

Supplementary material for

A case for a pragmatic oxygen-based approach to quantifying the biological contribution to the marine carbon sink.

Wolfgang Koeve and Ivy Frenger

Biogeochemical Modelling, GEOMAR Helmholtz Centre for Ocean Research Kiel,
Wischhofstraße 1-3, 24148 Kiel, Germany

* corresponding authors wkoeve@geomar.de and ifrenger@geomar.de

SI Tables

Tab. S1. Preindustrial state estimates of Cremin. Elemental ratios used ($R_{C:P}$, $R_{C:O_2}$) are those used in our model. See text and Online Methods for details.

Approach	Global integral (Pg C equivalents)
$AOU \cdot R_{C:O_2}$	1562
$DIC^{remin} \quad \$$	1290
$TOU \cdot R_{C:O_2} \quad \$$	1272
$C^{soft} = (PO_4 - PO_4^{pre}) \cdot R_{C:P} \quad \$$	1303
tracer based $\quad \$$ (average)	1288 \pm 16 (std)
$O_2^{dis,bio} \cdot R_{C:O_2}$	256
$C^{soft, tracer} + O_2^{dis,bio} \cdot R_{C:O_2}$	1544
$DIC^{dis,bio}$	520
$C^{soft, tracer} + DIC^{dis,bio}$	1808

$\quad \$$ estimates based on idealized model tracers DIC^{remin} , TOU , and PO_4^{pre} ; the term ‘tracer based’ refers to the mean of the three tracer based estimates; see Online Methods for technical definitions and references.

SI Figures (captions)

Fig. S1: Comparison of the preindustrial state estimates of O_2 -saturation (O_2^{sat} , mmol O_2 / m^3) and preformed O_2 (O_2^{pre} , mmol O_2 / m^3) of the default model variant. (a) Winter time surface O_2 -saturation (%) showing undersaturation in high latitude waters. (b) Wintertime O_2 disequilibrium ($O_2^{\text{dis}} = O_2^{\text{pre}} - O_2^{\text{sat}}$, mmol O_2 / m^3). (c) O_2^{sat} section through the Atlantic (30°W, from North to South), the Southern Ocean (60°S), and the Pacific Ocean (140°W, from South to North), (d) like (c), but for preformed O_2 , (O_2^{pre} , mmol O_2 / m^3)

Fig. S2: Decomposition of the O_2 disequilibrium. (a) total O_2 disequilibrium (O_2^{dis} , mmol O_2 / m^3), (b) biotic component ($O_2^{\text{dis,bio}}$, mmol O_2 / m^3), and (c) physical component ($O_2^{\text{dis,phys}}$, mmol O_2 / m^3). See online methods for details on the methodology applied to decompose O_2^{dis} .

Fig. S3: Forcing, climate response and global marine response of the default model experiment (ZECmip1000). (a) CO_2 -emissions forcing applied (Pg C / year), (b) change in atmospheric pCO_2 (ppm), (c) change in global mean surface air temperature (SAT, °C), (d) change in northern hemispheric (black) and southern hemispheric sea ice cover ($1e6 \text{ km}^2$). (e) change in net primary production (NPP, Pg C / year). (f) change in export production (Pg C / year). Note: I still have to add the preindustrial values

Fig. S4: Sections through the Atlantic (30°W, from North to South), the Southern Ocean (60°S), and the Pacific Ocean (140°W, from South to North), showing the regional distribution of transient changes of AOU (a, b), TOU (c, d), sum $TOU + O_2^{\text{dis,bio}}$ (e,f), and $O_2^{\text{dis,bio}}$ (g, h), all in mmol O_2 / m^3 . Panels a, c, e, g show differences between yr 100 and the initial state and panels b, d, f, h show differences between year 800 and the initial state. Warm colors indicate increase over time and cold colors decrease over time.

Fig. S5: Descriptive power of AOU during the climate change transient (default model variant), shown as $\Delta AOU / (\Delta TOU + \Delta O_2^{\text{dis,bio}}) * 100$ (%). Solid line indicates the period with ΔAOU lower than -1 Pmol O_2 , dashed line indicates period with ΔAOU between 0 and -1 Pmol O_2 .

Fig. S6: Transient changes of surface air temperature (a, SAT, °C) and Atlantic meridional overturning (b, AMOC, Sv) from experiments taken from Koeve et al., 2024 and explored in Fig. 3c. In short, experiments have been forced with prescribed atmospheric pCO_2 , initially increasing with 1% per year until reaching CO_2 doubling (all experiments). After reaching peak pCO_2 in year 70, experiments have been forced with decreasing atmospheric pCO_2 until reaching preindustrial pCO_2 again. Prescribed pCO_2 decreased with rates ranging from -1% per year (solid black lines) to -0.1% per year (light blue lines). See Koeve et al., 2024, SI for details.

Fig. S7: Three tracer based estimates of globally integrated ΔC^{remin} (BCP storage changes) from the default model variant using ZEC1000 forcing in fully coupled (COU*, solid lines) and biogeochemically-only coupled experiments (BGC*, dashed lines). TOU-tracer based estimate (black), DIC^{remin} tracer (red), and PO_4^{pre} (blue) based estimate. For details and equations see Online Methods.

Fig. S8: Forcing, climate response and global marine response of the biogeochemically only coupled model experiment (BGC*). (a) CO₂-emissions forcing applied (Pg C / year), (b) change in atmospheric pCO₂ (ppm), (c) change in global mean surface air temperature (SAT, °C), (d) change in northern hemispheric (black) and southern hemispheric sea ice cover (1e6 km²). (e) change in net primary production (NPP, Pg C / year). (f) change in export production (Pg C / year).

