



1 **Development and evaluation of the ECHAM6-iMAPLE**
2 **v1.0 coupled atmosphere-ecosystem model**

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20 **Abstract.** Land-atmosphere interactions play a fundamental role in regulating climate
21 variability, ecosystem productivity, and air quality through coupled exchanges of energy,
22 water, carbon, and reactive trace gases. However, many Earth system models adopt
23 simplified representations of vegetation physiological processes, leading to biases in
24 terrestrial carbon and water fluxes and increased uncertainties in climate simulations.
25 Here, we present ECHAM6-iMAPLE v1.0, a newly coupled modeling framework that
26 integrates the interactive Model for Air Pollution and Land Ecosystems (iMAPLE v1.0)
27 into the ECHAM6 atmospheric general circulation model. The coupled model is
28 evaluated against reanalysis, benchmark, and satellite datasets. Compared with the
29 original configuration, ECHAM6-iMAPLE substantially improves simulations of gross
30 primary productivity, evapotranspiration, and leaf area index, capturing their spatial
31 distributions and seasonal cycles more reasonably. These improvements arise from
32 well-constrained physiological parameters calibrated using extensive site-level
33 observations and a more realistic representation of key biophysical processes in
34 iMAPLE. With improved carbon and water fluxes, simulations of soil temperature, soil
35 moisture, and surface air temperature show reduced root mean square errors. Overall,
36 evaluations demonstrate that ECHAM6-iMAPLE provides a useful tool for
37 investigating atmosphere-ecosystem interactions and their implications for future
38 climate change projections.

39

40 **Keywords:** ECHAM6, iMAPLE, ecosystem, atmosphere, interaction, coupling

41



42 **1. Introduction**

43 Terrestrial ecosystems and climate systems are tightly linked through exchanges of
44 energy fluxes, water vapor, carbon, and trace gases, thereby regulating regional and
45 global climate patterns (Pielke et al., 1998; Green et al., 2017). Climate conditions
46 directly influence vegetation growth (Parmesan and Yohe, 2003; Fastovich et al., 2025),
47 distribution (Seidl et al., 2017; Nolan et al., 2018), and phenology (Higgins et al., 2023),
48 and atmospheric pollutants such as ozone can damage plant physiological processes by
49 reducing photosynthetic rates and stomatal conductance (Sitch et al., 2007; Cao et al.,
50 2024). Conversely, terrestrial ecosystems exert substantial feedbacks on the atmosphere
51 through physical and biogeochemical processes (Bonan, 2008; Heimann and Reichstein,
52 2008). For instance, terrestrial vegetation absorbs approximately 120 Pg C/year through
53 photosynthesis (Beer et al., 2010; Friedlingstein et al., 2025), while vegetation and soil
54 respiration return a comparable amount of CO₂ into the atmosphere (Ruehr et al., 2023).
55 Ecosystem also influences the partitioning of sensible and latent heat fluxes by
56 modulating transpiration and canopy conductance (Williams and Torn, 2015; Zeng et
57 al., 2017; Forzieri et al., 2020), thus affecting precipitation patterns and the
58 intensity/frequency of hydrological extremes (Zhou et al., 2021; Schumacher et al.,
59 2022; Zhang et al., 2025).

60 Models that couple atmospheric general circulation with vegetation dynamics
61 provide essential tools for investigating interactions between the atmosphere and
62 terrestrial ecosystems (Sellers et al., 1997; Berg et al., 2016). However, vegetation
63 models differ substantially in their parameterization schemes for representing complex
64 hydrological, biogeophysical, and biogeochemical processes. Uncertainties arising
65 from these parameterization differences can propagate through land-atmosphere
66 interactions, thereby influencing weather forecasts and climate projections (Wei et al.,
67 2010; Wei et al., 2018). Despite continuous development, many widely used vegetation
68 models still lack a comprehensive representation of key ecosystem processes that
69 regulate land-atmosphere exchanges of energy, water, and biogeochemical constituents
70 (Pitman, 2003; Blyth et al., 2021). For instance, vegetation dynamics such as seasonal
71 variations of leaf area are usually prescribed with empirical parameters rather than
72 being simulated through interactive responses to environmental drivers (Gulden et al.,
73 2007; Niu et al., 2011; Fisher and Koven, 2020). This limitation constrains the ability
74 of models to realistically capture ecosystem dynamics and their feedbacks to
75 atmosphere.



76 In this study, we present the coupled atmosphere-ecosystem model ECHAM6-
77 iMAPLE, developed by integrating the interactive Model for Air Pollution and Land
78 Ecosystems (iMAPLE) v1.0 (Yue et al., 2024) into the atmospheric general circulation
79 model ECHAM version 6.3 (<https://redmine.hammoz.ethz.ch/projects/hammoz>, last
80 access: 30 March 2026) (Tegen et al., 2019). ECHAM6 has been extensively developed
81 with multiple Earth system components and has participated in the Sixth Coupled
82 Model Intercomparison Project (CMIP6) (Stevens et al., 2013; Sidorenko et al., 2015;
83 Cao et al., 2018). The model has been widely used for global climate simulations and
84 projections owing to the robust dynamical core and comprehensive physical
85 parameterizations (Stevens et al., 2013; Siongco et al., 2017; Kretzschmar et al., 2019).
86 iMAPLE is a global dynamic vegetation model designed to simulate plant growth and
87 terrestrial carbon cycle interactively (Yue et al., 2024). It has demonstrated robust
88 capability in simulating major carbon and water fluxes through its participation in the
89 multi-model intercomparison project for Global Carbon Budget (Friedlingstein et al.,
90 2025). Here, we evaluate the dynamically simulated carbon and water fluxes from
91 ECHAM6-iMAPLE, and compare its performance with the original ECHAM6 model
92 in representing critical surface variables, including soil temperature, soil moisture, near-
93 surface air temperature, and precipitation. A detailed description of the model coupling
94 framework is provided in Section 2. Section 3 presents the simulation results and model
95 evaluations. Concluding remarks and future research perspectives are discussed in
96 Section 4.

