



1 **Assessing Air Pollution Drivers in Asia Through Multi-Species Data**

2 **Assimilation During NASA's ASIA-AQ Campaign**

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32



33 **Abstract.** Asia accounts for approximately half of global anthropogenic emissions of major pollutants, yet emission
34 inventories remain uncertain and ground-based monitoring is sparse across much of the region. To address these challenges,
35 this study applies a multi-species satellite data assimilation framework to estimate emissions and concentrations of key
36 chemical species during the NASA's ASIA-AQ campaign. The assimilation improves agreement with airborne observations
37 for O₃, NO_x, CO, and CH₂O, with the largest gains for CO (correlation increasing from 0.63–0.64 to 0.77) and CH₂O (biases
38 reduced by 41–70%). Domain-wide, the optimized emissions show increases of 15% for NO_x and 9% for CO, and a 52%
39 reduction in isoprene. Comparisons with multiple emission inventories reveal large discrepancies, with normalized standard
40 deviation ranging from 11% for NO_x in mainland China to 68% for CO in Taiwan. Over Thailand, the assimilation increases
41 fire emissions from 3.6 Tg (GFASv1.2) to 8.3 Tg, while FINNv2.7 produces estimates roughly twice as high, highlighting
42 persistent divergence among fire emission estimates. Source-receptor analysis reveals strong meteorological control on
43 transboundary pollutant: long-range transport contributes up to 62% of surface O₃ in Manila during strong monsoon conditions,
44 whereas Seoul exhibits local NO_x-saturated chemistry. During a strong transport episode, Indian emissions account for 72–
45 78% of the free tropospheric O₃ response over Taipei and Chiang Mai, and 24% over Seoul, highlighting an overlooked
46 transport pathway affecting Asian air quality. These results highlight the value of satellite data assimilation and the need for
47 improved inventories and coordinated action on local and transboundary pollution.

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51 **1 Introduction**

52 Asian countries in the 2020s account for approximately half of the global anthropogenic emissions of major pollutants such as
53 nitrogen oxides (NO_x), carbon monoxide (CO), and sulfur dioxide (SO₂) (Guizzardi et al., 2025), as well as carbon dioxide
54 (CO₂) (Friedlingstein et al., 2023). This dominance reflects the region's dramatic development over the past three decades,
55 characterized by surging energy consumption (Energy Institute, 2025), industrial expansion (UNIDO, 2025), and urbanization
56 (UN, 2025). These anthropogenic sources are compounded by biomass burning and biogenic emissions that vary seasonally
57 and interannually, modulated by land-use change and climate variability (Field et al., 2009; Reddington et al., 2014; Marlier
58 et al., 2015; Peñuelas and Staudt, 2010). Consequently, emissions exhibit complex spatiotemporal heterogeneity, complicating
59 efforts to attribute observed air quality (AQ) changes and to understand their underlying drivers across the region.
60

61 Air pollution governance across Northeast and Southeast Asia operates through a multi-layered architecture of United Nations
62 (UN) agencies and regional intergovernmental bodies. Partnerships such as the Asia Pacific Clean Air Partnership (APCAP)



63 by the UN Environment Programme (UNEP) and the North-East Asia Clean Air Partnership (NEACAP) under the UN
64 Economic and Social Commission for Asia and the Pacific (ESCAP) provide voluntary, science-based frameworks for policy
65 exchange (UNEP, 2018; UN.ESCAP, 2025b). The Association of Southeast Asia Nations (ASEAN)'s Agreement on
66 Transboundary Haze Pollution (AATHP), ratified by all ten member states, remains the region's only legally binding
67 transboundary air pollution framework. The implementation of AATHP is supported by the ASEAN Zero-Burning Policy
68 (ASEAN Secretariat, 2003) and the ASEAN Peatland Management Strategy targeting a haze-free ASEAN by 2030 (ASEAN
69 Secretariat, 2023). Building on these frameworks, UN ESCAP adopted the Regional Action Programme on Air Pollution
70 (RAPAP) in 2022 as the first region-wide action programme negotiated and agreed by governments to curb air pollution
71 (UN.ESCAP, 2025a). Effective AQ management ultimately depends on national implementation, which varies considerably
72 across the region.

73

74 At the national level, both China and South Korea have intensified efforts to mitigate air pollution through strengthened
75 regulatory frameworks. Since 2013, China has implemented rigorous air pollution prevention and control action plans,
76 achieving substantial reductions in NO_x and SO₂ emissions and ambient particulate matter (PM) concentrations (Zhang et al.,
77 2019); however, surface ozone (O₃) pollution continued to increase due to nonlinear photochemistry (Liu et al., 2023; Ren et
78 al., 2022). Similarly, South Korea's Clean Air Conservation Act in 2007 and the Special Act on the Improvement of Air
79 Quality in Seoul Metropolitan Area in 2015 have effectively mitigated severe PM pollution, yet the country continues to face
80 challenges with rising O₃ levels and remaining high PM_{2.5} levels driven by complex secondary chemistry (Choi et al., 2024;
81 Kim and Lee, 2018; Oak et al., 2025).

82

83 Countries in Southeast Asia have developed national AQ frameworks (Verma et al., 2023). The Philippines enacted the
84 Philippine Clean Air Act (Republic Act No. 8749) in 1999, establishing national ambient AQ guideline values for criteria
85 pollutants in 2000, with PM_{2.5} standards tightened to 35 µg/m³ (24-hour) in 2020. Euro 4 vehicle emission standards have been
86 mandatory since 2016, with Euro 5 scheduled for 2027. Limited studies indicate rising surface O₃ trends in the southern
87 Philippines (Salvador et al., 2021) and persistent PM_{2.5} exceedances in Metro Manila (Cruz et al., 2023; Hopke et al., 2008;
88 Pabroa et al., 2022; Tseng et al., 2021), suggesting that current policies have not fully mitigated AQ. In Thailand, national
89 ambient AQ standards were first issued in 1995, and PM_{2.5} standards were revised in 2023 to 37.5 µg/m³ (24-hour), with Euro
90 5 becoming mandatory for new light-duty vehicles in 2024. Air quality management in Thailand has historically emphasized
91 open burning control, yet Northern Thailand continues to experience severe seasonal pollution episodes driven by forest and
92 agricultural burning and transboundary haze, indicating that biomass burning emissions remain significant despite existing
93 policies (Phairuang et al., 2017, 2019; Song et al., 2022; Thongsame et al., 2025). PM_{2.5} levels remain above both national
94 standards and World Health Organization (WHO) guidelines (Nikam et al., 2021), and surface O₃ increased between 2000 and
95 2019 (Kittipornkul et al., 2023; Sukkhum et al., 2022).

96



97 Nevertheless, effective AQ management in the region is constrained by several fundamental challenges. First, the ground-
98 based monitoring network is insufficient to fully characterize regional AQ. Regional monitoring is supported by ground-based
99 networks such as the Acid Deposition Monitoring Network in East Asia (EANET) (EANET, 2021), remote sensing
100 observations from Geostationary Environment Monitoring Spectrometer (GEMS) (Kim et al., 2020) and the Pandora Network
101 (<https://www.pandonia-global-network.org>), and international research initiatives such as the Task Force on Hemispheric
102 Transport of Air Pollution (TF HTAP) (Huang et al., 2017) and the Model Inter-comparison Study for Asia (MICS-Asia)
103 (Itahashi et al., 2020), which further advance understanding of transboundary pollution. While some regions (e.g., eastern
104 China, South Korea, and the Bangkok Metropolitan Region in Thailand) possess dense networks, vast areas in Southeast Asia
105 lack sufficient observation sites, limiting our ability to monitor surface air quality (Martin et al., 2019; Sokhi et al., 2022). For
106 example, in the Philippines, despite maintaining over 100 monitoring stations nationwide, the network remains largely
107 concentrated in the National Capital Region, with limited real-time public reporting. In Thailand, the Pollution Control
108 Department operates over 180 continuous monitoring stations across 65 provinces through the Air4Thai network, including
109 approximately 90 in the Bangkok Metropolitan Region, yet coverage remains uneven across rural and northern provinces.

110
111 Second, bottom-up emission inventories, constructed using statistical activity data and emission factors for each sector, have
112 inherent limitations. Reported emissions differ substantially across inventories, by more than a factor of 2 for anthropogenic
113 emissions at the country level (Elguindi et al., 2020; Hoesly et al., 2018), factors of 2 to 3 for biomass burning emissions at
114 the regional annual scale (van der Werf et al., 2017; Wiedinmyer et al., 2023), and factors of 2 to 3 for biogenic emissions at
115 the global and regional annual scale (Guenther et al., 2012; Sindelarova et al., 2014, 2022). These uncertainties are particularly
116 pronounced in Southeast Asia, where biomass burning inventories are largely uncertain due to limited burned area detection
117 due to persistent cloud cover, as well as uncertain local activity data and emission factors (Kurokawa and Ohara, 2020;
118 Stavrou et al., 2014; Vongruang et al., 2017). In addition, most bottom-up inventories cannot be updated rapidly enough to
119 represent short-term and spatially heterogeneous changes driven by evolving environmental policies, dynamic economic shifts,
120 and climate variability.

121
122 Third, AQ in the region reflects both local emissions and substantial transboundary transport (Crawford et al., 2021).
123 Consequently, domestic mitigation strategies alone are sometimes insufficient. Achieving national AQ targets may require a
124 regional coordination and quantitative attribution of cross-border contributions.

125
126 To address the limitation of ground-based observations, satellite remote sensing provides continuous and wide-area
127 observations. Low Earth Orbit (LEO) sensors have provided a global coverage of key pollutants and their precursors. These
128 include the Ozone Monitoring Instrument (OMI) (Levelt et al., 2018), the Global Ozone Monitoring Experiment-2 (GOME-2)
129 (Chan et al., 2023), the MODerate resolution Imaging Spectroradiometer (MODIS) (Pagano and Durham, 1993), the Visible
130 Infrared Imaging Radiometer Suite (VIIRS) (Cao et al., 2014), the Measurements of Pollution in the Troposphere (MOPITT)



131 instrument (Deeter et al., 2017), and the TROPOspheric Monitoring Instrument (TROPOMI) (Veefkind et al., 2012). More
132 recently, GEMS, a geostationary satellite, provides hourly observations over Asia (Kim et al., 2020). To integrate these satellite
133 observations and to provide comprehensive information on atmospheric composition variability (Bocquet et al., 2015;
134 Bousserez et al., 2016; Lahoz and Schneider, 2014; Mizzi et al., 2016), chemical data assimilation systems have been
135 developed, enabling long-term chemical reanalysis (Inness et al., 2019; Miyazaki et al., 2015). Chemical data assimilation can
136 also be used to derive emission estimates while optimizing broad atmospheric composition fields. The Tropospheric Chemistry
137 Reanalysis version 2 (TCR-2) (Miyazaki et al., 2020b) was developed within the Multi-mOdel Multi-cOnstituent Chemical
138 data assimilation (MOMO-Chem) framework (Miyazaki et al., 2020a), which has successfully demonstrated the capacity to
139 provide chemically and dynamically consistent fields by correcting model biases through observational constraints.

140

141 While satellites are useful for estimating emissions, relating these measurements to surface AQ remains challenging due to
142 satellites' reduced sensitivity toward the surface and uncertainties in vertical distribution. Sub-orbital observations from
143 aircraft campaigns fill this critical gap by providing detailed vertical profiles of atmospheric composition. NASA's Airborne
144 and Satellite Investigation of Asian Air Quality (ASIA-AQ) campaign was conducted in February and March 2024 to address
145 these challenges in the Asian region. The campaign deployed the NASA DC-8 aircraft for 16 research flights across the
146 Philippines, Taiwan, South Korea, and Thailand, obtaining comprehensive vertical profiles of trace gases and aerosols. These
147 observations provide valuable data for validating satellite products and chemical reanalysis systems, quantifying emissions,
148 and improving our understanding of pollution sources and transport processes in Asia.

149

150 In this study, we utilize the MOMO-Chem framework to comprehensively assess the origins of Asian air pollutants during the
151 ASIA-AQ campaign. Tropospheric Chemistry Reanalysis version 3 (TCR-3), constructed through the assimilation of multi-
152 species satellite retrievals, including NO₂, SO₂, formaldehyde (CH₂O), CO, O₃ and nitric acid (HNO₃), provides a
153 comprehensive dataset of atmospheric composition and top-down emission estimates. Section 2 describes the methodology.
154 Section 3 validates the chemical reanalysis against extensive in situ measurements acquired during the ASIA-AQ campaign.
155 Section 4 characterizes emissions of air pollutant precursors constrained by satellite observations and compares them with
156 bottom-up inventories. Section 5 quantifies the contribution of transboundary transport to O₃ using a source–receptor analysis
157 based on the optimized emissions. Section 6 summarizes the main findings.

158

159 **2 Methodology**

160 **2.1. Tropospheric chemistry reanalysis**

161 The TCR-3 product, generated using the MOMO-Chem framework, is used in this study (Miyazaki et al., 2020a, 2020b). It
162 employs an ensemble-based data assimilation system based on the Local Ensemble Transform Kalman Filter (LETKF) (Hunt



163 et al., 2007). The state vector includes various chemical concentrations as well as surface emissions of CO, NO_x, SO₂, and
164 isoprene, as well as lightning production of NO_x, with background error covariance linking observed atmospheric
165 concentrations to emission adjustments for each model grid cell.

166
167 The forecast model used is MIROC-Chem (Watanabe et al., 2011; Sekiya et al., 2018), a global chemistry–climate model that
168 simulates coupled tropospheric and stratospheric chemistry, tracer transport, and wet and dry deposition. The model is
169 configured with a horizontal resolution of approximately 1.125° × 1.125° (T106) and 32 vertical levels extending from the
170 surface to 4.4 hPa. The model explicitly represents 92 chemical species and 262 chemical reactions, including the oxidation
171 of non-methane volatile organic compounds (NMVOCs). The meteorological fields simulated by the MIROC-AGCM were
172 nudged to 6-hourly ERA5 reanalysis (Hersbach et al., 2020). Compared to the previous reanalysis (TCR-2), TCR-3 utilizes
173 updated inventories as prior emissions. Biomass burning emissions used in this study are from the Global Fire Assimilation
174 System version 1.2 (GFASv1.2) (Kaiser et al., 2012). Anthropogenic emissions are from the Hemispheric Transport of Air
175 Pollution version 3 (HTAPv3) inventory (Crippa et al., 2023), using the 2018 base year, regridded to the model’s T106
176 horizontal resolution. Agricultural sector emissions are excluded from HTAPv3 to avoid double-counting with agricultural
177 fires in GFASv1.2. Biogenic isoprene emissions are from the Model of Emissions of Gases and Aerosols from Nature version
178 2 (MEGANv2) (Guenther et al., 2012). Soil NO_x emissions are prescribed from the Global Emissions Inventory Activity
179 (GEIA) (Graedel et al., 1993).

