



1 **Long-term evolution and type-specific rainy-season activity**

2 **of post-seismic geohazards in Southwest China**

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Abstract:

9 Long-term evolution of post-seismic geohazards represents a critical yet complex
10 challenge in earthquake-affected regions. This study systematically investigates the
11 spatiotemporal patterns of post-seismic geohazard evolution following three major seismic
12 events: the 2008 Wenchuan earthquake (Mw7.9), 2013 Lushan earthquake (Mw6.6), and 2017
13 Jiuzhaigou earthquake (Mw7.0), focusing on four representative counties (Wenchuan,
14 Beichuan, Lushan, and Jiuzhaigou County) in Southwest China. The results show that: (1)
15 post-seismic geohazard activity showed an intermittent clustered outbreak pattern; (2)
16 geohazards were mainly concentrated in valley-related geomorphic units, with county-level
17 differences characterized by reactivation–expansion, spatial inheritance, localization, and
18 contraction; (3) monthly records excluding earthquake-month records showed a clear rainy-
19 season concentration, with 78.1% of geohazards occurring from June to August; and (4)
20 Beichuan and Lushan showed clearer weakening trends, whereas Wenchuan and Jiuzhaigou
21 remained in low-level persistent activity stages and should receive continued monitoring
22 attention. These findings provide useful insights for long-term geohazard monitoring and risk
23 mitigation in earthquake-affected mountainous regions.

24 **Keywords:** Post-seismic geohazards; Intermittent clustered outbreak; Spatial evolution;
25 Rainy-season activity; Type-specific response

26 1. Introduction

27 Large earthquakes often trigger cascading geological hazards that pose long-term threats
28 to human settlements and critical infrastructure. The 2008 Wenchuan earthquake (Mw 7.9),
29 for example, generated over 56,000 landslides and 200,000 rock falls across an affected area



30 exceeding 110,000 km² (Huang & Li, 2009; Dai et al., 2011; Gorum et al., 2011; Fan et al.,
31 2018). Similar patterns were observed following the 2013 Lushan (Mw 6.6) and 2017
32 Jiuzhaigou (Mw 7.0) earthquakes, which induced tens of thousands of secondary geohazards,
33 particularly in mountainous regions (Xu & Xu, 2014; Fan et al., 2019). Globally, seismic
34 events such as the 2010 Maule (Mw 8.8), 2016 Kaikōura (Mw 7.8), and 2009 L'Aquila (Mw
35 6.3) earthquakes have demonstrated comparable cascading effects, triggering widespread
36 slope instabilities that persist for years (Moro et al., 2016; Massey et al., 2018; Basilone et al.,
37 2019; Serey et al., 2019; Tanyaş et al., 2022).

38 Beyond immediate coseismic effects, strong ground shaking induces profound geological
39 and ecological disturbances. Seismic waves fracture bedrock, weaken slope structures, and
40 generate deep-seated discontinuities that may later evolve into failure surfaces of rock masses
41 (Huang et al., 2012; He et al., 2020). Post-earthquake assessments in the Wenchuan region
42 revealed extensive rock mass fragmentation, with fracture densities increasing by factors of 2-
43 5 near fault zones (Parker et al., 2011; Zhang et al., 2015; Chen et al., 2020; Chen et al., 2022;
44 Xiong et al., 2022). Comparable degradation was documented after the 1999 Chi-Chi
45 earthquake (Mw 7.6), where persistent slope instability was linked to seismic fracturing (Lin
46 et al., 2003; Hovius et al., 2011; Shou et al., 2011). Additionally, vegetation as a key
47 stabilizing factor is severely disrupted. The Wenchuan earthquake destroyed 3.6×10^5 ha of
48 forest (Liu et al., 2010; Yang et al., 2018; Yunus et al., 2020; Lv et al., 2025).

49 The resulting rock masses were characterized by high porosity and low mechanical
50 properties, and remained a high potential of rainfall-triggered failures. Prolonged weathering,
51 freeze-thaw cycles, and gravitational stresses progressively weaken these materials and



52 induced recurrent instabilities (Iverson, 1997; Marc et al., 2015). For instance, landslide
53 activity persisted for over 15 years following the 2005 Kashmir earthquake (Mw 7.6), driven
54 by ongoing weathering of fractured bedrock (Bacha et al., 2022). Moreover, these rock
55 masses serve as primary sources for rainfall-induced debris flows, and extended the hazard
56 durations (Zhu et al., 2011; Zhang & Zhang, 2017; Jin et al., 2023).

57 Global observations consistently demonstrate that post-seismic geohazards often exhibit
58 protracted activity and could persist from several years to decades. Landslides triggered by
59 the 1906 San Francisco earthquake (Mw 7.9) reactivated for over 50 years in the California
60 Coast Ranges due to rainfall-induced remobilization of seismic debris (Keefer et al., 1987). In
61 some areas, geohazard activities returned to pre-earthquake level within years (Hovius et al.,
62 2011; Minato et al., 2012; Unjoh et al., 2012; Marc et al., 2015; Ishitsuka et al., 2017). While
63 in other regions, such as the Wenchuan region, the activity of post-seismic geohazard
64 persisted for over a decade (Fan et al., 2019). Yunus et al. (2020) and Chen et al. (2020)
65 estimated a return to pre-earthquake landslide activity levels within 18-25 years for the
66 Wenchuan region. Recovery to pre-earthquake geohazard baselines typically requires two key
67 processes: erosional removal of unstable materials and vegetation regeneration to restore
68 slope integrity (Chen et al., 2020; Yunus et al., 2020; Shen et al., 2020).

69 Rainfall is widely recognized as an important external trigger for post-seismic
70 geohazards, as it can exacerbate the instability of earthquake-fragmented rock masses.
71 Infiltration of rainfall increases pore-water pressure, reduces shear strength, and facilitates
72 slope failures (Shou et al., 2011; Marc et al., 2015; Li et al., 2018; Fan et al., 2019; Chen et
73 al., 2022; Shen et al., 2020). Following the 2015 Gorkha earthquake (Mw 7.8), monsoon rains



74 triggered over 5,000 landslides in a single season, predominantly in seismically weakened
75 zones (Martha et al., 2017; Tian et al., 2020).

76 Understanding the spatiotemporal evolution of post-seismic hazards is essential for long-
77 term risk mitigation. This work systematically investigates the long-term evolution of post-
78 seismic geohazards across four earthquake-affected counties in Southwest China: Wenchuan,
79 Beichuan, Lushan, and Jiuzhaigou County. The objectives were to: (1) characterize the
80 intermittent clustered-outbreak pattern of recorded post-seismic geohazards; (2) examine the
81 rainy-season concentration of non-earthquake-month geohazard activity and the type-specific
82 responses of landslides, rock falls, and debris flows; and (3) assess regional differences in
83 long-term post-seismic geohazard persistence and recession. These analyses provide a
84 practical basis for long-term monitoring, rainy-season preparedness, and differentiated risk
85 mitigation in earthquake-affected mountainous regions.

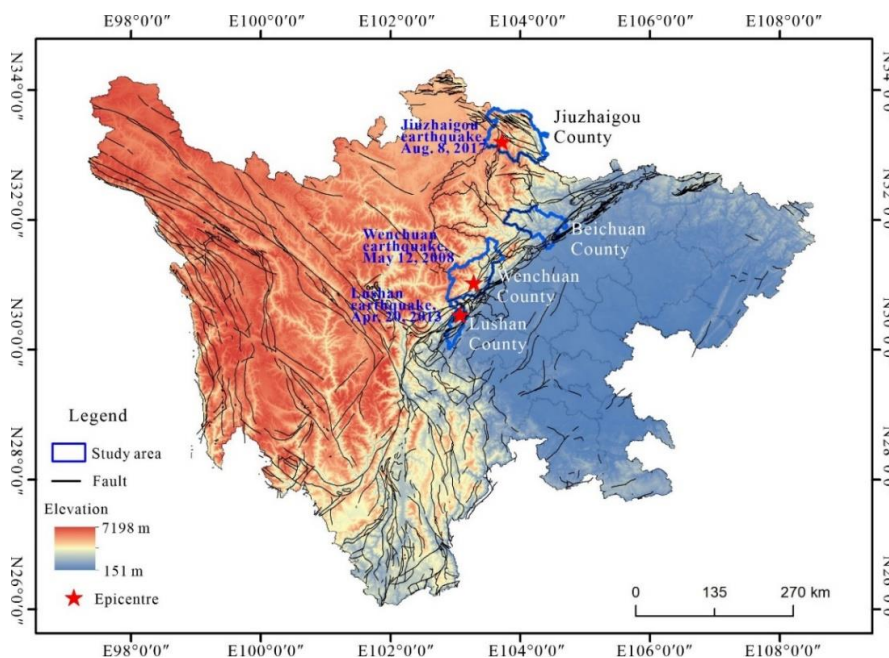
86 **2. Study area and geological setting**