Fig. S9: Comparison of computed O₂^{sat} (T,S) against explicit O₂^{sat} tracer. Section through the Atlantic (30°W, from North to South), the Southern Ocean (60°S), and the Pacific Ocean (140°W, from South to North) of (a) O₂^{sat}(T,S) - O₂^{sat}(tracer) and (b))O₂^{sat}(T,S) - O₂^{sat}(tracer)) / O₂^{sat}(tracer) *100

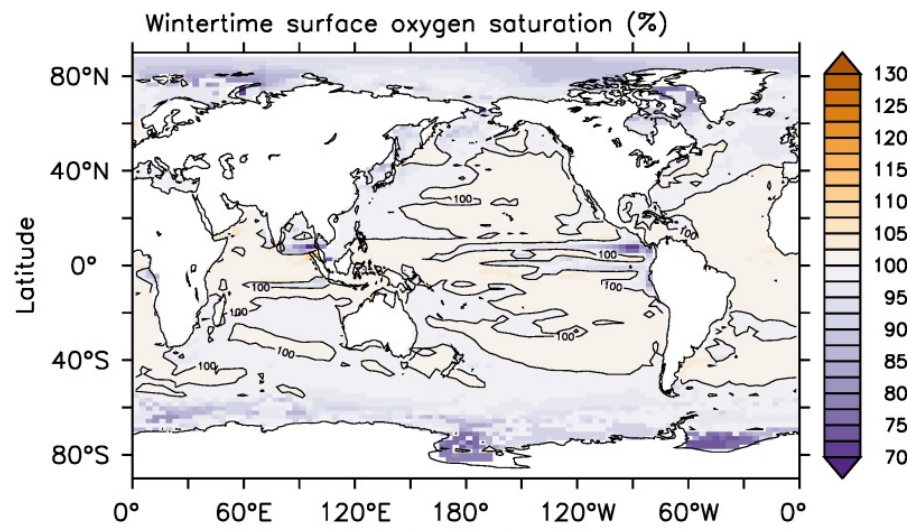
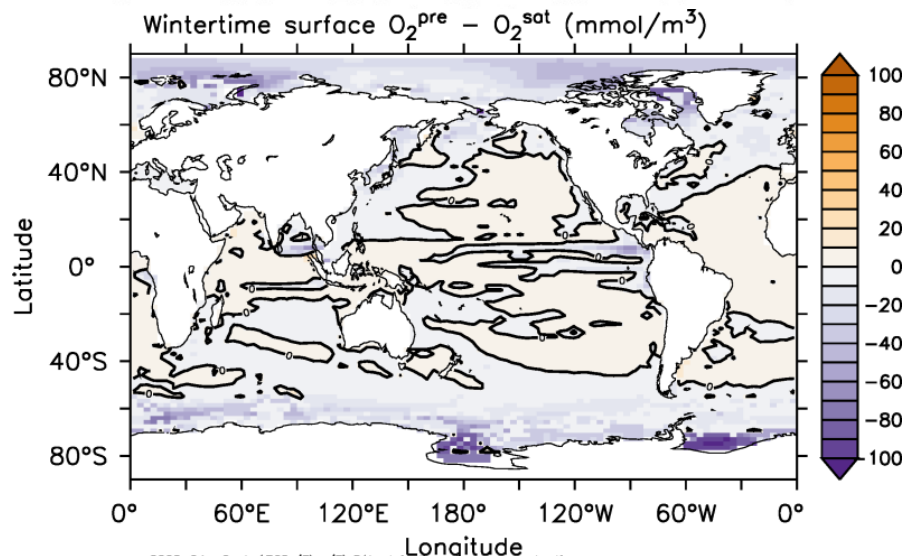
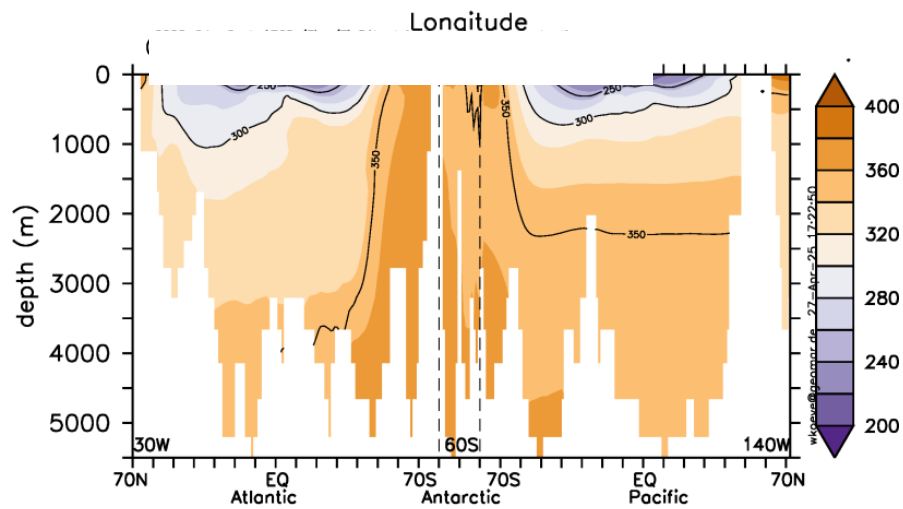


