

Responses to Reviewer #1

Note: The reviewers' comments are in black, and our responses are in blue. The changes made in the revised manuscript are highlighted in green. Figures in the manuscript are labeled using Arabic numerals (e.g., Fig. 1, Figure 1), and those in our responses to the reviewers are labeled using Arabic numerals with the prefix "R" (e.g., Fig. R1, Figure R1).

Reviewer #1:

The study investigates the regional differences (q) in the local dissipation efficiency of internal tides between the Luzon Strait and the Nansha Islands based on the LLC4320 simulation outputs. A physically-based empirical formulation and an XGBoost model are subsequently developed to parameterize q . This study contributes much to improving the understanding and parameterization of internal tide dissipation in the SCS. However, the manuscript still has some shortcomings, and the following points should be addressed further by the authors before publication.

Responses: We sincerely thank the reviewer for the valuable suggestions on our manuscript. We have carefully considered and fully incorporated all suggestions into the revised manuscript. We hope that these improvements will render the manuscript more refined and satisfy the reviewer.

1) Fig. 1: Since q is calculated from area-integrated conversion and dissipation, its magnitude may be sensitive to the choice of regional boundaries. Could the authors clarify the criteria used to determine the boundaries of the Luzon Strait and the Nansha Islands?

Responses: We are sorry for this information gap. The regional boundaries were determined according to the major topographic features and existing studies on internal tidal energetics. Specifically, the Luzon Strait domain is defined to encompass the Hengchun Ridge, the Lanyu Ridge, and the trough between them, which collectively constitute the well-known double-ridge where intense internal tides are generated (Buijsman et al., 2014; Tian et al., 2025; Wang et al., 2016). The Nansha Islands domain is selected to encompass the dominant topographic features (islands and reefs), which are located in the propagation paths of internal tides coming from the Luzon Strait (Tian et al., 2025; Xu et al., 2016). The interaction of the incoming internal tides with the rough islands and reefs promotes energy transfer to higher modes and hence enhances local dissipation (Vic et al., 2019). As a result, these two domains are chosen to represent a typical internal tide source region and a sink region, respectively. In the revised manuscript, we added a concise description about the choosing of these two domains.

References:

- Buijsman, M. C., Klymak, J. M., Legg, S., Alford, M. H., Farmer, D., MacKinnon, J. A., Nash, J. D., Park, J.-H., Pickering, A., and Simmons, H.: Three-dimensional double-ridge internal tide resonance in Luzon Strait, *J. Phys. Oceanogr.*, 44, 850–869, <https://doi.org/10.1175/JPO-D-13-024.1>, 2014.
- Tian, Y., Bai, X., Wang, C., and Liu, Z.: Tidal energetics in the eddying South China Sea from a high-resolution numerical simulation, *Prog. Oceanogr.*, 231, 103418, <https://doi.org/10.1016/j.pcean.2025.103418>, 2025.
- Vic, C., Naveira Garabato, A. C., Green, J. A. M., Waterhouse, A. F., Zhao, Z., Melet, A., de Lavergne, C., Buijsman, M. C., and Stephenson, G. R.: Deep-ocean mixing driven by small-scale internal tides, *Nat. Commun.*, 10, 2099, <https://doi.org/10.1038/s41467-019-10149-5>, 2019.

Wang, X., Peng, S., Liu, Z., Huang, R. X., Qian, Y., and Li, Y.: Tidal mixing in the South China Sea: an estimate based on the internal tide energetics, *J. Phys. Oceanogr.*, 46, 107–124, <https://doi.org/10.1175/JPO-D-15-0082.1>, 2016.

Xu, Z., Liu, K., Yin, B., Zhao, Z., Wang, Y., and Li, Q.: Long-range propagation and associated variability of internal tides in the South China Sea, *J. Geophys. Res. Oceans*, 121, 8268–8286, <https://doi.org/10.1002/2016JC012105>, 2016.

(Changes can be found in Lines 111–116 of the revised manuscript.)

2) Lines 100–108: The manuscript only retains the first five baroclinic modes in the modal decomposition. Please state why five modes are sufficient for the analysis presented in this study. Given that high-mode internal tides are significant and important over steep and rough topography, it would be helpful to either provide the proportion of total baroclinic kinetic energy captured by modes 1–5, or discuss the potential influence of unresolved higher modes on the diagnosed local dissipation.

Responses: We thank the reviewer for pointing out this information gap. As the reviewer suggested, we calculate the proportion of total depth-integrated baroclinic kinetic energy captured by the first five modes (modes 1–5) for the winter and summer months analyzed in this study. For each tidal band in each region, this proportion is calculated as the ratio of the area-integrated value of depth-integrated kinetic energy from modes 1–5 to the area-integrated value of total depth-integrated baroclinic kinetic energy. The numerator is calculated from the modal decomposition result (i.e., kinetic energy of modes 1–5), whereas the denominator is calculated directly from the full baroclinic velocity field before modal decomposition. The results show that in the Luzon Strait region and the Nansha Islands, modes 1–5 account for 85% and 89% of the total baroclinic kinetic energy, 87% and 99% of the diurnal baroclinic kinetic energy, and more than 98% of the semidiurnal baroclinic kinetic energy. These results indicate that the first five modes almost make up entire baroclinic kinetic energy during the chosen periods. We have incorporated these results into the *Methods* section of the revised manuscript. The potential impact of unresolved higher modes is also given in the revised manuscript. Modes 1–5 capture the dominant part of total baroclinic kinetic energy, although unresolved higher modes may still contribute less to local dissipation of internal tides over rough topography (Vic et al., 2019).

(Changes can be found in Lines 132–137 of the revised manuscript.)

3) Lines 40–44 and 122–125: q is defined as the fraction of locally dissipated internal tidal energy relative to locally generated internal tidal energy, which would conventionally be expected to fall between 0 and 1. However, subsequent results presented in the manuscript report $q > 1$ in the Nansha Islands and negative q values for some individual modes in the Luzon Strait. Please provide the definition and sign convention of q , and explicitly explain the physical meanings of $q > 1$ and $q < 0$ in the present energy-budget framework.

Responses: We are sorry for this confusion. The local dissipation efficiency, q , is a decisive dimensionless parameter defined as the fraction of locally generated internal tidal energy that dissipates near the generation site, with the remainder radiating away as low-mode internal tides; thus, q is restricted between 0 and 1 (St. Laurent et al., 2002). However, Wang et al. (2016) found that during summer the total local dissipation accounts for 267% of the local conversion in the broader SCS, in contrast to 38% in the LS. Tian et al. (2025) reported the corresponding values of about 203% in the broader SCS and 24% in the LS during summer. In such sink regions (such as the SCS), the local dissipation is furnished by local generation and incoming internal tidal energy from distant sources. Therefore, we define the local dissipation

ratio of internal tides, q^* , as the ratio of the area-integrated local dissipation to the area-integrated barotropic-to-baroclinic energy conversion, i.e., $q^* = \int_s ds \langle DIS_{bc} \rangle / \int_s ds \langle Conv \rangle$. Here, $\langle DIS_{bc} \rangle$ is inferred as $\langle DIS_{bc} \rangle = \langle Conv \rangle - \langle \nabla_h \cdot \mathbf{F} \rangle$, where $\langle \cdot \rangle$ denotes the tidal-period mean, $Conv$ is the barotropic-to-baroclinic energy conversion, and $\nabla_h \cdot \mathbf{F}$ is the horizontal divergence of the depth-integrated baroclinic energy flux. In this way, $q^* > 1$ occurs in the Nansha Islands when the area-integrated divergence of the baroclinic energy flux is negative, indicating a convergence of remotely generated internal tidal energy into the region. In this case, the inferred local dissipation exceeds the barotropic-to-baroclinic energy conversion. It can be seen that $q^* > 1$ does not imply that more than 100% of the locally generated internal tidal energy must be dissipated locally; instead, it means that the local dissipation is furnished by both the local barotropic-to-baroclinic energy conversion and the convergence of incoming internal tidal energy from the far field. For individual vertical modes, $q^* < 0$ occurs when the area-integrated divergence of the baroclinic energy flux of this mode exceeds its local barotropic-to-baroclinic energy conversion, yielding a negative inferred dissipation. The divergence of the baroclinic energy flux is furnished by both the local conversion and the net energy transfer from other vertical modes, allowing the divergence of the baroclinic energy flux to exceed the local barotropic-to-baroclinic energy conversion.

Reference:

St. Laurent, L. C., Simmons, H. L., and Jayne, S. R.: Estimating tidally driven mixing in the deep ocean, *Geophys. Res. Lett.*, 29, 2106, <https://doi.org/10.1029/2002GL015633>, 2002.

(Changes can be found in Lines 148–150 and 169–175 of the revised manuscript.)

4) Lines 90–92 and 125–127: The LLC4320 simulation outputs used in this study only span approximately 14 months, from September 2011 to November 2012. In the analysis, January and August are used to represent winter and summer, respectively. As a result, the summer-winter differences reported in this manuscript are based on a single annual cycle rather than a multiyear climatology. Could the authors point out this limitation explicitly and avoid interpreting these differences as climatological seasonal variations?

Responses: We thank the reviewer for this insightful comment. As the reviewer suggested, in the revised manuscript, January and August 2012 are now referred to as winter and summer months in the specifically simulated year, and the differences between them are interpreted as those between two individual months rather than as climatological seasonal variations.

(Changes can be found in Lines 176, 211–213 and 265–266 of the revised manuscript.)

5) Figs. 2–5: Since the first and final months of the dataset are incomplete, could the authors specify how the monthly mean values of q for these two months were treated in the time-series plots? If these mean values were calculated from incomplete monthly records, this should be stated clearly; alternatively, the corresponding data points can be removed or marked separately.

Responses: We are sorry for this information gap. In the original manuscript, the monthly mean values of q^* for September 2011 and November 2012 were calculated using available days within these two months. Now the information about these two months has been removed from the time-series plots in Figs. 2–5 of the revised manuscript. The *Methods* section has also been revised accordingly to clarify that incomplete months are excluded when computing the monthly mean values of q^* . The revised Figs. 2–4 are provided below as Figs. R1–R3. The revised Fig. 5, which also addresses Comment 6, is presented as Fig. R4 below the response to

Comment 6.

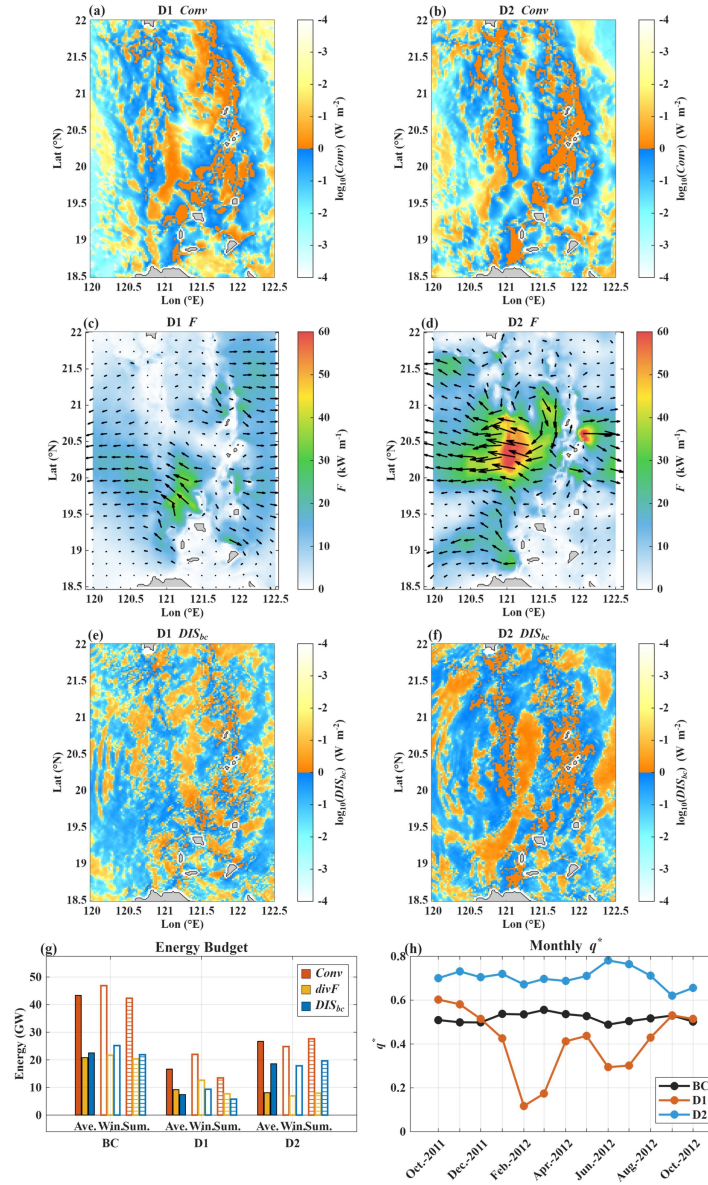


Figure R1. (a–f) Spatial distributions of the diurnal and semidiurnal internal tidal energy budget terms in the LS during summer. $\langle Conv \rangle$ is the tidal-period mean depth-integrated barotropic-to-baroclinic energy conversion, $\langle F \rangle$ is the tidal-period mean depth-integrated internal tidal energy flux, and $\langle DIS_{bc} \rangle$ is the tidal-period mean depth-integrated local internal tidal dissipation. (g) Bar chart of the budget terms. $\langle divF \rangle$ is the tidal-period mean divergence of the depth-integrated internal tidal energy flux. "Ave.", "Win.", and "Sum." stand for the annual-, winter-, and summer-mean values, respectively. BC, D1, and D2 represent the total internal tide, diurnal tide, and semidiurnal tide, respectively. (h) Time series of the monthly-mean local internal tidal dissipation ratio, q^* . In (a–b) and (e–f), the units for $Conv$ and DIS_{bc} are W m^{-2} , with orange (blue) representing positive (negative) values. In (c–d), the unit for F is kW m^{-1} , with colors representing magnitude and arrows indicating direction. In (g), the units for $Conv$, $\langle divF \rangle$, and DIS_{bc} are GW.

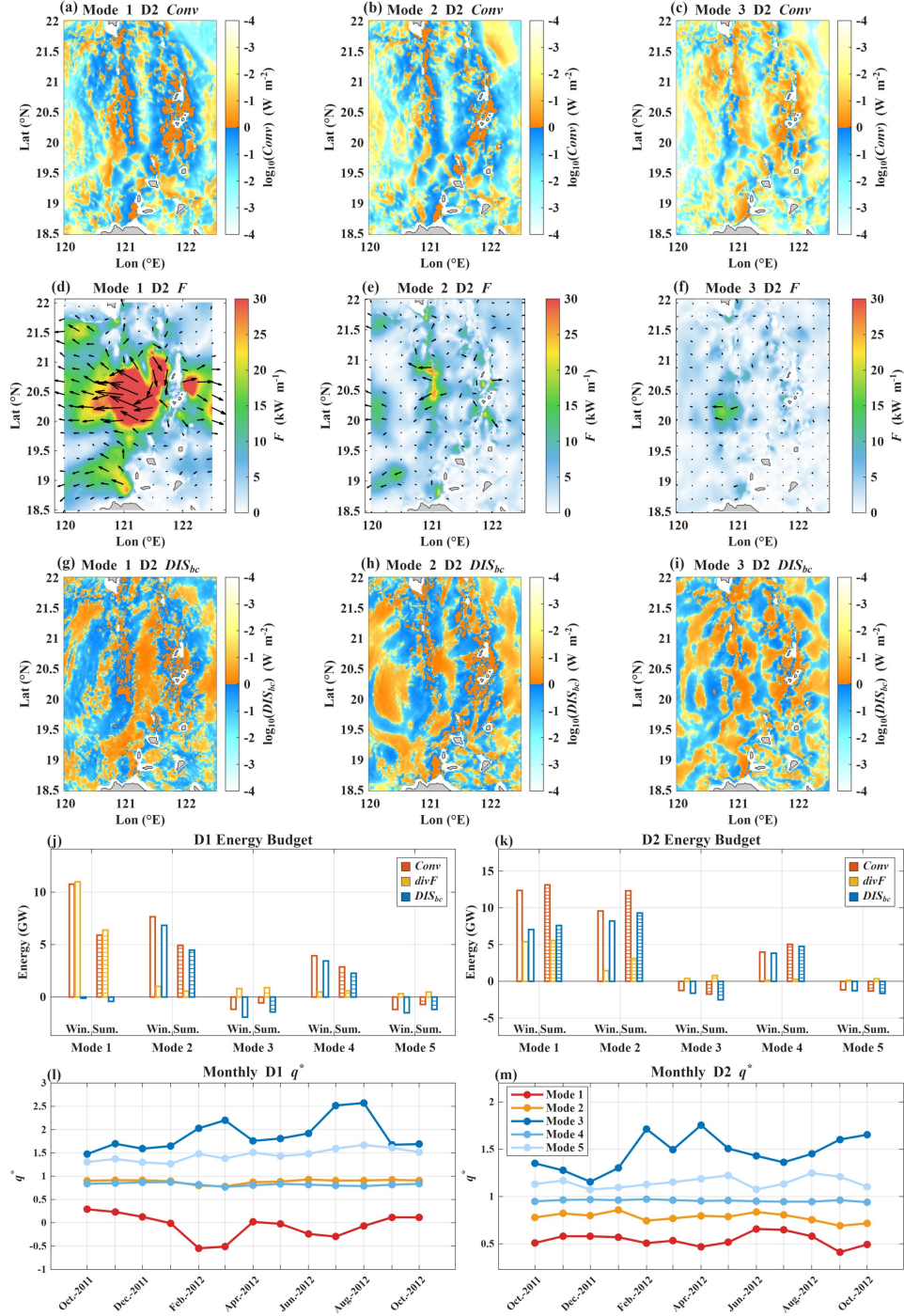


Figure R2. (a–i) Spatial distributions of the semidiurnal internal tidal energy budget terms for modes 1–3 in the LS during summer. (j–k) Bar charts of the energy budget terms for modes 1–5, and (l–m) time series of monthly-mean q^* for modes 1–5 for diurnal (D1) and semidiurnal (D2) tides, respectively. "Win." and "Sum." stand for winter and summer, respectively. In (a–c) and (g–i), the units for Conv and DIS_{bc} are W m⁻², with orange (blue) indicating positive (negative) values. In (d–f), the unit for F is kW m⁻¹, with colors indicating magnitude and arrows indicating direction. In (j–k), the units for Conv , $\text{div}F$, and DIS_{bc} are GW.

