



1 **Anthropogenic CO₂ Emissions in China Constrained by**
2 **OCO-2/3 XCO₂ Observations**

3
4 Shenpeng Qiu¹, Shuzhuang Feng^{1,2*}, Fei Jiang^{1,2}, Chong Wei³, Yusheng Shi⁴, Mengwei Jia^{1,2},
5 Honglin Zhuang¹, Hengmao Wang^{1,2}, Yongguang Zhang^{1,2}, Weimin Ju^{1,2}

6 ¹Jiangsu Provincial Key Laboratory for Advanced Remote Sensing and Geographic Information
7 Technology, International Institute for Earth System Science, Nanjing University, Nanjing, 210023,
8 China

9 ²Jiangsu Center for Collaborative Innovation in Geographical Information Resource Development
10 and Application, Nanjing, 210023, China

11 ³Shanghai Carbon Data Research Center, CAS Key Laboratory of Low-Carbon Conversion Science
12 and Engineering, Shanghai Advanced Research Institute, Chinese Academy of Sciences, Shanghai,
13 201210, China

14 ⁴State Key Laboratory of Remote Sensing Science, Aerospace Information Research Institute,
15 Chinese Academy of Sciences, Beijing, 100101, China

16

17

18 * Corresponding author: Prof. Shuzhuang Feng (fengsz@nju.edu.cn)



Abstract: Accurately quantifying anthropogenic CO₂ emissions is essential for evaluating carbon budget and mitigation strategies. However, traditional "bottom-up" emission inventories suffer from substantial uncertainties and update time lags, urgently requiring top-down constraints from atmospheric observations while accounting for confounding terrestrial biogenic interferences. In this study, we extended RegGCAS, a regional carbon assimilation system based on the WRF-CMAQ atmospheric chemical transport model and the Ensemble Kalman Filter algorithm. By assimilating column-averaged dry-air CO₂ mole fractions (XCO₂) from OCO-2/3 satellite observations, we inverted anthropogenic CO₂ emissions over mainland China during winter 2022-2023. The results revealed that the total national anthropogenic CO₂ emissions amounted to 2808.3 ± 157.0 Tg, 16.1% higher than the MEIC inventory. For key emission regions, emissions increased by 13.1% in the Beijing-Tianjin-Hebei region, whereas they decreased by 10.4% in the Yangtze River Delta. The system captured distinct urban-suburban emission adjustment differences in key regions, with reductions in city centers and increases in surrounding areas. It also reflected short-term emission fluctuations related to anthropogenic activity changes, such as the Spring Festival work stoppages. Evaluation demonstrates that assimilation effectively reduces prior emission errors by 68.0%. Validation shows that posterior simulation RMSE decrease by 5.8% against the assimilated OCO-2/3 XCO₂, and by 15.3%, 7.7%, and 25.2% against independent TCCON, ObsPack, and urban site observations, respectively, confirming the enhanced accuracy of the posterior emission estimates. This study provides a reliable inversion framework for tracking regional carbon dynamics and refining bottom-up emission inventories.

Keywords: Anthropogenic CO₂ emissions; Emission inversion; OCO-2/3; Data assimilation; RegGCAS

1. Introduction

Atmospheric CO₂ concentrations have increased by approximately 53.2% relative to preindustrial levels, reaching 426.5 ppm in 2025 (WMO, 2025). Anthropogenic CO₂ emissions, primarily from fossil fuel combustion and industrial processes, are the dominant driver of the observed rise in atmospheric CO₂ (Friedlingstein et al., 2024). As the world's largest emitter of anthropogenic CO₂, China urgently needs a scientific, high-spatiotemporal-resolution carbon emission accounting system to achieve its carbon peaking and carbon neutrality goals.



52 At present, the quantification of anthropogenic CO₂ emissions mainly relies on the
53 "bottom-up" emission inventory approach and the "top-down" atmospheric inversion
54 method (Janssens-Maenhout et al., 2020). The former, based on energy activity levels,
55 emission factors, and statistical data, serves as the fundamental method for current
56 emission estimation (IPCC, 2019). However, its estimates are highly susceptible to factors
57 such as statistical data quality, time lags in updates, and spatial disaggregation methods
58 (Hua et al., 2023; Zhang et al., 2022), leading to uncertainties that can reach up to 50% at
59 the provincial scale in China (Yang et al., 2025). In contrast, the "top-down" approach
60 combines prior emissions, CO₂ observations, and atmospheric transport models to
61 constrain surface emission fluxes through inversion or data assimilation, gradually
62 becoming an essential means for the independent evaluation and correction of emission
63 inventories (Friedlingstein et al., 2024).

64 In recent years, observation-constrained inversion studies of anthropogenic CO₂
65 emissions have steadily advanced. Early studies primarily focused on major point sources,
66 such as power plants, as well as high-emission regions including industrial clusters and
67 urban hotspots (Zheng et al., 2020). Since emissions in these areas are highly concentrated,
68 they produce distinct downwind XCO₂ enhancement signals in satellite or airborne
69 measurements, facilitating the estimation of emission intensities when combined with wind
70 field data and plume spatial characteristics (Nassar et al., 2021; Guo et al., 2023; Fuentes
71 Andrade et al., 2024). With the development of observational data and atmospheric
72 transport models (e.g., Lagrangian or Eulerian models), related research has further
73 expanded from specific major sources to the regional scale. Recent studies have utilized
74 satellite XCO₂ observations or ground-based CO₂ measurements to constrain fossil fuel
75 CO₂ emissions in major high-emission urban regions. For example, Lei et al. (2021) used
76 OCO-2 XCO₂ observations and a multi-model approach to estimate fossil fuel CO₂
77 emissions over Lahore, Pakistan, finding that ODIAC likely underestimates local
78 emissions and that robust urban emission trends require more high-quality satellite
79 overpasses. Lian et al. (2023) used continuous ground-based CO₂ observations and a
80 Bayesian inversion to estimate city-scale CO₂ emissions in the Paris metropolitan area,
81 finding that annual emissions declined across the region from 2016 to 2021. However,
82 compared to inversions targeting point sources, large-regional-scale emission inversions
83 remain relatively scarce (Chevallier and Broquet, 2025), and are predominantly focused on
84 quantifying biogenic carbon fluxes (Zhang et al., 2024a). At the regional scale, inversions
85 of anthropogenic CO₂ emissions are susceptible to the combined effects of biogenic carbon



86 fluxes and variations in background CO₂ concentrations (Oney et al., 2017; Schuh et al.,
87 2019). Consequently, effectively separating the anthropogenic emission signals from these
88 interferences remains a critical challenge.

89 Given the common sources of air pollutants and CO₂, some studies have indirectly
90 estimated anthropogenic CO₂ emissions by using observations to invert co-emitted species
91 related to fossil fuel combustion, such as NO_x or CO, and then combining them with CO₂-
92 to-NO_x or CO₂-to-CO emission ratios from prior inventories (Li et al., 2025; Kononov
93 et al., 2016). For example, Feng et al. (2024) utilized surface NO₂ observations to invert
94 fossil fuel CO₂ emissions in China, finding that "bottom-up" emission inventories generally
95 underestimate carbon emissions from the energy and power sectors in western China, while
96 significantly overestimating them in developed regions and key urban areas. Owing to the
97 relatively high observational coverage of proxy gases, this approach presents distinct
98 advantages in regional emission studies, as it can provide supplementary constraints when
99 direct CO₂ observations are insufficient. However, the estimates remain sensitive to
100 uncertainties and time lags in the emission ratios from bottom-up inventories (Juchem et
101 al., 2025).

102 During winter, biogenic carbon fluxes, wildfire carbon emissions, and ocean carbon
103 fluxes are significantly weaker than anthropogenic emissions. Taking the Yangtze River
104 Delta (YRD) as an example, biogenic sources accounted for only 0.2% of anthropogenic
105 emissions during winter (Zhang et al., 2024b). Therefore, winter is a more suitable period
106 for inverting anthropogenic CO₂ emissions to minimize interference from natural source
107 signals. Based on a regional carbon assimilation system RegGCAS (Feng et al., 2025),
108 Zhang et al. (2024b) inverted anthropogenic CO₂ emissions in the YRD during winter by
109 assimilating ground-based CO₂ observations. This approach can effectively reduce prior
110 emission uncertainties and has considerable application potential. However, as ground-
111 based CO₂ observations are sparsely distributed across only a few specific cities, it remains
112 challenging to provide effective constraints on anthropogenic emissions at the national
113 scale.

114 Given the sparse ground-based observations and biogenic signal interference in
115 regional anthropogenic CO₂ inversion, this study integrates OCO-2/3 XCO₂ observations
116 into RegGCAS to invert national anthropogenic CO₂ emissions during winter 2022-2023.
117 This work advances previous RegGCAS applications by expanding the observational
118 constraints from sparse ground-based data to satellite XCO₂, widening the inversion scope
119 from the regional YRD to the national scale, and transitioning from a single-source



120 optimization to a joint inversion of both anthropogenic and biogenic fluxes. The robustness
121 and reliability of the inversion are further evaluated using an Observing System Simulation
122 Experiment (OSSE), independent observations, and sensitivity experiments.

123 **2. Data and Methods**

124 **2.1 Data assimilation system**

125 RegGCAS is a regional carbon assimilation system developed as an extension of the
126 Regional multi-Air Pollutant Assimilation System (RAPAS) (Feng et al., 2023). Previous
127 work with this system has already assimilated ground-based CO₂ observations to optimize
128 anthropogenic CO₂ emissions (Zhang et al., 2024b) and TROPOMI observations to
129 constrain CH₄ emissions (Feng et al., 2025). This study further expanded RegGCAS by
130 incorporating OCO-2/3 XCO₂ observations to optimize regional anthropogenic CO₂
131 emissions. The system comprises a regional chemical transport model (WRF-CMAQ) and
132 an ensemble Kalman filter (EnKF) data assimilation module, used respectively to simulate
133 the spatiotemporal distribution of atmospheric CO₂ and to update anthropogenic emissions
134 under observational constraints. Within each data assimilation window, RegGCAS adopted
135 a "two-step" inversion scheme. In the forecast step, CMAQ simulated the CO₂
136 concentration field based on prior emissions, establishing the response relationship
137 between emissions and satellite XCO₂ observations. Subsequently, in the analysis step,
138 satellite XCO₂ observations were used to optimize the prior emissions and obtain posterior
139 emissions. These updated emissions were then passed as the prior emissions for the next
140 window, further driving CMAQ to generate the initial concentration field for that window.
141 This scheme enables observational information to be propagated across adjacent windows,
142 allowing iterative updates of emissions and thereby improving inversion accuracy and
143 stability (Feng et al., 2024).

