



# 1 **Distribution of cloud geometrical properties and types over the** 2 **Mediterranean: insights from a decade of CloudSat and 1 year of** 3 **EarthCARE measurements and ERA5 comparison.**

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15 **Abstract.** Space-based radar observations from the CloudSat mission are used to characterize cloud properties over the  
16 Mediterranean during 2007–2017. CloudSat’s Cloud Profiling Radar (CPR) provides vertically resolved cloud measurements,  
17 enabling analysis of cloud geometry and type by altitude and month. Low clouds (<1 km a.s.l.) occur up to four times more  
18 frequently over sea (32–39%) than land (8–12%). Over land, clouds are most common between 1–4 km, dominated by  
19 Stratocumulus (36%), with Altocumulus and Altostratus also present. Over sea, clouds peak at 0–2 km, with Stratocumulus  
20 most prevalent, especially in the West Mediterranean. High-level clouds peak at 8–11 km, reaching frequencies of ~30% across  
21 the region. Seasonally, Stratocumulus peaks in winter and autumn, while Cumulus maximizes in July–August, particularly in  
22 the East Mediterranean. Altostratus and Nimbostratus are least frequent in summer, whereas Altocumulus peaks over East  
23 Mediterranean in July–August. High clouds show maxima in spring–early summer and autumn, with minima in midsummer.  
24 CloudSat column water and ice are compared with ERA5 reanalysis. Both datasets show similar spatial and seasonal patterns,  
25 with higher values over the western and central Mediterranean. Liquid water differences (–0.004 to –0.008 kg/m<sup>2</sup>) indicate  
26 good agreement, while larger ice discrepancies (up to –0.024 ± 0.017 kg/m<sup>2</sup>) suggest underdetection of thin clouds in ERA5.  
27 Preliminary results from the EarthCARE mission are presented but are limited by shorter records. Overall, this study supports  
28 improved Climate Data Records and continuity between past, current and future satellite observations.

## 29 **1 Introduction**

30 Clouds are the primary modulators of the Earth’s radiation budget and still constitute the main source of uncertainty in model  
31 estimates of climate sensitivity (Randall et al., 2007). The processes governing their formation, evolution, geometrical and  
32 microphysical properties, as well as their radiative effects, are still not fully understood (IPCC, 2021). Different cloud types



33 have different radiative effects, which plays an indispensable role in modulating the global radiative budget. Main knowledge  
34 gaps and limitations in current state-of-the-art Global Climate Models (GCMs) are attributed to their inability to correctly  
35 describe the ice content in clouds, while the uncertainty is even higher for the quantification of ice and/or water content in  
36 Mixed Phase Clouds (Waliser et al., 2009; Komurcu et al., 2014; Barrett et al., 2017; McGraw et al., 2023). These limitations  
37 significantly affect the estimation of the cloud albedo in GCMs, altering the equilibrium climate sensitivity according to the  
38 different scenarios. Intercomparisons of contemporary climate models further reveal that these uncertainties in cloud ice and  
39 mixed-phase processes translate directly into biases in top-of-atmosphere radiative fluxes and cloud feedback diagnostics. The  
40 representation of supercooled liquid water, frozen hydrometeor mass, and their vertical distribution remain a primary source  
41 of spread in simulated cloud albedo and longwave cloud radiative effects, reinforcing the central role of cloud phase processes  
42 in determining model sensitivity to warming (Li, et al., 2023; Cesana et al., 2024). By identifying mixed-phase clouds and  
43 their microphysical properties, satellite observations can enhance the characterization of cloud processes and improve model  
44 accuracy (Forbes et al., 2016).

45 State-of-the-art methodologies use combined lidar-radar satellite observations to provide high-resolution vertical profiles of  
46 cloud properties and to evaluate or constrain cloud predictions in GCMs. For example, combined CloudSat and CALIPSO  
47 (Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation) observations provided globally cloud geometrical and  
48 microphysical properties from 2006 to 2017 (e.g., Li et al., 2015; raDAR/liDAR product (DARDAR-MASK product); Delanoë  
49 et al., 2008). Ceccaldi et al. (2013) developed and validated a synergistic radar–lidar cloud classification algorithm (DARDAR-  
50 MASK) that integrates CloudSat and CALIPSO measurements to distinguish cloud phases, including ice, liquid, and mixed-  
51 phase categories, demonstrating its utility for detailed cloud phase discrimination and vertical profiling from spaceborne active  
52 sensors (Ceccaldi et al., 2013; Delanoë and Hogan, 2010). The recent launch of the ESA's EarthCARE (Cloud, Aerosol and  
53 Radiation Explorer) mission, launched on 28 May 2024 (Wehr et al., 2023) is particularly important, as it is foreseen to  
54 continue the legacy of CloudSat by extending the global record of cloud observations. It is the largest and most complex Earth  
55 Explorer to date and is designed to ensure continuity of the active-sensor cloud record while providing enhanced sensitivity to  
56 cloud vertical structure (Feuillard et al., 2026). By extending spaceborne radar-lidar observations, EarthCARE offers improved  
57 insight into Mediterranean cloud climatology and strengthens the observational basis needed to evaluate and constrain climate  
58 model simulations (Enriquez-Alonso et al., 2016; Sanchez-Lorenzo et al., 2017).

59 Mediterranean Basin is a climate “hot spot,” characterized by a large variability of cloud systems, ranging from frontal and  
60 convective systems to cyclones (Miglietta et al., 2011), and has been the focus of several cloud studies. The study of Raveh-  
61 Rubin and Wernli (2015) offers a climatological analysis of large-scale wind and precipitation extremes in the Mediterranean  
62 from 1979 to 2012. They identify significant seasonal differences across subregions, with western Mediterranean precipitation  
63 extremes being more intense and occurring mainly in autumn, while East Mediterranean events predominantly occur in winter.  
64 Chaboureau and Claud (2006) analyzed the climatology of Mediterranean cloud systems and their links to large-scale  
65 circulation using data from the TIROS-N Operational Vertical Sounder. They reported that cloud systems are most frequently  
66 detected over western North Africa, the eastern coast of Spain, the Balearic Islands, the Gulf of Genoa, southern Italy, the



67 Aegean Sea, and Cyprus, with peak occurrence between March and October. Ioannidis et al. (2017) examined the seasonal  
68 variability of total cloud cover over the Mediterranean during 1948-2014 and identified three dominant modes of seasonal  
69 variation. Similarly, Levizzani et al. (2010) produced a 10-year climatology (1996-2005) of warm-season cloud patterns over  
70 Europe and the Mediterranean using Meteosat infrared observations.

71 While these studies have advanced our understanding of cloud occurrence and variability, they rely primarily on passive  
72 satellite observations, reanalyses, regional measurements, or overall cloud cover measurements, and therefore lack vertically  
73 resolved, long-term information from active sensors (e.g., radar–lidar), leaving gaps in understanding cloud vertical structure  
74 and seasonal phase variability across Mediterranean. More recent studies based on ground-based radar observations over  
75 selected subregions and limited time periods (e.g., Marinou et al., 2021) have provided detailed regional characterization of  
76 cloud types, but their spatial coverage remains restricted. Additionally, few studies have systematically compared multiannual  
77 satellite records from active missions such as CloudSat and CALIPSO with outputs from global climate models. Overall, a  
78 comprehensive analysis of cloud patterns over the entire Mediterranean region is lacking, especially regarding cloud vertical  
79 distributions and seasonal differences.

80 The aim of this study is threefold. First, we analyze the annual and seasonal vertical distribution of cloud properties across  
81 three subregions of the Mediterranean basin using spaceborne radar observations from CloudSat for the period 2007-2017.  
82 Second, we compare these observations with the ERA5 reanalysis to evaluate the representation of cloud vertical structure in  
83 GCMs. Third, we present recent measurements from the EarthCARE mission, which envisages to ensure observational  
84 continuity of CloudSat records while providing enhanced sensitivity to cloud vertical structure.

85 The paper is structured as follows: Section 2 outlines the methodology used in this study. Section 3 presents the CloudSat  
86 statistics (per height and per month), averaged over sea and land for three distinct Mediterranean regions, and the comparison  
87 of the cloud patterns derived from CloudSat with those obtained from the ERA5 reanalysis and the EarthCARE's patterns.  
88 Finally, Section 4 provides the summary and the main conclusions of the study.

## 89 **2. Data and Methodology**

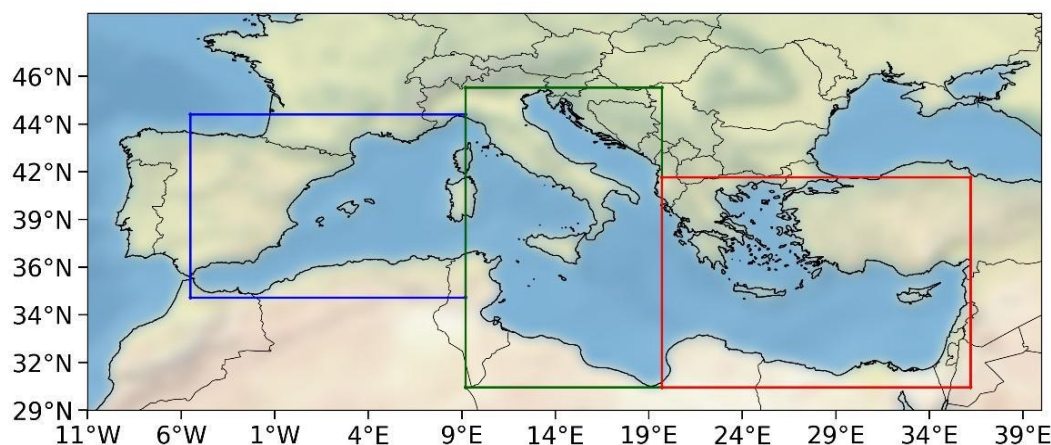
### 90 **2.1. Study area**

91 The Mediterranean basin exhibits distinct meteorological regimes across its subregions (Raveh-Rubin and Wernli, 2015  
92 Giuntoli et al., 2022). The West Mediterranean is often affected by convective activity associated with Atlantic weather  
93 systems and is characterized by a strong presence of mountains over land. The Central Mediterranean still experiences some  
94 deep convective activity, particularly over the sea. This is likely due to the warmer sea surface and the formation of mid-  
95 latitude cyclones. On the contrary, the East Mediterranean is dominated by low-level clouds due to the relatively stable and  
96 dry atmospheric conditions, especially in summer and autumn.