97

98 **2. Models, methods, and data**

99 **2.1 Descriptions of iMAPLE**

100 iMAPLE is an interactive model developed to investigate the interplay between air
101 pollution and terrestrial ecosystems (Yue et al., 2024). The model evolved from version
102 1.0 of the Yale Interactive terrestrial Biosphere (YIBs) model, with an expanded focus
103 on explicitly representing the interaction between atmospheric chemistry and land
104 ecosystems (Yue and Unger, 2015). Compared to YIBs, iMAPLE incorporates several
105 new processes, including dynamic fire emission (Pechony and Shindell, 2009; Li et al.,
106 2012), wetland methane emissions (Walter et al., 2001; Zhu et al., 2014), and
107 photosynthetic limitation under environmental stress (Arora et al., 2009). Furthermore,
108 iMAPLE integrates the process-based hydrological scheme Noah-MP, which enables
109 dynamical simulation of soil temperature and moisture and facilitates tight coupling



110 between the carbon and water cycles (Niu et al., 2011). These improvements
111 substantially enhance the model’s capacity to represent ecosystem-atmosphere
112 interactions. Below, we briefly describe the parameterizations of vegetation biophysics
113 and hydrological processes. A comprehensive model description is provided in Yue et
114 al. (2024) and Niu et al. (2011).

115

116 2.1.1 Vegetation biophysics of iMAPLE

117 The vegetation biophysics in iMAPLE utilizes the canonical Michaelis–Menten
118 enzyme-kinetics scheme to simulate photosynthetic processes for both C₃ and C₄ plant
119 functional types (Farquhar et al., 1980; Von Caemmerer and Farquhar, 1981). The total
120 leaf photosynthesis (A_{tot} , $\mu\text{mol m}^{-2} [\text{leaf}] \text{s}^{-1}$) is limited by one of three biochemical
121 processes:

$$122 \quad A_{tot} = \min(J_c, J_e, J_s) \quad (1)$$

123 Here, J_c represents the Rubisco-limited rate of carbon fixation catalyzed by the enzyme
124 ribulose 1,5-bisphosphate (RuBP) carboxylase/oxygenase (Rubisco). J_e denotes the
125 RuBP regeneration-limited rate driven by electron transport through the Calvin cycle
126 and associated thylakoid reactions. J_s is the limitation imposed by end-product
127 synthesis, reflecting the capacity for starch and sucrose formation to regenerate
128 inorganic phosphate required for photophosphorylation in C₃ plants, or
129 phosphoenolpyruvate (PEP) regeneration in C₄ plants. The rates J_c , J_e , and J_s are
130 parameterized as functions of environmental meteorological conditions and
131 photosynthetic capacity, including the maximum carboxylation rate (V_{cmax} , $\mu\text{mol m}^{-2}$
132 s^{-2}) (Collatz et al., 1991; Collatz et al., 1992a):

$$133 \quad J_e = \begin{cases} a_{leaf} \times PAR \times \alpha \left(\frac{C_i - \Gamma_*}{C_i + 2\Gamma_*} \right) & \text{for C3 plant,} \\ a_{leaf} \times PAR \times \alpha & \text{for C4 plant} \end{cases} \quad (2)$$

$$134 \quad J_c = \begin{cases} V_{cmax} \left(\frac{C_i - \Gamma_*}{C_i + K_c(1 + O_i/K_o)} \right) & \text{for C3 plant,} \\ V_{cmax} & \text{for C4 plant} \end{cases} \quad (3)$$

$$135 \quad J_s = \begin{cases} 0.5V_{cmax} & \text{for C3 plant,} \\ K_s \times V_{cmax} \times \frac{C_i}{P_s} & \text{for C4 plant} \end{cases} \quad (4)$$

136 Here, C_i and O_i denote the partial pressures (Pa) of CO₂ and oxygen inside the leaf,
137 while Γ_* (Pa) refers to the CO₂ compensation point. The Michaelis–Menten constants



138 for Rubisco carboxylation and oxygenation are represented by K_c (Pa) and K_o (Pa),
139 respectively. PAR ($\mu\text{mol m}^{-2} \text{s}^{-1}$) denotes the incident photosynthetically active
140 radiation, a_{leaf} is the light absorption coefficient of the leaf, and α represents the
141 intrinsic quantum efficiency. P_s is the ambient pressure, and K_s is a constant with a
142 value of 4000 (Lawrence et al., 2011). V_{cmax} is derived from its optimum value at 25 °C
143 (V_{cmax25}) through a Q10 function.

144 The canopy radiative transfer scheme employs an adaptive approach, typically
145 dividing vegetation into 2-16 vertical layers to resolve light stratification within the
146 canopy. At each layer, a two-leaf approach further partitions leaf area into sunlit and
147 shaded fractions, thereby capturing canopy light heterogeneity and distinguishing the
148 different light-use efficiency under diffuse and direct radiation (Spitters, 1986; Spitters
149 et al., 1986). Gross primary productivity (GPP) is then calculated by integrating
150 photosynthesis across all canopy layers:

$$151 \quad GPP = \int_0^{LAI} A_{tot} dL \quad (5)$$

152 After accounting for autotrophic respiration, net carbon assimilation is allocated to
153 leaves, stems, and roots to support plant growth and leaf development. Leaf area index
154 (LAI) is updated at 10-day intervals based on total leaf carbon and vegetation phenology.
155

156 2.1.2 Water cycle of iMAPLE

157 The hydrological cycle in iMAPLE adopts the Noah-MP module, which resolves
158 the grid-scale water balance among precipitation (P , $\text{Kg m}^{-2} \text{s}^{-1}$), evapotranspiration
159 (ET , $\text{Kg m}^{-2} \text{s}^{-1}$), runoff, and changes in terrestrial water storage (ΔTWS) at each grid:

$$160 \quad P = ET + \text{runoff} + \Delta TWS \quad (6)$$

161 ET is further divided into plant transpiration (E_{tr}), canopy evaporation (E_{can}), and
162 ground evaporation (E_{gro}):

$$163 \quad ET = E_{tr} + E_{can} + E_{gro} \quad (7)$$

164 The runoff comprises surface (R_{srf}) and subsurface (R_{sub}) parts:

$$165 \quad \text{runoff} = R_{srf} + R_{sub} \quad (8)$$

166 Terrestrial water storage (TWS) encompasses three major components: groundwater
167 storage (W_{gw}), which refers to water held in aquifers; soil water content (W_{soil}),