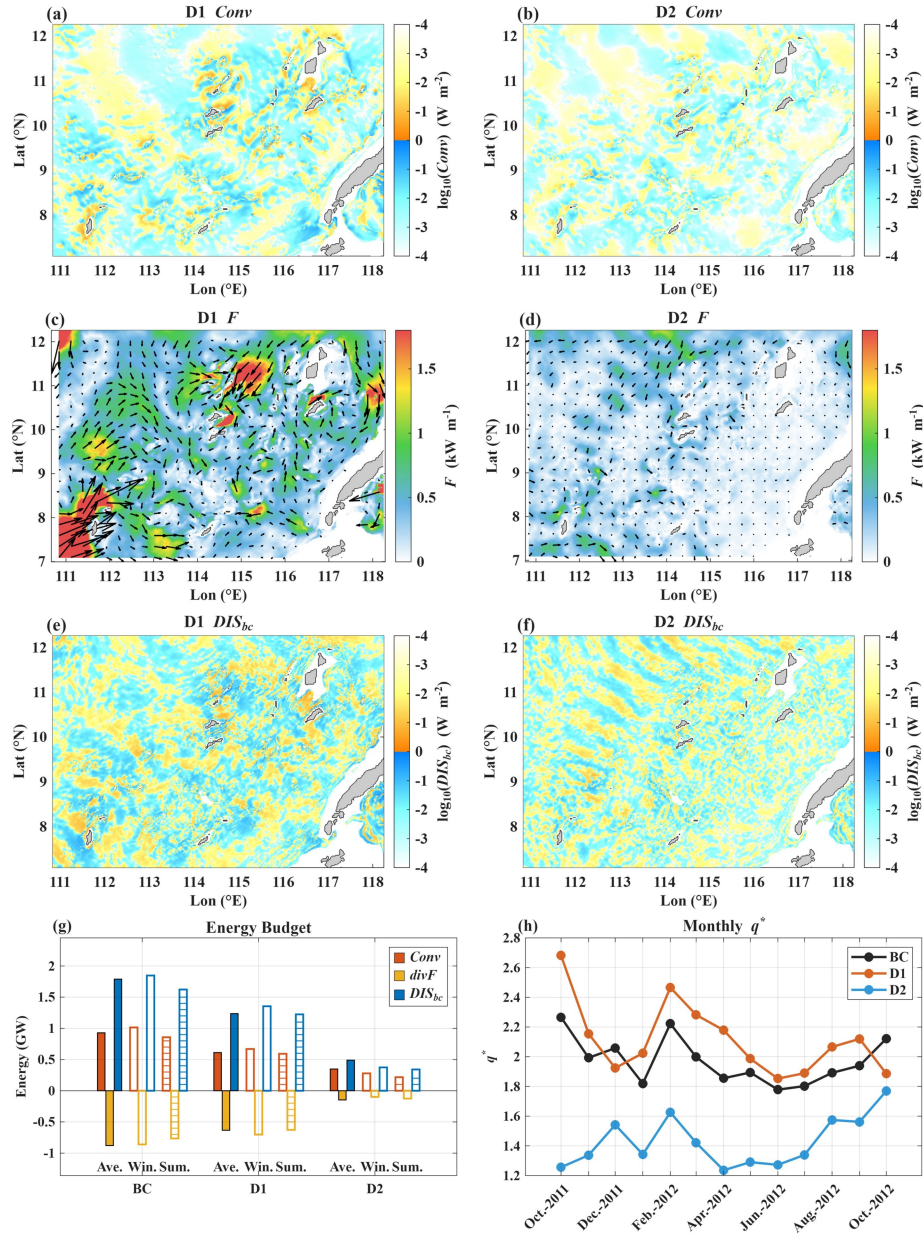


Figure R3. Same as Fig. R1, but for the Nansha Islands.

(Changes can be found in Lines 175–176 of the revised manuscript.)

6) Regarding Fig. 5j–m: These panels appear identical to Fig. 3j–m. Additionally, they are seemingly inconsistent with the results described in Lines 205–213, such as the reduced magnitude of energy-budget terms and the relatively large mode-1 q . Could the authors examine whether the proper data for the Nansha Islands have been plotted?

Responses: We sincerely apologize for this error. The corrected figure (Fig. 5 in the revised manuscript), presented as Fig. R4 below, shows that compared with the Luzon Strait, the Nansha Islands are characterized by smaller energy-budget terms, convergence of low-mode energy, and larger mode-1 q^* values that can exceed 4. In addition, we would like to clarify that when the area-integrated modal conversion is close to zero, several high-mode monthly q^* values can be extremely large. For the clarity of the figure, q^* values falling outside the displayed y-axis ranges

in Fig. R4l–m are not shown, as noted in the revised figure caption.

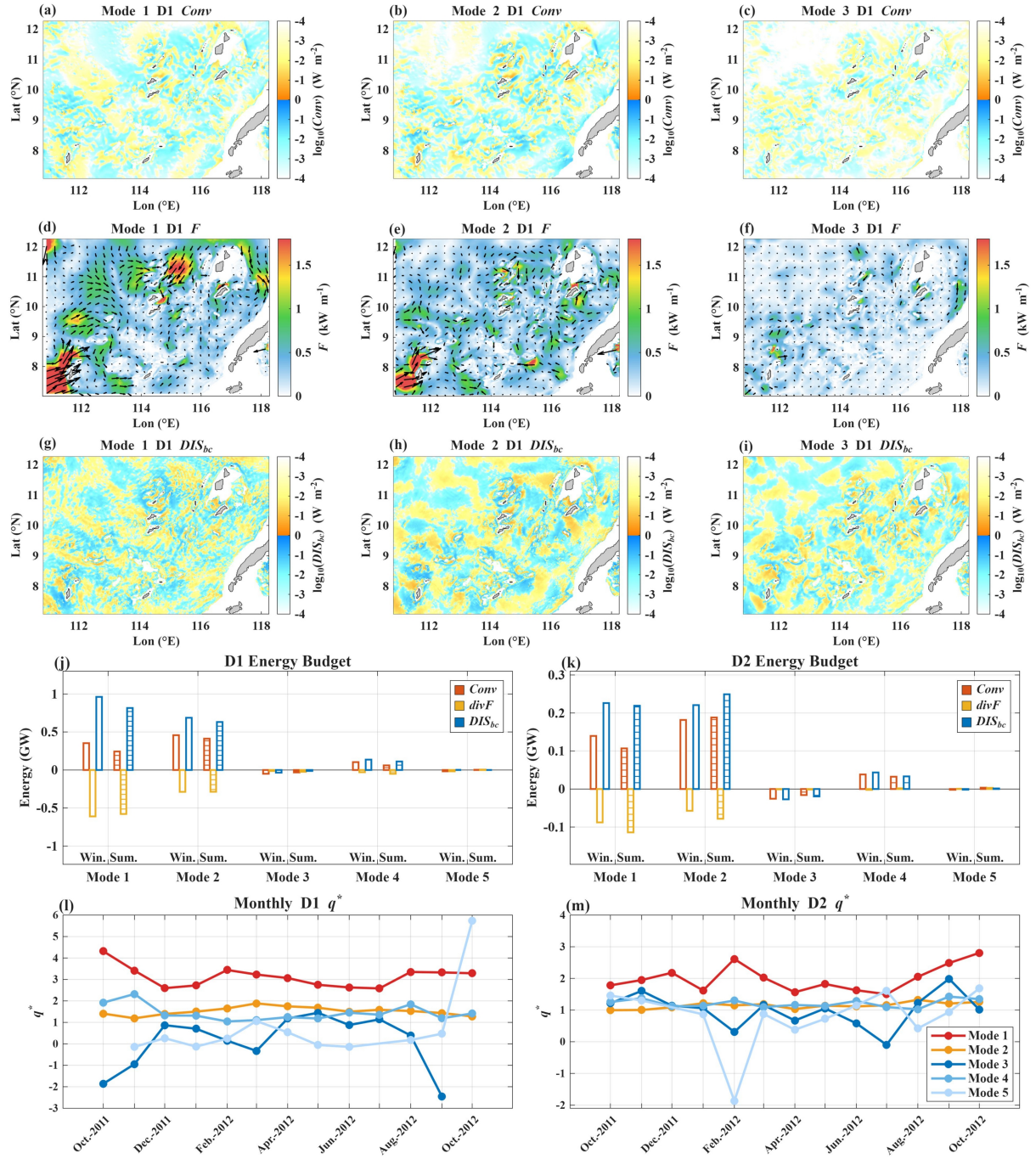


Figure R4. Same as Fig. R2, but for the Nansha Islands and (a–i) for diurnal internal tide. In (l–m), monthly mean values of q^* outside the displayed y-axis ranges are not shown for clarity.

(Changes can be found in Lines 259–267 of the revised manuscript.)

Responses to Reviewer #2

Note: The reviewers' comments are in black, and our responses are in blue. The changes made in the revised manuscript are highlighted in green. Figures in the manuscript are labeled using Arabic numerals (e.g., Fig. 1, Figure 1), and those in our responses to the reviewers are labeled using Arabic numerals with the prefix "R" (e.g., Fig. R1, Figure R1).

Reviewer #2:

General comment:

This manuscript uses hourly LLC4320 output to compare internal-tide energetics in the Luzon Strait, a strong generation region, and the Nansha Islands, a remote energy-sink region. The source-to-sink comparison, separation of diurnal and semidiurnal bands, modal analysis, and attempt to connect the diagnosed energy ratios with both a physics-based regression and XGBoost are useful. In particular, the contrast between strong local conversion in the Luzon Strait and substantial far-field energy import around the Nansha Islands is very interesting for relevant researchers.

The manuscript is generally well organized, and the main results appear publishable. I recommend minor revision, mainly about clarifying $q > 1$, distinguishing energy-budget residuals, testing robustness of modal analysis etc. These issues can largely be addressed using the existing output without redesigning the study.

Responses: We are grateful to the reviewer for the valuable suggestions on our manuscript. All comments have been carefully considered and fully addressed in the revised manuscript. We hope that these improvements will render the manuscript more refined and satisfy the reviewer.

Specific comments:

1. The conventional local dissipation efficiency q is a fraction of locally generated internal-tide energy and is therefore normally interpreted between zero and one. In the Nansha area, the reported quantity exceeds one because dissipation is normalized only by local conversion while the domain also imports energy. This is informative, but it is no longer an efficiency in the usual fractional sense. Maybe consider renaming it throughout as "local dissipation ratio"? Or for the Nansha region, you could report dissipation normalized by the total available energy supply, e.g., positive local conversion plus net incoming baroclinic flux? This would have a clearer interpretation as an efficiency.

Responses: We thank the reviewer for pointing out this issue. The term "local dissipation efficiency q " originates from the internal tidal mixing parameterization proposed by St. Laurent et al. (2002), where q is defined as the fraction of locally generated internal tidal energy that dissipates near the generation site, with the remainder radiating away as low-mode internal tides. It should be noted that this parameterization quantifies only near-field dissipation driven by high-mode internal tides, while explicitly excluding far-field dissipation associated with low-mode internal tides. By this design, q is a prescribed scalar between 0 and 1 as the reviewer commented. For example, St. Laurent et al. (2002) initially adopted a globally representative value of $q = 0.3 \pm 0.1$; subsequently, a spatially uniform value of 0.3 was taken in many ocean models (Saenko and Merryfield, 2005; Simmons et al., 2004). However, growing evidence indicates that near-field dissipation strongly depends on the modal partitioning of locally generated internal tidal energy, since high-mode energy dissipates rapidly near the generation sites whereas low-mode energy radiates over basin-scale distances (Vic et al., 2019). Consequently, q is not universal, it varies substantially among different topographies, ranging

from <0.4 over large-scale steep topography to >0.8 over small-scale rough topography (Vic et al., 2019; Lefauve et al., 2015).

The physical meaning of our defined parameter q differs from the aforementioned definition: our q quantifies the ratio of local dissipation to local generation, where the energy source for local dissipation consists of two parts (both locally generated internal tidal energy and incoming internal tidal energy from the far field). As recommended by the reviewer, we have renamed it as "local dissipation ratio q^* " in the revised manuscript to improve conceptual clarity and distinguish it clearly from the aforementioned "local dissipation efficiency q ". In this study, we apply the same definition of q^* to both the source region (Luzon Strait) and the sink region (Nansha Islands) to ensure a consistence. In source regions, locally generated internal tidal energy provides the principal energy supply for local dissipation; thus, q^* normally ranges from 0 to 1. In contrast, in sink regions, local generation is weak but dissipation is furnished by both local generation and incoming internal tidal energy from the far field; consequently, q^* exceeds 1.

As the reviewer suggested, normalizing local dissipation by the total available energy supply will yield q^* values between 0 and 1, even in sink regions. However, unlike local barotropic-to-baroclinic conversion, the incoming internal tidal energy is difficult to be estimated theoretically due to the complex wave-wave, wave-current, and wave-topography interactions, limiting the generalizability of q^* parameterization based on total energy supply. Therefore, local barotropic-to-baroclinic conversion is retained as our reference energy scale for defining q^* . This choice also enables an easy comparison with previous energy-budget studies in the South China Sea, in which dissipation was uniformly normalized only by local conversion (e.g., Tian et al., 2025; Wang et al., 2016).

References:

Lefauve, A., Muller, C., and Melet, A.: A three-dimensional map of tidal dissipation over abyssal hills, *J. Geophys. Res. Oceans*, 120, 4760–4777, <https://doi.org/10.1002/2014JC010598>, 2015.

Saenko, O. A. and Merryfield, W. J.: On the effect of topographically enhanced mixing on the global ocean circulation, *J. Phys. Oceanogr.*, 35, 826–834, <https://doi.org/10.1175/JPO2722.1>, 2005.

Simmons, H. L., Jayne, S. R., St. Laurent, L. C., and Weaver, A. J.: Tidally driven mixing in a numerical model of the ocean general circulation, *Ocean Modell.*, 6, 245–263, [https://doi.org/10.1016/S1463-5003\(03\)00011-8](https://doi.org/10.1016/S1463-5003(03)00011-8), 2004.

St. Laurent, L. C., Simmons, H. L., and Jayne, S. R.: Estimating tidally driven mixing in the deep ocean, *Geophys. Res. Lett.*, 29, 2106, <https://doi.org/10.1029/2002GL015633>, 2002.

Tian, Y., Bai, X., Wang, C., and Liu, Z.: Tidal energetics in the eddying South China Sea from a high-resolution numerical simulation, *Prog. Oceanogr.*, 231, 103418, <https://doi.org/10.1016/j.pcean.2025.103418>, 2025.

Vic, C., Naveira Garabato, A. C., Green, J. A. M., Waterhouse, A. F., Zhao, Z., Melet, A., de Lavergne, C., Buijsman, M. C., and Stephenson, G. R.: Deep-ocean mixing driven by small-scale internal tides, *Nat. Commun.*, 10, 2099, <https://doi.org/10.1038/s41467-019-10149-5>, 2019.

Wang, X., Peng, S., Liu, Z., Huang, R. X., Qian, Y., and Li, Y.: Tidal mixing in the South China Sea: an estimate based on the internal tide energetics, *J. Phys. Oceanogr.*, 46, 107–124,

<https://doi.org/10.1175/JPO-D-15-0082.1>, 2016.

(Changes can be found in Lines 8, 75–79, 166–167, 230–231, and 352–353, the title and the headings of Sections 3.1, 3.2, and 3.3 of the revised manuscript.)

2. The present dissipation estimate is obtained as a residual of the energy budget and should therefore be interpreted as an inferred dissipation. Have you compared this estimate with directly diagnosed dissipation from the model, using the available velocity/density gradients and vertical mixing coefficients? previous studies have reported difference between the two methods, or at least relevant discussions on this point should be included.

Responses: We thank the reviewer for this insightful comment. In the revised manuscript, we clarify that the present dissipation estimate (DIS_{bc}) represents an inferred dissipation from the energy budget equation rather than a directly diagnosed dissipation from the model.

