



Multi-scale dynamics and mesoscale drivers of the catastrophic September 2022 Marche (central Italy) flood

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Abstract. The catastrophic flood that hit the Marche region (central Italy) on 15 September 2022 has been analyzed using high-resolution Weather Research and Forecasting (WRF) model simulations. The convective rainfall event responsible for the flood was favored by the passage of warm, moist air masses that interacted with the Central Apennines. The convective cells, triggered by the rough orography of the area, formed quasi-stationary bands that caused high rainfall accumulation in limited areas. The formation of these rainbands was sustained by the interaction between the low-level flow and the local topography, which generated a persistent low-level convergence line extending downwind from Mount Amiata northeastward towards the Apennines. The observed vertical profile upstream of the orography reveals that the instability was initially suppressed by an inversion layer and was released only when the humid air arrived at low levels and eroded the inhibition. Although the model somewhat underestimates a peak of rainfall (simulating 160–180 mm versus over 400 mm observed), it correctly reproduces the position of the storm and its evolution, confirming the key role of the orography in anchoring the system and making it quasi-stationary. Furthermore, the simulations suggest that gravity waves generated by the mountains helped to sustain the vertical motion of the impinging air, and highlight the role of upper-level dynamics (passage of a jet streak). Sensitivity experiments reveal a strong dependence on initial and boundary conditions. A comparison of the control run (forced with Global Data Assimilation System (GDAS) analysis/forecasts) with a European Centre for Medium-Range Weather Forecasts (ECMWF)-driven simulation shows a strong sensitivity in the low-level dynamics. Although the large-scale upper-level forcing is almost identical, the ECMWF-driven forecast fails to reproduce the correct amount and distribution of rainfall due to a weaker surface pressure gradient, reduced moisture advection, and lower instability at the time of the event. Specifically, the simulation forced with the ECMWF data anticipates the arrival of the moist air mass, determining a timing mismatch between the low-level supply of moist air (and the consequent release of instability) and the upper-level forcing. This lack of synchronization shifts the rainfall away from the observed area.



1 Introduction

The Mediterranean basin is recognized as a primary climate change “hot-spot” and a region highly vulnerable to severe hydrometeorological extremes, including Heavy Precipitation Events (HPEs) and devastating flash floods (Drobinski et al., 2014). Acting as a transition zone between midlatitude and tropical climate regimes, the Mediterranean Sea serves as a reservoir of heat and moisture. When this moist air interacts with the complex and steep orography surrounding the basin, it frequently triggers highly energetic and stationary Mesoscale Convective Systems (MCSs). Recent findings (Khodayar et al., 2021) from the HyMeX program demonstrate that the severity of these events is mostly driven by large-scale atmospheric patterns. Synoptic structures, such as cut-off lows and Mediterranean cyclones, trigger intense air-sea interactions. Through evaporation and surface heat fluxes, the Mediterranean Sea moistens and destabilizes the lower atmosphere, building up potential energy. Low-level winds then transport this unstable marine air toward the coastlines. The complex topography surrounding the basin forces these incoming flows to rise, ultimately triggering intense convective storms.

In this context, the Mediterranean Sea is particularly susceptible to these dynamics due to the proximity of the Alps and Apennines to the warm surrounding sea (Miglietta and Davolio, 2022). A prominent example of this vulnerability is the Marche region, located in Central Italy (see Fig. 1). Geographically, the region is bounded by the Adriatic Sea to the east and the Central Apennines to the west, with their western slopes near the Tyrrhenian coast. Its topography is characterized by a series of narrow, steep river valleys oriented perpendicular to both the coastlines and the mountain barrier. When low-level moist atmospheric flows originating from the Mediterranean Sea impinge on the steep Apennine ridge, they undergo orographic lifting and make the region susceptible to severe weather events. Furthermore, the numerous small and medium-sized steep catchments exhibit a rapid hydrological response to heavy precipitation, frequently resulting in sudden flash floods (Lombardi et al., 2021; Bendia et al., 2026).

A recent example of this kind of dynamics is the flood that struck the Marche region on 15 September 2022. This event was characterized by a quasi-stationary, V-shaped MCS capable of generating extreme precipitation within just a few hours. Previous studies on similar events in Italy highlighted that the stationarity and intensity of such systems are often linked to low-level convergence lines, induced by both the synoptic-scale circulation and the local flow deflection by the terrain (Mastrangelo et al., 2011; Davolio et al., 2009), together with the advection of very moist air over even modest orography (e.g., Miglietta and Regano, 2008).

Compared to previous studies, the present research exploits the higher grid resolution of current numerical weather prediction models to better resolve small-scale terrain features. This improved resolution allows models to explicitly disentangle the meso-gamma scale, which plays a fundamental role in airflow dynamics near complex topography. Consequently, this helps us better reproduce localized, severe orographic precipitation events, which Sawyer (1956) proposed are governed by three core factors: the large-scale synoptic forcing, which dictates the characteristics of the impinging airmass, including wind velocity, stability, and humidity; the terrain interaction, which determines whether the flow goes over or around the orographic obstacle; microphysical processes, which control the efficiency of cloud development and whether moisture falls as precipitation.



The objective of this study is to investigate the atmospheric conditions and multi-scale dynamical mechanisms that drove the September 2022 Marche event using the Weather Research and Forecasting (WRF) model (Skamarock et al., 2019). In this context, the use of high-resolution numerical models, possibly coupled with hydrological models for early warning systems, becomes essential to manage hydrogeological risks in the Apennine district. It is worth noting that the scope of this research extends beyond the numerical reproduction of the rainfall; instead, it aims to better understand the mechanisms triggering this extreme event. Consequently, we focus on the simulation that provided the most accurate representation of the storm's localization and intensity, which has been used as the baseline for the physical analysis. This high-resolution simulation allows for the explicit representation of convective processes and their interaction with the complex orography, providing a tool for diagnosing and understanding extreme meteorological events.

The paper somewhat explores the predictability of localized heavy rainfall events, which is related to a fundamental challenge in numerical weather prediction: the scale-dependency of atmospheric predictability. The ECMWF-EPS (Ensemble Prediction System) initial perturbations are tuned for large-scale synoptic patterns. This presents a critical limitation when predicting localized heavy rainfall: different simulations, which are nearly indistinguishable at larger scales at initial times, may show some differences in mesoscale features which may dramatically affect the simulated evolution, since localized convective systems depend heavily on such small-scale dynamics. Therefore, different ensemble members can produce nearly identical synoptic forecasts while generating greatly different local precipitation totals. Zhang et al. (2002) showed that moist convective processes can cause a rapid upscale initial error growth in regions of moist instability so that the small amplitude error at small scales ultimately limits the predictability even at larger scales. Additionally, small changes in topography or moisture can cause massive differences in precipitation amount, location, and structure.

The paper is structured as follows: Sect. 2 introduces the data and the model used for the analysis; Sect. 3 describes the synoptic environment; Sect. 4 presents the observational analysis and the numerical simulations, sensitivity to initial and boundary conditions is shown in Sect. 5, while conclusions are drawn in Sect. 6.

2 Model Setup and Data

The numerical simulations in this study were performed using the Advanced Research WRF (WRF-ARW) model, version 4.6.1. WRF-ARW is a fully compressible, non-hydrostatic model that employs a hybrid sigma-pressure vertical coordinate system, designed to minimize errors over steep topography.

The WRF model was employed here to reproduce the Marche flood of September 2022, with the objective of understanding the atmospheric drivers and conditions that led to the event. Following a trial-and-error procedure, the configuration explained hereafter, initialized at 00:00 UTC on 15 September 2022 (24 h duration), was identified as optimal, yielding the best compromise in terms of precipitation localization and amount.

The simulation setup includes three one-way nested domains, as shown in Fig. 1.

- The parent domain (d01) covers a wide area with a horizontal grid spacing of 10 km (200×200 grid points).

