot2, corrplot, and trend packages) to generate multi-year time
211 series graphs showing seasonal dynamics and to compute monthly Pearson correlation matrices
212 revealing inter-index relationships stratified by growing-season month (Wei and Simko, 2017). We
213 then use QGIS for cartographic visualization and composition of publication-quality multi-panel
214 trend maps showing spatial patterns of drought change. This integrated framework yields actionable
215 findings regarding spatiotemporal drought patterns, long-term trends by region and season, and inter-
216 index correlations that directly inform drought early warning system design and climate-smart
217 resource management decisions.

218 2.4 Drought Indices Extraction

219 2.4.1 VCI

220 VCI, developed by Kogan (1990, 1995), quantifies vegetation stress relative to historical NDVI
221 extremes for each pixel and calendar period:

$$222 \quad VCI = \frac{NDVI_t - NDVI_{min}}{NDVI_{max} - NDVI_{min}} * 100 \dots \dots \dots (1)$$



where $NDVI_{min}$ and $NDVI_{max}$ are the long-term pixel-wise minimum and maximum NDVI computed across 2000–2025. VCI ranges from 0 (extreme drought stress) to 100 (optimal vegetation condition); values below 40 indicate drought (Kogan, 1995; Rhee et al., 2010).

2.4.2 TCI

TCI quantifies thermal stress relative to historical LST extremes:

$$TCI = \frac{LST_{max} - LST_t}{LST_{max} - LST_{min}} * 100 \quad (2)$$

where LST_{max} and LST_{min} are pixel-wise long-term extremes. Higher TCI values indicate cooler and more favorable thermal conditions; lower values indicate heat stress associated with drought (Kogan, 1995).

2.4.3 VHI

VHI integrates VCI and TCI into a composite drought severity metric (Kogan, 1997):

$$VHI = \alpha \times VCI + (1 - \alpha) \times TCI \quad (3)$$

where $\alpha = 0.5$, assigning equal weight to vegetation and thermal components as recommended by Kogan (1997) in the absence of site-specific calibration data. VHI values below 40 indicate drought conditions of varying severity.

2.4.4 TVDI

TVDI was computed following Sandholt et al. (2002) using the LST and NDVI triangle space:

$$TVDI = \frac{LST - LST_{min}(NDVI)}{LST_{dry}(NDVI) - LST_{min}(NDVI)} \quad (4)$$

where $LST_{min}(NDVI)$ represents the wet edge (minimum LST for a given NDVI) and $LST_{dry}(NDVI) = a + b \times NDVI$ is the dry edge, derived by fitting a linear regression to the upper boundary of the monthly LST/NDVI scatter plot within the basin for each time step. TVDI ranges from 0 (fully wet, no water stress) to 1 (extreme dryness), with higher values indicating greater surface moisture stress (Sandholt et al., 2002; Wang et al., 2004; Holzman et al., 2014).

2.4.5 SMCI

SMCI was derived from SMAP-based soil moisture data following Du et al. (2013):

$$SMCI = \frac{SM - SM_{min}}{SM_{max} - SM_{min}} \quad (5)$$

where SM is observed monthly soil moisture, and SM_{min} and SM_{max} are pixel-wise long-term minima and maxima. SMCI values near 100 indicate wet soil conditions, and values near 0 indicate severe soil moisture drought.

2.4.6 PCI



256 PCI captures meteorological drought and was computed from CHIRPS monthly precipitation
257 following Rhee et al. (2010):

$$258 \quad \text{PCI} = \frac{(P - P_{\min})}{(P_{\max} - P_{\min})} * 100 \dots \dots \dots (6)$$

259 where P is the monthly CHIRPS precipitation, and $P_{\min}$ and $P_{\max}$ are pixel-wise long-term extremes
260 over 2000–2025.

261 **2.4.7 NDWI**

262 NDWI was computed from MODIS surface reflectance bands following Gao (1996):

$$263 \quad \text{NDWI} = \frac{(\text{Band}_{\text{NIR}} - \text{Band}_{\text{SWIR}})}{(\text{Band}_{\text{NIR}} + \text{Band}_{\text{SWIR}})} \dots \dots \dots (7)$$

264 NDWI is sensitive to vegetation liquid water content and soil surface moisture. Values typically
265 range from −1 to +1, with higher values indicating greater moisture availability.

266 **2.5 Mann-Kendall Trend Analysis and Sen's Slope Estimation**

267 Monotonic temporal trends in each drought index time series were assessed pixel-by-pixel using the
268 non-parametric MK test (Mann, 1945; Kendall, 1975). The MK test is widely preferred in hydro-
269 climatic trend analysis due to its robustness to non-normality, seasonality, outliers, and missing data
270 (Hamed and Rao, 1998; Yue and Wang, 2002). The MK statistic S is defined as:

$$271 \quad S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sgn}(x_j - x_k) \dots \dots \dots (8)$$

272 where x_k and x_j are data values at time steps k and j, and $\text{sgn}(\cdot)$ is the sign function returning +1,
273 0, or −1. Positive (negative) S indicates an upward (downward) trend. Statistical significance was
274 assessed at the 95% confidence level ($p < 0.05$) using the standardized test statistic Z. The
275 normalized Kendall Tau (τ) was computed as:

$$276 \quad \tau = S / (n(n-1)/2) \dots \dots \dots (9)$$

277 and mapped spatially for each index as an indicator of trend direction and relative magnitude. The
278 magnitude of each detected trend was quantified using Sen's slope estimator (Sen, 1968):

$$279 \quad \beta = \text{median} \left((x_j - x_k) / (j - k) \right) \dots \dots \dots (10)$$

280 Sen's slope (expressed in index units per year) represents a robust, outlier-resistant estimate of the
281 trend magnitude and was mapped spatially alongside MK Tau to provide both directional and
282 quantitative information about long-term drought change.

283 **2.6 Inter-Index Pearson Correlation Analysis**

284 Monthly Pearson correlation coefficients I were computed between all pairwise combinations of the
285 seven drought indices using a 26-year time series (2000–2025) of basin-averaged monthly mean
286 values. Correlations were considered statistically significant at $p < 0.05$ based on t-tests applied
287 within each monthly matrix.

$$288 \quad r = \frac{\sum ((x_i - \bar{x})(y_i - \bar{y}))}{\sqrt{\sum (x_i - \bar{x})^2 \sum (y_i - \bar{y})^2}} \dots \dots \dots (11)$$



289 where r is the correlation coefficient, x_i and y_i are paired values, $\bar{x}$ and $\bar{y}$ are the mean values of
290 paired drought indices, respectively.

