

Improved representation of isoprene-derived secondary organic aerosol in CAM6-Chem reveals regional contrasts in its long-term changes over China

Wenxin Zhang¹, Man Yue^{1,2}, Xinyue Shao¹, Xinyi Dong^{1,3,4}, Minghuai Wang^{1,4}

5 ¹School of Atmospheric Science, Nanjing University, Nanjing, 210023, China

²Zhejiang Institute of Meteorological Sciences, Hangzhou, 310008, China

³Frontiers Science Center for Critical Earth Material Cycling, Nanjing University

⁴Joint International Research Laboratory of Atmospheric and Earth System Sciences & Institute for Climate and Global Change Research, Nanjing University, Nanjing, 210023, China

10 *Correspondence to:* Xinyi Dong (dongxy@nju.edu.cn)

Abstract. Isoprene-derived secondary organic aerosol (ISOA) is an important component of atmospheric organic aerosol, but its formation remains incompletely represented in global chemical models, creating uncertainty in ISOA changes and their drivers. In this study, we implemented an explicit ISOA formation scheme in Community Atmosphere Model version 6 with comprehensive tropospheric and stratospheric chemistry (CAM6-Chem), incorporating high-NO_x gas-phase epoxide precursors and explicitly representing branch-resolved heterogeneous uptake from both low-NO_x and high-NO_x pathways. Evaluation against ground-based observations shows that the updated model reasonably reproduces the concentrations and compositional structure of four explicitly represented ISOA subspecies. At the bulk aerosol level, the update alleviates the underestimation of SOA over China, improving the normalized mean bias from −76.7% to −51.6%. ISOA formation in China is governed by NO_x-dependent competition between the low- and high-NO_x pathways, with the former remaining dominant at the national scale. Long-term analysis for 2000–2019 shows a weak national-mean ISOA trend due to offsetting regional changes of opposite signs. The most pronounced increase occurs in Southwest China, where enhanced biogenic isoprene emissions emerge as the leading statistical predictor of the ISOA increase, with a regression-based relative contribution of 58.29%. The strongest decrease occurs in the Shaanxi–Gansu–Ningxia region, where declining sulfate is identified as the leading statistical predictor, with a regression-based relative contribution of 63.37%. These results highlight the regional heterogeneity of ISOA changes in China and the importance of jointly representing precursor supply and heterogeneous reaction conditions in simulating ISOA formation and trends.

1 Introduction

Isoprene (C₅H₈) is one of the most important biogenic volatile organic compounds (VOCs) in the atmosphere, with global emissions accounting for nearly half of the total biogenic VOCs emissions (Guenther et al., 2012). Owing to its very high chemical reactivity, isoprene is rapidly oxidized in the atmosphere by hydroxyl radicals (OH), ozone (O₃), and nitrate

radicals (NO_3) (Wennberg et al., 2018). The resulting oxidation products can further undergo multiphase chemical processes to form secondary organic aerosol (SOA) with semi-volatile and low-volatility characteristics.

The mechanisms by which isoprene forms SOA are highly sensitive to the ambient nitrogen oxides (NO_x) background (Surratt et al., 2010). Under low- NO_x conditions, the primary oxidation products of isoprene readily react with hydroperoxyl radicals (HO_2) or organic peroxy radicals (RO_2) to form isoprene hydroxyl hydroperoxides (ISOPROOH), as well as carbonyl and hydroxylated products (Wennberg et al., 2018). ISOPROOH is further oxidized by OH under low- NO_x conditions to produce isoprene epoxydiols (IEPOX), with a molar yield exceeding 75% (Paulot et al., 2009b). Hereafter, this low- NO_x sequence is referred to as the IEPOX pathway. IEPOX is highly water-soluble and reactive, and it can undergo acid-catalyzed multiphase processing to form SOA (Lin et al., 2012; Surratt et al., 2010). In these heterogeneous reactions, aerosol liquid water, sulfate, and nitrate can all act as key nucleophiles that add to the epoxide upon ring opening, producing low-volatility products. Specifically, aerosol liquid water-mediated pathways form 2-methyltetrols (2-MT) and related polyols, whereas sulfate- and nitrate-mediated nucleophilic addition produces organosulfates, organonitrates, and other oligomeric products that contribute substantially to SOA formation (Surratt et al., 2007, 2008, 2010). Under high- NO_x conditions, the primary oxidation products of isoprene preferentially react with nitric oxide (NO) to form important gas-phase intermediates such as methacryloyl peroxyxynitrate (MPAN) (Wennberg et al., 2018). Subsequent reactions of these intermediates with OH can produce epoxides, including methacrylic acid epoxide (MAE) and hydroxymethyl-methyl- α -lactone (HMML) (Lin et al., 2013; Nguyen et al., 2015). After partitioning into the particle phase, these epoxides undergo acid-catalyzed ring opening and subsequent multiphase processing, in which aerosol liquid water, sulfate, and nitrate act as important nucleophiles. These reactions ultimately yield SOA components dominated by 2-methylglyceric acid (2-MG), organosulfates, and organonitrates (Birdsall et al., 2014; Nguyen et al., 2015; Schwantes et al., 2019). Hereafter, this high- NO_x sequence is referred to as the HMML/MAE pathway. Although observational studies indicate that ambient concentrations of MAE and HMML are generally lower than those of IEPOX under low- NO_x conditions (Worton et al., 2013; Zhu et al., 2025), they remain key precursors for ISOA formation in high- NO_x environments and play a non-negligible role in the isoprene oxidation system. In regions strongly influenced by human activities (e.g., urban and industrial areas), NO_x concentrations are elevated while isoprene emissions remain substantial, making this pathway more influential for ISOA formation in such environments (Budisulistiorini et al., 2015; Rattanavaraha et al., 2016).

To improve model performance and simulation accuracy, it is essential to better represent ISOA formation across different NO_x regimes. The high- NO_x pathway remains insufficiently represented in global models, which can contribute to the underestimation of ISOA in polluted regions (Budisulistiorini et al., 2015; Hu et al., 2015; Lin et al., 2013). Observational evidence suggests that this pathway can contribute about 13–26% of ISOA in urban environments, highlighting its importance under polluted conditions (Ding et al., 2014; Zhang et al., 2022). Equally important is a product-level representation of IEPOX-derived ISOA ($\text{ISOA}_{\text{IEPOX}}$), rather than representing it as a single lumped product, because major

65 molecular tracers such as 2-MT and organosulfate/organonitrate products provide important constraints on ISOA
composition and particle-phase processing (Chen et al., 2024b; Lin et al., 2012). Therefore, explicitly representing key ISOA
subspecies from both low-NO_x and high-NO_x epoxide pathways provides a more complete description of ISOA formation
across pollution regimes, improves the simulation of ISOA composition, and enhances the model sensitivity to changes in
anthropogenic emissions.

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ISOA plays a key role at the interface of natural and anthropogenic emissions and has received wide attention in recent years
(Marais et al., 2016; Shrivastava et al., 2019, 2022). ISOA formation is jointly influenced by precursor emissions, oxidative
environment, and meteorological conditions, leading to significant temporal and spatial variability (Bardakov et al., 2021;
Carlton et al., 2009). Over the past two decades, rapid economic development across various regions and the continuous
75 implementation of different pollution control measures have led to significant changes in both anthropogenic pollutant
emissions and meteorological conditions, resulting in notable trends in ISOA concentrations (Silver et al., 2020; Su et al.,
2011; Zhang et al., 2025). Therefore, it is important to further investigate the key factors and mechanisms driving these
concentration changes. Existing studies on ISOA concentration changes have predominantly focused on the IEPOX pathway,
where sulfate has been identified as a key driver of variability. For instance, Marais et al. (2017) found that, during 1991–
80 2013, ISOA_{IEPOX} in the United States decreased significantly with declining sulfate, making sulfate the primary driver of the
long-term summertime reduction in surface organic aerosol (OA) (Marais et al., 2017). Similarly, long-term simulations by
Zheng et al. (2020) confirmed a 4.9% per year decrease in ISOA_{IEPOX} in the southeastern United States from 2000 to 2013,
attributed primarily to sulfate reductions (Zheng et al., 2020). This finding was consistent with Dong et al. (2022), who
showed that sulfate reductions in southern China corresponded to declines in ISOA_{IEPOX} (Dong et al., 2022). Zhang et al.
85 (2025) also reported that sulfate decreases across China were accompanied by a decline in ISOA_{IEPOX} (Zhang et al., 2025).
Liu et al. (2023) also observed that decreasing sulfate emissions contributed significantly to the decline in ISOA_{IEPOX} over
the continental United States based on long-term model simulations (Liu et al., 2023). However, some studies that consider
the HMML/MAE pathway report different results, indicating that the dominant drivers of ISOA can vary by region and
environment. For instance, Hu et al. (2025) showed that in Shanghai during 2015–2021, anthropogenic NO_x emissions were
90 the primary driver of ISOA reduction, with sulfur dioxide (SO₂) and aerosol acidity playing secondary roles (Hu et al., 2025).
Similarly, Budisulistiorini et al. (2015) found in the 2013 Southern Oxidant and Aerosol Study that human pollution
significantly influenced both the IEPOX and HMML/MAE pathways, with sulfate and NO₃ affecting their relative
importance in urban areas (Budisulistiorini et al., 2015).

95 Despite these advances, existing studies still rarely assess long-term ISOA changes by simultaneously considering pathway
competition under different NO_x emission backgrounds and the multiphase reaction environment. Given the large regional
contrasts in emissions, meteorological conditions, and oxidation environments across China (Ding et al., 2016; Li et al., 2018;
Wu et al., 2020; Yang et al., 2017), the dominant factors controlling ISOA formation and change are likely to differ by

region. To further explore this question, a model framework is needed that can represent both low-NO_x and high-NO_x epoxide pathways and distinguish their major reaction branches and products. The continuous development of SOA representation in large-scale models provides an important foundation for this purpose. Representations of biogenic SOA have evolved from simplified empirical or fixed-yield treatments used in early global models (Chung and Seinfeld, 2002; Chin et al., 2002; Carlton et al., 2010; Kim et al., 2015), through semi-empirical two-product absorptive partitioning models (Odum et al., 1996; Griffin et al., 1999), to the Volatility Basis Set (VBS) approach (Donahue et al., 2006; Robinson et al., 2007; Pye and Seinfeld, 2010), and more recently to explicit multiphase reactive uptake parameterizations for isoprene epoxides (Pye et al., 2013; Budisulistiorini et al., 2017; Jo et al., 2019, 2021; Ng et al., 2025). These developments allow ISOA formation to respond more directly to aerosol liquid water, acidity, sulfate, and other nucleophiles, and particle-phase reaction kinetics (Pye et al., 2013; Marais et al., 2016; Budisulistiorini et al., 2017; Jo et al., 2019, 2021). Building on this progress, this study extends the CAM6-Chem framework by incorporating high-NO_x epoxide precursors and subsequent heterogeneous reactions, distinguishing between H₂O-mediated nucleophilic addition and inorganic-nucleophile addition product classes, and evaluating how these updates affect ISOA composition and long-term ISOA trends over China. This framework further allows us to examine how ISOA trends are jointly controlled by precursor supply and heterogeneous reaction capacity across different regions. Nevertheless, important uncertainties remain in the current ISOA framework, including the simplified core-shell treatment of epoxide reactive uptake, uncertainties in aerosol phase state, organic-shell viscosity, phase separation, and diffusion limitations, as well as limited kinetic constraints for HMML/MAE uptake. Additional uncertainties arise from the assumptions of 100% ISOA yield and non-volatile epoxide-derived products, and from the possible omission of non-epoxide ISOA pathways in the current framework. Despite these remaining uncertainties, clarifying regional ISOA trends and controls can provide a scientific basis for evaluating how emission reductions jointly affect precursor supply and heterogeneous reaction environments, and for developing more targeted regional air-quality management strategies.

2 Data and Methods

2.1 Model Description

2.1.1 Model Configuration

This study employs the Community Atmosphere Model version 6 with comprehensive tropospheric and stratospheric chemistry (CAM6-Chem) within the Community Earth System Model version 2.1.0 (CESM2.1.0). Biogenic emissions are calculated online using the Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1), which is coupled to CESM (Emmons et al., 2020; Guenther et al., 2012). Anthropogenic emissions are taken from the Multi-resolution Emission Inventory for China (MEIC; <http://www.meicmodel.org>) (Li et al., 2017). In this study, intermediate-volatility organic compounds (IVOCs) and semi-volatile organic compounds (SVOCs) represent lumped lower-volatility

130 organic precursor classes defined by volatility rather than explicitly speciated molecular structures. Emissions of IVOCs and SVOCs are scaled from primary organic aerosol (POA) and non-methane volatile organic compounds (NMVOCs) emissions (Chang et al., 2022; Tilmes et al., 2019). The specific scaling equations and coefficients are described in the supplement of Zhang et al. (2025) (Eqs. A1–A3). Meteorological fields are constrained by the Modern-Era Retrospective analysis for Research and Applications (MERRA2) reanalysis data (Gelaro et al., 2017).

