



1 **Interaction between the East Asian summer monsoon and Westerlies**
2 **as shown by desert shrub-ring records from the Alxa Plateau,**
3 **Northwest China**

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15



16 **Abstract:**

17 The variations, driving mechanisms, and interactions between the mid-latitude westerlies
18 and the Asian summer monsoon represent a critical focus in global change research. Extensive
19 research has investigated climate change and its drivers across multiple timescales in the core
20 regions of both westerlies Asia and monsoonal Asia. However, studies are far fewer in the arid
21 inland deserts at the interactive margin of these two circulation systems, where precipitation
22 restrict the distribution of trees. To address this gap, we developed 26 shrub-ring width
23 chronologies of *Zygophyllum xanthoxylum* from four sub-regions of the Alxa Desert Plateau.
24 Our analysis reveals that shrub radial growth is primarily limited by aridity during the growing
25 season, driven by low precipitation. The four regional chronologies indicate substantial
26 spatiotemporal heterogeneity in dry–wet variability. By integrating these records with indices
27 of the westerlies and the East Asian summer monsoon (EASM), we demonstrate that
28 hydroclimatic variations are primarily modulated by the interaction between these two
29 circulations. The Alxa Plateau is predominantly influenced by the EASM, whose boundary can
30 extend approximately 400 km northwestward to the Yabulai–Nurgong–Lang mountains line in
31 the central plateau. This can cover the entire Tengger and Ulan Buh Deserts. Concurrently, the
32 westerlies can advance eastward to affect the Helan Mountains. Thus, the Alxa Plateau
33 constitutes a key transition zone and interaction arena between the westerlies and the EASM
34 within China’s arid interior. At the same time, the foehn effect in the mountains around the Alxa
35 Plateau also has a certain impact on atmospheric circulation. Climatic fluctuations in this region
36 critically influence local vegetation dynamics and desertification processes, underscoring its
37 importance for understanding past and future environmental change.

38

39 **Key words:** Alxa Plateau; Shrub-dendroclimatology; Interaction between Westerlies and East
40 Asian summer monsoon; *Zygophyllum xanthoxylum*.

41



42 **1. Introduction**

43 Two atmospheric circulation systems, the mid-latitude Westerlies and the Asian summer
44 monsoon, play key roles in the northern hemisphere climatic changes (An et al., 2012). The
45 eastern part of the arid and semi-arid region of Central and East Asia lies on the margin of the
46 East Asian monsoon domain and is influenced by both the westerly and monsoon circulations
47 (Wang, 2006; Liu et al., 2018) (Figure 1). From the perspective of large-scale atmospheric
48 systems, Chen et al. (2009) divided the Asian continent into two major climatic domains: the
49 “Westerlies Asia”, primarily controlled by the mid-latitude westerlies, and the “Monsoonal
50 Asia”, primarily governed by the monsoon circulation.

51 The arid northwestern China, located in the interior of the Eurasian continent, is one of the
52 driest areas globally at the same latitude. It exhibits the typical characteristics of a continental
53 climate, with a westerly circulation pattern. While primarily influenced by the westerlies, this
54 region is also affected by the East Asian monsoon. Precipitation decreases progressively from
55 the eastern and western edges towards the central arid and semi-arid zone of Eurasia,
56 accompanied by high variability and frequent droughts. Consequently, this area is considered
57 both sensitive to and ecologically vulnerable to global climate change (Chen et al., 2019a;
58 2019b; Zhang et al., 2023).

59 The East Asian summer monsoon (EASM) is a circulation system that creates distinct
60 differences in moisture levels (Tang et al., 2006a; Tang et al., 2006b). The fluctuating northern
61 boundary of the monsoon is known as the northern marginal zone (Xu and Qian, 2003; Hu and
62 Qian, 2007). According to Chen et al. (2018), the 300 mm isohyet is used as a climatic indicator
63 to define the northern boundary of the EASM in China. In northwestern China, this climatic
64 boundary roughly coincides with the Qilian, Helan and Daqing mountains.

65 The westerly zone is a vital link between the North Atlantic climatic region and the East
66 Asian monsoon region (Qu et al., 2004). Variations in the westerlies directly affect moisture
67 transport; a significant amount of the moisture in northwestern China in summer originates
68 from the interaction between the westerlies and the monsoon (Feng et al., 2004; Wang et al.,
69 2005; Li et al., 2008; Ma et al., 2011; Luo et al., 2026). Moreover, dust storm activity in
70 northern China is closely linked to changes in the location and strength of the mid-latitude
71 westerly jet (Zhong and Li, 2005; Li and Liu, 2015).



72 Against the background of global atmospheric circulation changes, the advance, retreat,
73 positional shifts, and interactions of the EASM and westerly circulation are inevitably affected.
74 These changes further influence climate variability in northwestern China, consequently
75 affecting ecological and environmental dynamics. Research on the climate and environment in
76 this region is of significant theoretical and practical importance, making it a key issue in global
77 change studies (Chen et al., 2019a; Chen et al., 2019b).

78 Dendroclimatology provides high-resolution, annually resolved proxy data on centennial
79 to millennial scales, which are invaluable for reconstructing monsoon and westerly circulation
80 histories, and for investigating the driving mechanisms of climate change (Liu et al., 2001; Liu
81 and Cai, 2003; Liu et al., 2008; Chen et al., 2013; Kang and Yang, 2015). However, previous
82 research has predominantly focused on the core zones influenced by the westerlies or the
83 monsoon. In these regions, dendroclimatic studies and investigations into climate forcing
84 mechanisms have largely been based on individual sampling sites. There remains a scarcity of
85 comprehensive, multi-site, and regional-scale research that systematically examines the
86 linkages with these major circulation systems. This gap is particularly pronounced in the
87 transition zone where the westerlies and the East Asian summer monsoon interact (Xiao et al.,
88 2024a;b).

89 Xiao et al. (2024a) conducted a preliminary investigation into the interaction mechanisms
90 between the EASM and the westerlies, using tree-ring data from *Picea crassifolia* at three sites
91 in the mountainous areas surrounding the arid Alxa Plateau in northwestern China. However,
92 research on the core interaction zone between these two circulation systems—the interior desert
93 region of the Alxa Plateau—remains notably scarce. This is largely because aridity limits the
94 distribution of tree species across the plateau, meaning that dendroclimatology studies here are
95 primarily based on shrubs.

96 This study uses tree-ring width chronologies from 26 *Zygophyllum xanthoxylum* sampling
97 sites, a dominant shrub species in the Alxa Plateau desert ecosystem, to analyze the climatic
98 response characteristics of its radial growth across four distinct regions. By integrating the
99 Westerly and EASM indices, we further investigate how these two major circulation systems
100 and their interactions drive regional climate variability. The study aims to clarify the
101 spatiotemporal heterogeneity of climate change in the desert and the surrounding mountains of



102 Alxa Desert Plateau, and to elucidate the interaction mechanisms between the two circulation
103 systems. This will provide a theoretical basis for climate change research and desertification
104 control in arid regions.

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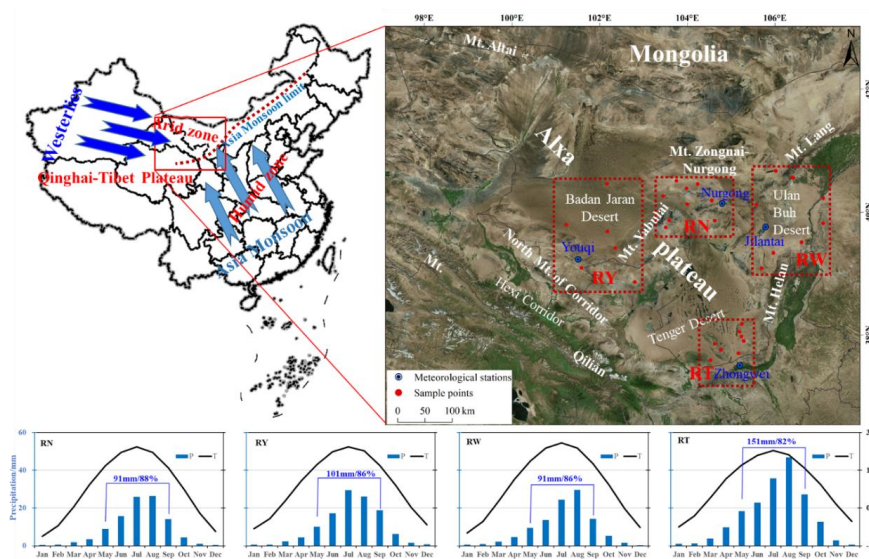
106 **2. Study area, Material and methods**

107 **2.1 Study area**

108 The Alxa Plateau is located to the north of the Hexi Corridor in China's Gansu Province,
109 and west of the Inner Mongolia Plateau, and south of the Gobi Altai Mountain (Figure 1).
110 Three deserts distribute in this area, including of the Tengger Desert, the Ulan Buh Desert, and
111 the Badain Jaran Desert. Several mountain ranges surround and span the Alxa Plateau. These
112 include the Helan Mountains to the east, the northern mountains of the Hexi Corridor to the
113 south, the outliers of the Altai Mountains to the north, and a series of denuded lower mountain
114 ranges of the Yabulai, Nurgong, Zongnai and Lang mountains in the middle. These mountains
115 not only block the desert's eastward and southward expansion (driven by high-pressure regions
116 from Mongolia), but also form the headwaters for rivers and streams that water the oases on the
117 plateau (Xiao et al. 2024a) (Figure 1).