97 In this study, the geographical and temporal characterisation of the cloud statistics in the Mediterranean basin [29°N - 47°N  
98 and 11°W - 40°E], is carried out by partitioning the domain in three sub-regions: West [34.9°N - 44.5°N, 5.5°W - 9.2°E],



99 Central [30.2°N - 45.9°N, 9.2°E - 19.7°E] and East [30.2°N - 41.2°N, 19.7°E - 36.2°E], as shown in Figure 1. Within each  
100 region, pixels over land and pixels over ocean are examined separately.  
101



102  
103 **Figure 1: CloudSat data analysis above West (blue), Central (green) and East (red) Mediterranean.**  
104  
105

## 106 2.2. CloudSat dataset

107 CloudSat's payload, the Cloud Profiling Radar (CPR) (Stephens et al., 2007, 2008), was the first spacebased 94-GHz (W-  
108 band) radar and provided vertical cloud profiles on a global scale. CloudSat was launched on 28 April 2006, and CPR  
109 instrument operations began on 2 June 2006. From that time until the end of its operational life (i.e., 20 December 2023), the  
110 CPR acquired the first continuous global time series of vertical cloud structure and vertical profiles of cloud liquid water and  
111 ice content.

112 CloudSat flew in the A-Train constellation of satellites. It operated at 94 GHz and had a sensitivity of -30 dBZ. Radar  
113 reflectivities were sampled every 240 m and had an effective vertical resolution of approximately 500 m. Profile spacing was  
114 about 1 km along track, corresponding to a sampling volume of roughly 1.8 km in the along-track direction and 1.5 km across  
115 track. The A-Train constellation was in a Sun-synchronous orbit with equator crossings at approximately 01:30 and 13:30 local  
116 time.

117 In this study we use the following CloudSat dataset (both ascending and descending overpasses): 1) the 2B-GEOPROF (R05)  
118 cloud mask and radar reflectivity products (Mace et al., 2006; Marchand et al., 2008), 2) the 2B-CLDCLASS-LIDAR (R05)  
119 product is being used for retrieving the heights of the cloud tops and bases and the cloud types (Sassen, K. et al., 2008) and 3)  
120 the 2B-CWC-RO (CloudSat Radar-Only Cloud Water Content) (R05) product is being used to retrieve the liquid and ice cloud  
121 water path products (Austin et al., 2009). In general, the combination of the CPR and the CALIPSO lidar measurements offer

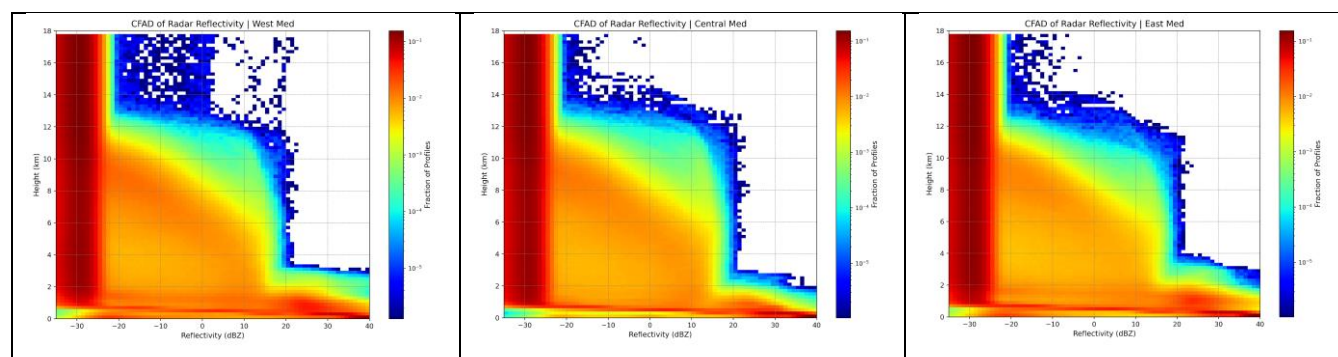


122 a more robust cloud type characterization (Sassen et al. 2008; Mace et al. 2009). Using the 'Navigation land sea flag' product,  
123 we present the different cloud properties over land and sea.

124 CloudSat algorithm classified clouds into eight categories: high-level clouds (Cirrus, Cirrostratus; Ci, Cs), altostratus (As),  
125 altocumulus (Ac), stratus (St), stratocumulus (Sc), cumulus (Cu), nimbostratus (Ns) and deep convective clouds (Cb). The  
126 classification is done based on different rules for hydrometeor vertical and horizontal scales, the maximum Ze measured by  
127 the CPR, indications of precipitation, and ancillary data including temperature T profiles and surface topography height (Sassen  
128 and Wang, 2007), as predicted by the ERA5 Reanalysis dataset of the European Centre for Medium-Range Weather Forecasts  
129 (ECMWF).

130 The frequency of occurrence of reflectivity values at different heights for all the clouds sampled in each domain is analysed to  
131 reveal regional differences in cloud properties, their vertical structure, and precipitation characteristics. Figure 2 shows the  
132 Contoured Frequency by Altitude Diagram (CFAD) of the total sample of radar reflectivity analysed in this study. The samples  
133 were binned by 1 dBZ at each level and 0.2 km height and were normalized by the total number of profiles at that altitude  
134 level. The mean cloud top (~14 km) is noted, as above that, reflectivity values are below the minimum detectable threshold.  
135 Moving downwards through the troposphere, the reflectivity quickly increases. In the upper troposphere (above 8 km), the  
136 reflectivity values for the three domains are generally low (-20 to -10 dBZ), indicating the presence of ice clouds or small  
137 cloud droplets. In the mid-troposphere (2 - 6 km), higher reflectivity values (~10 dBZ) are found, indicating regions of strong  
138 convection. An increase in reflectivity associated with melting is clearly visible between 1 to 2 km a.s.l. Below this, reflectivity  
139 increases rapidly towards the surface as the radar signal gets attenuated. The pattern is quite similar in the West and East  
140 Mediterranean, with only minor differences observed in the Central.

141



142

143 **Figure 2: A frequency by altitude diagram (CFAD), showing the frequency of occurrence of values of reflectivity at**  
144 **different heights for the total clouds sampled for West (a), Central (b) and East (c) Mediterranean. Data are binned at**  
145 **1 dBZ and 0.2 km height and then normalized by the total number of profiles in that height level. The mean cloud top**  
146 **(~14 km) is noted, and above that, reflectivity values are below the minimum detectable threshold.**

147



148 For the statistical analysis presented below, we used only CloudSat products with high confidence, defined as a  
149 CPR\_cloud\_mask value greater than or equal to 30.

150

### 151 **2.3 ERA5 Reanalysis dataset**

152 The ERA5 Reanalysis dataset, produced by the ECMWF/IFS, are available at a horizontal resolution of  $0.25^\circ \times 0.25^\circ$  with 137  
153 vertical levels (Vogelezang and Holtlag, 1996), offering a consistent representation of atmospheric conditions. ERA5 has  
154 been widely used for climate monitoring, weather analysis, and hydrometeorological studies (Wang et al., 2025). For the  
155 purpose of this study, we use hourly values of total column cloud and liquid water from ERA5 reanalyses data, accessed  
156 through ECMWF's MARS archive, and regridded through the request on a  $1^\circ \times 1^\circ$  spatial resolution. The data can be found  
157 in the CDS catalogue: Copernicus Climate Change Service, Climate Data Store, (2023): ERA5 hourly data on single levels  
158 from 1940 to present. Copernicus Climate Change Service (C3S) Climate Data Store (CDS). DOI: 10.24381/cds.adbb2d47  
159 (Accessed on 15-09-2025). The datasets are generated using a 4D-Var data assimilation system, that incorporates satellite and  
160 in situ observations into the ECMWF Integrated Forecasting System (IFS Cycle 41r2). Hersbach et al. (2020) provide detailed  
161 lists and timelines of observation types entering the ERA5 system. The total column liquid and ice water content cloud  
162 microphysical properties are used for comparison with CloudSat observations.

163

164

### 165 **2.4 EarthCARE dataset**

166 The Earth Clouds, Aerosols and Radiation Explorer (EarthCARE) mission (Illingworth et al., 2015) carries a CPR operating  
167 at 94 GHz (W-band), designed to provide vertically resolved observations of clouds and precipitation with enhanced sensitivity  
168 compared to CloudSat. More specifically, the EarthCARE CPR has higher sensitivity (5–6 dB more sensitive), better vertical  
169 sampling (100 versus 240 m), higher along-track resolution (500 versus 1100 m) and a smaller instantaneous field of view  
170 (IFOV, 800 versus 1400 m) and includes Doppler velocity measurements and improved detection in the lowest atmosphere  
171 levels (Mroz et al., 2023). This study utilizes the EarthCARE C-CLD product (Mroz et al., 2023) of BA baseline, with one  
172 year of measurements analyzed for the period from October 2024 to October 2025.

## 173 **3 Results**

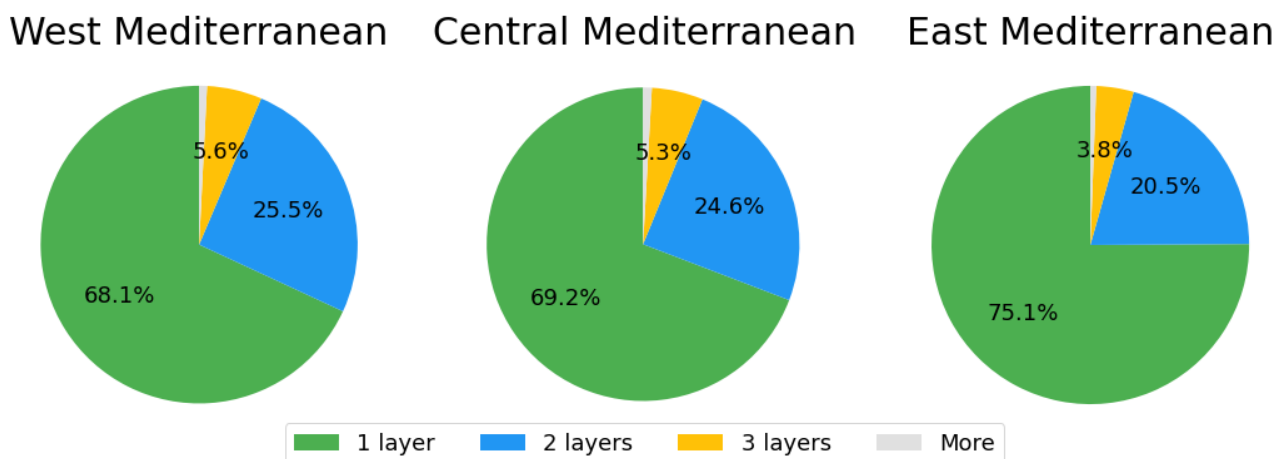
174 In this section, we present and analyze the frequency of occurrence of cloud types over the Mediterranean from CloudSat  
175 observations, focusing on both their vertical distribution and monthly variability. The analysis spans an 11-year period from  
176 January 2007 to December 2017. We further perform a statistical comparison between CloudSat observations and ERA5  
177 outputs, highlighting differences in spatial distribution and cloud water content. Additionally, first-year EarthCARE  
178 observations are presented and compared with ERA5, providing an assessment of the continuity of cloud patterns over the  
179 region.



