



1 **Hydrochemistry and modeling nitrate concentration in farmland**
2 **groundwater under different hydrological seasons by integrating**
3 **hybrid quantum-classical ML, virtual sample generation and**
4 **AlphaEarth Foundation**

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19

20 **Abstract**

21 Precise seasonal prediction of groundwater nitrate concentrations in intensive agricultural
22 areas faces challenges such as data sparsity, strong spatiotemporal heterogeneity, and complex
23 hydro-biogeochemical processes. To address these issues, this study proposes an integrated
24 prediction framework combining hybrid quantum-classical machine learning, advanced virtual
25 sample generation (t-SNE-GMM-KNN), and remote sensing foundation model semantic
26 embedding (AEF). Modeling was conducted across the 2022-2023 normal, dry, and wet seasons in
27 Xiong'an New Area. Hydrochemical types were dominated by Ca-Mg-HCO₃⁻, controlled by
28 mineral dissolution and evaporation. Nitrate concentrations were highest in the dry season (mean



29 42.93 mg L⁻¹), driven by evaporative concentration. Spatially, high-value zones shifted: southeast
30 (normal), central (dry), and northwest (wet). MixSIAR modeling based on isotopes indicated
31 domestic sewage and livestock manure (74.1%) as dominant sources. The t-SNE-GMM-KNN
32 strategy mitigated small-sample bias while preserving nonlinear structure. When virtual samples
33 were augmented to 10-fold, the Random Forest R² in the dry season increased from 0.284 to >0.85.
34 Furthermore, a hybrid quantum-classical Random Forest exhibited superior robustness for data
35 sparsity, achieving peak performance in the normal season (R²=0.962, RMSE=5.73 mg L⁻¹).
36 Additionally, using only AEF embeddings achieved screening-level accuracy (R² up to 0.860),
37 providing a feasible rapid survey scheme for extensive unmonitored regions. Correlation analysis
38 identified TDS and EC as persistent top predictors (r>0.8). This comprehensive framework offers
39 a robust solution for seasonal nitrate prediction and sustainable water management.

40 **Keywords:** Groundwater nitrate concentration; Hydrological seasons; Virtual sample generation;
41 Hybrid quantum-classical machine learning; AlphaEarth Foundation (AEF) embeddings; Nitrate
42 source apportionment.

43

44 1. Introduction

45 Nitrate (NO₃⁻) contamination in groundwater poses a serious threat to drinking water safety
46 and ecosystem health, particularly in intensively managed agricultural regions (Wang et al., 2021).
47 In China, groundwater nitrate pollution is a growing concern, national monitoring data from 2013
48 to 2017 revealed a nitrate exceedance rate exceeding 10%, with Hebei Province reporting an
49 alarming rate of 31.66% in 2017 (Li et al., 2019). Over recent decades, escalating nitrate
50 concentrations in surface and groundwater have been driven by intensified fertilizer use in
51 agriculture, along with discharges of industrial and domestic wastewater (Zhang et al., 2018).
52 Severe nitrate exceedances are especially prevalent in northern and northwestern China (Gu et al.,
53 2013), where key contributors include domestic and industrial effluents, nitrification of soil
54 organic nitrogen, and synthetic fertilizer application (Han et al., 2016). For instance, in the North
55 China Plain, shallow groundwater nitrate exceedance rates range from 9.5% to 34.1%, and a rising
56 trend persists at the regional scale, particularly in agricultural areas (Wang et al., 2018). In
57 monsoonal temperate regions, seasonal shifts in precipitation, evapotranspiration, and
58 groundwater recharge profoundly influence the transport, dilution, and accumulation of nitrate,



59 leading to pronounced intra-annual variability in its concentration and spatial distribution (Gao et
60 al., 2023; Zhu et al., 2025). Consequently, understanding and forecasting nitrate dynamics across
61 hydrological seasons is essential for informed groundwater management and pollution mitigation,
62 but remains a formidable challenge due to the nonlinearity, high dimensionality, and data scarcity
63 inherent in such systems (Deng et al., 2023).

64 Traditional monitoring and modeling approaches face three critical limitations. First, field
65 sampling campaigns though providing high-fidelity hydrochemical data are inherently sparse in
66 space and time, especially for large-scale or rapidly changing environments (Viswanathan et al.,
67 2022), which are time-consuming, labor-intensive, and costly, limiting the spatial and temporal
68 coverage of data (Cai et al., 2025). Second, while process-based models incorporate physical
69 mechanisms, they require extensive parameterization and are computationally prohibitive for
70 dynamic, multi-season forecasting at farm-to-regional scales (Feng et al., 2022). Hydrological
71 seasonal variations (normal, dry, and wet seasons) significantly influence the migration and
72 transformation of nitrogen in the soil-groundwater system (Chen et al., 2025). For instance,
73 concentrated rainfall during the wet season (accounting for 60%-80% of annual precipitation) can
74 promote the leaching of surface nitrogen into groundwater, leading to a 25-fold increase in stream
75 nitrate concentrations during storm events compared to baseflow (Sebestyen et al., 2014),
76 meanwhile, intense evaporation in the dry season leads to the accumulation of nitrate in shallow
77 aquifers, where concentrations can exceed the US EPA drinking water standard of 10 mg L⁻¹ by
78 2-3 times (Liu et al., 2025; Cox et al., 2016). These seasonal differences result in distinct
79 hydrochemical characteristics and nitrate concentration distributions, increasing the complexity of
80 prediction models (Wu et al., 2025). Third, even advanced machine learning (ML) techniques
81 such as Random Forest (RF), despite their robustness to nonlinearity and multicollinearity, still
82 rely heavily on sufficient representative samples to capture the multi-modal distribution and tail
83 behavior of environmental variables, particularly for heavy-tailed pollutants like NO₃⁻ (Luo et al.,
84 2022). Moreover, the small sample sizes obtained from discrete sampling often lead to data
85 sparsity and skewed distributions, reducing the model's generalization ability by 30%-50% when
86 applied to unmonitored areas and compromising the robustness and generalization ability of
87 machine learning (ML) models trained on such data (Thunyawatcharakul et al., 2025; Wang et al.,
88 2024).



89 To overcome these bottlenecks, recent efforts have explored virtual sample augmentation and
90 hybrid modeling. Gaussian Mixture Models (GMM) and deep generative frameworks (e.g., VAEs,
91 GANs) have shown promise in enriching training data, with GMM achieving an average similarity
92 of 83.0% between unmixed chemical spectra and ground truth in geochemical analysis (Farnia et
93 al., 2023; Tung et al., 2023), however, they often fail to preserve the non-linear manifold structure
94 of high-dimensional geochemical space or require large training sets, precisely what is lacking
95 (Zhou et al., 2025). Non-linear dimensionality reduction methods, such as t-SNE, excel at
96 revealing latent clusters corresponding to distinct hydrological processes, with a classification
97 accuracy of 92% for annual daily hydrograph clustering in mountainous watersheds, yet lack
98 explicit generative mechanisms (Wang et al., 2025; Tang et al., 2022). Meanwhile, the rise of
99 foundation models in Earth observation exemplified by Google's AlphaEarth Foundation (AEF),
100 offers unprecedented opportunities: its 64-dimensional semantic embeddings, derived from
101 multi-sensor satellite time series (including Sentinel-2, Landsat, and Sentinel-1), implicitly encode
102 land use, vegetation phenology, soil moisture, and anthropogenic footprints at 10 m resolution
103 (Tollefson, 2025). These features have been successfully applied in land use classification and
104 crop monitoring, but their potential for predicting groundwater nitrate concentrations, especially
105 across different hydrological seasons remains underexplored (Li et al., 2025). Quantum machine
106 learning (QML) further opens a new frontier. Parameterized Quantum Circuits (PQCs) can map
107 classical inputs into exponentially high-dimensional quantum Hilbert spaces, generating entangled
108 feature representations that reveal complex, non-linear patterns inaccessible to classical kernels
109 (Hong et al., 2025). For ozone concentration forecasting, a hybrid QML model achieved an R^2 of
110 94.12% for 1-hour forecasts and 75.62% for 6-hour forecasts, outperforming classical persistence
111 models by a forecast skill of 31.01-57.46% (Oliveira et al., 2025). Crucially, analytical quantum
112 feature extraction via Pauli-Z expectation values avoids the noisy sampling overhead of near-term
113 quantum hardware, reducing computational latency by ~80% compared to sampling-based
114 methods and making it viable for small-sample environmental modeling (Gujju et al., 2024;
115 Oliveira et al., 2025).

116 Furthermore, identifying the sources and controlling factors of nitrate pollution is crucial for
117 improving prediction accuracy and guiding targeted pollution control measures. Isotopic analysis
118 ($\delta^{15}\text{N-NO}_3^-$ and $\delta^{18}\text{O-NO}_3^-$) combined with the MixSIAR model has proven effective in



119 quantitatively apportioning nitrate sources (Tian et al., 2025). Meanwhile, Bayesian models and
120 SHapley Additive exPlanations (SHAP) analysis can reveal the key environmental variables
121 driving nitrate concentration changes, enhancing the interpretability of prediction models (Alam et
122 al., 2025). Despite these advancements, several gaps persist in the current research: (1) Few
123 studies have integrated hybrid quantum-classical ML with virtual sample augmentation to address
124 small-sample challenges in seasonal nitrate prediction; (2) The potential of AEF remote sensing
125 semantic features for groundwater nitrate prediction remains untested, particularly in comparison
126 with in-situ measured parameters; (3) The combined effects of hydrological seasonal variations,
127 nitrate source apportionment, and key environmental drivers on prediction model performance
128 require systematic investigation.

129 The North China Plain, an important agricultural production region in China, is characterized
130 by high nitrogen input intensity and significant seasonal hydrological variations, making it an area
131 prone to groundwater nitrate pollution (Liu et al., 2025). Conducting field-scale research on nitrate
132 pollution in this region is of great significance for the protection of regional water resources. The
133 Xiong'an New Area in China is a typical study area at the farmland scale in the North China Plain.
134 As a major agricultural area with high nitrogen input intensities and distinct seasonal hydrological
135 cycles, it faces significant groundwater nitrate pollution risks. This region's unique climatic
136 regime characterized by a dry spring, wet summer with concentrated precipitation, and a cold, dry
137 winter, creates marked seasonal disparities in groundwater recharge, evaporation, and pollutant
138 migration (Xu et al., 2022). A mechanistic understanding of how nitrate concentrations vary
139 across these hydrological seasons (normal, dry, wet) and their controlling factors is crucial for
140 regional water resource management.

141 To fill these gaps, this study aims to: (1) propose a novel virtual sample generation method
142 (t-SNE-GMM-KNN) to enhance small-sample datasets while preserving the non-linear structure
143 and multi-modal distribution of original data; (2) construct a hybrid quantum-classical random
144 forest model by integrating quantum feature encoding with classical random forest, improving the
145 model's ability to capture complex environmental relationships; (3) evaluate the predictive
146 performance of two input datasets (on-site measured water quality parameters and AEF remote
147 sensing semantic features) across normal, dry, and wet seasons under Leave-One-Out
148 Cross-Validation (LOOCV); (4) identify the dominant nitrate sources and key environmental



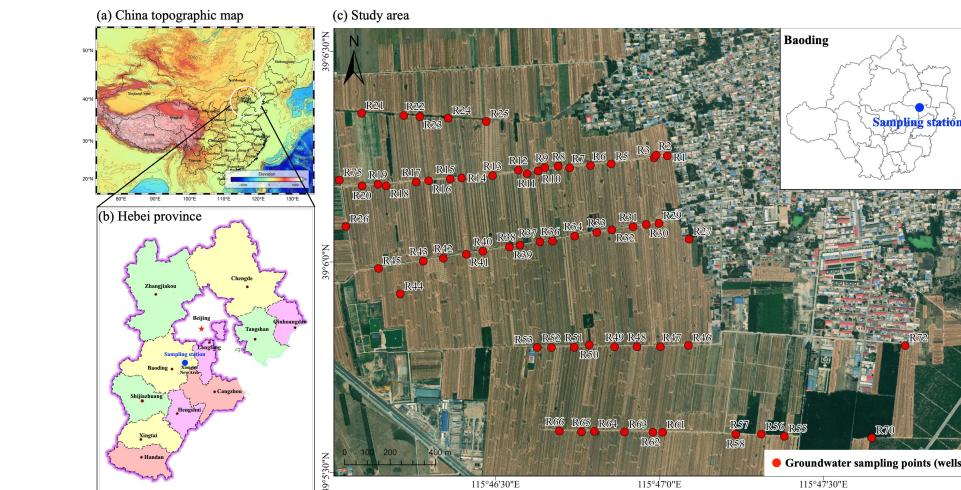
149 controlling factors using isotopic analysis, MixSIAR modeling, Bayesian analysis, and SHAP
150 interpretation; (5) establish a comprehensive and accurate prediction framework for groundwater
151 nitrate concentrations in intensive agricultural regions, providing scientific support for
152 groundwater pollution control and sustainable water resource management in the North China
153 Plain. The novelty of this study lies in the integration of hybrid quantum-classical machine
154 learning, advanced virtual sample augmentation, and remote sensing semantic features to address
155 the challenges of small-sample, high-dimensional, and seasonally variable nitrate prediction. The
156 findings are expected to advance the state-of-the-art in groundwater quality prediction and offer a
157 scalable approach for large-scale environmental monitoring in unmonitored areas.

158

159 2. Materials and Methods

160 2.1 Study area

161 The North China Plain is one of China's most important agricultural production bases. This
162 study focuses on the Xiong'an New Area, situated in the central part of Hebei Province, as a
163 representative research site within this plain. Located in the core region defined by Beijing,
164 Tianjin, and Baoding, it boasts an advantageous geographical position, with straight-line distances
165 of 105 km to both Beijing and Tianjin, and 30 km to Baoding. Its geographical coordinates range
166 from 38°43' to 39°10' N latitude and from 115°38' to 116°20' E longitude, covering an area of
167 approximately 1770 km² ([Xiong'an New Area Official Website, 2023](#)). The specific study area is
168 an unmanned farm located in Xieyeqiao Village, Nanzhang Town, Rongcheng County, within the
169 Xiong'an New Area ([Fig.1](#)). The farm covers an area of 3000 hectares and primarily cultivates two
170 main grain crops: wheat and corn. As the first mechanized unmanned farm in Xiong'an, it has
171 achieved full mechanization and intellectualization, enabling unmanned, precise, and standardized
172 operations throughout all stages of tillage, sowing, management, and harvesting.