87 The study area (Figure 1), including Lushan, Wenchuan, Beichuan, and Jiuzhaigou
88 counties in northwestern Sichuan Province, China, lies within the seismic zone affected by the
89 2008 Wenchuan earthquake (Mw 7.9). This region occupies a critical geotectonic position
90 along the transitional boundary between the eastern Tibetan Plateau and the Sichuan Basin. It
91 features complex structural geology that includes the prominent Longmenshan and Minjiang
92 fault systems. Since 2008, the study area has witnessed three significant seismic events: The
93 Wenchuan (Mw 7.9), Lushan (Mw 6.6), and Jiuzhaigou (Mw 7.0) earthquakes.

94 The Mw 7.9 Wenchuan earthquake on 12 May 2008 resulted from rupture along the
95 Yingxiu-Beichuan Fault, a NE-striking thrust fault with a significant right-lateral strike-slip



96 component within the Longmenshan Fault Zone. The coseismic rupture propagated over 240
97 km, producing a maximum vertical displacement of 6.2 m. On 20 April 2013, the Mw 6.6
98 Lushan earthquake ruptured the southern segment of the Longmenshan Fault Zone,
99 generating a 40 km long surface rupture dominated by thrust faulting mechanisms.
100 Subsequently, the Mw 7.0 Jiuzhaigou earthquake on 8 August 2017 occurred along a left-
101 lateral strike-slip fault within the eastern Kunlun fault system, demonstrating distinct
102 kinematic behavior compared to the preceding events.



103

104 Figure 1 Location of the four earthquake-affected counties in Southwest China.

105 Major seismic events induced widespread coseismic geohazards across the affected
106 region and caused substantial mountain fracturing and vegetation disturbance. These
107 processes generated unstable rock–soil masses with loose and fragmented structures in the
108 affected slopes (Figure 2). Subsequent rainfall may promote the reactivation or remobilization



109 of these earthquake-disturbed materials and contribute to the occurrence of post-seismic
110 geohazards (Figure 3). Therefore, the four counties provide a useful setting for examining the
111 long-term evolution of post-seismic geohazards, owing to the documented coseismic impacts
112 and the availability of long-term geohazard and rainfall records.



113
114 Figure 2 Earthquake-disturbed fractured rock masses and loose slope materials in the study

115 area.



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117 Figure 3 Representative post-seismic geohazards in the study area, including landslides, rock

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falls, and debris flows.

119 3. Data and analytical methods

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The geohazard inventory used in this study consists of statistical records from the four

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earthquake-affected counties between 2006 and 2023. These data were mainly derived from

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the long-term monitoring records of the Sichuan Provincial Geological Environment

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Monitoring Station and were checked and verified by professional geological survey teams.

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With continuous temporal coverage, the dataset can serve as a reliable and authoritative

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source for analyzing post-seismic geohazard activity. In this study, “geohazards” refers to

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three types of recorded slope-related hazards: landslides, rock falls, and debris flows.

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Landslides denote sliding failures of soil or rock–soil masses, rock falls denote rapid falling

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of rock blocks from steep slopes or cliffs, and debris flows denote water–sediment flows



129 along gullies or channels. These categories were retained from the source database and were
130 not further subdivided because consistent information on failure geometry and movement
131 style was unavailable. It should be noted that the inventory was primarily compiled for
132 geohazard prevention and risk management. Therefore, events threatening settlements, roads,
133 infrastructure, and production areas were preferentially recorded, whereas events in remote
134 uninhabited high-altitude areas were not systematically included. The dataset should therefore
135 be interpreted as a risk-oriented geohazard inventory rather than a complete geomorphic
136 inventory of all mass movements in the study area.

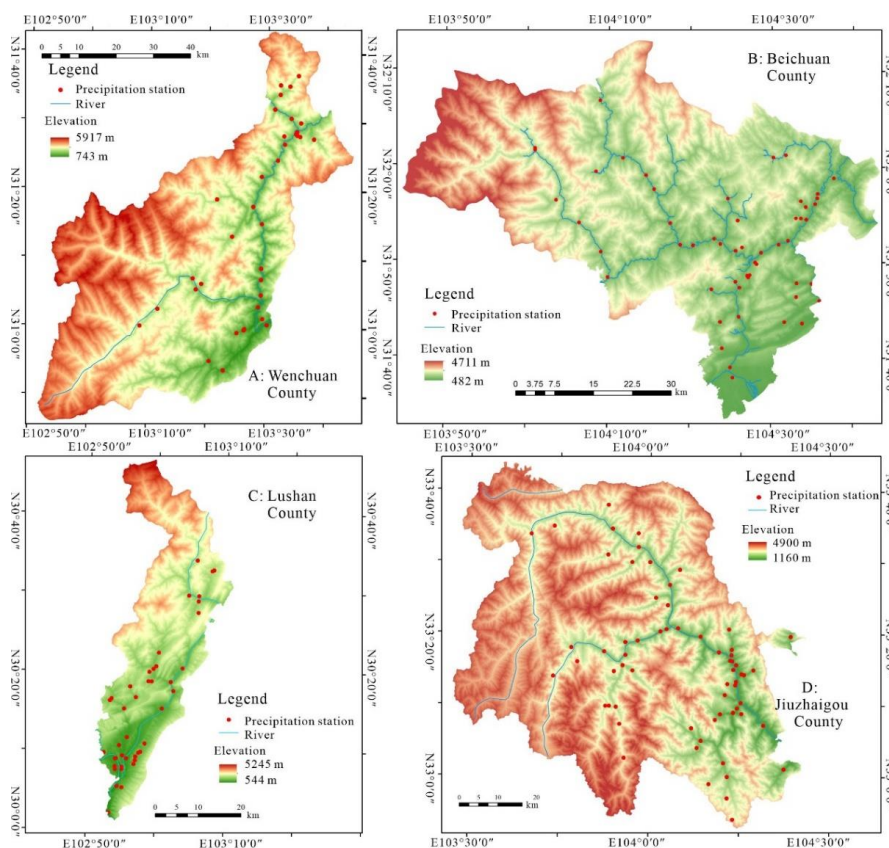
137 In this work, a total of 4,039 detailed data on geohazards were collected from the four
138 counties, among which 1,281 events were from Wenchuan County (including 537 landslides,
139 456 rock falls and 288 debris flows), 1,384 events from Beichuan County (including 1085
140 landslides, 95 rock falls and 204 debris flows), 493 events from Lushan County (including
141 164 landslides, 263 rock falls and 66 debris flows), and 881 events from Jiuzhaigou County
142 (including 270 landslides, 369 rock falls and 242 debris flows).

143 To reduce subjective identification, a clustered geohazard outbreak year was
144 operationally defined as a year in which the annual number of recorded geohazards exceeded
145 the county-specific mean annual count during the observation period. The corresponding
146 mean values were 75 for Wenchuan, 77 for Beichuan, 29 for Lushan, and 52 for Jiuzhaigou.
147 This criterion was used to identify relative high-activity years within each county, rather than
148 to define an absolute hazard-frequency threshold. The geohazard distribution maps were
149 generated in ArcGIS using the recorded event locations. The selected representative years
150 correspond to earthquake years, clustered outbreak years, or rainfall-rich years. These maps



151 show events recorded in the selected year only, rather than cumulative inventories from
152 previous years.

