



1 **The Dominance and Future Variations of Background Ozone in China under**
2 **Different Climate and Emission Scenarios**

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22 **Abstract**

23 Background ozone, defined as ozone present in the absence of domestic
24 anthropogenic emissions, has become increasingly dominant of the total surface ozone
25 in China as the domestic emissions decline. Using the GEOS-Chem model driven by
26 meteorological fields from GCAP2.0, this study investigates the changes in China
27 background ozone (CBO) from the present-day (PD) to the middle of the 21st century
28 (MC) under the SSP1-1.9 and SSP5-8.5 scenarios. CBO consistently dominates the
29 total surface ozone in China (CTO), accounting for about 90% of the annual mean
30 ozone in both PD and MC conditions. Under SSP1-1.9, CTO declines by 4.7 ppbv from
31 PD to MC, with CBO contributing 3.7 ppbv (79%) of this reduction. The decrease in
32 CBO is largely due to a sharp decline in transboundary transport of foreign pollution
33 ozone (FPO, -4.7 ppbv) resulting from global emission reductions, partially offset by a
34 climate-driven increase in natural background ozone (NBO, +1.0 ppbv). Under SSP5-
35 8.5, the annual mean CTO increases by 7.3 ppbv, of which 6.5 ppbv is attributed to the
36 rise in CBO. The increase is primarily driven by the enhancement of NBO, associated
37 with intensified photochemical production under warmer conditions. We also show that
38 in a low-emission future with controlled warming, the reduction in CBO strengthens
39 climate mitigation and creates a positive climate-ozone feedback, along with health co-
40 benefits. These contrasting trajectories reveal that background ozone is a dynamic
41 component shaped by global cooperation and climate change, a reality that must be
42 considered when setting air quality targets.

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45 **1. Introduction**

46 Tropospheric ozone, primarily formed through the photochemical oxidation of
47 volatile organic compounds (VOCs) and carbon monoxide (CO) in the presence of
48 nitrogen oxides ($\text{NO}_x = \text{NO} + \text{NO}_2$) under solar radiation, is a critical air pollutant with
49 significant implications for both regional and global air quality (Yin et al., 2021;
50 Vazquez Santiago et al., 2024; Zhou et al., 2025; Ni et al., 2026). It is also detrimental
51 to human health and ecosystems (Emberson, 2020; Wang et al., 2023), and plays a
52 pivotal role in atmospheric oxidizing capacity and climate dynamics (Ainsworth et al.,
53 2012; Ma et al., 2016; Lyu et al., 2023). The prevention or control of ozone pollution
54 has remained a significant challenge in China due to the complex photochemical
55 processes and the varying sources of ozone precursors (Ding et al., 2019; Chen et al.,
56 2023; Wang et al., 2025).

57 Background ozone, defined as ozone present in the absence of domestic
58 anthropogenic emissions, includes contributions from both natural source precursor
59 emissions (e.g., soil NO_x , lightning NO_x , and biogenic VOCs) and transboundary
60 transport (Jaffe et al., 2003, 2018; Yan et al., 2021). It is also defined as the Policy-
61 Relevant Background ozone, representing the portion of ozone that cannot be reduced
62 under current policy frameworks, and thus provides a baseline for assessing the realistic
63 potential of ozone mitigation (Wang et al., 2009). Studies indicated that the increasing
64 trend of background ozone had exacerbated the frequency and intensity of high-ozone
65 episodes in China, posing a growing challenge for effective air pollution mitigation
66 (Chen et al., 2023). Long-term and effective ozone pollution management must
67 consider not only local precursor emissions but also the growing influence of
68 transboundary transport and climate-sensitive natural sources, reflecting the central
69 importance of background ozone for developing air quality and climate policies that are
70 both effective and adaptable to a changing climate. Given the persistent challenge of
71 ozone pollution in China and its direct implications for public health and ecosystem
72 sustainability, understanding the evolution of background ozone in this region is a
73 matter of pressing policy relevance.

74 Several studies have recognized the critical role of background ozone and have



75 undertaken quantitative assessments of its source-specific contributions. Ni et al. (2018)
76 quantified the contribution of transboundary transport to near-surface ozone in China
77 using the GEOS-Chem model and showed that foreign anthropogenic emissions
78 contributed 2–11 ppbv (parts per billion by volume) to ozone concentrations during
79 spring. Meanwhile, Ye et al. (2024) examined the role climate-sensitive natural sources
80 in ozone formation, finding that lightning NO_x, soil NO_x, and biogenic VOCs contribute
81 summertime background ozone in China by 8.1, 6.4, and 3.9 ppbv, respectively. The
82 increasing dominance of background ozone under climate change, amplified by
83 regional transport and climate-sensitive natural sources, may offset the deficits
84 achieved through local emission mitigation efforts (Turnock et al., 2019; Skipper et al.,
85 2021). Given the growing influence of background ozone in shaping surface ozone
86 concentrations, particularly as domestic anthropogenic emissions decline, it is crucial
87 to understand how emissions and climate change shape background ozone levels.

88 To effectively address future risks of ozone pollution to human health and
89 ecosystems, it is essential to understand how ozone concentrations are likely to evolve
90 and what factors will govern those changes under different climate and emission
91 scenarios. The Shared Socioeconomic Pathways (SSPs) describe plausible future
92 evolutions of societal and environmental systems, detailing assumptions on population,
93 economy, technology, and environmental development (O'Neill et al., 2014; Riahi et
94 al., 2017). While local emission reductions are expected to lower ozone levels, studies
95 suggest that surface ozone may still exceed health standards due to increasing
96 background contributions. Using the GEOS-Chem model, Xu et al. (2024) found that
97 the maximum daily 8-hour average ozone in major polluted and agricultural regions of
98 China would surpass the WHO guideline of 100 µg m³ under most SSPs emissions
99 scenarios by 2050. Wang and Liao (2022) emphasized the role of transboundary
100 transport based on GEOS-Chem simulations, noting that emissions from South and
101 Southeast Asia contributed substantially to ozone levels in China, and such impacts
102 could be mitigated under low-emission scenarios (e.g., SSP1-1.9). Using the
103 Community Earth System Model with an ozone tagging method, Hou et al. (2023)
104 projected that biogenic sources could contribute up to 40% of summer surface ozone



105 concentrations over eastern China under a net-zero scenario, reflecting the growing role
 106 of natural emissions in shaping future ozone pollution patterns. These studies indicate
 107 that background ozone, driven by natural emissions and regional transport, is becoming
 108 an increasingly important factor in determining ozone pollution levels in a warmer
 109 climate.

110 While most previous research has focused on anthropogenic emission related-
 111 changing ozone concentrations, often overlooking the distinct evolution of background
 112 ozone. By using the chemical transport model GEOS-Chem, we investigate the
 113 temporal and spatial evolution of background ozone over China and its underlying
 114 drivers in present-day and middle of the 21st century under two diverse future scenarios,
 115 SSP1-1.9 and SSP5-8.5. Decomposing background ozone into natural and foreign-
 116 pollution components highlights how the interaction between global emission
 117 reductions and climate warming shapes the baseline that local policies must address.
 118 This understanding is crucial for informing the development of robust, climate-resilient
 119 environmental policies in China, with implications that extend to public health and
 120 climate feedbacks.

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122 **2. Data and methods**

123 **2.1 Model Description and Meteorological Inputs**

124 The chemical transport model GEOS-Chem version 14.2.3 is employed to
 125 simulate present-day ozone concentrations and project future ozone over China. As a
 126 global three-dimensional atmospheric chemistry model, it integrates a detailed
 127 mechanism of tropospheric ozone-NO_x-hydrocarbon chemistry mechanism (Bey et al.,
 128 2001; David et al., 2019) and has been extensively applied in studies of atmospheric
 129 composition (Colombi et al., 2015; Li et al., 2020; Hu et al., 2024). The chemical
 130 mechanism incorporates hundreds of species and reactions, enabling realistic
 131 simulation of ozone formation and removal processes under varying environmental
 132 conditions. To simulate the interaction between stratospheric and tropospheric ozone,
 133 the model utilizes the linearized ozone parameterization (McLinden et al., 2000). The
 134 model also integrates an advanced diagnostic system that quantifies the contributions



135 of chemical and physical processes, such as net chemical production, vertical
 136 convective transport, horizontal advection, dry deposition and diffusion, to ozone
 137 concentrations. Radiative forcing is quantified using the coupled radiative transfer
 138 module, the Rapid Radiative Transfer Model. Simulations from GEOS-Chem have been
 139 extensively evaluated using satellite, aircraft, and ground-based observations, and
 140 successfully applied in numerous regional and global studies related to tropospheric
 141 ozone (Shah et al., 2024; Wang et al., 2025; Zhai et al., 2024).

