Insights into the chemical and physical drivers of surface ozone in the Trans-Himalaya

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Abstract

The Trans-Himalaya, located between the Great Himalaya and the Tibetan Plateau, is ecologically fragile and one of the most climate-sensitive regions. However, lack of systematic in-situ measurements have hindered the evaluation of chemistry-climate models and constrained our understanding of the global-to-regional influences over this region. To address this, we present a comprehensive analysis by combining first yearlong (October 2024–September 2025) surface ozone (O_3) measurements at Leh (3.4 km above mean sea level) with satellite data and numerical simulations (global model: CAM4-Chem, regional model: WRF-Chem). Our measurements revealed a noontime O_3 build-up (~3–7 ppbv), a feature typically absent over high altitude mountains. WRF-Chem suggests that this feature is driven by photochemical production (NO_x-VOC transition regime for higher O_3) and entrainment. Despite partial compensation of nighttime loss by noontime build-up, the region acts as a net chemical sink. Seasonal cycle peaks in June (~65±6 ppbv; touching air quality threshold), unlike typical premonsoon peaks over Gangeatic and Himalayan regions. CAM4-Chem simulation with CAMS-derived stratospheric contribution, captured the O_3 seasonality, but overestimated the wintertime levels by 10–15 ppbv. Tagged O_3-tracer analysis shows significant contributions of anthropogenic NO_x emissions from South Asia (15–18 %) and other regions (32–37 %) to O_3 levels. Additionally, stratospheric inputs and biogenic emissions shape the seasonal cycle. This study provides insight into the chemistry and dynamics governing O_3 over the Trans-Himalaya across different time scales. These findings are invaluable for designing future field observations, model improvements, and policy planning for this region.

Keywords: Ladakh, free-troposphere, stratospheric intrusion, WRF-Chem, CAM-Chem, air quality, atmospheric composition, photochemistry, transport, source attribution

1. Introduction

Tropospheric ozone (O_3) plays a tripartite role: a primary controller of the atmosphere's oxidative capacity, a significant short-lived climate forcer, and a secondary air pollutant (Seinfeld and Pandis, 2016). High levels of ground level O_3 adversely impact human health, vegetation, crop yield and ecosystem (Chuwah et al., 2015; US EPA, 2013). Tropospheric O_3 also acts as a greenhouse gas, produced photochemically by methane (CH_4), carbon monoxide (CO), and non-methane hydrocarbons (NMHCs) in the presence of nitrogen oxides (NO_x). Thus, O_3 is considered an important link between air quality and climate. O_3 has a longer lifetime (6–27 days) in the free troposphere compared to the boundary layer (few hours to ~1 week), due to weaker loss processes associated with lower water vapour and negligible deposition (Prather and Zhu, 2024). Therefore, free-tropospheric O_3 levels can sustain more than those within the urban or rural local boundary layers, and hence such observations would be better representative of regional conditions. In addition, free tropospheric O_3 variations can provide insights into vertical transport including stratospheric intrusions as well as in situ or en-route O_3 formation through convectively transported precursors (Ansari et al., 2025; Nalam et al., 2025; Ojha et al., 2012; Sahu et al., 2009; Sinha et al., 2016). O_3 mixing ratios generally increase with altitude owing to more intense solar radiation and low concentrations of the titrating gases, for example, NO (Chevalier et al., 2007). The intrusions of O_3-rich stratospheric air also enhance O_3 levels more at higher altitudes compared to low-altitude plain regions (Cristofanelli et al., 2010). Therefore, studies on ozone variations at high-altitude remote stations are critical to obtain comprehensive insights into the roles of regional emissions, photochemistry, long-range transport, and stratospheric intrusions (Kumar et al., 2010; Sarangi et al., 2014).

Ladakh is a Union Territory (a region directly governed by the central government) and the northern-most part of India, situated in the Trans-Himalayan region between the Karakoram and Great Himalayan ranges. It is one of the coldest, driest and most sparsely populated regions in South Asia and is suggested to be one of the most climate-sensitive, highly vulnerable, and ecologically fragile regions of the world (Goodrich et al., 2019; Wester et al., 2019). Ladakh lies in one of the three sub-divisions of Hindu Kush Himalaya (HKH), i.e., north-western Himalaya and Karakoram range (32°–39° N; 71°–79° E). The International Centre for Integrated Mountain Development (ICIMOD) have reported that HKH is warming faster than the average rate of global warming with adverse impacts on the cryosphere, hydrology, ecology, vegetation, wildlife, and human society (Nair et al., 2024; Wester et al., 2019). Annual mean temperature has risen by more than 2 °C in Leh region over the period of 13 years (1995–2009; Chevuturi et al., 2018). This region is already experiencing substantial

consequences of the climate change, e.g., increased glacier retreating, water scarcity, heatwaves, landslides, flash floods, debris flows, and soil erosion, etc (Bhattacharjee and Dutta, 2023; Dubey et al., 2025; Khanday et al., 2024; Pandit, 2013; Ziegler et al., 2016). Being a high-altitude region, remote from major anthropogenic emissions, Ladakh experiences a cleaner background atmospheric conditions, as compared to the densely populated and global hotspot of emissions—the Indo-Gangetic Plain (IGP) in its immediate south. Due to spatial proximity plus strong atmospheric dynamics, the transport of trace gases and aerosols from IGP can potentially influence the Trans-Himalaya also, as observed for the Himalayan region (Das et al., 2022; Hassan et al., 2023; Ojha et al., 2012; Sarangi et al., 2014). Signatures of aged background aerosols as well as new-particle formation from gaseous precursors suggest a unique chemical environment over Trans-Himalaya (Ningombam et al., 2014).

Variations in O_3 over the high-altitude Ladakh region can provide invaluable insights into the regional background, long-range transport, stratospheric intrusions, besides into photochemistry intensified by higher solar radiation. This region also falls under one of the global hotspots of Stratosphere-Troposphere Exchange (STE; Škerlak et al., 2014). In other words, this region —one among a few places in the globe (e.g., Mt. Bachelor Observatory, Pico Mountain Observatory; (Okamoto and Tanimoto, 2016), is unique as the amalgamation of regional emission with the influx from free-troposphere and stratosphere is interacting with intense solar radiation at such a high elevation. Besides its role in photochemistry and atmospheric dynamics, O_3 monitoring is crucial from an air-quality perspective, as elevated O_3 levels pose significant respiratory risks to populations already exposed to low oxygen conditions at high altitudes (~3.4 km above mean sea level and higher). Moreover, even moderate O_3 concentrations can have prolonged ecological impacts on the fragile ecosystems of the Trans-Himalaya.

Despite the discussed significance, to the best of our knowledge, comprehensive studies combining systematic in situ measurements, satellite data, and chemistry-transport modeling are lacking over the Trans-Himalaya. Tourism and urbanization have been increasing in Ladakh particularly since it became a union territory in October 2019. Therefore, it is timely to initiate studies on air quality and chemistry-climate processes in this unique environment of the world. In this regard, here, we present the first detailed investigation into the chemistry and dynamics of O_3 over Ladakh by exploiting the potentials of field measurements, satellite data, and regional and global atmospheric chemistry-transport modeling.

The remainder of the manuscript is organised in three sections. In section 2, we describe the study region, in situ measurements, satellite data, models, and air quality indices used in the study. In section 3, we present the results and discussion on diurnal variations, daytime O_3 build-up, day-to-day

changes, air quality assessment, photochemical regime of O_3 formation, seasonal variation and its drivers, and long-term trends. In section 4, we present the summary and conclusions of our study.

2. Data and methodology

2.1 Study region and in situ measurements

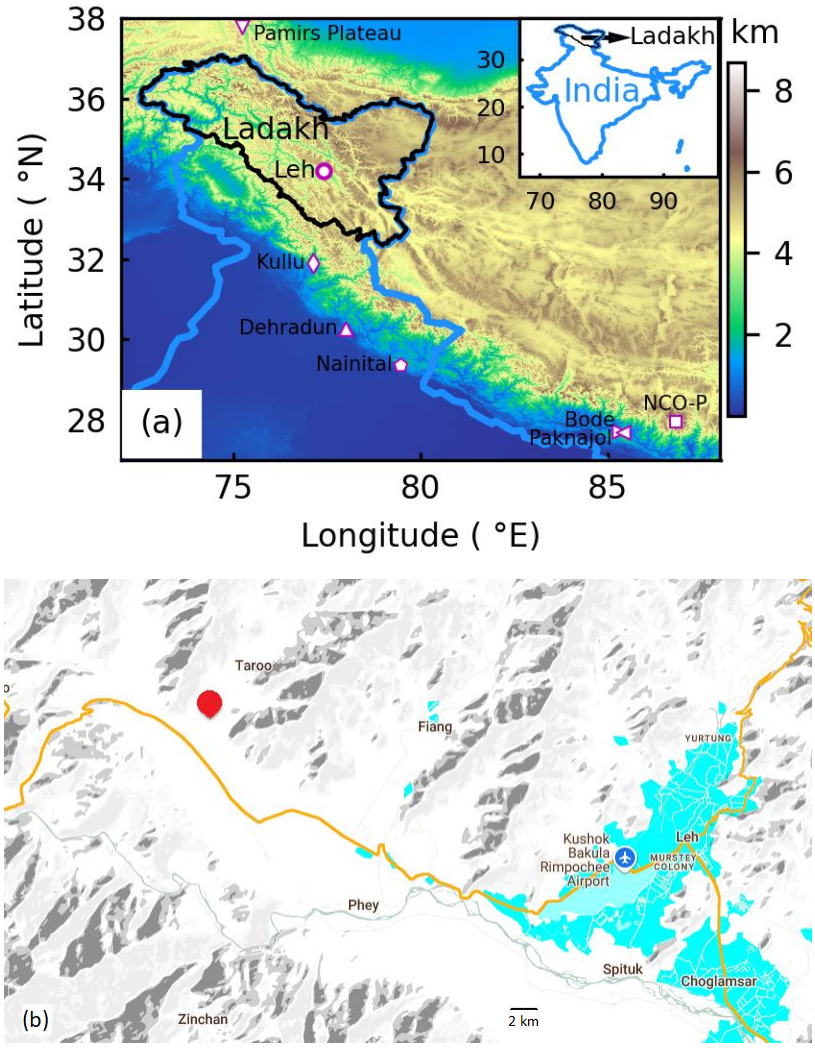


Figure 1: (a) Elevation map of the northern Indian Subcontinent. The location of the measurement site (Leh, Ladakh) is marked as a magenta circle. Other nearby sites are also shown which are used for comparison of O_3 variations. Inset in (a) shows a wider view, depicting Ladakh region in northern most part of India. (b) Zoomed map around measurement location, i.e., University of Ladakh (red symbol) and adjacent Leh town (big cyan patch). A highway is shown with an orange curve.

Surface O_3 measurements has been carried out since October 2024 at the University of Ladakh (34.185° N; 77.412° E; ~3.4 km amsl), which is ~16 km (aerial distance) away from the Leh town (figure 1b; modified from maps.google.com). No major local anthropogenic activity occurs within the radius

of ~~12–15~~2 km. Temperature, relative humidity, surface pressure and daily accumulated rain were observed in the range of -10–24 °C, 6–98 %, 651–668 hPa, 0–36.5 mm, respectively at Leh as per the India Meteorology Department. Temporal variations in these meteorological parameters are shown in figure S1. The synoptic winds at 600 hPa, obtained from the European Centre for Medium-Range Weather Forecasts (ECMWF) Reanalysis 5 (ERA5; Hersbach et al., 2020), show jumbled winds (owing to the complex topography) with downdrafts during winter and pre-monsoon. However, north-westerly or westerly winds prevailed without downdrafts during the monsoon and post-monsoon seasons (figure S2). Further details on the climate of the Leh region are available elsewhere (e.g., Chevuturi et al., 2018).

In situ O_3 measurements have been carried out using a factory-calibrated (NIST Traceable) ultraviolet (UV) photometric analyser (2B Tech, USA; Model: 106-L). Air samples were drawn from a height of ~10.6 m above the ground using a Teflon tube through a 5 µm non-reactive polytetrafluoroethylene hydrophobic dust filter. The analyser measures O_3 based on the absorption of UV radiation at ~245 nm applying the Beer–Lambert law. The lower detection limit, linearity, and minimum response time are 3 ppbv, ±1 %, and 20 s, respectively. The measurement uncertainty of the UV photometric analyser is reported to be approximately 5 % (Tanimoto et al., 2007). The instrument was continuously operated in a temperature controlled environment maintaining its permissible limits of temperature (0–50 °C). Inter-comparison of this instrument with a calibrated O_3 analyser (Thermo Scientific, USA) using a common inlet system showed an excellent agreement (figure S3), prior to its deployment at the site. Measurements were recorded at an averaging interval of 5 minutes and were converted to hourly averages for the analysis.

2.2 Satellite observations

To interpret the observed O_3 variability and assess long-term changes, the study utilized tropospheric columns of HCHO and NO_2, along with total column CO, derived from TROPOMI (TROPOspheric Monitoring Instrument) aboard the polar-orbiting Sentinel-5 Precursor satellite (equator crossing time: ~13:30 local time). NO_2, HCHO and CO are derived from the spectral measurements in 405–465 nm, around 2.3 µm and 328–359 nm, respectively. The retrieval algorithms and validations have been described in previous literature (Borsdorff et al., 2018; De Smedt et al., 2018; Landgraf et al., 2016; Van Geffen et al., 2022). Offline level-3 products of NO_2 and HCHO are available at spatial resolution of 3.5 × 5.5 km^2 and CO is available at spatial resolution of 7.5 × 5.5 km^2 (across × along track) since 28 June 2018. Data with quality assurance greater than or equal to 0.50 (0.75 for NO_2) were obtained from Google Earth Engine, https://code.earthengine.google.com.