180 **3.1. Cloud statistics**

181 The cloud layers distribution over the 3 regions are presented in Figure 3. In most of the cases, single-layer clouds are observed  
 182 (68-75% of the time), with 20-26% of the cases having double-layer clouds and less than 6% of the cases having three or more  
 183 cloud layers. A similar number of cloud layers is detected over the West and Central Mediterranean, with higher single-layer  
 184 clouds observed over the East Mediterranean.

185



186

187 **Figure 3. Cloud layers distribution for West (a), Central (b) and East (c) Mediterranean for the period 2007 – 2017.**

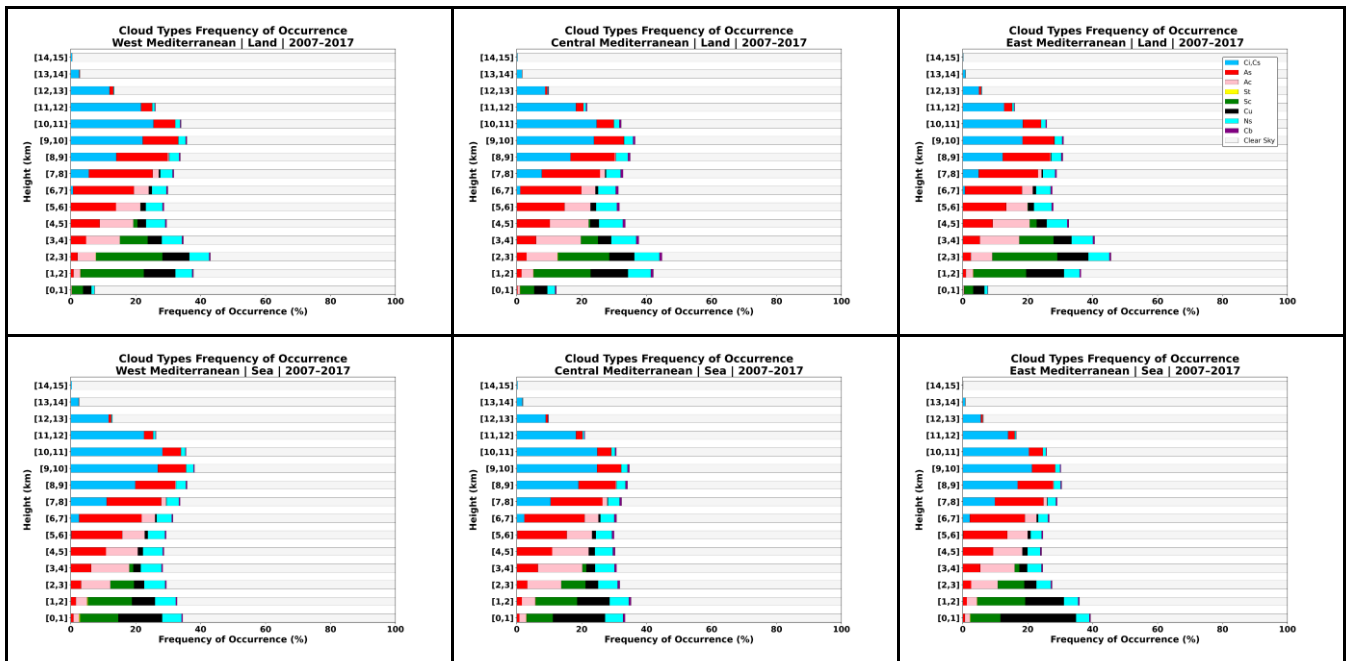
188 **3.2. Vertical frequency distribution of cloud types**

189 Figure 4 shows the vertical distribution of the different cloud types in the Mediterranean, separated over land (top panel) and  
 190 sea (lower panel), and over the West (left), Central (center), and East (right) Mediterranean, based on the decadal Cloudsat  
 191 dataset. All values are divided to the total profile number for each height.

192 From the statistics, it is evident that the occurrence of low clouds, i.e. below 1 km a.s.l., is almost four times more frequent  
 193 over sea than over land (32-39 % over sea, 8-12 % over land). In these altitudes, Cu is more frequently observed over sea (in  
 194 37 % of the cases), while Sc is more frequently observed over land (in 41 % of the cases). Cu are found with higher percentages  
 195 over the sea across all regions, with the East Mediterranean showing the highest overall occurrence. Over land, most clouds  
 196 occur between 1 and 4 km, where Sc clouds are more frequently observed (36 % of the cases), and where Ac and As begin to  
 197 form. Over sea, most clouds occur between 0-2 km. Stratocumulus are more frequently found over sea than over land across  
 198 all regions, with the West Mediterranean showing the highest overall occurrence. Stratus is the least frequent category across  
 199 all regions. As and Ac are the predominant cloud types at altitudes between 4 and 8 km, accounting for 24% of their joint  
 200 occurrence, while Cirrus (Ci) and Cirrostratus (Cs) are the dominant cloud types at altitudes of 9-11 km over land and 8-11  
 201 km over the ocean, accounting for 28% and 33% of occurrences, respectively. High-level clouds (Ci and Cs) show a  
 202 pronounced peak at 8-11 km across the entire Mediterranean region, with occurrence frequencies reaching up to 30%. Overall,



203 the presence of high-level clouds over the Mediterranean region is consistent with the global patterns reported by Li et al.  
 204 (2015). Thick, precipitating stratiform clouds (Ns) are observed mainly in altitudes between 1 to 9 km, at 8% to 19% of the  
 205 clouds, while thick, rain-bearing deep convective clouds (Cb) represent the second least frequent cloud category over the  
 206 Mediterranean. They occur more often over the Central Mediterranean (up to about 2%), both over land and sea.  
 207



208 **Figure 4: Vertical frequency distribution of cloud types, for West (left), Central (central) and East Mediterranean**  
 209 **(right), above Land (up) and Sea (down), during 2007 – 2017.**

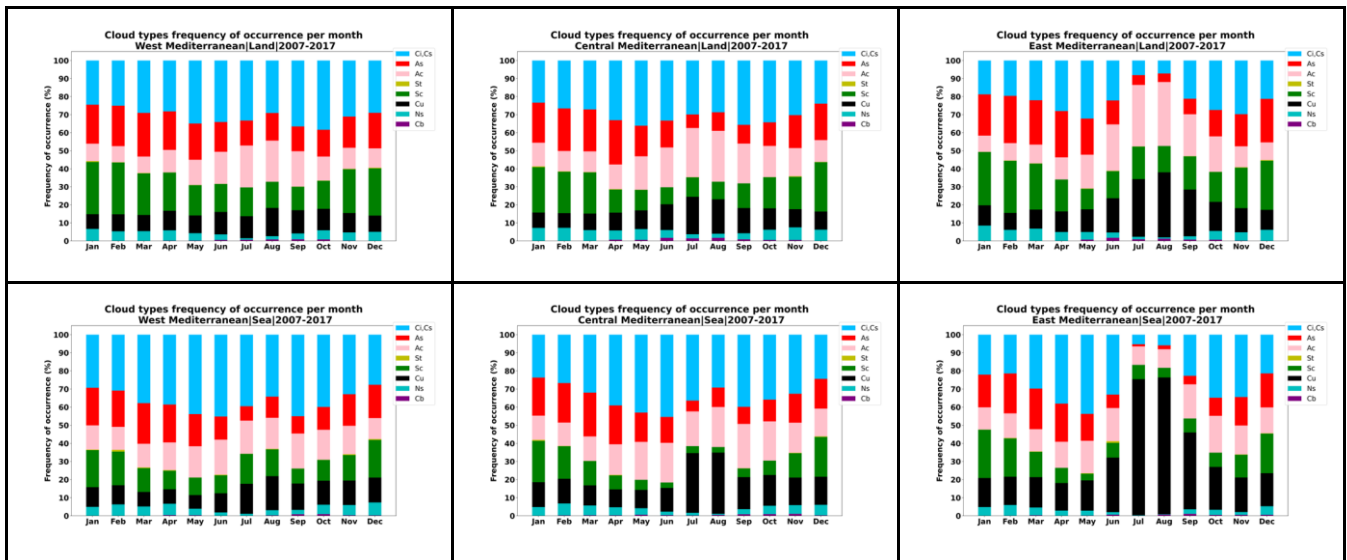
210  
 211 **3.3. Monthly frequency statistics of cloud types**

212 Figure 5 show the monthly frequency statistics of the different cloud types in the Mediterranean region, separated over land  
 213 (top panel) and sea (lower panel), over the West (left), Central (center), and East (right) Mediterranean, based on the decadal  
 214 Cloudsat dataset.