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Gas-phase chemistry follows MOZART-TS2 (Model of Ozone And Related chemical Tracers, Troposphere-Stratosphere V2), including comprehensive isoprene and monoterpenes chemistry (Schwantes et al., 2020). Aerosols are represented using the four-mode version of the Modal Aerosol Module (MAM4) (Liu et al., 2016). This study employs two different approaches to treat SOA formation. First, the VBS approach (Donahue et al., 2006; Hodzic et al., 2016) represents the gas-
140 particle interconversion of organic compounds. Organic compounds (including glyoxal, monoterpenes, sesquiterpenes, benzene, toluene, lumped xylenes, IVOCs, and SVOCs) are oxidized to form gas-phase SOA precursors in five volatility bins, with effective saturation concentrations (C^* , at 300 K) of 0.01, 0.1, 1.0, 10.0, and 100.0 $\mu\text{g m}^{-3}$. These precursors subsequently partition between the gas and particle phases (Tilmes et al., 2019). This VBS treatment follows the default CAM6-Chem SOA scheme and was not modified in this study. Second, SOA formation from isoprene gas phase products is
145 treated explicitly (Sect. 2.1.2). In addition to ISOA, all other SOA are simulated using the VBS approach. Photolysis rates for monoterpene-derived SOA were updated based on our previous work (Liu et al., 2023). Aerosol wet removal scheme uses the Cloud Layers Unified By Binormals (CLUBB) scheme to provide a unified treatment of shallow convection and stratiform clouds, coupled with the two-moment cloud microphysics scheme by Gettelman and Morrison (2015) (MG2) to represent aerosol activation and removal (Gettelman and Morrison, 2015). This study adopts the Zhang and McFarlane (1995)
150 (ZM95) parameterization for deep convective clouds and treats aerosol wet scavenging using empirical coefficients (Zhang & McFarlane, 1995).

The earlier CAM6-Chem configuration described by Zhang et al. (2025) represents ISOA formation mainly through the low-
NO_x pathway, in which ISOA_{IEPOX} is treated as a lumped product. This lumped low-NO_x representation limits the model's
155 ability to describe ISOA formation across different NO_x regimes and to distinguish product-level responses to changes in the particle-phase reaction environment. To address this limitation, Sect. 2.1.2 describes the explicit ISOA representation adopted in this study. This explicit representation introduces high-NO_x epoxide precursors and their subsequent heterogeneous uptake, allowing the model to characterize ISOA formation in urban and polluted regions where the high-NO_x pathway can contribute substantially to total ISOA. It also resolves reactive-uptake-derived ISOA into product classes
160 formed via H₂O-mediated nucleophilic addition and inorganic-nucleophile addition (involving sulfate, nitrate, and bisulfate ions), thereby improving the model's ability to represent ISOA compositional changes associated with aerosol liquid water, inorganic nucleophiles, and acidity. Overall, this explicit ISOA formation mechanism enhances the model's sensitivity to changes in anthropogenic emissions and climate conditions.

2.1.2 New processes added to isoprene chemistry

165 Under low-NO_x conditions, the MOZART-TS2 gas-phase mechanism is retained, in which ISOPOOH reacts with OH to produce IEPOX with a molar branching ratio of 0.85. Under high-NO_x conditions, following Lin et al. (2013), the gas-phase isoprene chemistry is expanded to explicitly represent MPAN and its subsequent reaction with OH (Lin et al., 2013). Methacrolein (MACR), the precursor of MPAN in the high-NO_x pathway, is formed following the default MOZART-TS2 mechanism. The reaction of MPAN with OH yields HMML and MAE, with molar branching ratios of 0.75 and 0.02, 170 respectively, based on the experimental results reported by Nguyen et al. (2015) (Table S2). No explicit threshold is imposed to separate low-NO_x and high-NO_x regimes. Pathway contributions are determined dynamically by competition between RO₂ reacting with NO and RO₂ reacting with HO₂. To avoid double-counting between the default lumped VBS representation of ISOA and the explicitly represented epoxide-derived ISOA pathways, the OH-initiated isoprene oxidation pathway in the VBS scheme is removed (see Sect. 2.1.1). This treatment focuses the updated mechanism on epoxide-derived 175 ISOA formation, although potential non-epoxide ISOA contributions originally included implicitly in the VBS scheme may not be explicitly resolved.

The reactive uptake parameterization and condensed-phase reaction constants used in this study were adopted from the IEPOX reactive uptake framework of Jo et al. (2019, 2021). Due to limited laboratory constraints on the multiphase kinetics 180 of HMML and MAE, HMML and MAE are treated as IEPOX-like epoxide precursors in the current implementation. This treatment is consistent with previous modeling practice. For example, Pye et al. (2013) treated HMML like MAE in terms of heterogeneous uptake and assumed the particle-phase reaction rate constants of MAE to be the same as those of IEPOX due to the lack of kinetic data. Zhang et al. (2022) also noted that most HMML-related reaction parameters in models are commonly assumed to be the same as those for IEPOX. Following these studies, HMML and MAE are assigned the same 185 solubility, diffusivity, and condensed-phase reaction constants as IEPOX in our implementation. Although a recent field-based study has estimated effective uptake parameters for IEPOX and lumped HMML/MAE, these constraints are not directly applicable here because they do not provide a complete set of standalone, compound-specific solubility, diffusivity, and branch-resolved condensed-phase reaction constants for HMML and MAE (Zhu et al., 2025). As a result, this IEPOX-like treatment introduces uncertainty in the absolute uptake rates of HMML and MAE and may affect the modeled 190 quantitative partitioning between high-NO_x and low-NO_x ISOA formation pathways. This study utilizes the Model for Simulating Aerosol Interactions and Chemistry (MOSAIC) aerosol module (Jo et al., 2019, 2021; Zaveri et al., 2008, 2021) to calculate the dynamic partitioning of H₂SO₄, HNO₃, HCl, and NH₃ across different modes and the associated particle-phase thermodynamics. Aerosol pH for each mode is calculated online using MOSAIC, as integrated into CESM by Zaveri et al. (2021) and Lu et al. (2021) (Lu et al., 2021; Zaveri et al., 2021). Additionally, following Jo et al. (2019), we used the 195 modified version of MOSAIC, which calculates submicron (aitken and accumulation modes) aerosol pH, excluding sea salt

(Jo et al., 2019). Subsequently, following the resistor model equation of Gaston et al. (2014), this study calculates the reactive uptake coefficient γ for IEPOX, HMML, and MAE (Gaston et al., 2014). The governing expression is:

$$\frac{1}{\gamma} = \frac{\omega R_p}{4D_{\text{gas}}} + \frac{1}{\alpha} + \frac{\omega R_p}{4RT H_{\text{org}} D_{\text{org}} (q_{\text{org}}^* F - 1)}, \quad (1)$$

Where γ represents the reactive uptake coefficient, ω is the mean molecular speed of epoxides (m s^{-1}), R_p is the particle radius (m), D_{gas} is the gas-phase diffusion coefficient of epoxides ($10^{-5} \text{ m}^2 \text{ s}^{-1}$), α is the mass accommodation coefficient (0.1), R is the universal gas constant ($8.2057 \times 10^{-2} \text{ L atm mol}^{-1} \text{ K}^{-1}$), T is temperature (K), H_{org} is the Henry's law coefficient in the organic layer ($2 \times 10^6 \text{ M atm}^{-1}$), and D_{org} is the diffusion coefficient of epoxides in the organic layer. The term F is calculated as:

$$F = \frac{\coth(q_{\text{org}}) + h(q_{\text{aq}}, q_{\text{org}}^*)}{1 + \coth(q_{\text{org}}) h(q_{\text{aq}}, q_{\text{org}}^*)}, \quad (2)$$

Where the function $h(q_{\text{aq}}, q_{\text{org}}^*)$ is given by:

$$h(q_{\text{aq}}, q_{\text{org}}^*) = -\tanh(q_{\text{org}}^*) \frac{\frac{H_{\text{aq}} D_{\text{aq}}}{H_{\text{org}} D_{\text{org}}} (q_{\text{aq}} \coth(q_{\text{aq}}) - 1) - (q_{\text{org}}^* \coth(q_{\text{org}}^*) - 1)}{\frac{H_{\text{aq}} D_{\text{aq}}}{H_{\text{org}} D_{\text{org}}} (q_{\text{aq}} \coth(q_{\text{aq}}) - 1) - (q_{\text{org}}^* \tanh(q_{\text{org}}^*) - 1)}, \quad (3)$$

Here, H_{aq} is the Henry's law coefficient in the aqueous core ($1.7 \times 10^7 \text{ M atm}^{-1}$), and D_{aq} is the diffusion coefficient of epoxides in the aqueous core ($10^{-9} \text{ m}^2 \text{ s}^{-1}$). The variables q_{org} , q_{aq} , and q_{org}^* are defined as:

$$q_{\text{org}} = R_p \sqrt{\frac{k_{\text{org}}}{D_{\text{org}}}}, \quad (4)$$

$$q_{\text{aq}} = R_c \sqrt{\frac{k_{\text{aq, total}}}{D_{\text{aq}}}}, \quad (5)$$

$$q_{\text{org}}^* = \frac{R_c}{R_p} q_{\text{org}}, \quad (6)$$

Where R_c is the inorganic aqueous core radius (m), k_{org} is the pseudo-first-order reaction rate constant of epoxides in the organic shell (s^{-1}), and $k_{\text{aq, total}}$ is the total pseudo-first-order aqueous-phase reaction rate constant (s^{-1}), calculated as the sum of all particle-phase reaction channels that consume each precursor:

$$k_{\text{aq, total}} = k_{\text{H}^+} [\text{H}^+] + k_{\text{nuc}} ([\text{SO}_4^{2-}] + [\text{NO}_3^-]) [\text{H}^+] + k_{\text{HSO}_4^-} [\text{HSO}_4^-], \quad (7)$$

The first term represents the acid-catalyzed ring-opening followed by nucleophilic addition of aerosol liquid water, leading to H_2O -mediated nucleophilic addition products. Here, k_{H^+} is the corresponding effective H_2O -mediated nucleophilic addition rate constant ($0.036 \text{ M}^{-1} \text{ s}^{-1}$), with aerosol liquid water treated as the solvent and included implicitly, and $[\text{H}^+]$ is the proton concentration (M). The second term represents acid-catalyzed ring opening followed by nucleophilic addition of

220 sulfate and nitrate ions, leading to organosulfates and organonitrates. Here, k_{nuc} is the third-order rate constant for sulfate- and nitrate-addition reactions ($2 \times 10^{-4} \text{ M}^{-2} \text{ s}^{-1}$), and $([\text{SO}_4^{2-}] + [\text{NO}_3^-])$ is the concentration of sulfate and nitrate ions (M). The third term represents concerted protonation and nucleophilic addition by bisulfate, leading to organosulfates. $k_{\text{HSO}_4^-}$ is the reaction rate constant due to the presence of bisulfate ($7.3 \times 10^{-4} \text{ M}^{-1} \text{ s}^{-1}$), and $[\text{HSO}_4^-]$ is the bisulfate concentration (M).

225 For each precursor considered here, including IEPOX, HMML, and MAE, $k_{\text{aq,total}}$ is used to calculate the reactive uptake coefficient and the total uptake flux. The resulting particle-phase ISOA production is then distributed into the H_2O -mediated nucleophilic addition and inorganic-nucleophile addition product classes according to the relative contribution of each reaction channel to $k_{\text{aq,total}}$. The H_2O -mediated nucleophilic addition class is represented by the $k_{\text{H}^+}[\text{H}^+]$ channel, whereas the inorganic-nucleophile addition class combines the sulfate-, nitrate-, and bisulfate-assisted channels, which are collectively represented by $k_{\text{nuc}}([\text{SO}_4^{2-}] + [\text{NO}_3^-])[\text{H}^+] + k_{\text{HSO}_4^-}[\text{HSO}_4^-]$. The corresponding branching ratios for the two
230 product classes are defined as $\text{BR}_{\text{H}_2\text{O}}$ and $\text{BR}_{\text{SO}_4/\text{NO}_3/\text{HSO}_4}$, respectively:

$$\text{BR}_{\text{H}_2\text{O}} = \frac{k_{\text{aq,H}_2\text{O}}}{k_{\text{aq,total}}} = \frac{k_{\text{H}^+}[\text{H}^+]}{k_{\text{aq,total}}} \quad (8)$$

$$\text{BR}_{\text{SO}_4/\text{NO}_3/\text{HSO}_4} = \frac{k_{\text{aq,SO}_4/\text{NO}_3/\text{HSO}_4}}{k_{\text{aq,total}}} = \frac{k_{\text{nuc}}([\text{SO}_4^{2-}] + [\text{NO}_3^-])[\text{H}^+] + k_{\text{HSO}_4^-}[\text{HSO}_4^-]}{k_{\text{aq,total}}} \quad (9)$$

This relative-rate branching treatment follows the approach of Budisulistiorini et al. (2017), in which ISOA_{IEPOX} speciation
235 between tetrols and organosulfates was determined from the relative rates of precursor conversion to these products. This study follows the approach of Jo et al. (2021), where the reactivity of epoxides in the organic shell was parameterized with the same reaction rate constant as in the aqueous phase ($k_{\text{org}} = k_{\text{aq,total}}$). This treatment follows the Jo et al. (2019, 2021) core-shell framework and represents a simplifying assumption, because phase separation, organic-shell viscosity, and differences in acidity between the inorganic core and organic shell may alter epoxide diffusion and reactivity (Schmedding et al., 2020; Farrell et al., 2025). Furthermore, considering the strong sensitivity of D_{org} to relative humidity (RH) in the
240 atmosphere, this study accounts for the RH dependence of D_{org} , as described in Table S3 of Zhang et al. (2018) (Zhang et al., 2018).

