



1 **Molecular characteristics and formation pathways of**
2 **organosulfur compounds: a comparative field study across**
3 **contrasting atmospheric environments**

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21 **Abstract.** Organosulfur compounds (OrgSs), especially organosulfates (OSs), are
22 ubiquitous aerosol components. However, the spatial, seasonal, and diurnal variations
23 of OrgS formation in polluted atmospheres remain poorly understood. Here, we
24 monitored particulate OrgSs at an urban site and a suburban site in Shanghai and
25 examined their molecular composition and formation pathways under contrasting
26 atmospheric conditions. A total of 1964, 1914, and 2689 OrgS molecular formulas were
27 detected in suburban summer, urban summer, and urban winter, respectively. More than
28 79% of sulfur-containing molecular formulas had $(4s + 3n)/o \leq 1$, indicating that OrgSs
29 were dominated by OSs and nitrooxy-OSs (NOSs). Compared with summer, wintertime
30 OrgSs exhibited lower O/C ratios but higher double-bond equivalence and aromaticity,
31 suggesting a stronger influence of anthropogenic emissions and more unsaturated
32 molecular structures. Although OrgSs were mostly present in aliphatic molecular
33 structures, an increase in the number of aromatic OSs in winter revealed an enhanced
34 contribution from anthropogenic sources. Isoprene/monoterpene-derived OSs peaked
35 during the daytime due to photochemical oxidation in summer, whereas monoterpene-
36 derived NOSs were markedly enhanced at night via nighttime NO_3 -initiated oxidation.
37 Non-metric multidimensional scaling analysis further revealed that OrgS composition
38 in summer was associated with temperature and O_3 during the day but shifted toward
39 RH-driven processing at night. In winter, inorganic nitrogen and sulfur species, aerosol
40 liquid water content, and particle acidity became more important in shaping OrgS
41 composition, suggesting enhanced aqueous-phase and acid-catalyzed formation. These
42 findings provide molecular-level insights into the sources and formation of atmospheric
43 OrgSs across contrasting environments.



44 1. Introduction

45 Organosulfur compounds (OrgSs) are ubiquitous in the atmosphere, spanning
46 diverse environments from remote regions to highly polluted urban areas (Cai et al.,
47 2020; Thomas et al., 2025; Yang et al., 2024; Zhang et al., 2025). They have been
48 recognized as an important component in ambient PM_{2.5} (up to 50%) (Lukacs et al.,
49 2009; Surratt et al., 2008; Tolocka & Turpin, 2012), and have critical influences on
50 aerosol physicochemical properties, such as acidity, hygroscopicity, and volatility (Fan
51 et al., 2022b; Hansen et al., 2015). OrgSs exhibit notable chemical stability and
52 environmental persistence, including organosulfates (OSs), sulfoxides, sulfonates, and
53 sulfones, among which OSs have been identified as the most abundant class (Cai et al.,
54 2020; Jiang et al., 2022; Wang et al., 2018b). Emerging evidence has underscored the
55 hygroscopicity, light absorption, and potential toxicity of OSs (Fan et al., 2022b;
56 Hansen et al., 2015; Jiang et al., 2025), further highlighting the critical need for in-
57 depth investigations into the composition, sources, and formation pathways of
58 atmospheric OrgSs.

59 Although OrgSs can be directly emitted from the combustion of fossil fuels and
60 biomass, they are mainly formed through heterogeneous or multiphase reactions of
61 sulfur dioxide (SO₂) and sulfate particles with anthropogenic and biogenic volatile
62 organic compounds (BVOCs) (Bruggemann et al., 2020; Tang et al., 2020b). The acid-
63 catalyzed ring-opening of epoxides in the presence of sulfuric acid seeds has been
64 widely invoked to explain OS formation (Cai et al., 2020; Jiang et al., 2022).
65 Additionally, OSs may be formed by sulfate radical-induced oxidation of unsaturated
66 compounds (Noziere et al., 2010; Schindelka et al., 2013), nucleophilic substitution of
67 tertiary organonitrates with sulfate (Darer et al., 2011), and sulfate esterification of
68 alcohols or epoxides (Minerath et al., 2008; Minerath & Elrod, 2009). Direct formation
69 of OSs from uptake of gaseous SO₂ by unsaturated fatty acids and organic peroxides
70 (Shang et al., 2016; Ye et al., 2018), as well as from reactions between sulfite/sulfate
71 ion radicals and unsaturated carbonyl compounds in the presence of Fe³⁺, have also
72 been reported (Huang et al., 2018). Furthermore, nighttime NO₃-initiated oxidation of
73 BVOCs is recognized as a crucial formation pathway of nitrooxy-organosulfates (NOSs)
74 (Hamilton et al., 2021; Surratt et al., 2008; Wang et al., 2018b). The presently proposed
75 formation pathways presumably explain the large variety and ubiquity of OSs. However,
76 the chemical composition of OrgSs in the actual atmosphere is highly complex, and the
77 atmospheric relevance of these processes has yet to be fully elucidated.

78 The formation of OSs is strongly influenced by anthropogenic sulfate, NO_x,
79 BVOCs, and ambient conditions, such as, relative humidity (RH) and aerosol physical



80 properties. As these factors vary across different atmospheric environments, field
81 observations are indispensable for validating mechanistic insights from chamber
82 experiments and uncovering previously unrecognized environmental factors. Shanghai,
83 a coastal megacity in the Yangtze River Delta (YRD) region of China, provides a unique
84 environment for investigating the formation and evolution of particulate OrgSs because
85 its atmospheric environment is shaped by strong local emissions, regional transport,
86 and pronounced seasonal meteorological contrasts. This region is characterized by high
87 emissions of VOCs and gaseous pollutants (e.g., NO_x and SO₂), together with generally
88 high RH and strong oxidizing capacity, which can foster complex biogenic-
89 anthropogenic interactions (Cai et al., 2020; Han et al., 2023b; Yang et al., 2023).
90 Furthermore, urban and suburban environments differ in emission intensity and
91 precursor composition. These spatial and seasonal contrasts create favorable conditions
92 for examining how anthropogenic–biogenic interactions regulate OrgS formation.
93 Although several studies have reported the concentrations and potential formation
94 mechanisms of biogenic VOC-derived OSs in the YRD, these OSs represented only a
95 small fraction of total particulate OrgSs in ambient aerosols (Wang et al., 2021b; Yang
96 et al., 2023). Therefore, a better understanding of the molecular characteristics,
97 formation mechanisms, and key influencing factors of OrgSs in Shanghai is crucial for
98 elucidating their role in particulate pollution and developing effective strategies to
99 mitigate secondary organic aerosol (SOA) concentrations.

100 In this study, we used ultra-high performance liquid chromatography coupled with
101 an Orbitrap mass spectrometer (UHPLC–Orbitrap MS) to characterize particulate OrgSs
102 in Shanghai, based on field campaigns conducted at a suburban site in summer and at
103 an urban site in both summer and winter. This strategic sampling captures the spatial,
104 seasonal, and diurnal variability of OrgSs under contrasting atmospheric conditions.
105 Our main objective was to unravel the spatiotemporal evolution, formation pathways,
106 and key drivers of OrgSs at the molecular level, by comparing their compositional and
107 structural characteristics across summer versus winter, day versus night, and low versus
108 high sulfate and NO_x levels. This work provides molecular-level insights into
109 organosulfur chemistry and is pivotal for assessing the role of anthropogenic pollutants
110 in driving SOA formation in biogenically influenced urban regions.

111 **2. Materials and methods**

112 **2.1 Aerosol sampling and site description**

113 Ambient PM_{2.5} sampling was conducted in both the urban center of Shanghai
114 (Fudan University, Yangpu Jiangwan Campus; 31.344°N, 121.518°E) and a suburban



115 area (Qingpu Dianshan Lake; 31.136°N, 121.092°E). The urban site is representative
116 of typical urban environments, characterized by high population density and heavy
117 traffic emissions. The suburban site, located approximately 62 km northeast of the
118 urban site, is influenced by a mix of regional pollutant transport and local biogenic
119 emissions. Sampling campaigns were carried out during summer (July 22 to August 5,
120 2023) and winter (January 9 to 20, 2024) at the urban site, and during summer (July 22
121 to August 2, 2023) at the suburban site. Sampling was conducted separately during
122 daytime (07:00–18:30) and nighttime (19:00–06:30) to capture the diurnal variations
123 in the formation of OrgSs. All PM_{2.5} samples were collected onto pre-baked quartz fiber
124 filters (Whatman Inc.) using high-volume samplers (TH-1000C, Tianhong, China) at a
125 flow rate of 1.05 m³/min. Additionally, 11.5-hour field blank samples were also
126 collected by turning off the sampling pump. A total of 60 PM_{2.5} samples and 4 field
127 blank samples were obtained from the urban site, while 30 PM_{2.5} samples and 2 field
128 blank samples were collected from the suburban site. All filter samples were stored at
129 –20°C before analysis. The real-time monitoring of O₃、NO、NO₂, and meteorological
130 parameters was also conducted during each sampling campaign (Table S1).

131 **2.2 Chemical analysis and prediction of aerosol acidity and ALW**

132 A 24 cm² aliquot was excised from each PM_{2.5} filter sample and subjected to
133 ultrasonic extraction three times, each with 3 mL of methanol for 30 min. The combined
134 extracts were filtered through a 0.45 μm PTFE syringe filter to remove insoluble
135 particles, then concentrated to near dryness under a gentle stream of nitrogen. The
136 residue was reconstituted in 80 μL of methanol containing internal standards (IS, 200
137 ppb D₁₇-Octyl sulfate). The resulting solution was centrifuged at 10,000 rpm for 10 min,
138 and the supernatant was collected for subsequent mass spectrometry analysis.

139 Sample extracts were analyzed using an ultra-high performance liquid
140 chromatography system coupled with an Orbitrap Exploris 120 mass spectrometer
141 (UHPLC-Orbitrap MS, Thermo Scientific, Germany). The electrospray ionization
142 source was operated in negative ionization mode with the following settings: capillary
143 temperature of 320°C, auxiliary gas flow at 10 units, sheath gas flow at 35 units, sweep
144 gas flow at 2 units, and spray voltage at 3.0 kV. The scan range was *m/z* 60–900, with
145 a resolution of 120,000 at *m/z* 200. Chromatographic separation was performed on an
146 Acquity UPLC HSS T3 column (1.8 μm, 100 mm × 2.1 mm, Waters) with a C18 guard
147 column (HSS T3, 1.8 μm). The flow rate was set to 0.3 mL/min, the column temperature
148 was maintained at 40°C, and the injection volume was 5 μL. The mobile phase used a
149 binary gradient, with solvent A as 0.1% formic acid in water and solvent B as 0.1%
150 formic acid in acetonitrile. The gradient program was as follows: 10% B was held for



151 2 min, linearly increased to 30% B over 2–25 min, further increased to 90% B over
152 25–50 min, held at 90% B for 10 min (50–60 min), then decreased to 10% B over
153 60–61 min, followed by a 9 min column re-equilibration at 10% B prior to the next
154 injection. External mass calibration was performed every three days using a commercial
155 calibration solution (Ultramark 1621, Thermo Fisher, Germany) covering an m/z range
156 of 74–1922. Due to the limited volume of the final extract (80 μL), duplicate injections
157 were not performed. Instead, the remaining sample solution from the first injection was
158 retained for potential targeted analyses rather than systematic re-injection. All
159 chromatographic peaks exhibited typical Gaussian shapes, indicating satisfactory
160 column performance and separation efficiency. Field blank samples were extracted and
161 analyzed following the same procedure to account for potential background
162 contamination.