118 The Alxa Plateau is located in the eastern margin of the inland, arid region of Central Asia.
119 Many studies showed that this area is the main interaction zone of the mid-latitude westerly
120 circulation and the Asian monsoon (Xiao et al., 2017; Chen et al., 2019b; Xiao et al., 2024a;b).
121 Its pronounced interannual variability effects on the vegetation cover changes (Ou and Qian,
122 2006; Tang et al., 2006a; 2006b; Li et al., 2013; Ren et al. 2026).

123 The Alxa Plateau has a typical continental climate. We have divided the study area into
124 four regions based on geographical distribution characteristics of the three major deserts and
125 the sampling points (Figure 1): the Yabulai-Nurgong low mountain area (RN), the Ulan Buh
126 Desert area (RW), the Badain Jaran Desert area (RY), and the Tengger Desert area (RT).



127

128 Figure 1. Location of tree-ring sampling sites (red points) and a climatic diagram (1955 –
 129 2017) of the study area in RY (Youqi meteorological station,1960-2017), RN (Nurgong
 130 meteorological station,1957-2017), RW (Jilantai meteorological station,1955-2017) and
 131 RT(Zhongwei meteorological station,1959-2017). The blue box and appended data in the
 132 climatic diagram indicate the total growing season (from May to September) precipitation in
 133 the study area and the proportion of total annual precipitation. The red dashed boxes indicate
 134 the four regions in the study area. This figure is modified after Xiao et al. (2024a).

135 Meteorological data for the study regions RN, RY, RW, and RT were obtained from the
 136 nearest stations at Nurgong, Youqi, Jilantai, and Zhongwei, respectively. The annual mean
 137 temperatures at these stations are 7.4 °C , 8.9 °C , 8.8 °C , and 9.0 °C , respectively
 138 (<http://data.cma.cn>). The multi-year average annual precipitation is 103.0 mm, 118.0 mm, 106.9
 139 mm, and 182.9 mm, respectively. Notably, 82 – 88% of the precipitation concentrated from
 140 May to September (Figure 1).

141 *Z. xanthoxylum* is a super-xerophilous and soil-stabilizing species endemic to Central
 142 Asia's deserts, and often serves as a dominant or constructive species in the gravelly riparian
 143 terraces, gently sloping hills, and pre-hill flood fans in the Alxa Desert (Xiao et al., 2012; Chen
 144 et al., 2016). *Z. xanthoxylum* has well-defined annual growth rings, making it suitable for tree-
 145 ring study (Xiao et al., 2019; Ding et al., 2021). Recently, shrub-ring data of this species have



146 been used to conduct climate reconstruction and ecological response studies in the Badain Jaran,
147 Tengger and Ulan Buh deserts, providing a new perspective for studying climate and
148 environmental changes in arid regions (Xiao et al., 2012; Xiao et al., 2019; Ding et al., 2021;
149 Peng et al., 2024; Ding et al., 2025).

150

151 **2.2 Data collection, processing, and analysis**

152 **2.2.1 Sample collection, processing, and chronology construction**

153 A total of 26 sampling sites of the shrub *Z. xanthoxylum* were collected from the Alxa
154 Plateau between 2007 and 2024 (Figure 1). The Supplementary Table 1 (S1) provides detailed
155 information on the location, sample size, chronology length of ring width, mean sensitivity,
156 credible chronology interval with a subsample signal strength (SSS) >0.85, and signal-to-noise
157 ratio for each chronology.

158 For each sampling sites, shrub disc treatment, cross-dating and ring width measurement
159 were carried out according to the standard method of dendrochronology, and the standard
160 chronologies was finally constructed ((Holmes, 1983; Cook, 1985; Xiao et al. 2019; Ding et
161 al., 2025)) .

162 A regional tree-ring chronology (Figure 2) was generated by calculating the weighted
163 mean of the standardized chronology indices from all sample sites within the region. The weight
164 for each site in a given year was determined by the proportion of its index value relative to the
165 sum of the indices from all sites for that given year. This method of constructing the chronology
166 can help enhance the strength of low-frequency signals to a certain extent.

167 **2.2.2 Atmospheric circulation indices and correlation analysis**

168 The East Asian summer monsoon index (EASMI defined by Li and Zeng (2005) for the
169 period 1950–2017) and the westerly circulation index (WCI annual mean, period 1951 – 2015)
170 (<https://cmdp.ncc-cma.net/cn/index.htm>) were same as Xiao et al.(2024a).

171 Response-function analysis was used to quantify the relationships between tree-ring
172 chronologies and climate variables, using the DENDROCLIM2002 software (Biondi and
173 Waikul, 2004). The climate variables included the total monthly precipitation, mean monthly
174 temperature measured at the nearest meteorological station in study area (Figure 1), and the
175 standardized precipitation evapotranspiration index (SPEI) over a 13-month period (from



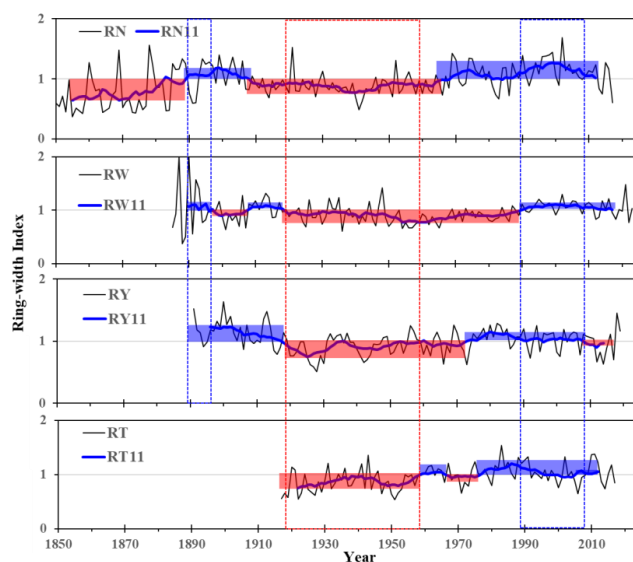
176 September of the previous year to September of the current year). The SPEI was calculated
177 based on observational data, using the SPEI package ([https://CRAN.R-](https://CRAN.R-project.org/package=SPEI)
178 [project.org/package=SPEI](https://CRAN.R-project.org/package=SPEI)) in R software, and used the FAO-56 Penman-Monteith equation for
179 potential evapotranspiration (Beguería et al., 2014).

180 Interannual and interdecadal (sliding moving average of 11 years) chrono-climatic and
181 chrono-cyclonic indices correlation and partial correlation analyses were performed using SPSS
182 19.0. Based on the characteristics of the two circulation indices, the sequences were classified
183 into three groups (low, medium, and high) using the mean ± 1 standard deviation as the
184 threshold. Correlation tests were conducted between these three grouped circulation indices and
185 the corresponding tree-ring indices.

186 3. Results and analysis

187 3.1 Regional chronologies and climate response characteristics

188 Regional chronologies exhibit high spatial heterogeneity in interannual variability (Figure
189 2). At the decadal scale (11-year moving average), consistent changes are evident in the four
190 regional chronologies around the 1890s, late 1910s to late 1950s, and late 1980s to late 2000s.
191 The period from late 1910s to late 1950s represents a low-growth phase, while the other two
192 periods correspond to consistent high-growth phases.



