

Measurement report: Dust Transport and Local Anthropogenic Emissions Differently Shape Atmospheric Ice-Nucleating Particles in an Industrial Urban Atmosphere

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Abstract. Atmospheric ice-nucleating particles (INPs) are vital for cloud formation, yet the importance of INPs from anthropogenic sources remains poorly understood. We conducted a month-long winter field campaign in Taiyuan (China), a heavily industrialized city, to quantify INP concentrations (N_{INP}) and ice nucleation active site density (n_s) of immersion mode INPs, alongside particle size distributions and chemical compositions. Our results indicate that N_{INP} ranged from 0.0532 to 13.4 L⁻¹ at and above -15 °C, corresponding to n_s values of 10⁵–10⁷ m⁻². During the identified desert dust event, both N_{INP} (7.47 L⁻¹; 95% confidence interval (CI): 6.64–8.41 L⁻¹) and n_s (1.57×10^7 m⁻²) were substantially higher than during the remaining observation periods without clear desert-dust intrusion (0.607 L⁻¹ and 9.52×10^5 m⁻², respectively), highlighting the strong influence of long-range transported desert dust. In contrast, during pollution periods, N_{INP} showed only weak correlations with urban aerosol components such as sulfate (SO₄²⁻), nitrate (NO₃⁻), and organic carbon (OC) ($|r| < 0.3$). Positive matrix factorization (PMF) identified five PM_{2.5} source factors: industrial emissions, dust-related particles, secondary aerosols, coal combustion and traffic emissions, and fireworks. Although these factors dominated the PM_{2.5} mass, they showed no significant covariation with N_{INP} . During this winter campaign, N_{INP} variability in Taiyuan was characterized by strong enhancement during the identified desert dust event and weak associations with the major PM_{2.5} source factors during the remaining observation periods.

Short summary

In clouds containing liquid water and ice (about 0 to -38 °C), ice-nucleating particles act as tiny “seeds” that help supercooled water freeze. We studied these particles in winter air in Taiyuan, an industrial city in northern China. The strongest increases occurred during a desert dust intrusion, while haze and pollution showed weak links to fine-particle pollution sources. This suggests urban freezing particles can be strongly affected by transported dust, not pollution intensity alone.

34 1 Introduction

35 Atmospheric ice-nucleating particles (INPs) play a critical role in cloud formation, cloud radiative properties and
36 precipitation through heterogeneous ice nucleation (Lohmann and Feichter, 2005; IPCC, 2014; Lohmann et al., 2016).
37 Although INPs constitute a minute fraction of ambient aerosols under natural conditions (approximately 1 in 10^5 ambient
38 particles in the free troposphere (DeMott et al., 2010)), their ability to modulate cloud properties and lifetimes has received
39 substantial scientific interest. Heterogeneous ice nucleation is generally understood to occur through four mechanisms:
40 deposition ice nucleation, condensation freezing, immersion freezing and contact freezing (Hoose and Möhler, 2012; Vali et
41 al., 2015). In addition, recent studies have suggested that under water-subsaturated conditions, pore condensation and freezing
42 (PCF) may provide a more realistic description of atmospheric ice nucleation than classical deposition nucleation in some
43 cases (Marcolli, 2014; David et al., 2019). Among these, immersion freezing has been widely reported to be the most important
44 ice nucleation mechanism for mixed-phase clouds (Ansmann et al., 2008; Murray et al., 2012; Westbrook and Illingworth,
45 2013).

46 The urban atmosphere is characterized by intensive anthropogenic activities, diverse aerosol emission sources, elevated
47 aerosol loadings, and complex chemical compositions. However, the key factors controlling atmospheric INP concentrations
48 in such environments, including aerosol composition, size, and source characteristics, remain uncertain (Kanji et al., 2017).
49 Previous studies in Beijing and Tokyo demonstrated that INP concentrations often show no significant correlations with black
50 carbon (BC) or $\text{PM}_{2.5}$ mass, or the number concentration of particles larger than $0.5\ \mu\text{m}$ during air pollution episodes (Chen
51 et al., 2018; Tobo et al., 2020; Zhang et al., 2022). This suggests that typical urban pollution aerosols are generally inefficient
52 INPs.

53 However, this observation contrasts with laboratory and field evidence showing that some anthropogenic particles can
54 exhibit immersion-mode ice-nucleating activity. These include certain fly ash, metallic/mineral particles, and other industrial
55 dusts associated with coal combustion and manufacturing activities (Toll et al., 2024). By contrast, the contribution of soot
56 particles to atmospheric INPs in the immersion mode has been suggested to be negligible in many cases (Kanji et al., 2020).
57 Moreover, biomass burning represents another potential source, with agricultural fires near Mexico City reported to induce
58 freezing at temperatures above $-15\ ^\circ\text{C}$, likely due to the presence of associated biological particles (McCluskey et al., 2014;
59 Jahn et al., 2020; Cabrera-Segoviano et al., 2022). The complexity of urban INPs is further exemplified by observations in
60 New Delhi, where INP concentrations correlated moderately ($r = 0.52$) with BC during fog episodes, indicating that particle
61 composition and meteorological conditions may modulate INP activity (Wagh et al., 2021). This gap between laboratory
62 potential and field observations indicates that urban INP variability may be governed by specific high-efficiency contributors
63 rather than total aerosol mass or number concentrations. Therefore, mineral dust and biological particles, which have long
64 been established as the most potent atmospheric INPs, require careful consideration even in heavily industrialized areas.

65 Mineral dust is widely recognized as the most important atmospheric INP due to its vast global emissions, estimated at up
66 to $\sim 5000\ \text{Tg yr}^{-1}$ (Engelstaedter et al., 2006). Its intrinsically high ice-nucleating activity at temperatures below $-8\ ^\circ\text{C}$ is

67 primarily attributed to specific mineral components such as K-feldspar, which has been systematically identified as a key
68 active species in laboratory settings (Atkinson et al., 2013; Augustin-Bauditz et al., 2014; Zolles et al., 2015). In the context
69 of East Asia, mineral dust transported from northwestern deserts can substantially enhance INP concentrations in downwind
70 urban areas, with ice-nucleating efficiencies comparable to those reported for Saharan dust (Price et al., 2018; Reicher et al.,
71 2019; Chen et al., 2021). Furthermore, the influence of dust in urban environments is not limited to long-range natural transport;
72 anthropogenic dust may also contribute to atmospheric INPs in some urban and peri-urban regions, although direct field
73 evidence remains limited (Chen et al., 2024).

74 In comparison to mineral dust, aerosols of biological origin, such as pollen, fungal spores, proteins, and other
75 macromolecular organics can nucleate ice at relatively warmer temperatures, often above -15°C (Pummer et al., 2015; Polen
76 et al., 2016; Kanji et al., 2017). While these bioaerosols may be less abundant by mass than mineral dust, their exceptional ice-
77 nucleating efficiency means that even a small population can dominate the urban INP budget at the early stages of cloud
78 glaciation.