180
181 The assimilated data are the v4.3 profile retrievals of O₃, CO, and HNO₃ from the Microwave Limb Sounder (MLS) (Livesey
182 et al., 2008; Santee et al., 2007), v9 MOPITT CO column measurements (Deeter et al., 2017), tropospheric NO₂ columns (van
183 Geffen et al., 2022), total SO₂ columns (Theys et al., 2021), and tropospheric CH₂O columns from TROPOMI (De Smedt et
184 al., 2021). We also assimilate VIIRS AOD products from the Deep Blue Aerosol L2 6-min Swath 6 km version 2
185 (AERDB_L2_VIIRS_SNPP and AERDB_L2_VIIRS_NOAA20) products. To reduce computational cost and minimize
186 representation errors in the data assimilation calculations, the super-observation approach is applied to the high-resolution
187 TROPOMI NO₂, SO₂, and CH₂O data (Rijdsdijk et al., 2025). The observation operator is constructed based on the prior profiles
188 and averaging kernels of the assimilated retrievals. To evaluate the impact of data assimilation, we also analyze the control
189 (CTL) simulation, which is performed without assimilation.

190
191 We conduct an additional sensitivity experiment for CO, in which Community Emissions Data System version v2021-04-21
192 (CEDS) (Hoesly et al., 2018) and Global Fire Emission Database version 4 (GFED4) (Giglio et al., 2013) are used as the prior
193 anthropogenic and biomass burning emission inventory, respectively, and TROPOMI CO column measurements (Borsdorff
194 et al., 2018) are assimilated in place of MOPITT CO. The results are discussed alongside the standard TCR-3 evaluation in
195 Sections 3 and 4.

196



197 **2.2 ASIA-AQ DC-8 airborne observations**

198 ASIA-AQ is a multinational field campaign led by NASA in collaboration with international partners in the Philippines,
 199 Taiwan, South Korea, and Thailand. Conducted in February and March 2024, the mission integrated airborne sampling,
 200 satellite retrievals, and ground-based monitoring to investigate emission sources, atmospheric chemistry processes, and
 201 transboundary transport mechanisms across Northeast and Southeast Asia. ASIA-AQ was designed to address several key
 202 objectives, including validation of satellite retrievals, particularly from GEMS, quantification and verification of emissions
 203 from key source regions, evaluation of chemical transport models, and characterization of aerosol and O₃ chemistry.

204

205 The campaign deployed the NASA DC-8 aircraft for 16 research flights across the four participating regions between February
 206 6 and March 27, 2024 (Figure 1). The aircraft measured over 100 trace gases and aerosol properties. For comparing with the
 207 TCR-3 reanalysis, this study focuses on measurements of O₃, NO_x, CO, and CH₂O. Figure 2 illustrates the time–altitude
 208 distributions of these species, alongside isoprene, highlighting the diverse pollution regimes sampled during the campaign,
 209 including urban emissions, biomass burning plumes, and biogenic-influenced air masses.

210

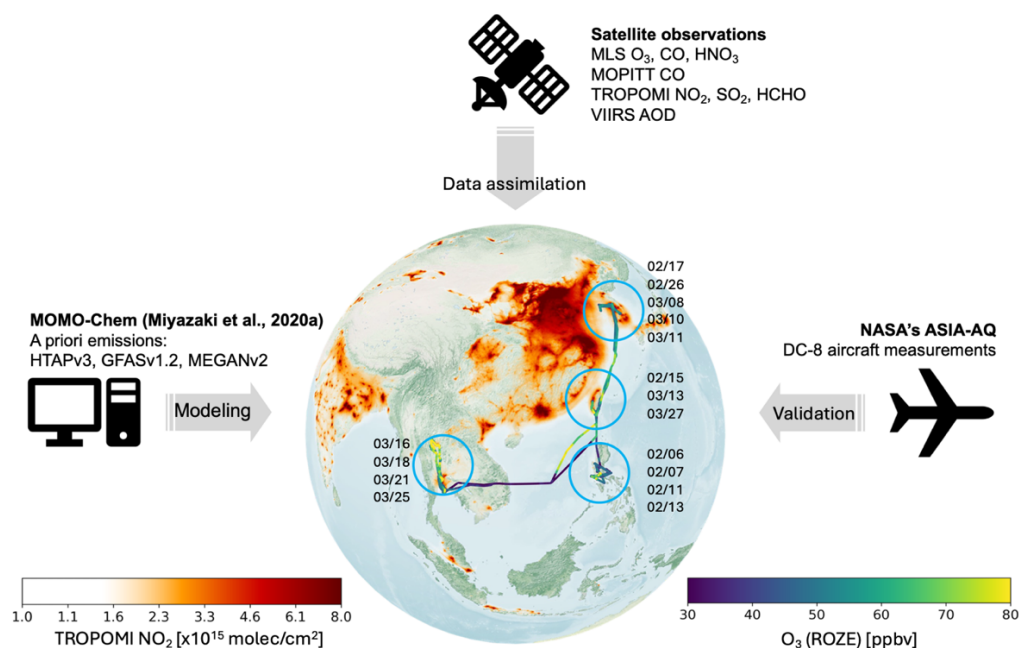
211 **Table 1.** DC-8 observation data from NASA’s ASIA-AQ campaign used for model evaluation. Data are from the 1-minute
 212 merged dataset downloaded on April 14, 2024.

Species	Instrument	Uncertainty	Institution	Reference
O ₃	Chemiluminescence instrument (NOxO3)	2-sigma accuracy of 5%*	National Center for Atmospheric Research (NCAR)	Ridley et al., 1992
	Rapid Ozone Experiment (ROZE)	1-sigma accuracy of 6.2%	NASA Goddard Space Flight Center	Hannun et al., 2020
NO _x	Chemiluminescence instrument (NOxO3)	2-sigma accuracy of 8% for NO and 11% for NO ₂	National Center for Atmospheric Research	Ridley et al., 1992
CO	Diode Laser Spectrometer (DACOM)	uncertainty of 2% (<1 ppm) to 5% (>1 ppm)	NASA Langley Research Center	Diskin et al., 2002
	Quantum Cascade Laser Absorption Spectrometer (MIRO)	the mean precision of 0.15 ppbv and the mean relative total uncertainty that can be up to 10%	Deutsches Zentrum für Luft- und Raumfahrt (DLR)	Knez et al., 2026
CH ₂ O	In Situ Airborne Formaldehyde instrument (ISAF)	accuracy of 10% + 10 pptv	NASA Goddard Space Flight Center	Cazorla et al., 2015
	Trace Organic Gas Analyzer with Time-of-Flight Mass Spectrometer (TOGA-TOF)	uncertainty of 35% and a lower limit of detection of 20 pptv	NSF NCAR Atmospheric Chemistry Observations & Modeling Laboratory	Jeong et al., 2025
Isoprene	Whole Air grab Samples analyzed via gas chromatography and mass spectrometry (WAS)	Precision of 3% and accuracy of 5% (Simpson et al., 2010)	University of California Irvine	Simpson et al., 2020



* The data additionally account for the uncertainty due to the unknown influence of water vapor inside the reaction vessel. Total uncertainties are provided as dependent variables in the data file and are shown as additional error bars in the vertical distributions in Fig. 5 (Section 3).

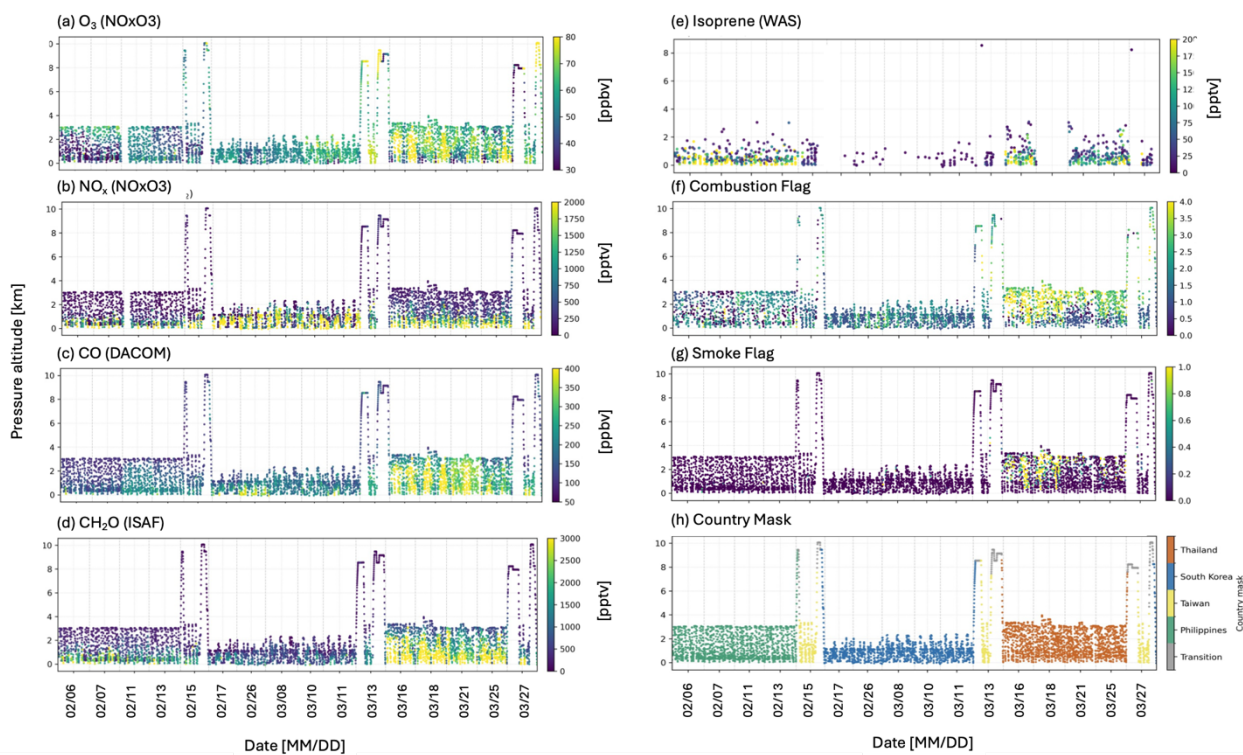
213
 214 To facilitate direct comparison with the two-hourly TCR-3 reanalysis data, the 1-minute averaged observational data in Table
 215 1 are further averaged into the two-hour windows and regrided to match the model's horizontal and vertical resolution.
 216 Additionally, we employ smoke and combustion classification flags developed by Miech et al. (2025) to qualitatively identify
 217 biomass burning influences. The campaign sampled distinct meteorological conditions and emission regimes across the four
 218 domains (Fig. 2). In the Philippines, the first two flights occurred during weak Northeast Monsoon conditions, while the latter
 219 two flights sampled during a strengthened Northeast Monsoon. Taiwan flights captured contrasting meteorological regimes,
 220 with the first flight conducted in pre-frontal conditions and the latter two in post-frontal environments. Thailand experienced
 221 weak synoptic flow during the first two flights, with the most severe pollution burdens observed as elevated concentrations of
 222 CO coincided with extensive biomass burning activity. A precipitation event on March 20 preceded the third Thailand flight
 223 on March 21, resulting in notably cleaner atmospheric conditions. Overall, meteorological conditions during the campaign
 224 were largely consistent with climatological norms for this season (Fig. S1). February was defined by a robust winter monsoon
 225 circulation driving northwesterly flow over Northeast Asia, while March marked a seasonal transition toward spring,
 226 characterized by weakened circulation and enhanced convection over the Maritime Continent.



227
 228 **Figure 1.** Schematic diagram of MOMO-Chem data assimilation framework integrating global multi-constituent satellite
 229 observations. The globe displays the 2024 annual mean TROPOMI NO₂ column densities, with four blue circles highlighting
 230 the four ASIA-AQ study regions: the Philippines, Taiwan, South Korea, and Thailand. The DC-8 flight tracks are overlaid on



231 the globe and are color-coded by O₃ (ROZE) mixing ratios across all altitudes. Flight dates are labelled adjacent to each study
 232 region.



233
 234 **Figure 2.** Time–altitude distributions of 1-min-averaged (a) O₃ from NO_xO₃, (b) NO_x from NO_xO₃, (c) CO from DACOM,
 235 (d) CH₂O from ISAF, and (e) isoprene from WAS aboard the DC-8 during ASIA-AQ, colored by observed mixing ratios.
 236 Panels (f) and (g) show the combustion and smoke flags, and panel (h) indicates the regional mask used to classify flight
 237 segments over the Philippines, Taiwan, South Korea, Thailand, and transition regions. The combustion flag (1 = higher
 238 combustion efficiency, 2 = mid, 3 = low, 4 = lowest) represents combustion characteristics inferred from CO/CO₂ enhancement
 239 ratios and can be associated with biomass-burning conditions. The smoke flag identifies the influence of biomass burning (0
 240 = lack of clearly defined biomass burning influence, 1 = clearly influenced).

241
 242 **3 Evaluation of TCR-3 against DC-8 measurements**

243 We evaluate the TCR-3 chemical reanalysis against the DC-8 measurements of O₃, NO_x, CO, and CH₂O. Reanalysis-
 244 measurement agreement is assessed using near-surface (below 1 km) spatial distributions (Fig. 3), time series along the flight
 245 tracks (Fig. 4), and vertical profiles (Fig. 5). Aggregate statistics for the entire campaign dataset are summarized in Table 2.
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Table 2. Summary statistics of the control simulation (CTL, without DA) and TCR-3 (with DA) evaluated against the DC-8 observations for the entire ASIA-AQ campaign. Observations are matched to their corresponding model grid cells in horizontal and vertical space and averaged within two-hour windows. Statistics are computed across all matched observation-model pairs spanning the full range of flight altitudes. R = correlation coefficient; NMB = Normalized Mean Bias (%); RMSE = Root Mean Square Error (ppbv).