As illustrated in Fig. 1, the uptake-derived ISOA product classes are mapped to four model species. ISOA formed from the
245 heterogeneous uptake of IEPOX is divided into two product classes: 2-MT (AIETET), produced through the H_2O -mediated nucleophilic addition channel, and organosulfate/organonitrate products (AIEOSN), produced through the inorganic-nucleophile addition channel. Similarly, ISOA formed from HMML and MAE is divided into 2-MG (AHMGA), produced through the H_2O -mediated nucleophilic addition channel, and organosulfate/organonitrate products (AHMOSN), produced through the inorganic-nucleophile addition channel. AIEOSN and AHMOSN are lumped product classes that include both
250 organosulfates and organonitrates but may be dominated by organosulfates. Recent kinetic evidence indicates that HSO_4^- is substantially less nucleophilic toward IEPOX than SO_4^{2-} (Cooke et al., 2024), while nitrate is weaker than sulfate in related

epoxide systems (Mael et al., 2015). Tertiary organonitrates can also be rapidly hydrolyzed by water to form polyols or undergo nucleophilic substitution by sulfate to form organosulfates (Darer et al., 2011). The internal organosulfate/organonitrate composition of these lumped classes is not explicitly resolved in the model. The ISOA yield from the reactive uptake of IEPOX, HMML, and MAE is assumed to be 100%, and ISOA derived from these epoxides is treated as non-volatile in the model. This assumption is in agreement with previous modeling studies (Budisulistiorini et al., 2017; Marais et al., 2016; Schmedding et al., 2019; Stadtler et al., 2018), which are based on field observations indicating that ISOA from these epoxides in the atmosphere exhibits very low volatility (Hu et al., 2016; Riva et al., 2019). Once ISOA from IEPOX, HMML, and MAE is formed, no further oxidation is represented in the model, consistent with previous modeling studies (Budisulistiorini et al., 2017; Marais et al., 2016; Schmedding et al., 2019). However, this assumption represents a limitation in the treatment of 2-methyltetrol sulfates (2-MTS), major IEPOX-derived organosulfates represented within the lumped AIEOSN species, because 2-MTS have been shown to undergo heterogeneous OH oxidation in laboratory-generated aerosol particles (Chen et al., 2020; Xu et al., 2024) and aqueous-phase OH oxidation in cloud, fog, and aerosol water mimics (Royer et al., 2026). Several oxidation products have also been detected in ambient fine aerosols (Chen et al., 2020).

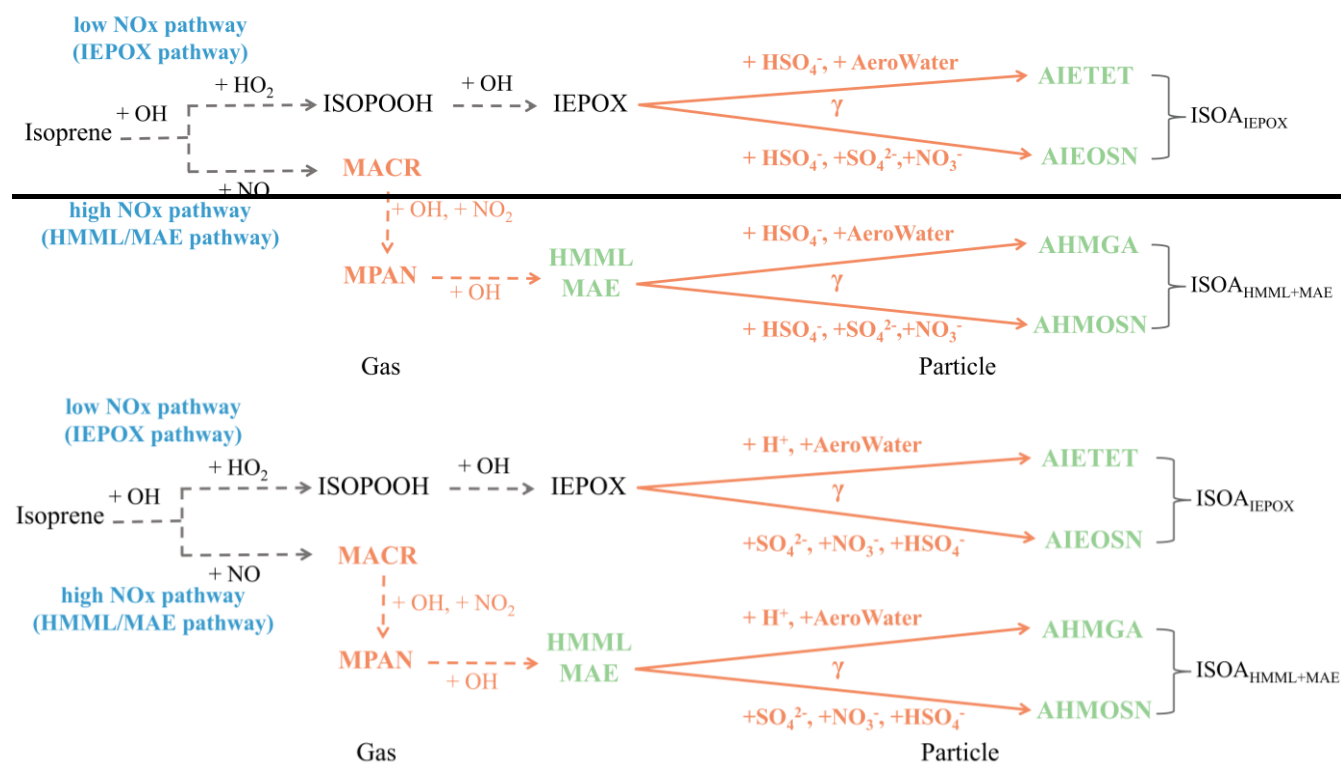


Figure 1: Schematic of the explicit isoprene-derived secondary organic aerosol (ISOA) formation mechanism implemented in CAM6-Chem in this study. Dashed arrows indicate gas-phase reactions, whereas solid arrows indicate particle-phase

heterogeneous uptake and product-formation reactions. Species explicitly represented as updated ISOA precursors or product classes in this mechanism are shown in green, and the explicitly represented ISOA formation pathways are shown in orange. AeroWater denotes aerosol liquid water. AIETET denotes IEPOX-derived 2-methyltetrols, and AHMGA denotes HMML/MAE-derived 2-methylglyceric acid. AIEOSN and AHMOSN denote lumped inorganic-nucleophile addition product classes from the IEPOX and HMML/MAE pathways, respectively. These classes include both organosulfates and organonitrates but may be dominated by organosulfates. ~~AIETET denotes IEPOX-derived 2-methyltetrols, AIEOSN denotes IEPOX-derived organosulfate/organonitrate products, AHMGA denotes HMML/MAE-derived 2-methylglyceric acid, and AHMOSN denotes HMML/MAE-derived organosulfate/organonitrate products.~~

2.2 Model experiments and observational data

This study simulates the years 2000, 2006, 2012, 2016, and 2019 at a horizontal resolution of 0.95° (latitude) $\times$ 1.25° (longitude). The vertical grid comprises 32 layers with a model top of around 40 km. For each simulated year, a 3-month spin-up is applied to minimize the influence of initial conditions, and a 50 h relaxation (nudging) timescale is used throughout the simulation.

This study uses ground-based observations reported by Ding et al. (2016) (Ding et al., 2016). The record spans October 2012 to September 2013 and provides annual and seasonal mass concentrations of two ISOA tracers, AIETET and AHMGA. The dataset comprises twelve sites, including five urban, three suburban, and four rural locations, and provides the geographic coordinates for each site. In total, 294 sets of field samples were compiled across the four seasons and summarized by site and by season for both tracers, enabling comparison with the model on consistent temporal and spatial scales. In addition, we used observations reported by Zhang et al. (2022), who conducted a 1-year measurement campaign from October 2013 to November 2014 at three urban sites from northern to southern China, namely Beijing, Hefei, and Kunming (Zhang et al., 2022). This dataset reports paired ISOA products formed through the IEPOX and HMML pathways, and provides site-resolved annual-average concentrations for key species, which are useful for evaluating the simulated ISOA composition and pathway partitioning over urban China. In addition, we used ground-based observational compilations assembled by Miao et al. (2021) and Chen et al. (2024a). These compilations summarize site-mean surface mass concentrations of OA, POA, and SOA across China for 2013–2019, together with the corresponding site location (Chen et al., 2024a; Miao et al., 2021). After selecting records within our study period and removing duplicates, a total of 39 measurements were retained for model evaluation.

3. Results and Discussions

3.1 Evaluation of model performance

By comparing the simulated results with ground-based observations, we evaluated the performance of the CAM6-Chem configuration using the explicit ISOA representation adopted in this study (Fig. 2). For all normalized mean biases (NMB) calculations in this section, model outputs were sampled at the corresponding observational site locations and averaged over the same observational periods before calculating the model-observation differences. The paired model and observational

305 values from different sites and time periods were then combined across all paired samples to calculate the NMB, thereby accounting for the site-specific locations and different temporal coverage of the observational datasets. Overall, the mechanism-updated model can more comprehensively capture the concentrations of key ISOA subspecies and better reproduce their relative contributions (Fig. 2(c)). Additionally, it alleviates the systematic underestimation of OA and SOA in the earlier CAM6-Chem configuration (Zhang et al., 2025). We evaluated the simulated AIETET and AHMGA against
310 ground-based observations from Ding et al. (2016) (Fig. 2(a)). The simulation using this explicit ISOA representation generally captures the concentration ranges and variability of these two subspecies, but still shows an overall positive bias. Based on monthly means in log space, the NMB values are 61.41% for AIETET and 89.18% for AHMGA, with corresponding Pearson correlation coefficients (r) of 0.75 and 0.54, respectively. Several factors may contribute to these positive biases. First, missing or simplified competing pathways may cause an over-allocation of epoxide precursors to the explicitly represented ISOA products. For example, a fraction of IEPOX-derived carbon may form structurally characterized $C_5H_{10}O_3$ reactive uptake products (Frauenheim et al., 2022). These semivolatile products can repartition from the particle phase to the gas phase and undergo gas-phase OH oxidation to form additional isoprene-derived SOA (Frauenheim et al., 2024). This competing product branch and its subsequent gas-phase oxidation are not explicitly represented in the model. Omission of this branch may therefore lead to excessive allocation of IEPOX-derived carbon to AIETET, contributing to its positive bias.
320 ~~First, missing or simplified competing loss pathways may cause an over-allocation of epoxide precursors to the explicitly represented ISOA products. For example, a fraction of IEPOX-derived products in the real atmosphere may be diverted to C5-alkene triols or other particle-phase products that are not explicitly represented here, thereby reducing the effective yield of AIETET (Pye et al., 2013).~~ In addition, previous studies have shown that AIETET and AHMGA are sensitive to factors such as pathway branching, mass transfer, reaction kinetics, and aerosol phase state. For example, uncertainties in the relative rates of water addition versus organosulfate formation can affect the predicted product distribution between 2-MT and organosulfate products (Budisulistiorini et al., 2017). Similarly, assumptions related to aerosol phase state, organic-shell viscosity, and phase separation can influence IEPOX reactive uptake and may affect the simulated abundance of IEPOX-derived tracers (Zhang et al., 2018, 2019; Chen et al., 2024b). In particular, the current core-shell uptake treatment assumes the same reaction rate coefficient in the organic shell and aqueous core. If phase separation
325 and viscous organic coatings impose stronger diffusion limitations than represented here, effective epoxide uptake may be reduced in the real atmosphere, and the model may overestimate ISOA formation, especially under conditions with high organic coatings or strong phase separation. For AHMGA, uncertainties in the multiphase uptake and reaction parameters of HMML/MAE may also contribute to the model bias, because compound-specific constraints for these high-NO_x epoxides remain limited (Zhang et al., 2022; Zhu et al., 2025). Accordingly, positive biases of several tens of percent have also been
330 reported in earlier modeling studies. For instance, Fahey et al. (2017) reported that, in a regional evaluation using the Community Multiscale Air Quality (CMAQ) model, the NMB of AHMGA reached 78.6% under their revised scheme (Fahey et al., 2017), whereas Chen et al. (2024b) found a pronounced overestimation of AIETET in a baseline configuration (NMB = 58%) in their evaluation of IEPOX multiphase parameterization (Chen et al., 2024b). Therefore, the magnitude of

the overestimation in this study is comparable to that reported previously, although the exact sources of bias may differ across models and configurations. It further suggests that incorporating a more complete set of competitive loss processes and additional constraints on multiphase chemistry may help reduce subspecies-level biases in future work.