79 Although considerable progress has been made in identifying both natural and anthropogenic sources of atmospheric INPs,
80 their formation pathways and source attributions in urban environments remain poorly constrained due to the complex mixture
81 of aerosol types and variable emission patterns. To address this gap, a comprehensive winter field campaign was conducted in
82 Taiyuan, a representative industrial city in northern China, over a one-month period. Taiyuan frequently experiences severe
83 particulate pollution and episodic dust events, positioning it as an ideal location to investigate the combined effects of
84 anthropogenic and natural emissions. In this study, the number concentration of immersion mode INP (N_{INP}) and the number
85 of active sites per unit surface area of INPs (n_s) were measured under mixed-phase cloud conditions. The freezing experiments
86 were continued until all droplets were frozen, but because most samples were already fully or nearly fully frozen by about
87 -25°C , the INP data presented and discussed in this study mainly span the temperature range from about -25°C to -5°C .
88 Source apportionment of $\text{PM}_{2.5}$ using a positive matrix factorization (PMF) model enabled the identification of major emission
89 sources and their temporal variations.

90 **2 Methods**

91 **2.1 Sampling**

92 Taiyuan is a major industrial city in northern China, often experiencing elevated concentrations and enrichment factors (EFs)
93 of $\text{PM}_{2.5}$ -bound heavy metals, influenced by local emissions, regional transport, and complex terrain (Li et al., 2021). A winter
94 observation campaign was conducted from 3 December 2023 to 14 January 2024, coinciding with the centralized heating
95 season characterized by intensive coal combustion. The sampling site (37.88°N , 112.55°E ; approximately 800 m a.s.l.) is
96 located on the rooftop (~ 20 m above ground level) of a four-story building at the Taiyuan Municipal Ecology and Environment
97 Bureau. Situated near Taoyuan 3rd Alley, this site is representative of a typical central urban environment. Although it is

removed from direct major industrial point sources, it remains suitable for monitoring urban fine particulate pollution driven by mixed anthropogenic activities.

Key atmospheric pollutants were continuously monitored at the station, including concentrations of water-soluble ions, heavy metals, elemental carbon (EC) and organic carbon (OC), gas concentrations (SO_2 , NO_x , O_3 , CO), particulate matter concentrations ($\text{PM}_{2.5}$ and PM_{10}), together with meteorological parameters (wind speed, wind direction, temperature, relative humidity, and atmospheric pressure). In addition, INP filter samples were collected using a two-channel sampler equipped with two parallel filter channels from 4 December 2023 to 5 January 2024. Polycarbonate filters (47 mm, Nuclepore Track-Etch Membrane, 0.2 μm pore size, Whatman) were used for sampling. The sampler was operated without a size-selective inlet. The sampling flow rate was maintained at 6 L min^{-1} by a mass flow controller (MFC), with a 24 h sampling duration. Considering flow fluctuations and filter switching time, the total sampled air volume for each filter was approximately 8000 L. Except for the filter samples, all observation data were recorded as hourly averages.

2.2 Instrumentation

2.2.1 Particle mass and size distribution measurement

Mass concentrations of particulate matter (PM) with aerodynamic diameter (d_a) smaller than 2.5 and 10 μm ($\text{PM}_{2.5}$ and PM_{10} , respectively) were measured by a tapered element oscillating microbalance (TEOM) monitor.

Two scanning mobility particle sizer (SMPS) systems were employed to measure particle number size distributions in the 2.02–514 nm range during the campaign, while the range of data used in this study was 3–453 nm. The first system consisted of a TSI Classifier 3080 with Differential Mobility Analyzer (DMA) 3085 and Condensation Particle Counter (CPC) 3776, while the second system included a TSI Classifier 3082 with DMA 3081 and CPC 3772. Sample humidity was regulated using a NafionTM membrane dryer (Perma Pure, MD-700-24F-3) to prevent water vapor condensation. Multiple charging corrections were applied during data processing (Wiedensohler, 1988).

An Optical Particle Counter (Aerosol Spectrometer and Dust Monitor 1.108, Grimm Aerosol Technik) was used to obtain number size distribution (PNSD) of particles ranging from 0.5 to 20 μm in optical-equivalent diameters. To merge the OPC data with the SMPS measurements and calculate particle surface area on the same diameter basis, the OPC size distribution was converted to mobility diameter (d_m). Estimating d_a from optical-equivalent diameter typically requires knowledge of particle complex refractive index and morphological parameters (e.g., dynamic shape factor) (Peters et al., 2006). Moreover, when converting to d_m , we assumed an effective particle density of 1.5 g cm^{-3} and a shape factor of unity. Previous sensitivity calculations showed that changing the particle density from 1.5 to 2.6 g cm^{-3} and the dynamic shape factor from 1.1 to 1.4 can lead to a factor of 1.07–2.36 difference in n_s (Chen et al., 2021). This suggests that diameter conversion may introduce an uncertainty of approximately a factor of 2 in n_s , particularly in the coarse-mode size range (Huang et al., 2021). A more exact uncertainty could not be assigned because particle refractive index, morphology, and particle-specific effective density were not independently measured. The OPC data were combined with the SMPS measurements to construct a merged particle size

130 distribution (SMPS: 3–453 nm; OPC: 0.5–20 μm), which was further used to estimate the total particle surface area for n_s
131 calculation.

132 2.2.2 Particle chemical composition

133 The water-soluble inorganic ions in $\text{PM}_{2.5}$ were measured in-situ using a Monitor for AeRosols and Gases in ambient Air
134 (MARGA 1S; Metrohm AG, Switzerland). The analysis focused on major inorganic ions, including NO_3^- , SO_4^{2-} , NH_4^+ , Cl^- ,
135 K^+ , Ca^{2+} , Na^+ and Mg^{2+} . In parallel, the MARGA simultaneously measured the gaseous pollutants including SO_2 , HNO_2 , HNO_3 ,
136 NH_3 , and HCl . EC and OC concentrations were determined using a Carbonaceous Aerosol Speciation System (CASS; Aerosol
137 d.o.o., Slovenia, EU) (Rumsey et al., 2014; Rumsey and Walker, 2016; Wan et al., 2022), following the thermal–optical
138 transmittance method. Elemental composition of aerosol samples was analyzed by online X-ray fluorescence spectrometry
139 (XRF) (Marguí et al., 2022), allowing quantification of a wide range of elements including light metals (K, Ca, etc.), metalloids
140 (As, Se, etc.), and heavy metals (Zn, Pb, Cu, Hg, etc.).

141 2.2.3 Ice-nucleating particle (INP) concentration

142 N_{INP} were determined using the Peking University Ice Nucleation Array (PKU-INA), a cold-stage-based ice nucleation
143 instrument (Chen et al., 2018, Chen et al., 2021). Each filter was fully immersed in 20 mL of double-distilled water (resistivity:
144 $18.2 \text{ M}\Omega \text{ cm}^{-1}$ at 25°C), and extracted using an ultrasonic shaker for 30 min. Ultrasonic extraction is an effective method for
145 removing particles from filters (Ardon-Dryer and Levin, 2014; Chen et al., 2018; Reicher et al., 2019; Chen et al., 2021),
146 although it may alter the bioactivity of proteins (De Leo et al., 2017). To minimize ultrasound-induced heating, the extraction
147 was performed in a water bath with ice. The resulting suspension was dispensed into 90 droplets (1 μL each) on a hydrophobic
148 glass slide. A metal spacer was placed on the slide to assist in droplet positioning and to suppress the Wegener–Bergeron–
149 Findeisen process (Jung et al., 2012). A top glass was placed on the spacer to seal the droplets. The cold stage was first cooled
150 from room temperature to 0°C at $20^\circ\text{C min}^{-1}$, and then at 1°C min^{-1} until all droplets were frozen. The image of each droplet
151 was recorded by a CCD camera every 6 seconds, corresponding to a temperature resolution of 0.1°C . These images were later
152 analyzed by customized software to identify the phase transition of each droplet by its brightness change upon freezing. High-
153 purity dry nitrogen was used as sheath gas to prevent condensation and frost formation. The temperature uncertainty at the
154 applied cooling rate was within $\pm 0.4^\circ\text{C}$ (Chen et al., 2018).