168 representing moisture within four soil layers ($N_{soil}=4$); and snow water equivalent
169 (W_{snow}), indicating the liquid water equivalent of accumulated snowpack:

$$170 \quad TWS = W_{gw} + W_{snow} + \sum_{i=1}^{N_{soil}=4} W_{soil} \quad (9)$$

171

172 2.2 Descriptions of ECHAM v6.3

173 ECHAM6 is the sixth generation of the ECHAM atmospheric general circulation
174 model, developed by the Max Planck Institute for Meteorology (MPI-M) (Giorgetta et
175 al., 2013; Stevens et al., 2013). Its dynamical core employs a hybrid numerical
176 framework that combines spectral and finite-difference methods to solve the primitive
177 equations (Arakawa, 1966; Bourke et al., 1977). Horizontally, ECHAM6 utilizes a
178 truncated spherical harmonic expansion to calculate the nonlinear and parameterized
179 terms of dynamical fields on a Gaussian grid. Vertically, the model employs a hybrid
180 sigma-pressure coordinate system implemented on a Lorenz grid (Phillips, 1957;
181 Simmons and Burridge, 1981).

182 Compared to previous ECHAM versions, ECHAM6 incorporates substantial
183 improvement in key physical processes, including a more sophisticated representation
184 of land surface processes, updated radiation parameterizations, refined surface albedo
185 calculations, and revised convection triggering criteria (Stier et al., 2005; Crueger et al.,
186 2018; Tegen et al., 2019). The model supports multiple configurations, with horizontal
187 resolutions spanning from T31 (3.75°) to T255 (0.7°) and 47 or 95 vertical layers
188 (Stevens et al., 2013). In this study, we adopt the T63L47 configuration to enable
189 coupling with the iMAPLE model. The T63 spectral grid is characterized by an
190 approximate grid spacing of 1.875° in longitude and latitude. The L47 vertical
191 configuration comprises 47 layers extending from the surface to 0.01 hPa.

192 ECHAM6 employs JSBACH version 3.10 to simulate land vegetation processes
193 (Goll et al., 2017). Compared to the earlier versions, JSBACH v3.10 incorporates a soil
194 water diffusion scheme that provides a more physically consistent representation of
195 terrestrial water budgets. In addition, modules for the dynamical simulation of LAI,
196 canopy radiative transfer, and leaf-level photosynthesis have been integrated into the
197 current versions. Here, we provide a brief description of JSBACH v3.10 to facilitate
198 the comparisons of its physical processes with those of iMAPLE v1.0.

199



200 2.2.1 Vegetation biophysics of JSBACH

201 JSBACH utilizes a tiling scheme to represent land surface heterogeneity and
202 incorporates dynamic vegetation components consisting of 12 plant functional types
203 (PFTs) and two bare-ground classes (Reick et al., 2013). For vegetation productivity,
204 JSBACH adopts parameterizations from the Biosphere Energy Transfer Hydrology
205 (BETHY) module, including canopy radiative transfer calculations and two distinct
206 photosynthesis simulations (Sellers et al., 1992; Knorr, 2000).

207 Leaf photosynthesis in BETHY is implemented in two sequential steps. First,
208 potential productivity is calculated without considering soil water deficit, which
209 generates stomatal conductance under unstressed conditions. The stomatal conductance
210 derived in this step is then passed to the soil hydrological scheme to estimate potential
211 transpiration and associated water loss. Second, stomatal conductance is adjusted
212 according to soil water availability, and photosynthesis is recalculated to obtain actual
213 productivity. The A_{tot} in JSBACH v3.10 can be expressed as:

$$214 \quad A_{tot} = \min(J_c, J_e) \quad (10)$$

215 In JSBACH, photosynthesis of C_3 plants is represented using the Farquhar model
216 (Farquhar et al., 1980), consistent with iMAPLE. For C_4 plants, JSBACH utilizes the
217 Collatz model (Collatz et al., 1992b). The primary difference between the Collatz and
218 Farquhar formulations lies in their parameterization of the carboxylation rate (J_c) and
219 the electron transport rate (J_e):

$$220 \quad J_c = k * C_i \quad (11)$$

$$221 \quad J_e = \frac{1}{2\theta_s} \left(V_{p,max} + J_i - \sqrt{(V_{p,max} + J_i)^2 - 4\theta_s * V_{p,max} * J_i} \right) \quad (12)$$

$$222 \quad J_i = \alpha_i * I \quad (13)$$

$$223 \quad I = (R(0) + R_1(0)) * fAPAR \quad (14)$$

224 Here k is specific parameters for carboxylase, C_i the partial pressures of CO_2 and
225 oxygen inside the leaf (in Pa), θ_s is a curve parameter for J_e . $V_{p,max}$ is the maximum
226 electron rate for C_4 plant. J_i is the rate of light-dependent potential electron transport. I
227 is the radiation absorbed by leaves. α_i is the integrated C_4 quantum efficiency. $R(0)$
228 and $R_1(0)$ represent the direct radiation and downward diffuse radiation at the top of
229 the canopy, respectively. $fAPAR$ is the fraction of absorbed PAR .

230 The canopy radiative transfer in JSABCH is based on a two-stream approximation
231 (Meador and Weaver, 1980; Dickinson, 1983; Sellers et al., 1992). This approach



232 assumes (i) a uniform vertical distribution of leaves within the canopy, (ii) horizontal
233 homogeneity in radiation fields, (iii) purely vertical radiative fluxes, and (iv) identical
234 leaf reflectance and transmittance for direct and diffuse radiation across the entire PAR
235 spectrum. For GPP calculation, the procedure is consistent with that applied in iMAPLE.
236 In the second computational step, water-stressed photosynthetic rates are integrated
237 across all canopy layers to obtain ecosystem productivity. LAI is dynamically simulated
238 using the phenological model of LoGro-P (Dalmonech and Zaehle, 2013; Dalmonech
239 et al., 2015).