At the bulk OA and SOA levels, we further evaluated model performance against observations from Miao et al. (2021) and Chen et al. (2024a) (Chen et al., 2024a; Miao et al., 2021) by comparing the simulated results from the earlier CAM6-Chem configuration described by Zhang et al. (2025) with those from the CAM6-Chem configuration using the explicit ISOA representation adopted in this study (Fig. 2(b)). The results show that the explicit ISOA representation implemented in this study substantially reduces the underestimation of SOA, with the NMB improving from -76.7% to -51.6%. This improvement reflects the enhanced representation of ISOA formation introduced in this study, since other SOA formation pathways were not modified relative to the earlier CAM6-Chem configuration. The underestimation of total OA is also alleviated, with the NMB improving from -47.0% to -25.7%. Although the available observations do not support a robust spatiotemporal decomposition of this improvement, the model results suggest that the explicit ISOA representation may affect bulk OA/SOA simulations more strongly in regions with higher anthropogenic NO_x emissions (Fig. S2(c)), where high-NO_x pathway products contribute more substantially to ISOA (Fig. 3(b) and Fig. 3(c)). Beyond this bulk evaluation, we further assessed the relative composition of modeled ISOA product classes using ground-based measurements of ISOA tracers from Beijing, Hefei, and Kunming (Zhang et al., 2022) (Fig. 2(c)). The comparison shows that the explicit ISOA representation adopted in this study can reproduce the relative contributions of the major observed ISOA species and yields a reasonable compositional distribution consistent with the available observations. In particular, the simulation using this explicit ISOA representation reproduces the dominant contribution of AIETET and the relatively small contributions of inorganic-nucleophile addition class products, including AIEOSN and AHMOSN. This result indicates that the product-speciation treatment captures the observed dominance of H₂O-mediated nucleophilic addition products in the available molecular tracer composition, which is also consistent with previous field observations showing relatively small contributions from inorganic-nucleophile-addition products (He et al., 2018). Although 2-MTS can contribute substantially to fine-particle organic carbon in some regions (Hettiyadura et al., 2019), available observations of 2-MTS and 2-methylglyceric acid sulfate in China have limited spatial coverage, and the two compounds are not separately resolved within the lumped AIEOSN and AHMOSN classes, respectively. Accordingly, Fig. 2(c) provides a product-class-level comparison of relative contributions rather than a species-specific evaluation of absolute concentrations. Residual discrepancies in individual product fractions may also reflect uncertainties in the reactive uptake parameterization, possible missing transformation or loss processes, and the representation of the heterogeneous reaction medium, including the effective abundance and accessibility of different inorganic nucleophiles. This improvement is important because it demonstrates that the explicitly represented reaction pathways and product-speciation treatment are chemically meaningful and enhance the model's capability to reproduce both the magnitude and composition of ambient ISOA, providing a more robust modeling foundation for the subsequent quantitative attribution of ISOA changes and their driving factors.

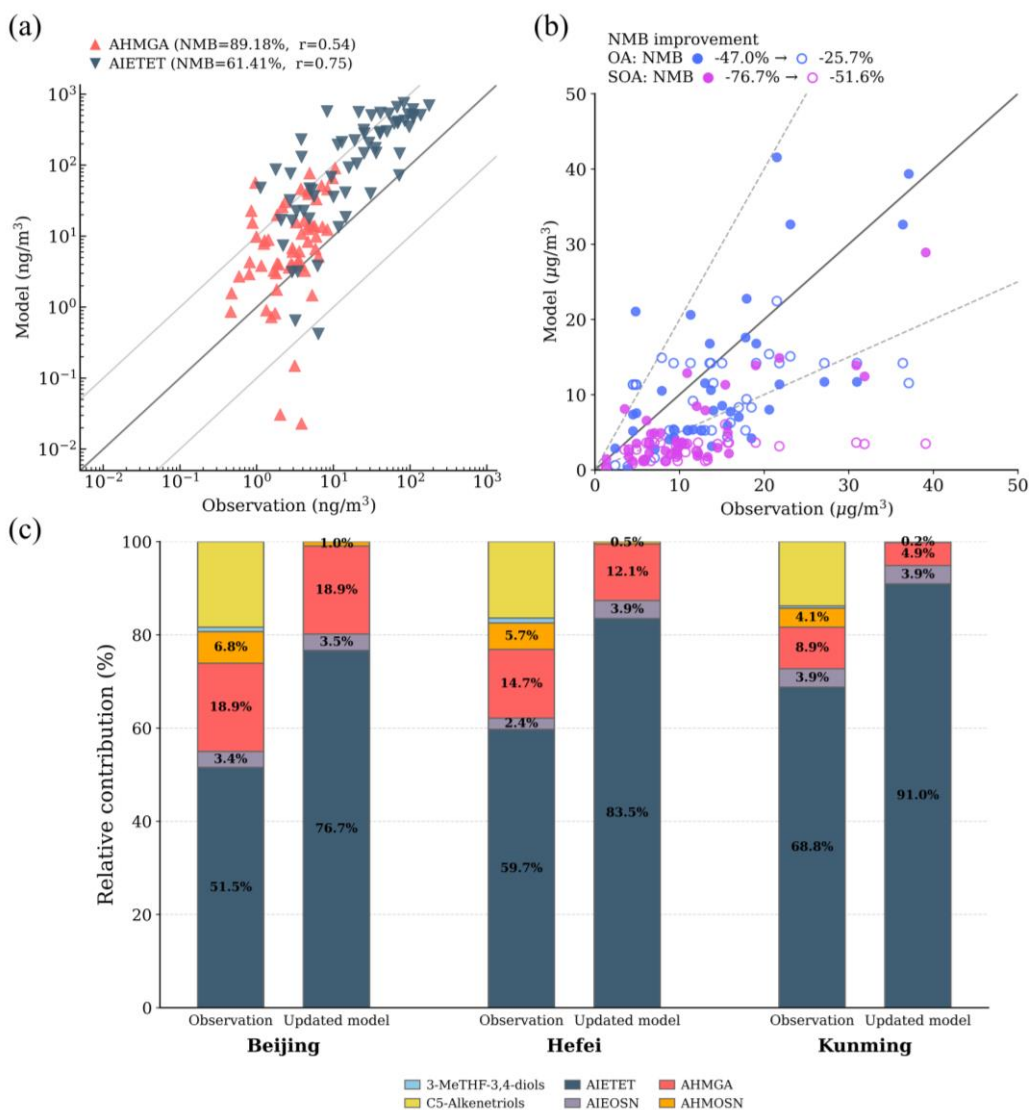


Figure 2: (a) Evaluation of simulated 2-methyltetrols (AIETET) and 2-methylglyceric acid (AHMGA) concentrations against ground-based observations from Ding et al. (2016) (ng m^{-3}). Statistical metrics were calculated using monthly means in log space. The dark gray solid line denotes the 1:1 line, and the light gray solid lines denote the factor-of-10 range. (b) Evaluation of simulated organic aerosol (OA) and secondary organic aerosol (SOA) concentrations before and after the mechanism update against ground-based observations compiled from Miao et al. (2021) and Chen et al. (2024a) ($\mu\text{g m}^{-3}$). Filled circles represent simulations before the update, while open circles represent simulations after the update. Blue and magenta markers denote OA and SOA, respectively. The inset annotation summarizes the changes in normalized mean bias (NMB) from the original to the updated simulation. The dark gray solid line denotes the 1:1 line, and the light gray dashed lines denote the factor-of-two range. (c) Comparison of ISOA product composition between ground-based observations adapted from Zhang et al. (2022) and the updated model simulation in this study. The grouped stacked bars show the relative contribution of each ISOA product to total measured or simulated ISOA at Beijing, Hefei, and Kunming. For each city, the left bar represents the observed composition and the right bar represents the updated model.

3.2 Spatial and seasonal characteristics of ISOA and pathway contributions

The simulations show pronounced regional contrasts in the spatial distribution of ISOA over China (Fig. 3(a)). Total ISOA forms a major hotspot in Southwest China ($>1000 \text{ ng m}^{-3}$), remains relatively high in Southeastern China, and is generally lower over North China and Northwestern China. In terms of composition, AIETET is the dominant subspecies and accounts for the largest national fraction (88.9%). AHMGA is the second-largest contributor (6.9%). The spatial patterns of these two subspecies are broadly consistent with the hotspot of total ISOA (Fig. S1). In contrast, AIEOSN (3.9%) and AHMOSN (0.3%) contribute less to the national burden on average. They are more pronounced in typical high-emission regions in Eastern China. This enhancement is especially evident over major urban and industrial clusters such as the North China Plain, the Yangtze River Delta, and the Pearl River Delta (Fig. 3(c)). In these NO_x-rich regions, the relative contributions of AHMGA and AHMOSN increase markedly. Together they account for approximately 17–21% of the total ISOA (Fig. 3(b)). These results indicate that the explicit ISOA representation adopted in this study enables a more comprehensive characterization of ISOA composition and concentration levels by resolving uptake-derived ISOA into different product classes. It is also noteworthy that the concentration differences of AHMGA and AHMOSN between polluted regions and the southwestern hotspot are smaller than those of the low-NO_x pathway products (Fig. S1). This feature indicates a relatively more spatially distributed pattern for these high-NO_x pathway products. Overall, the subspecies collectively characterize the spatial heterogeneity of ISOA. Low-NO_x pathway components dominate the national burden and define the primary hotspot. High-NO_x pathway components contribute less to the national average burden, but their relative contributions are more evident over eastern urban and industrial regions (Fig. 3(c)).

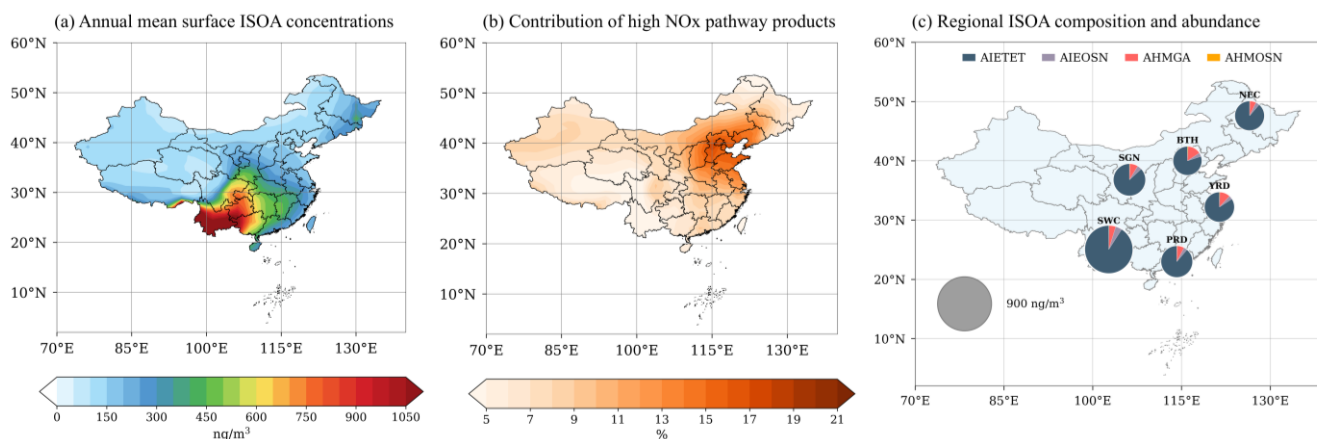


Figure 3: (a) Annual mean surface concentrations of isoprene-derived secondary organic aerosol (ISOA) over China, averaged over the five simulation years (2000, 2006, 2012, 2016, and 2019) (unit: ng m^{-3}). (b) Spatial distribution of the contribution of high-NO_x pathway products, including 2-methylglyceric acid (AHMGA) and organosulfate/organonitrate products from the high-NO_x pathway (AHMOSN), to annual mean total ISOA over China, averaged over the five simulation years (unit: %). (c) Composition of surface ISOA subspecies in six regions of China, including Southwest China (SWC), the Beijing–Tianjin–Hebei region (BTH), the Yangtze River Delta (YRD), the Pearl River Delta (PRD), the Shaanxi–Gansu–Ningxia region (SGN), and Northeast China

(NEC), based on annual mean concentrations averaged over the five simulation years. Pie charts show the fractional contributions of individual ISOA subspecies, and the circle size represents the corresponding surface ISOA concentration (unit: ng m^{-3}).