155 The cumulative number of ice-active sites per unit droplet volume above temperature T was calculated according to Vali et
156 al. (2015):

$$157 \quad K(T) = -\frac{\ln(1-f_{\text{ice}}(T))}{V} \quad (\text{cm}^{-3} \text{ of water}), \quad (1)$$

158 where $f_{\text{ice}}(T)$ is the fraction of droplets frozen at T , and V is the volume of each pipetted droplet (1 μL in this study). Combined
159 with the total sampling volume, the $N_{\text{INP}}(T)$ per unit volume of sampled air is calculated as

$$N_{\text{INP}}(T) = -\frac{\ln(1-f_{\text{ice}}(T))}{V_{\text{air}}} \text{ (L}^{-1} \text{ air)}, \quad (2)$$

where V_{air} is the total volume of sampled air per droplet converted to standard conditions (0 °C, 1013 hPa) during each sample collection period. To quantify and compare the ice-nucleating activity of aerosol particle with different sizes, the n_s (Vali et al., 2015), i.e., the cumulative ice-nucleation-active-site surface density (Connolly et al., 2009; Niedermeier et al., 2011; Hoose and Möhler, 2012), is calculated from the $N_{\text{INP}}(T)$ as

$$n_s(T) = \frac{N_{\text{INP}}(T)}{A} \text{ (m}^{-2}\text{)}, \quad (3)$$

where A is the total surface area of the particles per unit volume of sampled air, derived from size distribution measurements of the particles. For $n_s(T)$ calculation, the merged SMPS–OPC particle number size distributions were converted to the same standard-condition air-volume basis as N_{INP} (0 °C, 1013 hPa) using concurrent temperature and pressure data. The particle surface area was then estimated from the converted size distributions expressed in mobility diameter.

The main source of uncertainty in the experimental results stems from the statistical representativeness of the droplets analyzed relative to the entire particle suspension. Due to the low ambient concentrations of INPs, the number of INPs present in the resulting particle suspension is typically small. As each droplet has a volume of only 1 μL , individual droplets may not adequately reflect the true particle distribution in the bulk extract. Moreover, the total number of droplets tested per sample (90 in this study) imposes a statistical limitation. To account for this, confidence intervals for the estimated INP concentration per droplet (and per unit sample volume) were calculated following the statistical approach described in previous studies (Barker, 2002; O’Sullivan et al., 2018), as shown in Eq. (4):

$$\mu(T) \pm \frac{(Z_{\alpha/2})^2}{2n} \pm Z_{\alpha/2} \left[4\mu + \frac{(Z_{\alpha/2})^2}{2n} \right]^{0.5} / (4n)^{0.5} \quad (4)$$

where $\mu(T)$ represents the number of INPs per droplet, n is the total number of droplets analyzed, and $Z_{\alpha/2}$ denotes the standard normal score corresponding to the desired confidence level (1.96 for a 95 % confidence interval).

For each aerosol sample, the parameters ($f_{\text{ice}}(T)$, $K(T)$, $N_{\text{INP}}(T)$, $n_s(T)$, and the confidence intervals for $N_{\text{INP}}(T)$) were calculated individually. The $f_{\text{ice}}(T)$ was derived from cold-stage measurements, V_{air} was obtained from recorded sampling volumes, and A was calculated from the standard-condition combined particle size distributions measured by the SMPS and OPC. The uncertainty in n_s arises from both the statistical uncertainty in N_{INP} and the uncertainty in the particle surface area concentration used for normalization. The latter is mainly associated with the SMPS–OPC merging and diameter conversion in the coarse-mode range, as discussed in Section 2.2.1.

2.2.4 Air mass backward trajectory analysis

Backward trajectory analysis and visualization were performed using the MeteoInfoMap program (Wang, 2014), which is based on the HYSPLIT4 (Hybrid Single Particle Lagrangian Integrated Trajectory) model developed by the NOAA (National Oceanic and Atmospheric Administration, USA) Air Resources Laboratory. The analysis used GDAS (Global Data

190 Assimilation System, 1°, global) meteorological data from NOAA (Stein et al., 2015). Trajectories were calculated with a
 191 duration of 72 h and an interval of 6 h between each trajectory, arriving at an altitude of 850 m above sea level (a.s.l.), which
 192 approximately represents the actual sampling inlet height, considering a ground elevation of about 800 m a.s.l. and a rooftop
 193 sampling height of about 20 m above ground level.

194 2.2.5 Positive matrix factorization (PMF)

195 The Positive Matrix Factorization (PMF) model was applied to analyze the sources of PM_{2.5}. PMF is a widely used receptor
 196 model for source apportionment. It can identify the composition and contribution of various sources. The model assumes that
 197 species emitted from the same source exhibit strong correlations, enabling the extraction of meaningful factors from measured
 198 concentration data (Karagulian et al., 2015; US EPA, 2015).

199 Given an $i \times j$ matrix $\mathbf{X}$ consisting of concentrations of j chemical species across i samples, PMF decomposes the matrix
 200 into two positive matrices—factor contributions ($\mathbf{G}$) and factor profiles ($\mathbf{F}$)—according to Equation (5):

$$201 \quad X_{ij} = \sum_{k=1}^p \mathbf{G}_{ik} \mathbf{F}_{kj} + \mathbf{E}_{ij} \quad (5)$$

202 where p is the number of factors, i denotes the sample index, and j represents the species index. X_{ij} is the measured
 203 concentration of species j in sample i ; $\mathbf{G}_{ik}$ represents the contribution of factor k to sample i ; $\mathbf{F}_{jk}$ is the mass fraction of species
 204 j in factor k ; and $\mathbf{E}_{ij}$ denotes the residuals. Both $\mathbf{G}_{ik}$ and $\mathbf{F}_{jk}$ are constrained to be non-negative.

205 To obtain the optimal solution, PMF minimizes an objective function Q (Eq. 6) using uncertainty estimates u_{ij} and non-
 206 negativity constraints on $\mathbf{G}$ and $\mathbf{F}$:

$$207 \quad Q = \sum_{i=1}^m \sum_{j=1}^n \left(\frac{X_{ij} - \sum_{k=1}^p \mathbf{G}_{ik} \mathbf{F}_{kj}}{u_{ij}} \right)^2 \quad (6)$$

208 The uncertainty u_{ij} is estimated following the EPA PMF 5.0 User Guide (US EPA, 2015) as:

$$209 \quad u_{ij} = \begin{cases} \frac{5}{6} \times \text{MDL} & (c \leq \text{MDL}) \\ \sqrt{(ef \times X_{ij})^2 + (0.5 \times \text{MDL})^2} & (c > \text{MDL}) \end{cases} \quad (7)$$

210 where c is the measured concentration of a given species, MDL is the method detection limit, and ef is the analytical
 211 uncertainty, typically between 0.05 and 0.2. In this study, $ef = 0.1$ was applied.