291 **3. Results**

292 **3.1 Temporal Dynamics of Drought Indices (2000–2025)**

293 **3.1.2 TCI**

294 TCI time series showed broadly similar seasonal patterns to VCI, with July and August recording
295 the highest values (65–80) and April the lowest (10–25), as shown in Figure 2a. However, TCI
296 exhibited lower interannual variability than VCI in most months, suggesting that land surface
297 temperatures in the Tekeze Basin are less year-to-year variable than vegetation greenness, likely
298 because LST is smoothed by thermal inertia and is not solely controlled by rainfall. A notable feature
299 is the visible upward tendency in July and August TCI from approximately 2018 onward, consistent
300 with the spatially dominant positive trend (82.2% positive coverage) identified in the Mann-Kendall
301 analysis.

302 **3.1.1 VCI**

303 Basin-averaged VCI time series (2000–2025) revealed pronounced seasonal stratification across the
304 six growing-season months (Figure 2b). July and August consistently recorded the highest VCI
305 values (60–85), corresponding to peak Kiremt monsoon activity and maximum greenness, while
306 April and May exhibited the lowest values (10–20), frequently falling below the drought threshold
307 of 40 and indicating moderate to severe vegetation stress during the early growing season.
308 Interannual variability was highest in June (the Kiremt onset month), reflecting the sensitivity of
309 vegetation greenness to the timing and intensity of seasonal rains. Exceptional drought years with
310 basin-wide VCI minima below 15 across all months include 2002–2003, 2008–2009, and 2015–
311 2016, corresponding to documented El Niño events and positive Indian Ocean Dipole (IOD) phases.
312 VCI values showed a gradual increasing tendency in July and August from approximately 2015
313 onward, consistent with the predominantly positive long-term trend found in the spatial analysis.

314 **3.1.3 VHI**

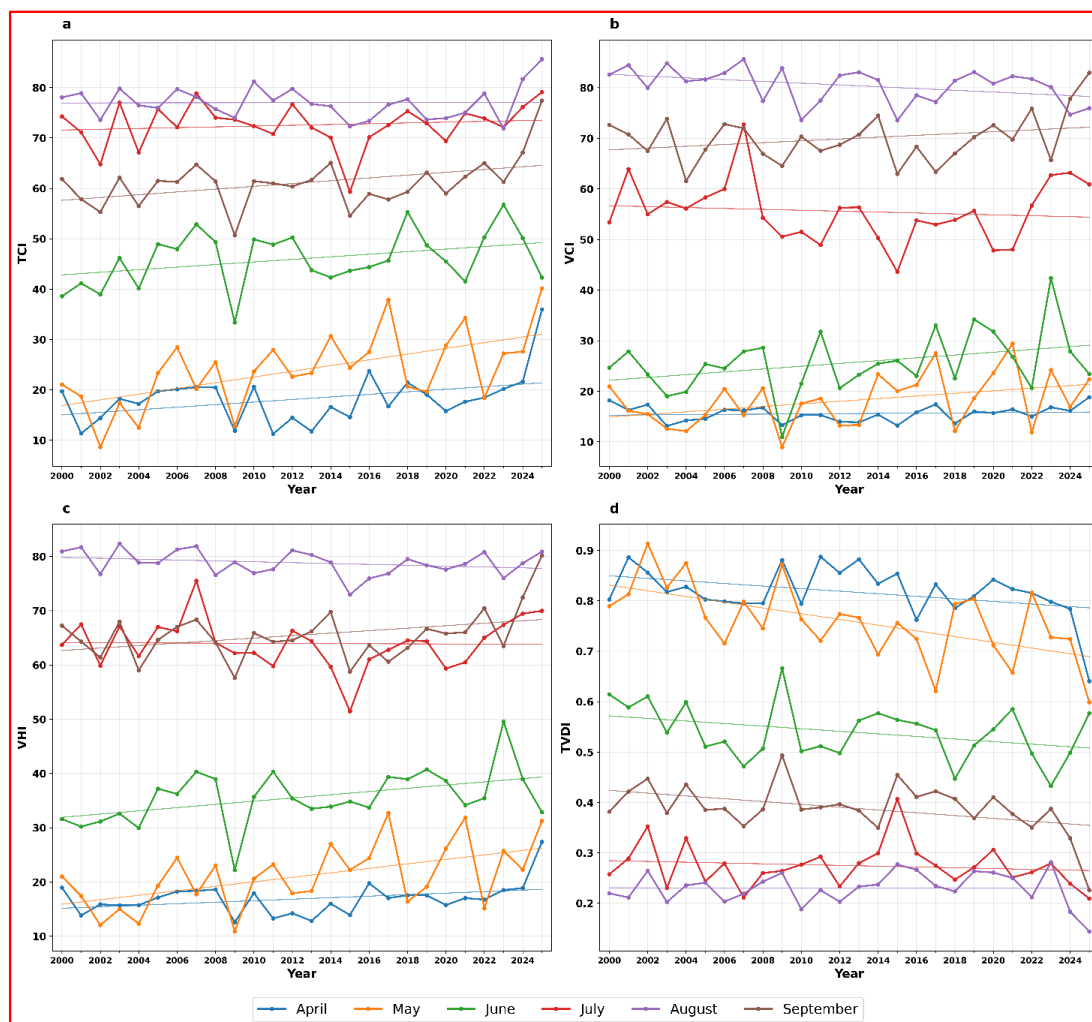
315 VHI, as the composite of VCI and TCI, closely mirrored VCI temporal patterns across all months
316 (Figure 2c). April values were consistently below 25 across almost all years, representing the most
317 severe composite drought stress period of the growing season. Major drought years (2002–2003,
318 2009, 2015–2016) were clearly identifiable as multi-month episodes with VHI below 30 across
319 April–June. The August time series showed a modest positive trend since approximately 2012,
320 consistent with the watershed rehabilitation and land restoration programs implemented in Tigray
321 during the 2000s and 2010s.

322 **3.1.4 TVDI**

323 TVDI exhibited an inverse seasonal pattern relative to the vegetation indices (Figure 2d). April and
324 May recorded the highest TVDI values (0.75–0.85), reflecting maximum land surface dryness at the
325 onset of the growing season before Kiremt rains have recharged soil moisture. TVDI decreased
326 progressively from June onward, reaching minima of 0.18–0.28 in July and August at peak canopy



cover and soil moisture. The year 2009 stood out as an extreme year with elevated TVDI across all months. There was no clear basin-mean directional trend in TVDI, although the pixel-wise spatial analysis revealed a dominant basin-wide drying trend (Section 3.2).