Species	Instrument	Simulation	R	NMB (%)	RMSE (ppbv)
O ₃	NO _x O ₃	CTL	0.50	-17.4	17.1
		TCR-3	0.59	-13.9	15.3
	ROZE	CTL	0.47	7.31	12.5
		TCR-3	0.58	11.9	12.6
NO _x	NO _x O ₃	CTL	0.72	-14.7	2.60
		TCR-3	0.66	-29.7	2.62
CO	DACOM	CTL	0.64	-32.1	107
		TCR-3 (MOPITT)	0.77	-15.4	76.5
		TCR-3 (TROPOMI)	0.77	-0.5	87.5
	MIRO	CTL	0.60	-32.4	114
		TCR-3 (MOPITT)	0.75	-15.6	80.4
		TCR-3 (TROPOMI)	0.77	1.0	90.3
CH ₂ O	ISAF	CTL	0.75	54.6	1.78
		TCR-3	0.77	16.6	1.14
	TOGA-TOF	CTL	0.77	70.1	1.77
		TCR-3	0.79	29.4	1.16

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254 3.1 CO

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Observed CO shows pronounced regional variability, with highest concentrations over Thailand (mean ~310–316 ppbv, reaching 1,048 ppbv during biomass burning episodes), followed by Taiwan (~199–206 ppbv), South Korea (~195 ppbv), and the Philippines (~153–155 ppbv). They exhibit gradual decreases with altitude, with concentrations declining from peak boundary layer values to ~100–200 ppbv at 3–4 km. The surface elevated layers extend to 2–3 km over Thailand. TCR-3 shows substantial improvements for CO across all regions. The control simulation (CTL) exhibits a consistent low bias throughout the domain, with a Normalized Mean Bias (NMB) of -32.1% and -32.4% against DACOM and MIRO CO, respectively (Table 2). TCR-3 reduces the NMB to -15.4% and -15.6% while increasing the correlation coefficient (R) from 0.60–0.64 to 0.75–0.77 domain-wide. Particularly significant improvements occur over Thailand, where the correlation strengthens markedly (from 0.07 to 0.50 below 1 km against DACOM) and CTL substantially underestimates elevated CO.



264 TCR-3 better captures the observed high surface peaks in Northern Thailand exceeding 800 ppbv (Fig. 3), as well as their
265 temporal variability during episodic burning events in March (Fig. 4), reducing the surface NMB from -28% to -8% . Modest
266 improvements are also evident over the Philippines (where R increases from 0.48 to 0.62 and NMB improves from -35% to
267 -22% below 1 km) and Taiwan (R increases from 0.76 to 0.81), while improvements are less pronounced over South Korea.

268
269 TROPOMI TCR-3, the additional sensitivity experiment described in Section 2.1, yields a near-zero NMB (-0.5% and $+1.0\%$
270 against DACOM and MIRO, respectively) compared to -15.4% and -15.6% in the standard MOPITT TCR-3, with comparable
271 R (0.77) but higher RMSE (88–90 ppbv vs. 77–80 ppbv) (Table 2). Notably, TROPOMI TCR-3 reduces the underestimation
272 of elevated CO at ~ 2 km altitude over Thailand (NMB of $\sim +3$ to $+4\%$ above 2 km compared to $\sim -18\%$ in the standard TCR-
273 3). This improvement is associated with increased CO emissions in Northwest Thailand and Myanmar constrained by
274 TROPOMI (Section 4), which alters the spatial distribution of biomass burning emissions and their subsequent transport aloft.
275 However, TROPOMI TCR-3 simultaneously overestimates the DC-8 near-surface concentrations, with mean values reaching
276 ~ 449 ppbv against observed ~ 337 – 345 ppbv ($\sim 30\%$ overestimation) compared to ~ 318 ppbv in the standard TCR-3 (Fig. 5).
277 This trade-off (better free-tropospheric performance at the cost of surface overestimation) is particularly evident in Northern
278 Thailand, and may suggest that the model places too large a fraction of these emissions near the surface, or that vertical mixing
279 in the model does not efficiently redistribute them aloft. This also points to uncertainties in the vertical distribution of biomass
280 burning emissions, particularly their injection heights, as a key limiting factor in reproducing the observed vertical CO profile
281 (Pan et al., 2026).

282

283 3.2 CH₂O

284 Observed CH₂O concentrations show strong regional gradients, with the highest values over Thailand (mean 2.3–2.6 ppbv,
285 ranging up to 14.9 ppbv), followed by the Philippines (1.0–1.4 ppbv), Taiwan (1.0–1.2 pp), and South Korea (~ 0.6 ppbv).
286 CH₂O decreases with altitude more steeply than CO due to its shorter atmospheric lifetime, yielding strong vertical gradients
287 below 2 km (Fig. 5) that reflect surface emissions and secondary production from VOC oxidation near the surface. CTL
288 overestimates the boundary layer (below 1 km) CH₂O across the domain, with NMBs of $+89\%$ and $+85\%$ against ISAF and
289 TOGA-TOF. When aggregated over all altitudes, CTL shows domain-wide NMB of $+54.6\%$ against ISAF and $+70.1\%$ against
290 TOGA-TOF (Table 2). Data assimilation reduces these biases to $+16.6\%$ and $+29.4\%$, respectively, consistent with decreased
291 isoprene emissions (Section 4), with greater improvements over the Philippines (surface NMB reduced from $+74\%$ to $+22\%$)
292 and Thailand (surface NMB reduced from $+87\%$ to $+43\%$). Over Taiwan, surface NMB decreases from $+109\%$ to $+48\%$.
293 Above 2 km over Thailand, the CTL positive bias decreases from $+27\%$ to $+7\%$ in TCR-3, while the correlation increases
294 from 0.58 to 0.64.

295



296 **3.3 O₃**

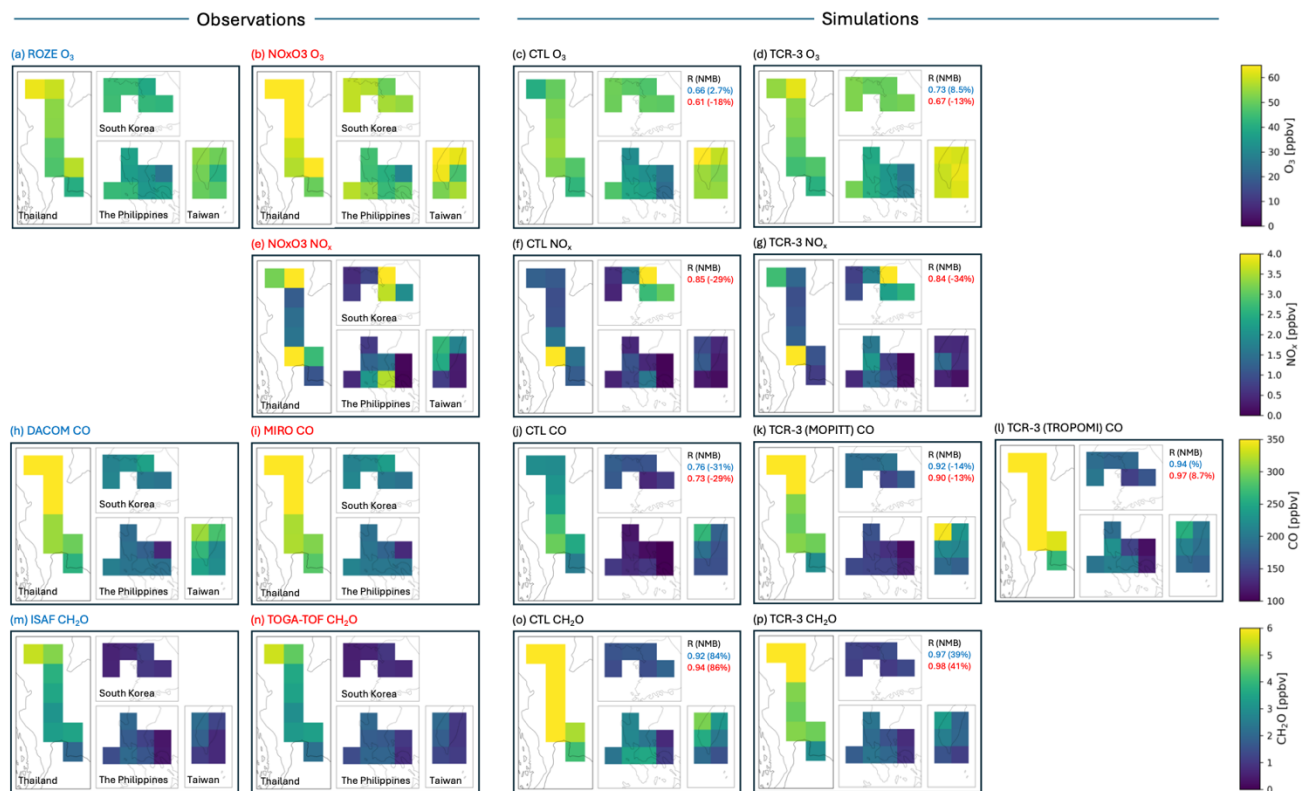
297 Observed O₃ concentrations range from 38–48 ppbv over the Philippines to 52–57 ppbv over South Korea and Taiwan, and
298 51–66 ppbv over Thailand. O₃ generally increases with altitude over the Philippines and South Korea (Fig. 5), but decreases
299 over Taiwan. The highest O₃ occurs during biomass burning episodes over Thailand, where concentrations exceed 127 ppbv.
300 Validation of O₃ is complicated by substantial differences between the two measurement instruments. Against NO_xO₃, CTL
301 shows NMB of –17.4% and R of 0.50, while TCR-3 reduces the bias to –13.9% and improves R to 0.59 (Table 2). Relative to
302 ROZE, the positive bias increases from +7.3% in CTL to +11.9% in TCR-3, although R still improves from 0.47 to 0.58.
303 Except over the Philippines, the assimilation generally increases O₃ throughout the troposphere. In the lower troposphere at
304 low latitudes, these adjustments are primarily driven by precursor emission optimization, while assimilation of MLS O₃ in the
305 UTLS exerts a large influence in the extratropic (Miyazaki et al., 2020; Selina et al., 2025). Thailand shows the most substantial
306 improvements. Below 1 km, R increases from 0.39 to 0.67 against NO_xO₃ and from 0.42 to 0.66 against ROZE. Above 2 km,
307 R improves from 0.26 to 0.61 against both instruments, with NMB improving from –24% to –13%. Episodic enhancements
308 associated with biomass burning plumes over Thailand in late March are better captured by TCR-3, though peak concentrations
309 remain slightly underestimated (Fig. 4). Over South Korea, Taiwan, and the Philippines, R increased by 0.05–0.10 in the
310 boundary layer.

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312 **3.4 NO_x**

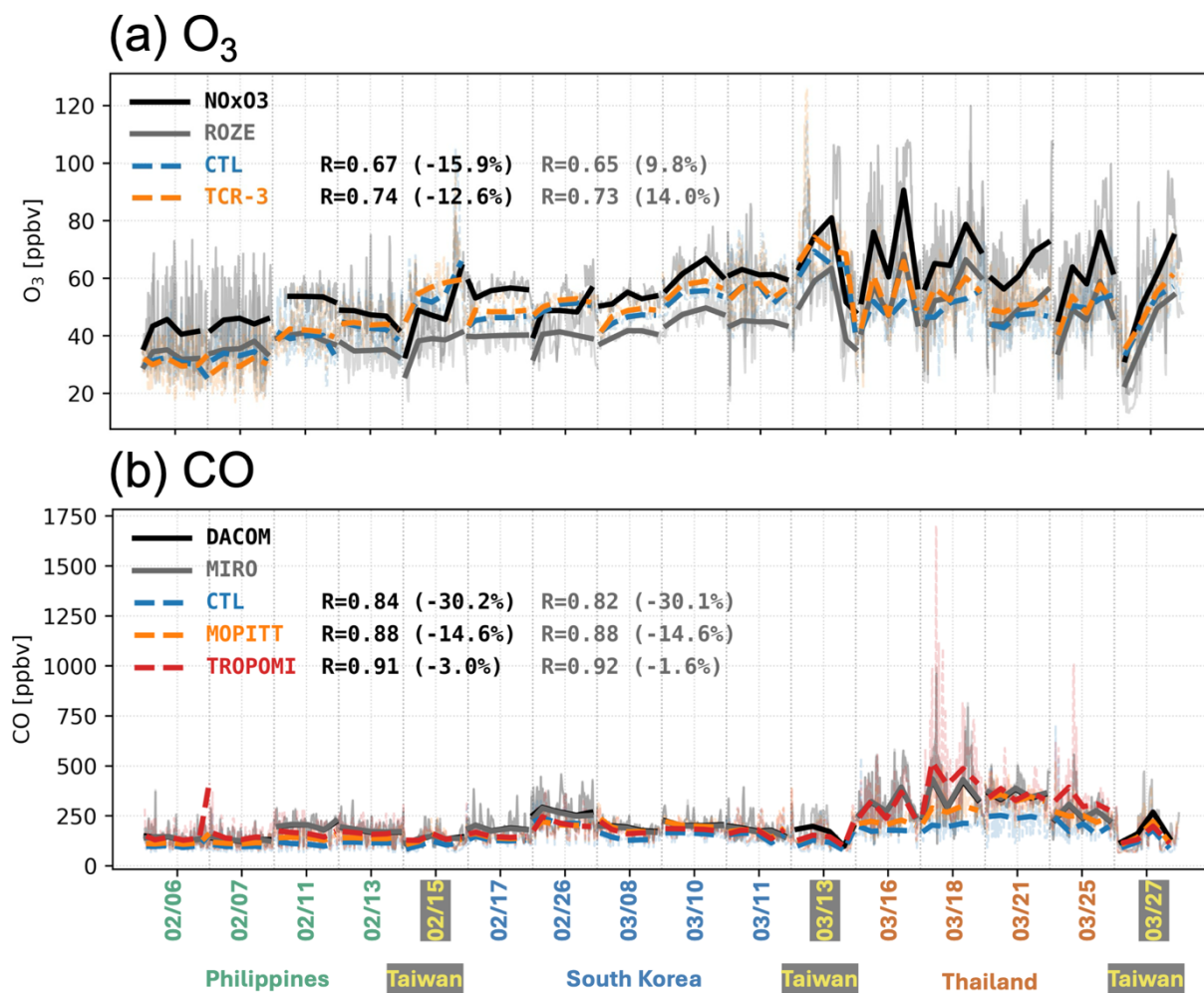
313 South Korea shows the highest surface NO_x associated with urban and industrial emissions, followed by Thailand, Taiwan,
314 and the Philippines. Although the overall benefit of assimilation is not evident when evaluated against NO_xO₃ measurements
315 aggregated over all altitudes (Table 2), the impact varies substantially by region and altitude. In Taiwan, assimilation reduces
316 the negative bias, improving NMB from –55% to –45% below 1 km and from –62% to –50% above 2 km. In the Philippines
317 above 2 km, NMB improves from –53% to –27%, though R decreases from 0.62 to 0.30. Over South Korea below 1 km, CTL
318 shows a slight positive bias (+25% NMB) that shifts to a negative bias of –18% in TCR-3. Over Thailand, assimilation
319 increases the negative bias from –36% to –43% below 1 km, while reducing the larger negative biases from –65% to –44%
320 and increasing R from 0.11 to 0.48 above 2 km. Underestimation of boundary layer NO_x has been commonly reported by
321 previous studies (Miyazaki et al., 2012; Skoulidou et al., 2021). Increasing model resolution remains essential to resolve urban
322 plumes and their non-linear chemistry (Valin et al., 2011; Sekiya et al., 2021). Evaluation against NO₂ measurements from
323 NO_xO₃ shows a similar pattern (not shown). CTL shows a near-neutral domain-wide bias (+2% NMB) with R of 0.68, while
324 TCR-3 introduces a negative bias (–19% NMB) with R reducing to 0.62, consistent with the NO_x results. In contrast to NO_x,
325 whose RMSE increases slightly with assimilation (Table 2), RMSE for NO₂ improves from 3.0 to 2.6 ppbv.