153 Precipitation data were obtained from a network of 293 rainfall monitoring stations
154 distributed across the study area, including 61 stations in Wenchuan County, 88 in Beichuan
155 County, 54 in Lushan County, and 90 in Jiuzhaigou County. Station rainfall data were used to
156 calculate annual and monthly rainfall series for the four counties. The spatial distribution of
157 these rainfall stations is illustrated in Figure 4. Monthly geohazard records were further
158 grouped into landslides, rock falls, and debris flows to examine the type composition of
159 clustered outbreak months.



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Figure 4 Locations of precipitation stations in the four counties.



162 **4. Results**

163 **4.1 Annual evolution of geohazards during the observation period**

164 [Figure 5](#) presents the annual counts of geohazards and the corresponding annual rainfall
165 in the four counties from 2006 to 2023.

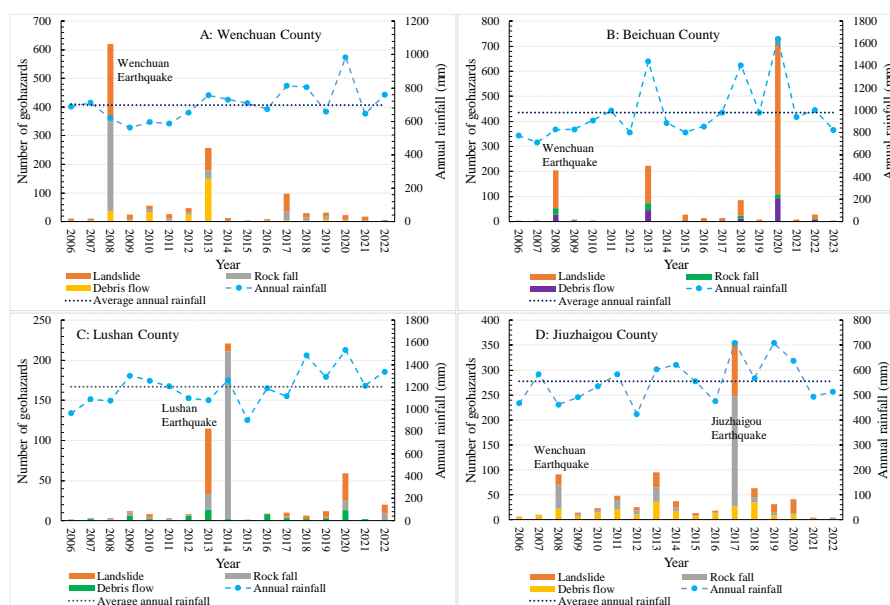
166 In Wenchuan and Beichuan counties, the 2008 Wenchuan earthquake corresponded to
167 clear annual peaks, with 621 and 205 geohazard events, respectively. In Lushan County, only
168 three events were recorded in 2008, indicating a limited response to the Wenchuan
169 earthquake, whereas the 2013 Lushan earthquake corresponded to a sharp increase to 115
170 events. In Jiuzhaigou County, the annual count increased to 91 events in 2008, indicating an
171 evident response to the Wenchuan earthquake, and reached 358 events in 2017, corresponding
172 to the Jiuzhaigou earthquake.

173 According to the county-specific mean-count criterion defined in Section 3, the clustered
174 geohazard outbreak years were identified as 2008, 2013, and 2017 for Wenchuan; 2008, 2013,
175 2018, and 2020 for Beichuan; 2013, 2014, and 2020 for Lushan; and 2008, 2013, 2017, and
176 2018 for Jiuzhaigou. Between these clustered outbreak years, relatively low-activity periods
177 were observed. For example, Wenchuan decreased to 25–55 events during 2009–2012 and 5–
178 32 events after 2017; Beichuan decreased to 1–7 events during 2009–2012; Lushan fluctuated
179 between 1 and 12 events during 2015–2019; and Jiuzhaigou fluctuated between 15 and 48
180 events during 2009–2012.

181 Overall, the four counties experienced two to three clustered geohazard outbreaks after
182 the major earthquakes. These outbreaks were separated by relatively low-activity periods
183 rather than continuous high activity, showing an intermittent outbreak pattern. In addition,



184 some clustered geohazard outbreak years coincided with years of above-average rainfall,
185 preliminarily suggesting a possible rainfall correspondence. The geohazard type compositions
186 also differed among the four counties: Wenchuan contained comparable numbers of
187 landslides, rock falls, and debris flows; Beichuan was dominated by landslides; and rock falls
188 accounted for a relatively large proportion in Jiuzhaigou and Lushan.



189
190 Figure 5 Long-term evolutions of geohazards and annual rainfalls in the four counties from
191 2006 to 2023.

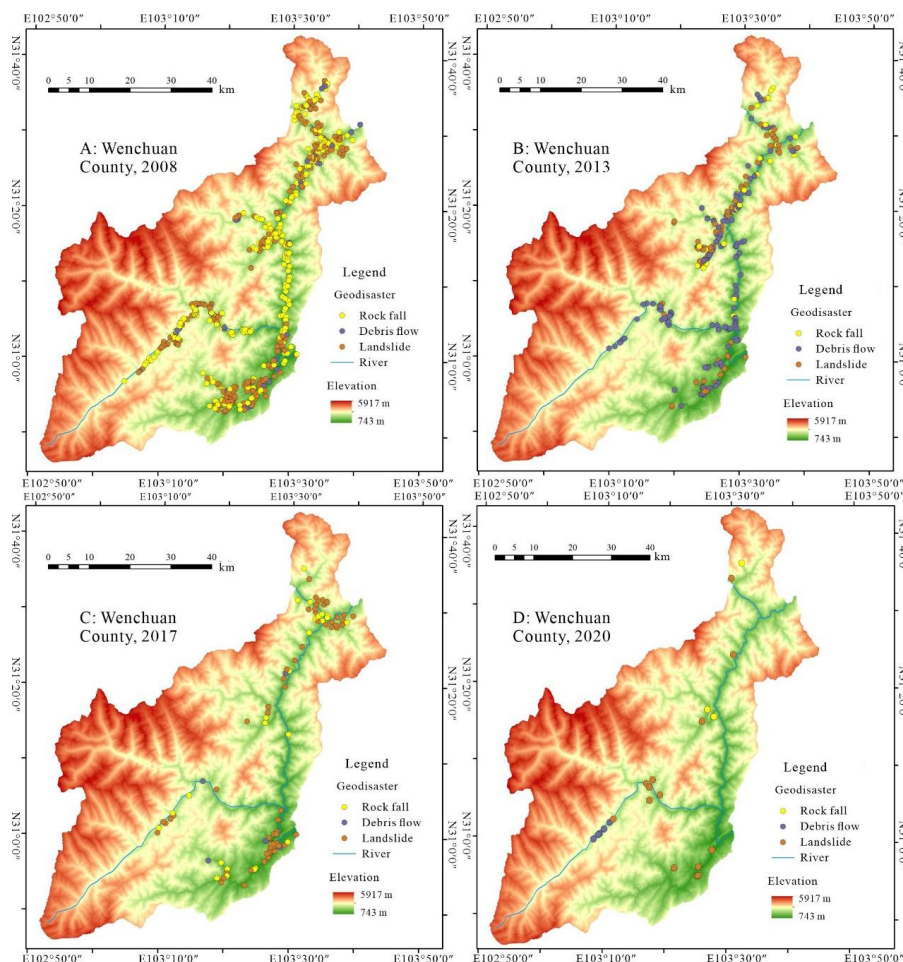
192 4.2. Spatial distribution of post-seismic geohazards in representative years

193 To further clarify the spatiotemporal evolution of post-seismic geohazards, this section
194 focuses on their spatial characteristics in representative years. Based on the annual evolution
195 identified in Section 4.1, earthquake years, clustered outbreak years, and rainfall-rich years
196 were selected to examine where geohazards were concentrated and whether their spatial
197 patterns changed during different stages of post-seismic evolution.



198 **4.2.1 Wenchuan County**

199 The 2008 Wenchuan earthquake induced a large number of coseismic geohazards, mainly
200 rock falls and landslides, which were recorded primarily along major valleys (Figure 6-A). In
201 2013, geohazards were still mainly recorded along valley slopes and drainage areas, but
202 debris flows became more prominent (Figure 6-B). In 2017 and 2020, the number of events
203 further decreased, while most recorded geohazards remained located in valley-related areas
204 (Figure 6-C). Notably, although annual rainfall was high in 2020, only a small number of
205 geohazards were recorded (Figure 6-D). Overall, Wenchuan showed a relatively stable valley-
206 related distribution pattern in the available records, whereas the activity intensity and
207 dominant geohazard types changed through time.