To consider the possible influences of cloudiness on atmospheric chemistry, we have also utilized data from the Indian geostationary satellite, INSAT-3DS (Indian National Satellite - 3D second repeat). Level-2B cloud mask (L2B CMK) is derived from three infrared channels of the imager (Mid-wave infrared: 3.9 (3.80–4.00) µm, Thermal infrared-1: 10.8 (10.3–11.3) µm, Thermal infrared-2: 12 (11.5–12.5) µm) and data is available at 30-minute intervals with a spatial resolution of 4 km. Details of the algorithm can be found here: http://mosdac.gov.in/doc/INSAT_3D_ATBD_MAY_2015.pdf. We considered 25 pixels with a central pixel lying over the study site during the morning hours (8:30–14:00 IST—Indian Standard Time) and summed up the number of cloudy pixels on a daily basis. These daily sum values were used to obtain the monthly averages of the number of cloudy pixels.

2.3 Reanalysis datasets

Surface O_3 and stratospheric ozone tracer ~~(O_3s)~~ from the Copernicus Atmosphere Monitoring Service (CAMS; Inness et al., 2019), available at 3-hour interval with a spatial resolution of 0.75° × 0.75°, have been used to analyse the seasonal pattern, stratospheric contribution, and long-term trend during the 2015–2024 period. This is the fourth generation ECMWF global reanalysis of atmospheric composition (EAC4). Monthly mean vertical wind from ERA5—ECMWF Reanalysis 5 (hourly interval; 0.25° × 0.25°; Hersbach et al., 2020) is used to understand the prevailing downdraft/updraft. Both CAMS and ERA5 utilise the ECMWF Integrated Forecasting System (IFS) with 4D-var data assimilation. CAMS employs a fully integrated chemistry framework based on the Carbon Bond (CB05) mechanism, comprising 54 species and 126 reactions (Huijnen et al., 2010), with assimilation of satellite observations. Further details of the evaluation of these reanalyses can be found elsewhere (Harithasree et al., 2024; Inness et al., 2019; Park et al., 2020). The CAMS time series at grid, 33.75° N; 78.25° E, closest to the study region is considered for the analysis.

2.4. Global model simulations and source tagging

We have utilized output from dual-source tagging-enabled global model CAM4-Chem (Community Atmosphere Model with Chemistry version-4; Butler et al., 2020) available for a 19-year period of 2000–2018, as described in Ansari et al., (2025) and Nalam et al., (2025). Hourly surface O_3 mixing ratios along with its tagged contributions have been analysed at the grid cell, closest to the observation site. The model calculates O_3 contributions from both NO_x as well as VOC (including CO and CH_4) through two separately tagged simulations. Geographic definitions of specific source regions are shown in figure S4. Monthly varying gridded anthropogenic emissions of NO_x, CO, SO_2, NH_3, OC, BC and NMVOCs, input to the model, are based on the HTAPv3.1 global mosaic inventory for the 2000–2018 period (Crippa et al., 2023). Biogenic emissions are included from the CAMS-GLOB-BIO-v3.0 (Sindelarova et al., 2022) while biogenic soil NO_x is specified as in Tilmes et al. (2015). Biomass burning

emissions are prescribed from GFED-v4 inventory (Van Der Werf et al., 2010). The model has a horizontal resolution of 1.9°×2.5°, with 56 vertical levels extending up to model top at approximately 1.86 hPa. The model setup and evaluation results assessing the same global simulations have been discussed in recent studies over North America and Europe (Ansari et al., 2025), East and Southeast Asia (Lu et al., 2025), and the global troposphere (Nalam et al., 2025). In this analysis, stratospheric O_3 (O_3s) contribution to surface O_3 has been considered from CAMS reanalysis as this adjustment improved the agreement of CAM4-Chem simulation with observations. For this, CAM4-Chem estimated O_3s has been subtracted from CAM4-Chem estimated O_3 and then CAMS derived O_3s has been added.

Here, O_3 source attribution is performed using the TOAST 1.0 tagging system developed by Butler et al., (2018). The system introduces additional tagged species into the model's chemical mechanism and physical process routines, allowing them to be transported and processed alongside the original chemical species that represent the total baseline emissions. Two parallel simulations are conducted: one tracking the contributions of NO_x emissions and the other tracking those of volatile organic compounds (VOCs). In each simulation, tagged species represent emissions from a specific source region or sector, with the corresponding tag indicated by a variable suffix (see table S1 and figure S4 for the regional and sectoral definitions). For example, NO_NAM, NO_SAS, and NO_EAS represent NO emitted from North America, South Asia, and East Asia, respectively. As tagged species participate in chemical reactions, their source labels are propagated through the reaction pathways and ultimately assigned to the resulting O_3 molecules. Thus, O_3 formed from NO_x originating in South Asia is represented as O_3_SAS in the NO_x-tagging simulation, while O_3 formed through the oxidation of VOCs originating in the same region is represented as O_3_SAS in the VOC-tagging simulation.

Separate simulations are necessary because O_3 formation involves interactions between NO_x and VOCs that may originate from different regions or sectors. A single O_3 molecule can therefore be associated with both a NO_x source and a VOC source; the two simulations independently quantify these contributions. The tagging system also distinguishes between NO_2 formed through the oxidation of NO and NO_2 formed through O_3 titration. When a tagged NO molecule (e.g., NO_NAM) reacts with an O_3 molecule carrying a different regional tag (e.g., O_3_SAS), the resulting NO_2 is assigned the O_3 molecule's tag (NO_2_SAS), rather than the NO molecule's tag (NO_2_NAM). This convention preserves the source attribution of O_3 consumed through titration and distinguishes it from NO_2 produced through NO oxidation. Further details of the tagging methodology and simulation configurations are provided by Ansari et al. (2025) and Butler et al. (2018).

Note that since the latest emissions for 2024–2025 are unavailable, we have used simulation results from previous years, 2009–2018, as climatological mean.

2.5 Regional model simulations

We have performed regional-scale numerical simulations of meteorological parameters and atmospheric composition at spatial resolution of 12 km x 12 km, by applying the Weather Research and Forecasting model coupled with Chemistry (WRF-Chem; Grell et al., 2005; Skamarock et al., 2019) version-4.5.1. Simulations have been conducted for 2 full months, February representing winter and September representing monsoon. Meteorological fields input to the model are based on ERA5 reanalysis (Hersbach et al., 2020). For parametrizing cumulus physics and boundary layer processes, the Grell 3D (Grell and Dévényi, 2002) and Yonsei University (Hong et al., 2006) schemes, respectively, have been utilized. To limit errors in large-scale meteorological and dynamical variations, the temperature, horizontal winds, and water vapour have been nudged at all vertical levels with nudging coefficients of 3 x 10^{-4} s^{-1} (Otte, 2008; Sharma et al., 2017). Here, nudging coefficient is kept lower than previous studies over South Asia (6 x 10^{-4} s^{-1}; Kumar et al., 2012a, b), to minimise the potential suppression of fine-scale local circulations simulated by the WRF dynamics.

Anthropogenic emissions are based on the EDGAR (Emissions Database for Global Atmospheric Research) inventory for year 2015 (Crippa et al., 2020; Mogno and Marvin, 2022) and fire emissions are based on FINN v2.5 (Fire INventory from NCAR (National Center for Atmospheric Research); Wiedinmyer et al., 2023) for year 2021. Natural biogenic emissions are computed online using the MEGAN (Model of Emissions of Gases and Aerosols from Nature Guenther et al., 2006). Gas-phase chemistry is represented by the MOZART (Model for Ozone and Related chemical Tracers) mechanism (Emmons et al., 2010) and aerosols are calculated using the MOSAIC (Model for Simulating Aerosol Interactions and Chemistry; Zaveri et al., 2008). Initial and lateral boundary conditions for the chemical fields have been provided from the WACCM (Whole Atmosphere Community Climate Model) global simulations (Gettelman et al., 2019). We have noted a systematic overestimation of O_3 in remote environments of India and that reducing inflow from the global model by about 40 % leads to better agreement between WRF-Chem output versus measurements. Comparison of model simulated surface O_3 with measured O_3 shows a good agreement (r=0.4–0.8) with smaller negative bias of 7–9 ppbv (figure S5). This shows the model's capability to capture the observed variations reasonably and hence it can be used to gain deeper insights into the various processes contributing to observed variations. The tendency terms in the WRF-Chem (Barth et al., 2012; Girach et al., 2017) have been analysed to quantify roles of various chemical and dynamical processes in influencing O_3 variations. More detailed description of the WRF-Chem model (Grell et al., 2005) and its evaluations and diverse

applications over Indian subcontinent can be found elsewhere (Ghude et al., 2025; Girach et al., 2017; Kumar et al., 2012a, b; Sharma et al., 2017).

Here, EDGAR inventory for the year 2015 has been used due to the lack of up-to-date emission changes over this region in inventories. As discussed in section 3.7, the anthropogenic emissions are increasing over the study region and hence actual emission could be different. Nevertheless, here we use WRF-Chem simulation only to infer the main processes influencing the diurnal and day-to-day changes in O_3.

2.6 Indices for air quality assessment

Maximum Daily 8-h Average ozone (MDA8) is related to short-term health impact with a threshold value of 100 µg m^{-3} set by the World Health Organization (WHO, 2021) as well as by India's National Ambient Air Quality Standards (NAAQS; https://cpcb.nic.in/uploads/National_Ambient_Air_Quality_Standards.pdf; last access: 03 January 2026). MDA8 is a maximum value of 8-h running averages obtained using hourly O_3 on a daily basis. Hourly O_3 mixing ratio (ppbv) is converted to concentration (µg m^{-3}) using measured pressure and temperature (figure S1). To quantify health impacts due to long-term O_3 exposure, another metric OSDMA8 (ozone season daily maximum 8 h average) has been estimated, which is basically the annual maximum of six-month running mean of MDA8 (Murray et al., 2020). The risk to vegetation/crops and ecosystems can be assessed based on the estimation of four different metrics: M7 (7-hour Mean, averaged during daytime, 09:00–16:00 h), AOT40 (Accumulated exposure to Ozone above a threshold of 40 ppbv), SUM06 (the sum of all hourly O_3 concentrations ≥0.06 ppm) and W126 (hourly concentrations weighed by a sigmoidal weighting function). AOT40, SUM06 and W126 are calculated for daylight hours, 07:00–18:00 h (Fuhrer et al., 1997; Hogsett et al., 1988; Lee et al., 1988; Lefohn and Runeckles, 1987; Mills et al., 2007). SUM06 and W126 better represent the crop loss as found in the experimental studies in the US (Kaylor et al., 2023). Threshold value or safe limit for AOT40 is 1 ppmv h as defined by UNECE (United Nations Economic Commission for Europe), 1996 (Fuhrer et al., 1997; https://www.umweltbundesamt.de/system/files/medien/4038/dokumente/manual_complete_english.pdf). There are no thresholds defined for M7, W126 and SUM06 as lower levels of O_3 can also have adverse impacts on vegetation. Formulae for these exposure indices are given in supplementary text S1.

3. Results and Discussions

3.1 Diurnal variations and daytime O_3 build-up

ΔO_3 for a given day is obtained by subtracting the daily mean from the hourly values of the corresponding day. The mean diurnal patterns of ΔO_3 for different months are shown in figure 2. The average diurnal variations for different seasons are shown in figure S6. The diurnal patterns of ΔO_3 are observed to be quite unique in this region, nevertheless, similar across different months, with three important features: (a) nighttime (22–7 h) decline, (b) build up after sunrise to the noontime (7–14 h) followed by decline in the afternoon (13–17 h), and (c) increase in the evening (17–20 h). Noontime peak values are comparable or lower than midnight/nighttime values. Since daytime values are generally lower than nighttime values by 0–3 ppbv, there appears to be a net chemical loss on a daily scale. However, to a certain extent, this shows that O_3 loss occurring during nighttime (20–6 h) is partially compensated by daytime photochemical production. The day-to-day variability, as represented by the standard deviation (shaded region in figure 2) is seen to be larger (3–5 ppbv) during monsoon, as compared to other seasons (1–4 ppbv). This is likely due to sporadic mixing of different air masses owing to changes in meteorological conditions, especially day-to-day changes in convection strengths during July to September (Romatschke and Houze, 2011).

The rate of increase from morning (7–9 h) to afternoon (12–14 h) is estimated as the slope of linear fit to five successive hourly values from the morning minimum value, which ranged between 0.6–1.5 ppbv h^{-1}. The photochemical O_3 build-up, estimated as the noontime maximum minus morning minimum, ranged from 2.6–6.8 ppbv. O_3 concentration starts increasing in the morning 1.5–2 h after sunrise. The O_3 curve attains its peak 0–2 h after the local noon (zenith). The evening increase is mostly before sunset (figure 2). Interestingly, the evening O_3 peak is higher than the noontime peak for most of the months, except during winter.