Fig. S1

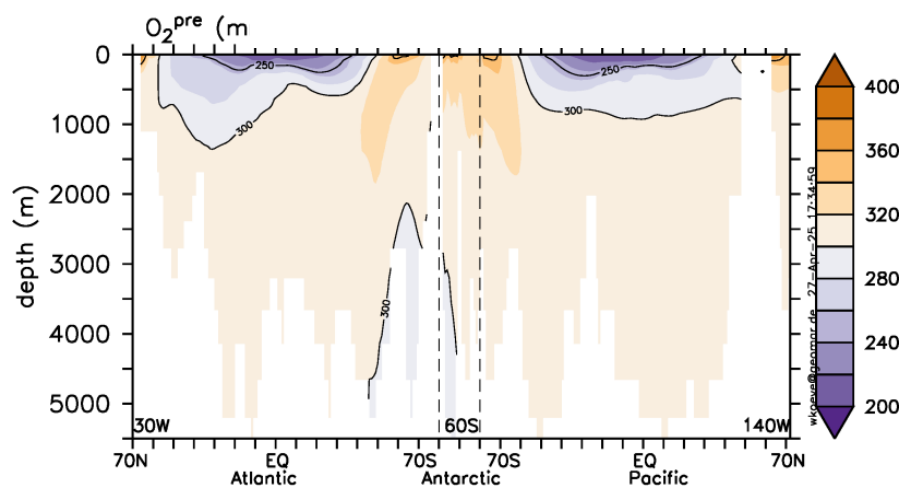
(a)



(b)



(c)



(d)

Fig. S2