208

209 Figure 6 Spatial distribution of recorded geohazards in representative years in Wenchuan

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County.

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4.2.2 Beichuan County

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In Beichuan County, the 2008 Wenchuan earthquake triggered a clear coseismic

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geohazard response, with events mainly distributed along valley flanks (Figure 7-A). In 2013,

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the number of geohazards exceeded that of the coseismic events in 2008, and the events were

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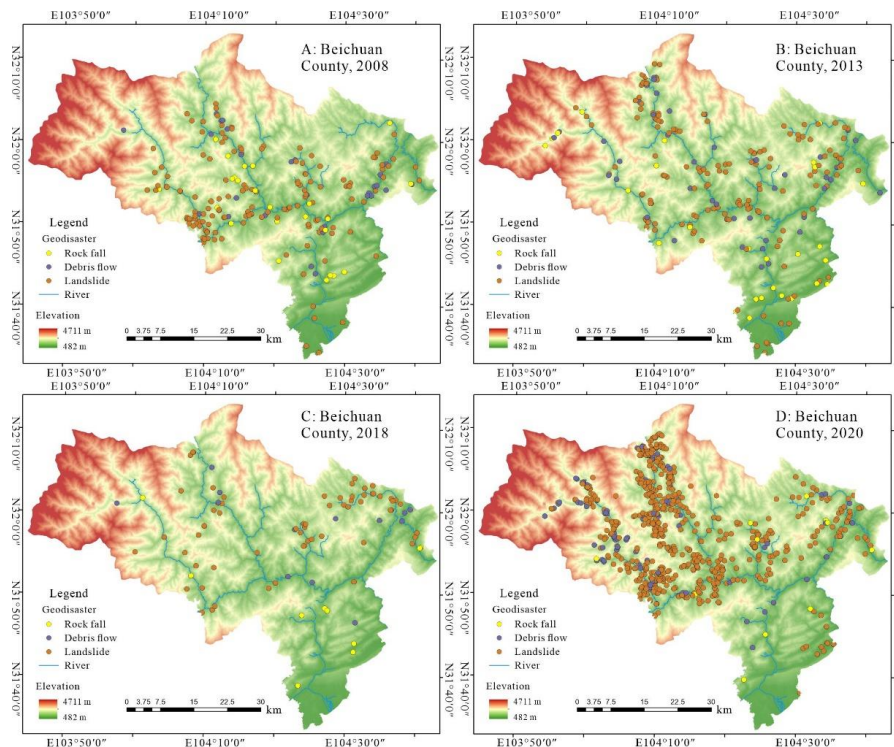
distributed over a wider area without a pronounced regional concentration (Figure 7-B). This

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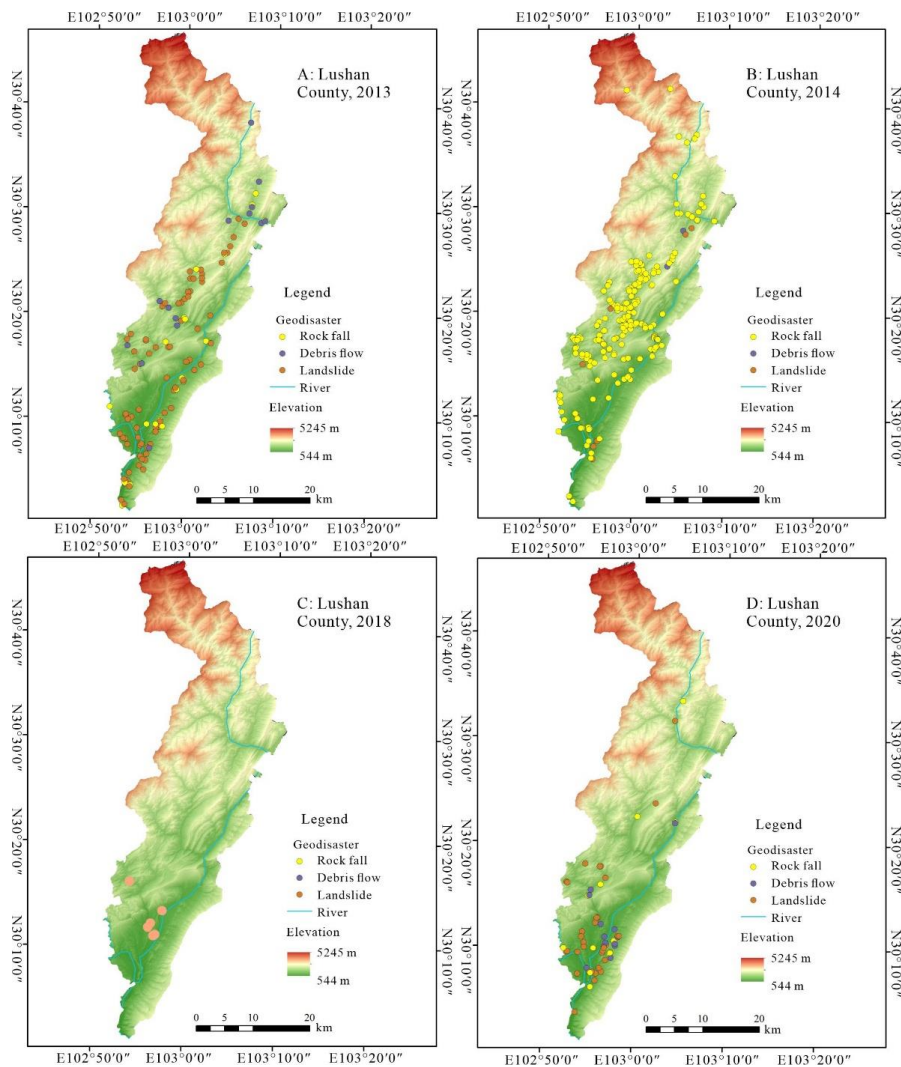
suggests that post-seismic activity in Beichuan was not confined to the initial coseismic



217 concentration zones, but involved more widespread reactivation of earthquake-disturbed
218 slopes. In 2018, the number of geohazards decreased, although the distribution remained
219 relatively scattered, with most events located in the northern part of the county (Figure 7-C).
220 In 2020, Beichuan experienced the strongest clustered geohazard outbreak during the
221 observation period. These events were mainly concentrated in the northern and western
222 regions, showing a spatial pattern substantially different from the 2008 coseismic distribution
223 (Figure 7-D). This indicates that later post-seismic activity became more spatially selective
224 and was concentrated in specific subregions where unstable slope materials remained
225 susceptible to failure. Overall, geohazard activity in Beichuan County was dominated by
226 landslides throughout the observation period; its post-seismic evolution was therefore mainly
227 expressed as a change from relatively widespread distribution to localized concentration.



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242 Figure 8 Spatial distribution of recorded geohazards in representative years in Lushan County.