The complete form of the depth-integrated baroclinic energy equation (Niwa and Hibiya, 2004) is $TEN_{bc} = -\nabla_h \cdot \mathbf{F} + Conv + ADV_{bc} + DIS_{bc}$, where $Conv$ is the depth-integrated barotropic-to-baroclinic energy conversion, $\mathbf{F}$ is the depth-integrated baroclinic energy flux, and TEN_{bc} , ADV_{bc} , and DIS_{bc} denote the tendency, the advection, and the dissipation of baroclinic energy, respectively (Wang et al., 2016). After a tidal-period averaging and neglecting the time-mean tendency and advection terms, the dissipation term simplifies as $\langle DIS_{bc} \rangle = \langle Conv \rangle - \langle \nabla_h \cdot \mathbf{F} \rangle$, where $\langle \cdot \rangle$ denotes the tidal-period averaging, a positive $\langle DIS_{bc} \rangle$ denotes the dissipation of baroclinic energy. This simplified equation has been widely applied in energy-budget studies of the South China Sea (e.g., Tian et al., 2025; Wang et al., 2016) and lays the basis for the tidal-mixing parameterization of St. Laurent et al. (2002). In this framework, $\langle Conv \rangle$ constrains the total energy available for turbulent mixing near internal tide generation sites. However, not all of the locally generated energy flux contributes to local dissipation, since a portion radiates away as low-mode waves. To account for this, St. Laurent et al. (2002) introduced a "local dissipation efficiency" (q) to quantify the fraction of locally generated internal tidal energy likely to dissipate near its source by turbulent processes.

As the reviewer noted, besides inferring from the energy budget equation, the dissipation can also be diagnosed directly from the simulation using the velocity gradients and the horizontal and vertical eddy viscosity coefficients:

$$\varepsilon_{visc} = \rho A_h \left[\left(\frac{\partial u}{\partial x} \right)^2 + \left(\frac{\partial u}{\partial y} \right)^2 + \left(\frac{\partial v}{\partial x} \right)^2 + \left(\frac{\partial v}{\partial y} \right)^2 \right] + \rho A_v \left[\left(\frac{\partial u}{\partial z} \right)^2 + \left(\frac{\partial v}{\partial z} \right)^2 \right],$$

where ρ is the reference density, A_h and A_v are the horizontal and vertical eddy viscosities, (u, v) is the horizontal baroclinic velocity (Nikurashin and Legg, 2011). However, in the LLC4320 simulation, A_h is determined internally by a parameterization scheme and is not provided as an output field; therefore, in this study it is a pity that we cannot exactly diagnose dissipation from the model output.

We acknowledge that there exist both connections and differences between the inferred dissipation and the diagnosed dissipation: the former not only includes the latter (i.e., explicit numerical dissipation), but also additionally covers other physical processes such as bottom friction dissipation. Fortunately, their quantitative discrepancy is not critical. For example, assuming a constant A_h , Fan et al. (2024) compared the diagnosed dissipation from the LLC4320 simulation with the inferred dissipation from the energy budget equation. They found that these two dissipation fields exhibit comparable magnitudes and broadly consistent spatial patterns across the northern South China Sea, despite localized discrepancies attributed to the

constant A_h assumption.

In the revised manuscript, we add a discussion about the inferred dissipation and the diagnosed dissipation.

References:

Fan, L., Sun, H., Yang, Q., and Li, J.: Numerical investigation of interaction between anticyclonic eddy and semidiurnal internal tide in the northeastern South China Sea, *Ocean Sci.*, 20, 241–264, <https://doi.org/10.5194/os-20-241-2024>, 2024.

Nikurashin, M. and Legg, S.: A mechanism for local dissipation of internal tides generated at rough topography, *J. Phys. Oceanogr.*, 41, 378–395, <https://doi.org/10.1175/2010JPO4522.1>, 2011.

Niwa, Y. and Hibiya, T.: Three-dimensional numerical simulation of M_2 internal tides in the East China Sea, *J. Geophys. Res.*, 109, C04027, <https://doi.org/10.1029/2003JC001923>, 2004.

(Changes can be found in Lines 152–165 of the revised manuscript.)

3. The four lower panels in Fig. 5 appear identical to the corresponding Luzon Strait panels in Fig. 3. Their magnitudes and curves also contradict lines 205–213, which state that the Nansha modal energy terms are much smaller and that mode-1 $q^* > 2$, reaching approximately 5. Please recheck these panels, and verify numerical statement that depends on them.

Responses: We sincerely apologize for this error. The four lower panels in Fig. 5 were improperly duplicated from the corresponding panels of Fig. 3. These panels have now been replaced with the correct figures for the Nansha Islands, and the statements based on them have been thoroughly rechecked. The corrected figure (Fig. R1) shows that the modal energy-budget terms in the Nansha Islands are significantly smaller than those in the Luzon Strait, and the energy budget is dominated by modes 1 and 2. Mode-1 baroclinic energy flux converges into the Nansha Islands, and mode-1 q^* values are larger than those for mode 2. Diurnal mode-1 q^* ranges from 2 to 5, and semidiurnal mode-1 q^* ranges from 1 to 3. For modes 3–5, monthly q^* values occasionally become extremely large, particularly when the area-integrated modal conversion is close to zero. The q^* values outside the displayed y-axis ranges in panels (l) and (m) of Fig. R1 are omitted for visual clarity, as stated in the revised figure caption.

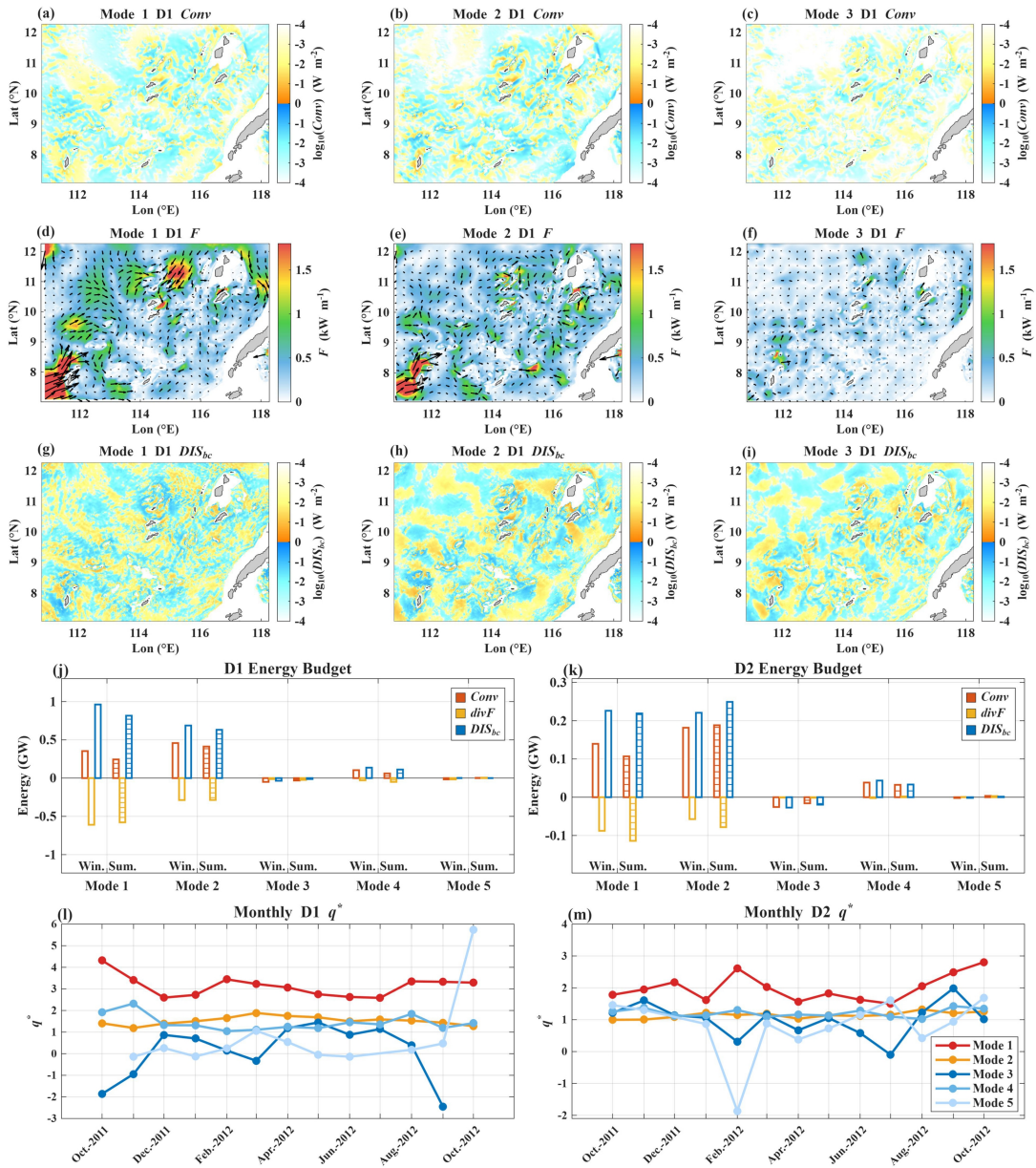


Figure R1. Same as Fig. 3, but for the Nansha Islands and (a–i) for diurnal internal tide. In (l–m), monthly mean values of q^* outside the displayed y-axis ranges are not shown for clarity.

(Changes can be found in Lines 259–267 of the revised manuscript.)

4. The paper attributes local dissipation to "high modes," yet only modes 1–5 are retained. Could you please compare the sum of modes 1–5 with the undecomposed modes to see the fraction for higher modes (>5).

Responses: We thank the reviewer for this insightful comment. In our analysis, only modes 1–5 are retained after modal decomposition; the baroclinic kinetic energy not captured by these five modes originates from unretained higher modes (modes >5). To quantify the residual contribution of these unretained modes, we compare the depth-integrated baroclinic kinetic energy reconstructed from the sum of modes 1–5 with the total baroclinic kinetic energy prior to modal decomposition, for each tidal band and each region individually, using data for January and August 2012. In the Luzon Strait, unretained modes (>5) account for 15% of the total baroclinic kinetic energy, and the contribution is 13% for the diurnal component and less

than 2% for the semidiurnal component. In the Nansha Islands, the corresponding values are 11%, 1%, and < 2%. These results confirm that modes 1–5 capture the dominant part of the baroclinic kinetic energy across all tidal bands and regions examined, and the unretained modes (>5) make only a small contribution. In the revised manuscript, modes 3–5 are referred to by explicit mode numbers (e.g., modes 3–5) rather than using the confusing term "high modes".

(Changes can be found in Lines 13–14, 19–20, and 132–137 of the revised manuscript.)

5. L30–36 The introduction would benefit from distinguishing radiation, loss of phase coherence, scattering, and irreversible dissipation. Alford et al. (2019) provide an independent estimate of basin-scale mode-1 attenuation and emphasize that apparent flux decay can include both true energy loss and loss of stationarity. This is relevant to the manuscript's nonlocal framing.

Responses: We thank the reviewer for this constructive suggestion. As the reviewer suggested, the introduction has been revised to distinguish radiation, loss of phase coherence, scattering, and irreversible dissipation, which play different roles in the internal tide energy budget.

Radiation transports internal tidal energy out of its source region through the baroclinic energy flux. Loss of phase coherence reduces the stationary (harmonically detectable) component of the signal but not necessarily the total internal tidal energy. Topographic scattering redistributes incident internal tidal energy across different propagation directions and vertical modes, rather than dissipating it (i.e., without net energy loss). Irreversible dissipation removes energy from the internal wave field through wave breaking and small-scale turbulence. Among these four processes, only irreversible dissipation serves as a true sink of internal tidal energy.

Alford et al. (2019) estimated the spatial attenuation of the mode-1 M_2 internal tide energy flux, which consists of both a stationary component detected by multisatellite altimetry and a nonstationary component corrected using an eddy-resolving tidal model. Across the North Pacific, the total fluxes yield e -folding scales of approximately 750–3,000 km. This apparent decrease in the total flux includes actual loss of mode-1 energy, through irreversible dissipation or scattering to higher modes, as well as loss of phase coherence (i.e., reduced stationarity). Such an apparent decrease therefore cannot be interpreted directly as equivalent irreversible energy loss.

The apparent decay of the mode-1 M_2 energy flux emphasized by Alford et al. (2019) is directly relevant to the nonlocal framing of the present study, in which the internal tidal energy dissipated in a sink region (e.g., the Nansha Islands) is supplied by the energy radiated from the far field (such as the Luzon Strait). For example, a substantial portion of the internal tidal energy generated in the Luzon Strait radiates out through the baroclinic energy flux and propagates toward the Nansha Islands, where topographic scattering redistributes the incoming energy from low to higher modes and facilitates their subsequent breaking and local dissipation. Crucially, the observed flux decay along this propagation pathway combines true energy dissipation with loss of phase coherence rather than dissipation alone. Because the energy budgets in the present study are calculated from hourly bandpass-filtered model fields rather than from phase-locked harmonic components, the de-phased energy remains within the target frequency band as part of the diagnosed flux rather than being attributed to the residual dissipation, as long as its frequency falls within the corresponding bandpass frequency band.

Reference:

Alford, M. H., Simmons, H. L., Marques, O. B., and Girton, J. B.: Internal tide attenuation in the North Pacific, *Geophys. Res. Lett.*, 46, 8205–8213, <https://doi.org/10.1029/2019GL082648>, 2019.

(Changes can be found in Lines 32–45 of the revised manuscript.)

7. Lines 40–46 define higher q as stronger near-field dissipation at a generation site. This interpretation cannot be transferred directly to Nansha, where $q > 1$ arises from imported energy.

Responses: We are sorry for this confusion. In the revised manuscript, we clearly distinguish two related but conceptually distinct quantities in terms of both the physical meanings and the symbols: (1) the "local dissipation efficiency" q (restricted between 0 and 1) used in prior parameterizations; (2) our newly defined "local dissipation ratio" q^* (can be larger than 1). Both quantities normalize local dissipation by local energy generation, but they differ fundamentally in how "local dissipation" is sourced: q considers only dissipation driven by locally generated internal tide energy, whereas q^* includes all dissipation fueled by both locally generated internal tide energy and incoming internal tide energy from the far field. In the revised introduction, a lower q indicates stronger far-field radiation of low-mode internal tides and a higher q indicates stronger near-field dissipation within an internal tide source region. In the *Methods* section, the inferred q^* is not necessarily restricted between 0 and 1, because the dissipation, furnished by both locally generated and imported internal tidal energy, is normalized by the local barotropic-to-baroclinic conversion alone, as in the Nansha Islands.

(Changes can be found in Lines 52–55, 78–79, and 169–171 of the revised manuscript.)

8. L69–71, here the introduction identifies background-flow shear as a possible control on q . Wang and Legg (2023, 2025) examined how baroclinic eddy shear affects internal-tide refraction, modal energy transfer, and dissipation, and therefore directly connect vertical shear to dissipation of internal tides.

Responses: We thank the reviewer for sharing these valuable references. Following the reviewer's suggestion, these references (Wang and Legg, 2023, 2025) have been added to the revised manuscript, together with a brief summary of their key findings. Specifically, these studies show that the shear current associated with strong baroclinic eddies can transfer part of the incident mode-1 internal tide energy into higher modes and trap these high modes within the eddies, thereby making the eddies hotspots for internal tide dissipation. This mechanism establishes a direct link between vertical shear and internal tide dissipation.

References:

Wang, Y. and Legg, S.: Enhanced dissipation of internal tides in a mesoscale baroclinic eddy, *J. Phys. Oceanogr.*, 53, 2293–2316, <https://doi.org/10.1175/JPO-D-23-0045.1>, 2023.

Wang, Y. and Legg, S.: Agulhas rings locally enhance dissipation of internal tides, *Sci. Adv.*, 11, eadq5963, <https://doi.org/10.1126/sciadv.adq5963>, 2025.

(Changes can be found in Lines 84–90 of the revised manuscript.)

9. L199–203, Kelly et al. (2013) is a relevant independent reference for the redistribution of incident mode-1 energy among reflection, transmission, and scattering to higher modes at large topographic features. It would also help the authors distinguish scattering from final irreversible dissipation.

Responses: We thank the reviewer for sharing this valuable reference. The reference (Kelly et

al., 2013) has been added into our revised manuscript. In Section 3.2, topographic scattering is now explicitly distinguished from the irreversible dissipation that follows the breaking of the scattered high-mode waves.

Reference:

Kelly, S. M., Jones, N. L., Nash, J. D., and Waterhouse, A. F.: The geography of semidiurnal mode-1 internal-tide energy loss, *Geophys. Res. Lett.*, 40, 4689–4693, <https://doi.org/10.1002/grl.50872>, 2013.

(Changes can be found in Lines 251–258 of the revised manuscript.)

10. Abstract: Please distinguish total-band q , constituent-specific q , and modal q . The abstract reports total q occasionally exceeding 2.5, whereas the text later reports mode-1 values approaching 5.