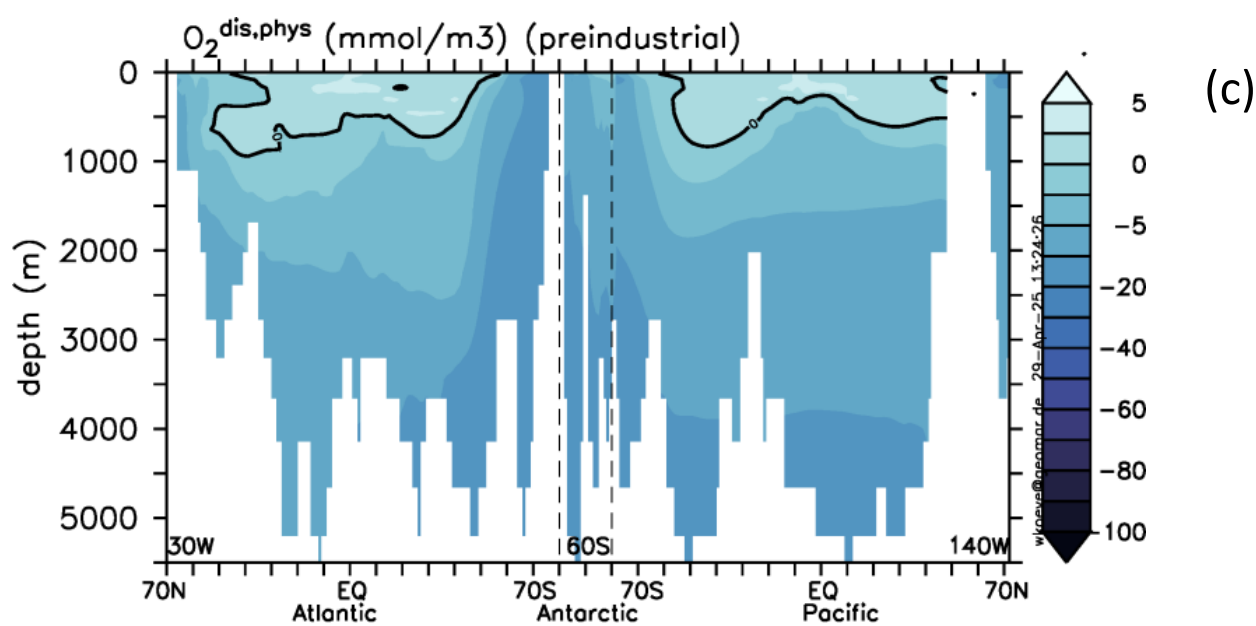
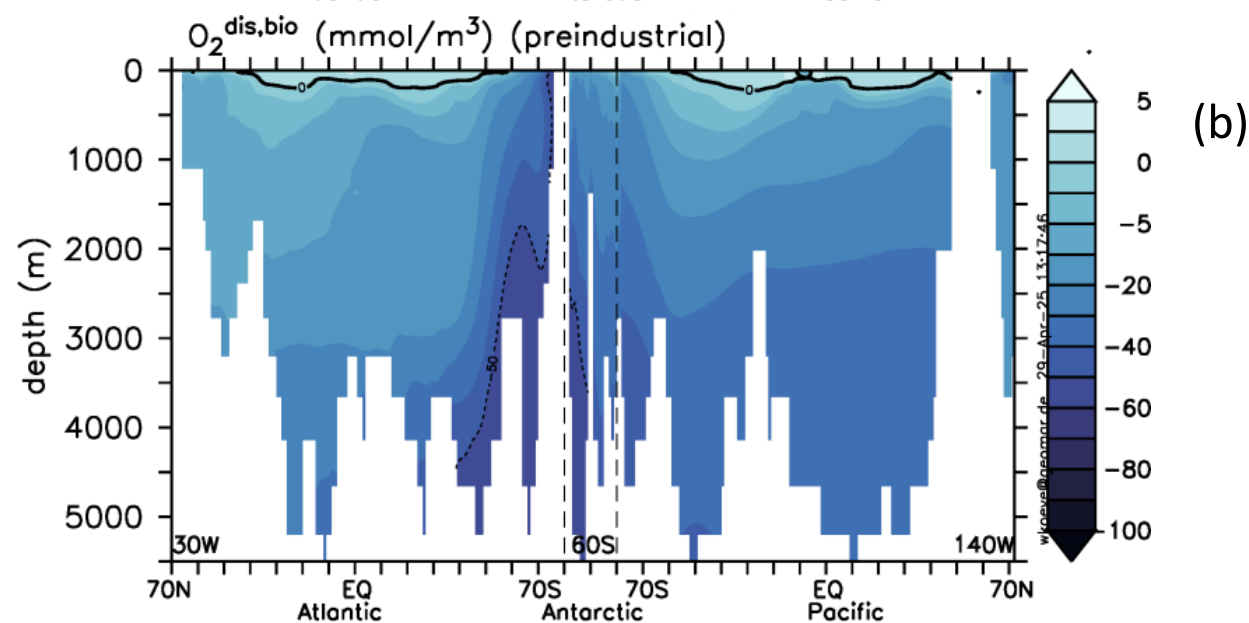
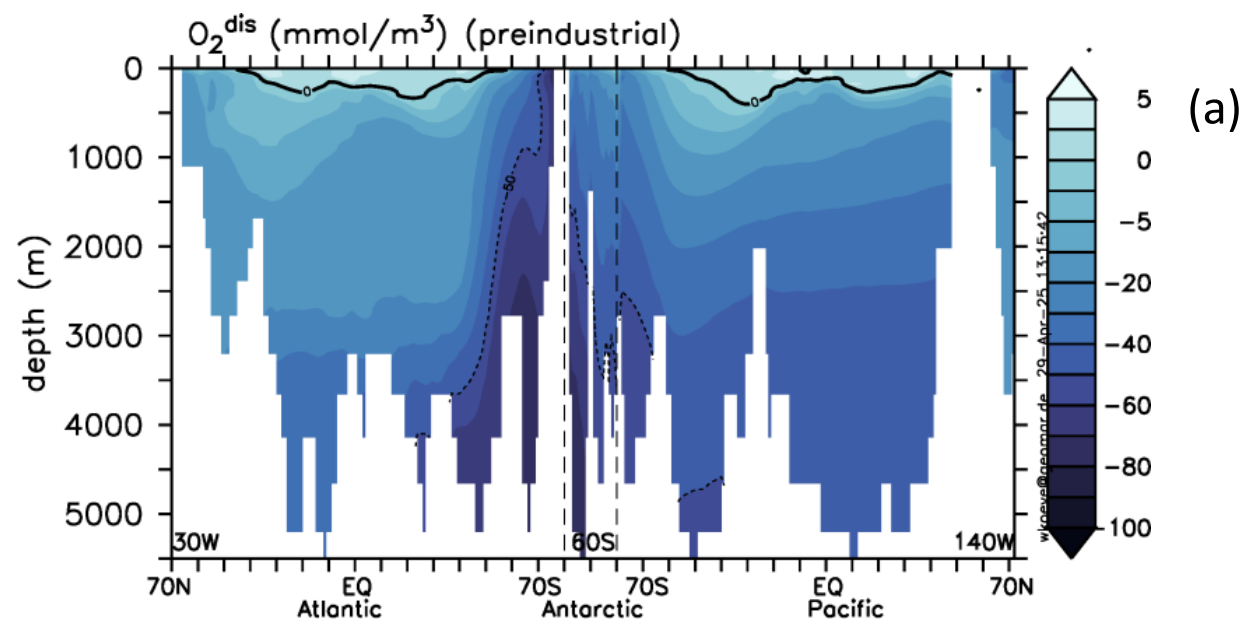