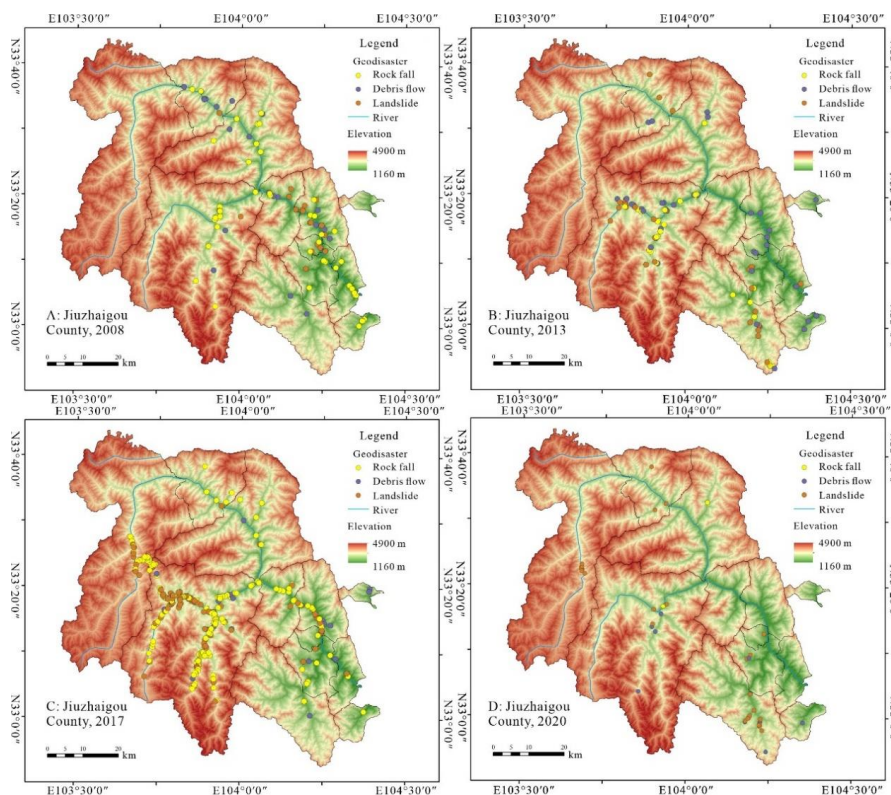
243 4.2.4 Jiuzhaigou County

244 In Jiuzhaigou County, the 2008 Wenchuan earthquake was associated with an increase in
245 geohazard activity, with events mainly occurring within mountain valleys and dominated by
246 rock falls (Figure 9-A). In 2013, the first post-seismic clustered outbreak was recorded, and
247 the events were still mainly concentrated in the central and southern valley areas (Figure 9-B).



248 In 2017, the Jiuzhaigou earthquake triggered a sharp increase in coseismic geohazards, which
249 were dominated by rock falls and mainly distributed in the central and western high-mountain
250 valleys (Figure 9-C). In 2020, the number of geohazards decreased to 41, and the events were
251 mainly concentrated in the southern region (Figure 9-D), indicating that later activity did not
252 completely disappear but became more localized.

253 The four county cases indicate that post-seismic geohazard evolution showed both
254 common spatial constraints and clear regional differences. A common feature is that recorded
255 geohazards were mainly distributed along valleys, valley slopes, drainage areas, and high-
256 mountain valley settings, reflecting the importance of valley-related geomorphic units in
257 earthquake-affected mountainous regions. However, the spatial evolution differed among
258 counties. These differences show that post-seismic geohazard evolution was expressed not
259 only by changes in annual event counts, but also by spatial persistence, localization, and
260 changes in dominant geohazard types.



261

262 Figure 9 Spatial distribution of recorded geohazards in representative years in Jiuzhaigou
263 County.

264 4.3 Monthly type-specific geohazard activity and rainfall

265 Based on the annual series shown in Figure 5, several clustered outbreak years coincided
266 with years of above-average rainfall, whereas this correspondence was not consistent across
267 all counties and years. For example, clustered outbreaks in Wenchuan County in 2013 and
268 2017 and in Beichuan County in 2013 and 2020 occurred during years with rainfall above the
269 local multi-year mean. In contrast, some rainfall-rich years, such as Wenchuan in 2020 and
270 Lushan in 2018, did not show comparable increases in recorded geohazards. Therefore,
271 monthly rainfall and monthly type-specific geohazard records from 2006 to 2023 were further



272 analyzed to examine the intra-annual timing, rainfall correspondence, and geohazard type
273 composition of high-activity years in the four counties. The results are shown in [Figure 10](#).

274 The monthly records captured the direct influence of the three major earthquake months.
275 In May 2008, Wenchuan, Beichuan, and Jiuzhaigou recorded 599, 151, and 72 geohazard
276 events, respectively, indicating the pronounced impact of the Wenchuan earthquake on these
277 counties. Rock falls accounted for the largest proportion in Wenchuan and Jiuzhaigou,
278 whereas landslides were more prevalent in Beichuan. Lushan recorded only one event,
279 suggesting a limited direct response to the Wenchuan earthquake. In April 2013, the Lushan
280 earthquake produced a distinct monthly peak in Lushan County, with landslides representing
281 the main recorded type. In August 2017, the Jiuzhaigou earthquake generated a clear monthly
282 peak in Jiuzhaigou County, where rock falls constituted the dominant type.

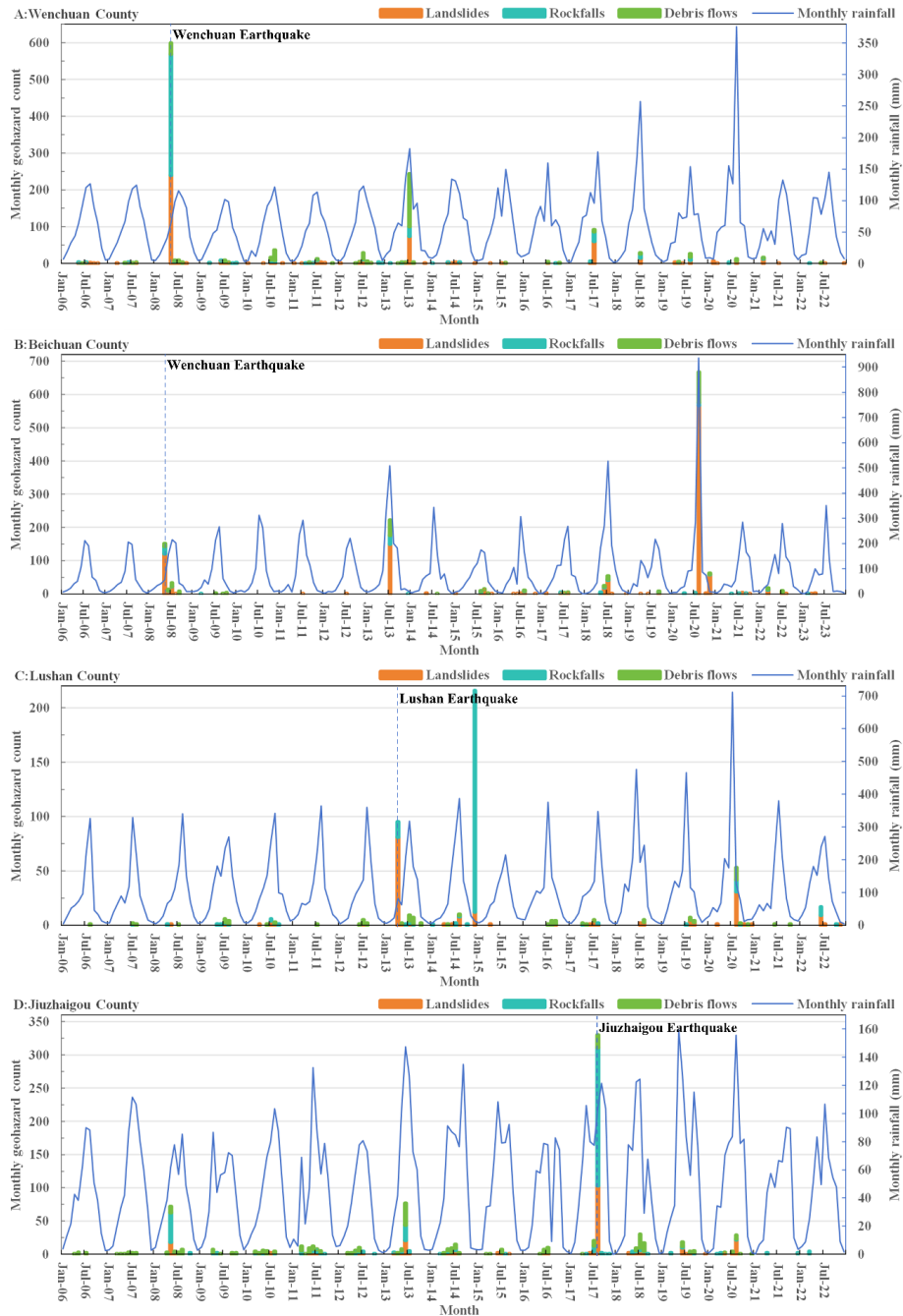
283 After excluding the months of the major earthquakes, most identified clustered outbreak
284 months occurred in June, July, or August, corresponding to the rainy season, and clear
285 differences in geohazard type composition were observed among the four counties. In
286 Wenchuan County, the July 2013 geohazard peak was dominated by debris flows, whereas the
287 dominant type shifted to landslides during the July 2017 peak. In Beichuan County, the
288 clustered outbreaks in July 2013 and August 2020 were both dominated by landslides. Lushan
289 County showed a different pattern: the December 2014 peak was mainly composed of rock
290 falls and occurred outside the main rainy season. This non-rainy-season rock-fall peak may
291 reflect a delayed response of earthquake-damaged rock masses under continued weathering,
292 freeze–thaw action, or gravitational adjustment. The August 2020 peak in Lushan was
293 dominated by landslides. In Jiuzhaigou County, the June 2013 peak showed a slightly higher



294 contribution from debris flows. These differences indicate that earthquake-month geohazard
295 peaks were commonly dominated by rock falls and landslides, whereas rainy-season peaks
296 were more frequently associated with debris flows or landslides.