415 In addition to the pronounced spatial heterogeneity, ISOA also shows strong seasonal variability (Fig. 4). Based on the 5-year monthly means, surface ISOA increases markedly during the warm season and reaches a maximum in summer. Summer ISOA concentrations are approximately 3 to 5 times higher than those in winter (Fig. 4(a)). The compositional information in Fig. 4(a) further shows that the warm-season enhancement of ISOA is dominated by $\text{ISOA}_{\text{IEPOX}}$, particularly AIETET. In contrast, AIEOSN and $\text{ISOA}_{\text{HMML+MAE}}$ remain minor components of the national mean ISOA, although they also exhibit

420 higher concentrations during the warm season. Overall, the seasonal cycle of total ISOA at the national scale is primarily driven by AIETET. Biogenic isoprene emissions exhibit a similar seasonal cycle. They increase rapidly from spring and peak in summer, which is generally consistent with the timing of the ISOA maximum (Fig. 4(a)). The standardized seasonal cycle further shows that the standardized OH signal also peaks in summer, indicating that enhanced precursor supply and oxidation capacity jointly support the summer ISOA maximum (Fig. 4(b)). However, the standardized sulfate and aerosol

425 liquid water anomalies show different seasonal patterns from total ISOA. The standardized sulfate and aerosol liquid water anomalies are positive in winter and become negative in summer, while the standardized aerosol pH shows positive anomalies in late spring and early summer (Fig. 4(b)). Therefore, the summer ISOA maximum occurs even when sulfate and aerosol liquid water are below their annual-mean levels in the standardized seasonal cycle, suggesting that the strong increases in isoprene emissions and OH oxidation capacity outweigh the less favorable multiphase conditions during summer.

430 The early warm-season ISOA enhancement from April to May further illustrates this combined-control behavior, but with a different balance among the controlling factors. It occurs before the annual maximum of isoprene emissions and during the rising phase of the standardized seasonal OH signal, but at a time when biogenic isoprene emissions have already increased substantially from spring, OH oxidation capacity is strengthening, and the multiphase reaction environment has not yet reached its least favorable summer state. Thus, the May enhancement likely reflects the combined influence of increasing

435 precursor supply, enhanced oxidation, and still favorable multiphase conditions, rather than the dominance of a single factor. Overall, these features indicate that the seasonal cycle of total ISOA cannot be explained by one controlling factor alone, but instead reflects the combined effects of emission strength, oxidation intensity, and multiphase chemical conditions.

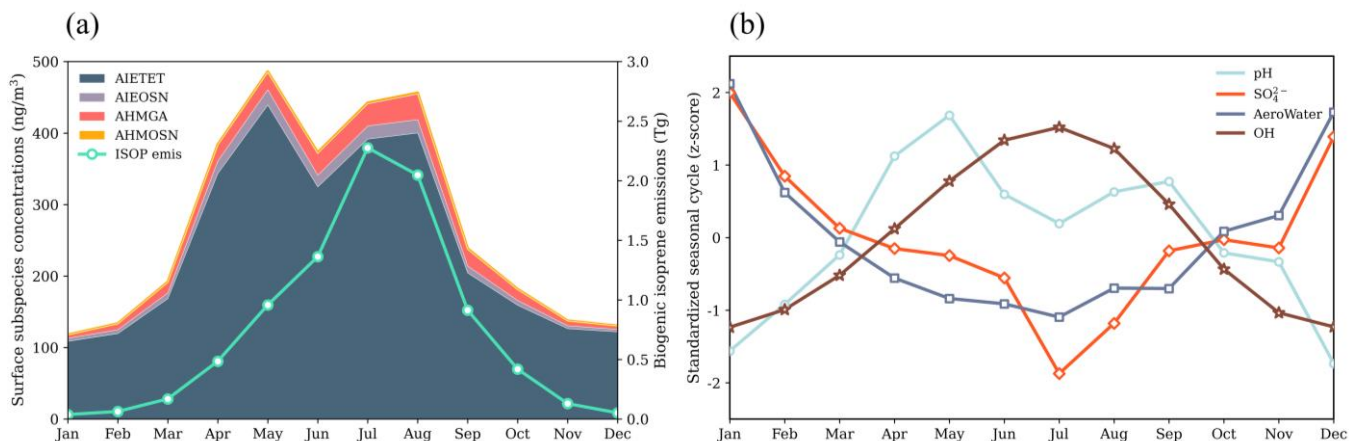


Figure 4: (a) Monthly variations of surface isoprene-derived secondary organic aerosol (ISOA) subspecies (left y axis; unit: ng m^{-3}) and biogenic isoprene emissions (ISOP emis; right y axis; unit: Tg) in China, based on the climatological monthly means averaged over the five simulation years (2000, 2006, 2012, 2016, and 2019). (b) Standardized monthly variations of background chemical and aerosol variables relevant to ISOA formation, including surface aerosol pH, surface sulfate (SO_4^{2-}) concentrations, surface aerosol liquid water (AeroWater) concentrations, and surface hydroxyl radical (OH) concentrations in China, based on the climatological monthly means averaged over the five simulation years.

ISOA formation is strongly influenced by the NO_x background, so competition exists between the IEPOX pathway and the HMML/MAE pathway. Overall, the HMML/MAE pathway contributes less to ISOA than the IEPOX pathway, but its relative importance varies across regions and seasons.

We use ratio-based metrics to quantify the relative contributions of the two pathways. For the national mean $\text{ISOA}_{\text{HMML+MAE}}/\text{ISOA}_{\text{IEPOX}}$, the annual mean ratio is about 0.07 (Fig. 5(c)), indicating that the IEPOX pathway dominates at the national scale. A further analysis for densely populated regions such as the Beijing–Tianjin–Hebei region shows a higher ratio that still remains well below 1 (Fig. S7). This indicates that the IEPOX pathway remains the primary source of ISOA even in NO_x -rich regions.

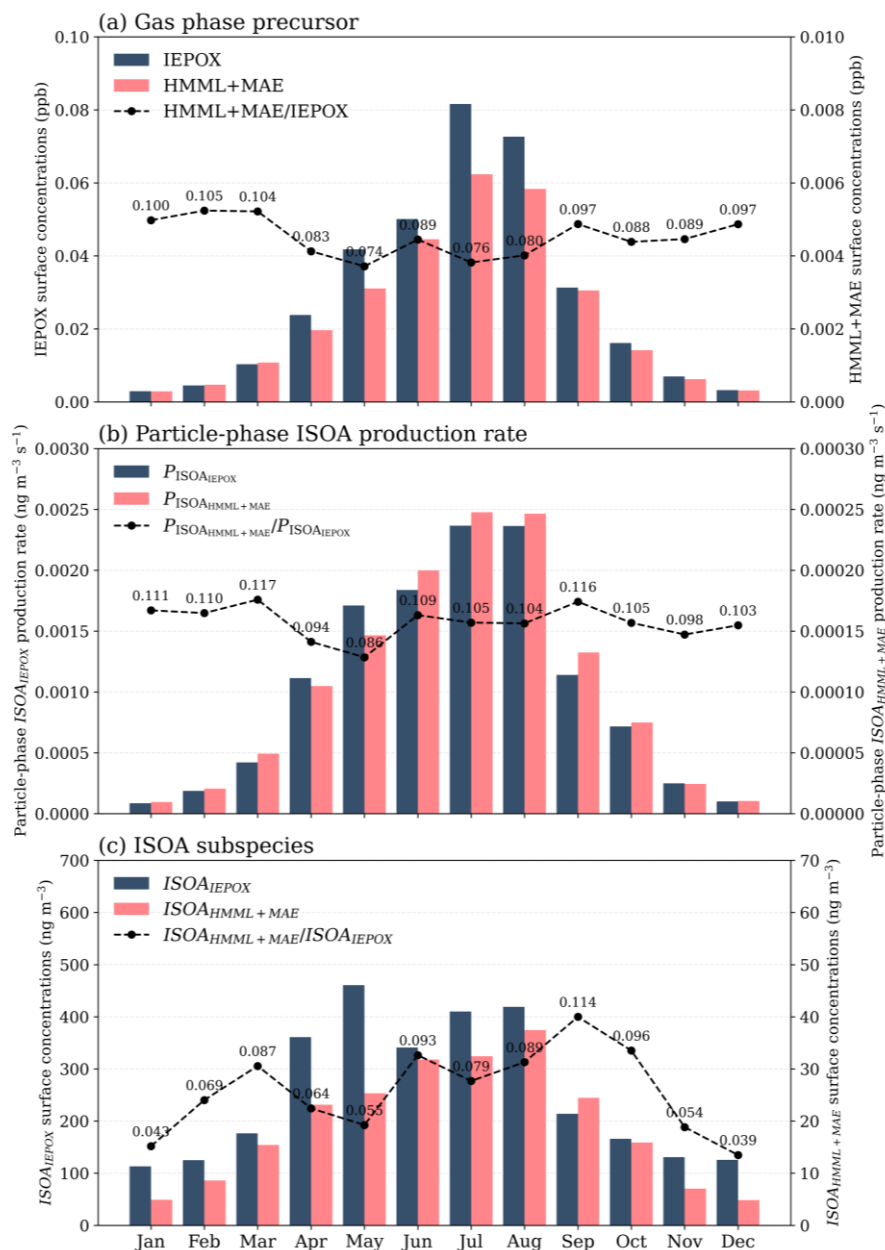


Figure 5: Monthly variations of IEPOX-related and HMML+MAE-related gas-phase precursors, particle-phase production rates, and ISOA concentrations in China, based on the climatological monthly means averaged over the five simulation years (2000, 2006, 2012, 2016, and 2019). (a) Surface concentrations of isoprene epoxydiols (IEPOX; dark-blue bars; left y axis; unit: ppb) and hydroxymethyl-methyl- α -lactone plus methacrylic acid epoxide (HMML + MAE; pink bars; right y axis; unit: ppb). (b) Particle-phase ISOA production rates diagnosed from tagged heterogeneous reactions for the IEPOX pathway ($P_{ISOA_{IEPOX}}$; dark-blue bars; left y axis; unit: $ng\ m^{-3}\ s^{-1}$) and the HMML+MAE pathway ($P_{ISOA_{HMML+MAE}}$; pink bars; right y axis; unit: $ng\ m^{-3}\ s^{-1}$). (c) Surface concentrations of ISOA derived from IEPOX ($ISOA_{IEPOX}$; dark-blue bars; left y axis; unit: $ng\ m^{-3}$) and ISOA derived from HMML and MAE ($ISOA_{HMML+MAE}$; pink bars; right y axis; unit: $ng\ m^{-3}$). In each panel, the black dashed line denotes the ratio of the HMML+MAE-related quantity to the corresponding IEPOX-related quantity, with monthly ratio values labeled above the markers.

The national mean $\text{ISOA}_{\text{HMML}+\text{MAE}}/\text{ISOA}_{\text{IEPOX}}$ is not constant throughout the year. It becomes relatively higher in late summer and early autumn and reaches its annual maximum in September (Fig. 5(c)). To interpret this feature, we further compare the gas-phase precursors and particle-phase production rates diagnosed from the tagged heterogeneous reactions for the two pathways. The ratio of the gas-phase epoxide precursors HMML + MAE to IEPOX shows only a modest seasonal variation and generally remains within about 0.074 to 0.10 throughout the year (Fig. 5(a)). The ratio of tagged particle-phase ISOA production rates, $\text{P}_{\text{HMML}+\text{MAE}}/\text{P}_{\text{IEPOX}}$, exhibits a similar seasonal pattern, with an annual mean of about 0.10 (Fig. 5(b)). These results indicate that neither the HMML+MAE-to-IEPOX precursor ratio nor the corresponding tagged particle-phase ISOA production-rate ratio shows a comparably strong increase in September. Therefore, the September peak in the $\text{ISOA}_{\text{HMML}+\text{MAE}}/\text{ISOA}_{\text{IEPOX}}$ ratio is not primarily driven by an enhanced relative precursor abundance or production rate of the HMML/MAE pathway, but rather by the stronger decrease in $\text{ISOA}_{\text{IEPOX}}$ relative to $\text{ISOA}_{\text{HMML}+\text{MAE}}$ from August to September, as shown in Fig. 5(c). As a result, the $\text{ISOA}_{\text{HMML}+\text{MAE}}/\text{ISOA}_{\text{IEPOX}}$ ratio increases from 0.089 in August to 0.114 in September, corresponding to an absolute increase of 0.025 in the dimensionless ratio, or a relative increase of 28.1%, and reaches its annual maximum in September. Overall, the HMML/MAE pathway becomes relatively more important during late summer to early autumn, but the IEPOX pathway still dominates ISOA formation on the national average.