Responses: We are sorry for this confusion. In the revised abstract, the q^* values specifically refer to those of the total internal tide. The q^* value of the total internal tide fluctuates between 0.3 and 0.7 in the Luzon Strait and is generally larger than 1 in the Nansha Islands, occasionally exceeding 2.5. The q^* values of individual tidal constituents and vertical modes are reported in the *Results* section. In the LS, the semidiurnal q^* ranges from 0.6 to 0.8 and the mode-1 semidiurnal q^* ranges from 0.3 to 0.7. In the Nansha Islands, the diurnal q^* ranges from 1.5 to 3.0 and the mode-1 diurnal q^* ranges from 2.0 to 5.0.

(Changes can be found in Lines 13–18 of the revised manuscript.)

11. Table 1. revise caption. The caption promises "associated errors," but no coefficient errors or confidence intervals are shown.

Responses: We are sorry for this error. Table 1 presents the fitted coefficients together with two metrics R^2 and RMSE, which are indicators of the overall model performance rather than errors or uncertainties associated with individual coefficients. To improve clarity, both the table caption and the corresponding sentence in the manuscript have been revised accordingly.

(Changes can be found in Lines 293–296 and 319 of the revised manuscript.)

12. Figure 8. Only D2 results are shown, although performance claims are made for both bands.

Responses: We are sorry for this information gap. Figure 8 has been expanded to show the time-series comparisons between the predicted q^* and actual q^* for each combination of region and tidal band, namely, (a) LS-D1, (b) LS-D2, (c) Nansha-D1, and (d) Nansha-D2. The revised Fig. 8 is provided below as Fig. R2.

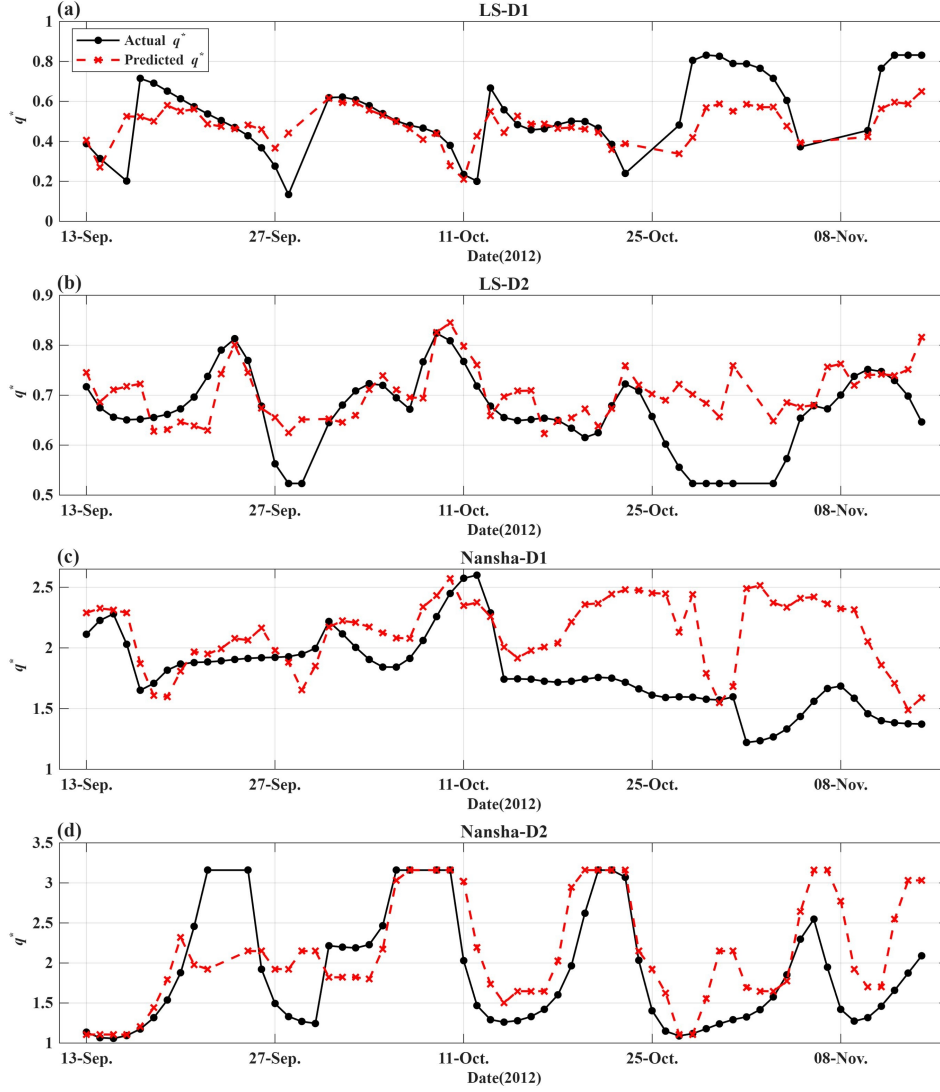


Figure R2. Time series of the predicted q^* from the XGBoost model and the actual q^* for the diurnal (D1) and semidiurnal (D2) tides in the LS and the Nansha Islands.

(Changes can be found in Lines 340–342 of the revised manuscript.)

13. Please discuss uncertainty associated with approximately 2-km horizontal resolution, unresolved wave breaking, numerical mixing, and the residual method.

Responses: We are sorry for this information gap. A new paragraph outlining the limitations of our study has been added at the end of the *Summary* section.

The horizontal grid spacing of LLC4320 output is approximately 2 km in the study region, at which the major topographic features and the low-mode internal tides are explicitly resolved. However, the small-scale topographic features, the high-mode internal tides with shorter horizontal wavelengths, the internal wave breaking and the associated turbulence remain subgrid-scale processes and hence are not resolved. Consequently, internal tide generation driven by the unresolved roughness is missing in the simulation, and the energy cascading to the subgrid-scale processes is instead dissipated via the parameterized viscosity and diffusivity, the bottom drag, and the numerical damping.

Because the present dissipation estimate (DIS_{bc}) is obtained as the residual of the simplified

baroclinic energy budget equation, it conflates multiple physical and numerical processes such as the parameterized viscosity, diffusivity, and bottom drag, the numerical mixing and damping. The full energy budget cannot be closed with the archived model output, and these individual contributions cannot be quantified separately. Therefore, DIS_{bc} represents an inferred dissipation integrated over the study region rather than a diagnosed dissipation from the model at individual topographic features. Since DIS_{bc} appears in the numerator of q^* , its uncertainty propagates directly into the estimate of q^* . Crucially, the relative uncertainty of q^* is amplified where the area-integrated local conversion is weak. Such uncertainty is particularly pronounced over rough topography and for higher-mode internal tides (Vic et al., 2019). Future work with higher-resolution simulations and diagnostics of the individual dissipative terms is expected to separate and quantify these contributions.

(Changes can be found in Lines 376–390 of the revised manuscript.)

14. Equations and notation. Please define all symbols upon first use and ensure consistent notation, signs. Also clarify the sign conventions for $Conv$, $(\text{div } F)$, and (DIS_{bc}) .

Responses: We are sorry for this information gap. All equations and their accompanying descriptions have been checked and revised. Each symbol is explicitly defined at its first appearance, the notation and sign are applied consistently throughout the revised manuscript, and the sign is clearly stated in the *Methods* section. Specifically, a positive $Conv$ denotes conversion from barotropic to baroclinic energy, that is, local internal tide generation; a negative $Conv$ denotes the reverse (baroclinic-to-barotropic) conversion. A positive divergence of F indicates outward propagation of internal tidal energy from the local region; a negative divergence of F indicates convergence of energy flux from the far field. A positive DIS_{bc} denotes net dissipation of baroclinic energy inferred from the energy budget equation; a negative DIS_{bc} denotes a net gain of baroclinic energy. Such negative DIS_{bc} values can occur for individual vertical modes when the divergence of baroclinic energy flux for the mode exceeds its local conversion, indicating energy input from other vertical modes via inter-mode transfer.

(Changes can be found in Lines 122–123, 147–150, and 152–155 of the revised manuscript.)

Local dissipation **ratio** of internal tides at key topographic features in the South China Sea

Zuqing Yuan¹★, Hui Sun¹★, Qingxuan Yang^{1,2}, Fenyuan Han¹, Jianing Li¹

¹State Key Laboratory of Physical Oceanography, Ocean University of China, Qingdao, China.

5 ²Laboratory for Ocean Dynamics and Climate, Qingdao Marine Science and Technology Center, Qingdao, China.

★These authors contributed equally to this work.

Correspondence to: Qingxuan Yang (yangqx@ouc.edu.cn)

Abstract. The local dissipation ratio of internal tides, q^* , is a critical parameter in tidal mixing parameterizations. However, the conventionally adopted constant value of 0.3 in large-scale ocean models neglects its significant spatiotemporal variability.

10 Based on the MITgcm LLC4320 simulation, the internal tidal energy budgets at the Luzon Strait (a source region, LS) and the Nansha Islands (a sink region, Nansha) in the South China Sea (SCS) are analyzed. Results indicate that the barotropic-to-baroclinic energy conversion in the LS reaches approximately 45 GW, with semidiurnal constituents accounting for roughly 60%, due to the resonance over the double-ridge topography. The q^* value of the total internal tide in the LS fluctuates between 0.3 and 0.7, primarily modulated by the local dissipation of modes 3–5. Local internal tide generation around the Nansha

15 Islands is less than 1.5 GW; however, this region experiences significant convergence of internal tidal energy flux, so that the q^* value of the total internal tide is generally larger than 1 and occasionally exceeds 2.5. Modal analysis confirms that the intensified dissipation over the Nansha Islands originates predominantly from topographic scattering and breaking of mode-1 internal tides from the far field. Parameterizations for q^* are developed based on both physical factors and data-driven algorithms, both of which successfully capture the macroscopic clustering of q^* . In the LS, q^* is modulated by near-field factors

20 such as the barotropic tidal forcing and the local dissipation of modes 3–5. Conversely, q^* around the Nansha Islands is primarily contributed by mode-1 internal tidal energy coming from the far field, highlighting the joint modulation of local extreme dissipation by far-field beam interference and nonlinear topographic scattering.

Diapycnal mixing in the ocean is a key dynamical process that maintains the global ocean stratification, shapes the meridional overturning circulation, and regulates the transport of heat, energy, and nutrients (Munk and Wunsch, 1998; Wunsch and Ferrari, 2004). In the deep ocean, the interaction of barotropic tides with rough topography generates internal tides, which are widely regarded as the primary mechanical energy source furnishing deep-ocean mixing (Egbert and Ray, 2000; Garrett and Kunze, 2007; Ferrari and Wunsch, 2009). It is estimated that approximately 1 TW of barotropic tidal energy is converted into internal tides, accounting for roughly half of the total energy required to maintain global abyssal mixing (Egbert and Ray, 2003; Nycander, 2005). However, after generation, internal tidal energy undergoes radiation, loss of phase coherence, topographic scattering, and irreversible dissipation, each playing a distinct role in the internal tide energy budget. Radiation transports internal tidal energy away from the generation region through the baroclinic energy flux (St. Laurent and Garrett, 2002). Loss of phase coherence reduces the stationary (i.e., harmonically detectable) component without necessarily reducing the total internal tidal energy. Topographic scattering redistributes incident internal tidal energy across different propagation directions and vertical modes rather than dissipating it directly (i.e., without net energy loss); however, by transferring energy to higher, more dissipative modes, it indirectly promotes subsequent wave breaking (MacKinnon et al., 2017). Irreversible dissipation removes internal tidal energy from the internal wave field as energy cascades nonlinearly to progressively smaller scales, ultimately triggering wave breaking and turbulence, which drive diapycnal mixing (Waterhouse et al., 2014). Usually, these processes are intertwined. For instance, Alford et al. (2019) estimated the spatial attenuation of the mode-1 M_2 internal tide energy flux, which consists of both a stationary component detected by multisatellite altimetry and a nonstationary component corrected using an eddy-resolving tidal model. Across the North Pacific, the total fluxes yield e -folding scales of approximately 750–3,000 km. This apparent decrease in the total flux includes actual loss of mode-1 energy, through irreversible dissipation or scattering to higher modes, as well as loss of phase coherence (i.e., reduced stationarity). Therefore, accurately quantifying the spatiotemporal distribution of the local internal tidal dissipation is crucial for improving the performance of ocean circulation and climate models (Melet et al., 2016; de Lavergne et al., 2020).

In state-of-the-art climate-scale ocean models, the mixing induced by internal tidal breaking cannot be explicitly resolved due to limitations in grid resolution and must therefore be parameterized (Jayne, 2009). The local dissipation efficiency, q , is a decisive dimensionless parameter defined as the fraction of locally generated internal tidal energy that dissipates near the generation site, with the remainder radiating away as low-mode internal tides; thus, q is restricted between 0 and 1; it directly modulates the spatial distribution of internal tidal energy (St. Laurent et al., 2002). In an internal tide source region, a lower q indicates that a larger fraction of the locally generated internal tidal energy propagates toward the far field as low-mode internal tides; a higher q indicates stronger near-field dissipation, leading to enhanced local mixing (Klymak et al., 2011; Vic et al., 2019). For a long time, to simplify calculations, ocean circulation models have typically prescribed q as a globally uniform constant of 0.3 (Simmons et al., 2004; Saenko and Merryfield, 2005). However, recent studies have demonstrated significant spatial heterogeneity in q . Lefauve et al. (2015) noted its topographic dependence, and Vic et al. (2019) further elucidated that

60 q is regulated by the vertical modal structure of internal tides via a global semi-analytical model. Over small-scale rough topography, high-mode internal tides are predominantly generated, yielding high q values (> 0.8); whereas over large-scale steep topography, low-mode internal tides prevail, leading to low q values (< 0.4). Ignoring the spatiotemporal variability of q causes an order-of-magnitude miscalculation of the mechanical energy supplying internal tidal mixing, thereby significantly reducing the model's capacity to simulate the evolution of water masses and the intensity of global meridional overturning circulation (Cimoli et al., 2019).

65 As the largest marginal sea in the western Pacific and one of the most well-known hotspots for internal tides, the South China Sea (SCS) is an ideal study area for quantifying internal tidal energy budgets and exploring the spatiotemporal evolution of the dissipation parameter q due to its complex dynamical environment. The SCS not only encompasses the Luzon Strait (LS), a globally recognized strong internal tidal source region (Niwa and Hibiya, 2004; Jan et al., 2007), but also contains complex topographies such as the Nansha Islands (Nansha) and the Dongsha Islands, both located within or adjacent to the basin, which serve as key sink regions for internal tidal energy. Internal tidal energy budget analysis by Wang et al. (2016) 70 revealed that the total local dissipation accounts for 38% of the local conversion in the LS during summer, whereas in the broader SCS it is 267% of the local conversion. Tian et al. (2025) reported that during summer about 24% of the baroclinic tidal energy generated in the LS is dissipated locally, and in the broader SCS the total local dissipation is 203% of the local conversion. When low-mode internal tides propagate over rough topography far from the source regions, topographic scattering effects promote the energy transfer to high modes, thereby enhancing the dissipation there (Klymak et al., 2010). 75 Therefore, we define the local dissipation ratio of internal tides, q^* , as the area-integrated total local dissipation inferred from the baroclinic energy budget divided by the area-integrated barotropic-to-baroclinic conversion. q^* is normalized by the same local conversion as the conventional q , but its numerator may be furnished by both local conversion and the convergence of internal tidal energy from the far field. Consequently, in source regions (such as the Luzon Strait) q^* has the same meaning as q and normally ranges from 0 to 1, whereas in sink regions (such as the Nansha Islands) it may exceed 1.

80 Although the generation and propagation of internal tides in the SCS have been extensively explored in previous studies (Zhao et al., 2010; Tian et al., 2025), research on q and its parameterizations remains significantly inadequate. First, existing studies mostly focus on estimating the spatial distribution of q globally and its magnitude in internal tide generation regions. There is a lack of analysis of the temporal variation of q^* and of systematic comparison of the differences in q^* between source and sink regions for internal tidal energy. Second, although it is theoretically established that q depends on the internal tidal 85 modal structure and may be related to topographic gradients and the shear intensity of background flow fields (Polzin, 2004; Dunphy and Lamb, 2014), most global ocean circulation models still use empirical constants to set q , neglecting its spatiotemporal variability. Wang and Legg (2023, 2025) examined how baroclinic eddy shear affects internal tide refraction, modal energy transfer, and dissipation, and showed that the shear current associated with strong baroclinic eddies can transfer part of the incident mode-1 internal tide energy into higher modes and trap these high modes within the eddies, thereby making 90 the eddies hotspots for internal tide dissipation. Therefore, exploring the spatiotemporal variations of q^* under different

dynamical environments and developing physically consistent parameterizations are essential for accurately simulating internal tidal mixing.

In this study, we intend to diagnose the internal tidal energy budgets, quantify the temporal variations and regional differences of the local dissipation ratio q^* , and establish parameterizations of q^* for two representative internal tidal regions, a source region (LS) and a sink region (Nansha Islands) in the SCS. In Sect. 2, we introduce the utilized data from the MITgcm LLC4320 simulation output, the methods for vertical modal decomposition and internal tidal energy budget analysis, and the eXtreme Gradient Boosting (XGBoost) model employed to predict q^* . Section 3 presents the results of internal tidal energy budgets and q^* in the LS and the Nansha Islands, respectively. Specifically, the tidal energy budgets and q^* are analyzed for each vertical mode separately. Then, the performance of the q^* parameterizations based on both the physical-based empirical formulation and the XGBoost model is demonstrated. Finally, our results are summarized in Sect. 4.