215 Stratocumulus (Sc) exhibits a seasonal pattern, with higher occurrences between January-March and October-December, due  
 216 to stable marine boundary layer processes. Cumulus (Cu) also exhibits a clear seasonality, with higher occurrences in July and  
 217 August over both land and sea across all three regions, due to the strong evaporation. Occurrences are particularly high in the  
 218 East Mediterranean, exceeding 30% over the land and 70% over sea. Altostratus (As) are less frequent in June, July, and  
 219 August, reaching their lowest percentages in July and August over the East Mediterranean. Altocumulus (Ac) exhibits a  
 220 different seasonal pattern, with peak occurrences in July and August over the East Mediterranean, over land. Nimbostratus  
 221 (Ns) shows a minimum between June and September, with near-zero occurrences over the sea in the East Mediterranean. High



222 clouds (Ci and Cs) exhibit peak frequencies between spring and early summer (April–June) and in autumn (September–  
223 October) over both land and sea. Ci and Cs lowest percentages are shown over July and August for the East Mediterranean.  
224 The rain-bearing deep convective clouds (Cb) peaks in summer and early autumn over land and for the Central and East  
225 Mediterranean.  
226



227 **Figure 5: Monthly frequency of occurrence of different cloud types, for West Med (left), Central Med (central) and**  
228 **East Med (left), above Land (up) and Sea (down), for the period 2007 – 2017.**

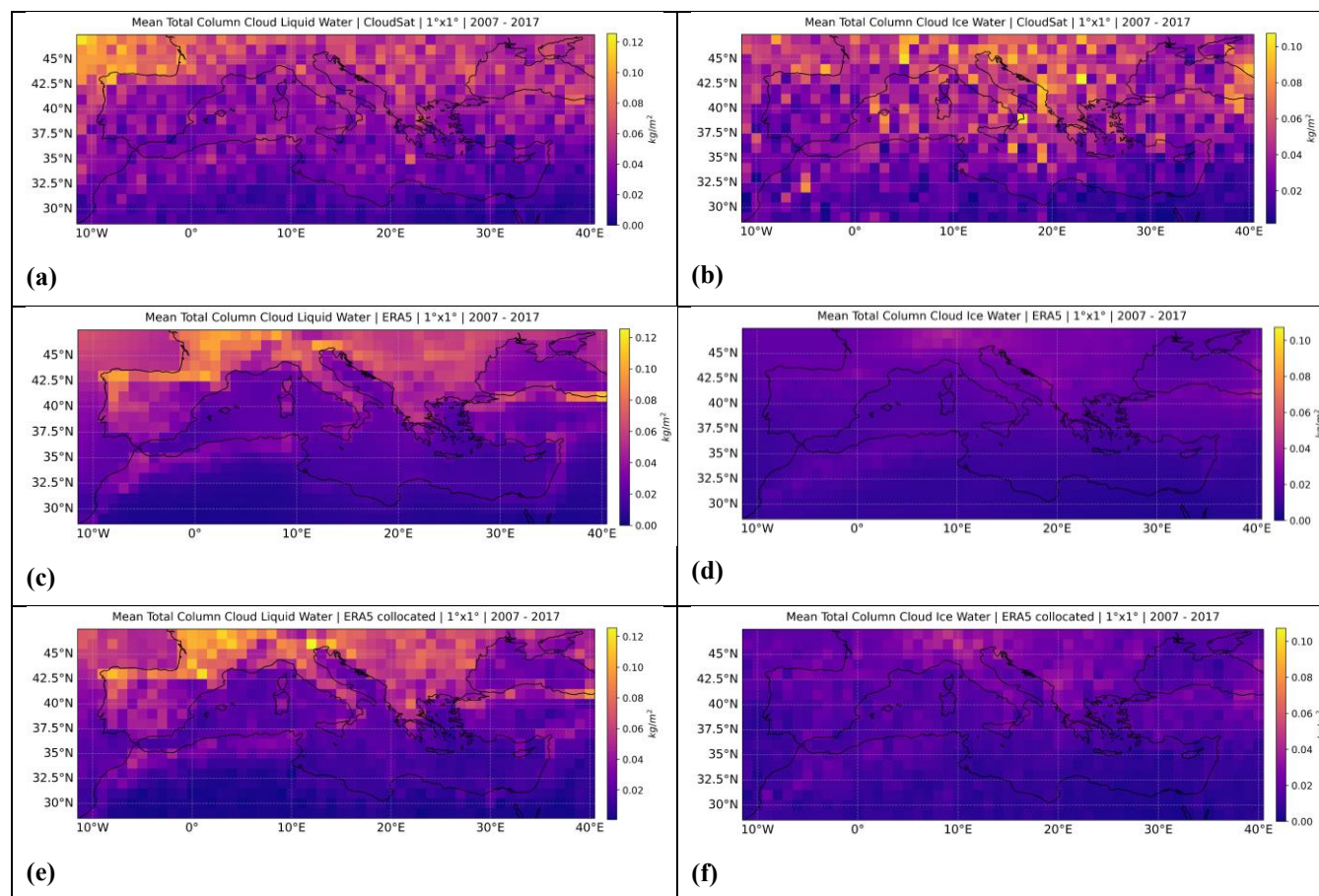
### 229 3.4. CloudSat observations vs ERA5 reanalysis: Spatiotemporal patterns

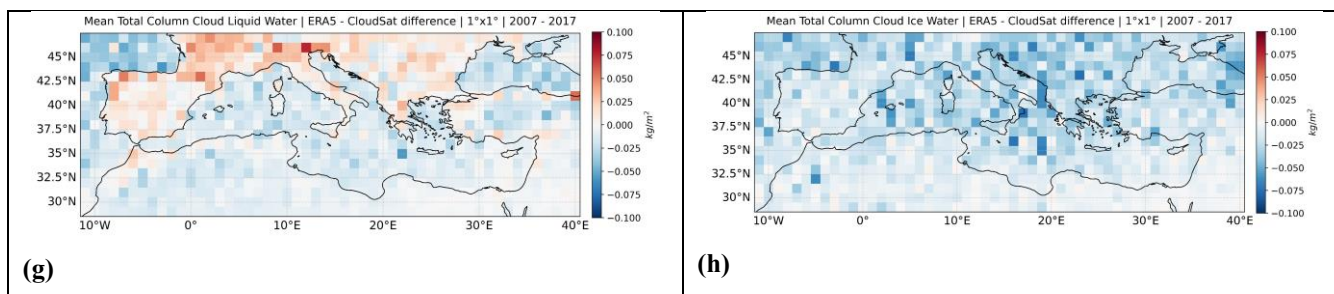
230 To compare CloudSat with ERA5 and evaluate the reanalysis of cloud predictions, we re-grid CloudSat dataset to the  $1^\circ \times 1^\circ$   
231 ERA5 spatial resolution over the Mediterranean domain. Figure 6 presents the mean total column of cloud liquid water (left  
232 panels) and cloud ice water (right panels) over the Mediterranean region, based on: (a, b) CloudSat observations, (c, d) ERA5  
233 reanalysis over the full dataset, (e, f) ERA5 reanalysis dataset limited to CloudSat spatiotemporal collocated data and (g, h)  
234 the corresponding differences between the ERA5 reanalysis dataset limited to CloudSat spatiotemporal collocated data and the  
235 CloudSat dataset.

236 To handle “problematic” grid cells that exhibited significantly lower counts (or zero counts) compared to their neighboring  
237 cells in the collocation process, we defined grid cells with few counts as “empty.” This threshold is considerably lower than  
238 the maximum count observed in some cells (approximately 300,000). The values of these low-count (empty) cells were then  
239 filled by replacing them with the average of the surrounding cells within a  $3 \times 3$  cell window, with the empty cell at the center.  
240 We notice that the ERA5 dataset limited to CloudSat collocated data, captures the same patterns observed in the ERA5 full  
241 dataset, although the values are observed noisier. This highlights the representativeness of the Cloudsat decadal climatology  
242 to represent the total column cloud liquid/ice water properties in the domain. We observe fair agreement in the regional



243 distribution of the total column water, with higher values above land, and especially mountainous areas (e.g. Seiki et al., 2019).  
244 Total column water is also increased (in both datasets) over northern latitudes.  
245 Overall, the largest deviation from the observations occurs over the sea, with the most significant discrepancy found in the  
246 total column of ice water content. More specifically, for total column liquid water, CloudSat values on the ERA5 grid range  
247 from  $0.038 \pm 0.015 \text{ kg/m}^2$  in the west to  $0.029 \pm 0.014 \text{ kg/m}^2$  in the east, while ERA5 collocated with CloudSat reports slightly  
248 lower values ( $0.034 \pm 0.020$  to  $0.022 \pm 0.014 \text{ kg/m}^2$ ). The resulting differences ( $-0.004$  to  $-0.008 \text{ kg/m}^2$ ) indicate good overall  
249 agreement between the two datasets. ERA5 systematically underestimates total column ice water compared to CloudSat. While  
250 CloudSat on the ERA5 grid reports  $0.037 \pm 0.015$ ,  $0.039 \pm 0.022$ , and  $0.029 \pm 0.016 \text{ kg/m}^2$  in the West, Central, and East  
251 Mediterranean, respectively, ERA5 collocated with CloudSat shows lower values ( $0.017 \pm 0.004$ ,  $0.015 \pm 0.008$ , and  $0.012 \pm$   
252  $0.006 \text{ kg/m}^2$ ). The resulting differences (ERA5 - CloudSat) range from  $-0.017 \pm 0.012$  to  $-0.024 \pm 0.017 \text{ kg/m}^2$ , with the largest  
253 underestimation over the central region. This could be attributed to the fact that many thin clouds are missed or underdetected  
254 in the ERA5 reanalysis (Yeo et al., 2022). To the best of our knowledge, this is the first study where the ERA5 cloud properties  
255 are evaluated against CloudSat measurements in the Mediterranean region.  
256



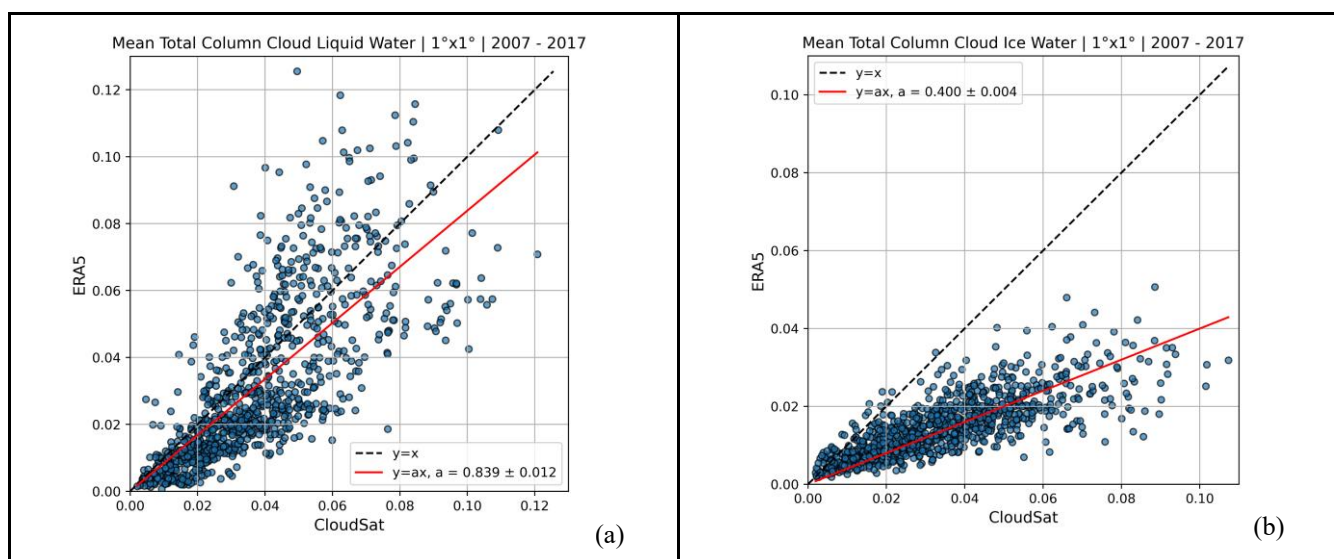


257

258 **Figure 6. The mean total column cloud liquid water (left panels) and cloud ice water (right panels) over the**  
 259 **Mediterranean region derived from CloudSat regridded to ERA5 observations (a, b), the corresponding averages from**  
 260 **the ERA5 reanalysis over the full dataset (c, d), the averages from ERA5 sampled only along CloudSat overpasses (e,**  
 261 **f) and the corresponding percentage differences between ERA5 and CloudSat (g, h).**