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Figure 3. Spatial distributions of near-surface concentrations (below 1 km) during ASIA-AQ on the TCR-3 grid at $1.125^\circ \times 1.125^\circ$. Panels are organized by species and dataset. For O_3 , panels (a) ROZE and (b) NO_xO_3 show observations, followed by simulations in (c) CTL and (d) TCR-3. For NO_x ($NO + NO_2$), panel (e) NO_xO_3 shows observations, followed by (f) CTL and (g) TCR-3. For CO, panels (h) DACOM and (i) MIRO show observations, followed by (j) CTL, (k) TCR-3 constrained with MOPITT, and (l) TCR-3 constrained with TROPOMI. For CH_2O , panels (m) ISAF and (n) TOGA-TOF show observations, followed by (o) CTL and (p) TCR-3. Correlation coefficient R and normalized mean bias NMB are calculated for each simulation against the corresponding observations; when two observational datasets are available, statistics relative to the first observation are shown in blue and those relative to the second observation are shown in red. For the spatial distribution analysis of CO in Taiwan, only DACOM observations were used.



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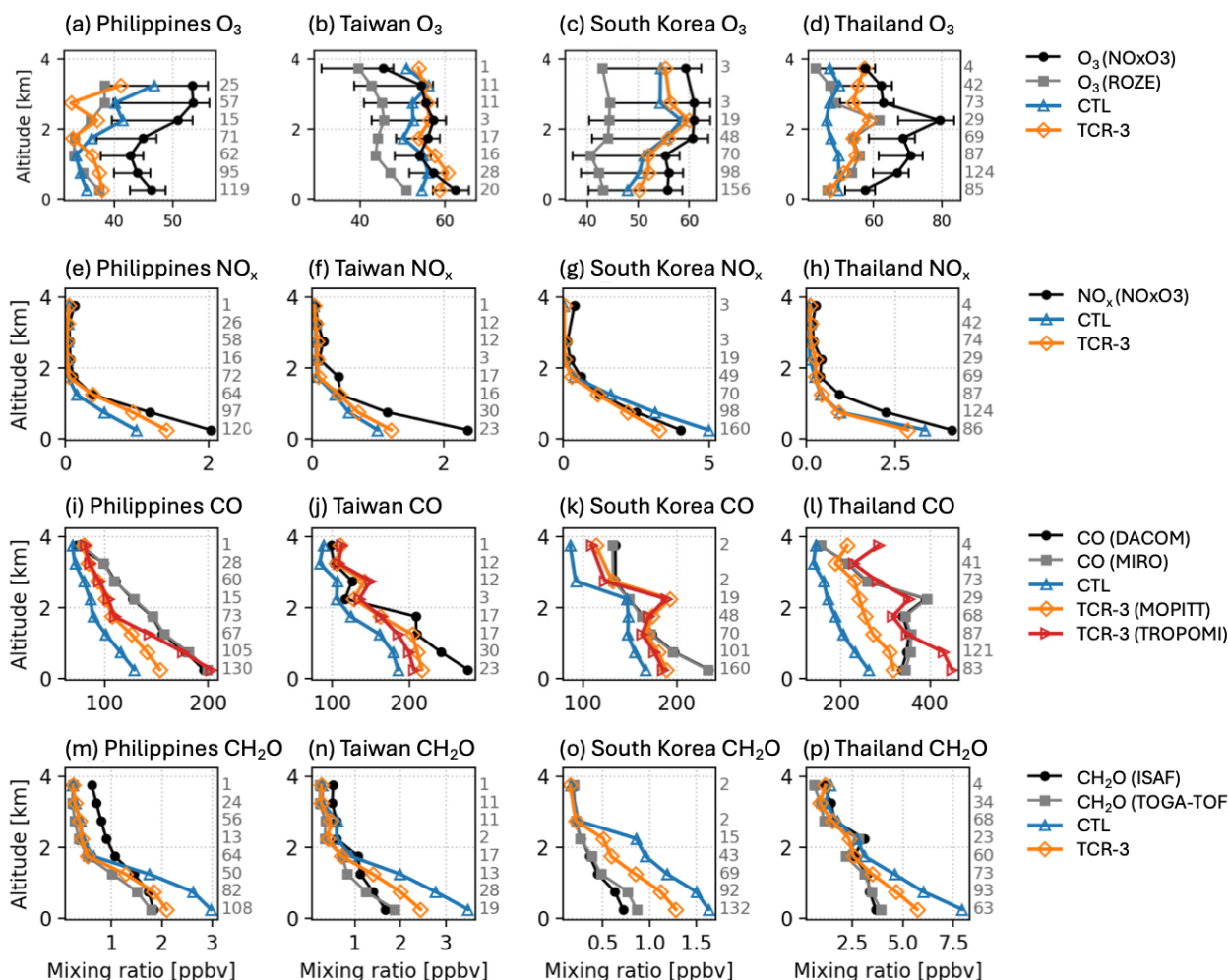
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Figure 4. Temporal evolution of (a) O₃ and (b) CO during ASIA-AQ DC-8 flights (February–March 2024) shown in sequential-ordinal representation. Thick lines show two-hour temporal averages for observations (NOxO3/ROZE for O₃; DACOM/MIRO for CO) and model simulations (CTL and TCR-3 for O₃; CTL and TCR-3 constrained by MOPITT and TROPOMI for CO). Thin background lines show observations and model output regridded to model horizontal and temporal (two-hourly) resolution along the flight track, illustrating spatial variability within each two-hour period. Data are plotted as a discontinuous timeseries, preserving the sampling order within each day while including visual breaks between dates to reflect the non-continuous observations. Correlation coefficients (R) and normalized mean bias (NMB%) are calculated from the two-hourly averaged data. Grey vertical lines denote calendar date boundaries.



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Figure 5. Vertical profiles of O_3 (a–d), NO_x (e–h), CO (i–l), and CH_2O (m–p) over four ASIA-AQ regions: the Philippines (a, e, i, m), Taiwan (b, f, j, n), South Korea (c, g, k, o), and Thailand (d, h, l, p). Black and grey curves with filled markers represent 1-min-averaged DC-8 observations, the blue curves with open triangles indicate the CTL simulation, and the orange curves with open diamonds show the TCR-3 posterior simulation. Specifically for the CO panels, the red curves with open right-pointing triangles show the TCR-3 (TROPOMI) posterior simulation. Numbers on the right y-axes indicating the number of data points. For O_3 , additional error bars on the black curves represent the total uncertainty of the NO_xO_3 instrument, accounting for the unknown influence of water vapor inside the reaction vessel. For the vertical distribution analysis of CO in Taiwan, only DACOM observations were used.



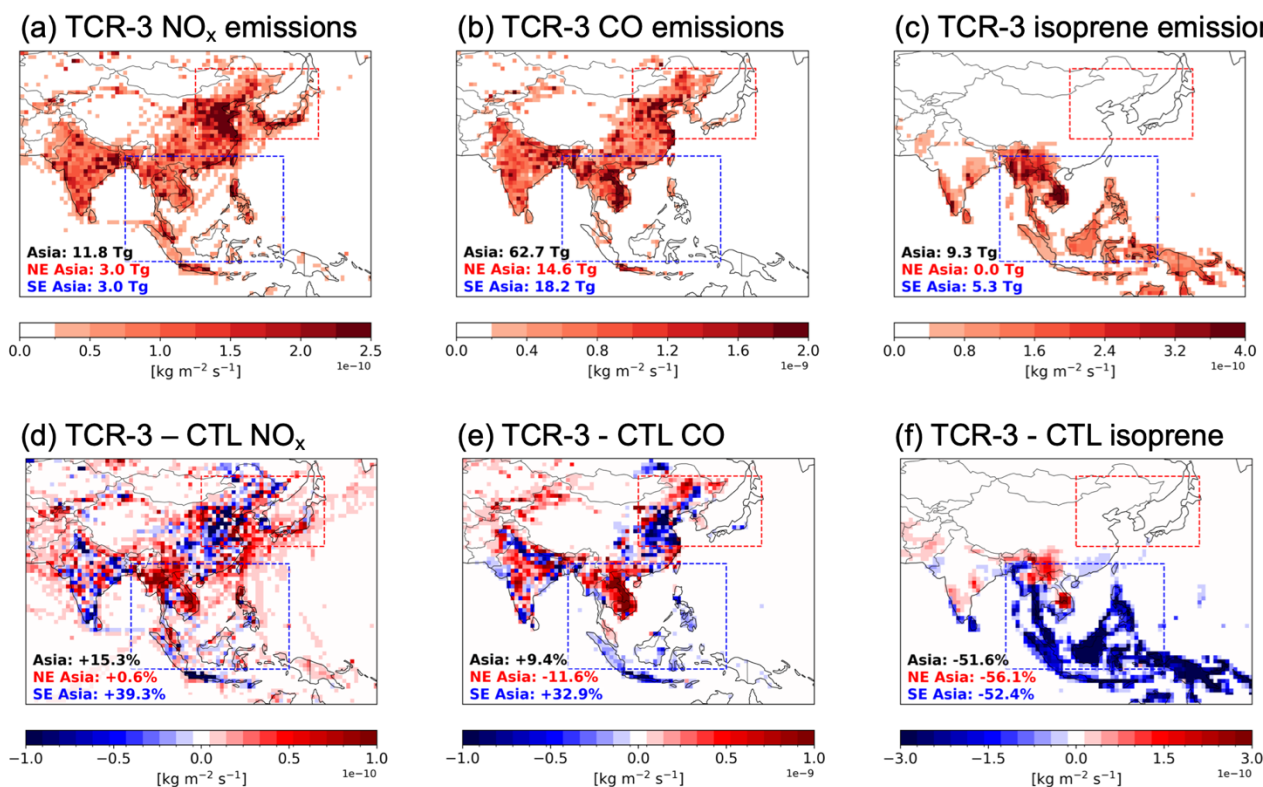
358 **4 Estimated NO_x, CO, and Isoprene emissions**

359 Figure 6 shows the optimized total emissions from TCR-3 and their adjustments relative to the prior emissions for NO_x, CO,
360 and isoprene during February–March 2024. We primarily focus on O₃ and its precursors in this analysis. The regional total
361 NO_x emissions increase by 15% in TCR-3 relative to the prior, reflecting increases across continental Southeast Asia (+39%)
362 in contrast to decreases in mainland China (−9%). At the national scale, the largest relative increases are found in Myanmar
363 (+143%), Laos (+115%), and Cambodia (+95%). Among the four ASIA-AQ flight regions, South Korea shows a broad
364 decrease (−18%). Taiwan (+14%), Thailand (+18%), and the Philippines (+36%) show spatially heterogeneous adjustments,
365 with reductions over major urban areas (e.g., southwest Taiwan, Bangkok, southern Manila, Cebu) and increases elsewhere
366 modulating the urban-rural gradients. Over East Asia, total ship NO_x emissions also increase by approximately 59 %.

367
368 The regional total CO emissions increase by 9%, reflecting increases in continental Southeast Asia (+33%). Over China,
369 assimilation decreases CO emissions by 18% in the eastern industrial corridor (28–42°N, 110–122°E; from 13.2 to 11.4 Tg)
370 and increases them by 27% in southwestern China near the Myanmar/Laos border (21–30°N, 97–108°E; from 2.3 to 4.3 Tg),
371 collocated with biomass burning areas. Increases exceeding +100 % are found in Laos and Cambodia. Unlike NO_x, CO
372 emissions increase over South Korea (+34%). Taiwan shows similar increases (+25%), whereas Philippines shows decreases
373 (−41%). Thailand exhibits large increases (+52%), driven primarily by enhanced emissions outside Bangkok in regions
374 impacted by biomass burning.

375
376 Isoprene emissions decrease by 52% domain-wide, with pronounced reductions across Maritime Southeast Asia (−70%), the
377 Philippines (−63%), Taiwan (−71%), Thailand (−40%), and South Korea (−69%). These reductions are consistent with
378 previous studies showing MEGAN overestimates isoprene emissions over Southeast Asian tropical forests by factors of 2–4
379 (Langford et al., 2010; Stavrou et al., 2014). In contrast, mainland China shows an increase (+60%). The adjustments
380 underscore the substantial uncertainties in bottom-up isoprene inventories over Asian tropical regions, with important
381 implications for O₃ production, OH budgets, and secondary organic aerosol formation (Yoon et al., 2025). At the same time,
382 the top-down isoprene estimates remain subject to considerable uncertainty. They are sensitive to the representation of isoprene
383 chemistry and the broader chemical environment within the model (Choi et al., 2025; Wells et al., 2020; Yang et al., 2023). In
384 addition, potential biases in the assimilated TROPOMI CH₂O retrievals may influence the optimized emissions and require
385 further investigation (Oomen et al., 2024).

386



387
 388 **Figure 6.** Spatial distributions of NO_x (left), CO (center), and isoprene emissions (right) averaged over February–March 2024
 389 from the TCR-3 (upper panels) and their differences relative to the prior emissions (lower panels). Domain-total emissions are
 390 indicated in teragrams (Tg) in panels (a–c), and percentage changes relative to the prior emissions are shown in panels (d–f).
 391

392 4.1 Anthropogenic emissions

393 We compare our top-down estimates with five bottom-up inventories: HTAPv3 and HTAPv3.2 (Crippa et al., 2023; Guizzardi
 394 et al., 2025), Air Quality North East Asia version 4 (AQNEAv4), the Emissions Database for Global Atmospheric Research
 395 version 8.1 (EDGARv8.1, https://edgar.jrc.ec.europa.eu/dataset_ap81, Crippa et al., 2020), CEDSv2025-04-18 (Hoesly et al.,
 396 2018), and the Copernicus Atmosphere Monitoring Service Global Anthropogenic Emissions version 6.2 (CAMSGLOB-
 397 ANTv6.2) and mosaic version 1.0 (CAMSGLOB-ANTm1.0) (<https://permalink.aeris-data.fr/CAMSGLOB-ANT>, Denier
 398 van der Gon et al., 2025; Soulie et al., 2024). For this comparison, the top-down total emissions at each grid cell are partitioned
 399 into anthropogenic, biomass-burning, and biogenic components using the a priori fractional contributions from each sector



400 (see Section 2.1). The resulting TCR-3 sectoral emissions are then downscaled to $0.1^\circ \times 0.1^\circ$ resolution using spatial proxies
401 derived from the a priori emissions. Ship emissions are not included in the region-specific analysis. Regional totals
402 (anthropogenic + biomass burning) of TCR-3 NO_x and CO emissions for February-March 2024 are summarized in Table S1.