163 Organic carbon (OC) and elemental carbon (EC) were determined using a
164 thermal/optical carbon analyzer (DRI model 2001, Desert Research Institute, USA)
165 with the well-established IMPROVE-A thermal/optical reflectance protocol. Organic
166 matter (OM) was calculated by multiplying the OC by 1.6 (Liu et al., 2018; Turpin &
167 Lim, 2001; Xing et al., 2013). Water-soluble inorganic ions (e.g., SO_4^{2-} , NO_3^- , Cl^- ,
168 NH_4^+ , K^+ , Na^+ , Ca^{2+} , and Mg^{2+}) were quantified by an ion chromatograph (940
169 Professional IC Vario, Metrohm) following procedures described by Cai et al. (2020).

170 The aerosol liquid water content (ALWC) and acidity (pH) were calculated using
171 the ISORROPIA-II thermodynamic model. ISORROPIA-II was run in forward mode,
172 assuming the particles were in a "metastable" state. The input parameters included
173 environmental relative humidity (RH), temperature (T), and the inorganic components
174 of the particle phase (SO_4^{2-} , NO_3^- , Cl^- , NH_4^+ , K^+ , Na^+ , Ca^{2+} , and Mg^{2+}) (Guo et al.,
175 2015; Hennigan et al., 2015; Song et al., 2018b; Weber et al., 2016). In this study, no
176 gaseous NH_3 data were available at the sampling sites. Wang et al. (2020b) showed that
177 the estimated aerosol pH values, when $\text{NH}_3(\text{g})$ was not input into the model, had an
178 average deviation of -0.97 pH units. Earlier studies also indicated that when $\text{NH}_3(\text{g})$
179 was not included in the model, a system bias of -1 pH unit needed to be considered
180 when calculating aerosol pH values (Guo et al., 2015). Therefore, due to the lack of
181 $\text{NH}_3(\text{g})$ data, a correction of 1 pH unit was applied when calculating aerosol pH values
182 for both summer and winter.

183 2.3 Data processing and statistical analysis

184 Raw data were acquired using Xcalibur software (V2.2; Thermo Scientific). The
185 HPLC-HRMS data were processed with MZmine-2.33 software
186 (<http://mzmine.github.io>) to obtain the m/z ratios, formulas, retention times, and peak



187 areas of detected organic compounds (Cai et al., 2024; Hu et al., 2016; Wang et al.,
188 2017). The molecular formula assignment was mainly based on detected m/z values
189 with a mass tolerance of ± 2 ppm. The compounds assigned as $C_{1-40}H_{1-100}O_{1-20}N_{0-4}S_{0-2}$
190 with $s = 1$ or 2 will be collectively referred to as organosulfur compounds, including
191 CHOS ($n = 0$) and CHONS ($n = 1$ or 2).

192 For chemical formula $C_cH_hO_oN_nS_s$, the double bond equivalent (DBE) value,
193 representing the degree of unsaturation, can be calculated by Eq. (1):

$$194 \quad \text{DBE} = (2c + 2 + n - h)/2 \quad (1)$$

195 The aromaticity equivalent (X_c) was introduced by Yassine et al. (2014) for the
196 identification and characterization of monocyclic and polycyclic aromatic compounds.
197 The X_c for compounds containing only carbon, hydrogen, nitrogen, oxygen, and sulfur
198 can be calculated following Eq. (2):

$$199 \quad X_c = \frac{3[\text{DBE} - (i \times o + j \times s)] - 2}{\text{DBE} - (i \times o + j \times s)} \quad (2)$$

200 where i and j correspond to the fractions of oxygen and sulfur atoms involved in the π -
201 bond structures of a compound, respectively. Here, $i = j = 0.5$ was applied for the
202 calculation of X_c values of compounds detected in ESI⁻, as carboxylic compounds with
203 $p = q = 0.5$ are preferably ionized in negative mode (Tong et al., 2016; Wang et al.,
204 2017). If $\text{DBE} \leq i \times o + j \times s$, then X_c was defined as zero. The ranges of $2.50 \leq X_c <$
205 2.71 , $2.71 \leq X_c < 2.80$, $2.80 \leq X_c < 2.83$, and $2.83 \leq X_c$ are proposed as unambiguous
206 minimum criteria for the presence of mono-, di-, tri-, and polycyclic aromatics,
207 respectively (Yassine et al., 2014).

208 Given the large number of detected OrgSs, it is impractical to evaluate the
209 ionization efficiencies of each compound. In this study, qualitatively comparing the
210 variation trends of different compounds detected by HR-MS is a commonly used and
211 valid approach for characterizing the overall molecular composition of OA across
212 different sites or seasons (Cai et al., 2020; Wang et al., 2024b; Zhang et al., 2024).
213 Accordingly, despite the potential influence of matrix effects, analyzing and comparing
214 the number or signal intensity percentages of different OrgS categories can, to some
215 extent, reflect the variability in the molecular composition of OrgSs under varying
216 atmospheric conditions. To further assess the impact of matrix effects in different
217 samples, we spiked all samples with 200 ppb of D₁₇-Octyl sulfate as an internal standard
218 (IS). The IS intensity exhibited minimal variability across different samples, indicating
219 that matrix effects on ionization efficiency can be considered negligible. The relative
220 abundance of each compound refers to its chromatographic peak area, and all detected
221 abundances were blank-corrected. To facilitate more accurate comparisons, the relative
222 abundance of each OrgS compound was defined as the ratio of its absolute abundance



223 to that of IS (set as 100%) within the same sample.

224 To evaluate the associations between environmental variables and OrgSs in
225 suburban summer, urban summer, and urban winter, nonmetric multidimensional
226 scaling (NMDS) analysis was conducted based on Bray–Curtis distances in R using the
227 vegan package (Jiang et al., 2022). For all three sample sets, the OrgS compounds were
228 dimensionally reduced to two components (NMDS1 and NMDS2), and all stress values
229 were less than 0.15. The selected environmental parameters, including meteorological
230 parameters (T and RH), gaseous pollutants (NO₂, CO, SO₂, O₃, and NO), chemical
231 tracers (EC, K⁺, and Cl⁻), inorganic nitrogen and sulfur species (SO₄²⁻, NO₃⁻, and NH₄⁺)
232 as well as ALWC and pH, were fitted onto the ordination plot to evaluate the key
233 environmental drivers influencing the molecular distributions of OrgSs. The
234 significance (*p*-values) of the correlations between the environmental parameters and
235 the NMDS dimensions was assessed using 999 permutations. Only the factors
236 significantly correlated with the NMDS dimensions were reserved and could be
237 considered as the possible drivers associated with molecular distribution. Score and
238 loading plots were constructed according to NMDS variables from each OrgSs
239 compound (gray triangles and circles). The potential driving factors associated with the
240 molecular distribution of OrgSs were indicated by arrows, with their direction and angle
241 representing the relationship between the variables and each dimension. Spearman
242 correlation between the sum-normalized peak areas of individual molecules and some
243 important environmental variables and chemical tracers was performed in R, and then
244 VK diagrams were plotted for each variable based on the Spearman correlation
245 coefficients (Kellerman et al., 2014). Molecules found in at least four samples were
246 adopted for correlation analysis. A false discovery rate-adjusted *p*-value was applied to
247 avoid errors arising from using a large dataset.

248 **3 Results and discussion**

249 **3.1 Molecular Characteristics of OrgSs**

250 By using UHPLC-Orbitrap MS analysis, a total of 1964, 1914, and 2689
251 organosulfur formulas were detected in suburban summer, urban summer, and urban
252 winter aerosol samples, respectively. The intensity and number of detected OrgSs were
253 significantly higher in winter than in summer, which aligns with the seasonal variations
254 in PM_{2.5} and organic matter mass concentrations (Table S1). The elevated number of
255 OrgS species detected in urban winter further indicates a highly complex molecular
256 diversity during the cold season. The identified OrgSs can be classified into six
257 categories: CHOS₁, CHOS₂, CHON₁S₁, CHON₁S₂, CHON₂S₁, and CHON₂S₂. CHOS₁



258 and $\text{CHON}_{1-2}\text{S}_1$ account for ~88% of the molecular number and ~96% of the total
259 intensity of OrgSs (Table S2), suggesting their dominant roles in OrgS composition.
260 Notably, the intensity contribution of CHOS_1 in suburban summer was lower than those
261 in urban summer and winter, whereas $\text{CHON}_{1-2}\text{S}_1$ showed an opposite trend. This
262 disparity underscores the distinct influences of geographical location and seasonality
263 on OrgS composition, as discussed in detail below. Furthermore, as many as 79%–92%
264 of OrgS species contained enough oxygen atoms to enable the assignment of $-\text{OSO}_3\text{H}$
265 and $-\text{ONO}_2$ groups (e.g., $(4s + 3n)/o \leq 1$) in their formulas, suggesting that they were
266 potential OSs or nitrooxy-OSs (NOSs), which is consistent with previous studies (Cai
267 et al., 2020; Wang et al., 2016). The assignments of most OSs were further conducted
268 based on their sulfur-containing fragment ions (e.g., $\text{SO}_3^+ m/z$ 79.96, $\text{HSO}_3^- m/z$ 80.92,
269 and $\text{HSO}_4^- m/z$ 96.96) in collisional activation mass spectrometry (Hettiyadura et al.,
270 2015; Riva et al., 2016a).