Fig. S3

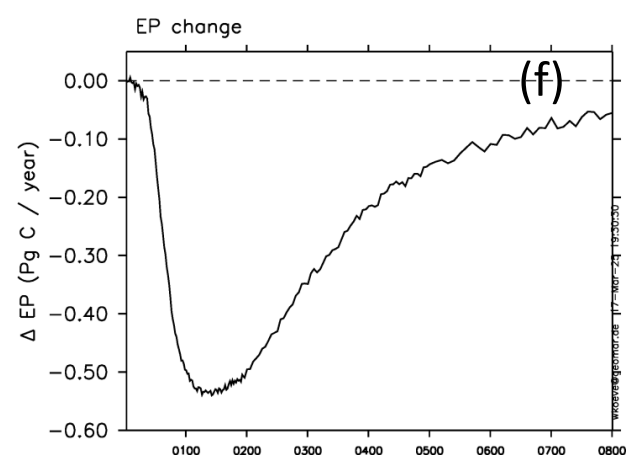
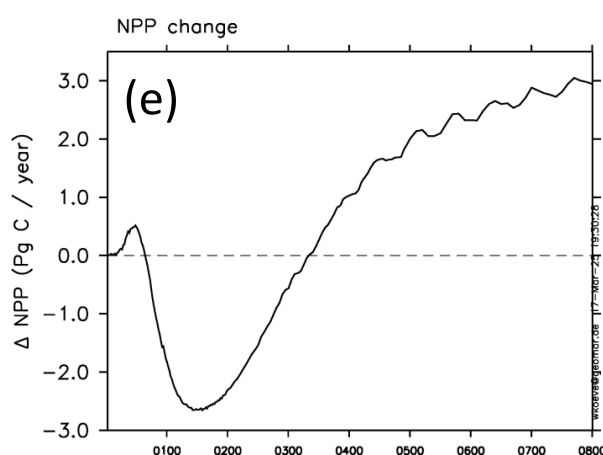
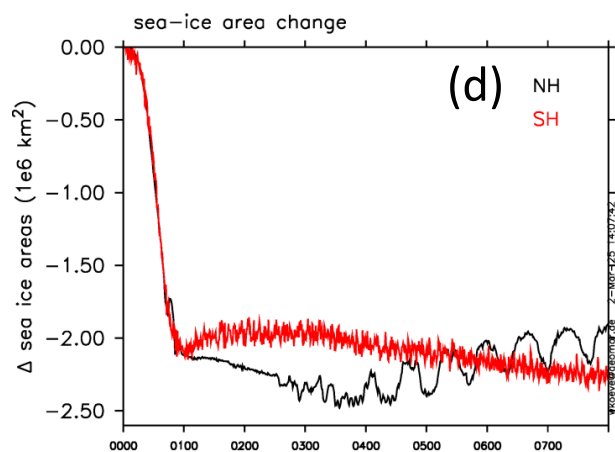
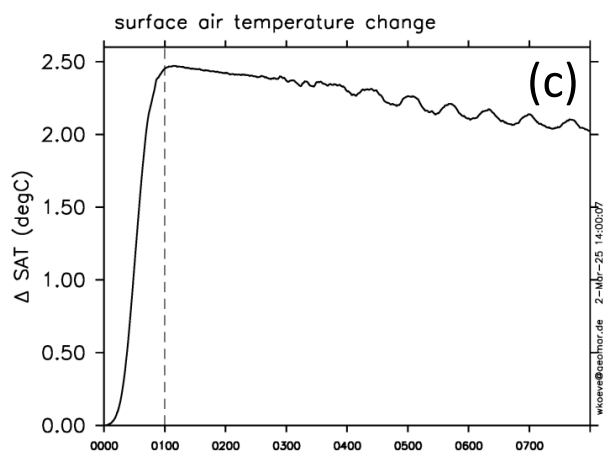
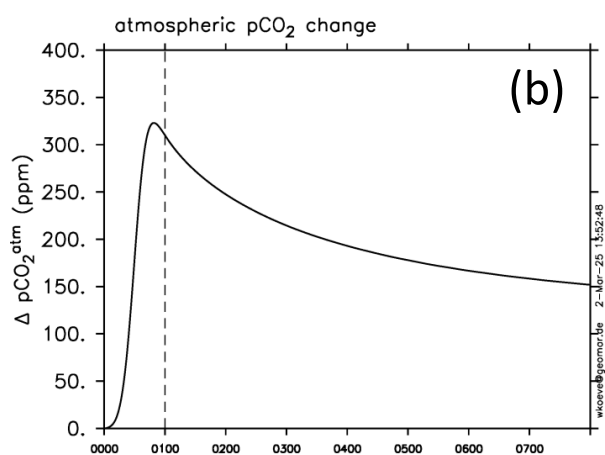
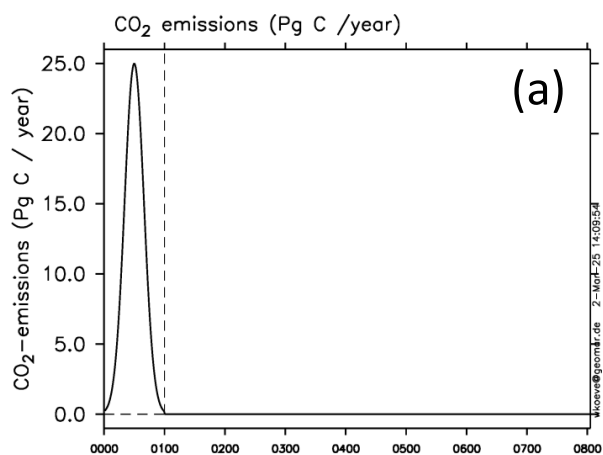