212 EPA PMF 5.0 software was employed for source apportionment of PM_{2.5} in the study region. The specific datasets and their
 213 sources used for PMF analysis are described in Section 3.4. To determine the optimal number of factors, 4-7 factors were
 214 tested with 20 iterative runs each, and the solution with the most physically interpretable source profiles and lowest residuals
 215 was selected for final analysis (US EPA, 2015). Consequently, a five-factor solution was determined to be the optimal choice.

217 **3.1 Overview of INP temporal variability and aerosol characteristics**

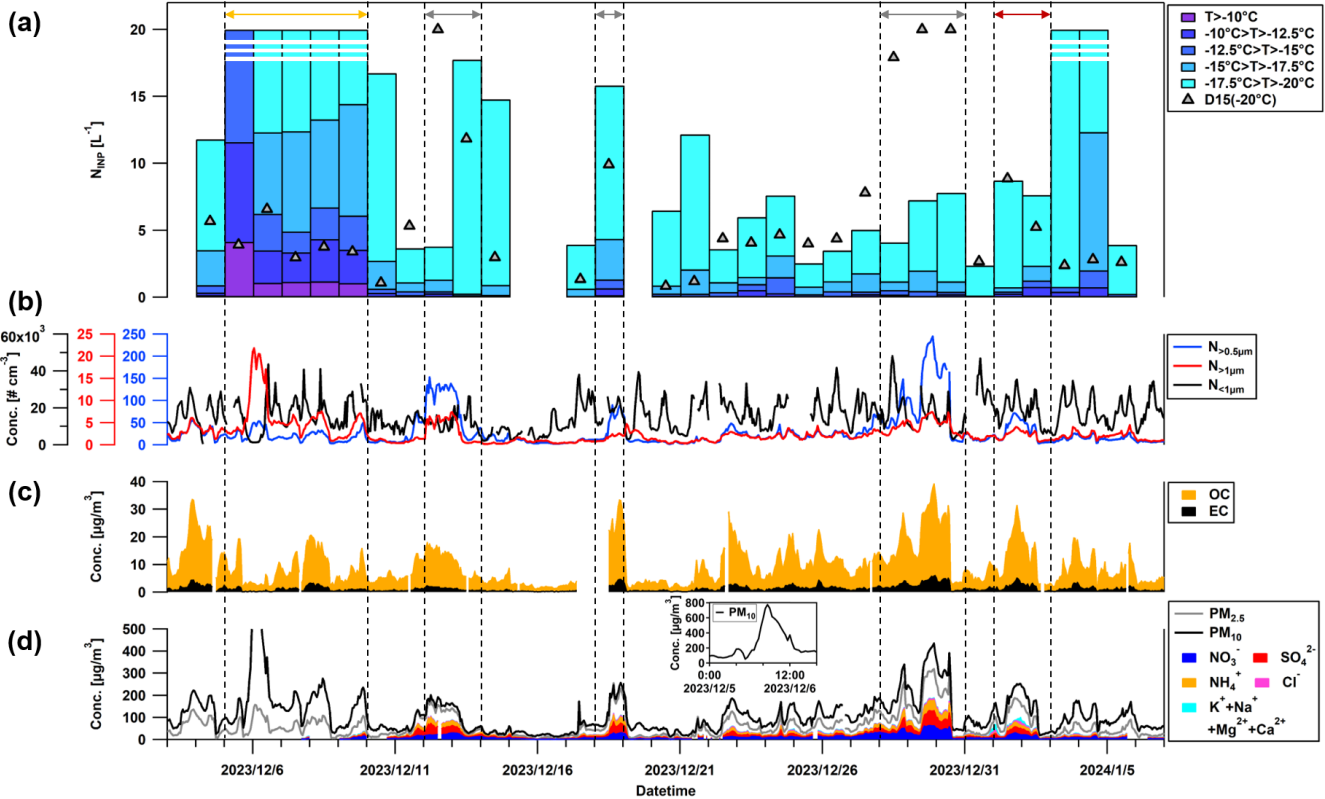
218 Figure 1 presents the temporal evolution of atmospheric INP concentrations and relevant aerosol parameters during the
219 observation period in Taiyuan, including size-segregated particle number concentrations, carbonaceous components (OC and
220 EC), as well as mass concentrations of PM_{2.5} and PM₁₀, and water-soluble ions. Daily N_{INP} at different temperatures are shown
221 in Fig. 1a (data for 15, 16, and 19 December are missing). At and above −15 °C, the cumulative N_{INP}(T), representing the total
222 INP concentration active at and above this temperature, ranged from 10^{−2} to 10¹ L^{−1} with a campaign mean of 1.75 L^{−1}. The
223 corresponding 95% confidence intervals for N_{INP} at selected temperatures are provided in Table S3. Correspondingly, the n_s
224 values at and above −15 °C (see Fig. 3) spanned 10⁵ to 10⁷ m^{−2}, with a campaign mean of 3.42 × 10⁶ m^{−2}.

225 Several distinct events can be identified throughout the campaign, as highlighted in Fig. 1 by dashed boundary lines and
226 colored arrows. The first event (5-9 December, marked by the orange arrow in Fig. 1a) was characterized by a sharp increase
227 in coarse-mode particle concentrations, represented here by N_{>1μm} derived from the merged SMPS–OPC size distribution (Figs.
228 1b, S1b, and S2). During this episode, the mean N_{>1μm} increased to 5.94 cm^{−3}, which was approximately 2.7 times higher than
229 the average concentration during the remaining observation periods (2.21 cm^{−3}). A peak daily concentration of 12.08 cm^{−3} was
230 recorded on 6 December, accompanied by a steep rise in PM₁₀ that reached nearly 800 μg m^{−3} (Figs. 1b and 1d).

231 72-hour back trajectory analysis by HYSPLIT model confirmed that air masses during this period predominantly originated
232 from the Taklamakan, Gurbantünggüt, or Badain Jaran Deserts in northwestern China, with trajectory paths highly overlapping
233 the desert regions (Fig. S3). This pattern is consistent with desert dust transport events previously observed in northern China
234 (Chen et al., 2021), further supporting the identification of this period as a desert dust intrusion event. Wind rose analysis (Fig.
235 S6) shows that both the dust-event and non-desert-dust periods were dominated by northerly flow, whereas the dust-event
236 period exhibited a more concentrated northerly flow pattern. This serves as supplementary meteorological support for the event
237 classification. Given the well-established high ice-nucleating activity of mineral dust, this episode is identified as a desert dust-
238 intrusion event responsible for the pronounced enhancement in N_{INP}, especially at temperatures > −12.5 °C. The remaining
239 observation periods are not assumed to be free of mineral dust influence; rather, no similarly clear desert dust intrusion
240 signature was identified outside this event. Further discussion of its characteristics and implications is provided in Section 3.2.

241 Other episodes (13-14, 18, and 28-30 December, marked by the gray arrows in Fig. 1a) were characterized by simultaneous
242 increases in sulfate, nitrate, ammonium, and organic carbon, accompanied by elevated PM_{2.5} and PM₁₀ levels (Fig. 1c and Fig.
243 1d). This indicates the influence of air pollution driven by regional secondary aerosol accumulation under stagnant winter
244 meteorology. Moderate N_{INP} increases were observed during the earlier two episodes (13–14 and 18 December), suggesting
245 that secondary aerosol formation may indirectly modulate ice-nucleating activity through coating or mixing with pre-existing
246 ice-active particles. The late-December event showed no corresponding N_{INP} enhancement despite high PM levels, indicating
247 that aerosol mass concentration alone may not determine INP abundance. This contrast emphasizes that particle composition

248 and surface properties, rather than overall pollution intensity, govern INP activity during pollution episodes characterized by
 249 enhanced secondary inorganic and organic aerosol accumulation (Chen et al., 2018; Wagh et al., 2021; Zhang et al., 2022).