262

263 A direct comparison between column cloud liquid water (Figure 7a) and cloud ice water (Figure 7b) over the Mediterranean  
 264 region, between CloudSat regridded to ERA5 and ERA5 sampled along CloudSat observations, is presented below. Figure 7b  
 265 shows that ERA5 systematically detects lower column ice water values, as indicated by the higher slope. This underestimation  
 266 could be related to limitations in ERA5 cloud microphysics and vertical resolution (Yeo et al., 2022), which make it difficult  
 267 to capture thin cirrus layers. In contrast, column liquid water is represented more accurately (Figure 7a), suggesting that ERA5  
 268 performs better for low-level clouds than for high-altitude ice clouds.

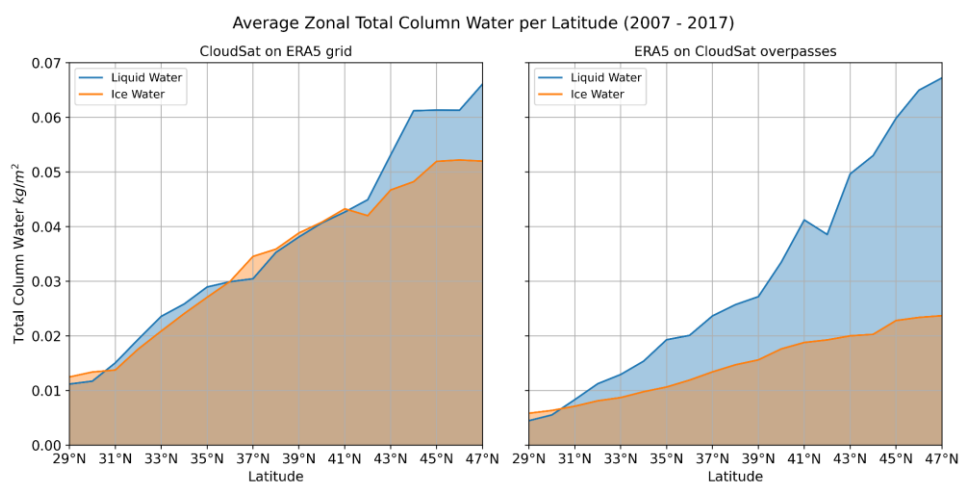


269 **Figure 7. Scatterplots of the column cloud liquid water (a) and cloud ice water (b) over the Mediterranean region, as**  
 270 **derived from CloudSat and ERA5 observations.**

271



272 Figure 8 presents the multiyear annual zonal-mean liquid water path (LWP; blue shading) and ice water path (IWP; brown shading) from (a) CloudSat and (b) the ERA5 dataset, for the overall Mediterranean domain. From Cloudsat, both LWP and  
 273 IWP exhibit similar patterns, with values increasing toward the northern latitudes. The LWP increases in magnitude as we  
 274 move towards higher latitudes, while the IWP remains almost stable. Higher IWP and LWP are found in CloudSat observations  
 275 than in the ERA5 reanalysis dataset (consistent to what is observed in Figure 6). The zonal-mean patterns derived from  
 276 CloudSat over the Mediterranean in this study are consistent with the global zonal-mean findings of Stephens et al. (2018),  
 277 although the absolute values differ from their reported statistics. More specifically, Stephens et al. (2018) showed higher annual  
 278 zonal-mean IWP values, which may be attributed to the different period used (they used the CloudSat 2C-ICE 2006 - 2010  
 279 dataset.  
 280 dataset.



281  
 282 **Figure 8: A multiyear annual zonal-mean liquid water path (blue shading) and ice water path (brown shading, from**  
 283 **CloudSat 2C-ICE dataset (a) and ERA5 dataset (b) for the period 2007 - 2017.**

284

### 285 3.5. CloudSat observations vs ERA5 reanalysis: Monthly patterns

286 Figure 9 shows the monthly variations of liquid water column (red and orange) and ice (green and blue) and based on CloudSat  
 287 observations and ERA5 reanalysis, respectively. The results show that ice and liquid water column values are generally higher  
 288 over land than over sea in both CloudSat and ERA5 datasets. The same monthly pattern is observed in both CloudSat  
 289 observations and ERA5 modeled fields, with lower values during the summer period (i.e., June, July, August). In general, the  
 290 East Mediterranean exhibits lower ice and liquid water column values, compared to the other two regions.

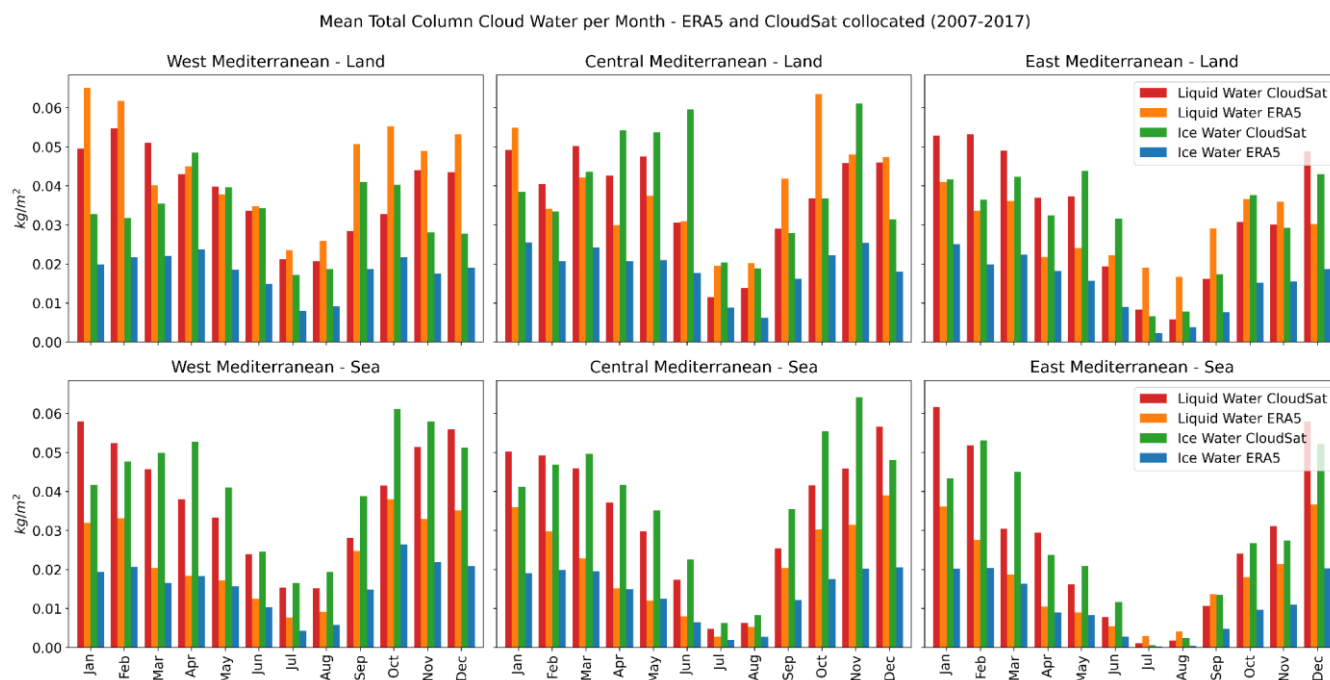
291 More specifically, the CloudSat observations show total column of liquid water values reaching up to 0.055 kg/m<sup>2</sup>, with a  
 292 significant decrease during the summer period (for all regions), with the minimum value recorded at 0.0057 kg/m<sup>2</sup> in August  
 293 over the Eastern Mediterranean. The ERA5 reanalysis shows a similar pattern, with total column liquid water values reaching  
 294 up to 0.065 kg/m<sup>2</sup> and exhibiting a comparable decrease during the summer period. Overall, the highest differences between



295 the Cloudsat and ERA5 collocated datasets is observed in the total column liquid water above the sea (up to 0.043 kg/m<sup>2</sup> over  
296 the Central Mediterranean).

297 Concerning total column of ice water, CloudSat observations reach up to 0.061 kg/m<sup>2</sup> over both land and sea in the Central  
298 Mediterranean in November. Lower values are consistently found in the ERA5 reanalysis over both land and sea across all  
299 months, remaining below 0.03 kg/m<sup>2</sup>.