403
404 EDGARv8.1 provides globally consistent emissions for the 1970–2022 period by applying standardized methodologies and
405 default emission factors across all countries. EDGAR serves as a foundational baseline for several other inventories compared
406 in this study (HTAP, CEDS, and CAMS-ALOG-ANT) and therefore these inventories are not fully independent of each other.
407 HTAPv3 covers 2000–2018 using EDGARv6.1 as a baseline, integrated with regional data such as REASv3.2.1 for East and
408 Southeast Asia (Kurokawa and Ohara, 2020). HTAPv3.2 extends this coverage to 2020, updates the baseline to EDGARv8,
409 and incorporates MEICv1.4 for China (Wu et al., 2024). AQNEAv4, developed by Seoul National University for the ASIA-
410 AQ campaign, integrates 2019 national data for East Asia with local datasets for Taiwan (base year 2024) and Thailand (base
411 year 2022), alongside MIXv2.3 (Li et al., 2024) for background regions. CEDSv2025-04-18 covers 1980–2023. CAMS-GLOB-
412 ANTv6.2, part of the Copernicus Atmosphere Monitoring Service (CAMS), covers 2000–2025. We use CAMS-GLOB-
413 ANTv1.0 for 2024, which utilizes MEICv1.4 for China (Wu et al., 2024). Sectoral contributions are further examined using
414 the IPCC 2006 categories: Power (1A1), Industry (1A2, 1B, 2), Residential + Commercial + Others (Buildings; 1A4, 1A5),
415 Transport (1A3), and Agriculture (3A, 3C, 5A, 4).

416
417 Figures 7 and 8 present time series of average February–March anthropogenic NO_x and CO emissions for 2000–2025, placing
418 the 2024 TCR-3 estimates in the context of historical inventory trends. The spread among inventories is wider for CO than
419 NO_x, likely reflecting the stronger sensitivity of CO emissions to combustion conditions, technology types, and emission factor
420 assumptions. Figure 9 provides a detailed sectoral breakdown for each inventory using the closest year to the 2024 ASIA-AQ
421 period. For NO_x, transport is the dominant sector across most regions and inventories, with exceptions in mainland China and
422 India, where industry and power lead, respectively, in several datasets. For CO, the dominant sector varies more widely by
423 region and inventory. Among these inventories, CEDS attributes a larger fraction of national CO totals to the power sector (8–
424 48%) compared to other inventories (1–7%, except AQNEAv4 in South Korea at 12%), accompanied by comparatively small
425 industry sector contributions. EDGAR shows lower transport contributions for both CO and NO_x across most regions, with
426 CO transport fractions of only 3–4% in mainland China, India, and the Philippines compared to 6–19%, 11–26%, and 20–
427 35%, respectively, in other datasets.

428
429 Following the approach of Crippa et al. (2023) and Guizzardi et al. (2025), we quantify the range of estimates by calculating
430 the coefficient of variation (CV; standard deviation divided by mean) across five inventories with the same base year (2019).
431 Figures 10 and S2 show the spatial distribution of inter-inventory CV at $0.1^\circ \times 0.1^\circ$ resolution by sector for NO_x and CO,
432 respectively. Figure S3 provides the CV aggregated by region for both NO_x and CO. Across sectors, the power sector exhibits
433 the largest inter-inventory variability, with CVs exceeding 120% over most regions and surpassing 200% in parts of India. The



434 industry sector also shows substantial variability, with CVs typically in the range of 90-120%. In contrast, the buildings sector
435 has lower variability, although moderate CVs are evident in several ASIA-AQ study regions, including Thailand, South Korea,
436 and the Philippines. Transport displays a more spatially heterogeneous pattern, with higher CVs occurring outside major
437 transport corridors, indicating greater disagreement in the representation of diffuse and non-highway emission sources (Fig.
438 S4).

439
440 The intercomparison of emissions is discussed separately for each region below.

441 442 **Mainland China**

443 Inventories consistently show declining NO_x and CO emissions after 2012 (Fig. 7, 8). For NO_x, trends before 2020 are
444 comparable across HTAP, EDGAR, and CEDS (−92 to −115 Gg month^{−1} yr^{−1}), while CAMS-GLOB-ANT shows a
445 substantially weaker decline (−27 Gg month^{−1} yr^{−1}). This discrepancy is primarily driven by the power sector, which exhibits a
446 much smaller decreasing trend in CAMS-GLOB-ANT. Inter-inventory agreement is relatively strong in 2019 (CV = 11%; Fig.
447 S3). Beginning in 2021, EDGAR and CEDS indicate increasing NO_x emissions, whereas CAMS-GLOB-ANT shows a rebound
448 in 2021 followed by a renewed decline. Among regional inventories, AQNEA (2019) yields the lowest estimates, while
449 CAMS-GLOB-ANTm1.0, which incorporates MEICv1.4 for China, produces higher emissions than CAMS-GLOB-ANTv6.2.
450 TCR-3 estimates for 2024 fall at the upper end of the inventory range.

451
452 For CO, emission reductions between 2012 and 2020 are larger in magnitude (−192 to −898 Gg month^{−1} yr^{−1}) than after that
453 period, with CAMS-GLOB-ANT again showing the slowest decline. These reductions are mainly driven by the industry and
454 buildings sectors. Although inter-inventory spread is substantial in the early 2000s, it narrows over time, reaching a CV of
455 19% by 2019. AQNEA (2019) provides the highest CO estimates among inventories, and both TCR-3 products exceed CAMS-
456 GLOB-ANT inventories in 2024.

457
458 Transport exhibits large inter-inventory variability outside major highway corridors for both NO_x and CO (Figs. 10, S2, S4).
459 This likely reflects the limited representation of area sources in global inventories compared to HTAP and AQNEA, suggesting
460 a potential underestimation of emissions from secondary roads and non-road machinery.

461 462 **India**

463 All inventories show increasing NO_x emissions prior to 2020 (21–32 Gg month^{−1} yr^{−1} since 2000), largely driven by the power
464 sector, with moderate inter-inventory agreement in 2019 (CV = 11%). HTAP reports higher NO_x emissions than other
465 inventories, potentially reflecting different assumptions regarding the penetration of low-NO_x burner technologies. TCR-3
466 follows the general increasing trajectory, similar to CAMS-GLOB-ANT, and lies between lower estimates (AQNEA, EDGAR,
467 CEDS) and higher estimates (HTAP).



468

469 For CO, inventory trends diverge after 2014 (CV = 15% in 2019). EDGAR and HTAPv3.2 continue to increase, while HTAPv3
470 shows a flattening trend. In contrast, CEDS and CAMS-GLOB-ANT exhibit decreasing emissions after 2014. Industry
471 dominates the continued increases in HTAPv3.2 and EDGAR, whereas reductions in buildings and transport drive the declining
472 trends in HTAPv3, CAMS-GLOB-ANT, and CEDS. TCR-3 CO estimates for 2024 exceed all bottom-up inventories and align
473 with the extrapolated upper range represented by HTAPv3.2.

474

475 **The Philippines**

476 NO_x emissions show a consistent increasing trend after 2009, primarily driven by the transport sector, with additional
477 contributions from the power sector. The growth rates before 2020 in CEDS and HTAP (2.8–3.8 Gg month⁻¹ yr⁻¹) are
478 approximately 2–3 times higher than those in EDGAR and CAMS-GLOB-ANT (1.4–1.7 Gg month⁻¹ yr⁻¹). EDGAR and
479 CAMS-GLOB-ANT generally report lower emissions than CEDS (highest), HTAP, and AQNEA. These differences are
480 largely associated with the transport sector. Inter-inventory spread in 2019 is moderate (CV = 25%). TCR-3 NO_x estimates for
481 2024 exceed all bottom-up inventories by 30–52% (Fig. 9).

482

483 CO emissions also increase after 2009, with relatively small inter-inventory spread in 2019 (CV = 14%). AQNEA provides
484 the lowest CO estimates for that year. The two TCR-3 products show substantial divergence: TCR-3 (MOPITT) is 30–52%
485 lower than bottom-up inventories, whereas TCR-3 (TROPOMI) falls within the range of bottom-up estimates.

486

487 **Taiwan**

488 Taiwan shows the largest inter-inventory spread for both NO_x and CO (CV = 49% and 68%, respectively), driven primarily
489 by CAMS-GLOB-ANT, which yields elevated estimates for both species owing to large contributions from the transport sector
490 (CO and NO_x) and the industry sector (CO). CEDS shows the steepest NO_x decline over the period (–1.6 Gg month⁻¹ yr⁻¹),
491 followed by EDGAR (–1.1) and HTAP (–0.7). HTAP begins ~40% lower in 2000 and converges with CEDS by 2018–2019,
492 as does AQNEA by 2019. TCR-3 in 2024 falls within the broad bottom-up spread, broadly consistent with the EDGAR
493 trajectory.

494

495 For CO, the large inter-inventory spread and divergence in long-term trends are dominated by CAMS-GLOB-ANT, with
496 transport sector differences as the principal driver. AQNEA lies near the center of the 2019 spread. TCR-3 (MOPITT) aligns
497 closely with AQNEA, while TCR-3 (TROPOMI) sits at the lower end near CEDS.

498

499 **South Korea**

500 All inventories show a NO_x decline since 2015, though the rates diverge substantially (–1.3 to –5.5 Gg month⁻¹ yr⁻¹). CEDS
501 and HTAP exhibit broadly similar trends, as do CAMS-GLOB-ANT and EDGAR, although the two latter inventories diverge



502 in their absolute levels throughout the record, with EDGAR consistently the highest. The dominant declining sector also differs.
503 Industry drives the reduction in EDGAR, whereas power dominates in CAMS-GLOB-ANT. The inter-inventory spread in
504 2019 is moderate ($CV = 30\%$), with AQNEA at the lower end of the range. By 2023, CAMS-GLOB-ANT converges toward
505 the lower-bound estimates following a steep decline attributable to reductions in transport and industry. TCR-3 in 2024 falls
506 near the middle of the spread. Transport is the dominant NO_x source across all inventories.

507
508 For CO, the inter-inventory spread is substantially larger ($CV = 61\%$), driven by EDGAR and CAMS-GLOB-ANT reporting
509 values 2-4 times higher than other inventories, primarily due to large industrial emission contributions absent from regional
510 inventories and CEDS. The two inventories also show divergent temporal behaviour: EDGAR and CAMS-GLOB-ANT exhibit
511 increasing CO emissions between 2010 and 2015, attributable to industry and buildings sectors, whereas the remaining
512 inventories show a continuous decline since 2000, dominated by reductions in transport. AQNEA falls between the lower-
513 bound estimates (HTAP, CEDS) and the higher-bound estimates (EDGAR, CAMS-GLOB-ANT). TCR-3 lies at the lower end
514 of the spread, near CEDS and HTAP.

515

516 **Thailand**

517 All inventories indicate increasing NO_x emissions in Thailand since 2000, with trend rates ranging from 0.9 to 2.1 Gg month⁻¹
518 yr⁻¹. The inter-inventory spread remains relatively modest in 2019 ($CV = 16\%$), with AQNEA reporting the lowest emissions,
519 near CAMS-GLOB-ANT. TCR-3 aligns more closely with the higher-end, upward-trending HTAP and CEDS estimates.

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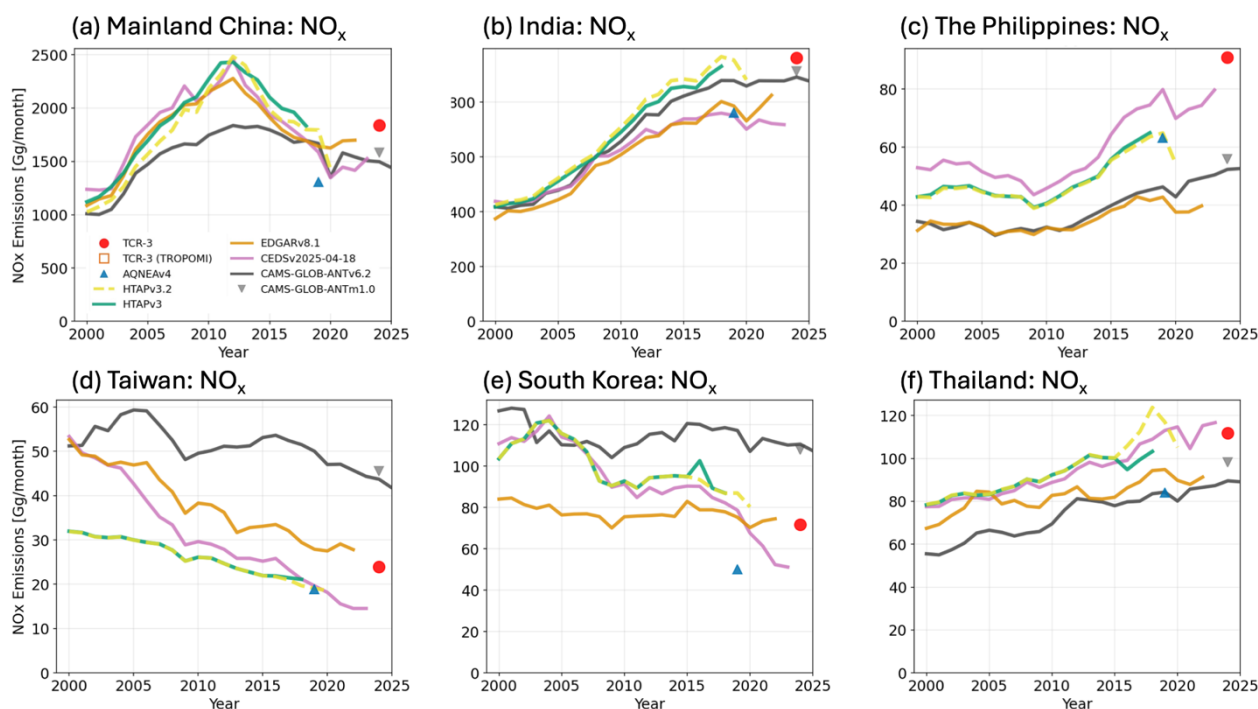
521 For CO, all inventories transitioned from increasing trends in the early 2000s to decreasing trends after the mid-2010s (2012–
522 2015), with the exception of CAMS-GLOB-ANT, which shows near-zero trend after 2016. This divergence drives a
523 substantially wider inter-inventory spread by 2019 ($CV = 48\%$). AQNEA reports the lowest CO in 2019. TCR-3 (MOPITT)
524 falls near the middle of the spread, while TCR-3 (TROPOMI) lies at the lower end, close to the CEDS trajectory.