142 Meteorological inputs for the GEOS-Chem simulations in this study are provided
 143 by the Global Change and Air Pollution version 2.0 (GCAP 2.0) framework, which
 144 supplies meteorological fields under various CMIP6 (Coupled Model Intercomparison
 145 Projection Phase 6) climate scenarios across past, present, and future periods. More
 146 detailed information on GCAP 2.0 can be found in Murray et al. (2021). GCAP 2.0
 147 provides representative time slices, including pre-industrial (1851–1860), present-day
 148 (2001–2014), mid-century (2040–2049), and end-of-century (2090–2099), which are
 149 generated under the historical condition and multiple future SSPs scenarios. The
 150 meteorological fields include key atmospheric variables such as temperature, wind
 151 components, humidity, and precipitation, which are provided at a horizontal resolution
 152 of 2° latitude $\times$ 2.5° longitude with an hourly temporal frequency. These datasets have
 153 been widely applied in studies of climate-chemistry interactions and scenario-based air
 154 quality assessments (Kang et al., 2025; Zhao et al., 2024; Murray et al., 2024). The
 155 GCAP 2.0 data during the present-day period (2011–2014, PD) and the mid-century
 156 period (2046–2049, MC) under SSP1-1.9 and SSP5-8.5 are applied in this study to
 157 represent current and projected future atmospheric conditions, respectively. SSP1-1.9
 158 represents a sustainability-oriented pathway with aggressive mitigation measures and a
 159 target radiative forcing of 1.9 W m^{-2} by 2100, while SSP5-8.5 reflects a fossil-fueled
 160 development trajectory with weak environmental regulation and a high radiative
 161 forcing of 8.5 W m^{-2} .

162 **2.2 Emission and Boundary Conditions**

163 Anthropogenic emissions used in the study are derived from the CMIP6
 164 framework. The emissions inventory includes monthly data of aerosols, and ozone



precursors (e.g., NO_x, CO, and VOCs), provided at a spatial resolution of 0.5°, and has been preprocessed for compatibility with the GEOS-Chem model (<http://atmos.earth.rochester.edu/input/gc/ExtData>). Historical anthropogenic emissions are sourced from the Community Emissions Data System (CEDS), while future emissions follow SSPs scenarios developed by integrated assessment models. Surface boundary conditions for long-lived species, such as methane (CH₄) and nitrous oxide, are prescribed as global mean surface volume mixing ratios using historical reconstructions from Meinshausen et al. (2017) and future concentrations defined under the SSPs framework (Riahi et al., 2017).

GEOS-Chem accounts for a range of natural emission sources, some of which are calculated online within the model. Lightning NO_x emissions are parameterized following Murray et al. and Ott et al. (2010) based on convective cloud-top height and flash rates. Soil NO_x emissions are computed based on the scheme by Hudman et al. (2012), which are sensitive to surface temperature, soil moisture, cloud cover, wind speed, and radiation. Biogenic VOCs emissions are computed by the Model of Emission of Gases and Aerosols from Nature (MEGAN v2.1; Guenther et al., 2012), with fluxes dynamically varying according to photosynthetically active radiation, surface temperature, and soil moisture. These climate-sensitive natural emission processes provide important background sources of ozone precursors and are included consistently across all simulation scenarios.

2.3 Experimental Design

To investigate the spatial and temporal variability of background ozone over China and to quantify the relative contributions of natural sources and transboundary transport, we conduct a series of GEOS-Chem simulations. The baseline series of simulations (BASE) adopt the above emission and meteorological settings for both PD (2011–2014) and MC (2046–2049) periods following the SSP1-1.9 and SSP5-8.5 pathways. The resulting ozone concentrations over China are referred to as total ozone (CTO). To isolate the contribution of background ozone, a set of sensitivity simulations (NO_CHN) is conducted, in which anthropogenic emissions of ozone precursors from China, including NO_x, CO, and NMVOCs, are excluded in the simulations. The resulting ozone



195 concentrations, driven exclusively by natural sources and the transboundary transport
 196 of ozone or its precursors from sources outside China, are referred to as China
 197 background ozone (CBO). A set of global sensitivity simulations (NO_ANT) is also
 198 conducted, in which all anthropogenic emissions of ozone precursors are turned off
 199 worldwide. The resulting ozone concentrations represent the natural background ozone
 200 (NBO). A summary of simulations is provided in Table 1.

201 The difference between CTO and CBO represents the net impact of domestic
 202 anthropogenic emissions, which is defined as China pollution ozone (CPO). Likewise,
 203 the difference between the CBO and the NBO reflects the contribution from sources
 204 outside China and is referred to as foreign pollution ozone (FPO). Based on this
 205 modeling configuration, surface ozone concentrations in the simulations can be
 206 approximately decomposed according to the following relationship:

$$207 \quad \text{CTO} = \text{CBO} + \text{CPO}, \text{CBO} = \text{NBO} + \text{FPO}.$$

208 Simulations are conducted on a horizontal grid of 2° latitude $\times$ 2.5° longitude
 209 resolution and 40 vertical layers extending from the surface to 0.1 hPa. While the
 210 relatively coarse resolution limits the ability of the model to resolve urban-scale
 211 chemical gradients, it is appropriate for examining large-scale background ozone
 212 patterns across China. To ensure the stability and reliability of the simulated ozone
 213 responses to changes in climate and natural emissions, all GEOS-Chem simulations are
 214 conducted over a four-year period, preceded by an 18-month spin-up to minimize the
 215 influence of initial conditions.

216

217 **3. Model evaluation**

218 Figure. 1 compares the spatial distribution of seasonal mean surface ozone
 219 concentrations from China National Environmental Monitoring Center (CNEMC) and
 220 GEOS-Chem simulations in 2014. CNEMC provides air quality monitoring data since
 221 2013 and this study selects 2014 as the baseline year for model evaluation, as CNEMC
 222 provides much more data in 2014 than 2013 and the historical GCAP 2.0 meteorological
 223 data ends in 2014. The model successfully reproduces the observed seasonal variation
 224 of ozone that peaks in summer and reaches their lowest levels in winter. It also captures



the spatial variability across China, with correlation coefficients ranging from 0.30 to 0.69 ($p < 0.01$ for all seasons). The normalized mean bias (NMB) analysis reveals consistent positive biases across all seasons, with NMB values of +23%, +16%, +19%, and +17% in winter, spring, summer, and autumn, respectively, indicating a model tendency to overestimate surface ozone concentrations. The bias is consistent with previous studies using GEOS-Chem over China and is likely attributable to uncertainties in anthropogenic emission inventories, incomplete chemical mechanisms, and the model's limited spatial resolution (Gong and Liao, 2019; Lu et al., 2019; Yang et al., 2022). Nevertheless, the model remains robust for capturing large-scale spatial and seasonal ozone patterns.

Given the limited availability of long-term surface ozone observations in China, satellite retrievals of tropospheric column ozone (TCO), obtained by combining data from the Aura Ozone Monitoring Instrument (OMI) and the Microwave Limb Sounder (MLS), serve as an important complement for evaluating model performance (https://acd-ext.gsfc.nasa.gov/Data_services/cloud_slice/new_data.html). Figure. 2 compares the annual mean TCO from the GEOS-Chem BASE simulation with satellite retrievals from OMI/MLS during 2011–2014. The model successfully reproduces the overall spatial distribution of observed TCO, including the characteristic west-to-east gradient, but slightly overestimates values, especially in regions with high pollution. Over eastern China (east of 100°E), the observed mean TCO is 40.9 DU, while the model simulates a higher value of 43.2 DU. In western China (west of 100°E), the observed and simulated values are 34.3 DU and 35.1 DU, respectively.

While some discrepancies between model simulations and observations are evident due to spatial resolution limitations, emission inventory uncertainties, and simplified chemistry, the biases remain within an acceptable range. Despite a general tendency of GEOS-Chem to overestimate both surface and tropospheric ozone, the model reproduces key spatiotemporal patterns. The comprehensive validation against multiple observational datasets underscores the credibility of the GEOS-Chem model and justifies its application in subsequent diagnostic and attribution analysis.