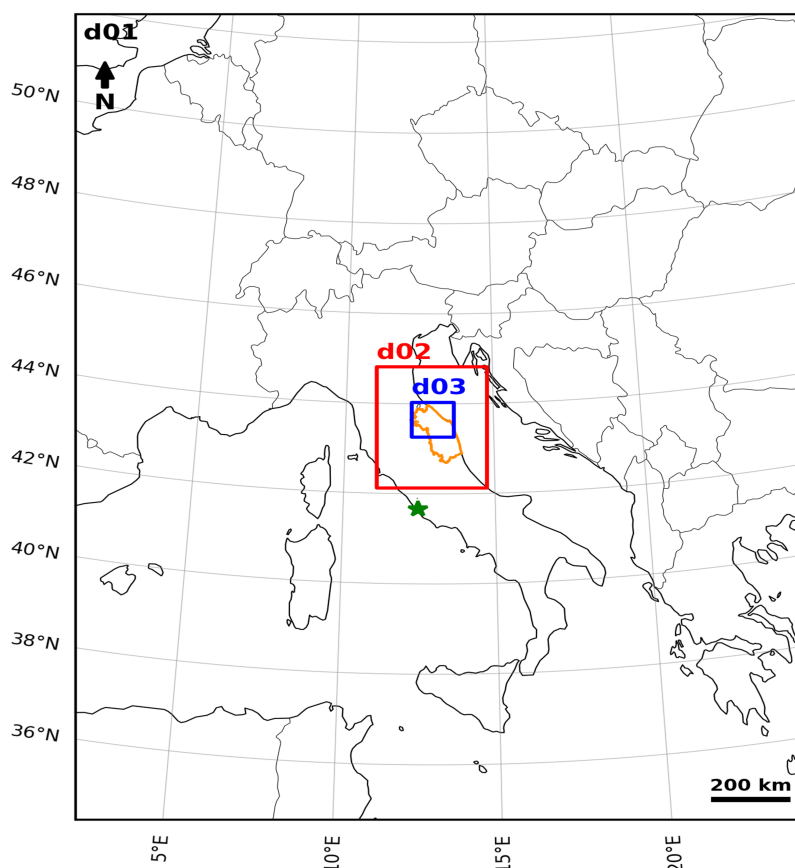


Figure 1. Configuration of the WRF model domains. The orange outline indicates the Marche region, while the green star marker indicates the location of the Pratica di Mare station.

- The intermediate domain (d02) focuses on Central Italy with a grid spacing of 2 km (151×151 grid points).
- The innermost domain (d03) covers in detail the portion of the Marche region affected by the highest precipitation peaks, having a grid spacing of 0.66 km (172×130 grid points).

Vertically, the model atmosphere is discretized into 41 levels, extending from the surface up to a top pressure of 50 hPa.

90 Initial and lateral boundary conditions were derived from the NCEP GDAS/FNL 0.25° Global Tropospheric Analyses and Forecast Grids. Boundary conditions were updated every 3 h using the analysis fields (f00), available every 6 h, as well as the short-term 3 h forecasts (f03) for the intermediate times. This high-frequency update indicates the diagnostic nature of this work, aiming to reconstruct the event dynamics using the setup emerging as the best available in terms of atmospheric boundary forcing.

95 A summary of the parameterization schemes adopted here is provided in Table 1; they have been selected following the configuration used in Tartaglione (2025). A crucial distinction in the setup concerns the cumulus parameterization. For the



Table 1. Summary of WRF physics parameterizations used in the simulation.

Parameter type	Selected Scheme
Microphysics	WRF Single-Moment 5-class scheme (Hong et al., 2004)
Longwave Radiation	RRTMG Longwave scheme (Iacono et al., 2008)
Shortwave Radiation	RRTMG Shortwave scheme (Iacono et al., 2008)
Surface Layer	Revised MM5 scheme (Jiménez and Dudhia, 2012)
Land Surface	Noah–MP LSM (Niu et al., 2011)
Planetary Boundary Layer	YSU scheme (Hong et al., 2006)
Cumulus (d01 only)	Tiedtke scheme (Tiedtke, 1989)

coarser domain (d01), the Tiedtke scheme (Tiedtke, 1989) was activated to parameterize sub-grid scale convection. Conversely, for the intermediate and innermost domains (d02 and d03), a convection-permitting configuration was adopted.

3 Synoptic Environment and Precursors

100 The synoptic environment leading to the extreme event of 15 September 2022 was characterized by a combination of atmospheric patterns that facilitated the accumulation of high potential energy and moisture over the central part of the Italian peninsula in the previous days (Torcasio et al., 2023; Tartaglione, 2025).

The thermal structure (Fig. 2d–f) reveals a pronounced temperature gradient moving eastward over the western Mediterranean. The Marche region was positioned at the boundary between a significant ridge extending from the south, bringing
 105 warm air from North Africa, and a trough on its northwestern side, drawing in colder air from Northern Europe. By 15 September (panels c and f), the synoptic situation over the central Mediterranean evolved from a meridionally elongated pattern into a predominantly zonal flow pattern, with isohypses oriented perpendicular to the Apennine chain (Tartaglione, 2025). The low-level warming and moisture accumulated during the previous days, due to the persistence of the anticyclonic configuration, combined with the advection of moist and cooler mid-level air, increased the potential instability in the area. This evolution is
 110 typical of severe convective storms in the Mediterranean region (e.g., Avolio et al., 2020; González-Alemán et al., 2023) and ultimately facilitated the triggering of a stationary V-shaped system.

The upper-level dynamics also played a key role in sustaining the convective system. The rapid evolution of the geopotential at 300 hPa on 15 September (see Fig. S1 in the Supplement) reveals that the jet streak reached Northern Italy in the afternoon, placing the Marche region below the right entrance region (right-rear quadrant). This configuration is well known to
 115 induce upper-level divergence and low-level convergence (Uccellini and Johnson, 1979), generating vertical lift that provides a favorable dynamic environment for deep convection development.

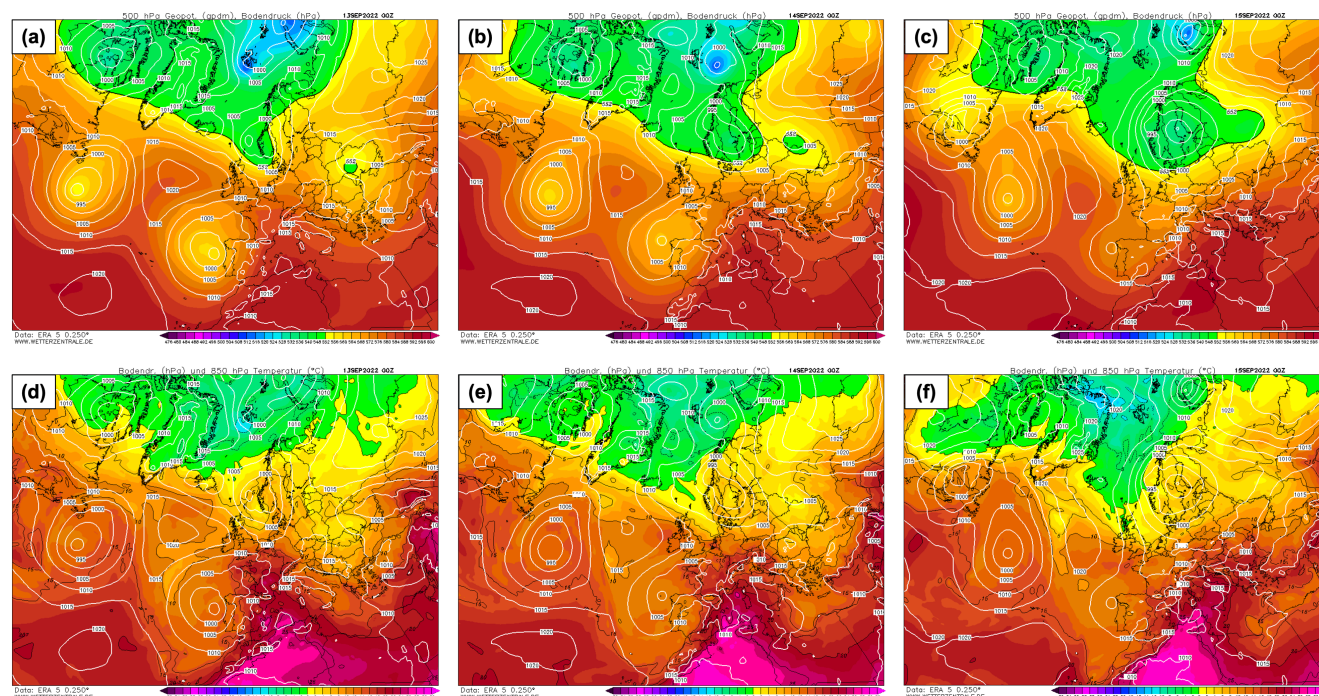


Figure 2. Synoptic maps at 00:00 UTC for 13 September (left column: a, d), 14 September (center column: b, e), and 15 September 2022 (right column: c, f). Top row (a–c): 500 hPa geopotential height (gpdam, shading) and sea level pressure (hPa, white lines). Bottom row (d–f): 850 hPa temperature ($^{\circ}\text{C}$, shading) and mean sea level pressure (hPa, white lines). The maps are based on ERA5 reanalysis data (0.25° resolution). Source: Wetterzentrale (<https://www.wetterzentrale.de/>).

4 Results and Discussion

4.1 Observational Analysis of the Event

The event of 15 September 2022 was characterized by the development of several convective systems that were triggered near the western border of the Marche region (Tartaglione, 2025). These systems organized into “multicell line storms” with a distinct west-to-east orientation. This specific alignment was crucial, as it allowed for continuous and intense rainfall to persist over the same geographical areas for several hours, significantly amplifying the flooding severity.

Consequently, the system remained anchored over the Marche Apennine orography for several hours. Remote sensing platforms captured the intensity and the structure of the event: the H-SAF satellite estimation (Fig. 3a) highlights the V-shaped structure of the storm associated with very intense instantaneous precipitation rates, as well as its highly localized character, while the Civil Protection radar network (Fig. 3b, c) shows reflectivity values reaching the “very strong” intensity class over a narrow band where they persisted for several hours.