173

174 **Fig.1.** Study area map showing the sampling location. ((b) Based on the standard map (Approval
175 No. Ji S (2025) 009) from the Department of Natural Resources of Hebei Province; base map is
176 unmodified.)

177 Cultivated land constitutes a large proportion of the total area in the Xiong'an New Area and
178 is predominantly dryland. Traditional fertilization in the region involves high application rates of
179 nitrogen and manure. As a representative farm within this area, the study site also follows this
180 conventional practice, making it susceptible to the impacts of high fertilization intensity. The
181 annual nitrogen fertilizer application rate at the study site ranges from 540 to 660 kg (N) $\text{ha}^{-1} \text{ yr}^{-1}$,
182 primarily supplied as urea (46% N). The extensive application of chemical fertilizers and manure
183 consequently elevates the risk of nitrogen pollution in groundwater. Furthermore, the rural
184 population is relatively densely distributed, contributing to pollution from domestic sewage
185 discharge in the vicinity. The climate is classified as a warm-temperate, monsoonal, continental
186 semi-humid climate. Springs are dry and rainless, summers are humid with abundant precipitation,
187 autumns are cool and dry, and winters are cold with minimal snowfall. The mean annual air
188 temperature in Xiong'an New Area is 12.6°C, exhibiting relatively minor inter-annual fluctuations.
189 The mean annual precipitation is 480.8 mm, which is highly concentrated from June to September.
190 The average annual sunshine duration is 2335.2 hours, with longer periods in spring and summer
191 and shorter ones in autumn and winter. The average frost-free period lasts 204 days. The mean
192 annual wind speed is 1.7 m s^{-1} , with the highest average occurring in April and the lowest in
193 January, August, and December. The multi-year average evaporation is 1661.1 mm (Liao et al.,



194 2020). The soil texture is dominated by silty loam, and the 2-8.5 m soil layer contains interlayers
195 with high clay content such as clay and silty clay, reflecting the characteristics of vadose zone
196 sediments in the central plain under geomorphic sedimentation. Nitrogen in the thick vadose zone
197 is dominated by organic nitrogen, accounting for approximately 97% of the total nitrogen content.
198 The shallow vadose zone at 3-6 m stores the largest amount of nitrate, accounting for about half of
199 the total nitrate reserves in the North China Plain (Li et al., 2025; Zhang et al., 2007).
200 Groundwater in the study area is primarily hosted in Quaternary unconsolidated porous aquifers,
201 with sampled wells ranging from 70 to 120 m in depth (Bai et al., 2023). The primary source of
202 groundwater recharge in the study area's farmland is atmospheric precipitation, while the main
203 discharge pathway is artificial extraction for agricultural irrigation. Irrigation followed crop
204 phenological stages. Wheat underwent muddy water irrigation at pre-sowing, overwintering,
205 regreening, and jointing stages, and maize received a single post-sowing muddy water irrigation.
206

207 2.2 Data collection and measurements

208 2.2.1 Field sampling data and laboratory analysis

209 Field investigations and the collection of hydrochemical and isotopic samples were
210 conducted in the study area from 2022 to 2023. A total of 66, 65, and 50 groundwater samples
211 were collected in October 2022, April 2023, and August 2023, respectively. All groundwater
212 samples were obtained from existing agricultural irrigation wells within the study area. Prior to
213 sample collection, each well was purged by pumping. Sampling commenced only after the
214 pumped volume exceeded three times the well's casing volume and on-site parameters had
215 stabilized (i.e., showing minor fluctuations around a constant value rather than a continuous rising
216 or falling trend), a procedure implemented to ensure the representativeness of the samples. At each
217 sampling point, one 1000 mL and two 100 mL samples were collected. Before final collection, the
218 sample bottles were rinsed three times with the water to be sampled. Immediately after collection,
219 the samples were sealed and stored in a portable cooler for transport to the laboratory for
220 subsequent analysis. Furthermore, the precise geographical location of each sampling point was
221 recorded using a GPS device.

222 In-situ physicochemical parameters were measured using a Hach HQ400 multi-parameter
223 water quality meter (Li et al., 2022). The measured parameters included water temperature (T, °C),



224 pH, total dissolved solids (TDS, mg L⁻¹), dissolved oxygen (DO, mg L⁻¹), electrical conductivity
225 (EC, μ S cm⁻¹), and oxidation-reduction potential (ORP, mV). The concentration of HCO₃⁻ was
226 determined within 24 hours of sample collection using the dilute sulfuric acid-methyl orange
227 titration method (Huang et al., 2012). Prior to the determination of cations and anions, water
228 samples were filtered through 0.45 μ m membrane filters. Major cations (K⁺, Ca²⁺, Na⁺, Mg²⁺)
229 were analyzed using an inductively coupled plasma optical emission spectrometer (Avio 500).
230 Major anions (NO₃⁻, Cl⁻, SO₄²⁻) were analyzed using an ion chromatograph (ICS-2100). The
231 analytical precision for cations and anions was controlled within \pm 0.2 mg L⁻¹, and the charge
232 balance error was maintained within 5% to ensure reliability. The concentrations of nitrite
233 nitrogen and ammonia nitrogen were determined using a flow injection analyzer (Smartchem 200,
234 AMS Alliance) and measured using dual wavelength spectrophotometry and the indophenol blue
235 method (Kim et al., 2019; Sun et al., 2022). The limits of detection for nitrite nitrogen and
236 ammonium nitrogen were both 0.01 mg L⁻¹. For the analysis of stable hydrogen and oxygen
237 isotopes, water samples were filtered through 0.22 μ m membrane filters and measured using an
238 LGR liquid water isotope analyzer (TIWA-45-EP). The analytical precisions for δ ²H, δ ¹⁷O, and
239 δ ¹⁸O were \pm 0.15‰, \pm 0.02‰, and \pm 0.02‰, respectively (Hamidi et al., 2023). The isotopic
240 compositions of nitrate (δ ¹⁸O-NO₃⁻ and δ ¹⁵N-NO₃⁻) were determined using a MAT-253 mass
241 spectrometer coupled with an elemental analyzer (Li et al., 2022). To ensure analytical precision,
242 standard references, reagent blanks, and duplicate samples were employed. Furthermore,
243 international standards USGS 34 and USGS 35 were used for δ ¹⁸O quality control, while USGS
244 32 and USGS 34 were used for δ ¹⁵N quality control. All isotope results are reported in per mil
245 (δ , ‰).

246

247 2.2.2 Google AlphaEarth Foundation

248 To facilitate comparisons with predictions based on in-situ field sampling data and to validate
249 the accuracy of predicting groundwater nitrate concentration using remote sensing data, this study
250 incorporates the Google AlphaEarth Foundation (AEF) dataset. AEF is a collection of
251 high-dimensional surface semantic embedding features generated via pre-training on multi-source
252 remote sensing data (Brown et al., 2025). By fusing imagery from Sentinel-2, Landsat, and other
253 Earth observation satellites, this dataset constructs a 64-dimensional vector representation



254 (denoted as A00-A63) at a global scale with an annual temporal resolution and a 10 m spatial
255 resolution (Alvarez et al., 2025). These embeddings implicitly encode complex environmental
256 semantics, such as land cover types, vegetation dynamics, soil moisture, and the intensity of
257 human activity, and have been successfully applied in tasks including land use classification, crop
258 monitoring, and environmental risk modeling (Tollefson et al., 2025).

259 The primary processing workflow involved spatially sampling the 64-dimensional AEF
260 vectors at a 10 m resolution using the GOOGLE/SATELLITE_EMBEDDING/V1/ANNUAL
261 product on the Google Earth Engine (GEE) platform, based on the geographic coordinates of the
262 field sampling points. To ensure data quality, only samples exhibiting exact matches between the
263 GEE extraction and the actual field data points were retained. Given the redundancy within the
264 initial 64-dimensional AEF features, Principal Component Analysis (PCA) based on Singular
265 Value Decomposition (SVD) was employed for feature compression. Specifically, SVD was
266 performed on the centered feature matrix to select the minimum number of principal components
267 accounting for at least 95% of the cumulative explained variance (Ilyas et al., 2025). The
268 orthogonalized, low-dimensional principal component scores were subsequently used as model
269 inputs. This approach preserves the vast majority of the semantic information from the original
270 embeddings while significantly mitigating the risk of overfitting. Ultimately, the PCA-reduced
271 AEF features served as the input variables for the model.

272

273 2.3 MixSIAR model and isotopic composition of nitrate sources

274 The MixSIAR model uses prior information such as the number of end - members, errors,
275 and distribution characteristics, and iterates based on the Markov Chain Monte Carlo (MCMC)
276 method to quantitatively restore the contribution fraction of each end-member to the mixed sample
277 (Stock et al., 2018). At present, this analytical method has been widely applied in the quantitative
278 analysis of nitrate pollution sources in water bodies. The calculation principle of the model is as
279 follows:

$$X_{ij} = \sum_{k=1}^K P_k (S_{jk} + C_{jk}) + \varepsilon_{ij} \quad (1)$$

$$S_{jk} \sim N(\mu_{jk}, \omega_{jk}^2) \quad (2)$$

$$C_{jk} \sim N(\lambda_{jk}, \tau_{jk}^2) \quad (3)$$

$$\varepsilon_{ij} \sim N(0, \sigma_j^2) \quad (4)$$



284 In the formula, X_{ij} is the value of isotope j in the i -th sample ($i = 1, 2, \dots, 20, j = 1, 2$); P_k is the
285 contribution rate of the k -th pollution source; S_{jk} is the value of isotope j in the k -th pollution
286 source, where μ is the mean and ω is the variance of the normal distribution; C_{jk} is the
287 fractionation coefficient, where λ is the mean and τ is the variance of the normal distribution; ε_{ij}
288 is the residual, with 0 as the mean and σ as the variance of the normal distribution.

289 In this study, the MixSIAR model is used to calculate the five potential sources of NO_3^- in
290 water bodies, namely precipitation (NP), soil organic nitrogen (SON), synthetic NH_4^+ fertilizer
291 (NHF), synthetic NO_3^- fertilizer (NOF), and domestic sewage & manure (DSM). The end-
292 member values of the five sources are selected as shown in [Table 1](#) (Mao et al., 2023; Gao et al.,
293 2023; Torres-Martínez et al., 2021).

294 **Table 1.** Summary statistics of $\delta^{18}\text{O}$ and $\delta^{15}\text{N}$ for potential nitrate sources.

Sources	$\delta^{18}\text{O-NO}_3^-$		$\delta^{15}\text{N-NO}_3^-$	
	Mean	SD	Mean	SD
NP	57.2	6.9	0.6	1.5
NHF	-4.1	2.7	-2.1	0.7
NOF	21.7	2.9	0.2	2.3
SON	-2.7	4.4	3.8	1.8
DSM	6.1	1.6	17.4	3.9

295
296 2.4 t-SNE-GMM-KNN: based on nonlinear structure modeling in feature space
297 To address the challenges of overfitting and poor generalization performance in small-sample
298 modeling, which arise from data sparsity and skewed distributions, this study proposes a
299 three-stage virtual sample generation strategy termed t-SNE-Gaussian Mixture Sampling with
300 KNN Inverse mapping. This method aims to preserve the non-linear manifold structure and
301 multi-modal distribution characteristics of the original high-dimensional feature space while
302 generating physically plausible and statistically consistent synthetic samples. The specific
303 workflow is as follows:

304 1. Data standardization

305 All input features are standardized using Z-score standardization to eliminate scale



306 differences and enhance the stability of the subsequent dimensionality reduction ([Jamshidi et al., 2022](#)).

308 2. t-SNE non-linear dimensionality reduction

309 t-Distributed Stochastic Neighbor Embedding (t-SNE) is employed to map the
310 high-dimensional feature space into a low-dimensional latent space (d=2) ([Islam et al., 2023](#)). To
311 balance the preservation of local and global structures, the perplexity is set to 10, and PCA
312 initialization is used to ensure reproducibility. t-SNE effectively reveals the clustered structure of
313 samples on the low-dimensional manifold, reflecting the differentiation of underlying
314 environmental processes within hydrological seasons ([Liu et al., 2021](#)).

315 3. GMM clustering and optimal component selection

316 In the t-SNE-reduced low-dimensional space, a Gaussian Mixture Model (GMM) is
317 constructed to characterize the probability density distribution of the data ([Jia et al., 2022](#)). The
318 GMM assumes that the data are generated from a linear combination of several Gaussian
319 distributions. The weights, means, and covariance matrices of each Gaussian component are
320 estimated via the Expectation-Maximization (EM) algorithm, thereby accurately capturing the
321 complex distribution patterns of the data ([Yan et al., 2023](#)). To avoid subjectively setting the
322 number of clusters, the Bayesian Information Criterion (BIC) is used to automatically optimize the
323 number of components, K, within the range ([Ghodba et al., 2025](#)):

$$324 \quad BIC(K) = -2 \log L + p_K \log n \quad (5)$$

325 where L is the model's likelihood, k is the total number of free parameters for a K-component
326 model, and n is the sample size. The value of K corresponding to the minimum BIC is selected as
327 the optimal number of components, ensuring a balance between goodness-of-fit and model
328 complexity.

329 4. Virtual sample generation and inverse mapping

330 Based on the optimal GMM, a specified number of virtual points are randomly sampled from
331 its joint probability distribution. This generation process naturally inherits the multi-modality and
332 covariance structure of the original data. To reconstruct the low-dimensional virtual samples back
333 into the original feature space, a k-Nearest Neighbors regression model is trained ([Niu et al., 2025](#)).
334 This model uses the t-SNE coordinates as input and the standardized original features as output,
335 approximating the inverse of the non-linear t-SNE mapping. Finally, the virtual sample set in its



336 original physical units is obtained by applying inverse standardization.