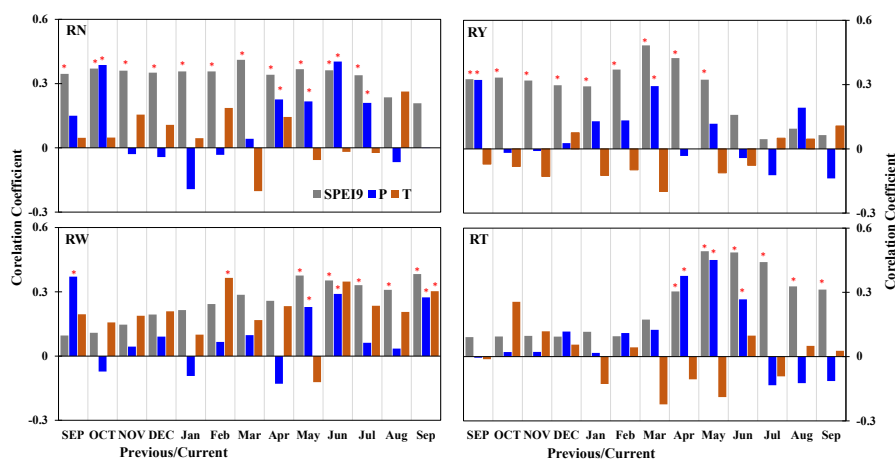
193

194 Figure 2. Spatial heterogeneity of *Z. xanthoxylum* radial growth in the Alxa Desert.



195 Note: Red bands denote periods of reduced growth in the 11-year moving average chronology,
 196 while blue bands indicate periods of enhanced growth. Red dashed boxes highlight intervals of
 197 reduced growth that show multi-regional agreement across the chronologies, while blue dashed
 198 boxes mark periods of enhanced growth with similar multi-regional consistency.

199 Correlation analysis between the four regional chronologies and the climate variables
 200 reveals spatially heterogeneous correlation patterns. The climate variables include monthly
 201 mean temperature, monthly precipitation, and SPEI at different scales (1962–2017). Correlation
 202 analysis of SPEI across scales from 1 to 12 months revealed the greatest significance at the 9-
 203 month scale (Figure 3).



204
 205 Figure 3. Correlation between the four regional chronologies and monthly mean temperature
 206 (red), monthly precipitation (blue), and 9-month scale SPEI (grey). * Pearson's r correlation,
 207 significant at $P < 0.05$.

208 Overall, the four regional chronologies exhibit a positive correlation with the SPEI from
 209 September of the previous year to September of the current year, always achieving significant
 210 levels in the growing season. The four regional chronologies also significantly correlate with
 211 the average 9-month scale SPEI from April to August. A predominantly positive correlation was
 212 observed between the four regional chronologies and monthly precipitation during the growing
 213 season, reaching significant levels, particularly in the RN, RW, and RT, from April to June, with
 214 most months achieving significance levels. A significant positive correlation with monthly
 215 mean temperature was evident in the RW in February and September of the current year, while



216 the other three chronologies exhibited weak negative correlations with monthly mean
217 temperatures during the growing season.

218 These findings indicate that radial growth variation of *Z. xanthoxylum* on the Alxa Desert
219 Plateau is primarily constrained by climatic aridity, resulting from low precipitation during the
220 growing season.

221

222 **3.2 Interannual and decadal characteristics of the influence of two circulation systems on** 223 **the radial growth of desert shrubs**

224 To investigate the impact of two major circulation systems on the radial growth of desert
225 shrubs on the Alxa Plateau, correlation analyses were conducted at the interannual scale, using
226 the entire regional chronology and grouped circulation indices.

227 The results indicate that no significant correlations were observed among the four regional
228 chronologies or their groupings at the interannual scale (Figure 4). Analysis of the two major
229 circulation indices revealed relatively consistent correlations between the medium index group
230 and the four regional chronology indices: negative with WCI, and positive with EASMI. The
231 low WCI index group showed positive correlations with RN and RT, and negative correlations
232 with RW and RY. The high WCI index group exhibited positive correlations with the four
233 regional chronologies. The low EASMI index group exhibited negative correlations with the
234 four regional chronologies. For the high EASMI index group, negative correlations were
235 observed with all four regional chronologies, except RT.

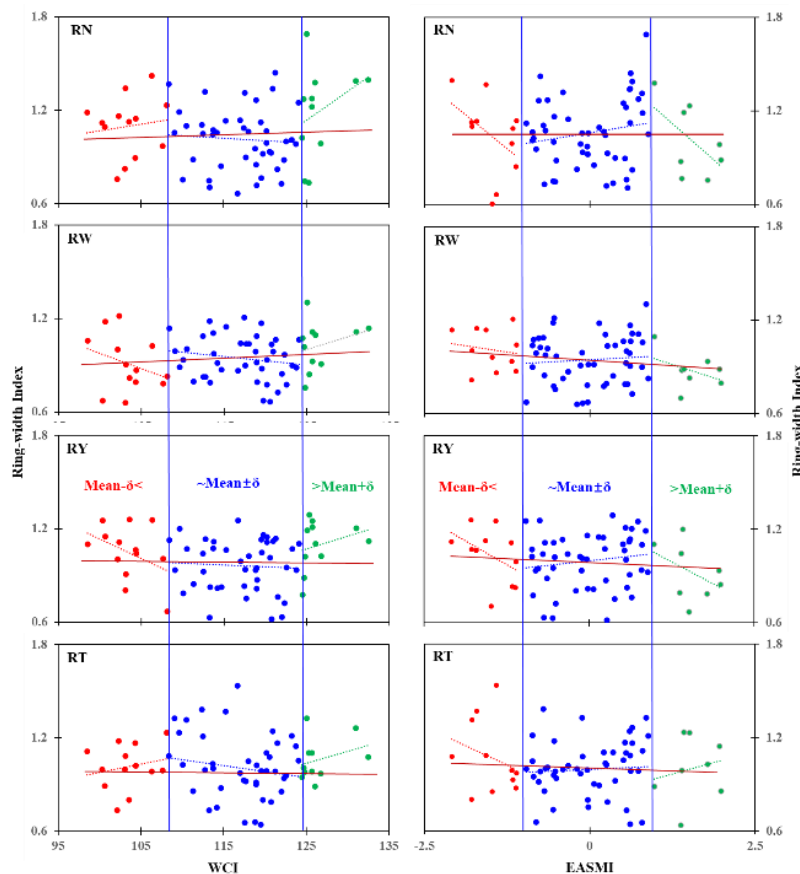
236 Overall, shrub diameter growth in the medium index group showed a negative correlation
237 with WCI, and a positive correlation with EASMI. For the high and low index groups, however,
238 shrub growth showed inconsistent correlations with WCI and EASMI.

239 Correlation analysis between the two circulation indices and the four regional
240 chronologies at the interdecadal scale (Figure 5) indicates a negative correlation between the
241 four chronologies and the EASMI ($p < 0.01$), especially in the Ulan Buh Desert and the
242 Nurgong mountain. For the WCI, the Ulan Buh Desert (RW) and Nurgong (RN) regions showed
243 a significant positive correlation. Conversely, the Tengger Desert (RT) and Badain Jaran Desert
244 (RY) showed insignificant negative correlations.



245 Further partial correlation analysis of the two circulation indices and four regional
 246 chronologies at the decadal scale revealed that all four chronologies exhibited significantly
 247 negative correlations with EASMI when controlling for the WCI variable. Controlling for the
 248 EASMI variable produced significant positive correlations between the RN and RW in the
 249 northern Alxa Plateau and the WCI, while the RY and RT in the southern Alxa Plateau showed
 250 negative correlations, with only the RY reaching statistical significance (see Table 1).

251 As shown in Figure 5, the WCI exhibits four discernible phases on a decadal scale: a
 252 gradual strengthening phase (1956–1963), followed by a weakening phase (1964–1979), a
 253 renewed strengthening period (1980–1994), and a subsequent weakening phase (1995–2011).
 254 The highest index values occurred primarily during two intervals: 1962–1967 and 1989–2007.
 255 In contrast, the lowest values were observed between 1971–1985 and 1956–1958.