271 3.1.1 CHOS species

272 CHOS species predominated in both suburban and urban samples, accounting for
273 65%–77% of the total OrgS intensity (Table S2). The majority (83%–88%) of CHOS
274 formulas were assigned with $4s/o \leq 1$, suggesting that these compounds are potential
275 OSs (Lin et al., 2012b). Clear seasonal differences were observed in their molecular
276 characteristics. In urban samples, CHOS species in summer exhibited higher O/C_w and
277 H/C_w ratios ($\text{O}/\text{C}_w = 0.53 \pm 0.11$, $\text{H}/\text{C}_w = 1.81 \pm 0.04$) than those in winter ($\text{O}/\text{C}_w = 0.45$
278 ± 0.06 , $\text{H}/\text{C}_w = 1.62 \pm 0.03$), whereas the average DBE_w of wintertime CHOS
279 compounds (2.56 ± 0.47) exceeded that in summer (2.18 ± 0.44). These results indicate
280 that wintertime CHOS species were characterized by the lower oxidation state and
281 higher unsaturation degree and aromaticity, likely associated with intensified
282 anthropogenic emissions in winter and enhanced photochemical processing in summer.
283 The enhanced anthropogenic influence in winter can be further supported by the
284 aromaticity equivalent (X_c), a parameter determining the presence of monoaromatics
285 ($X_c \geq 2.50$) and polyaromatics ($X_c \geq 2.71$) (Yassine et al., 2014). As depicted in Figure
286 1b,c,f, the intensity fraction of CHOS compounds with $X_c \geq 2.5$ was greater in winter
287 ($18 \pm 8\%$) than in urban summer ($10 \pm 5\%$). Aromatic CHOS compounds were
288 dominated by phenyl OrgSs with $2.50 \leq X_c < 2.71$, accounting for $70 \pm 14\%$ and $73 \pm$
289 9% of the total aromatic CHOS peak intensity in summer and winter, respectively,
290 which possibly suggests significant influences from primary anthropogenic emissions
291 (Cui et al., 2019; Ma et al., 2014; Song et al., 2018a). Furthermore, the intensity fraction
292 of polyaromatic CHOS compounds ($X_c \geq 2.71$) increased from summer ($3 \pm 2\%$ in
293 urban) to winter ($5 \pm 2\%$), suggesting enhanced combustion emission during the cold



294 season. This interpretation is consistent with the nearly twofold higher concentrations
295 of combustion-related tracers (e.g., EC and CO) in winter than in summer. In contrast,
296 higher temperatures and O₃ levels in summer likely facilitated stronger biogenic VOC
297 emissions and more intense photochemical oxidation. All of those highlight seasonal
298 differences in emission sources and secondary transformation processes of CHOS
299 species in urban Shanghai.

300 Compared with urban summer, CHOS species in suburban summer exhibited
301 slightly higher average O/C_w (0.58 ± 0.14) and H/C_w (1.88 ± 0.06) ratios, and a lower
302 DBE_w value (2.06 ± 0.46), indicating a relatively higher degree of oxygenation and
303 saturation in suburban CHOS species. Correspondingly, the intensity fractions of
304 phenyl CHOS and polyaromatic CHOS in suburban summer were $7\% \pm 5\%$ and $2\% \pm$
305 1% , respectively, both slightly lower than those in urban summer. This spatial disparity
306 is consistent with the weaker anthropogenic influence at the suburban site, where the
307 average concentrations of EC and CO were approximately half of those observed at the
308 urban site (Table S1). Conversely, the higher average O₃ level at the suburban site
309 suggests stronger oxidative processing. These results imply that suburban CHOS
310 composition was more strongly influenced by the oxidation of biogenic VOCs under
311 summer conditions. Regionally, the H/C_w ratios of CHOS in this study were comparable
312 to or exceeded those reported for ambient aerosols from diverse locations worldwide
313 (Table S3) (Jiang et al., 2022; Lin et al., 2012a; Ning et al., 2025; O'Brien et al., 2014;
314 Rincón et al., 2012; Wang et al., 2024a; Willoughby et al., 2014), suggesting that the
315 OrgSs in Shanghai are relatively enriched with saturated structures. However, the O/C_w
316 ratios of CHOS compounds identified in this study were slightly higher than that
317 reported for PM_{2.5} in Anshan (0.45) (Ning et al., 2025) and comparable to values
318 measured in the Pearl River Delta (0.52 ± 0.07) (Jiang et al., 2022), but markedly lower
319 than those reported for aerosols from California ($0.82 - 0.93$) and polluted organic
320 aerosols from Mainz ($0.78 - 1.44$) and Chinese ($1.17 - 1.48$) cities (O'Brien et al., 2014;
321 Wang et al., 2021a; Wang et al., 2018a; Wang et al., 2019). Such regional discrepancies
322 suggest that CHOS in Shanghai is likely driven by emissions from different sources and
323 subsequently undergoes complex atmospheric oxidation processes. This also implies
324 that, owing to pronounced spatiotemporal heterogeneity, CHOS in Shanghai may
325 exhibit a distinct molecular composition relative to other regions, highlighting the need
326 for further investigations into OrgS sources and their molecular-level distributions.

327 The most abundant class of CHOS species identified in all samples had 4 – 7 O
328 atoms, with SiO₅ being the most prevalent (Fig. 1d and S1a). This characteristic was
329 most evident in the urban winter samples. The elevated oxygen content implies the
330 presence of additional oxidized functional groups, such as hydroxyl and carbonyl. As



331 shown in Figures 1a–c, the main distribution of CHOS compounds falls within a DBE
332 range of 1 to 10 and a carbon number range of 2 to 22. Based on established DBE
333 classifications (Jiang et al., 2022; Lin et al., 2012b), the detected CHOS species were
334 tentatively classified into three categories: saturated aliphatic-derived species ($\text{DBE} \leq$
335 1; $51 \pm 10\%$), biogenic-derived species ($\text{DBE}=2/3$; $23 \pm 12\%$), and aromatic-derived
336 species ($\text{DBE} \geq 4$; $26 \pm 15\%$), collectively accounting for the total intensity of assigned
337 CHOS compounds (Fig. 1e). CHOS compounds with $\text{DBE} \leq 3$ predominantly
338 contained 8 to 22 carbon atoms, accounting for $59 \pm 13\%$ of the total CHOS intensity.
339 The C_8 – C_{22} compounds are likely associated with the oxidation products of
340 monoterpenes (C_{10}) and sesquiterpenes (C_{15}), as well as dimeric and trimeric BVOC
341 oxidation products (Daellenbach et al., 2019; Kourtchev et al., 2016). However,
342 anthropogenic sources have also been proposed for C_8 – C_{22} CHOS compounds,
343 including the photooxidation of long-chain alkanes from vehicle emissions (Riva et al.,
344 2016a; Tao et al., 2014) and heterogeneous reaction between SO_2 and unsaturated fatty
345 acids in ambient particles (Shang et al., 2016; Zhu et al., 2019). For example,
346 $\text{C}_{10}\text{H}_{18}\text{O}_6\text{S}$ and $\text{C}_{10}\text{H}_{18}\text{O}_7\text{S}$ could be formed via photooxidation of both biogenic (α -
347 pinene) and anthropogenic (cyclodecane) precursors (Riva et al., 2016a; Surratt et al.,
348 2008). These findings underscore the challenge of unambiguous source apportionment
349 based solely on molecular formula. In addition to the dominant C_8 – C_{22} compounds, low
350 molecular weight CHOS compounds (C_2 – C_7) with $\text{DBE} \leq 3$ were also observed, which
351 are likely isoprene derivatives or fragmentation products generated during atmospheric
352 oxidation processes (Hatch et al., 2011; Riva et al., 2016b). High molecular weight
353 compounds ($\text{C} > 22$) contributed minimally to the total CHOS signal. Among the CHOS
354 compounds, five species (e.g., $\text{C}_9\text{H}_{16}\text{O}_7\text{S}$, $\text{C}_{10}\text{H}_{20}\text{O}_5\text{S}$, $\text{C}_{12}\text{H}_{26}\text{O}_4\text{S}$, $\text{C}_{14}\text{H}_{28}\text{O}_5\text{S}$, and
355 $\text{C}_{16}\text{H}_{34}\text{O}_5\text{S}$) were identified as the dominant components in OA. $\text{C}_9\text{H}_{16}\text{O}_7\text{S}$ has been
356 proposed to be mainly derived from the photooxidation of limonene/limonaketone in
357 the presence of acidified sulfate seed aerosol (Surratt et al., 2008). The remaining four
358 species, characterized by saturated aliphatic structures with $\text{DBE} \leq 1$, suggest that their
359 formation likely involves atmospheric oxidation and subsequent sulfation of
360 unsaturated fatty acids or long-chain alkanes (Riva et al., 2016a; Shang et al., 2016; Tao
361 et al., 2014).

362 3.1.2 CHONS species

363 Although CHONS species contributed less to the total OrgS intensity than CHOS
364 species, their nitrogen-containing molecular structures suggest their potential
365 importance in the atmospheric nitrogen cycle, acting as important reservoirs for organic
366 nitrogen. As shown in Table S2, a total of 895, 787, and 1277 CHONS molecular



367 formulas were assigned in suburban summer, urban summer, and urban winter aerosol
368 samples, respectively, accounting for 35%, 23% and 26% of the total intensity of OrgSs,
369 respectively, which underscores that CHONS species were an important component of
370 particulate OrgSs in suburban atmospheric environments. In winter, the average MW_w ,
371 O/C_w , and DBE_w values of CHONS compounds were 322 ± 9.08 , 0.65 ± 0.17 , and 6.02
372 ± 2.60 , respectively (Table S2), which were higher than those of CHOS compounds
373 (MW_w : 285 ± 15.4 , O/C_w : 0.45 ± 0.06 , DBE_w : 2.56 ± 0.47), probably due to the presence
374 of additional nitrate ($-ONO_2$) functional groups (Wang et al., 2019). This trend was
375 also observed in suburban and urban summer samples. Similar to CHOS species, the
376 DBE_w values of CHONS species in urban winter were 1.23 and 1.41 times higher than
377 those observed in urban and suburban summer, respectively, while their average O/C_w
378 values were lower in urban winter. Moreover, CHONS species in suburban summer
379 were characterized by the highest H/C_w and O/C_w ratios, and the lowest DBE_w and X_{C_w}
380 values. These results indicate that the molecular distributions of CHONS species were
381 likely shaped by seasonal and spatial differences in emission sources and atmospheric
382 oxidative capacity. Additionally, the results from the comparison between the average
383 H/C_w and O/C_w ratios of CHONS compounds and those reported previously were
384 consistent with those observed for CHOS species (Table S4). For example, the O/C_w
385 ratios of CHONS compounds here were lower than those reported in six Chinese
386 megacities (Wang et al., 2024a).