337 5. Physical constraints and quality control

338 For the target variable, NO_3^- , a non-negativity constraint ($\text{NO}_3^- \geq 0 \text{ mg L}^{-1}$) is imposed to
339 prevent non-physical solutions that may arise from the regression approximation. Other variables,
340 such as pH and ORP, are allowed to fluctuate within reasonable ranges without hard clipping to
341 retain the model's flexibility. The consistency between the virtual and measured samples is
342 validated by comparing their statistical characteristics, including mean, standard deviation,
343 coefficient of variation, range of extreme values, and boxplot distributions. This comparison
344 confirms that the generated data are highly consistent with the original data in terms of statistical
345 properties, without introducing systematic bias or outliers.

346 The advantages of this method are as follows: ① t-SNE excels at capturing local
347 neighborhood relationships, effectively separating implicit subgroups under different hydrological
348 conditions. ② The GMM provides a probabilistic generative framework, supporting reasonable
349 extrapolation for heavy-tailed distributions and extreme values. ③ The KNN-based inverse
350 mapping circumvents the need for large training datasets, which is a limitation of traditional
351 autoencoders, making it particularly suitable for small-sample scenarios (Tang et al., 2022;
352 Razavi-Termeh et al., 2024).

353

354 2.5 Machine learning methods

355 2.5.1 Random forest

356 In this study, Random Forest (RF) was adopted as the baseline model. As an ensemble
357 learning method that leverages bootstrap sampling and random feature selection, RF builds
358 numerous decision trees and integrates their predictions (Abderzak et al., 2025). This approach
359 effectively suppresses overfitting and improves generalization performance, making it especially
360 well-suited for environmental data modeling scenarios involving small samples, high
361 dimensionality, non-linearity, and multicollinearity (Boddu et al., 2025). Hyperparameters were
362 configured based on a preliminary grid search and domain expertise: $n_{\text{estimators}}=100$,
363 $\text{max_depth}=5$, $\text{min_samples_split}=6$, $\text{min_samples_leaf}=3$, and $\text{max_features}=\sqrt{p}$. For the
364 interpretation of driving mechanisms, feature importance was quantified by the mean decrease in
365 Gini impurity (Gini Importance) to identify the critical hydrogeochemical indicator factors (Kaur



366 et al., 2025).

367

368 2.5.2 Hybrid quantum-classical random forest

369 Based on the random forest, a Hybrid Quantum-Classical Random Forest (QCRF) model was
370 constructed, integrating quantum feature enhancement with classical random forests. The core
371 idea of the model is: utilizing a Parameterized Quantum Circuit (PQC) to perform quantum feature
372 encoding on standardized input features, generating high-dimensional quantum features with
373 non-linear entanglement properties (Naresh et al., 2025). These are then concatenated with the
374 original features to construct an enhanced hybrid feature space, which is finally fed into a random
375 forest regressor for modeling (Lamichhane et al., 2025).

376 (1) Quantum Feature Encoding

377 Quantum state transformation of classical data is achieved based on the Z-feature map in
378 quantum computing (Vedavyasa et al., 2025). The ZFeatureMap maps the classical
379 high-dimensional feature space into a quantum Hilbert space through single-qubit Z-gate
380 operations and two-qubit CZ-gate entanglement operations (Khalil et al., 2025). Its core advantage
381 lies in obtaining quantum features via analytical calculation of quantum state vectors, thereby
382 avoiding noise interference introduced by quantum sampling and ensuring feature stability. The
383 ZFeatureMap provided by Qiskit is used as the feature encoding circuit, with its Hamiltonian form
384 given by (Tehrani et al., 2024):

$$385 U_{ZMap}(x) = \prod_{k=1}^R \left[\bigotimes_{i=1}^d H_i \cdot \exp \left(-i \sum_{S \subseteq \{1, \dots, d\}} \phi_S(x) \bigotimes_{j \in S} Z_j \right) \right] \quad (6)$$

386 where d is the selected number of principal factors, R is the number of repetition layers, and ϕ_S
387 represents data-dependent rotation angles (using linear embedding).

388 For each sample x , construct the corresponding quantum state $|\psi(x)\rangle$, and analytically
389 calculate the Pauli-Z expectation value for each qubit (Liao et al., 2024):

$$390 \langle Z_i \rangle = \langle \psi(x) | Z_i | \psi(x) \rangle = P(q_i=0) - P(q_i=1) \quad (7)$$

391 Here, the number of i is equal to the number of predictor variables. This method requires no
392 quantum hardware sampling, completely avoiding the interference of measurement noise and shot
393 noise on small-sample modeling, thus ensuring the determinism and reproducibility of feature
394 generation.



395 (2) Feature Fusion and Modeling

396 Concatenate the original n-dimensional raw features with the n-dimensional quantum $\langle Z \rangle$
397 features to form a 2n-dimensional hybrid feature vector $x_{\text{aug}} = [x_{\text{raw}}; \langle Z \rangle]$ (Cowlessur et al.,
398 2025). That is, original features+quantum encoded features. Using this as input, construct a
399 random forest regression model:

400
$$\hat{y} = \frac{1}{M} \sum_{m=1}^M \text{Tree}_m(x_{\text{aug}}) \quad (8)$$

401 The hyperparameter settings are the same as those for the classical random forest method
402 described above.

403

404 2.6 Evaluation methods and prediction process

405 2.6.1 SHAP analysis

406 This study adopts the SHapley Additive exPlanations (SHAP) method for local and global
407 explainability analysis (Merabet et al., 2025). Through three typical visualization methods, namely
408 summary plot, dependence plot, and waterfall plot, the following are revealed respectively: (1)
409 The overall ranking and distribution of feature importance across all samples (global perspective);
410 (2) The nonlinear relationship or interaction effects between a single predictor variable and the
411 predicted nitrate concentration (conditional dependence); (3) The contribution decomposition of
412 each feature in the prediction result of a representative sample (local attribution) (Alam et al.,
413 2025).

414 The SHAP value is mathematically defined as: the marginal contribution of feature j to the
415 model output offset from the baseline mean (Li et al., 2024), and its form is:

416
$$\phi_j = \sum_{S \subseteq F \setminus \{j\}} \frac{|S|!(|F|-|S|-1)!}{|F|!} [f(S \cup \{j\}) - f(S)] \quad (9)$$

417 where F is the set of all features, S is a subset not containing feature j , and f is the model output.
418 By averaging the absolute SHAP values $|\phi_j|$ over all samples, a feature importance measure with a
419 game-theoretic foundation, unbiased and robust, can be obtained (Hollmann et al., 2025).

420

421 2.6.2 Leave-One-Out Cross-Validation (LOOCV) and model evaluation indicators

422 Given the limited sample size in each hydrological season, this study adopts Leave-One-Out
423 Cross-Validation (LOOCV) for model performance evaluation to maximize the use of training



424 data and reduce evaluation bias (Austin et al., 2025). The LOOCV process is: each time, one
425 sample is left out as the validation set, and the remaining n-1 samples are used for training. After
426 repeating nn times, the average of the evaluation indicators is taken as the final result (Ren et al.,
427 2021).

428 The coefficient of determination R^2 , root mean square error (RMSE), and mean absolute
429 error (MAE) are used to quantitatively describe the model accuracy and error characteristics (Gul
430 et al., 2025):

$$R^2 = 1 - \frac{\sum_{i=1}^n (y_i - \hat{y}_i)^2}{\sum_{i=1}^n (y_i - \bar{y})^2} \quad (10)$$

$$RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (y_i - \hat{y}_i)^2} \quad (11)$$

$$MAE = \frac{1}{n} \sum_{i=1}^n |y_i - \hat{y}_i| \quad (12)$$

434 where y_i is the measured value of the i-th sample, \hat{y}_i is the model's predicted value, \bar{y} is the
435 mean of the measured values, and n is the number of samples in the test set.

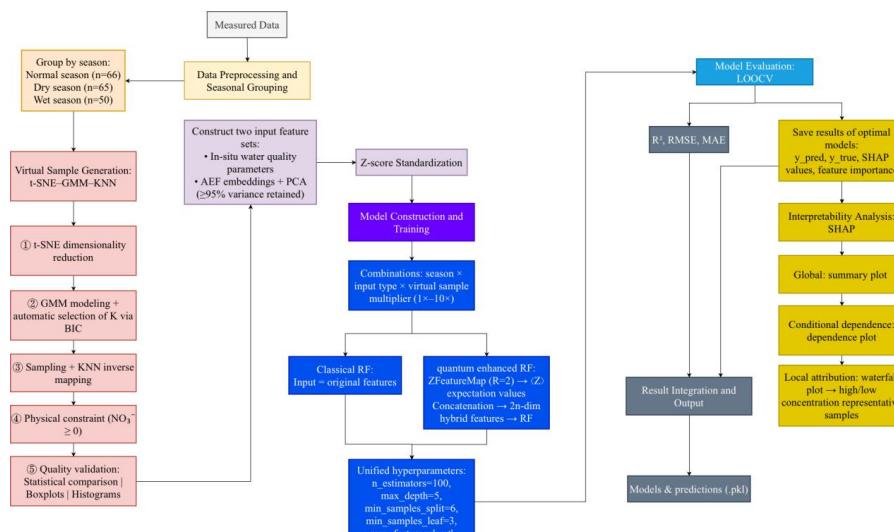
436

437 2.6.3 Standardized prediction workflow

438 To systematically evaluate the nitrate concentration prediction capabilities of different input
439 variables and modeling strategies across various hydrological seasons, and to validate the
440 effectiveness of virtual sample augmentation for small-sample modeling, this study established a
441 standardized prediction pipeline (Fig.2). The specific steps are as follows: (1) Data Preprocessing
442 and Grouping: Observed samples were partitioned by seasons. Z-score normalization was applied
443 separately to two types of input features: field water quality parameters and AlphaEarth
444 Foundation (AEF) features reduced via Principal Component Analysis (PCA). (2) Virtual Sample
445 Generation and Validation: A t-SNE-GMM-KNN strategy was employed to generate virtual
446 samples. Their physical plausibility and distribution consistency were rigorously verified using
447 statistical indicators, box plots, and histograms. (3) Model Training: Under unified
448 hyperparameters, classical Random Forest (RF) and quantum-enhanced RF models were
449 constructed. The latter generates $\langle Z \rangle$ quantum features via Parameterized Quantum Circuits (PQC)
450 encoding, which are concatenated with original features to form $2*n$ input features. Models were
451 trained using two distinct input datasets and combinations of "original samples + 1~10× virtual
452 samples." (4) Model Evaluation: The Leave-One-Out Cross-Validation (LOOCV) strategy was



453 adopted to calculate R^2 , RMSE, and MAE. Visual diagnostics were performed using
 454 observed-predicted scatter plots, residual plots, and box plots. (5) Interpretability Analysis:
 455 Multi-scale interpretation was conducted based on the SHAP framework, including summary plots
 456 (global importance ranking), dependence plots (nonlinear response and interaction effects), and
 457 waterfall plots (local attribution). The driving mechanisms were cross-verified with results from
 458 Bayesian models and Pearson correlation analysis. This workflow encompasses the full process
 459 from data augmentation, modeling, and evaluation to attribution, providing a reproducible and
 460 highly transparent solution for precise groundwater nitrate prediction under conditions of small
 461 samples, multiple seasons, and multi-source inputs.



462
 463 **Fig.2.** Process diagram for constructing prediction framework.

464

465 3. Results

466 3.1 Seasonal hydrochemical controls of nitrate distribution in farmland groundwater

467 3.1.1 Hydrochemical parameters

468 Regarding the basic physical parameters, the pH value was weakly alkaline during the normal
 469 water period with minimal variation, whereas it was near-neutral in the dry and wet seasons (Table
 470 2). A minimum value of 5.77 occurred in the wet season, indicating the presence of acidic water
 471 bodies. Temperature (T) exhibited significant seasonal variation but remained relatively stable
 472 within each season (CV≈0.06). EC, salinity, and TDS showed consistent patterns, all peaking



473 during the dry season and reaching their lowest levels in the wet season. The redox indicators
474 displayed high volatility. The mean DO was slightly higher in the wet season, while the mean
475 ORP was consistently low across all seasons, with extremely large standard deviations and
476 coefficients of variation. In terms of ionic composition, the average concentrations of Ca^{2+} and
477 Mg^{2+} were highest during the dry season, and Na^+ also peaked in this period. The concentration of
478 K^+ was relatively low. Among the anions, HCO_3^- concentration was highest in the wet season,
479 while the average concentrations of Cl^- and SO_4^{2-} were both at their maximum during the dry
480 season. The average concentration of NO_3^- was higher in the dry season than in other periods and
481 lowest in the wet season. The concentrations of nitrite (NO_2^-) and ammonium (NH_4^+) were much
482 lower than that of nitrate. Concerning the distribution of the indicators, most variables were
483 right-skewed. Notably, extreme values were present for NO_3^- in the dry season (maximum=358.58
484 mg L^{-1} , mean=42.93 mg L^{-1}), Cl^- in the dry season (maximum=241.36 mg L^{-1} , mean=24.90 mg L^{-1}), and F^- in the normal period (maximum=13.17, mean=3.70).

486 **Table 2.** Statistical summary of chemical and field measurement parameters.