144 **2.1.1 Atmospheric transport model**

145 In this study, the Weather Research and Forecasting (WRF v4.0) model (Skamarock and
146 Klemp, 2008) and the Community Multiscale Air Quality (CMAQ v5.0.2) model (Byun
147 and Schere, 2006) were employed to simulate the meteorological fields and CO₂ transport
148 processes, respectively. The CMAQ modeling domain covers mainland China and its
149 surrounding regions (Fig. 1a), with a horizontal resolution of 27 km and 225 × 165 grid



cells. Vertically, 25 sigma-pressure layers were set from the surface up to 100 Pa. To better characterize the urban underlying surface, the MODIS land cover product (MCD12C1, version 6.1) for 2022 was used to update the urban and built-up land information in the model. Meteorological initial and lateral boundary conditions for WRF were derived from the National Centers for Environmental Prediction Final Analysis (NCEP FNL) data at $1^\circ \times 1^\circ$ resolution and 6 h intervals. The initial field for the first assimilation window and the lateral boundary conditions were both derived from the inversion-optimized results of the Global Carbon Assimilation System (GCAS) v2 (Jiang et al., 2021). Considering that CO₂ is a long-lived gas and chemically inert over the study period, gas-phase chemistry and aerosol-related processes were disabled in CMAQ, retaining only physical processes (e.g., transport and diffusion). This configuration allows the simultaneous simulation of CO₂ concentration ensembles while reducing computational cost.

2.1.2 EnKF assimilation algorithm

The ensemble Kalman filter (EnKF) employs a Monte Carlo approach, using a set of ensemble states to represent the probability distribution of the true state (Evensen, 2003). Through the assimilation of observations, the model states are updated iteratively, minimizing the deviation between simulations and observations. In this study, the Ensemble Square Root Filter (EnSRF) method was used to optimize anthropogenic CO₂ emissions (Whitaker and Hamill, 2002).

The EnSRF first applied a zero-mean Gaussian perturbation to the background state vector X^b to construct an ensemble of model states, with the perturbation magnitude determined by the state uncertainty. Subsequently, the background error covariance matrix P^b was calculated based on these ensemble samples:

$$P^b = \frac{1}{N-1} \sum_{i=1}^N (X_i^b - \bar{X}^b) (X_i^b - \bar{X}^b)^T \quad (1)$$

where N is the ensemble size; X_i^b represents the i th ensemble sample; and $\bar{X}^b$ is the ensemble mean. The background error covariance matrix P^b was used to represent the uncertainty of the background field and its error correlation structure, providing the basis for subsequent state updates. In the forecast step, WRF-CMAQ advanced each ensemble member to the next time step, propagating the emission perturbations to the CO₂ concentration field, thereby establishing the response relationship between concentrations



180 and emissions. In the analysis step, the observation vector y was further introduced to
181 update the background state, yielding the analyzed state ensemble mean $\overline{X^a}$, which was
182 taken as the optimal estimate of the emissions at the current time:

$$183 \quad \overline{X^a} = \overline{X^b} + K(y - H\overline{X^b}) \quad (2)$$

$$184 \quad K = P^b H^T (H P^b H^T + R)^{-1} \quad (3)$$

185 where R is the observation error covariance matrix; and K is the Kalman gain matrix
186 that determines the relative weights of background and observational information.

187 OCO-2/3 provides XCO_2 , whereas the model outputs three-dimensional CO_2
188 concentrations. Therefore, an observation operator H is required to map the model state
189 into the satellite observation space. Specifically, the CO_2 vertical profile of the
190 corresponding grid was first extracted from CMAQ based on the latitude, longitude, and
191 time of the observation. The model-layer concentrations were then interpolated to the OCO
192 retrieval layers using pressure-level information, yielding a simulated profile x_i^m
193 consistent with the satellite vertical layering. On this basis, the simulated XCO_2 was
194 calculated as follows:

$$195 \quad XCO_2^m = XCO_2^a + \sum_{i=1}^n w_i A_i (x_i^m - x_i^a) \quad (4)$$

196 where XCO_2^m is the simulated XCO_2 processed by the observation operator; XCO_2^a is the
197 prior XCO_2 from the satellite product; x_i^m and x_i^a denote the modeled CO_2
198 concentration and satellite prior profile at the i th retrieval layer, respectively; A_i and w_i
199 represent the averaging kernel and pressure weight for that layer; and n is the number of
200 satellite retrieval layers.

201 The system used 50 ensemble members and a 3-day assimilation window. This is
202 because it takes time for emission perturbations to propagate to the concentration field.
203 Appropriately extending the window allows more satellite observations to be incorporated,
204 thereby enhancing the constraints on emissions. However, if the window is set too long,
205 transport errors and spurious correlations within the ensemble filter will gradually
206 accumulate, which may degrade the assimilation performance (Feng et al., 2023).
207 Furthermore, to address the problem of spurious long-distance correlations in the EnKF
208 (Houtekamer and Mitchell, 2001), this study applied covariance localization using the
209 Gaspari-Cohn localization function (Gaspari and Cohn, 1999). This function is a piecewise
210 continuous fifth-order polynomial approximation of a normal distribution (Miyazaki et al.,



211 2017), allowing the covariance to decrease gradually with distance and thereby reducing
212 the unreasonable influence of remote observations on local analyses. Taking into account
213 the ensemble size, the data assimilation window length, and the transport characteristics of
214 CO₂, the localization scale was set to 500 km (Flowerdew, 2015).

215 **2.2 Prior emissions**

216 In the initial data assimilation window, prior anthropogenic CO₂ emissions were prescribed
217 using the Multi-resolution Emission Inventory for China (MEIC) v1.4 2019 inventory (Fig.
218 1a) for mainland China (Li et al., 2017a) and the MIX Asian anthropogenic emission
219 inventory for other regions (Li et al., 2017b). In each subsequent window, the prior
220 anthropogenic CO₂ emissions were set to the posterior emissions from the previous
221 window and updated progressively, so that the sensitivity to errors in the original inventory
222 gradually decreased. To mitigate the interference of terrestrial biosphere fluxes on
223 anthropogenic emission estimates, biogenic carbon fluxes were integrated into the
224 assimilation system for joint optimization. Prior biogenic carbon fluxes were derived from
225 the net ecosystem production (NEP) simulated by the Boreal Ecosystem Productivity
226 Simulator (BEPS) model (Wang et al., 2021) at $0.073^{\circ} \times 0.073^{\circ}$ resolution. Prior
227 uncertainties for anthropogenic emissions and biogenic fluxes were assigned grid-cell
228 values of 40% and 80%, respectively, accounting for the errors from statistical data, spatial
229 disaggregation, and biophysical modeling (Super et al., 2024). Additionally, the system
230 incorporated biomass burning emissions and ocean carbon fluxes to represent regional-
231 scale CO₂ background variations. Specifically, fire carbon emissions were taken from the
232 Global Fire Emissions Database (GFED) v4.1s (Randerson et al., 2017) at $0.25^{\circ} \times 0.25^{\circ}$
233 and 3-hour resolution. Ocean carbon fluxes were obtained from the simulations of GCAS
234 v2 (Jiang et al., 2021) at $1^{\circ} \times 1^{\circ}$ and 3-hour resolution.

235 **2.3 Observation data**

236 This study assimilated bias-corrected Level 2 Lite XCO₂ retrievals from OCO-2 v11.2r and
237 OCO-3 v10.4r, both processed with the Atmospheric Carbon Observations from Space
238 (ACOS) algorithm. OCO-2 operates in a sun-synchronous polar orbit, providing regular
239 global coverage, whereas OCO-3, mounted on the International Space Station, provides
240 more flexible observing modes, including target and Snapshot Area Map observations.
241 Consequently, the two missions differ in spatial coverage, overpass time, viewing geometry,



242 and data density (Eldering et al., 2019). Despite these sampling differences, quality-filtered
243 and bias-corrected XCO₂ retrievals from OCO-2 and OCO-3 are broadly consistent.
244 (Taylor et al., 2023) reported a mean difference of 0.2 ppm over land between collocated
245 OCO-3 and OCO-2 observations, indicating negligible systematic biases. Thus, combining
246 OCO-2 and OCO-3 XCO₂ data increases observational coverage and strengthens
247 constraints for regional-scale CO₂ emission inversions.

248 During preprocessing, we retained only Land Nadir and Land Glint (LNLG)
249 observations to match the terrestrial inversion domain and reduce retrieval uncertainties
250 over non-land surfaces. Invalid or missing XCO₂ values were removed, and only
251 high-quality retrievals (xco2_quality_flag = 0) were used to mitigate the influence of
252 clouds, aerosols, and instrument noise. The spatial distribution and counts of the
253 assimilated observations are shown in Fig. 3. For the EnSRF assimilation, the XCO₂
254 uncertainty from the satellite products was used as the observation error for each
255 observation.

256 **2.4 Evaluation method and data**

257 To evaluate the performance of the RegGCAS and the accuracy of the posterior
258 anthropogenic CO₂ emissions, this study conducted analyses from two aspects: the
259 consistency evaluation using assimilated observations and validation using independent
260 observations (Bocquet et al., 2015). The consistency evaluation used OCO-2/3 XCO₂
261 observations to compare the agreement between the model simulations (before and after
262 assimilation) and the assimilated observations, thereby testing the internal consistency of
263 the assimilation system (Whitaker and Hamill, 2002). For independent validation, we used
264 non-assimilated ground-based CO₂ observations from public observational networks and
265 an additional urban observation site. The public datasets included XCO₂ observations from
266 the Total Carbon Column Observing Network (TCCON) at Xianghe (XH) and Hefei (HF),
267 and CO₂ concentrations from the ObsPack GLOBALVIEWplus v10.1 dataset at five sites:
268 DSI, LLN, TAP, AMY, and WLG. In addition, in situ CO₂ observations from the Shanghai
269 Zhangjiang urban site (ZJ; 31.18°N, 121.59°E), measured by a Picarro G2301 cavity ring-
270 down spectrometer, were used to further evaluate posterior CO₂ simulations over the YRD
271 urban region. Located on the rooftop of the Shanghai Advanced Research Institute, Chinese
272 Academy of Sciences, in Zhangjiang Hi-Tech Park, this site is surrounded by educational
273 and commercial districts and lies near urban expressways, thus providing a critical
274 assessment of a region with intense anthropogenic activity. Together, these observations



275 provided independent constraints on the simulated atmospheric CO₂ from both column and
276 near-surface perspectives. The spatial distribution of all validation sites is shown in Fig. 2a.
277 The evaluation metrics included mean bias (BIAS), mean absolute error (MAE), root mean
278 square error (RMSE), and correlation coefficient (CORR).

279 To corroborate the emission adjustments derived from OCO-2/3 XCO₂ assimilation,
280 this study additionally employed a NO_x proxy-based method to infer anthropogenic CO₂
281 emissions (Feng et al., 2024). The method assumes a CO₂-NO_x co-emission relationship
282 during fossil fuel combustion, so that optimized NO_x emissions can indicate the spatial
283 adjustment of anthropogenic CO₂. For comparability, the same MEIC inventory and study
284 period were used. Firstly, the RAPAS system assimilated hourly surface NO₂ observations
285 from China's national air quality monitoring stations to constrain MEIC NO_x and produce
286 posterior estimates. Cell-wise CO₂-to-NO_x emission ratios from the MEIC were then
287 applied to the posterior NO_x to infer anthropogenic CO₂ emissions. Finally, the resulting
288 CO₂ was compared with the posterior from direct XCO₂ assimilation on a uniform 27-km
289 grid, serving as an auxiliary reference for evaluating the robustness and uncertainty of the
290 key-region emission adjustments.