250
 251 **Figure 1.** Time series of (a) temperature-resolved N_{INP} , (b) size-resolved particle number concentrations, (c) mass concentrations of OC and
 252 EC, and (d) mass concentrations of $PM_{2.5}$, PM_{10} and major water-soluble ions (K^+ , Na^+ , Mg^{2+} , Ca^{2+} , Cl^- , NH_4^+ , SO_4^{2-} , NO_3^-) measured by
 253 MARGA during the observation period. In panel (a), gray triangles ($D15(-20^{\circ}C)$) represent values calculated using the parameterization of
 254 DeMott et al. (2015), shown here mainly as a dust-based reference for comparison with the observed day-to-day temporal variation in INP
 255 concentrations. Special periods are indicated by dashed boundary lines across all panels and colored arrows at the top of panel (a): the orange
 256 arrow denotes the desert dust period, the gray arrows denote pollution periods, and the red arrow denotes the fireworks event. In panel (a),
 257 white horizontal lines indicate samples for which all droplets had already frozen before the colder temperature intervals were reached. Under
 258 these conditions, the cumulative N_{INP} exceeded the quantifiable range of the assay and therefore could not be plotted as a finite value. In
 259 panel (b), $N_{<1\mu m}$ refers to particles from 3 nm to 1 μm , whereas $N_{>0.5\mu m}$ and $N_{>1\mu m}$ denote the concentrations integrated over 0.5–20 μm and
 260 1–20 μm , respectively. The inset in panel (d) shows the full PM_{10} peak on 5–6 December. Ion data before approximately 7 December 2023
 261 were unavailable and are therefore not shown in panel (d).

262 Around 2 January 2024 (marked by the red arrow in Fig. 1a), a short-lived pollution event occurred, accompanied by sharp
 263 increases in PM mass and concurrent peaks in K, OC, EC (Figs. 1c and 1d), which are consistent with firework emissions
 264 during New Year celebrations. Despite elevated particle concentrations and combustion-related aerosol components, N_{INP}
 265 remained low, suggesting that freshly emitted anthropogenic particles, including black carbon and other combustion products,
 266 are relatively inefficient as INPs, consistent with previous observations that even locally high concentrations of combustion
 267 aerosols contribute little to atmospheric ice-nucleating activity (Adams et al., 2020; Zhang et al., 2022).

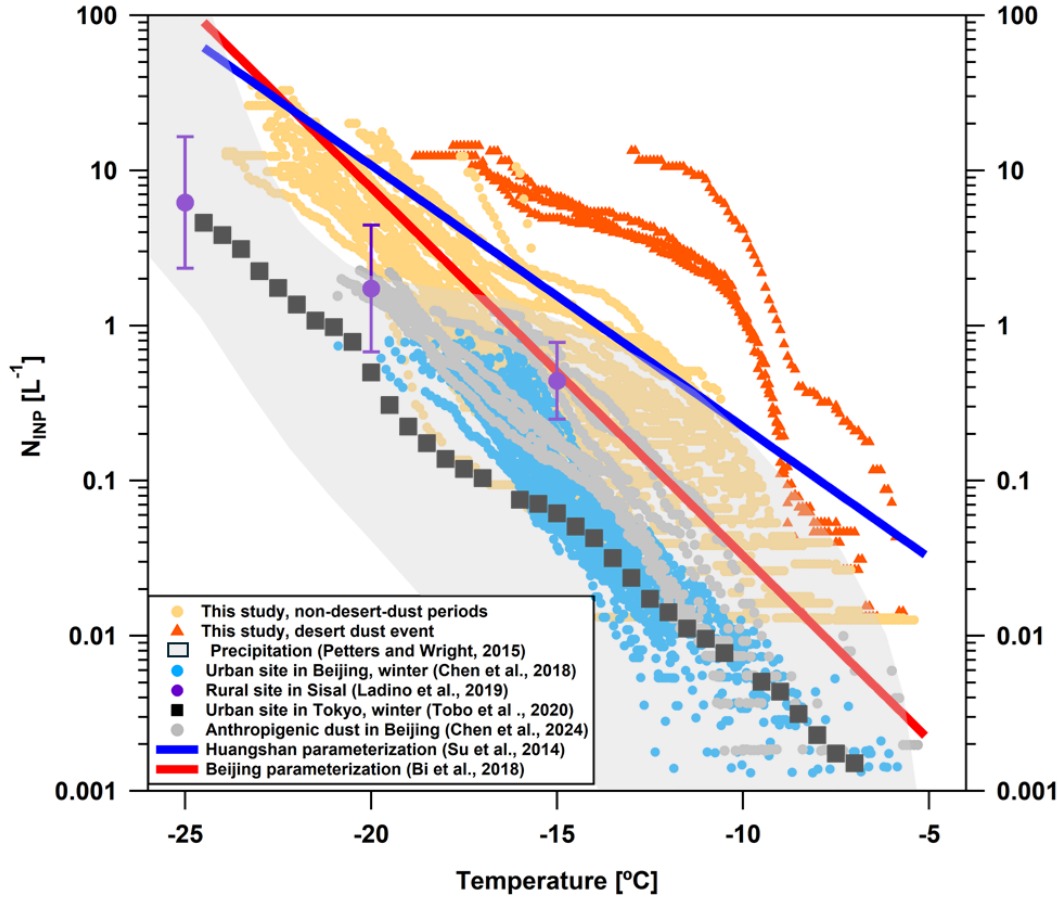
268 3.2 Comparison with previous urban studies and parameterizations

269 Figure 2 compares the measured N_{INP} in this study with values reported previously. All samples in Taiyuan activated before
270 -15°C , with the most active samples initiating freezing near -5°C . The N_{INP} varied by more than two orders of magnitude
271 among different samples, particularly in the -10°C to -15°C range, where values ranged from below 0.01 L^{-1} to above 10 L^{-1} .
272 Such variability primarily reflects the influence of the desert dust intrusion identified earlier (highlighted in darker orange in
273 Fig. 2), during which enhanced N_{INP} values were observed due to long-range transport of natural mineral particles. However,
274 the enhanced warm-temperature INP activity during this episode may not be fully attributable to mineral dust alone. During
275 East Asian dust events, Chen et al. (2021) showed that heat-sensitive INPs made substantial contributions at relatively warm
276 temperatures. Similarly, Hu et al. (2023) reported enhanced warm-temperature INP activity during spring dust transport in
277 Beijing and suggested, based on heat-treatment analysis, that transported dust particles may be mixed with heat-sensitive,
278 potentially biological components. Together, these studies support the possibility of similar heat-sensitive contributions during
279 the Taiyuan dust episode. In the present study, however, no heat-treatment analysis was performed; therefore, such a
280 contribution can only be inferred and cannot be directly assessed.