240

241 2.2.2 Water cycle of JSBACH

242 The soil hydrology module in JSBACH discretizes the soil column into five layers
243 extending to a depth of 10 meters (Hagemann and Stacke, 2015). Vertical soil water
244 movement is described using the Richards equation, while surface runoff and drainage
245 are calculated based on the Arno scheme (Richards, 1931; Todini, 1996). Transpiration
246 is regulated through root-zone soil layers, and carbon-water coupling is implemented
247 to represent interactions among hydrological and biogeochemical processes. The total
248 soil water content $\partial h_{tot}^{(i)}$ of each tile (i) can be expressed as:

$$249 \quad \rho_w \frac{\partial h_{tot}^{(i)}}{\partial t} = (1 - f_v)P - E_{bs}^{(i)} - E_{tr}^{(i)} + M_{sn}^{(i)} + M_{snc}^{(i)} - R_{srf}^{(i)} - R_d^{(i)} \quad (15)$$

250 where ρ_w denotes the density of water. f_v represents the fraction of precipitation (P)
251 intercepted by the canopy, E_{bs} is bare soil evaporation, and E_{tr} denotes transpiration.
252 M_{sn} and M_{snc} represent snowmelt at the surface and within the canopy, respectively.
253 R_{srf} is surface runoff and R_d is subsurface drainage.

254

255 2.3 Coupling between ECHAM6 and iMAPLE

256 Although both JSBACH and iMAPLE participated in the multi-model
257 intercomparison project for Global Carbon Budget, iMAPLE demonstrated superior
258 performance relative to JSBACH in simulating key carbon (e.g., GPP) and water (e.g.,
259 ET) fluxes (Friedlingstein et al., 2025). In addition, iMAPLE incorporates a more
260 comprehensive representation of biogeochemical processes, thereby enhancing the
261 simulation of interactions between atmospheric chemistry and terrestrial ecosystems.
262 In this study, iMAPLE was coupled to the ECHAM6 global circulation model (Fig. 1)



263 and configured to run in parallel with JSBACH, enabling a systematic comparison of
264 their process representations and model performance.

265 ECHAM6 integrates atmospheric dynamic processes at an interval of 7.5 minutes,
266 which is much shorter than the 60-minute interval used in iMAPLE. As a result,
267 simulated climatic variables from ECHAM6 are passed to iMAPLE every 8 ECHAM6
268 time steps. To ensure seamless information exchange, iMAPLE adopts the same core
269 functional modules as ECHAM6 for input/output (I/O) operations, memory
270 management, parallel computation, and calendar handling. At each coupling interval,
271 ECHAM6 provides hourly meteorological variables to iMAPLE, including
272 precipitation, surface air temperature, wind speed, humidity, surface pressure, CO₂
273 concentrations, direct and total radiation. In return, iMAPLE feeds back simulated soil
274 temperature, soil moisture, and *ET* to ECHAM6, thereby affecting the land surface
275 energy balance and atmospheric general circulation. Extensive numerical tests were
276 conducted to ensure integration stability and computational efficiency of the coupled
277 modeling framework.

278 The land cover dataset used in iMAPLE was developed by integrating satellite-
279 based observations from Moderate Resolution Imaging Spectroradiometer (MODIS)
280 and Advanced Very High Resolution Radiometer (AVHRR) (Defries et al., 2000;
281 Hansen et al., 2003), which include nine PFTs representing major global terrestrial
282 ecosystems. In contrast, JSBACH employs a flexible land-cover library that allows
283 users to specify the number and properties of PFTs. To reduce uncertainties arising from
284 differences in boundary conditions, the iMAPLE land cover dataset was harmonized
285 and mapped onto the JSBACH PFT categories (Table 1). Comparisons of simulated
286 GPP and LAI from ECHAM6-JSBACH showed improved performance when using the
287 updated PFT derived from iMAPLE (Table 2). Therefore, this updated land cover
288 configuration of ECHAM6-JSBACH is adopted for subsequent comparison with
289 ECHAM6-iMAPLE. We conducted a set of 15-year simulations spanning from 2000 to
290 2014 to evaluate and compare the performance of ECHAM6-iMAPLE and ECHAM6-
291 JSBACH. The first five years were used for model spin-up, and the remaining ten years
292 were averaged for analysis and performance assessment.

293

294 **2.4 Data for model evaluations**

295 Observational and reanalysis datasets from 2005 to 2014 were used to evaluate and
296 compare the performance of the ECHAM6-iMAPLE model and ECHAM6-JSBACH



297 models. Specifically, Global LAnd Surface Satellite (GLASS) products were used to
298 assess LAI. For GPP and ET, we used the benchmark product of FLUXCOM datasets
299 (Tramontana et al., 2016; Jung et al., 2020), which generate global gridded flux
300 estimates using machine learning algorithms trained on flux measurements from
301 FLUXNET eddy covariance sites. We also used the Modern-Era Retrospective analysis
302 for Research and Applications, Version 2 (MERRA-2) product (Gelaro et al., 2017) for
303 the evaluations of soil temperature, soil moisture, surface air temperature, and
304 precipitation. All datasets were interpolated to the T63 spectral resolution for consistent
305 comparisons, corresponding to a horizontal resolution of approximately 1.875° in both
306 latitude and longitude.