291 **2.5 OSSE experiment**

292 To evaluate the reliability and constraint capability of RegGCAS, this study designed an
293 OSSE experiment. The overall experimental procedure was generally consistent with the
294 emission inversion constrained by real satellite observations, except that the observations
295 and prior emissions were idealized. Specifically, the MEIC emission inventory was first
296 used as the true emissions to drive the WRF-CMAQ model for forward simulations. Then,
297 based on the spatiotemporal sampling characteristics of OCO-2/3, CO₂ concentrations at
298 the corresponding locations were extracted from the simulation results to construct pseudo-
299 observations. Subsequently, the MEIC emissions were uniformly scaled down by 30% to
300 serve as the prior, and the inversion was conducted under the same assimilation
301 configuration as the real inversion (Liu et al., 2017). Finally, by comparing the posterior
302 emissions with the true emissions, we evaluated the capability of RegGCAS to invert
303 anthropogenic CO₂ emissions under idealized observational conditions.



304 **3. Results and discussion**

305 **3.1 Anthropogenic CO₂ emissions**

306 Fig. 1(a-c) presents the spatial distribution of anthropogenic CO₂ emissions over mainland
307 China during winter before and after assimilation, as well as their deviations from the
308 MEIC prior. Overall, high-emission areas were primarily distributed in densely populated
309 and economically active regions in eastern coastal and central-eastern China, including
310 Beijing-Tianjin-Hebei (BTH), YRD, the Pearl River Delta, and the Sichuan Basin.
311 Conversely, emissions were relatively low in the northwestern, southwestern, and
312 northeastern regions. Compared with the MEIC, the posterior emissions exhibited an
313 overall upward adjustment at the national scale. The total posterior anthropogenic CO₂
314 emissions during winter amounted to 2808.3 ± 157.0 Tg, 16.1% higher than the MEIC,
315 indicating a certain underestimation in the MEIC inventory.

316 The adjustments of posterior CO₂ emissions exhibited distinct spatial heterogeneity,
317 with emission increases mainly located in northern China, particularly in the North China
318 Plain. This may be related to the underestimation of winter heating emissions in the prior
319 inventory. Winter heating in northern China relies heavily on fossil fuels such as coal, and
320 residential coal combustion, along with decentralized heating, contributes substantially to
321 regional emissions (Zhang et al., 2021). However, these emission sources are relatively
322 dispersed, leading to considerable uncertainties in activity level statistics, and a portion of
323 these emissions may not be fully incorporated into the inventory (Li et al., 2017a). Previous
324 studies have indicated that rural residential coal consumption in the North China Plain may
325 be underreported in official statistics, resulting in an underestimation of winter coal
326 combustion emissions in existing inventories (Peng et al., 2019). This justifies the overall
327 upward adjustment of emissions in this region after assimilation.

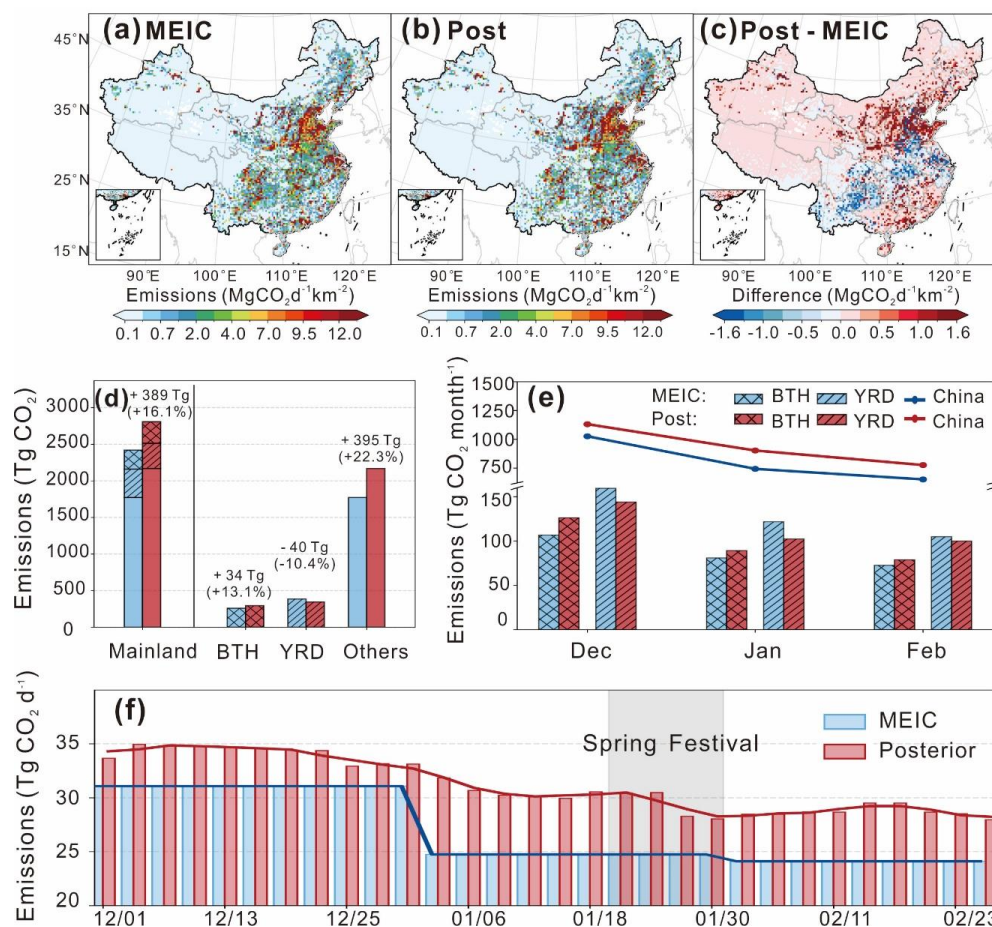


Fig. 1. Anthropogenic CO₂ emissions over mainland China before and after assimilation. (a) MEIC inventory ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$); (b) posterior emissions ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$); (c) posterior-minus-prior differences ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$); (d) winter totals over mainland China, BTH, and YRD (TgCO₂); (e) monthly totals (TgCO₂ month⁻¹); (f) winter daily means (TgCO₂ d⁻¹), with gray shading for the Spring Festival.

As two major hotspots for anthropogenic emissions, the BTH and YRD regions accounted for 10.8% and 16.0% of the national total, respectively. For the BTH region, most emissions were mainly distributed in the densely populated urban and industrial areas in the east and south. The total posterior emissions were 294.2 ± 28.9 Tg, 13.1% higher than the MEIC prior, with adjustments primarily occurring in high-emission areas (Fig. 2b-c). For the YRD region, significant downward adjustments of posterior emissions were primarily located in the eastern coastal belt and the major urban clusters (Fig. 2e-f). The total posterior emissions for the YRD were 345.9 ± 46.3 Tg, 10.4% lower than the MEIC,



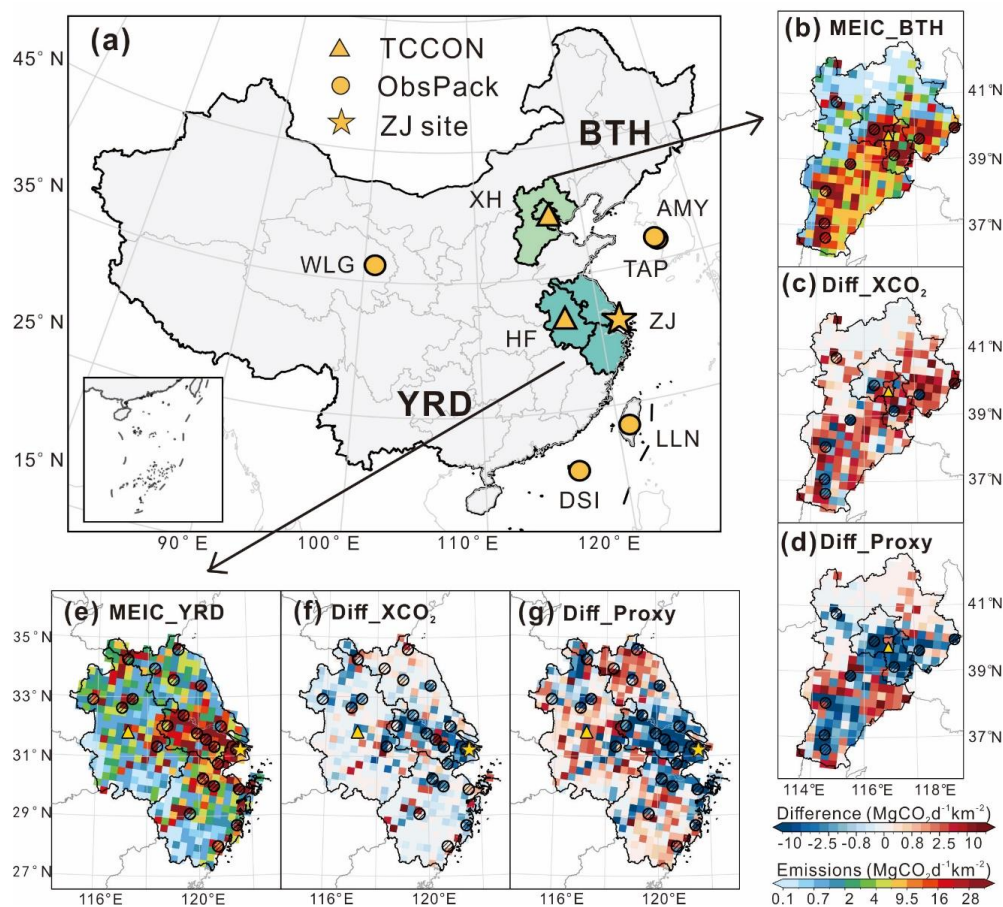
342 suggesting a potential overestimated in the MEIC for this region. This discrepancy likely
343 arose from overestimated emission factors. As noted by Feng et al. (2024), economically
344 developed areas such as the YRD generally possessed technological levels and energy
345 utilization efficiencies that surpassed the national average, resulting in relatively lower
346 carbon emission intensities per unit of economic output. Consequently, applying uniform
347 emission factors in the inventory failed to accurately reflect actual emission levels, thereby
348 leading to positively biased prior emissions. In both regions, emissions in most key city
349 centers were adjusted downward, while emissions in suburban and surrounding areas
350 increased. This spatial adjustment pattern likely reflected biases in the spatial
351 disaggregation of the emission inventory. Gridded inventories like MEIC typically rely on
352 proxy variables (e.g., population density and road networks) for spatial downscaling
353 (Zheng et al., 2021). However, these proxy variables are overwhelmingly concentrated in
354 urban centers and struggle to accurately represent high-emission point sources (e.g.,
355 factories and power plants) located in urban peripheries or suburban industrial belts. Thus,
356 these inventories tend to allocate excessive emissions to central urban areas, thereby
357 underestimating actual emissions in the peripheries. It is worth noting that, although
358 satellite XCO₂ observations are less sensitive to local emission changes than ground-based
359 observations (Parker et al., 2023), this study, constrained solely by OCO-2/3 XCO₂
360 observations, still identified a distinct emission adjustment pattern between city centers
361 and their peripheries in high-emission regions. This demonstrates that RegGCAS can not
362 only constrain total regional emissions but also to some extent reflect spatial allocation
363 biases within the urban clusters of these key emission regions.