4. Background ozone from present-day to the mid-21st century

To establish a baseline for future scenario analysis, the spatial distribution of annual mean background ozone over China under the PD condition is firstly shown in Fig. 3. Spatially, CBO exhibits a distinct west-to-east decreasing gradient, with concentrations higher than 50 ppbv in western China and lower than 30 ppbv over the North China Plain (NCP) in eastern China. The simulated background ozone concentrations in this study were evaluated against available data from the literature (Chen et al., 2025, Fig. S1 and Tab. S1). The favorable regional comparison confirms the reliability of our simulations for the 2011–2014 period. On a national scale, the averaged CBO concentration is 42.7 ppbv, constituting 90% of the CTO (47.5 ppbv) in PD. The modeled ratio aligns well with previous studies using various models and methodologies, which have reported CBO contributions between 71% and 94% (Wang et al., 2011; Lu et al., 2019; Ye et al., 2024; Sahu et al., 2021). This spatial pattern of CBO largely resembles that of CTO, confirming that CBO is the dominant component of the ozone burden in most regions of China. In contrast, the CPO is relatively modest, averaging 4.8 ppbv nationwide. Building on the sensitivity simulations, the CBO is further decomposed into NBO (33.3 ppbv) and FPO (9.4 ppbv). NBO accounts for 78% of CBO, and its spatial pattern closely aligns with that of CBO, underscoring the key role of natural sources in ozone pollution in China. FPO is markedly elevated along the eastern coastline and the southwestern borders, clearly pinpointing transboundary transport as an important contributor to CBO (22%).

The spatial distribution and national averaged ratio of annual mean CBO to CTO is presented in Fig. 4. As the substantial emission reduction under SSP1-1.9, the ratio of CBO to CTO increases from 89.9% in PD to 91.2% in MC, while the ratio does not change under SSP5-8.5 (89.7%). The dominance of CBO is particularly evident in western China under SSP1-1.9, where anthropogenic precursor emissions are minimal. Despite the decrease in CTO under SSP1-1.9 and increase in CTO under SSP5-8.5 in mid-21st century relative to PD, the relative contribution of CBO remains consistently high, highlighting its critical role in shaping current and future ozone pollution in China.

Figure. 5 illustrates the changes in annual mean CTO and its major source



284 components in MC under the SSP1-1.9 and SSP5-8.5 scenarios relative to PD. Under
 285 the SSP1-1.9 scenario (Fig. 5a-e), CTO decreases across most of China except over the
 286 NCP, with a national mean reduction of 4.7 ppbv. The future ozone reduction under
 287 SSP1-1.9 is primarily driven by a 3.7 ppbv decline in CBO, which accounts for
 288 approximately 79% of the total change. The strong reductions occur over western and
 289 border regions, reflecting the effects of coordinated global mitigation efforts in
 290 suppressing background ozone levels. CPO also decreases by 1.1 ppbv, reflecting the
 291 impact of domestic precursor emission controls under the SSP1-1.9 scenario.
 292 Nevertheless, a regional increase in CPO is identified over NCP, causing the elevated
 293 level of CTO over this region. NBO exhibits a modest increase (+1.0 ppbv), associated
 294 with enhanced photochemical activity under a warmer climatic condition. Meanwhile,
 295 FPO declines more substantially (−4.7 ppbv), leading to an overall reduction in CBO.
 296 The decrease in FPO exceeds the climate-driven increase in NBO, reflecting the global
 297 benefits of coordinated emission mitigation strategies.

298 Under the SSP5-8.5 scenario (Fig. 5f-j), CTO exhibits a significant nationwide
 299 growth by 7.3 ppbv in MC relative to PD, with approximately 90% attributed to
 300 elevated CBO (6.5 ppbv). NBO is the dominant contributor to the increase in CBO,
 301 with a projected rise of 6.7 ppbv, indicating a substantial climate-driven enhancement
 302 in natural emissions in a much warmer future. FPO shows negligible change on the
 303 national scale (−0.2 ppbv), yet localized enhancements are observed over southwestern
 304 China, which contribute to the increased concentrations in that region. CPO exhibits a
 305 modest increase of 0.8 ppbv under SSP5-8.5, with relatively greater changes occurring
 306 in eastern China, contributing the CTO enhancement. These results indicate that, under
 307 a high-emission and unleashed-climate future, surface ozone increases in China will be
 308 primarily dominated by background sources.

309 To further elucidate the regional characteristics of future changes in background
 310 ozone under the diverse scenarios, we examine the seasonal mean near-surface ozone
 311 in China and its contributors under future emission and climate scenarios in Figs. 6–8.
 312 The absolute concentrations of different components in the PD and MC under the
 313 SSP119 and SSP585 scenarios are provided in Figs. S2–S4. Compared to the PD



condition, CTO under the SSP1-1.9 scenario shows a clear seasonal pattern, with the most pronounced reduction occurring in summer (−10.2 ppbv), followed by spring (−5.1 ppbv), autumn (−4.0 ppbv). The reduction in summer is largely driven by a substantial decrease in CPO, which declines by 6.9 ppbv, reflecting the effectiveness of stringent domestic emission controls implemented under the SSP1-1.9 sustainability scenario. In spring and autumn, the decline in CTO is primarily attributed to the reduction in CBO, which accounts for 92% and 83% of the decreases in CTO, respectively, with the most significant reductions occurring in western China and border regions. In winter, the nationally averaged CTO exhibits a slight increase of 0.3 ppbv, accompanied by pronounced spatial heterogeneity. The increase is most evident over the NCP, whereas most regions in western China show a decline. The enhanced CTO over the NCP is primarily attributed to a rise in CPO, likely resulting from weakened NO titration due to declining NO_x emissions, which facilitates ozone accumulation under VOC-limited conditions (Hou et al., 2023; Kang et al., 2025). A similar but weaker increase in ozone is also observed in spring and autumn over the NCP, which may be partly attributed to the potential bias in ozone prediction under the low emission scenario in most global models (Ni et al., 2026). Nevertheless, it does not largely affect the focus of background ozone evaluation in this study.

The future changes in seasonal CBO can be further decomposed into the changes in NBO and FPO (Figs. 8a-h). The rise in NBO (0.5–1.5 ppbv) in MC under SSP1-1.9 relative to PD, mainly over eastern China, reflects strengthened ozone production from natural sources due to climate-driven increases in natural emissions (e.g., biogenic VOCs and soil NO_x emissions). In contrast, FPO decreases considerably in all seasons (−3.8 ~ −6.2 ppbv), especially over western China and border regions where the influence of transboundary transport is most evident, peaking in spring with strong westerly jets. As a result, the modest rise in NBO is offset by the substantial decline in FPO, leading to a net decrease in CBO under SSP1-1.9.

Under the SSP5-8.5 scenario, CTO shows a marked nationwide increase relative to PD, with seasonal enhancements ranging from 6.4 ppbv in winter to 8.1 ppbv in summer (Fig. 7). CBO is the dominant contributor to the increased CTO, constituting



344 87–91% of the total enhancement. The increase in CBO is particularly pronounced
 345 central and western China, suggesting that regions with limited anthropogenic
 346 emissions may face growing challenges from background ozone pollution in a much
 347 warmer future. The increase in ozone from domestic anthropogenic sources remains
 348 relatively small, ranging from +0.7 to +1.0 ppbv, with spatial increases mainly confined
 349 in urbanized eastern China.

350 The increase in CBO under SSP5-8.5 is primarily driven by the enhanced NBO
 351 (Fig. 8i–p), which exhibits a rise across the country, with national increases of 5.8–7.4
 352 ppbv. This widespread growth in the NBO highlights that climate-sensitive precursors,
 353 including both increased natural emissions and rising CH₄ concentrations, have become
 354 the primary drivers of ozone increase under the strong warming scenario. The
 355 contribution of FPO shows insignificant changes, ranging from 0.1 to 0.4 ppbv,
 356 although localized increases are still showed in southwestern China, which further
 357 contribute to the increase in CBO. These findings suggest that under the SSP5-8.5,
 358 future ozone pollution in China will be increasingly dominated by natural sources,
 359 posing greater challenges for effective mitigation through domestic emission controls
 360 alone.

361 The vertical distribution of ozone provides insights into how its components
 362 respond differently to changes in climate and emissions at various altitudes. Figs. 9–11
 363 compare vertical differences in ozone components between MC and PD under SSP1-
 364 1.9, SSP5-8.5, with a focus on eastern China (averaged over 110°E–122.5°E), a hotspot
 365 of human exposure to ozone pollution. Under the SSP1-1.9 scenario, the wintertime
 366 ozone increase in the lower troposphere is primarily due to a rise in CPO from domestic
 367 anthropogenic emissions near the surface (Fig. 9). This contrasts with the upper-level
 368 decreases driven by a reduction in CBO, which highlights the distinct vertical impacts
 369 of local chemical production versus long-range transport from foreign regions. A
 370 similar pattern of near-surface CTO increase is also evident over the NCP during spring
 371 and autumn. In summer, CTO decreases significantly by more than 10 ppbv from the
 372 surface to the middle of the troposphere, primarily due to the reduced CPO, which
 373 reflects the effectiveness of domestic emission controls under high photochemical



activity. The vertical extension of NBO increase implies a climate-related enhancement of natural precursor emissions that affects ozone throughout the lower troposphere (Fig. 11). However, the NBO increase does not suffice to offset the reduction in transboundary contributions. The concurrent increase in NBO and decrease in FPO further illustrates the interplay between climate-amplified natural source contributions and successful global coordination of emission controls, highlighting the dual challenges and co-benefits of integrated climate-air quality management.

