



Assessing site-to-site greenhouse gas measurement consistency in the UK and Ireland atmospheric monitoring network using whole and synthetic air reference materials

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Abstract. Atmospheric greenhouse gas (GHG) measurements are essential for assessing climate change and verifying national emission inventories. As global networks rely on high-precision observations of potent GHGs such as carbon dioxide (CO₂), methane (CH₄) and nitrous oxide (N₂O), ensuring accuracy between observation stations is essential. Maintaining accuracy within a network requires the use of whole air reference materials (RMs) with well-defined calibration scales, such as those
20 maintained by the National Oceanic and Atmospheric Administration (NOAA) for the World Meteorological Organisation (WMO). With the global expansion of atmospheric monitoring networks there is a growing need for readily available RMs, stimulating the development of complementary alternatives such as synthetic air RMs traceable to the International System of Units (SI).

In the United Kingdom (UK) and Ireland, the atmospheric GHG monitoring network, comprising long-term atmospheric
25 monitoring stations, is equipped with high precision in situ GHG analysers. These observations are used to verify bottom-up assessments of national emissions. Here, a case study is presented which provides the first direct, network-wide evaluation of whole air versus synthetic air RMs under routine field conditions, allowing an assessment of the implications of SI-traceable RMs within an active national atmospheric monitoring system. We assess the UK and Ireland's network accuracy using WMO-NOAA scale-traceable whole air RMs and SI-traceable synthetic air RMs, through a blind multi-site round-robin
30 intercomparison. Between 2021 and 2022, measurements were made of CO₂, CH₄ and N₂O in whole and synthetic air RMs at six sites using a range of instruments and the results were used to assess biases between sites. The measured mole fractions for each cylinder, at each site were then compared with the assigned value on the WMO-NOAA scale and the SI-traceable values for the whole air and synthetic air RMs, respectively. Weighted residual analysis reveals that whole air RMs generally



meet the WMO/Global Atmospheric Watch (GAW) compatibility goals for CO₂ and CH₄, while N₂O remains more difficult
35 due to its small atmospheric variability and instrument performance. Synthetic SI-traceable RMs have associated absolute
uncertainties on their assigned values for CH₄ that are comparable to the extended WMO compatibility goals. However,
absolute uncertainties for CO₂ and N₂O remain significantly larger than both the compatibility and extended compatibility
goals, which is consistent with previous studies. Overall, these results indicate that while scale-traceable whole air RMs remain
essential for achieving the highest level of network compatibility required for atmospheric monitoring, SI-traceable synthetic
40 RMs could play a valuable complementary role, particularly for CH₄. However, it is first critical that improved characterisation
of matrix effects, mole fraction dependant instrument responses, and long-term stability, are achieved for them to become
suitable for high-precision atmospheric monitoring.

1 Introduction

Atmospheric greenhouse gas (GHG) measurements are crucial for understanding climate change (Petit et al., 1999;
45 Ramanathan and Feng, 2009; IPCC, 2023), therefore, the World Meteorological Organisation (WMO) has identified them as
critical for global monitoring. As of 2023, mean global surface temperatures are 1.6°C warmer on land compared to the pre-
industrial era (IPCC, 2023), with climate projections indicating an increase of over 2°C by 2100 (IPCC, 2023). Global warming
is driven by the rise in GHGs mole fractions, primarily due to human activities (IPCC, 2023), with carbon dioxide (CO₂),
methane (CH₄) and nitrous oxide (N₂O) being the most significant contributors due to their strong radiative forcing and long
50 atmospheric lifetimes. Atmospheric CO₂, CH₄ and N₂O mole fractions are at levels higher than any point in the last 800,000
years (<https://gml.noaa.gov/ccgg/data/index.html>, last accessed January 2026; IPCC, 2023).

As a consequence, individual nations and governments took initiatives to monitor and regulate their emissions to aid in the
global efforts. National GHG emissions inventories are reported annually to the United Nations Framework Convention on
Climate Change (UNFCCC) and are estimated in compliance with Intergovernmental Panel on Climate Change (IPCC)
55 methodological guidance. This methodology uses activity data, emission factors and facility-level measurements, known as
bottom-up methods (Ciais et al., 2014; Gurney et al., 2016; Nisbet and Weiss, 2010). However, for some sources, the outputs
from these calculations are very uncertain; therefore, verification of inventories is recommended (Buendia et al., 2019). Some
of the most common verification methods are inverse techniques, which use atmospheric measurements complemented with
atmospheric transport modelling tools (Ganesan et al., 2014; Manning et al., 2021, 2011), known as a top-down approach. The
60 IPCC has set out the best practices for the use of top-down emissions estimates to verify inventories (IPCC, 2023). Expanding
the atmospheric measurement network would improve top-down sensitivity to emissions and reduce the associated uncertainty.
As a result, many national and regional GHG monitoring networks are currently expanding.