307

308 **3. Results**

309 **3.1 Terrestrial carbon cycle**

310 We evaluated the performance of ECHAM6-iMAPLE and ECHAM6-JSBACH in
311 simulating terrestrial carbon fluxes (Fig. 2). Both models reasonably capture the spatial
312 distribution of GPP consistent with FLUXCOM observations, characterized by high
313 values in tropical regions, moderate values in boreal forests, and low values in arid
314 areas. ECHAM6-iMAPLE simulates a global annual total GPP of $126.9 \text{ Pg C yr}^{-1}$,
315 slightly higher than that of $124.1 \text{ Pg C yr}^{-1}$ estimated by the FLUXCOM benchmark. In
316 contrast, ECHAM6-JSBACH produces a much higher global GPP of $134.2 \text{ Pg C yr}^{-1}$,
317 with pronounced positive biases mainly over tropical forest regions, including the
318 Indian subcontinent, the Amazon Basin, and Central Africa (Fig. 2c). Relative to
319 ECHAM6-JSBACH, ECHAM6-iMAPLE exhibits improved performance (Fig. 2b),
320 reflected by a higher correlation coefficient ($R = 0.75$, $p < 0.01$) and a lower root mean
321 square error ($\text{RMSE} = 1.51 \text{ g C m}^{-2} \text{ day}^{-1}$). This improvement can be primarily
322 attributed to the calibration of vegetation photosynthetic parameters in iMAPLE based
323 on site-level datasets (Yue and Unger, 2015), as well as the constraint of plant
324 phenology through satellite-based observations (Yue et al., 2015).

325 Consistent with the GPP pattern, simulated LAI from ECHAM6-iMAPLE (Fig. 2e)
326 shows a higher R of 0.86 ($p < 0.01$) and a lower RMSE of $0.73 \text{ m}^2 \text{ m}^{-2}$ against GLASS
327 retrievals (Fig. 2d) compared to ECHAM6-JSBACH (Fig. 2f). Regionally, ECHAM6-
328 iMAPLE successfully reproduces the elevated LAI in tropical rainforests, in good
329 agreement with observational patterns. Globally, ECHAM6-iMAPLE underestimates
330 LAI by about 9%, primarily due to notable underestimation across high-latitude regions



331 of the Northern Hemisphere (Fig. 2e). In contrast, ECHAM6-JSBACH shows a higher
332 RMSE of 0.79 relative to GLASS, with the largest negative biases occurring in the
333 central Africa and the Southeast Asia region (Fig. 2f).

334

335 **3.2 Terrestrial water flux**

336 We evaluated the simulated water fluxes from ECHAM6-iMAPLE and ECHAM6-
337 JSBACH (Fig. 3). Both models successfully reproduce the observed ET patterns,
338 characterized by hotspots over tropical rainforest and low values in arid and/or cold
339 regions. Simulated ET from ECHAM6-iMAPLE (Fig. 3b) exhibits a higher spatial
340 correlation with the FLUXCOM benchmark (Fig. 3a) ($R=0.86$, $p<0.01$) and a lower
341 RMSE (15.2 mm mon^{-1}) compared to ECHAM6-JSBACH (Fig. 3c). Globally, the area-
342 weighted mean ET simulated by ECHAM6-iMAPLE is 45.5 mm mon^{-1} , closely
343 matching the observed value of 45.9 mm mon^{-1} , whereas ECHAM6-JSBACH
344 underestimates ET at 42.1 mm mon^{-1} . The poorer performance of ECHAM6-JSBACH
345 is mainly attributed to substantial underestimations in arid and semi-arid regions,
346 particularly over the western United States and Australia (Fig. 3c).

347 For water use efficiency (WUE), defined as the ratio of GPP to ET, ECHAM6-
348 iMAPLE exhibits high values in the northern middle-to-high latitudes and moderate
349 values in the tropics (Fig. 3e). It achieves a spatial R of 0.52 ($p < 0.01$) and a RMSE of
350 $0.77 \text{ g C/kg H}_2\text{O}$ against the FLUXCOM benchmark (Fig. 3d). In contrast, ECHAM6-
351 JSBACH exhibits substantially poorer performance, with a lower spatial R of 0.20 and
352 a higher RMSE of $0.95 \text{ g C/kg H}_2\text{O}$ (Fig. 3f), which is primarily associated with
353 underestimated ET across mid-to-high northern latitudes and Central Africa. At the
354 global scale, the area-weighted mean WUE simulated by ECHAM6-iMAPLE is 1.43 g
355 $\text{C/kg H}_2\text{O}$, closer to the FLUXCOM estimate of $1.55 \text{ g C/kg H}_2\text{O}$ and representing a
356 clear improvement over the $1.35 \text{ g C/kg H}_2\text{O}$ simulated by ECHAM6-JSBACH (Fig.
357 3f).

358

359 **3.3 Seasonal variations**

360 In addition to the spatial pattern, ECHAM6-iMAPLE effectively captures the
361 seasonal cycles of terrestrial carbon and water fluxes (Fig. 4). Consistent with
362 observations, simulated GPP and ET both reach their annual maxima during the
363 vegetation growing season, with the most pronounced peak occurring in July (Figs 3a
364 and 3c). Although the LAI simulated by ECHAM6-iMAPLE exhibits a one-month lag



365 relative to the observations, the temporal correlation remains high ($R=0.96$), indicating
366 a good representation of seasonal dynamics (Fig. 4b). In contrast, ECHAM6-JSBACH
367 predicts the annual peaks of terrestrial carbon and water fluxes in June, one month
368 earlier than observed peaks. Moreover, the corresponding R between ECHAM6-
369 JSBACH simulations and observations are consistently lower than those derived from
370 ECHAM6-iMAPLE (Fig. 4). The divergence in peak timing and the systematically
371 lower R values collectively demonstrate that ECHAM6-iMAPLE provides a more
372 accurate representation of the seasonal dynamics of terrestrial carbon and water cycle
373 processes compared to ECHAM6-JSBACH.