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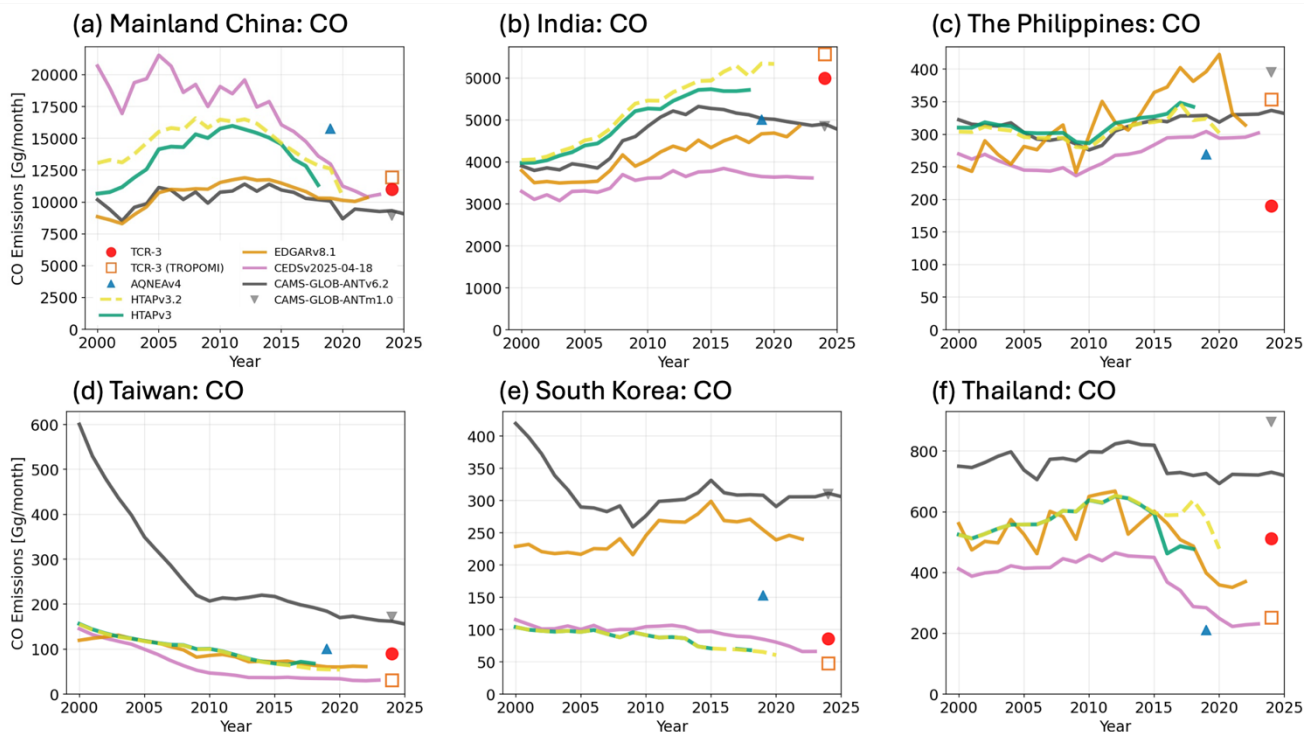
526 The diversity among bottom-up inventories underscores the need for continued improvement of individual products, while our
527 top-down estimates can provide quantitative guidance on the magnitude of revisions required, albeit with their own
528 uncertainties. Bottom-up inventories diverge because they differ in activity data, emission factors, and the spatial-temporal
529 allocation of emissions (Solazzo et al., 2021). Activity statistics such as fuel consumption are relatively well documented for
530 major economies but often rely on extrapolation or expert judgment in developing regions (Deng et al., 2022). Emission factors
531 depend on technology, combustion conditions, and abatement equipment, and are inconsistently constrained by limited field
532 measurements (Liu et al., 2015). Allocation approaches range from population-based proxies in global inventories to facility-
533 level point sources in regional ones (Crippa et al., 2024; Gurney et al., 2019). Methodological definitions add further
534 discrepancies: the use of Net versus Gross Calorific Values introduces 5 to 10% differences in fuel-related emissions (IPCC,
535 2008; Janssens-Maenhout et al., 2019), and global inventories applying Tier 1 defaults (e.g., EDGAR) systematically diverge



536 from national reports using Tier 2 or Tier 3 country-specific methods (e.g., HTAP, AQNEA) (Oda et al., 2018). The detailed
 537 comparison presented here, involving multiple bottom-up inventories and top-down constraints, helps identify priority areas
 538 for improvement and provides useful guidance for future inventory development.
 539



540
 541 **Figure 7.** Time series of average February–March anthropogenic NO_x emissions across six Asian regions. The panels display
 542 monthly emissions [Gg/month] for (a) mainland China, (b) India, (c) the Philippines, (d) Taiwan, (e) South Korea, and (f)
 543 Thailand. Each region compares multiple emission inventories: TCR-3 (red circle), AQNEAv4 (blue triangle), HTAPv3.2
 544 (yellow dashed line), HTAPv3 (green solid line), EDGARv8.1 (orange solid line), CEDSV2025-04-18 (purple solid line),
 545 CAMS-GLOB-ANTv6.2 (dark grey solid line), and CAMS-GLOB-ANTm1.0 (grey downward triangle).
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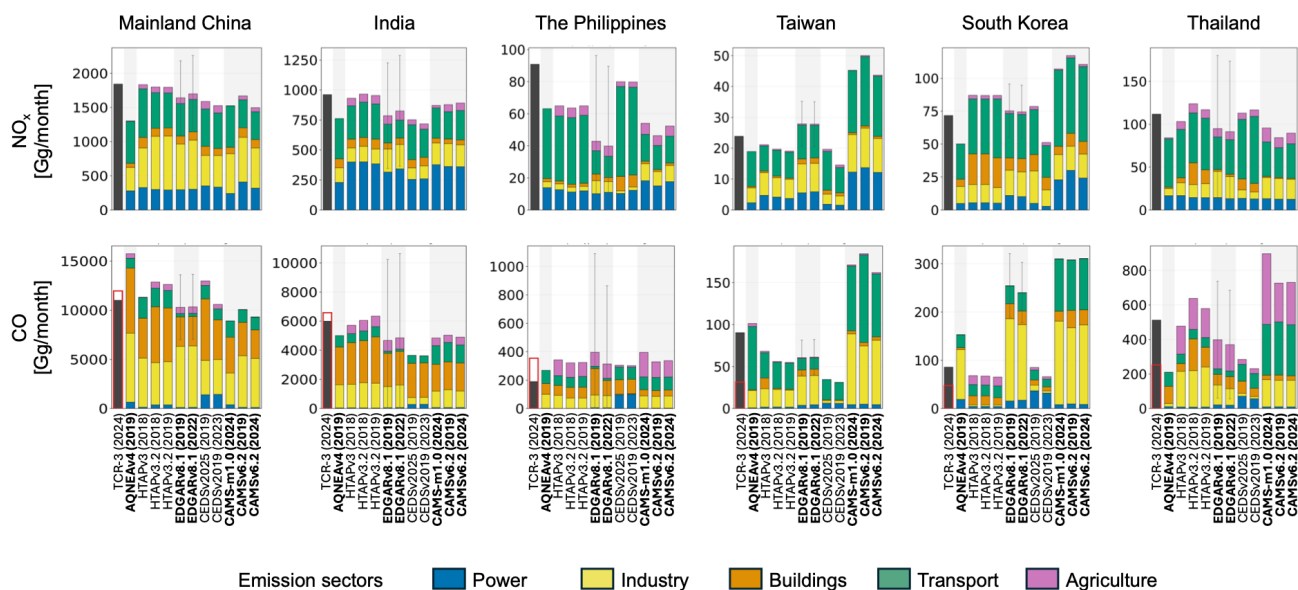
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Figure 8. Same as Fig. 7 but for anthropogenic CO emissions. Each region compares multiple emission inventories, including TCR-3 (MOPITT) (red circle) and TCR-3 (TROPOMI) (open orange square).



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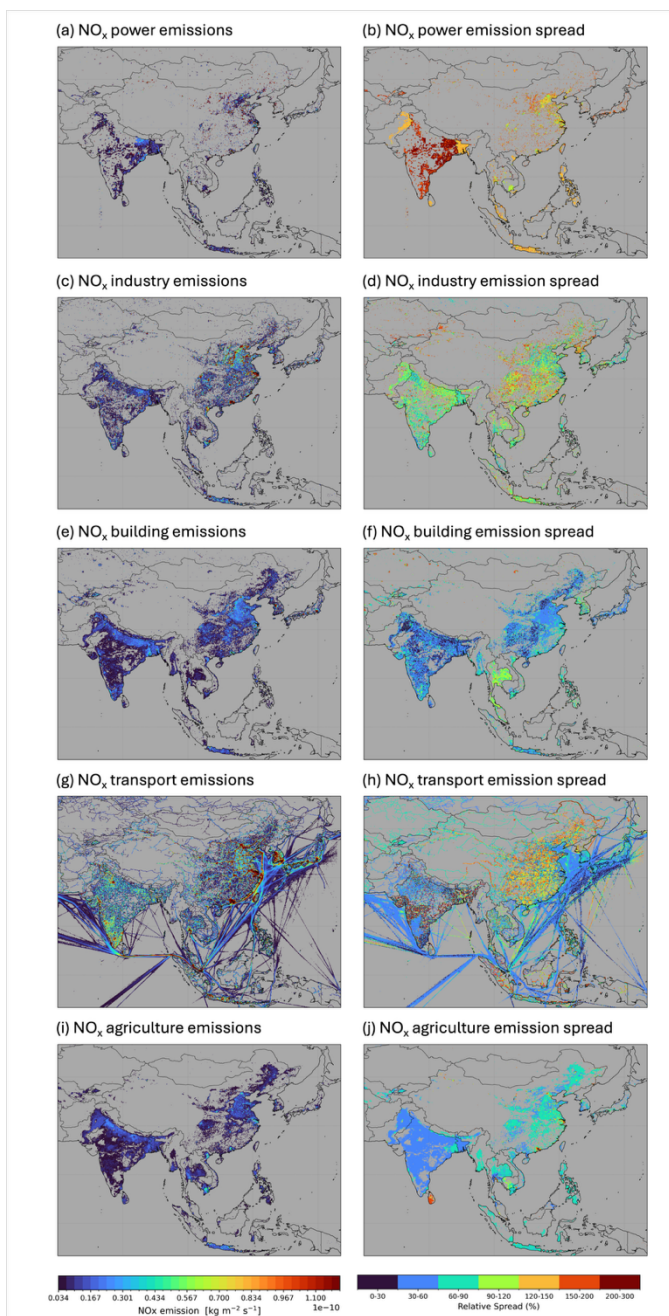
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Figure 9. Regional total anthropogenic emissions of NO_x (top row) and CO (bottom row) during February–March 2024, expressed in Gg/month for mainland China, India, the Philippines, Taiwan, South Korea, and Thailand. Bars are grouped by inventory: TCR-3 (2024), AQNEAv4 (2019), HTAPv3 (2018), HTAPv3.2 (2018 and 2019), EDGARv8.1 (2019 and 2022), CEDSv2025-04-18 (2019 and 2023), CAMS-GLOB-ANTm1.0 (2024), and CAMS-GLOB-ANTv6.2 (2019 and 2024). CAMS-GLOB-ANTm1.0 and v6.2 are the only inventories covering 2024, the target year of TCR-3 and ASIA-AQ. For CO panels, open red squares on the TCR-3 (2024) bars indicate the TCR-3 (TROPOMI) values. Stacked color indicates the sectoral contribution: Power (blue), Industry (yellow), Buildings (orange), Transport (green), and Agriculture (purple). TCR-3 (2024) does not provide sectoral information and is shown in dark grey. The TCR-3 prior excludes agricultural emissions to avoid double-counting with biomass burning emissions. Error bars represent 1-sigma EDGAR uncertainties from Table S3 of Crippa et al. (2019), ranging from 26–125% for NO_x and 26–175% for CO depending on the region. Alternating white and light-grey background shading visually separate inventory groups.



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Figure 10. Mean sectoral NO_x emissions (a,c,e,g,i) and Inter-inventory coefficient of variation (CV, %) (b,d,f,h,j) for Feb-Mar 2019 at $0.1^\circ \times 0.1^\circ$ resolution by sector: Power (a, b), Industry (c, d), Buildings (e, f), Transportation (g, h), and Agriculture (i, j). Mean and CV are calculated from five inventories with base year 2019: AQNEAv4, HTAPv3.2, EDGARv8.1, CEDSv2025-04-18, and CAMS-GLOB-ANTv6.2. CV is shown only for grid cells where mean emissions exceed a minimum threshold ($0.034e-10 \text{ kg m}^{-2} \text{ s}^{-1}$).