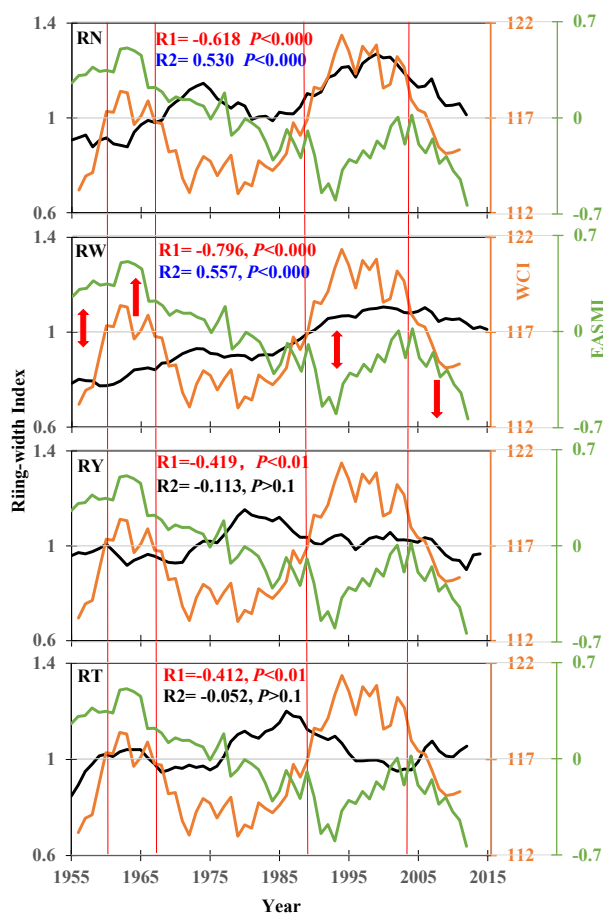


256

257 Figure 4. Correlation between the *Z. xanthoxylum* radial growth and the East Asian summer



258 monsoon index (EASMI) and Westerlies Index (WCI).
 259 Note: Linear fits between the chronologies and circulation indices are shown for the whole
 260 series (red solid line) and for low, medium, and high index groups (red, blue, and green dashed
 261 lines, respectively). None of these fits were significant ($p > 0.05$).



262
 263 Figure 5. Correlations between the four regional chronologies (black) on the Alxa Plateau and
 264 the EASMI (green) and WCI (brown) at the decadal scale (11-year moving average), alongside
 265 features of their interaction. R1 and R2 represent the respective correlation coefficients and
 266 significance levels between the chronologies and EASMI/WCI. Red arrows indicate periods
 267 when both circulation systems were typically strong (upward), weak (downward), or when one
 268 was strong and the other was weak (bidirectional).



269 The EASMI exhibited a strengthening phase from 1955 to 1963, followed by a weakening
 270 phase from 1964 to 1993, another strengthening phase from 1994 to 1982, and then a weakening
 271 phase from 1983 to 2011. During the entire study period, the strongest EASMI occurred from
 272 1955 to 1978, while the weakest values were observed from 1990 to 1998 and from 2011 to
 273 2012.

274 Table 1. Partial correlation analysis of the regional chronologies on the Alxa Plateau with the
 275 EASMI and WCI

Control variables	RN	RW	RY	RT
WCI	-0.609***	-0.794***	-0.469***	-0.432***
EASMI	0.458***	0.545***	-0.274*	-0.198

276 Note: Two-tailed correlation test (n=53): ***($p < 0.001$), *($p < 0.05$);

277 Four patterns can be identified based on the intensity characteristics of the two circulation
 278 systems over the decadal scale: both circulations are strong (e.g., 1960–1966); both circulations
 279 are weak (e.g., 1980–1988 and 2005–2015); the WCI is strong and the EASMI is weak (e.g.,
 280 1990–1998); and the WCI is weak and the EASMI is strong (e.g., 1955–1959) (Figure 5).

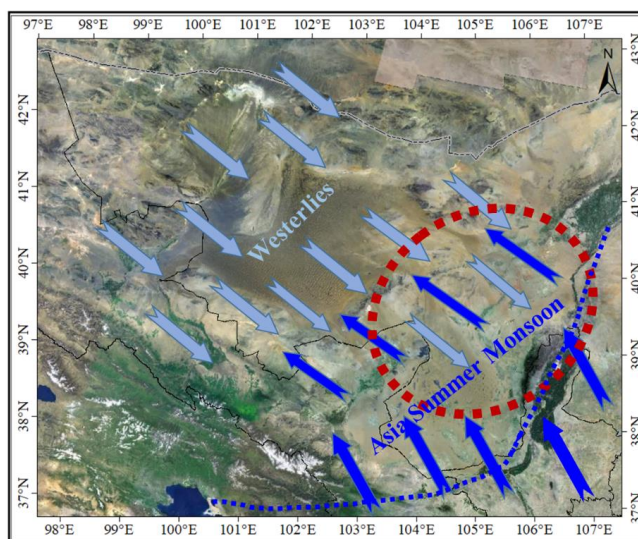
281 When both circulation indices were strong, the chronology indices of the RN, RW, and RY
 282 were relatively low, while the RT indices were relatively high. Under the pattern of a strong
 283 WCI and weak EASMI, the chronology indices of the RN and RW were relatively high, while
 284 those of the RY and RT were relatively low. When both circulation indices were weak, the
 285 values across the four chronologies were inconsistent. For example, during 1980–1988, the
 286 indices of the RN, RW, and RT showed an upward trend, whereas the RY exhibited a downward
 287 trend. Conversely, during 2005–2015, the indices in the RN, RW, and RY declined, while an
 288 upward trend was observed in the RT. When the WCI was weak and the EASMI was strong,
 289 the chronology indices in the RN, RW, and RT were relatively low, while the RY indices were
 290 close to the mean level.

291 The above results indicate that the radial growth of *Z. xanthoxylum* in different regions of
 292 the Alxa Plateau varies spatially in response to the interplay between the two circulation
 293 systems. When the EASMI is weak and the WCI is strong, radial growth is enhanced in the



294 northern parts of the plateau (RN and RW), but suppressed in the southern parts (RY and RT).
295 Conversely, when the EASMI is strong and the WCI is weak, radial growth is generally
296 inhibited across all four regions. When both circulation systems are strong, growth is favorable
297 only in the RT region, and unfavorable in the other three regions. When both circulation systems
298 are weak, radial growth is suppressed in the northern regions (RN and RW), but promoted in
299 the southern regions (RY and RT).

300 Overall, the study area is predominantly influenced by the EASM. However, the Westerlies
301 exert a significant impact on the central-southern, central-northern, and north-eastern parts of
302 the Alxa Plateau, except in the south-eastern region (RT), where its influence is weaker.
303 Consequently, we propose that the interaction zone between the EASM and the Westerlies is
304 primarily located in the central and northern parts of the Alxa Plateau (see Figure 6).
305 Furthermore, the northern boundary of the EASM extends approximately 400 km
306 northwestwards to the Yabulai–Nurgong–Lang Mountains line, covering the entire Tengger and
307 Ulan Buh deserts in the process. Meanwhile, the Westerlies can advance eastwards, influencing
308 areas up to the Helan Mountains.



309
310 Figure 6. Schematic diagram showing the interaction between the EASM (dark blue arrows)
311 and the Westerlies (light blue arrows) on the Alxa Plateau. The blue dashed lines show the



312 conventional EASM boundary, while the red circles highlight the primary area of interaction
313 between the EASM and Westerlies. This figure is modified after Xiao et al. (2024a).

314

315 **4. Discussion and Conclusions**

316 **4.1 Climate changes of deserts indicated by regional chronologies.**

317 Similar to the findings of other dendroclimatic studies in arid and semi-arid regions (Xiao
318 et al., 2012; Xiao et al., 2019; Xiao et al., 2024a; Ding et al., 2025), the radial growth of *Z.*
319 *xanthoxylum* in the four regions of the Alxa Plateau is primarily limited by low precipitation
320 and high temperatures. There is a significant positive correlation between radial growth and the
321 growing-season drought index. Consequently, the four regional chronologies primarily reflect
322 regional moisture variations (Figure 4).