374

375 **3.4 Soil properties and near-surface climate**

376 Soil temperature and moisture play critical roles in regulating vegetation growth
377 (Davidson and Janssens, 2006; Green et al., 2019). Both ECHAM6-iMAPLE (Fig. 5b)
378 and ECHAM6-JSBACH (Fig. 5c) predict high soil temperatures in tropical regions,
379 especially in northern Africa and the Indian Peninsula. The two models exhibit
380 extremely high spatial correlations with MERRA2 reanalysis data ($R=0.99$ and 0.98 ,
381 respectively). However, ECHAM6-iMAPLE more accurately reproduces the global
382 mean soil temperature and shows a lower RMSE of 2.16 K, whereas ECHAM6-
383 JSBACH shows a slight warm bias with a higher RMSE of 2.49 K. For soil moisture,
384 ECHAM6-JSBACH shows substantial overestimation in tropical regions and the high-
385 latitude Northern Hemisphere, leading to a 39.1% overestimation of the global mean
386 value compared to MERRA2 (Fig. 5f). In contrast, ECHAM6-iMAPLE closely
387 matches the global mean soil moisture from MERRA2, with a lower RMSE of 0.07 and
388 a moderate spatial R of 0.76 (Fig. 5e).

389 Changes in soil temperature and moisture influence surface meteorology through
390 land-atmosphere coupling processes (Seneviratne et al., 2006). We therefore compared
391 surface air temperature and precipitation simulated by ECHAM6-iMAPLE and
392 ECHAM6-JSBACH (Fig. 6). Both models reproduce the large-scale spatial distribution
393 of surface air temperature, albeit with overall warm biases (Figs 5b and 5c). However,
394 ECHAM6-iMAPLE performs better than ECHAM6-JSBACH, showing a smaller
395 global mean warm bias (0.34°C vs. 0.76°C) and a lower RMSE (3.33°C vs. 3.47°C).
396 This improvement is likely associated with the more accurate simulation of soil
397 temperature in ECHAM6-iMAPLE (Fig. 5b), highlighting the role of land surface
398 processes in constraining near-surface climate. For precipitation, however, both models



399 substantially overestimate the global mean amount with regional overestimation in Asia
400 but underestimation in central Africa (Figs 5e and 5f). Although coupling with iMAPLE
401 improves the simulation of ET (Fig. 3b) and soil moisture (Fig. 5e), the complex and
402 nonlinear feedbacks within the hydrological cycle may instead magnify biases in
403 precipitation estimates.

404

405 **4. Conclusion and Discussion**

406 The terrestrial biosphere and atmosphere are tightly coupled through exchanges of
407 energy, water, carbon, and reactive species (Green et al., 2017). These interactions
408 regulate climate variability, ecosystem resilience, and air pollutants such as ozone and
409 aerosols (Zeng et al., 1999; Yue and Unger, 2014; Zhou et al., 2024). However, climate
410 models often incompletely represent vegetation biophysical and physiological
411 processes, leading to biased carbon and water fluxes and increased uncertainties in
412 weather predictions (Koster et al., 2011; Guo et al., 2012). To address this limitation,
413 we coupled the iMAPLE vegetation model to the ECHAM6 atmospheric general
414 circulation model, and evaluated its performance against reanalysis, benchmark, and
415 satellite datasets. Compared to the original ECHAM6-JSBACH configuration, the
416 updated ECHAM6-iMAPLE model substantially improves simulations of terrestrial
417 carbon and water fluxes, more accurately capturing their spatial patterns and seasonal
418 variations.

419 These improvements stem from both rigorous parameter calibration and a more
420 realistic representation of key biophysical processes in iMAPLE. Critical physiological
421 parameters (e.g., V_{cmax}) are tightly constrained using *in situ* measurements from 201
422 FLUXNET sites (Yue et al., 2024), and phenological schemes have been validated
423 against thousands of ground-based records and multiple satellite retrievals (Yue et al.,
424 2015). Moreover, iMAPLE distinguished the biophysical effects of direct and diffuse
425 radiation (Yue and Unger, 2017), allowing improved quantifications of canopy radiation
426 transfer and photosynthetic carbon assimilation dynamics. Unlike JSBACH, which
427 applies a fixed carbon allocation scheme among leaf, wood, and reserve pools, iMAPLE
428 adopts a dynamic framework that allocates carbon to root, stem, and leaf respiration in
429 proportion to realistic LAI variations. Together, these advances enhance the ability of
430 ECHAM6-iMAPLE in simulating vegetation dynamics under climate change.

431 Despite these advances, several limitations remain in the current version of
432 ECHAM6-iMAPLE. First, iMAPLE does not include dynamic nitrogen and



433 phosphorus cycles, introducing uncertainties in simulated photosynthetic responses to
434 elevated CO₂ (Gruber and Galloway, 2008; Zaehle et al., 2011). Second, the vegetation
435 model prescribes land cover and neglects competitions among PFTs, preventing
436 simulation of vegetation shifts under anthropogenic or natural disturbances. Third, not
437 all vegetation properties simulated by iMAPLE are fully coupled to ECHAM6. For
438 instance, dynamically simulated LAI is not consistently used by ECHAM6 to update
439 land surface parameters or ecosystem emissions. In addition, iMAPLE considers the
440 variations of soil temperature and soil moisture only within upper two meters, limiting
441 the full prediction of deeper soil processes in the coupled system.

442 In future work, we will further improve ECHAM6-iMAPLE in several key aspects.
443 First, incorporating a process-based nitrogen cycle will substantially improve
444 simulations of terrestrial carbon sinks and reduce uncertainties in carbon flux estimates.
445 Second, we will strengthen the representation of natural source and sink processes. A
446 major advantage of iMAPLE lies in its dynamic simulation of natural emissions,
447 including wildfires, biogenic volatile organic compounds (BVOCs), and soil emissions.
448 In addition, dynamically simulated stomatal conductance influences the dry deposition
449 of atmospheric constituents. Future efforts will focus on improving the real-time
450 simulation and evaluation of these source-sink processes within the coupled framework.
451 Third, we plan to incorporate an interactive module of atmospheric chemistry into
452 ECHAM6-iMAPLE to enable fully coupled interactions among climate, ecosystems,
453 and atmospheric chemistry. This development will enhance capabilities for climate
454 projection and air pollution forecasting, and provide a useful tool for quantifying of the
455 long-term impacts of human activities on the Earth system through multi-sphere
456 interactions and feedbacks.

457

458 ***Code availability.***

459 The code and model description for ECHAM6-iMAPLE version 1 is available at
460 <https://doi.org/10.6084/m9.figshare.31877242> (Fu et al., 2026).