Periods		pH	T	EC	DO	ORP	Salt	TDS	Depth	K^+	Ca^{2+}
	Unit		°C	$\mu\text{s cm}^{-1}$	mg L^{-1}	mv	ppt	mg L^{-1}	m	mg L^{-1}	mg L^{-1}
Normal season n=66	Max	8.60	16.70	1110.00	10.57	369.60	0.53	724.00	20.64	36.66	69.73
	Min	7.61	13.40	349.00	2.20	-58.30	0.11	225.00	18.26	1.04	13.38
	Mean	8.16	14.90	549.32	6.31	4.57	0.21	357.21	18.91	2.92	37.28
	SD	0.11	0.71	173.56	1.94	56.27	0.09	113.34	0.62	4.45	13.06
Dry season n=65	Max	8.21	18.80	1134.00	8.82	144.20	0.54	737.00	18.95	3.52	43.86
	Min	6.97	14.10	343.00	1.85	-110.40	0.11	227.00	17.88	0.67	4.81
	Mean	7.35	15.59	658.68	6.34	4.55	0.27	427.65	18.43	1.93	16.89
	SD	0.45	0.92	185.20	1.64	42.10	0.10	120.55	0.27	0.56	8.35
Wet season n=50	Max	8.97	21.00	977.00	9.61	195.00	0.41	635.00	18.87	3.37	51.54
	Min	5.77	15.50	371.00	3.34	-112.20	0.12	243.00	17.61	1.14	17.27
	Mean	7.34	17.01	535.24	6.98	17.57	0.20	347.92	18.27	1.75	25.70



	SD	0.52	0.94	153.65	1.32	47.59	0.08	100.39	0.40	0.34	6.83
	CV	0.07	0.06	0.29	0.19	2.71	0.42	0.29	0.02	0.19	0.27
		Na ⁺	Mg ²⁺	HCO ₃ ⁻	Cl ⁻	SO ₄ ²⁻	F	NO ₃ ⁻	NO ₂ ⁻	NH ₄ ⁺	
	Unit	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	mg L ⁻¹	
Normal season	Max	98.85	92.78	192.15	87.09	40.49	13.17	161.17	0.45	0.20	
n=66	Min	3.81	3.80	22.88	1.03	1.79	0.12	2.39	0.04	0.00	
	Mean	24.01	31.78	96.65	20.53	15.53	3.70	33.67	0.11	0.04	
	SD	12.46	17.18	44.51	17.84	10.02	2.02	35.83	0.08	0.04	
	CV	0.52	0.54	0.46	0.87	0.64	0.55	1.06	0.72	1.00	
Dry season	Max	123.20	138.90	289.29	241.36	123.75	0.57	358.58	10.38	0.84	
n=65	Min	16.28	17.97	5.10	1.40	2.00	0.21	0.10	0.57	0.05	
	Mean	48.52	58.17	42.19	24.90	15.91	0.33	42.93	3.35	0.16	
	SD	16.48	25.86	71.11	35.13	18.67	0.08	56.35	1.98	0.11	
	CV	0.34	0.44	1.69	1.41	1.17	0.26	1.31	0.59	0.69	
Wet season	Max	36.21	53.72	207.40	51.93	34.15	0.45	98.36	4.31	0.41	
n=50	Min	10.95	12.76	83.88	1.56	5.88	0.09	4.15	1.51	0.02	
	Mean	21.78	22.57	132.49	15.60	15.16	0.20	27.14	2.62	0.10	
	SD	4.90	8.13	27.12	13.60	7.47	0.08	23.86	0.61	0.08	
	CV	0.23	0.36	0.20	0.87	0.49	0.39	0.88	0.23	0.83	

487

488 3.1.2 Type of water

489 During the dry season, the data points are highly concentrated in the zone of calcium-type
 490 cations and bicarbonate-type anions, indicating that the groundwater is primarily controlled by the
 491 dissolution of carbonate rocks (Fig.3). In the wet season, although the Ca-Mg-HCO₃⁻ type remains
 492 dominant, some samples shift towards the sulfate and chloride types, reflecting the leaching input
 493 effect of surface pollutants (such as agricultural fertilizers and domestic sewage) brought by
 494 rainfall infiltration. By the normal season, the hydrochemical types exhibit the widest distribution,
 495 presenting a mixed type with coexisting bicarbonate and chloride types. Overall, the groundwater
 496 hydrochemical characteristics in the study area are jointly controlled by precipitation-evaporation
 497 dynamics and carbonate weathering.

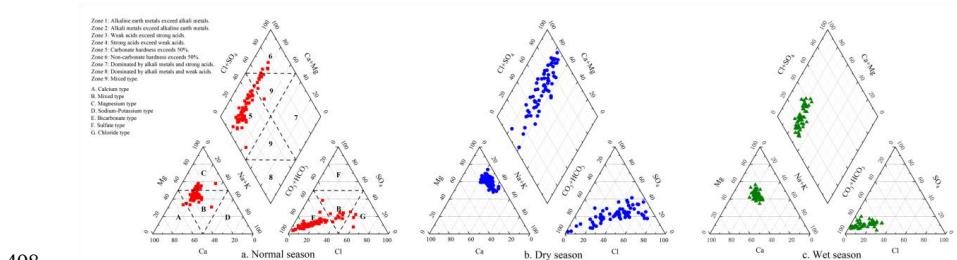
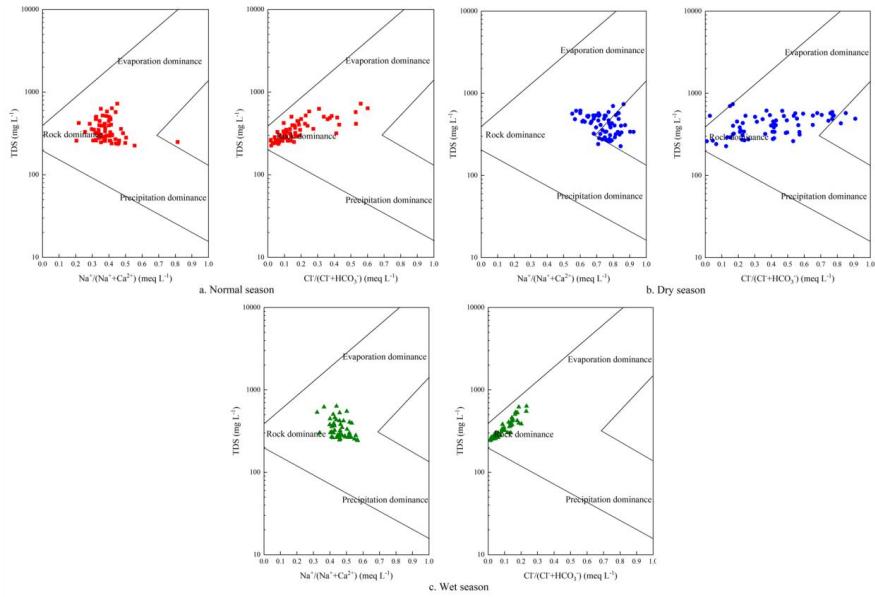


Fig.3. Piper diagram classifying the hydrochemical facies of the analyzed groundwater.

3.1.3 Sources and controlling factors of ions in groundwater

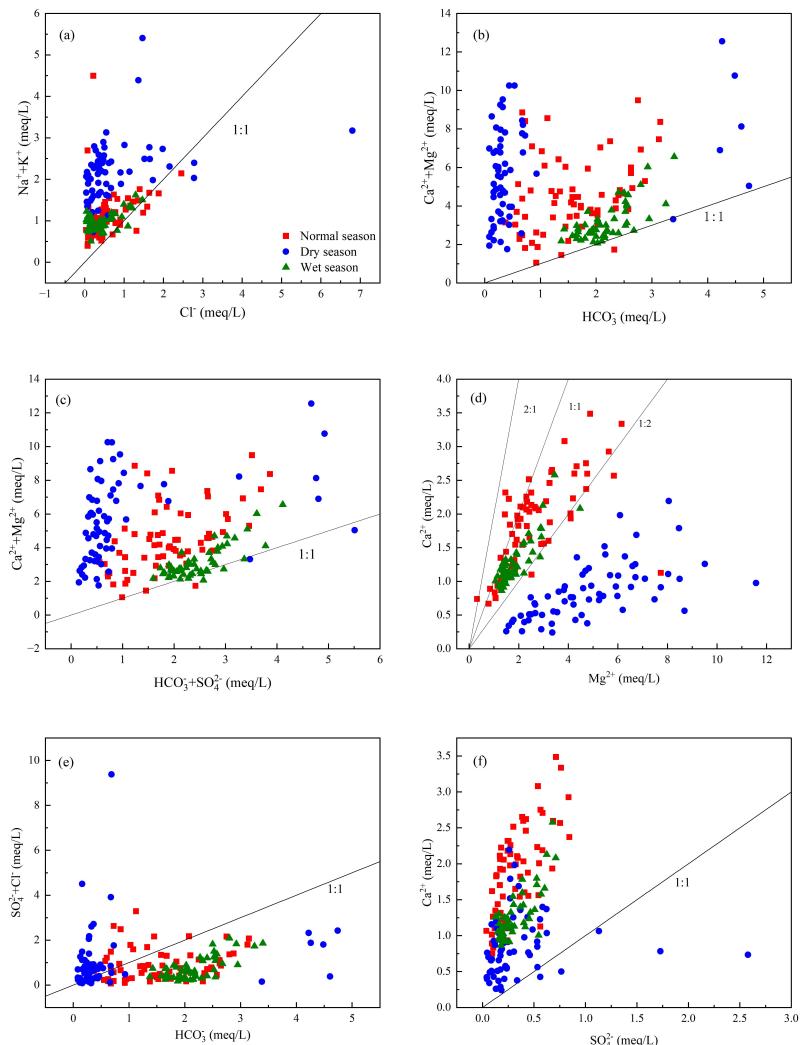
The Gibbs diagram shows that the groundwater in the study area is primarily controlled by rock weathering during the normal, dry, and wet seasons, indicating the dominance of water-rock interaction (Fig. 4). The ratio of $\gamma(\text{Na}^+ + \text{K}^+)$ to γCl^- (Fig. 5a) shows that the vast majority of sample points plot above the 1:1 line, indicating that Na^+ and K^+ are primarily sourced from the dissolution of evaporite rocks. In the relationships between $\gamma(\text{Ca}^{2+} + \text{Mg}^{2+})$ and γHCO_3^- , and between $\gamma(\text{Ca}^{2+} + \text{Mg}^{2+})$ and $\gamma(\text{HCO}_3^- + \text{SO}_4^{2-})$ (Fig. 5b-c), samples from all periods plot above the 1:1 line, confirming that Ca^{2+} and Mg^{2+} mainly originate from the dissolution of carbonate minerals. Furthermore, the γCa^{2+} - γMg^{2+} relationship (Fig. 5d) helps identify the types of mineral dissolution. Samples from the dry season are concentrated below the 1:2 line, indicating a dominance of magnesium-poor mineral dissolution, with cation exchange causing a relative depletion of Ca^{2+} . Samples from the normal and wet seasons are stably distributed between the 1:1 and 1:2 lines, reflecting that dolomite dissolution has reached equilibrium while calcite remains in a state of non-equilibrium dissolution, continuously supplying Ca^{2+} . In the relationship between $\gamma(\text{SO}_4^{2-} + \text{Cl}^-)$ and γHCO_3^- (Fig. 5e), the distribution of sample points on both sides of the 1:1 line suggests that groundwater ions have dual contributions from both evaporite and carbonate rocks. Conversely, in the γCa^{2+} versus γSO_4^{2-} relationship (Fig. 5f), samples generally plot above the 1:1 line, which excludes gypsum as a primary source of Ca^{2+} and indicates that Ca^{2+} is mainly derived from the dissolution of carbonate minerals. Therefore, the chemical composition of groundwater in the study area is primarily controlled by the dissolution of carbonate minerals, and is also influenced by hydrological seasonal variations and cation exchange processes.



522

523

Fig.4. Gibbs diagrams of the groundwater samples.



524

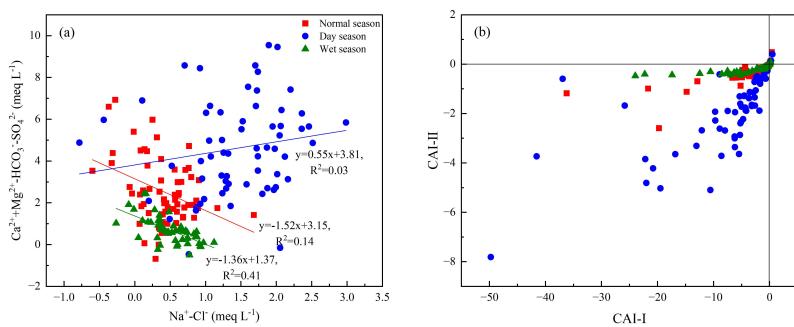
525

Fig.5. Plots of ion ratio relationship.

526 The Chloro-Alkaline Index method was employed to analyze the cation exchange and
527 adsorption between groundwater and sediments. A CAI value less than zero indicates the
528 occurrence of cation exchange, with more negative values reflecting stronger exchange intensity. .
529 Furthermore, the relationship between $[\gamma(\text{Ca}^{2+}) + \gamma(\text{Mg}^{2+}) - \gamma(\text{HCO}_3^-) - \gamma(\text{SO}_4^{2-})]$ and $[\gamma(\text{Na}^+) -$
530 $\gamma(\text{Cl}^-)]$ can be used to further investigate the cation exchange processes in the groundwater. During



531 the dry season, the slope was 0.55, suggesting the presence of extremely weak cation exchange in
532 the water body (Fig. 6a). With the exception of samples from the dry season, sampling points from
533 the normal and wet seasons were plotted near a line with a slope of -1, with respective slopes of
534 -1.52 and -1.36. This trend is consistent with the conclusions drawn from the Chloro-Alkaline
535 Index, providing further evidence that cation exchange and adsorption occurred in the
536 groundwater during the normal and wet seasons. The ion exchange process was more active
537 during the rainy season ($R^2=0.41$), leading to the enrichment of Na^+ in the groundwater of the area.
538 In contrast, most groundwater samples from the dry season showed no evidence of cation
539 exchange and adsorption.



