



Advancing meteorite impact chronology with in situ mica Rb-Sr dating

Bruno V. Ribeiro^{1,2*}, Jayden Squire³, Christopher L. Kirkland¹, Juraj Farkas³, Cecilia Loyola³,

Victoria Cousins¹, Victor Gostin³, Charles Verdel⁴

¹School of Earth & Planetary Science (Timescales of Mineral Systems Group), Curtin University, Perth, Western Australia 6102, Australia

²John de Laeter Centre, Curtin University, Perth, Western Australia 6102, Australia

³School of Physics, Chemistry and Earth Sciences, Adelaide University, Adelaide, SA 5005, Australia

⁴Northern Territory Geological Survey, NT Government, Alice Springs, Northern Territory, 0870, Australia

*Corresponding author

Dr Bruno Vieira Ribeiro

bruno.vieiraribeiro@curtin.edu.au



1 **ABSTRACT**

2 Meteorite impacts are highly energetic processes that drives deformation on Earth under extreme
3 pressure and temperature conditions, far exceeding those of typical crustal processes.
4 Consequently, it may promote environmental perturbations, blossoming/extinction of life and even
5 produce suitable conditions for the formation of mineral resources. Thus, constraining the timing
6 of impacts is pivotal to shedding light on its role in Earth's biogeodynamics, yet less than ~20%
7 of the impacts worldwide are precisely dated. Here, we present novel in situ mica Rb–Sr isotopes
8 from the Australian Acraman and Gosses Bluff impact sites collected via LA-ICP-MS/MS,
9 including single- and multi-collector instruments to expand the chronological toolbox for dating
10 meteorite impacts. Whilst monazite (980 ± 28 Ma) and apatite (1448 ± 79 Ma) yield older ages
11 compared to the expected Acraman impact age of 588 ± 35 Ma, in situ Rb–Sr from muscovite-
12 bearing domains from the Acraman ejecta layer (580 ± 8 Ma; multi-collector age) and associated
13 fine-grained zircon (598 ± 16 Ma) are consistent with the expected impact age. Similarly, apatite
14 (132 ± 14 Ma) and mica-bearing domains (137 ± 9 Ma; multi-collector age) are comparable to the
15 proposed impact age of 133 ± 3 Ma, whilst zircon yields mostly discordant data. The Rb–Sr results
16 comparison between single- and multi-collector ICP-MS/MS has shown that the latter yielded
17 significantly more precise Sr measurements, likely due to measurement of Sr isotopes with high
18 resistor Faraday cups ($10^{13} \Omega$) resulting in improved signal-to-noise ratio, consequently yielding
19 more precise isotopic ratios and isochron ages. Our findings show that in situ Rb–Sr dating of
20 micas formed during impact metamorphism offers the means for determining the timing of
21 meteorite impact events. This approach effectively addresses the textural complexities commonly
22 present in impact-related rocks, which are often overlooked by bulk isotope-dilution techniques,
23 and delivers accuracy sufficient to establish the age of impacts. The effectiveness of in situ Rb–
24 Sr dating of impact-related micas may be attributed to their higher resistance to hydrothermal
25 alteration compared to Ar isotopes, which often yield complex degassing spectra and younger
26 apparent ages such as for the Acraman impact. This study shows that in situ Rb–Sr isotopes of
27 newly grown micas in impact-related rocks refines the chronology of impacts, with potential to
28 increase the number of dated impacts globally with a low-cost, speedy technique with minimum
29 sample preparation. Such task is crucial for understanding the role of meteorite impacts, including
30 their potential influence on environmental crises, mass extinctions, and the emergence of life.



31 **1. Introduction**

32 The Earth is one of the several planetary bodies within the solar system that is subject to
33 interaction with extraterrestrial bodies, producing meteorite impacts on the Earth's surface if these
34 bodies survive the extreme atmospheric conditions on Earth. Some meteorite impacts are known
35 for their destructive potential, such as the Chicxulub impact in the Yucatán Peninsula (Mexico),
36 potentially resulting in global consequences that could reshape ecosystems and the environment
37 with mass extinctions (Artemieva and Morgan, 2017; Hildebrand et al., 1991; Schulte et al., 2010).
38 Conversely, while locally and immediately catastrophic, these changes can create favorable
39 environmental conditions for the blossoming of primordial life (Cockell et al., 2002; Collins et al.,
40 2012; Osinski et al., 2020). 2020) and hosting economically significant mineral systems (Hayward
41 et al., 2005; Reimold et al., 2005; Tuchscherer and Spray, 2002). Importantly, knowing the timing
42 of the impact formation is essential to unravel its role in earth systems evolution and the Earth's
43 biological history (Jourdan et al., 2012, 2009; Osinski et al., 2022)

44 Establishing the timing of impact structures remains a significant challenge, with fewer than
45 20% of known examples yielding radiometric ages precise to within 2% uncertainty (French and
46 Koeberl, 2010; Quintero et al., 2021). The last revision of the Australian impacts has shown that
47 only five out of 31 confirmed impact craters/structures have well-constrained radiometric age
48 (Quintero et al., 2021). Obtaining such constraints is complicated by the enormous energy
49 released during impacts, subjecting rocks to extreme pressure–temperature deformation
50 conditions far exceeding those of typical crustal processes (Collins et al., 2012; Timms et al.,
51 2017). Rock deformation at these conditions generates profound mineralogical and chemical–
52 isotopic complexities in key gold-standard geochronometers such as zircon and monazite, often
53 resulting in incomplete isotopic resetting and intricate age spectra of difficult interpretation (e.g.,
54 Guerrero et al., 2025; Herrmann et al., 2024; McGregor et al., 2019). Thus, there is a pressing
55 need to broaden the geochronological toolbox by developing novel approaches to easily



56 accessible rock-forming minerals like micas and feldspars, which may provide robust temporal
57 constraints on meteorite impact events.

58 Here, we test a novel approach to date fine-grained micas from shock-related rocks from the
59 Australian Acraman and Gosses Bluff (Tnorala) impacts via in situ Rb–Sr geochronology carried
60 out using single- and multi-collectors instruments for the first time. This comparison allowed us to
61 test the accuracy and precision of each analytical approach to expand the geochronological
62 toolbox currently used for dating terrestrial impacts.

63 **2. Background**

64 The Acraman impact structure displays a circular structure of ~30 km diameter within the
65 Proterozoic volcanics in the Gawler Ranges (Williams, 1992, 1986). The volcanics rocks that
66 underwent impact deformation exhibit shatter cones and shock lamellae (planar deformation
67 feature) in quartz, suggesting shock pressures higher than 15 GPa (Williams, 1992, 1986). The
68 Acraman impact has been suggested to be a trigger for environmental crisis in the Late Vendian
69 spanning 600–543 Ma, a known period of increased diversification of fossil assemblages prior to
70 the Cambrian explosion (McCall, 2006), based on its size and estimated impact energy (Williams
71 and Wallace, 2003). Yet, chronological data from the impact rocks has either yielded the
72 protoliths' ages, or younger orogenic resetting events (Compston et al., 1987; Schmieder et al.,
73 2015), whilst step-heating K–Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ analyses of quartz-feldspar bearing melt rocks from
74 the center of the Acraman structure yielded complex degassing patterns and age estimates
75 younger than ca. 450 Ma (Baldwin et al., 1991; Jourdan, 2012; Schmieder et al., 2015), interpreted
76 as the timing of rock alteration post-impact. Currently, the Acraman structure is stratigraphically
77 constrained between 625–570 Ma (Hill et al., 2007). The suggested age of the Acraman event
78 comes from the distinct ejecta layer that is prominent in parts of the Adelaide Superbasin and the
79 Officer Basin (Gostin et al., 1986), with the host unit yielding a whole-rock Rb–Sr age of 588 ± 35
80 (Webb et al., 1983). However, the robustness of this age remains to be assessed since it has only
81 been mentioned in a conference abstract.



82 Gosses Bluff is a world-known impact structure located within the northern part of the
83 Paleozoic Amadeus Basin in Central Australia (Northern Territory). This structure represents an
84 eroded complex crater of an estimated ~24 km diameter that underwent extensive erosion
85 exposing a circular rim of ~5 km in diameter (Milton et al., 1972; Milton and Sutter, 1987). The
86 circular structure is hosted in flat-lying sedimentary sequences (including sandstones, shales and
87 limestones), displaying abundant evidence of impact metamorphism including shatter cones and
88 quartz sandstones with obliterated sedimentary fabric containing quartz displaying planar
89 deformation features (Crook and Cook, 1966; Dietz, 1967). Impact melt breccias composed of
90 mosaic fine-grained quartz and micas with locally rounded quartz grains (i.e., partly recrystallized
91 sandstone) and devitrified glassy rocks (including melt-pellets) displaying fine-grained quartz,
92 micas and feldspars are common in the Gosses Bluff structure (Milton and Sutter, 1987). Pioneer
93 chronological constraints produced via total fusion K-Ar of sanidine-rich suevite (133 ± 3 Ma) and
94 zircon fission track (130 ± 6 Ma) from a sandstone beneath the melt breccia suggested a Lower
95 Cretaceous age for the formation of the Gosses Bluff impact structure (Milton et al., 1972).
96 Subsequently, (Milton and Sutter, 1987) proposed an older impact formation date of 142.5 ± 0.8
97 Ma based on total fusion $^{40}\text{Ar}/^{39}\text{Ar}$ dating of pumiceous suevite clasts from the melt breccia
98 composed of fine-grained sanidine, potentially linking the formation of the Gosses Bluff impact
99 with the Jurassic–Cretaceous stratigraphic boundary. Yet, the older age should be treated with
100 caution given the discordant and complex ^{39}Ar spectrum. Similar complex $^{40}\text{Ar}/^{39}\text{Ar}$ spectrum were
101 produced from the analyses of similar rocks (Jourdan, 2012), which were interpreted to reflect a
102 ^{40}Ar Fickian diffusion profile converging towards c. 145 Ma interpreted as a minimum age for the
103 Gosses Bluff impact. Despite the multiple attempts, a temporal link between this impact with the
104 Jurassic-Cretaceous stratigraphic boundary remain unresolved.

105 Notably, the Acraman and the Gosses Bluff impact structures were tentatively dated via total
106 fusion K-Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ analyses of whole-rock specimens, dismissing the mineralogical and
107 microstructural diversity of such rocks resulting in complex results. Here, we attempt to address



108 this problem by applying novel in situ Rb–Sr geochronology by targeting mica-bearing domains
109 in ejecta layer from the Acraman and melt-pellets from the Gosses Bluff impact structures to
110 tentatively determine the timing of impact deformation. This chronological approach accounts for
111 the variable textural components on such complex rocks, opening a new field in the chronology
112 of impact processes.

113 **3. Methods**

114 3.1. Automate mineralogy mapping

115 The samples were scanned with a TESCAN Integrated Mineral Analyzer (TIMA) at Curtin
116 University's John de Laeter Centre to characterize the mineral assemblage, focusing on micas in
117 the Acraman and Gosses Bluff samples. The TIMA is an automated scanning electron
118 microscopy system that integrates quantitative EDS analysis and a statistical approach that can
119 obtain up to 500 spectra per second (Hrstka et al., 2018), which is further combined with
120 backscatter electron (BSE) images providing high-resolution mineralogy, chemistry and textural
121 data. The analyzes were carried out with an accelerating voltage of 25 kV, beam intensity of 19.8,
122 probe current of 7.15 nA, spot size of 124 nm, field size of 1000 μm and working distance of
123 15 mm. A pure Mn standard was used for EDS spectra standardization.

124 3.2. In situ Rb–Sr geochronology

125 Impact-related samples from the Acraman and Gosses Bluff were dated via in situ Rb–Sr
126 geochronology through different analytical approaches to evaluate instrument sensitivity and
127 further implications to isotopic ratio uncertainties, including single-quadrupole tandem-ICP-
128 MS/MS and multi-collector tandem-ICP-MS/MS. The Rb–Sr isochrons and ages were calculated
129 using the maximum likelihood model in IsoplotR (Vermeesch, 2018) with the ^{87}Rb decay constant
130 from Villa et al. (2015). Dates and uncertainties are presented at 2-standard error (2SE).

131 3.2.1. Single-collector ICP-MS/MS (Adelaide Microscopy, Adelaide University)

132 The protocol employed at Adelaide Microscopy (University of Adelaide) includes RESolution
133 ArF excimer LA system (193 nm) coupled to an Agilent 8900 triple quadrupole in MS/MS mode



134 with N₂O (Hogmalm et al., 2017; Zack and Hogmalm, 2016). The analyses were carried out using
135 a 67 μm circular beam diameter, 5 Hz repetition rate and a laser fluence of ~3.5 J.cm⁻² measured
136 above the cell prior to the analytical session. The international reference glass NIST SRM 610
137 with known ⁸⁷Rb/⁸⁶Sr and ⁸⁷Sr/⁸⁶Sr ratios (Woodhead and Hergt, 2001) was employed as primary
138 reference material for normalization and instrumental drift correction, while MDC phlogopite
139 (Redaa et al., 2021a) is used as the matrix matched secondary reference material for ⁸⁷Rb/⁸⁶Sr
140 correction. Entire Creek biotite and Robins Folly muscovite (Milton et al., 1972; Milton and Sutter,
141 1987) were employed as secondary reference materials and treated as unknowns to account for
142 secondary age checks.

143 Data reduction was carried out using LADR to determine drift-corrected ⁸⁷Rb/⁸⁶Sr and ⁸⁷Sr/⁸⁶Sr
144 isotopic ratios (Redaa et al., 2021a) employing NIST SRM 610 (Woodhead and Hergt, 2001) as
145 primary reference materials. Matrix correction was carried out offline using the MDC phlogopite
146 (reference age of 519.4 ± 6.5; Redaa et al., 2021a). Secondary reference materials such as the
147 Robins Folly muscovite and Entire Creek biotite yielded matrix-corrected isochron ages of 320 ±
148 10 (MSWD = 1.3, N = 25) and 311 ± 8 (MSWD = 0.1, N = 24), respectively, in agreement with
149 reference ages.