571

572 **4.2 Biomass burning emissions**

573 We compare our top-down biomass burning emissions with GFASv1.2 (Kaiser et al., 2012), GFED4 (Giglio et al., 2013), the
574 Fire INventory from NCAR version 2.7 (FINNv2.7) (Wiedinmyer et al., 2023; Maneenoi et al., submitted), and the Quick Fire
575 Emissions Dataset version 2.6 R1 (QFEDv2.6r1) (Koster et al., 2015). GFASv1.2 assimilates Moderate Resolution Imaging
576 Spectroradiometer (MODIS) fire radiative power (FRP) observations to estimate dry matter combustion rates and applies
577 biome-specific emission factors to derive trace gas and aerosol fluxes. GFED4 combines MODIS burned area maps with active
578 fire data from the Tropical Rainfall Measuring Mission (TRMM) Visible and Infrared Scanner (VIRS) and the Along-Track
579 Scanning Radiometer (ATSR) family of sensors. FINNv2.7 combines Visible Infrared Imaging Radiometer Suite (VIIRS) and
580 Himawari-9 Advanced Himawari Imager (AHI) active fire detections with land cover data, emission factors from the
581 NEIVAv1.0 database (Binte Shahid et al., 2024), and fuel loadings to produce daily emissions. Compared to FINNv2.5,
582 FINNv2.7 applies a corrected Himawari pixel size consistent with VIIRS and uses a proximity filter to remove duplicated and
583 nearby detections, leading to reduced emissions overall (Maneenoi et al., submitted; Fig. S5). QFEDv2.6r1 derives emissions
584 from FRP observations using biome-specific coefficients optimized to reproduce observed aerosol optical depth patterns.
585

586 The largest emitters are Laos, Myanmar, and India, followed by mainland China and Cambodia (Fig. 11). Substantial
587 differences are evident among the three inventories. FINNv2.7 yields CO emissions larger by a factor of 2–5 across Southeast
588 Asia relative to GFASv1.2, GFED4, and QFEDv2.6r1. Also, FINNv2.7 derives emissions from satellite fire detections
589 including VIIRS and hourly Himawari-9 sensors, providing greater sensitivity to small, fragmented, and local evening crop
590 burning that are common in Southeast Asia (Liu et al., 2020; Marvin et al., 2024; Maneenoi et al., submitted). In contrast, FRP
591 based inventories that rely on biome dependent or globally uniform conversion factors (GFAS and QFED) tend to produce
592 lower emissions, particularly where agricultural burning dominates, cloud and haze obscure active fires, or combustion occurs
593 at low intensity.
594

595 For the top-down CO emission estimates, we use two different configurations: one based on MOPITT CO assimilation with
596 daily GFASv1.2 emissions as the prior, and the other based on TROPOMI CO assimilation with monthly GFED4 emissions
597 as the prior. Because both the assimilated observations and the prior emissions influence the inferred emissions, dedicated
598 sensitivity experiments in which these factors are varied separately would be required to isolate their individual roles. In the
599 present study, these two configurations are used to characterize the range of uncertainty in the estimated emissions arising
600 from different assimilation settings (Fig. S6). The inversion increases the biomass burning fraction of total CO from 14.4%
601 (GFASv1.2 prior) to 24.0% for TCR-3 (MOPITT) and from 17.3% (GFED4 prior) to 36.9% for TCR-3 (TROPOMI). TCR-3
602 (TROPOMI) yields biomass burning CO emissions roughly twice those of TCR-3 (MOPITT) over Asia (31 vs 15 Tg), despite
603 a 15% lower prior, indicating that TROPOMI detects substantially larger fire-related CO enhancements (Fig. S7). Across the
604 major biomass burning regions, TCR-3 (MOPITT) generally exceeds GFASv1.2 by a factor of 1.5–3 (excepts over Indonesia).
605 TCR-3 (TROPOMI) substantially exceeds GFED4 across all major regions by factors of 2–8. FINNv2.7 exceeds TCR-3
606 (MOPITT) by a factor of 1.3–5, but agrees closely with TCR-3 (TROPOMI) across the major burning regions.



607

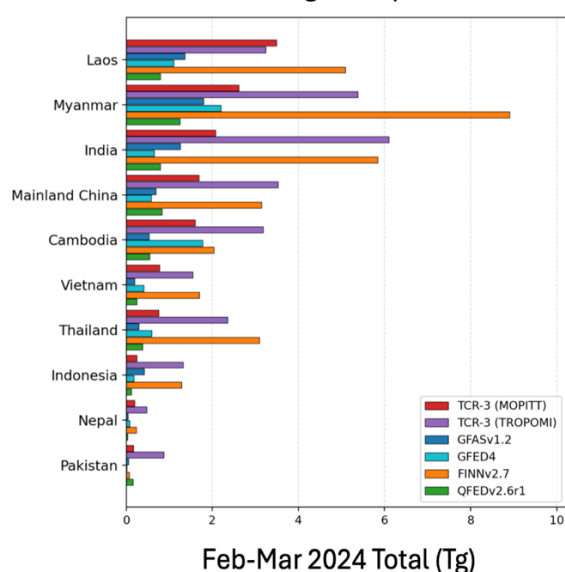
608 Thailand experienced the most severe AQ conditions during the ASIA-AQ period, driven by intense biomass burning in the
 609 north. Figure 12 shows daily CO fire emissions for two subregions: northern Thailand and surrounding regions (NTHA, above
 610 16°N) and southern Thailand and surrounding regions (STHA, below 16°N). In the NTHA region, emissions peak in early to
 611 mid-March, coinciding with the ASIA-AQ observation period (March 13–27, shaded). The STHA region shows much lower
 612 fire emissions across all inventories throughout the period.

613

614 TCR-3 (MOPITT) and TCR-3 (TROPOMI) increase biomass burning CO emissions over the NTHA and STHA domains
 615 relative to their respective prior inventories, from 3.6 (GFASv1.2) to 8.3 Tg and 5.6 (GFED4) to 14.5 Tg, respectively, roughly
 616 a factor of 2–3. FINNV2.7 emissions are substantially higher than both TCR-3 products at 17.9 Tg. The biomass burning
 617 fraction of total CO emissions increases substantially with assimilation in both configurations (Fig. S7), with TCR-3
 618 (TROPOMI) showing particularly large enhancements concentrated in North and West Thailand. TCR-3 (MOPITT) captures
 619 the day-to-day evolution of fire activity, increasing emissions relative to GFASv1.2. TCR-3 (TROPOMI), by contrast, uses
 620 monthly GFED4 as its prior and therefore does not resolve day-to-day fire variability. The assimilated TROPOMI CO column
 621 nevertheless produces general enhancements in emissions during early-to-mid March, particularly over NTHA. The
 622 contrasting emission estimates between the two TCR-3 configurations are reflected in their surface CO performance over
 623 Thailand, where TCR-3 (MOPITT) reduces the surface negative bias while TCR-3 (TROPOMI) overestimates surface
 624 concentrations despite better capturing elevated CO aloft, as discussed in Section 3.1.

625

Biomass burning CO top-10 countries



626

627 **Figure 11.** Total CO emissions (Tg) from the ten highest-emitting regions within our study domain (Fig. 6) where the order
 628 follows TCR-3 (MOPITT). Bars represent six different emission inventories: TCR-3 (MOPITT) (red), TCR-3 (TROPOMI)
 629 (purple), GFASv1.2 (blue), GFED4 (cyan), FINNV2.7 (orange), and QFEDv2.6r1 (green).

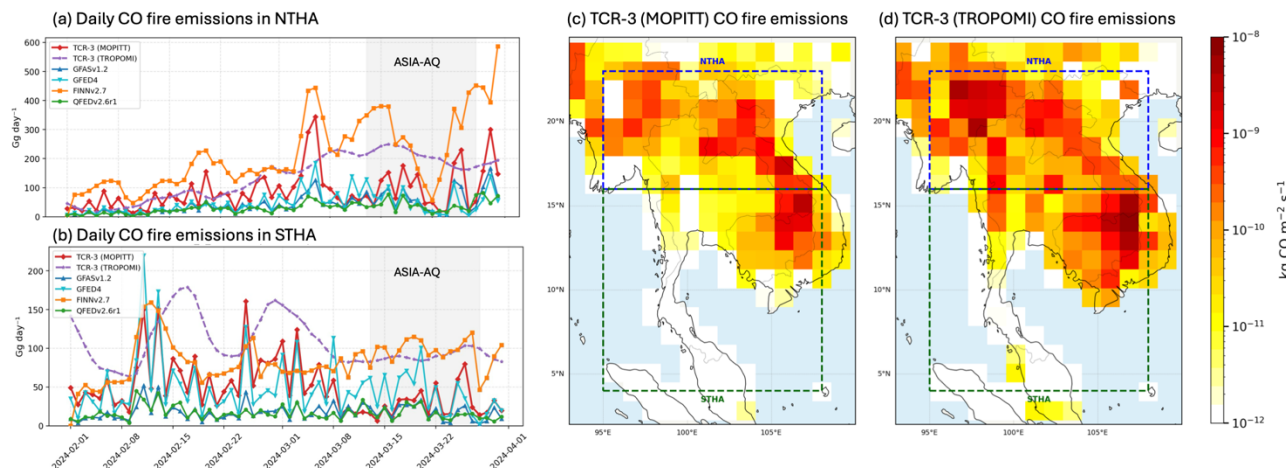


Figure 12. Daily CO fire emissions (Gg/day) in (a) northern Thailand and surrounding regions (NTHA; north of 16°N), and (b) southern Thailand and surrounding regions (STHA; south of 16°N) during February–March 2024. The shaded grey period (March 13–27) indicates the ASIA-AQ observation period in Thailand. Lines represent four biomass-burning emission inventories: TCR-3 (MOPITT) (red), TCR-3 (TROPOMI) (purple), GFASv1.2 (blue), GFED4 (cyan), FINNv2.7 (orange), and QFEDv2.6r1 (green). Spatial distribution of (c) TCR-3 (MOPITT) and (d) TCR-3 (TROPOMI) CO fire emissions ($\text{kg CO m}^{-2} \text{s}^{-1}$) averaged over February–March 2024, with the NTHA (blue box) and STHA (green box) subregions marked.

5. Local and transboundary transport

To understand the sources contributing to O_3 pollution across the ASIA-AQ regions and to quantify the influence of transboundary transport, we conduct a series of emission perturbation experiments for CO, NO_x , and isoprene. These experiments employ two scenarios: a 20% emission reduction to evaluate the surface O_3 response to near-future mitigation strategies, and a 100% emission reduction (zero-out) to maximize the signal of specific sources. Perturbations are applied individually to ten source regions (Fig. S8): South Korea, Philippines, Taiwan, Thailand, mainland China, India, Myanmar, Laos, Vietnam, and Rest of Southeast Asia. While assessing the impacts of specific policy measures would require more targeted and sector-resolved sensitivity experiments, these perturbations are designed to provide a first-order, domain-wide quantification of the relative importance of regional source contributions.

5.1 Surface O_3

Figure 13 shows the daily surface O_3 response in five receptor cities (Seoul, Chiang Mai, Taipei, Manila, and Bangkok) to 20% emission reductions applied separately to each source region, representing a near-future mitigation scenario. The 100% emission reduction (zero-out) results are shown in Fig. S9 for reference. A comparison between the two experiments reveals substantial nonlinearity in the O_3 response, particularly over Seoul and Taipei, where the 100% response is not proportional to



653 the 20% response. For example, over Seoul on February 11, removing 100% of mainland China emissions produces an O₃
654 response of ~-30 ppbv, whereas a 20% reduction yields only ~-2 ppbv, far less than one-fifth of the zero-out response,
655 reflecting NO_x-saturated photochemical conditions and the nonlinear sensitivity of O₃ to precursor emissions. In general,
656 foreign emission contributions appear larger in the 100% sensitivity simulation, showing that the 20% reduction results
657 presented here provide a conservative and policy-relevant estimate of transboundary influences under realistic mitigation
658 scenarios.

659

660 Over Manila, the relative importance of transported pollution varies strongly with meteorological regimes. The DC-8 flight
661 dates over the Philippines captured two distinct regimes. The first two flights (February 6–7) occurred under a weak Northeast
662 (NE) Monsoon, when local emissions dominate O₃ formation, The latter two flights (February 11, and 13) coincided with an
663 enhanced NE Monsoon (northeasterly) that enabled substantial long-range transport, accounting for up to 62% of the total O₃
664 response, with the largest contribution from mainland China (18%). This contrast underscores the controlling role of synoptic-
665 scale circulation in modulating O₃ over Manila.

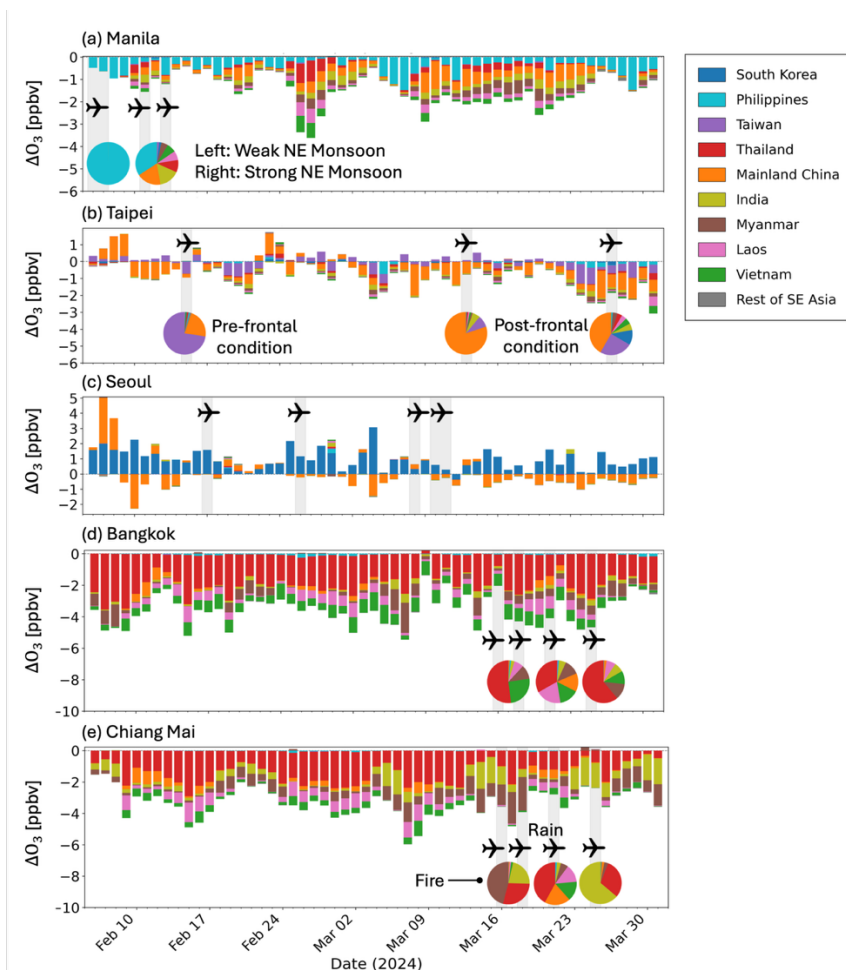
666

667 Taipei and Seoul show O₃ increases and decreases in response to emission reductions. Over Taipei, the DC-8 flight dates
668 corresponded to contrasting meteorological conditions: pre-frontal conditions on February 15, when local emissions dominated
669 O₃ formation, versus post-frontal northwesterly flow on March 13 and 27, when transport from mainland China dominated.
670 Over Seoul, O₃ increases in response to local emission reductions throughout the period, reflecting NO_x-saturated conditions.
671 Seoul's response to Chinese emissions varies according to wind directions, typically showing decreases when transport
672 originates from the North China Plain but increases when influenced by less polluted source regions.

673

674 Bangkok and Chiang Mai are strongly influenced by local emissions, but exhibit distinct transboundary signals. Over Bangkok,
675 transport from Vietnam becomes important during southerly maritime flow. In contrast, Chiang Mai is primarily influenced
676 by inflow from Myanmar, Laos, and India. Chiang Mai experiences the most severe biomass burning impacts during periods
677 flagged for enhanced combustion and smoke (Figure 2). During these fire episodes, Myanmar becomes the dominant external
678 source, with contributions exceeding those from local Thailand emissions. During March 21–23, a mid-latitude synoptic
679 system over northern Thailand brought more diverse foreign influences, including mainland China.