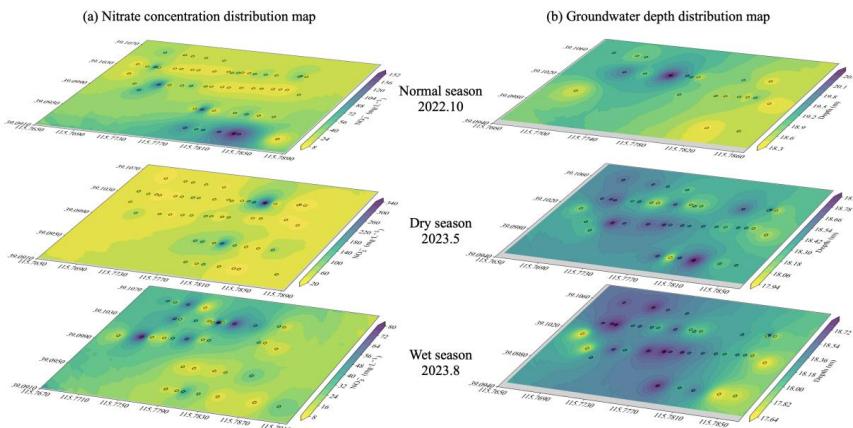
540
541 **Fig.6.** Relationship diagram of groundwater ($\text{Ca}^{2+} + \text{Mg}^{2+} - \text{SO}_4^{2-} - \text{HCO}_3^-$) and ($\text{Na}^+ - \text{Cl}^-$) along with
542 CAI-1 and CAI-2 correlation diagrams.
543

544 3.1.4 Spatial distribution dynamics of groundwater depth and nitrate driven by seasonal
545 hydrological processes

546 The spatial distribution of groundwater depth reflects the regional hydraulic gradient and
547 groundwater flow direction, whereas the spatial variability of nitrate concentration is closely
548 associated with flow paths, pollution source inputs, and hydrological processes (Fig.7). During the
549 normal season, the groundwater depth distribution is relatively uniform. The eastern region of the
550 farm, characterized by shallower depths, serves as a recharge zone, with groundwater flowing
551 towards the deeper western region. At this time, nitrate concentration are relatively dispersed, with
552 high-concentration zones located in the southeastern part of the farm. In the dry season, the
553 groundwater depth becomes shallower, and the flow direction shifts from the eastern and western



554 sides towards the central area. During this period, nitrate concentration reach their annual peak
555 (mean: 42.93 mg L⁻¹). The distribution of nitrate exhibits a higher degree of spatial coincidence
556 with the groundwater flow direction, indicating that enhanced evaporative concentration during
557 the dry season leads to the further accumulation of flow-transported pollutants in the discharge
558 zone. In the wet season, the groundwater depth further decreases, and groundwater flows from the
559 southeastern region towards the northwestern region. nitrate concentration drop to their annual
560 minimum (mean: 27.14 mg L⁻¹). High-concentration areas are distributed in the northwest,
561 overlapping with regions of deeper groundwater depth. It is inferred that precipitation infiltration
562 during the rainy season dilutes the groundwater nitrate; as dilution is the dominant process during
563 infiltration, the nitrate concentration exhibits a decreasing trend along the groundwater flow
564 direction.



565
566 **Fig. 7.** Spatial distribution of nitrate concentration and groundwater depth in different seasons.

567
568 3.2 Groundwater recharge sources and pollution source identification

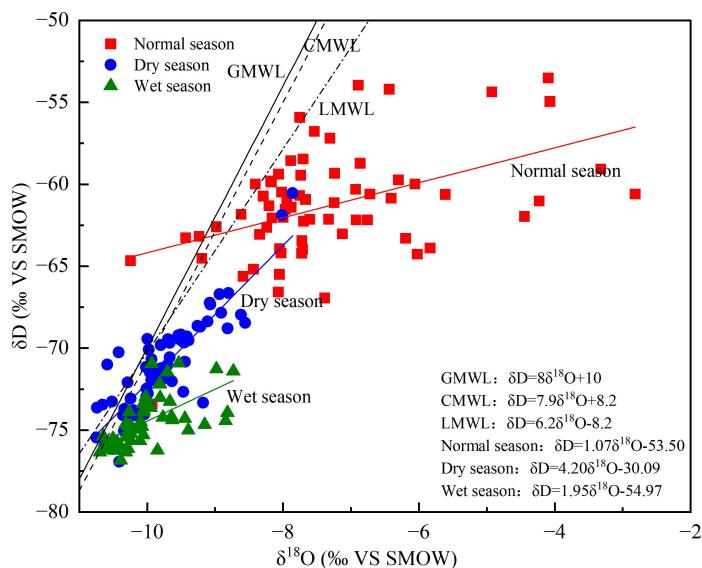
569 3.2.1 Stable hydrogen and oxygen isotope composition of water

570 During the normal season, the mean values of groundwater δD and $\delta^{18}\text{O}$ were -61.31‰ and
571 -7.31‰, with ranges of -73.40‰ to -53.52‰ and -10.25‰ to -2.82‰, respectively. The d-excess
572 values ranged from -38.07‰ to 17.30‰, with a mean of -2.80‰. In the dry season, the mean
573 groundwater δD and $\delta^{18}\text{O}$ values were -71.05‰ and -9.74‰, with ranges of -76.93‰ to -60.55‰
574 and -10.75‰ to -7.86‰, respectively. The $\delta^{17}\text{O}$ values ranged from -5.60‰ to -2.79‰, averaging



575 -5.01‰, while the d-excess varied from 0.01‰ to 13.69‰, with a mean of 6.89‰. During the wet
576 season, the mean groundwater δD and $\delta^{18}\text{O}$ were -74.43‰ and -9.99‰, with ranges of -76.84‰ to
577 -70.91‰ and -10.70‰ to -8.73‰, respectively. The $\delta^{17}\text{O}$ values were between -5.72‰ and
578 -4.72‰, with a mean of -5.28‰, and the d-excess ranged from -3.67‰ to 9.80‰, averaging
579 5.50‰. The d-excess during the dry season was the highest among the three periods, while it was
580 the lowest during the normal period, indicating significant variations in d-excess across different
581 seasons. The $\delta^{17}\text{O}$ values in the dry season were higher than those in the wet and normal periods,
582 which is a direct reflection of the impact of precipitation variations on the isotopic composition of
583 the water body.

584 The isotopic values of precipitation δD and $\delta^{18}\text{O}$ ranged from -97.78‰ to -20.22‰ and from
585 -13.48‰ to -1.96‰, with mean values of -55.36‰ and -7.60‰, respectively. Overall, the $\delta^{18}\text{O}$
586 and δD values of precipitation in the study area fall within the global ranges of -50‰ to 10‰ and
587 -350‰ to 50‰. The Local Meteoric Water Line (LMWL) for the study area is defined by the
588 equation: $\delta D = 6.2 \delta^{18}\text{O} - 8.2$ (Fig.8). Specifically, the equations for the normal, dry, and wet
589 seasons are: $\delta D = 1.07 \delta^{18}\text{O} - 53.50$, $\delta D = 4.20 \delta^{18}\text{O} - 30.09$, and $\delta D = 1.95 \delta^{18}\text{O} - 54.97$, respectively.
590 The slope of the annual LMWL is lower than that of the Global Meteoric Water Line (GMWL)
591 proposed by Craig in 1964 ($\delta D = 8 \delta^{18}\text{O} + 10$), as well as lower than the China Meteoric Water
592 Line (CMWL) ($\delta D = 7.9 \delta^{18}\text{O} + 8.2$). The stable hydrogen and oxygen isotopic characteristics of
593 groundwater samples from the three periods all exhibit a discrete, linear distribution and plot
594 below both GMWL and LMWL. This phenomenon reveals that the water isotopes have undergone
595 strong fractionation during evaporation in the normal, wet, and dry seasons. Furthermore, the
596 stable hydrogen and oxygen isotope data for the dry and normal seasons are mainly concentrated
597 in the lower-left region of the plot, indicating relative isotopic depletion during these two periods.
598 During the normal period, the δD and $\delta^{18}\text{O}$ values exhibit a high degree of dispersion and are
599 widely distributed in the upper-central part of the scatter plot. This reflects that the stable
600 hydrogen and oxygen isotopes are relatively enriched and have a wide range of variation during
601 the normal season.



602

603 **Fig.8.** $\delta^{18}\text{O}/\delta\text{D}$ relationship of groundwater samples.

604

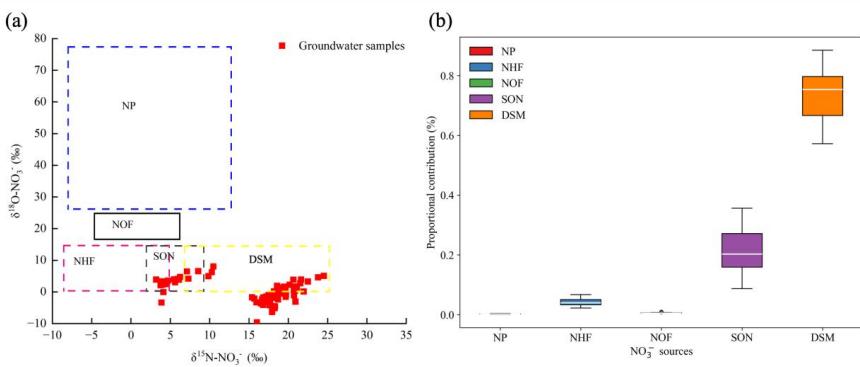
605 3.2.2 Identification of nitrate sources using isotopes and MixSIAR model

606 During the normal water period, the nitrogen and oxygen isotopic compositions in
607 groundwater exhibit a wide range of variation. The $\delta^{15}\text{N-NO}_3^-$ values range from 5.6‰ to 24.52‰
608 (average: 18.22‰), while the $\delta^{18}\text{O-NO}_3^-$ values range from -6.33‰ to 6.23‰ (average: 0.22‰).
609 In the low water period, the range of $\delta^{15}\text{N-NO}_3^-$ values expands to 3.2‰-21.96‰ (average:
610 12.19‰), and the $\delta^{18}\text{O-NO}_3^-$ values range from -9.58‰ to 8.04‰ (average: 0.65‰). Previous
611 studies have established characteristic $\delta^{18}\text{O-NO}_3^-$ ranges for different nitrate sources: atmospheric
612 deposition (23‰-75‰), nitrate fertilizers (18‰-24‰), and products of nitrification (-10‰-10‰).
613 The data points are predominantly concentrated within the zone of animal manure and domestic
614 wastewater, indicating that nitrate is primarily derived from these sources, with soil nitrogen as a
615 secondary contributor.

616 The MixSIAR model was employed to quantitatively apportion the sources of groundwater
617 nitrate nitrogen. According to the average contributions from each source, the five pollution
618 sources in the study area were ranked as follows: DSM (74.1%) > SON (20.9%) > NHF (4.2%) >



619 NOF (0.6%) > NP (0.2%) (Fig.9). This indicates that the primary contributor to groundwater
620 NO_3^- -N in the study area was manure and sewage, followed by soil nitrogen. The influences of
621 atmospheric precipitation and chemical fertilizers on groundwater nitrate were negligible. The
622 quantitative results from the MixSIAR analysis are consistent with the qualitative findings,
623 confirming that manure and sewage, along with soil nitrogen, are the dominant sources of nitrate
624 pollution in the study area.



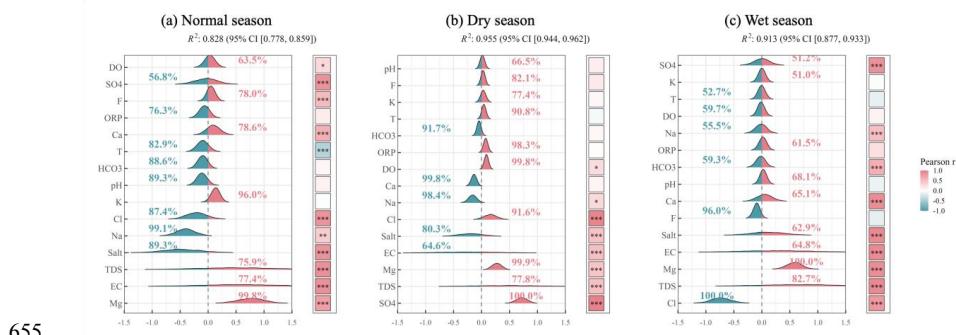
625
626 **Fig.9.** (a) distributions of $\delta^{15}\text{N-NO}_3^-$ and $\delta^{18}\text{O-NO}_3^-$ values in the study area. (b) proportional
627 contributions of the main NO_3^- sources evaluated by the MixSIAR model. Note: boxplots denote
628 the 25th, 50th and 75th percentiles.

629
630 3.3 Bayesian model analysis and correlation analysis

631 During the normal season, Bayesian model indicated the central role of Mg^{2+} , which is
632 consistent with its strong positive correlation ($r=0.75$) in the correlation matrix (Fig.10). Na^+
633 exhibited a significant negative effect, whereas it only showed a weak positive correlation ($r=0.39$)
634 in the correlation matrix. This suggests that variations in Na^+ concentration are more reflective of
635 hydrological processes, such as evaporative concentration, rather than direct involvement in the
636 chemical transformation of NO_3^- . Although the correlation matrix revealed strong correlations
637 between NO_3^- and both TDS and EC ($r > 0.8$), their respective probabilities of direction (pd) in the
638 Bayesian model were both below 80%. This further confirms that their influence is primarily
639 manifested indirectly through collinearity with other ions. In the dry season, Bayesian model
640 identified SO_4^{2-} as the primary positive driver of NO_3^- , a finding that is in strong agreement with
641 the high positive correlation ($r=0.96$) between SO_4^{2-} and NO_3^- observed in the correlation matrix.



642 Concurrently, Bayesian model indicated significant negative effects for both Na^+ and Ca^{2+} , which
 643 contrasts with their weak positive correlations with NO_3^- in the correlation matrix. This
 644 discrepancy likely arises because the elevated concentrations of Na^+ and Ca^{2+} are attributed to
 645 evaporative concentration, whereas the increase in NO_3^- stems from anthropogenic inputs,
 646 indicating no direct causal relationship between them. During the wet season, Bayesian model
 647 identified Cl^- and Mg^{2+} as the most critical driving factors, with clear directional effects and high
 648 confidence. This aligns with the trends observed in the correlation matrix, where NO_3^- correlated
 649 negatively with Cl^- ($r=-0.78$) and positively with Mg^{2+} ($r=0.84$), thereby validating their direct
 650 influence on NO_3^- concentrations during the wet season. While the correlation matrix also showed
 651 high positive correlations between NO_3^- and both TDS and EC, their posterior distributions in the
 652 Bayesian model were wider and their pd values were lower. This suggests that their impact is
 653 likely a result of high collinearity with key variables such as Mg^{2+} and Cl^- , rather than an
 654 independent effect.



655
 656 **Fig.10.** Factor effects and Pearson coefficients of physicochemical variables on NO_3^- at different
 657 periods. The left part of each subplot shows the relative importance and posterior distribution of
 658 each environmental variable to NO_3^- after the Bayesian model operation. The red area represents
 659 the probability density of the positive effect, and the blue area represents the probability density of
 660 the negative effect. The percentage values beside the distribution represent the Probability of
 661 Direction (pd). The right part of each subplot is the heat map of the correlation analysis.