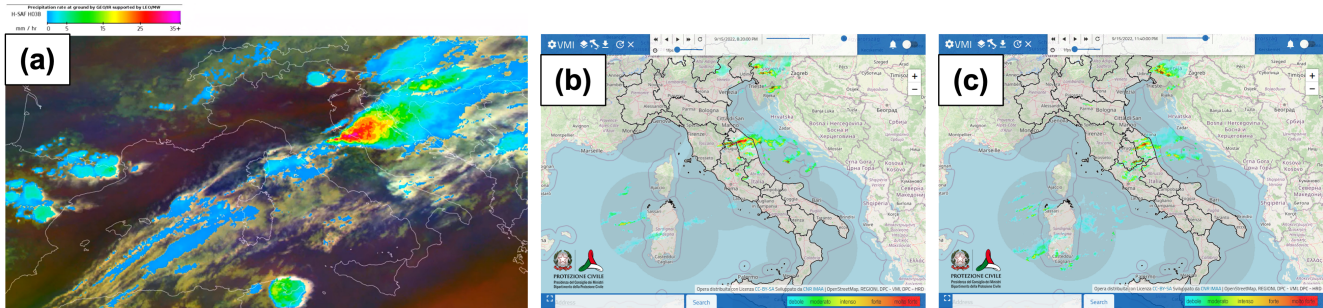


Figure 3. (a) Satellite instantaneous precipitation rate estimation derived from the H-SAF (Satellite Application Facility on Support to Operational Hydrology and Water Management) H03B product (© EUMETSAT) on 15 September at 17:30 UTC; (b) and (c) Vertical Maximum Intensity (VMI) radar reflectivity fields provided by the Italian Department of Civil Protection on 15 September at 18:20 UTC and 21:40 UTC, respectively.

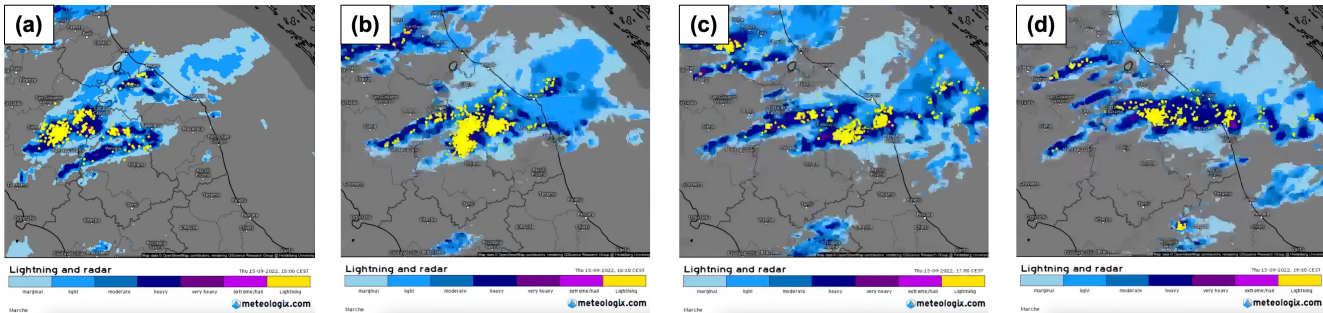


Figure 4. Temporal evolution of lightning activity (yellow dots; data from the EUCLID/Siemens BLIDS network) and radar reflectivity (provided by the Italian Department of Civil Protection) over the Marche region on 15 September 2022 at: (a) 13:00 UTC, (b) 14:00 UTC, (c) 15:00 UTC, and (d) 16:00 UTC. Source: <https://meteologix.com/it>.

The convective character of the event is further evidenced by the lightning activity. High-resolution lightning detection data indicate a high frequency of strokes concentrated along the rainbands. The temporal evolution of the lightning strikes is consistent with the intensification of the radar reflectivity cores (Fig. 4).

The vertical profile, as derived from the radiosoundings launched in Pratica di Mare (WMO ID: 16245; see Fig. 1 for its location), near the Tyrrhenian coast, provides critical insights into the upstream thermodynamic environment preceding the intense rainfall event (Fig. 5). The sequence of profiles from 14 to 15 September 2022 illustrates a transition toward an environment more favorable to the development of intense convection.

In the days before the event, the region was characterized by high-pressure conditions. Large-scale subsidence associated with the anticyclonic regime induced adiabatic compression of the air mass, resulting in a prominent temperature inversion and the development of dry layers in the lower troposphere. This process established a “Loaded Gun” sounding profile, i.e. a



Table 2. Summary of thermodynamic indices from Pratica di Mare radiosoundings.

Date/Time (UTC)	CAPE (MU) [J kg^{-1}]	CIN (MU) [J kg^{-1}]	PW [mm]	Mixing Ratio [g kg^{-1}]
14/09/2022 12:00	1730	−103	36	14.7
15/09/2022 12:00	3092	−73	41	16.3

configuration originally identified by a “Type 1” sounding by Fawbush and Miller (1954), featuring high convective potential inhibited by a low-level inversion and dry air.

140 On 14 September at 12:00 UTC (Fig. 5a), the atmosphere exhibited significant potential energy, with a Most Unstable Convective Available Potential Energy (MUCAPE) of 1730 J kg^{-1} . However, this instability was effectively inhibited by a robust capping inversion, reflected by a convective inhibition (MUCIN) value of -103 J kg^{-1} . Despite a remarkable precipitable water (PW) content of 36 mm, the “cap” prevented any convective development.

A turning point occurred between 00:00 and 12:00 UTC on 15 September (Table 2). The synoptic-scale flow induced a strong low-level moisture advection from the south-southwest, which acted as the primary trigger for the event. This moisture inflow is visible in Fig. 5b, as the surface mixing ratio rose from 14.7 g kg^{-1} to 16.3 g kg^{-1} , and the total column PW increased from 36 mm to 41 mm. In the lowest layers (up to $\approx 940 \text{ hPa}$), the dew point profile closely follows the mixing ratio isopleths on both days. This behavior indicates a nearly constant mixing ratio with height, suggesting vigorous turbulent mixing within the planetary boundary layer (PBL) that efficiently distributes moisture uniformly throughout the mixed layer. The intrusion of warm, moist air in the lower levels combined with the mid-level cooling increased the instability (MUCAPE = 3092 J kg^{-1}), while it reduced the capping inversion (MUCIN = -73 J kg^{-1}).

These modifications, enhanced in the following hours by an additional moistening of the mid-levels and arrival of cooler air (as shown in the following sections), eroded the capping inversion, while the lifting provided by the Apennine ridge and the low-level convergence — the specific roles of which will be further evaluated and discussed in the simulation results section — pushed the parcels up to the level of free convection (LFC). The evolution of these thermodynamic parameters is summarized in Table 2.

4.2 Simulated Thermodynamic Structure

First, to characterize the evolution of the thermodynamic properties of the air-mass feeding the convective system and validate the simulation, we tested the model’s ability to reproduce the vertical structure of the atmosphere. Figure 6 presents the Skew-T diagram derived from the WRF model run at the location of Pratica di Mare at 12:00 UTC, on 15 September. A comparison with the observed sounding discussed in Sect. 4.1 reveals a qualitative agreement in the thermodynamic structure. The model correctly reproduces the “Loaded Gun” signature characterized by a warm, moist boundary layer (dew point temperature $\approx 20^\circ \text{C}$), capped by a stable inversion layer around 950–850 hPa, and surmounted by a drier layer aloft. The quantitative values of the instability indices confirm the potential for severe weather: the simulated MUCAPE is 1731 J kg^{-1} , the MUCIN is



4.3 Physical Mechanisms and Simulation Analysis

The WRF model simulation allows for a detailed analysis of the physical mechanisms initiating and maintaining the stationary convection.

4.3.1 Precipitation Analysis

Figure 7 compares the observed 24 h accumulated precipitation (from 00:00 UTC on 15 September to 00:00 UTC on 16 September), retrieved from the Italian Department of Civil Protection (Dewetra platform), with the WRF simulated accumulation.

The simulation qualitatively captures the main convective signature: two distinct precipitation bands extending from the hinterland of Central Italy to the Adriatic coast. These elongated structures are consistent with the lightning activity and radar reflectivity patterns shown in Fig. 4, as well as with the observational precipitation data derived from the Department of Civil Protection. Indeed, the model correctly identifies the presence of stationary convergence lines (as will be further detailed in Fig. 8) capable of generating persistent precipitation over the same area for several hours.

The simulation exhibits a similar spatial agreement with the observations, mostly reproducing the localization of the precipitation bands. However, the quantitative accumulation is underestimated: while observed peaks were extreme (419 mm and 384.4 mm in 24 h at Cantiano and Frontone, respectively), the simulated maxima remained near 180 mm.