2 Data and Methods

2.1 Data

The data used in this study are the LLC4320 simulation outputs from the Massachusetts Institute of Technology general circulation model (MITgcm), which solves the Navier-Stokes equations under the hydrostatic and Boussinesq approximations. The model is forced by the complete luni-solar tidal potential including 16 principal tidal constituents (such as M_2 , S_2 , K_1 , and O_1) and by the 6-hourly atmospheric reanalysis dataset provided by the European Centre for Medium Range Weather Forecasts (ECMWF) (Marshall et al., 1997; Arbic et al., 2018). An Arakawa-C grid mesh is used in the simulation with a horizontal resolution of $1/48^\circ$ (approximately 2 km in the SCS). In the vertical, the model has 90 z -coordinate levels, with layer thickness increasing from 1 m at the surface to around 480 m near the seabed. The hourly output of three-dimensional velocity, potential temperature, salinity, and sea surface height from 13 September 2011 to 15 November 2012 is used, with specific focus on two key topographic regions, the LS and the Nansha Islands (Fig. 1). The Luzon Strait domain is defined to encompass the Hengchun Ridge, the Lanyu Ridge, and the inter-ridge trough, which collectively constitute the double-ridge where intense internal tides are generated (Buijsman et al., 2014; Tian et al., 2025; Wang et al., 2016). The Nansha Islands domain is selected to encompass the dominant topographic features (islands and reefs), which are located in the propagation paths of internal tides coming from the LS (Tian et al., 2025; Xu et al., 2016). The interaction of the incoming internal tides with the rough islands and reefs promotes energy transfer to higher modes and hence enhances local dissipation (Vic et al., 2019).

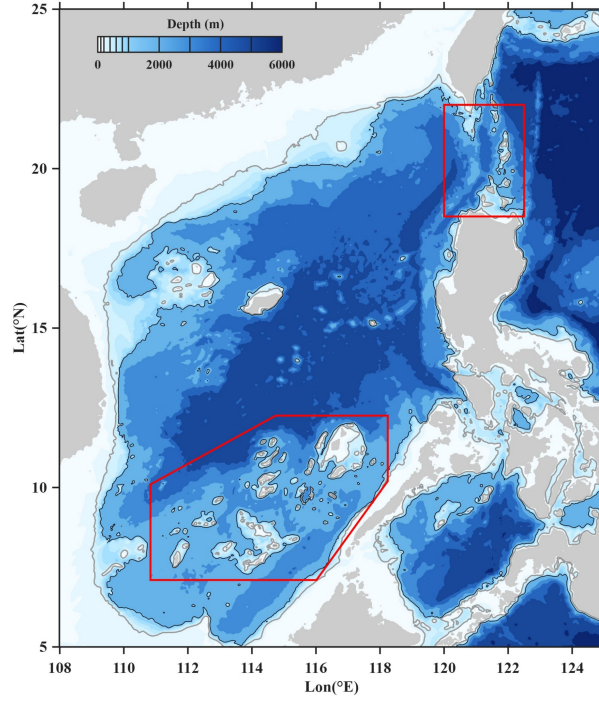


Figure 1. Topography of the SCS. Gray and black contours represent the 200-m and 1000-m isobaths, respectively. Red boxes represent the key topographic regions, including the LS and the Nansha Islands.

120 2.2 Methods

The subtidal fields ($\bar{\mathbf{u}}$ and $\bar{p}$) are obtained by applying a 3-day moving average to the total fields, and the perturbation fields ($\mathbf{u}', p'$) are obtained by subtracting the subtidal fields from the total ones. The barotropic velocity $\mathbf{u}'_{bt} = (u'_{bt}, v'_{bt})$ is defined as the depth-averaged velocity, and the baroclinic velocity is calculated as $\mathbf{u}'_{bc} = \mathbf{u}' - \mathbf{u}'_{bt}$. A bandpass filter is then applied to extract the diurnal (D1, 0.90–1.10 cpd) and semidiurnal (D2, 1.87–2.10 cpd) tidal constituents. When diagnosing the internal tidal energy budgets and q^* for individual vertical modes, a vertical modal decomposition method is employed. The modal structure function $\phi_n(z)$ can be obtained by solving the Sturm-Liouville equation (Gill, 1982):

$$\frac{d^2 \phi_n(z)}{dz^2} + \frac{N^2(z)}{c_n^2} \phi_n(z) = 0, \quad (1)$$

where boundary conditions are given as $\phi_n(0) = \phi_n(-H) = 0$, and H is the water depth. Here, $N(z)$ is the buoyancy frequency, c_n is the eigenspeed, and n is the mode number. Due to the orthogonality of the modes, the baroclinic velocity $\mathbf{u}'_{bc}$ and the pressure perturbation p' can be expressed as $\mathbf{u}'_{bc}(z, t) = \sum_{n=1}^N U_n(t) \phi_n(z)$, $p'(z, t) = \sum_{n=1}^N P_n(t) \phi_n(z)$, where $\phi_n(z) = d\phi_n(z)/dz$ is the vertical structure for horizontal velocity and pressure, $U_n(t)$ and $P_n(t)$ are the coefficients of individual modes, and $N = 5$ is the total number of modes used in this study. Taking January and August 2012 as representative cases, in the LS and the Nansha Islands, modes 1–5 account for 85% and 89% of the total baroclinic kinetic

energy, 87% and 99% of the diurnal baroclinic kinetic energy, and more than 98% of the semidiurnal baroclinic kinetic energy, respectively. Modes 1–5 capture the dominant part of total baroclinic kinetic energy, while the remaining fraction (mode 6 and above) does not exceed 15% of the total baroclinic kinetic energy, although these unretained higher modes may still contribute less to local dissipation of internal tides over rough topography (Vic et al., 2019).

To evaluate the generation and dissipation of internal tidal energy, the depth-integrated baroclinic energy equation is used here. After tidal-period averaging and neglecting the tendency and advection terms, the energy equation reduces to (Niwa and Hibiya, 2004):

$$\langle DIS_{bc} \rangle = \langle Conv \rangle - \langle \nabla_h \cdot \mathbf{F} \rangle, \quad (2)$$

where $\langle \cdot \rangle$ denotes the tidal-period mean. $Conv$ is the depth-integrated barotropic-to-baroclinic energy conversion, indicating the local internal tide generation, which can be calculated as:

$$Conv = \int_{-H}^{\eta} \rho' g w'_{bt} dz, \quad (3)$$

where η is the sea level displacement, and w'_{bt} is the vertical velocity induced by the barotropic tides. This vertical velocity w'_{bt} is given by $w'_{bt} = u'_{bt} \left(\sigma \frac{\partial D}{\partial x} + \frac{\partial \eta}{\partial x} \right) + v'_{bt} \left(\sigma \frac{\partial D}{\partial y} + \frac{\partial \eta}{\partial y} \right) + (\sigma + 1) \frac{\partial \eta}{\partial t}$ (Mellor, 2004), where D is the total water depth, $D = H + \eta$, and $\sigma = (z - \eta)/D$. A positive $Conv$ denotes conversion from barotropic to baroclinic energy, that is, local internal tide generation, whereas a negative $Conv$ denotes the reverse (baroclinic-to-barotropic) conversion. $\mathbf{F}$ is the depth-integrated baroclinic energy flux, and $\nabla_h \cdot \mathbf{F}$ is the horizontal divergence of $\mathbf{F}$. A positive $\nabla_h \cdot \mathbf{F}$ indicates outward propagation of internal tidal energy from the local region; a negative $\nabla_h \cdot \mathbf{F}$ indicates convergence of energy flux from the far field. $\mathbf{F}$ is given as:

$$\mathbf{F} = \int_{-H}^{\eta} \mathbf{u}'_{bc} p' dz, \quad (4)$$

where p' is calculated as $p'(z, t) = -\frac{1}{H} \int_{-H}^{\eta} \int_z^{\eta} g \rho'(\hat{z}, t) d\hat{z} dz + \int_z^{\eta} g \rho'(\hat{z}, t) d\hat{z}$. DIS_{bc} is the depth-integrated local dissipated internal tidal energy, inferred as the residual of the simplified depth-integrated baroclinic energy equation (Eq. (2)) rather than a diagnosed dissipation from the model. A positive DIS_{bc} denotes net dissipation of baroclinic energy inferred from the energy budget; a negative DIS_{bc} denotes a net gain of baroclinic energy. In some studies, the dissipation can also be diagnosed from the simulation using the velocity gradients and the horizontal and vertical eddy viscosity coefficients (Nikurashin and Legg, 2011). However, in the LLC4320 simulation, horizontal eddy viscosity is determined internally by a parameterization scheme and is not provided as an output; therefore, we cannot exactly diagnose dissipation from the model output. There exist both connections and differences between the inferred dissipation and the diagnosed dissipation: the former not only includes the latter (i.e., explicit numerical dissipation), but also covers other physical processes such as bottom friction dissipation. Fortunately, their quantitative discrepancy is not critical. For example, assuming a constant horizontal eddy viscosity, Fan et al. (2024) compared the diagnosed dissipation from the LLC4320 simulation with the inferred dissipation from the energy budget equation. They found that these two dissipation fields exhibit comparable magnitudes and broadly consistent spatial

patterns across the northern SCS, despite localized discrepancies attributed to the constant horizontal eddy viscosity assumption.

The local internal tidal dissipation ratio, q^* , is defined as the ratio of the area-integrated DIS_{bc} to $Conv$, i.e., $q^* = \int_s ds(DIS_{bc}) / \int_s ds(Conv)$. Here, the spatial integration domain s corresponds to the red-boxed regions shown in Fig. 1. Based on the above calculations, the internal tidal energy budgets and q^* for different tidal components and different vertical modes can be estimated. In this way, q^* is not necessarily restricted between 0 and 1. $q^* > 1$ occurs when the area-integrated divergence of the baroclinic energy flux is negative, indicating a convergence of far-field internal tidal energy. In this case, the inferred local dissipation exceeds the barotropic-to-baroclinic energy conversion. For individual vertical modes, $q^* < 0$ occurs when the area-integrated divergence of the baroclinic energy flux of this mode exceeds its local barotropic-to-baroclinic energy conversion, yielding a negative inferred dissipation. The divergence of the baroclinic energy flux is furnished by both the local conversion and the net energy transfer from other vertical modes, and may therefore exceed the local barotropic-to-baroclinic energy conversion. Incomplete months (September 2011 and November 2012) are excluded when computing the monthly mean values of q^* . January and August 2012 are selected as winter and summer months in the simulated year, respectively.

To develop a more robust parameterization of q^* , this study not only adopts a physical-based empirical formulation but also uses the XGBoost model jointly, which can make up for the limitations of traditional physical-based empirical formula in complex dynamic environments. The model inputs include 6 physical features, namely the internal tidal shear averaged within the bottommost 500 m, stratification, barotropic tidal kinetic energy, mode-1 internal tidal energy, the proportion of mode-1 internal tidal energy in the total internal tidal energy, and the mode-1 internal tidal group velocity. The data from the LS and the Nansha Islands are sequentially combined into a training set, a validation set, and a test set with proportions of 70%, 15%, and 15%, respectively. Considering the differences in the physical processes governing q^* between the internal tidal source regions and sink regions, this study constructs two separate models, one for the LS and one for the Nansha Islands. The Mean Absolute Error (MAE) is adopted as the objective loss function, i.e., $MAE = \frac{1}{n} \sum_{i=1}^n |\hat{y}_i - y_i|$, where y_i and $\hat{y}_i$ denote the actual and predicted values, respectively, and n is the number of samples. An early stopping iteration setting of 500 rounds and 800 rounds are applied for models of the LS and Nansha Island, respectively.

3 Results

3.1 Local dissipation ratio of internal tides in the LS

The internal tidal energy budget in the LS is characterized by intense generation, strong outward radiation, and relatively low q^* . The barotropic-to-baroclinic conversion in the LS ranges from 40 to 50 GW, of which 15–25 GW radiates out of the LS (Fig. 2g), yielding a q^* value of approximately 0.5 (Fig. 2h). The semidiurnal wavelength better matches the alignment of the double ridges in the LS, resulting in a larger semidiurnal tide generation (20–30 GW) than diurnal tide generation (10–25

GW) (Fig. 2g). The semidiurnal dissipation (10–20 GW) and q^* values (0.6–0.8) are also higher than those of the diurnal tide, whose dissipation ranges from 1 to 10 GW and q^* from 0.1 to 0.6 (Fig. 2g–h). The semidiurnal conversion, dissipation, and q^* are larger in summer than in winter (Fig. 2g–h). The diurnal q^* reaches minima of 0.1 and 0.3 in February–March and June–July, respectively (Fig. 2h).

The summer spatial distributions demonstrate that internal tide generation and dissipation are concentrated over and around the double ridges (Fig. 2a–b and 2e–f). Substantial internal tidal energy propagates westward and northwestward from the western ridge into the SCS (Fig. 2c–d). The semidiurnal component plays a dominant role in the internal tide generation and dissipation (Fig. 2a–b, 2e–g). The q^* values ranging from 0.49 to 0.57 in the LS are comparable to those in other representative internal tide generation regions; for example, You et al. (2023) estimated a value of about 0.58 for the M_2 internal tide at the Aleutian Ridge, suggesting that q^* values in submarine ridges serving as internal tide source regions share a similar magnitude.

The internal tidal energy budgets and q^* in the LS exhibit differences among individual tidal components and vertical modes. Both diurnal and semidiurnal internal tide generation are dominated by modes 1–2, with mode 1 accounting for roughly 50% of the total generation and exhibiting an energy flux divergence greater than those of other modes (Fig. 3j–k). The divergence of diurnal mode-1 energy flux is even larger in magnitude than the local generation (Fig. 3j), resulting in negative q^* values (Fig. 3l), implying that the internal tidal energy generated in the LS radiates outward as low-mode internal tides intensively. Tian et al. (2025) indicated that the diurnal tide has a larger decay scale than the semidiurnal tide, which favors more outward radiation of the diurnal tide as mode 1. **By contrast, although the semidiurnal energy budget is also dominated by modes 1–2, the proportion of mode-1 energy flux divergence is relatively smaller, and its difference between January and August 2012 is weaker than that of the diurnal tide (Fig. 3j–k).** Regarding local dissipation, q^* for modes 3–5 of both the diurnal and semidiurnal components ranges from 0.8 to 2.7, significantly larger than that of modes 1–2 (< 0.8) (Fig. 3l–m).

The strongly positive generation regions for modes 1–2 internal tides are concentrated at the crests of the double ridges, whereas the high-generation regions for mode 3 are distributed along the ridge flanks (Fig. 3a–c). The energy budget of the mode-1 internal tide is characterized by substantial energy propagating northwestward and westward into the SCS, as well as eastward into the Pacific Ocean (Fig. 3d). The majority of the mode-1 semidiurnal energy flux enters the SCS, with peak values reaching 25 kW m^{-1} (Fig. 3d). Local dissipation gradually weakens and becomes spatially dispersed with increasing mode number (Fig. 3g–i). At the Mid-Atlantic Ridge and the Hawaiian Ridge, Vic et al. (2018) indicated that a dominant proportion of mode-1 internal tide generation facilitates the outward radiation of substantial energy as low-mode internal tides, yielding small values of the dissipation ratio. This result confirms the robust far-field radiation capacity of internal tides and the smaller q^* values of mode 1 relative to those of higher modes in the LS.