662
 663 3.4 Model performance evaluation
 664 3.4.1 Virtual data analysis



665 To address the modeling bias arising from limited measured samples, this study constructed
666 virtual datasets at 1-10 times the original scale based on a strategy combining t-SNE
667 dimensionality reduction, GMM clustering sampling, and KNN inverse mapping, to enhance the
668 robustness of model training. Taking the 10x virtual dataset as an example, the statistical
669 characteristics (Table 3) show that the virtual samples effectively reproduced the central tendency
670 and dispersion of the original data. For the normal season, the mean NO_3^- concentration was 30.41
671 mg L^{-1} (vs. observed mean of 33.67), with a standard deviation of 28.78 (vs. 35.83) and a
672 coefficient of variation (CV) of 0.95 (vs. 1.06). In the dry season, the maximum value of the
673 virtual samples reached 178.09 mg L^{-1} , while this did not fully replicate the extreme high values
674 (observed maximum of 358.58 mg L^{-1}), it effectively expanded the range of the heavy-tailed
675 distribution. For the wet season, although the CV for all indicators was slightly lower than the
676 measured values, their ranges (8.47-80.37 vs. 4.15-98.36 mg L^{-1}) still showed a high degree of
677 overlap, indicating that no systematic distortion was introduced.

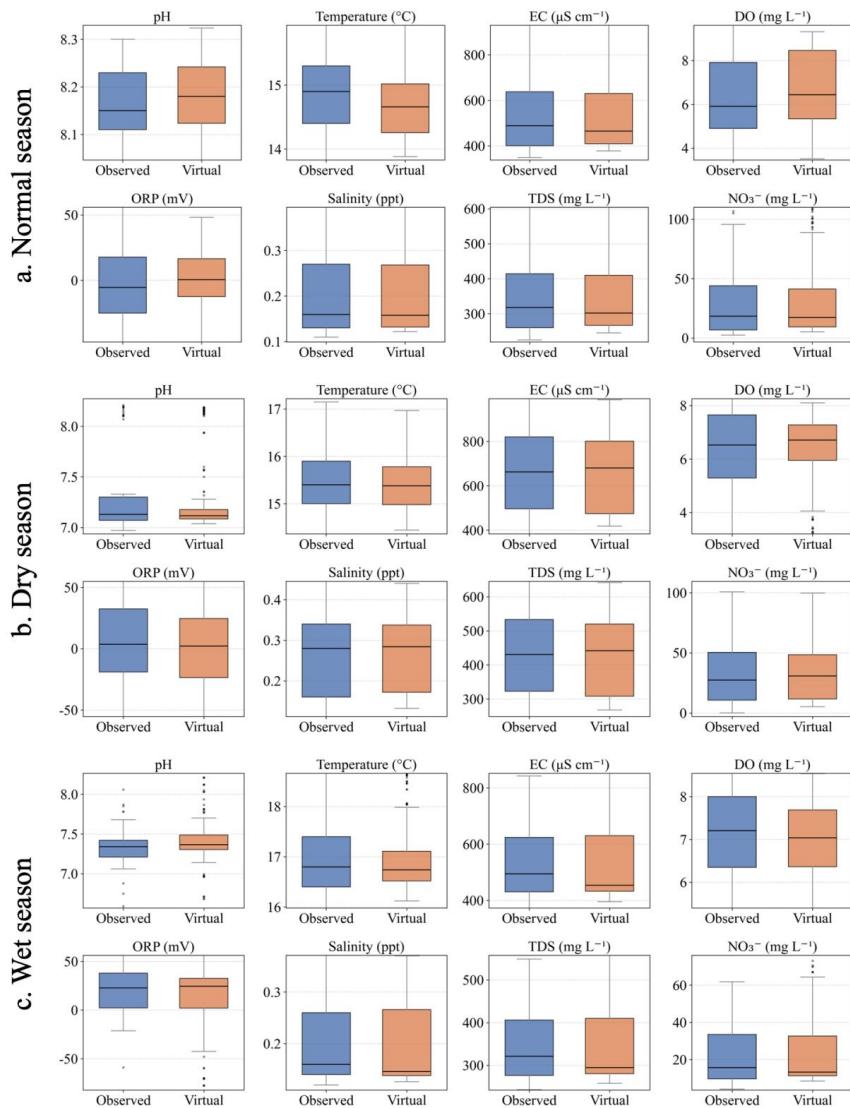
678 **Table 3.** Statistical characteristics of different virtual samples.

Periods		pH	T	EC	DO	ORP	Salt	TDS	NO_3^-
	Unit		$^{\circ}\text{C}$	$\mu\text{s cm}^{-1}$	mg L^{-1}	mv	ppt	mg L^{-1}	mg L^{-1}
Normal season n=660	Max	8.32	16.2	963.6	9.31	101.02	0.42	628.6	124.03
	Min	7.95	13.88	378.2	3.52	-48.28	0.12	245	5.10
	Mean	8.18	14.72	527.31	6.66	2.09	0.20	342.72	30.41
	SD	0.07	0.54	148.31	1.66	25.97	0.08	96.92	28.78
Dry season n=650	Max	8.19	17.59	987.8	8.10	69.18	0.44	641.8	178.09
	Min	7.04	14.44	417.6	2.68	-59.14	0.132	267.8	5.19
	Mean	7.35	15.49	661.79	6.414	0.52	0.27	429.90	37.75
	SD	0.44	0.69	170.65	1.24	29.61	0.09	111.16	33.13
Wet season n=500	Max	8.21	18.88	872.8	8.54	56.4	0.37	569	80.37
	Min	6.31	16.12	395.8	5.34	-77.64	0.13	257.6	8.47
	Mean	7.38	16.93	535.7	7.06	13.75	0.20	348.24	25.85



SD	0.33	0.62	131.02	0.78	28.85	0.08	86.05	19.88
CV	0.04	0.04	0.24	0.11	2.10	0.38	0.25	0.77

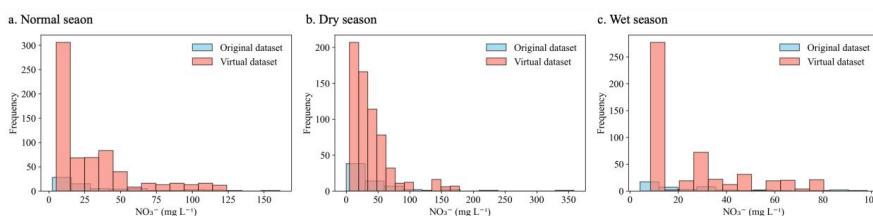
679 Standardized multivariate boxplots ([Fig.11](#)) visually confirm that the median, interquartile
680 range (IQR), whisker length, and outlier distribution of the virtual data for each period were
681 highly similar to the measured data, demonstrating that the central tendency and dispersion
682 characteristics were well-preserved. Hydrological seasonal characteristics, such as high
683 EC/TDS/Cl/NO₃⁻ in the dry season and low, concentrated NO₃⁻ in the wet season, were also
684 accurately preserved. Although the number of some newly added outliers slightly increased, they
685 all fell within physically reasonable ranges, with no non-physical solutions, such as negative
686 concentrations or out-of-bounds pH values, occurring. [Fig.12](#) presents a comparison of nitrate
687 concentration frequency distributions between the original and synthetic datasets across normal,
688 dry, and wet periods. The distributional comparison indicates that the proposed t-SNE + GMM +
689 KNN inverse mapping synthetic sample generation strategy maintains the core features of the
690 NO₃⁻ distribution for each hydrological period, while simultaneously improving sample
691 representation in sparse areas and intervals of high variability. Therefore, the t-SNE + GMM
692 method effectively captured the non-linear structure and extreme value information of the original
693 data, and can provide reliable data support for subsequent model training.



694

695

Fig.11. Box plots of the observed and virtual variable data at different periods.



696

697

Fig.12. Comparison of nitrate concentration distribution in the original and virtual datasets under



698 different periods.

699

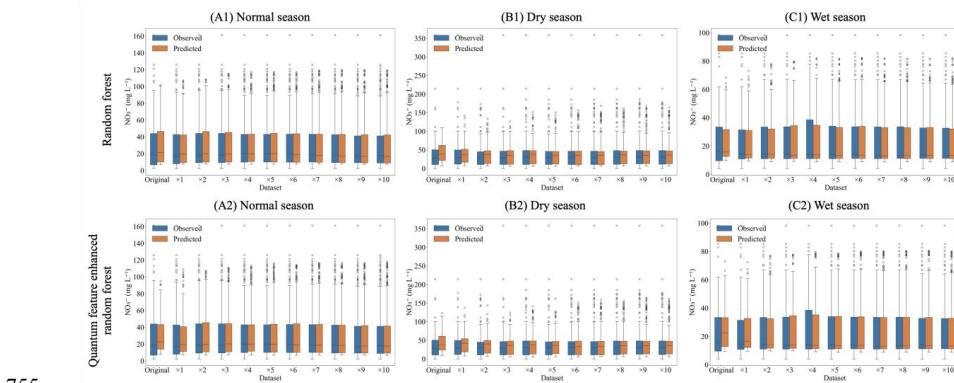
700 3.4.2 Prediction based on on-site measured water quality data

701 During the normal season, the R^2 values for the baseline Random Forest and the
702 quantum-enhanced RF models were 0.673 and 0.660, respectively, indicating high prediction
703 errors ([Table S1](#)). As the number of virtual samples was increased from 1-fold to 10-fold the size
704 of the original dataset, the R^2 of both models steadily improved to above 0.958, with the
705 quantum-enhanced RF model performing better and ultimately achieving an R^2 of 0.9622. When
706 the number of virtual samples exceeded 500, the performance gains began to plateau. For the dry
707 season, the modeling performance with the original data was the poorest, which is correlated with
708 the high variability of NO_3^- concentrations during this period ([Table S2](#)). This suggests that with a
709 limited sample size, models are susceptible to interference from outliers, and a small number of
710 measured samples is insufficient to support effective model learning. After introducing virtual
711 samples, the model performance improved significantly. A mere 1-fold augmentation of the
712 sample size increased the R^2 to 0.527 (RF). When augmented to 8-fold, the R^2 reached 0.854.
713 Although the quantum-enhanced model slightly underperformed the classical RF in the initial
714 stages (≤ 2 -fold augmentation), their performances converged at higher augmentation levels, both
715 achieving high accuracy. This indicates that virtual samples effectively mitigated the modeling
716 challenges posed by data sparsity and skewed distributions. In contrast, the modeling performance
717 with original data was optimal during the wet season, attributed to the generally lower NO_3^-
718 concentrations and their smaller spatial variability ([Table S3](#)). The use of virtual samples further
719 elevated the prediction accuracy to an exceptionally high level. A 4-fold augmentation yielded an
720 R^2 of 0.962. After augmentation to 10-fold, the R^2 of the RF model stabilized at 0.977, with the
721 RMSE dropping to as low as 3.03 mg L⁻¹. The overall performance of the quantum-enhanced RF
722 was consistent with the classical RF, with only slight fluctuations within a very small error range,
723 showing that when data quality is high and the relationships are more linear, the marginal gains
724 from quantum feature encoding are limited.

725 During the normal season, as illustrated in [Fig. 13\(A1\)-\(A2\)](#), a deviation was observed
726 between the predicted and observed values for both models when utilizing only the 66 original
727 samples. The prediction results exhibited high dispersion, and the median deviated markedly from



728 the observed median. This is consistent with the low R^2 values, indicating errors inherent in
729 small-sample modeling. With an increase in the number of virtual samples, the distribution of
730 predicted values gradually converged towards the observed values, and the interquartile range
731 (IQR) and whiskers of the boxplots progressively narrowed, indicating a substantial enhancement
732 in model stability and accuracy. When the virtual samples were expanded tenfold (to 660 virtual
733 samples), the boxplots of the predicted values highly overlapped with those of the observed values,
734 consistent with the reduction of the RMSE to 6.02 mg L^{-1} . In the final stage, the
735 quantum-enhanced model slightly outperformed the classical RF model, achieving an R^2 of 0.9622.
736 In the dry season (Fig. 13(B1)-(B2)), on the original dataset, the predicted values from both
737 models were generally overestimated due to the extremely high and skewed distribution of NO_3^-
738 concentrations. Consequently, the predicted boxplots were positioned entirely above the actual
739 values, yielding an R^2 of only 0.28. The introduction of virtual samples led to a substantial
740 improvement in model performance. Starting from single-fold augmentation, the median and
741 range of the predicted boxplots began to converge towards the observed values, at augmentation
742 levels of 8-fold and higher, the predicted values effectively captured the distributional
743 characteristics of the high-concentration intervals. Although the classical RF model slightly
744 outperformed the quantum-enhanced model at low augmentation levels, their performances
745 converged as the sample size further increased. This demonstrates that the virtual sample
746 generation strategy effectively alleviates modeling challenges caused by data sparsity and extreme
747 values. During the wet season (Fig. 13(C1)-(C2)), the predicted values of both models already
748 exhibited good consistency with the observed values on the original dataset, with the predicted
749 boxplots substantially overlapping the observed ones. With the incorporation of virtual samples,
750 the prediction accuracy and stability of the models were further enhanced. The IQR and whiskers
751 of the boxplots continued to narrow, and the predicted values became more concentrated within
752 the true distribution range of the observed values. Following tenfold data augmentation, the
753 agreement between predicted and observed values was exceptionally high, with an R^2 reaching
754 0.977 and an RMSE decreasing to 3.03 mg L^{-1} , demonstrating excellent predictive performance.



755

756 **Fig.13.** Comparison of observed and predicted NO_3^- concentrations across data augmentation

757 levels for random forest and quantum feature-enhanced random forest models in normal, dry, and
758 wet Seasons.

759

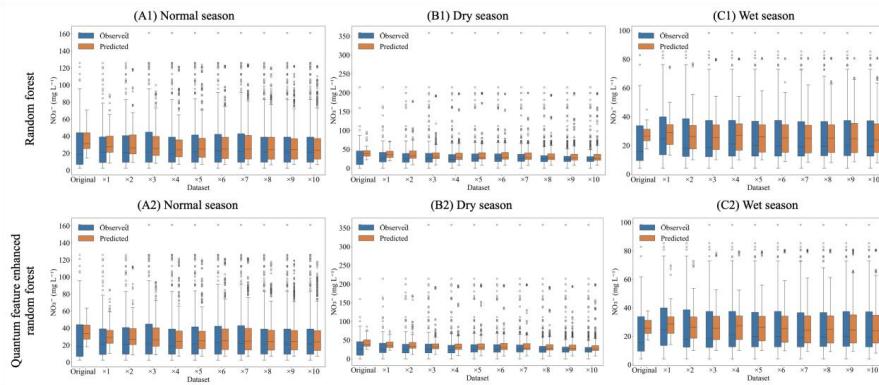
760 3.4.3 Prediction based on AlphaEarth Foundation Embeddings

761 To explore the potential of remote sensing semantic embedding features in predicting
762 groundwater nitrate concentrations, this section employs the 64-dimensional surface semantic
763 vectors derived from the Google AlphaEarth Foundation (AEF) dataset as model input variables.
764 We reduced the variables through principal component analysis to preserve $\geq 95\%$ variance.