364 This study also presents the results derived from the proxy-based method. For the
365 BTH, emission adjustments identified were primarily located in key cities and their
366 surroundings, with downward adjustments in most key city centers. This finding was
367 generally consistent with the direct assimilation results of XCO₂ observations (Fig. 2d).
368 However, the overall CO₂ emissions derived from this method were lower and exhibited a
369 broader extent of downward adjustments, which may be related to the CO₂-to-NO_x ratios
370 used in the conversion. Due to the time lag in updating "bottom-up" inventories, this study
371 adopted the ratios from the 2019 inventory. Previous studies have demonstrated a
372 continuous increase in the CO₂-to-NO_x ratios over East Asia in recent years (Li et al., 2024;
373 Li et al., 2023), and this trend may be even more pronounced in the BTH region, where
374 winter pollution control requirements were more stringent (Xu et al., 2023; Liu et al., 2020).
375 Therefore, continuing to rely on these relatively outdated CO₂-to-NO_x ratios likely led to



376 an underestimation of actual CO₂ emissions. For the YRD, our results were similar to those
377 from the proxy-based approach, with emission adjustments predominantly downward and
378 concentrated mainly in key urban clusters in the east, particularly in southern Jiangsu
379 Province (Fig. 2g). Overall, the two methods showed relatively consistent emission
380 adjustment patterns in the key regions, which to some extent supports the reliability of the
381 results in this study.

382 In terms of monthly variations, anthropogenic CO₂ emissions over mainland China,
383 as well as in the BTH and YRD regions, exhibited a continuous monthly declining trend
384 (Fig. 1e). The national posterior emissions remained consistently higher than the prior, with
385 monthly increases of 12.5%, 20.6%, and 17.2% during the winter, indicating that
386 assimilation exerted a more pronounced upward effect on emissions. For the BTH,
387 posterior emissions were higher than the prior, with the most significant upward adjustment
388 (18.4%) in December. For the YRD, posterior emissions were consistently lower than the
389 prior, recording the most substantial decrease (16.0%) in January. In the 3-day assimilation
390 window (Fig. 1f), posterior emissions dropped markedly during the Spring Festival holiday,
391 with a reduction of approximately 2-3 Tg CO₂ d⁻¹, and gradually rebounded thereafter. This
392 indicates that, compared to the prior emissions that only reflect monthly variations, the
393 posterior results better capture the short-term emission fluctuations associated with reduced
394 anthropogenic activities during the Spring Festival and the subsequent post-holiday
395 recovery.



396

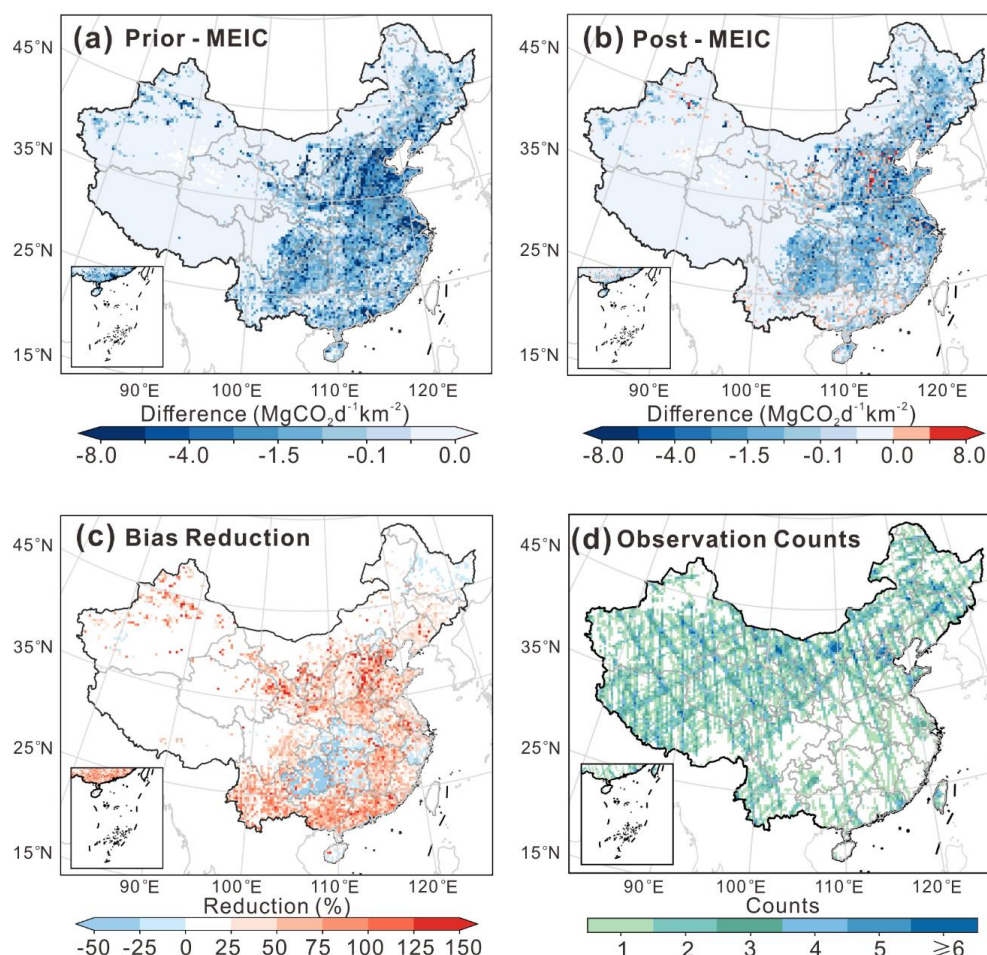
397 Fig. 2. Spatial distribution of emissions and adjustments ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$) in the BTH and YRD
 398 regions. (a) BTH and YRD locations. Yellow circles denote ObsPack sites; (b-d) BTH: posterior
 399 emissions, and emissions adjustments derived from direct XCO_2 assimilation (Diff_XCO₂) and the
 400 proxy-based method (Diff_Proxy), respectively; (e-g) same as (b-d) but for YRD. The yellow
 401 triangle denotes the XH site (b-d) and HF site (e-g). The yellow five-pointed star denotes the ZJ
 402 site. Hatched circles denote key city centers.

403 3.2 Emission evaluation

404 We first utilized an OSSE to evaluate the reliability and constraint capability of the
 405 RegGCAS. The results revealed that assimilation can effectively utilize pseudo-
 406 observations, bringing the emission estimates closer to the MEIC. In terms of regional
 407 totals, the national prior emissions were underestimated by 30%, which decreased to 9.6%
 408 after assimilation, representing a 68.0% reduction in prior bias. Notably, in the densely



409 observed BTH region, the total posterior emissions were nearly identical to the true
410 emissions, with a 98.7% reduction in prior bias. In the YRD region, the prior bias was also
411 reduced by 58.0% after assimilation. In terms of spatial distribution (Fig. 3b), the range of
412 negative biases was significantly reduced after assimilation, with emission biases
413 approaching zero in many regions. Meanwhile, certain residual negative biases persisted
414 locally, suggesting that under the current OCO observational coverage, the constraints in
415 some regions may still be insufficient, preventing full recovery of true emissions across all
416 grids. This also implied that the posterior emissions depicted in Fig. 1(b) were still
417 underestimated. However, in some regions (e.g., Guizhou, Hubei, and Hunan), adjustments
418 still deviated from the true emissions (Fig. 3c), primarily due to inadequate local
419 observation coverage (Fig. 3d). Observations in these regions are relatively sparse. Once
420 emissions were over-adjusted in an individual assimilation window, and subsequent
421 observations were insufficient to correct the adjustment, the bias persisted in the following
422 windows.



423

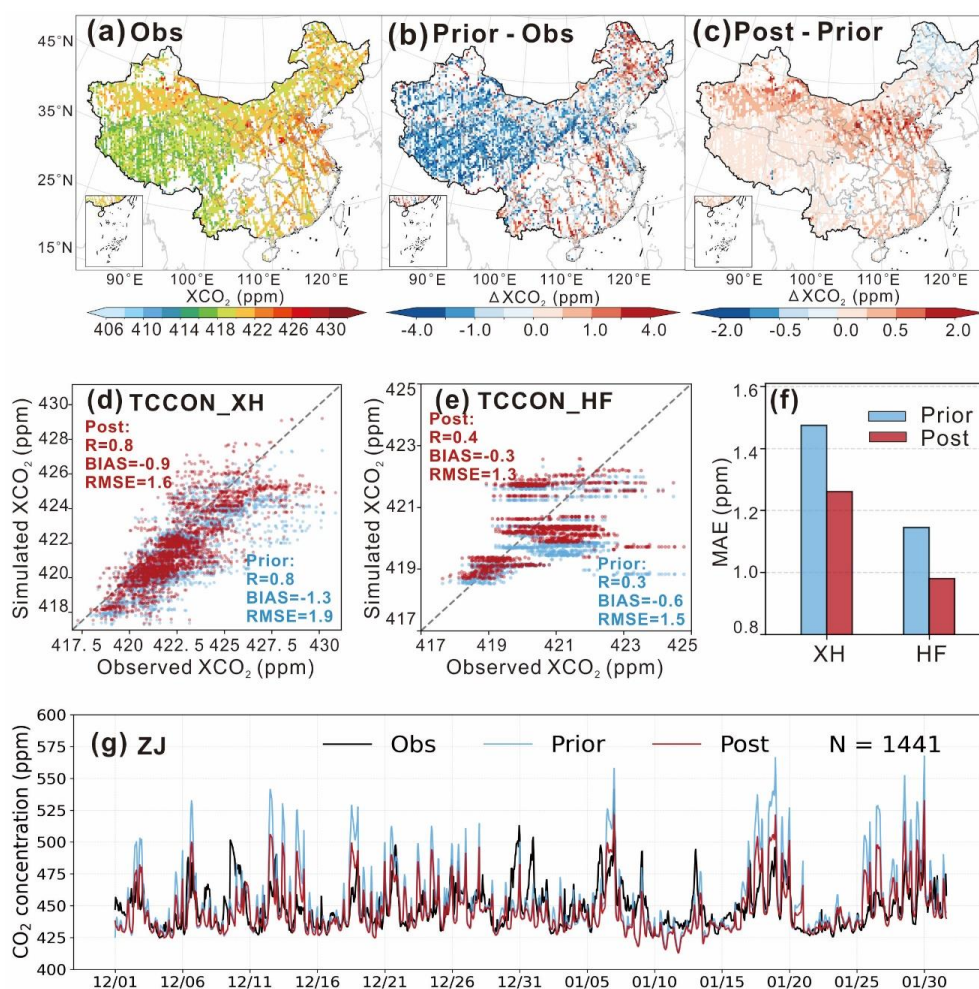
424 Fig. 3. OSSE results. (a) prior-minus-MEIC differences ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$); (b) posterior-minus-
425 MEIC differences ($\text{MgCO}_2 \text{d}^{-1} \text{km}^{-2}$); (c) prior bias reduction (%). Reduction = (Posterior - Prior) /
426 (MEIC - Prior) $\times 100$; (d) OCO observation counts.