281 For the remaining non-desert-dust periods, N_{INP} in Taiyuan was generally higher than wintertime urban observations in
282 Beijing measured by INDA (Chen et al., 2018) and in Tokyo (Tobo et al., 2020) over much of the overlapping temperature
283 range. This difference may be related to the semi-arid setting of Taiyuan and its proximity to the Loess Plateau and
284 northwestern dust source regions, where mineral and resuspended dust particles can provide a more persistent background of
285 coarse-mode aerosol than in the Beijing and Tokyo observations. Consistently, particles larger than $1\text{ }\mu\text{m}$ remained at
286 appreciable concentrations even outside the identified desert dust event, although no clear desert-dust intrusion signature
287 comparable to the 5–9 December episode was identified. Differences in sampling size range should also be considered, because
288 the Taiyuan INP filters were collected without a size-selective inlet and can be regarded as approximately total suspended
289 particles (TSP), whereas Chen et al. (2018) collected $\text{PM}_{2.5}$ samples in Beijing. The anthropogenic-dust samples collected in
290 Beijing in summer (Chen et al., 2024) overlap mainly with the lower part of the Taiyuan data. These comparisons suggest the
291 higher Taiyuan N_{INP} levels likely reflect a combination of regional dust influence, local dust-associated coarse particles,
292 seasonal differences, and sampling-size-range differences, rather than a single controlling factor.

293 The linear N_{INP} parameterizations derived from field measurements in Beijing (Bi et al., 2018) and Huangshan (Su et al.,
294 2014) also capture the overall magnitude of the present measurements. Additionally, the gray shaded area denotes the N_{INP}
295 range derived from precipitation samples summarized by Petters and Wright (2015), which is used here as a broad reference
296 for background atmospheric INP abundance. At temperatures warmer than -15°C , the Taiyuan N_{INP} values fall mostly within
297 this range during the non-desert-dust periods. At colder temperatures, however, the measured N_{INP} values exceed this range
298 substantially, suggesting contributions from additional ice-active components active at lower temperatures. Taken together,
299 these features indicate that episodic desert dust transport was associated with the most prominent INP enhancements during

300 this campaign, whereas the remaining variability during the non-desert-dust periods likely reflects a more complex mixture of
 301 sources and particle properties that cannot be uniquely resolved by the present dataset.



302
 303 **Figure 2.** Comparison of N_{INP} obtained in this study with those reported in previous studies. Pale-yellow circles represent N_{INP} values from
 304 this study during non-desert-dust periods, while orange triangles denote data collected during the identified desert dust event. Blue circles
 305 denote wintertime urban INP observations in Beijing measured by INDIA during November–December 2016 (Chen et al., 2018), and black
 306 squares denote wintertime urban INP observations in Tokyo during December 2016–February 2017 (Tobo et al., 2020). Purple circles
 307 indicate data from Sisal, Mexico (Ladino et al., 2019), and gray circles show anthropogenic-dust samples collected in Beijing (Chen et al.,
 308 2024). Solid red and blue lines correspond to parameterizations for Beijing and Huangshan (Bi et al., 2018; Su et al., 2014). The light-gray
 309 shaded area shows N_{INP} ranges from precipitation samples (Petters and Wright, 2015).

310 3.3 Characterization of dust INPs

311 During the desert dust event (5–9 December; marked by the orange arrow and dashed boundary lines in Fig. 1a), both N_{INP}
 312 and n_s increased markedly. At and above -15 °C, N_{INP} reached 13.4 L^{-1} and n_s peaked at $3.11 \times 10^7 \text{ m}^{-2}$, while their mean
 313 values rose from 0.607 L^{-1} to 7.47 L^{-1} and from $9.52 \times 10^5 \text{ m}^{-2}$ to $1.57 \times 10^7 \text{ m}^{-2}$, respectively. These coherent increases in
 314 concentration and surface-active-site density indicate that the episode was characterized by both an increased abundance of
 315 INPs and substantially enhanced intrinsic ice-nucleating activity. The n_s values discussed here were calculated from the

316 standard-condition merged SMPS–OPC size distribution, thereby accounting for contributions from coarse particles in the
317 surface-area estimate. Figure 1a also compares the measurements with the DeMott et al. (2015) parameterization (D15), which
318 links INP concentrations to the number concentration of particles larger than 0.5 μm and was originally developed for desert
319 dust aerosol. In this study, the D15 values were calculated using the observed particle number concentrations larger than 0.5
320 μm and are therefore shown mainly as a dust-based reference. Accordingly, the comparison is most meaningful for the
321 identified dust episode, whereas for the remaining periods it should be interpreted with caution. While D15 captures the correct
322 order of magnitude, it tends to underestimate the INP peak during the desert dust event (5–9 December) and overestimate INP
323 concentrations during the pollution episode (28–30 December). This discrepancy suggests that coarse-particle abundance alone
324 may not fully constrain INP variability in Taiyuan. Particles larger than 0.5 μm in Taiyuan outside the identified desert dust
325 event may differ in source characteristics and ice-nucleating efficiency from those of the highly active desert dust represented
326 by the D15 parameterization. The apparent overestimation by D15 during pollution episodes implies that these coarse particles
327 may possess lower ice-nucleating efficiency than the highly active desert dusts represented by the parameterization. Therefore,
328 while size-based parameterizations provide a useful approximation, incorporating source-specific or compositional factors
329 could further improve INP predictability in such complex industrial-urban environments.

330 Figure 3 compares n_s as a function of temperature for desert dust-affected samples (red) with non-desert-dust samples (blue).
331 Across the full temperature range, dust-affected samples exhibit systematically higher n_s values than non-desert-dust samples.
332 Between $-10\text{ }^{\circ}\text{C}$ and $-15\text{ }^{\circ}\text{C}$, desert dust-affected samples exhibited n_s values approximately one order of magnitude higher
333 than the most active non-desert-dust samples. In the colder range ($-20\text{ }^{\circ}\text{C}$ to $-15\text{ }^{\circ}\text{C}$), most dust-affected samples fall within
334 about one order of magnitude of the desert-dust parameterizations proposed by Niemand et al. (2012) and Chen et al. (2021),
335 indicating ice-nucleating behavior consistent with long-range transported mineral dust. Considering that Taiyuan is located on
336 the western edge of the North China Plain while Beijing lies farther downwind, dust arriving in Beijing may have experienced
337 a longer transport history and more atmospheric processing. As shown in Fig. 3, the Beijing dust parameterization broadly
338 overlaps with the n_s values of the dust samples in this study across the measured temperature range. This comparison is
339 consistent with Chen et al. (2023), who reported observational evidence that atmospheric chemical modification did not
340 suppress the ice nucleation activity of East Asian dust. However, because the present study did not directly characterize the
341 aging state of the dust particles or the specific aging processes involved, this comparison is used only to suggest that strong
342 ice-nucleating activity was retained during the Taiyuan dust episode, rather than to demonstrate or quantify the effect of
343 atmospheric aging.