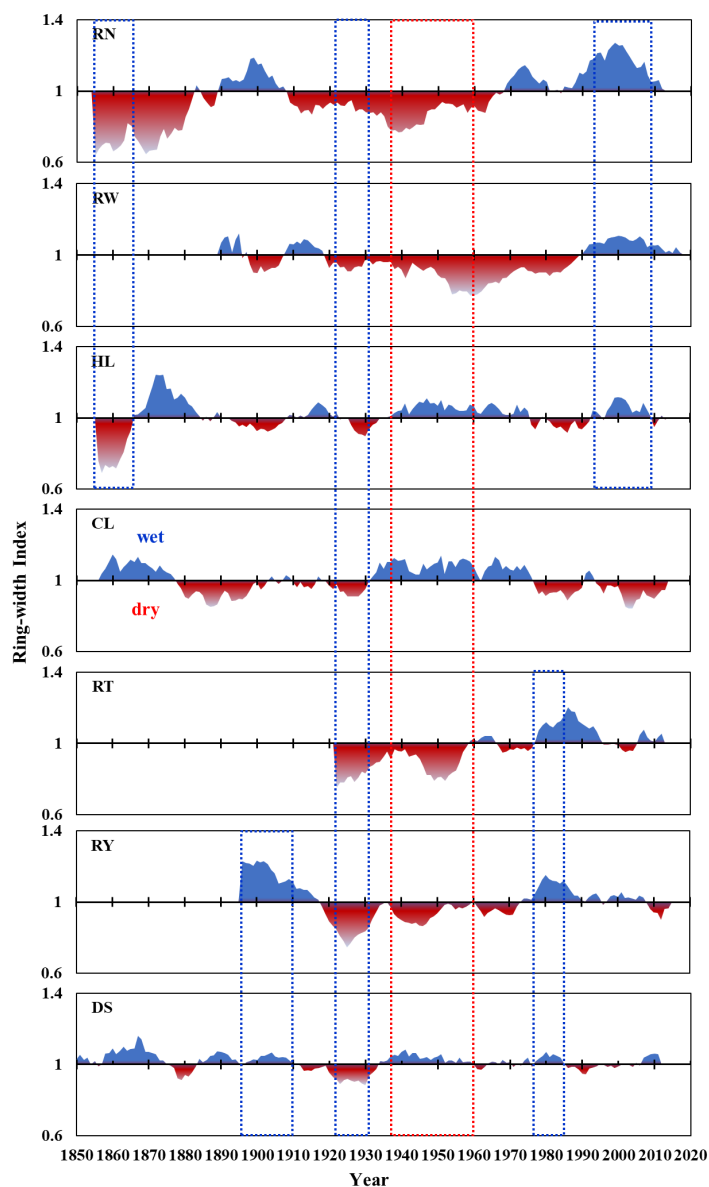
323 On an interannual scale, the four regions exhibit coherent dry/wet patterns, as well as
324 considerable spatiotemporal heterogeneity (Figure 2). Arid conditions dominate for a larger
325 proportion of the time across the four regions. This is particularly evident in northern Alxa
326 Plateau, where the longest continuous drought occurred from the 1910s to the 1980s, with the
327 most severe drought happening during the 1850s–1860s. A relatively consistent humid phase
328 occurred mainly from the late 1980s to the late 2000s across the four regions, though the timing,
329 duration, and intensity of this phase varied between regions. These characteristics reflect the
330 inherently arid nature of the inland desert climate.

331 On a broader (interdecadal) scale, an extreme drought period in the 1920s – 1930s was
332 shared by much of northern China (Liang et al., 2006; Fang et al., 2009; Fang et al., 2010). On
333 a decadal scale, a comparison of the four desert chronologies from the Alxa Plateau and three
334 chronologies from the surrounding mountains (Xiao et al., 2024a) indicates that all of them
335 experienced a sustained dry phase from the 1920s to the early 1930s (Figure 7). This drought
336 event persisted for approximately half a century.

337 A humid phase during the 1990s–2000s was revealed by both the two northern
338 chronologies of Alxa (RN and RW) and the chronologies from nearby Helan Mountains. This
339 humid period lasted longer and was relatively more pronounced in the desert regions than in
340 the mountains, especially in the northwestern Alxa (RN). The RN region also recorded a
341 synchronous dry period with the Helan Mountains around the 1860s. In RN, this drought



342 persisted until the late 1880s, whereas in the Helan Mountains, it transitioned to a humid phase
343 in the mid-1860s, which continued until the mid-1880s (Figure 7). During the 1900s and 1980s,
344 the two southern chronologies from the Alxa Desert (RY and RT) indicated periods of increased
345 humidity, which coincided with the northern mountains of the Hexi Corridor (DS). Once again,
346 the desert regions exhibited higher moisture levels.



347

348 Figure 7. Comparison of wet/dry variations between the desert chronologies of the Alxa Plateau



349 (RN, RW, RY, RT) and tree-ring records from the surrounding mountains (HL: Helan
350 Mountains; CL: Changling Mountains; DS: Dongda Mountains; data from Xiao et al., 2024a).
351 Blue boxes denote periods of consistent moisture variation, while red boxes indicate opposing
352 trends between the desert and mountain sites.

353

354 From the 1930s to the late 1950s, the Alxa Desert experienced a dry phase, while the
355 surrounding mountains were in a humid phase, reflecting an opposite climatic tendency
356 between the two types of terrain. This opposite climatic condition persisted for nearly half a
357 century, or even longer, in this area. The dry periods captured in the desert chronologies are
358 consistent with earlier single-site climate reconstructions (Xiao et al., 2012; Xiao et al., 2019;
359 Ding et al., 2025). These humid periods were also documented in the monsoon-influenced
360 mountainous areas east of the Alxa Plateau, as evidenced by local historical records (Yuan,
361 1994) and tree-ring chronologies from Mt. Xinglong (Fang et al., 2009; Chen et al., 2013) and
362 Mt. Guiqing (Fang et al., 2010).

363

364 **4.2 Influence of atmospheric circulations and their interaction on climate change in the** 365 **Alxa Plateau**

366 The westerly jet axis is situated near 40° N, and its zonal variation generates two pairs of
367 cross-coupled positive and negative anomaly centers, one over the Iranian Plateau and the
368 Mongolian Plateau, and the other over the Eastern Europe and Eastern China. The Alxa Plateau
369 lies at the transitional margin between the Mongolian Plateau and Eastern China anomaly
370 centers. Northward shifts of the westerly jet are generally associated with a slight increase in
371 precipitation over the region, whereas southward shifts correspond to reduced precipitation
372 (Sun et al., 2003; Guo et al., 2025).

373 When the westerlies intensify, their core shifts southwards, leading to increased
374 precipitation in the central and northern parts of the Alxa Plateau. Conversely, a weakening of
375 the westerlies results in reduced precipitation. At the same time, the cold, dry westerly airflow
376 lowers regional temperatures, thereby reducing vegetation evapotranspiration and alleviating
377 drought stress. Consequently, the radial growth of *Z. xanthoxylum* in the northern Alxa (RN and
378 RW) shows a significant positive correlation with the WCI, whereas the two chronologies in



379 the southern Alxa (RY and RT) show a negative correlation (with RY reaching statistical
380 significance) (Table 1). This finding is further supported by the significant positive correlation
381 observed on a decadal scale between the WCI and the radial growth of *P. crassifolia* in the
382 Helan Mountains, east to the Alxa Plateau (Xiao et al., 2024a).

383 Research indicates that some of the moisture carried by the westerlies moves southwards
384 after traversing northern Xinjiang, reaching northern Qinghai, the Hexi Corridor, northern
385 Ningxia, and northern Shaanxi (Li et al., 2012). Western China lies within the zone of the
386 prevailing westerlies, with the majority of moisture transported by these winds (Chen et al.,
387 2008; 2009; 2010; Liu et al., 2017). Analyses indicate that the WCI is a reliable indicator of
388 moisture transport and the moisture budget in northwestern China, primarily in summer. This
389 is based on reanalysis data from the American National Environmental Protection
390 Committee/National Center for Atmospheric Research (Feng et al., 2004; Wang et al., 2005; Li
391 et al., 2008).

392 When the EASM intensifies, its warm, moist airflows are obstructed by high mountain
393 ranges such as the Helan Mountains and the eastern Qilian Mountains at the monsoon boundary.
394 This orographic lifting leads to precipitation on the windward side, while the descending air on
395 the leeward side becomes warm and dry. This further exacerbates the high temperatures and
396 aridity of the Alxa Plateau. Consequently, the radial growth of *Z. xanthoxylum* in the four
397 regions of the plateau shows a significant negative correlation with the EASMI (Table 1).
398 Dendroclimatological studies of *P. crassifolia* in the Helan and Changling Mountains, the
399 southeastern part outside the Alxa, corroborate this mechanistic interpretation by reporting a
400 positive correlation between radial growth and the EASMI on a decadal scale. This confirms
401 the interception of warm, humid monsoon airflows by the surrounding mountains (Xiao et al.,
402 2024a). This phenomenon is referred to as the foehn effect within a rain shadow area (where
403 foehn refers to a type of dry, warm, and down-slope wind). Wind tunnel experiments and
404 simulation indicate that the tropical semi-arid climate and the formation of desertified savanna
405 in western Hainan Island, China, are primarily caused by the foehn effect associated with
406 multiple mountain ranges in the rain shadow region (Li et al., 2006). Meteorological
407 observations on both sides of the Southern Alps in New Zealand have also confirmed the
408 occurrence of lee-side foehn effects resulting from orographic blocking (Mcgowan and Andrew,



409 1994). The Alxa Plateau is surrounded by high mountain ranges, including the Qilian
410 Mountains, the North Mountains of the Hexi Corridor, the extensions of the Altai Mountains,
411 and the Helan Mountains. These ranges have an average elevation of between 3,000 and 4,000
412 meters. These ranges block both major atmospheric circulation systems, creating significant
413 foehn effects, which subsequently influence climate change and vegetation growth in the desert
414 regions of the Alxa Plateau.