330

331 **Figure 2.** Annual time series (2000–2025) of basin-averaged monthly mean values for four drought indices
332 in the Tekeze Basin, Ethiopia, across the growing season months April (pink), May (olive), June (green), July
333 (teal), August (blue), and September (magenta). (a) VCI; 0–100 scale; threshold of 40 denotes drought onset,
334 (b) TCI; 0–100 scale, (c) VHI; composite of VCI and TCI; 0–100 scale, and (d) TVDI; 0–1 scale; higher
335 values indicate greater land surface dryness.

336 3.1.5 SMCI

337 SMCI displayed the most extreme seasonal amplitude of all seven indices, with August values
338 consistently reaching 80–95 (near-optimal soil moisture) and April values falling to 5–25 (severe
339 soil moisture drought), reflecting the strongly unimodal rainfall regime of the Tekeze Basin (Figure



340 3a). Interannual variability was highest in June and September, the Kiremt onset and cessation
341 months, when shifts in the rainy season timing translate directly into large differences in soil water
342 availability. The April SMCI time series showed a modest but visible declining tendency post-2015,
343 consistent with negative SMCI trends detected in the spatial analysis and potentially indicating a
344 progressive deepening of early growing-season moisture deficits.

345 **3.1.6 PCI**

346 PCI exhibited the highest interannual variability of all indices, particularly in April and May where
347 the coefficient of variation exceeded 45% in most years (Figure 3b). July and August recorded the
348 highest PCI values in wet years (70–100), while April PCI frequently approached 0 across dry years,
349 indicating near-complete absence of rainfall relative to the long-term maximum. Major drought
350 years were clearly identifiable as multi-month PCI minima: 2002–2003, 2008–2009, and 2015–
351 2016.

352 **3.1.7 NDWI**

353 NDWI values remained consistently low throughout the study period (–0.05 to +0.16), reflecting the
354 dominantly semi-arid land surface character of the basin (Figure 3c). A visible positive trend
355 emerged across all months from approximately 2010 onward, with the most pronounced increases
356 in August and September post-2020. This positive NDWI trajectory is the strongest long-term
357 directional signal visible in the basin-mean time series and aligns closely with the spatial trend
358 analysis showing 86.4% positive NDWI coverage across the basin.

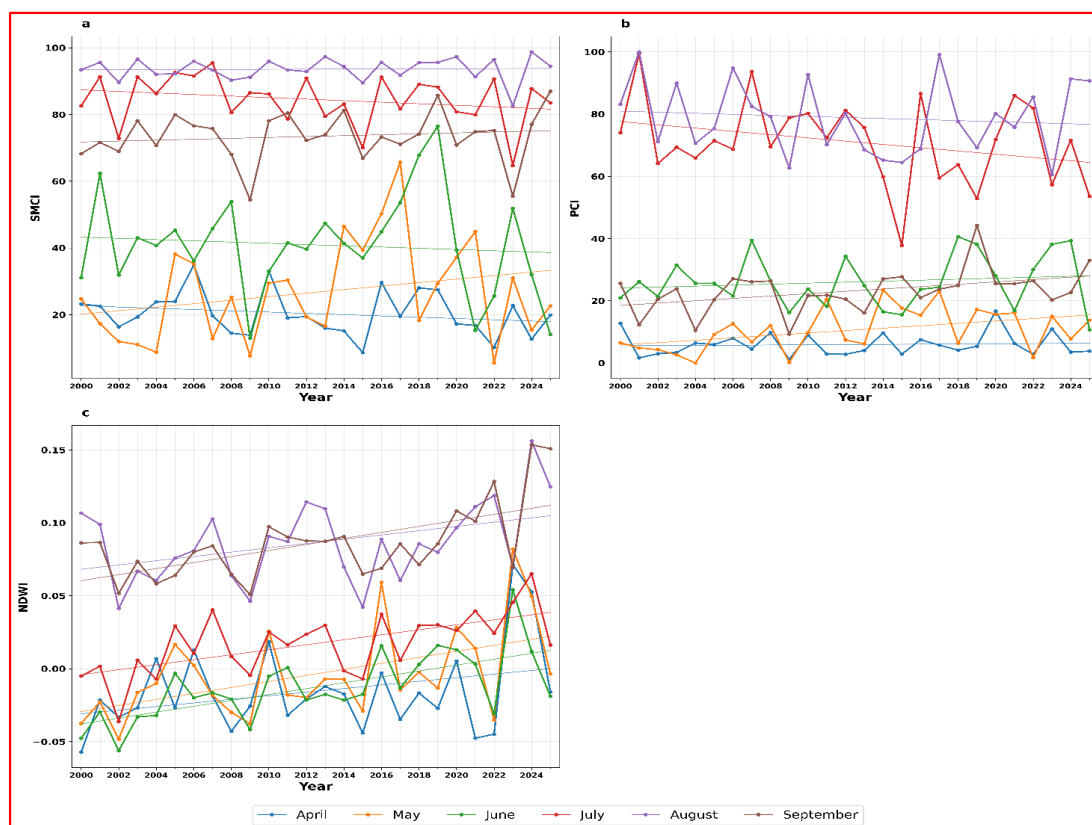


Figure 3. Annual time series (2000–2025) of basin-averaged monthly mean values for three additional drought indices in the Tekeze Basin, Ethiopia, across growing season months (color coding as in Figure 2). (a) SMCI; 0–100 scale, (b) PCI; 0–100 scale, (c) NDWI; –1 to +1 scale.

3.2 Spatiotemporal Drought Trend Maps

3.2.1 VCI, TCI, and VHI Trends

The spatial distribution of MK Tau (Figure 4a) and Sen's slope (Figure 4b) for VCI revealed predominantly positive trends across 71.5% of the basin area. Positive MK Tau values (indicating improving vegetation condition) were most pronounced in the central and southeastern portions of the basin at mid-elevations, with Tau values reaching up to +0.636. Negative trends (Tau as low as –0.557) were concentrated in the northern sub-basin, particularly in the Tigray highlands near the Eritrean border and in isolated pockets in the western lowland margins. Sen's slope for VCI ranged from –0.006 to +0.008 VCI units per year. TCI showed even more dominant positive trends, with 82.2% positive coverage and Tau ranging from –0.447 to +0.573 (Figure 4c), while Sen's slope (Figure 4d) spanned –0.003 to +0.004 TCI units per year. VHI spatial trends closely mirrored VCI and TCI combined patterns, with 74.8% positive coverage and MK Tau (Figure 4e) ranging from –0.470 to +0.605, and Sen's slope (Figure 4f) from –0.004 to +0.006 VHI units per year. The spatial coherence between VCI, TCI, and VHI trend patterns indicates expected internal consistency of the multi-index dataset.

