



1 **Understanding Spatial and Temporal Evolution of Progressive**
2 **Excavation Damage Zone in Granitic Rocks – Development of a Novel**
3 **Mine-by and Laboratory Experimental Framework at the BedrettoLab**

4 List of Authors: Pooya Hamdi¹, Jonas Dickmann¹, Michal Kruszewski¹, Antonio Pio Rinaldi², Linus Villiger², Alexis
5 Shakas³, Mathew Perras⁴, Navid Bahrani⁵, Patrick Bianchi², Peter Achtziger-Zupančič⁶, Steven Gaines⁷, Kai Bröker⁸,
6 Xiaodong Ma⁹, Paul A Selvadurai³, Virginie Durand¹⁰, Kavan Khaledi⁶, Mohammadreza Jalali⁷, Yousef Golabchi¹¹, Sam
7 Harris⁴, Alba Zappone³, Mathilde Wimez³, Marian Hertrich³, Stefan Wiemer², Florian Amann^{1,6}

8 Corresponding author: Pooya Hamdi, email: hamdi@lih.rwth-aachen.de

9 ¹ Chair of Engineering Geology & Hydrogeology, RWTH Aachen University, Aachen, Germany

10 ² Swiss Seismological Service, ETH Zurich, Zurich, Switzerland

11 ³ Department of Earth Sciences, ETH Zurich, Sonneggstrasse 5, Zurich, 8092, Switzerland

12 ⁴ Trent School of the Environment, Trent University, Peterborough, Ontario, Canada

13 ⁵ Department of Civil and Resource Engineering, Dalhousie University, Canada

14 ⁶ Fraunhofer Research Institution for Energy Infrastructures and Geotechnologies, IEG, Aachen, Germany

15 ⁷ Natural Resources Canada, Ontario, Canada

16 ⁸ Center for Hydrogeology and Geothermics, University of Neuchâtel, Neuchâtel, Switzerland.

17 ⁹ University of Science and Technology of China, School of Earth and Space Sciences, Hefei, China

18 ¹⁰ Geoazur Laboratory, Valbonne, France

19 ¹¹ Department of Civil Engineering, Lassonde School of Engineering, York University, Canada

20

21 **Abstract**

22 The brittle damage around tunnels in the form of spalling is common in massive crystalline rocks
23 subjected to high in-situ stresses. Understanding the short-term development and long-term
24 evolution of the Excavation Damage Zone (EDZ) is particularly important for deep geological
25 repositories (DGRs) for heat-generating nuclear waste, as the EDZ may provide preferential
26 pathways for radionuclide migration. This paper presents an overview of the PRECODE mine-by
27 experiment at BedrettoLab, a newly developed underground research laboratory (URL) in
28 southern Switzerland designed to investigate brittle fracturing and EDZ evolution in crystalline
29 rocks under repository-relevant conditions. The PRECODE tunnel was excavated in Rotondo
30 Granite using both controlled drill-and-blast (D&B) and line-drilling and rock-breaking (LDRB)
31 methods to evaluate the influence of excavation technique on EDZ characteristics. The
32 experiment combines multidisciplinary in-situ monitoring including acoustic emission,
33 distributed acoustic and strain sensing, hydraulic monitoring, electrical resistivity tomography,
34 and terrestrial laser scanning with laboratory-based fracture mechanics and time-dependent
35 testing under varying environmental conditions. The main objectives are to characterize short-
36 term EDZ formation, investigate long-term crack growth and stress corrosion processes, quantify
37 permeability evolution, and assess the interaction between excavation-induced damage and
38 nearby fault zones. Preliminary results indicate clear differences in AE activity and EDZ
39 development between excavation methods, while observations from the mature EDZ of the more
40 than 40-year-old Bedretto Tunnel suggest ongoing time-dependent brittle fracturing. The coupled
41 in-situ and laboratory datasets provide a unique basis for improving predictive models and long-
42 term safety assessments for nuclear waste repositories in crystalline rocks.

43 **1. Introduction**

44 The brittle damage in the form of spalling and rock bursting is common in massive rocks under
45 high stress conditions (Martin et al., 1997; Diederichs, 2007; Kaiser et al., 2010; Perras &
46 Diederichs, 2016). Spalling is the development of visible extension fractures under a compressive
47 stress regime that may or may not be violent (Diederichs, 2007). Strain bursting, in contrast,
48 refers to the violent rupture of a rock volume under high stress levels, during which rock fragments



49 of varying sizes are ejected from the tunnel wall at high velocity, releasing stored energy (Kaiser
50 et al., 1996). Understanding the shape, extent and properties of the Excavation Damage Zone
51 (EDZ), and its evolution over time is of foremost relevance for safety analysis of underground
52 infrastructures, particularly deep geological repositories for nuclear waste disposals. Within the
53 EDZ, permeability is often increased (Kelsall et al., 1984; Rutqvist et al., 2009) and may provide
54 preferred and interconnected flow pathways for the migration of radionuclides. In contrast to
55 clayshales and salt rocks, crystalline rocks do not possess natural self-sealing or self-healing
56 properties (Tsang et al., 2005; Bossart et al., 2017), and damage in the form of brittle fracturing
57 tends to persist and may propagate over time under unfavorable conditions.

58 In-situ observations around tunnels in massive, brittle rocks revealed that the stress-induced
59 fracturing begins at stress levels well below the rock's laboratory unconfined compressive
60 strength (Martin & Chandler, 1994; Martin, 1997; Diederichs, 2003; Diederichs et al., 2010). The
61 stress level at which brittle fracturing is initiated around tunnels in the short-term (i.e. during the
62 construction) has been found to be 35-50% of the rock strength determined at the laboratory
63 scale (Martin et al., 1997). This stress level coincides with the crack initiation (CI) damage
64 threshold that can be determined from laboratory samples using non-destructive methods such
65 as strain measurement or acoustic emission monitoring (Eberhardt et al., 1999; Diederichs, 2003;
66 Amann et al., 2011; Perras et al., 2024). In the near-field of a tunnel, where the confining stress is
67 low, macroscopic and interconnected fractures may form, as the differential stress exceeds CI.
68 The macroscopic fracture zone is limited in depth by the spalling limit that represents the
69 boundary between extensional type failure and shear-dominated failure, and is defined by the
70 ratio between the maximum and minimum principal stresses. A variety of empirical and fracture
71 mechanics studies (i.e. Diederichs, 2007) on crystalline rocks showed that the spalling limit
72 ranges between $\sigma_1/\sigma_3=1/20-1/15$.

73 For confining stresses beyond the spalling limit, the rock fails in shear. While this rock mechanical
74 concept for the short-term formation of an EDZ around tunnels in crystalline rocks is generally
75 accepted and applied in the nuclear waste industry (Martin & Christiansson, 2009; Siren et al.,
76 2015), a variety of different concepts and methodologies exist for the assessment of the long-
77 term EDZ development. Since the seminal work of Bieniawski (1967), it is commonly assumed
78 that the long-term strength is associated with a stress level that marks the transition between
79 stable and unstable crack growth (crack damage threshold, CD). At this damage threshold,
80 cracks that were initiated individually begin to interact with each other and lead to rapid failure
81 under continued loading (Diederichs, 2007). This stress level is typically found at around 80-90%
82 of the short-term strength. A vast amount of static fatigue tests at different environmental
83 conditions and stress levels ranging from the short-term strength to stress levels below the crack
84 damage threshold has been performed (mainly in the radioactive waste community; e.g. Peng,
85 1971; Schmidtke & Lajtai, 1985) and demonstrated, that time-dependent failure may also occur
86 at stress levels below the CD threshold. These test results (i.e. time to failure at various stress
87 levels) have been extrapolated to long service times of tunnels and are often used to predict the
88 long-term development of the excavation damage and disturbance zone around repository
89 tunnels (Damjanac & Fairhurst, 2010; Paraskevopoulou et al., 2018).

90 Considering the lifespan of deep geological repositories (on the order of more than 100,000
91 years), understanding the long-term strength of the host rock, in addition to its short-term
92 strength, and the factors controlling its time-dependent behavior is crucial for assessing the
93 integrity of underground infrastructure used to host nuclear waste from the construction stage
94 through to the operational phase. Time-dependent and progressive failure (Terzaghi, 1962)
95 involves gradual internal structural changes within a rock mass resulting from incremental, slow



96 damage accumulation and localized weakening (e.g., rock-bridge failure), which can eventually
97 lead to large-scale failure. The mechanisms and rate of progressive rock failure are influenced by
98 a combination of external forces (e.g., temperature fluctuations, water infiltration, seismic
99 activity, loading and unloading, etc.) and internal properties (e.g., rock composition, structural
100 discontinuities, porosity, rock bridge characteristics, etc.). Processes of long-term crack
101 development in crystalline rocks are often associated with stress corrosion and sub-critical crack
102 growth in chemically reactive environments (Nara & Kaneko, 2006; Nara et al., 2010). This is
103 specifically relevant for radioactive waste disposal with elevated rock temperatures and, for
104 example, high calcium ion concentrations due to cementitious materials (Nara et al., 2010). The
105 development of subcritical cracks stands as a primary contributor to time-dependent fracturing
106 in rock which impacts the long-term strength of the rock. The rate of crack growth depends on
107 various parameters including the type of applied load (tensile or compressive), magnitude of
108 stress relative to the short-term strength, as well as the environmental factors such as
109 temperature, relative humidity and water chemistry.

110 Achieving a comprehensive understanding of brittle fracturing in crystalline rocks necessitates
111 evaluating the spatial and temporal development of damage across various scales, from the
112 micro-scale to the macro-scale (Figure 1). This paper presents an overview of the development of
113 the PRECODE mine-by tunnel and laboratory experimental framework at the BedrettoLab, a new
114 underground research laboratory (URL) in Switzerland.

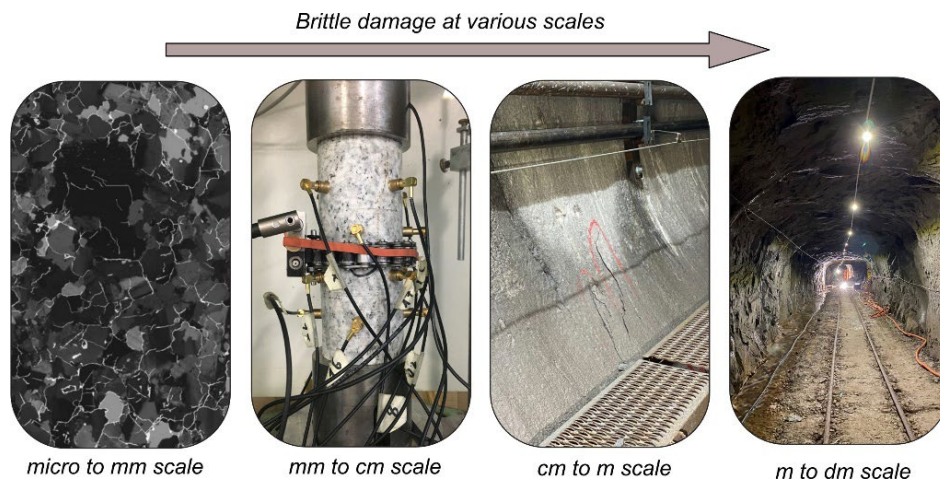


Figure 1: Illustration of the brittle damage in crystalline rocks at various scales

2. State-of-the-art across Scales

2.1 Laboratory Scale

119 Early studies have shown that brittle rock failure, particularly under low confining pressure, is
120 controlled by extensional fractures rather than shear fractures (Cook, 1965; Hoek & Bieniawski,
121 1965; Bieniawski, 1967a, 1967b; Tapponnier & Brace, 1976; Horii & Nemat-Nasser, 1985; Martin,
122 1997). At the laboratory scale, the brittle failure of intact rock subjected to uniaxial or triaxial
123 compression typically comprises four distinct stages: crack closure, elastic region, stable crack
124 growth, and unstable crack growth (Martin & Chandler, 1994; Reches & Lockner, 1994; Eberhardt
125 et al., 1999). Non-destructive monitoring methods, including axial and lateral strain
126 measurements and acoustic emission (AE) monitoring, have been used to identify these stages



127 and characterize the progressive damage evolution in brittle rocks (Martin, 1993; Eberhardt,
128 1998; Eberhardt et al., 1999; M. S. Diederichs et al., 2004; Cai et al., 2004).