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Figure 13. Regional contributions to daily surface O₃ concentrations in five cities based on 20% combined emission reduction sensitivity experiments. Panels show results for (a) Manila, (b) Taipei, (c) Seoul, (d) Bangkok, and (e) Chiang Mai. Stacked bars represent the daily change in surface O₃ (ΔO_3 , ppbv) when CO, NO_x, and isoprene emissions from each source region are simultaneously reduced by 20%. Colors indicate different source regions: South Korea (blue), Philippines (teal), Taiwan (purple), Thailand (red), mainland China (orange), India (olive), Myanmar (brown), Laos (pink), Vietnam (green), and Rest of Southeast Asia (grey). Negative values indicate O₃ decreases resulting from emission reductions, while positive values indicate nonlinear O₃ responses to precursor emission changes. Grey shading with airplane icons marks DC-8 flight dates during ASIA-AQ. Pie charts show the relative contributions from each region during selected flight periods: weak versus strong northeast monsoon conditions in Manila (a), pre-frontal versus post-frontal conditions in Taipei (b), and biomass burning versus cleaner periods in Bangkok and Chiang Mai (d-e). No pie chart is shown for Seoul (c) due to the nonlinear response where local emission reductions (South Korea) lead to O₃ increases rather than decreases.



694 5.2 Influences from India

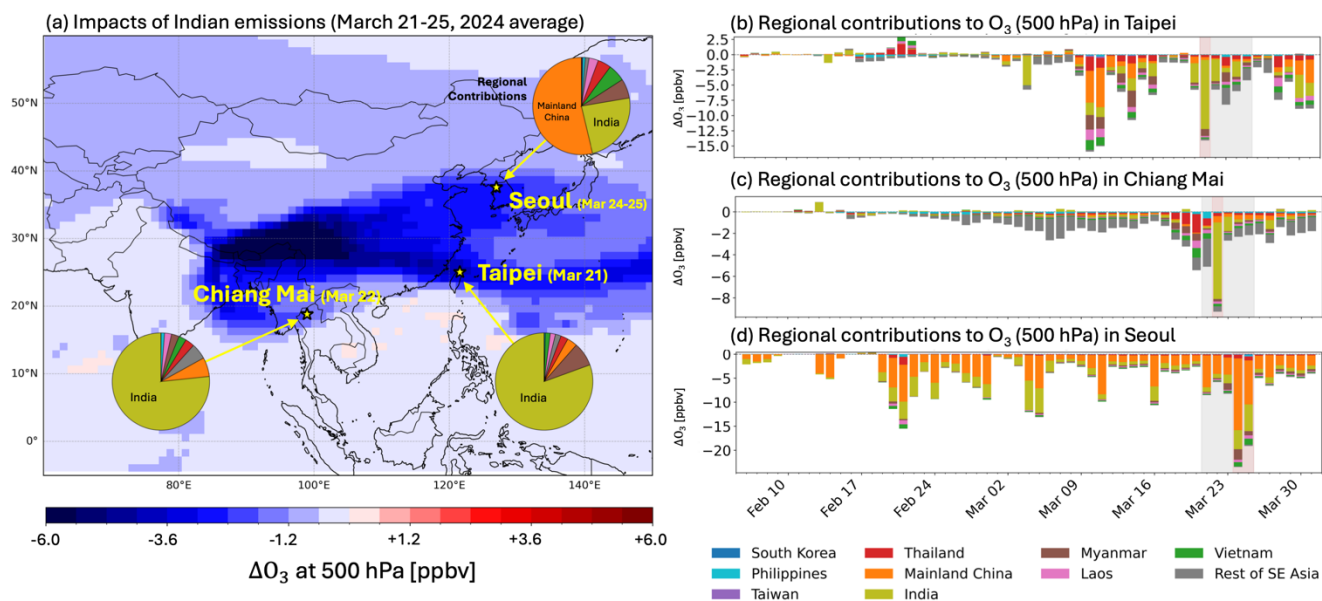
695 Although India lies outside the ASIA-AQ domain, rapid economic growth and associated emission increases raise concerns
696 for both domestic air quality and downwind regions (Ding et al., 2022; Guizzardi et al., 2025; Kurokawa and Ohara, 2020;
697 Soulie et al., 2024). Despite national mitigation efforts such as the National Clean Air Programme (NCAP) (Guttikunda et al.,
698 2023), substantial health burdens and high levels of O₃ and its precursors remain (Gopikrishnan et al., 2025; Gopikrishnan and
699 Kuttippurath, 2024). Beyond domestic impacts, Indian emissions can influence regional atmospheric composition through
700 frequent deep convection, which lofts pollutants into the free troposphere and facilitates long-range transport toward Southeast
701 and East Asia (Lawrence and Lelieveld, 2010; Lelieveld et al., 2001).

702
703 While Indian emissions play a relatively minor role in surface O₃ across most of the ASIA-AQ domain, their influence becomes
704 substantial aloft. Figure 14 illustrates the Indian emission contribution to mid-tropospheric O₃ at 500 hPa, showing the O₃
705 response to 100% emission reductions averaged over March 21–25, 2024, when the Indian influence was particularly strong.
706 The response extends from the Indian subcontinent across the Bay of Bengal into Southeast Asia and southern China, reaching
707 as far as Seoul. The spatial pattern suggests that polluted air is being lofted by deep convection over Northeast India and the
708 Himalayan foothills and subsequently transported eastward in the mid-troposphere.

709
710 Downwind O₃ enhancements may reflect both O₃ produced over India and the export of precursors, such as peroxyacetyl
711 nitrate (PAN), which releases NO_x and promotes O₃ production during transport. Observations of PAN (Payne et al., 2022a)
712 from the Cross-Track Infrared Sounder (CrIS) (Shephard and Cady-Pereira, 2015) reveal elevated concentrations along the
713 simulated transport pathway (Fig. S10). These elevated PAN levels provide independent evidence for the efficient long-range
714 export of both O₃ and its precursors. The thermal stability and persistence of PAN under cold upper-tropospheric conditions
715 facilitate its transport over long distances, potentially extending to trans-Pacific scales where subsequent subsidence and
716 warming release NO_x to produce O₃ in downwind receptor regions.

717
718 The time series for Taipei, Chiang Mai, and Seoul suggest that the Indian emission contribution to 500 hPa O₃ peaks during
719 March 21–25, 2024. At peak influence, Indian emissions account for 78%, 72%, and 24% of the total O₃ response at 500 hPa
720 in Taipei (March 21), Chiang Mai (March 25), and Seoul (March 24–25), respectively. These large contributions at considerable
721 distances from the source underscore the efficiency of convective lofting and mid-tropospheric transport in delivering Indian
722 emissions to remote receptors. Although the surface O₃ impact across the ASIA-AQ domain is generally limited, Indian
723 emissions can make a substantial contribution to mid-tropospheric O₃ during active convective period, with potential surface
724 impacts through vertical mixing during convective events or frontal passages.

725



726
 727 **Figure 14.** Impact of Indian emissions on mid-tropospheric O₃ (500 hPa) over East and Southeast Asia. (a) Spatial distribution
 728 of ΔO₃ at 500 hPa averaged over March 21–25, 2024, resulting from 100% reduction of Indian CO, NO_x, and isoprene
 729 emissions. Yellow stars mark three cities where Indian impacts are examined: Seoul, Taipei, and Chiang Mai. Pie charts show
 730 the relative regional contributions to 500 hPa O₃ during peak Indian influence periods (dates in parentheses). (b-d) Time series
 731 of daily changes in 500 hPa O₃ (ΔO₃, ppbv) from 100% combined emission reductions in each source region for (b) Taipei,
 732 (c) Chiang Mai, and (d) Seoul. Stacked bars show contributions from different regions using the same color scheme as Fig.
 733 13. Grey shading indicates March 21–25, when Indian impacts were widespread across the domain. Red shading highlights
 734 peak Indian influence periods at each city, corresponding to the dates shown in panel (a).

735 **6. Conclusions**

736 To advance understanding of air pollutants drivers across East and Southeast Asia, we utilize the TCR-3 chemical reanalysis,
 737 produced using the MOMO-Chem data assimilation framework, which integrates multiple satellite observations during the
 738 ASIA-AQ campaign (February–March 2024). The comprehensive ASIA-AQ measurements provide a unique opportunity to
 739 evaluate chemical reanalysis performance, assess bottom-up emission inventories, and quantify the value of satellite observing
 740 systems, aimed at supporting air quality management in Asia.

741
 742 Evaluations against DC-8 aircraft measurements demonstrates that the data assimilation improves agreement for O₃, NO_x, CO,
 743 and CH₂O across the campaign domain. The largest gains are for CO and CH₂O. CO correlations increase from 0.63–0.64 to
 744 0.77, while NMB is reduced from –32–33% to –15–16% across the domain. The most substantial regional improvement occurs



745 over Thailand, where the correlation increases from 0.07 to 0.50 below 1 km and surface NMB improves from -28% to -8%,
746 reflecting improved representation of episodic biomass burning enhancements. Modest but consistent improvements are also
747 evident over the Philippines and Taiwan. For CH₂O, the positive model biases are reduced from 55–70% to 17–29% across
748 the domain, with the largest corrections over the Philippines (NMB from +74% to +22% below 1 km), Taiwan (+109% to
749 +48%), and Thailand (+87% to +43%), consistent with the reduced isoprene emission by data assimilation.

750
751 The optimized emissions emission estimates reveal substantial discrepancies with existing bottom-up inventories. TCR-3 NO_x
752 exceeds both global (EDGAR, CEDS, and CAMS-GLOB-ANT) and regional (HTAP, AQNEA) inventories by 30–52% over
753 the Philippines and lies at the upper end of bottom-up estimates over mainland China, suggesting potential underestimation of
754 transport emissions in bottom-up inventories. For anthropogenic CO emissions, EDGAR and CAMS-GLOB-ANT estimates
755 over South Korea and Taiwan are 2–4 times higher than both TCR-3 and regional inventories, where other inventories lack
756 large industrial sector contributions. In India, TCR-3 anthropogenic CO exceeds all bottom-up inventories, aligning with the
757 extrapolated upper range of HTAPv3.2, suggesting continued industry-driven emission growth beyond recent inventory
758 updates.

759
760 Biomass burning emissions show even larger discrepancies. Northern Thailand experienced the most severe air quality
761 conditions during the ASIA-AQ period, with biomass burning CO emissions reaching 5.3 and 8.1 Tg for TCR-3 (MOPITT)
762 and TCR-3 (TROPOMI), respectively, out of regional totals of 15.0 and 30.5 Tg across Northeast and Southeast Asia. Across
763 the broader region, bottom-up inventories range from 5.8 Tg (QFEDv2.6r1) to 34.0 Tg (FINNv2.7), highlighting persistent
764 uncertainties in fire emission estimates. The 53% domain-wide reduction in isoprene emissions inferred by TCR-3 support
765 earlier findings that MEGAN overestimates biogenic fluxes in tropical Southeast Asia (Choi et al., 2025; Sindelarova et al.,
766 2014), with important implications for O₃ formation. Overall, the substantial emission adjustments indicated by TCR-3
767 underscore the limitations of bottom-up inventories in capturing rapidly evolving emission patterns across Asia.

768
769 In addition to local emissions, transboundary transport plays a highly variable and often dominant role in shaping air quality
770 across the ASIA-AQ domain. Over Manila, long-range transport accounts for up to 62% of the surface O₃ response during
771 enhanced northeast monsoon conditions (Feb 11), with mainland China as the largest single foreign contributor (~18%). Over
772 Bangkok, Vietnamese emissions become important during southerly maritime flow, while Chiang Mai is primarily influenced
773 by inflow from Myanmar, Laos, and India, with Myanmar emerging as the dominant external source during intense fire
774 episodes in March, exceeding even local Thailand contributions. Notably, Indian emissions exert a substantial influence in the
775 free troposphere, contributing 78%, 72%, and 24% of the O₃ response over Taipei, Chiang Mai, and Seoul during March 21–
776 25, likely reflecting deep convective lofting of precursors over the Indian subcontinent and subsequent eastward transport.
777 These results highlight the importance of three-dimensional transport pathways and cross-border precursor contributions,
778 emphasizing the need for coordinated international air quality management strategies.



779

780 This study demonstrates the value of integrating multi-platform satellite observations with aircraft campaigns like ASIA-AQ
781 to address three key scientific challenges: sparse ground-based monitoring, uncertain bottom-up inventories, and complex
782 transboundary transport. Emerging satellite measurements provide significant opportunities for future improvements.
783 Geostationary instruments such as GEMS enable hourly observations of NO₂ and CH₂O over Asia (Kwon et al., 2019; Lee et
784 al., 2024; Seo et al., 2024), while Himawari-9 provides high-resolution aerosol data; together, these could further constrain the
785 diurnal variability in emissions and photochemistry that our analysis has shown to be important (Yang et al., 2024). CrIS
786 retrievals of O₃ (Boynard et al., 2025), CO (Fu et al., 2016; Worden et al., 2022), PAN (Payne et al., 2022b), ammonia (NH₃)
787 (Cady-Pereira et al., 2024), and isoprene (Fu et al., 2019; Wells et al., 2020, 2022), and other NMVOCs (Brewer et al., 2024;
788 Wells et al., 2025) can further improve emission estimates and chemical processes. In particular, NH₃, a key precursor of
789 inorganic aerosols whose emissions remain highly uncertain (Cao et al., 2020), is not addressed in this study. Fully exploiting
790 these observations will require continued advances in data assimilation methods and model process representation.

791

792 **Data availability**

793 Tropospheric Chemistry Reanalysis are available at <https://doi.org/10.25966/9qgv-fe81> (Miyazaki et al., 2019) and on the
794 NASA GES DISC website ([https://disc.gsfc.nasa.gov/information/data-
795 release?title=Release%20of%20TROPES%20Chemical%20Reanalysis%20Products](https://disc.gsfc.nasa.gov/information/data-release?title=Release%20of%20TROPES%20Chemical%20Reanalysis%20Products)). ASIA-AQ DC-8 airborne
796 observations are open access and available in the ASIA-AQ Data Archive ([https://doi.org/10.5067/SUBORBITAL/ASIA-
797 AQ/DATA001](https://doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001), ASIA-AQ Science Team, 2024).

798

799 **Author contributions**

800 Conceptualization: JC and KM. Methodology & Software: JC, KM, and TS. Investigation: JC, KM, TS, RB, AR, LKn, GSD,
801 JAM, JPD, YC, AF, CC, ECA, JMSc, RAH, GMW, ED, AS, and DRB. Formal Analysis & Visualization: JC. Resources &
802 Data Curation: JC, KM, TS, HE, PR, LKE, NM, MC, RMH, SJS, JHW, and LKu. Writing – Original Draft: JC. Writing –
803 Review & Editing: All coauthors.

804



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816

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