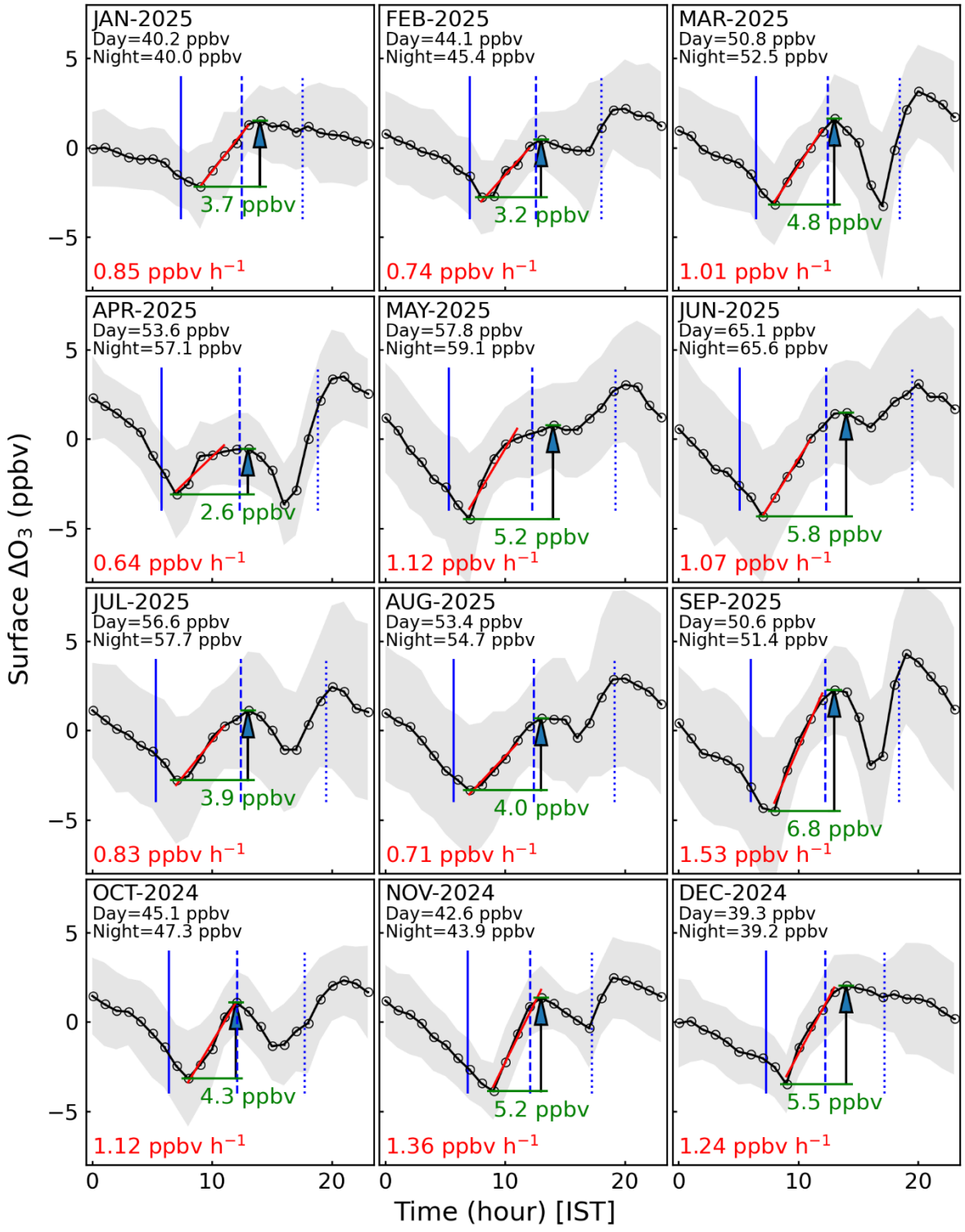
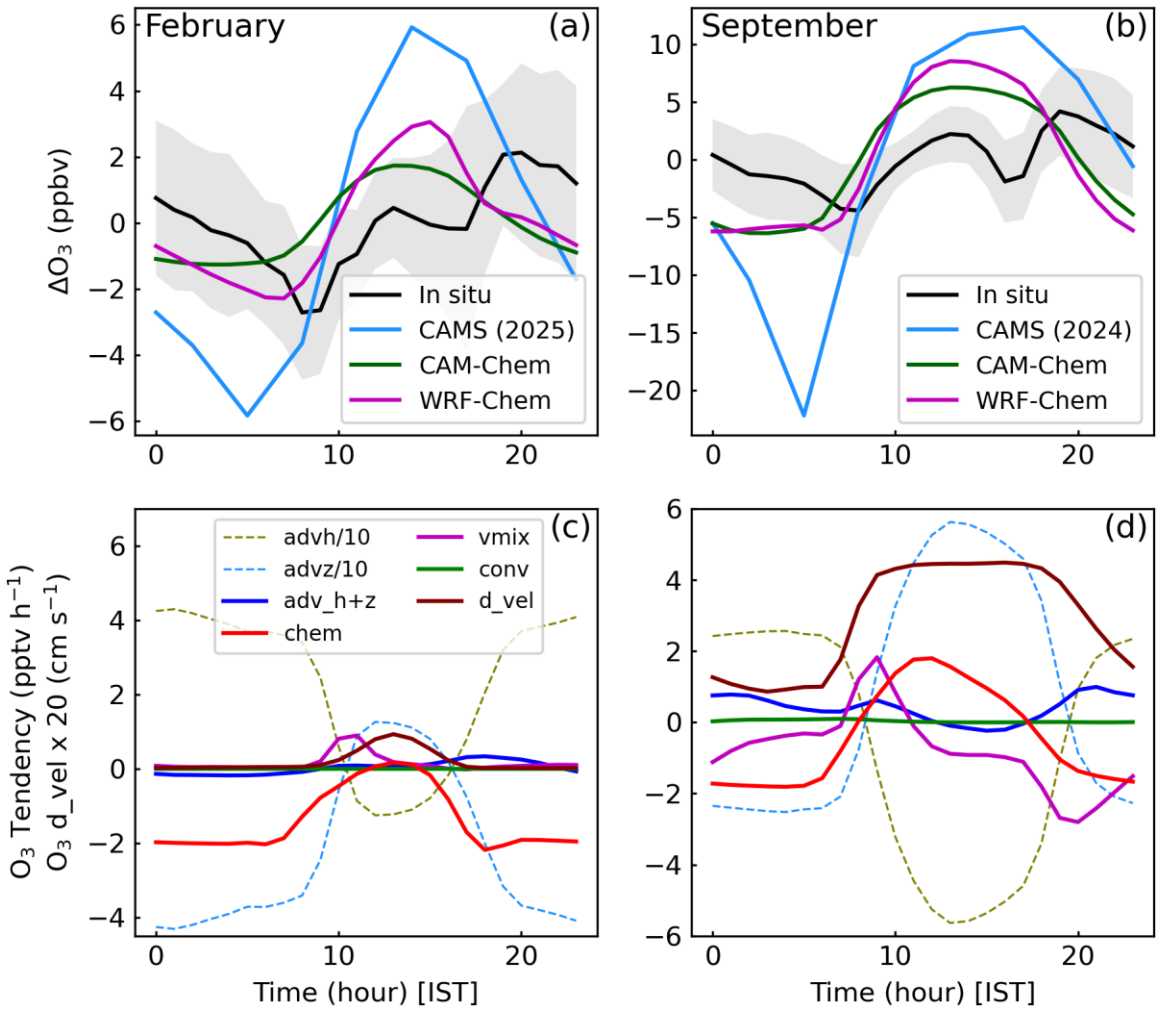


Figure 2: Monthly mean diurnal variations of ΔO_3 over the Trans-Himalaya. The shaded grey region shows the standard deviation. The red line shows the linear fit over the 5-h period from the morning minimum. The slope representing net O_3 production rate (ppbv h^{-1}) is given in red colour font. Total daytime build-up, noontime maximum minus morning minimum, is marked in green colour. Vertical blue lines represent sunrise (continuous line), local noon (dashed line) and sunset (dotted line). Mean

values of daytime and nighttime O_3 are also mentioned. Here, daytime is defined as 2 hours after sunrise until 2 hours before sunset. Likewise, night is defined.

It is important to note that the study region experiences a photochemical O_3 build-up after sunrise, a feature typically unseen over other high-altitude sites such as Nainital (1.9 km amsl, Sarangi et al., 2014), Mount Abu (24.6° N, 72.7° E, 1.7 km amsl, Naja, 2003), Mount Fuji (35.4° N, 138.7° E, 3.7 km amsl; Tsutsumi et al., 1994), Ooty (11.4° N, 76.7° E, 2.5 km amsl; Udayasoorian et al., 2013) and Ponmudi (8.76° N, 77.11° E, 1 km amsl; Ajayakumar et al., 2024). In particular, the ΔO_3 diurnal patterns show decline after the sunrise or remain nearly stable throughout the day over high-altitude sites in the Himalaya (Nainital and NCO-P; figure S6). Afternoon time O_3 enhancements at Nainital was attributed to transport of pollutants from adjoining low-altitude IGP (Indo-Gangetic Plain) through boundary layer mixing. This comparison suggests that unlike the central/eastern Himalaya, there are significant precursors over the Trans-Himalaya which can result in observed photochemical O_3 production in the presence of intense solar radiation. Note that due to its higher elevation, the Trans-Himalaya region receives more intense solar radiation (7 kWh m^{-2} d^{-1}) than central Himalayas or other lower elevation sites (5.6 kWh m^{-2} d^{-1}) in India (Santra, 2016). Note that here O_3 production rate is slower (0.6–1.5 ppbv h^{-1}) than that in polluted valleys and foothills of the Himalaya (4.7–10.1 ppbv h^{-1}; Ojha et al., 2019). For further insight into photochemistry, we analysed co-located satellite observations to identify O_3 formation regime which are discussed later in section 3.4.


February
(a)
September
(b)
ΔO3 (ppbv)
In situ
CAMS (2025)
CAMS (2024)
CAM-Chem
WRF-Chem
(c)
(d)
advh/10
advz/10
adv_h+z
chem
vmix
conv
d_vel
O3 Tendency (pptv h−1)
O3 d_vel x 20 (cm s−1)
Time (hour) [IST]

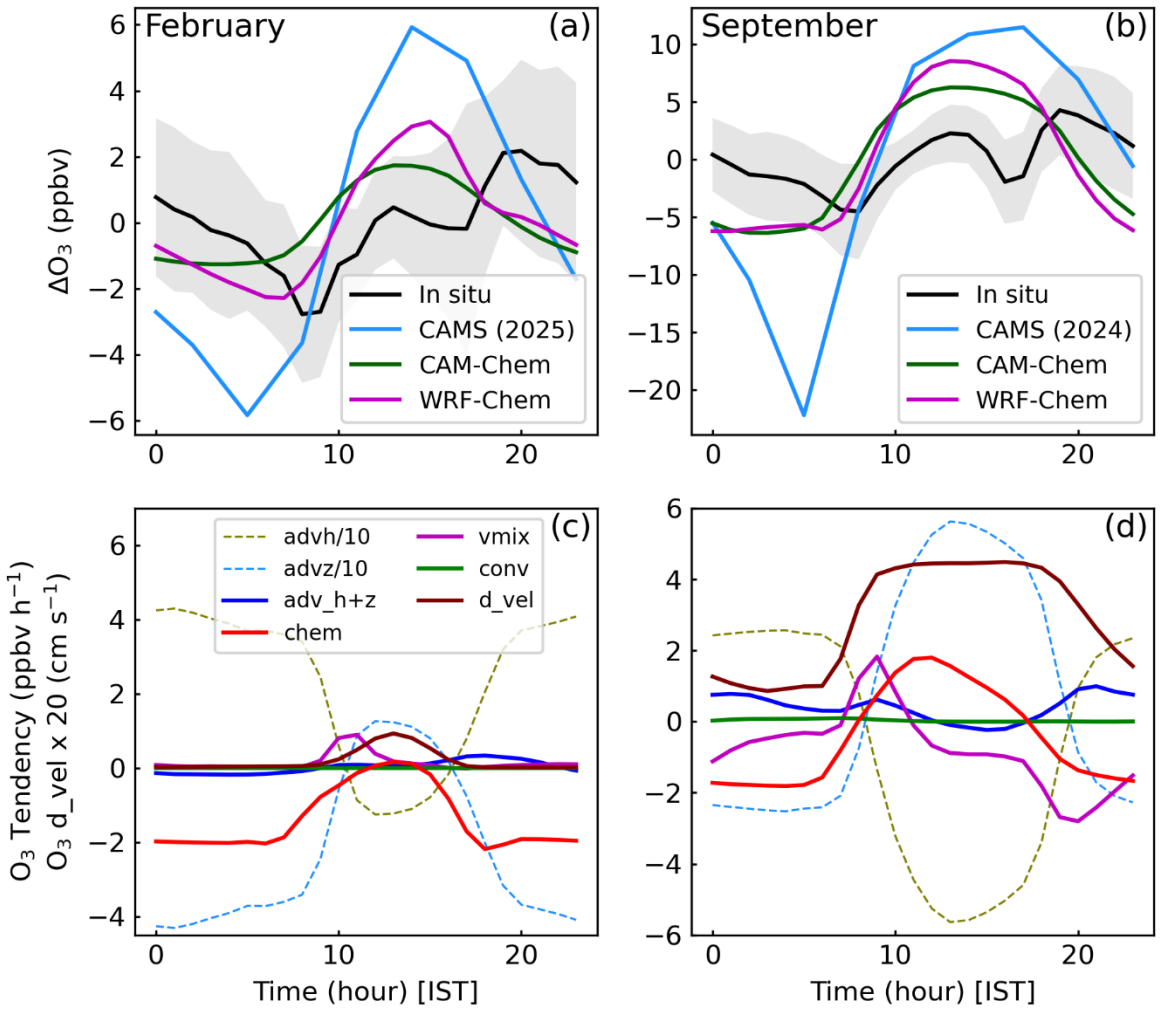


Figure 3: (a–b) Comparison of observed diurnal ΔO_3 variation at Leh with simulations from global (CAMS and CAM4-Chem) and regional (WRF-Chem) models. WRF-Chem inferred deposition velocity (d_vel) and mean tendencies due to different processes: horizontal advection (advh), vertical advection (advz), convection (conv), chemistry (chem), and vertical mixing (vmix) are shown in (c–d).