129 The crack closure represents the first stage of the brittle fracturing process and is characterized
130 by progressive closure of pre-existing microcracks, resulting in a nonlinear trend in the axial-
131 lateral strain relationship before the onset of linear elastic deformation (Eberhardt, 1998).
132 Following crack closure, the rock undergoes linear elastic deformation. The elastic properties
133 such as Young's modulus and Poisson's ratio can be derived. Crack initiation is the start of micro-
134 fracturing. During compression test, this threshold can be identified when the lateral or
135 volumetric strain responses depart from linearity or as the point at which a systematic increase
136 in acoustic emission activity is first observed (Eberhardt, 1998). Cracks at this stage initiate in
137 isolation and stop forming when loading stops or decreases (Eberhardt, 1998). Experimental
138 results and in-situ observations have shown that crack initiation typically occurs when the
139 applied stress level reaches approximately 30-50% of the peak uniaxial compressive strength
140 (UCS) (Scholz, 1968; Lockner et al., 1991; Martin & Chandler, 1994; Martin, 1997; Eberhardt,
141 1998). With continued loading, crack growth remains stable until the start of unstable crack
142 growth (CD threshold) (Reches & Lockner, 1994). During compression test, the CD threshold
143 marks when the volumetric strain reversal occurs (Amann et al., 2011). Using AE measurements,
144 this threshold is determined by a marked increase in acoustic emission activity (Eberhardt,
145 1998). Unstable crack growth proceeds until microcracks coalesce into larger fractures, leading
146 to peak strength and ultimate failure of the rock samples.

147 Recent advances in laboratory-scale rock mechanics have substantially improved the
148 characterization of brittle fracturing processes in crystalline rocks under stress conditions
149 relevant to deep geological repositories for nuclear waste. More recently, the integration of fiber-
150 optic sensing techniques, including Distributed Strain Sensing (DSS), has enabled continuous,
151 spatially resolved strain monitoring along the outer periphery of laboratory specimens, providing
152 further insight into the initiation and propagation of brittle damage zones at the specimen scale
153 (Hegger et al., 2022; Salazar et al., 2022; Salazar Vásquez et al., 2024; Bianchi et al., 2024; Chen
154 et al., 2025). The monitoring of surface strain fields using digital image correlation (DIC) has also
155 provided a more robust understanding of failure in brittle rock subjected to uniaxial stressing (e.g.
156 Dutler et al., 2018). Significant progress has also been made in understanding time-dependent
157 brittle fracturing and long-term strength degradation through relaxation and static fatigue
158 (Paraskevopoulou et al., 2016, 2018), subcritical crack growth and stress corrosion experiments
159 (Nara, Takada, et al., 2010; Hamdi et al., 2026). Laboratory studies using double-torsion and
160 fracture toughness testing under controlled temperature, relative humidity, and pore-fluid
161 chemistry have shown that environmental conditions strongly affect crack growth rates and time-
162 to-failure in crystalline rocks (Brantut et al., 2013; Nara et al., 2013). Adsorption-induced
163 softening of rock caused by humidity and/or water saturation have been shown to significantly
164 affect rock performance within environments expected in many underground excavation sites
165 (Wu et al., 2023). Finally, advances in microstructural and multiscale analysis have strengthened
166 the mechanistic understanding of brittle fracture processes and their upscaling to repository-
167 scale behavior. High-resolution microscopy, micro-CT imaging, and micro-indentation testing
168 have revealed the grain-scale controls on crack initiation, the influence of mineralogical
169 heterogeneity, and the role of pre-existing microstructures such as grain boundaries and healed
170 fractures (Nasser et al., 2005; Renard et al., 2019; M. Li et al., 2020; Winhausen et al., 2021)

171 **2.2 Underground Research Laboratories**

172 The definition of the EDZ to describe the extent and characteristics of the damage zones around
173 tunnels has evolved over time (Tsang et al., 2005; Siren et al., 2015; Perras & Diederichs, 2016).



In an attempt to standardize the EDZ definitions, Tsang et al. (2005) provided a comprehensive description of various damage zones around tunnels applicable for nuclear waste repositories. The characteristics of the damage surrounding tunnels depend on various parameters, such as tunnel shape, rock strength, initial stress conditions, excavation method, and the geometrical and mechanical properties of pre-existing discontinuities (Figure 2). The extent of the construction-induced damage (CDZ) can vary and depends on the excavation method used, i.e., drill & blast (D&B) or mechanical excavation, and can be minimized or potentially eliminated (Read, 2004; Martino et al., 2007; Jonsson et al., 2009). The highly damaged zone (HDZ), unlike the CDZ, is a direct result of stress redistribution due to the opening an excavation, and its extent is controlled by the geometry of the tunnel, structural features, the magnitude and orientation of stress. Visible stress-induced fractures develop in the HDZ, forming spalled blocks, outbreaks and interconnected macro cracks. Spalling to form a v-shaped notch is a typical example of the brittle damage within the HDZ (Read, 2004) due to the maximum tangential stress being concentrated at two opposing locations around the excavation. Beyond that, the inner EDZ (EDZ_i), characterized by connected damage, gradually transitions to the outer EDZ (EDZ_o), with partially connected to isolated damage (Bossart et al., 2002). The damage in this zone is often characterized by the development of microcracks and can be detected, for instance, through borehole/core p-wave velocity measurements (Falls & Young, 1998). Beyond the inner and outer EDZ, there will be an excavation influence zone (EIZ) that is characterized by elastic and reversible changes of strain (Siren et al., 2015).

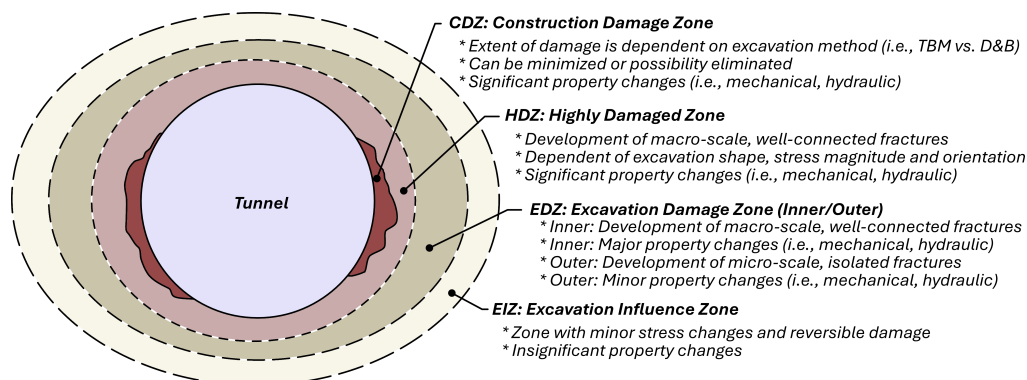


Figure 2: Definition of various damage zones around tunnels and their associated characteristics (modified after Siren et al., 2015; Perras & Diederichs, 2016)

In the context of nuclear waste disposal studies, several in-situ underground research laboratories (URLs) have been developed worldwide with the primary objective of investigating rock behavior at the repository scale, with an emphasis on the safety and feasibility of long-term storage. The focus on the surrounding rocks as a geological barrier is to understand the long-term thermo-hydro-mechanical (THM) behavior of the rock mass as a result of interaction with the nuclear waste. Some of the main URLs in crystalline rocks include: Äspö Hard Rock Laboratory (HRL) in Sweden, Onkalo Posiva Oy facility in Finland, AECL's Underground Research Laboratory (URL), Grimsel Test Site in Switzerland, Mizunami URL (MIU) in Japan, Bukov Underground Research Facility in Czech Republic and Beishan underground URL in China. Many of these underground research laboratories have been constructed in a variety of rock types, at different depths, and using diverse excavation geometries and methods under varying environmental conditions, enabling systematic evaluation of their influence on EDZ characteristics. The following paragraphs summarize some of the key findings from these URLs.



210 The Mine-by Experiment in Canada at AECL's URL aimed to study the evolution of the progressive
211 failure and development of the EDZ around a mechanically-excavated circular tunnel that was
212 heavily instrumented prior to the excavation. Through analysis of several thousand microseismic
213 and acoustic emission (MS/AE) events around Mine-by Experiment, Collins (1997) found that the
214 majority of the events are concentrated in the roof and floor of the excavation where the
215 differential stress is the highest. The brittle failures in the form of V-shape notches were
216 developed, primarily in the roof and floor of the tunnel. The stress level at which the brittle
217 fracturing was initiated around the tunnel was found to be 50-60% of the rock strength
218 determined at the laboratory scale (Martin et al., 1997). The influence of the excavation shape and
219 method on the extent of the EDZ was assessed within an Excavation Stability Study in the
220 Canadian Mine-by Tunnel experiment. Using AE monitoring, Falls and Young (1998) compared the
221 extent of the EDZ developed around two tunnels with similar geometry and stress levels but with
222 different excavation methods, i.e. drill and blast (D&B) and tunnel boring machine (TBM). They
223 found that the extent of the EDZ around the TBM tunnel was substantially less compared to the
224 D&B tunnel, with the rate of AE counts being about ten times higher in the D&B tunnel compared
225 to the TBM tunnel. Similar results were also obtained within ZEDEX project at Äspö HRL, Sweden
226 and Onkalo site, Finland (Siren et al., 2015). Within the ZEDEX, the AE monitoring showed a
227 distinct difference in observed EDZ extent caused by excavation methods i.e., D&B vs. TBM, with
228 the depth of the EDZ for the D&B section estimated to be up to 1 meter from the tunnel wall, while
229 for the TBM section was limited to a few centimeters.

230 The hydraulic tests conducted prior and post excavation within ZEDEX project indicated that the
231 preferential pathways for radionuclide transport were limited to the EDZ zone and not enhanced
232 permeability was observed beyond that. The ZEDEX results indicated that the transition from the
233 EDZ, which has irreversible changes, to the EIZ is gradual, with no distinct boundary that can be
234 marked between the two zones through measurement methods (Bäckblom & Martin, 1999;
235 Olsson et al., 2004). Results from a series of borehole hydraulic tests conducted as part of the
236 EDZ experiment at the Grimsel Test Site indicated the presence of a zone approximately 1 m thick
237 around the tunnel with permeability higher than that of the intact rock matrix, reflecting the
238 development of an EDZ (Frieg et al., 2012).

239 Upon nuclear waste emplacement, a heat transfer will occur from the canisters across bentonite
240 buffer material towards the rock, with the canister surface temperature reaching up to 150 °C and
241 a temperature of ~ 90 °C at the bentonite-rock interface (Bossart et al., 2017), depending on the
242 specific design and layout of the repository. The Heated Failure Tests in the Canadian Mine-by
243 tunnel studied the effect of thermally-induced stress on initiation of new cracks and propagation
244 of the fractures. The Mine-by Heated Experiment showed the effect of thermally-induced stress
245 on initiation of new cracks and propagation of the fractures that are in the critical-equilibrium
246 state. Falls and Young (1998) reported a substantial increase in the number of AE counts around
247 the region that was heated from ambient temperature to 85 °C. The in-situ monitoring of the AE
248 development at the ASPE Experiment within the HRL in Sweden also indicated the development
249 of thermally-induced fractures and growth of the EDZ caused by an increase in temperature
250 (Andersson & Martin, 2009; Andersson et al., 2009).

251 **2.3 Numerical Modelling of Spalling in Brittle Rocks**

252 Over the past few decades, various numerical modeling approaches have been developed to
253 simulate brittle failure and the characteristics of EDZ around underground openings. These
254 include continuum methods (Martin et al., 1999; Hajiabdolmajid et al., 2002; Diederichs, 2007;
255 Sanipour et al., 2022; Rógenes et al., 2024), discontinuum methods (Potyondy & Cundall, 2004;



256 Shin et al., 2007; Rasmussen & Min, 2024; Potyondy & Fu, 2026), and hybrid methods (Lisjak et
257 al., 2015; Hamdi et al., 2017; Vazaios et al., 2019; Potyondy et al., 2020).