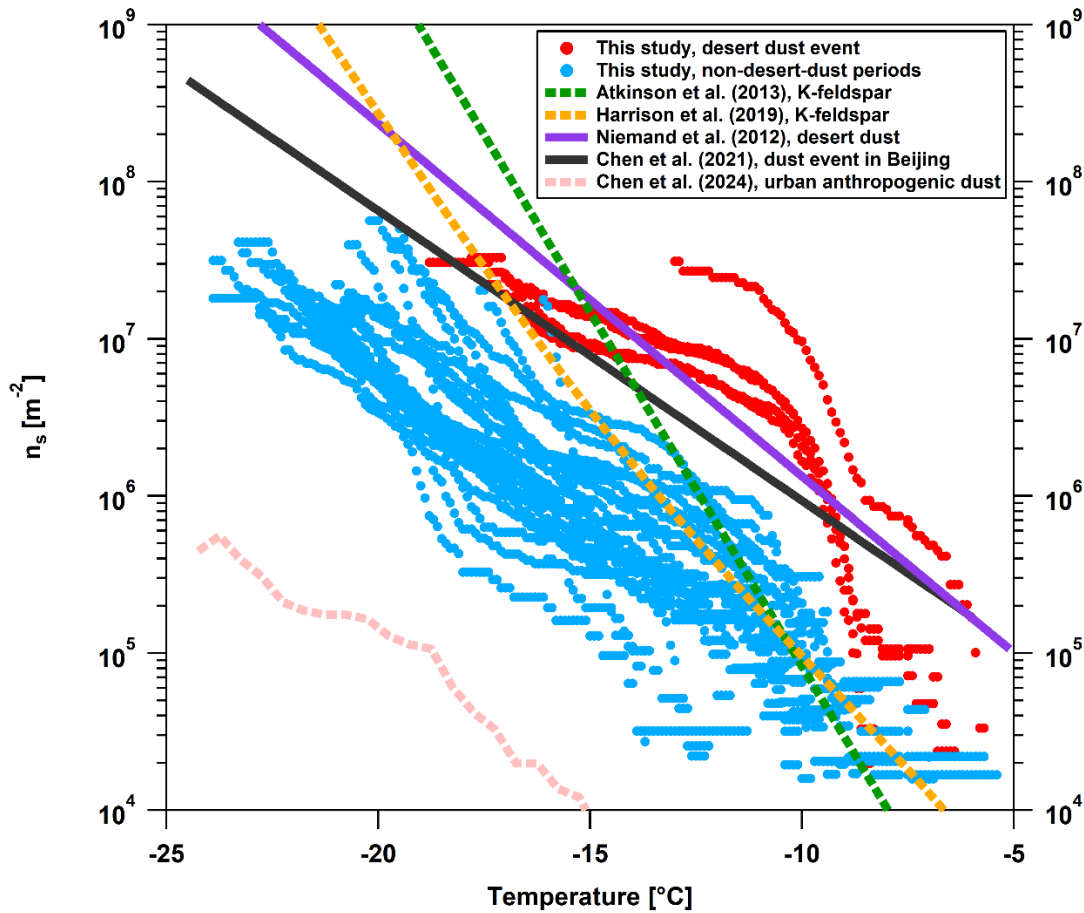


Figure 3. Comparison of n_s obtained in this study with those reported in previous studies. Blue and red markers represent the n_s values of samples from this study during non-desert-dust periods and the identified desert dust event, respectively. The solid lines show n_s parameterizations for dust or individual mineral components reported in previous studies (Atkinson et al., 2013; Chen et al., 2021; Harrison et al., 2019; Niemand et al., 2012). The light-pink dashed line denotes the median trend of n_s values for urban anthropogenic dust particles in Beijing, digitized from Chen et al. (2024).

It is also noteworthy that at relatively warm temperatures (approximately $T > -12\text{ }^{\circ}\text{C}$), n_s values of the ambient desert dust samples substantially exceed the parameterizations for pure K-feldspar (Atkinson et al., 2013; Harrison et al., 2019), suggesting that mineralogy alone cannot account for the elevated warm-temperature activity. Previous laboratory and field studies have shown that mineral dust particles can act synergistically with biological macromolecules or cellular fragments, leading to pronounced enhancement of heterogeneous freezing at warmer temperatures (Augustin-Bauditz et al., 2016; O’Sullivan et al., 2016; Yahya et al., 2019). The enhanced n_s observed at $T > -12\text{ }^{\circ}\text{C}$ in this study may therefore reflect the combined influence of mineral dust and potentially co-transported biological INP components, rather than mineral dust alone. The TSP-like sampling configuration may have facilitated retention of such coarse ice-active components, but it cannot by itself explain the selective enhancement observed during the dust event.

359 3.4 Anthropogenic particle pollution and INP variability during non-desert-dust periods

360 The Pearson correlation coefficients between N_{INP} at different temperatures and various pollutants were calculated for non-
361 desert-dust periods. As shown in Table S1, only Ca, Co, Ni, and Au exhibited statistically significant correlations ($p < 0.05$; n
362 $= 24$). The mean concentrations of Co, Ni, and Au were all below 10 ng m^{-3} , suggesting that their correlations are of limited
363 physical relevance. By contrast, Ca averaged about 405 ng m^{-3} but was notably correlated only with N_{INP} at $-10 \text{ }^{\circ}\text{C}$. As a
364 typical crustal element, Ca is closely linked to particles larger than $1 \text{ }\mu\text{m}$ (Chen et al., 2024). Excluding the influence of long-
365 range dust transport, this observation likely points to contributions from local or regional dust-associated particles, such as
366 resuspended soil and fugitive dust.

367 The weak correlations for all pollutants highlight that individual chemical species provide limited explanatory power for
368 INP variability during non-desert-dust periods. This reflects the fact that INPs constitute only a minor and compositionally
369 distinct subset of the total aerosol population, and thus bulk pollutant concentrations are not expected to correlate strongly with
370 N_{INP} . To further examine whether the dominant $\text{PM}_{2.5}$ source factors were associated with N_{INP} variability during the non-
371 desert-dust periods, PMF analysis was applied to the $\text{PM}_{2.5}$ chemical composition dataset, resolving five major factors (industry,
372 dust-related particles, secondary aerosols, coal combustion and traffic, and fireworks). The corresponding factor profiles (Fig.
373 S4) are consistent with the assigned source categories, including crustal tracers in the dust-related factor and Cu–Ba–K
374 enrichment in the fireworks factor. The inter-species correlation patterns shown in Fig. S5 also support these factor assignments.
375 Because the PMF analysis was based on $\text{PM}_{2.5}$ whereas the INP filter samples can be regarded as approximately TSP, the PMF
376 results are used here only to assess covariation with N_{INP} , rather than to provide direct source apportionment for INPs. Data
377 from the identified desert dust event (5–9 December) are not shown in Fig. 4, because the PMF– N_{INP} correlation analysis
378 focused on non-desert-dust periods. The dust-related factor exhibits characteristic signatures of dust-associated particles,
379 including high loadings of Ca and Fe. During the non-desert-dust periods, this factor is interpreted as mainly reflecting local
380 dust-associated particles, such as road dust, fugitive dust from industrial and construction activities, and resuspended soil. This
381 dust-related $\text{PM}_{2.5}$ factor showed no significant correlation with N_{INP} during the non-desert-dust periods (Table 1), indicating
382 that local dust-associated $\text{PM}_{2.5}$ variability did not explain N_{INP} variability outside the identified long-range desert dust event.
383 The reconstructed daily $\text{PM}_{2.5}$ contributions (Fig. 4) provide the temporal context for the subsequent PMF– N_{INP} correlation
384 analysis, showing that secondary aerosols and coal/traffic emissions dominated the $\text{PM}_{2.5}$ mass during much of the non-desert-
385 dust period, whereas fireworks emissions became prominent during the New Year period.