Within these networks, measurements can only be combined reliably when network compatibility is maintained, meaning that
biases between the same measurements across a network remain below levels that would affect their scientific interpretation.
65 These biases are set out by the WMO/Global Atmospheric Watch (GAW) programme and are referred to as “compatibility



goals” (WMO, 2022). Combining GHG measurement results from different monitoring sites to gain both a global and regional understanding of the atmospheric distribution is challenging, particularly when attempting to work towards the WMO/GAW compatibility goals. A variety of factors can introduce differences between measurements made at different monitoring sites, including sampling artefacts, local environmental interferences, and difference in calibration procedures (Hall et al., 2014).
70 Assessing network compatibility, therefore, requires comparing measurements of the same air, ideally under identical conditions, and must consider the entire sampling system and analysis chain (e.g. two laboratories measuring the same discrete air sample).

One of the most significant factors contributing to measurement differences is the calibration scale which they are reported against (Hall et al., 2014). Greenhouse gas mole fractions are typically reported relative to internationally maintained
75 calibration scales, which are realised through carefully prepared reference standards stored in compressed gas cylinders. These scales consist of well-characterised reference artefacts spanning the relevant mole-fraction range. To minimise potential biases or interferences, the composition (matrix) of these reference materials (RMs) are designed to closely match that of ambient atmospheric air; for this reason, natural (whole) air collected under background conditions is commonly used (Hall et al., 2014). The most effective manner in which to meet the WMO/GAW network compatibility goals is for all laboratories to use
80 calibration RMs that are traceable to the same primary scale. The WMO has assigned the National Oceanic and Atmospheric Administration (NOAA) as the Central Calibration Laboratory (CCL) of the GAW programme, making it responsible for the preparation, maintenance, and dissemination of network standards for CO₂, CH₄, N₂O, carbon monoxide (CO) and sulphur hexafluoride (SF₆) measurements. Several independently maintained calibration scales exist: Scripps Institution of Oceanography (SIO), USA (CO₂, N₂O, SF₆, and many other trace species), Tohoku University, Japan (CO₂, CH₄, CO, N₂O,
85 and SF₆), National Institute for Environmental Studies, Japan (CO₂, CH₄, CO, N₂O, and SF₆) and the University of Heidelberg, Germany (SF₆) (WMO, 2022). These independent scales play an important role in understanding long-term trends and in evaluating uncertainties associated with calibration systems. In addition, it remains essential to quantify the relationship among these scales to ensure that differences between them do not introduce bias into atmospheric measurements.

When laboratories adopt a single reference scale, it is important to consider both the robustness of this approach and the
90 demands it places on the supporting infrastructure. As monitoring networks grow and measurement techniques advance, there is increasing need for larger quantities of calibration standards. International System of Units (SI) traceable RMs are produced gravimetrically under controlled laboratory conditions using individual high purity target and matrix gases. The contributions to the absolute uncertainty include the mass and purity of the component gases, and a full uncertainty budget is associated with the mole-fraction assignment (Brewer et al., 2018). However, SI traceability also means that all contributors to uncertainty
95 propagate into the final measurement, which can limit the achievable uncertainty. The scale approach, in contrast, uses a high-precision instrument as a comparator to measure the composition of RMs relative to the established calibration scale, such as the WMO-NOAA scales. This comparison typically provides a smaller relative uncertainty than gravimetric assignment as measurement of all RMs are made relative to a single, well-maintained scale and so the comparison is not impacted by any uncertainty associated with preparation (Brewer et al., 2018). At present, traceability to established calibration scales represents



100 the most effective and reliable means to achieving the WMO/GAW compatibility goals across a network, with no alternative
approach yet demonstrating equivalent performance. However, the use of a scale ultimately relies on the continuous
maintenance of a central artefact that is depleted with use. This makes the use of laboratory-prepared (synthetic air) GHG RMs
increasingly valuable for meeting the growing demand, despite their higher mole fraction uncertainty.

To support the development and evaluation of SI-traceable RMs, the International Bureau of Weights and Measures (BIPM)
105 coordinate key comparison exercises. This provides a means to verify the consistency of RMs prepared by independent
National Measurement Institutes (NMIs) within the range of their SI-traceable uncertainties. Over the past decade, there has
been a decrease in the uncertainty in the reference values reported from key comparisons of several GHGs (Flores et al., 2019,
2015; Lee et al., 2017; Viallon et al., 2023). The absolute mole fraction uncertainty of the SI-traceable RMs, however, still
need to be reduced by a factor of three to four to match the relative uncertainty achievable with the scale-based approach. The
110 current levels of uncertainty have enabled several key developments. They supported the update to the WMO-CO₂-X2019
scale (Flores et al., 2019; Hall et al., 2021), identified minor biases in the WMO-NOAA CH₄ scale and initiated corrective
measures (Flores et al., 2015). They also demonstrated internal consistency between standards traceable to independent N₂O
scales (Viallon et al., 2023) and confirmed consistency between CO in air RMs while highlighting challenges with stability
(Lee et al., 2017).