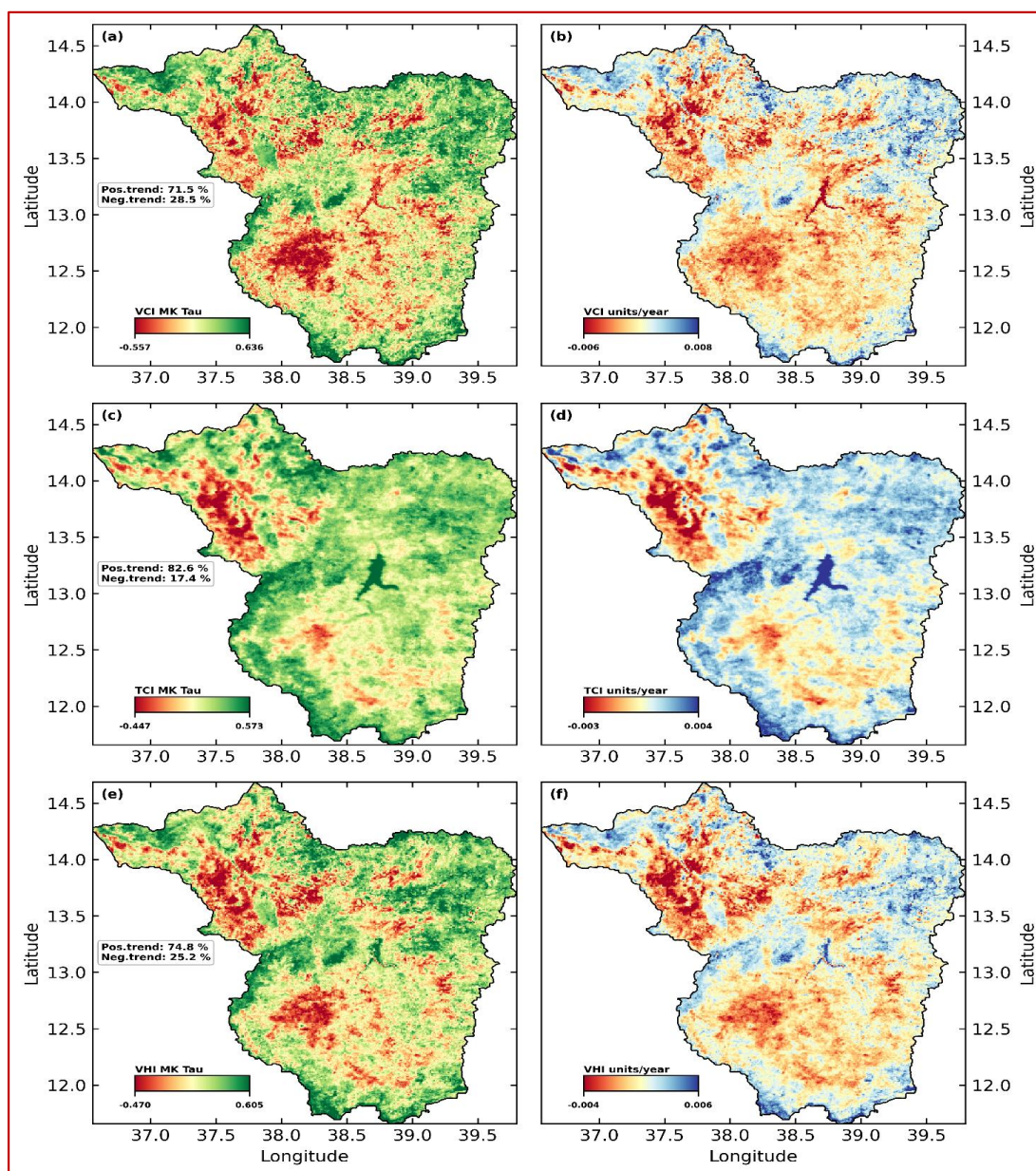


Figure 4. Spatial distribution of Mann-Kendall Tau (τ ; left column) and Sen's slope in units per year (right column) for three vegetation-based drought indices in the Tekeze Basin, computed from 2000–2025 monthly growing-season time series. (a) VCI: MK Tau, (b) VCI: Sen's slope, (c) TCI: MK Tau, (d) TCI: Sen's slope, (e) VHI: MK Tau, and (f) VHI: Sen's slope. Green tones in MK Tau maps indicate improving (less drought) trends; red tones indicate deteriorating (more drought) trends. Blue tones in Sen's slope maps indicate positive (improving) change; red-orange tones indicate negative (worsening) change. Black outline denotes the Tekeze Basin boundary.



386 3.2.2 PCI Trends

387 PCI exhibited the most spatially heterogeneous trend pattern of all indices, with 54.7% negative
388 trend coverage predominantly in the northern and northeastern sub-basins. MK Tau (Figure 5g)
389 ranged from -0.304 to $+0.360$, and Sen's slope (Figure 5h) from -0.019 to $+0.022$ PCI units per
390 year. The positive PCI Sen's slope reached the highest magnitude of any index ($+0.022$), reflecting
391 large absolute interannual variability in precipitation.

392 3.2.3 TVDI Trends

393 TVDI exhibited a strikingly different trend pattern from the vegetation indices: 82.2% of the basin
394 area showed negative MK Tau values, indicating increasing land surface dryness (higher TVDI)
395 over the study period. This represents the strongest and most spatially consistent negative trend
396 signal detected across all seven indices. MK Tau (Figure 5i) ranged from -0.613 to $+0.375$, with the
397 most negative values concentrated in the basin interior and the Afar lowlands. Positive TVDI trends
398 (wetting tendencies) were confined to limited areas in the northwestern highlands. Sen's slope
399 (Figure 5j) ranged from -0.005 to $+0.002$ TVDI units per year. This finding that land surface dryness
400 is increasing while vegetation condition is improving reveals an important decoupling between
401 above-ground greenness and surface energy balance dryness that has significant implications for
402 drought assessment (see Section 5.1).

403 3.2.4 SMCI Trends

404 SMCI showed 55.3% negative trend coverage, indicating that soil moisture conditions have
405 deteriorated across more than half the basin. The strongest negative SMCI trends (MK Tau as low
406 as -0.929 , the most negative of all indices in this study) were found in the central basin and the
407 eastern lowlands (Figure 5k). The extremely wide Sen's slope (Figure 5l) range (-0.145 to $+0.135$
408 SMCI units per year) reflects high spatial variability in soil moisture trend trajectories, driven by
409 topographic and land-cover gradients. Positive SMCI trends were observed in portions of the
410 northwestern highlands and along the Tekeze main river corridor, potentially reflecting reservoir
411 effects and rehabilitation-induced surface hydrological changes.