Fig. S4

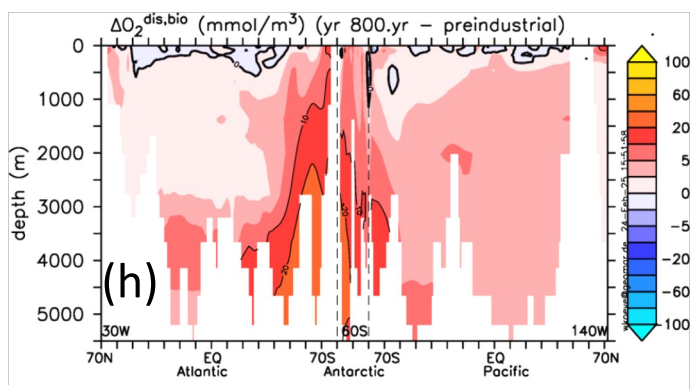
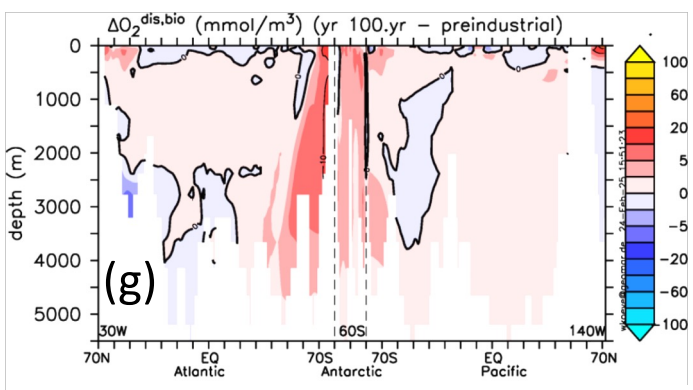
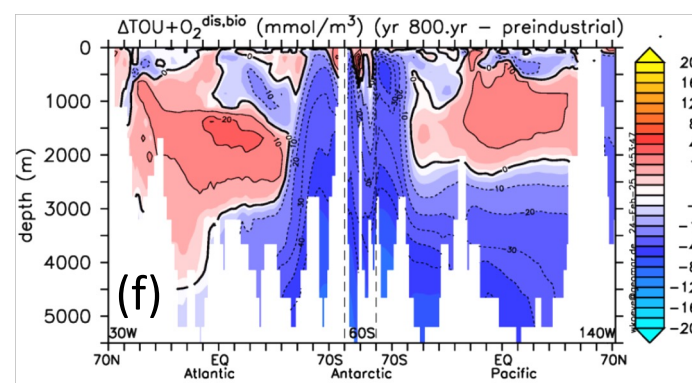
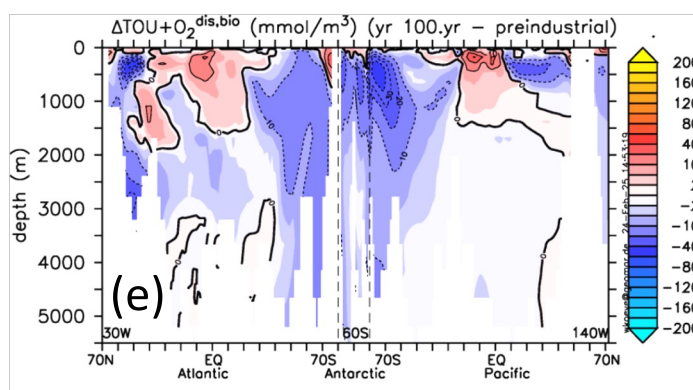
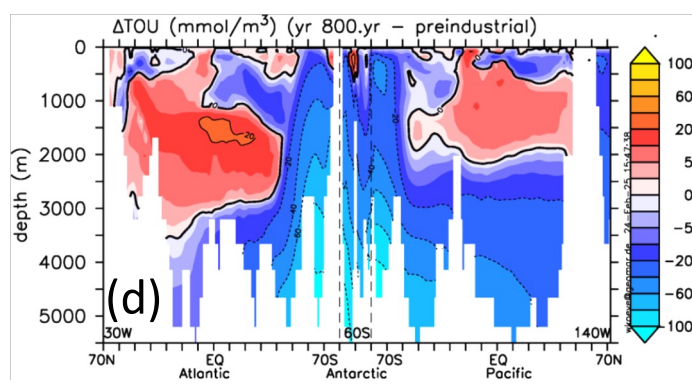
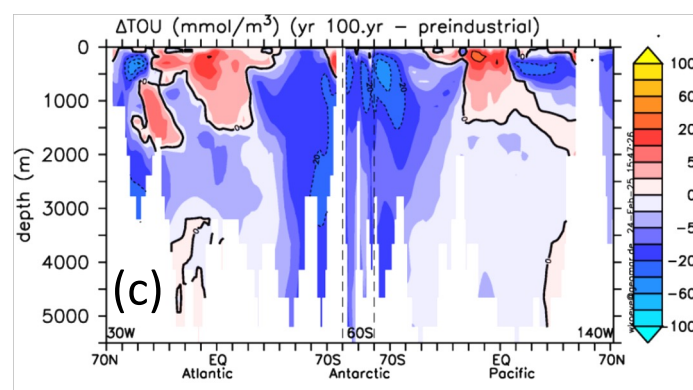
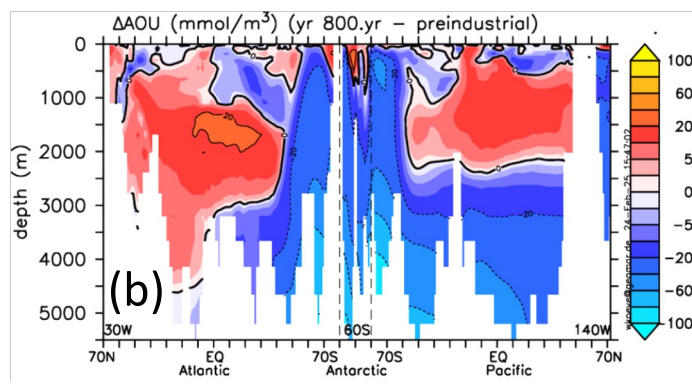
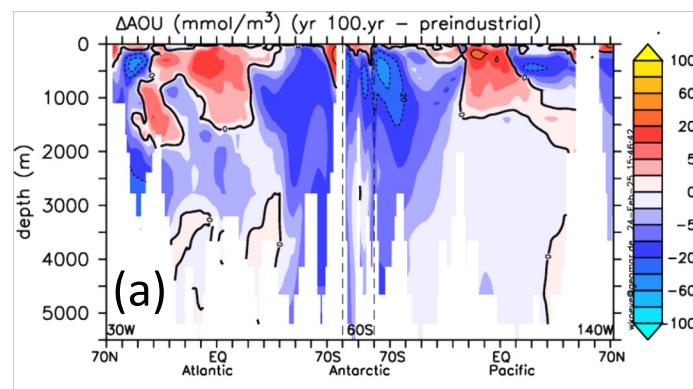