387 CHONS species were primarily dominated by those with more than 6 oxygen
388 atoms (45–75% in total) on individual days (Fig. S2a). In summer, the number of
389 CHONS species containing $O \geq 10$ was substantially higher than in winter, likely
390 attributable to enhanced photochemical oxidation processes during the summer (Cai et
391 al., 2020; Fan et al., 2022b). Furthermore, the suburban site exhibited a notably higher
392 number of CHONS species with $O \geq 10$ compared to the urban site during the summer.
393 This spatial discrepancy may be associated with reduced NO titration effects, higher O_3
394 exposure (Table S1), and more abundant biogenic precursor emissions in the suburban
395 atmosphere (Surratt et al., 2008; Wang et al., 2020b). The identified CHONS
396 compounds spanned $N_1S_1O_1 - N_1S_1O_{12}$ and $N_2S_1O_1 - N_2S_1O_{12}$, with $N_1S_1O_7$ class
397 exhibiting the highest relative abundance. Their DBE and carbon number distributions
398 closely resembled those of CHOS species (Fig. 2a–c). Although CHONS compounds
399 containing two N atoms were also identified, their relatively low intensity suggests a
400 minor contribution compared to those with one N atom. The average DBE_w of
401 $CHON_2S_1$ species was approximately 3 – 4 times higher than that of $CHON_1S_1$
402 compounds at both sites, implying that $CHON_2S_1$ probably contains numerous aromatic
403 NOSs, whereas $CHON_1S_1$ species are predominantly composed of NOSs featuring long



404 aliphatic carbon chains with low degrees of oxidation and unsaturation (Fig. 2a–c and
405 Table S2). Based on the (DBE – N) value, which serves as a more reliable indicator of
406 aromaticity (Lin et al., 2012b), CHONS compounds were dominated by olefinic species
407 (DBE – N = 2/3), followed by saturated aliphatic (DBE – N ≤ 1) and aromatic (DBE –
408 N ≥ 4) compounds. Notably, the most abundant saturated aliphatic and olefinic CHONS
409 species fell within the C8 – C12 range with O ≥ 7 (Fig. 2e). C₁₀H₁₇NO₇S (*m/z* =
410 294.0653, DBE=3; Fig. 2c and S3) was the most abundant CHONS formulas across
411 most samples. This species has been proven to be formed via atmospheric oxidation of
412 α-pinene (Surratt et al., 2008). Nevertheless, recent evidence has confirmed its presence
413 in coal combustion-generated aerosols (Song et al., 2018a), suggesting this compound
414 likely has mixed sources in polluted environments. The tentatively assigned NOSs
415 contributed comparably to the total number of CHONS species across both sites,
416 accounting for 56% – 60%, which was notably lower than that of OSs in CHOS species
417 (83% – 88%) (Table S2). The intensity fraction of NOSs in CHONS species was higher
418 at the urban site (89 ± 10%) than at the suburban site (73 ± 10%) during summer. This
419 disparity can plausibly be attributed to elevated NO_x levels in the urban environment
420 (Table S1), which promote NOS formation via reactions with CHOS species or other
421 organic precursors in the presence of acidic aerosols (Fan et al., 2022a; Hamilton et al.,
422 2021; Surratt et al., 2008). This result also revealed that other N- or S-containing
423 heteroatom compounds in CHONS species, such as those derived from the oxidation of
424 pyrrole and thiophene (Jiang et al., 2019), contributed a smaller intensity fraction at the
425 urban site (11 ± 2%) than at the suburban site (27 ± 4%).

426 3.2 Potential precursor assignment of organosulfur species

427 The detected OSs were classified into four groups based on their potential
428 precursors, including BVOC-derived OSs (e.g., isoprene/monoterpene/sesquiterpene-
429 derived OSs and other BVOC-derived OSs from the precursors of green leaf volatiles),
430 anthropogenic VOCs (AVOC)-derived OSs from the precursors of PAHs and
431 anthropogenically emitted alkanes, multiple-sources-derived OSs, and OSs with
432 unidentified precursors. The potential precursors of multiple-sources-derived OSs
433 include carbonyl compounds (e.g., glyoxal and glycolaldehyde), unsaturated fatty acids,
434 and long-chain alkenes emitted from biogenic and/or anthropogenic sources (Fu et al.,
435 2008; Passananti et al., 2016; Shang et al., 2016). The precursor assignment of OSs was
436 performed by matching the measured OS molecular formulas to the published OSs
437 whose precursors have been verified. This approach has been widely used because its
438 feasibility relies on the high mass resolution of HR-MS, which enables reliable mass
439 peak detection and the assignment of formulas within a narrow mass tolerance (Cai et



440 al., 2020; Wang et al., 2019; Yang et al., 2023). Details of these OS formulas with the
441 determined precursors are listed in Tables S5–S9.

442 Figure 3e shows the intensity distribution of OSs derived from various precursors
443 among the three campaigns. A total of 270 OS molecular formulas could be assigned to
444 potential precursors, accounting for approximately 70% of the total CHOS intensity.
445 Among them, BVOC-derived OSs contributed $32 \pm 13\%$ in suburban summer, $30 \pm 12\%$
446 in urban summer, and $28 \pm 6.7\%$ in urban winter, whereas AVOC-derived OSs
447 accounted for $16 \pm 4.6\%$, $18 \pm 4.7\%$, and $20 \pm 2.4\%$, respectively. These results are
448 consistent with the molecular signatures discussed above, suggesting relatively stronger
449 biogenic contributions in suburban and urban summer, but enhanced anthropogenic
450 influences in urban winter. Monoterpene-derived OSs constituted the dominant fraction
451 of BVOC-derived OSs. Several highly abundant formulas of monoterpene-derived OSs,
452 including $C_9H_{16}O_6S$ (m/z 251.0595), $C_{10}H_{18}O_5S$ (m/z 249.0802), and $C_9H_{16}O_7S$ (m/z
453 267.0544), and $C_{10}H_{18}O_7S$ (m/z 281.0701), have been frequently reported in ambient
454 aerosols and are commonly attributed to acid-catalyzed reactions of monoterpene
455 oxidation products (Bruggemann et al., 2020; Surratt et al., 2008; Wang et al., 2020b).
456 By contrast, isoprene-derived OSs contributed a much smaller fraction to the total
457 CHOS intensity than monoterpene-derived OSs.

458 For identified AVOC-derived OSs, aromatic-derived OSs contributed only minor
459 fractions to the total CHOS intensity in suburban summer ($1.3 \pm 0.4\%$), urban summer
460 ($1.5 \pm 0.6\%$), and urban winter samples ($1.9 \pm 0.7\%$), whereas long-chain alkane-
461 derived OSs accounted for much larger fractions ($15 \pm 4.5\%$, $17 \pm 4.6\%$, and $18 \pm 2.4\%$,
462 respectively). Polycyclic aromatic hydrocarbons have been recognized as important
463 precursors of aromatic OSs based on laboratory and field evidence (Kundu et al., 2013;
464 Riva et al., 2015). Aromatic OSs with benzyl and polycyclic aromatic carbon backbones,
465 such as $C_6H_6O_4S$, $C_7H_6O_4S$, $C_7H_8O_4S$, $C_8H_8O_4S$, and $C_9H_{12}O_5S$ derived from the
466 photooxidation of naphthalene and 2-methylnaphthalene (Riva et al., 2015), have been
467 widely observed in urban and semi-rural fine aerosols worldwide (Bruggemann et al.,
468 2020; Jiang et al., 2022; Wang et al., 2019; Yang et al., 2023), and were also detected
469 in our samples. Nevertheless, despite their frequent occurrence, only a few aromatic
470 OSs with relatively low intensity could currently be unambiguously classified. In
471 contrast, aromatic CHOS species with $X_c \geq 2.5$ accounted for $16 \pm 3.8\%$ and $19 \pm 2.9\%$
472 of total CHOS species by number in suburban and urban summer, respectively, and
473 increased to $22 \pm 3.4\%$ in urban winter, suggesting enhanced anthropogenic emissions
474 in winter. The intensity fractions of aromatic CHOS in Shanghai (7%–18%) were
475 comparable to those reported for Guangzhou (9%–20%) (Jiang et al., 2022), further
476 underscoring a non-negligible influence of anthropogenic emissions on the molecular



477 composition of atmospheric OrgSs in Chinese megacities. In addition, multiple-source
478 OSs showed a slight increase in their contribution to total CHOS intensity from
479 suburban summer ($22 \pm 7.2\%$) to urban summer ($23 \pm 6.8\%$) and further to urban winter
480 ($27 \pm 6.9\%$), probably hinting a stronger influence of anthropogenic emissions on their
481 sources and formation processes.

482 The precursors for $28 \pm 15\%$ of OS molecular formulas cannot yet be confidently
483 recognized. Among these, subgroup B, defined as organosulfates with $C > 8$, $DBE < 3$,
484 and $3 < O < 7$ for CHOS, contributed $6.7 \pm 3.3\%$, $16 \pm 8.3\%$, and $8.7 \pm 2.5\%$ to the total
485 CHOS intensity in suburban summer, urban summer, and urban winter, respectively.
486 Subgroup B is characterized by a high molecular weight, long alkyl chains, and a low
487 degree of unsaturation and oxidation. Tao et al. (2014) speculated that the precursors of
488 subgroup B could be long-chain alkanes from traffic emissions. Long-chain alkanes can
489 undergo rapid photooxidation under typical urban conditions (Lim & Ziemann, 2009;
490 Yee et al., 2013), yielding hydroxylated or carbonylated intermediates that were further
491 esterified to form alkyl OSs in acidic aerosol. Riva et al. (2016a), through smog-
492 chamber photooxidation experiments of long-chain alkanes, revealed that alkane-
493 derived OSs can also form via gaseous epoxide precursors with subsequent acid-
494 catalyzed reactive uptake onto sulfate aerosols and/or heterogeneous reactions of
495 hydroperoxides (Fig. 6d). Moreover, the formation of OSs from heterogeneous
496 reactions of gaseous SO_2 with unsaturated fatty acids (USFA) was also important for
497 these highly saturated OSs (Shang et al., 2016; Zhu et al., 2019). Large amounts of
498 USFA have been observed in high-temperature cooking emissions, especially Chinese-
499 style cooking with more frying, which is among the important emission sources in
500 urban areas (Zhao et al., 2015; Zhu et al., 2019). Accordingly, subgroup B contributed
501 more to total CHOS at the urban site ($16 \pm 8.3\%$) than at the suburban site ($6.7 \pm 3.3\%$)
502 during the summer, indicating a greater influence of urban hydrocarbon emissions.
503 Moreover, the intensity fraction of subgroup B showed positive correlations with RH
504 (Fig. S3a), suggesting that humidity-driven heterogeneous or multiphase processing
505 may facilitate the formation of these long-chain OSs. Although aromatic VOCs have
506 been widely studied as drivers of particulate air pollution in densely populated regions
507 like Shanghai (Gao et al., 2019; Han et al., 2023a), our results suggest that long-chain
508 hydrocarbons and their corresponding OS products also play an important role in the
509 formation of particulate organosulfurs in urban atmospheres.