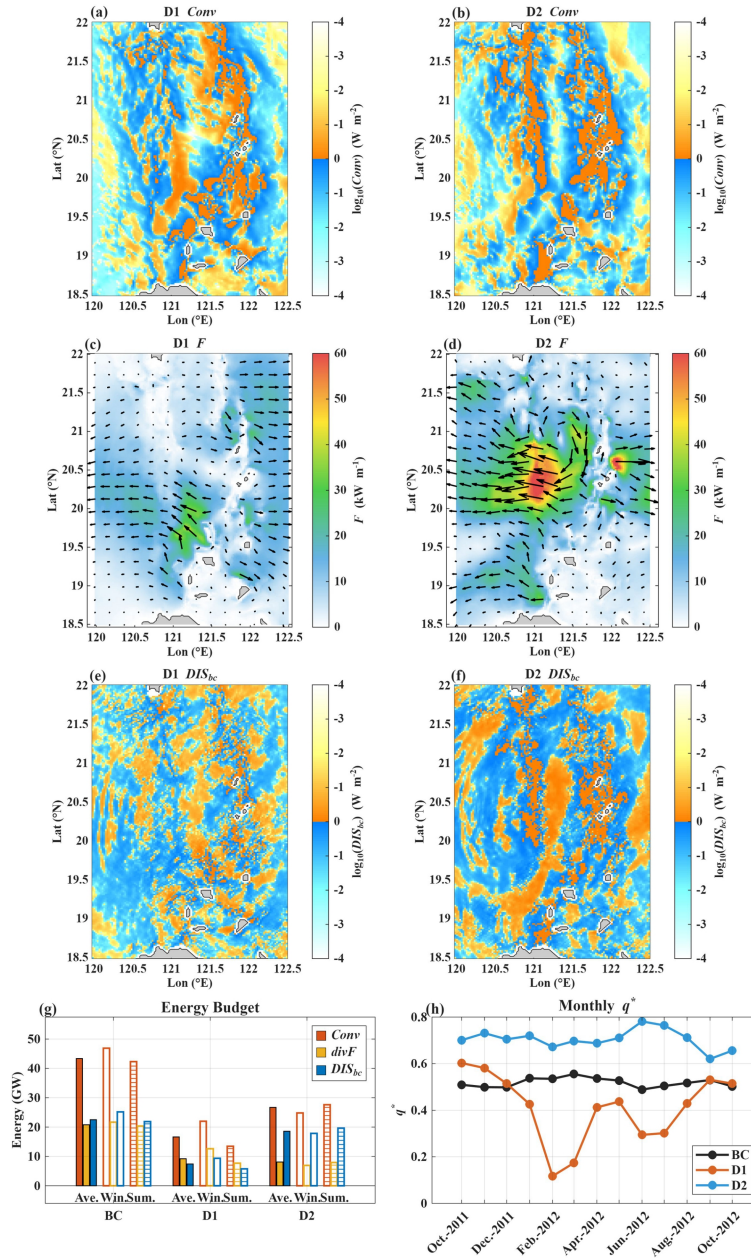
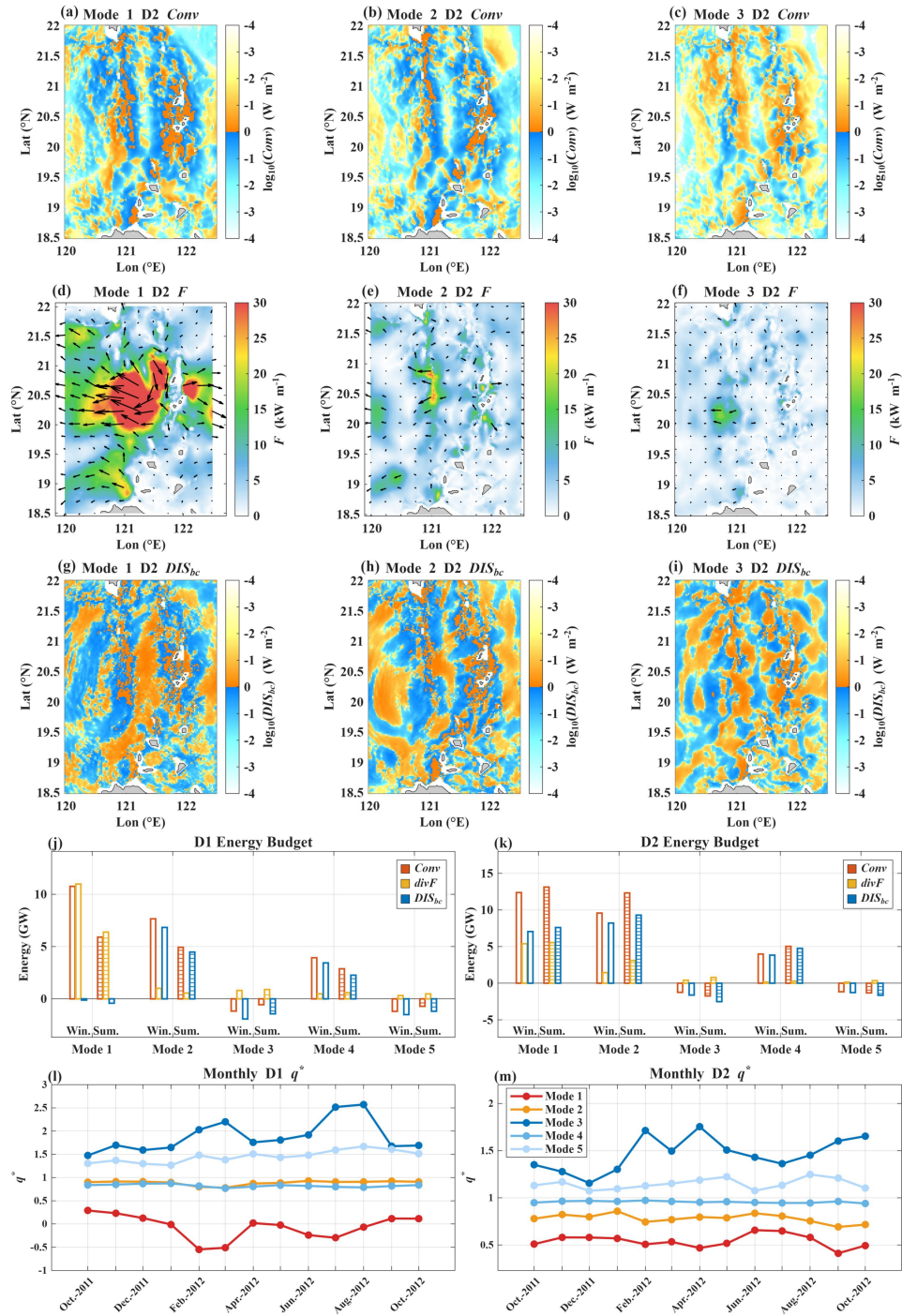


Figure 2. (a–f) Spatial distributions of the diurnal and semidiurnal internal tidal energy budget terms in the LS during summer. $\langle Conv \rangle$ is the tidal-period mean depth-integrated barotropic-to-baroclinic energy conversion, $\langle F \rangle$ is the tidal-period mean depth-integrated internal tidal energy flux, and $\langle DIS_{bc} \rangle$ is the tidal-period mean depth-integrated local internal tidal dissipation. (g) Bar chart of the budget terms. $\langle divF \rangle$ is the tidal-period mean divergence of the depth-integrated internal tidal energy flux. “Ave.,” “Win.,” and “Sum.” stand for the annual-, winter-, and summer-mean values, respectively. BC, D1, and D2 represent the total internal tide, diurnal tide, and semidiurnal tide, respectively. (h) Time series of the monthly-mean local internal tidal dissipation ratio, q^* . In (a–b) and (e–f), the units for $Conv$ and DIS_{bc} are $W m^{-2}$, with orange (blue) representing positive (negative) values. In (c–d), the unit for F is $kW m^{-1}$, with colors representing magnitude and arrows indicating direction. In (g), the units for $Conv$, $divF$, and DIS_{bc} are GW.



235 **Figure 3.** (a–i) Spatial distributions of the semidiurnal internal tidal energy budget terms for modes 1–3 in the LS during summer. (j–k) Bar charts of the energy budget terms for modes 1–5, and (l–m) time series of monthly-mean q^* for modes 1–5 for diurnal (D1) and semidiurnal (D2) tides, respectively. “Win.” and “Sum.” stand for winter and summer, respectively. In (a–c) and (g–i), the units for $Conv$ and DIS_{bc} are $W m^{-2}$, with orange (blue) indicating positive (negative) values. In (d–f), the unit for F is $kW m^{-1}$, with colors

240 indicating magnitude and arrows indicating direction. In (j–k), the units for *Conv*, *divF*, and *DIS_{bc}* are GW. In (l–m), monthly mean values of q^* outside the displayed y-axis ranges are not shown for clarity.

3.2 Local dissipation ratio of internal tides in the Nansha Islands

In contrast to the LS, the internal tidal energy budget of the Nansha Islands is characterized by low local generation, strong energy convergence, and a high value of q^* . The local generation of internal tides in the Nansha Islands is weak, fluctuating between 0.6 and 1.3 GW, whereas the incoming internal tidal energy flux can reach 1.5 GW (Fig. 4g). This energy budget (dominated by far-field input) directly elevates q^* values to 1.0–3.0 (Fig. 4h). Both the local generation and the energy flux of the diurnal tide dominate those of the semidiurnal tide (Fig. 4g), maintaining the q^* of the diurnal tide in the range of 1.5–3.0, which is larger than the q^* of the semidiurnal tide varying between 1.0 and 2.0 (Fig. 4h).

Both the generation and local dissipation of internal tides in the Nansha Islands exhibit a scattered patchy pattern, with high values concentrated around reefs (Fig. 4a–b and 4e–f). All the energy budget terms of the diurnal tide are larger than those of the semidiurnal tide (Fig. 4a–f), with a significant incoming diurnal energy flux of approximately 1.5 kW m^{-1} at the southwestern and northern boundaries of the Nansha Islands (Fig. 4c). Kelly et al. (2013) estimated that at continental margins, incident mode-1 internal tidal energy is partitioned among reflection, transmission, and scattering to higher modes. Lahaye et al. (2020) demonstrated that incident low-mode internal tides can be scattered to higher modes through interactions with steep topography. Subsequently, high-mode internal tides are susceptible to breaking near their generation sites (Siyanbola et al., 2024). Modal scattering redistributes internal tidal energy rather than dissipating it; the breaking of the scattered high modes constitutes the irreversible dissipation. It is therefore reasonable that far-field internal tidal beams trigger intense topographic scattering, which supplies sufficient energy to sustain a high local dissipation ratio ($q^* > 1$) despite weak local generation in this region.

In terms of vertical modes, the internal tidal energy budget terms in the Nansha Islands are significantly smaller in magnitude than those in the LS, with strong low-mode energy convergence and high values of q^* . The energy budget in the Nansha Islands is dominated by modes 1 and 2 (Fig. 5j–k). Unlike in the LS, here the local generation of mode 1 is weaker than that of mode 2 (Fig. 5j–k). This energy budget, dominated by far-field input of mode-1 energy flux, directly results in the q^* values of mode 1 exceeding those of mode 2, as shown in Fig. 5l–m. The energy budget terms for modes 3–5 are all near zero with highly variable q^* (Fig. 5j–m), indicating that local generation into these modes contributes little to the regional dissipation. Diurnal tides dominate in the Nansha Islands, with all energy budget terms exceeding those of the semidiurnal tide and showing pronounced differences between January and August 2012 (Fig. 5j–k). The q^* values of the mode-1 diurnal tide can exceed 4 (Fig. 5l).

In the Nansha Islands, all internal tidal energy budget terms decrease with increasing mode number, except for the mode-2 energy conversion (Fig. 5a–i). The mode-3 energy budget terms are generally less than one-third of those for modes 1–2. Two dominant incoming energy fluxes exist in the Nansha Islands, originating from the northeast and the southwest,

respectively. With increasing mode number, the energy flux from the southwestern shelf slope of the SCS decays more rapidly than that from the northeastern source, suggesting that the reflection from the southwestern shelf slope of the SCS may furnish local dissipation in the Nansha Islands. Siyanbola et al. (2024) pointed out that in weak generation regions, intense local dissipation depends on sufficient convergence and slope reflection of far-field low-mode internal tidal beams. This is consistent with our results that the elevated mode-1 dissipation and q^* (> 1) in the Nansha Islands are potentially dominated by the incoming mode-1 energy flux, particularly the reflection from the southwestern shelf slope of the SCS.

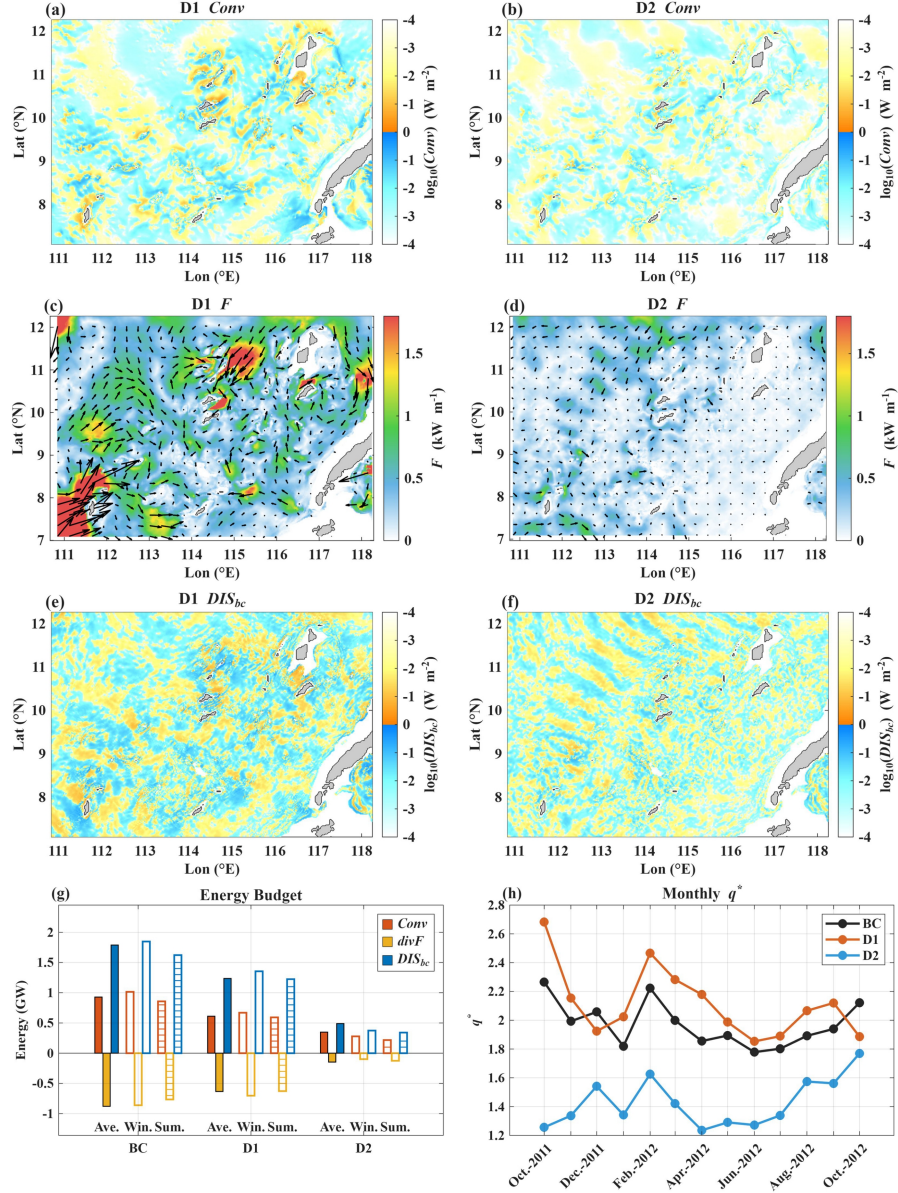
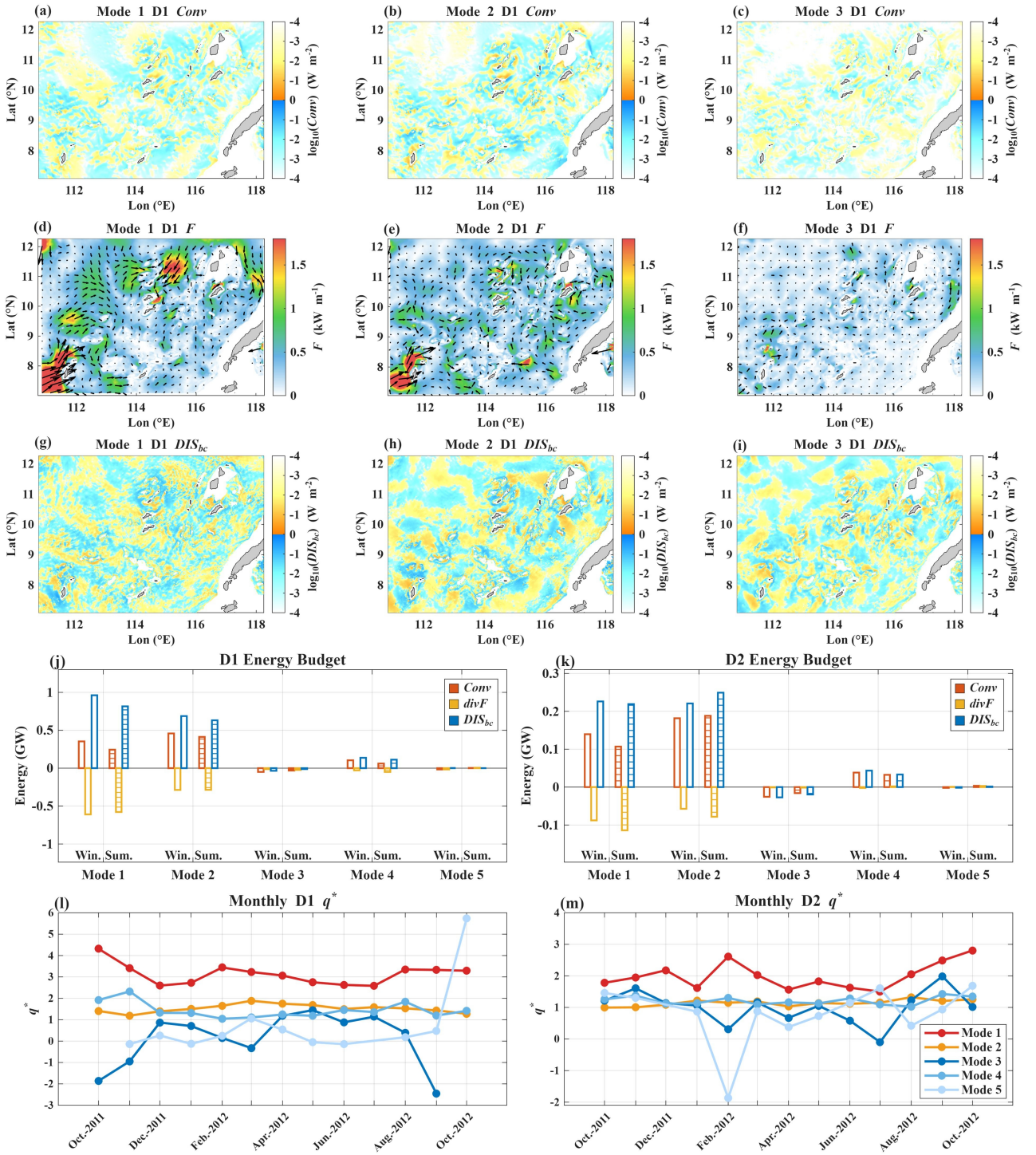


Figure 4. Same as Fig. 2, but for the Nansha Islands.