An analysis of the temporal evolution provides further insight into this behavior. Figure 7c compares the observed cumulative rainfall with the model output at six specific rain gauge stations: Arcevia, Cantiano, Colle, Frontone–Monte Acuto, Jesi, and Montefano. The validated station observations used for this comparison were retrieved from the official platform of the Civil Protection Department of the Marche Region (<http://app.protezionecivile.marche.it/>). To account for the difference between the station locations and the model grid point, the simulated precipitation is evaluated using two distinct approaches: a direct point interpolation at the station coordinates, and a neighborhood method that extracts the maximum accumulation at each time step within a 10×10 km box centered on the station. This spatial tolerance mitigates location approximation errors by capturing the most intense rainfall in the vicinity of the sensors.

Looking at the temporal dynamics, observational data indicate that the main convective phase lasted approximately 6 h, from 14:00 to 20:00 UTC. The model captures this temporal window reasonably well, simulating heavy rainfall between 15:00 and 19:00 UTC, although the onset of the event is delayed by about 1–2 h across the domain. In terms of intensity, the 10×10 km simulation performs well during the initial peak between 16:00 and 17:00 UTC, reaching maximum hourly rates of 50 mm h^{-1} , which is comparable to the observed 62 mm h^{-1} at Cantiano. However, after 17:00 UTC, the simulated precipitation rates decay much more rapidly than the observed rainfall, and the absence of a weaker second precipitation event in the evening is responsible for the underestimation of the 24 h total accumulation.

Sensitivity tests using different forcing datasets (e.g., ERA5, GFS, ECMWF), as detailed in Sect. 5, highlighted marked differences in the spatial distribution and amount of precipitation, emphasizing a strong sensitivity to initial and boundary conditions and the intrinsic low predictability of this case study.

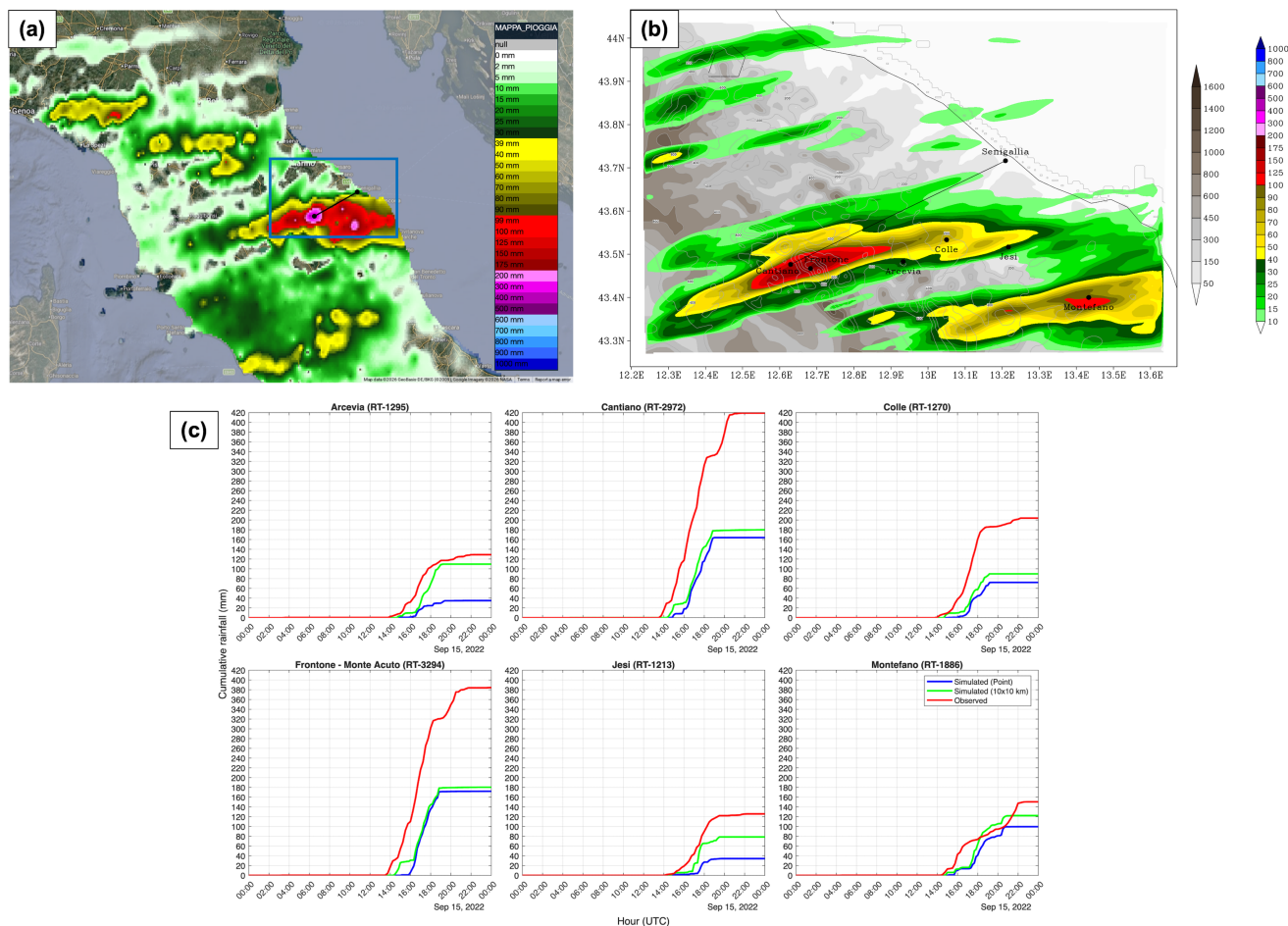


Figure 7. Comparison of 24 h accumulated rainfall (mm) ending at 00:00 UTC on 16 September 2022. (a) Isohyet map derived from the Italian Department of Civil Protection; the blue rectangle indicates the d03 domain. (b) 24 h accumulated precipitation simulated by the WRF model over the d03 domain. (c) Observed (red) and WRF simulated (blue: point; green: 10×10 km neighborhood) cumulative rainfall profiles at six stations: Arcevia (RT-1295), Cantiano (RT-2972), Colle (RT-1270), Frontone–Monte Acuto (RT-3294), Jesi (RT-1213), and Montefano (RT-1886).

4.3.2 Triggering Mechanism and Thermodynamic Instability

The interaction between the low-level flow and the complex topography is shown in Fig. 8a, where the 10 m wind field shows the effect of the Apennines in deflecting the impinging air. This disturbance produces a convergence line at 950 hPa (Fig. 8b) oriented SW–NE. Its leading edge originates downwind of Mount Amiata (approx. 42.9° N, 11.6° E) and extends
205 northeastward, reaching the Apennine ridge. This convergence line is located upstream of the main rainband, thus it reasonably contributed to the uplift of the incoming moist air masses up to the LFC.