765 During the normal season, modeling based on original samples yielded poor performance.
766 The R^2 for Random Forest and quantum-enhanced RF were 0.167 and 0.119, respectively, with
767 RMSE values as high as 32.89 and 33.82 mg L $^{-1}$ (Table S4). These results suggest that, given the
768 limited sample size, relying solely on AEF embedding features is insufficient to fully capture the
769 hydrological processes characteristic of this period. Model performance improved with the
770 introduction of virtual samples. When the virtual samples were expanded to ten times the size of
771 the original dataset, the R^2 of the RF model increased to 0.860, and the RMSE decreased to 10.73
772 mg L $^{-1}$. Similarly, the quantum-enhanced RF achieved an R^2 of 0.844, exhibiting a consistent
773 overall trend. A comparison of boxplots (Fig. 14A) reveals that the initial predictions severely
774 overestimated the low-to-medium concentration ranges while underestimating the high-value tails.
775 As the sample size expanded, the predicted boxplots progressively converged toward the observed
776 distribution. The agreement between the median and interquartile range (IQR) improved
777 significantly, confirming that virtual samples effectively enhanced the capability of AEF features



778 to represent non-linear patterns.



779

780 **Fig.14.** Comparison of observation and prediction of NO₃⁻ concentration by random forest and
781 quantum featture-enhanced random forest models at data enhancement levels in normal, dry and
782 wet seasons: based on AlphaEarth Foundation as the input variable.

783 In the dry season, modeling with the original dataset resulted in a negative R², reflecting the
784 extremely weak generalization ability of AEF features in scenarios characterized by high
785 variability and heavy-tailed distributions (Table S5). Following the introduction of a single-fold of
786 virtual samples, the R² rose to 0.039. Upon an 8-fold expansion, the RF R² reached 0.641
787 (RMSE=21.05 mg L⁻¹), at a 10-fold expansion, it further improved to 0.674 (RMSE=19.86 mg
788 L⁻¹). The quantum-enhanced RF slightly outperformed the standard RF at high expansion levels,
789 indicating that quantum encoding offers certain advantages in mitigating the influence of extreme
790 values and enhancing model robustness (Fig. 14B). The prediction distribution plots indicate that
791 the initial model failed entirely to identify the high-concentration clustering characteristics of NO₃⁻
792 during the dry season. After data augmentation, the predicted boxplots progressively covered the
793 true high-value intervals, and the trend of tail extension gradually aligned with the observed data.

794 In the wet season, although modeling with the original data still yielded a negative R², the
795 performance improvement was the most rapid (Table S6). An R² of 0.5 was achieved with only a
796 2-fold expansion of virtual samples. At 5-fold expansion, it reached 0.685, and after a 10-fold
797 expansion, the RF R² stabilized at 0.784 (RMSE=8.27 mg L⁻¹), while the quantum-enhanced RF
798 reached 0.781. The boxplots show that the predicted values, initially severely dispersed and
799 systematically biased, rapidly converged to the dense intervals of the observed values, with the
800 final IQR and whisker ranges showing a high degree of overlap.

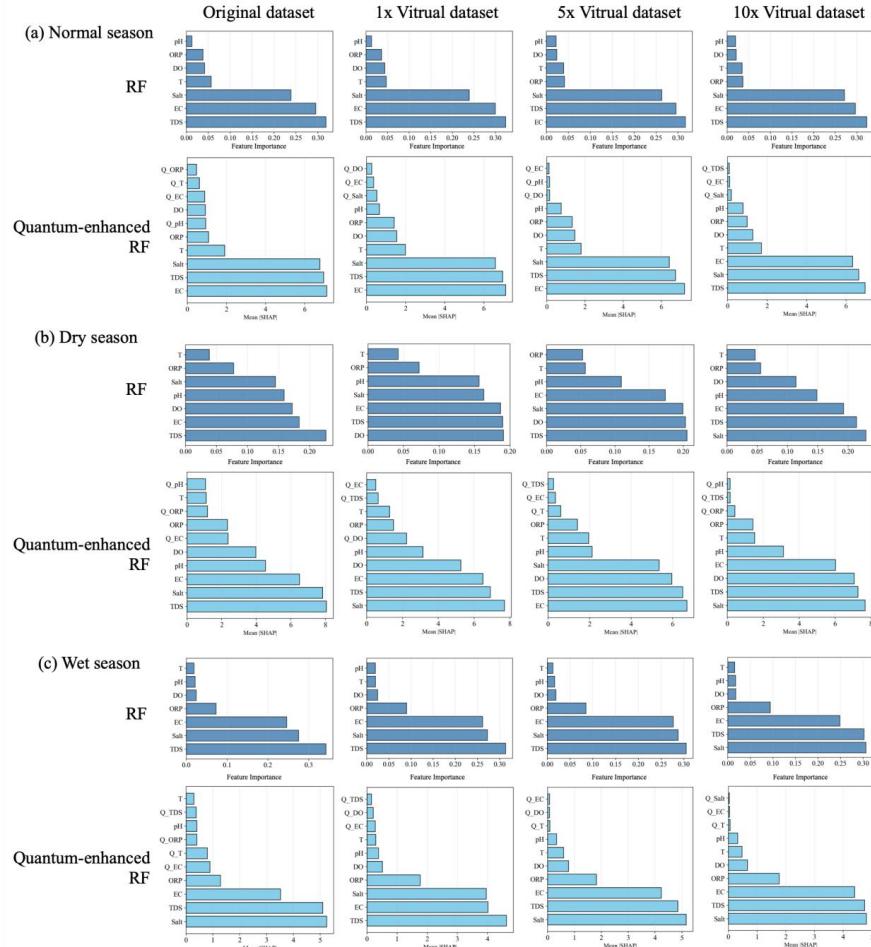


801 Compared to modeling results based on in-situ observation data, the predictive performance
802 based on AlphaEarth Foundation embedding features was generally lower. Under the same virtual
803 sample augmentation multiplier, the maximum R^2 for the normal, dry, and wet seasons were
804 approximately 10.27%, 17.37%, and 19.33% lower, respectively. This indicates that measured
805 water quality parameters more directly reflect the key processes of nitrogen migration and
806 transformation. However, given that AEF can be obtained globally without the need for field
807 sampling, it offers a feasible alternative for the rapid screening of groundwater nitrate risks in
808 large-scale unmonitored areas.

809

810 3.5 Feature importance analysis

811 Fig.15 illustrates the feature importance rankings of the RF and quantum-enhanced RF
812 models when using in-situ measured water quality parameters as inputs across different
813 hydrological seasons. The dominant predictive factors vary across different seasons, and the
814 virtual sample augmentation strategy influences both the stability of feature importance and model
815 performance. There are distinct differences in the key driving factors for each season, which aligns
816 with the results of the Bayesian models and correlation analysis. In the normal season, TDS, EC,
817 Salt, and DO are the most important predictive variables, with their importance significantly
818 higher than that of other parameters. In the dry season, TDS, EC, Salt, and pH exhibit the highest
819 importance. In the wet season, the importance of TDS, EC, Salt, and ORP is most prominent. With
820 the increase in the number of virtual samples, the ranking of feature importance tends to stabilize.
821 For instance, in the dry season, when the sample size increased from the original 65 to 715, the
822 importance of TDS and EC continued to rise and eventually stabilized. Comparing the RF and
823 quantum-enhanced models, quantum enhancement did not fundamentally alter the ranking of
824 feature importance; however, it slightly increased the importance of certain variables or made
825 them more stable, demonstrating the effectiveness of quantum feature encoding as a means of
826 information enhancement.



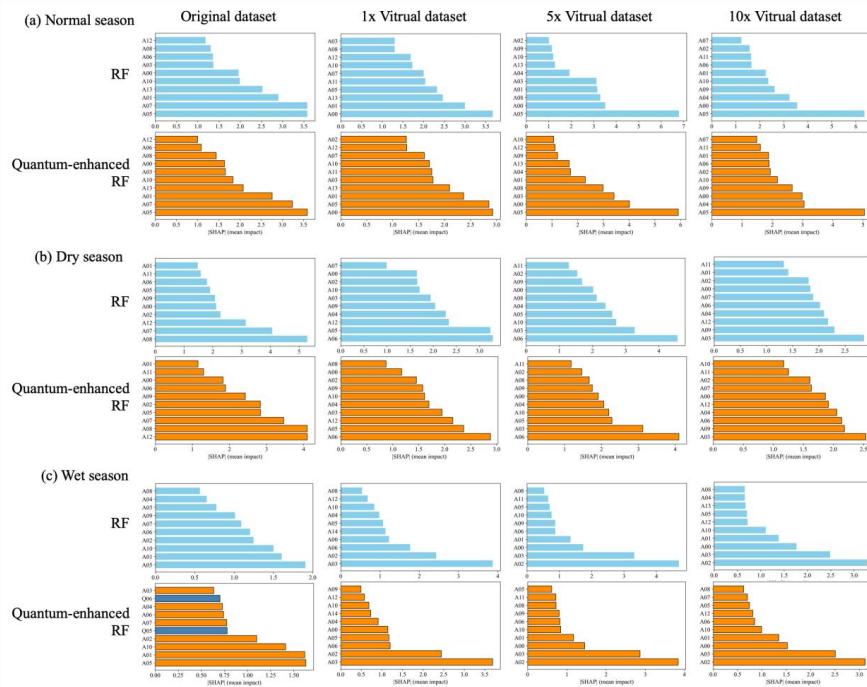
827

828 **Fig.15.** Input feature importance of classical and quantum-enhanced Random
 829 nitrate prediction under different data augmentation strategies: based on Gini index and SHAP
 830 Model.

831 [Fig.16](#) presents the feature importance of the RF and quantum-enhanced RF models when
 832 using only the 64-dimensional AEF semantic embedding vectors as inputs. Since the AEF features
 833 themselves are highly abstract, we cannot assign them specific physical meanings; however, we
 834 can infer which remote sensing semantic information is critical for predicting nitrate concentration
 835 through their importance rankings. Compared to in-situ measured parameters, the importance of
 836 AEF features fluctuates significantly more across different seasons and data volumes, lacking a
 837 consistent core feature set. This reflects that although AEF embeddings contain rich



838 environmental semantic information, their direct correlation with groundwater nitrate
 839 concentration is relatively weak, necessitating the learning of large amounts of data to establish a
 840 robust mapping relationship. In the normal season, features such as A05, A07, and A00 exhibit
 841 relatively high importance. These features may encode seasonal information related to land use
 842 types, soil moisture, or vegetation cover. In the dry season, features such as A08, A06, and A05
 843 are relatively more important. These features may be associated with surface dryness, bare surface
 844 area, or the intensity of human activity, correlating with the spatial distribution of
 845 high-concentration NO_3^- pollution sources during the dry season. In the wet season, the
 846 importance of features like A02, A03, and A05 is prominent. These features may be related to
 847 surface runoff, vegetation growth status, or soil water content, reflecting the driving effect of
 848 rainfall on pollutant migration during the wet season.



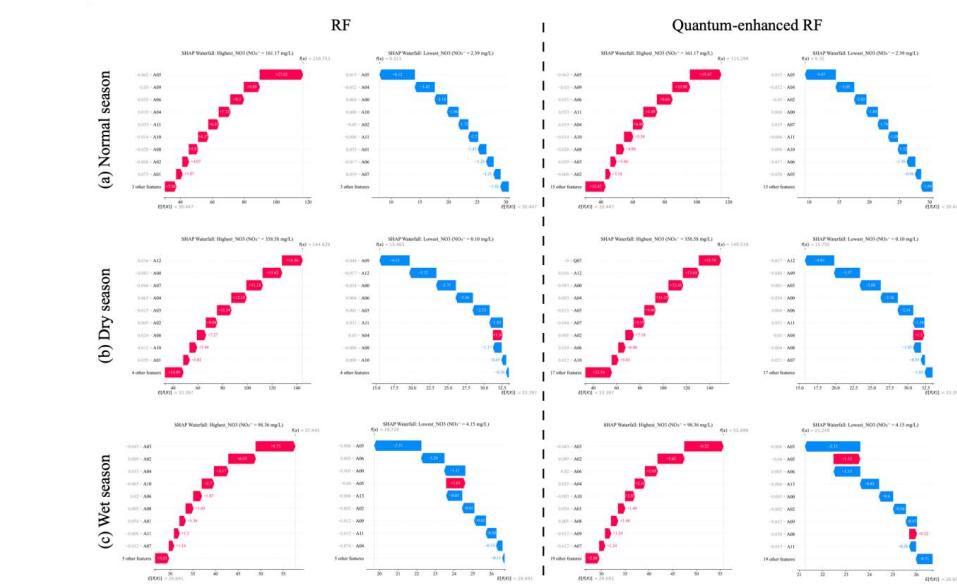
849
 850 **Fig.16.** Importance of AEF input features in seasonal nitrate prediction using classical and
 851 quantum enhanced Random Forest under different data augmentation strategies.

852 The introduction of virtual samples is crucial for stabilizing the importance of AEF features.
 853 On the original dataset, the feature importance ranking was chaotic and unstable; as virtual
 854 samples increased, the ranking gradually became clearer, and the importance of certain core



855 features was highlighted. This once again demonstrates the effectiveness of the virtual sample
856 generation strategy for small-sample modeling. When using AEF features, the feature importance
857 distribution of the quantum-enhanced model is similar to that of the RF model, but it occasionally
858 assigns slightly higher weights to certain features. This suggests that quantum feature encoding
859 may assist the model in extracting more discriminative information from the high-dimensional,
860 complex remote sensing semantic space, thereby slightly optimizing the feature selection process.