150 3.2.2. Single-collector ICP-MS/MS (GeoHistory Facility; Curtin University)

151 The analytical setup in the GeoHistory Facility includes a RESOlution 193 nm ArF excimer
152 laser with a Laurin Technic S155 sample cell coupled to an Agilent 8900 triple quadrupole mass
153 spectrometer in MS/MS mode using the N₂O in the reaction cell (Hogmalm et al., 2017). The
154 analyses were carried out using a 90 μm circular beam diameter, 7 Hz repetition rate and a laser
155 fluence of ~2.8 J.cm⁻² measured above the cell prior to the analytical session. A 'squid' mixing
156 device (Laurin Technic) was used to smooth the aerosol pulses between the laser and mass
157 spectrometer. Instrument tuning was performed using NIST SRM 610 glass in single quadrupole
158 (no-gas mode) to optimize plasma conditions, minimize oxide interferences and maximize ⁸⁵Rb
159 sensitivity. After shifting to MS/MS mode, the ICP-MS was tuned to maximizing sensitivity for SrO⁺



160 reaction products. Ablation was performed in a He atmosphere (320 ml.min⁻¹), mixed with Ar (1
161 l.min⁻¹), and a low flow of N₂ (4 ml.min⁻¹) was added to the carrier gas before the ICP torch to
162 enhance sensitivity (Hu et al., 2008).

163 Data reduction was performed in Lolite 4 (Paton et al., 2011a) using an in-house data reduction
164 scheme adapted from Redaa et al. (2021). The ⁸⁷Sr/⁸⁶Sr and ⁸⁷Rb/⁸⁶Sr ratios were drift-corrected
165 and calibrated against NIST SRM 610 with recommended isotopic ratios from Woodhead and
166 Hergt (2001). and further matrix correction was carried using an isochronous calibration based on
167 the Rb-Sr from LaPosta biotite (e.g., Giuliani et al., 2024). Unknowns were additionally bracketed
168 with Mount Dromedary biotite (Rösel and Zack, 2022). Matrix-corrected Mount Dromedary
169 returned an isochron date of 99 ± 3 Ma (n = 32, MSWD = 1.4) considering an initial ⁸⁷Sr/⁸⁶Sr of
170 0.70446 (c.f., Rösel and Zack, 2022).

171 3.2.3. Multi-collector ICP-MS/MS (GeoHistory Facility, Curtin University)

172 The approach utilized at the GeoHistory Facility (Curtin University) includes a Coherent LR
173 laser ablation system (193 nm) equipped with a Lauren Technic S155 sample cell connected to
174 a Thermo Fisher Neoma multi-collector inductively coupled plasma mass spectrometer (ICP-
175 MS/MS) equipped with a reaction cell. The analyses were carried out across two analytical
176 sessions using a 90 µm circular beam diameter, 7 Hz repetition rate and a laser fluence of ~2.8
177 J.cm⁻² measured above the cell prior to the analytical session. A 'squid' mixing device (Laurin
178 Technic) was used to smooth the aerosol pulses between the laser and mass spectrometer.
179 Ablation was performed in a He atmosphere (320 ml.min⁻¹), mixed with Ar (1 l.min⁻¹), and a low
180 flow of N₂ (3 ml.min⁻¹) was added to the carrier gas before the ICP torch to enhance sensitivity
181 (Hu et al., 2008).

182 The mass spectrometer was run in MS/MS mode using SF₆ as reaction gas to facilitate
183 simultaneous measurement of Rb and Sr isotopes (Bevan et al., 2021; Cruz-Uribe et al., 2023;
184 Dauphas et al., 2022). All relevant masses from ⁸⁵Rb to ⁸⁸Sr¹⁹F were measured for 140 integrations



185 for 0.5 s with a 30 s baseline acquisition, followed by a 30 s sample ablation and 10 s washout.
186 The $^{86}\text{Sr}^{19}\text{F}$ (H3 cup) and $^{87}\text{Sr}^{19}\text{F}$ (H4 cup) analytes were measured in Faraday cups connected
187 to $10^{13} \Omega$ amplifiers to enhance signal to noise ratios, whereas all other analytes were measured
188 in Faraday cups connected to $10^{11} \Omega$ amplifiers. The peak center was positioned on the left
189 shoulder of the time-resolved peak to reduce potential interferences from SF_3O and SF_3OH_2 on
190 masses m/z 105 and 107.

191 Data reduction was performed in Lolite 4 (Paton et al., 2011) using an in-house data reduction
192 scheme. The $^{87}\text{Rb}/^{86}\text{Sr}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ were drift-corrected and calibrated against NIST SRM 612
193 (Woodhead and Hergt, 2001), and further matrix correction was carried using an isochronous
194 calibration based on the Rb-Sr from Mount Dromedary biotite (e.g., Giuliani et al., 2024). The
195 matrix offset as a correction factor to the $^{87}\text{Rb}/^{86}\text{Sr}$ ratios of samples and LaPosta biotite
196 (secondary reference material). The matrix-corrected LaPosta returned an isochron date of $87 \pm$
197 4 Ma ($n = 13$, $\text{MSWD} = 0.2$; session 1) and $92 \pm 3 \text{ Ma}$ ($n = 26$, $\text{MSWD} = 0.4$; session 2) in
198 agreement with reference values (Rösel and Zack, 2022).

199 3.3. In situ U–Pb geochronology

200 We obtained apatite, zircon and monazite U–Pb data from the Acraman and Gosses Bluff
201 impact to compare with new mica Rb–Sr ages.

202 The setup includes a Coherent LR laser ablation system (193 nm) equipped with a Lauren
203 Technic S155 sample cell connected to an Agilent 8900 for the measurement of U–Pb isotopes.
204 A ‘squid’ mixing device (Laurin Technic) was used to smooth the aerosol pulses between the laser
205 and mass-spectrometer. The sample cell was flushed with ultrahigh purity He ($320 \text{ ml}\cdot\text{min}^{-1}$) and
206 N_2 ($1.0 \text{ ml}\cdot\text{min}^{-1}$), both of which were passed through an inline Hg trap. The low flow rate of N_2
207 was added to the carrier gas before the ICP torch to enhance sensitivity (Hu et al., 2008). The
208 samples and reference materials were ablated using a laser fluence of $\sim 2.7 \text{ J}\cdot\text{cm}^{-2}$, repetition rate
209 of 10 Hz and a circular beam of 23 μm for apatite and zircon, and 15 μm for monazite. Each
210 analysis consisted of 30 s of baseline acquisition, followed by 20 s sample ablation and 10 s fast



211 washout. Instrument tuning was performed using NIST SRM 610 glass to optimize plasma
212 conditions, maximize instrument sensitivity and minimize oxide interferences.

213 Unknowns were bracketed by a suite of reference materials. Madagascar MAD-2 apatite
214 (Thomson et al., 2012), zircon GJ-1 (Jackson et al., 2004) and Stern monazite (Horstwood et al.,
215 2016) were employed as primary reference materials for U-Pb calibration. McClure Mountain
216 (Schoene and Bowring, 2006) and FC-Duluth apatite (Apen et al., 2024; Schmitz et al., 2003;
217 Thomson et al., 2012), zircon 91500 (Woodhead and Hergt, 2005) and Plesovice (Slama et al.,
218 2008), and monazite 44069 (Aleinikoff et al., 2006) and Managotry (Montel et al., 1996) monazite
219 were employed as secondary reference materials and treated as unknown. Data reduction was
220 performed in Iolite4 (Paton et al., 2011) using an exponential with linear equation to account for
221 U-Pb downhole fractionation. The U–Pb diagrams and dates were calculated using IsoplotR
222 (Vermeesch, 2018), with ages and uncertainties being presented at the 2-standard error (2SE).

223 **4. Results**

224 4.1. Sample description

225 The ejecta layer sample from the Acraman impact displays feldspar- and fine-grained mica
226 domains (Fig. 2A). The feldspar domain is composed of coarse-grained albite containing fine-
227 grained, anhedral quartz inclusions, which are clasts of the Yardea Dacite, Gawler Ranges
228 Volcanics. The fine-grained mica domain consists of elongated white-mica sheets, identified as
229 muscovite through comparison of its EDS chemical composition with simulated data (Fig. 2B). In
230 addition, fine-grained quartz, albite and ilmenite also occurs in minor abundance (Fig. 2C, D). The
231 grain-size and polymineralic nature of the mica domain could have direct implication for in situ
232 Rb–Sr dating, imposing limitations to date individual mica sheets (Fig. 2E). Nevertheless, fine-
233 grained muscovite predominate in this texture, surpassing other minerals generally lack significant
234 amounts of Rb and Sr, except for minor occurrences of very fine-grained feldspar. Thus, the Rb–
235 Sr isotopic signal is likely to reflect the direct response of micas in such domain.



236 The Gosses Bluff melt-pellet features complex mineral textures with quartz-feldspar-mica
237 intergrowth, and patches composed of ultrafine-grained mica targeted for in situ Rb–Sr
238 geochronology (Fig. 3A). The ultrafine-grained micas are Fe-Mg-bearing classified as aliettite
239 based on the comparison between its chemistry and simulated EDS composition (Fig. 3B).
240 Simply, aliettite is a rare phyllosilicate linked to the dehydration of interstratified talc-trioctahedral
241 smectite (Brigatti and Poppi, 1987), already reported in chondrites and impact-related settings
242 (Bland et al., 2014). In high-spatial resolution backscatter images, the aliettite domains display
243 extremely fine-grained, acicular micas with complex texture and potential chemical zoning
244 (especially in Fe, Mg and Ca; major elements) with subordinate ilmenite (Fig. 3C). Some aliettite-
245 bearing domains are sufficiently clear (i.e., monomineralic) to allow for individual laser spots (Fig.
246 3D). However, the intricate intergrowth with quartz-feldspar presents challenges in obtaining
247 uncontaminated isotopic analyses, as it often results in mixing with a representative proportion of
248 feldspar (Fig. 3E).

249 4.2. Single-collector triple quadrupole Rb–Sr geochronology

250 Muscovite from the Acraman sample yielded single Rb–Sr isochrons based on the data
251 acquired by the Adelaide and Curtin University. We collected 46 Rb–Sr data at the Adelaide
252 University, which yield an inverse isochron age of 562 ± 34 Ma (MSWD = 0.9) with an initial
253 $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7279 ± 0.0074 (Fig. 4A). Overall, the individual $^{87}\text{Rb}/^{86}\text{Sr}$ ratios have an average
254 internal uncertainty of 5.20 ± 2.80 % (standard deviation hereafter), and corresponding $^{87}\text{Sr}/^{86}\text{Sr}$
255 measurements with tighter average internal uncertainty of 1.99 ± 0.75 %. At Curtin University,
256 we collected 53 Rb–Sr data yielding an inverse isochron age of 584 ± 23 Ma (MSWD = 1.5) with
257 an initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7189 ± 0.0067 (Fig. 4A). This dataset yielded more precise $^{87}\text{Rb}/^{86}\text{Sr}$ (2.67
258 ± 1.38 %) and $^{87}\text{Sr}/^{86}\text{Sr}$ (1.22 ± 0.55 %) likely due to the difference in laser parameters especially
259 beam diameter.

260 The Rb–Sr data obtained at the Adelaide University of the aliettite domains from the Gosses
261 Bluff melt-pellet yield average $^{87}\text{Rb}/^{86}\text{Sr}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ internal uncertainty of 1.78 ± 1.32 % and



262 $^{87}\text{Sr}/^{86}\text{Sr}$ internal uncertainty of 1.13 ± 0.42 %. These data yield an isochron age of 91 ± 42 Ma
263 (MSWD = 1.6, N = 56) with an initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7387 ± 0.0030 (Fig. 4C). We note some
264 overdispersed data compared to the main isochron trend, and a high homogeneity in Rb/Sr ratio
265 limiting the precision of the isochron age. The data obtained at Curtin University yielded
266 comparatively less precise $^{87}\text{Rb}/^{86}\text{Sr}$ uncertainties (average of 3.74 ± 1.66 %; with some outliers
267 with over 10 % internal uncertainty), but more precise $^{87}\text{Sr}/^{86}\text{Sr}$ (0.41 ± 0.17 %) compared to the
268 Adelaide's dataset. The Curtin data yield an inverse isochron date of 169 ± 23 Ma (MSWD = 1.3)
269 with an initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7174 ± 0.0004 (Fig. 4D).

270 4.3. Multi-collector Rb–Sr geochronology

271 The protocol utilized at Curtin University employing the ThermoFisher Neoma multi-collector
272 yielded comparatively more precise isotopic ratios to the single-collector triple quadrupole
273 protocol, yielding an overall internal $^{87}\text{Rb}/^{86}\text{Sr}$ precision of 3.66 ± 2.12 %, and significantly
274 improved $^{87}\text{Sr}/^{86}\text{Sr}$ measurements (0.81 ± 0.30 %) compared to the single-collector triple
275 quadrupole results. Across two analytical sessions, the Neoma data from the Acraman sample
276 yielded the most precise isochron ages. The data collected in session 1 yielded an inverse
277 isochron age of 575 ± 12 Ma (MSWD = 1.2, N = 42) with an initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7188 ± 0.0026 ,
278 whereas the data collected in session 2 yielded an inverse isochron age of 588 ± 12 Ma (MSWD
279 = 0.9, N = 40) with an initial $^{87}\text{Sr}/^{86}\text{Sr}$ 0.7187 ± 0.0023 . Combined, the completed dataset yields
280 an inverse isochron age of 580 ± 8 Ma (MSWD = 1.2, N = 82) with an initial $^{87}\text{Sr}/^{86}\text{Sr}$ of $0.7190 \pm$
281 0.0017 , hereafter considered the most reliable and precise age for the Acraman ejecta layer.