461

462 ***Author contributions.***

463 XY conceived the idea of coupling the two models. WF and CT contributed equally
464 to this work. WF and CT contributed to the coupling process, simulations, and result
465 analysis. YZ, YH, JH, and HC helped with code improvements. All authors contributed
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467

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472

473 ***Competing interests.***

474 The authors declare that they have no conflict of interest.

475

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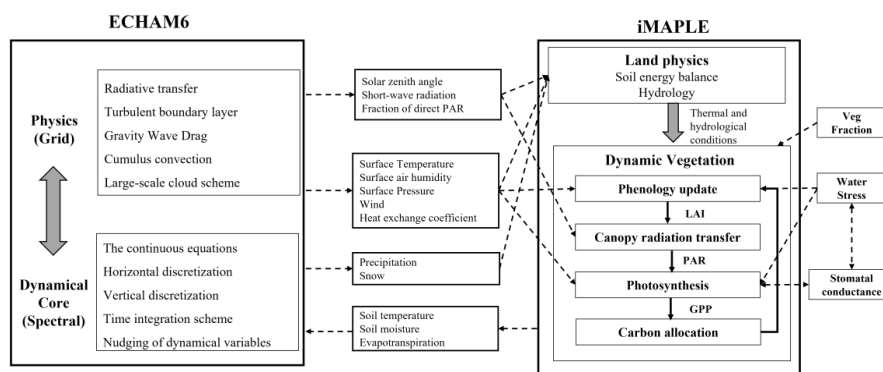
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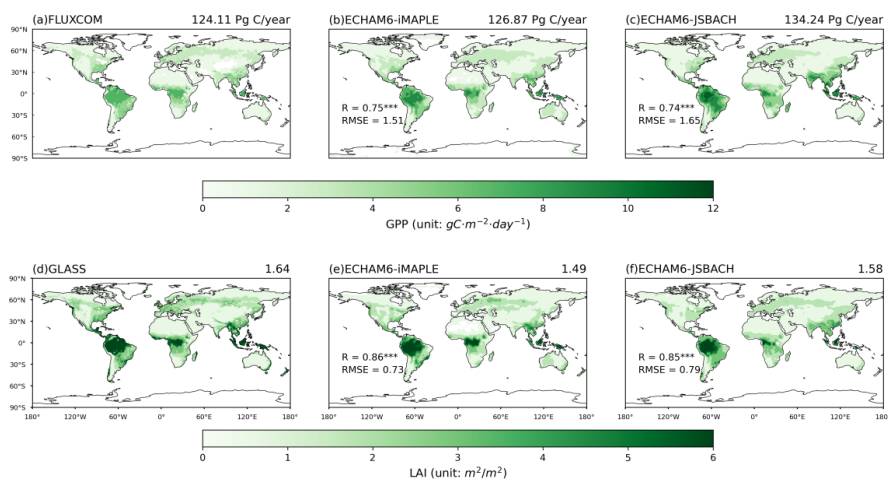


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803 **Figure 1.** The integrated modeling framework for ECHAM6-iMAPLE v1.0:

804 Atmosphere-Ecosystem Coupled Model.

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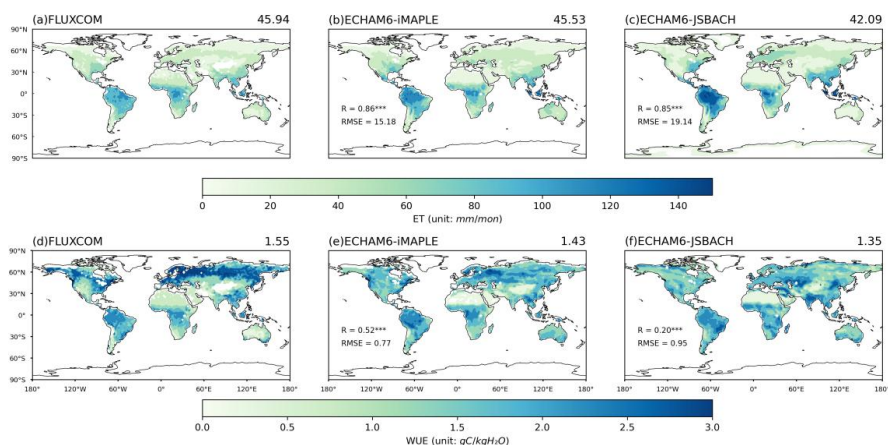


806

807 **Figure 2.** Spatial distributions of (a) gross primary productivity (GPP, $\text{g Cm}^{-2} \text{ day}^{-1}$)
808 from FLUXCOM and (d) leaf area index (LAI, m^2/m^2) from GLASS, compared with
809 simulations from (b, e) ECHAM6-iMAPLE and (c, f) ECHAM6-JSBACH. Both model
810 simulations and benchmark/observational datasets are averaged for the period of 2005-
811 2014. The spatial correlation coefficient (R) and root mean square errors (RMSE)
812 between simulations and the corresponding observations are shown in the lower left
813 corner of the simulation panels.

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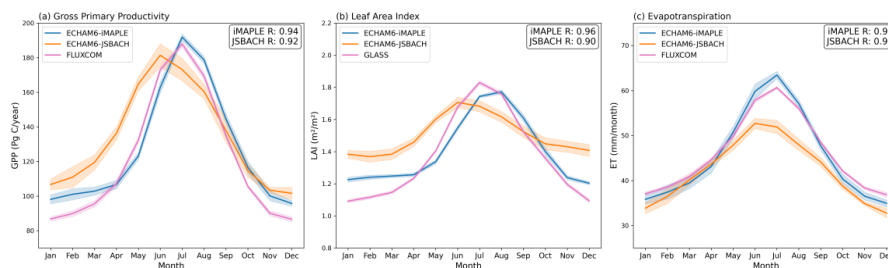


816

817 **Figure 3.** Spatial distributions of (a) evapotranspiration (ET, mm/month) and (d) water
818 use efficiency (WUE, g C/kg H₂O) from FLUXCOM, compared with simulations from
819 (b, e) ECHAM6-iMAPLE and (c, f) ECHAM6-JSBACH. Both model simulations and
820 the FLUXCOM benchmark datasets are averaged for the period of 2005-2014. The
821 spatial R and RMSE between simulations and benchmark are shown in the lower left
822 corner of simulation panels. The area-weighted mean ET and WUE are indicated at the
823 top of each panel.