427 The results of the assimilated observation consistency analysis (Fig. 4a-c) indicate
428 that, at the national scale, the prior simulated XCO_2 broadly exhibited negative biases,
429 particularly in the North and Northwest China. The adjustments in posterior simulations
430 relative to the prior simulations were generally consistent with the posterior emission
431 adjustments, and were mainly distributed in northern China, where anthropogenic sources
432 were intensive and observations were relatively dense. After assimilation, the posterior
433 simulation biases were significantly reduced. For example, the posterior simulated
434 concentrations in North and Northwest China were increased, significantly improving the



underestimation in the prior simulations. Similarly, posterior simulations effectively alleviated the overestimation in Heilongjiang Province in Northeast China by decreasing concentrations. Statistical results show that after assimilation, the national BIAS was reduced from -0.6 ppm to -0.4 ppm, while the MAE and RMSE decreased by 6.8% and 5.8%, respectively, indicating an improved consistency between the simulations and observations. For key emission regions, the BTH showed notable improvement, with BIAS decreasing from approximately -0.7 ppm to -0.3 ppm, a reduction of 56.5%, and MAE and RMSE were reduced by 14.4% and 12.5%, respectively. In contrast, the YRD region exhibited relatively small changes in statistical metrics, which is largely attributed to the relatively insufficient satellite observational coverage in this region. Nevertheless, the posterior simulation MAE was still reduced by 3.2%.

The evaluation results at the TCCON sites show that the posterior simulations were consistently improved compared to the prior simulations (Fig. 4h-i). Specifically, at the XH site, BIAS was reduced by 28.8%, MAE and RMSE decreased by 14.3% and 15.6%, respectively, and the CORR increased from 0.76 to 0.80. At the HF site, BIAS decreased from -0.6 ppm to -0.3 ppm, a reduction of approximately 50.0%, MAE and RMSE reduced by 14.0% and 14.9%, respectively, and the CORR increased from 0.31 to 0.41. Overall, with a 15.3% average RMSE reduction at both stations, posterior simulations outperformed prior simulations after assimilation.



455

456 Fig. 4. Evaluation of XCO₂ simulation based on satellite observations and independent site data
 457 (ppm). (a-c) OCO-2/3 XCO₂ observations, prior-minus-observation differences, and posterior-
 458 minus-observation differences; (d-e) Comparison of simulated and observed XCO₂ at the XH and
 459 HF sites; (f) MAE of prior and posterior simulations at two sites; (g) Time series of observed near-
 460 surface CO₂ concentrations and prior and posterior simulations at the Shanghai ZJ site.

461 Table 1 presents the independent validation results at the ObsPack ground-based sites.
 462 At the TAP site, MAE decreased from 2.1 ppm to 1.8 ppm, RMSE dropped by 17.4%, and
 463 the CORR increased from 0.89 to 0.94. The DSI, LLN, and WLG sites also exhibited
 464 certain improvements, with their posterior MAE and RMSE decreasing by approximately
 465 2.5%-9.0% and 3.3%-5.0%, respectively. Although posterior MAE at the AMY site was
 466 slightly higher than the prior result, its RMSE still decreased by 8.9%. The varying degrees



of improvement among sites are mainly due to differences in site environments and their sensitivity to anthropogenic emission changes. TAP, DSI, and AMY are all located in coastal or offshore environments, where observed concentrations are highly susceptible to the combined effects of regional transport and local meteorological processes (Lin et al., 2018). WLG and LLN are high-altitude background stations with relatively weak influences from surrounding local anthropogenic sources, making them more sensitive to large-scale background variations and long-range transport. Consequently, in addition to the emissions, their simulation errors are closely related to lateral boundary inputs, background field, and transport errors (Guo et al., 2020; Ou-Yang et al., 2025). Overall, most sites showed improvement after assimilation, with BIAS decreasing by 8.7% and RMSE by 7.7%, indicating that the assimilation of OCO XCO₂ observations can effectively improve the estimation of anthropogenic CO₂ emissions and enhance the capability to simulate CO₂ concentrations.

Table 1. Validation statistics for the prior and posterior simulated CO₂ concentrations at the ObsPack ground-based sites and the ZJ site.

Site	Prior				Posterior			
	BIAS (ppm)	MAE (ppm)	RMSE (ppm)	CORR	BIAS (ppm)	MAE (ppm)	RMSE (ppm)	CORR
TAP	-0.35	2.11	2.65	0.89	-0.32	1.77	2.19	0.94
DSI	-1.53	2.10	2.79	0.65	-1.27	1.91	2.65	0.66
LLN	-0.06	1.19	1.51	0.47	0.14	1.16	1.46	0.53
WLG	-1.01	1.21	1.63	0.70	-0.79	1.13	1.57	0.68
AMY	0.94	1.90	2.47	0.74	1.03	1.94	2.25	0.75
ZJ	4.34	15.23	21.95	0.56	-2.38	11.90	16.42	0.58

In addition, the near-surface CO₂ comparison at the Shanghai ZJ urban site provides an independent evaluation of the posterior emissions under strong urban influence. The prior simulation exhibits a positive bias relative to observations, especially during high CO₂ episodes (Fig. 4g), which is consistent with the overestimated prior emissions in this region. Following optimization, the posterior simulation effectively mitigates this



487 overestimation. Although the posterior simulation retains a residual negative bias,
488 primarily attributable to the spatial representation error between grid-mean model
489 concentrations and point-scale urban measurements heavily impacted by intense local
490 emissions (Tolk et al., 2008), the MAE and RMSE decrease by 21.9% and 25.2%,
491 respectively, while the CORR slightly increases from 0.56 to 0.58. This marked reduction
492 in overestimation provides compelling evidence that the system can robustly constrain
493 emissions even in urban regions characterized by dense anthropogenic sources.

494 **3.3 System sensitivity analysis**

495 To evaluate the sensitivity of the inversion results to different system configurations, this
496 study conducted a series of sensitivity experiments focusing on factors including the
497 background fields, observation settings, prior anthropogenic emission, biogenic carbon
498 fluxes, localization scales, as well as background and observation errors (Table 2). Among
499 these, the localization scale primarily influences the spatial range of observational
500 information (Hunt et al., 2007), and the background and observation errors affect the
501 assimilation results by altering the relative weights between the prior and observations
502 (Evensen, 2003). The results demonstrated that the posterior total deviations relative to the
503 BASE were consistently constrained within $\pm 3.5\%$ for a subset of sensitivity experiments,
504 including those on different localization scales (LOC_300, LOC_800), background error
505 levels (ERR_BKG0.2, ERR_BKG0.8), and observation error magnitudes (ERR_OBS0.5,
506 ERR_OBS2). This demonstrated that the two-step inversion strategy can effectively
507 mitigate the sensitivity of the inversion results to system parameters, indicating the
508 robustness of the system.

509 When assimilating only OCO-2 (OBS_OCO2) or only OCO-3 (OBS_OCO3), the
510 posterior totals decreased by 1.1% and 10.4% relative to BASE, respectively. Based on the
511 OSSE results, the BASE experiment still underestimated emissions under current
512 observations, while both single-satellite assimilation experiments yielded even lower
513 emission results than BASE. This indicates that a single satellite has inherent deficiencies
514 in spatial coverage and sampling uniformity, making it insufficient to fully constrain
515 regional emissions and necessitating the joint assimilation of multiple observational
516 datasets.

517

518



Table 2. Configurations of the sensitivity experiments and their posterior emission differences relative to the BASE experiment.

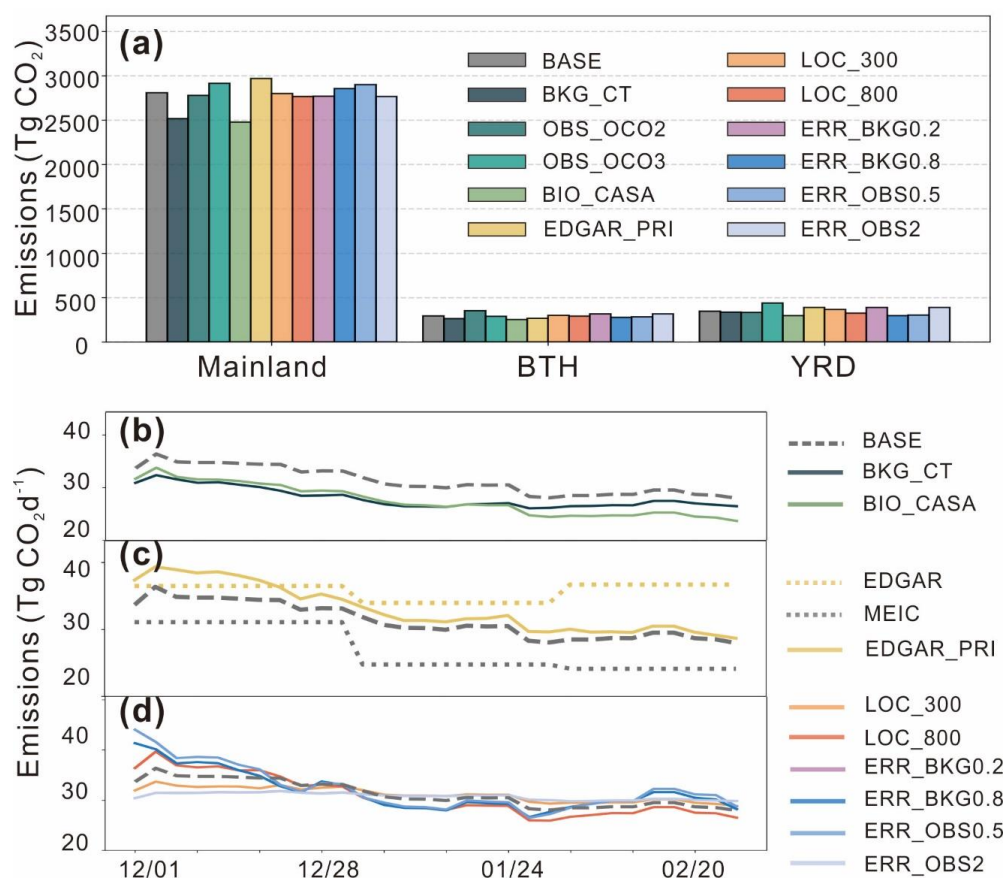
Name	BKG	Anth	Bio	Obs	Loc (km)	BKG_Error (%)	OBS_Error (factor)	Diff (%)
BASE	GCAS	MEIC	BEPS	O2+O3	500	40	1.0	0.0
BKG_CT	Carbon Tracker	MEIC	BEPS	O2+O3	500	40	1.0	-10.3
OBS_OCO2	GCAS	MEIC	BEPS	O2	500	40	1.0	-1.1
OBS_OCO3	GCAS	MEIC	BEPS	O3	500	40	1.0	-10.4
BIO_CASA	GCAS	MEIC	CASA	O2+O3	500	40	1.0	-11.7
EDGAR_PRI	GCAS	EDGAR	BEPS	O2+O3	500	40	1.0	+5.7
LOC_300	GCAS	MEIC	BEPS	O2+O3	300	40	1.0	-0.3
LOC_800	GCAS	MEIC	BEPS	O2+O3	800	40	1.0	-1.5
ERR_BKG0.2	GCAS	MEIC	BEPS	O2+O3	500	20	1.0	-1.4
ERR_BKG0.8	GCAS	MEIC	BEPS	O2+O3	500	80	1.0	+1.7
ERR_OBS0.5	GCAS	MEIC	BEPS	O2+O3	500	40	0.5	-1.5
ERR_OBS2	GCAS	MEIC	BEPS	O2+O3	500	40	2.0	+3.3

Note: BKG denotes inversion background field; Anth represents prior anthropogenic CO₂ emissions; Bio denotes biogenic CO₂ fluxes; Obs denotes assimilated OCO satellite observations. O2 and O3 refer to OCO-2 and OCO-3 data, and O2+O3 represents their joint assimilation. Loc is the assimilation localization scale. BKG_Error represents background error. OBS_Error denotes the prescribed observation error scaling factor. Diff indicates the relative difference in posterior anthropogenic emissions between sensitivity tests and the BASE case.