510 In contrast to the relatively well-established precursor assignments for CHOS
511 compounds, only a limited number of CHONS species could be linked to specific
512 precursors (Tables S5–S9). Therefore, we further evaluated the potential precursor
513 origins of CHONS species based on their Van Krevelen distributions, the precursor-



514 based classification of NOSs, and molecular fingerprints from representative source
515 samples. As shown in Figure 4a–c, CHONS compounds, particularly NOSs, were
516 predominantly distributed in the unsaturated aliphatic region ($H/C \approx 1.5\text{--}2.0$,
517 $(O-3S-2N)/C \approx 0.1\text{--}0.6$), suggesting that their precursors are mainly derived from the
518 oxidation products of BVOCs, such as isoprene and monoterpenes (Hamilton et al.,
519 2021; Lin et al., 2012b; Surratt et al., 2008), with additional contributions from traffic-
520 related alkenes and long-chain hydrocarbon oxidation intermediates (Riva et al., 2016a;
521 Tao et al., 2014). The distributions of suburban and urban samples in summer were
522 broadly similar, indicating comparable dominant precursors and formation mechanisms.
523 In contrast, urban winter samples exhibited a more constrained distribution toward
524 lower $(O-3S-2N)/C$ values, consistent with a stronger influence of NO_3 -initiated
525 nighttime oxidation under high NO_x , low temperature, and reduced photochemical
526 activity (Brown & Stutz, 2012; Hamilton et al., 2021). Figure 5e shows that seasonal
527 variability was more influential than spatial differences in shaping NOS composition.
528 Monoterpene-derived NOSs predominated in most samples (11%–78%), underscoring
529 monoterpenes as important precursors of NOSs, which aligns with previous studies (Lin
530 et al., 2012b; Surratt et al., 2008). In contrast, the relative contribution of isoprene-
531 derived NOSs decreased markedly in urban winter, likely due to reduced isoprene
532 emissions under low-temperature conditions (Guenther et al., 2012; Li et al., 2011).
533 Furthermore, monoterpene-derived NOSs exhibited noticeable diurnal variability,
534 especially in summer, with generally higher contributions at night (Fig. 4e and S4),
535 which is consistent with an important role of nighttime NO_3 -initiated oxidation in NOS
536 formation (Brown & Stutz, 2012; Ng et al., 2017).

537 To further constrain the possible precursor origins of CHONS, we compared their
538 molecular characteristics with those reported for representative source samples in
539 previous studies, including coal combustion organic aerosols (CCOAs), biomass
540 burning organic aerosols (BBOAs), vehicle emissions, tunnel aerosols, and emissions
541 from nonroad excavators and ships (Cui et al., 2019; Tang et al., 2020a). As shown in
542 Figure 4a–d, substantial overlap in CHONS species was observed between our field
543 samples and previously reported source samples, particularly in the high H/C region.
544 However, ambient samples extended toward more oxidized regions, indicating that
545 CHONS species are predominantly formed via secondary processes rather than direct
546 primary emissions. The most abundant CHONS species in Shanghai were dominated
547 by unsaturated aliphatic compounds (Fig. 4a–d), whose overall characteristics more
548 closely resembled those of tunnel aerosol samples that had likely undergone
549 atmospheric aging processes. By contrast, CHONS compounds in fresh vehicle
550 emissions were abundant in aromatics, with 59% of the identified CHONS species



551 showing $X_c \geq 2.50$ (Table S11). Likewise, a large number of aromatic and highly
552 unsaturated CHONS species were detected in both BBOAs and CCOAs, but their
553 molecular characteristics differed substantially from the dominant CHONS components
554 observed in Shanghai. Notably, although approximately 36%, 40%, and 44% of
555 CHONS species by number could be classified as aromatic organosulfur compounds in
556 suburban summer, urban summer, and urban winter, respectively, most of these species
557 exhibited relatively low abundances (Fig. 4a–d and Table S11). Thus, although
558 combustion sources can directly emit a large variety of CHONS compounds, the low-
559 oxidation and aromatic CHONS species that dominate the source samples contributed
560 only weakly to the MS intensity of the ambient CHONS pool observed here. These
561 results suggest that the CHONS in Shanghai may be less influenced by primary
562 emissions and more likely driven by secondary formation processes, such as secondary
563 formation via combustion-emitted precursors. Similar features were also observed in
564 the comparison between CHOS species detected in our field samples and source
565 samples (Fig. 3 and Table S10).

566 Taken together, the precursor assignment and source comparison highlight the
567 important role of secondary formation from both biogenic and anthropogenic
568 precursors in shaping the molecular composition of particulate OrgSs in Shanghai.
569 Given that isoprene- and monoterpene-derived OSs and NOSs are representative
570 BVOC-derived species and are sensitive to variations in precursor emissions, oxidant
571 regimes, NO_x/NO_3 chemistry, and sulfate-related multiphase processing, their spatial,
572 seasonal, and diurnal variability was further examined.

573 3.3 Spatiotemporal variability of isoprene/monoterpene-derived OSs and NOSs

574 During summer, isoprene/monoterpene-derived OSs at both suburban and urban
575 sites exhibited distinct diurnal variation, with significantly higher relative abundance
576 during the day than at night ($p < 0.05$; Fig. 5a,b). This pattern highlights the dominance
577 of daytime photochemical processes, likely driven by enhanced biogenic emissions
578 under high temperatures and intense solar radiation, followed by $\bullet\text{OH}$ -initiated
579 oxidation (Surratt et al., 2010; Xu et al., 2015). In contrast, no significant diurnal
580 variation was observed in winter urban samples ($p > 0.05$), suggesting that the
581 formation of these OS species in winter was less controlled by a single photochemical
582 pathway and more influenced by multiple pathways under low-temperature, weak-
583 radiation, and high- NO_x conditions. Notably, isoprene-derived OSs showed only weak
584 seasonal differences, whereas monoterpene-derived OSs were more abundant in winter
585 than in summer (Fig. 5a,b). This appears to suggest that enhanced nighttime and/or
586 heterogeneous production in winter may partially compensate for the decline in



587 photochemical efficiency. Recent studies suggested that OSs could also be formed
588 through heterogeneous or multiphase reactions between SO₂ and monoterpene-derived
589 SOA or unsaturated fatty acids in the particle phase (Fig. 6b) (Yao et al., 2019; Zhu et
590 al., 2019). Since SO₂ levels during our winter campaign was 1.6–1.8 times higher than
591 those during the summer campaigns (Table S1), such formation pathways involving
592 SO₂ likely led to the higher abundance of monoterpene-derived OSs in winter. Such
593 compensation may also be aided by additional inputs of isoprene and monoterpenes
594 from anthropogenic activities, such as biomass burning, which could help sustain OS
595 formation under cold conditions (Akagi et al., 2011; Gilman et al., 2015; Yang et al.,
596 2025). Overall, these results suggest that photochemistry was the dominant driver of
597 isoprene/monoterpene-derived OS formation in summer, whereas wintertime formation
598 was governed by a more complex, multi-pathway regime involving nocturnal NO₃•
599 chemistry and heterogeneous or multiphase processes.

600 The diurnal patterns of monoterpene/isoprene-derived NOSs differed markedly
601 from those of OSs. The day–night variations of isoprene-derived NOSs were not
602 obvious across the three field campaigns (Fig. 5c), suggesting that their formation may
603 be jointly influenced by daytime •OH oxidation and nighttime NO₃• oxidation. In
604 contrast, monoterpene-derived NOSs exhibited obvious nighttime enhancements ($p <$
605 0.05 ; Fig. 5d), with significantly higher relative abundances at night than during the
606 day in both suburban and urban summer samples ($p < 0.01$). This trend underscores that
607 •NO₃-initiated nighttime oxidation is likely an important pathway for their formation
608 (Fig. 6e), consistent with previous field and laboratory studies highlighting the
609 important role of NO₃• oxidation of monoterpenes in NOS formation (Hamilton et al.,
610 2021; Surratt et al., 2008; Yang et al., 2023). Elevated nighttime NO_x levels likely
611 further favored monoterpene-NOS formation via nighttime NO₃• oxidation (Table S1).
612 Moreover, the lower daytime relative abundance of monoterpene-derived NOSs could
613 be exacerbated by daytime photolysis and decomposition processes (He et al., 2011;
614 Wang et al., 2023). Spatial and seasonal variations of monoterpene-derived NOSs
615 further highlight these drivers (Fig. 5d and S4). In summer, monoterpene-derived NOSs
616 were more abundant in suburban than in urban samples, likely reflecting a stronger
617 biogenic influence and the greater importance of monoterpene precursors at the
618 suburban site in the YRD (Ma et al., 2022; Wang et al., 2020a). Notably, their overall
619 relative abundance in urban samples was higher in winter than that in summer,
620 suggesting that a more favorable oxidation environment under high-NO_x conditions
621 may promote the production efficiency of monoterpene-derived NOSs, partly
622 compensating for the typically lower biogenic monoterpene emissions in winter (Wang
623 et al., 2020b). Additionally, previous studies have shown that monoterpenes can also be



624 emitted from biomass burning, besides their well-known biogenic emission (Akagi et
625 al., 2011; Gilman et al., 2015), which may further contribute to the accumulation of
626 monoterpene-derived NOSs during the cold season.

627 **3.4 Formation Pathways of OrgSs and the influencing factors**

628 Atmosphere OrgSs in Shanghai originate from diverse sources, with secondary
629 formation serving as the dominant contributor. Despite the identification of multiple
630 reaction pathways for OS formation, their formation mechanisms in the real atmosphere
631 remain incompletely understood. Among the known pathways, the acid-catalyzed ring-
632 opening reaction of epoxides derived from the oxidation of VOCs is widely recognized
633 as a kinetically feasible and key mechanism for the atmospheric formation of OSs (Fig.
634 6a) (Bruggemann et al., 2020; Cai et al., 2020). For instance, previous studies have
635 demonstrated that high concentrations of isoprene-derived OSs observed in Beijing and
636 Guangzhou were mainly produced through the acid-catalyzed ring-opening of isoprene
637 epoxydiol intermediates (Jiang et al., 2022; Wang et al., 2018b). Furthermore, NOSs
638 can be formed via the nighttime NO_3 -initiated oxidation of monoterpenes on acidic
639 sulfate seed particles (Ng et al., 2017; Surratt et al., 2008). The formation and molecular
640 distribution of OS and NOS products are generally governed by multiple environmental
641 factors, including precursor abundance (e.g., organic precursors and anthropogenic
642 pollutants such as NO_x and SO_2), aerosol acidity, RH, and oxidant levels. To disentangle
643 how these complex variables drive the molecular distribution of OrgSs across different
644 spatiotemporal settings (suburban summer, urban summer, and urban winter), NMDS
645 analysis of OrgSs was conducted (Fig. 7). Overall, samples from different seasons and
646 urban/suburban sites exhibit distinct separation in the two-dimensional ordination space,
647 indicating that the molecular composition of OrgSs is highly sensitive to changes in
648 precursor sources, oxidation conditions, and the particulate-phase reaction environment.