282 The Gosses Bluff melt-pellet yielded complex Rb–Sr data dispersion due to the textural mixing
283 between alietite and feldspar (Fig. 3A, E). Data from the mixing spots (44 out of 73 from both
284 sessions) are aligned with a reference isochron age of ca. 160 Ma, similar to the isochron age
285 obtained from the Rb–Sr data obtained with Agilent 8900 at Curtin University (Fig. 4D). The Rb–
286 Sr data from clean laser spots (i.e., with minimum feldspar mixing) yielded inverse isochron ages
287 of 139 ± 14 Ma (MSWD = 1.2, N = 17; session 1) with initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7124 ± 0.0016 and 136



288 ± 11 Ma (MSWD = 1.5, N = 12; session 2) with initial $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7236 ± 0.0015 (Fig. 5B). The
289 complete dataset yields an inverse isochron age of 137 ± 9 Ma (MSWD = 1.6, N = 29) with initial
290 $^{87}\text{Sr}/^{86}\text{Sr}$ of 0.7237 ± 0.0012 , which we interpret as the best age estimate for the alietite
291 crystallization in the Gosses Bluff melt-pellet sample.

292 4.4. U–Pb geochronology

293 From the nine monazite U–Pb data we collected from the Acraman ejecta layer, only two are
294 near-concordant yielding a weighted $^{206}\text{Pb}/^{238}\text{U}$ mean age of 980 ± 28 Ma (Fig. 6A). Fine-grained
295 apatite texturally associated with the mica-bearing domain in the ejecta layer yield a lower
296 intercept age in the Tera-Wasserburg of 1448 ± 79 Ma (MSWD = 2.0, N = 13; Fig. 6B) with an
297 initial $^{207}\text{Pb}/^{206}\text{Pb}$ of 0.901 ± 0.0029 . Whilst both monazite and apatite yield distinctively older
298 ages compared to the suggested Acraman impact age, fine-grained zircon embedded in the
299 ejecta layer micaceous matrix yield a lower intercept age of 598 ± 16 Ma (MSWD = 0.5, N = 6;
300 Fig. 6A) with an initial $^{207}\text{Pb}/^{206}\text{Pb}$ of 0.867 ± 0.016 as expected for the Neoproterozoic (Stacey
301 and Kramers, 1975).

302 Zircon from the Gosses Bluff melt-pellet yield significantly overdispersed, discordant U–Pb
303 data with only five near-concordant data smearing from ca. 1540 Ma down to ca. 510 Ma (Fig.
304 6C), imposing significantly limitation to constrain an impact age. Conversely, fine-grained apatite
305 yield a well-defined mixing line between radiogenic and common-Pb components in Tera-
306 Wasserburg space with a lower intercept age of 132 ± 14 Ma (MSWD = 0.4, N = 23; Fig. 6D) and
307 initial $^{207}\text{Pb}/^{206}\text{Pb}$ of 0.856 ± 0.013 as expected for the Neoproterozoic (Stacey and Kramers,
308 1975).

309 5. Discussion

310 5.1. Comparison between in situ single- and multi-collector Rb–Sr geochronology

311 In situ Rb–Sr geochronology has been widely propelled by the new advances in mass
312 spectrometers with collision-reaction cell that enable oxidizing reactions and consequently the
313 measurement of ^{87}Rb – ^{87}Sr interference-free (Cheng et al., 2008; Hogmalm et al., 2017; Redaa et



314 al., 2021a; Zack and Hogmalm, 2016). This includes single-collector triple quadrupole instruments
315 including the Thermo Fisher iCap TQ and Agilent 8900, and the Thermo Fisher Neoma and
316 NuPlasma Sapphire multi-collector mass spectrometers. Although single-collector triple
317 quadrupole instruments have proven its ability to successfully date mica and feldspar from a range
318 of rocks and textural settings (Li et al., 2020; Tillberg et al., 2020; Subarkah et al., 2022; Ribeiro
319 et al., 2022, 2023b, 2023a, 2025; Liebmann et al., 2022; Zametzer et al., 2022; Kirkland et al.,
320 2023; Larson et al., 2023; Almeida et al., 2023; Kellett et al., 2024; Nicomedes et al., 2025;
321 Bedoya et al., 2025; Caxito et al., 2025, among many others), the development and optimization
322 of analytical routines using multi-collector MS/MS instruments are still in its infancy (Bevan et al.,
323 2021; Cruz-Uribe et al., 2023; Dauphas et al., 2022; Huang et al., 2025) – especially when referring to
324 the comparison between those two instruments.

325 The protocols used at Curtin University employed identical laser parameters for Rb–Sr
326 analyses with both the single-collector Agilent 8900 MS/MS and the multi-collector Thermo Fisher
327 Neoma MS/MS, thereby enabling an independent comparison of instrument performance. Such
328 comparison is often carried out using homogeneous materials (e.g., NIST SRM glasses), to
329 exclude internal variability as a factor on measurement precision. As such, the Neoma MS/MS
330 results yielded significantly more precise and accurate $^{87}\text{Rb}/^{86}\text{Sr}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ ratios from NIST
331 SRM 612 compared to the single-collector Agilent 8900 MS/MS (Fig. 7A), likely due to better
332 performance of the Neoma MS/MS in measuring ^{84}SrF , ^{86}SrF and ^{87}SrF with high resistor Faraday
333 cups of $10^{13} \Omega$ (e.g., Huang et al., 2025) improving signal-to-noise ration and resulting in more
334 precise isotopic measurements compared to the ion counter measurements in peak hoping mode
335 on the Agilent 8900 triple quadrupole. This is more significant for minerals with high Sr
336 concentration such as feldspar and some mica variations (e.g., white-mica and some biotite)
337 (Bevan et al., 2021; Rankenburg et al., 2024), often yielding 10 times better internal precision on
338 Sr isotopic measurements. For instance, we observed that Mount Dromedary biotite yielded more



339 precise $^{87}\text{Rb}/^{86}\text{Sr}$ (1.05 ± 0.50 %, internal precision) and $^{87}\text{Sr}/^{86}\text{Sr}$ (0.85 ± 0.20 %) ratios compared
340 to Agilent 8900 ($^{87}\text{Rb}/^{86}\text{Sr}$ of 1.43 ± 0.27 %, and $^{87}\text{Sr}/^{86}\text{Sr}$ of 1.10 ± 0.17 %) (Fig. 7B, C).
341 Conversely, low-Sr micas such as LaPosta biotite are better suited to be analyzed via ion-counted
342 equipped instruments like the Agilent 8900 triple quadrupole, yielding more precise isotopic
343 measurements compared to the Neoma MS/MS (Fig. 7B, C).

344 In summary, these results indicate that choosing the best analytical approach for unknown
345 samples rely on knowing the estimated Sr concentration in the mineral of interest. Overall, the
346 Neoma MS/MS is better suited to high-Sr, low Rb/Sr ratio specimens like alkali feldspar and some
347 white-mica and clay mineral varieties (e.g., sericite, muscovite, illite and glauconite) (Bevan et al.,
348 2021; Huang et al., 2025; Kirkland et al., 2025), resulting in tighter constrained isochron ages in
349 spite of the narrow spread of $^{87}\text{Rb}/^{86}\text{Sr}$ ratios.

350 5.2. Assessing the chronology of the Acraman and Gosses Bluff impacts

351 The Acraman and Gosses Bluff impacts has been continuously investigated in respect to its
352 potential significance with global climate and environmental events, requiring precise chronology
353 to establish such correlations. Various dating approaches including whole-rock Rb–Sr (isotope
354 dilution), whole-rock and minerals (e.g., feldspar) total fusion K–Ar and $^{40}\text{Ar}/^{39}\text{Ar}$, and zircon U–Pb
355 and fission track have been applied to an array of impact-related samples including suevite and
356 melt breccias (e.g., Jourdan, 2012; Milton et al., 1972; Milton and Sutter, 1987; Webb et al., 1983;
357 Williams and Schmidt, 2021), but the impact age of the Acraman and Gosses Bluff remains
358 debatable (Jourdan, 2012).

359 5.2.1. Acraman impact structure

360 Schmieder et al. (2015) carried out the latest and most comprehensive investigation of the
361 Acraman chronology yet by focusing on U–Pb geochronology of unshocked and shocked zircon
362 from impact-melt rocks and step-heating $^{40}\text{Ar}/^{39}\text{Ar}$ analysis of melted lithic dacite clasts and
363 feldspar separates from shatter cones. Unshocked to moderately shocked zircon (with planar
364 deformation features) from impact-melt rocks and from ejecta layers returned Mesoproterozoic



365 ages (c. 1.6 Ga) reflecting the magmatic age of the Gawler Volcanic Rages (Compston et al.,
366 1987; Schmieder et al., 2015). ZrO₂-bearing granular zircon, often serving as gold-standard
367 chronometer to constrain the timing of impact metamorphism (Erickson et al., 2021, 2020; Osinski
368 et al., 2022; Timms et al., 2017), defined an ancient Pb-loss trend towards ca. 1114 Ma. Notably,
369 this age is far from the Ediacaran ejecta-stratigraphic age bracket spanning 635–540 Ma
370 considered the most reliable age interval for the timing of the Acraman impact, suggesting that
371 granular zircon experienced partial diffusional Pb diffusion potentially attributed to impact
372 metamorphism. Successive attempts in obtaining step-heating ⁴⁰Ar/³⁹Ar ages from impact-melt
373 rocks and K-bearing mineral separates like alkali feldspar failed to yield plateau ages due to ³⁹Ar
374 and ³⁷Ar recoil redistribution (Jourdan, 2012), or yielded (mini)-plateau ages younger than 460–
375 450 Ma (Baldwin et al., 1991; Jourdan, 2012; Schmieder et al., 2015). Such young ⁴⁰Ar/³⁹Ar ages
376 are interpreted to reflect post-impact hydrothermal alteration when compared to the ejecta-
377 stratigraphic constraints, supported by the presented of secondary minerals (e.g., zeolite in vugs
378 and chlorite and baryte in voids). To date, the ejecta layer whole-rock Rb–Sr isochron age of 588
379 ± 35 Ma may be considered the most reliable impact age (Webb et al., 1983), although its
380 accuracy and statistical robustness are unknown.

381 Considering the Neoma superior precision for analysing the muscovite from the Acraman
382 ejecta layer, the combined Neoma Rb–Sr isochron age of 580 ± 8 Ma is interpreted as the best
383 age estimate for the Acraman impact. This age is supported by the new zircon U–Pb lower
384 intercept age of 598 ± 16 Ma (although less precise) from the ejecta layer, and with the recent in
385 situ Rb–Sr age of 585 ± 15 Ma from the Acraman ejecta layer obtained via single-collector triple
386 quadrupole ICP-MS/MS (Collins et al., 2026). These results show that the known complications
387 in dating the Acraman impact (e.g., inheritance or partial reset preserved in apatite and monazite
388 and hydrothermal alteration as shown by the ⁴⁰Ar/³⁰Ar data) could be overcome through in situ
389 laser beam techniques with sufficient lateral resolution to precisely date newly-grown mica



390 textures with more resilient isotopes to post-impact hydrothermal alteration and weathering such
391 as Rb–Sr (and potentially K–Ca yet to be explored).

392 5.2.2. Gosses Bluff impact crater

393 Early investigations on the Gosses Bluff chronology have set an important framework to
394 constraint the timing of meteorite impact directly using impact-related melt-rocks. Total fusion K–
395 Ar dating of melt-rock (133 ± 3 Ma) and zircon fission track dating (134 ± 6 Ma) defined the best
396 estimate for the timing of the Gosses Bluff impact (Milton et al., 1972; Milton and Sutter, 1987).
397 Further attempts to refine this age with high-precision $^{40}\text{Ar}/^{39}\text{Ar}$ dating from pumiceous suevite
398 clast and single sanidine grains from melt-rock yielded complex, discordant $^{39}\text{Ar}^*$ degassing
399 patterns failing to yield plateau ages, rather displaying an increase age spectrum converging
400 towards ca. 145 Ma (Jourdan, 2012; Milton and Sutter, 1987) (Fig. 8B). The older apparent age
401 has been interpreted as minimum age for the Gosses Bluff impact (Jourdan, 2012).

402 In situ Rb–Sr data collected using single-collector triple quadrupole instruments did not
403 produce accurate and precise ages, likely because of the complex feldspar-alietite intergrowths
404 (Fig. 3E) and the instrument's lower mass resolution compromising isotopic ratio precision and
405 obscured variability arising from the textural complexity. The Neoma MS/MS yielded significantly
406 more precise Sr isotopic measurements, allowing the distinction between isochrons defined by
407 pure alietite (137 ± 9 Ma) and mixing feldspar-alietite domains (ca. 160 Ma; Fig. 5B). The
408 younger isochron age from the pure alietite domains is consistent with new apatite U–Pb lower
409 intercept age of 132 ± 14 Ma, likely reflecting apatite growth from the impact melt, and with melt-
410 rock K–Ar and zircon fission track ages (Milton et al., 1972; Milton and Sutter, 1987). Whilst the
411 zircon from the Gosses Bluff melt-pellet yielded mostly discordant U–Pb data and an array of
412 concordant data between 1540–510 Ma, potentially reflecting an inherited component or partial
413 reset, preliminary results from granular, shock-recrystallized zircons suggest mean U–Pb age of
414 139 ± 4 Ma (Cousins et al., 2025).



415 Collectively, the published and new ages (this study) from the Gosses Bluff points to an impact
416 age of 133 ± 3 Ma (Fig. 8B), indicating that older $^{40}\text{Ar}/^{39}\text{Ar}$ data with complex degassing spectra
417 could represent excess and/or inherited $^{40}\text{Ar}^*$ likely producing older apparent ages (e.g., Jourdan
418 et al., 2007).