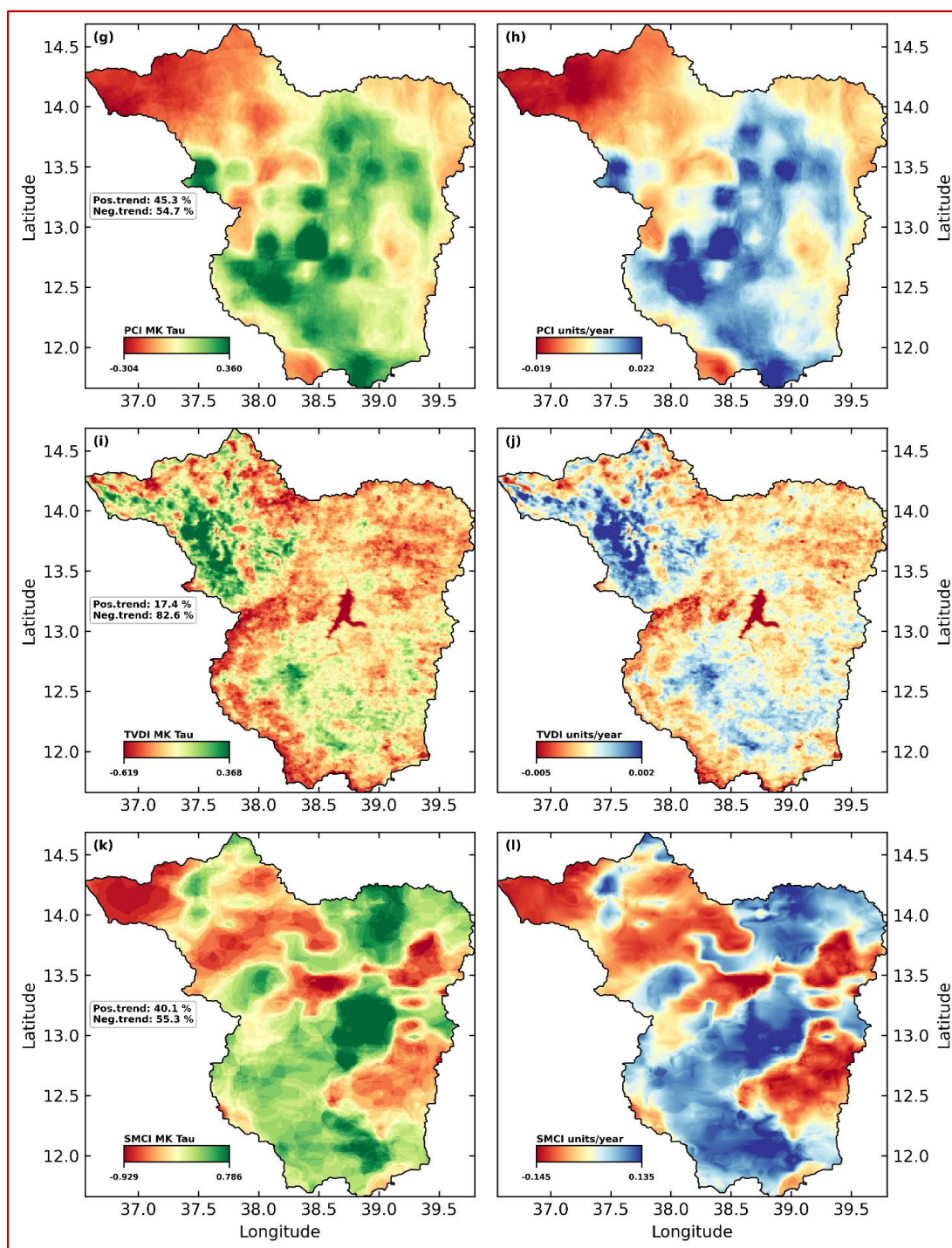




Figure 5. Spatial distribution of Mann-Kendall Tau (τ ; left column) and Sen's slope in units per year (right column) for three additional drought indices in the Tekeze Basin. (g) PCI: MK Tau, (h) PCI: Sen's slope, (i) TVDI: MK Tau, (j) TVDI: Sen's slope, (k) SMCI: MK Tau, and (l) SMCI: Sen's slope. Note: Green tones in MK Tau maps indicate improving (less drought) trends; red tones indicate deteriorating (more drought) trends. Blue tones in Sen's slope maps indicate positive (improving) change; red-orange tones indicate negative (worsening) change. Black outline denotes the Tekeze Basin boundary.

3.2.5 NDWI Trends

NDWI showed the highest proportion of positive trends of all seven indices, with 86.4% positive MK Tau coverage across the basin. MK Tau (Figure 6m) ranged from -0.455 to $+0.723$, and Sen's slope (Figure 6n) from -0.005 to $+0.003$ NDWI units per year. The small negative Sen's slope pockets were confined to the northern sub-basin, consistent with negative VCI and PCI trends in the same area.

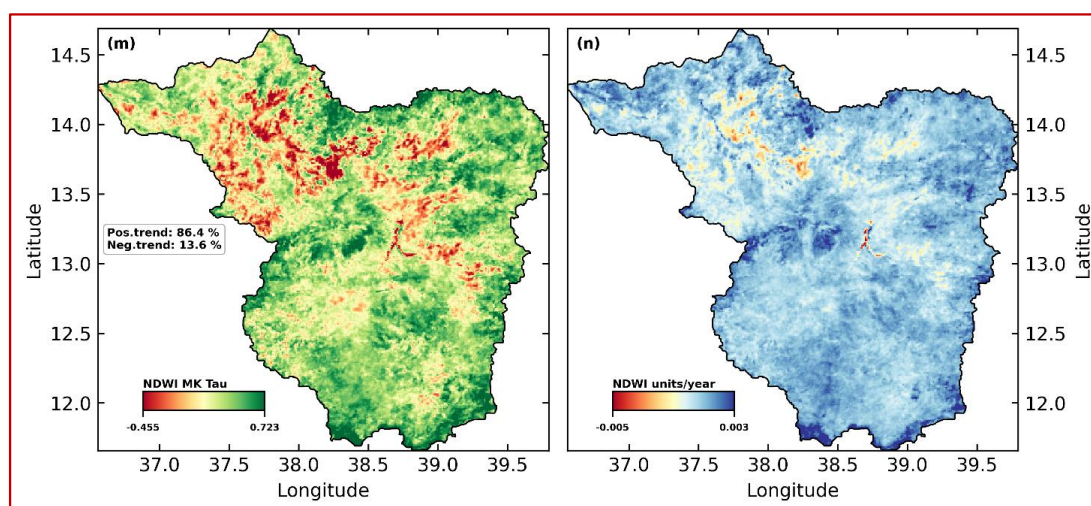


Figure 6. Spatial distribution of Mann-Kendall Tau (τ ; left column) and Sen's slope in units per year (right column) for four additional drought indices in the Tekeze Basin. (m) NDWI: MK Tau, (n) NDWI: Sen's slope. Note: Green tones in MK Tau maps indicate improving (less drought) trends; red tones indicate deteriorating (more drought) trends. Blue tones in Sen's slope maps indicate positive (improving) change; red-orange tones indicate negative (worsening) change. Black outline denotes the Tekeze Basin boundary.