415 The southeastern Tengger Desert lies within a gap between the southern Helan Mountains
416 and the eastern Qilian Mountains. This creates a topographic funnel effect that intensifies the
417 interaction between the weakened edges of the two circulation systems. On a decadal scale,
418 when the two circulation systems show contrasting intensities (one strong and one weak), *Z.*
419 *xanthoxylum* radial growth in this region tends to be lower (Figure 5). Conversely, when both
420 circulations are either strong or weak, radial growth tends to be higher. This pattern suggests
421 that a more balanced state between the two circulation systems may reduce or counteract
422 regional aridity, thereby favoring shrub growth. Decadal-scale correlation and partial
423 correlation analyses further indicate that this area is primarily influenced by the EASM (Table
424 1).

425 A study of Holocene lake-level evolution in the ancient Zhuye Lake, on the central Alxa
426 Plateau, showed that lake-level change was subject to the combined effects of EASM and the
427 arid climate of Central Asia (Li, 2009). The radial growth of *Z. xanthoxylum* in the RY and RT
428 is generally similar. However, due to its more westerly location, the RY is influenced by both
429 the EASM and the Westerlies, as indicated by a significant partial correlation with the WCI
430 (Table 1). Dendroclimatological studies from nearby mountains further illustrate this: on a
431 decadal scale, the radial growth of *P. crassifolia* in the Changling Mountains (on the southern
432 edge of the Tengger Desert) shows a positive correlation with the EASMI and a non-significant
433 negative correlation with the WCI. In contrast, spruce growth in the Dongda Mountains
434 (northern mountains of the Hexi Corridor, and southern edge of the Badain Jaran Desert)
435 correlates positively with the EASMI, yet exhibits a significant negative correlation with the
436 WCI (Xiao et al., 2024a).

437 Paleoclimate records from Qinghai Lake on the northern Tibetan Plateau indicate that the
438 cold, dry Westerlies and the warm, humid EASM exhibited an anti-phase relationship on



439 glacial–interglacial to millennial scales during the late Quaternary. Another study of the nearby
440 Toson Lake area also showed that the out-of-phase moisture transport and interaction between
441 the Westerlies and EASM drive asynchronous lake evolution between Central and East Asia
442 over centennial-millennial timescales (Li et al., 2022). Their interplay likely represented a
443 dominant climatic mode in eastern Central Asia for much of this period (An et al., 2012). This
444 anti-phase pattern is also evident in the instrumental indices of the two circulation systems over
445 the past five decades (Figure 5). This interaction mechanism has also been confirmed by studies
446 on the arid regions of Central Asia (Luo et al., 2026). These indices are constructed based on
447 the seasonal to interannual positional shifts, boundary variations, and intensity changes of the
448 atmospheric circulations (Zhang et al., 2003; Li and Zeng, 2005; Li et al., 2008; Chen and Chen,
449 2017; Chen et al., 2018). Consequently, the tree-ring chronologies from the four regions of the
450 Alxa Plateau predominantly show inverse correlations with the two circulation indices on both
451 interannual and decadal scales (Figure 5; Table 1).

452 Research using proxy indicator cycles has shown that the Alxa Plateau and its
453 surroundings area are influenced by large-scale ocean-atmosphere oscillations on interannual
454 and interdecadal scales, such as the North Atlantic Oscillation, Pacific Decadal Oscillation, El
455 Niño–Southern Oscillation, and solar activity (Gou et al., 2015a; 2015b; Liu et al., 2016; Wang
456 et al., 2017). The intensification of the EASM in humid monsoon regions and arid inland
457 monsoon margins may be associated with rising temperatures in the Northern Hemisphere (An
458 et al., 2012). However, all of the aforementioned large-scale climate and ocean-atmosphere
459 changes affect the EASM and westerly circulation through different pathways (Li, 2009). These
460 influences, in turn, have various effects on the northwestern margin of the EASM and the
461 interaction zone between the two major atmospheric circulations, often with a lagged response
462 (Ou and Qian, 2006). Consequently, over the past five decades, the phase relationship and
463 interaction between the two circulations have exhibited additional strong-weak combinatorial
464 patterns in response to global-scale climatic changes (Figure 5). This complexity has
465 contributed to the heterogeneity of the spatiotemporal response of *Z. xanthoxylum* radial growth
466 to climate across the four desert regions of the Alxa Plateau.

467 In conclusion, this dendroclimatological study, based on tree-ring width chronologies from
468 26 *Z. xanthoxylum* shrub sites across four desert regions of the Alxa Plateau (RN: Nurgong;



469 RW: Ulan Buh; RY: Badain Jaran; RT: Tengger), demonstrates that the radial growth of this
470 desert shrub is primarily limited by drought conditions resulting from low precipitation during
471 the growing-season. The dry/wet variations reflected in the four chronologies exhibit notable
472 spatiotemporal heterogeneity, which is largely governed by the interaction between the EASM
473 and Westerlies. While the entire study area is predominantly influenced by the EASM, and the
474 northern and southwestern parts are also affected by the westerly jet. Consequently, the
475 interaction between the westerly jet and the EASM occurs mainly over the central and northern
476 Alxa Plateau, thereby driving regional vegetation dynamics and desertification processes.

477

478 **Data availability.** All data for this paper are available upon request.

479 **Author contributions.** All authors approved the manuscript and agreed on its submission. SX:
480 conceptualization, methodology, funding acquisition, investigation, resources, writing (original
481 draft, review and editing). XP: investigation, data collection and writing review. AD:
482 investigation, data collection, and sample processing. QT: investigation. YR and JW:
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484 **Competing interests.** The contact author has declared that none of the authors has any
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493 **Review statement.**

494 **References**

495 An, Z. S., Colman, S. M., Zhou, W. J., Li, X. Q., Brown, E. T., Jull, A. T., Cai, Y. J., Huang, Y.
496 S., Lu, X. F., and Chang, H.: Interplay between the Westerlies and Asian monsoon recorded
497 in Lake Qinghai sediments since 32 ka, *Sci Rep*, 2(1), 619,
498 doi:<https://doi.org/10.1038/srep00619>, 2012.



- 499 Beguería, S., Vicente-Serrano, S. M., Reig, F., and Latorre, B.: Standardized precipitation
500 evapotranspiration index (SPEI) revisited: parameter fitting, evapotranspiration models,
501 tools, datasets and drought monitoring, *Int. J. Climatol.*, 34(10), 3001–3023,
502 doi:<https://doi.org/10.1002/joc.3887>, 2014.
- 503 Biondi, F., and Waikul, K.: DENDROCLIM2002: A C++ program for statistical calibration of
504 climate signals in tree-ring chronologies, *Comput. Geosci.*, 30(3), 303–311,
505 doi:<https://doi.org/10.1016/j.cageo.2003.11.004>, 2004.
- 506 Chen, F., Wei, W. S., Yuan, Y. J., Yu, S. L., Shang, H. M., Zhang, T. W., Zhang, R. B., Wang, H.
507 Q., and Qin, L.: Variation of annual precipitation during 1768-2006 in Gansu inferred from
508 multi-site tree-ring chronologies, *Journal of Desert Research*, 33(5), 1520, 2013.
- 509 Chen, F., Yuan, Y. J., Zhang, T. W., and Linderholm, H. W.: Annual precipitation variation for
510 the southern edge of the Gobi Desert (China) inferred from tree rings: linkages to climatic
511 warming of twentieth century, *Nat. Hazards*, 81(2), 939–955,
512 doi:<https://doi.org/10.1007/s11069-015-2113-z>, 2016.
- 513 Chen, F. H., Chen, J. H., Holmes, J., Boomer, I., Austin, P., Gates, J. B., Wang, N. L., Brooks,
514 S. J., and Zhang, J. W.: Moisture changes over the last millennium in arid central Asia: a
515 review, synthesis and comparison with monsoon region, *Quat. Sci. Rev.*, 29(7-8), 1055–
516 1068, doi:<https://doi.org/10.1016/j.quascirev.2010.01.005>, 2010.
- 517 Chen, F. H., Chen, J. H., Huang, W., Chen, S. Q., Huang, X. Z., Jin, L. Y., Jia, J., Zhang, X. J.,
518 An, C. B., and Zhang, J. W.: Westerlies Asia and monsoonal Asia: Spatiotemporal
519 differences in climate change and possible mechanisms on decadal to sub-orbital
520 timescales, *Earth-Sci. Rev.*, 192, 337–354,
521 doi:<https://doi.org/10.1016/j.earscirev.2019.03.005>, 2019a.
- 522 Chen, F. H., Fu, B. J., Xia, J., Wu, D., Wu, S. H., Zhang, Y. L., Sun, H., Liu, Y., Fang, X. M.,
523 and Qin, B. Q.: Major advances in studies of the physical geography and living
524 environment of China during the past 70 years and future prospects, *Sci. China Earth Sci.*,
525 49(11), 1659–1696, doi:<https://doi.org/10.1007/s11430-019-9522-7>, 2019b.
- 526 Chen, F. H., Sun, J. H., and Huang, W.: A discussion on the westerly-dominated climate model
527 in mid-latitude Asia during the modern interglacial period, *Earth Sci. Front.*, 16(6), 23–32,
528 2009.