419 5.3. Advances and limitations of in situ mica Rb–Sr geochronology in meteorite impacts

420 In situ mica Rb–Sr geochronology has been successfully applied to dating complex textures
421 in deformed rocks (e.g. Kirkland et al., 2023; Larson et al., 2023; Ribeiro et al., 2023a, 2023b;
422 Tollefson and Larson, 2026; Tollefson et al., 2025), yet the behavior of Rb–Sr isotopes in shocked
423 micas and its potential to constraint the timing of impact processes remain underexplored (Ribeiro
424 et al., 2026).

425 The Rb–Sr results from the Acraman and Gosses Bluff impacts have shown newly grown
426 micas in ejecta layer and melt-rocks have the potential to preserve the timing of impact processes,
427 whilst apparently less susceptible to post-impact hydrothermal alteration compared to Ar isotopes
428 (Jourdan, 2012; Milton and Sutter, 1987). Additionally, the lateral-spatial resolution achieved by
429 modern laser ablation system provide the means to resolve complex textural relationships,
430 intrinsic to impact-related rocks, aid in obtaining clean analytical data. Combined with high-
431 precision reaction-cell instruments, such as the Thermo Fisher Neoma multi-collector, in situ Rb–
432 Sr dating of complex mica textures can overcome challenges posed by inherited or altered
433 isotopic systems and obtain chronologies that more accurately reflect the timing of meteorite
434 impacts such as inherited and/or excess $^{40}\text{Ar}^*$ (Jourdan, 2012; Jourdan et al., 2007). Since micas
435 are widely abundant on terrestrial rocks (Bailey, 1984), in situ Rb–Sr also circumvents the
436 challenges of dating accessory phases (< 1 vol.%) such as apatite, zircon and monazite. Despite
437 being powerful to precisely define the timing of impact via U–Pb system in some cases (Erickson
438 et al., 2021, 2020; Kenny et al., 2019; Kirkland et al., 2025), those accessory minerals often yield
439 complex datasets due to inheritance components and/or partial isotopic resetting as observed in
440 this study and in many other impact sites (Bohor et al., 1993; Erickson et al., 2021, 2017; Kirkland



441 et al., 2025; McGregor et al., 2019; Ribeiro et al., 2026; Schmieder et al., 2015). In such cases,
442 newly grown micas not only are far more abundant and easier to date via in situ Rb–Sr techniques,
443 but it can also provide high-quality isotopic data to pinpoint the timing of impact processes at
444 relevant precision. As this field advances, the reliability and applicability of in situ Rb–Sr
445 geochronology may provide valuable insights into both terrestrial and planetary impact events
446 (e.g., Luu et al., 2026).

447 Despite these encouraging new results, in situ Rb–Sr results of pre-impact micas from the
448 Yarrabubba impact structure has shown that these isotopes are resilient to impact-driven resetting
449 despite the formation of shock-related microstructures that could enhance isotope diffusion
450 (Ribeiro et al., 2026). Yet they preserve the protolith’s crystallization age rather than the impact
451 age, likely due to the short timescale of the thermal peak phase associated with impact
452 metamorphism (Ribeiro et al., 2026). We recommend that future mica dating studies on impact-
453 related rocks focus on newly grown textures, as they are more likely to provide accurate impact
454 ages.

455 5.4. Further applications for in situ Rb–Sr dating in planetary sciences

456 K-bearing minerals including alkali feldspar and a wide variety of micas (white-mica, biotite,
457 phlogopite) and clays (illite-smectite) are common on Earth’s terrestrial rocks, but also on
458 extraterrestrial rocks (Foland, 1998). In addition to being datable via K–Ar, $^{40}\text{Ar}/^{39}\text{Ar}$, they are also
459 potential targets for in situ Rb–Sr dating opening a new research avenue to investigate impact-
460 melt formation, deformation and syn- to post-impact hydrothermal alteration on Earth and
461 extraterrestrial materials (e.g., Papanastassiou and Wasserburg, 1972a, 1972b). The use of Rb–
462 Sr isotopes, chemically separated via isotope dilution and independently measured in solution
463 mode, has been successfully employed to explore the timing and nature of aqueous alteration of
464 Martian meteorites (Beard et al., 2013), defining the time and duration of lunar crust formation
465 (Carlson et al., 2014) and the overall isotopic evolution of Martian meteorites in combination with
466 other radiogenic isotopes (Borg et al., 2005).



467 Although isotope dilution Rb–Sr isotopic measurements provide precise isotopic ratios from
468 critical, but low-abundance Rb and Sr minerals encountered in extraterrestrial rocks including
469 orthopyroxene, plagioclase and chromite (e.g., Papanastassiou and Wasserburg, 1972b; Beard
470 et al., 2013), the development and optimization of in situ Rb–Sr analysis isotopic measurements
471 via multi-collector ICP-MS/MS holds significant promise for exploring extraterrestrial materials
472 with accessible, speedy, low-cost technique (Dauphas et al., 2022; Nie et al., 2024). The
473 capability to target specific mineral phases could provide critical insights into the timing of
474 planetary surface evolution, the chronology of impact events across the solar system, and the
475 associated effects on planetary environments and potential habitability. As analytical techniques
476 continue to improve, in situ Rb–Sr geochronology is poised to become a valuable tool for
477 unraveling complex geological histories both on Earth and beyond.

478 **6. Conclusions**

479 This study demonstrates the significant advances and ongoing challenges of in situ Rb–Sr
480 geochronology for precisely dating impact-related processes in both terrestrial and extraterrestrial
481 materials. Newly grown micas in ejecta layers and melt-rocks have proven to be robust recorders
482 of meteorite impact timing, offering greater resistance to post-impact alteration compared to
483 traditional Ar isotope systems. Modern analytical techniques, including high-resolution laser
484 ablation and multi-collector mass spectrometry, enable the resolution of complex mineral textures
485 and improve the accuracy of isotopic measurements. The resilience of pre-impact micas to
486 isotopic resetting highlights the importance of targeting newly formed mineral domains for reliable
487 age determinations. Further, the application of in situ Rb–Sr dating to planetary science opens
488 new avenues for understanding planetary surface evolution, impact histories, and related
489 hydrothermal processes. As analytical methods continue to evolve, in situ Rb–Sr geochronology
490 is poised to provide critical insights into the timing and effects of impact events across our solar
491 system.

492



493 **Acknowledgements**

494 This work is supported by the Timescales of Mineral Systems Group (Curtin
495 University). Instruments in the John de Laeter Centre, Curtin University, are supported by the
496 Australian Research Council, AuScope, and the Australian Government via the National
497 Collaborative Research Infrastructure Strategy. This research was undertaken using the Tescan
498 Clara FESEM (ARC LE190100176) at the John de Laeter Centre, Curtin University.

499

500 **Author's contributions**

501 Bruno V. Ribeiro: Conceptualization, formal analysis, data curation, writing – original draft

502 Jayden Squire: Conceptualization, formal analysis, data curation, writing – original draft

503 Christopher Kirkland: Conceptualization, validation, writing – original draft

504 Juraj Farkas: Conceptualization, validation, writing – original draft

505 Cecilia Loyola: Formal analysis, data curation, writing – review and editing

506 Victoria Cousins: Formal analysis, data curation, writing – review and editing

507 Victor Gostin: Validation, writing – review and editing

508 Charles Verdel: Validation, writing – review and editing

509

510 **Competing interests**

511 Authors declare that they have no competing interests.

512

513

514

515

516

517

518



519 References

- 520 Aleinikoff, J.N., Schenck, W.S., Plank, M.O., Srogi, L., Fanning, C.M., Kamo, S.L., Bosbyshell, H., 2006.
521 Deciphering igneous and metamorphic events in high-grade rocks of the Wilmington Complex,
522 Delaware: Morphology, cathodoluminescence and backscattered electron zoning, and SHRIMP U-
523 Pb geochronology of zircon and monazite. *Geol. Soc. Am. Bull.* 118, 39–64.
- 524 Almeida, V.V., Ribeiro, B.V., Rodrigues, J.B., Cawood, P.A., Neto, I.C., Silveira, F.V., Faleiros, F.M.,
525 Kirscher, U., 2023. Linking kimberlite magmatism in the Brazilian Platform with Pangea break-up
526 events using in situ Rb Sr in phlogopite. *Lithos* 107486.
527 <https://doi.org/10.1016/j.lithos.2023.107486>
- 528 Apen, F.E., Gaynor, S.P., Schoene, B., Cottle, J.M., 2024. Evaluating reference materials and common-Pb
529 corrections for high-resolution apatite U Pb geochronology. *Chem. Geol.* 661, 122191.
530 <https://doi.org/10.1016/j.chemgeo.2024.122191>
- 531 Artemieva, N., Morgan, J., 2017. Quantifying the Release of Climate-Active Gases by Large Meteorite
532 Impacts With a Case Study of Chicxulub. *Geophys. Res. Lett.* 44.
533 <https://doi.org/10.1002/2017GL074879>
- 534 Bailey, S.W., 1984. *Micas: Reviews in Mineralogy*. Mineralogical Society of America.
- 535 Baldwin, S.L., McDougall, I., Williams, G.E., 1991. K/Ar and ⁴⁰Ar/³⁹Ar analyses of meltrock from the
536 Acraman impact structure, Gawler Ranges, South Australia. *Australian Journal of Earth Sciences* 38,
537 291–298. <https://doi.org/10.1080/08120099108727973>
- 538 Beard, B.L., Ludois, J.M., Lapen, T.J., Johnson, C.M., 2013. Pre-4.0 billion year weathering on Mars
539 constrained by Rb–Sr geochronology on meteorite ALH84001. *Earth Planet. Sci. Lett.* 361, 173–182.
540 <https://doi.org/10.1016/j.epsl.2012.10.021>
- 541 Bedoya, A., Glorie, S., Hand, M., Kirkland, C.L., De Vries Van Leeuwen, A.T., 2025. In situ Rb–Sr insights in
542 the cooling history of the Petermann Orogeny, Central Australia. *Geoscience Frontiers* 16, 102080.
543 <https://doi.org/10.1016/j.gsf.2025.102080>
- 544 Bevan, D., Coath, C.D., Lewis, J., Schwieters, J., Lloyd, N., Craig, G., Wehrs, H., Elliott, T., 2021. In situ Rb–
545 Sr dating by collision cell, multicollection inductively-coupled plasma mass-spectrometry with pre-
546 cell mass-filter, (CC-MC-ICPMS/MS). *J. Anal. At. Spectrom.* 36, 917–931.
547 <https://doi.org/10.1039/D1JA00006C>
- 548 Bland, P.A., Collins, G.S., Davison, T.M., Abreu, N.M., Ciesla, F.J., Muxworthy, A.R., Moore, J., 2014.
549 Pressure–temperature evolution of primordial solar system solids during impact-induced
550 compaction. *Nat. Commun.* 5, 5451. <https://doi.org/10.1038/ncomms6451>
- 551 Bohor, B.F., Betterton, W.J., Krogh, T.E., 1993. Impact-shocked zircons: discovery of shock-induced
552 textures reflecting increasing degrees of shock metamorphism. *Earth Planet. Sci. Lett.* 119, 419–
553 424. [https://doi.org/10.1016/0012-821X\(93\)90149-4](https://doi.org/10.1016/0012-821X(93)90149-4)
- 554 Borg, L.E., Edmunson, J.E., Asmerom, Y., 2005. Constraints on the U-Pb isotopic systematics of Mars
555 inferred from a combined U-Pb, Rb-Sr, and Sm-Nd isotopic study of the Martian meteorite Zagami.
556 *Geochim. Cosmochim. Acta* 69, 5819–5830. <https://doi.org/10.1016/j.gca.2005.08.007>
- 557 Brigatti, M.F., Poppi, L., 1987. Natural and monoionic alietite: hydration and dehydration states. *Clay*
558 *Miner.* 22, 187–197. <https://doi.org/10.1180/claymin.1987.022.2.07>