In contrast to SSP1-1.9, the SSP5-8.5 scenario leads to widespread ozone increases over eastern China below the middle of the troposphere that are consistent across all seasons (Fig. 10). The pervasive vertical increase in CTO under SSP5-8.5 is primarily attributed to the enhanced CBO, signaling the growing dominance of background processes over local anthropogenic contributions. NBO exhibits a top-down enhancement pattern, with stronger increases in the mid-troposphere than near the surface, suggesting that downward transport from higher atmospheric layers may significantly contribute to the rise in background ozone.

5. Mechanistic drivers of the changing background ozone

As mentioned above, the future changes in CBO are controlled by transboundary transport related to the changes in foreign anthropogenic emissions under SSP1-1.9 and background sources in a warmer climate under SSP5-8.5. In this section, mechanistic drivers of the changing CBO are further illustrated by analyzing the future changes in anthropogenic emissions (Fig. 12), meteorological fields (Fig. S5) and the associated physical and chemical processes (Fig. 13).

Under SSP1-1.9, anthropogenic emissions show widespread and substantial reductions, with global annual emissions of NO_x , CO, and NMVOCs decreasing by 52.5 Tg yr^{-1} (−68%), 409.4 Tg yr^{-1} (−68%), and 88.2 Tg yr^{-1} (−74%) in MC relative to PD, respectively. Apart from China, regions such as South Asia, Southeast Asia, Europe, and North America also show considerable reductions, which collectively weaken the transboundary transport of ozone and precursor gases into China under the low emission scenario. Under the SSP5-8.5 scenario, changes in global anthropogenic precursor emissions are relatively modest. Significant increases of NO_x emissions are



also observed in South Asia, explaining the weak increase in FPO over southwestern China (Fig. 5j). CO emissions show an overall decline, with a smaller reduction compared to SSP1-1.9, and global NMVOC emissions decrease slightly by approximately 6.7 Tg yr^{-1} .

Changes in key meteorological factors in MC relative to PD show that under SSP1-1.9, surface temperature over eastern China increases by approximately 1 K, accompanied by enhanced surface shortwave radiation and reduced relative humidity. These changes promote higher photochemical reaction rates and stimulate more natural precursor emissions, driving an increase in net photochemical production of ozone by $+52.5 \text{ Gg yr}^{-1}$ in China. While changes in circulation-related meteorological parameters, such as wind patterns, are relatively modest, the significant reduction in anthropogenic precursor emissions from surrounding regions weakens the background ozone supply to China, leading to a substantial negative contribution from horizontal advection (-84.8 Gg yr^{-1}). This weakening of transboundary contribution (FPO) is consistent with the finding in Wang and Liao (2022), which reported a decrease of approximately 1.6 ppbv in the transport contribution from South and Southeast Asia to ozone over China between 2015 and 2050 under SSP1-1.9 with fixed climate conditions but varied anthropogenic emissions. Therefore, the decrease in CBO under SSP1-1.9 is primarily driven by large-scale reductions in global anthropogenic emissions and the associated weakening of transboundary influence, while enhanced chemical production under the favorable meteorological conditions partly offsets this effect.

Under SSP5-8.5, surface temperature undergoes sharp increases across China. However, in eastern China, where photochemical activity is most pronounced, anomalies in surface shortwave radiation and relative humidity are minor. Despite the modest response of several meteorological variables, temperature is widely recognized as a key meteorological driver modulating future ozone changes over China (Kang et al., 2025). In the context of stronger warming, the budget diagnosis reveals that the net chemical production increases by 70.6 Gg yr^{-1} , which indicates an enhancement in NBO and emerges as the dominant process driving the rise in CBO under SSP5-8.5. Horizontal advection contributes negatively (-38.4 Gg yr^{-1}), suggesting that transport



434 acts to weaken the rise in tropospheric background ozone, but the overall response
435 remains dominated by chemical production. The process-based budget diagnostics are
436 consistent with the sensitivity attribution, jointly indicating the future changes in CBO
437 are dominated by the weakened foreign anthropogenic influence under SSP1-1.9 and
438 enhanced climate-driven chemistry under SSP5-8.5.

439 In the simulations, future changes in CBO can also be affected by the time-varying
440 CH₄ concentrations. To isolate the impact of CH₄, an additional set of simulations
441 (NO_CH₄low) with the CH₄ concentrations fixed at the pre-industrial level (1850 year)
442 is conducted, yielding the pre-industrial natural background ozone (PNB). Compared
443 to the PD level of 28.5 ppbv in China (Fig. S6), PNB rises by 2.6 and 3.8 ppbv under
444 SSP1-1.9 and SSP5-8.5, respectively, indicating that climate change alone, even
445 without future methane growth, can enhance CBO. Therefore, while controlling CH₄
446 remains essential for mitigating background ozone, climate adaptation strategies should
447 also account for the persistent contribution from non-methane natural sources under
448 warming conditions, including enhanced biogenic VOCs emissions, soil NO_x release,
449 lightning activity, and stratosphere-to-troposphere exchange that contribute to elevated
450 background ozone levels.

451 **6. Conclusion and Discussion**

452 Background ozone defines the baseline for regional pollution and the air quality
453 improvements achievable through domestic emission controls, making its
454 understanding crucial for informing ozone standards and guiding control strategies. In
455 this study, future changes in near-surface ozone concentrations across China in mid-
456 21st century (MC) under the SSP1-1.9 and SSP5-8.5 scenarios relative to the present-
457 day (PD) condition are quantified based on GEOS-Chem model, focusing on the
458 dominant role of background ozone and its evolving source composition. The analysis
459 reveals that CBO remains a major contributor to ozone in China, accounting for about
460 90% of annual mean CTO in both present-day and future climate and emission status.
461 The persistent dominance of background ozone points to the need for a strategic shift
462 from local measures toward regional coordination, global cooperation, and climate
463 mitigation, as the potential for further air quality improvements through local emission



464 controls alone is increasingly limited.

465 Under the SSP1-1.9 scenario, CTO decreases by 4.7 ppbv relative to the PD,
466 driven by a 3.7 ppbv reduction in CBO (accounting for 79% of the total decline). The
467 decrease in CBO can be further attributed to changes in the contributions from NBO
468 and FPO. FPO declines significantly by 4.7 ppbv, reflecting the effective suppression
469 of transboundary transport related to global emission reductions, which is the major
470 factor driving the reduction in CBO. In contrast, NBO increases slightly by 1.0 ppbv,
471 driven by enhanced photochemical production due to climate warming and increased
472 shortwave radiation. The budget diagnostics identify the reduced horizontal transport
473 (-84.8 Tg yr^{-1}) as the key mechanism driving CBO reduction, confirming that
474 weakened transboundary transport from surrounding regions, because of global
475 emission cuts, is the primary cause of the decline in CBO. These findings underscore
476 the effectiveness of globally coordinated mitigation in reducing CBO, despite the
477 climate-driven NBO increases. The seasonal CTO changes follow a consistent pattern
478 across most regions and seasons, except for a modest winter increase in the NCP driven
479 by the weakened NO_x titration.

480 Under the SSP5-8.5 scenario, CTO increases by 7.3 ppbv, with approximately 90%
481 of this rise attributed to CBO, underscoring its continued dominance in driving ozone
482 levels. Unlike the SSP1-1.9 scenario, the increase in CBO is primarily driven by a
483 substantial rise in NBO (6.7 ppbv), reflecting the enhanced photochemistry and the
484 significant contribution of climate-sensitive natural sources to ozone concentration
485 under the intensified warming scenario. An increase in net photochemical production
486 of 70.6 Tg yr^{-1} confirms the dominance of chemical processes. Under a high-emission,
487 warming future, ozone pollution in China will be increasingly governed by natural
488 sources, suggesting the requirement of integrated climate-air quality management in the
489 future.

490 We also calculate the population-weighted CTO/CBO (Fig. S8 and Text S1) and
491 changes in radiative forcing at the top of the atmosphere, which are more related to
492 human health and climate change, respectively. Even under SSP1-1.9, the population-
493 weighted CTO will increase, attributed to the increase in CPO over the densely



494 populated eastern China, highlighting the importance of coordinated regional
 495 mitigation. The population-weighted CTO and CBO rise sharply under SSP5-8.5,
 496 indicating that climate-driven growth can elevate population health risks. The overall
 497 decline in CBO under SSP1-1.9 produces a negative change in radiative forcing of
 498 -0.12 W m^{-2} over China, while the increase in CBO under SSP5-8.5 leads to a positive
 499 radiative forcing change of $+0.11 \text{ W m}^{-2}$. Climate mitigation suppresses background
 500 ozone levels, creating a co-benefit that reinforces warming reduction, whereas
 501 unchecked warming drives ozone increases that exacerbate health risks and impose an
 502 additional burden on the climate system.