Due to the long atmospheric lifetime of CO₂, background concentrations contribute substantially to the total CO₂ concentrations (Mustafa et al., 2021). Consequently, when biases exist in the background fields, the inversion system tends to compensate for them by adjusting anthropogenic emissions (Pisso et al., 2019). When using the Carbon Tracker background field (BKG_CT), the total posterior emissions decreased by 10.3% relative to BASE, indicating that discrepancies in background concentrations exert a relatively substantial impact on the emission inversion. In the biogenic carbon flux perturbation experiment (BIO_CASA), posterior anthropogenic emissions decreased by 11.7% compared to the BASE. Compared to BEPS, the Carnegie-Ames-Stanford Approach (CASA) product exhibited stronger positive fluxes at the national scale, particularly in the central-eastern, North China, and Northeast China regions, thereby elevating the natural background CO₂ levels. To match satellite XCO₂ observations, the assimilation system offsets this higher biogenic background by producing correspondingly lower posterior anthropogenic emissions.



540 When employing EDGAR as the prior inventory (EDGAR_PRI), posterior
541 anthropogenic emissions increased by 5.7% relative to BASE. Although the emissions in
542 the EDGAR inventory were higher than those in the MEIC, prior inventory differences had
543 a limited impact on inversion results. For instance, the prior difference between EDGAR
544 and MEIC increased from 16.2 Tg in the first window to 37.8 Tg in the last window,
545 whereas the corresponding posterior difference narrowed from 11.1 Tg to 2.1 Tg. This
546 indicated that observational constraints effectively mitigated the influence of prior
547 inventory differences on the posterior results.



548

549 Fig. 5. Sensitivity experiment results. (a) total anthropogenic CO₂ emissions over mainland China
550 and key regions under different sensitivity experiments (TgCO₂); (b-d) temporal variations of
551 national anthropogenic CO₂ emissions under different system configurations (TgCO₂ d⁻¹).

552 Comparing the inversion results across sensitivity experiments over time (Fig. 5b-d),
553 it can be observed that in the first assimilation window, the regional emission estimates



554 varied across experiments, with a standard deviation of 4.4 Tg and a range of 13.6 Tg. As
555 assimilation advanced, the dispersion gradually decreased. By the final window, the
556 standard deviation and range had been reduced by 52.3% and 39.7%, respectively. This
557 indicated that although different system assumptions introduce certain discrepancies during
558 the initial phase of assimilation, the inversion results gradually converged as more
559 observations were assimilated, ultimately achieving good consistency. Based on all
560 sensitivity experiments, the uncertainty of regional total emission estimates was 5.6%, thus
561 verifying the high reliability of the posterior emission estimates.

562 **4. Summary and conclusions**

563 Accurately quantifying anthropogenic CO₂ emissions is essential for evaluating carbon
564 budgets and formulating effective mitigation strategies. However, structural uncertainties
565 and spatial allocation biases in traditional "bottom-up" inventories significantly hinder our
566 understanding of anthropogenic carbon forcing. To independently evaluate and optimize
567 these inventories, this study extended RegGCAS, a regional carbon assimilation system
568 based on the WRF-CMAQ atmospheric transport model and the EnKF algorithm. By
569 jointly assimilating OCO-2 and OCO-3 XCO₂ observations, we inverted anthropogenic
570 CO₂ emissions over mainland China during the winter of 2022-2023.

571 Our results yielded national total posterior emissions of 2808.3 ± 157.0 Tg, which is
572 16.1% higher than the MEIC inventory, indicating a systematic underestimation in winter
573 prior emissions. These emission adjustments exhibited significant regional heterogeneity.
574 Specifically, emissions in the BTH region were revised upward by 13.1%, largely due to
575 the underestimation of winter heating emissions. Conversely, emissions in the YRD were
576 adjusted downward by 10.4%, reflecting positively biased emission factors in MEIC.
577 Spatially, the posterior estimates corrected local emission intensities, effectively resolving
578 the emission gradients between urban centers and surrounding suburban areas. Temporally,
579 the assimilated results captured pronounced seasonal and short-term variations, most
580 notably the emission decline during the Spring Festival, demonstrating the system's
581 enhanced capability to respond to short-term anthropogenic activity fluctuations. System
582 reliability was validated through an OSSE, which showed that assimilation reduced prior
583 emission biases by approximately 68%. Furthermore, independent validations
584 demonstrated that the posterior simulations reduced the RMSEs against assimilated OCO-
585 2/3 XCO₂, TCCON XCO₂, ObsPack surface CO₂, and dedicated high-precision urban site



586 observations by 5.8%, 15.3%, 7.7%, and 25.2%, respectively, confirming the enhanced
587 accuracy of the posterior estimates and simulated concentrations.

588 Compared to previous studies that primarily focused on specific point sources or
589 relied on sparsely distributed ground-based observations in limited urban areas, this study
590 advances the field by utilizing satellite XCO₂ constraints to achieve a targeted inversion of
591 anthropogenic emissions at a national scale. Our findings align with and expand upon
592 existing knowledge. The upward adjustment in the BTH region corroborates evidence of
593 underreported coal consumption, while the downward adjustment in the YRD reflects
594 higher regional energy efficiencies that uniform emission factors fail to capture. Moreover,
595 the distinct urban-suburban correction pattern highlights the limitations of proxy-based
596 downscaling in current gridded inventories, which frequently allocate excessive emissions
597 to urban cores while underestimating major industrial sources in suburban belts. Our results
598 are also consistent with previous proxy-based inversion studies relying on co-emitted NO_x;
599 however, direct XCO₂ assimilation overcomes the inherent uncertainties associated with
600 the lagged updates of prior CO₂-to-NO_x ratios.

601 Despite the robustness of this assimilation framework, certain limitations must be
602 acknowledged. As indicated by the OSSE results, existing OCO footprints provide
603 inadequate observation coverage over certain regions (e.g., Guizhou), leading to residual
604 local emission biases. Furthermore, sensitivity analyses revealed that while the inversion
605 is resilient to most internal parameter settings, it exhibits moderate sensitivity to biogenic
606 carbon fluxes and background CO₂ concentrations. Because of the long atmospheric
607 lifetime of CO₂, discrepancies in background fields can inadvertently force the assimilation
608 system into excessive anthropogenic emission adjustments. Finally, although this study
609 intentionally focused on the winter season, the inherent structural biases identified within
610 the bottom-up inventories are likely applicable throughout the year. Future efforts should
611 prioritize the joint assimilation of multi-source surface and satellite observations, along
612 with carbon isotope measurements, to better decouple natural and anthropogenic emissions,
613 thereby extending reliable top-down emission constraints to an annual scale.

614 Ultimately, this study demonstrates that satellite XCO₂ assimilation provides an
615 effective top-down constraint to correct anthropogenic CO₂ emission biases, enhancing the
616 capability of models to capture short-term temporal fluctuations and fine-scale spatial
617 gradients. Crucially, the identification of these systematic regional disparities provides
618 valuable guidance for refining the compilation and proxy-based downscaling methods of
619 future bottom-up emission inventories. Furthermore, the implementation of RegGCAS



620 provides a reliable, operational regional inversion framework. Such an independent and
621 verifiable top-down carbon monitoring system is essential for informing robust, science-
622 based climate mitigation policies.

623 **Code availability**

624 The code of the RegGCAS system is available to the community and can be accessed upon request from
625 Shuzhuang Feng (fengsz@nju.edu.cn) at Nanjing University.

626 **Data availability**

627 The inversion results of this study are available online (<https://doi.org/10.5281/zenodo.20605680>)

628 **Author contributions**

629 SQ and SF designed the research. SQ and HZ processed the data. SQ performed model simulations,
630 analyzed the results, and wrote the manuscript. FJ, CW, YS, MJ, HW, YZ, and WJ discussed and revised
631 the manuscript.

632 **Competing interest**

633 The author declares no conflicts of interest relevant to this study.

634 **Acknowledgments**

635 This work is supported by the National Key Research and Development Program of China
636 (Grant Nos. 2023YFB3907404 and 2022YFB3904801), the National Natural Science
637 Foundation of China (Grant No. 42305116), and the Natural Science Foundation of Jiangsu
638 Province of China (Grant No. BK20230801). We also gratefully acknowledge the High-
639 Performance Computing Center (HPCC) of Nanjing University for supporting the
640 numerical calculations in this paper on its blade cluster system.



References

- Bocquet, M., Elbern, H., Eskes, H., Hirtl, M., Žabkar, R., Carmichael, G. R., Flemming, J., Inness, A., Pagowski, M., Pérez Camacho, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J., Baklanov, A., Carnevale, C., Grell, G., and Seigneur, C.: Data assimilation in atmospheric chemistry models: current status and future prospects for coupled chemistry meteorology models, *Atmospheric Chemistry and Physics*, 15, 5325-5358, <https://doi.org/10.5194/acp-15-5325-2015>, 2015.
- Byun, D. and Schere, K. L.: Review of the Governing Equations, Computational Algorithms, and Other Components of the Models-3 Community Multiscale Air Quality (CMAQ) Modeling System, *Applied Mechanics Reviews*, 59, 51-57, <https://doi.org/10.1115/1.2128636>, 2006.
- Chevallier, F. and Broquet, G.: Atmospheric inversions for carbon dioxide and methane at a national scale, *Natl Sci Rev*, 12, nwaf063, <https://doi.org/10.1093/nsr/nwaf063>, 2025.
- Eldering, A., Taylor, T. E., O'Dell, C. W., and Pavlick, R.: The OCO-3 mission: measurement objectives and expected performance based on 1 year of simulated data, *Atmos. Meas. Tech.*, 12, 2341-2370, [10.5194/amt-12-2341-2019](https://doi.org/10.5194/amt-12-2341-2019), 2019.
- Evensen, G.: The Ensemble Kalman Filter: theoretical formulation and practical implementation, *Ocean Dynamics*, 53, 343-367, <https://doi.org/10.1007/s10236-003-0036-9>, 2003.
- Feng, S., Jiang, F., Zhang, Y., Chen, H., Zhuang, H., Wang, S., Bai, S., Wang, H., and Ju, W.: High-resolution regional inversion reveals overestimation of anthropogenic methane emissions in China, *Atmospheric Chemistry and Physics*, 25, 15121-15143, <https://doi.org/10.5194/acp-25-15121-2025>, 2025.
- Feng, S., Jiang, F., Wang, H., Liu, Y., He, W., Wang, H., Shen, Y., Zhang, L., Jia, M., Ju, W., and Chen, J. M.: China's fossil fuel CO₂ emissions estimated using surface observations of coemitted NO₂, *Environ Sci Technol*, 58, 8299-8312, <https://doi.org/10.1021/acs.est.3c07756>, 2024.
- Feng, S., Jiang, F., Wu, Z., Wang, H., He, W., Shen, Y., Zhang, L., Zheng, Y., Lou, C., Jiang, Z., and Ju, W.: A Regional multi-Air Pollutant Assimilation System (RAPAS v1.0) for emission estimates: system development and application, *Geoscientific Model Development*, 16, 5949-5977, <https://doi.org/10.5194/gmd-16-5949-2023>, 2023.
- Flowerdew, J.: Towards a theory of optimal localisation, *Tellus A: Dynamic Meteorology and Oceanography*, 67, 25257, <https://doi.org/10.3402/tellusa.v67.25257>, 2015.
- Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Landschützer, P., Le Quéré, C., Li, H., Luijkx, I. T., Olsen, A., Peters, G. P., Peters, W., Pongratz, J., Schwingshackl, C., Sitch, S., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., Arneeth, A., Arora, V., Bates, N. R., Becker, M., Bellouin, N., Berghoff, C. F., Bittig, H. C., Bopp, L., Cadule, P., Campbell, K., Chamberlain, M. A., Chandra, N., Chevallier, F., Chini, L. P., Colligan, T., Decayeux, J., Djeutchouang, L. M., Dou, X., Duran Rojas, C., Enyo, K., Evans, W., Fay, A. R., Feely, R. A., Ford, D. J., Foster, A., Gasser, T., Gehlen, M., Gkritzalis, T., Grassi, G., Gregor, L., Gruber, N., Gürses, Ö., Harris, I., Hefner, M., Heinke, J., Hurtt, G. C., Iida, Y., Ilyina, T., Jacobson, A. R., Jain, A. K., Jarníková, T., Jersild, A., Jiang, F., Jin, Z., Kato, E., Keeling, R. F., Klein Goldewijk, K., Knauer, J., Korsbakken, J. I., Lan, X., Lauvset, S. K., Lefèvre, N., Liu, Z., Liu, J., Ma, L., Maksyutov, S., Marland, G., Mayot, N., McGuire, P. C., Metzl, N., Monacchi, N. M., Morgan, E. J., Nakaoka, S.-I., Neill, C., Niwa, Y., Nützel, T., Olivier, L., Ono, T., Palmer, P. I., Pierrot, D., Qin, Z., Resplandy, L., Roobaert, A., Rosan, T. M., Rödenbeck, C., Schwinger, J., Smallman, T. L., Smith, S. M., Sospedra-Alfonso, R., Steinhoff, T., Sun, Q., Sutton, A. J., Séférian, R., Takao, S., Tatebe, H., Tian, H., Tilbrook, B., Torres, O., Tourigny, E., Tsujino, H., Tubiello, F., van der Werf, G., Wanninkhof, R., Wang, X., Yang, D., Yang,