- 559 Carlson, R.W., Borg, L.E., Gaffney, A.M., Boyet, M., 2014. Rb-Sr, Sm-Nd and Lu-Hf isotope systematics of
560 the lunar Mg-suite: the age of the lunar crust and its relation to the time of Moon formation.
561 Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences
562 372, 20130246. <https://doi.org/10.1098/rsta.2013.0246>
- 563 Caxito, F.A., Ribeiro, B. V., Alkmim, F., Deramchi, A., Kirkland, C.L., 2025. Near-Synchronous Cambrian
564 Evolution of the Transbrasiliano-Kandi-4°50' Shear Zone in South America and Africa. Terra Nova.
565 <https://doi.org/10.1111/ter.70017>
- 566 Cheng, P., Koyanagi, G.K., Bohme, D.K., 2008. On the chemical resolution of the $87\text{Rb}+$ (s0)/ $87\text{Sr}+$ (s1)
567 isobaric interference: A kinetic search for an optimum reagent. Anal. Chim. Acta 627, 148–153.
568 <https://doi.org/10.1016/j.aca.2008.03.057>
- 569 Cockell, C.S., Lee, P., Osinski, G., Horneck, G., Broady, P., 2002. Impact-induced microbial endolithic
570 habitats. Meteorit. Planet. Sci. 37, 1287–1298. [https://doi.org/10.1111/j.1945-](https://doi.org/10.1111/j.1945-5100.2002.tb01029.x)
571 [5100.2002.tb01029.x](https://doi.org/10.1111/j.1945-5100.2002.tb01029.x)
- 572 Collins, A.S., Blades, M.L., Subarkah, D., Cooke, H., Edwards, L., Forbister, C., Jolly, J., Lloyd, J.C., Gilbert,
573 S.E., Löhr, S., Farkaš, J., Gostin, V., 2026. Age of the Acraman impact ejecta layer in the adelaide
574 superbasin and Implications on clay-mineral provenance from the Rb–Sr systematics of middle
575 Ediacaran shales. Precambrian Res. 435, 108015.
576 <https://doi.org/10.1016/j.precamres.2026.108015>
- 577 Collins, G.S., Melosh, H.J., Osinski, G.R., 2012. The Impact-Cratering Process. Elements 8, 25–30.
578 <https://doi.org/10.2113/gselements.8.1.25>
- 579 Compston, W., Williams, I.S., Jenkins, R.J.F., Gostin, V.A., Haines, P.W., 1987. Zircon age evidence for the
580 Late Precambrian Acraman ejecta blanket. Australian Journal of Earth Sciences 34, 435–445.
581 <https://doi.org/10.1080/08120098708729424>
- 582 Cousins, V.K., Cavosie, A.J., Timms, N.E., Quintero, R.R., Kirkland, C.L., Evans, N.J., McDonald, B., 2025.
583 Constraining the Age of Impact: A Microstructural Investigation of Zircon from the Gosses Bluff
584 Impact Structure, Australia, in: 87th Annual Meeting of the Meteoritical Society. Perth, Western
585 Australia.
- 586 Crook, K.A.W., Cook, P.J., 1966. Gosses bluff — diapir, crypto-volcanic structure or astrobleme?*. Journal
587 of the Geological Society of Australia 13, 495–516. <https://doi.org/10.1080/00167616608728628>
- 588 Cruz-Urbe, A.M., Craig, G., Garber, J.M., Paul, B., Arkula, C., Bouman, C., 2023. Single Spot Rb-Sr
589 Isochron Dating of Biotite by LA-MC-ICP-MS/MS. Geostand. Geoanal. Res. 47, 795–809.
590 <https://doi.org/10.1111/ggr.12518>
- 591 Dauphas, N., Hopp, T., Craig, G., Zhang, Z.J., Valdes, M.C., Heck, P.R., Charlier, B.L.A., Bell, E.A., Harrison,
592 T.M., Davis, A.M., Dussubieux, L., Williams, P.R., Krawczynski, M.J., Bouman, C., Lloyd, N.S.,
593 Tollstrup, D., Schwieters, J.B., 2022. In situ 87Rb – 87Sr analyses of terrestrial and extraterrestrial
594 samples by LA-MC-ICP-MS/MS with double Wien filter and collision cell technologies. J. Anal. At.
595 Spectrom. 37, 2420–2441. <https://doi.org/10.1039/D2JA00135G>
- 596 Dietz, R.S., 1967. Shatter Cone Orientation at Gosses Bluff Astrobleme. Nature 216, 1082–1084.
597 <https://doi.org/10.1038/2161082a0>
- 598 Erickson, T.M., Kirkland, C.L., Jourdan, F., Schmieder, M., Hartnady, M.I.H., Cox, M.A., Timms, N.E., 2021.
599 Resolving the age of the Houghton impact structure using coupled $40\text{Ar}/39\text{Ar}$ and U-Pb
600 geochronology. Geochim. Cosmochim. Acta 304, 68–82. <https://doi.org/10.1016/j.gca.2021.04.008>



- 601 Erickson, T.M., Kirkland, C.L., Timms, N.E., Cavosie, A.J., Davison, T.M., 2020. Precise radiometric age
602 establishes Yarrabubba, Western Australia, as Earth's oldest recognised meteorite impact
603 structure. *Nat. Commun.* 11, 300. <https://doi.org/10.1038/s41467-019-13985-7>
- 604 Erickson, T.M., Timms, N.E., Kirkland, C.L., Tohver, E., Cavosie, A.J., Pearce, M.A., Reddy, S.M., 2017.
605 Shocked monazite chronometry: integrating microstructural and in situ isotopic age data for
606 determining precise impact ages. *Contributions to Mineralogy and Petrology* 172, 11.
607 <https://doi.org/10.1007/s00410-017-1328-2>
- 608 Foland, K.A., 1998. Potassium-argon dating method, in: *Geochemistry*. Kluwer Academic Publishers,
609 Dordrecht, pp. 522–525. https://doi.org/10.1007/1-4020-4496-8_259
- 610 French, B.M., Koeberl, C., 2010. The convincing identification of terrestrial meteorite impact structures:
611 What works, what doesn't, and why. *Earth. Sci. Rev.* 98, 123–170.
612 <https://doi.org/10.1016/j.earscirev.2009.10.009>
- 613 Giuliani, A., Oesch, S., Guillong, M., Howarth, G.H., 2024. Mica Rb-Sr dating by laser ablation ICP-MS/MS
614 using an isochronous calibration material and application to West African kimberlites. *Chem. Geol.*
615 649, 121982. <https://doi.org/10.1016/j.chemgeo.2024.121982>
- 616 Gostin, V.A., Haines, P.W., Jenkins, R.J.F., Compston, W., Williams, I.S., 1986. Impact Ejecta Horizon
617 Within Late Precambrian Shales, Adelaide Geosyncline, South Australia. *Science* (1979). 233, 198–
618 200. <https://doi.org/10.1126/science.233.4760.198>
- 619 Guerrero, D., Reimold, W.U., Hauser, N., Kenny, G., Whitehouse, M., Lambert, P., 2025. Shock
620 deformation and U-Pb isotope systematics in zircon from impactites of the Rochechouart impact
621 structure: Impact age and zircon provenance. *Geochim. Cosmochim. Acta* 402, 338–358.
622 <https://doi.org/10.1016/j.gca.2025.05.049>
- 623 Hayward, C.L., Reimold, W.U., Gibson, R.L., Robb, L.J., 2005. Gold mineralization within the
624 Witwatersrand Basin, South Africa: evidence for a modified placer origin, and the role of the
625 Vredefort impact event. *Geological Society, London, Special Publications* 248, 31–58.
626 <https://doi.org/10.1144/GSL.SP.2005.248.01.02>
- 627 Herrmann, M., Kenny, G.G., Martell, J., Whitehouse, M.J., Alwmark, C., 2024. The first U–Pb age for
628 shocked zircon from the Mien impact structure, Sweden, and implications for metamictization-
629 induced zircon texture formed during impact events. *Meteorit. Planet. Sci.* 59, 211–241.
630 <https://doi.org/10.1111/maps.14116>
- 631 Hildebrand, A.R., Penfield, G.T., Kring, D.A., Pilkington, M., Camargo Z., A., Jacobsen, S.B., Boynton, W.
632 V., 1991. Chicxulub Crater: A possible Cretaceous/Tertiary boundary impact crater on the Yucatán
633 Peninsula, Mexico. *Geology* 19, 867. [https://doi.org/10.1130/0091-
634 7613\(1991\)019<0867:CCAPCT>2.3.CO;2](https://doi.org/10.1130/0091-7613(1991)019<0867:CCAPCT>2.3.CO;2)
- 635 Hill, A.C., Haines, P.W., Grey, K., Willman, S., 2007. New records of Ediacaran Acraman ejecta in
636 drillholes from the Stuart Shelf and Officer Basin, South Australia. *Meteorit. Planet. Sci.* 42, 1883–
637 1891. <https://doi.org/10.1111/j.1945-5100.2007.tb00547.x>
- 638 Hogmalm, K.J., Zack, T., Karlsson, A.K., Sjoqvist, A.S.L., Grabe-Schonberg, D., 2017. In-situ Rb-Sr and K-Ca
639 dating by LA-ICP-MS/MS: An evaluation of N₂O and SF₆ as reaction gases. *J. Anal. At. Spectrom.* 32,
640 305–313.
- 641 Horstwood, M.S.A., Košler, J., Gehrels, G., Jackson, S.E., McLean, N.M., Paton, C., Pearson, N.J.,
642 Sircombe, K., Sylvester, P., Vermeesch, P., Bowring, J.F., Condon, D.J., Schoene, B., 2016.



- 643 Community-Derived Standards for LA-ICP-MS U-(Th)-Pb Geochronology – Uncertainty Propagation,
644 Age Interpretation and Data Reporting. *Geostand. Geoanal. Res.* 40, 311–332.
645 <https://doi.org/10.1111/j.1751-908X.2016.00379.x>
- 646 Hrstka, T., Gottlieb, P., Skála, R., Breiter, K., Motl, D., 2018. Automated mineralogy and petrology -
647 applications of TESCAN Integrated Mineral Analyzer (TIMA). *Journal of Geosciences* 47–63.
648 <https://doi.org/10.3190/jgeosci.250>
- 649 Hu, Z., Gao, S., Liu, Y., Hu, S., Chen, H., Yuan, H., 2008. Signal enhancement in laser ablation ICP-MS by
650 addition of nitrogen in the central channel gas. *J. Anal. At. Spectrom.* 23, 1093.
651 <https://doi.org/10.1039/b804760j>
- 652 Huang, C., Wang, H., Xie, L., Xu, L., Wu, S., Yang, Y., Yang, J., 2025. High-precision Rb-Sr isotope analysis
653 with Neoma MS/MS: Enhancing in situ geochronology by laser ablation. *Spectrochim. Acta Part B*
654 *At. Spectrosc.* 224, 107117. <https://doi.org/10.1016/j.sab.2025.107117>
- 655 Jackson, S.E., Pearson, N.J., Griffin, W.L., Belousova, E.A., 2004. The application of laser ablation-
656 inductively coupled plasma-mass spectrometry to in situ U–Pb zircon geochronology. *Chem. Geol.*
657 211, 47–69.
- 658 Jourdan, F., 2012. The $^{40}\text{Ar}/^{39}\text{Ar}$ dating technique applied to planetary sciences and terrestrial impacts.
659 *Australian Journal of Earth Sciences* 59, 199–224. <https://doi.org/10.1080/08120099.2012.644404>
- 660 Jourdan, F., Reimold, W.U., Deutsch, A., 2012. Dating Terrestrial Impact Structures. *Elements* 8, 49–53.
661 <https://doi.org/10.2113/gselements.8.1.49>
- 662 Jourdan, F., Renne, P.R., Reimold, W.U., 2009. An appraisal of the ages of terrestrial impact structures.
663 *Earth Planet. Sci. Lett.* 286, 1–13. <https://doi.org/10.1016/j.epsl.2009.07.009>
- 664 Jourdan, F., Renne, P.R., Reimold, W.U., 2007. The problem of inherited $^{40}\text{Ar}^*$ in dating impact glass by
665 the $^{40}\text{Ar}/^{39}\text{Ar}$ method: Evidence from the Tswaing impact crater (South Africa). *Geochim.*
666 *Cosmochim. Acta* 71, 1214–1231. <https://doi.org/10.1016/j.gca.2006.11.013>
- 667 Kellett, D.A., Larson, K.P., Skipton, D.R., 2024. Integration of white mica in situ $^{87}\text{Rb}/^{87}\text{Sr}$ with in situ
668 and step-heat $^{40}\text{Ar}/^{39}\text{Ar}$ dates in orogenic settings. *Lithos* 482–483, 107687.
669 <https://doi.org/10.1016/j.lithos.2024.107687>
- 670 Kenny, G.G., Schmieder, M., Whitehouse, M.J., Nemchin, A.A., Morales, L.F.G., Buchner, E., Bellucci, J.J.,
671 Snape, J.F., 2019. A new U-Pb age for shock-recrystallised zircon from the Lappajärvi impact crater,
672 Finland, and implications for the accurate dating of impact events. *Geochim. Cosmochim. Acta* 245,
673 479–494. <https://doi.org/10.1016/j.gca.2018.11.021>
- 674 Kirkland, C.L., Erickson, T.M., Johnson, T.E., Prave, T., Strachan, R.A., Daggitt, M.L., Hartnady, M.I.H.,
675 Reddy, S.M., Taylor, R.J.M., Ribeiro, B.V., Rankenburg, K., Kirscher, U., Fischer, S., 2025. A one-
676 billion-year-old Scottish meteorite impact. *Geology* 53, 621–625.
677 <https://doi.org/10.1130/G53121.1>
- 678 Kirkland, C.L., Olierook, H.K.H., Danišik, M., Liebmann, J., Hollis, J., Ribeiro, B. V., Rankenburg, K., 2023.
679 Dating mylonitic overprinting of ancient rocks. *Commun. Earth Environ.* 4, 47.
680 <https://doi.org/10.1038/s43247-023-00709-5>
- 681 Larson, K.P., Button, M., Shrestha, S., Camacho, A., 2023. A comparison of $^{87}\text{Rb}/^{87}\text{Sr}$ and $^{40}\text{Ar}/^{39}\text{Ar}$
682 dates: Evaluating the problem of excess ^{40}Ar in Himalayan mica. *Earth Planet. Sci. Lett.* 609,
683 118058. <https://doi.org/10.1016/j.epsl.2023.118058>