503 This study quantifies the contributions of anthropogenic and background sources
 504 to ozone levels through emission shutdown sensitivity experiments, offering a clear and
 505 interpretable framework for isolating source contributions under changing climate and
 506 emission scenarios. Although widely used in source attribution studies, the approach
 507 may still be subject to uncertainties due to nonlinear interactions, particularly within
 508 the complex tropospheric chemistry. To evaluate the uncertainty, additional
 509 perturbation experiments are conducted with 20% and 80% reductions in domestic
 510 emissions. The results indicate that the nonlinear residual is approximately 0.5 ppbv
 511 (Fig. S9), which accounts for about 10% of the ozone response under the 100%
 512 reductions simulation. This suggests that the impact of nonlinear effects is relatively
 513 small, further confirming the robustness and reliability of the emission shutdown
 514 approach employed in this study. Future work could further explore source-specific and
 515 nonlinear responses under evolving climate and emission conditions using advanced
 516 attribution techniques such as tagged tracers. In addition, as climate warming continues
 517 and anthropogenic emissions decline, natural sources are expected to play an
 518 increasingly important role in modulating background ozone. These findings highlight
 519 the need to better quantify the contributions of individual natural sources to ozone
 520 changes, thereby supporting the development of climate-resilient and source-specific
 521 air quality management strategies.



522 **Author contributions.** YY designed the research; ML performed the simulations and
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525

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529

530 **Competing interests.** At least one of the (co-)authors is a member of the editorial board
531 of Atmospheric Chemistry and Physics.

532

533 **Code and data availability.** The observed surface ozone concentrations in 2014 are derived from
534 the China Ministry of Ecology and Environment (<https://quotsoft.net/air/>). The monthly mean
535 tropospheric ozone data from OMI–MLS are downloaded from https://acd-ext.gsfc.nasa.gov/Data_services/cloud_slice/new_data.html. The GEOS-Chem model is available
536 at http://wiki.seas.harvard.edu/geos-chem/index.php/GEOS-Chem_14.2.3. The climate outputs
537 from GCAP 2.0 can be downloaded from
538 <http://atmos.earth.rochester.edu/input/gc/ExtData/GCAP2/CMIP6/>. Population density across
539 China was obtained from the Gridded Population of the World version 4 (GPWv4) dataset
540 (<https://www.earthdata.nasa.gov/data/projects/gpw>).

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802 **Table 1.** Experimental design in this study.

Simulation	Emission settings	Ozone output
BASE	Full-chemistry simulation including all anthropogenic emissions and online natural emissions.	CTO
NO_CHN	Same as BASE, but with anthropogenic emissions of ozone precursors removed within China.	CBO; CTO – CBO = CPO
NO_ANT	Same as BASE, but with anthropogenic emissions of ozone precursors removed globally.	NBO; CBO – NBO = FPO
NO_CH ₄ low	Same as NO_ANT, but with CH ₄ fixed at the year-1850 level.	PNB

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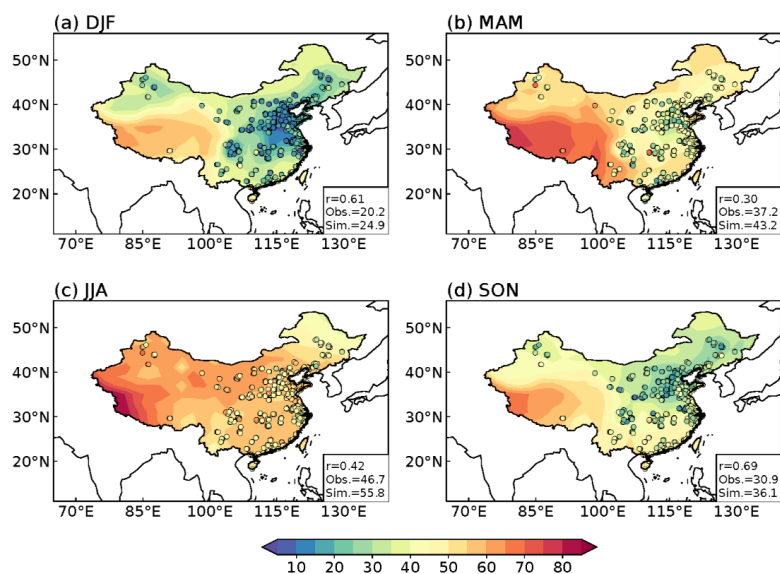


Figure 1. Spatial distribution of seasonal mean surface ozone concentrations (ppbv) over China in (a) winter (December-January-February, DJF), (b) spring (March-April-May, MAM), (c) summer (June-July-August, JJA), and (d) autumn (September-October-November, SON) in 2014 from GEOS-Chem simulation (colored background) and CNEMC observations (circles). Spatial correlation coefficients (r) between observed and simulated ozone concentrations and regional mean values from observation (Obs.) and simulation (Sim.) over the observational sites are indicated in bottom right of each panel.

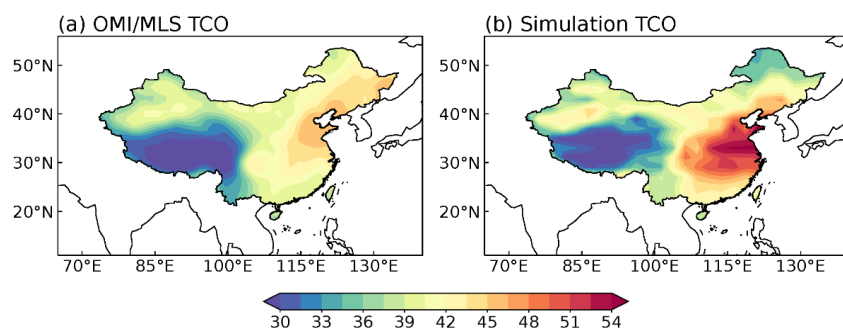


Figure 2. Spatial distribution of annual mean tropospheric column ozone (TCO, units: DU) over China during 2011–2014 based on (a) satellite retrievals from OMI/MLS and (b) GEOS-Chem BASE simulation.

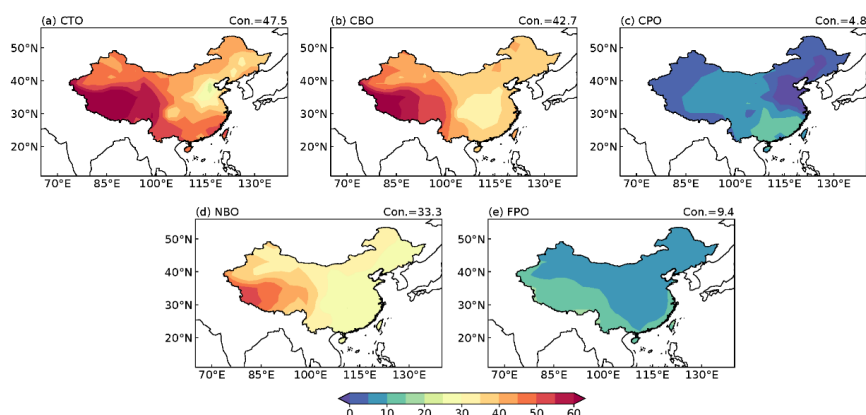
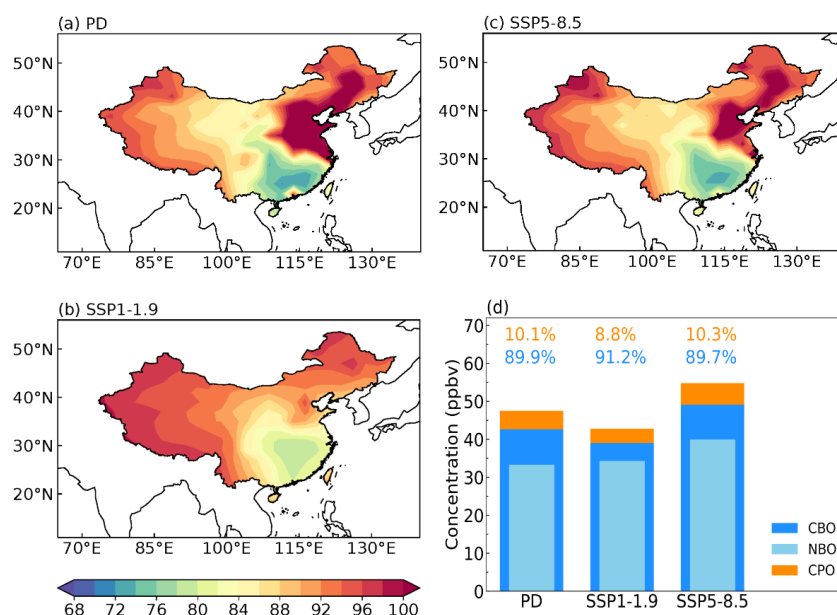


Figure 3. Spatial distributions of simulated annual mean concentrations (ppbv) of (a) China total ozone (CTO), (b) background ozone (CBO), (c) China pollution ozone (CPO), (d) natural background ozone (NBO), and (e) foreign pollution ozone (FPO) over China in PD. The regional mean value in China is shown in top right of each panel.