- 682 X., Yu, Z., Yuan, W., Yue, X., Zaehle, S., Zeng, N., and Zeng, J.: Global Carbon Budget 2024, *Earth System*
683 *Science Data*, 17, 965-1039, <https://doi.org/10.5194/essd-17-965-2025>, 2024.
- 684 Fuentes Andrade, B., Buchwitz, M., Reuter, M., Bovensmann, H., Richter, A., Boesch, H., and Burrows, J. P.:
685 A method for estimating localized CO₂ emissions from co-located satellite XCO₂ and NO₂ images,
686 *Atmospheric Measurement Techniques*, 17, 1145-1173, <https://doi.org/10.5194/amt-17-1145-2024>, 2024.
- 687 Gaspari, G. and Cohn, S. E.: Construction of correlation functions in two and three dimensions, *Quarterly*
688 *Journal of the Royal Meteorological Society*, 125, 723-757, <https://doi.org/10.1002/qj.49712555417>, 1999.
- 689 Guo, M., Fang, S., Liu, S., Liang, M., Wu, H., Yang, L., Li, Z., Liu, P., and Zhang, F.: Comparison of atmospheric
690 CO₂, CH₄, and CO at two stations in the Tibetan Plateau of China, *Earth and Space Science*, 7,
691 e2019EA001051, <https://doi.org/10.1029/2019ea001051>, 2020.
- 692 Guo, W., Shi, Y., Liu, Y., and Su, M.: CO₂ emissions retrieval from coal-fired power plants based on OCO-2/3
693 satellite observations and a Gaussian plume model, *Journal of Cleaner Production*, 397, 136525,
694 <https://doi.org/10.1016/j.jclepro.2023.136525>, 2023.
- 695 Houtekamer, P. L. and Mitchell, H. L.: A Sequential ensemble Kalman filter for atmospheric data assimilation,
696 *Monthly Weather Review*, 129, 123-137, [https://doi.org/10.1175/1520-0493\(2001\)129<0123:ASEKFF>2.0.CO;2](https://doi.org/10.1175/1520-0493(2001)129<0123:ASEKFF>2.0.CO;2), 2001.
- 698 Hua, Y., Zhao, X., and Sun, W.: Estimation of anthropogenic CO₂ emissions at different scales for assessing
699 SDG indicators: Method and application, *Journal of Cleaner Production*, 414, 137547,
700 <https://doi.org/10.1016/j.jclepro.2023.137547>, 2023.
- 701 Hunt, B. R., Kostelich, E. J., and Szunyogh, I.: Efficient data assimilation for spatiotemporal chaos: a local
702 ensemble transform Kalman filter, *Physica D: Nonlinear Phenomena*, 230, 112-126,
703 <https://doi.org/10.1016/j.physd.2006.11.008>, 2007.
- 704 IPCC: 2019 Refinement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories,
705 Intergovernmental Panel on Climate Change, Switzerland, 2019.
- 706 Janssens-Maenhout, G., Pinty, B., Dowell, M., Zunker, H., Andersson, E., Balsamo, G., Bézy, J. L., Brunhes, T.,
707 Bösch, H., Bojkov, B., Brunner, D., Buchwitz, M., Crisp, D., Ciais, P., Counet, P., Dee, D., Denier van der Gon,
708 H., Dolman, H., Drinkwater, M. R., Dubovik, O., Engelen, R., Fehr, T., Fernandez, V., Heimann, M., Holmlund,
709 K., Houweling, S., Husband, R., Juvyns, O., Kentarchos, A., Landgraf, J., Lang, R., Löschner, A., Marshall, J.,
710 Meijer, Y., Nakajima, M., Palmer, P. I., Peylin, P., Rayner, P., Scholze, M., Sierk, B., Tamminen, J., and Veefkind,
711 P.: Toward an operational anthropogenic CO₂ emissions monitoring and verification support capacity,
712 *Bulletin of the American Meteorological Society*, 101, E1439-E1451, <https://doi.org/10.1175/bams-d-19-0017.1>, 2020.
- 714 Jiang, F., Wang, H., Chen, J. M., Ju, W., Tian, X., Feng, S., Li, G., Chen, Z., Zhang, S., Lu, X., Liu, J., Wang, H.,
715 Wang, J., He, W., and Wu, M.: Regional CO₂ fluxes from 2010 to 2015 inferred from GOSAT XCO₂ retrievals
716 using a new version of the Global Carbon Assimilation System, *Atmospheric Chemistry and Physics*, 21,
717 1963-1985, <https://doi.org/10.5194/acp-21-1963-2021>, 2021.
- 718 Juchem, H., Maier, F., Levin, I., Jordan, A., Pöhler, D., Rosendahl, C., Della Coletta, J., Preunkert, S., and
719 Hammer, S.: Challenges and benefits of using NO_x as a quantitative proxy for fossil fuel CO₂ in an urban area
720 based on radiocarbon measurements, *Atmospheric Chemistry and Physics*, 25, 18373-18388,
721 <https://doi.org/10.5194/acp-25-18373-2025>, 2025.
- 722 Konovalov, I. B., Berezin, E. V., Ciais, P., Broquet, G., Zhuravlev, R. V., and Janssens-Maenhout, G.: Estimation



- 723 of fossil-fuel CO₂ emissions using satellite measurements of "proxy" species, *Atmospheric Chemistry and*
724 *Physics*, 16, 13509-13540, <https://doi.org/10.5194/acp-16-13509-2016>, 2016.
- 725 Lei, R., Feng, S., Danjou, A., Broquet, G., Wu, D., Lin, J. C., O'Dell, C. W., and Lauvaux, T.: Fossil fuel CO₂
726 emissions over metropolitan areas from space: A multi-model analysis of OCO-2 data over Lahore, Pakistan,
727 *Remote Sensing of Environment*, 264, 112625, <https://doi.org/10.1016/j.rse.2021.112625>, 2021.
- 728 Li, H., Qiu, J., and Zheng, B.: Air-pollution-satellite-based CO₂ emission inversion: system evaluation,
729 sensitivity analysis, and future research direction, *Atmospheric Chemistry and Physics*, 25, 1949-1963,
730 <https://doi.org/10.5194/acp-25-1949-2025>, 2025.
- 731 Li, H., Zheng, B., Lei, Y., Hauglustaine, D., Chen, C., Lin, X., Zhang, Y., Zhang, Q., and He, K.: Trends and drivers
732 of anthropogenic NO_x emissions in China since 2020, *Environ Sci Ecotechnol*, 21, 100425,
733 [10.1016/j.ese.2024.100425](https://doi.org/10.1016/j.ese.2024.100425), 2024.
- 734 Li, M., Liu, H., Geng, G., Hong, C., Liu, F., Song, Y., Tong, D., Zheng, B., Cui, H., Man, H., Zhang, Q., and He, K.:
735 Anthropogenic emission inventories in China: a review, *National Science Review*, 4, 834-866,
736 <https://doi.org/10.1093/nsr/nwx150>, 2017a.
- 737 Li, M., Zhang, Q., Kurokawa, J.-i., Woo, J.-H., He, K., Lu, Z., Ohara, T., Song, Y., Streets, D. G., Carmichael, G.
738 R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S., Liu, F., Su, H., and Zheng, B.: MIX: a mosaic Asian
739 anthropogenic emission inventory under the international collaboration framework of the MICS-Asia and
740 HTAP, *Atmospheric Chemistry and Physics*, 17, 935-963, <https://doi.org/10.5194/acp-17-935-2017>, 2017b.
- 741 Li, S., Wang, S., Wu, Q., Zhang, Y., Ouyang, D., Zheng, H., Han, L., Qiu, X., Wen, Y., Liu, M., Jiang, Y., Yin, D.,
742 Liu, K., Zhao, B., Zhang, S., Wu, Y., and Hao, J.: Emission trends of air pollutants and CO₂ in China from 2005
743 to 2021, *Earth System Science Data*, 15, 2279-2294, <https://doi.org/10.5194/essd-15-2279-2023>, 2023.
- 744 Lian, J., Lauvaux, T., Utard, H., Bréon, F.-M., Broquet, G., Ramonet, M., Laurent, O., Albarus, I., Chariot, M.,
745 Kotthaus, S., Haeffelin, M., Sanchez, O., Perrussel, O., Denier van der Gon, H. A., Dellaert, S. N. C., and Ciais,
746 P.: Can we use atmospheric CO₂ measurements to verify emission trends reported by cities? Lessons from
747 a 6-year atmospheric inversion over Paris, *Atmospheric Chemistry and Physics*, 23, 8823-8835,
748 <https://doi.org/10.5194/acp-23-8823-2023>, 2023.
- 749 Lin, X., Ciais, P., Bousquet, P., Ramonet, M., Yin, Y., Balkanski, Y., Cozic, A., Delmotte, M., Evangeliou, N., Indira,
750 N. K., Locatelli, R., Peng, S., Piao, S., Saunio, M., Swathi, P. S., Wang, R., Yver-Kwok, C., Tiwari, Y. K., and Zhou,
751 L.: Simulating CH₄ and CO₂ over South and East Asia using the zoomed chemistry transport model LMDz-
752 INCA, *Atmospheric Chemistry and Physics*, 18, 9475-9497, <https://doi.org/10.5194/acp-18-9475-2018>,
753 2018.
- 754 Liu, S., Chu, Y., and Hu, J.: Clean air actions and air quality improvements — Beijing-Tianjin-Hebei and
755 surrounding areas, China, 2013-2019, *China CDC Weekly*, 2, 418-421, [10.46234/ccdcw2020.107](https://doi.org/10.46234/ccdcw2020.107), 2020.
- 756 Liu, X., Mizzi, A. P., Anderson, J. L., Fung, I. Y., and Cohen, R. C.: Assimilation of satellite NO₂ observations at
757 high spatial resolution using OSSEs, *Atmospheric Chemistry and Physics*, 17, 7067-7081, [10.5194/acp-17-7067-2017](https://doi.org/10.5194/acp-17-7067-2017), 2017.
- 759 Miyazaki, K., Eskes, H., Sudo, K., Boersma, K. F., Bowman, K., and Kanaya, Y.: Decadal changes in global
760 surface NO_x emissions from multi-constituent satellite data assimilation, *Atmospheric Chemistry and*
761 *Physics*, 17, 807-837, <https://doi.org/10.5194/acp-17-807-2017>, 2017.
- 762 Mustafa, F., Bu, L., Wang, Q., Yao, N., Shahzaman, M., Bilal, M., Aslam, R. W., and Iqbal, R.: Neural-network-
763 based estimation of regional-scale anthropogenic CO₂ emissions using an Orbiting Carbon Observatory-2