Fig. S5

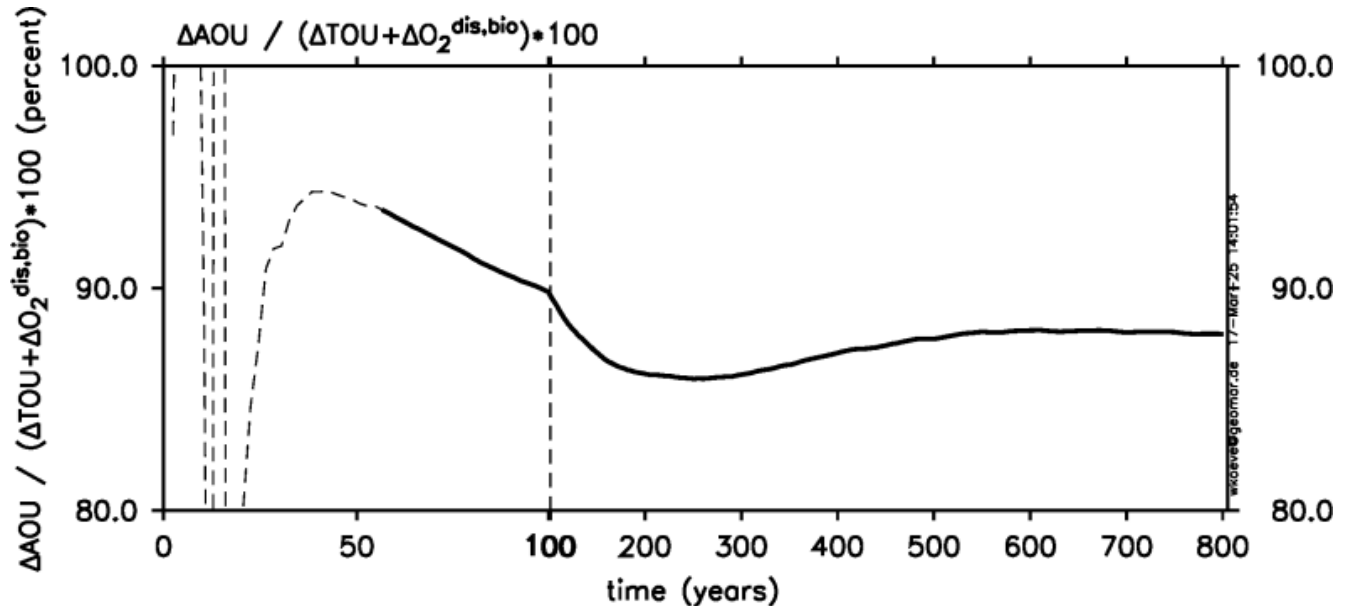


Fig. S6

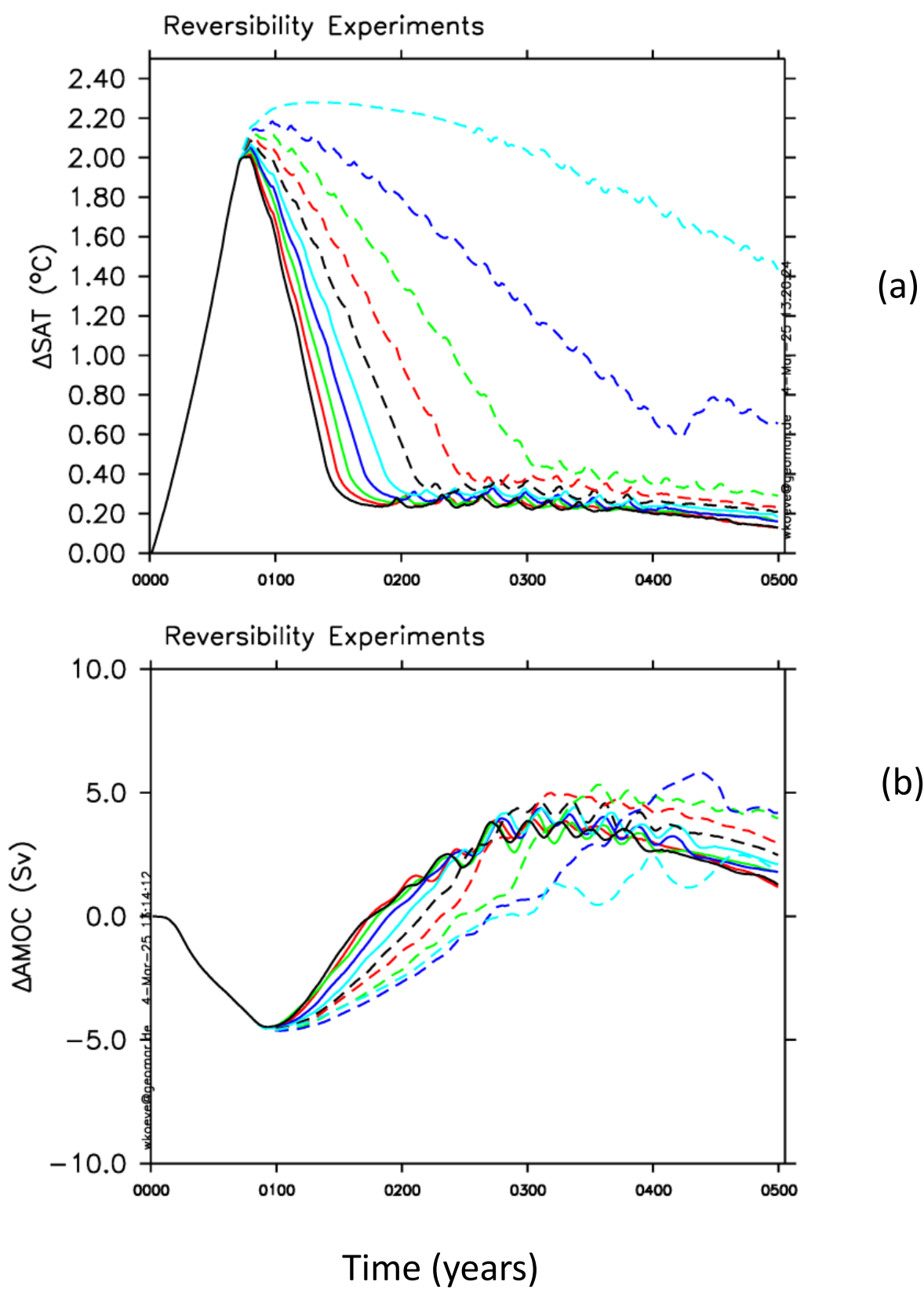