3.3 Monthly Inter-Index Correlation Analysis

3.3.1 Overall Correlation Structure

Monthly Pearson correlation analysis (Figure 7) revealed consistent and physically interpretable inter-index relationships across all six growing-season months. The most robust positive relationship was between TCI and VCI ($r = 0.48$ to 0.97 across months), confirming the tight mechanistic coupling between thermal stress reduction and vegetation greenness improvement, both driven primarily by soil moisture availability. The strongest negative relationships were between TVDI and VHI ($r = -0.54$ to -0.97), and between TVDI and VCI ($r = -0.48$ to -0.92), which is an inverse relationship between land surface dryness and vegetation health.



440 3.3.2 April Correlations

441 In April, the earliest and most drought-stressed month, TCI–VCI correlation was moderate ($r = 0.48$)
442 and TVDI and VHI were strongly negative ($r = -0.97$). The extreme TVDI and VHI coupling in
443 April reflects the dominance of thermal-moisture stress in controlling composite vegetation health
444 during the late dry season, before the Belg rains have significantly recharged soil water. PCI showed
445 the weakest correlations with other indices in April ($r = 0.19$ – 0.35), consistent with the short lag
446 between rainfall onset and vegetation or soil moisture response.

447 3.3.3 May and June Correlations

448 From May to June, as the Belg rains progress and soils begin to wet up, PCI and SMCI correlations
449 strengthened substantially (from $r = 0.19$ in April to $r = 0.72$ in June), indicating that precipitation is
450 increasingly translating into measurable soil moisture recharge by June. The VCI and PCI correlation
451 reached $r = 0.82$ in June. TVDI and VHI and TVDI and VCI relationships remained strongly negative
452 through June ($r = -0.86$ and -0.84 in June, respectively), and TCI and VCI strengthened to $r = 0.86$.

453 3.3.4 July and August Correlations

454 July and August yielded the strongest overall inter-index correlations, consistent with peak Kiremt
455 monsoon activity and maximum seasonal moisture differentiation between wet and dry years. TCI
456 and VCI correlation in July reached $r = 0.55$ – 0.82 , and VCI and VHI reached $r = 0.92$. The SMCI–
457 PCI correlation was 0.73 in July and 0.57 in August, reflecting the time-integrated effect of
458 cumulative rainfall on soil water stores. TVDI and VHI weakened slightly to $r = -0.63$ in August
459 relative to its April peak of -0.97 , suggesting that the relative influence of thermal stress on
460 composite vegetation health diminishes when water availability is at its seasonal maximum.

461 3.3.5 September Correlations

462 In September, the Kiremt cessation month, VCI and VHI correlations remained very strong ($r =$
463 0.96), and TCI and VCI was 0.84 , indicating continued tight biophysical coherence during the late
464 rainy season. PCI–SMCI correlation was 0.54 , and NDWI and SMCI was 0.42 , consistent with soil
465 water stores still being replenished by residual rains and reflected in both canopy water content
466 (NDWI) and root-zone moisture (SMCI). PCI and TVDI showed its strongest negative correlation
467 in September ($r = -0.54$), confirming that precipitation is a primary control on land surface dryness
468 during the late growing season.