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827 **Figure 4.** Comparison of the seasonal cycle of (a) GPP (Pg C/year), (b) LAI (m²/m²),

828 (c) ET (mm/month) (c). The pink lines represent observational or benchmark datasets.

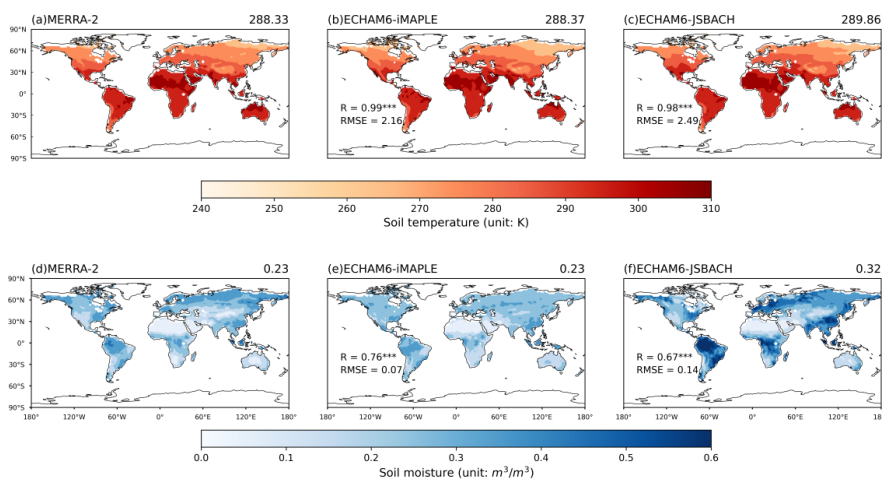
829 The blue and orange lines represent simulations from the ECHAM6-iMAPLE and

830 ECHAM6-JSBACH models, respectively. The shaded areas indicate the interannual

831 variability (one standard deviation) over the period 2005-2014.

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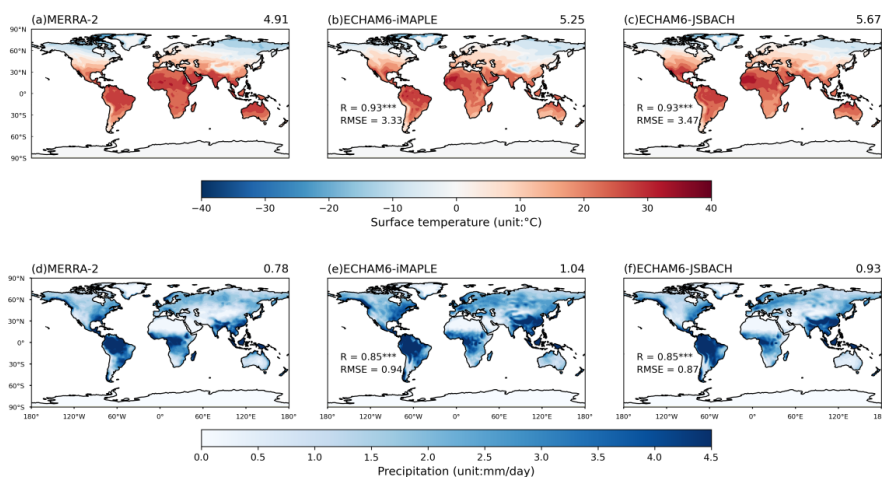


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835 **Figure 5.** Spatial distributions of (a) soil temperature (ST, K) and (d) soil moisture (SM,
836 m^3/m^3) from MERRA2 reanalyses, compared with simulations from (b, e) ECHAM6-
837 iMAPLE and (c, f) ECHAM6-JSBACH. Both model simulations and MERRA2
838 reanalyses are averaged for the period of 2005-2014. The spatial R and RMSE between
839 simulations and reanalyses are shown in the lower left corner of simulation panels. The
840 area-weighted mean ST and SM are indicated at the top of each panel.

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843

844 **Figure 6.** Spatial distributions of (a) surface temperature (T , °C) and (d) precipitation
845 (PREC, mm/day) from MERRA2 reanalyses, compared with simulations from (b, e)
846 ECHAM6-iMAPLE and (c, f) ECHAM6-JSBACH. Both model simulations and
847 MERRA2 reanalyses are averaged for the period of 2005-2014. The spatial R and
848 RMSE between simulations and observations are shown in the lower left corner of
849 simulation panels. The area-weighted mean T and PREC are indicated at the top of each
850 panel.

851



852 **Table 1** Projection of PFTs from JSBACH to iMAPLE

JSBACH	iMAPLE
Glacier and Tropical evergreen trees	evergreen broadleaf
Tropical deciduous trees	
Extra-tropical evergreen trees	evergreen needleleaf
Extra-tropical evergreen trees	deciduous broadleaf
Raingreen shrubs	arid adapted shrub
Deciduous shrubs	cold shrub
C ₃ grass and pasture	C ₃ grass
C ₄ grass and pasture	C ₄ grass
C ₃ crop	C ₃ crop
C ₄ crop	C ₄ crop

853

854



855 **Table 2** Comparison of Gross Primary Productivity (GPP) and Leaf Area Index (LAI)
856 between observational datasets and simulations from ECHAM6-iMAPLE and
857 ECHAM6-JSBACH with original/updated PFTs.

	GPP (Pg C/year)	LAI (m ² /m ²)
Observation	124.11	1.64
ECHAM6-iMAPLE	126.87	1.49
ECHAM6-JSBACH (original PFTs)	142.13	1.56
ECHAM6-JSBACH (updated PFTs)	134.24	1.58

858