258 Initial attempts to model the EDZ around Mine-by URL showed that the conventional constitutive
259 laws with simultaneous mobilization of cohesion and friction angle failed to reproduce the depth
260 and shape of EDZ (Martin et al., 1999; Hajiabdolmajid et al., 2002). Using Hoek-Brown constitutive
261 model, Martin et al. (1999) showed the depth of failure around the URL test tunnel can only be
262 simulated using elastic models when the Hoek-Brown parameters m and s are modified ($m = 0$, s
263 $= 0.11$). Since that time, several other continuum methods have been proposed, with two
264 common approaches being the Cohesion Weakening Frictional Strengthening (CWFS)
265 (Hajiabdolmajid et al., 2002) and the Damage Initiation and Spalling Limit (DISL) (Diederichs,
266 2007). While, these methods can be able to simulate the depth and shape of failure around
267 underground openings (Perras & Diederichs, 2016), they often fail to accurately predict the extent
268 of the EDZ. This limitation arises from the continuum assumption, which prevents the explicit
269 representation of fracture initiation and propagation, resulting in failure near excavation
270 boundaries being predominantly represented as shear deformation. More recently, incorporation
271 of the grain-like discontinuum concept (i.e., Voronoi Tessellated Model, VTM) into continuum
272 models have shown promising potential to simulate brittle fracturing by incorporating
273 heterogeneities with continuum methods (Sanipour, 2021; Rógenes et al., 2024).

274 Discontinuum methods (i.e., discrete element method, bonded-block models or bonded-particle
275 models) simulate the rock domain as an assembly of blocks/particles that are joined together at
276 their contact interfaces Potyondy & Cundall, 2004). These models explicitly account for material
277 discontinuities and the heterogeneous grain-scale structure of rock and can simulate the
278 progressive brittle failure through breakage of bondage. Since the seminal work of Potyondy &
279 Cundall (2004) to model the observed failure around Mine-by URL in Canada using PFC software,
280 numerous studies have investigated the capability of discontinuum methods to replicate brittle
281 fracture processes. Using the Grain-Based Model (GBM) in UDEC, Shin et al. (2007) successfully
282 reproduced both the depth and shape of spalling around the Mine-by tunnel and demonstrated
283 the importance of grain geometry and heterogeneity in simulating spalling damage. More
284 recently, Potyondy & Fu (2026), reviewed the historical development of discontinuum bonded-
285 particle models and discussed their application to the simulation of spalling failure around
286 underground excavations.

287 Hybrid methods, such as the Finite-Discrete Element Method (FDEM), combine the advantages
288 of both continuum and discontinuum approaches by simulating the transition from a continuous
289 to a discontinuous state through the explicit modelling of fracture initiation, propagation, and
290 coalescence (Munjiza, 2004). In recent years, several modelling benchmarks have focused on
291 simulating the EDZ around the Mine-by URL tunnel in Canada. Using FDEM approach, Hamdi et
292 al. (2017) investigated the impact of micro-heterogeneity on the in-situ strength reduction of the
293 massive Lac du Bonnet granite and attempted to reproduce the shape and depth of spalling
294 around the Canadian Mine-by Tunnel. Using Irazu software based on FDEM method, Vazaios et
295 al. (2019) simulated the V-shaped notch formation around the URL test tunnel. According to their
296 results, a reduced uniaxial compressive strength (to 120 MPa) is needed to properly capture the
297 depth of EDZ observed around the tunnel. More recently, Potyondy et al. (2020) used coupled
298 PFC3D/FLAC3D with various grain shapes (spherical, Voronoi, and tetrahedral) to assess the
299 capability of the combined continuum-discontinuum model to simulate the spalling process at
300 the URL test tunnel. Their results demonstrated that representing the rock mass using tetrahedral
301 grain blocks provided the best agreement with the in-situ spalling depth.



3. Open Science Questions

Despite many advances in understanding the brittle fracturing in terms of its temporal and spatial evolution and across multiple scales, there still remain a number of open, fundamental questions. One key open question concerns the estimation of the long-term strength of brittle rocks at laboratory and repository scale. It is commonly assumed that the long-term strength is associated with a stress level that marks the transition between stable and unstable crack growth (CD threshold). A vast amount of static fatigue tests at different environmental conditions and stress levels ranging from the short-term strength to stress levels below the crack damage threshold has been performed (Peng & Johnson, 1972; Schmidtke & Lajtai, 1985; Paraskevopoulou et al., 2018; Wang et al., 2026) and demonstrated, that time-dependent failure may also occur at stress levels below the CD threshold. These test results (i.e. time to failure at various stress levels) have been extrapolated to long service times of tunnels and are often used to predict the long-term development of the excavation damage and disturbance zone around repository tunnels. The majority of previous attempts to assess long-term strength, such as time-to-failure analysis, have relied on laboratory tests that were subjected to stress levels above the CI threshold. A different concept is needed to address time-dependent fracture propagation processes at stress levels close to and below CI. As the confining stress increases, time-dependent fracture formation and propagation occur at stress levels above the CI threshold but significantly below the crack damage threshold. Thus, static fatigue tests with low differential loads, conducted at relevant confining stresses of 4-6 MPa (representative of a tunnel at 1000-1500 m depth), need to be performed to characterize time-dependent processes in this zone. Commonly, time-to-failure tests are conducted under unconfined conditions, and thus, the influence of confining stress is often not addressed. From a rock mechanical perspective, a systematic experimental analysis of this aspect is currently lacking.

It is well established that time-dependent subcritical crack growth and stress corrosion are primary mechanisms controlling the long-term strength of brittle crystalline rocks (Atkinson, 1984; Brantut et al., 2013; Nara, 2015). However, the observational window available for studying time-dependent crack growth, particularly at repository scale, is orders of magnitude shorter than the expected lifespan of deep geological repositories, which extends over hundreds of thousands of years. Most data collected from in-situ experiments at existing URLs span only approximately 5–20 years, limiting our ability to fully characterize the long-term evolution of the excavation damage zone (EDZ), particularly beyond the initial post-excavation phase (Read, 2004; Jonsson et al., 2009; Mellanen et al., 2010; Frieg et al., 2012). This gap is further intensified by the influence of environmental factors such as relative humidity, temperature, and pore-water chemistry, which are known to significantly affect subcritical crack growth rates and permeability evolution in crystalline rocks. Extending observational capabilities to capture EDZ characteristics over longer timescales under natural stress and environmental perturbations is therefore essential to improve our understanding of long-term rock mass behavior and to strengthen the safety assessment of deep geological nuclear waste repositories.

Understanding the impact of fault activation and earthquakes on the integrity of deep geological nuclear waste repositories is of critical importance for long-term safety assessment. Tectonic fault reactivation whether driven by natural stress evolution, glacial loading and unloading, or anthropogenic activities can lead to coseismic and aseismic slip, stress redistribution, and permeability enhancement in the surrounding rock mass, potentially compromising engineered and geological barriers. Earthquake-induced deformation may alter the hydraulic properties of faults and adjacent damage zones, creating preferential pathways for radionuclide migration and affecting repository isolation over geological timescales (Sibson, 1990; Faulkner et al., 2010;



349 Cappa & Rutqvist, 2011; Yang et al., 2021). Laboratory experiments, in-situ observations, and
350 numerical modeling studies have shown that even small stress perturbations can trigger slip on
351 critically stressed faults, emphasizing the sensitivity of fault systems in crystalline host rocks
352 (Scholz, 1968; Ortlepp, 2000; Y. Li et al., 2025). Consequently, international repository programs
353 increasingly emphasize the need to characterize fault mechanics, seismic hazard to evaluate
354 repository robustness under both seismic and aseismic fault activity scenarios (i.e., SKB Report,
355 2011).

356 Even the massive rock mass suitable as deep geological repository inherently may contain pre-
357 existing microfractures, healed joints, veins, and mineral fabric anisotropies that influence their
358 mechanical response. Experimental, field and numerical studies have shown that stress-induced
359 fracturing around underground openings are strongly interacting with pre-existing discontinuities,
360 which may act as stress concentrators and crack initiators and mechanical contrast with the host
361 rock (Kranz, 1983; M. Diederichs, 2003; Hamdi et al., 2015). Laboratory observations indicate that
362 microcrack initiation and propagation often preferentially occur along grain boundaries, tip of
363 microcracks, or healed fractures, leading to earlier onset of damage and altered failure paths
364 compared to truly flaw-free material (Eberhardt et al., 1999). At the tunnel scale, the interaction
365 between stress-induced fractures and natural fracture networks governs the characteristics of
366 EDZ (i.e., geometry, depth, and connectivity), with important implications for permeability
367 evolution and long-term repository integrity.

368 **4. Bedretto Tunnel**

369 **4.1 Bedretto Tunnel Description**

370 Located near the Gotthard Pass in the Swiss Central Alps, the Bedretto Tunnel is a NW-SE
371 sidetrack tunnel that connects the Furka Base Tunnel to the Bedretto Valley. The tunnel was
372 constructed 1970s and 1980s using conventional drill & blast method, with the main purpose of
373 transporting excavated material during construction of the Furka Base Tunnel. This horseshoe-
374 shaped tunnel is 5.2 km long with dimensions of 2.8x2.8 meters. The tunnel axis is oriented
375 approximately N317° W, with the southern portal (Bedretto Valley, TM0) at an elevation of about
376 1479.5 meters above sea level (a.s.l.), and the northern portal (Furka Tunnel, TM5220) at an
377 elevation of about 1505.2 meters a.s.l. The tunnel passes through mountainous terrain
378 characterized by highly variable overburden depths. Starting from the Bedretto Valley, the
379 overburden depth gradually increases, reaching a maximum depth of approximately 1.6 km under
380 Pizzo Rotondo at around TM3140. It then decreases to a depth of approximately 1.1 km at
381 TM4200, before increasing again to approximately 1.3 km at TM5200. Apart from some sporadic
382 steel-set lining support at challenging sections, the Bedretto Tunnel has been mostly left
383 unsupported since its first construction.

384 **4.2 Geological and Structural Setting**

385 The Bedretto Tunnel is situated within the Gotthard Massif in the Central Swiss Alps and crosses
386 three main geological units: 1) Tremola series, 2) Prato series and 3) Rotondo granite (Figure 3).
387 Starting from the South Portal (Tunnel Meter, TM0), the tunnel passes through metamorphic rocks
388 of the Helvetic domain. Between TM0 to TM 434, the tunnel is within Tremola series. Between
389 TM434 to TM 1138, the tunnel is located within Prato series (Val Nalps Gneiss). Beyond this point,
390 the tunnel is situated within the Rotondo Granite, extending to TM5220 (Keller & Schneider, 1982).
391 The Tremola series comprises a sequence of metasedimentary rocks, predominantly featuring
392 mica gneisses, mica schists, hornblende schists, and hornblende gneisses (Rast, 2020). The
393 Prato series is part of the Val Nalps Gneiss Complex and is characterized by banded hornblende
394 schists and amphibolites in the northern and southern part and mica gneisses and schists with



lenses of amphibolites in the central part (Hafner, 1958; Rast, 2020). The Rotondo granite that encompasses majority of the Bedretto Tunnel is characterized as an equigranular to porphyritic, medium to coarse grained granite with a mineral content of quartz (25 – 35%), alkali feldspar (20 – 40%), plagioclase (10 – 25%), biotite (3 – 8%) and garnet, phengite, chlorite, epidote, apatite and zircon as accessory minerals (Hafner, 1958; V. Lützenkirchen & Loew, 2011; Rast, 2020).

Geological structures including fracture sets and fault zones with various persistence and thickness ranging from sub-meters to tens of meters are frequently observed along the tunnel. Extensive surface and subsurface mapping of the geological structures has recently been conducted to analyze the geometrical and geomechanical characteristics of the structures in the study area (Lützenkirchen, 2002; Lützenkirchen & Loew, 2011; Jordan, 2019; Achtziger-Zupančič et al., 2024; Ceccato et al., 2024). The NE-SW and N-S structures, both steeply dipping, are the two dominant structures at the Bedretto Tunnel.