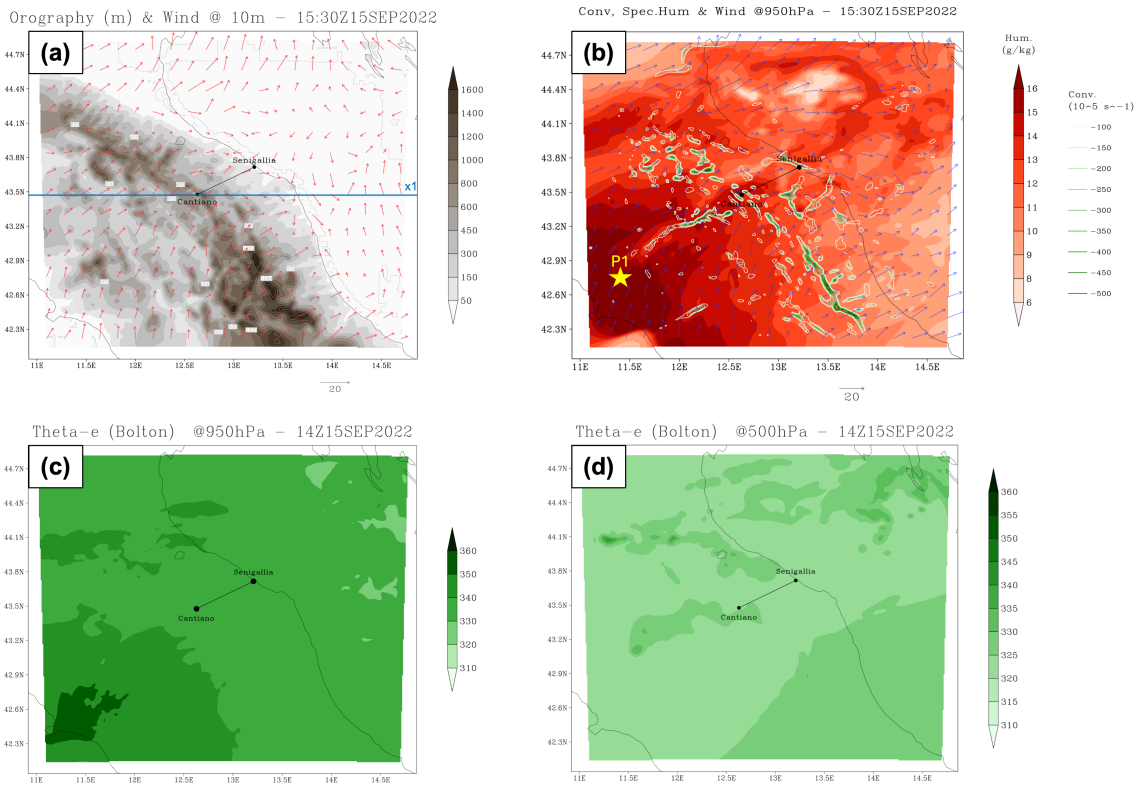


Figure 8. Mesoscale analysis over the d02 domain on 15 September 2022. Top row (15:30 UTC): (a) 10 m wind vectors and orography; the blue straight line identifies the location of the cross-section shown in Fig. 9a, b. (b) 950 hPa convergence (green contours), specific humidity (red shaded), and wind vectors. The yellow marker indicates the location selected for the thermodynamic analysis: P1 (42.7° N, 11.4° E), representative of the upstream southerly inflow near Mt. Amiata. Bottom row (14:00 UTC): Equivalent potential temperature (θ_e), calculated using the approximation by Bolton (1980), at (c) 950 hPa (green shaded) and (d) 500 hPa (green shaded).

The persistence of the pattern for several hours contributed to the continuous regeneration of the system. The latter was also sustained by the rich moisture supply. The specific humidity field at 950 hPa (Fig. 8b) reveals a distinct tongue of moist air (specific humidity $> 15 \text{ g kg}^{-1}$) advected directly into the precipitation area by the southwesterly flow.

From a thermodynamic perspective, the equivalent potential temperature (θ_e) maps in Fig. 8c–d clearly illustrate a condition of potential instability ($\partial\theta_e/\partial z < 0$), with the highest values of 950 hPa θ_e in the southwestern corner of the domain (up to 351 K), just upstream of the rainband, and lower values (of about 315–320 K) at 500 hPa, nearly uniform over the region.

4.3.3 Orographic Forcing and Flow Dynamics

The vertical structure is further detailed in the vertical cross-sections presented in Fig. 9, taken at a fixed latitude of 43.48° N crossing the convective line. The potential temperature (on 15 September at 15:00 UTC, Fig. 9a) reveals undulations of the isentropic surfaces, associated with the stable (dry) stratification of the lower troposphere in response to the orographic forcing,

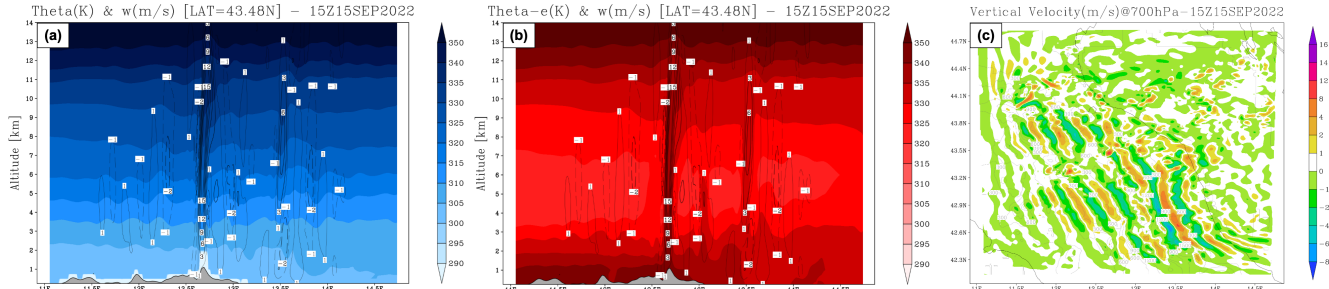


Figure 9. Thermodynamic structure and orographic forcing at 15:00 UTC over the d02 domain on 15 September 2022. (a) and (b) Vertical cross-sections along latitude 43.48° N (transect in Fig. 8a) showing (a) θ (K) with vertical velocity w contours (solid for positive, dashed for negative) and (b) θ_e (K). (c) Vertical velocity (w) field at 700 hPa.

suggesting the presence of gravity waves triggered by the interaction of the flow with the orography (Fig. 9c). These flow perturbations are locally interrupted by large values of vertical velocity (w) contours highlighting the intense updrafts associated with the convective cells (Fig. 9b). Furthermore, θ_e values (Fig. 9b) highlight the presence of convective instability, with a
220 core of high- θ_e air in the lower levels surmounted by lower values above. The orographically induced wave patterns likely contributed to enhancing the vertical lift, allowing low-level air parcels to reach the level of free convection, thus favoring the triggering and persistence of the convective cells just downstream of the crests in the gravity wave patterns.

The simulated vertical velocity field at 700 hPa (Fig. 9c) reveals a well-defined pattern of alternating ascending and descending motions, with a characteristic wavelength estimated at approximately 10–15 km. These features, positioned above
225 and on the leeward side of the main orographic barriers, and perpendicular to the main WSW midlevel flow, are a characteristic signature of intense wind responsible for orographic gravity waves triggered as air parcels are forced upward by the terrain and subsequently oscillate under the influence of (dry) stable conditions.

To better investigate the interaction between the flow and the Apennine orography, the Froude number (Fr) was computed upstream of the orography in the d02 domain. Fr is here defined using the bulk properties of the low-level flow impinging on
230 the mountains as:

$$Fr = \frac{U}{N \cdot h} \quad (1)$$

where h is a characteristic mountain height, roughly set here to 1000 m, which is considered representative of the Apennine barrier, while both the mean horizontal wind speed (U) and the Brunt-Väisälä frequency (N) are computed in the 1000–900 hPa layer.

235 The temporal evolution of the Froude number upstream of Mt. Amiata and the Apennines shows values consistently above the critical threshold of 1 (indicating a “flow-over” regime, as opposed to “flow-around” for $Fr < 1$; Smith, 1989). These values range from 1.5–2 in the morning (09:00 UTC) to 2.5–3 in the afternoon (15:30 UTC), thus increasing due to the progressive intensification of the wind and the advection of moister and warmer (i.e., less stable) air.



Therefore, the upstream flow was definitely able to surmount the orographic barrier rather than flow around it.

240 4.3.4 Sensitivity to Orographic Forcing

To further investigate the specific role of the Apennine orography in sustaining convection, and to justify the complete removal of the terrain as a theoretical bound, two sensitivity experiments were performed by reducing the orography of central Apennines (setting terrain height to 50% and 0%) within the area defined by coordinates 42.0–45.0° N and 10.5–15.0° E. Results (see Fig. 10a and b) indicate that the precipitation maintains a similar orientation as in the control simulation, but the rainbands are shifted slightly northeastward. Interestingly, the atmospheric response is highly non-linear: while in the 0% experiment the rainfall intensity is severely reduced, with peak accumulations decreased by approximately 60%, the 50% reduction scenario maintains peak accumulations close to the control run (reduction of the peak by approximately 25%). An analysis of the low-level fields (see Fig. 10c and d) shows that the reduction or absence of the mountain barrier allows the moist southerly inflow, characterized by high specific humidity, to penetrate further inland and eastward undisturbed. The orography modulates the moisture advection, anticipating the moisture maximum over the city of Cantiano in the two simulations with reduced/absent orography. Convection is still triggered by the convergence between this warm, moist flow and the westerly currents; however, the convergence line is shifted northeastward, anchored along the northern edge of the high θ_e tongue.

This result reaffirms that the mesoscale flow dynamics determine the storm's triggering and positioning, whereas the orography is essential for anchoring the system, generating gravity waves, and providing the additional vertical lift required to amplify precipitation to extreme values.