Qualitatively, the models reproduce the diurnal ΔO_3 pattern, except missing the late evening-time enhancement. Quantitatively, diurnal amplitude in CAM4-Chem as well as WRF-Chem models matches with the observations (~3 ppbv) during February. Both models do show an increase in diurnal amplitude in September but overestimate its magnitude (12–15 ppbv) than observed values (7 ppbv). However, chemistry appears much more intense in the CAMS reanalysis particularly during September. Our results suggest that simulations from CAM4-Chem and WRF-Chem may be analysed for deriving further insights into the factors governing O_3 magnitude and variations. Few other limitations which we are reporting here, although are not quite surprising, would be useful for future improvements and fine-tuning of the models. It is likely that the evening surge is associated with mesoscale dynamics, transport of O_3 rich air from higher altitudes due to mountain-valley circulation

(Sarangi et al., 2014) or through other orographic influences unresolved even at high resolutions (Singh et al., 2021) in complex Himalayas. Alternatively, the afternoon dip (13–17 h) could have been associated with local-scale dynamics, which could not be represented adequately in the model.

The tendency terms from WRF-Chem simulations have been analysed to infer the main processes governing O_3 diurnal variation during two contrasting seasons (figure 3c–d). The daytime build-up during winter is linked with decreased ~~photo~~chemical loss and increased contribution through vertical mixing in the morning hours (figure 3c). Small contributions from the advection are seen in the evening hours. Contrary to winter, larger photochemical production drives daytime increase of O_3 during monsoon season (figure 3d). Interestingly, while vertical mixing positively contributes to enhanced O_3 in the morning hours, it reduces O_3 in the evening hours in the model. Positive advective contributions help to sustain O_3 levels during the monsoon throughout the day, except near noontime. The deposition velocity of O_3 is elevated during daytime, especially in the monsoon, but it is negligible during nighttime in both seasons. The results suggest higher depositional loss of O_3 in summer as compared to the winter. A slow decrease during nighttime is due to chemical loss plus vertical mixing as seen by negative chemical tendencies during both seasons (figure 3c–d). Thus, the model fields provided key insights into the relative roles of different processes shaping O_3 diurnal pattern, besides indicating the factors that may need improvements to reduce the model-measurement differences such as in late evening or night.

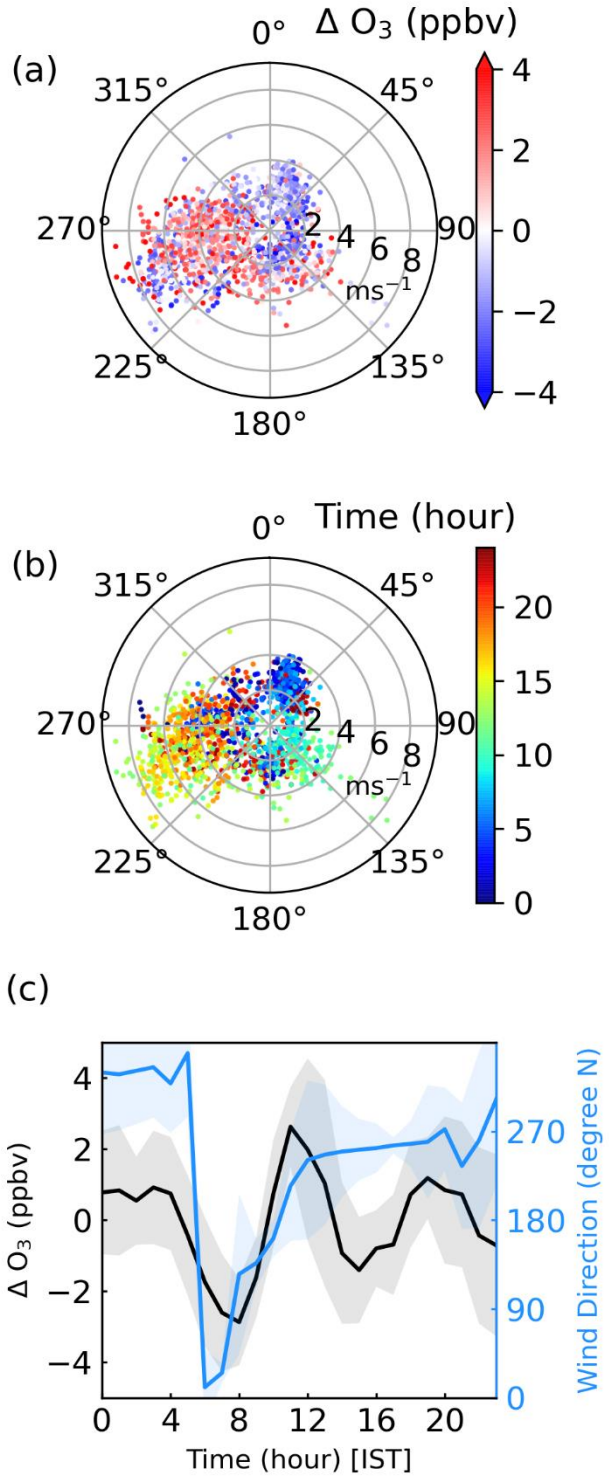


Figure 4: Wind rose colour coded with ΔO_3 and time (a–b) and diurnal variations in ΔO_3 and wind direction (c) over the study region during 23–28 October 2024. Shaded regions in (c) represent standard deviation.

The diurnal pattern of O_3 is also linked with the wind direction, which would transport different airmasses during the course of the day. To infer this effect, a few days of collocated measurements of

wind and ΔO_3 (23–28 October 2025) have been analysed (figure 4). While the lower O_3 in the morning hours are associated with mainly wind direction 0–60° N, higher values are associated with the wind from 90–315° N (Northwest, southwest and southeast sectors) during noon to late night hours. Further, the standard deviation (figure 4c) estimated following Yamartino, 1984 suggests significant variability in the wind direction. The climatologies of satellite observed precursors, NO_2 and CO, also show relatively higher levels over valley regions of the Trans-Himalaya (figure 5a and c), indicating local emissions. Relatively higher anthropogenic activities including vehicular emission (note the national highway in valley region, south of the measurement site; figure 1b) is evident from the relatively higher concentration of NO_2 and CO (figure 5d and f). Therefore, the noontime higher O_3 or noontime build-up corresponding to wind directions 90–315° N (figure 4), is likely to be contributed by emissions in the adjoining areas (figure 5d, f; south and southwest of the measurement site). Tropospheric column NO_2 over the observational site is in the range of ~0.2–1.7 molecules cm^{-2}, which translates to the boundary layer mixing ratio of 0.1–0.9 ppbv with assumption that 90% of tropospheric NO_2 is within 1 km of boundary layer height. This would contribute to the observed daytime build-up of O_3. In addition, there exists afternoon dip in O_3 over the same wind sector which could be due to titration by fresh emissions. The lower O_3 in the morning hours is likely to have minimal local influence as precursors are lower in the north of the study site (0–60° N; figure 4a–b, 5d and f). The higher NO_2 in the east or southeast of the study site represents Leh town and its direct influence on the study site is unlikely due to mountains in-between (figure 1b). Additionally, the lower HCHO levels (proxy of VOCs) would limit the in situ photochemical production of O_3 over the site. Nevertheless, further observations in the Trans-Himalaya related to chemical composition and dynamics would be needed for pinpointing the role of local emissions and fine-scale dynamical processes.

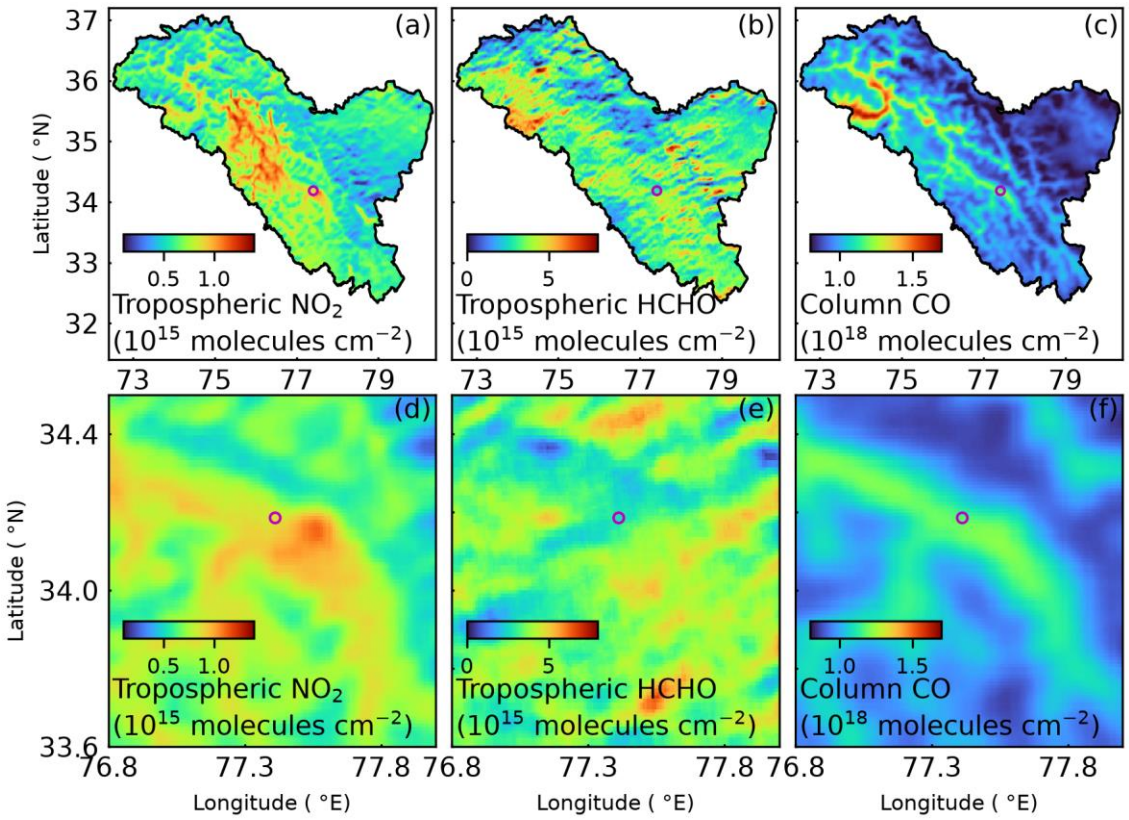


Figure 5: Mean distribution of TROPOMI observed NO2, HCHO and CO over the study region during October 2018–September 2026. Magenta circle marks the study site. Figure d–f is the zoomed view around the study site.

3.2 Day-to-day variations and contributing processes

The day-to-day changes in noontime (13 h) values of O_3 and different tendency terms, as derived from the WRF-Chem simulation, are shown in figure 46. For this, change in O_3 is obtained by subtracting values at 13 h on $(n-1)^{th}$ day from value at the same time on n^{th} day. Similarly, tendencies shown here are obtained the same way from WRF-chem simulated accumulated tendencies. Generally, the tendencies (difference between consecutive values of cumulative tendency at hourly interval) oscillate around a baseline reflecting embedded diurnal changes and the sum of all tendencies (conv_ten, chem_ten, vmix_ten, and adv_ten) represent the hourly net change in O_3. Here, day-to-day changes in O_3 have been analysed by estimating the tendencies at the daily scale (ppbv d^{-1}; figure 6). To assess the role of different tendencies in changing O_3, correlation coefficient and normalised covariance are also estimated between tendencies versus O_3 change as mentioned in the figure. Day-to-day changes in noontime O_3 are smaller in February (up to 10 ppbv d^{-1}) compared to September (up to 20 ppbv d^{-1}), as seen in figures ~~4b~~ 6b and i. Chemical tendencies are typically negative over this region (figure ~~3d~~6d, k), being -30 to -40 ppbv d^{-1} in February but 0 to -10 ppbv d^{-1} in September (except

during convective episode; 8–11 September). This indicates that the region acts as a net chemical sink of O_3 during both the seasons, being a stronger sink during the winter season. The gradual nighttime (20–6 h IST) declines seen in the observations (figure 2) also support the role of titration of O_3. Monthly sum of hourly changes in observed O_3 turns out to be negative during most of the months, except in January, March, June, and August (figure S7). This is consistent with the model-based inference that the region acts as a net chemical sink. Dynamics (convection plus advection) compensated more than the stronger chemical losses during 8–10 September, leading to overall O_3 increase. Little variations in chemical losses in other days but strong anti-correlation in this period led to an overall negative correlation between chemical tendencies and O_3 change in September.