3.3 Trend and attribution of ISOA

This section first presents the spatial distribution of long-term trends in surface ISOA over China during 2000–2019, along with the corresponding trends in its key controlling factors, and identifies the regions with the most pronounced ISOA changes. It then uses multiple linear regression to diagnose the dominant drivers of simulated ISOA variability. Sulfate and aerosol liquid water are key components of the multiphase reaction environment: sulfate can provide inorganic nucleophiles for particle-phase reactions, whereas aerosol liquid water provides an aqueous medium for heterogeneous processing (Budisulistiorini et al., 2017; Eddingsaas et al., 2010). Proton availability further catalyzes the ring-opening of epoxide groups, a key step in ISOA formation (Gaston et al., 2014; Pye et al., 2013). Biogenic isoprene emissions provide the direct gas-phase precursor supply for ISOA formation, whereas anthropogenic NO_x emissions influence the NO_x-dependent oxidation chemistry. Based on these considerations, simulated monthly ISOA concentration was used as the dependent variable, and proton concentration (H^+), sulfate (SO_4^{2-} , $\mu\text{g m}^{-3}$), biogenic isoprene emissions (ISOP emis, g m^{-2}), anthropogenic NO_x emissions (Anthro NO_x emis, g m^{-2}), and aerosol liquid water (AeroWater, $\mu\text{g m}^{-3}$) were included as predictors in the regression analysis.

Our nationwide analysis indicates that the long-term evolution of surface ISOA over China is relatively weak in the national mean. Overall, total surface ISOA decreases by 10.23 ng m^{-3} per 20 years (-3.67% per 20 years) from 2000 to 2019. This net change arises from competing trends among the major ISOA subspecies, with the largest decreases occurring in AIETET, followed by AIEOSN, while concurrent increases in other subspecies partly offset these declines. Over the same period, the key controlling factors also show systematic but contrasting national-scale changes, with biogenic isoprene emissions and

anthropogenic NO_x emissions increasing by 5.38% and 48.75% per 20 years, respectively, while anthropogenic SO₂ emissions, sulfate, and aerosol liquid water decrease by 75.12%, 43.15%, and 5.54% per 20 years, respectively (Figs. 6 and 7). These changes imply enhanced precursor supply under rising NO_x but a weakened heterogeneous reaction medium at the national scale. All trends were derived from linear regressions based on annual data from 2000 to 2019. The weak national-mean trend does not imply spatial uniformity and may instead reflect the cancellation of regionally heterogeneous trends with opposite signs. For this reason, the following analysis focuses on the spatial distribution of ISOA trends and their key driving factors at the regional scale.

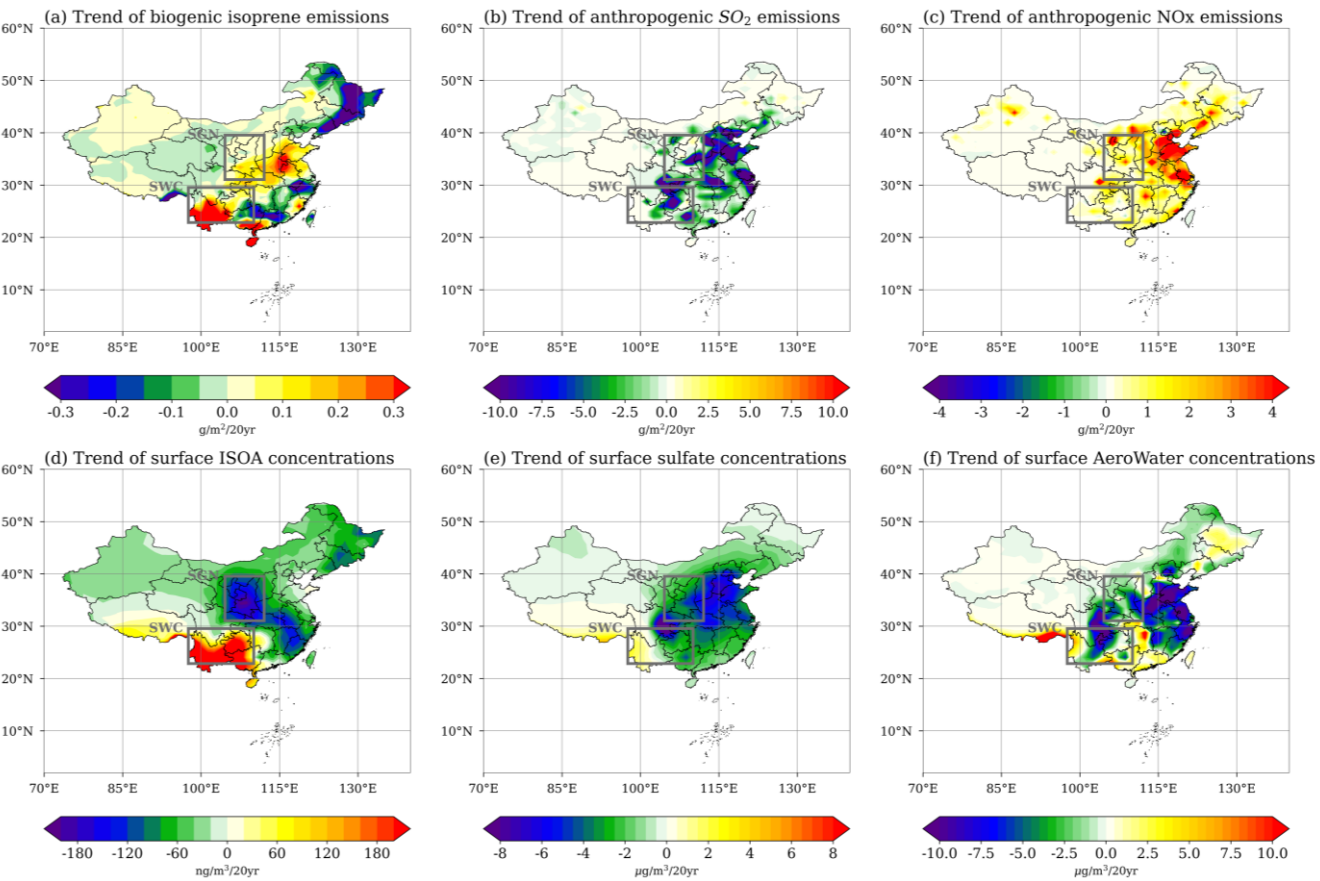


Figure 6: Long-term trends in annual mean surface concentrations of biogenic isoprene emissions (a; unit: g m⁻² per 20 years), anthropogenic sulfur dioxide (SO₂) emissions (b; unit: g m⁻² per 20 years), and anthropogenic nitrogen oxides (NO_x) emissions (c; unit: g m⁻² per 20 years), and of surface isoprene-derived secondary organic aerosol (ISOA) concentrations (d; ng m⁻³ per 20 years), surface sulfate concentrations (e; μg m⁻³ per 20 years), and surface aerosol liquid water (AeroWater) concentrations (f; μg m⁻³ per 20 years), in China from 2000 to 2019. The two boxed regions denote Southwest China (SWC) and the Shaanxi-Gansu-Ningxia region (SGN), respectively.

The spatial distribution of long-term trends confirms that the weak national-mean change in surface ISOA is largely caused by strong regional contrasts with opposite signs (Fig. 6(d)). The simulations show that the strongest increasing trend is found

in Southwest China (SWC) (Fig. 6(d)), where the annual mean surface ISOA concentration increases significantly by 218.01 ng m⁻³ per 20 years, corresponding to an increase of 25.27% per 20 years (Fig. 7(a)). In contrast, the strongest decreasing trend occurs in the Shaanxi–Gansu–Ningxia region (SGN) (Fig. 6(d)), where the annual mean surface ISOA concentration decreases significantly by 142.21 ng m⁻³ per 20 years, corresponding to a decline of 41.87% per 20 years (Fig. 7(a)). As these two regions represent the opposite extremes of the national trend pattern, we focus on them in the following analysis to investigate their ISOA changes and the differences in the dominant controlling factors.

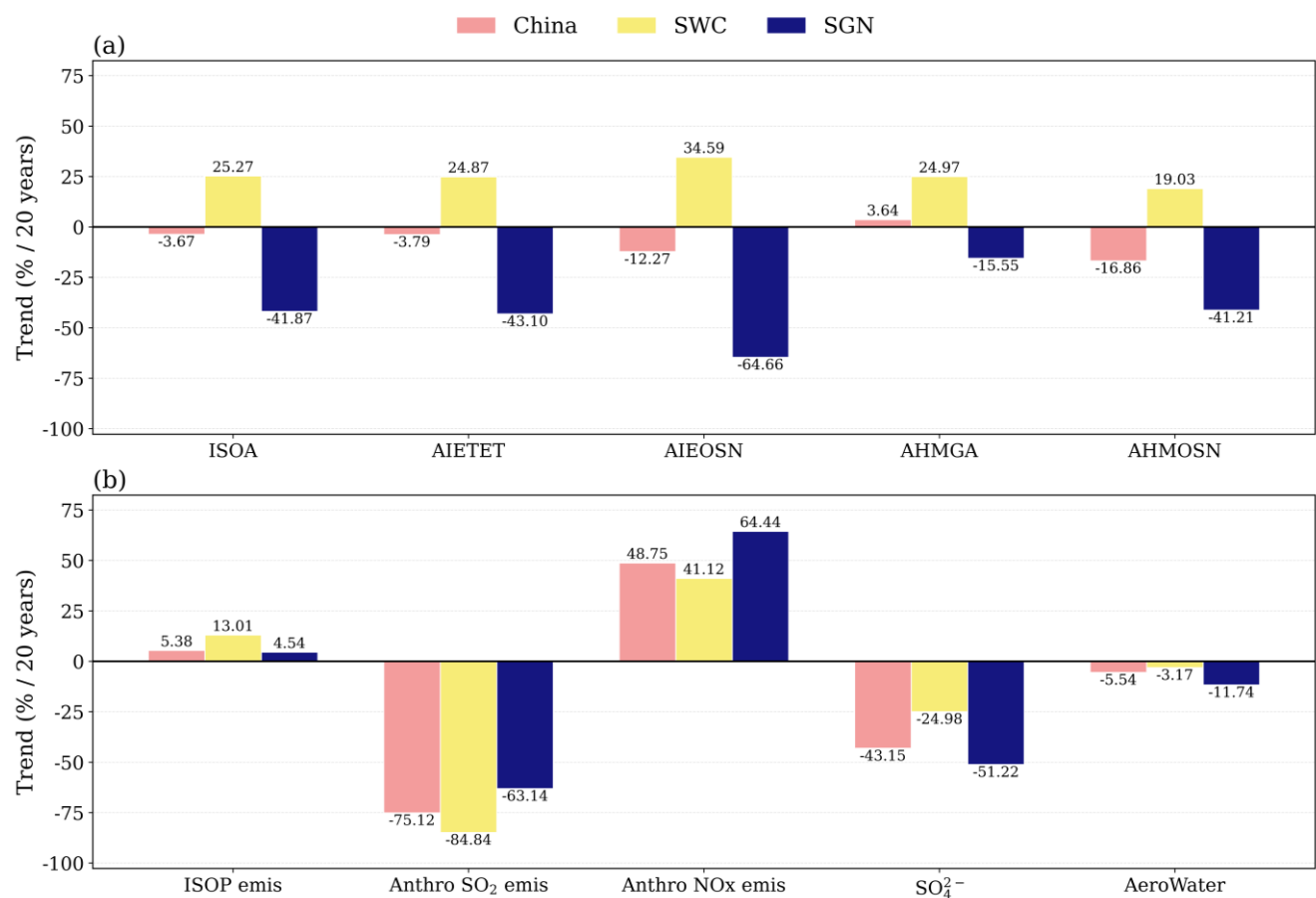


Figure 7: Long-term trends in China (pink bar), Southwest China (SWC; bright yellow bar), and the Shaanxi–Gansu–Ningxia region (SGN; dark blue bar) of (a) surface isoprene-derived secondary organic aerosol (ISOA) and its major subspecies, and (b) biogenic isoprene emissions (ISOP emis), anthropogenic sulfur dioxide (SO₂) emissions (Anthro SO₂ emis), anthropogenic nitrogen oxides (NO_x) emissions (Anthro NO_x emis), surface sulfate (SO₄²⁻) concentrations, and surface aerosol liquid water (AeroWater) concentrations. Bars indicate relative trends over 20 years (% / 20 years).

It is noteworthy that, despite broadly similar temporal trends in the key controlling factors over the two regions, surface ISOA exhibits opposite long-term changes. As shown in Fig. 7(a), ISOA increases in SWC but decreases in SGN. The controlling variables generally evolve in similar directions in both regions, although the magnitudes of their changes differ.