297 Beyond these high-activity months, the overall non-earthquake-month records also
298 showed a clear rainy-season concentration. Using June–August as the main rainy-season
299 window, the three geohazard types showed different seasonal patterns after excluding
300 earthquake-month records. During the observation period, 78.1% of non-earthquake-month
301 geohazards occurred from June to August. Among these rainy-season events, landslides
302 accounted for the largest proportion, at 59.1%, followed by debris flows at 29.0% and rock
303 falls at 11.9%. This indicates that landslides and debris flows were the main geohazard types
304 during the rainy season, whereas rock falls showed a weaker rainy-season association. This
305 type-specific difference suggests that landslides and debris flows were more closely related to
306 rainy-season activity, while rock falls were less seasonally constrained.

307 In summary, the monthly records show that clustered outbreak years had different intra-
308 annual structures. Earthquake-month peaks were mainly associated with immediate slope
309 failures, with rock falls or landslides as the dominant types. After excluding earthquake-
310 month records, most clustered outbreaks occurred during June–August, and 78.1% of non-
311 earthquake-month geohazards were recorded in this period. These rainy-season events were
312 dominated by landslides and debris flows, indicating a stronger rainy-season association for
313 these two types.



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Figure 10 Monthly counts of different geohazard types and monthly rainfall in the four counties from 2006 to 2023



317 **5. Discussion**

318 Since 2006, the study area encompassing four counties has experienced three major
319 seismic events and recurrent post-seismic geohazards. The available event records provide an
320 opportunity to examine how recorded geohazard activity evolved through clustered outbreak
321 years, low-activity intervals, and regional differences in persistence and recession. The
322 fundamental reason for the long-term frequent occurrence of post-earthquake geohazards lies
323 in that abundant rock masses with fragmented structures, poor physical and mechanical
324 properties, and low stability are distributed widely across the affected areas (Iverson, 1997;
325 Marc et al., 2015). Long-term exposure to freeze-thaw cycles, physico-chemical weathering,
326 rainfall, and long-term action of gravitational stress are the major factors for triggering the
327 post-seismic geohazards (Zhu et al., 2011; Zhang & Zhang, 2017; Jin et al., 2023).

328 Global observational evidence consistently indicates that post-seismic geohazards
329 display protracted activity, often persisting for decades. The notable examples include
330 landslide activity lasting over 15 years in the 2005 Kashmir earthquake (Mw 7.6) region
331 (Bacha et al., 2022), and landslides triggered by the 1906 San Francisco earthquake (Mw 7.9)
332 reactivated for more than 50 years in the California Coast Ranges (Keefer et al., 1987).
333 Similarly, post-seismic geohazard frequencies in the 2008 Wenchuan earthquake area
334 remained 3-5 times higher than pre-seismic levels even 12 years after the earthquake (Fan et
335 al., 2019). In the region affected by the Wenchuan earthquake is supposed to return to pre-
336 earthquake landslide activity levels within 18-25 years (Yunus et al., 2020; Chen et al., 2020).
337 Landslide rates in areas affected by the Chi-Chi earthquake and certain Japanese seismic
338 events have been shown to decay to background levels within 1-6 years (Hovius et al., 2011;



339 [Marc et al., 2015](#)).

340 The present study focuses on long-term county-scale event records to identify
341 intermittent clustered outbreak years, spatial differences, regional persistence–recession
342 differences, and monthly type-specific rainy-season activity. Based on our work, the long-
343 term evolution of post-seismic geohazards in the four-county study area exhibits the following
344 characteristics:

345 (1) Annual and monthly records both show that post-seismic geohazard activity did not
346 remain continuously high after the major earthquakes. Instead, it exhibited an intermittent
347 clustered-outbreak pattern, in which high-activity periods were separated by relatively low-
348 activity intervals. The persistent low-level activity during these intervals may reflect a stage
349 of continuous adjustment of earthquake-disturbed slopes and loose materials. After a clustered
350 outbreak, part of the readily mobilized loose material may be transported out of slopes and
351 channels, while the remaining fractured rock–soil masses continue to undergo weathering,
352 freeze–thaw action, rainfall infiltration, and gravitational deformation. In some areas, such as
353 Beichuan, the number of post-seismic clustered geohazard events even exceeded the
354 corresponding coseismic event count. This further indicates that post-seismic geohazard
355 activity did not necessarily weaken continuously after the main earthquake. Instead, its long-
356 term evolution was characterized by alternating high-activity and low-activity periods,
357 reflecting the nonlinear adjustment of earthquake-disturbed slopes.

358 (2) The spatial results reveal both common geomorphic constraints and different spatial
359 evolution modes of post-seismic geohazards. A common feature is that recorded geohazards
360 were mainly distributed along valleys, valley slopes, drainage areas, and high-mountain



361 valley settings. These valley-related geomorphic units provide steep slopes, free-face
362 conditions, runoff pathways, and loose materials, and they are also the main corridors where
363 settlements, roads, and other exposed elements are distributed. More importantly, the four
364 counties showed different post-seismic spatial evolution modes. Beichuan represents a typical
365 reactivation–expansion mode. After the 2008 coseismic concentration along valley flanks, the
366 2013 activity became more widely distributed and even exceeded the 2008 coseismic event
367 count, suggesting broader reactivation of earthquake-disturbed slopes. Wenchuan showed
368 clear spatial inheritance of the coseismic geohazard distribution, with later events still mainly
369 concentrated in the valley-related areas where coseismic geohazards had occurred, rather than
370 expanding into new regions. Lushan showed a delayed rock-fall-dominated peak in 2014,
371 followed by a clear decrease in event counts and spatial contraction. Jiuzhaigou experienced a
372 new coseismic peak in 2017, followed by reduced but persistent local activity. These
373 differences indicate that post-seismic spatial evolution may appear as reactivation and
374 expansion, inherited valley-related activity, spatial contraction, or localized persistence in
375 different earthquake-affected regions.

376 (3) After excluding earthquake-month records, monthly geohazard activity was clearly
377 concentrated in the rainy season from June to August, with 78.1% of events occurring during
378 these three months, indicating a close association between rainfall and post-seismic geohazard
379 activity. Among these rainy-season events, landslides and debris flows were the dominant types,
380 together accounting for 88.1%. More importantly, the dominant geohazard types changed
381 significantly after the earthquakes. Earthquake-month peaks were commonly dominated by
382 rock falls and landslides, reflecting the immediate response of slopes to strong ground shaking.



383 In contrast, later rainy-season peaks were more frequently characterized by landslides and
384 debris flows. For landslides, rainfall infiltration may increase soil moisture and pore-water
385 pressure, reduce effective stress and shear strength, and promote the failure or reactivation of
386 earthquake-weakened slopes. For debris flows, rainfall-generated runoff may erode and entrain
387 loose materials stored on slopes, in gullies, and in channels, transforming earthquake-disturbed
388 sediments into debris-flow activity (Shou et al., 2011; Marc et al., 2015; Li et al., 2018; Fan et
389 al., 2019; Chen et al., 2022; Shen et al., 2020; Zhu et al., 2011; Zhang & Zhang, 2017; Jin et
390 al., 2023). By contrast, rock falls are more strongly controlled by fractured rock masses, steep
391 slopes, discontinuity-controlled instability, and earthquake-induced rock damage. The
392 December 2014 rock-fall peak in Lushan provides a typical example: it occurred outside the
393 main rainy season and was more likely related to delayed instability of earthquake-damaged
394 rock masses than to a typical rainy-season response. Therefore, rainfall may promote different
395 forms of post-seismic geohazard activity by weakening earthquake-disturbed slopes and
396 remobilizing loose materials during long-term post-seismic evolution.