4.3.5 Moisture Flux Analysis

To assess the moisture availability fueling the precipitation, the Integrated Vapor Transport (IVT) and Precipitable Water (PW) were diagnosed. These parameters are relevant for the identification of atmospheric rivers (e.g., Ralph and Dettinger, 2011), which are often observed in correspondence with heavy rain events in the southern Alpine region (Davolio et al., 2020). The magnitude of the IVT vector was calculated by vertically integrating the specific humidity (q) and horizontal wind components (u, v) from 1000 hPa to 300 hPa:

$$IVT = \frac{1}{g} \sqrt{\left(\int_{1000}^{300} qu \, dp \right)^2 + \left(\int_{1000}^{300} qv \, dp \right)^2} \quad (2)$$

Similarly, Precipitable Water (PW) was computed by integrating the specific humidity from the surface to 300 hPa:

$$PW = \frac{1}{g} \int_{sfc}^{300} q \, dp \quad (3)$$

Figure 11a illustrates the sustained moisture flux directed towards the Marche region at 15:00 UTC, providing the fuel for heavy precipitation. The highest water vapor transport is concentrated in the area of highest precipitation, stressing the

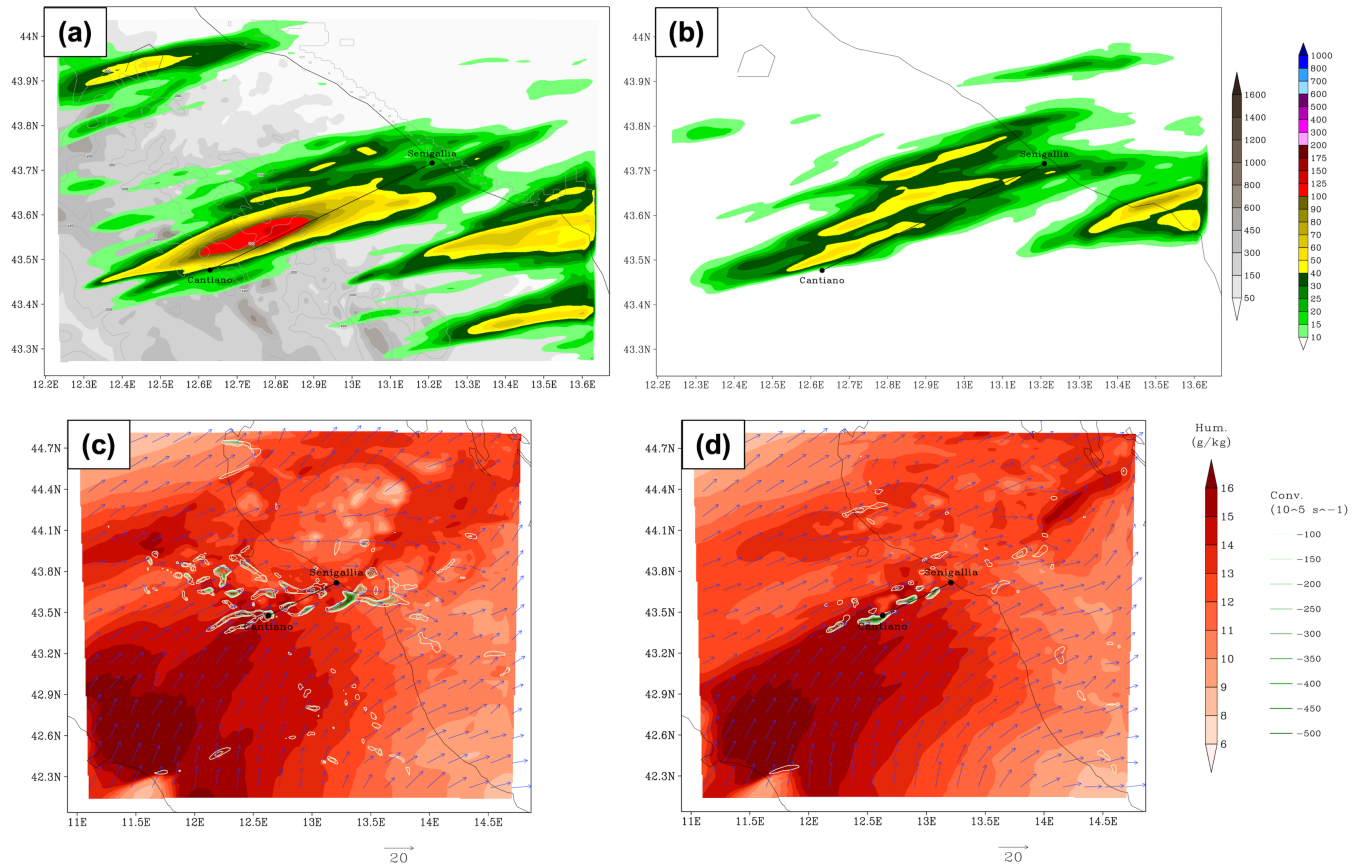


Figure 10. Top row: 24 h accumulated precipitation simulated by the WRF model over the d03 domain (ending at 00:00 UTC on 16 September 2022) for the 50% (a) and 0% (b) orography tests. Bottom row: corresponding low-level convergence (green contours), specific humidity (red shaded) and wind vectors at 950 hPa for the 50% (c) and 0% (d) orography tests.

importance of the high humidity content, characterized by PW values exceeding 40–50 mm along the coast and the Apennine foothills (Fig. 11b), for favoring convection, enhancing CAPE and reducing CIN.

4.3.6 Thermodynamic Environment and Instability

270 To further analyze the evolution of the airmass reaching the Apennines, the thermodynamic vertical profiles were analyzed at the location P1 (42.7° N, 11.4° E, see the star marker in Fig. 8b), representative of the upstream flow near Mt. Amiata. At 05:00 UTC (Fig. 12a), the upstream profile exhibits moderate-to-strong instability but this is suppressed by a significant inhibition, similar to the observed sounding in Pratica di Mare at 12:00 UTC (Fig. 5b). By 16:00 UTC, the simulated profile indicates the marked erosion of the thermal inversion mainly due to the arrival of cooler air in the low-to-mid levels (Fig. 12b).

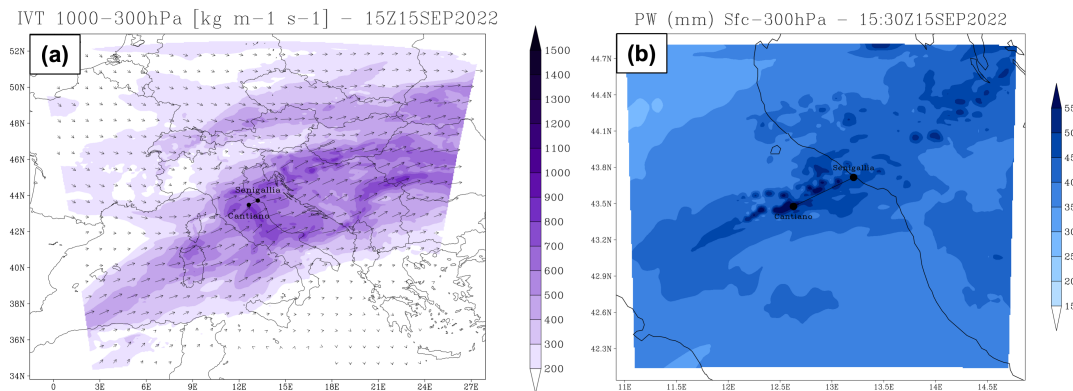


Figure 11. Moisture analysis on 15 September 2022: a) Integrated Vapor Transport (IVT) [$\text{kg m}^{-1} \text{s}^{-1}$] at 15:00 UTC (d01) and (b) Precipitable Water (PW) [mm] at 15:30 UTC (d02).

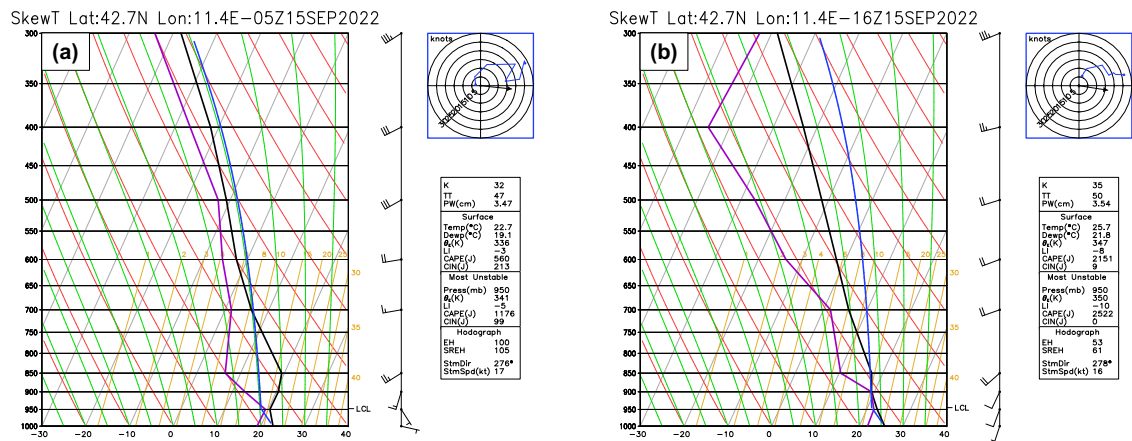


Figure 12. Simulated upstream vertical profile at the location P1 (near Mt. Amiata) at (a) 05:00 UTC and (b) 16:00 UTC on 15 September 2022.

275 The potential for severe convection is confirmed by the values and evolution of the Most Unstable CAPE and CIN fields from 08:00 UTC to 15:30 UTC (Fig. 13). A distinct tongue of high instability, with MUCAPE values exceeding $3000\text{--}3500 \text{ J kg}^{-1}$, progressively extends from the Tyrrhenian Sea towards the Central Apennines, spatially collocated with the high- θ_e tongue observed in the low levels. The temporal evolution indicates that these high instability values progressively built up from the morning, while the convective inhibition — initially present — was gradually eroded in the afternoon, allowing for the release
280 of the accumulated energy.