- 684 Li, S.S., Santosh, M., Farkaš, J., Redaa, A., Ganguly, S., Kim, S.W., Zhang, C., Gilbert, S., Zack, T., 2020.
685 Coupled U-Pb and Rb-Sr laser ablation geochronology trace Archean to Proterozoic crustal
686 evolution in the Dharwar Craton, India. *Precambrian Res.* 343.
687 <https://doi.org/10.1016/j.precamres.2020.105709>
- 688 Liebmann, J., Kirkland, C.L., Kelsey, D.E., Korhonen, F.J., Rankenburg, K., 2022. Lithological fabric as a
689 proxy for Rb–Sr isotopic complexity. *Chem. Geol.* 608, 121041.
690 <https://doi.org/10.1016/j.chemgeo.2022.121041>
- 691 Luu, T.-H., Peters, D., Gérard, Y., Crawshaw, H., Moynier, F., 2026. Precise in situ ⁸⁷Rb–⁸⁷Sr dating of
692 terrestrial and extraterrestrial samples with Nu Sapphire XD LA-CRC-MC-ICP-MS/MS. *J. Anal. At.*
693 *Spectrom.* <https://doi.org/10.1039/D6JA00061D>
- 694 McCall, G.J.H., 2006. The Vendian (Ediacaran) in the geological record: Enigmas in geology’s prelude to
695 the Cambrian explosion. *Earth. Sci. Rev.* 77, 1–229.
696 <https://doi.org/10.1016/j.earscirev.2005.08.004>
- 697 McGregor, M., McFarlane, C.R.M., Spray, J.G., 2019. In situ multiphase U–Pb geochronology and shock
698 analysis of apatite, titanite and zircon from the Lac La Moinerie impact structure, Canada.
699 *Contributions to Mineralogy and Petrology* 174, 62. <https://doi.org/10.1007/s00410-019-1598-y>
- 700 Milton, D.J., Glikson, A.Y., Manwaring, E.A., Moss, F.J., Sedmik, E.C.E., Van Son, J., Barlow, B.C., Brett, R.,
701 Young, G.A., 1972. Gosses Bluff Impact Structure, Australia. *Science* (1979). 175, 1199–1207.
702 <https://doi.org/10.1126/science.175.4027.1199>
- 703 Milton, D.J., Sutter, J.F., 1987. Revised age for the Gosses Bluff impact structure, Northern Territory,
704 Australia, based on ⁴⁰Ar/³⁹Ar dating. *Meteoritics* 22, 281–289. [https://doi.org/10.1111/j.1945-](https://doi.org/10.1111/j.1945-5100.1987.tb00625.x)
705 [5100.1987.tb00625.x](https://doi.org/10.1111/j.1945-5100.1987.tb00625.x)
- 706 Montel, J.-M., Foret, S., Veschambre, M., Nicollet, C., Provost, A., 1996. Electron microprobe dating of
707 monazite. *Chem. Geol.* 131, 37–53. [https://doi.org/10.1016/0009-2541\(96\)00024-1](https://doi.org/10.1016/0009-2541(96)00024-1)
- 708 Nicomedes, I.S., Ribeiro, B. V., Caxito, F.A., Kirkland, C.L., Victoria, A., Quadir, Z., 2025. In situ Rb–Sr
709 geochronology of slickensides reveals reactivation of cratonic margins post-Gondwana assembly. *J.*
710 *Struct. Geol.* 191, 105305. <https://doi.org/10.1016/j.jsg.2024.105305>
- 711 Nie, N.X., Dauphas, N., Zhang, Z.J., Hopp, T., Sarantos, M., 2024. Lunar soil record of atmosphere loss
712 over eons. *Sci. Adv.* 10. <https://doi.org/10.1126/sciadv.adm7074>
- 713 Osinski, G.R., Cockell, C.S., Pontefract, A., Sapers, H.M., 2020. The Role of Meteorite Impacts in the
714 Origin of Life. *Astrobiology* 20, 1121–1149. <https://doi.org/10.1089/ast.2019.2203>
- 715 Osinski, G.R., Grieve, R.A.F., Ferrière, L., Losiak, A., Pickersgill, A.E., Cavosie, A.J., Hibbard, S.M., Hill,
716 P.J.A., Bermudez, J.J., Marion, C.L., Newman, J.D., Simpson, S.L., 2022. Impact Earth: A review of
717 the terrestrial impact record. *Earth. Sci. Rev.* 232, 104112.
718 <https://doi.org/10.1016/j.earscirev.2022.104112>
- 719 Papanastassiou, D.A., Wasserburg, G.J., 1972a. The Rb-Sr age of a crystalline rock from Apollo 16. *Earth*
720 *Planet. Sci. Lett.* 16, 289–298. [https://doi.org/10.1016/0012-821X\(72\)90205-1](https://doi.org/10.1016/0012-821X(72)90205-1)
- 721 Papanastassiou, D.A., Wasserburg, G.J., 1972b. Rb-Sr age of a Luna 16 basalt and the model age of lunar
722 soils. *Earth Planet. Sci. Lett.* 13, 368–374. [https://doi.org/10.1016/0012-821X\(72\)90111-2](https://doi.org/10.1016/0012-821X(72)90111-2)
- 723 Paton, C., Hellstrom, J., Paul, B., Woodhead, J., Hergt, J., 2011a. *lolite*: Freeware for the visualisation and
724 processing of mass spectrometric data. *J. Anal. At. Spectrom.* 26, 2508.
725 <https://doi.org/10.1039/c1ja10172b>



- 726 Paton, C., Hellstrom, J., Paul, B., Woodhead, J., Hergt, J., 2011b. lolite: Freeware for the visualisation and
727 processing of mass spectrometric data. *J. Anal. At. Spectrom.* 26, 2508–2518.
- 728 Quintero, R.R., Cavosie, A.J., Cox, M.A., Miljković, K., Dugdale, A., 2021. Australian impact cratering
729 record: Updates and recent discoveries, in: *Large Meteorite Impacts and Planetary Evolution VI*.
730 Geological Society of America, pp. 41–68. [https://doi.org/10.1130/2021.2550\(02\)](https://doi.org/10.1130/2021.2550(02))
- 731 Rankenburg, K., Ribeiro, B.V., Evans, N.J., McDonald, B., 2024. Comparing CRC-MC-ICPMS to QQQ-
732 ICPMS/MS for in situ Rb-Sr applications, in: *Goldschmidt*. Chicago.
- 733 Redaa, A., Farkaš, J., Gilbert, S., Collins, A.S., Wade, B., Löhr, S., Zack, T., Garbe-Schönberg, D., 2021a.
734 Assessment of elemental fractionation and matrix effects during *in situ* Rb–Sr dating of phlogopite
735 by LA-ICP-MS/MS: implications for the accuracy and precision of mineral ages. *J. Anal. At.*
736 *Spectrom.* 36, 322–344. <https://doi.org/10.1039/D0JA00299B>
- 737 Redaa, A., Farkaš, J., Gilbert, S., Collins, A.S., Wade, B., Löhr, S., Zack, T., Garbe-Schönberg, D., 2021b.
738 Assessment of elemental fractionation and matrix effects during *in situ* Rb–Sr dating of phlogopite
739 by LA-ICP-MS/MS: implications for the accuracy and precision of mineral ages. *J. Anal. At.*
740 *Spectrom.* 36, 322–344. <https://doi.org/10.1039/D0JA00299B>
- 741 Reimold, W.U., Koeberl, C., Gibson, R.L., Dressler, B.O., 2005. Economic Mineral Deposits in Impact
742 Structures: A Review. *Impact Tectonics* 479–552. https://doi.org/10.1007/3-540-27548-7_20
- 743 Ribeiro, B.V., Finch, M.A., Cawood, P.A., Faleiros, F.M., Murphy, T.D., Simpson, A., Glorie, S., Tedeschi,
744 M., Armit, R., Barrote, V.R., 2022. From microanalysis to supercontinents: Insights from the Rio Apa
745 Terrane into the Mesoproterozoic SW Amazonian Craton evolution during Rodinia assembly.
746 *Journal of Metamorphic Geology* 40, 631–663. <https://doi.org/10.1111/jmg.12641>
- 747 Ribeiro, B.V., Jakobsson, V., Kirkland, C.L., Spikings, R.A., Zametzer, A., Rickard, W.D.A., Centrella, S.,
748 Timms, N.E., Cavosie, A.J., Danišić, M., 2026. Microstructural and geochemical response of
749 muscovite to impact metamorphism. *Geology* 54, 179–183. <https://doi.org/10.1130/G53895.1>
- 750 Ribeiro, B. V., Kirkland, C.L., Finch, M.A., Faleiros, F.M., Reddy, S.M., Rickard, W.D.A., Hartnady, M.I.H.,
751 2023a. Microstructures, geochemistry, and geochronology of mica fish: Review and advances. *J.*
752 *Struct. Geol.* 104947. <https://doi.org/10.1016/j.jsg.2023.104947>
- 753 Ribeiro, B. V., Kirkland, C.L., Kelsey, D.E., Reddy, S.M., Hartnady, M.I.H., Faleiros, F.M., Rankenburg, K.,
754 Liebmann, J., Korhonen, F.J., Clark, C., 2023b. Time-strain evolution of shear zones from
755 petrographically constrained Rb–Sr muscovite analysis. *Earth Planet. Sci. Lett.* 602, 117969.
756 <https://doi.org/10.1016/j.epsl.2022.117969>
- 757 Rösel, D., Zack, T., 2022. LA-ICP-MS/MS Single-Spot Rb-Sr Dating. *Geostand. Geoanal. Res.*
758 <https://doi.org/10.1111/ggr.12414>
- 759 Schmieder, M., Tohver, E., Jourdan, F., Denyszyn, S.W., Haines, P.W., 2015. Zircons from the Acraman
760 impact melt rock (South Australia): Shock metamorphism, U–Pb and $40\text{ Ar}/39\text{ Ar}$ systematics, and
761 implications for the isotopic dating of impact events. *Geochim. Cosmochim. Acta* 161, 71–100.
762 <https://doi.org/10.1016/j.gca.2015.04.021>
- 763 Schmitz, M.D., Bowring, S.A., Ireland, T.R., 2003. Evaluation of Duluth Complex anorthositic series (AS3)
764 zircon as a U–Pb geochronological standard: new high-precision isotope dilution thermal ionization
765 mass spectrometry results. *Geochim. Cosmochim. Acta* 67, 3665–3672.
766 [https://doi.org/10.1016/S0016-7037\(03\)00200-X](https://doi.org/10.1016/S0016-7037(03)00200-X)



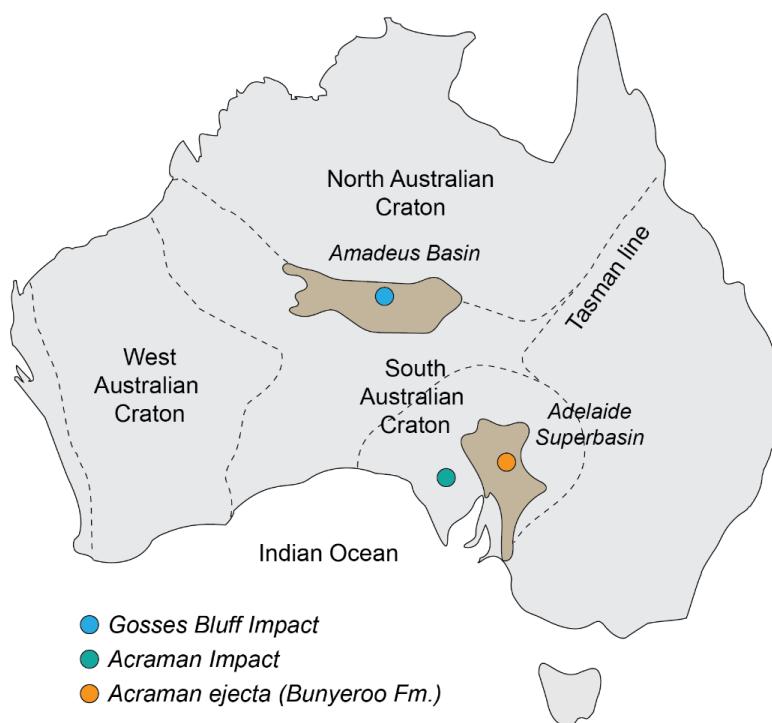
- 767 Schoene, B., Bowring, S.A., 2006. U-Pb systematics of the McClure Mountain syenite:
768 Thermochronological constraints on the age of the $^{40}\text{Ar}/^{39}\text{Ar}$ standard MMhb. *Contributions to*
769 *Mineralogy and Petrology* 151, 615–630. <https://doi.org/10.1007/s00410-006-0077-4>
- 770 Schulte, P., Alegret, L., Arenillas, I., Arz, J.A., Barton, P.J., Bown, P.R., Bralower, T.J., Christeson, G.L.,
771 Claeys, P., Cockell, C.S., Collins, G.S., Deutsch, A., Goldin, T.J., Goto, K., Grajales-Nishimura, J.M.,
772 Grieve, R.A.F., Gulick, S.P.S., Johnson, K.R., Kiessling, W., Koeberl, C., Kring, D.A., MacLeod, K.G.,
773 Matsui, T., Melosh, J., Montanari, A., Morgan, J. V., Neal, C.R., Nichols, D.J., Norris, R.D., Pierazzo,
774 E., Ravizza, G., Rebolledo-Vieyra, M., Reimold, W.U., Robin, E., Salge, T., Speijer, R.P., Sweet, A.R.,
775 Urrutia-Fucugauchi, J., Vajda, V., Whalen, M.T., Willumsen, P.S., 2010. The Chicxulub Asteroid
776 Impact and Mass Extinction at the Cretaceous-Paleogene Boundary. *Science* (1979). 327, 1214–
777 1218. <https://doi.org/10.1126/science.1177265>
- 778 Slama, J., Kosler, J., Condon, D.J., Crowley, J.L., Gerdes, A., Hanchar, J. M. Horstwood, M. S. A. Morris,
779 G.A., Nasdala, L., Norberg, N., Schaltegger, U. Schoene, N., Tubrett, M. N. Whitehouse, M.J., 2008.
780 Plesovice zircon - a new natural reference material for U-Pb and Hf isotopic microanalysis. *Chem.*
781 *Geol.* 249, 1–35.
- 782 Stacey, J.S., Kramers, J.D., 1975. Approximation of terrestrial lead isotope evolution by a two-stage
783 model. *Earth Planet. Sci. Lett.* 26, 207–221. [https://doi.org/10.1016/0012-821X\(75\)90088-6](https://doi.org/10.1016/0012-821X(75)90088-6)
- 784 Subarkah, D., Blades, M.L., Collins, A.S., Farkaš, J., Gilbert, S., Löhr, S.C., Redaa, A., Cassidy, E., Zack, T.,
785 2022. Unraveling the histories of Proterozoic shales through *in situ* Rb-Sr dating and trace element
786 laser ablation analysis. *Geology* 50, 66–70. <https://doi.org/10.1130/G49187.1>
- 787 Thomson, S.N., Gehrels, G.E., Ruiz, J., Buchwaldt, R., 2012. Routine low-damage apatite U-Pb dating
788 using laser ablation–multicollector–ICPMS. *Geochemistry, Geophysics, Geosystems* 13.
- 789 Tillberg, M., Drake, H., Zack, T., Kooijman, E., Whitehouse, M.J., Åström, M.E., 2020. *In situ* Rb-Sr dating
790 of slickenfibres in deep crystalline basement faults. *Sci. Rep.* 10, 562.
791 <https://doi.org/10.1038/s41598-019-57262-5>
- 792 Timms, N.E., Erickson, T.M., Pearce, M.A., Cavosie, A.J., Schmieder, M., Tohver, E., Reddy, S.M., Zanetti,
793 M.R., Nemchin, A.A., Wittmann, A., 2017. A pressure-temperature phase diagram for zircon at
794 extreme conditions. *Earth. Sci. Rev.* 165, 185–202. <https://doi.org/10.1016/j.earscirev.2016.12.008>
- 795 Tollefson, K., Larson, K., 2026. Investigation of the effects of ductile deformation on *in situ* biotite and
796 muscovite Rb-Sr geochronology. *Geol. Mag.* 163, e22.
797 <https://doi.org/10.1017/S0016756826100685>
- 798 Tollefson, K.T., Larson, K.P., Moynihan, D.P., Cleven, N.R., 2025. *In situ* mica Rb–Sr geochronology and
799 microstructural analysis to evaluate the timing and sense of shear zone movement: Shannon and
800 d’Abbadie Faults, Yukon, Canada. *Can. J. Earth Sci.* 62, 1320–1341. <https://doi.org/10.1139/cjes-2024-0144>
- 801
- 802 Tuchscherer, M.G., Spray, J.G., 2002. Geology, Mineralization, and Emplacement of the Foy Offset Dike,
803 Sudbury Impact Structure*. *Economic Geology* 97, 1377–1397.
804 <https://doi.org/10.2113/gsecongeo.97.7.1377>
- 805 Vermeesch, P., 2018. IsoplotR: A free and open toolbox for geochronology. *Geoscience Frontiers* 2 9,
806 1479–1493.
- 807 Villa, I.M., De Bièvre, P., Holden, N.E., Renne, P.R., 2015. IUPAC-IUGS recommendation on the half life of
808 ^{87}Rb . *Geochim. Cosmochim. Acta* 164, 382–385. <https://doi.org/10.1016/j.gca.2015.05.025>