- 764 (OCO-2) dataset over East and West Asia, *Atmospheric Measurement Techniques*, 14, 7277–7290,
765 <https://doi.org/10.5194/amt-14-7277-2021>, 2021.
- 766 Nassar, R., Mastrogiacomo, J.-P., Bateman-Hemphill, W., McCracken, C., MacDonald, C. G., Hill, T., O'Dell, C.
767 W., Kiel, M., and Crisp, D.: Advances in quantifying power plant CO₂ emissions with OCO-2, *Remote Sensing*
768 *of Environment*, 264, 112579, <https://doi.org/10.1016/j.rse.2021.112579>, 2021.
- 769 Oney, B., Gruber, N., Henne, S., Leuenberger, M., and Brunner, D.: A CO-based method to determine the
770 regional biospheric signal in atmospheric CO₂, *Tellus B: Chemical and Physical Meteorology*, 69, 1353388,
771 <https://doi.org/10.1080/16000889.2017.1353388>, 2017.
- 772 Ou-Yang, C. F., Wang, J. L., Lin, C. C., Chiu, C. Y., Liu, W. T., Lan, X., Neff, D., Miller, J. B., Michel, S. E., Vaughn,
773 B. H., Chiu, Y. C., Yu, J. Y., Cheng, C. C., Schnell, R. C., White, J. W. C., and Lin, N. H.: Long-term measurements
774 of CO₂, CH₄, and isotopic ratios of CO₂ in the Western Pacific: trends, variations, and implications, *Journal*
775 *of Geophysical Research: Atmospheres*, 130, e2024JD042255, <https://doi.org/10.1029/2024jd042255>, 2025.
- 776 Parker, H. A., Laughner, J. L., Toon, G. C., Wunch, D., Roehl, C. M., Iraci, L. T., Podolske, J. R., McKain, K., Baier,
777 B., and Wennberg, P. O.: Inferring the vertical distribution of CO and CO₂ from TCCON total column values
778 using the TARDISS algorithm, *Atmospheric Measurement Techniques*, 16, 2601–2625,
779 <https://doi.org/10.5194/amt-16-2601-2023>, 2023.
- 780 Peng, L., Zhang, Q., Yao, Z., Mauzerall, D. L., Kang, S., Du, Z., Zheng, Y., Xue, T., and He, K.: Underreported
781 coal in statistics: A survey-based solid fuel consumption and emission inventory for the rural residential
782 sector in China, *Applied Energy*, 235, 1169–1182, <https://doi.org/10.1016/j.apenergy.2018.11.043>, 2019.
- 783 Pisso, I., Patra, P., Takigawa, M., Machida, T., Matsueda, H., and Sawa, Y.: Assessing Lagrangian inverse
784 modelling of urban anthropogenic CO₂ fluxes using in situ aircraft and ground-based measurements in the
785 Tokyo area, *Carbon Balance Manag*, 14, 6, <https://doi.org/10.1186/s13021-019-0118-8>, 2019.
- 786 Randerson, J. T., van der Werf, G. R., Giglio, L., Collatz, G. J., and Kasibhatla, P. S.: Global Fire Emissions
787 Database, Version 4.1 (GFED v4) [dataset], <https://doi.org/10.3334/ORNLDAAAC/1293>, 2017.
- 788 Schuh, A. E., Jacobson, A. R., Basu, S., Weir, B., Baker, D., Bowman, K., Chevallier, F., Crowell, S., Davis, K. J.,
789 Deng, F., Denning, S., Feng, L., Jones, D., Liu, J., and Palmer, P. I.: Quantifying the impact of atmospheric
790 transport uncertainty on CO₂ surface flux estimates, *Global Biogeochem Cycles*, 33, 484–500,
791 <https://doi.org/10.1029/2018GB006086>, 2019.
- 792 Skamarock, W. C. and Klemp, J. B.: A time-split nonhydrostatic atmospheric model for weather research and
793 forecasting applications, *Journal of Computational Physics*, 227, 3465–3485,
794 <https://doi.org/10.1016/j.jcp.2007.01.037>, 2008.
- 795 Super, I., Scarpelli, T., Droste, A., and Palmer, P. I.: Improved definition of prior uncertainties in CO₂ and CO
796 fossil fuel fluxes and its impact on multi-species inversion with GEOS-Chem (v12.5), *Geoscientific Model*
797 *Development*, 17, 7263–7284, <https://doi.org/10.5194/gmd-17-7263-2024>, 2024.
- 798 Taylor, T. E., O'Dell, C. W., Baker, D., Bruegge, C., Chang, A., Chapsky, L., Chatterjee, A., Cheng, C., Chevallier,
799 F., Crisp, D., Dang, L., Drouin, B., Eldering, A., Feng, L., Fisher, B., Fu, D., Gunson, M., Haemmerle, V., Keller,
800 G. R., Kiel, M., Kuai, L., Kurosu, T., Lambert, A., Laughner, J., Lee, R., Liu, J., Mandrake, L., Marchetti, Y.,
801 McGarragh, G., Merrelli, A., Nelson, R. R., Osterman, G., Oyafuso, F., Palmer, P. I., Payne, V. H., Rosenberg,
802 R., Somkuti, P., Spiers, G., To, C., Weir, B., Wennberg, P. O., Yu, S., and Zong, J.: Evaluating the consistency
803 between OCO-2 and OCO-3 XCO₂ estimates derived from the NASA ACOS version 10 retrieval algorithm,
804 *Atmos. Meas. Tech.*, 16, 3173–3209, [10.5194/amt-16-3173-2023](https://doi.org/10.5194/amt-16-3173-2023), 2023.



- 805 Wang, M., Wang, S., Zhao, J., Ju, W., and Hao, Z.: Global positive gross primary productivity extremes and
806 climate contributions during 1982-2016, *Sci Total Environ*, 774, 145703,
807 <https://doi.org/10.1016/j.scitotenv.2021.145703>, 2021.
- 808 Whitaker, J. S. and Hamill, T. M.: Ensemble data assimilation without perturbed observations, *Monthly*
809 *Weather Review*, 130, 1913-1924, [https://doi.org/10.1175/1520-0493\(2002\)130<1913:EDAWPO>2.0.CO;2](https://doi.org/10.1175/1520-0493(2002)130<1913:EDAWPO>2.0.CO;2),
810 2002.
- 811 World Meteorological Organization (WMO): State of the Global Climate 2024, World Meteorological
812 Organization, Geneva WMO-No. 1368, 2025.
- 813 Xu, J., Zhang, Z., Zhao, X., and Cheng, S.: Downward trend of NO₂ in the urban areas of Beijing-Tianjin-Hebei
814 region from 2014 to 2020: Comparison of satellite retrievals, ground observations, and emission inventories,
815 *Atmospheric Environment*, 295, 119531, [10.1016/j.atmosenv.2022.119531](https://doi.org/10.1016/j.atmosenv.2022.119531), 2023.
- 816 Yang, H., Wu, K., Shen, H., Janssens-Maenhout, G., Crippa, M., Guizzardi, D., and Zhou, M.: A comparative
817 analysis of China's anthropogenic CO₂ emissions (2000-2023): insights from six bottom-up inventories and
818 uncertainty assessment, *Atmospheric Chemistry and Physics*, 25, 18111-18127,
819 <https://doi.org/10.5194/acp-25-18111-2025>, 2025.
- 820 Zhang, L., Jiang, F., He, W., Wu, M., Wang, J., Ju, W., Wang, H., Zhang, Y., Sitch, S., and Chen, J. M.: Improved
821 estimates of net ecosystem exchanges in mega-countries using GOSAT and OCO-2 observations,
822 *Communications Earth & Environment*, 5, 737, <https://doi.org/10.1038/s43247-024-01910-w>, 2024a.
- 823 Zhang, S., Lei, L., Sheng, M., Song, H., Li, L., Guo, K., Ma, C., Liu, L., and Zeng, Z.: Evaluating anthropogenic
824 CO₂ bottom-up emission inventories using satellite observations from GOSAT and OCO-2, *Remote Sensing*,
825 14, 5024, <https://doi.org/10.3390/rs14195024>, 2022.
- 826 Zhang, Z., Feng, S., Chen, Y., Liu, Q., Ju, W., Xiao, W., Huang, C., Wang, Y., Wang, H., Jia, M., Wang, X., and
827 Jiang, F.: Development of a regional carbon assimilation system and its application for estimating fossil fuel
828 carbon emissions in the Yangtze River Delta, China, *Sci Total Environ*, 957, 177720,
829 <https://doi.org/10.1016/j.scitotenv.2024.177720>, 2024b.
- 830 Zhang, Z., Zhou, Y., Zhao, N., Li, H., Tohniyaz, B., Mperejekumana, P., Hong, Q., Wu, R., Li, G., Sultan, M.,
831 Zayan, A. M. I., Cao, J., Ahmad, R., and Dong, R.: Clean heating during winter season in Northern China: A
832 review, *Renewable and Sustainable Energy Reviews*, 149, 111339,
833 <https://doi.org/10.1016/j.rser.2021.111339>, 2021.
- 834 Zheng, B., Chevallier, F., Ciais, P., Broquet, G., Wang, Y., Lian, J., and Zhao, Y.: Observing carbon dioxide
835 emissions over China's cities and industrial areas with the Orbiting Carbon Observatory-2, *Atmospheric*
836 *Chemistry and Physics*, 20, 8501-8510, <https://doi.org/10.5194/acp-20-8501-2020>, 2020.
- 837 Zheng, B., Cheng, J., Geng, G., Wang, X., Li, M., Shi, Q., Qi, J., Lei, Y., Zhang, Q., and He, K.: Mapping
838 anthropogenic emissions in China at 1 km spatial resolution and its application in air quality modeling, *Sci*
839 *Bull (Beijing)*, 66, 612-620, <https://doi.org/10.1016/j.scib.2020.12.008>, 2021.