- 529 Chen, F. H., Yu, Z. C., Yang, M. L., Ito, E., Wang, S. M., Madsen, D. B., Huang, X. Z., Zhao,
530 Y., Sato, T., and Birks, H. J. B.: Holocene moisture evolution in arid central Asia and its
531 out-of-phase relationship with Asian monsoon history, *Quat. Sci. Rev.*, 27(3-4), 351–364,
532 doi:<https://doi.org/10.1016/j.quascirev.2007.10.017>, 2008.
- 533 Chen, H. S., and Chen, J. K.: Classification of East Asian summer monsoon indices and their
534 basic physical features, *Transactions of Atmospheric Sciences*, 40(3), 299–309,
535 doi:<https://doi.org/10.13878/j.cnki.dqkxxb.20160323001>, 2017.
- 536 Chen, J., Huang, W., Jin, L. Y., Chen, J. H., Chen, S. Q., and Chen, F. H.: A climatological
537 northern boundary index for the East Asian summer monsoon and its interannual
538 variability, *Sci. China Earth Sci.*, 48(1), 93–101, doi:[https://doi.org/10.1007/s11430-017-](https://doi.org/10.1007/s11430-017-9122-x)
539 9122-x, 2018.
- 540 Cook, E. R.: A time series analysis approach to tree ring standardization (dendrochronology,
541 forestry, dendroclimatology, autoregressive process), Ph.D. thesis, The University of
542 Arizona, 1985.
- 543 Ding, A. J., Xiao, S. C., Peng, X. M., and Tian, Q. Y.: Shrub-ring reconstruction of precipitation
544 in the southeast of Tengger Desert, northwest China and its relationship with dust storm
545 events, *Dendrochronologia*, 126426, doi:<https://doi.org/10.1016/j.dendro.2025.126426>,
546 2025.
- 547 Ding, A. J., Xiao, S. C., Tian, Q. Y., and Han, C.: Correcting eccentric growth rings using basal
548 area increment: a case study for a desert shrub in Northwestern China, *Tree-Ring Res.*,
549 77(1), 1–9, doi:<https://doi.org/10.3959/TRR2020-4>, 2021.
- 550 Fang, K. Y., Gou, X. H., Chen, F. H., D'arrigo, R., and Li, J. B.: Tree-ring based drought
551 reconstruction for the Guiqing Mountain (China): linkages to the Indian and Pacific
552 Oceans, *Int. J. Climatol.*, 30(8), 1137–1145, doi:<https://doi.org/10.1002/joc.1974>, 2010.
- 553 Fang, K. Y., Gou, X. H., Chen, F. H., Yang, M. X., Li, J. B., He, M. S., Zhang, Y., Tian, Q. H.,
554 and Peng, J. F.: Drought variations in the eastern part of northwest China over the past two
555 centuries: evidence from tree rings, *Clim. Res.*, 38(2), 129–135,
556 doi:<https://doi.org/10.3354/cr00781>, 2009.
- 557 Feng, W., Wang, K. L., and Jiang, H.: Influences of westerly wind interannual change on water
558 vapor transport over northwest China summer, *Plateau Meteorology*, 23(2), 271–27, 2004.



- 559 Gou, X. H., Deng, Y., Gao, L. L., Chen, F. H., Cook, E., Yang, M. X., and Zhang, F.: Millennium
560 tree-ring reconstruction of drought variability in the eastern Qilian Mountains, northwest
561 China, *Clim. Dyn.*, 45(7), 1761–1770, doi:<https://doi.org/10.1007/s00382-014-2431-y>,
562 2015a.
- 563 Gou, X. H., Gao, L. L., Deng, Y., Chen, F. H., Yang, M. X., and Still, C.: An 850-year tree-ring-
564 based reconstruction of drought history in the western Qilian Mountains of northwestern
565 China, *Int. J. Climatol.*, 35(11), 3308–3319, doi:<https://doi.org/10.1002/joc.4208>, 2015b.
- 566 Guo, J. N., Xie, X. A., and Wang, A. Q.: A comparative study on the relationship between Asian
567 westerly jet position and summer precipitation in mid-latitude regions of East Asia and
568 Central Asia, *J. Earth Sci. Environ.*, 16(01), 53–65,
569 doi:<https://doi.org/10.7515/JEE222089>, 2025.
- 570 Holmes, R. L.: Computer-assisted quality control in tree-ring dating and measurement, 1983.
- 571 Hu, H. R., and Qian, W. H.: Confirmation of the northern edge of the East Asian summer
572 monsoon, *Prog. Nat. Sci.*, 17(1), 57–65, 2007.
- 573 Kang, S. Y., and Yang, B.: Precipitation variability at the northern fringe of the Asian summer
574 monsoon in Northern China and its possible mechanism over the past 530 years, *Quat.*
575 *Sci.*, 35(5), 1185–1193, doi:<https://doi.org/10.11928/j.issn.1001-7410.2015.05.14>, 2015.
- 576 Li, D. L., Shao, P. C., and Wang, H.: The position variations of the north boundary of East Asia
577 subtropical summer monsoon in 1951–2009, *Journal of Desert Research*, 33(5), 1511,
578 2013.
- 579 Li, G. Q., Wang, X. Y., Zhang, X. J., Yan, Z. F., Liu, Y. L., Yang, H., Wang, Y. X., Jonell, T. N.,
580 Qian, J. K., Gou, S. Y., et al.: Westerlies-Monsoon interaction drives out-of-phase
581 precipitation and asynchronous lake level changes between Central and East Asia over the
582 last millennium, *Catena*, 218, 106568, doi:<https://doi.org/10.1016/j.catena.2022.106568>,
583 2022.
- 584 Li, J. L., Li, Z. R., Yang, J. C., Shi, Y. Z., and Fu, J.: Analyses on spatial distribution and
585 temporal variation of atmosphere water vapor over northwest China in summer of later 10
586 years, *Plateau Meteorology*, 31(6), 1574–1581, 2012.



- 587 Li, J. P., and Zeng, Q. C.: A new monsoon index, its interannual variability and relation with
588 monsoon precipitation, *Clim. Environ. Res.*, 10(3), 351–365,
589 doi:<https://doi.org/10.3878/j.issn.1006-9585.2005.03.09>, 2005.
- 590 Li, S., Liu, X. W., and Sun, W.: Study of influence of rain shadow region on study
591 desertification in West of Hainan island based on wind tunnel experiments, *Journal of*
592 *desert research*, 26(2), 165–171, 2006.
- 593 Li, W. L., Wang, K. L., Fu, S. M., and Jiang, H.: The interrelationship between regional westerly
594 index and the water vapor budget in northwest China, *Journal of Glaciology and*
595 *Geocryology*, 30(1), 28–34, 2008.
- 596 Li, X., and Liu, X. D.: Relation of spring dust-storm activities in northern China and changes
597 of upper westerlies, *Plateau Meteorology*, 34(5), 1292–1300,
598 doi:<https://doi.org/10.7522/j.issn.1000-0534.2014.00067>, 2015.
- 599 Li, Y.: The pollen records from lake sediments and climate & lake model in the Marginal area
600 of Asian monsoon, Ph.D. thesis, Lanzhou University, Lanzhou, China, 2009.
- 601 Liang, E. Y., Liu, X. H., Yuan, Y. J., Qin, N. S., Fang, X. Q., Huang, L., Zhu, H. F., Wang, L.,
602 and Shao, X. M.: The 1920s drought recorded by tree rings and historical documents in
603 the semi-arid and arid areas of northern China, *Clim. Change*, 79(3), 403–432,
604 doi:<https://doi.org/10.1007/s10584-006-9082-x>, 2006.
- 605 Liu, W. C., Zhang, Q., and Fu, C.: Variation characteristics of precipitation and its affecting
606 factors in northwest China over the past 55 years, *Plateau Meteorology*, 36(6), 1533–1545,
607 2017.
- 608 Liu, Y., and Cai, Q. F.: The East Asia summer monsoon precipitation variations since AD 1840,
609 a case study of the tree-ring records from China and North Korean, *Sci. China Earth Sci.*,
610 33(6), 543–549, 2003.
- 611 Liu, Y., Cai, Q. F., Liu, W. G., Yang, Y. K., Sun, J. Y., Song, H. M., and Li, X. X.: Monsoon
612 precipitation variation recorded by tree-ring $\delta^{18}\text{O}$ in arid Northwest China since AD 1878,
613 *Chem. Geol.*, 252(1–2), 56–61, doi:<https://doi.org/10.1016/j.chemgeo.2008.01.024>, 2008.
- 614 Liu, Y., Cai, Q. F., Ma, L. M., and An, Z. S.: Tree ring precipitation records from Baotou and
615 the East Asia summer monsoon variations for the last 254 years, *Earth Sci. Front.*, 8(1),
616 91–97, 2001.