In the SWC region, surface ISOA exhibits a significant increasing trend (Fig. 6(d)), and the increases among individual ISOA subspecies are broadly comparable (Fig. 7(a)). This enhancement is consistent with signals of regional climate change. Regional temperature increases by 0.24% per 20 years and is accompanied by a pronounced increase in biogenic isoprene emissions of 13.01% per 20 years (Fig. 6(a), 7(b)). At the same time, anthropogenic SO₂ emissions decrease by 84.84% per 20 years (Fig. 6(b), 7(b)), accompanied by a corresponding decline in sulfate of 24.98% per 20 years (Fig. 6(e), 7(b)). Aerosol liquid water also shows a slight decrease of 3.17% per 20 years (Fig. 6(f), 7(b)). These results indicate that, in SWC, the increase in isoprene emissions offsets the adverse influence of declining reaction-medium levels. As a result, surface ISOA still shows a marked increase. In contrast, the substantial decrease in ISOA over SGN mainly results from reductions in products formed through the IEPOX pathway (Fig. 7(a), S4). AIETET and AIEOSN decrease by 127.63 ng m⁻³ per 20 years and 9.83 ng m⁻³ per 20 years, respectively. This decreasing trend is consistent with the emission-reduction signal associated with anthropogenic pollution-control measures. Following the implementation of these policies, anthropogenic SO₂ emissions decreased by 63.14% per 20 years (Fig. 6(b), 7(b)), leading to a parallel decrease in sulfate of 51.22% per 20 years (Fig. 6(e), 7(b)). Aerosol liquid water, an important component of the reaction medium, also declines by 11.74% per 20 years (Fig. 6(f), 7(b)) and further reinforces the downward trend in ISOA. Although biogenic isoprene emissions increase slightly by 4.54% per 20 years (Fig. 6(a), 7(b)), this increase is insufficient to offset the overall weakening of the reaction medium and the associated formation environment. Consequently, surface ISOA shows a significant decline in SGN. Overall, the small change in the national-mean ISOA trend mainly results from pronounced regional heterogeneity and offsetting trends, and the dominant factors driving ISOA concentration changes differ substantially across regions.

To quantitatively diagnose the contributions of influential factors and better understand why ISOA trends differ between SWC and SGN, we constructed a multiple linear regression framework linking simulated monthly ISOA concentrations to key controlling factors. This regression framework was used in two complementary ways. Both analyses were based on regional monthly anomalies derived by first spatially averaging the original monthly gridded data over SWC and SGN to obtain regional-mean monthly time series and then removing the mean seasonal cycle. First, these monthly anomalies were standardized to zero mean and unit standard deviation to compare the relative statistical importance of different predictors on a common scale, as summarized in Text S1 and Table S1. This standardized analysis primarily diagnoses the leading predictors of month-to-month ISOA variability. Second, for the trend attribution diagnostic presented in Table 1, we used the same monthly anomalies without standardization, retaining the original physical units of each predictor. Multiple linear regression was subsequently applied to these unstandardized anomalies to obtain sensitivity coefficients in physical units. These sensitivity coefficients (α) describe how regional ISOA responds to changes in each predictor under the regression framework. For each predictor, its contribution to the simulated 20-year ISOA trend (C/20yr) was then estimated by multiplying its regression sensitivity by its own 20-year trend (dX/20yr). The relative contribution of each factor was further expressed as C/dY (%), where dY denotes the simulated 20-year ISOA trend in the corresponding region. The results show that the long-term increase in ISOA in SWC is most strongly associated with increasing biogenic isoprene emissions,

resulting in an ISOA increase of 112.1 ng m^{-3} over 20 years, which corresponds to 58.29% of the total ISOA trend (Table 1). Sulfate has a positive sensitivity coefficient, but its concentration decreases over the study period, leading to a negative regression-based ISOA contribution of -74.18 ng m^{-3} over 20 years, which offsets the positive ISOA trend in SWC by 38.56%. In contrast, increasing anthropogenic NO_x emissions are associated with a positive ISOA contribution of 32.86 ng m^{-3} over 20 years, although the corresponding regression coefficient is not statistically significant. Contributions from aerosol liquid water and $\log_{10}(\text{H}^+)$ are relatively small, with regression-based ISOA changes of -6.39 and 4.34 ng m^{-3} over 20 years, respectively (Table 1). In SGN, the long-term decrease in ISOA is mainly associated with declining sulfate, which contributes -84.38 ng m^{-3} over 20 years and accounts for 63.37% of the total ISOA decrease (Table 1). Despite increasing over the study period, anthropogenic NO_x emissions are associated with an additional negative regression-based ISOA contribution of -45.25 ng m^{-3} over 20 years, corresponding to 33.98% of the total ISOA decrease, although this coefficient is not statistically significant and should therefore be interpreted cautiously. By comparison, increasing biogenic isoprene emissions partly offset the regional ISOA decline, resulting in a regression-based ISOA increase of 8.62 ng m^{-3} over 20 years, which offsets the total ISOA decrease by 6.48%. Aerosol liquid water also slightly offsets the ISOA decline, with a regression-based contribution of 5.66 ng m^{-3} over 20 years, corresponding to an offset of 4.25%, whereas $\log_{10}(\text{H}^+)$ contributes only weakly to the decrease, with a contribution of -0.72 ng m^{-3} over 20 years (Table 1). We further evaluated potential predictor correlations and collinearity by calculating pairwise Pearson correlation coefficients among the predictors used in the regression analysis (Fig. S9). The results indicate that some predictors are not fully independent. In both SWC and SGN, sulfate, aerosol liquid water, and $\log_{10}(\text{H}^+)$ show moderate to strong positive correlations, with correlation coefficients of 0.70–0.81 in SWC and 0.75–0.88 in SGN. These correlations are physically expected because aerosol liquid water is thermodynamically calculated in MOSAIC based on inorganic aerosol composition, including sulfate, as well as meteorological conditions. Sulfate also contributes to aerosol acidity and aerosol liquid water formation, while aerosol liquid water can further influence aqueous-phase uptake and processing. Therefore, sulfate, aerosol liquid water, and aerosol acidity are treated here as physically connected components of the aerosol chemical environment. Accordingly, the regression-based relative contributions associated with these covarying predictors should be interpreted as diagnostic indicators of statistical associations with ISOA variability, rather than as a strict decomposition of independent causal effects, because they may partly reflect coupled variations among sulfate, aerosol liquid water, and aerosol acidity. This caveat is less important for biogenic isoprene emissions and anthropogenic NO_x emissions, as both predictors show weak correlations with the other predictors in both regions. Overall, these results suggest that the leading statistical predictors of long-term ISOA change differ between the two regions. The ISOA increase in SWC is primarily associated with enhanced biogenic isoprene emissions, while the ISOA decrease in SGN is most strongly associated with declining sulfate.

Driver	Southwest China (SWC)					Shaanxi–Gansu–Ningxia region (SGN)				
	dX/20yr	α	p_ α	C/20yr	C/dY (%)	dX/20yr	α	p_ α	C/20yr	C/dY (%)
ISOP emis (g/m ² /mon)	0.028	4074	0.001	112.1	58.29	0.005	1618	0.043	8.622	-6.480
SO ₄ ²⁻ (μg/m ³)	-1.529	48.53	0.126	-74.18	-38.56	-3.525	23.94	0.049	-84.38	63.37
Anthro NOx emis (g/m ² /mon)	0.033	1006	0.632	32.86	17.08	0.081	-553.0	0.271	-45.25	33.98
AeroWater (μg/m ³)	-2.004	3.188	0.622	-6.387	-3.320	-2.834	-1.997	0.554	5.658	-4.250
log ₁₀ (H ⁺)	-0.109	-39.70	0.901	4.336	2.250	-0.015	46.48	0.618	-0.715	0.540

Table 1: Regression-based attribution of long-term ISOA trends in Southwest China (SWC) and the Shaanxi–Gansu–Ningxia region (SGN). The regression is based on deseasonalized monthly anomalies from the simulation years, and the resulting sensitivity coefficients are combined with the changes in individual drivers over 2000–2019 to estimate their contributions to the long-term ISOA trend. For each driver, the table lists its 20-year change (dX/20 yr), regression sensitivity coefficient (α), significance level (p_ α), attributed contribution to the ISOA trend (C/20 yr), and relative contribution (C/dY; %). Bold values indicate the largest relative contributions in each region.

We also examined the interannual evolution of the four major ISOA subspecies in SWC and SGN over 2000–2019, and the detailed results are presented in Fig. 8. The two regions show clearly different interannual changes in ISOA subspecies, consistent with the contrasting regional ISOA trends discussed above. In SWC, all four ISOA subspecies increased from 2000 to 2006 (Fig. 8(a)). This increase is consistent with the simultaneous increases in biogenic isoprene emissions (Fig. 8(b)), epoxide precursor concentrations (Fig. 8(c)), and particle-phase reaction conditions, including sulfate and aerosol liquid water (Fig. 8(d)), during the same period. After 2006, most ISOA subspecies generally decrease toward 2016, whereas AHMGA shows a delayed maximum in 2012, mainly attributable to enhanced anthropogenic NOx emissions and HMML + MAE precursor concentrations, together with the peak in aerosol liquid water that favors H₂O-mediated heterogeneous uptake. This is followed by a rebound in the normalized values of all four ISOA subspecies from 2016 to 2019 (Fig. 8(a)). Notably, this rebound occurs despite the continued decreases in anthropogenic SO₂ emissions, sulfate, and aerosol liquid water after 2012, suggesting that precursor supply and epoxide precursor concentrations play a stronger role in driving the later-stage ISOA recovery in SWC. This feature is also consistent with the multiple linear regression results discussed above, which identify biogenic isoprene emissions as the leading statistical indicator of ISOA variability and long-term changes in SWC. In contrast, ISOA subspecies in SGN generally peak in 2006 and decline thereafter, showing a more synchronous temporal evolution across different product classes. This decline in ISOA subspecies after 2006 occurs even though biogenic isoprene emissions and gas-phase epoxide precursor concentrations increase over the same period, suggesting that declining sulfate and aerosol liquid water can constrain heterogeneous uptake and particle-phase product formation in SGN. These patterns suggest that the temporal evolution in SGN is more strongly linked to changes in the aerosol reaction medium. This

interpretation is also consistent with the multiple linear regression results discussed above, which identify sulfate as the leading statistical indicator and the dominant regression-based contributor to the long-term ISOA decrease in SGN. These regional differences indicate that the mechanisms governing ISOA changes are not spatially uniform and provide additional process-based support for the contrasting ISOA trends identified in SWC and SGN.

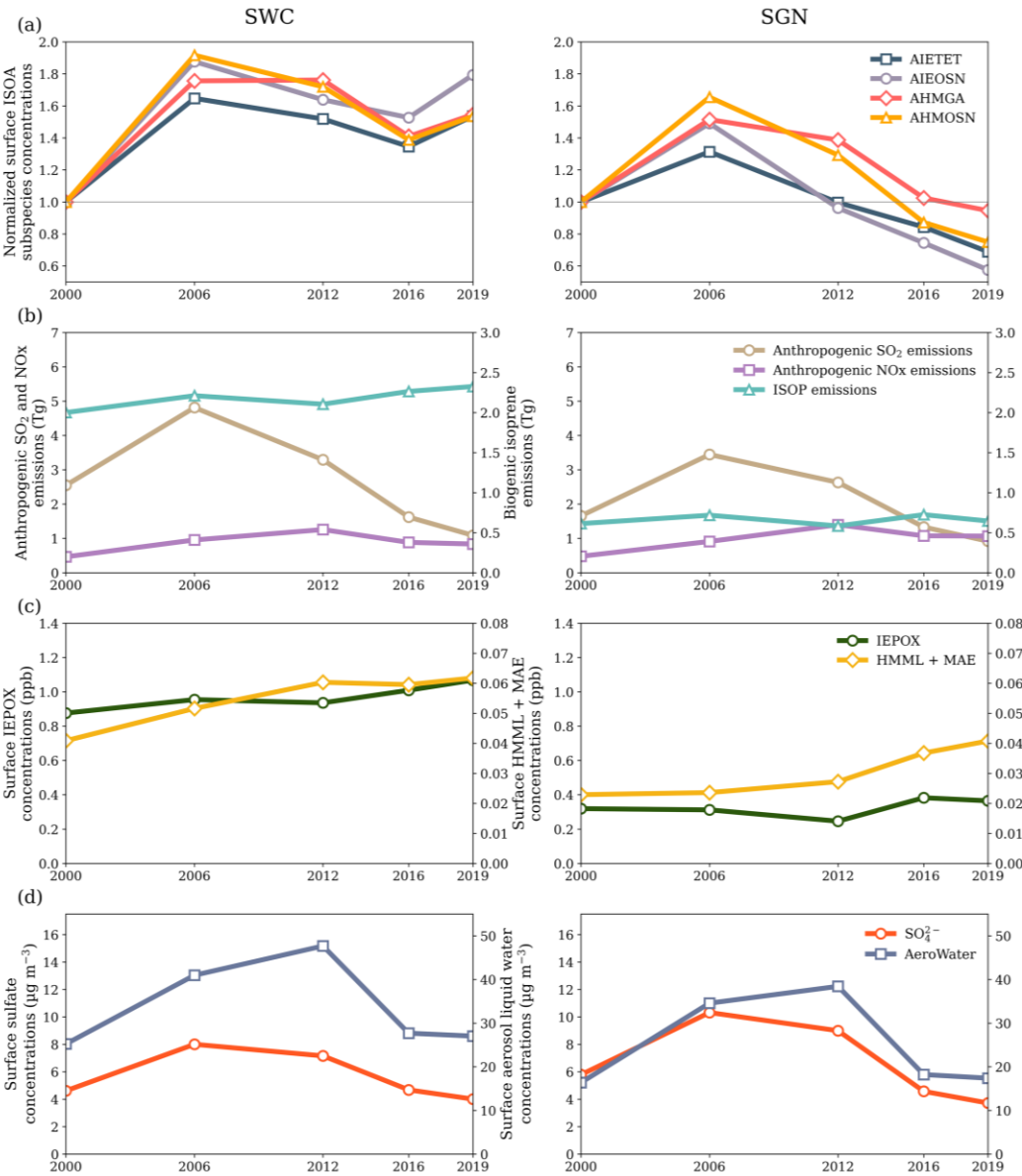


Figure 8: Time series for Southwest China (SWC; left column) and the Shaanxi–Gansu–Ningxia region (SGN; right column), showing (a) normalized surface isoprene-derived secondary organic aerosol (ISOA) subspecies concentrations, including AIETET (dark blue), AIEOSN (gray-purple), AHMGA (rose red), and AHMOSN (earthy yellow). The normalized values are calculated as the ratio of the value in each year to that in 2000; (b) anthropogenic sulfur dioxide (SO₂) emissions (tan; left y axis; unit: Tg) and