829

830 **Figure 4.** Spatial distributions of the background-to-total ozone ratio (CBO/CTO, %)
 831 in (a) PD and MC under (b) SSP1-1.9 and (c) SSP5-8.5 scenarios, and national average
 832 concentrations of CTO and its components (d). Bars represent CPO (orange), CBO
 833 (blue), and NBO (light blue) contributions. Percentages above bars indicate relative
 834 contributions of CPO and CBO to CTO.

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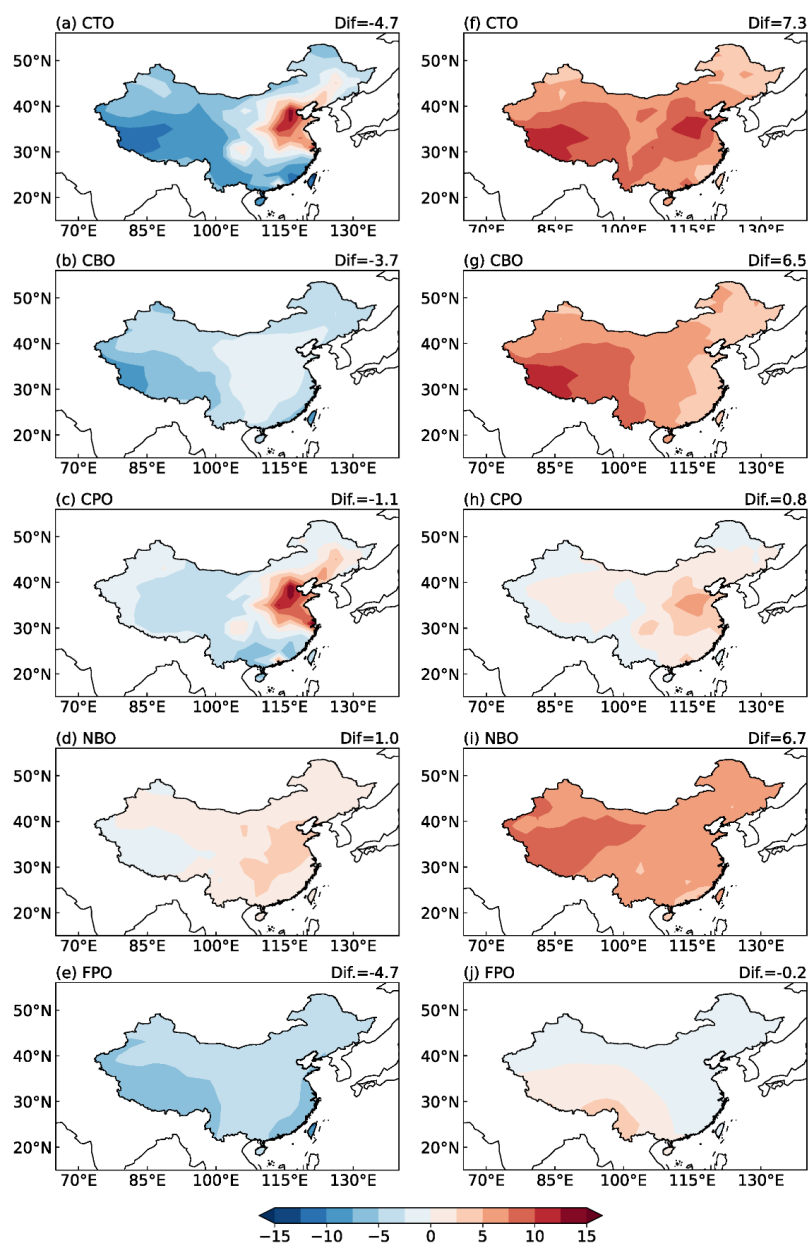
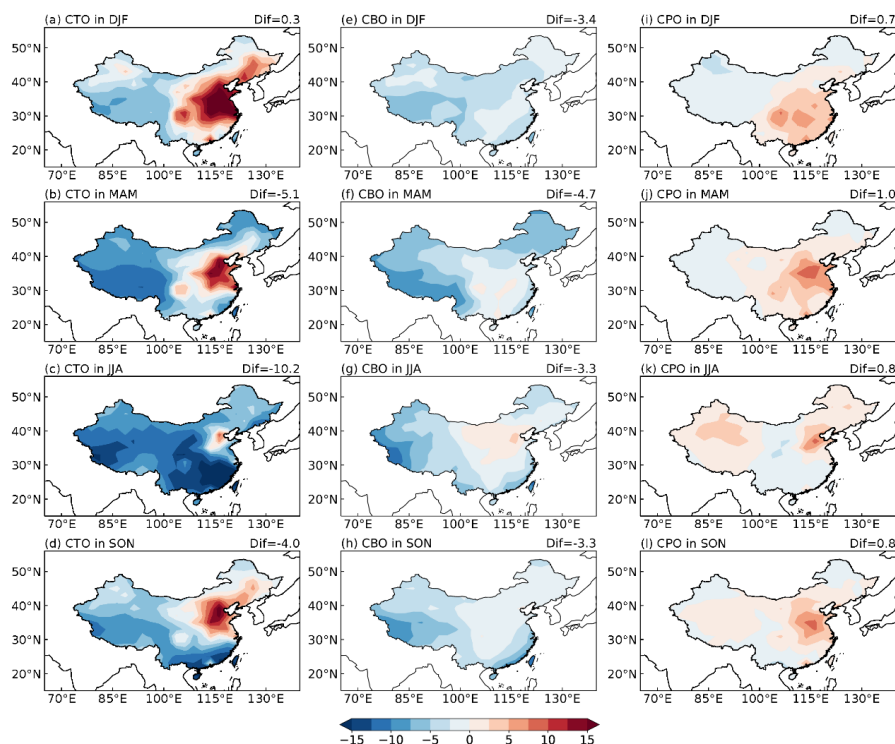


Figure 5. Spatial distributions of projected changes in annual mean surface ozone and its source components (CTO, CBO, CPO, NBO, and FPO) across China (Units: ppbv) in MC under (a–e) SSP1-1.9 and (f–j) SSP5-8.5 relative to PD. The value in the upper-right corner of each panel represents the regional mean in China.



843

844 **Figure 6.** Spatial distribution of the differences in seasonal mean CTO, CBO, and CPO
 845 between MC under SSP1-1.9 and PD (Units: ppbv). The value in the upper-right corner
 846 of each panel represents the regional mean in China.

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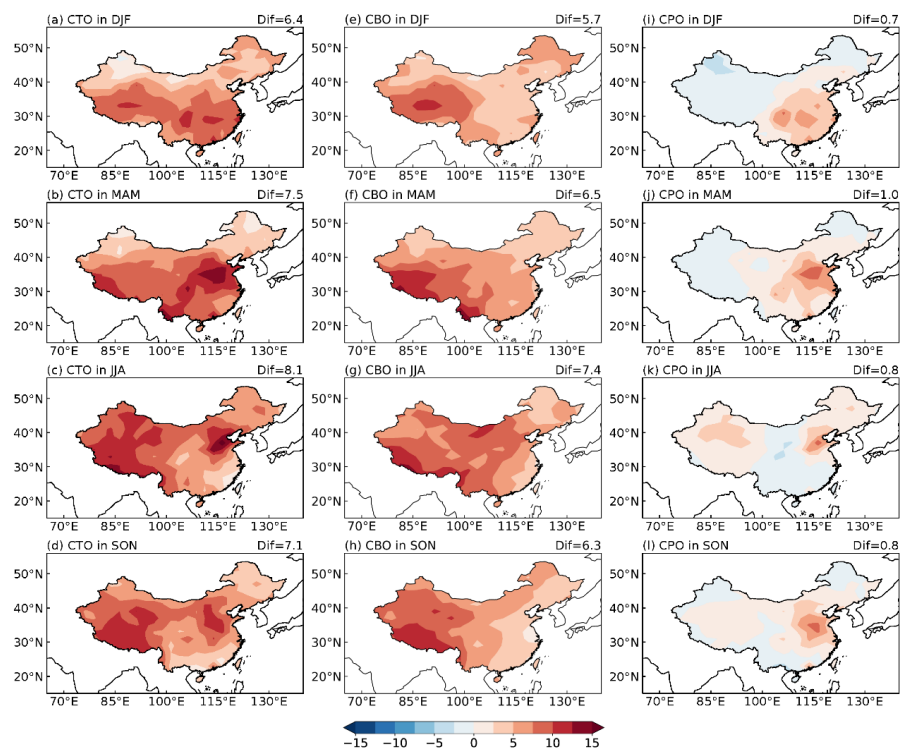


Figure 7. Spatial distribution of the differences in seasonal mean CTO, CBO, and CPO between MC under SSP5-8.5 and PD (Units: ppbv). The value in the upper-right corner of each panel represents the regional mean in China.