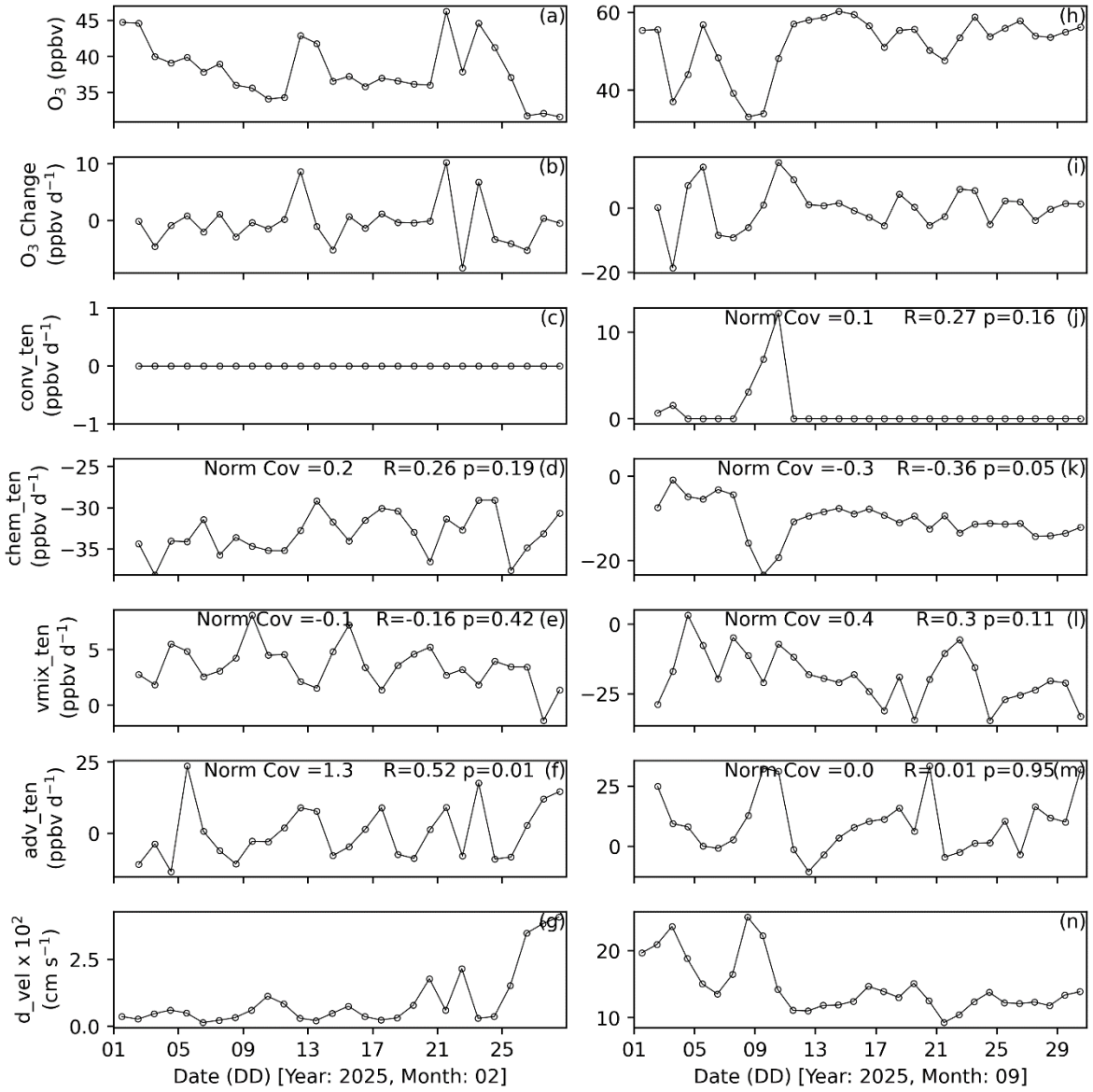


Figure 46: Noontime O_3 (13 h IST), its changes with respect to the previous day, various tendencies and deposition velocity during representative months of winter (a–g) and summer monsoon (h–n)

seasons. The y-axis labels, conv_ten, chem_ten, vmix_ten, adv_ten, and d_vel, represent O_3 tendencies (ppbv d^{-1}) due to convection, chemistry, vertical mixing, advection, and O_3 deposition velocity (cm s^{-1}). Correlation coefficient (R) and normalised covariance are between O_3 change (ppbv d^{-1}; b and i) and tendencies (c–f, j–m). Here, covariance is normalised with respect to variance in O_3 change (b and i). The p-value is also marked along with R.

The advection tendencies are generally positive and cause an increase in O_3. Transport due to advection has a greater bearing on O_3 change during winter, as inferred from the higher correlation and covariance (figure ~~4f~~6f; R=0.5, normalised covariance 1.3). In contrast, no single process dominates day-to-day O_3 changes during monsoon. The contributions from multiple processes govern the O_3 changes during monsoon.

Convection is not a key factor in general, except during a large increase in O_3 during 8–11 September when it exceeded 10 ppbv d^{-1} (figure ~~4c~~6c, j). Deposition velocity also shows significant variations (February: 0.01±0.01 cm s^{-1}; September: 0.15±0.04 cm s^{-1}) with enhancements during convective episodes (e.g., around 9 September 2025).

3.3 Air quality impacts

Our first systematic O_3 measurements can provide insights into O_3 impacts on the public health and vegetation over the Leh region. In this regard, several standard and widely applied exposure metrics have been estimated and compared with the thresholds (figure ~~5~~7). All O_3 exposure indices are observed to be higher during May–July, peaking in the month of June. MDA8 values ranged from 52–104 µg m^{-3} (37–80 ppbv) during the study period. MDA8 values are very close to the threshold of 100 µg m^{-3}, with 33 days experiencing values above 90 µg m^{-3} and ~~once~~ thrice exceeding the threshold during May–July. OSDMA8 is 82 µg m^{-3} (or 61 ppbv), which is significantly higher than the minimum risk level of 32.4 ppbv (Turnock et al., 2025). The AOT40 values are above the threshold of 1 ppmv h except during November–February. Annual cumulative values of AOT40, SUM06 and W126 are estimated to be 44.9, 44.6, and 40.9 ppmv h, respectively. Magnitudes of these indices are comparable with those over the US and South Korea (Kaylor et al., 2023; Lee et al., 2023). AOT40, SUM06 and W126 have been reported to range between 0–5, 1–20, and 1–15 ppmv h, respectively, over China during the crop growing season of winter (Jiao et al., 2025), which are comparable or slightly higher than the present study region. Increase in AOT40 between 13 to 17 ppmv h is also reported over the Indian region during wheat growing season (January ~~–~~ March) of 2005–2021 (Anagha and Kuttippurath, 2026). With increasing global baseline O_3 and anthropogenic emissions across South Asia, our results indicate likelihood of more frequent exceedance of air quality standards in this region.

In this regard, O_3 formation regime (Section 3.4) and long-term trends (Section 3.7) have been analysed for informing the policies.

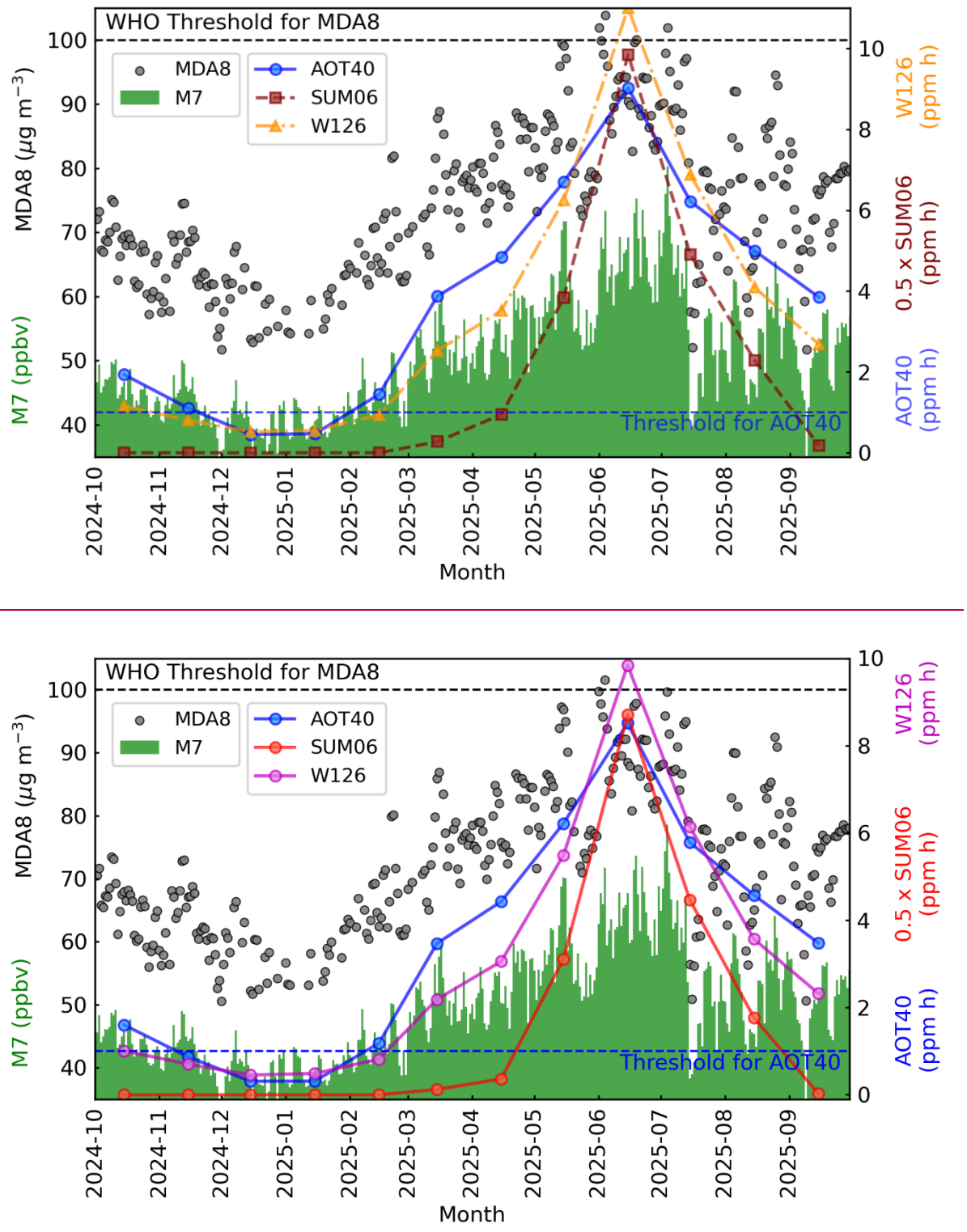


Figure ~~5~~7: Variations in various ozone exposure metrics during the study period. Threshold values of MDA8 (100 µg m^{-3}) and AOT40 (1 ppm h) are indicated by horizontal dashed lines.

3.4 O_3 formation regime

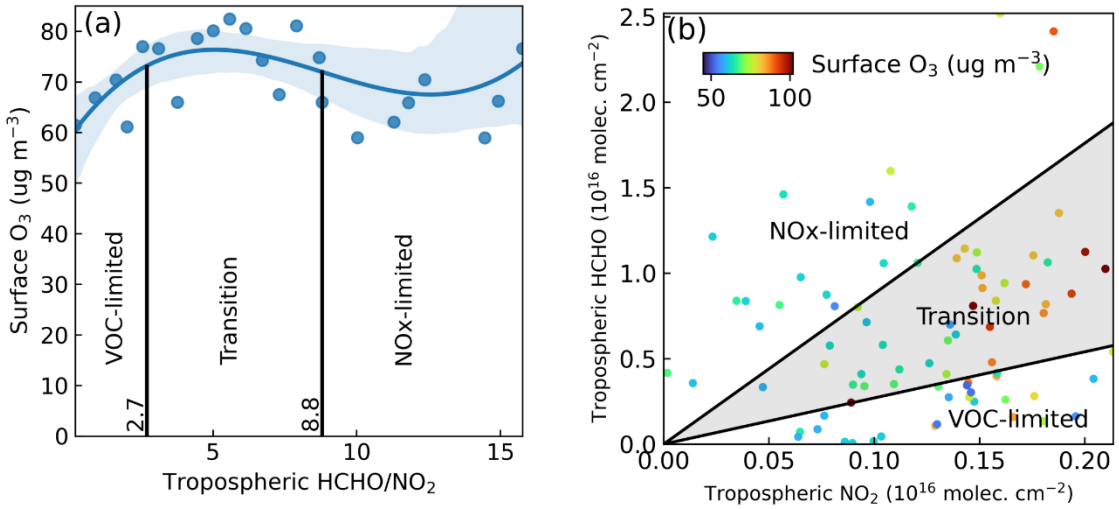


Figure ~~6~~8: (a) Variations of surface O_3 with collocated tropospheric $HCHO/NO_2$ ratio over the study region. Blue curve represents the 3rd order polynomial fit with 95 % confidence level (shaded region). (b) Scatterplot of tropospheric HCHO versus NO_2 colour-coded with surface O_3. Black lines in (a) and (b) demarcate O_3 formation regimes, as labelled.

TROPOMI observations of tropospheric HCHO and NO_2 have been used to determine the threshold values of the $HCHO/NO_2$ ratio to demarcate NO_x and VOC-limited regimes suitable for the study region, following the literature (Jin et al., 2017; Martin et al., 2004). TROPOMI observations at a single pixel strictly co-located with ground-site are considered for the spatial co-location. Surface O_3 values are averaged over a ±15-minute time window with respect to the satellite overpass (i.e., 30 minutes centred on the overpass time). The $HCHO/NO_2$ ratio is divided into 400 bins, and corresponding averages of O_3 are fitted with a 3rd order polynomial as shown in figure ~~6a~~8a. Since the peak of the fitted curve is the central value in the transition regime, i.e., ridge of O_3 isopleths, a slope from -3 to +3 of the fitted curve demarcates the VOC-limited versus NO_x-limited regimes (Ren et al., 2022), corresponding to $HCHO/NO_2$ ratios of 2.6 and 8.8, respectively (figure ~~6a~~8a). Therefore, the values less than 2.7 represent a VOC-limited regime, whereas the values higher than 8.8 represent NO_x-limited regime. Most of the higher values of O_3 lie in the transition regime (figure ~~6b~~8b) suggesting that the O_3 formation corresponding to higher values is contributed by both NO_x and VOCs and hence the effective mitigation will require reductions in both NO_x and VOCs. In addition, since the O_3 levels are particularly higher during May–July (figure ~~5~~7), when NO_2/CO ratio suggests prevalence of "relatively" fresh airmasses (compared to winter season; see figure S8), reduction in regional emissions in addition to local emissions would be essential to contain the O_3 level within safe limits.