300



301

302 **Figure 9: Monthly mean total column of the liquid and ice cloud water, for West Med (left), Central Med (central) and**  
303 **East Med (left), above Land (up) and Sea (down) for the period 2007 – 2017 based on CloudSat and ERA5 observations.**

304

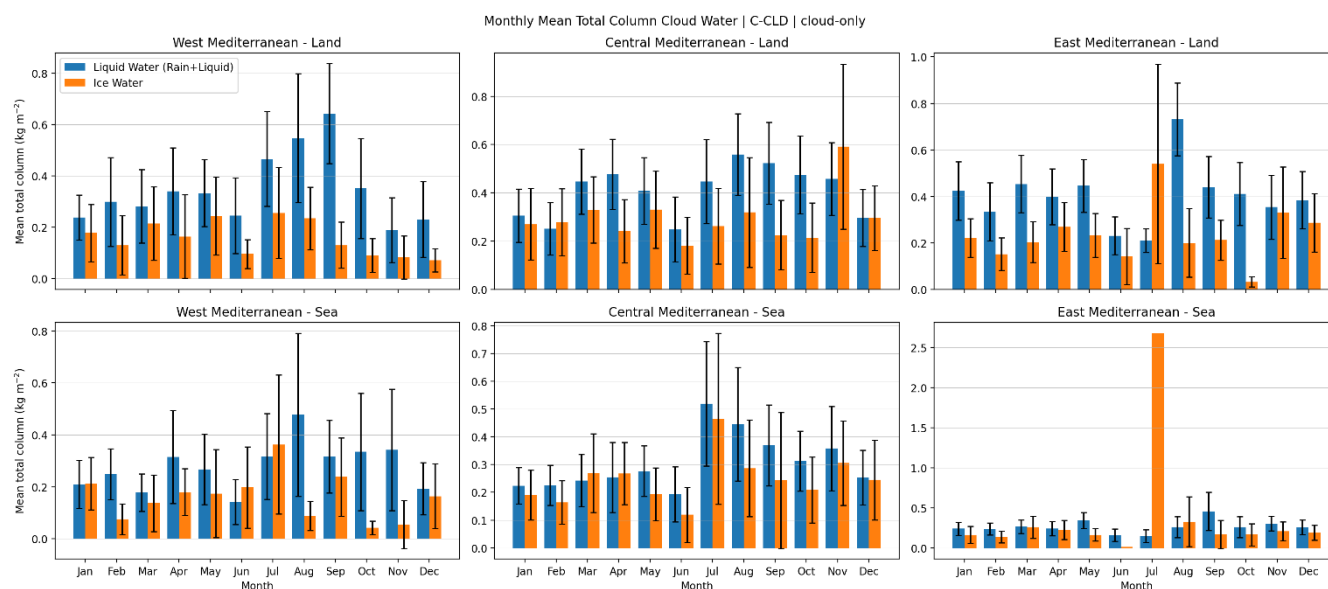
### 305 3.6. EarthCARE observations vs ERA5 reanalysis: Seasonal patterns

306 Figure 10 presents the liquid (blue) and ice (orange) water columns, along with the standard error, derived from the 1 year of  
307 EarthCARE observations (October 2024-October 2025). The results show that total column liquid water values are generally  
308 higher over land and exceed those retrieved from CloudSat (reaching above 0.7 kg/m<sup>2</sup> over the East Mediterranean). Similarly,  
309 total column ice water values are also higher than the corresponding CloudSat retrievals. The highest value is observed in July,  
310 reaching up to 2.7 kg/m<sup>2</sup> over the Eastern Mediterranean.

311 It is evident that one year of EarthCARE observations does not capture the monthly patterns observed in the decadal CloudSat  
312 statistics. However, it should be noted that additional measurements are needed before drawing robust conclusions.



313 In general, higher values of both liquid and ice water paths are observed in the one-year EarthCARE dataset. The absolute  
 314 differences compared to CloudSat can be attributed to the specific meteorological conditions prevailing during the analyzed  
 315 year, differences in sampling between the two spaceborne radars (e.g., observations at different local times), as well as  
 316 differences in the applied algorithms.  
 317

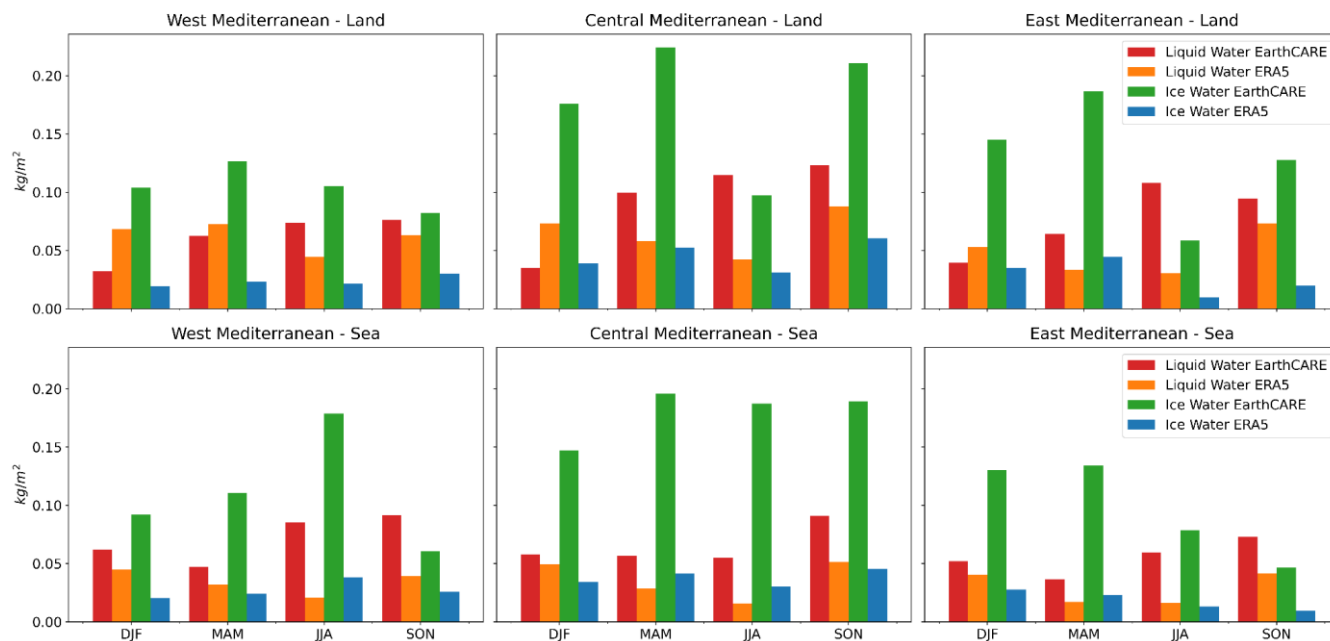


318  
 319 **Figure 10: Monthly mean total column of the liquid and ice cloud water, along with the standard error, for West Med**  
 320 **(left), Central Med (central) and East Med (left), above Land (up) and Sea (down) for the period 2024 – 2025 based on**  
 321 **EarthCARE observations.**

322  
 323 A comparison between the one year of EarthCARE observations and ERA5 is also performed. The results are presented on a  
 324 seasonal basis. Figure 11 illustrates the seasonal variations of liquid (red and orange) and ice water columns (green and blue)  
 325 based on the 1-year EarthCARE observations and ERA5 reanalysis, respectively. The seasonal patterns indicate that both ice  
 326 and liquid water column values are generally higher over land than over sea in the EarthCARE and ERA5 datasets, consistent  
 327 with CloudSat findings from the decadal dataset. However, larger values are found for the 1-year EarthCARE observations.  
 328 Larger differences are observed in both the ice and liquid water column values over the Central and East Mediterranean, with  
 329 the highest values generally occurring in the ice water column over land (up to 0.17 kg/m<sup>2</sup>). These large differences are  
 330 observed during all seasons. However, a critical factor affecting the reliability of the comparisons is the data availability,  
 331 therefore, additional measurements will be incorporated in future analyses.  
 332



Mean Total Column Cloud Water per Season - ERA5 and EarthCARE collocated October 2024 - October 2025



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**Figure 11: Seasonal mean total column of the liquid and ice cloud water, for West Med (left), Central Med (central) and East Med (left), above Land (up) and Sea (down) for the period October 2024 – October 2025 based on EarthCARE and ERA5 observations.**

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#### 4 Conclusions

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In this study, we examine vertical distributions and patterns of cloud types and properties above the Mediterranean Sea, using a decadal synergistic CloudSat/CALIPSO vertical dataset, and we investigate the capability of ERA5 reanalysis dataset to accurately capture the observed mean IWC and LWC values above the area. In addition, one year of EarthCARE observations is discussed in terms of consistency with the identified monthly patterns. The key findings of this work are summarized below:

- The occurrence of low clouds, with cloud bases below 1 km above the ground, is almost four times more frequent over sea areas compared to land. In this layer, Cu and Sc dominate over sea and land, respectively. Over land, most cloud bases occur between 1 and 3 km, where mid-level clouds (Ac and As) begin to form. They are found with higher percentages over the sea than over land across all regions, with the Central Mediterranean showing the highest overall occurrence. High clouds are common between 6 and 14 km, with a peak at 10-11 km over the entire Med region, with higher percentages over the West and Central Mediterranean (up to 30%). Cb is the less prominent category and more frequent over the central Med (with percentages up to 2%). Over land, deep convective clouds are less frequent but still present, particularly in the West Mediterranean.



- 350 • In the West and Central Mediterranean, mid-level clouds (particularly As) are significant in winter and autumn and  
351 more frequent over the sea, while low-level clouds are relatively less frequent over land compared to the sea, with  
352 peaks in autumn (i.e., October - November) and early winter. Sc dominates above the West Mediterranean during  
353 autumn (i.e., September - November) and winter (i.e., December - February), due to stable marine boundary layer  
354 processes. The frequency of occurrence of the high clouds above the East Med is low during the summer period and  
355 As are less than Ac.
- 356 • The comparison of the monthly variations of ice and water columns using CloudSat observations and ERA5 reanalysis  
357 show a good agreement in the regional distribution of total column water, with higher values observed over land,  
358 particularly in mountainous regions. Total column water is also increased over northern latitudes and mountainous  
359 areas. The resulting differences ( $-0.004$  to  $-0.008$  kg/m<sup>2</sup>) in the liquid water column indicate good overall agreement  
360 between the two datasets (ERA5 – CloudSat). In contrast, the differences in the ice water column range from  $-0.017$   
361  $\pm 0.012$  to  $-0.024 \pm 0.017$  kg/m<sup>2</sup>, with the largest underestimation occurring over the Central Mediterranean. A  
362 number of thin clouds are missing or underdetected in the ERA5 reanalysis dataset.
- 363 • The one year of EarthCARE observations of the monthly variations of ice and water columns based on the C-CLD  
364 product is also presented. The higher values observed compared to the decadal CloudSat statistics, may be attributed  
365 to the specific meteorological conditions prevailing during the analyzed year, differences in the sampling  
366 characteristics of the two space-borne radars (including observations at different local times), or discrepancies arising  
367 from the applied retrieval algorithms. Future analyses will incorporate additional EarthCARE measurements to  
368 evaluate the consistency of liquid and ice water distribution patterns across the Mediterranean region.