649 For suburban summer samples, Cl^- was positively correlated with the NMDS1
650 dimension, whereas tracers indicating combustion emissions (EC and CO), biomass
651 burning (K^+), and secondary inorganic nitrogen species (NO_3^-), were mainly distributed
652 on the negative direction of NMDS1 (Fig. 7a). This suggests that the NMDS1
653 dimension may reflect the integrated influence of different air mass sources on the
654 molecular composition of OrgSs. It should be noted that Cl^- can originate from both
655 combustion sources (e.g., coal smoldering) and marine aerosols (Wu et al., 2024).
656 Given that the sampling site was situated in a coastal region and that Cl^- exhibits an
657 opposite trend to typical combustion tracers like EC, Cl^- is more likely to represent the
658 influence of marine air masses or sea-salt input. Therefore, the divergent trends of Cl^-
659 and EC along the NMDS1 axis may highlight the distinct impacts of anthropogenic



660 versus marine sources on the molecular composition of OrgSs in suburban summer.
661 Notably, the suburban summer samples exhibited a pronounced diurnal separation
662 along the NMDS2 dimension. Daytime samples primarily clustered along the positive
663 axis of NMDS2, aligning with T and O₃; conversely, nighttime samples aggregated in
664 the negative direction, consistent with RH and NO₂. These results indicate that the
665 molecular composition of OrgSs during the daytime was more strongly governed by
666 high temperature and photochemical oxidation processes. In summer, elevated daytime
667 temperatures facilitate the emission of volatile organic compounds (VOCs), especially
668 BVOCs, while stronger photochemistry and enhanced O₃ levels further promote the
669 oxidation of VOCs and the subsequent formation of OSs. In contrast, RH and NO₂
670 emerge as the predominant factors influencing nighttime OrgS composition. On one
671 hand, RH affects OSs formation by influencing the uptake of gaseous precursors like
672 SO₂ by VOCs (Li et al., 2024; Nestorowicz et al., 2018; Ye et al., 2018). On the other
673 hand, NO₂ and its associated nighttime oxidative system (e.g., NO₃ radical) may jointly
674 influence the formation of both OSs and NOSs by modifying the oxidation pathways
675 and product branching of organic precursors, as well as through competition or coupling
676 with heterogeneous processes involving sulfate and acidic particles (Fan et al., 2022b;
677 Hamilton et al., 2021).

678 For urban summer samples, primary combustion sources related to anthropogenic
679 activities also appeared to play an important role in shaping the molecular composition
680 of OrgSs, whereas the influence of marine air masses was comparatively weak. In
681 contrast to the prominent role of inorganic nitrogen in suburban summer samples, the
682 molecular distribution of OrgSs in urban summer is markedly influenced by SO₄²⁻ (Fig.
683 7b). This observation aligns with previous studies in Shanghai and other polluted
684 regions, which have highlighted the importance of sulfate-related processes in OS
685 formation (Wang et al., 2018b; Yang et al., 2023). SO₄²⁻ acts not only as a critical
686 reactant but also as a regulator of particle acidity and ALWC, which in turn affects the
687 reactive uptake and heterogeneous conversion efficiency of precursors (Wang et al.,
688 2021b). The distinct influences of different anthropogenic emission species on OrgSs
689 distribution between urban and suburban summer underscore the varying impact of
690 human activities on OS formation across different environments. Nevertheless, the key
691 drivers of the diurnal variation in OrgSs composition were generally consistent between
692 the two sites. Daytime samples cluster with T and O₃ along the positive axis of NMDS2,
693 while nighttime samples aggregate with RH along the negative axis. This result
694 suggests that site type had only a limited influence on the diurnal distribution of OrgSs
695 in summer. The main processes governing the diurnal variability of OrgSs appeared to
696 be broadly similar across the two sites, mainly involving enhanced photochemical



697 oxidation during the daytime and RH-dependent multiphase processing at night.

698 For urban winter samples, the NMDS results suggest that the factors influencing
699 the molecular distribution of OrgSs were more complex than those in summer (Fig. 7c).
700 In addition to the impact of primary anthropogenic emissions, inorganic nitrogen and
701 sulfur species emerge as significant drivers, clustering along the negative axis of
702 NMDS1. ALWC is also distributed in the negative NMDS1 direction, suggesting that
703 it was another key parameter associated with the wintertime molecular distribution of
704 OrgSs. ALWC is governed not only by RH, but also by the loading of hygroscopic
705 inorganic components such as NO_3^- and SO_4^{2-} . Under the humid and polluted
706 conditions typical of winter, the accumulation of these species can significantly enhance
707 ALWC and thereby favor aqueous-phase processing (Wang et al., 2021b). For OS
708 formation, elevated ALWC may promote the mass transfer and reactive uptake of SO_2
709 and oxidized organic intermediates into the particle phase, while it also provides a
710 medium for subsequent aqueous-phase transformations. These results suggest that
711 ALWC is an important particle-phase parameter influencing the formation and
712 molecular composition of organosulfur compounds in winter. Aerosol pH appeared to
713 be another pivotal factor. Lower pH values not only favor the reactive uptake of
714 epoxides and other oxidative intermediates onto acidic sulfate particles, but may also
715 facilitate aqueous-phase esterification via acid-catalyzed mechanisms (Wang et al.,
716 2018b). Therefore, the combination of elevated ALWC and acidic particle conditions
717 likely provided a favorable multiphase reaction environment for the secondary
718 formation and molecular evolution of urban OrgSs in winter. Overall, while the
719 molecular composition of OrgSs in summer was primarily controlled by photochemical
720 oxidation and humidity-driven heterogeneous processes, wintertime OrgSs appeared to
721 be more strongly influenced by the combined effects of enhanced primary emissions,
722 the accumulation of secondary inorganic species, and the synergistic regulation of the
723 particle-phase reaction environment by ALWC and pH.

724 **4. Conclusions and atmospheric implications**

725 This study provides a molecular-level characterization of atmospheric
726 organosulfur compounds (OrgSs) in Shanghai under contrasting atmospheric
727 environments, including suburban summer, urban summer, and urban winter. A total of
728 1964, 1914, and 2689 OrgS molecular formulas were detected in suburban summer,
729 urban summer, and urban winter aerosol samples, respectively, indicating substantially
730 higher molecular diversity in winter than in summer. OrgSs were dominated by CHOS_1
731 and CHON_1S_1 species, which together accounted for most of the detected formulas and



732 signal intensity. As many as 79%–92% of OrgS species had the elemental composition
733 of $(4s + 3n)/o \leq 1$ in their formulas, suggesting that they were potential OSs or nitrooxy-
734 OSs. The molecular characteristics of both CHOS and CHONS exhibited pronounced
735 seasonal variability. Compared with summer, wintertime OrgSs showed lower O/C
736 ratios but higher DBE values and aromaticity, suggesting a larger contribution from
737 anthropogenic emissions and more unsaturated structures under winter conditions.
738 Although aliphatic structures dominated the OrgS pool overall, the increased abundance
739 of aromatic and polyaromatic species in winter further pointed to enhanced combustion-
740 related influences during the cold season.

741 Potential precursor analysis showed that BVOC-derived OSs were the largest
742 identified group, followed by OSs derived from anthropogenic VOCs and mixed-source
743 precursors. Monoterpene-derived OSs dominated the biogenic fraction, while the
744 intensity fraction of isoprene-derived OSs increased in summer, especially at the
745 suburban site. In contrast, aromatic and long-chain alkane-derived OSs increased in
746 winter, highlighting a stronger anthropogenic influence. CHONS species showed
747 weaker similarity to primary source profiles and appeared to be more strongly affected
748 by secondary formation, although combustion emissions may still provide important
749 precursors in polluted environments.

750 The diurnal and seasonal behaviors of isoprene/monoterpene-derived OSs and
751 NOSs revealed distinct controlling mechanisms. In summer, isoprene/monoterpene-
752 derived OSs peaked during the daytime due to photochemical oxidation, whereas
753 monoterpene-derived NOSs were significantly enhanced at night via NO_3 -initiated
754 oxidation. The NMDS results further demonstrated that the molecular composition of
755 OrgSs was highly sensitive to seasonal, spatial, and diurnal changes in precursor
756 sources, oxidant levels, and particle-phase reaction conditions. In summer, daytime
757 OrgSs were mainly associated with temperature and O_3 , indicating the importance of
758 photochemical oxidation, whereas nighttime composition was more closely linked to
759 RH and NO_2 , suggesting a strong role of humidity-driven multiphase processing and
760 nighttime oxidation. The similar day–night separation observed at the suburban and
761 urban sites implies that the dominant mechanisms controlling summer diurnal
762 variability were broadly consistent across different functional zones. In winter, the
763 molecular distribution of OrgSs became more strongly associated with inorganic
764 nitrogen and sulfur species, ALWC, and aerosol acidity, indicating that aqueous-phase
765 and acid-catalyzed processes were more important under polluted winter conditions.
766 Together, these observations suggest that the formation of atmospheric OrgSs in
767 Shanghai is jointly regulated by precursor availability, oxidation pathways, and the
768 physicochemical properties of the particle phase. Overall, this study provides new



769 molecular-level evidence that improves our understanding of the sources, formation
770 pathways, and environmental controls of atmospheric OrgSs, with implications for
771 interpreting aerosol evolution and for better representing organosulfur chemistry in
772 atmospheric models.

773 **Data availability**

774 The data presented in this work will be uploaded to an open archive prior to publication.

775 **Author contributions**

776 D.C. and X.G. conceived and designed the study. S. Z. and Y.C. carried out the field
777 sampling. D.C., S.R., and Y. F. performed the UHPLC-Orbitrap MS analysis and data
778 processing. X.G. and J. C. acquired funding and undertook project management. D.C.
779 prepared the manuscript with contributions from all co-authors.