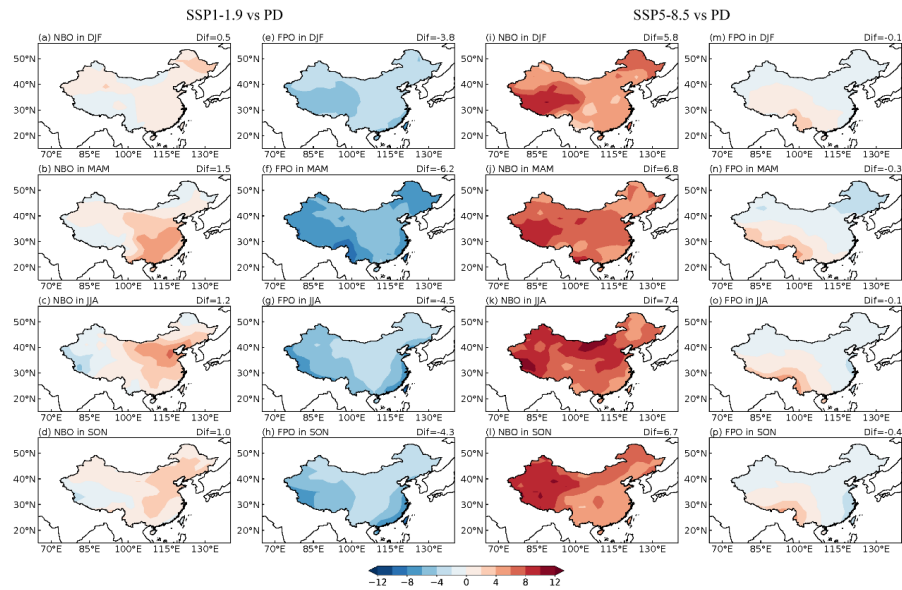


Figure 8. Spatial distribution of the differences in seasonal mean NBO and FPO between MC under SSP1-1.9 (a-h) and SSP5-8.5 (i-p) and PD (Units: ppbv). The value in the upper-right corner of each panel represents the regional mean in China.

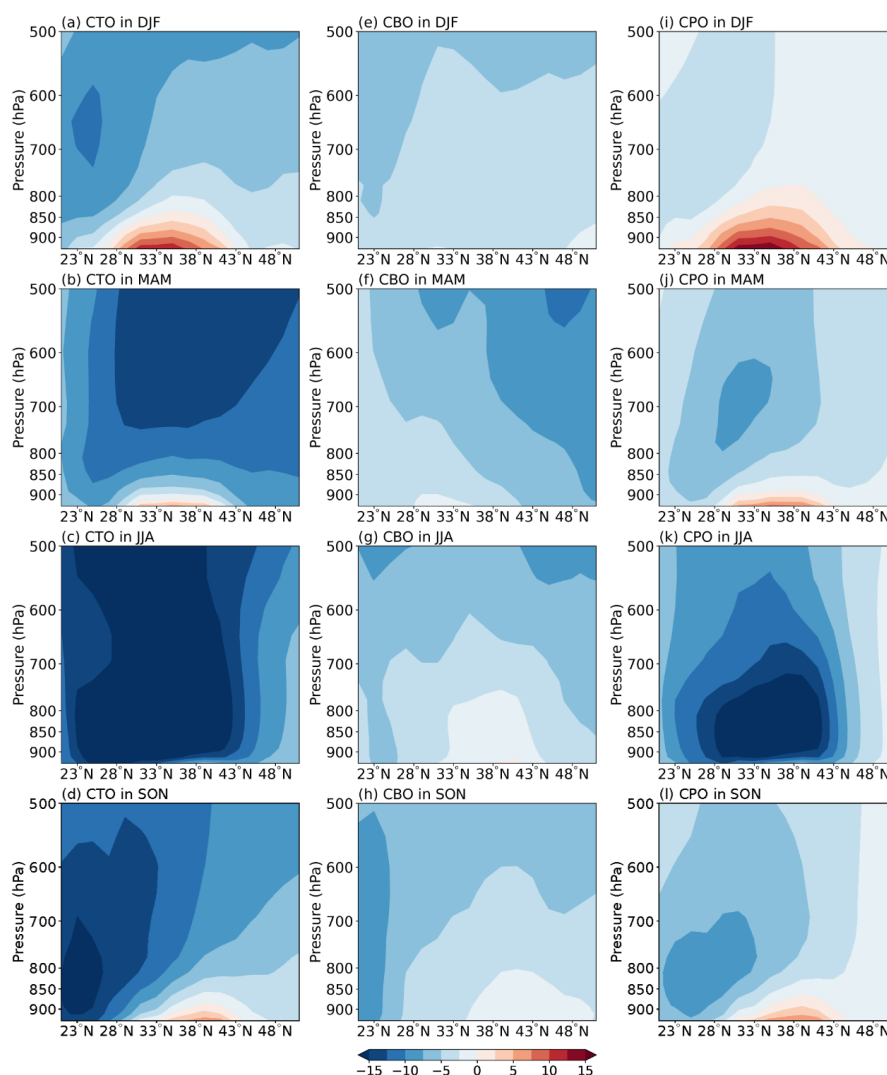


Figure 9. Pressure–latitude cross-section averaged over 110–122.5°E for differences in simulated seasonal mean CTO, CBO, and CPO between MC under SSP1-1.9 and PD (Units: ppbv).

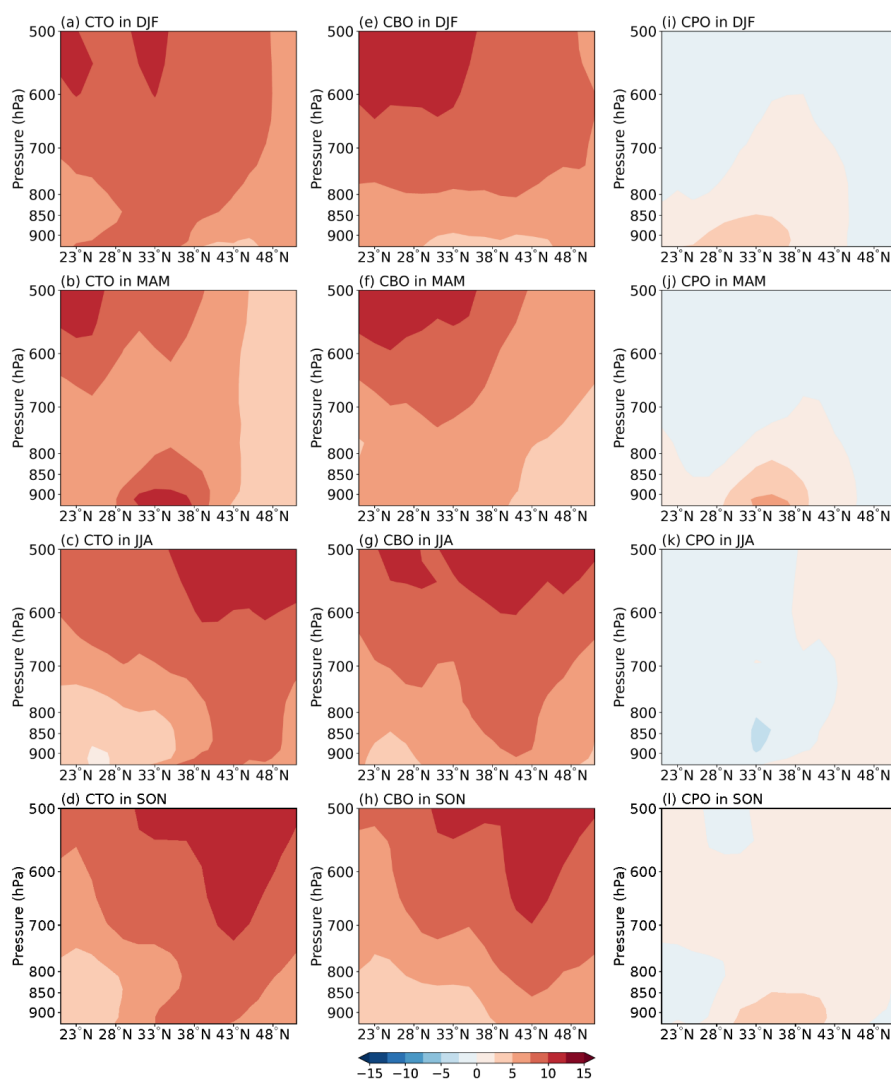


Figure 10. Pressure–latitude cross-section averaged over 110–122.5°E for differences in simulated seasonal mean CTO, CBO, and CPO between MC under SSP5-8.5 and PD (Units: ppbv).