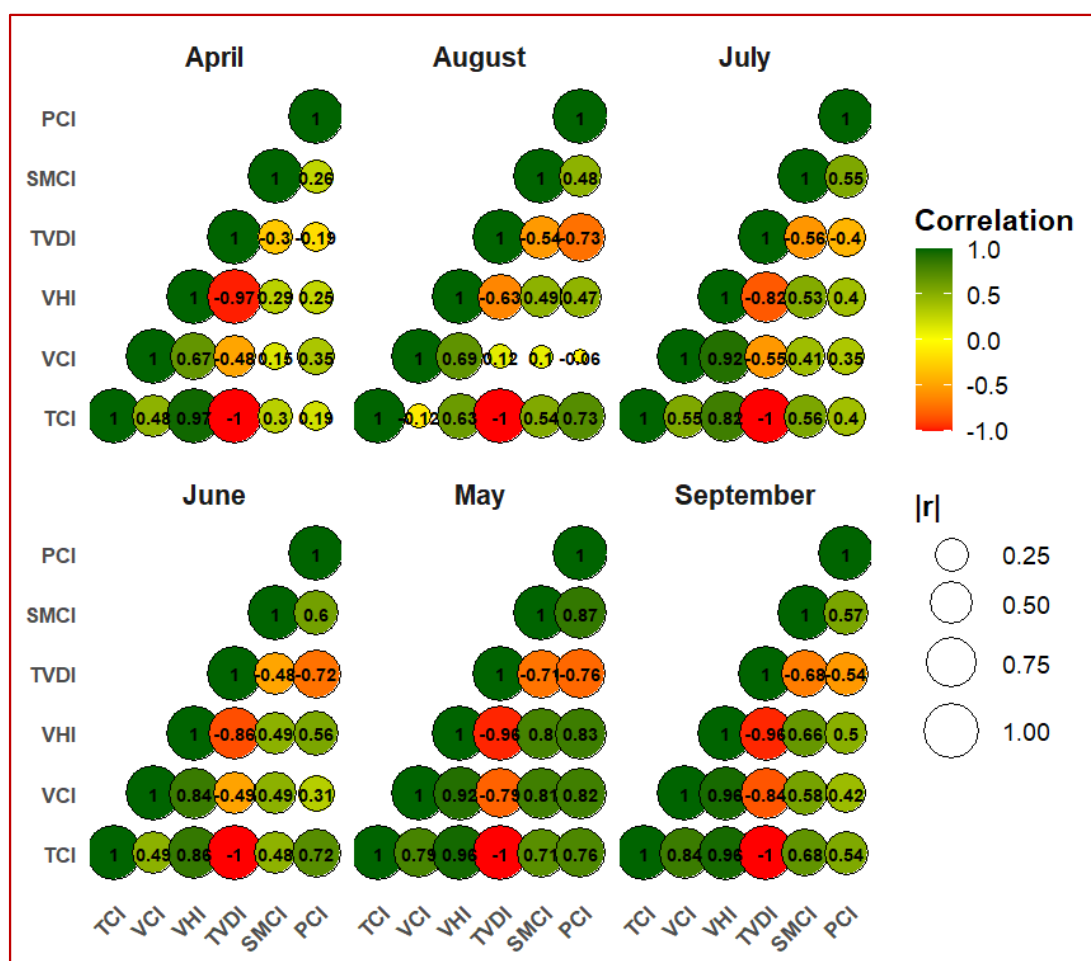


Figure 7. Monthly Pearson correlation matrices (corrplots) for all pairwise combinations of seven drought indices (VCI, TCI, VHI, TVDI, SMCI, PCI) in the Tekeze Basin, computed from 26-year (2000–2025) basin-averaged monthly time series. Results are shown for six growing-season months: April, May, June, July, August, and September. Circle color indicates the direction and magnitude of correlation: dark green = strong positive correlation (r approaching +1.0); red/orange = strong negative correlation (r approaching -1.0); intermediate colors indicate weaker correlations. Circle size is proportional to the absolute correlation coefficient $|r|$. All correlations with $p < 0.05$ (two-tailed t-test) are shown; non-significant correlations would appear smaller.

4. Discussion

5.1 Complementarity of Multi-Index Drought Assessment

A fundamental finding of this study is that the seven drought indices capture distinct and only partially overlapping dimensions of drought in the Tekeze Basin, validating the necessity of multi-index frameworks argued by AghaKouchak et al. (2015), Wilhite et al. (2014), and Yao et al. (2020).



483 The most striking divergence is between positive trends in vegetation-based indices (VCI, TCI, VHI,
484 NDWI) and negative trends in soil moisture and land surface dryness indices (TVDI, SMCI). This
485 apparent paradox vegetation improving while soil moisture and surface dryness worsen has been
486 increasingly documented in dryland systems globally and has been attributed to: (1) CO₂
487 fertilization, which enables vegetation to maintain canopy greenness at reduced soil moisture
488 through improved intrinsic water use efficiency (Liu et al., 2020); (2) species-level compositional
489 shifts toward drought-tolerant vegetation types that maintain greenness under drier conditions; (3)
490 the disproportionate statistical influence of a few extreme wet years on VCI/NDWI long-term means,
491 while median soil moisture conditions deteriorate; and (4) the selective response of above-ground
492 greenness to watershed rehabilitation programs, which may increase leaf area index without
493 proportionally restoring soil water-holding capacity or surface energy balance (Descheemaeker et
494 al., 2022; Brandt et al., 2017). These mechanisms are not mutually exclusive and are likely acting in
495 concert across the Tekeze Basin. The lesson for drought monitoring practitioners is clear: an
496 assessment based solely on VCI or VHI would generate an optimistic but potentially misleading
497 characterization of drought trends in this basin, while exclusive reliance on TVDI or SMCI would
498 overstate deterioration. Only a multi-index framework reveals the full complexity.

499 **4.2 Spatial Heterogeneity and Sub-Basin Vulnerability**

500 The spatial trend maps expose marked within-basin heterogeneity that would be entirely invisible in
501 station-based or catchment-averaged drought assessments. The northern Tekeze sub-basin,
502 particularly the Tigray highlands and the Eritrean border zone, consistently emerged as the most
503 drought-vulnerable area, exhibiting negative trends simultaneously in VCI, VHI, and PCI. This
504 spatial convergence of degrading trends across multiple indices strongly suggests a persistent and
505 worsening drought hotspot in the northern basin, consistent with documented declines in Kiremt
506 season rainfall in northern Ethiopia reported by Seleshi and Zanke (2004), Conway et al. (2009),
507 Gebremeskel et al. (2019), and Haile et al. (2020), and likely exacerbated by the 2020–2022 Tigray
508 conflict, which disrupted land management and rehabilitation activities. In contrast, the central basin
509 showed predominantly positive multi-index trends, potentially reflecting the positive hydrological
510 and vegetation effects of the Tekeze Hydropower Reservoir (post-2009), ongoing watershed
511 enclosure programs, and elevation-dependent rainfall redistribution.

512 **4.3 Seasonal Dynamics and Growing Season Implications**

513 The temporal analysis consistently identified April and May as the most drought-stressed months
514 across all indices, with VCI, TCI, VHI, SMCI, PCI, and NDWI at or near their seasonal minima and
515 TVDI at its seasonal maximum during this pre-monsoon period. This finding has direct implications
516 for agricultural planning: smallholder farmers in the Tekeze Basin who plant rain-fed crops in March
517 to April are exposed to the highest drought risk at the planting and early establishment stage, before
518 Kiremt rains materialize. Integrating PCI and SMCI into drought early warning during March–May
519 when these indices can detect emerging moisture deficits weeks before vegetation stress becomes
520 visible in NDVI-based products would provide a critical 2–4 week lead time advantage for timely
521 agricultural interventions and livelihood protection decisions (Rhee et al., 2010; AghaKouchak et
522 al., 2015).

523 **4.4 Inter-Index Relationships and Physical Mechanisms**



524 The strong and consistent TCI and VCI positive correlations (up to $r = 0.97$ in April) confirm the
525 tight thermal-vegetation coupling in this semi-arid system: as soil moisture increases with advancing
526 rains, both canopy transpiration (cooling LST, raising TCI) and vegetation greenness (raising NDVI,
527 raising VCI) respond simultaneously (Sandholt et al., 2002; Wang et al., 2004). The strong TVDI
528 and VHI negative relationship (up to $r = -0.97$ in April and May) demonstrates that TVDI is a
529 reliable inverse proxy for composite vegetation health in this basin. The progressive strengthening
530 of PCI–SMCI correlations from April ($r \approx 0.19$) to June ($r \approx 0.72$) reveals the seasonal buildup of
531 the precipitation–soil moisture linkage as cumulative Kiremt rains progressively fill soil water
532 stores. The relatively weaker PCI and TVDI negative correlations compared to SMCI and TVDI
533 across all months confirm that TVDI responds primarily to the integrated soil moisture state rather
534 than instantaneous precipitation events, a distinction important for designing multi-index early
535 warning systems (Du et al., 2013).