861 [Fig.17](#) presents a local feature attribution analysis for representative samples predicting the
862 highest and lowest NO_3^- concentrations using SHAP waterfall plots. Regardless of whether the
863 classical RF or the quantum-enhanced RF model is used, samples predicting high NO_3^-
864 concentrations are driven by a set of features with positive contributions (red bars). In the normal
865 season, for the highest NO_3^- sample with a predicted value of 161.17 mg L^{-1} , features A05, A09,
866 and A06 contributed the highest positive values, with A05 making the largest contribution and
867 serving as the key factor driving the prediction to a high level. In the dry season, for the sample
868 with a predicted value as high as 358.58 mg L^{-1} , features A12, A00, and A07 were the main
869 positive driving factors, with A12 contributing most prominently. In the wet season, for the
870 sample with a predicted value of 98.36 mg L^{-1} , features A03, A02, and A04 provided the main
871 positive contributions, with A03 contributing the most. For samples predicting low NO_3^-
872 concentrations, model decisions mainly rely on features with negative contributions (blue bars).
873 The role of these features is to pull the predicted value down from the baseline ($E[f(X)]$). In the
874 normal season, for the lowest NO_3^- sample with a predicted value of 2.39 mg L^{-1} , features A05,
875 A04, and A00 exhibited strong negative contributions, with A05 showing the largest negative
876 contribution. In the dry season, for the sample with a predicted value of only 0.10 mg L^{-1} , features
877 A09, A12, and A06 were the main negative driving factors, with A09 contributing the most
878 negatively. In the wet season, for the sample with a predicted value of 4.15 mg L^{-1} , features A03,
879 A06, and A00 provided the main negative contributions, with A03 contributing the most
880 negatively.



881

882 **Fig.17.** SHAP waterfall-based feature attribution comparison between classical and
 883 quantum-enhanced Random Forest across different seasons.

884

885 4. Discussion

886 4.1 Nitrogen sources, migration, and transformation

887 The Piper diagram indicates that the hydrochemical type is predominantly Ca-Mg-HCO₃⁻.
 888 The Gibbs diagram and ion ratios confirm that the hydrochemical background is dominated by
 889 carbonate rock dissolution, with weak cation exchange. Concentrated precipitation during the
 890 rainy season leads to dilution and infiltration, reducing the NO₃⁻ concentration to 27.14 mg L⁻¹.
 891 This indicates that surface manure leaches into the groundwater with rainfall (Sun et al., 2024).
 892 During this period, cation exchange is enhanced, improving the aquifer's temporary retention
 893 capacity for NO₃⁻ (Wang et al., 2025). Isotopic evidence and MixSIAR source apportionment
 894 consistently indicate that the primary sources of current nitrate pollution are domestic sewage and
 895 manure (DSM, 74.1%) and soil organic nitrogen (SON, 20.9%), whereas the contributions from
 896 chemical fertilizers and precipitation are minimal. This suggests that in the study area, the direct
 897 leaching of fertilizer nitrogen is not the dominant pathway, rather, fertilizer nitrogen remains in the
 898 soil-vadose zone and enters groundwater through long-term water drive (Wang et al., 2025).
 899 Given that the vadose zone thickness in the North China Plain generally exceeds 10 m, the



900 currently elevated NO_3^- levels are more likely derived from historical fertilizer residues and the
901 long-term infiltration of manure, particularly as the farmlands in the study area are situated near
902 rural residential areas (Wu et al., 2024). The mean $\delta^{15}\text{N}$ - NO_3^- values range from 12.2‰ to 18.2‰,
903 which far exceeds the typical range for chemical fertilizers but closely matches that of manure and
904 soil organic nitrogen. This confirms that the nitrogen has undergone microbial mineralization and
905 nitrification processes (Li et al., 2022). SON is converted to NO_3^- through ammonification
906 followed by nitrification under aerobic conditions, while ammonium from DSM also enters the
907 groundwater via nitrification (Liu et al., 2023).

908 The seasonal variation in nitrate concentrations is essentially driven by scarce precipitation
909 and strong evaporation during the dry season. This leads to a decline in groundwater levels and a
910 reduction in flow velocity, creating a positive migration potential gradient. Consequently, NO_3^-
911 accumulates in the discharge areas along with the groundwater flow. Furthermore, cation
912 exchange is inhibited, weakened Na^+ adsorption and relative Ca^{2+} depletion indicate a decrease in
913 the aquifer's retention capacity for NO_3^- , making accumulation the dominant process (Ahmed et al.,
914 2013). In contrast, concentrated precipitation during the wet season triggers rapid infiltration,
915 raising groundwater levels and increasing flow velocity, during which cation exchange becomes
916 active (Zhang et al., 2023). The groundwater is generally oxidizing, as evidenced by the extremely
917 low concentrations of NO_2^- ($<0.11 \text{ mg L}^{-1}$) and NH_4^+ ($<0.16 \text{ mg L}^{-1}$). This indicates that the
918 majority of the area is an oxidative environment conducive to the stable existence of NO_3^- . The
919 $\delta^{18}\text{O}$ - NO_3^- values range from -9.58‰ to 8.04‰, falling within the typical nitrification interval.
920 This excludes significant denitrification, confirming that the transformation process is dominated
921 by nitrification while denitrification is limited (Zhang et al., 2025).

922 |
923 4.2 Virtual sample generation effectively mitigates small-sample bias and reveals the model's
924 sensitive response to seasonal heterogeneity

925 Model overfitting and insufficient generalization resulting from small-sample data are
926 prevalent challenges in the field of environmental forecasting (Zhu et al., 2023). The
927 t-SNE-GMM-KNN virtual sample generation strategy proposed in this study demonstrates that the
928 generated virtual samples are highly consistent with the original data in terms of statistical
929 characteristics, such as mean, standard deviation, and coefficient of variation, and successfully



930 reproduce hydrochemical differences across different seasons. The substantial improvement in
931 model performance following virtual sample expansion clearly confirms that data sparsity, rather
932 than insufficient model capacity, is the core bottleneck in seasonal nitrate modeling (Saha et al.,
933 2023). Furthermore, even with the incorporation of generated virtual samples, the magnitude of
934 prediction performance gains exhibits seasonal divergence. During the dry season, characterized
935 by highly right-skewed NO_3^- concentrations, the model benefits most significantly. With 10-fold
936 expansion, the R^2 value surges from 0.28 to over 0.85. Conversely, in the wet season, although the
937 absolute performance gain is smaller due to dominant dilution effects and low concentrations,
938 excellent predictive accuracy is still achieved. This phenomenon aligns with fundamental
939 hydrological principles: strong evaporation and concentration during the dry season intensify the
940 spatial heterogeneity of pollutant accumulation and process nonlinearity, necessitating richer
941 samples to characterize tail behaviors (Li et al., 2025). In contrast, the dilution effects caused by
942 rainfall leaching during the wet season tend to homogenize the system, reducing its dependency
943 on sample size (Bigler et al., 2024). The t-SNE–GMM–KNN strategy proposed in this study
944 outperforms traditional oversampling methods (e.g., SMOTE) or deep generative models (e.g.,
945 VAE) in preserving multimodal structures and heavy-tailed covariance; the latter often ignore the
946 manifold geometric properties of high-dimensional geochemical spaces or inherently rely on large
947 amounts of training data (Udu et al., 2025), which is precisely what is lacking in the scenario of
948 this study. Compared to common methods such as Gaussian Mixture Models (GMM) and
949 Generative Adversarial Networks (GANs), the core advantages of this strategy are reflected in
950 three aspects: first, t-SNE dimensionality reduction accurately captures sample clustering
951 structures driven by different hydrological processes, providing a reliable foundation for
952 subsequent distribution modeling; second, the number of GMM clusters is automatically
953 optimized based on the Bayesian Information Criterion (BIC), avoiding biases arising from
954 subjective settings; and third, KNN inverse mapping enables reconstruction from low-dimensional
955 to high-dimensional space without the need for large-scale training data, making it more suitable
956 for small-sample scenarios (Silva et al., 2023; Kurniawan et al., 2024; Peng et al., 2025).

957
958 4.3 Performance analysis of hybrid quantum-classical model

959 Quantum Machine Learning offers a novel approach to capturing complex non-linear



960 relationships through feature mapping in high-dimensional quantum Hilbert spaces. The hybrid
961 quantum-classical Random Forest yields slight performance improvements in scenarios where
962 original data is scarce or the distribution is highly skewed (Lamichhane et al., 2025). When
963 classical feature representation capacity approaches saturation, the Z-feature mapping based on
964 Parameterized Quantum Circuits (PQC) can expose entangled non-linear patterns in the Hilbert
965 space, thereby enhancing feature discriminability. The gain from this enhancement tends to
966 converge after sufficient virtual sample expansion. In this study, quantum features were generated
967 by analytically calculating the Pauli-Z expectation value $\langle Z \rangle$, completely circumventing hardware
968 noise interference associated with quantum sampling. This renders the quantum-enhanced RF
969 practically feasible for small-sample environmental tasks. However, the performance
970 improvement of the quantum-enhanced RF is not absolute; in scenarios with high data quality and
971 significant linear relationships during the wet season, the marginal gain of quantum features is
972 limited. Conversely, in the dry season, characterized by sparse data and numerous extreme values,
973 quantum encoding demonstrates stronger stability by reducing measurement noise interference
974 (Ranga et al., 2024). This phenomenon indicates that the advantages of hybrid quantum-classical
975 modeling are concentrated in scenarios with data complexity and limited information. Its essence
976 lies in expanding the model's representational capacity through quantum feature enhancement,
977 rather than replacing the core logic of classical models. This exploration verifies the potential
978 value of quantum machine learning in addressing small-sample problems in earth sciences, even if
979 its absolute advantage may not be as pronounced as in pure quantum algorithms (Adhikari, 2022).
980

981 4.4 Potential and limitations of using AlphaEarth Foundation Embeddings for large-scale
982 monitoring

983 Modeling performance using only AEF embeddings as input generally yields an R^2
984 approximately 10-20% lower than that achieved using measured water quality parameters. The
985 core reason for this discrepancy lies in the fact that water quality parameters directly reflect the
986 immediate state of the groundwater chemical environment and are directly related to nitrate
987 transport and transformation processes, whereas surface remote sensing semantics provide only an
988 indirect characterization (Alam et al., 2025). After 10-fold virtual expansion, the AEF model still
989 achieves R^2 values of >0.67 in the dry season, >0.85 in the normal season, and >0.78 in the wet



990 season, proving its feasibility as a rapid large-scale screening tool, particularly in unmonitored
991 areas. Seasonal shifts in feature importance (dominated by A05/A00 in the normal season,
992 A08/A06 in the dry season, A02/A03 in the wet season) suggest potential physical interpretations.
993 A05/A00 may encode crop residue or soil organic matter information, A08/A06 may characterize
994 the degree of bare soil exposure, and A02/A03 may reflect vegetation growth status or surface
995 runoff potential. These inferences align highly with MixSIAR source apportionment and Bayesian
996 driving factors. Although causal inference remains indirect, the global coverage and annual update
997 characteristics of AEF make it a powerful supplement rather than a substitute for large-scale
998 monitoring.

999

1000 5. Conclusions

1001 This study develops an integrated prediction framework combining hybrid quantum-classical
1002 machine learning, advanced virtual sample augmentation (t-SNE–GMM–KNN), and remote
1003 sensing foundation model embedding (AlphaEarth Foundation, AEF). The framework is designed
1004 to systematically address three core challenges in predicting groundwater nitrate concentrations in
1005 agricultural areas across different hydrological seasons: small sample bias, seasonal heterogeneity,
1006 and input data scarcity.

1007 Hydrological seasonality acts as the dominant controlling factor for the spatiotemporal
1008 variability of nitrates. Nitrate concentrations peak during the dry season (mean: 42.93 mg L⁻¹),
1009 driven primarily by evaporative concentration and pollutant accumulation effects. In contrast,
1010 concentrations reach a minimum in the wet season (mean: 27.14 mg L⁻¹) due to dilution by
1011 precipitation. The groundwater hydrochemical type is consistently Ca-Mg-HCO₃⁻ across all
1012 seasons, controlled predominantly by carbonate mineral dissolution. TDS, EC, and salinity remain
1013 consistently top-ranked across all seasons, with additional season-specific drivers including Mg²⁺
1014 and Na⁺ (normal season), SO₄²⁻ (dry season), and Cl⁻ (wet season). Stable hydrogen and oxygen
1015 isotope analysis reveals strong evaporative fractionation of groundwater. MixSIAR analysis
1016 quantitatively apportioned nitrate sources: domestic sewage and manure (DSM) contribute 74.1%,
1017 soil organic nitrogen (SON) 20.9%, while synthetic fertilizers (NHF+NOF=4.8%) and
1018 atmospheric deposition (0.2%) are negligible, strongly indicating that legacy nitrogen stored in the
1019 thick vadose zone, rather than in-season fertilizer leaching, sustains current pollution.



1020 The proposed t-SNE-GMM-KNN virtual sample strategy effectively alleviates the bottleneck
1021 associated with small-sample modeling. By preserving the nonlinear manifold structure and
1022 multimodal distribution characteristics of the high-dimensional hydrochemical space, this method
1023 significantly enhances the model's ability to fit heavy-tailed distributions. Model performance
1024 improves significantly with virtual sample expansion. Using measured parameters as inputs, a
1025 10-fold augmentation increased the coefficient of determination (R^2) for the dry season from 0.284
1026 to >0.85, while stabilizing it at >0.95 for the normal and wet seasons. This confirms that data
1027 sparsity is the fundamental constraint limiting performance. Although performance gains are
1028 limited with high-quality data, the quantum-enhanced Random Forest demonstrates superior
1029 stability compared to classical models in small-sample, highly skewed scenarios, validating the
1030 feasibility and value of quantum feature enhancement strategies in environmental small-sample
1031 learning. The overall prediction performance using measured hydrochemical parameters surpasses
1032 that of AEF remote sensing semantic embeddings (R^2 is approximately 10-20% higher), as the
1033 former directly reflects subsurface nitrogen migration and transformation processes. Following
1034 10-fold virtual sample augmentation, the AEF model also achieves usable accuracy, with feature
1035 importance exhibiting seasonal shifts.

1036

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1040

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1049



1050

1051 Competing interests

1052 The contact author has declared that none of the authors has any competing interests.

1053

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