369 The findings of this study can be used to assess cloud microphysical parameterizations in numerical weather prediction  
370 (NWP) models, contributing to enhanced forecasts of water and ice. Additionally, the derived statistics on clouds' high-  
371 resolution vertical distributions and thermodynamic phase can be combined with the Cloudnet cloud products and lidar-  
372 retrieved aerosol properties to study aerosol-cloud interactions in the understudied region of the Mediterranean region. In  
373 our future research, we will utilize the full dataset of ESA's Earth Cloud Aerosol and Radiation Explorer (EarthCARE)  
374 CPR measurements to derive cloud statistics, aiming to harmonize and bridge past, current and upcoming missions, and  
375 to produce Climate Data Records on cloud properties for the Mediterranean region.

#### 377 **Data availability.**

378 The Cloudsat dataset is available at <https://cloudsat.atmos.colostate.edu/>

379 ERA5 reanalysis data, accessed from ECMWF's MARS archive, © European Centre for Medium-Range Weather Forecasts  
380 (ECMWF), 2025. Licensed under CC-BY-4.0. Dataset available through: Copernicus Climate Change Service, Climate Data  
381 Store, (2023): ERA5 hourly data on single levels from 1940 to present. Copernicus Climate Change Service (C3S) Climate  
382 Data Store (CDS). DOI: [10.24381/cds.adbb2d47](https://doi.org/10.24381/cds.adbb2d47) (Accessed on 15-09-2025).




383 The EarthCARE Level-2 products (C-FMR and C-CD) used in this study are publicly accessible from the ESA Earth Online  
384 gateway at <https://earth.esa.int/eogateway/catalog/earthcare-esa-l2-products> (last access: 7 November 2025) and via the ESA  
385 MAAP where the processing is also conducted. (<https://portal.maap.eo.esa.int/>)

386 **Author contributions.** All authors have contributed to this work. KAV drafted the manuscript and supervised the data  
387 processing. EM and AB conceptualized the study. IK and VS processed the CloudSat dataset, while DK and IT handled the  
388 EarthCARE dataset. EM, PK, EG, AB and VA reviewed the manuscript.

389 **Competing interests.** The authors declare that they have no conflict of interest.

390

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## 396 References

397 Austin, R.T., A.J. Heymsfield, and G. L. Stephens, Retrieval of ice cloud microphysical parameters using the CloudSat  
398 millimeter-wave radar and temperature, *J. Geophys. Res.*, 114, D00A23, doi:10.1029/2008JD010049, 2009.

399

400 Barrett, Andrew I., Robin J. Hogan, and Richard M. Forbes., Why are mixed-phase altocumulus clouds poorly predicted by  
401 large-scale models? Part 1. Physical processes." *Journal of Geophysical Research: Atmospheres* 122.18 (2017): 9903-9926.,  
402 2017.

403

404 Battaglia, A., Kollias, P., Dhillon, R., Roy, R., Tanelli, S., Lamer, K., et al., Spaceborne cloud and precipitation radars: Status,  
405 challenges, and ways forward. *Reviews of Geophysics*, 58, e2019RG000686. <https://doi.org/10.1029/2019RG000686>, 2020.

406

407 Copernicus Climate Change Service, Climate Data Store, (2024): ERA5 post-processed daily-statistics on single levels from  
408 1940 to present. Copernicus Climate Change Service (C3S) Climate Data Store (CDS), DOI: 10.24381/cds.4991cf48  
409 (Accessed on 23-12-2025)

410



- 411 Copernicus Climate Change Service, Climate Data Store, (2023): ERA5 hourly data on single levels from 1940 to present.  
412 Copernicus Climate Change Service (C3S) Climate Data Store (CDS). DOI: 10.24381/cds.adbb2d47 (Accessed on 23-12-  
413 2025)
- 414
- 415 Ceccaldi, M., Delanoë, J., Hogan, R. J., Pounder, N. L., Protat, A., From CloudSatCALIPSO to EarthCare: Evolution of the  
416 DARDAR cloud classification and its comparison to airborne radar-lidar observations. *Journal of Geophysical Research:*  
417 *Atmospheres*, 2013, 118 (14), pp.7962- 7981. [ff10.1029/jgrd.50579ff](https://doi.org/10.1029/2013JD019599), 2013.
- 418
- 419 Cesana, G.V., Ackerman, A.S., Fridlind, A.M. et al. Observational constraint on a feedback from supercooled clouds reduces  
420 projected warming uncertainty. *Commun Earth Environ* 5, 181 (2024). <https://doi.org/10.1038/s43247-024-01339-1>
- 421
- 422 Chaboureau, J.-P., and C. Claud (2006), Satellite-based climatology of Mediterranean cloud systems and their association with  
423 large-scale circulation, *J. Geophys. Res.*, 111, D01102, doi:[10.1029/2005JD006460](https://doi.org/10.1029/2005JD006460), 2006.
- 424
- 425 Delanoë, J., & Hogan, R. J. A variational scheme for retrieving ice cloud properties from combined radar, lidar, and infrared  
426 radiometer. *Journal of Geophysical Research*, 113, D07204. <https://doi.org/10.1029/2007jd009000>, 2008.
- 427
- 428 Delanoë, J., and R. J. Hogan (2010), Combined CloudSat-CALIPSO-MODIS retrievals of the properties of ice clouds, *J.*  
429 *Geophys. Res.*, 115, D00H29, doi:[10.1029/2009JD012346](https://doi.org/10.1029/2009JD012346).
- 430
- 431 Engelmann, R., Kanitz, T., Baars, H., Heese, B., Althausen, D., Skupin, A., Wandinger, U., Komppula, M., Stachlewska, I.S.,  
432 Amiridis, V., et al. The automated multiwavelength Raman polarization and water-vapor lidar PollyXT: The neXT  
433 generation. *Atmos. Meas. Tech.* 9, 1767–1784, 2016.
- 434
- 435 Enriquez-Alonso, A., Sanchez-Lorenzo, A., Calbó, J. et al. Cloud cover climatologies in the Mediterranean obtained from  
436 satellites, surface observations, reanalyses, and CMIP5 simulations: validation and future scenarios. *Clim Dyn* 47, 249–269,  
437 <https://doi.org/10.1007/s00382-015-2834-4>, 2016.
- 438
- 439 Feuillard, N., Toledo Bittner, F., Pfitzenmaier, L., Ribaud, J.-F., Delanoë, J., Haeffelin, M., and Dupont, J.-C.: Validation of  
440 EarthCARE CPR reflectivity using the ACTRIS cloud radar network, *EGUsphere* [preprint],  
441 <https://doi.org/10.5194/egusphere-2026-925>, 2026.
- 442
- 443 Forbes, R., Geer, A., Lonitz, K., Ahlgrimm, M., Reducing systematic errors in cold-air outbreaks, *Meteorology section of*  
444 *ECMWF Newsletter No. 146 – Winter 2015/16*, pp. 17–22, 2016.



- 445  
446 Giuntoli, I., Fabiano, F. & Corti, S. Seasonal predictability of Mediterranean weather regimes in the Copernicus C3S  
447 systems. *Clim Dyn* **58**, 2131–2147, <https://doi.org/10.1007/s00382-021-05681-4>, 2022.
- 448  
449 Hersbach H, Bell B, Berrisford P, et al. The ERA5 global reanalysis. *Q J R Meteorol Soc.* 2020; 146: 1999–2049.  
450 <https://doi.org/10.1002/qj.3803>.
- 451  
452 Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J., Nicolas, J., Peubey, C., Radu, R., Rozum,  
453 I., Schepers, D., Simmons, A., Soci, C., Dee, D., Thépaut, J.-N. (2023): ERA5 hourly data on single levels from 1940 to present.  
454 Copernicus Climate Change Service (C3S) Climate Data Store (CDS), DOI: 10.24381/cds.adbb2d47 (Accessed on 15-09-  
455 2025)
- 456  
457 Illingworth, A. J., Barker, H. W., Beljaars, A., Ceccaldi, M., Chepfer, H., Clerbaux, N., Cole, J., Delanoë, J., Domenech, C.,  
458 Donovan, D. P., Fukuda, S., Hirakata, M., Hogan, R. J., Huenerbein, A., Kollias, P., Kubota, T., Nakajima, T., Nakajima, T.  
459 Y., Nishizawa, T., Ohno, Y., Okamoto, H., Oki, R., Sato, K., Satoh, M., Shephard, M. W., Velázquez-Blázquez, A.,  
460 Wandinger, U., Wehr, T., and van Zadelhoff, G.-J.: The EarthCARE Satellite: The Next Step Forward in Global Measurements  
461 of Clouds, Aerosols, Precipitation, and Radiation, *B. Am. Meteorol. Soc.*, 96, 1311–1332, [https://doi.org/10.1175/BAMS-D-](https://doi.org/10.1175/BAMS-D-12-00227.1)  
462 [12-00227.1](https://doi.org/10.1175/BAMS-D-12-00227.1), 2015.
- 463 Ioannidis, E., Lolis, C.J., Bartzokas, A., Seasonal Variability of Total Cloud Cover in the Mediterranean Region for the Period  
464 1948–2014. In: Karacostas, T., Bais, A., Nastos, P. (eds) *Perspectives on Atmospheric Sciences*. Springer Atmospheric  
465 Sciences. Springer, Cham. [https://doi.org/10.1007/978-3-319-35095-0\\_89](https://doi.org/10.1007/978-3-319-35095-0_89), 2017.
- 466  
467 IPCC 2021: Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment  
468 Report of the Intergovernmental Panel on Climate Change, edited by: Masson-Delmotte, V., Zhai, P., Pirani, A., Connors, S.  
469 L., Péan, C., Berger, S., Caud, N., Chen, Y., Goldfarb, L., Gomis, M. I., Huang, M., Leitzell, K., Lonnoy, E., Matthews, J. B.  
470 R., Maycock, T. K., Waterfield, T., Yelekçi, O., Yu, R., and Zhou, B.: Cambridge University Press, Cambridge, United  
471 Kingdom and New York, NY, USA, <https://doi.org/10.1017/9781009157896>, in press, 2021.
- 472  
473 Komurcu, M., T. Storelvmo, I. Tan, U. Lohmann, Y. Yun, J. E. Penner, Y. Wang, X. Liu, and T. Takemura, Intercomparison  
474 of the cloud water phase among global climate models, *J. Geophys. Res. Atmos.*, 119, 3372–3400,  
475 doi:10.1002/2013JD021119, 2014.
- 476  
477 Levizzani V., Pinelli F., Pasqui M., Melani S., Laing AG., Carbone RE, A 10-year climatology of warm-season cloud patterns  
478 over Europe and the Mediterranean from Meteosat IR observations, *Atmospheric Research* 97.4 (2010): 555-576, 2010.