780 **Competing interests**

781 The authors declare that they have no conflict of interest.

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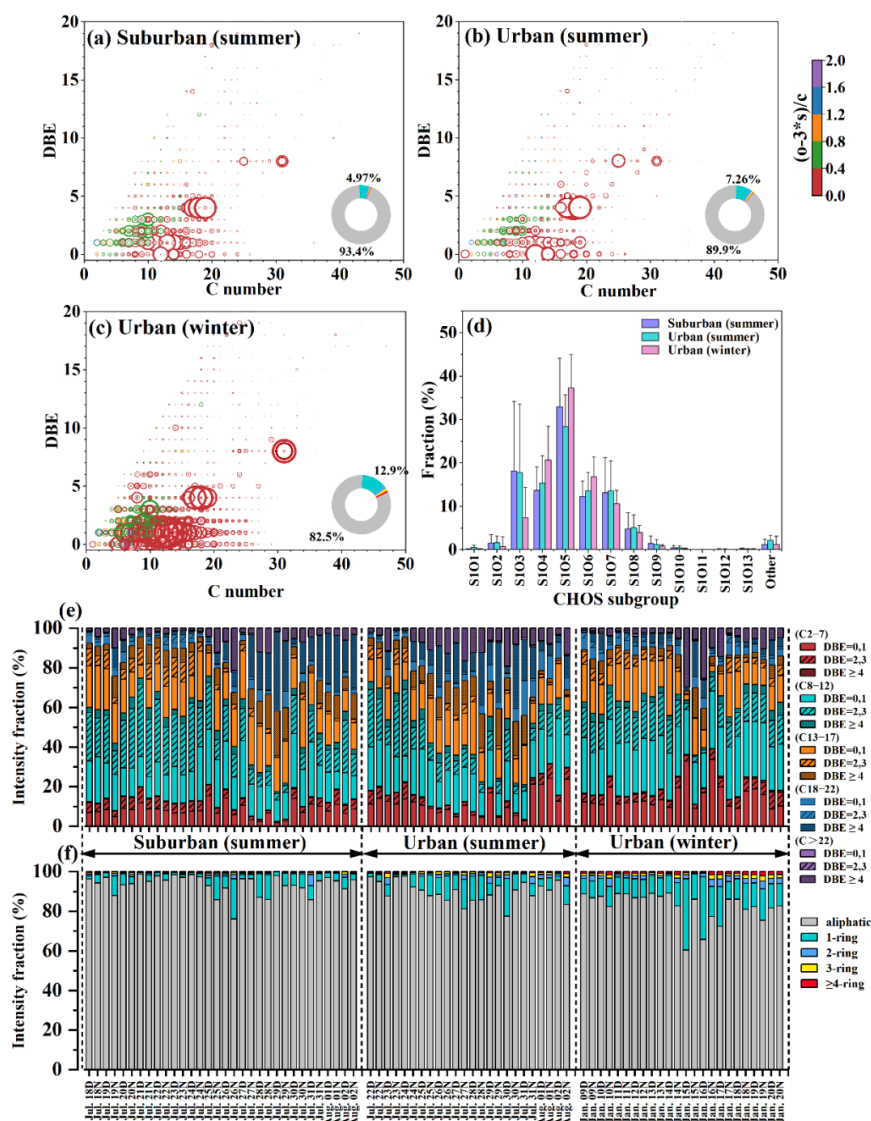


Figure 1 Molecular distribution of CHOS compounds detected by Orbitrap MS for the PM_{2.5} samples collected in the three field campaigns. (a–c) Double-bond equivalent (DBE) vs C number for all the CHONS compounds in the suburban (summer) and urban (summer, winter). The circle area is proportional to the square root of the relative abundance of individual molecules, and the color bar denotes the degree of oxidation. (d) Classification of CHOS species into different subgroups according to the numbers of S and O atoms in their molecules. (e) Percentages of the intensity of each subgroup divided based on the DBE value and the length of carbon skeleton in the formulas. (f) Intensity percentages of each subgroup which were divided based on the X_c value of formulas.

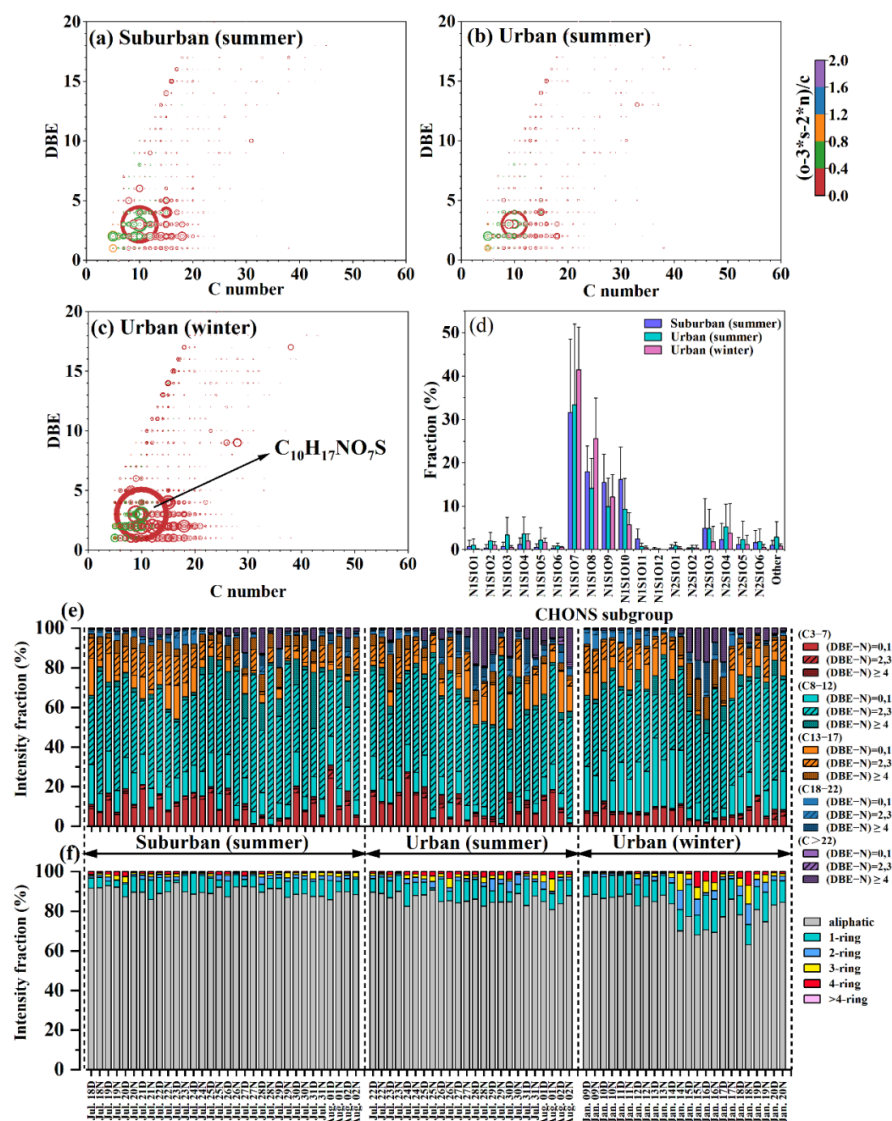


Figure 2 Molecular distribution of CHONS compounds detected by Orbitrap MS for the PM_{2.5} samples collected in the three field campaigns. (a–c) Double-bond equivalent (DBE) vs C number for all the CHONS compounds in the suburban (summer) and urban (summer, winter). The circle area is proportional to the square root of the relative abundance of individual molecules, and the color bar denotes the degree of oxidation. Several of the most abundant CHONS species listed in descending order by their average intensities in (a) are C₉H₁₅NO₈S, C₁₀H₁₇NO₇S, C₁₀H₁₇NO₈S, C₁₀H₁₇NO₉S, C₁₀H₁₇NO₁₀S, and C₁₅H₂₅NO₇S (Figure S4). (d) Classification of CHONS species into different subgroups according to the numbers of S and O atoms in their molecules. (e) Percentages of the intensity of each subgroup divided based on the (DBE–N) value and the length of carbon skeleton in the formulas. (f) Intensity percentages of each subgroup which were divided based on the X_c value of formulas.

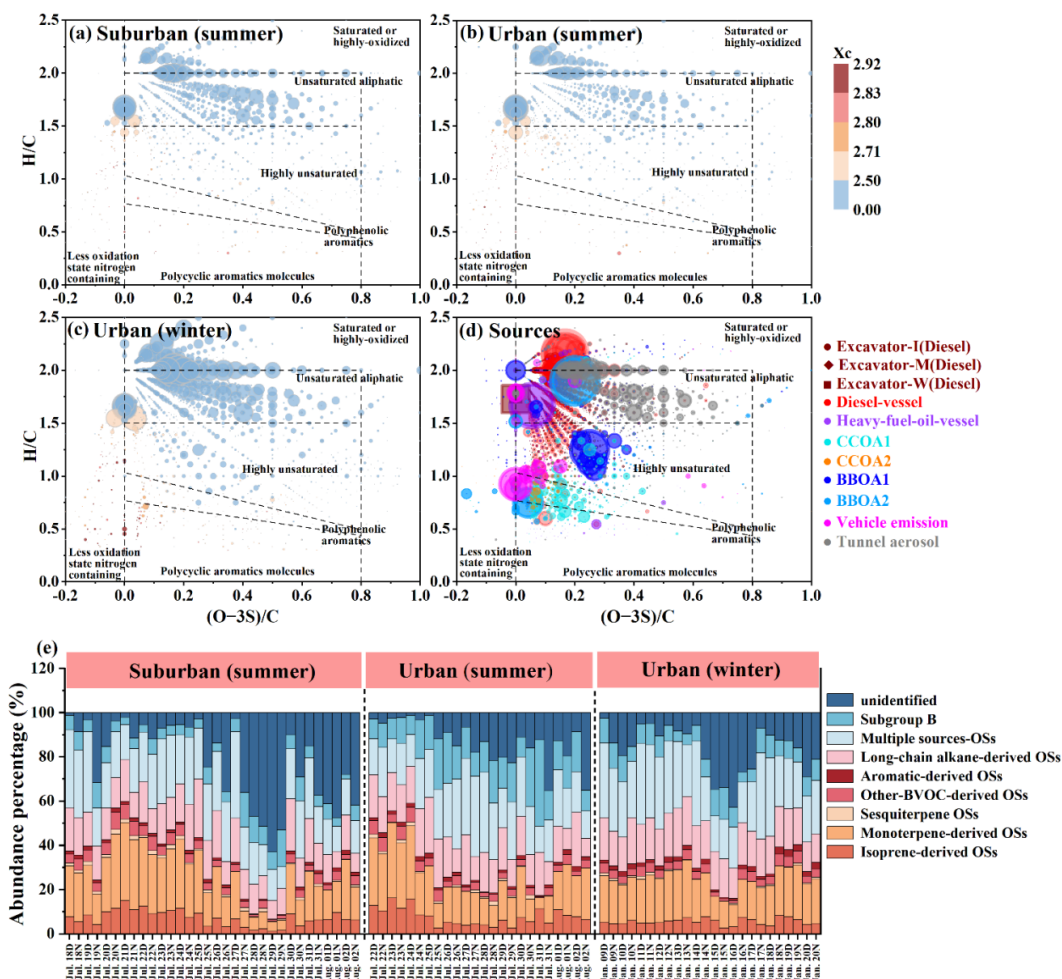


Figure 3 (a–d) Van Krevelen diagrams of CHOS for the field samples collected in Shanghai (a–c) and source samples (d) obtained from Cui et al. (2019) and Tang et al. (2020), including biomass burning organic aerosols (BBOAs), coal combustion organic aerosols (CCOAs), vehicle emissions, tunnel aerosols, and off-road engine emissions (excavator and vessel). Excavator-I, -M, and -W denote the operation modes of idling, moving, and working respectively. The circle area is proportional to the square root of the relative abundance of individual molecules in Figure 3 (a–c). The marker size denotes the percentages of MS intensity to the total identified CHOS compounds in Figure 3d. (e) Spatial, seasonal, and diurnal variations of potential precursors of detected OSs to the total intensity of identified CHOS species; subgroup B denotes OSs with unidentified precursors that satisfy $C > 8$, $DBE < 3$ and $3 < O < 7$, whereas “unidentified” denotes NOSs with unidentified precursors that fall outside these criteria.