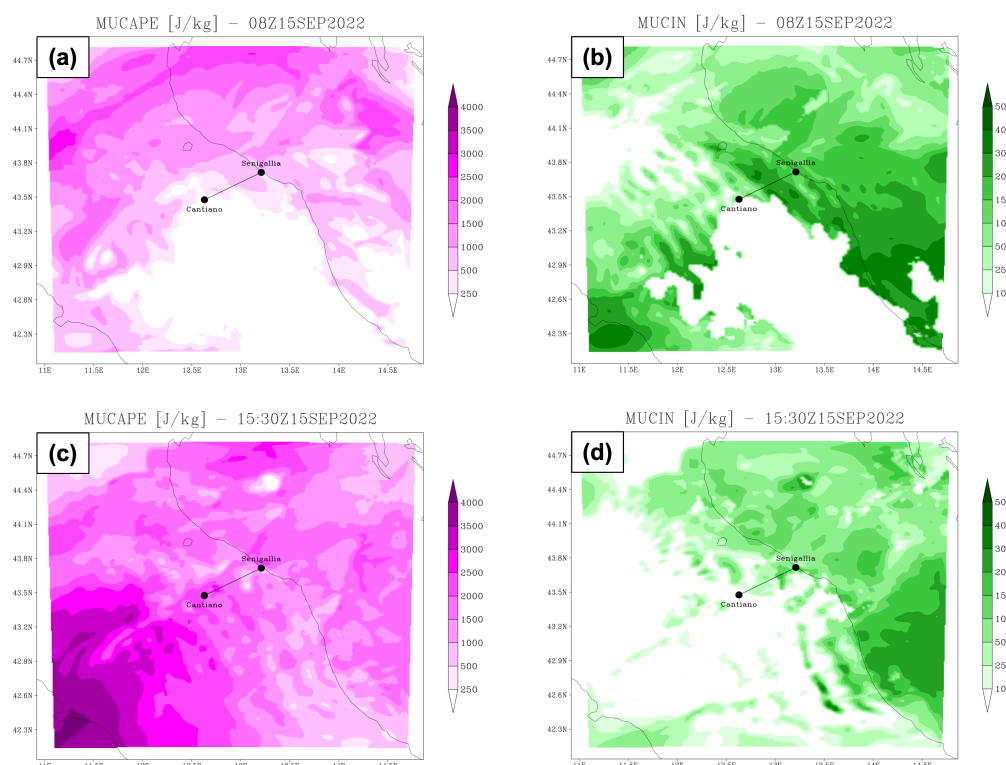


Figure 13. MUCAPE and MUCIN over the d02 domain on 15 September 2022. Top row: (a) Most unstable CAPE and (b) CIN at 08:00 UTC. Bottom row: (c) Most unstable CAPE and (d) CIN at 15:30 UTC.

5 Sensitivity to Initial and Boundary Conditions

To evaluate the model's sensitivity to initial conditions (I.C.), boundary conditions (B.C.), and physical parameterizations, several sensitivity experiments were performed. The configurations are summarized in Table 3, and the resulting 24 h accumulated precipitation is shown in Fig. 14. Figure 14a represents the “reference simulation” (REF) (initialized at 00:00 UTC, 15 September and forced with NCEP/GDAS analysis and forecast fields as explained in Sect. 2), which is the control run extensively discussed in the previous sections. The authors believe that this sensitivity analysis is relevant to understand the spread in the model outputs at a relatively local scale, which can be important in a civil protection warning system framework. Indeed, dramatic differences emerge in terms of precipitation amount and distribution, as shown in Fig. 14.

A comparison between the REF simulation (Fig. 14a) and the run initialized 6 h later (Fig. 14b) reveals a significant decrease in the accumulated rainfall over Cantiano and the loss of the secondary precipitation band located further south. This result illustrates the key role of the initial conditions in correctly modulating the rainfall in the areas affected by the heavy rainfall event. Conversely, the third run, using boundary conditions updated only every 6 h (Fig. 14c) increases the rainfall intensity and extends the area affected by intense rainfall, but erroneously displaces the precipitation peak north of Cantiano. Overall,



these changes appear relatively minor from a warning perspective, as the high impact of the event is clearly identified in the region, although the location is not perfectly represented. Furthermore, the choice of the parameterization schemes plays a crucial role in the spatial distribution and intensity of the precipitation; as shown in Fig. 14d, changing the model physics may significantly change the precipitation: this particular run shifts the rainfall pattern northward and decreases the rainfall amount. These results suggest a significant sensitivity to the initial conditions and the model physics, making a limited area model ensemble approach (including both the sensitivity to initial conditions and model physics) highly recommended to correctly evaluate the possibility of severe events affecting a limited area. This is particularly true for this specific event for which most real-time simulations missed the severity of the rainfall.

Using the NCEP GFS historical archive for both I.C. and B.C. (Fig. 14e) — thus changing both the initial and boundary conditions with respect to REF — results in a primary precipitation band that is less elongated in the southwest–northeast direction, with lower accumulation peaks compared to the REF simulation; conversely, the southern rainband appears more intense and longer than in REF. To isolate the impact of the boundary conditions alone, an additional simulation was performed using the REF initial conditions (GDAS dataset) but driven by GFS forecast boundary conditions (Fig. 14f). Compared to Fig. 14e, this configuration produces a less elongated northern rainband and more concentrated precipitation near Cantiano, while the southern rainband appears weaker. Also considering the result in Fig. 14c, the impact of the boundary condition appears a secondary one, although not negligible.

Finally, Fig. 14g and h demonstrate that driving the model with high-resolution (HRES) ECMWF operational data and ERA5 reanalysis, respectively, fails to accurately reproduce the accumulated precipitation for this case. The ECMWF-driven run completely shifts the precipitation pattern to the north of the Adriatic coast (offshore).

To understand this behavior, it is important to clarify the nature of the forcing datasets. The ECMWF initial and boundary conditions were provided as operational data on 20 standard pressure levels, whereas GDAS analyses included additional data and were available on 34 pressure levels. This coarser vertical resolution of the ECMWF fields might limit the accurate representation of the vertical thermodynamic structure, including the boundary layer and upper-troposphere dynamics. Additionally, we checked for potential differences in the Sea Surface Temperature (SST) fields between the two datasets at the initial time of the simulation. Although GDAS exhibits slightly higher temperatures in the western Mediterranean and Alboran Sea, these features appear too far from our target domain to fully explain the different simulation outcomes. In contrast, the ERA5-driven run fails to generate the elongated banded structure, producing a weaker and incorrectly localized peak.

5.1 Identification of the Physical Drivers

To identify the physical factors responsible for the different simulation behavior, a comparative analysis of selected meteorological fields was conducted between the reference simulation (REF) and one of the runs missing the correct localization and intensity of the rainbands (i.e., the ECMWF-driven simulation). By comparing these two very different cases, the changes in the driving physical mechanisms are more pronounced.

The main discrepancies between the two simulations are localized in the lower troposphere. Specifically, in the afternoon the REF simulation exhibits a stronger meridional wind component over the area of interest compared to the ECMWF-driven



Table 3. Summary of the WRF sensitivity experiments. All simulations share the same domains and grid spacing (10 km, 2 km, 0.66 km). Δt_{BC} indicates the update frequency of the lateral boundary conditions. All simulations share the same physical parameterizations, except for the “Modified physics” experiment (panel d). This specific run employs the Thompson microphysics scheme (Thompson et al., 2008), the Kain-Fritsch cumulus parameterization (Kain, 2004), the Dudhia shortwave and RRTM longwave radiation schemes (Dudhia, 1989; Mlawer et al., 1997), and the Unified Noah Land Surface Model (Tewari et al., 2004).