- 479
- 480 Li, J., Huang, J., Stamnes, K., Wang, T., Lv, Q., and Jin, H.: A global survey of cloud overlap based on CALIPSO and CloudSat  
481 measurements, *Atmos. Chem. Phys.*, 15, 519–536, <https://doi.org/10.5194/acp-15-519-2015>, 2015.
- 482
- 483 Li, J.-L. F., Xu, K.-M., Tsai, Y.-C., Lee, W.-L., Jiang, J. H., Yu, J.-Y., et al. (2023). Evaluation of radiatively active frozen  
484 hydrometeors mass in CMIP6 global climate models using CloudSat-CALIPSO observations. *Journal of Geophysical*  
485 *Research: Atmospheres*, 128, e2023JD039200. <https://doi.org/10.1029/2023JD039200>
- 486
- 487 Mace, G., Marchand, R., Zhang, Q., and Stephens, G. Global hydro-meteor occurrence as observed by CloudSat: Initial  
488 observations from summer 2006, *Geophys. Res. Lett.*, 2007, 34, L09808, doi:10.1029/2006GL029017, 2007.
- 489
- 490 Mace, G. G., Zhang, Y., Marchand, R., Stephens, G., & Ackerman, T., A description of hydrometeor layer occurrence statistics  
491 derived from the first year of merged CloudSat and CALIPSO data. *Journal of Geophysical Research: Atmospheres*, 114(D8),  
492 D00A26. <https://doi.org/10.1029/2007JD009755>, 2009.
- 493
- 494 Madonna, F., Amodeo, A., Boselli, A., Cornacchia, C., Cuomo, V., D’Amico, G., Giunta, A., Mona, L., Pappalardo, G. CIAO:  
495 The CNR-IMAA advanced observatory for atmospheric research. *Atmos. Meas. Tech.* 4, 1191–1208, 2011.
- 496
- 497 Marchand, R., G Mace, G, Ackerman, T., and G. Stephens, Hydro-meteor detection using CloudSat an Earth-orbiting 94-GHz  
498 cloud radar, *J. Atmos. Oceanic Technol.*, 25(4), 519–533, 2008.
- 499
- 500 Manninen, A.J., O’Connor, E.J., Vakkari, V., Petäjä, T. A generalised background correction algorithm for a Halo Doppler  
501 lidar and its application to data from Finland. *Atmos. Meas. Tech.* 9, 817–827, 2016.
- 502
- 503 Marinou, E., Voudouri, K.A., Tsikoudi, I., Drakaki, E., Tsekeri, A., Rosoldi, M., Ene, D., Baars, H., O’Connor, E., Amiridis,  
504 V., Meleti, C. Geometrical and Microphysical Properties of Clouds Formed in the Presence of Dust above the Eastern  
505 Mediterranean. *Remote Sens.* 13, 5001. <https://doi.org/10.3390/rs13245001>, 2021.
- 506
- 507 McGraw, Z., Storelvmo, T., Polvani, L. M., Hofer, S., Shaw, J. K., & Gettelman, A. (2023). On the links between ice  
508 nucleation, cloud phase, and climate sensitivity in CESM2. *Geophysical Research Letters*, 50, e2023GL105053.  
509 <https://doi.org/10.1029/2023GL105053>
- 510



511 Miglietta MM., Moscatello A., Conte D., Mannarini G., Lacorata G., Rotunno R., Numerical analysis of a Mediterranean  
512 ‘hurricane’ over south-eastern Italy: Sensitivity experiments to sea surface temperature, *Atmospheric Research*, Volume 101,  
513 Issues 1–2, 2011, Pages 412-426, ISSN 0169-8095, <https://doi.org/10.1016/j.atmosres.2011.04.006>, 2011.

514

515 Mroz, K., Treserras, B. P., Battaglia, A., Kollias, P., Tatarevic, A., and Tridon, F.: Cloud and precipitation microphysical  
516 retrievals from the EarthCARE Cloud Profiling Radar: the C-CLD product, *Atmos. Meas. Tech.*, 16, 2865–2888,  
517 <https://doi.org/10.5194/amt-16-2865-2023>, 2023.

518

519 Nabat, P., Somot, S., Cassou, C., Mallet, M., Michou, M., Bouniol, D., Decharme, B., Drugé, T., Roehrig, R., and Saint-  
520 Martin, D.: Modulation of radiative aerosols effects by atmospheric circulation over the Euro-Mediterranean region, *Atmos.*  
521 *Chem. Phys.*, 20, 8315–8349, <https://doi.org/10.5194/acp-20-8315-2020>, 2020.

522

523 Randall, D. A., Wood, R. A., Bony, S., Colman, R., Fichet, T., Fyfe, J., Kattsov, V., Pitman, A., Shukla, J., Srinivasan, J.,  
524 Stouffer, R. J., Sumi, A., & Taylor, K. E. (2007). *Climate Change 2007: The Physical Science Basis. Contribution of Working*  
525 *Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change* (S. Solomon et al., Eds.).  
526 Cambridge University Press, 2007.

527

528 Raveh-Rubin, S. and Wernli, H., Large-scale wind and precipitation extremes in the Mediterranean: a climatological analysis  
529 for 1979–2012. *Q.J.R. Meteorol. Soc.*, 141: 2404-2417. <https://doi.org/10.1002/qj.2531>, 2015.

530

531 Sanchez-Lorenzo, A., Enriquez-Alonso, A., Calbó, J. *et al.* Fewer clouds in the Mediterranean: consistency of observations  
532 and climate simulations. *Sci Rep* 7, 41475, <https://doi.org/10.1038/srep41475>, 2017.

533

534 Sassen, K., Wang, Z., and Liu, D., Global distribution of cirrus clouds from CloudSat/Cloud-Aerosol Lidar and Infrared  
535 Pathfinder Satellite Observations (CALIPSO) measurements. *Journal of Geophysical Research: Atmospheres*, 113(D8),  
536 D00A12. <https://doi.org/10.1029/2008JD009972>, 2008.

537

538 Sassen, K., and Wang, Z.: Classifying clouds around the globe with the CloudSat radar: 1-year of results, *Geophys. Res.*  
539 *Lett.*, 35, L04805, doi:[10.1029/2007GL032591](https://doi.org/10.1029/2007GL032591), 2008.

540

541 Sassen K., and Wang Z.: Level 2 cloud scenario classification product process description and interface control document,  
542 version 5.0, 2007, available at <http://cloudsat.cira.colostate.edu/data/CDlist.php?go=list&path=/2B-CLDCLASS>.

543



- 544 Seiki, T., Kodama, C., Satoh, M., Hagihara, Y., & Okamoto, H., Characteristics of ice clouds over mountain regions detected  
545 by CALIPSO and CloudSat satellite observations. *Journal of Geophysical Research: Atmospheres*, 124, 10858–  
546 10877. <https://doi.org/10.1029/2019JD030519>, 2019.
- 547
- 548 Stephens, G.L., and N.B. Wood, Properties of Tropical Convection Observed by Millimeter-Wave Radar Systems, *Mon. Wea.*  
549 *Rev.*, 135, 821-842, doi:10.1175/MWR3321.1, 2007.
- 550
- 551 Stephens, G.L., D.G. Vane, S. Tanelli, E. Im, S. Durden, M. Rokey, D. Reinke, P. Partain, G.G. Mace, R. Austin, T. L'Ecuyer,  
552 J. Haynes, M. Lebsock, K. Suzuki, D. Waliser, D. Wu, J. Kay, A. Gettelman, Z. Wang, and R. Marchand, CloudSat mission:  
553 Performance and early science after the first year of operation, *J. Geophys. Res.*, 113, D00A18, doi:10.1029/2008JD009982,  
554 2008.
- 555
- 556 Stephens, G., D. Winker, J. Pelon, C. Trepte, D. Vane, C. Yuhas, T. L'Ecuyer, and M. Lebsock, 2018: CloudSat and CALIPSO  
557 within the A-Train: Ten Years of Actively Observing the Earth System. *Bull. Amer. Meteor. Soc.*, 99, 569–581,  
558 <https://doi.org/10.1175/BAMS-D-16-0324.1>, 2018.
- 559
- 560 The PRE-TECT Experimental Campaign. Available online: <http://PRE-TECT.space.noa.gr/> (accessed on 25 November 2025).
- 561
- 562 Vogelezang, D. H. P., and A. A. M. Holtslag, 1996: Evaluation and model impacts of alternative boundary-layer height  
563 formulation. *Bound.-Layer Meteor.*, 81, 245–269, doi:10.1007/BF02430331.
- 564
- 565 Voudouri, K. A.; Giannakaki, E.; Komppula, M.; and Balis, D. Variability in cirrus cloud properties using a PollyXT Raman  
566 lidar over high and tropical latitudes, *Atmos. Chem. Phys.*, 20, 4427–4444, <https://doi.org/10.5194/acp-20-4427-2020>, 2020.
- 567
- 568 Waliser, Duane E., et al. "Cloud ice: A climate model challenge with signs and expectations of progress." *Journal of*  
569 *Geophysical Research: Atmospheres* 114.D8, 2009.
- 570
- 571 Wang, Z., Bugliaro, L., Gierens, K., Hegglin, M. I., Rohs, S., Petzold, A., Kaufmann, S., and Voigt, C.: Machine learning for  
572 improvement of upper-tropospheric relative humidity in ERA5 weather model data, *Atmos. Chem. Phys.*, 25, 2845–2861,  
573 <https://doi.org/10.5194/acp-25-2845-2025>, 2025.
- 574
- 575 Wehr, T., Kubota, T., Tzeremes, G., Wallace, K., Nakatsuka, H., Ohno, Y., Koopman, R., Rusli, S., Kikuchi, M., Eisinger, M.,  
576 Tanaka, T., Taga, M., Deghaye, P., Tomita, E., and Bernaerts, D.: The EarthCARE mission – science and system overview,  
577 *Atmos. Meas. Tech.*, 16, 3581–3608, <https://doi.org/10.5194/amt-16-3581-2023>, 2023.



578

579 Yeo, H., Kim, M.-H., Son, S.-W., Jeong, J.-H., Yoon, J.-H., Kim, B.-M., and Kim, S.-W.: Arctic cloud properties and  
580 associated radiative effects in the three newer reanalysis datasets (ERA5, MERRA-2, JRA-55): Discrepancies and possible  
581 causes, Atmospheric Research, Volume 270, 2022, 106080, ISSN 0169-8095, <https://doi.org/10.1016/j.atmosres.2022.106080>,  
582 2022.