397 (4) The four counties indicate two contrasting trajectories: Beichuan and Lushan showed
398 clearer weakening trends after repeated clustered outbreaks, whereas Wenchuan and Jiuzhaigou
399 still remained in a low-level persistent activity stage. In Beichuan, several post-seismic
400 clustered geohazard events occurred after the 2008 Wenchuan earthquake, and their event
401 counts even exceeded the 2008 coseismic records. However, the subsequent decrease in annual
402 event counts and the concentration of later activity in the northern and western parts of the
403 county suggest that post-seismic activity gradually became more localized, possibly reflecting
404 progressive slope stabilization or substantial removal of readily mobilized materials. Lushan



405 also experienced post-seismic clustered activity after the 2013 earthquake, especially the
406 delayed rock-fall peak in 2014 and the landslide-dominated activity in 2020. The relatively
407 smaller magnitude of the Lushan earthquake, together with the later decrease in event counts
408 and clear spatial contraction, reflects a more evident weakening trend in this county. By contrast,
409 in Wenchuan, the clustered outbreaks in 2013 and 2017 were much smaller than the 2008
410 coseismic peak, and later events were still mainly concentrated in the valley-related areas
411 initially affected by the 2008 Wenchuan earthquake. This spatial continuity suggests that areas
412 with previous coseismic failures may have a higher probability of recurrent clustered activity,
413 because they combine favorable geomorphic conditions with long-lasting earthquake-disturbed
414 materials. [Fan et al. \(2018\)](#) also reported that loose and fragmented materials remained on
415 slopes in the Wenchuan earthquake-affected region even after a decade of post-earthquake
416 evolution, providing an important material basis for recurrent post-seismic geohazards. In
417 Jiuzhaigou, the time elapsed since the 2017 earthquake is relatively short, and post-seismic
418 activity decreased but still persisted, indicating that this county may still be in a low-activity
419 adjustment stage after the coseismic peak. Therefore, from a monitoring perspective, Wenchuan
420 and Jiuzhaigou should remain key areas for continued observation. Beichuan and Lushan
421 showed clearer recession signals, but localized residual risks should still be considered,
422 especially in areas where previous clustered geohazard activity overlaps with valley-related
423 geomorphic settings.

424 Several limitations should be noted. The inventory used in this study is risk-oriented and
425 lacks complete geohazard records from high-altitude and uninhabited areas. Because consistent
426 information on hazard boundaries, spatial extent, and volumes was unavailable, this study used



427 event counts to characterize long-term geohazard activity. In addition, because rainfall was
428 analyzed at annual and monthly scales, the observed relationship represents rainfall–event
429 correspondence rather than rainfall-triggering thresholds. Despite these limitations, the
430 available county-scale records are still useful for identifying long-term temporal patterns,
431 spatial differences, type-specific rainy-season activity, and regional persistence–recession
432 characteristics of recorded post-seismic geohazards.

433 **6. Conclusions**

434 Wenchuan County, Beichuan County, Lushan County, and Jiuzhaigou County have
435 experienced three strong earthquakes since 2006, including the 2008 Wenchuan earthquake
436 (Mw7.9), the 2013 Lushan earthquake (Mw6.6), and the 2017 Jiuzhaigou earthquake
437 (Mw7.0). These earthquakes triggered a large number of coseismic geohazards and generated
438 abundant loose and fragmented rock masses, providing important material sources for
439 clustered outbreaks of post-seismic geohazards. Taking these four counties as examples, this
440 study investigated the long-term evolution characteristics of post-seismic geohazards and
441 reached the following main conclusions.

442 (1) Post-seismic geohazards mainly showed an intermittent clustered-outbreak pattern.

443 The alternation between high-activity outbreaks and low-activity intervals suggests a non-
444 linear adjustment process of earthquake-disturbed slopes and loose materials.

445 (2) Spatially, recorded geohazards commonly concentrated in valley-related geomorphic
446 units, while different counties showed distinct evolution characteristics, including
447 reactivation, spatial inheritance, localization, and contraction.

448 (3) Monthly records excluding earthquake-month records showed a clear rainy-season



449 concentration, with 78.1% of geohazards occurring from June to August. Landslides and
450 debris flows accounted for 88.1% of these rainy-season events, indicating that the dominant
451 geohazard types shifted toward rainfall-associated landslides and debris flows.

452 (4) In terms of evolution trends, Beichuan and Lushan showed relatively clear weakening
453 tendencies, whereas Wenchuan and Jiuzhaigou still remained in low-level persistent activity
454 stages and should receive continued monitoring attention.

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460 **Conflict of interest**

461 The authors declare that they have no conflicts of interest.

462 **Data availability statement**

463 The datasets used and analyzed during the current study are available from the
464 corresponding author on reasonable request.

465 **References**

- 466 1. Bacha, A., Shafique, M., Werff, H., 2022. Spatio-temporal landslide inventory and
467 susceptibility assessment using Sentinel-2 in the Himalayan mountainous region of
468 Pakistan. *Environ Monit Assess*, 194:845, <https://doi.org/10.1007/s10661-022-10514-w>



- 469 2. Basilone, L., Bonfardeci, A., Romano, P., Sulli, A., 2019. Natural Laboratories for Field
470 Observation About Genesis and Landscape Effects of Palaeo-Earthquakes: a Proposal for
471 the Rocca Busambra and Monte Barracu Geosites (West Sicily). *Geoheritage*, 11:821-837,
472 DOI: 10.1007/s12371-018-0334-8
- 473 3. Chen, M., Tang, C., Li, M., et al., 2022. Changes of surface recovery at coseismic
474 landslides and their driving factors in the Wenchuan earthquake-affected area. *Catena*,
475 210:105871, <https://doi.org/10.1016/j.catena.2021.105871>
- 476 4. Chen, M., Tang, C., Xiong, J., et al., 2020. The long-term evolution of landslide activity
477 near the epicentral area of the 2008 Wenchuan earthquake in China. *Geomorphology*,
478 367:107317, <https://doi.org/10.1016/j.geomorph.2020.107317>
- 479 5. Dai, F.C., Xu, C., Yao, X., Xu, L., Tu, X.B., Gong, Q.M., 2011. Spatial distribution of
480 landslides triggered by the 2008 Ms 8.0 Wenchuan earthquake, China. *Journal of Asian*
481 *Earth Sciences*, 40:883-895, DOI: 10.1016/j.jseaes.2010.04.010
- 482 6. Fan, X.M., Juang, C., Wasowski, J., et al., 2018. What we have learned from the 2008
483 Wenchuan Earthquake and its aftermath: A decade of research and challenges.
484 *Engineering Geology*, 241:25-32, <https://doi.org/10.1016/j.enggeo.2018.05.004>
- 485 7. Fan, X.M., Scaringi, G., Korup, O., et al., 2019. Earthquake-Induced Chains of Geologic
486 Hazards: Patterns, Mechanisms, and Impacts. *Reviews of Geophysics*, 57:421-503,
487 <https://doi.org/10.1029/2018RG000626>
- 488 8. Gorum, T., Fan, X., Westen, C., et al., 2011. Distribution pattern of earthquake-induced
489 landslides triggered by the 12 May 2008 Wenchuan earthquake. *Geomorphology*,
490 133:152-167, DOI: 10.1016/j.geomorph.2010.12.030