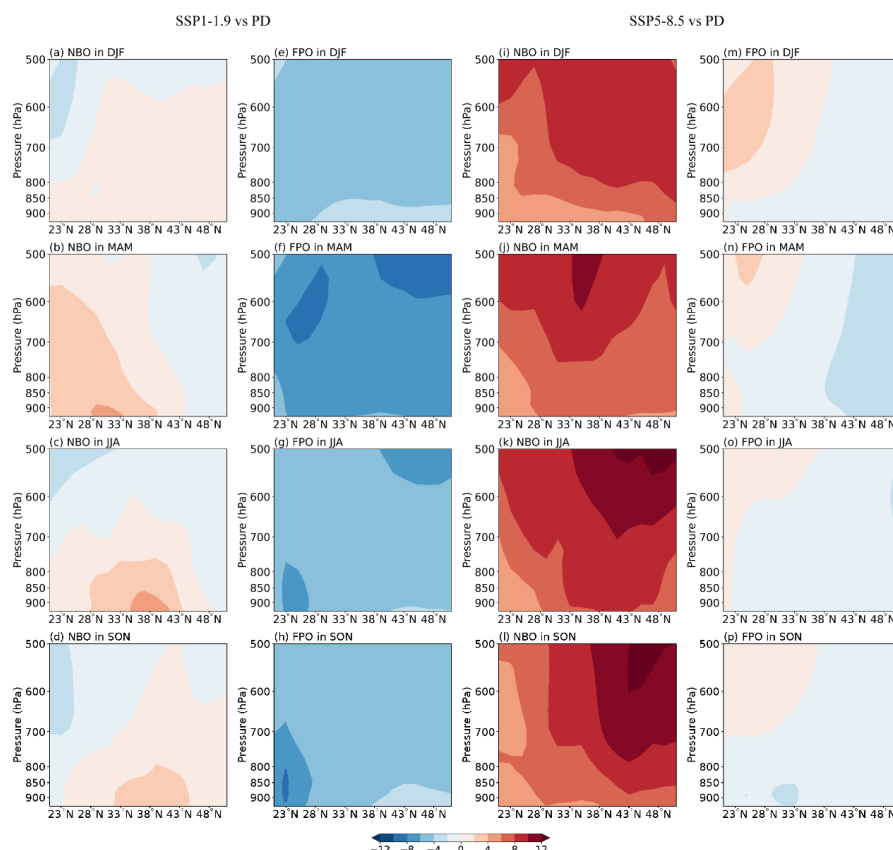
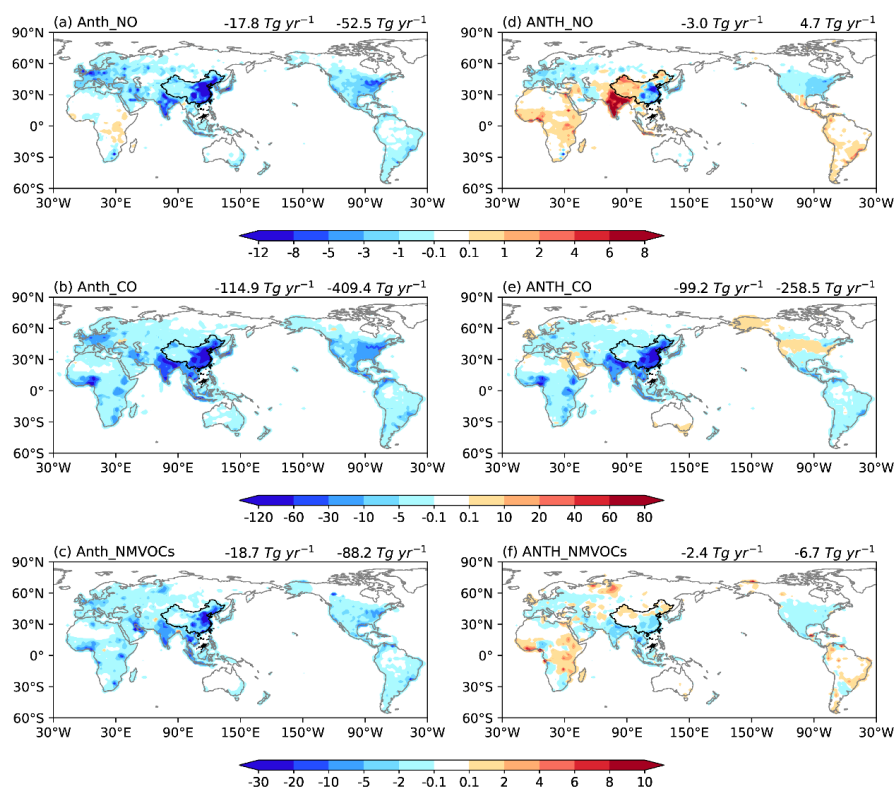
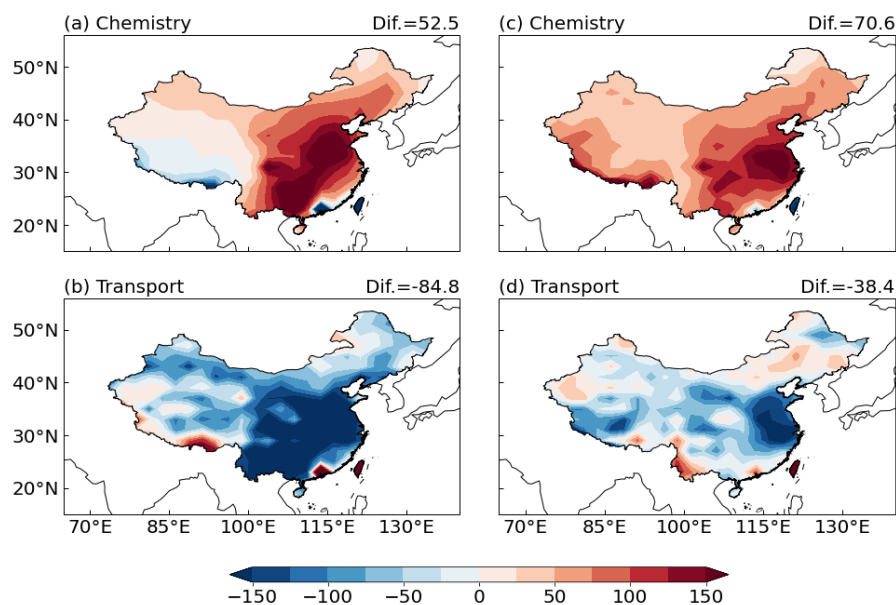


Figure 11. Pressure–latitude cross-section averaged over 110–122.5°E for differences in simulated seasonal mean NBO and FPO between MC under SSP1-1.9 and SSP5-8.5 and PD (Units: ppbv).



875
 876 **Figure 12.** Spatial distribution of anthropogenic emission rates changes in (a) NO_x, (b)
 877 CO, and (c) NMVOCs in MC under the SSP1-1.9 (left) and SSP5-8.5 (right) scenarios
 878 relative to PD (Units: $10^{-11} \text{ kg m}^{-2} \text{ s}^{-1}$). The values at the top of each panel are changes
 879 in annual total anthropogenic emissions for China and globe (Units: Tg yr^{-1}).
 880
 881



882
 883 **Figure 13.** Spatial distribution of the difference in tropospheric ozone mass budgets of
 884 the (a, c) net chemical production (Chemistry), and (b, d) horizontal advection
 885 (Transport) between MC under SSP1-1.9 (left) and SSP5-8.5 (right) relative to PD.
 886 These results are derived from the NO_CHN experiments. The value in the upper-right
 887 corner of each panel represents the regional mean in China (units: Gg yr⁻¹).