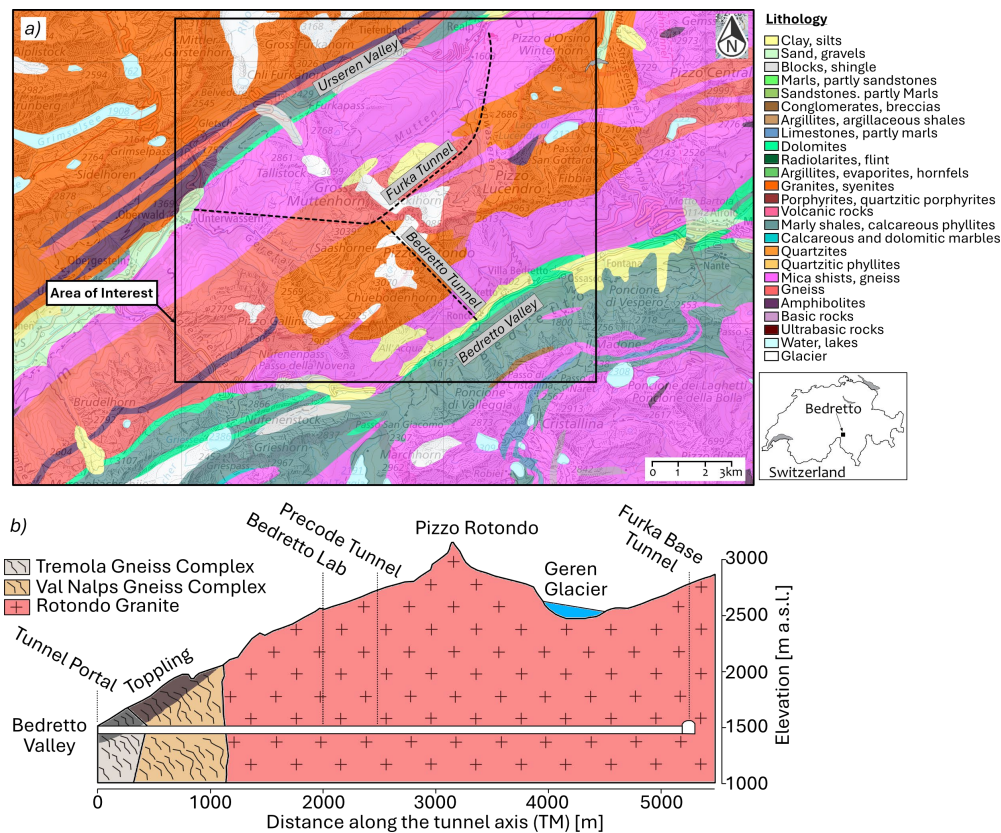


Figure 3: a) Lithology of the Bedretto Valley and surrounding area (source: SwissTopo), b) cross section along Bedretto Tunnel, with highlighted location of BedrettoLab and PRECODE Tunnel (modified after Keller & Schneider, 1982)

4.3 Hydro-Mechanical Laboratory Characterization

At a first approximation, the Rotondo granite within the Bedretto tunnel can be considered as homogeneous and isotropic medium, although a weak foliation due to ductile deformation during the Variscan orogeny is identifiable in various locations (Lützenkirchen & Loew, 2011). The petrophysical, mechanical and hydraulic properties of the Rotondo granite have been extensively



studied under framework of the ongoing projects at the BedrettoLab (David et al., 2020; Ma et al., 2022; Perras et al., 2024; Hamdi et al., 2026; Kruszewski et al., 2026). A comprehensive laboratory testing program has been conducted as part of the PRECODE project, in order to identify the primary brittle fracturing thresholds needed to predict the extent of the EDZ in the Rotondo granite around the mine-by tunnel. These include uniaxial and triaxial compression tests, indirect and direct tensile, permeability, fracture toughness and petrophysical tests. Table 1 provides an overview of the average petrophysical and mechanical properties of the Rotondo Granite derived from samples collected in the Bedretto Tunnel.

424

425

Table 1: Average petrophysical and mechanical properties of the Rotondo Granite obtained from the Bedretto Tunnel (modified after Ma et al., 2022; Perras et al., 2024)

427

	Parameter	Units	Dynamic	Static
Mechanical Properties	Young's Modulus	GPa	26.5	52.5
	Shear Modulus	GPa	10.8	-
	Poisson's Ratio	-	0.2	0.22
	Indirect Tensile Strength (BTS)	MPa	-	7.0
	Uniaxial Compressive Strength (UCS)	MPa	-	180.9
	Crack Initiation (CI)	MPa	-	64.5
	Crack Damage (CD)	MPa	-	117.0
Petrophysical Properties	P-wave Velocity	km/sec	3.4	-
	S-wave Velocity	km/sec	2.0	-
	Connected Porosity	(%)	-	1.36
	Dry Density	Kg/m ³	-	2606
Hydromechanical Properties	Permeability	m ²		4.4e ⁻¹⁸

428

4.4 Stress Measurement

The Bedretto valley has a complex tectonic history, with three distinct metamorphic phases (Labhart, 2005), Alpine uplift, erosion, glaciation/deglaciation, ductile and brittle deformation, resulting in a highly variable distribution of field stress tensor. Due to the very low level of seismic activity in the Ticino region, including the Bedretto area, the use of focal mechanism solutions to infer the in situ stress field is limited (Sue et al., 2007). In recent years, a series of in situ stress measurement campaigns, mainly based on the mini-frac method, have been conducted at BedrettoLab to quantify the magnitude and orientation of the in-situ stress field along the tunnel (for more detail see: Ma et al., 2019; Bröker & Ma, 2022; Bröker et al., 2024). Overall, the stress state in the vicinity of the Bedretto Lab (TM2000) is found to be transitional between normal and strike-slip faulting ($S_{hmin} < S_{Hmax} \leq S_v$), with the estimated average stress magnitudes to be $S_{hmin} = 14.6 \pm 1.4$ MPa, $S_{Hmax} = 25.4 \pm 2.3$ MPa and $S_v \approx 26.5$ MPa (Bröker & Ma, 2022) between TM1750 and TM2250 with approximate overburden depth of 1030 m.

441

4.5 Excavation Damage Zone of the Existing Bedretto Tunnel

Frequent occurrences of buckling in the gneiss and spalling in the granitic zones were documented during the construction of the Furka and Bedretto tunnels (Amberg, 1983). Since its completion in 1982, the tunnel has remained largely unlined, providing a unique opportunity to investigate the long-term evolution of spalling and the excavation damage zone in crystalline rock. Several spalling characterization campaigns using conventional and remote sensing methods revealed that the depth and orientation of the excavation damage zone (EDZ) vary along the tunnel in response to changes in topography and structures. The depth of spalling along the



tunnel varies with the overburden, reaching values of up to 1.1–1.5 m in the most affected sections (Alcaino-Olivares et al., 2024; Hamdi et al., 2024, 2025). Huber (2004) divided the failure observed at the Bedretto Tunnel into three main categories of: 1) spall shears, 2) kink bands and 3) tension cracks. The spall shears are defined as the “onion-skin” surface parallel fractures that are formed in the area of maximum stress concentration with their thickness varying from a few millimeters to centimeters. The kink bands are triangular shaped fractures that are generally formed at the bottom and top of the spalled notches. The tension cracks are formed due to volumetric dilation in the slabs subjected to bending.

Continuous laser scanning of the tunnel wall conducted in 2024 between TM2400 and TM5200 (where the overburden ranges from 1.2 to 1.6 km) revealed that the excavation damage zone (EDZ) is asymmetrical, with its shape, position, and depth varying significantly along the tunnel. Systematic brittle fracturing and spalling have been observed along the tunnel particularly in the areas with higher magnitude of overburden between TM2000–5200. In addition, a number of sections along the tunnel shows extensive overbreaks as a result of interaction between stress-induced fractures and pre-existing geological structures. The asymmetric shape of the EDZ indicates that the principal stresses are deviated from horizontal and vertical planes and are also not parallel to the tunnel axis, agreeing with current modelling results, in-situ stress measurements, and the regional stress field.

4.6 Environmental Condition at Bedretto

Keller & Schneider (1982) measured the temperature variation along the Bedretto Tunnel after completion of the tunnel construction. The rock temperature in the vicinity of the Bedretto Valley Portal (TM200) was 8°C, increasing to 18°C at TM2300 and remaining relatively unchanged until TM4200. The relative humidity at the Bedretto tunnel varies along the tunnel and may fluctuate seasonally, but the measurements have shown a high relative humidity, exceeding 80%. It is important to note that between 1984 and 2017, there was no door between the Furka Base Tunnel and the Bedretto Tunnel. As a result, the repeated passing of high-speed trains through the Furka Tunnel caused instantaneous cyclic changes in relative humidity between 80–100%, providing a more corrosive environment and allowing for a higher potential for subcritical crack growth during that time frame.

5. Development of a Novel Experimental Framework

5.1 PRECODE Project: Opportunities, Goals, and Objectives

The Bedretto Underground Laboratory (BedrettoLab) was established by ETH Zürich in 2018. During the initial phase, the in-situ experiments were conducted within the tunnel segment between TM2000 and TM2100, where the tunnel was widened to approximately 6 meters. The overburden in this segment ranges from roughly 1000 to 1030 meters. In the most recent phase, a new tunnel was excavated at TM2428, parallel to a fault zone on the southeast side of the existing Bedretto tunnel, with a length of approximately 100 m. This site currently hosts a number of large-scale in-situ experiments, with two of the relevant experiments being: (1) Fault Activation and Earthquake Rupture (FEAR), in which an isolated fault zone will be stimulated to seismogenic failure (Meier et al., 2026), and (2) Progressive Excavation Damage Zone Evolution During and Post Mine-By Tunneling (PRECODE), aimed at studying the development of the EDZ around crystalline rocks relevant for deep geological nuclear waste repositories. The primary objectives of the PRECODE project are to:

- Improve our understanding of the short-term rock mass behaviour due to tunnel advance with advanced monitoring systems that were not available for past granitic rock mine-by



- 494 experiments (fiber optic strain sensing, distributed acoustic/strain sensing, micro-seismic
495 sensors, LiDAR scanning, etc.),
- 496 • Improve our understanding of the influence of natural repository perturbations (humidity,
497 temperature, seismicity, etc.) on progressive microcrack development within the EDZ over
498 the long-term,
 - 499 • Determine changes in EDZ's permeability with time under open (worst case) adit conditions
 - 500 • Analyze the impact of tunneling on dislocations across fault zones and related changes in
501 hydraulic properties
 - 502 • Quantify the impact of activation of nearby faults on the EDZ characteristics and potential
503 permeability enhancement

504 The research addresses brittle fracturing across multiple spatial scales (micro-, laboratory-, and
505 in-situ scales) and temporal scales (short-, medium-, and long-term), using a range of
506 complementary approaches organized into three major work packages:

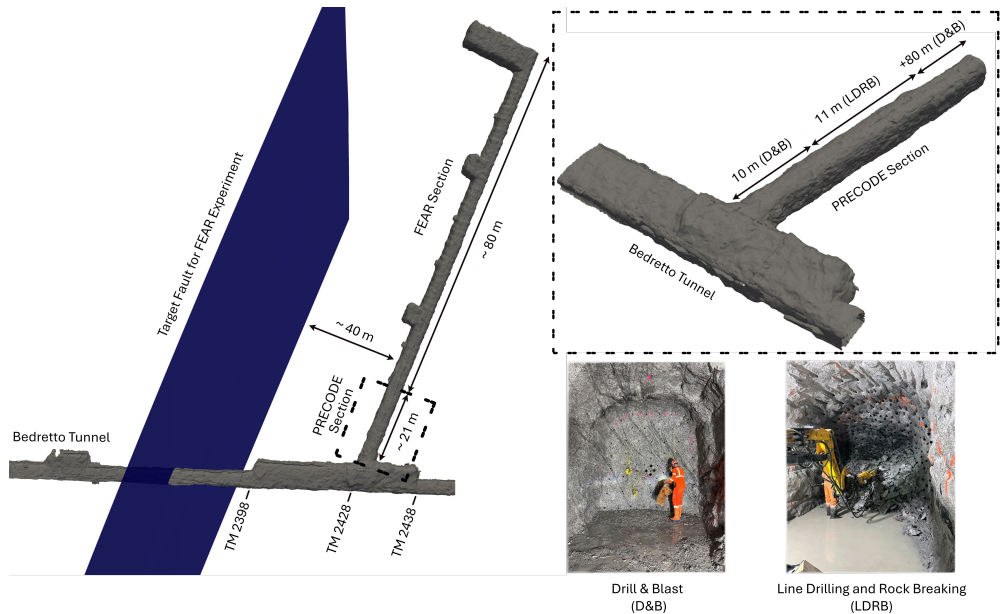
- 507 (1) **Multi-disciplinary in-situ monitoring:** with a focus on detecting and quantifying the
508 spatial and temporal development of the EDZ around the tunnel using various
509 approaches, including passive and active seismic monitoring, hydraulic monitoring,
510 geophysical monitoring, and deformation monitoring,
- 511 (2) **Comprehensive laboratory testing:** with a focus on deriving the short-term and long-
512 term strength of the Rotondo Granite under various stress states (e.g., uniaxial and triaxial
513 compression, indirect and direct tension, fatigue, relaxation) and environmental
514 conditions (temperature, relative humidity, and water chemistry),
- 515 (3) **Advanced numerical modelling:** Using the seismo-hydro-mechanical dataset obtained
516 from in-situ monitoring and laboratory experiments, to evaluate the capability of
517 modelling approaches to predict the geometry and depth of failure, as well as the extent
518 of the EDZ around the test tunnel.

519 5.2 Construction of PRECODE Mine-by Tunnel

520 Within the PRECODE Project, a new 100 m long side tunnel was excavated between 2024-2025
521 as a branch from the existing Bedretto tunnel (Figure 4). The azimuth of the new tunnel is S245°W
522 and is 72° angled from the Bedretto Tunnel with an azimuth of N317°W. Before and during the
523 mine-by excavation, the vicinity of the tunnel was densely instrumented with various sensors to
524 monitor the evolution of the EDZ over a period of a few years. The tunnel was excavated at TM
525 2428 with approximate overburden depth of 1300 m, sufficient to create stress-induced brittle
526 fracturing in the form of spalling. To assess the impact of construction-induced damage (CDZ) on
527 the development of the EDZ, two excavation methods were employed during the construction of
528 the PRECODE Tunnel: 1) controlled drill and blast (D&B), and 2) line drilling and rock breaking
529 (LDRB), internally known as “soft excavation”. The comparison between the LDRB and D&B
530 sections allowed quantification of the EDZ resulting from stress redistribution alone and the EDZ
531 resulting from both the construction method and stress redistribution. The first 10 m of the tunnel
532 was built using D&B. The D&B excavation was carried out in 2024, with an average advance rate
533 of 1 m/day. The subsequent 11 m were excavated using LDRB to minimize the influence of the
534 excavation method on the development of the EDZ. The LDRB procedure involved drilling a series
535 of perimeter boreholes along the intended tunnel profile to decouple the outer perimeter from the
536 inner rock mass. The inner section was then sequentially removed using a rock splitter and a rock
537 hammer. After 4 m of advance using LDRB, excavation was paused to allow for the drilling and
538 instrumentation of a series of radial boreholes equipped with hydraulic and deformation
539 monitoring sensors after that the LDRB was continued for another 7 m (more details in Section

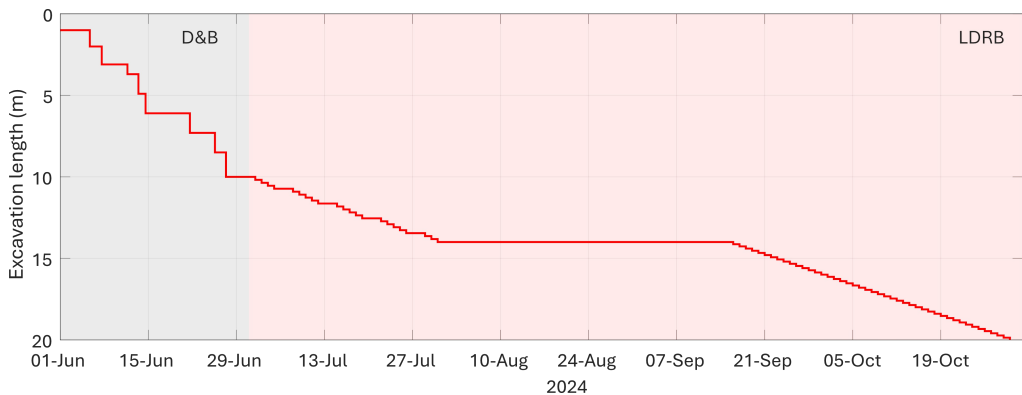


540 5.3). The construction of the LDRB section was significantly more time-demanding than the D&B
541 section, with an average advance rate of 0.1-0.25 m/day. The excavation of the PRECODE section
542 (the first 21 m) was completed in October 2024 (Figure 5). The excavation was then continued for
543 an additional ~80 m using D&B method to create an Earthquake Physics Testbed (ETB) for the
544 FEAR experiment (Meier et al., 2026).



545

546 Figure 4: Left: Isometric view of the Bedretto Tunnel and the newly-excavated PRECODE/FEAR Tunnel,
547 Right: As-built geometry of the PRECODE Section, showing D&B and LDRB sections



548

549 Figure 5: Excavation advancement for the PRECODE section

550 5.3 Multi-Disciplinary In-situ Monitoring

551 Comprehensive multidisciplinary surface and borehole monitoring was implemented before and
552 during the construction of the PRECODE section to investigate the spatial and temporal evolution
553 of the EDZ, the impact of tunneling, post-construction permeability enhancement, the influence
554 of excavation methods on EDZ characteristics, and impact of fault-activation on EDZ. A series of
555 longitudinal and radial boreholes were drilled and instrumented prior to the start of tunnel



excavation, enabling monitoring of the development of the EDZ and its impact on the surrounding rock mass during excavation and in the post-excavation phase (Shakas et al., 2026). In addition to borehole monitoring, a series of surface-based measurements including terrestrial laser scanning (TLS), surface ground-penetrating radar (GPR), electrical resistivity tomography (ERT) and fracture mapping were carried out, complementing the borehole observations. Table 3 and Figure 6 present details of the boreholes drilled prior to and during tunnel construction, along with the corresponding installed sensors.

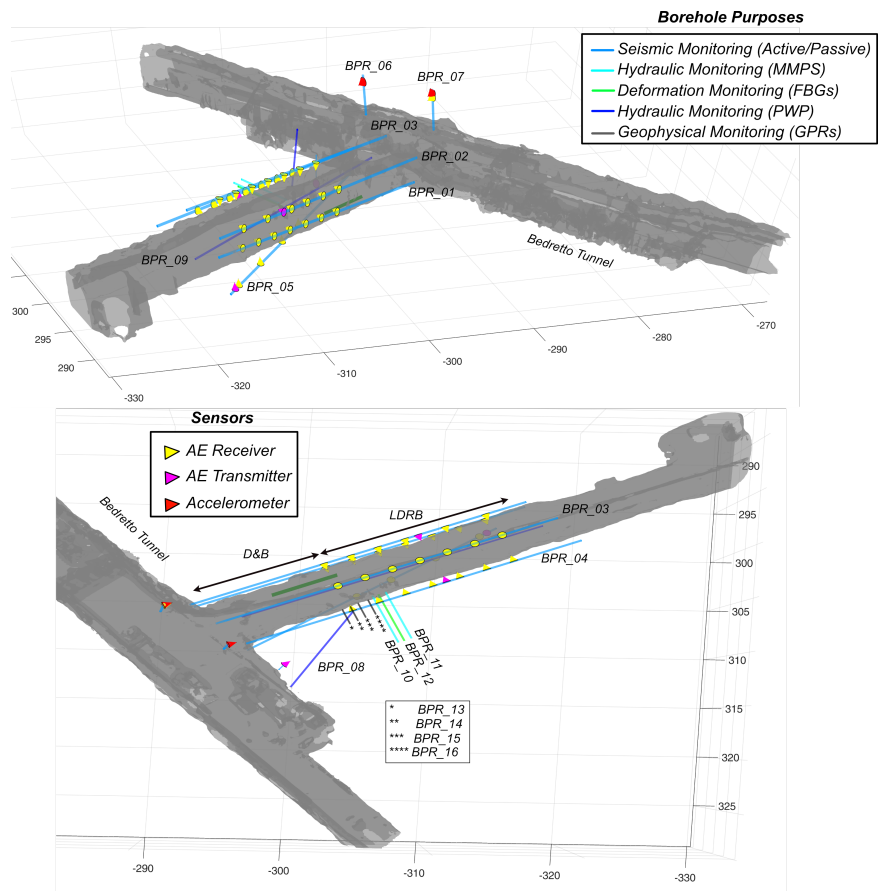
Table 2: Overview of drilled boreholes and their respective configurations around PRECODE Mine-by Tunnel

Borehole Name	Borehole Type	Length (m)	Dip from Hz. (- deg down)	direction (deg)	Diameter (mm)	Purpose
BPR_01	Longitudinal	25	2	245	96	Passive and active seismic monitoring
BPR_02		25	2	245	96	
BPR_03		25	2	245	96	
BPR_04		25	2	245	96	
BPR_05		22	-6	233	96	
BPR_06		3	80	245	96	
BPR_07		3	80	245	96	
BPR_08	Longitudinal	8.5	-5	231	76	Hydraulic monitoring
BPR_09	Longitudinal	23	-4	245	76	Characterization hole
BPR_10	Radial	5	-2	335	76	Hydraulic monitoring
BPR_11		5	-2	335	76	
BPR_12	Radial	5	-5	335	96	Deformation monitoring
BPR_13	Radial	2	-10	335	45	Geophysical monitoring
BPR_14		2	-10	335	45	
BPR_15		2	-10	335	45	
BPR_16		2	-10	335	45	

Active and Passive Seismic Monitoring: The seismic monitoring network was installed prior to initial drill-and-blast excavation phase and was designed to track the transient evolution of microcracking in the damage zone during and post excavation. The network was designed with increased sensitivity during the four- and seven-meter soft excavation phases, while sensitivity was lower during the initial drill-and-blast phase. Additionally, as excavation progressed, sensors were excluded from the location procedure once the direct rays from the microcracks to the sensors passed through the newly excavated void space. The network comprised a total of 31 in situ acoustic emission (AE) sensors, placed in five longitudinal boreholes approximately one meter from the tunnel wall (Figure 6, yellow cones). Four of these boreholes (BPR_01 to BPR_04) were dry, drilled with a 2° upward dip, and housed 27 AE sensors (GMuG-BLw-7-70) equipped with curved front surfaces. These sensors were deployed in shuttles, where spring cylinders ensured contact with the borehole walls. The fifth borehole (BPR_05), drilled with a 6° downward dip beneath the tunnel, contained four AE sensors (GMuG-BLc-30-60). In addition to the AE sensors, the borehole was equipped with distributed acoustic sensing (DAS) and distributed strain sensing (DSS) cables and was subsequently grouted. The AE sensors, manufactured by Gesellschaft für Materialprüfung und Geophysik GmbH (GMuG), operate within a bandwidth of 1 to 100 kHz, with peak sensitivity at 70 kHz. For calibration, two steeply upward-dipping boreholes



582 (BPR_06 and BPR_07) at the tunnel entrance were instrumented with one AE sensor (GMuG-BLW-
583 7-70), two one-component (1C) accelerometer (Wilcoxon 736T), and a three-component (3C)
584 accelerometers (Endevco 45A). The accelerometers are indicated by red cones in Figure 6.
585 Additionally, three seismic sources were distributed across three different boreholes for active
586 seismic surveys (Figure 6, pink cones). The data acquisition systems of the network acquired
587 continuous data at a sampling rate of 200kHz. The continuous data was processed using the
588 SeisComp seismological software package (Helmholtz-Centre Potsdam-GFZ German Research
589 Centre For Geosciences & GEMPA GmbH, 2008). Due to elevated noise levels during excavation
590 and other tunnel operations, AE data was analyzed exclusively during periods when no personnel
591 were present in the tunnel.