Here, it is worth mentioning that while satellite-retrieved HCHO/NO_2 ratios do provide insights into the chemical regimes, the inferences could have uncertainties. In this regard, particularly important factors include (a) higher retrieval noise under low concentrations over mountain region, which could result in uncertainty by >100% in HCHO/NO_2 ratio, (b) inherent chemistry assumption leading to 20% standard error due to inaccurate description of the ROx–HOx cycle, (c) column-to-boundary layer translation, and (d) limited representativeness due to coarse spatial resolution missing sharp gradients in concentration across satellite pixel (Acdan et al., 2023; Souri et al., 2023). Few studies have reported positive biases of ~26% and 9.5% in TROPOMI HCHO and NO_2 respectively for cleaner regions (HCHO < 2.5×10^{15} and NO_2 < 2×10^{15} molecules cm^{-2}) (De Smedt et al., 2021; Pasternak et al., 2026). Thus, further in-situ measurements would be needed for validation of the results derived from satellite retrievals for this environment having strong heterogeneity in topography and tracer's distribution.

3.5 Seasonal variation: role of stratospheric intrusion and photochemistry

Figure ~~7~~ 9 shows the seasonal variations of observed O_3 and related parameters derived from reanalysis and satellite observations over the study region. O_3 levels are lower during winter (39.2±3.9 ppbv in December), increase in spring and attain maximum in June, just before the summer monsoon (65.2±5.8 ppbv). Note that since summer monsoon onset is late, July rather than June, here monsoon period is from July to September (Bhan et al., 2015). The amplitude of the observed O_3 seasonal cycle is 26 ppbv. CAMS derived O_3s climatology shows its peak (11.4 ppbv) during June explaining about 18 % of the seasonal maximum of O_3 over this region (figure ~~7a~~9a). This result is broadly consistent with more stratospheric intrusions over the Tibetan Plateau/China in the month of June (Liao et al., 2025). The stratospheric contribution is also substantial in winter (8–10 ppbv) but lower during other seasons (5–~~–~~6 ppbv). Vertical wind at 650 hPa (close to the observed surface pressure of 651–668 hPa; figure S1) from ERA5 shows that the study region experiences downdraft consistently, except during the monsoon (see background colours in figure ~~7a~~9a). Strong downdrafts during winter with higher potential vorticity (5–10×10^{-6} K kg^{-1} m^2 s^{-1}; figure S9) provide conducive conditions for downward transport of O_3-rich air, in line with the analysis of CAMS O_3s. During June 2025 when O_3 peaked, potential vorticity remained above 2×10^{-6} K kg^{-1} m^2 s^{-1} suggesting the presence of stratospheric influence. The spatial distribution of O_3s (not shown here) suggests that its enhanced contribution during June is not due to direct downward transport, instead it is due to horizontal transport from the Tibetan Plateau region. Stratospheric intrusions have been shown to increase surface O_3 over Tibetan Plateau (e.g., Liao et al., 2025) with which this model-based inference is consistent.

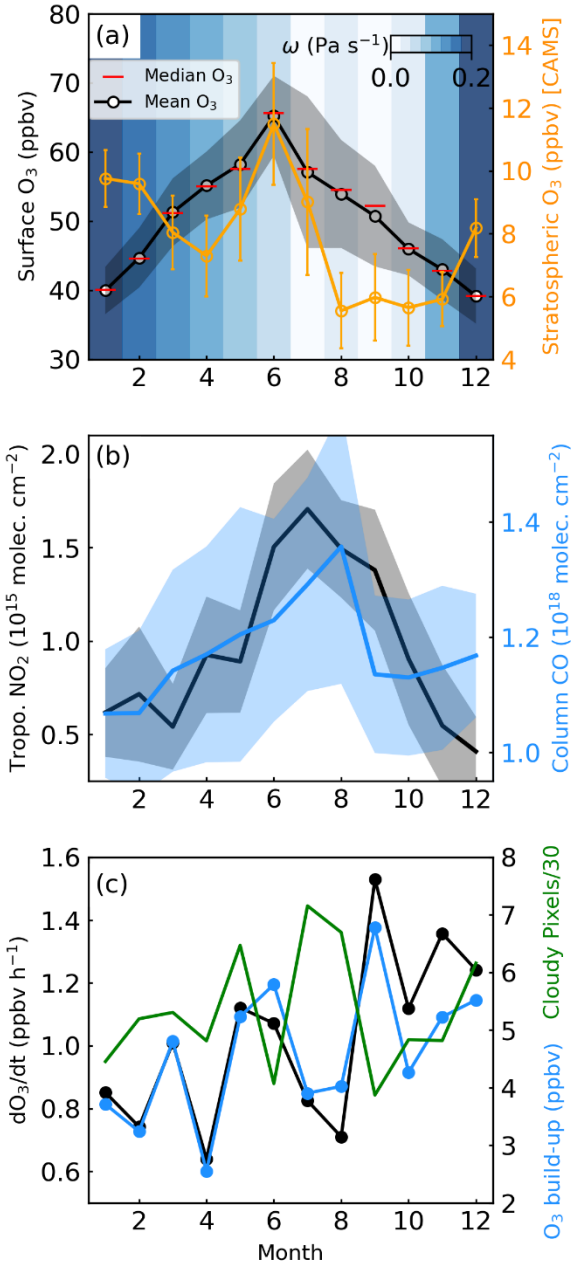


Figure ~~7~~9: Seasonal variation in (a) Observed O_3, vertical wind, and stratospheric O_3 tracer (O_3s), (b) satellite-derived tropospheric column NO_2 and total CO column, and (c) observed rate of O_3 increase (dO_3/dt), daytime O_3 build-up, and number of cloudy pixels derived from INSAT-3DS (8:30–14:00 IST) over the Leh region. All parameters are for the measurement period (October 2024–September 2025), except O_3s for which climatology (2015–2024) is shown. The shaded region in (b) shows the standard deviation. Here, O_3 build-up is the same as in figure 2.

Tropospheric NO_2 and total CO columns from TROPOMI over the study region show higher values (NO_2: 1.4–1.7×10^{15} molecules cm^{-2}; CO: 1.2–1.4×10^{18} molecules cm^{-2}) during monsoon (figure ~~7b~~9b).

NO_2/CO ratio peaks during June-July (figure S8) suggesting presence of photochemically "fresh" airmass likely transported from nearby regions, besides local emissions. The tourist activities are generally higher during summer (May–September) peaking in June, as compared to winter (Dolma, 2020). Larger abundances of the precursors in presence of intense solar radiation provide conducive conditions for photochemical O_3 production during June. Figure ~~7c~~ 9c shows that higher rate of increase in O_3 (dO_3/dt: 1.1–1.5 ppbv h^{-1}) and O_3 build-up (6–7 ppbv; figure 2) during June and September but lower values during July-August (0.7–0.8 ppbv h^{-1}; ~ 4 ppbv) despite availability of precursors (figure ~~7b~~9b). The average number of cloudy pixels during morning hours (8:30–14:00 IST) obtained from INSAT-3DS shows more frequent cloudy conditions during July-August, which suppress the photochemistry leading to weaker noontime build-up. In addition, rain events during July-August (figure S1) would have washed out the precursors, inhibiting the O_3 production.

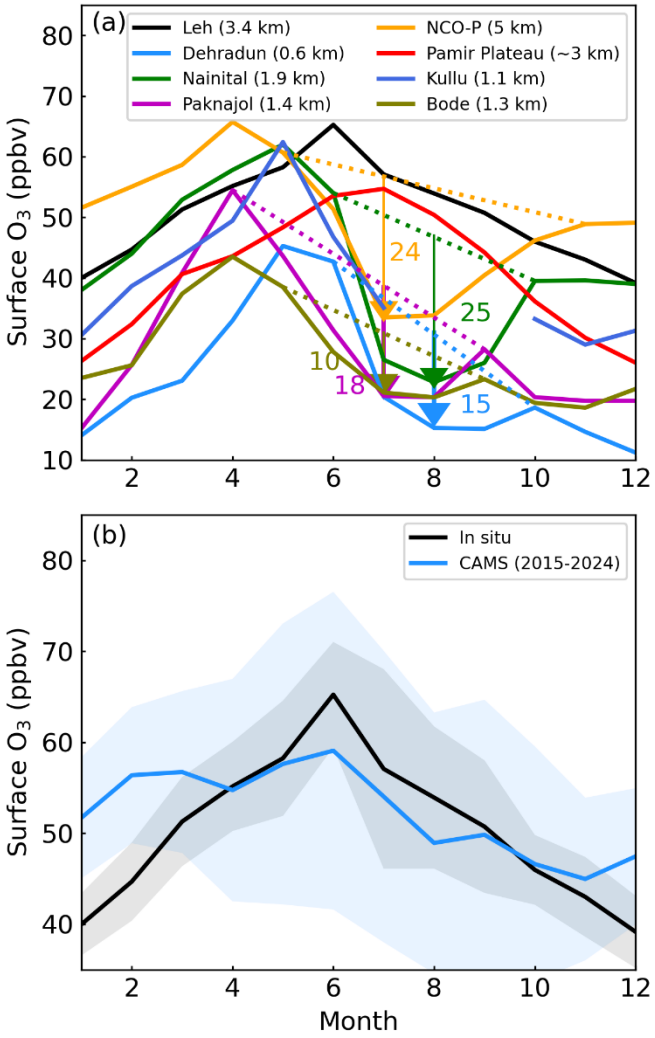


Figure ~~8~~10: (a) Comparison of O_3 seasonal cycle over Leh with observations in the Himalaya, its foothills and Pamir Plateau. The altitudes of the sites are shown in the legend. Numbers in different

colours show the approximate drop in O_3 concentration during monsoon with respect to the hypothetical reference curve (dotted line) connecting pre- and post-monsoon values. (b) Seasonal variation of O_3 from measurements and CAMS reanalysis over the study region. Stations shown in (a) are marked in figure 1a.

O_3 seasonal variation observed at Leh is compared with those at seven stations across Himalaya (Kullu, Dehradun, Nainital, Paknajol, Bode, NCO-P, and Pamir Plateau; marked in figure 1a), as shown in figure ~~8a~~10a. Monthly mean values are adopted from the literature (Cristofanelli et al., 2010; Harithasree et al., 2024; Mahata et al., 2018; Putero et al., 2015; Sarangi et al., 2014; Sharma et al., 2013; Yongde and Xinchun, 2025). O_3 at Leh peaks in June, unlike other Himalayan sites where it peaks during pre-monsoon season (April-May) except over the Pamir Plateau (Yongde and Xinchun, 2025). Similar to the aforementioned regions, the pre-monsoon peak is ubiquitous in the Gangatic region or northern Indian subcontinent, owing to the regional pollution and intense solar radiation (Lal et al., 2017; Ojha et al., 2022). It is important to note that central Himalaya and its foothills are influenced by transport of precursors from the IGP during pre-monsoon, causing a peak in the seasonal pattern (Kumar et al., 2010; Sarangi et al., 2014). The O_3 level reduces by a large magnitude of 10–25 ppbv during summer monsoon, with reference to pre- and post-monsoon values (dotted line in figure ~~8a~~10a), over most of the sites. The dotted line is derived by connecting pre-monsoon to post-monsoon months. Pre-monsoon month is chosen just before sharper monsoonal decline (April, May or June) and post-monsoon month (September, October or November) is chosen as corresponding to the peak after the monsoonal trough. Th~~is~~e influence of the monsoon is absent or negligibly small over Leh as well as Pamir Plateau. In addition to the difference in the timing of the seasonal peak, the absence of strong monsoonal reduction is another key feature of the Trans-Himalaya.

Figure ~~8b~~ 10b shows the monthly mean variations of observed O_3 concentrations during 2024 ~~–2015~~ 2025 against the climatological mean O_3 (2015–2024) derived from CAMS reanalysis for the study region. The CAMS-derived seasonal climatology agrees well with the observations except during winter. We do not have further parameters from the reanalysis for a deeper analysis of various contributing factors. To find the contributing factors responsible for the observed O_3 variation, we analysed the results from the global chemistry-climate model as discussed in the following section.

3.6 Source attribution inferred from global model simulations

Figure ~~9~~ 11 shows the observed and CAM4-Chem simulated seasonal variation of O_3 concentrations over the Trans-Himalaya. The model broadly captures the seasonal pattern (r=0.94) with a positive bias of 1–15 ppbv. The lowest bias is seen during the seasonal peak in June. Winter time overestimation (here, 10–16 ppbv) is commonly seen in most of the global models (Ansari et al., 2025;

Gao et al., 2025 and references therein), including CAMS (figure ~~8b~~10b). A winter-time positive bias has also been reported for other regions in previous studies (Liu et al., 2022; Miyazaki et al., 2025). This high bias could be attributed to factors such as the inappropriate representation of temperature-dependent chemical and physical processes, and photolysis rates impacted by clouds. Unresolved topography could also contribute to bias as air masses are trapped in the shallow boundary layer during winter and the associated dynamics may not be captured at the coarse resolutions of the model. A recent study (Wild and Ryan, 2026) used an ensemble of atmospheric model runs with perturbed process parameters and found that the photochemical kinetic parameters and deposition rates, particularly over water and unvegetated surfaces (which largely define our study region), have been the largest source of uncertainty in simulating surface O_3 (Emmerichs et al., 2021). More detailed investigations, especially model intercomparison exercises with perturbed process parameters, and not only with perturbed emissions, are required to pinpoint the causes of model biases (Carslaw et al., 2025).