- 809 Webb, A.W., Coats, R.P., Fanning, C.M., Flint, R.B., 1983. Geochronological framework of the Adelaide
810 Geosyncline, in: Geological Society of Australia Abstracts.
- 811 Williams, G.E., 1992. Acraman: A major impact structure from the Neoproterozoic of Australia. pp. 209–
812 224. <https://doi.org/10.1130/SPE293-p209>
- 813 Williams, G.E., 1986. The Acraman Impact Structure: Source of Ejecta in Late Precambrian Shales, South
814 Australia. *Science* (1979). 233, 200–203. <https://doi.org/10.1126/science.233.4760.200>
- 815 Williams, G.E., Schmidt, P.W., 2021. Dating the Acraman asteroid impact, South Australia: the case for
816 deep drilling the ‘hot shock’ zone of the central uplift. *Australian Journal of Earth Sciences* 68, 355–
817 367. <https://doi.org/10.1080/08120099.2020.1808066>
- 818 Williams, George E., Wallace, Malcolm W., 2003. The Acraman asteroid impact, South Australia:
819 magnitude and implications for the late Vendian environment. *J. Geol. Soc. London*. 160, 545–554.
820 <https://doi.org/10.1144/0016-764902-142>
- 821 Woodhead, J.D., Hergt, J.M., 2005. A Preliminary Appraisal of Seven Natural Zircon Reference Materials
822 for In Situ Hf Isotope Determination. *Geostand. Geoanal. Res.* 29, 183–195.
- 823 Woodhead, J.D., Hergt, J.M., 2001. Strontium, Neodymium and Lead Isotope Analyses of NIST Glass
824 Certified Reference Materials: SRM 610, 612, 614. *Geostand. Geoanal. Res.* 25, 261–266.
825 <https://doi.org/10.1111/j.1751-908X.2001.tb00601.x>
- 826 Zack, T., Hogmalm, K.J., 2016. Laser ablation Rb/Sr dating by online chemical separation of Rb and Sr in
827 an oxygen-filled reaction cell. *Chem. Geol.* 437, 120–133.
828 <https://doi.org/10.1016/j.chemgeo.2016.05.027>
- 829 Zametzer, A., Kirkland, C.L., Barham, M., Hartnady, M.I.H., Bath, A.B., Rankenburg, K., 2022. Episodic
830 alteration within a gold-bearing Archean shear zone revealed by in situ biotite Rb–Sr dating.
831 *Precambrian Res.* 382, 106872. <https://doi.org/10.1016/j.precamres.2022.106872>
832
- 833
- 834
- 835
- 836
- 837
- 838
- 839
- 840
- 841
- 842



843 **FIGURES**

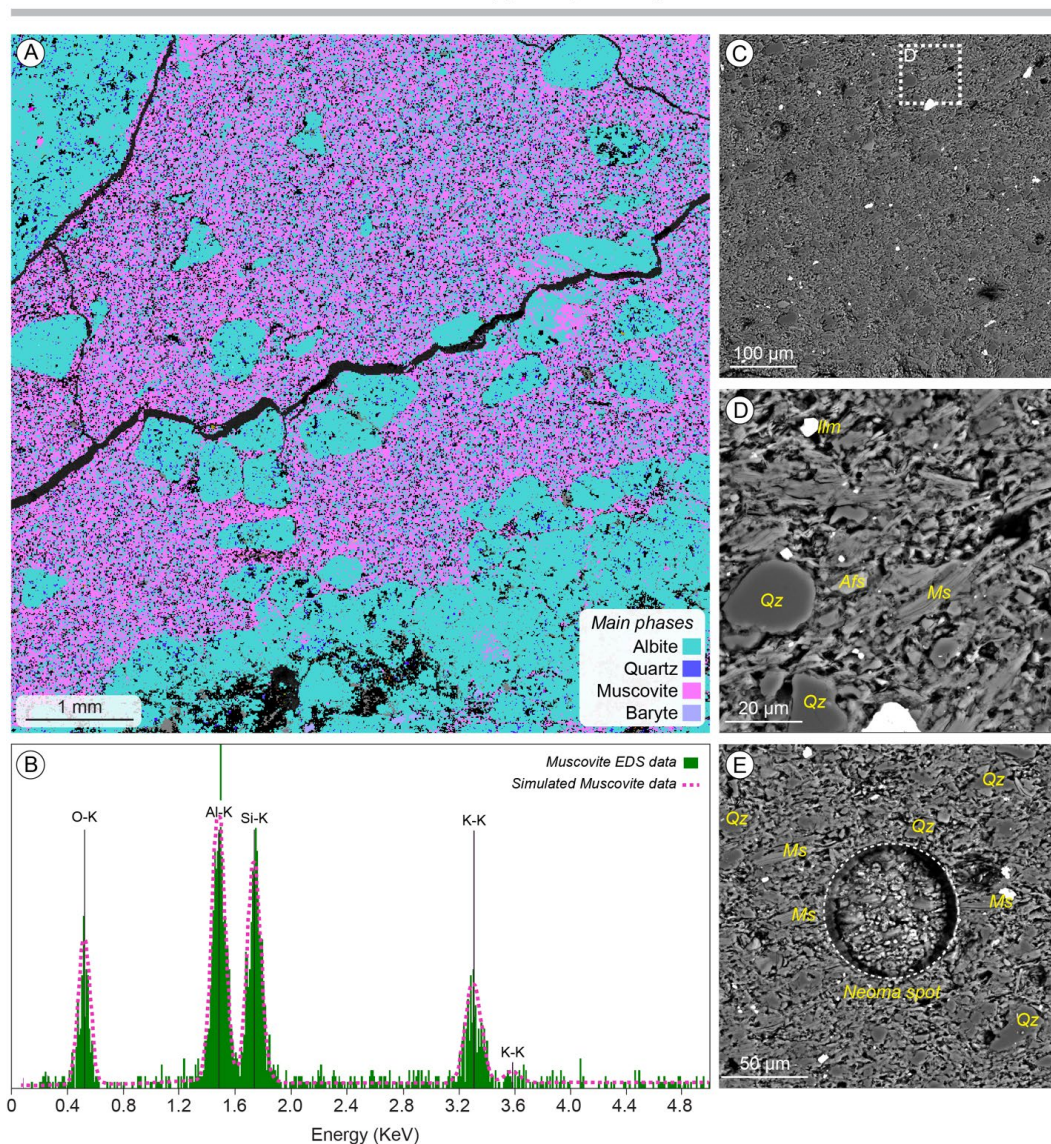


844

845 Figure 1. Geological context of the Gosses Bluff and Acraman impact sites in respect to the
846 Australian main geological provinces, including the extent of the ejecta layer of the Acraman
847 impact (Bunyerroo Formation).

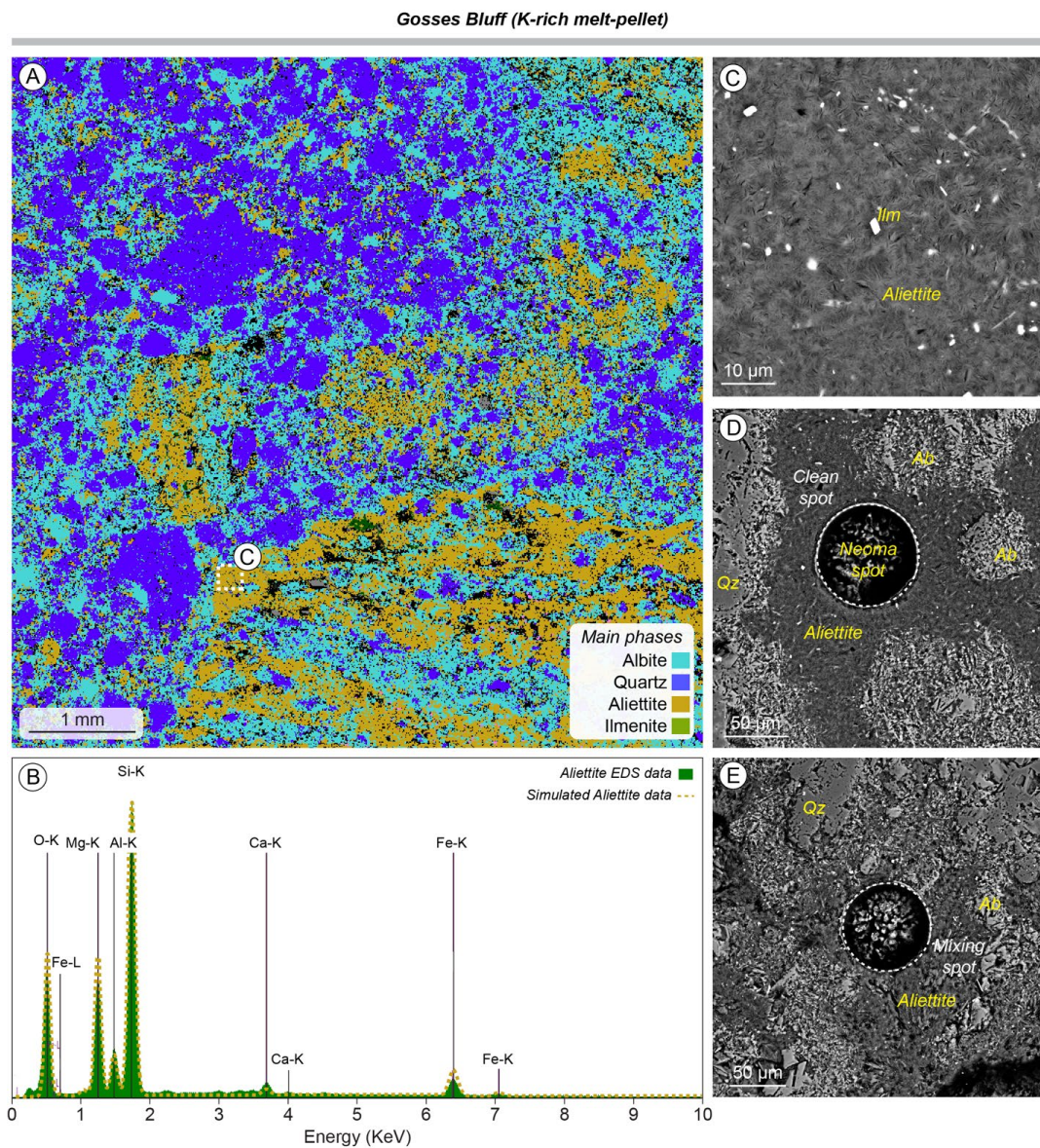


Acraman (ejecta layer matrix)



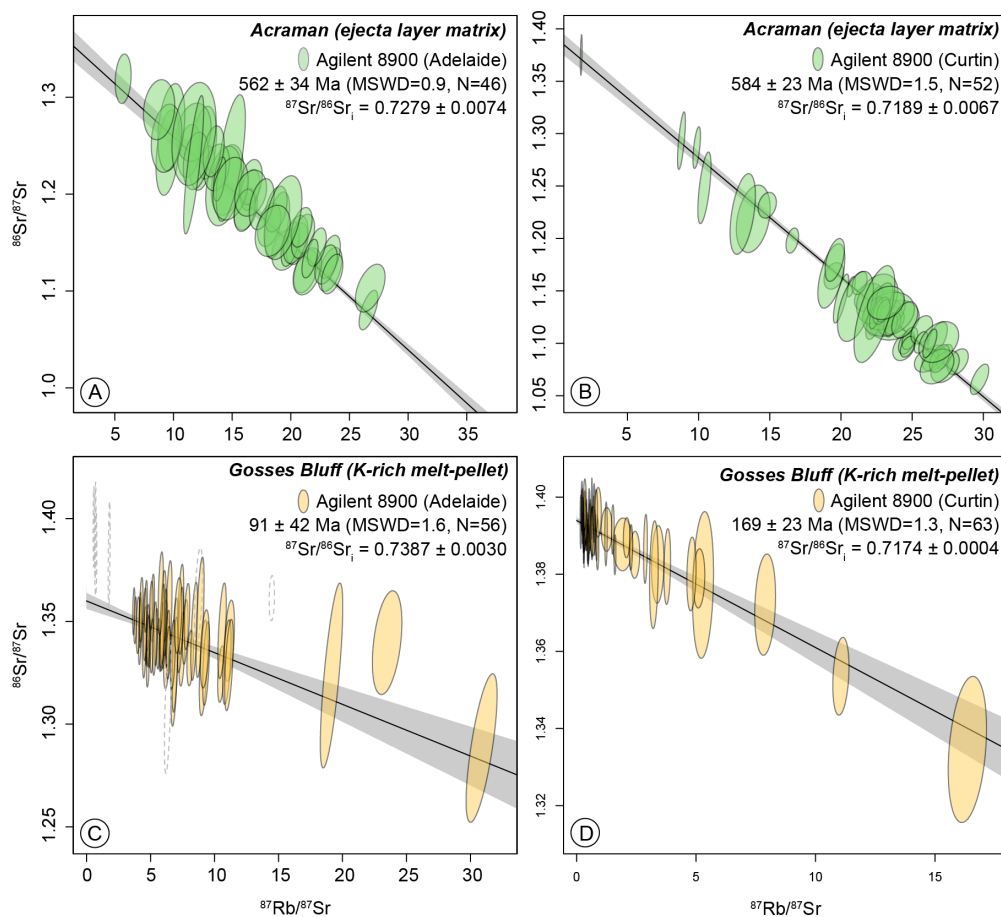
848

849 Figure 2. A) TIMA-based mineral phase map of the Acraman sample; B) Muscovite EDS spectra
850 compared to simulated data, highlighting the peaks of major components (K-energy level); C-E)
851 electron backscattered photomicrographs of the fine-grained muscovite domain. The laser-
852 Neoma spot in panels (D) and (E) are 90 μm in diameter. Mineral abbreviations follow Whitney
853 and Evans (2010).