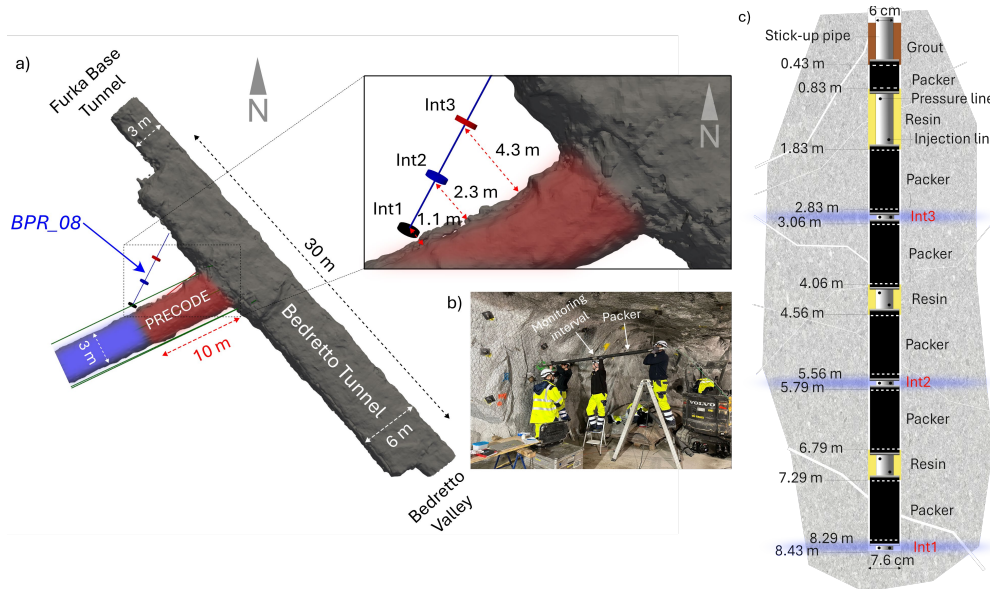


592
593 Figure 6: Multidisciplinary monitoring scheme of the PRECODE Section

594 **Pore Pressure Monitoring:** The BPR_08 borehole, with a diameter of 76 mm and a total length of
595 8.4 m, was drilled diagonally toward the PRECODE tunnel along an azimuth of S231°W with a
596 downward inclination of approximately 5°. This borehole trajectory was chosen to enable pore
597 pressure and permeability measurements at three monitoring intervals located at distances of
598 4.3 m, 2.3 m, and 1.1 m from the tunnel wall. Following coring, comprehensive borehole
599 geophysical logging was performed, including ground-penetrating radar (GPR), acoustic
600 televiewer (ATV), and optical televiewer (OTV) surveys. These logs were used to select monitoring
601 sections within intact, unfractured parts of the borehole, ensuring that the measurements



602 represent rock mass properties rather than fracture-dominated flow. The borehole was then
603 equipped with a multi-packer system to hydraulically isolate the monitoring intervals using
604 inflatable rubber packers. To further improve sealing, particularly in the presence of pre-existing
605 or excavation-induced open fractures and to prevent cross-flow between adjacent sections,
606 three additional borehole segments with lengths of 1.0 m, 0.5 m, and 0.5 m were filled with resin.
607 Pore pressure is monitored in real-time in three isolated intervals with lengths of 0.14 m, 0.23 m,
608 and 0.23 m, respectively. Pressure measurements started immediately after installation of the
609 monitoring system on 21 March 2024 using Keller pressure sensors operating at a sampling
610 frequency of 5.56×10^{-4} Hz, with an accuracy of ± 0.1 kPa, a resolution of 0.01 kPa, and a pressure
611 rating of 10 MPa. Hydraulic pulse injection tests using clean water with overpressures of
612 approximately 0.3 MPa, as well as pulse withdrawal tests, were conducted in the isolated
613 monitoring intervals before, during and after PRECODE tunnel excavation.



614
615 Figure 7: a) Instrumentation of the BPR_08 pressure monitoring borehole during the PRECODE
616 experiment; b) an image of the packer system during installation, c) the detail of the three monitoring
617 intervals

618 **Active Hydraulic Monitoring:** Active hydraulic testing was conducted in two radial, sub-parallel
619 boreholes (BPR_B_10 at TM 10.75 and BPR_B_11 at TM 11.75; Table 3 & Figure 6), drilled
620 immediately after excavation of the first 4 m of the LDRB part of the PRECODE tunnel (TM 14). The
621 boreholes were drilled sub-horizontally ($\sim 2^\circ$ downward) into the NW wall of the mechanically
622 excavated tunnel section at a height of approximately 1.20 m above the tunnel floor. In August
623 2024, two modular multi-packer system (MMPS; manufactured by Solexperts AG) were installed
624 in each of the ~ 5 m long boreholes. In each borehole, four test intervals were hydraulically
625 isolated by four inflated rubber packers with sleeve lengths of 0.25 m and 0.50 m (Figure 8), with
626 identical configurations between both boreholes. Interval spacing decreases toward the tunnel
627 wall to achieve high spatial resolution, enabling detailed characterization of the EDZ. The deepest
628 interval (Int1) extends from the deepest packer (Pa1) to the borehole end (Figure 8). The interval
629 positions and lengths are summarized in Table 4.

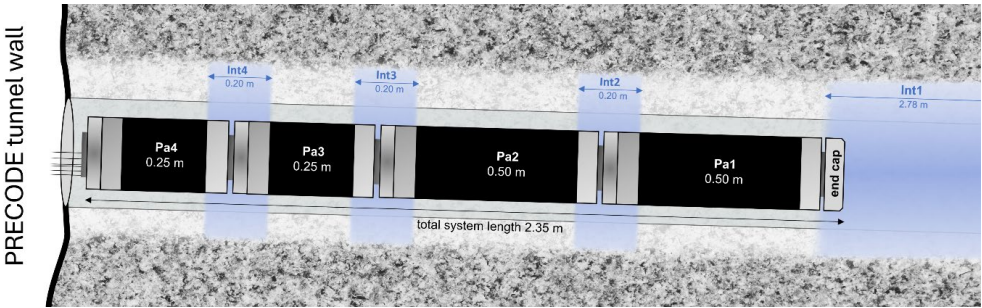


630 The MMPS installations were designed to remain in place during continued tunnel excavation and
631 after completion, allowing continuous pore-pressure monitoring and repeated hydraulic testing
632 at fixed interval positions to characterize the spatial and temporal evolution of the EDZ. A series
633 of hydraulic and pneumatic in-situ tests was performed and repeated over several field
634 campaigns to evaluate interval permeabilities. These tests include pulse injection and withdrawal
635 tests, constant head tests, and constant rate tests. Gas tests were limited to the unsaturated
636 intervals, closest to the tunnel wall (Int4), where atmospheric or sub-atmospheric pressures were
637 measured, indicating suction effects ($P_{int} \leq P_{atm}$).

638 Table 3: Borehole interval positions in BPR_B_10 and -11

Interval Name	Start depth [m]	End depth [m]	Interval length [m]
Int1	2.28	5.06	2.78
Int2	1.58	1.78	0.20
Int3	0.88	1.08	0.20
Int4	0.43	0.63	0.20

639 Hydraulic testing was performed using an electric pressure pump (*REMS E-Push 2*) and a
640 flowboard equipped with three *Bronkhorst Cori-Flow* flowmeters covering a flow-rate range from
641 0.017 to 667 ml/min. At a control board attached to the tunnel side wall, the packers and borehole
642 intervals are connected via steel lines and monitored using manometers and digital pressure
643 sensors (*Keller, type PAA-33X*). The system is designed to measure interval pressures of up to 10
644 MPa at a sampling frequency of 1 Hz. Interval pressures were continuously monitored and made
645 remotely accessible. Pneumatic tests were conducted using N_2 for injection from a gas bottle and
646 a dedicated gas flowboard equipped with three *Vögtlin* thermal mass flow meters (*red-y-smart*
647 *series*) covering a flow-rate range from 0.25 ml_n/min to 100 l_n/min.

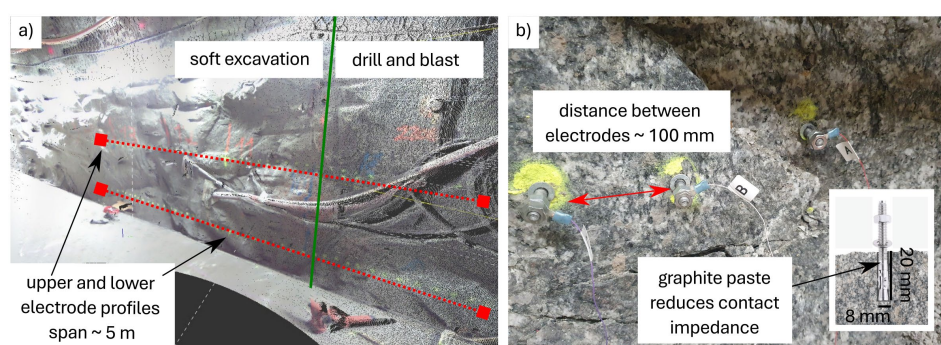


648
649 Figure 8: Schematic illustration of the modular multi-packer system (MMPS) installed in boreholes
650 BPR_B_10 and BPR_B_11. The configuration is identical for both boreholes. Packer sleeve lengths (Pa1-
651 Pa4) and the resulting hydraulically isolated test intervals are indicated (Int1-Int4)

652 **Geophysical Monitoring:** Geophysical monitoring consists of repeated borehole and tunnel
653 measurements. The objective of the geophysical monitoring is to perform a time lapse of EDZ
654 evolution and allow for comparison of excavation method on the shape and extent of the EDZ. In
655 both the radial and longitudinal boreholes, we collected repeated optical/acoustic televiewer
656 (OPTV/ATV from Advanced Logic Technologies) data to characterize the EDZ around and ahead
657 of tunnel face. The televiewer probes were run with a spacing of 5 mm. In addition, we repeated
658 borehole GPR with antennas from *GeoScanners AB* having central frequencies of 500 MHz and 1



659 GHz. The profiles were collected with a trace spacing of 8 mm using a triggering wheel and cable
 660 markers for precise location. For time-lapse imaging, we installed two electrical resistivity arrays
 661 along the tunnel wall as shown in Figure 9a. We designed the profile geometry to traverse from
 662 the drill and blast to the soft excavation part of the experiment, with the transition roughly at the
 663 middle of the profile. Each profile consists of 50 electrodes, separated by 100 mm (Figure 9b) and
 664 therefore spans roughly 5 m. The profiles are horizontal and separated by a vertical distance of
 665 0.5 m. To install the electrodes, we drilled into the rock at ~ 20 mm depth using 8 mm drill bits.
 666 We then cleaned the hole, filled it up with graphite paste and pushed permanent, stainless steel
 667 bolts inside (Figure 9b). We tested various contact methods and the graphite paste was ideal to
 668 minimize the contact impedance between the metal and rock, while also reducing the risk of
 669 corrosion over time. The mechanical coupling of the bolts was firm due to their expandable
 670 nature, however, some loose parts of the rock resulted in a total of 8 failed electrodes along the
 671 two profiles. The setup is permanent, and we connected to each electrode takeouts from a
 672 custom-made multi-core cable attached to a *GeoTOM* electrical resistivity meter. We setup the
 673 acquisition so that Wenner, Schlumberger and Dipole-Dipole arrays are repeated every 3-days.