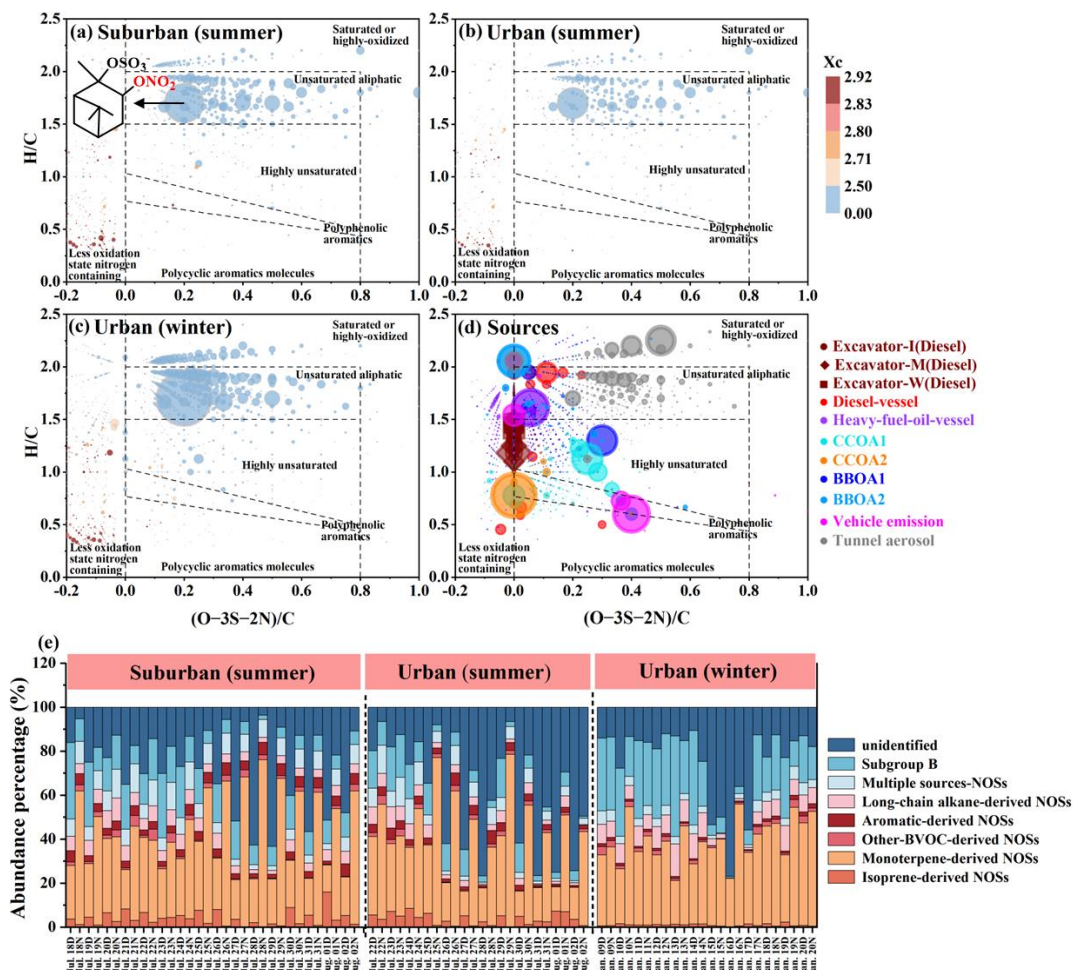


Figure 4 (a–d) Van Krevelen diagrams of CHONS for the field samples collected in Shanghai (a–c) and source samples (d) obtained from Cui et al. (2019) and Tang et al. (2020), including biomass burning organic aerosols (BBOAs), coal combustion organic aerosols (CCOAs), vehicle emissions, tunnel aerosols, and off-road engine emissions (excavator and vessel). Excavator-I, -M, and -W denote the operation modes of idling, moving, and working respectively. The circle area is proportional to the square root of the relative abundance of individual molecules in Figure 4 (a–c). The marker size denotes the percentages of MS intensity to the total identified CHONS compounds in Figure 4d. (e) Spatial, seasonal, and diurnal variations of potential precursors of detected NOSs to the total intensity of identified CHONS species; subgroup B denotes NOSs with unidentified precursors that satisfy $C > 8$, $DBE < 3$ and $6 < O < 10$, whereas “unidentified” denotes NOSs with unidentified precursors that fall outside these criteria.

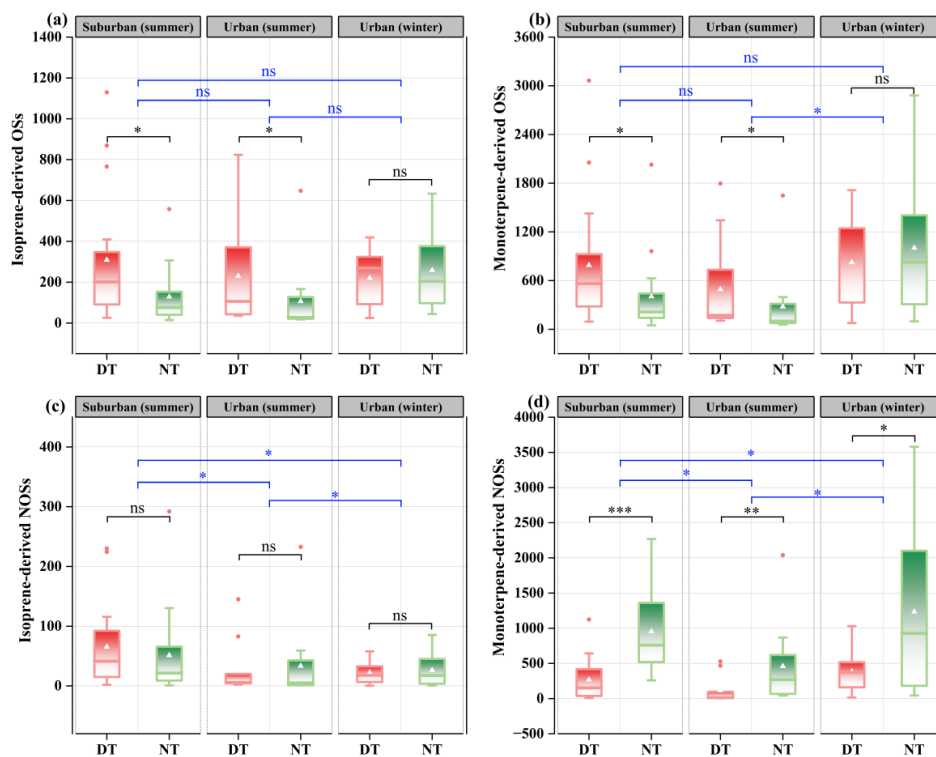


Figure 5. Spatial, seasonal, and diurnal variability of isoprene- and monoterpene-derived OSs (a, b) and NOSs (c, d). The y-axis represents the relative abundances of monoterpene/isoprene-derived OSs and NOSs, defined as peak areas normalized to the internal standard. The whisker plot boundaries are between the 10th and 90th percentiles, the box displays the 25th, 50th (median), and 75th percentiles, and the triangles display the mean values. Asterisks denote statistical significance (* $p < 0.05$, ** $p < 0.01$, *** $p < 0.001$), while “ns” indicates no significance.

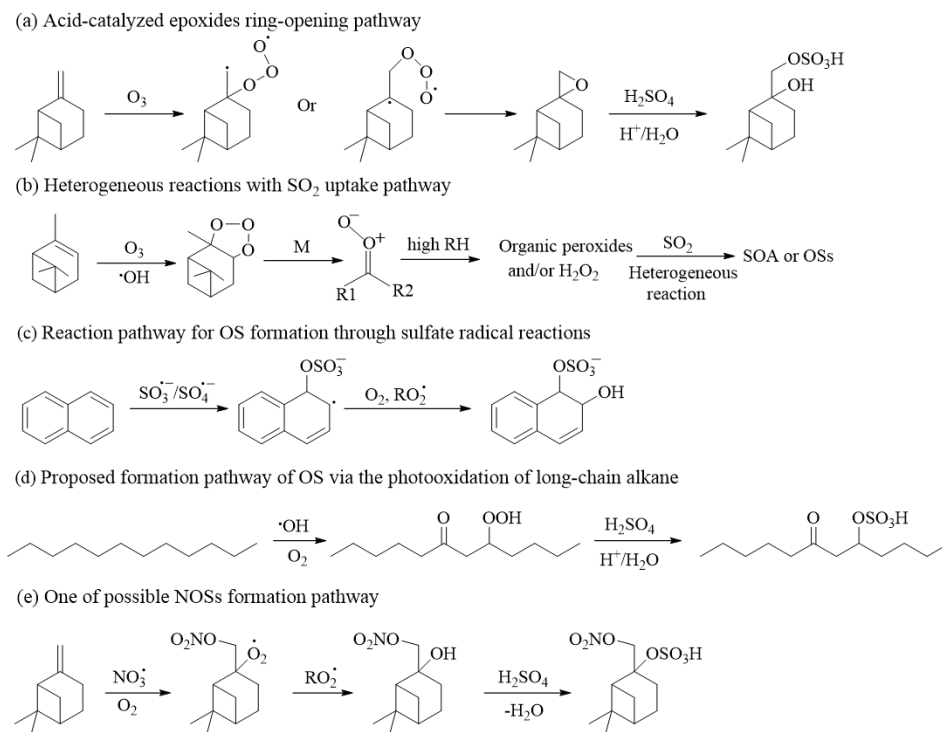


Figure 6. Potential formation mechanisms of OSs and NOSs in Shanghai (Brüggemann et al., 2020; Brüggemann et al., 2020; Surratt et al., 2008; Ye et al., 2018). (a) Proposed OSs formation pathway for acid-catalyzed ring-opening of epoxides. (b) Proposed OSs formation pathway for ozonolysis of unsaturated hydrocarbon in the presence of SO₂ at high relative humidity. (c) Reaction pathway for OSs formation through sulfate radical reactions. (d) One of the possible OSs formation pathways by long-chain alkanes (Riva et al., 2016). (e) One of the possible NOSs formation pathways.