Panel	Sensitivity	Forcing Dataset	Start Time (UTC)	Δt_{BC} [h]	Remarks
a	REF	NCEP GDAS/FNL (0.25°)	15/09/2022 00:00	3	Control simulation
b	I.C.	NCEP GDAS/FNL (0.25°)	15/09/2022 06:00	3	Delayed initialization
c	B.C.	NCEP GDAS/FNL (0.25°)	15/09/2022 00:00	6	Lower boundary update frequency
d	Physics	NCEP GDAS/FNL (0.25°)	15/09/2022 00:00	3	Modified physics (see caption)
e	I.C. / B.C.	NCEP GFS (0.25°)	15/09/2022 00:00	3	Historical forecast archive
f	B.C.	I.C. GDAS / B.C. GFS	15/09/2022 00:00	3	Mixed forcing datasets
g	I.C. / B.C.	HRES ECMWF (0.1°)	15/09/2022 00:00	3	High-resolution operational analysis
h	I.C. / B.C.	ERA5 (0.25°)	15/09/2022 00:00	1	Hourly reanalysis

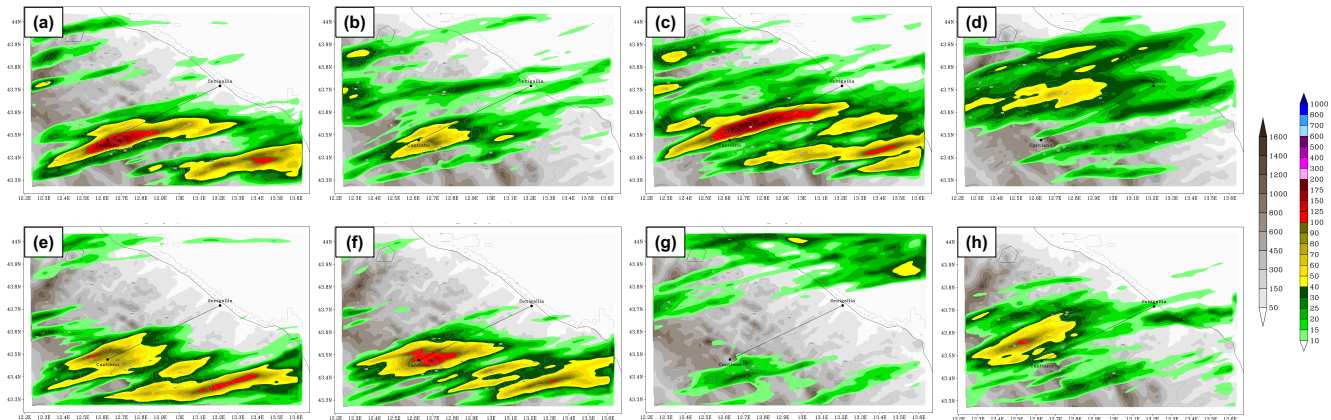


Figure 14. Comparison of 24 h accumulated rainfall for the sensitivity experiments over the d03 domain. The panels (a) through (h) correspond to the configurations detailed in Table 3.

run. The REF simulation also shows an enhancement in the low-level equivalent potential temperature (θ_e), Precipitable Water (PW), and Integrated Vapor Transport (IVT) during the afternoon, due to a greater transport of warm and moist air over the Marche region (see Fig. S2 in the Supplement). Conversely, the ECMWF-driven run anticipates the maxima of these fields during the morning, causing a temporal misalignment between the large-scale forcing and the mesoscale triggers.

This discrepancy in the low-level wind pattern is driven by differences in the Mean Sea Level Pressure (MSLP) fields (Fig. 15a–c). Specifically, the REF simulation (Fig. 15a) exhibits a stronger pressure gradient, established between a distinct

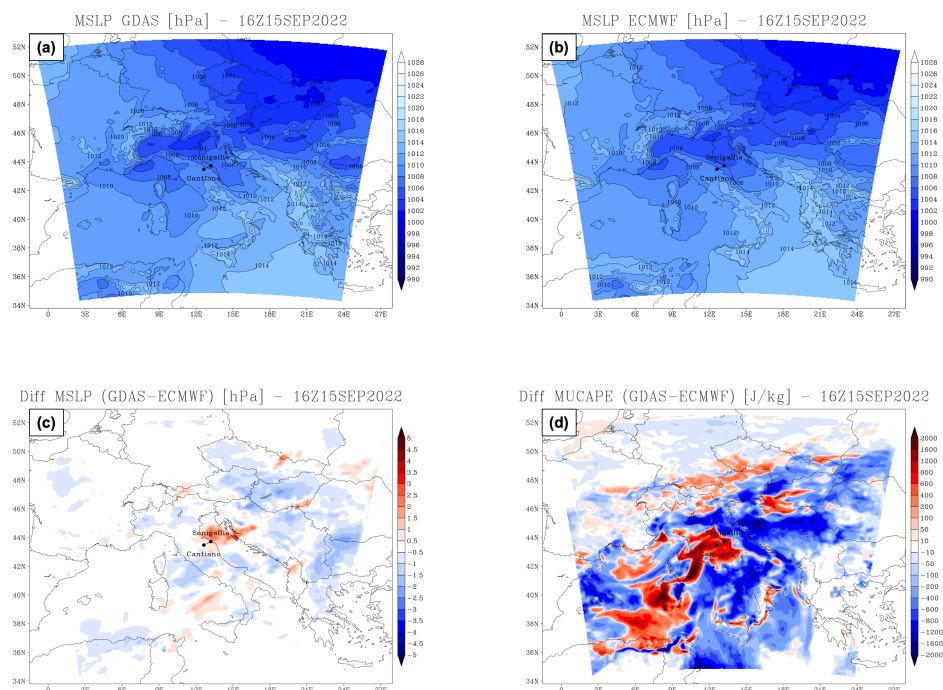


Figure 15. Mean Sea Level Pressure (MSLP) [hPa] at 16:00 UTC on 15 September 2022 for (a) the REF simulation and (b) the ECMWF simulation. Differences (REF minus ECMWF) at the same time are shown for (c) MSLP [hPa] and (d) Most Unstable CAPE (MUCAPE) [J kg^{-1}].

low-pressure minimum near Venice and higher values over the northern Adriatic coast. This specific configuration drives the southerly flow over central Italy, while the same gradient is less apparent in the ECMWF-driven simulation (Fig. 15b).

These dynamical differences affect the thermodynamic environment. Figure 15d shows the spatial difference in MUCAPE between the REF and ECMWF simulations. The REF run produces higher instability over the Central Apennines and the Marche region, while the ECMWF simulation shifts the peaks eastward.

Finally, no significant differences were observed in large-scale features of the mid-troposphere, particularly in the 500 hPa geopotential height and 500 hPa θ_e fields. This proves that the large-scale synoptic forcing does not affect the predictability of this case. Conversely, the failed convective triggering in the ECMWF run depends on the mesoscale dynamics in the lower levels and their interaction with the complex Apennine orography.

6 Conclusions

The analysis of the flood of 15 September 2022 in the Marche region highlights that the catastrophic event emerged from the simultaneous occurrence of different atmospheric forcings across multiple scales. Through the use of high-resolution WRF



model simulations and observational data, this study identifies the key ingredients that caused the initiation and maintenance of the stationary convective system.

At the synoptic scale, the continuous transport of humid air masses from the western Mediterranean acted as the primary fuel, sustaining convection for several hours. While these dynamic factors drove the intensity of the phenomenon, the Apennine topography played the critical role of anchoring the convective system. The analysis reveals a primary convergence line that partially overlapped with the main precipitation band. Other isolated convergence nuclei were present inland and co-located with localized rainfall peaks. This suggests that orographic forcing played a crucial role: the terrain deformed the horizontal wind field, forced the air to converge, and consequently triggered convection through strong vertical uplift. In addition, the interaction between the flow and the terrain triggered orographic gravity waves leeward of the mountain barrier, locally reinforcing vertical motions.

The comparison between the model output and the observations offers critical information on the capabilities and limitations of numerical simulations. Although the WRF model suggests a qualitative spatial agreement in reproducing the dynamical chain of events and the geometry of the precipitation band, including its orientation and position, it showed evident limitations in quantifying the exact rainfall accumulation. Nevertheless, the model's ability to correctly identify the structural precursors — such as the capping inversion, the convergence line, and the moisture transport — confirms the predictive skill of these systems not merely for quantitative rainfall prediction, but for the early identification of high-risk hydrogeological scenarios.

Sensitivity experiments emphasized the role of initial and boundary conditions, as well as physical parameterizations, in accurately predicting the event. While the reference GDAS-driven simulation and the runs with the same initial conditions but different boundary conditions captured the mesoscale features of the event, the model run forced with ECMWF data resulted in a significant northeastward displacement of the precipitation band. A comparative analysis identified the primary drivers of this discrepancy: despite an almost identical large-scale synoptic forcing in the mid-troposphere, the ECMWF-driven run failed to reproduce the correct low-level mesoscale dynamics for this case. Specifically, an underestimated meridional pressure gradient led to a weaker low-level wind field and altered the moisture transport. This induced a temporal mismatch between the thermodynamic instability peak and the dynamic triggers, preventing the system from releasing the accumulated potential energy at the correct location and timing. This demonstrates that the predictability of such extreme Mediterranean storms is constrained by the model's ability to resolve the dynamics of the lower troposphere and their timing in interacting with the local topography. This is particularly relevant in the context of climate change, as the warmer sea surface temperature (SST) in the Mediterranean area is affecting the air-sea interaction, hence the thermodynamics and dynamics of the lower troposphere.

Data availability. The WRF model outputs and the corresponding namelists generated during this study are available from the corresponding author upon reasonable request. The NCEP GDAS/FNL and GFS forcing datasets are publicly accessible through the NCAR Research Data Archive (<https://rda.ucar.edu/>). The ERA5 reanalysis data were obtained from the Copernicus Climate Data Store (<https://cds.climate.copernicus.eu/>). The ECMWF operational high-resolution (HRES) data were retrieved from the ECMWF Meteorological Archival and Retrieval System (MARS).



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380 M.B.; Supervision: M.M.M., E.A.; Writing (original draft preparation): M.B.; Writing (review and editing): M.B., M.M.M., E.A.

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