- 617 Liu, Y., Sun, C. F., Li, Q., and Cai, Q. F.: A *Picea crassifolia* tree-ring width-based temperature
618 reconstruction for the Mt. Dongda region, Northwest China, and its relationship to large-
619 scale climate forcing, *PLoS One*, 11(8), e0160963,
620 doi:<https://doi.org/10.1371/journal.pone.0160963>, 2016.
- 621 Liu, Y. Z., Wu, C. Q., Jia, R., and Huang, J. P.: An overview of the influence of atmospheric
622 circulation on the climate in arid and semi-arid region of Central and East Asia, *Sci. China
623 Earth Sci.*, 48(9), 1141–1152, doi:<https://doi.org/10.1007/s11430-017-9202-1>, 2018.
- 624 Luo, Y., Xu, C. C., Li, Z. Y., Li, L., Wang, Q., and Zhang, Q. Y.: Synergistic enhancement of
625 Asian monsoons to westerlies intensifies the drying trend of arid Central Asia over the last
626 20 years, *Gondwana Res.*, 149, 337-350, doi:<https://doi.org/10.1016/j.gr.2025.08.021>,
627 2026.
- 628 Ma, M. J., Pu, Z. X., Wang, S. O., and Zhang, Q.: Characteristics and numerical simulations of
629 extremely large atmospheric boundary-layer heights over an arid region in north-west
630 China, *Bound.-Layer Meteor.*, 140(1), 163–176, doi:<https://doi.org/10.1007/s10546-011-9608-2>, 2011.
- 632 Mcgowan, A. H., and Andrew, P. S.: Regional and local scale characteristics of Foehn wind
633 events over the South Island of New Zealand, *Meteorology And Atmospheric Physics*, 58,
634 151-164, doi:<https://doi.org/10.1007/BF01027562>, 1994.
- 635 Ou, T. H., and Qian, W. H.: Vegetation variations along the monsoon boundary zone in East
636 Asia, *Chinese Journal of Geophysics - Chinese edition*, 49(3), 627–636, 2006.
- 637 Peng, X. M., Zhang, B. W., Wang, W. S., Ding, A. J., and Xiao, S. C.: Different drought indices
638 showed different variations and applicability to dendrochronological studies,
639 *Dendrochronologia*, 84, 126179, doi:<https://doi.org/10.1016/j.dendro.2024.126179>, 2024.
- 640 Qu, W. J., Zhang, X. Y., Wang, D., Shen, Z. X., Mei, F. M., Cheng, Y., and Yan, L. W.: The
641 important significance of westerly wind study, *Marine Geology & Quaternary Geology*,
642 24(1), 125–132, 2004.
- 643 Ren, Y., Xiao, S. C., Peng, X. M., Ding, A. J., Tian, Q. Y., Wang, J. K.: Shrub dendrochronology
644 reveals the last 150 years of spatiotemporal patterns of desertification in the Tengger
645 Desert. *Catena* 264 (2026) 109844. <https://doi.org/10.1016/j.catena.2026.109844>



- 646 Sun, D. H., An, Z. S., Su, R. X., Lu, H. Y., and Sun, Y. B.: The dust deposition records of the
647 evaluation of Asia monsoon and Westerly circulation in north China in the last 2.6 Ma, *Sci.*
648 *China Earth Sci.*, 33(6), 497–504, doi:<https://doi.org/10.1007/s11430-025-1596-y>, 2003.
- 649 Tang, X., Qian, W. H., and Liang, P.: Climatic features of boundary belt for East Asian summer
650 monsoon, *Plateau Meteorology*, 25(3), 375–381, 2006a.
- 651 Tang, X., Sun, G. W., and Qian, W. H.: Study on the north edge of the Asia Summer Monsoon,
652 Atmosphere Press, 2006b.
- 653 Wang, J. L., Yang, B., Ljungqvist, F. C., Luterbacher, J., Osborn, T. J., Briffa, K. R., and Zorita,
654 E.: Internal and external forcing of multidecadal Atlantic climate variability over the past
655 1,200 years, *Nat. Geosci.*, 10(7), 512–517, doi:<https://doi.org/10.1038/ngeo2962>, 2017.
- 656 Wang, J. S.: Study on regional climate response to global warming in the arid central-east Asia
657 over the past 100a, Ph.D. thesis, Lanzhou University, Lanzhou, 2006.
- 658 Wang, K. L., Jiang, H., and Zhao, H. Y.: Atmospheric water vapor transport from westerly and
659 monsoon over the Northwest China, *Advances in Water Science*, 16(3), 432–438,
660 doi:<https://doi.org/10.14042/j.cnki.32.1309.2005.03.021> 2005.
- 661 Xiao, S. C., Chen, X. H., and Ding, A. J.: Study process of climate changes, environment
662 evolution and its driving mechanism in the last two centuries in the Alxa Desert, *Journal*
663 *of Desert Research*, 37(6), 1102–1110, doi:[https://doi.org/10.7522/j.issn.1000-](https://doi.org/10.7522/j.issn.1000-694X.2017.00002)
664 [694X.2017.00002](https://doi.org/10.7522/j.issn.1000-694X.2017.00002), 2017.
- 665 Xiao, S. C., Peng, X. M., Tian, Q. Y., Ding, A. J., Xie, J. L., and Su, J. R.: Interaction between
666 the East Asian summer monsoon and westerlies as shown by tree-ring records, *Clim. Past.*,
667 20(7), 1687–1701, doi:<https://doi.org/10.5194/cp-20-1687-2024>, 2024a.
- 668 Xiao, S. C., Su, J. R., Peng, X. M., and Tian, Q. Y.: Review on the study of monsoon-westerly
669 interaction in the inland arid none of northwestern China, *Plateau Meteorology*, 43(6),
670 1355–1363, doi:<https://doi.org/10.7522/j.issn.1000-0534.2024.00028>, 2024b.
- 671 Xiao, S. C., Xiao, H. L., Dong, Z. B., and Peng, X. M.: Dry/wet variation recorded by shrub
672 tree-rings in the central Badain Jaran Desert of northwestern China, *J. Arid Environ.*, 87,
673 85–94, doi:<https://doi.org/10.1016/j.jaridenv.2012.06.013>, 2012.
- 674 Xiao, S. C., Yan, C. Z., Tian, Y. Z., Si, J. H., Ding, A. J., Chen, X. H., Han, C., and Teng, Z. Y.:
675 Regionalization for desertification control and countermeasures in the Alxa Plateau, China,



- 676 Journal of Desert Research, 39(5), 182–19, doi:<https://doi.org/10.7522/j.issn.1000->
677 694X.2019.00068, 2019.
- 678 Xu, Y., and Qian, W. H.: Research on East Asia monsoon: a review, *Acta Geographica Sinica*,
679 (z1), 138–146, 2003.
- 680 Yuan, L.: Hazards history in northwestern China, Gansu people’s press, Lanzhou, China, 1994.
- 681 Zhang, Q., Yang, J. H., Wang, P. L., Yu, H. P., Yue, P., Liu, X. Y., Lin, J. J., Duan, X. Y., Zhu,
682 B., and Yan, X. Y.: Progress and prospect on climate warming and humidification in
683 Northwest China, *Chin. Sci. Bull.*, 68(14), 1814–1828, doi:<https://doi.org/10.1360/TB->
684 2022-0643, 2023.
- 685 Zhang, Q. Y., Tao, S. Y., and Chen, L. T.: The inter-annual variability of east asian summer
686 monsoon indices and its association with the pattern of general circulation over east Asia,
687 *Acta Meteorol. Sin.*, (05), 559–56, 2003.
- 688 Zhong, H. L., and Li, D. L.: Relationship between sand-dust storm in northern China in April
689 and westerly circulation, *Plateau Meteorology*, 24(1), 104–111, 2005.