- 491 9. He, J., Qi, S., Wang, Y., et al., 2020. Seismic response of the Lengzhuguan slope caused
492 by topographic and geological effects. *Engineering Geology*, 265, DOI:
493 10.1016/j.enggeo.2019.105431
- 494 10. Huang, R., Li, W., 2009. Development and distribution of geohazards triggered by the
495 5.12 Wenchuan Earthquake in China. *Science in China Series E: Technological Sciences*,
496 52:810-819, DOI: 10.1007/s11431-009-0117-1
- 497 11. Huang, R., Pei, X., Fan, X., et al., 2012. The characteristics and failure mechanism of the
498 largest landslide triggered by the Wenchuan earthquake, May 12, 2008, China.
499 *Landslides*, 9:131-142, DOI: 10.1007/s10346-011-0276-6
- 500 12. Hovius, N., Meunier, P., Lin, C., et al., 2011. Prolonged seismically induced erosion and
501 the mass balance of a large earthquake. *Earth and Planetary Science Letters*, 304:347-355,
502 DOI: 10.1016/j.epsl.2011.02.005
- 503 13. Ishitsuka, K., Matsuoka, T., Nishimura, T., et al., 2017. Ground uplift related to
504 permeability enhancement following the 2011 Tohoku earthquake in the Kanto Plain,
505 Japan. *Earth, Planets and Space*, 69:81, DOI: 10.1186/s40623-017-0666-7
- 506 14. Iverson, RM., 1997. The physics of debris flows. *Reviews of Geophysics*, 35:245-296,
507 DOI: 10.1029/97RG00426
- 508 15. Jin, W., Cui, P., Zhang, G., 2023. Evaluating the post-earthquake landslides sediment
509 supply capacity for debris flows. *Catena*, 220:106649,
510 <https://doi.org/10.1016/j.catena.2022.106649>
- 511 16. Keefer, D., Wilson, R., Mark, R., et al., 1987. Real-Time Landslide Warning During
512 Heavy Rainfall. *Science*, 238:921-925, DOI: 10.1126/science.238.4829.921



- 513 17. Li, C., Wang, M., Liu, K., 2018. A decadal evolution of landslides and debris flows after
514 the Wenchuan earthquake. *Geomorphology*, 323:1-12,
515 <https://doi.org/10.1016/j.geomorph.2018.09.010>
- 516 18. Lin, C., Shieh, C., Yuan, B., et al., 2003. Impact of Chi-Chi earthquake on the occurrence
517 of landslides and debris flows: example from the Chenyulan River watershed, Nantou,
518 Taiwan. *Engineering Geology*, 71:49-61, DOI: 10.1016/S0013-7952(03)00125-X
- 519 19. Liu, Y., Liu, R. & Ge, Q., 2010. Evaluating the vegetation destruction and recovery of
520 Wenchuan earthquake using MODIS data. *Nat Hazards*, 54:851–862,
521 <https://doi.org/10.1007/s11069-010-9511-z>
- 522 20. Lv, J., He,, X., Bao, Y. et al., 2025. Spatiotemporal pattern of post-earthquake vegetation
523 recovery in a mountainous catchment in southwestern China. *Nat Hazards*, 121:3023-
524 3046, <https://doi.org/10.1007/s11069-024-06918-1>
- 525 21. Marc, O., Hoviuse, N., Meunier, P., et al., 2015. Transient changes of landslide rates after
526 earthquakes. *Geology*, 43:883-886, DOI: 10.1130/G36961.1
- 527 22. Martha, T., Roy, P., Mazumdar, R., et al., 2017. Spatial characteristics of landslides
528 triggered by the 2015 Mw 7.8 (Gorkha) and Mw 7.3 (Dolakha) earthquakes in Nepal.
529 *Landslides*, 14:697-704, DOI: 10.1007/s10346-016-0763-x
- 530 23. Massey, C., Townsend, D., Rathje, E., et al., 2018. Landslides Triggered by the 14
531 November 2016 Mw 7.8 Kaikōura Earthquake, New Zealand. *Bulletin of the*
532 *Seismological Society of America*, 108:1630-1648, DOI: 10.1785/0120170305



- 533 24. Minato, S., Tsuji T., Ohmi S., and Matsuoka T., 2012. Monitoring seismic velocity change
534 caused by the 2011 Tohoku-oki earthquake using ambient noise records. *Geophys. Res.*
535 *Lett.*, 39:L09309, DOI: 10.1029/2012GL051405.
- 536 25. Moro, M., Chini M., Saroli, M., et al., 2011. Analysis of large, seismically induced,
537 gravitational deformations imaged by high-resolution COSMO-SkyMed synthetic
538 aperture radar. *Geology*, 39:527-530, DOI: 10.1130/G31748.1
- 539 26. Parker, R., Densmore, A., Rosser, N., et al., 2011. Mass wasting triggered by the 2008
540 Wenchuan earthquake is greater than orogenic growth. *Nature Geoscience*, 4:449-452,
541 DOI: 10.1038/NGEO1154
- 542 27. Serey, A., Piñero-Feliciangeli, L., Sepúlveda, S.A., et al., 2019. Landslides induced by the
543 2010 Chile megathrust earthquake: a comprehensive inventory and correlations with
544 geological and seismic factors. *Landslides*, 16:1153-1165, DOI: 10.1007/s10346-019-
545 01150-6
- 546 28. Shen,P., Zhang,L, Fan, Zhu, R., Zhang, S., 2020. Declining geohazard activity with
547 vegetation recovery during first ten years after the 2008 Wenchuan earthquake.
548 *Geomorphology*, 352:106989, <https://doi.org/10.1016/j.geomorph.2019.106989>
- 549 29. Shou, K., Hong, C., Wu, C., et al., 2011. Spatial and temporal analysis of landslides in
550 Central Taiwan after 1999 Chi-Chi earthquake. *Engineering Geology*, 123:122-128, DOI:
551 10.1016/j.enggeo.2011.03.014
- 552 30. Tanyaş, H., Görüm, T., Fadel, I., et al., 2022. An open dataset for landslides triggered by
553 the 2016 Mw 7.8 Kaikōura earthquake, New Zealand. *Landslides*, 19:1405-1420, DOI:
554 10.1007/s10346-022-01869-9



- 555 31. Tian, Y., Owen, L., Xu, C., et al., 2020. Landslide development within 3 years after the
556 2015 Mw 7.8 Gorkha earthquake, Nepal. *Landslides*, 17:1251-1267, DOI:
557 10.1007/s10346-020-01366-x
- 558 32. Unjoh, S., Kaneko, M., Kataoka, S., et al., 2012. Effect of earthquake ground motions on
559 soil liquefaction. *Soils and Foundations*, 52:830-841, DOI: 10.1016/j.sandf.2012.11.006
- 560 33. Xiong, J., Tang, C., Tang, H., et al., 2022. Long-term hillslope erosion and landslide–
561 channel coupling in the area of the catastrophic Wenchuan earthquake. *Engineering
562 Geology*, 305:106727, <https://doi.org/10.1016/j.enggeo.2022.106727>
- 563 34. Xu, C., and Xu, X., 2014. The spatial distribution pattern of landslides triggered by the 20
564 April 2013 Lushan earthquake of China and its implication to identification of the
565 seismogenic fault. *Chinese Science Bulletin*, 59:1416-1424, DOI: 10.1007/s11434-014-
566 0202-0
- 567 35. Yang, W., Qi, W., Zhou, J., 2018. Decreased post-seismic landslides linked to vegetation
568 recovery after the 2008 Wenchuan earthquake. *Ecological Indicators*, 89:438-444,
569 <https://doi.org/10.1016/j.ecolind.2017.12.006>
- 570 36. Yunus, A., Fan, X., Tang, X., et al., 2020. Decadal vegetation succession from MODIS
571 reveals the spatio-temporal evolution of post-seismic landsliding after the 2008 Wenchuan
572 earthquake. *Remote Sensing of Environment*, 236:111476,
573 <https://doi.org/10.1016/j.rse.2019.111476>
- 574 37. Zhang, S., Zhang, L., 2017. Impact of the 2008 Wenchuan earthquake in China on
575 subsequent long-term debris flow activities in the epicentral area. *Geomorphology*,
576 276:86-103, <http://dx.doi.org/10.1016/j.geomorph.2016.10.009>



- 577 38. Zhang, S., Zhang, L., Lacasse, S., Nadim, F., 2015. Evolution of Mass Movements near
578 Epicentre of Wenchuan Earthquake, the First Eight Years. Scientific RepoRts, 6:36154,
579 DOI: 10.1038/srep36154
- 580 39. Zhu, J., Ding, J., Liang, J., 2011. Influences of the Wenchuan Earthquake on Sediment
581 Supply of Debris Flows. J. Mt. Sci., 8: 270– 277, DOI: 10.1007/s11629-011-2114-7
- 582