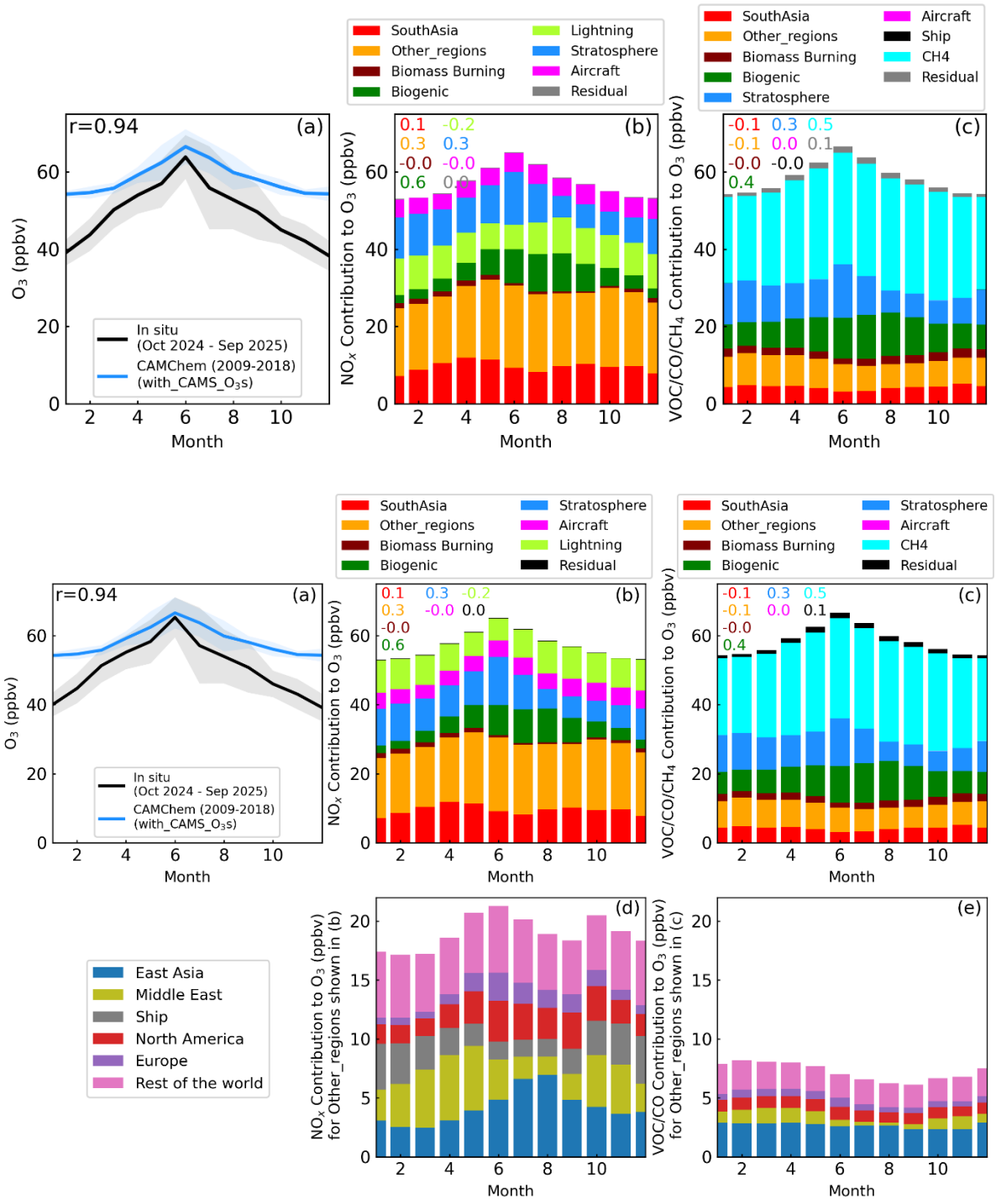


Figure ~~9~~11: (a) Seasonal variations of surface O_3 from in situ measurements and from CAM4-Chem derived climatology (2009–2018) with stratospheric contribution updated by CAMS O_3s, and CAM4-Chem derived contributions to O_3 from (b) NO_x sources and (c) VOC/CO/CH_4 sources. Covariances of different sources normalised with variance in modelled O_3 are marked with respective colours as in the legends in (b) and (c). The shaded region in (a) shows the standard deviations. Note that stratosphere in (b) and (c) refer to stratospheric O_3, rather than contribution from stratospheric NO_x or other precursors. Figure (d) and (e) provides further splits for different regions outside South Asia

shown in figure (b) and (c). The definition of geographic regions is shown in figure S4. Here, "Rest of the world" in (d) and (e) includes contributions from all the regions (figure S4), other than those shown in legends.

O_3 contributions are highly distributed across diverse NO_x sources: global anthropogenic, South Asian, biomass burning, biogenic, lightning, and stratospheric O_3, highlighting the background characteristic of this region. Interestingly, based on the NO_x tagging (figure ~~9b~~11b), the largest contribution to O_3 in CAM4-Chem model is seen from the anthropogenic sources outside the South Asia (figure S4), accounting for 30.9–36.6 % (17.1–21.3 ppbv). Other major contributions were from anthropogenic NO_x emitted in South Asia (12.9–20.1 %; 7.1–11.9 ppbv), lightning (9.5–17.3 %; 6.3–9.4 ppb), and stratospheric intrusion (9.4–20.7 %; 5.6–13.8 ppbv). Smaller contributors were global aircraft NO_x emission (7.1–9.8 %; 3.9–5.3 ppbv), biogenic emissions from vegetation (3.9–16.3 %; 2.1–9.8 ppbv), and biomass burning (0.8–2.6 %; 0.5–1.4 ppbv). Note that the O_3 contribution from global shipping NO_x is included here in the "other regions" segment. The residual is estimated to be negligible (0.14–0.19 ppbv). The NO_x contribution from the regions outside South Asia (17.1–21.3 ppbv, i.e., orange segment in figure 11b), is further bifurcated into different regions (figure 11d), as follows: East Asia: 4–12% (2–7 ppbv), Middle East: 3–10% (2–6 ppbv), Ocean: 2–8% (1–4 ppbv), North America: 3–5% (2–3 ppbv), Europe: 1–4% (1–2 ppbv), and Rest of the world: 8–11 % (5–6 ppbv). This suggests that NO_x emission from East Asia and Middle East have greater influences over the Trans-Himalaya. The seasonal pattern of contributions from East Asia, Europe, and North America are similar, i.e., lower during winter and higher during summer. Whereas for Middle East and international shipping, contribution is lower in summer and higher in spring/winter. It is interesting to note that the contribution from East Asia during July-August (7–8 ppbv) is comparable to those from South Asia (8–10 ppbv), and East Asian contribution peaks when South Asian contribution is lower during July-August.

Based on VOC/CO/CH_4 tagging (figure ~~9c~~11c), largest contribution of 40.1–50.3 % (21.9–21.2 ppbv) is from oxidation of CH_4, followed by biogenic (11.3–19.0 %; 6.2–11.4 ppbv), anthropogenic NMVOC emissions from foreign regions outside South Asia (10.3–14.9 %; 6.0–8.1 ppbv), stratospheric intrusion (9.4–20.7 %; 5.6–13.8 ppbv). Minor contributors were South Asian anthropogenic NMVOC emissions (4.8–9.5 %; 3.1–5.1 ppbv), and biomass burning (2.3–4.4 %; 1.6–2.4). The contribution of VOC emissions from aircrafts and ships is negligible (<0.03 ppbv) and the residual is small, 1.4–2.5 % (0.8–1.5 ppbv). Contribution from outside South Asia is 10–15% (6–8 ppbv) and further bifurcation of different regional sources (figure 11e) shows that East Asia (4–5%; 2–3 ppbv), Middle East (0–2%; 0–1 ppbv), and North America (1–2%; ~1 ppb) have smaller contributions. This suggests that much of the foreign-transported O_3 is due to foreign anthropogenic NO_x but not due to foreign anthropogenic

VOC+CO (figure 11d and e). Season-wise contributions from various factors are shown as pie charts in figure S10. Among these different sources, the O_3 contributions from CH_4 oxidation, from biogenic emissions, stratospheric contribution, and from anthropogenic NO_x emissions outside South Asia show larger variations (covariance normalised with variance of O_3: 0.3–0.6; see coloured text in figure ~~10b~~11b–c), and hence these four contributors largely shape the observed seasonal O_3 pattern. While stratospheric intrusion in June, 20.7 % (13.8 ppbv), largely drives the peak in O_3, biogenic emissions help in maintaining O_3 levels during the monsoon season. In other words, global anthropogenic influences plus contributions of natural sources (biogenic + stratospheric intrusion) in conjunction with distinct meteorology of the region sustain substantial O_3 even during the monsoon, a quite unique feature in South Asia (figure ~~8a~~10a). Since the model shows larger bias during winter, interpretations of source attributions need~~s~~ to be made cautiously.

3.7 Long-term trend in O_3 and its precursors

We have analysed CAM4-Chem long-term simulations to know how O_3 and its sources have changed over the study region for 1.5 decades, during 2003–2018. Figure ~~10~~ 12 shows a time series of de-seasonalised O_3 and contributions to O_3 from most of the sources as discussed in the previous section. The minor contributors to the study regions (i.e., biomass burning and VOCs from ships) are not shown here. Here, de-seasonalisation is carried out by removing the mean seasonal cycle from monthly average values. Linear regression is shown where the trend is statistically significant ($p<0.01$). O_3 increased over the study region until 2013 with a rate of 0.6 ppbv y^{-1} and slightly declined afterwards and the trend is statistically insignificant for the period of 2014–2018 (figure ~~10a~~12a). The increasing trend during 2003–2013 is attributed to increasing contributions from NO_x from South Asia (0.19 ppbv y^{-1}) and outside South Asia (0.23 ppbv y^{-1}), VOC+CO from South Asia (0.11 ppbv y^{-1}), CH_4 (0.17 ppbv y^{-1}) and stratospheric O_3 (0.31 ppbv y^{-1}). These increasing contributions were partially offset by a small declining contribution from biogenic NO_x (-0.08 ppbv y^{-1}) during 2003–2013. The increase and decrease in NO_x contributions from the regions outside South Asia before and after 2013 are due to increasing NO_x emission in the northern hemisphere until 2013 and decline afterwards (Ansari et al., 2025; Lu et al., 2025). Interestingly, stratospheric O_3 also showed a decline during 2014–2018, which is unexpected and needs separate investigations. Thus, slight decline or stationary levels of O_3 during 2014–2018 is due to counterbalance between increasing trend in NO_x and VOC+CO over South Asia and aircraft NO_x emissions; and decreasing trend in NO_x emissions outside South Asia and stratospheric contributions.

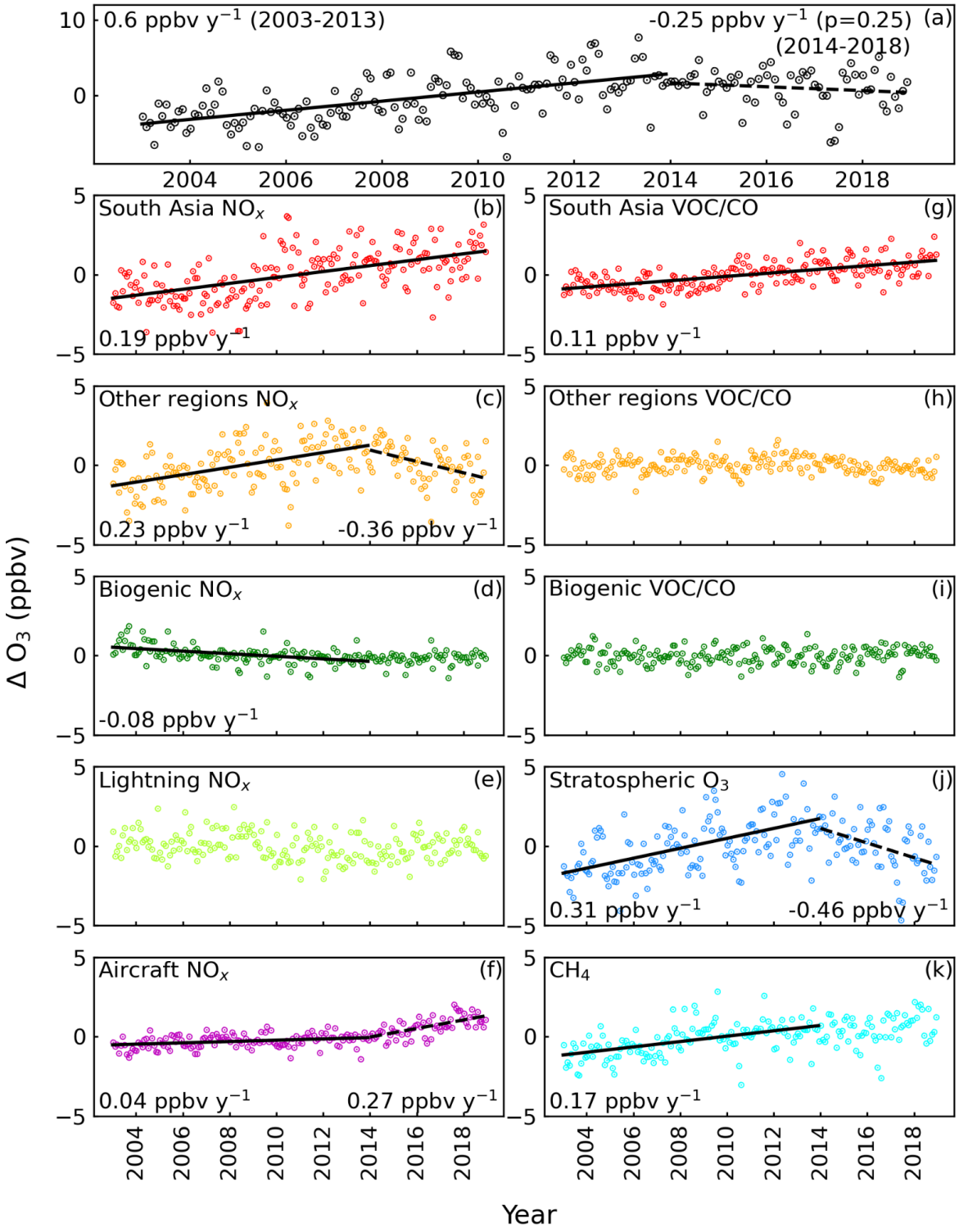


Figure ~~10~~12: Time series of CAM4-Chem simulated de-seasonalised O_3 (ΔO_3) and contributions from its sources NO_x and VOC/CO/CH_4 as labelled in each panel during 2003–2018. Black lines show linear regressions and corresponding trend values are shown in respective panels for 2003–2013 (continuous line) and 2014–2018 (dashed line). All trends are statistically significant (p-value <0.01),

except for a trend in O_3 during 2014–2018 shown in panel (a). Colours of the scatters are chosen the same as in the figure 9b and c.