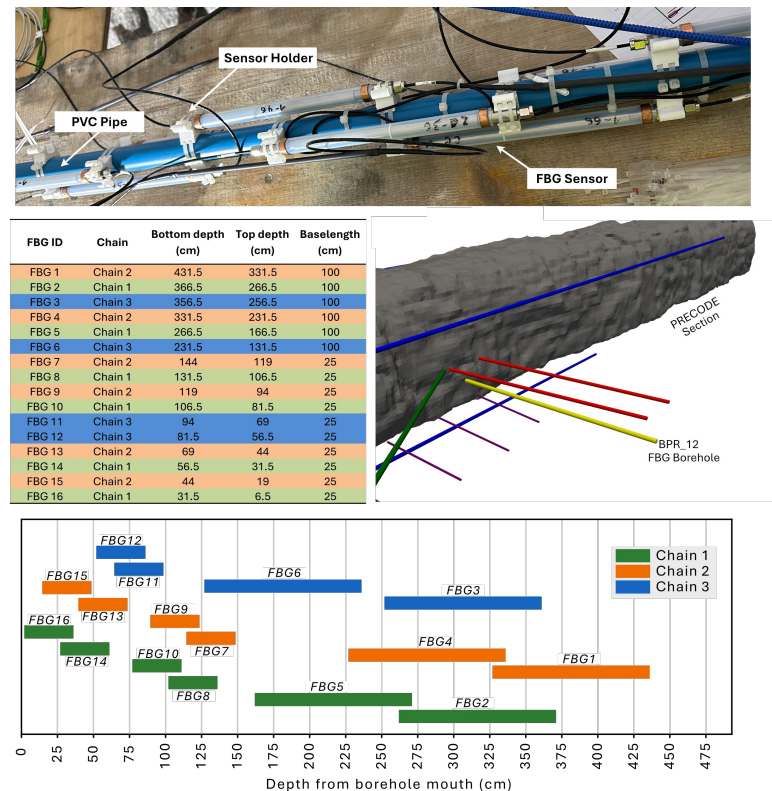
674

675 Figure 9: a) Illustration of the electrical resistivity profiles installed on the tunnel wall, that transverse
 676 between soft excavation and drill-and-blast. The image is derived from a laser scan (ref) and shows the
 677 roughness of the tunnel wall. b) A closeup picture of three electrodes, indicating the relevant dimensions
 678 and wall roughness

679 **Deformation Monitoring:** A radial borehole (BPR_12) was drilled in the tunnel sidewall
 680 approximately 1 m above the tunnel floor in an area prone to spalling concentration. The borehole
 681 was instrumented with a dense array of Fiber Bragg Grating (FBG) sensors and Distributed Strain
 682 Sensing (DSS) to monitor the evolution of deformation during and after excavation. The borehole
 683 was 5 m long, inclined downward at 5°, and had a diameter of 96 mm. In total, 16 FBG sensors
 684 (model: os3600 manufactured by Luna) were installed, of which 10 had a baselength of 25 cm
 685 and the remaining 6 had a baselength of 100 cm. The os3600 sensor has $\pm 2,500 \mu\epsilon$ strain limits
 686 with strain sensitivity of $\sim 1.2 \text{ pm}/\mu\epsilon$ with operating temperature range of -40 to 80°C . The cables
 687 were pre-strained to allow for monitoring both extensional (opening) and contractional (closing)
 688 deformation. For the FBG sensors, emphasis was placed on high redundancy, ensuring that
 689 multiple sensors covered each cross-section at nearly every depth, with higher redundancy near
 690 the tunnel sidewall, where the EDZ concentration was greatest. The FBG sensors were arranged
 691 sequentially with overlapping coverage on a PVC pipe, with the first sensor located 6.5 cm inside
 692 the borehole mouth, allowing deformation to be captured immediately at the tunnel sidewall. To
 693 provide additional redundancy and to accommodate wavelength limitations, the FBG sensors
 694 were divided into three separate chains, each with fiber access from both ends. A robust V13
 695 cable was installed in a loop configuration to provide redundancy, with the loop formed using a



696 Mini-Loop at the bottom of the borehole. The detailed depth of all sensors and other objects and
697 additional information regarding the FBG sensors can be found in Figure 10. Upon placement of
698 the sensor's chain in the borehole, cementation was carried out using a hand-driven mixing and
699 pumping system. In order to optimally fill the cavity of the borehole and guarantee a good bonding
700 to the rock, a grout mixture containing bentonite, cement and water was mixed and injected from
701 the bottom of the borehole to the borehole mount.



702
703 Figure 10: FBGs layout in BPR_12 borehole to monitor the development of deformation, Top: An example
704 of the FBG array, Middle: list of installed FBGs and isometric view of the borehole, Bottom: schematic
705 illustration of the installed FBG within the borehole and their overlaps

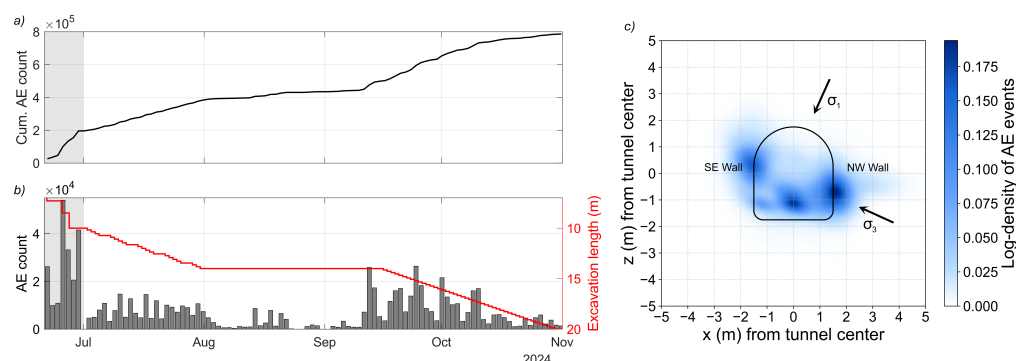
706 6. Preliminary Results

707 6.1 Brittle Damage Characterization through Acoustic Emission Monitoring

708 Acoustic emission (AE) monitoring started in June 2024, during the final phase of drill-and-blast
709 (D&B) excavation, and continued until the end of the line drilling and rock breaking phase in early
710 November 2024. To minimize false positives, AE data processing was suspended during active
711 excavation and other tunnel operations. The results of the AE monitoring as a function of time and
712 excavation progress are shown in Figure 11. Figure 11a and Figure 11b indicate that significantly
713 more AE events were recorded during the D&B phase than during excavation using the LDRB
714 method. During the D&B phase, peaks in AE activity closely correspond to individual blasting
715 rounds. During July 2024, when excavation was performed with the LDRB method, work was
716 limited to weekdays, with no excavation on weekends. Following a break between August and
717 early September, excavation resumed in mid-September and continued through weekends,



718 which coincides with an increase in AE activity. Figure 11c shows the log-density distribution of
719 AE events recorded during both excavation phases. A clear concentration of events occurs along
720 the tunnel sidewalls, where stress concentrations are expected to be highest. The spatial
721 distribution of AE events around the tunnel cross-section suggests that the maximum principal
722 stress (σ_1), likely corresponding to the vertical stress in the PRECODE tunnel, may be inclined by
723 approximately 15° relative to the vertical. The highest density of AE events occurs within
724 approximately 1.5 m of the tunnel wall, which likely represents the maximum extent of the outer
725 excavation damage zone (EDZ), including irreversibly micro-damaged rock with and without
726 significant dilation. The highest concentration of AE events is observed within the first 0.5-0.6 m
727 from the tunnel wall, likely marking the transition between the inner and outer EDZ.



728

729 Figure 11: AE monitoring during PRECODE tunnel excavation: a) cumulative AE count during PRECODE
730 excavation (shaded area indicates drill and blast excavation); b) AE count during PRECODE tunnel
731 excavation with excavation lengths marked as red solid line; c) Log-density of AE events registered during
732 the PRECODE tunnel excavation in an x-z plane. The ideal tunnel geometry is marked for reference.

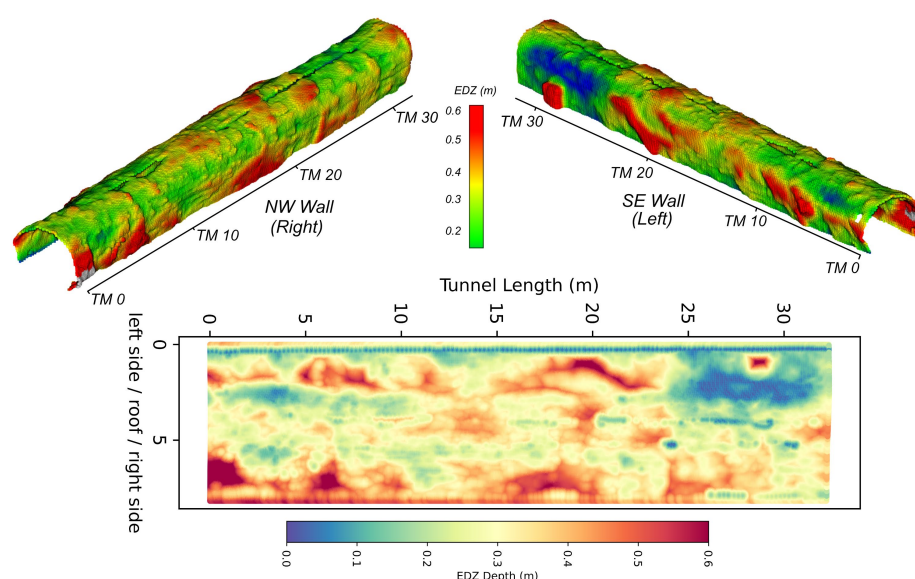
733 6.2 Short-Term Development of Excavation Damage Zone around PRECODE Mine-by 734 Tunnel

735 Upon completion of the PRECODE mine-by tunnel, a number of laser scanning and outbreak
736 mapping was conducted to assess the characteristics of EDZ around the tunnel. The point
737 clouds, with average pixel spacing of 2 mm, were acquired using terrestrial laser scanner (LiDAR,
738 Optech Teledyne). All scans were registered together and the characteristics of the EDZ including
739 depth, shape and orientation were analyzed using cloud-to-cloud distance algorithm in Cloud
740 Compare software (CloudCompare, 2023), by comparing the current profile of the tunnel against
741 the designed horseshoe-shaped tunnel.

742 Figure 12 shows the location and extent of short-term EDZ around the tunnel after completion of
743 excavation in June 2025. The maximum extent of the brittle failure in the form of spalling is
744 approximately 60 cm and is concentrated in the lower NW wall (lower right) and the upper SE wall
745 (upper left). The location of the EDZ concentration agrees with the seismic monitoring data, the
746 overall understanding of the stress regime at the BedrettoLab (Ma et al., 2019; Bröker & Ma, 2022;
747 Bröker et al., 2024) and the results of the stress modelling (Hamdi et al., 2025), where the normal
748 faulting is likely to be the background stress regime in the area of the PRECODE Section with the
749 principal stresses being slightly deviated from horizontality and verticality. The non-symmetrical
750 formation of the EDZ also suggests that the trajectory of the tunnel is not parallel to one of the
751 principal stresses. The location of the overbreaks and spalling developed in the characterization
752 borehole (BPR_09 in Figure 6) drilled parallel to the trajectory of the tunnel also agreed with the
753 spalling location around the tunnel. The rock around the PRECODE Section is considered to be
754 massive to moderately fractured rock mass with sporadic distribution of discontinuities. What is



755 also clearly observed is the interaction between the pre-existing fractures and the stress-induced
756 fractures in some areas of the mine-by tunnel, an example being the slab failure in the SE wall at
757 TM 20 and similar overbreak in the opposite wall (NW) (Figure 12). The comparison between the
758 construction induced damage (CDZ) between D&B and LDRB area shows that the extent of the
759 CDZ is approximately 15-20 cm higher in the D&B section as a result of the excavation method. It
760 is also shown that the LDRB method is efficient in reducing the CDZ due to its minimal impact to
761 the surrounding rocks.