536 4.5 Methodological Contributions

537 This study makes a methodological contribution by demonstrating the effective integration of
538 Google Earth Engine Python Colab for large-scale pixel-wise trend computations, R software for
539 rigorous statistical analysis and visualization, and QGIS for publication-quality cartographic outputs
540 a combined workflow that is reproducible, open-source, and transferable to other data-scarce basins
541 across Africa. The GEE Python Colab environment proved particularly powerful for pixel-wise MK
542 and Sen's slope computation across 82,000 km² at 250 m resolution without requiring local
543 computational infrastructure, enabling a level of spatial detail and temporal scope not achievable
544 with traditional local-computing workflows. This framework is fully extensible to additional drought
545 indices, additional basins, and future climate projections.

546 4.6 Limitations and Future Research Directions

547 Several limitations of this study should be acknowledged. First, SMAP soil moisture data are
548 available only from April 2015 at standard quality, requiring gap-filling or surrogate estimation for
549 2000–2015 SMCI computation, which may introduce uncertainty into early-period trend estimates.
550 Second, MODIS 16-day composites cannot fully capture sub-monthly drought onset dynamics,
551 potentially missing the full intensity of short-duration drought events. Third, the pixel-wise MK
552 trend analysis does not correct for spatial autocorrelation between adjacent pixels, which may inflate
553 the effective degrees of freedom and affect the interpretation of statistical significance thresholds.
554 Fourth, the analysis does not include a formal quantitative attribution of detected trends to climate
555 variability versus land use change versus land rehabilitation interventions, a direction that warrants
556 future investigation using multivariate attribution frameworks. Future studies should: (1) integrate
557 ground-based soil moisture, streamflow, and agricultural yield data for multi-source validation of
558 the remote sensing indices; (2) investigate the role of large-scale climate drivers (ENSO, IOD,
559 AMO) in modulating interannual drought variability across the basin; (3) extend the multi-index
560 framework to include projected future drought changes using CMIP6 regional climate model
561 outputs; and (4) develop a composite drought vulnerability index for the Tekeze Basin integrating
562 the seven indices and socioeconomic exposure data for targeted policy applications.

563

564



565 5. Conclusions

566 This study delivered the first comprehensive multi-index, pixel-scale spatiotemporal drought
567 analysis in the Tekeze Basin, Ethiopia, integrating seven satellite-derived drought indices (VCI, TCI,
568 VHI, TVDI, SMCI, PCI, NDWI) computed from MODIS, SMAP, and CHIRPS datasets across the
569 full growing season (April to September) for the period 2000–2025. Index computation and trend
570 analysis were performed in Google Earth Engine Python Colab; statistical charts and correlation
571 matrices were produced in R; and spatial visualization was conducted in QGIS. The principal
572 conclusions are:

573 **(1) Strong seasonal drought gradients:** April and May consistently exhibited the most severe
574 drought stress across all indices (VCI $\approx$ 10–20; SMCI $\approx$ 5–25; TVDI $\approx$ 0.75–0.85), while July and
575 August were the most favorable months, highlighting the pre-monsoon period as the highest-risk
576 window for agricultural drought impacts.

577 **(2) Divergence between vegetation and moisture trends:** Vegetation-based indices showed
578 predominantly improving trends (VCI: 71.5% positive; TCI: 82.2%; VHI: 74.8%; NDWI: 86.4%),
579 while soil moisture and dryness indices showed dominant negative trends (TVDI: 82.2% negative;
580 SMCI: 55.3%; PCI: 54.7%), revealing a decoupling between above-ground greenness and below-
581 ground moisture status that can only be detected through multi-index analysis.

582 **(3) Persistent northern sub-basin drought hotspot:** The northern Tekeze sub-basin (Tigray
583 highlands) showed negative trends simultaneously in VCI, VHI, and PCI, the most severe and multi-
584 dimensionally consistent drought deterioration in the basin, warranting priority attention in drought
585 early warning and resilience investment planning.

586 **(4) Strong biophysical inter-index coupling:** TCI and VCI correlations (r up to 0.97) and TVDI
587 and VHI correlations (r down to -0.97) confirmed tight thermal-vegetation linkages across all
588 months, with inter-index correlations strengthening progressively from April to August as the
589 Kiremt monsoon advances. PCI and SMCI linkages strengthened from April ($r \approx 0.19$) to June ($r \approx$
590 0.72), reflecting the seasonal soil moisture recharge dynamics.

591 **(5) Value of multi-index integration:** No single drought index adequately captured all dimensions
592 of drought in the Tekeze Basin. The seven-index framework revealed complementary drought
593 signals undetectable by any individual index and provides a robust foundation for multi-dimensional
594 drought early warning system design in Ethiopia and comparable dryland basins across East Africa.

595 The open-source, cloud-computing-based multi-index drought monitoring framework developed in
596 this study is transferable to other data-scarce semi-arid basins in sub-Saharan Africa and provides a
597 timely methodological reference for drought researchers and operational agencies engaged in
598 climate-smart natural hazard management.

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604 Data Availability Statement

605 All input datasets used in this study are publicly accessible: MODIS MOD13Q1 and MOD11A2
606 products are available via NASA Earthdata (<https://earthdata.nasa.gov>). NASA-USDA SMAP soil
607 moisture data are available at NASA Earthdata and the GEE data catalog. CHIRPS v2.0 precipitation
608 data are available from the Climate Hazards Center, University of California, Santa Barbara
609 (<https://www.chc.ucsb.edu/data/chirps>). Analysis scripts (GEE Python Colab notebooks and R code)
610 are available from the corresponding author upon reasonable request.

611 Author Contributions

612 (Bedasa Asefa) Conceptualization, methodology, visualization (QGIS), writing original draft,
613 writing review, and editing. (Jianjun Zhao) Conceptualization, methodology, writing, review, and
614 editing, supervision. (*Hongyan Zhang*) Data acquisition and validation, MODIS and CHIRPS
615 dataset processing, quality assurance/quality control, literature review, writing, review, and editing.
616 (*Zhengxiang Zhang*) Study area characterization, interpretation of spatiotemporal patterns, writing,
617 review, and editing. (*Xiaoyi Guo*) Funding acquisition, project administration, institutional support,
618 supervision, writing, review, and editing. (Mebrahtom Zerom) formal analysis, data curation,
619 Google Earth Engine Python scripting, Mann-Kendall trend analysis. (Yosef Ermias) R statistical
620 analysis, correlation matrix computation, temporal graphing, and interpretation of results. (*Andeptop*
621 *Sule*) hydrological and climatic context, field knowledge of the Tekeze Basin,

622 Competing Interests

623 The authors declare that they have no conflict of interest.

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