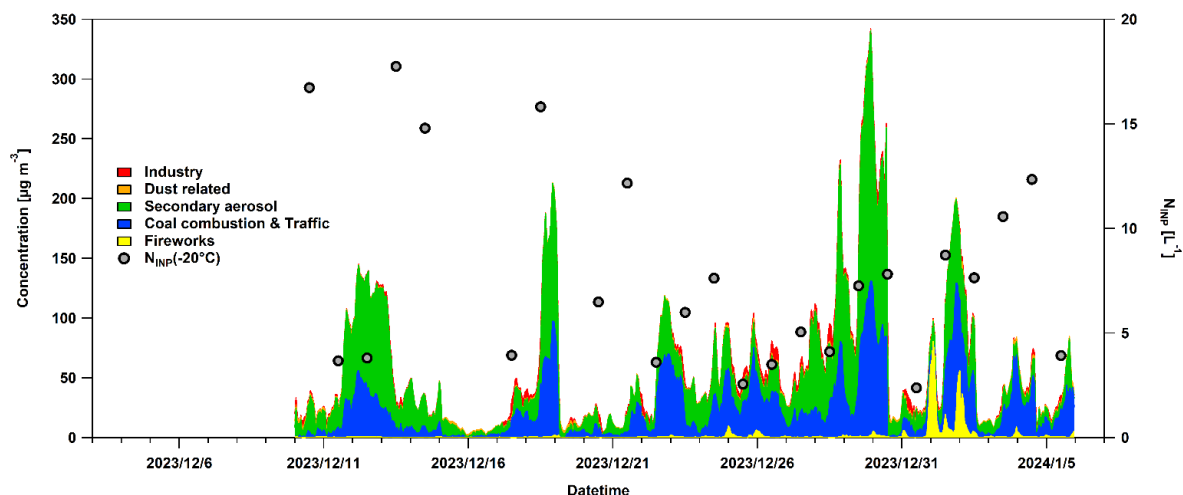


Figure 4. Daily PM_{2.5} source contributions resolved using the PMF model, shown together with the corresponding N_{INP} at -20 °C during the non-desert-dust periods. Data from the identified desert dust event are not shown. The samples collected on December 13, 14, 21 and January 3, 4 were fully frozen before -20 °C, exceeding the quantifiable range of the assay.

To evaluate whether any of these major PM_{2.5} sources covaried with INP, Pearson correlations were calculated between N_{INP} at -20 °C and the time series of the PMF factors during non-desert-dust periods (Table 1). All correlations were weak ($|r| < 0.3$) and statistically insignificant ($p > 0.05$), indicating that the dominant PM_{2.5} source factors did not show clear covariation with INP abundance. Importantly, even during periods with strong increases in secondary inorganic and organic aerosols, such as 13–14 December, 18 December, and 28–30 December, N_{INP} did not rise proportionally (Fig. 4). This behavior suggests that the observed pollution episodes, which were largely associated with fine-mode secondary aerosol accumulation, were not accompanied by corresponding INP enhancements, consistent with previous urban observations showing weak relationships between INPs and bulk pollution indicators (Chen et al., 2018; Zhang et al., 2022).

Table 1. Pearson correlation coefficients (r) and corresponding p -values between N_{INP}(-20 °C) and concentration contributions of PMF-resolved factors during non-desert-dust periods.

Aerosol Source	Pearson r value	p value
Industry	-0.264	0.274
Dust related	0.209	0.391
Secondary aerosol	0.145	0.554
Coal & Traffic	0.035	0.886
Fireworks	0.108	0.661

These results suggest that, during non-desert-dust periods of this campaign, N_{INP} variability was not clearly explained by any of the major PMF-resolved source categories. This is consistent with previous observations at urban sites in different regions, where INPs remain scarce and episodic relative to total aerosol loading, and where their variability is often not well explained by the major anthropogenic aerosol components. Similar analyses were also conducted at -10, -12.5, -15, and -17.5 °C (Table S2). No statistically significant correlations were found between N_{INP} and any PMF-resolved factor at the examined temperatures ($p > 0.05$), indicating that this result was not specific to the choice of -20 °C.

406 4 Conclusions

407 This study provides a detailed observational assessment of atmospheric INPs in Taiyuan, a heavily industrialized city in
408 North China. By combining freezing assays, aerosol chemical composition, back-trajectory analysis, and PMF analysis, the
409 work characterizes the concentration range, temperature dependence of INP activation, and variability of INPs in an urban
410 environment where such information has been largely absent. The results provide observational constraints on wintertime INP
411 abundance and variability in an industrial urban atmosphere.

412 A key finding of this winter campaign is that the strongest INP enhancements in Taiyuan were associated with episodic
413 long-range desert dust transport. During the 5–9 December desert dust event, N_{INP} and n_s increased by up to an order of
414 magnitude and freezing initiated at markedly warmer temperatures. Back trajectories confirmed that these highly active
415 samples were associated with air masses originating from the major desert regions in northwestern China. Comparison with
416 previously reported dust parameterizations indicates that the transported desert dust retained strong ice-nucleating activity
417 during the dust transport episode. The enhanced warm-temperature activity may also reflect the possible contribution of co-
418 transported biological components, although this cannot be directly verified in the absence of heat-treatment analysis.

419 In contrast, during the non-desert-dust periods of this campaign, N_{INP} remained relatively low and showed weak correlations
420 with common urban aerosol species. PMF analysis resolved five major $\text{PM}_{2.5}$ source factors: industrial emissions, dust-related
421 particles, secondary aerosols, coal combustion and traffic emissions, and fireworks, with secondary aerosols and coal/traffic
422 contributions dominating the particle mass. However, none of these PMF-resolved factors showed statistically significant
423 covariation with N_{INP} during the non-desert-dust periods. This suggests that INP abundance during these periods could not be
424 clearly explained by the dominant contributors to $\text{PM}_{2.5}$ mass.

425 Overall, the observations from this winter campaign indicate that the strongest INP enhancements in Taiyuan were
426 associated with long-range desert dust transport, whereas the remaining INP variability during the non-desert-dust periods
427 reflected more complex influences that were not resolved by the major $\text{PM}_{2.5}$ source factors considered here. These results
428 provide observational constraints for understanding wintertime urban INPs and for improving their representation in chemical
429 transport and climate models.

430 Data availability

431 The data supporting this study are available from Zenodo at <https://doi.org/10.5281/zenodo.20232172> (Yang et al., 2026). The
432 dataset includes INP freezing data, temperature-resolved N_{INP} and confidence intervals, particle size distributions and surface
433 area data, auxiliary observations, PMF source contribution data, and HYSPLIT backward trajectory outputs.

434 **Author contributions**

435 JY, JingC and ZW designed experiments and methodology. JY, WF, JingC and ZF conducted the field observation. JY, ZW,
436 WF, ZF, TZ, YQ, JunW and RM performed aerosol chemical and size distribution analyses; specifically, YQ was responsible
437 for PMF (Positive Matrix Factorization) analysis, while JunW and RM assisted with offline sampling and chemical component
438 analysis. JY, JieC and JingC calibrated the PKU-INA. TZ and WF provided meteorological data. MH, ZW, YL and NT
439 supervised the study. JY, ZW, JieC and JingC prepared the article with input from all co-authors.

440 **Competing interests**

441 The authors declare that they have no conflict of interest.

442 **Acknowledgements**

443 We thank Ling Mu, Jiajie Li, Ying Wei, and Chenhui Li (Taiyuan University of Technology) for collecting the offline
444 membrane samples.

445 **Financial support**

446 This work was supported by the National Natural Science Foundation of China (NSFC, grant No. 42375093), the Exploratory
447 Innovation Project of the National Institute of Metrology, China (AKYCX2512), and the cooperative research programs of the
448 Institute of Nature and Environmental Technology, Kanazawa University, Japan (25001, 25015).

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