In view of the significantly higher values of O_3, nearing air quality thresholds (section 3.3), we have also analyzed the long-term trend in surface O_3 (from CAMS reanalysis) and precursors (from TROPOMI considering region: 77.41°–77.70° E; 34.08°–34.23° N) during the past few years, as shown in figure ~~11~~13. CAMS O_3 shows a positive trend of 0.69 ppbv yr^{-1} (1.4 % y^{-1} with respect to an annual mean level of 50.6 ppbv during 2015). This trend value is larger than previously reported trend over the Indian region at a similar vertical level (0.1 ppbv y^{-1} during 1994-2019; Gaudel et al., 2024). The magnitude of the trend over the study region is also higher as compared to the observations of surface O_3, 0.1–0.4 ppbv y^{-1} (1973–2014) at Thiruvananthapuram (8.5° N; 77.9° E, Nair et al., 2018). Trend at 600 hPa over three locations in India (Thiruvananthapuram, Delhi and Pune) ranged -0.1 to 2.3 % y^{-1} during 1989–2001 (Saraf and Beig, 2004). If only June values of CAMS O_3 are considered where it closely matches with observations, an increasing trend of 0.74 ppbv y^{-1} is found, which is however statistically insignificant during 2015–2024.

Satellite-derived NO_2 and HCHO, show statistically significant trends of 0.45 × 10^{14} molecules cm^{-2} y^{-1} (~7 % y^{-1} with respect to annual mean 6.44× 10^{15} molecules cm^{-2} for year 2019) and 0.15 × 10^{15} molecules cm^{-2} y^{-1} (~4 % y^{-1} with respect to annual mean 6.44× 10^{15} molecules cm^{-2} for year 2019), respectively, during 2018–2025 (figure ~~11b~~ 13b and c). However, since the rate of increase in NO_2 is much higher than HCHO, the HCHO/NO_2 ratio is declining at the rate of 0.03 y^{-1} (figure S11), leading to a reduction from 0.65 in 2019 to 0.45 in 2025. This decline is pushing the O_3 formation regime towards the VOC-limited regime. If this scenario or trend continues further, local production of O_3 is expected to reduce over Leh according to figure ~~6a~~8a, however, regional contribution through transport would decide the net increase/decrease scenario. The observed increase in NO_2 could be partly attributed to the recent increase in anthropogenic ~~the tourism~~ activities including the tourism ~~in~~ over the region. Number of tourists in Ladakh have increased over the recent years: 0.28, 0.33, 0.28, 0.52, 0.38, and 0.33 million in 2017, 2018, 2019, 2023, 2024, and 2025, respectively (https://www.ijmra.us/project%20doc/2020/IJRSS_DECEMBER2020/IJRSS12Dec20-21086.pdf, https://www.ibef.org/states/ladakh-presentation, https://lokniwas.ladakh.gov.in/press-details/369). Satellite (Resourcesat-2 LISS-IV) based estimation of build-up area over Leh town (77.55–77.63° E; 34.08–34.20° N) shows increase from 1249.9 h (12.5 km^2) in 2016 to 1755.3 h (17.5 km^2) in 2025 (https://vedas.sac.gov.in/vapps/lama/main.html?app=urban; https://vedas.sac.gov.in/vapps/lama/assets/LAMA-interim-report.pdf#page=16). Therefore, local anthropogenic emissions would have contributed to the observed increasing trend in precursors (figure 13). However, CO does not show consistent increase or decline over the study region (figure

11d), and since it has a longer lifetime of a week to 2 months, the regional variation of CO needs to be investigated further.

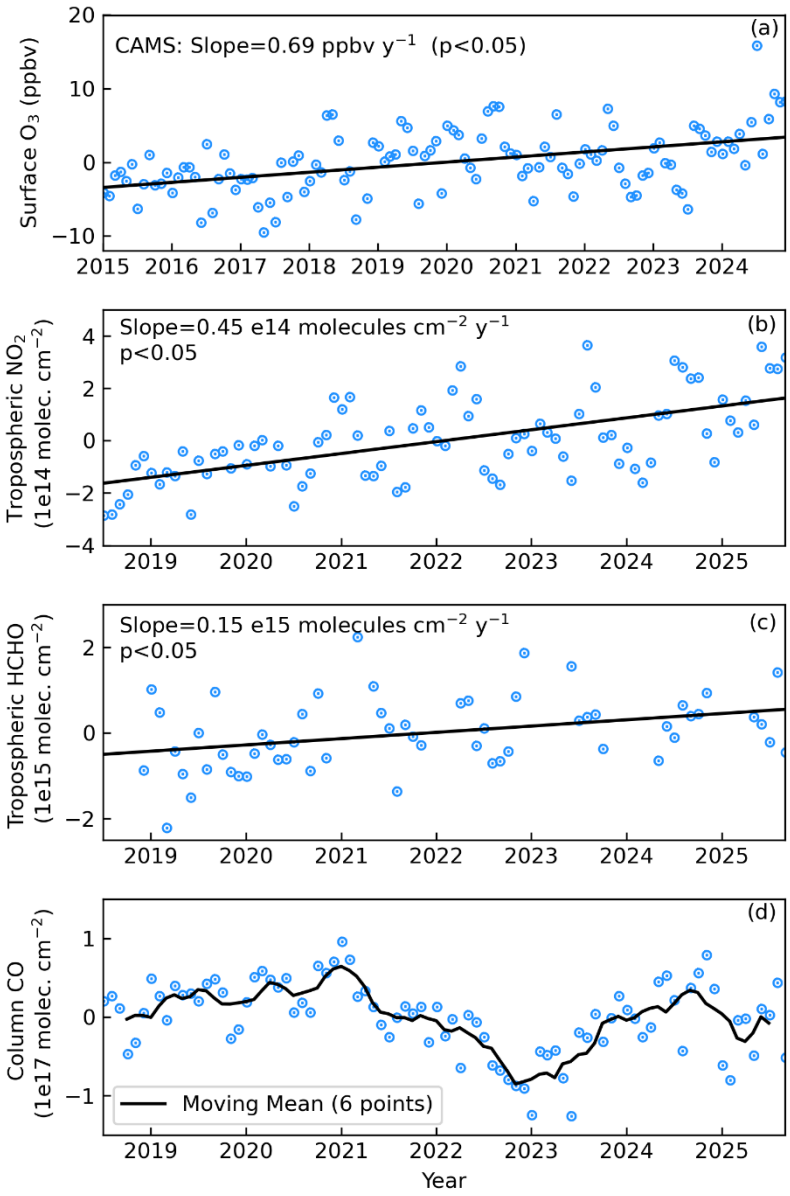


Figure ~~11~~13: Long-term trend in de-seasonalised monthly mean surface O_3 from reanalyses (2015–2025) and tropospheric NO_2, HCHO and column CO from TROPOMI (2018–2025) over the study region. Trend magnitude, i.e., slope of linear fit, is marked in the top-left corner along with p-value.

4. Summary and conclusions

The dynamics of surface O_3 over the Trans-Himalayan region has been comprehensively investigated ~~for the first time~~ by combining year-long in situ measurements, analysis of satellite observations, and results from global and regional models. The key findings are summarised as follows.

- Trans-Himalaya is observed to be a unique environment exhibiting a daytime O_3 build-up of 2.6–6.8 ppbv (with rate of 0.6–1.5 ppbv h^{-1}), a feature not observed at high-altitude mountain regions. Despite a daytime build-up partially offsetting the nighttime losses, the region remains a net chemical sink. While photochemistry and vertical entrainment are major drivers of diurnal pattern during winter, advection is an additional contributor during monsoon season.

- WRF-Chem-derived tendencies suggest that advection is largely responsible for day-to-day changes in O_3 during winter, whereas variabilities during monsoon arise from multiple competing processes rather than a single dominant factor.

- The O_3 seasonal cycle peaked in June (65.2 ± 5.8 ppbv) unlike the ubiquitous pre-monsoon peak (April-May) observed over the Himalaya and Gangetic regions. The monsoonal dip (July to September) is much smaller compared to those observed over the Himalaya region. MDA8 is close to and occasionally crosses the WHO threshold of 100 µg m^{-3} during May-Julyne along with higher values of vegetation-related O_3 exposure indices. The photochemistry, fueled by the higher levels of precursors, falling in the NO_x-VOC transition regime leads to higher O_3 values during May-JuneJuly.

- Surface O_3 varied from 39–65 ppbv with seasonal amplitude of 26 ppbv, being lower during winter. The CAM4-Chem O_3 climatology, after revising stratospheric contribution from CAMS, agrees with the observed seasonal pattern, however, model overestimated wintertime levels by 10–15 ppbv. Nearly half of the O_3 in Trans-Himalaya is contributed by anthropogenic NO_x emissions from South Asia (15–18 %) and outside South Asia (32–37 %) throughout the year. From a VOC/CO/CH_4 perspective, nearly half of O_3 is driven by oxidation of CH_4. Among various sources, the following four major contributors largely shape the observed seasonal pattern over the Trans-Himalaya: CH_4 oxidation, biogenic emissions of NO_x and VOCs, stratospheric O_3, and anthropogenic NO_x emission from regions outside South Asia. The contribution of NO_x emission (17–21 ppbv) is greater than VOC/CO contributions (6–8 ppbv) from the regions outside South Asia. Besides these regional and global photochemical contributions, seasonal peak in O_3 is significantly influenced by stratospheric inputs.

- The study can serve as a reference for designing future field experiments and improvement of global and regional models over the Himalayan region. In particular, measurements of NO_x, CH_4 and other VOCs, studying the dynamics enhancing O_3 in the evening hours, and reducing wintertime biases in the model could be crucial. This study also represents the need for continuation of O_3 measurements for air quality monitoring as its levels are close to the threshold.

Overall, the study highlights the varying roles of chemistry, dynamics and regional-to-global anthropogenic sources in controlling the O_3 magnitude and variations over the Trans-Himalaya. The rising levels of O_3 and its precursors—with surface O_3 at 1.4 % y^{-1}, NO_2 at ~7 % y^{-1} and HCHO at ~4 % y^{-1}—highlight the impact of rapid urbanization and/or tourism, and underscore the need of monitoring of O_3 and its precursors along with high-resolution chemistry-transport modeling over this climate-sensitive region of Trans-Himalaya. Comprehensive measurements of atmospheric composition along with meteorological parameters ~~and continuous O_3 monitoring~~ in this data-sparse region ~~can provide~~would be crucial for validating satellite-based insights and improving models. This is important as O_3 impacts are found substantial in this region although the~~into unexplored atmospheric chemistry and transport processes, which are hard to obtain through other~~ public health-centric O_3 monitoring has been~~observations~~ focused mostly ~~only~~ in urban and rural areas ~~centres~~ across India. Observed higher values of O_3 are sensitive to NO_x as well as VOCs, hence monitoring of various NO_x and VOCs besides other precursors is essential for better quantification of the gradient between local emissions and background levels. With new observation-based insights and evaluation of models, future field studies can be designed more strategically and improve the model simulations to further advance the understanding of atmospheric chemistry and dynamics of the Trans-Himalayan region.

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Author contributions: IAG and MP conceptualised and designed the study. IAG and SA performed the measurements, with support from AV, LKS, BPS and SS. TA, TB, and NO contributed to the model results and related inputs. IAG led the study and prepared the first draft of the manuscript. All co-authors contributed to the discussions, reviews, and editing.

Data availability

In situ measured surface O_3 (monthly means) is provided in supplementary (Table ~~S1~~S2). Hourly interval O_3 (measured and WRF-Chem modelled) and monthly O_3 (CAM4-Chem) are available at https://doi.org/10.5281/zenodo.23916482~~will be made available on Zenodo at the acceptance of the paper~~.

Declaration of competing interest

One of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

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