Fig. S7

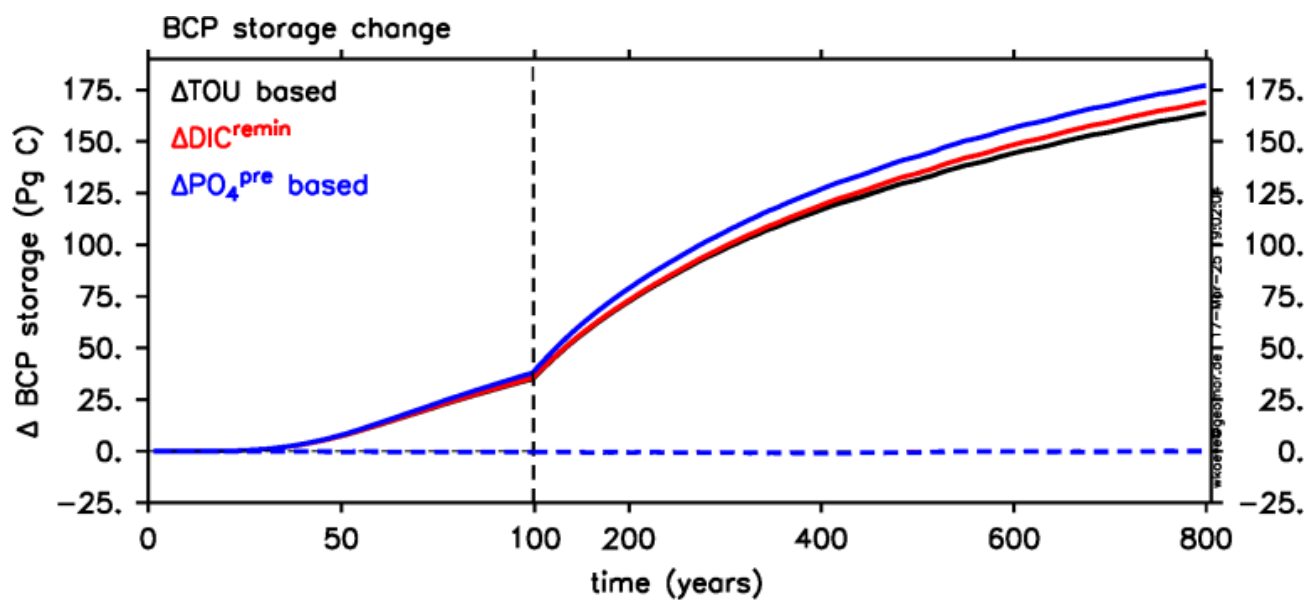


Fig. S8

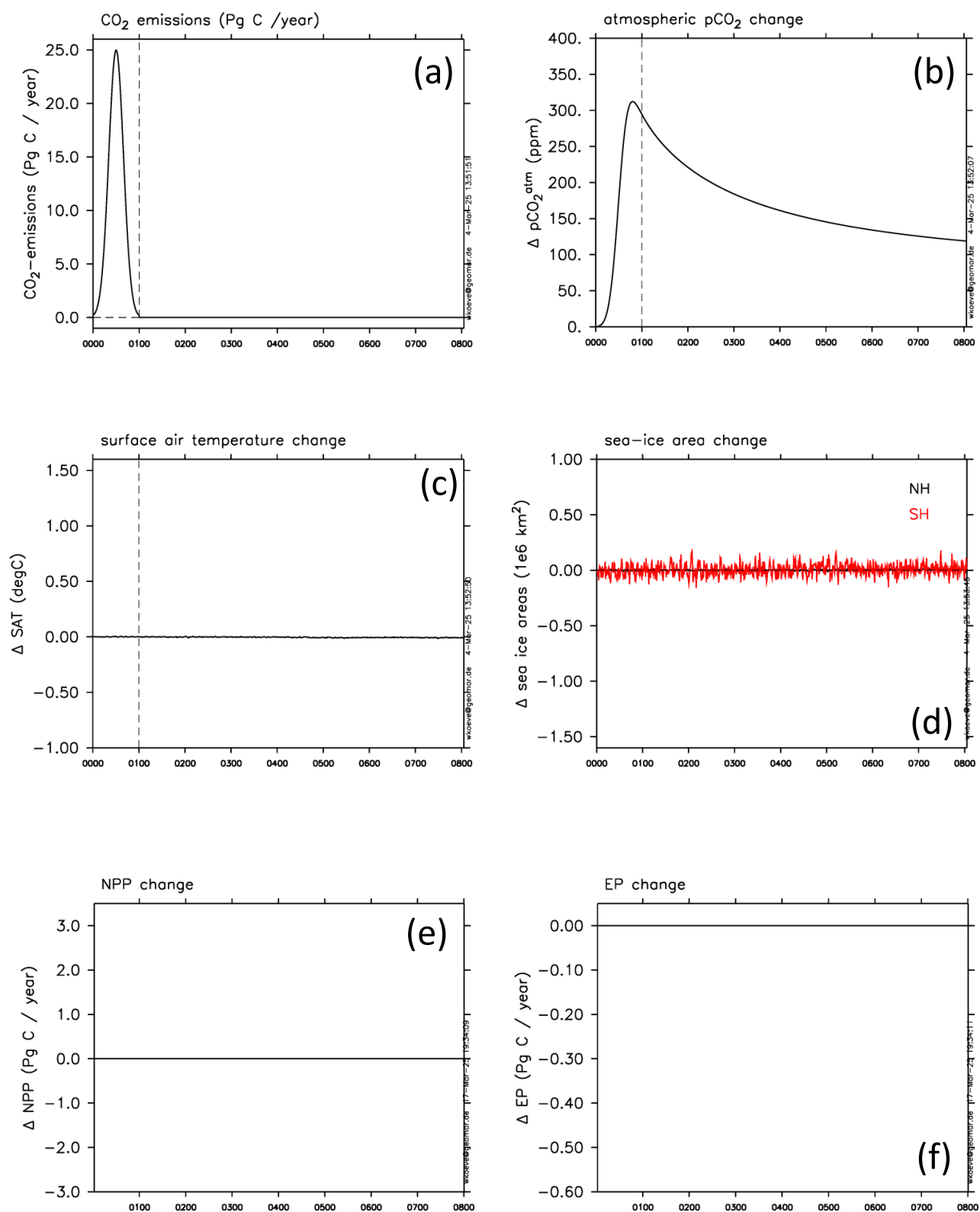


Fig. S9