762

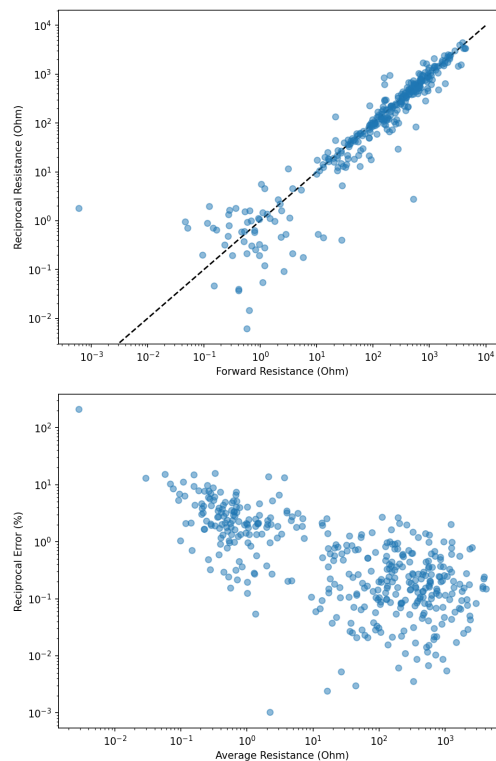
763 Figure 12: The calculated depth of EDZ around the PRECODE Tunnel obtained from TLS, Top: Isometric
764 view of the EDZ depth on the 3D point clouds, Bottom: EDZ depth projected on the unwrapped PRECODE
765 tunnel

766 6.3 Construction-Induced Damage

767 The electrical resistivity measurements collected along the PRECODE tunnel were examined to
768 investigate how the excavation process, and subsequent damage of the rock mass due to
769 different excavation method, may reflect in the electrical resistivity of the rock. Here, the
770 hypothesis is that an increased excavation damage zone may result in lower resistivity zones,
771 because the granitic host rock is highly impermeable and therefore highly resistive. Any cracks in
772 the rock will result in preferential flow paths, especially after the space between the cracks has
773 been filled by water. Since the measurements shown below have been collected almost a year
774 after the tunnel was excavated, we expect that the latter takes place and the damaged zone will
775 appear as less electrically resistive. ERT measurements are used to image the characteristics of
776 the inner and outer EDZ zones and to assess the potential impact of the excavation method on
777 their properties. The measurement process considered Wenner, Schlumberger and Dipole-
778 Dipole electrode arrays for 31 electrodes in total. These electrodes span the drill and blast (first
779 16) and soft excavation (following 15) zones. The transition occurs in the middle of the profile,
780 precisely at 1.5 m. The first electrode is denoted as the origin. To assess the precision of the
781 campaign, we collect reciprocal measurements, where the current injection electrodes (A&B) are
782 swapped with the potential measurement electrodes (M&N). Theory suggests that such a
783 swapping should result in identical readings, unless the host rock behaves non-linearly or if there
784 is a change in the medium (e.g., water infiltration or drying of the rock) occurring in the time



785 between these two measurements (Parasnis, 1988). For the latter, we do not expect any influence
786 as these measurements were collected within a few hours. Figure 13 shows the comparison
787 between the forward and reciprocal measurements for the profile. The scatter plot suggests that
788 the measurements are highly correlated, especially for high resistance values, as expected for
789 good repeatability (Tso et al., 2017). The second panel shows how the reciprocal error, defined as
790 $\frac{|R_f - R_r|}{R_{av}}$ where the average resistance is computed from $R_{av} = \frac{(R_f + R_r)}{2}$. Here, R_f refers to the
791 forward measurement and R_r refers to the reciprocal measurement. A total of 855
792 measurements are considered, and after filtering 527 measurements remain for the inversion.



793

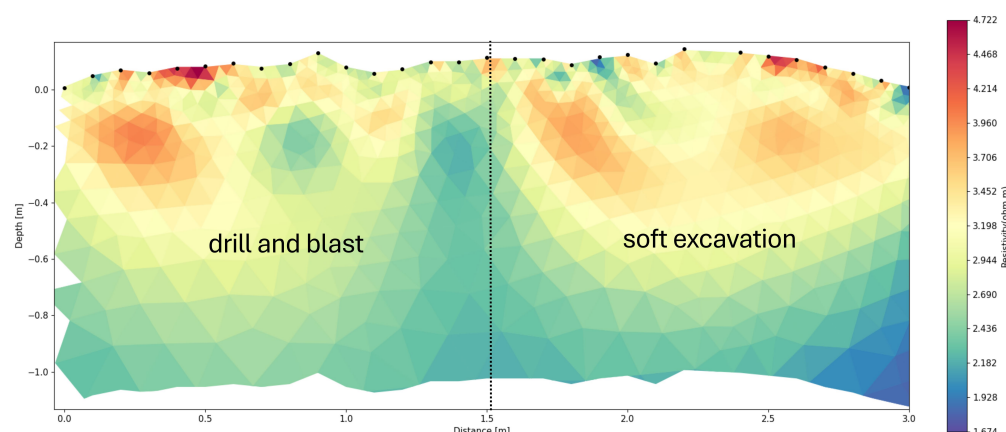
794

795 Figure 13: The left panel shows a comparison of forward and reciprocal electrical resistivity
796 measurements for the profile under examination, alongside the perfect match line. The forward and
797 reciprocal measurements overlap well, especially for high resistivities, suggesting high precision. The
798 right panel shows the reciprocal error compared to the average resistance for the measurements.
799 Reciprocal errors above 10% are not included in the data inversion.

800 To obtain a geometry for the forward and inverse modeling of the electrical resistivity data we
801 used precise laser scanning information outlined in Section 6.2. From the laser scans we
802 extracted the topography of the tunnel wall, and by determining where the electrodes lie, we
803 extracted the geometry of the profile. The deviation from a flat topography is considerable; in
804 some cases, up to 10 cm of topography is observed. It was therefore essential to account for the
805 irregular shape of the tunnel wall when forward modeling the electrical response. For the
806 subsequent modeling we only assume a 2D cross section. The inversion was run with ResIPy
807 (Blanchy et al., 2020) on a triangular mesh, and convergence is reached after roughly 10
808 iterations.



809 The results of the inversion are presented in Figure 14. Overall, higher resistivities are observed in
810 the soft excavation region compared to the D&B part, aligning with the initial hypothesis that the
811 soft excavation keeps the rock intact and therefore results in higher electrical resistance. A high
812 resistivity zone in the initial section of the drill and blast section may correlate with a block of
813 intact rock. In addition, the very high resistivity sections on the surface of the inversion mesh, at
814 0.5 m and 2.6 m, may be related to bad coupling of electrodes or loose parts of the rock that do
815 not couple well to the remaining rock mass.



816

817 Figure 14: Inversion results for electrical resistivity tomography spanning the drill and blast, and soft
818 excavation zones. The results show some prominent very high resistivity zones in the shallow few
819 centimeters, likely resulting from loose slabs of rock along the wall. The information at depth reveals a
820 higher resistivity zone for the soft excavation, suggesting more intact rock compared to the drill and blast
821 zone.

822 7. Conclusion and Outlook

823 The BedrettoLab was established in the Swiss Central Alps within the old Bedretto Tunnel. The
824 facility hosts a wide range of meso- to hectometer-scale experiments addressing key challenges
825 related to geothermal energy, induced seismicity, and deep geological disposal of nuclear waste.
826 The PRECODE mine-by experiment at BedrettoLab builds upon insights gained from existing URLs
827 to advance our understanding of time-dependent brittle fracturing in crystalline rocks relevant to
828 the disposal of heat-generating nuclear waste. By combining investigations of the mature EDZ in
829 the Bedretto Tunnel with intensive, multidisciplinary monitoring of EDZ development around the
830 PRECODE tunnel, a unique opportunity is created to study the spatial and temporal evolution of
831 the EDZ under natural repository conditions.

832 Within the PRECODE project, a new 100 m long mine-by tunnel was excavated between 2024 and
833 2025 as a branch of the existing Bedretto Tunnel at an overburden depth of approximately 1300
834 m. The high in-situ stresses associated with this overburden were sufficient to induce brittle
835 fracturing in the form of spalling. To capture the development of the Excavation Damage Zone
836 (EDZ), the tunnel surroundings were densely instrumented before and during excavation using a
837 multidisciplinary monitoring system comprising active and passive seismic monitoring, hydraulic
838 monitoring, deformation monitoring using Fiber Bragg Grating (FBG) and Distributed Strain
839 Sensing (DSS), as well as geophysical monitoring through electrical resistivity tomography (ERT)
840 and ground-penetrating radar (GPR). In addition, two excavation methods: controlled drill-and-
841 blast (D&B) and line drilling and rock breaking (LDRB) were employed to evaluate the influence of
842 excavation-induced damage on EDZ development. The comparison between the LDRB and D&B



sections provides a unique opportunity to distinguish between damage resulting primarily from stress redistribution and damage associated with both stress redistribution and excavation-induced disturbance, thereby improving our understanding of the mechanisms governing EDZ formation and evolution in crystalline rocks.

Acoustic emission (AE) monitoring revealed significantly higher AE activity during the D&B phase than during excavation using the LDRB method, indicating a greater degree of excavation-induced damage associated with drill-and-blast excavation. A pronounced concentration of AE events was observed along the tunnel sidewalls, where stress concentrations were highest. The spatial distribution of AE activity suggests that the maximum principal stress, likely corresponding to the vertical stress component at the PRECODE site, is inclined by approximately 15° relative to the vertical. Most AE events were located within approximately 1.5 m of the tunnel wall, which likely represents the maximum extent of the outer EDZ. The highest event density was concentrated within the first 0.5–0.6 m from the tunnel wall, potentially marking the transition between the inner and outer EDZ. These findings are consistent with observations from terrestrial laser scanning and geophysical monitoring using ERT. The ERT inversions indicate higher electrical resistivities in the LDRB section compared to the D&B section, supporting the hypothesis that the LDRB method reduces excavation-induced damage and preserve a greater degree of rock mass integrity.

The proximity of the PRECODE mine-by tunnel to the FEAR project, in which a natural fault zone will be stimulated under controlled conditions, provides a unique opportunity to investigate the impact of fault activation and induced seismicity on tunnel stability and EDZ evolution. Furthermore, PRECODE offers an exceptional framework for bridging laboratory and in-situ observations, particularly in improving the characterization of subcritical and time-dependent crack growth processes and their implications for long-term barrier integrity. Ongoing laboratory experiments, including fatigue, relaxation, and fracture mechanics tests conducted under controlled environmental conditions, will provide valuable insight into the long-term behavior of Rotondo Granite and the mechanisms governing stress corrosion and progressive damage accumulation. In parallel, the comparison of hydraulic properties measured within the mature (>40-year-old) EDZ of the Bedretto Tunnel and the newly excavated PRECODE tunnel will enable quantification of long-term permeability evolution under natural environmental perturbations. Together, these investigations will contribute to a more comprehensive understanding of the spatial and temporal evolution of excavation-induced damage in crystalline rocks and provide critical data for improving predictive models and long-term safety assessments of deep geological repositories for heat-generating nuclear waste.

Code and Data Availability

The tunnel geometry data, including high-resolution point clouds, surface meshes, and borehole information, are publicly available through the ETH data repository at ETH Research Collection Dataset (<https://doi.org/10.3929/ethz-c-000801363>). Additional raw and processed datasets generated and analyzed during this study are available from the corresponding author upon reasonable request.

Author Contributions

This collective work has been carried out under umbrella of the PRECODE Project and authors have contributed to different part of the manuscript. Specifically, PH, JD, MK, APR, LV, AS, PB, PA conducted the mechanical, hydrogeological, seismological, geophysical and geological



887 characterization, data collection and analysis. MP, SG, YG, SM carried out the laboratory testing
888 and analysis. KB and XM carried out the stress measurement and analysis. NB and KK provided
889 support in the numerical modelling part. PS, VD, MJ, AZ contributed to the design and installation
890 of the in-situ monitoring. MH and MW oversaw the operational and logistical aspects of the
891 experiments. FA, SW supervised the project. All co-authors contributed to the writing of the
892 paper.

893 **Competing interests.**

894 The contact author has declared that neither they nor their co-authors have any competing
895 interests.

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