280 Figure 5. Same as Fig. 3, but for the Nansha Islands and (a-i) for diurnal internal tide.

3.3 Parameterizations for the local dissipation ratio of internal tides

3.3.1 Parameterization constrained by physical factors

To develop parameterizations for q^* that are applicable to dynamically distinct regions, 3 key physical parameters are chosen based on the features of diurnal and semidiurnal tides in the LS and the Nansha Islands. The depth-averaged internal tidal shear (S^2) within the bottommost 500 m represents the local internal wave intensity; the barotropic tidal kinetic energy (E_{BT}) quantifies the local tidal forcing strength; and the mode-1 internal tidal energy ($E^{(1)}$) is used to represent the potential for radiating low-mode energy to the far field. A dimensionless power-law function is fitted to the 429-day time series of q^* via multiple regression, with the specific formula expressed as

$$q_{\text{para}}^* = q_0 \left(\frac{S^2}{S_0^2} \right)^{a_1} \left(\frac{E_{bt}}{E_{bt0}} \right)^{a_2} \left(\frac{E^{(1)}}{E_0^{(1)}} \right)^{a_3} + q_b, \quad (5)$$

where S_0^2 , E_{bt0} and $E_0^{(1)}$ are the time-mean reference values of each variable within each region; q_0 is the amplitude coefficient, q_b is a constant representing the background state; and a_1 , a_2 and a_3 are the power-law exponents corresponding to the internal tidal shear, barotropic tidal kinetic energy, and mode-1 internal tidal energy, respectively.

Table 1 presents the fitted coefficients together with the coefficients of determination (R^2) and root mean square error (RMSE) for this physical parameterization of q^* for diurnal and semidiurnal tides in both the LS and the Nansha Islands. R^2 and RMSE are indicators of the overall model performance rather than errors or uncertainties associated with individual coefficients. In the LS, R^2 for the diurnal and semidiurnal tides are both 0.35, with RMSE of 0.14 and 0.08, respectively. The fitted parameters characterize the dynamical process modulating q^* in the LS. For both diurnal and semidiurnal components, a_2 representing the sensitivity to barotropic tidal kinetic energy is positive (0.40 and 0.37, respectively), whereas a_3 representing the sensitivity to mode-1 internal tidal energy is negative (-0.31 and -0.40, respectively). Stronger local barotropic forcing therefore enhances internal tide generation and local dissipation, leading to a positive modulation of q^* . Conversely, when mode-1 internal tide energy accumulates substantially, the internal tide tends to propagate to the far field as coherent beams, thereby reducing q^* relatively. The background state constant q_b in the LS (0.05 and 0.15, respectively) is smaller than the amplitude coefficient q_0 (0.44 and 0.55, respectively). As shown in Fig. 6, the parameterization captures the mean state of q^* in the LS, although it exhibits some smoothing toward extreme values, manifesting as overestimation at low q^* and underestimation at high q^* .

The parameterized q^* in the Nansha Islands reveals a dynamical process different from that of the LS. As shown in Table 1, the R^2 values for diurnal and semidiurnal tides are 0.26 and 0.48, respectively. For both diurnal and semidiurnal components, a_2 for barotropic tidal kinetic energy is negative (-0.98 and -2.13, respectively), whereas a_3 for mode-1 energy is positive (0.91 and 2.05, respectively). This indicates that when weakened local barotropic forcing reduces internal tide generation, q^* tends to increase, provided that the far-field low-mode energy remains elevated and drives intense local dissipation. a_1 for the internal

tidal shear is negative for both diurnal and semidiurnal components in the LS (−0.27 and −0.02, respectively), whereas in the Nansha Islands it is positive for the diurnal component (0.57) and negative for the semidiurnal component (−0.48). The negative values do not indicate that enhanced internal tidal shear suppresses local dissipation; rather, the regionally averaged shear cannot capture local dissipation around the double ridges in the LS and the reefs in the Nansha Islands. In contrast to the

315 LS, the background constants q_b for diurnal and semidiurnal tides in the Nansha Islands (1.47 and 1.06, respectively) substantially exceed the corresponding amplitude coefficients q_0 (0.69 and 0.39, respectively). This suggests that topographic scattering and reflection by the widespread seamounts and reefs, together with the incoming low-mode internal tidal energy, sustain q^* at a high level above 1.

Table 1. Fitted coefficients together with R^2 and RMSE of the physical parameterization for q^* in the LS and the Nansha Islands.

Region-Band	q_0	q_b	a_1	a_2	a_3	R^2	RMSE
LS-D1	0.44	0.05	-0.27	0.40	-0.31	0.35	0.14
LS-D2	0.55	0.15	-0.02	0.37	-0.40	0.35	0.08
Nansha-D1	0.69	1.47	0.57	-0.98	0.91	0.26	0.67
Nansha-D2	0.39	1.06	-0.48	-2.13	2.05	0.48	0.48

320

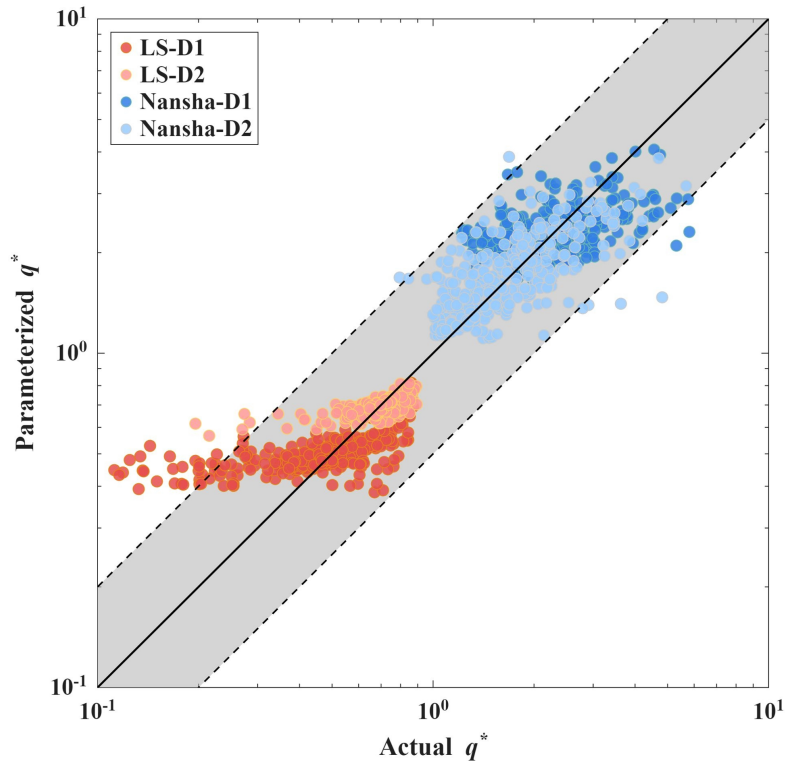


Figure 6. Scatter plot of parameterized q^* versus actual q^* for the LS and the Nansha Islands. The black line denotes the 1:1 reference line, and the two outer black dashed lines represent the 2:1 and 1:2 bounds.

325 **3.3.2 Parameterization based on the XGBoost model**

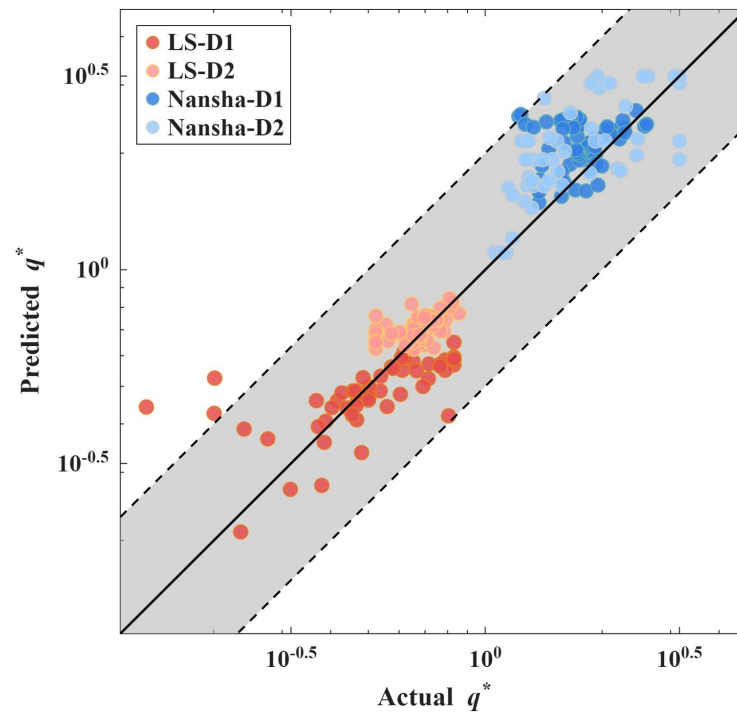
The XGBoost model achieves solid predictive performance on the combined testing set spanning both regions and tidal components, with an R^2 of 0.70 and an RMSE of 0.39 (Table 2). When predicting the semidiurnal component independently, the R^2 reaches 0.71. As shown in Fig. 7, the great overall performance of the model primarily stems from its ability to achieve clear macroscopic clustering across regions, separating the data into a low- q^* cluster for the LS and a high- q^* cluster for the Nansha Islands.

However, the predictive performance of the XGBoost model degrades markedly when predicting individual regions and tidal components separately. Table 2 shows that the R^2 values for diurnal and semidiurnal tides in the LS decrease to 0.39 and 0.11, respectively. The R^2 for the semidiurnal tide in the Nansha Islands is 0.30, and that for the diurnal tide is negative. As shown in Fig. 7, the model predictions for the LS are relatively concentrated and clustered along the 1:1 reference line, whereas those for the Nansha Islands exhibit greater dispersion. This degradation likely arises from the fact that extreme q^* values are highly sensitive to the topographic scattering of far-field low-mode internal tides and nonlinear wave-wave interactions.

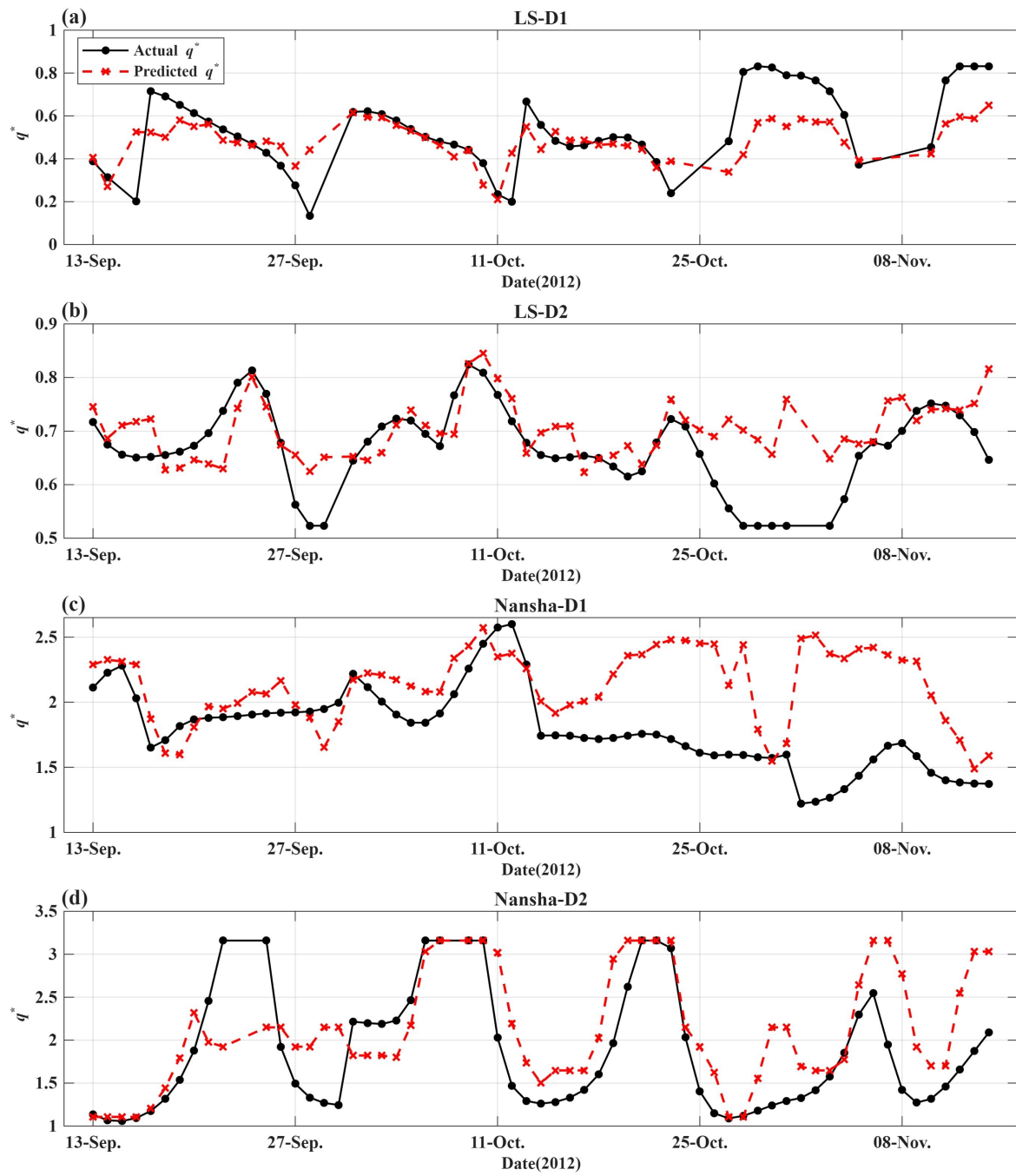
As noted by Zhao et al. (2010), long-range internal tides form spatially complex interference patterns through superposition from multiple sources, with phase differences further modulated by mesoscale refraction. These far-field inputs and mesoscale processes are not captured by the local area-averaged input features employed here, which inevitably smooth out interference signals. **Consequently, the XGBoost model broadly captures the phase of q^* variability over the spring-neap tidal cycles and predicts the temporal evolution of q^* well for the diurnal tide in the LS and the semidiurnal tide in the Nansha Islands (Fig. 8), although its predictions of q^* exhibit a degree of smoothing.**

Table 2. Performance of the XGBoost model on the testing set.

Region	Band	RMSE	R^2
LS, Nansha	D1, D2	0.39	0.70
LS, Nansha	D1	0.38	0.69
LS, Nansha	D2	0.41	0.71
LS	D1	0.14	0.39
LS	D2	0.08	0.11
Nansha	D1	0.50	-1.69
Nansha	D2	0.57	0.30



345 **Figure 7.** Scatter plot of the predicted q^* from the XGBoost model versus the actual q^* for the LS and the Nansha Islands. The black line represents the 1:1 reference line, and the two outer black dashed lines represent the 2:1 and 1:2 bounds.



350 **Figure 8.** Time series of the predicted q^* from the XGBoost model and the actual q^* for the diurnal (D1) and semidiurnal (D2) tides in the LS and the Nansha Islands.

4 Summary

Based on the LLC4320 simulation, we reveal differences in the local internal tidal dissipation ratio (q^*) between the source region (represented by the LS) and the sink region (represented by the Nansha Islands). In the LS, the q^* value for the total internal tides ranges from 0.3 to 0.7, which is consistent with earlier results. The mode-1 energy flux radiates intensively out of the LS, resulting in mode-1 q^* values smaller than those of higher modes. This confirms that in generation regions of internal tides, much of the baroclinic energy propagates to the far field as low modes rather than dissipating locally. Conversely, in the Nansha Islands, q^* value for the total internal tides is generally larger than 1. In particular, the q^* values of the mode-1 diurnal and semidiurnal tides are consistently larger than those of higher modes. The reason for $q^* > 1$ is that the far-field internal tidal energy flux entering the Nansha Islands exceeds the locally generated one. Thus, the primary source for local dissipation is not the locally generated internal tide, but rather far-field internal tidal beams propagating from the LS.