anthropogenic nitrogen oxides (NO_x) emissions (purple; left y axis; unit: Tg), and biogenic isoprene emissions (blue; right y axis; unit: Tg); (c) surface isoprene epoxydiols (IEPOX) concentrations (dark green; left y axis; unit: ppb) and surface hydroxymethyl-methyl- α -lactone plus methacrylic acid epoxide (HMML + MAE) concentrations (yellow; right y axis; unit: ppb); and (d) surface sulfate (SO₄²⁻) concentrations (red-orange; left y axis; unit: $\mu\text{g m}^{-3}$) and surface aerosol liquid water (AeroWater) concentrations (blue-gray; right y axis; unit: $\mu\text{g m}^{-3}$).

Taken together, the contrasting ISOA trends in SWC and SGN highlight the need for a coupled precursor-multiphase reaction framework. The SWC case represents a more precursor-driven regime, in which enhanced biogenic isoprene emissions and NO_x-dependent oxidation promote epoxide precursor formation and enhance ISOA production despite less favorable aerosol multiphase reaction conditions. In contrast, the SGN case represents a more reaction-medium-limited regime, in which declining sulfate and aerosol liquid water suppress heterogeneous uptake and particle-phase product formation, even when precursor supply remains relatively high. This regime-based interpretation emphasizes the balance between precursor supply, NO_x-dependent oxidation chemistry, and aerosol multiphase reaction conditions, as summarized schematically in Fig. 9.

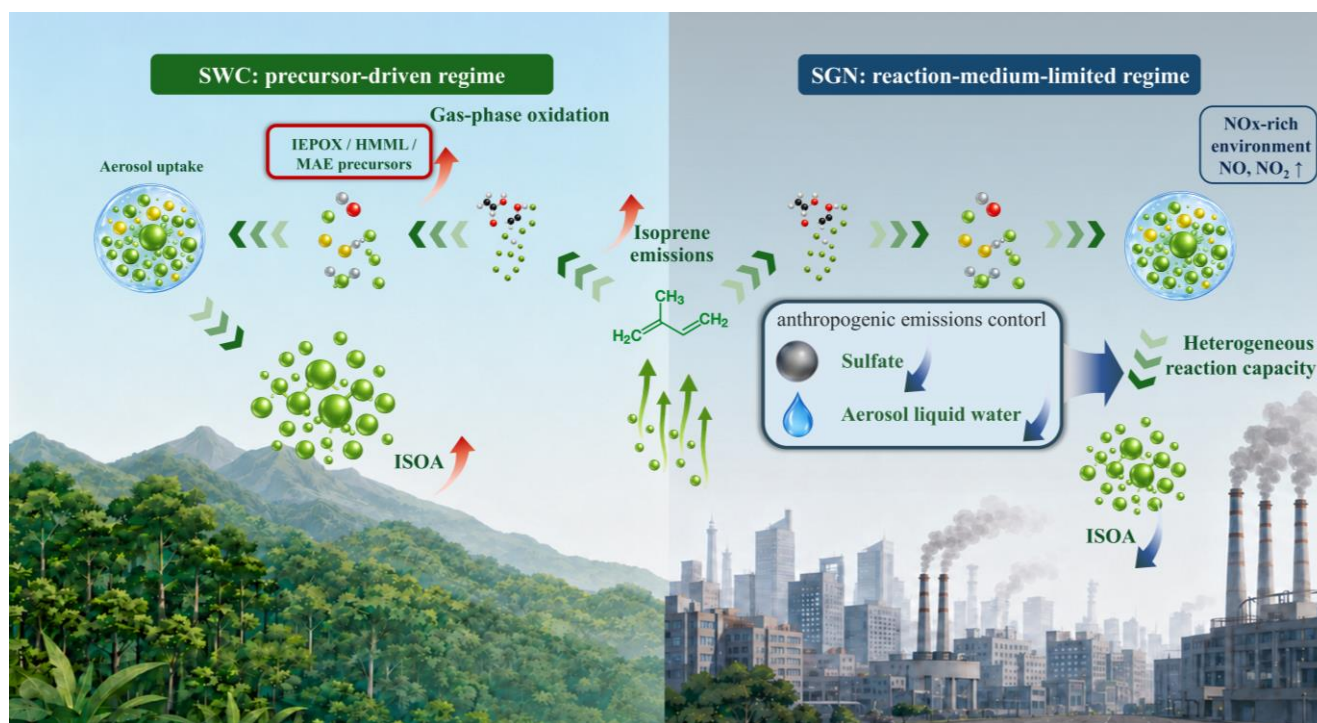


Figure 9. Schematic illustration of contrasting regional controls on surface isoprene-derived secondary organic aerosol (ISOA) formation in Southwest China (SWC; left) and the Shaanxi-Gansu-Ningxia region (SGN; right). The background scenes and selected graphical elements were created with the assistance of OpenAI's image-generation tools in ChatGPT (OpenAI, 2026).

650 4. Conclusions and Discussions

This study implements an explicit ISOA formation mechanism in CAM6-Chem that introduces representations of high-NO_x epoxide precursors and their heterogeneous uptake, while resolving reactive-uptake-derived ISOA into H₂O-mediated nucleophilic addition and inorganic-nucleophile-addition product classes. Evaluation against ground-based observations shows that the explicit ISOA mechanism better reproduces the concentrations and compositional structure of major ISOA subspecies, although positive biases remain for AIETET and AHMGA. At the bulk aerosol level, this mechanism substantially alleviates the underestimation of SOA and OA over China, with the NMB improving from -76.7% to -51.6% for SOA and from -47.0% to -25.7% for OA. In addition, the explicit simulation better captures the observed multi-component nature of ambient ISOA and better reproduces the relative contributions of subspecies associated with the newly introduced pathways. These results demonstrate that the explicit ISOA mechanism is chemically meaningful and improves the model's ability to represent both the magnitude and composition of ambient ISOA.

Using this explicit ISOA framework, we find that ISOA formation in China reflects competition between the IEPOX pathway and the HMML/MAE pathway under different NO_x backgrounds. On the national average, the IEPOX pathway remains the dominant source of ISOA, with an annual mean $\text{ISOA}_{\text{HMML+MAE}}/\text{ISOA}_{\text{IEPOX}}$ ratio of only about 0.07. The relative importance of the HMML/MAE pathway increases in NO_x-rich regions and during late summer to early autumn, but it remains secondary to IEPOX even under these conditions. Overall, these results indicate that NO_x enhances the contribution of the HMML/MAE pathway regionally and seasonally, while the IEPOX pathway continues to dominate ISOA formation at the national scale.

Against this national-scale background, spatial analysis showed that the long-term national-mean change in surface ISOA over China during 2000–2019 was weak, largely because strong regional trends with opposite signs offset each other. The most pronounced increase occurred in SWC, whereas the strongest decrease occurred in SGN. Regression-based attribution further showed that the ISOA increase in SWC was most strongly associated with enhanced biogenic isoprene emissions, while the ISOA decrease in SGN was mainly associated with declining sulfate. These contrasts were also reflected in the normalized interannual evolution of major ISOA subspecies. In SWC, the normalized values of all four ISOA subspecies rebounded from 2016 to 2019 despite the post-2012 decreases in sulfate and aerosol liquid water, suggesting an important role of precursor supply. In SGN, by contrast, ISOA subspecies generally peaked in 2006 and declined thereafter, even though biogenic isoprene emissions and gas-phase epoxide precursor concentrations increased after 2006, highlighting the stronger constraint from weakened aerosol multiphase reaction conditions. Together, these results demonstrate that long-term ISOA changes in China are governed by different regional balances between precursor availability and aerosol multiphase reaction conditions.

Future model development should further address several simplifying assumptions in the current epoxide reactive uptake parameterization. As described in Sect. 2.1.2, the model follows a core-shell treatment in which the organic shell and aqueous core are assumed to have the same reaction rate coefficient. Because phase separation and increased organic-shell viscosity can limit epoxide diffusion and reactive uptake, this assumption may overestimate epoxide-derived ISOA under strongly phase-separated or highly viscous conditions (Zhang et al., 2019; Schmedding et al., 2020; Chen et al., 2024b; Farrell et al., 2025). HMML and MAE are also represented using the same solubility, diffusivity, and condensed-phase reaction parameters as IEPOX because compound-specific laboratory constraints remain limited. This IEPOX-like treatment introduces uncertainty in the absolute uptake rates of high-NO_x epoxides and in the quantitative $\text{ISOA}_{\text{HMML+MAE}}/\text{ISOA}_{\text{IEPOX}}$ ratio. The assumptions of 100% ISOA yield from epoxide reactive uptake and non-volatile epoxide-derived ISOA introduce additional uncertainty, although they are consistent with previous global ISOA modeling studies (Marais et al., 2016; Jo et al., 2019, 2021). Lower effective yields or partial volatility, especially for products such as 2-MT and 2-MGA, could reduce simulated ISOA concentrations and their sensitivity to sulfate, aerosol liquid water, and acidity. Consistent with this need, gas-particle partitioning treatments for 2-MT in regional models suggest a useful direction for future improvement in global models (Pye et al., 2018; Shrivastava et al., 2022). Finally, the removal of ISOA formation from the VBS scheme may also remove important non-epoxide contributions to ISOA that were previously represented implicitly, such as multifunctional organic peroxides under low-NO_x conditions and organic nitrates under high-NO_x conditions (Paulot et al., 2009a; Schwantes et al., 2019). Taken together, these assumptions mean that the long-term trends and regression-based attributions diagnosed here should be interpreted within the adopted epoxide-focused chemical framework, rather than as exact causal contributions independent of model structure and parameterization.

Beyond the uncertainties associated with parameterizations already included in the current framework, some potentially important ISOA processes remain absent from the model. These include isoprene cloud-water chemistry, which may provide an additional ISOA source and influence simulated aerosol composition, and photolytic loss, since ISOA may photolyze faster than monoterpene-derived SOA (Zawadowicz et al., 2020), thereby affecting its simulated lifetime and burden. Incorporating these missing formation and loss processes, together with more composition-resolved observations to constrain pathway-specific chemistry, will be important for future model refinement. Future global-scale simulations and attribution analyses would also be useful for evaluating whether the updated ISOA mechanism has broader implications for other high-isoprene regions, where climate change and human activities may shift the relative roles of biogenic precursor supply, anthropogenic emissions, and the heterogeneous aerosol reaction environment in shaping future ISOA responses.

Code and data availability

The Community Earth System Model (CESM) is an open-source framework available at <https://www.cesm.ucar.edu/models/cesm2/download>. The modified code for the explicit isoprene chemistry scheme

715 implemented in this study is available upon reasonable request. Ground-based measurements for OA and SOA were obtained from the supplementary materials of published articles by Miao et al. (2021) and Chen et al. (2024a). Ground-based measurements of AIETET and AHMGA were obtained from Ding et al. (2016). Ground-based measurements of ISOA products used to evaluate ISOA composition and pathway partitioning were obtained from Zhang et al. (2022).

Author contributions

720 XD, MW, and WZ designed the study. WZ developed the model code, performed the simulations, produced the figures, and wrote the manuscript draft. MY and XS contributed to the model simulations. All the authors contributed to the discussion and editing of the manuscript.

Competing interests

At least one of the (co-)authors is a member of the editorial board of Atmospheric Chemistry and Physics.

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