854

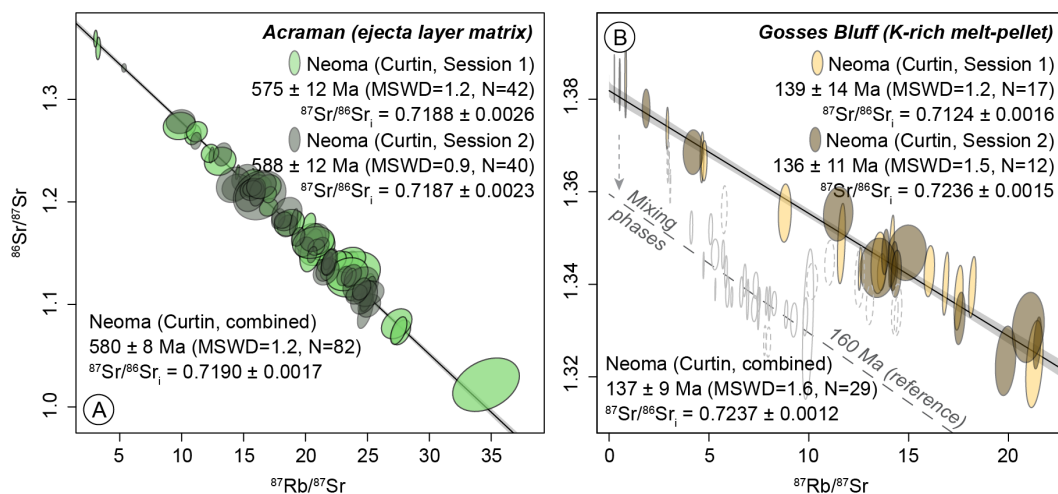
855 Figure 3. A) TIMA-based mineral phase map of the Gosses Bluff sample; B) Aliettite EDS spectra
856 compared to simulated data, highlighting the peaks of major components (K-energy level); C-E)
857 electron backscattered photomicrographs of the fine-grained aliettite domains. The laser-Neoma
858 spot in panels (D) and (E) are 90 µm in diameter. Mineral abbreviations follow Whitney and Evans
859 (2010).



860

861 Figure 4. In situ single-collector Rb-Sr data from the Acraman (A, B) and Gosses Bluff (C, D)
862 impact samples. Left panels display the Rb–Sr data obtained at the University of Adelaide, whilst
863 the right panels display the data obtained at Curtin University. Uncertainties are presented at 2-
864 standard error.

865



866

867 Figure 5. In situ multi-collector Rb-Sr data from the Acraman and Gosses Bluff impact samples.

868 Uncertainties are presented at 2-standard error.

869

870

871

872

873

874

875

876

877

878

879

880

881

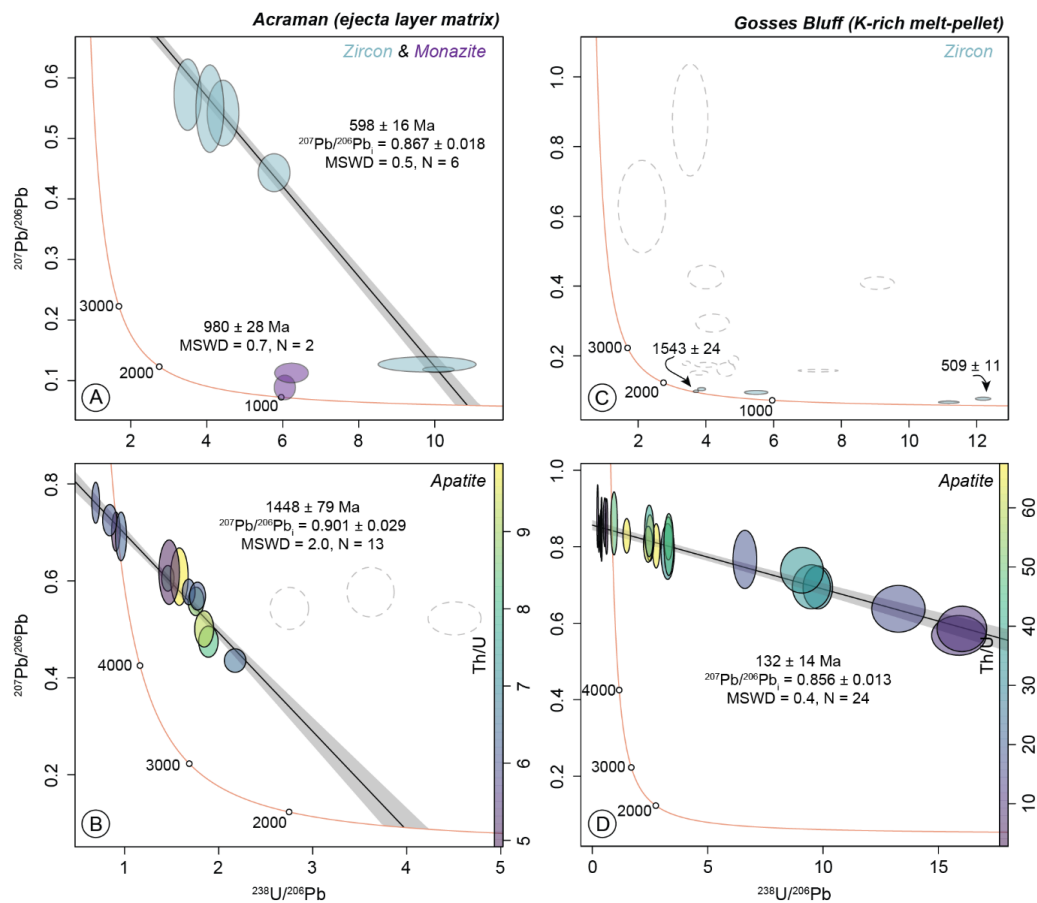
882

883



884

885



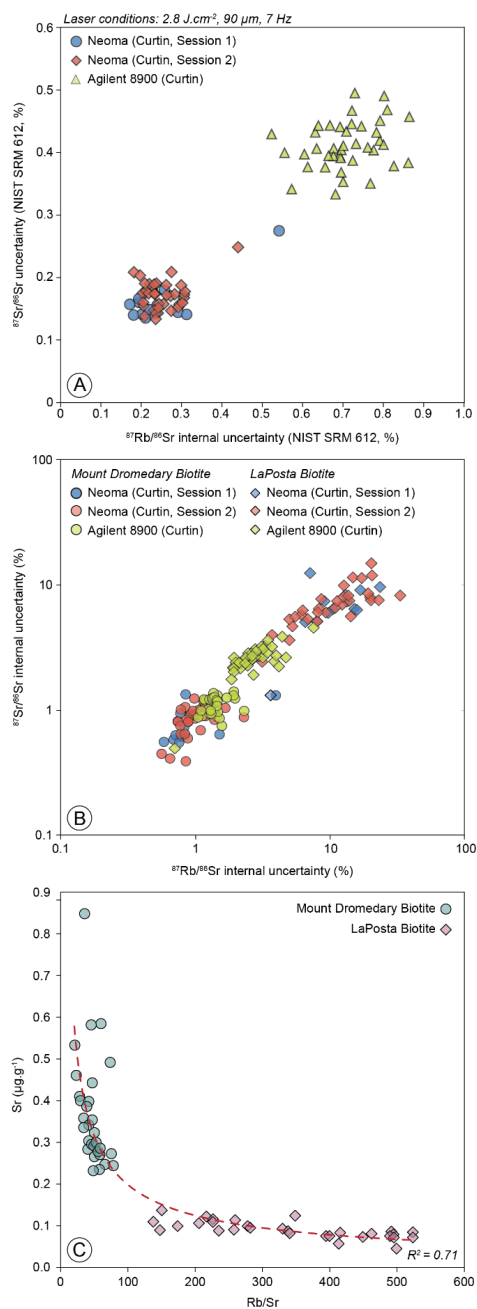
886

887 Figure 6. A) Monazite and zircon, and B) apatite U–Pb data from the Acraman ejecta layer. C)

888 Zircon and (D) apatite U–Pb data from the Gosses Bluff K-rich melt pellet. Uncertainties are

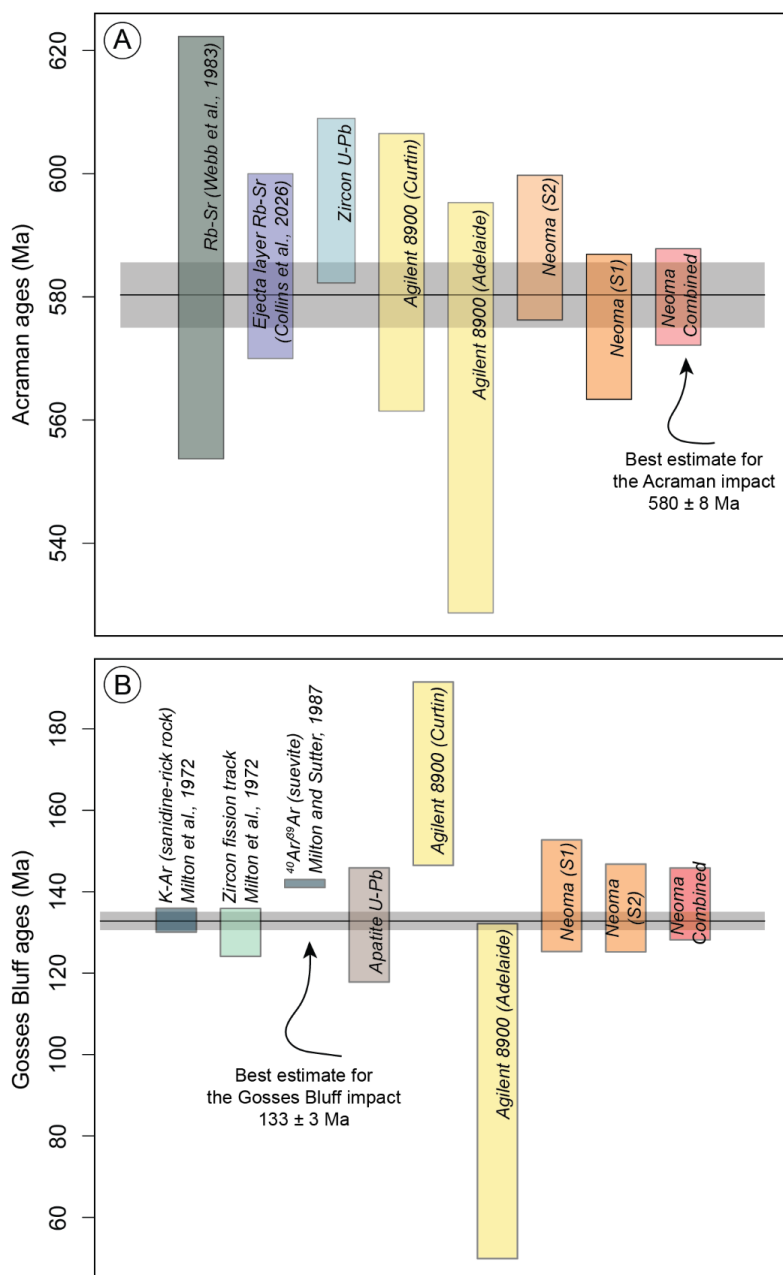
889 presented at 2-standard error.

890



891

892 Figure 7. Analytical comparison between single- and multi-collector mass spectrometers,
893 including (A) the measurement of $^{87}\text{Sr}/^{86}\text{Sr}$ and $^{87}\text{Sr}/^{86}\text{Sr}$ measurements on NIST SRM 612, and
894 (B) natural mica reference materials (Mount Dromedary and LaPosta biotite).



895

896 Figure 8. Comparison between new ages from the Acraman (A) and Gosses Bluff (B) impact and
 897 available literature data from both sites (Milton et al., 1972; Webb et al., 1983; Milton and Sutter,
 898 1987; Collins et al., 2026).