The physical-based empirical formulation and the XGBoost model used to scale q^* in this study successfully capture the macroscopic clustering of the mean states of q^* in both the LS and the Nansha Islands. They also reproduce some of the periodic temporal variation of q^* , with an R^2 value reaching 0.70. Moreover, the physical parameterization results reveal regional differences in the dominant mechanisms modulating q^* . In the LS, while intense barotropic forcing generates energetic internal tides and enhances local dissipation, the substantial accumulation of mode-1 internal tidal energy favors the radiation of coherent beams into the far field, thereby reducing q^* . In the Nansha Islands, the negative power-law coefficient (a_2) for barotropic tidal kinetic energy, along with the positive coefficient (a_3) for mode-1 internal tidal energy, indicates that local dissipation is dominated by far-field energy input. Furthermore, the scattering and reflection of internal tides by the widespread seamounts and reefs play a dominant role in sustaining high q^* in the Nansha Islands. On the other hand, the XGBoost model yields a consistent periodic temporal change and accurately predicts the mean q^* value in both source and sink regions, despite some quantitative discrepancies. However, those far-field inputs and mesoscale processes are not captured by the local area-averaged input features, resulting in a degree of smoothing in the predicted q^* . This insight suggests that future parameterizations of q^* and internal tide mixing must move beyond the traditional framework that relies solely on local generation; instead, they should explicitly and thoughtfully incorporate the nonlocal processes, e.g., the far-field propagation of internal wave field experiencing interference and scattering.

The present study has several limitations. The horizontal grid spacing of LLC4320 is approximately 2 km in the study regions. At this resolution, the major topographic features and low-mode internal tides are explicitly resolved, whereas small-scale topographic features, the high-mode internal tides with short horizontal wavelengths, the internal wave breaking and the associated turbulence remain subgrid-scale processes and hence are not resolved. Therefore, internal tide generation driven by the unresolved roughness is absent from the simulation. Energy loss associated with unresolved scales may enter the model energy budget through parameterized viscosity and diffusivity and numerical damping, while bottom drag provides an additional explicit energy sink. In addition, because the present dissipation estimate (DIS_{bc}) is obtained as the residual of the

simplified baroclinic energy budget equation, it may conflate dissipation associated with parameterized viscosity, diffusivity, and bottom drag; numerical mixing and damping. The archived model output does not permit an independent closure of the full energy budget, and these contributions cannot be quantified separately. Therefore, DIS_{bc} represents an inferred dissipation integrated over the study regions rather than diagnosed dissipation from the model at individual topographic features. Since DIS_{bc} appears in the numerator of q^* , its uncertainty propagates directly into the estimate of q^* . Crucially, the relative uncertainty of q^* is amplified where the area-integrated local conversion is weak. Such uncertainty is particularly pronounced over rough topography and for higher-mode internal tides (Vic et al., 2019). Future work with higher-resolution simulations and diagnostics of the individual dissipative terms is expected to separate and quantify these contributions.

Data availability

The MITgcm LLC4320 data that support the findings of this study are openly available at https://data.nas.nasa.gov/ecco/data.php?dir=/eccodata/llc_4320.

Author contributions

The study was conceived and designed by QY. Data preparation, material collection, and analysis were performed by ZY and HS. ZY and HS prepared the manuscript with contributions from all co-authors.

Competing interests

The contact author has declared that none of the authors has any competing interests.

Disclaimer

Publisher's note: Copernicus Publications remains neutral with regard to jurisdictional claims made in the text, published maps, institutional affiliations, or any other geographical representation in this paper. While Copernicus Publications makes every effort to include appropriate place names, the final responsibility lies with the authors. Views expressed in the text are those of the authors and do not necessarily reflect the views of the publisher.

Acknowledgements

405 Financial support

This work was supported by the National Natural Science Foundation of China (Grants 42376012, 42506011, 42076012 and 42006012).

References

- Alford, M. H., Simmons, H. L., Marques, O. B., and Girtton, J. B.: Internal tide attenuation in the North Pacific, *Geophys. Res. Lett.*, 46, 8205–8213, <https://doi.org/10.1029/2019GL082648>, 2019.
- Arbic, B. K., Alford, M., Ansong, J., Buijsman, M., Ciotti, R., Farrar, J., Hallberg, R., Henze, C., Hill, C., Luecke, C., Menemenlis, D., Metzger, E., Müller, M., Nelson, A., Nelson, B., Ngodock, H., Ponte, R., Richman, J., Savage, A., and Zhao, Z.: A primer on global internal tide and internal gravity wave continuum modeling in HYCOM and MITgcm, in *New Frontiers in Operational Oceanography*, edited by: Chassignet, E. P., Pascual, A., Tintoré, J., and Verron, J., GODAE Ocean View, 307–392, <https://doi.org/10.17125/gov2018.ch13>, 2018.
- Buijsman, M. C., Klymak, J. M., Legg, S., Alford, M. H., Farmer, D., MacKinnon, J. A., Nash, J. D., Park, J.-H., Pickering, A., and Simmons, H.: Three-dimensional double-ridge internal tide resonance in Luzon Strait, *J. Phys. Oceanogr.*, 44, 850–869, <https://doi.org/10.1175/JPO-D-13-024.1>, 2014.
- Cimoli, L., Caulfield, C. P., Johnson, H. L., Marshall, D. P., Mashayek, A., Naveira Garabato, A. C., and Vic, C.: Sensitivity of deep ocean mixing to local internal tide breaking and mixing efficiency, *Geophys. Res. Lett.*, 46, 14622–14633, <https://doi.org/10.1029/2019GL085056>, 2019.
- de Lavergne, C., Vic, C., Madec, G., Roquet, F., Waterhouse, A. F., Whalen, C. B., Cuypers, Y., Bouruet-Aubertot, P., Ferron, B. and Hibiya, T.: A parameterization of local and remote tidal mixing, *J. Adv. Model. Earth Sy.*, 12, e2020MS002065, <https://doi.org/10.1029/2020MS002065>, 2020.
- Dunphy, M. and Lamb, K. G.: Focusing and vertical mode scattering of the first mode internal tide by mesoscale eddy interaction, *J. Geophys. Res. Oceans*, 119, 523–536, <https://doi.org/10.1002/2013JC009293>, 2014.
- Egbert, G. D. and Ray, R. D.: Significant dissipation of tidal energy in the deep ocean inferred from satellite altimeter data, *Nature*, 405, 775–778, <https://doi.org/10.1038/35015531>, 2000.
- Egbert, G. D. and Ray, R. D.: Semi-diurnal and diurnal tidal dissipation from TOPEX/Poseidon altimetry, *Geophys. Res. Lett.*, 30, 1907, <https://doi.org/10.1029/2003GL017676>, 2003.
- Fan, L., Sun, H., Yang, Q., and Li, J.: Numerical investigation of interaction between anticyclonic eddy and semidiurnal internal tide in the northeastern South China Sea, *Ocean Sci.*, 20, 241–264, <https://doi.org/10.5194/os-20-241-2024>, 2024.
- Ferrari, R. and Wunsch, C.: Ocean circulation kinetic energy: reservoirs, sources, and sinks, *Annu. Rev. Fluid Mech.*, 41, 253–282, <https://doi.org/10.1146/annurev.fluid.40.111406.102139>, 2009.

- 435 Garrett, C. and Kunze, E.: Internal tide generation in the deep ocean, *Annu. Rev. Fluid Mech.*, 39, 57–87,
<https://doi.org/10.1146/annurev.fluid.39.050905.110227>, 2007.
- Gill, A.: Atmosphere-ocean dynamics, in: 1st Edn., Academic Press, New York, 680 pp., ISBN 978-0-12-283522-3, 1982.
- Jan, S., Chern, C., Wang, J., and Chao, S.: Generation of diurnal K_1 internal tide in the Luzon Strait and its influence on surface
 tide in the South China Sea, *J. Geophys. Res.*, 112, C06019, <https://doi.org/10.1029/2006JC004003>, 2007.
- 440 Jayne, S. R.: The impact of abyssal mixing parameterizations in an ocean general circulation model, *J. Phys. Oceanogr.*, 39,
 1756–1775, <https://doi.org/10.1175/2009JPO4085.1>, 2009.
- Kelly, S. M., Jones, N. L., Nash, J. D., and Waterhouse, A. F.: The geography of semidiurnal mode-1 internal-tide energy loss,
Geophys. Res. Lett., 40, 4689–4693, <https://doi.org/10.1002/grl.50872>, 2013.
- Klymak, J. M., Legg, S., and Pinkel, R.: A simple parameterization of turbulent tidal mixing near supercritical topography, *J.*
 445 *Phys. Oceanogr.*, 40, 2059–2074, <https://doi.org/10.1175/2010JPO4396.1>, 2010.
- Klymak, J. M., Alford, M. H., Pinkel, R., Lien, R., Yang, Y. J., and Tang, T.: The breaking and scattering of the internal tide
 on a continental slope, *J. Phys. Oceanogr.*, 41, 926–945, <https://doi.org/10.1175/2010JPO4500.1>, 2011.
- Lahaye, N., Gula, J., and Roullet, G.: Internal tide cycle and topographic scattering over the North Mid-Atlantic Ridge, *J.*
Geophys. Res. Oceans, 125, e2020JC016376, <https://doi.org/10.1029/2020JC016376>, 2020.
- 450 Lefauve, A., Muller, C., and Melet, A.: A three-dimensional map of tidal dissipation over abyssal hills, *J. Geophys. Res.*
Oceans, 120, 4760–4777, <https://doi.org/10.1002/2014JC010598>, 2015.
- MacKinnon, J. A., Zhao, Z., Whalen, C. B., Waterhouse, A. F., Trossman, D. S., Sun, O. M., St. Laurent, L. C., Simmons, H.
 L., Polzin, K., Pinkel, R., Pickering, A., Norton, N. J., Nash, J. D., Musgrave, R., Merchant, L. M., Melet, A. V., Mater,
 B., Legg, S., Large, W. G., Kunze, E., Klymak, J. M., Jochum, M., Jayne, S. R., Hallberg, R. W., Griffies, S. M., Diggs,
 455 S., Danabasoglu, G., Chassignet, E. P., Buijsman, M. C., Bryan, F. O., Briegleb, B. P., Barna, A., Arbic, B. K., Ansong,
 J. K., and Alford, M. H.: Climate process team on internal wave-driven ocean mixing, *Bull. Amer. Meteor. Soc.*, 98,
 2429–2454, <https://doi.org/10.1175/BAMS-D-16-0030.1>, 2017.
- Marshall, J., Adcroft, A., Hill, C., Perelman, L., and Heisey, C.: A finite-volume, incompressible Navier Stokes model for
 studies of the ocean on parallel computers, *J. Geophys. Res.*, 102(C3), 5753–5766, <https://doi.org/10.1029/96JC02775>,
 460 1997.
- Melet, A., Legg, S., and Hallberg, R.: Climatic impacts of parameterized local and remote tidal mixing, *J. Climate*, 29, 3473–
 3500, <https://doi.org/10.1175/JCLI-D-15-0153.1>, 2016.
- Mellor, G. L.: Users guide for a three-dimensional primitive equation numerical ocean model, Princeton Univ., Princeton, NJ,
 08544–10710, 2004.
- 465 Munk, W. H. and Wunsch, C.: Abyssal recipes II: energetics of tidal and wind mixing, *Deep-Sea Res.*, 45, 1977–2010,
[https://doi.org/10.1016/S0967-0637\(98\)00070-3](https://doi.org/10.1016/S0967-0637(98)00070-3), 1998.
- Nikurashin, M. and Legg, S.: A mechanism for local dissipation of internal tides generated at rough topography, *J. Phys.*
Oceanogr., 41, 378–395, <https://doi.org/10.1175/2010JPO4522.1>, 2011.

- Niwa, Y. and Hibiya, T.: Three-dimensional numerical simulation of M_2 internal tides in the East China Sea, *J. Geophys. Res.*, 109, C04027, <https://doi.org/10.1029/2003JC001923>, 2004.
- Nycander, J.: Generation of internal waves in the deep ocean by tides, *J. Geophys. Res.*, 110, C10028, <https://doi.org/10.1029/2004JC002487>, 2005.
- Polzin, K.: A heuristic description of internal wave dynamics, *J. Phys. Oceanogr.*, 34, 214–230, [https://doi.org/10.1175/1520-0485\(2004\)034<0214:AHDOIW>2.0.CO;2](https://doi.org/10.1175/1520-0485(2004)034<0214:AHDOIW>2.0.CO;2), 2004.
- Saenko, O. A. and Merryfield, W. J.: On the effect of topographically enhanced mixing on the global ocean circulation, *J. Phys. Oceanogr.*, 35, 826–834, <https://doi.org/10.1175/JPO2722.1>, 2005.
- Simmons, H. L., Jayne, S. R., St. Laurent, L. C., and Weaver, A. J.: Tidally driven mixing in a numerical model of the ocean general circulation, *Ocean Modell.*, 6, 245–263, [https://doi.org/10.1016/S1463-5003\(03\)00011-8](https://doi.org/10.1016/S1463-5003(03)00011-8), 2004.
- Siyanbola, O. Q., Buijsman, M. C., Delpech, A., Barkan, R., Pan, Y., and Arbic, B. K.: Interactions of remotely generated internal tides with the U.S. West Coast continental margin, *J. Geophys. Res. Oceans*, 129, e2023JC020859, <https://doi.org/10.1029/2023JC020859>, 2024.
- St. Laurent, L. C. and Garrett, C.: The role of internal tides in mixing the deep ocean, *J. Phys. Oceanogr.*, 32, 2882–2899, [https://doi.org/10.1175/1520-0485\(2002\)032<2882:TROITI>2.0.CO;2](https://doi.org/10.1175/1520-0485(2002)032<2882:TROITI>2.0.CO;2), 2002.
- St. Laurent, L. C., Simmons, H. L., and Jayne, S. R.: Estimating tidally driven mixing in the deep ocean, *Geophys. Res. Lett.*, 29, 2106, <https://doi.org/10.1029/2002GL015633>, 2002.
- Tian, Y., Bai, X., Wang, C., and Liu, Z.: Tidal energetics in the eddying South China Sea from a high-resolution numerical simulation, *Prog. Oceanogr.*, 231, 103418, <https://doi.org/10.1016/j.pocean.2025.103418>, 2025.
- Vic, C., Naveira Garabato, A. C., Green, J. A. M., Spingys, C., Forryan, A., Zhao, Z., and Sharples, J.: The lifecycle of semidiurnal internal tides over the Northern Mid-Atlantic Ridge, *J. Phys. Oceanogr.*, 48, 61–80, <https://doi.org/10.1175/JPO-D-17-0121.1>, 2018.
- Vic, C., Naveira Garabato, A. C., Green, J. A. M., Waterhouse, A. F., Zhao, Z., Melet, A., de Lavergne, C., Buijsman, M. C., and Stephenson, G. R.: Deep-ocean mixing driven by small-scale internal tides, *Nat. Commun.*, 10, 2099, <https://doi.org/10.1038/s41467-019-10149-5>, 2019.
- Wang, X., Peng, S., Liu, Z., Huang, R. X., Qian, Y., and Li, Y.: Tidal mixing in the South China Sea: an estimate based on the internal tide energetics, *J. Phys. Oceanogr.*, 46, 107–124, <https://doi.org/10.1175/JPO-D-15-0082.1>, 2016.
- Wang, Y. and Legg, S.: Enhanced dissipation of internal tides in a mesoscale baroclinic eddy, *J. Phys. Oceanogr.*, 53, 2293–2316, <https://doi.org/10.1175/JPO-D-23-0045.1>, 2023.
- Wang, Y. and Legg, S.: Agulhas rings locally enhance dissipation of internal tides, *Sci. Adv.*, 11, eadq5963, <https://doi.org/10.1126/sciadv.adq5963>, 2025.
- Waterhouse, A. F., MacKinnon, J. A., Nash, J. D., Alford, M. H., Kunze, E., Simmons, H. L., Polzin, K. L., St. Laurent, L. C., Sun, O. M., Pinkel, R., Talley, L. D., Whalen, C. B., Huussen, T. N., Carter, G. S., Fer, I., Waterman, S., Naveira Garabato,

- A. C., Sanford, T. B., and Lee, C. M.: Global patterns of diapycnal mixing from measurements of the turbulent dissipation rate, *J. Phys. Oceanogr.*, 44, 1854–1872, <https://doi.org/10.1175/JPO-D-13-0104.1>, 2014.
- Wunsch, C. and Ferrari, R.: Vertical mixing, energy, and the general circulation of the oceans, *Annu. Rev. Fluid Mech.*, 36, 281–314, <https://doi.org/10.1146/annurev.fluid.36.050802.122121>, 2004.
- Xu, Z., Liu, K., Yin, B., Zhao, Z., Wang, Y., and Li, Q.: Long-range propagation and associated variability of internal tides in the South China Sea, *J. Geophys. Res. Oceans*, 121, 8268–8286, <https://doi.org/10.1002/2016JC012105>, 2016.
- You, J., Xu, Z., Li, Q., Zhang, P., Yin, B., and Hou, Y.: M_2 Internal tide energetics and behaviors in the subpolar North Pacific, *J. Phys. Oceanogr.*, 53, 1269–1290, <https://doi.org/10.1175/JPO-D-22-0032.1>, 2023.
- Zhao, Z., Alford, M. H., MacKinnon, J. A., and Pinkel, R.: Long-range propagation of the semidiurnal internal tide from the Hawaiian Ridge, *J. Phys. Oceanogr.*, 40, 713–736, <https://doi.org/10.1175/2009JPO4207.1>, 2010.