115 This study presents a coordinated round-robin intercomparison designed to assess atmospheric GHG measurement consistency
within the UK and Ireland monitoring network. First, a suite of whole air and synthetic air RMs containing CO₂, CH₄ and N₂O
mole fractions that span typical atmospheric ranges were circulated sequentially between multiple atmospheric monitoring
sites between 2021 and 2022. These RMs were measured using routine instrumentation, calibration procedures and sampling
configurations in place at each site, allowing direct assessment of site-to-site and instrument-to-instrument differences under
120 field conditions. Using the whole and synthetic air RMs on the instrumentation at six sites allowed the identification of
systematic offsets and the evaluation of internal network consistency independent of assigned reference values. Second, the
measurements of the whole air RMs (traceable to the WMO-NOAA scales) and the synthetic air RMs (traceable to the SI) and
their uncertainties were compared to their assigned values and uncertainties to assess any biases. Together, these analyses
provide insight into the extent to which SI-traceable RMs can complement established calibration scales, while maintaining
125 the levels of compatibility required for atmospheric GHG monitoring networks.

2 Greenhouse gas monitoring networks

2.1 The UK Deriving Emissions linked to Climate Change (UK DECC network)

The UK DECC network became operational in 2012 (Adcock et al., 2023; Kikaj et al., 2025; Stanley et al., 2018; Stavert et
al., 2019). It currently consists of four tall tower measurement stations in England (Bilsdale (BSD; temporarily inactive),
130 Ridgehill (RGL), Tacolneston (TAC), and Heathfield (HFD)) and a background site at Mace Head (MHD) in Ireland, (see Fig.
1). There are also affiliated sites at Weybourne Atmospheric Observatory (WAO) and Jodrell Bank Observatory (JBO) in



England, as well as Invergowrie (IVG) in Scotland. The aim of the network is to provide high precision and high frequency in-situ mole fraction measurements of GHGs to improve UK emissions estimates.

2.2 Integrated Carbon Observation System (ICOS)

135 The Integrated Carbon Observation System (ICOS) is a European wide research infrastructure (<https://www.icos-ri.eu>, last
access: 15 January 2026) which provides harmonised and high-precision data on the carbon cycle and GHG emissions
(<https://www.icos-ri.eu>, last access: 15 January 2026). There are currently two ICOS sites within the UK (WAO and RGL, see
Fig. 1). Mace Head is currently in the labelling process required to become an ICOS site. The ICOS monitoring stations are
standardised, using similar instruments, sampling set-ups and calibration methods (Yver Kwok et al., 2015; Yver-Kwok et al.,
140 2021). Data are routinely quality checked by station PI's and assessed by the Atmospheric Monitoring Station Assembly
(MSA) twice a year (Yver-Kwok et al., 2021).



145 **Figure 1: The UK and Ireland atmospheric monitoring network. The network includes atmospheric monitoring sites at Bilsdale (BSD), Jodrell Bank Observatory (JBO), Weybourne Atmospheric Observatory (WAO), Tacolneston (TAC), Ridgehill (RGL), and Heathfield (HFD) in England, Invergowrie (IVG) in Scotland and Mace Head (MHD) in Ireland. Sites labelled in black took part in the round robin intercomparison. The sites labelled in grey did not take part in the round robin intercomparison. The figure is modified from Kikaj et al., 2025.**



3 Methods

150 3.1 Instrumentation

3.1.1 Cavity ring down spectrometer

The use of advanced optical instrumentation for atmospheric GHG monitoring has grown rapidly over the last decade. This is due primarily to the precision and stability of these instruments and their straightforward deployment requirements in terms of maintenance and expertise. They do not require a carrier gas, can provide continuous measurements, and have an improved
155 linearity response for N₂O and CO relative to commonly used methods based on gas chromatography (GC).

The cavity ring-down spectroscopy (CRDS) instruments (Picarro, Inc., Santa Clara, CA, USA) measure mole fractions of CO₂, CH₄ and N₂O at intervals of approximately 3 seconds (Stanley et al., 2018; Winderlich et al., 2010). The shape of the spectral lines are sensitive to temperature and pressure changes and so the instruments control the temperature and pressure of the sample.

160 The CRDS measurements are also sensitive to water vapour interference through dilution and pressure broadening effects, so the sample must either be dried prior to measurement, or these effects must be corrected for. Within the UK and Ireland atmospheric monitoring network, these corrections are derived individually for each CRDS analyser. This is done by measuring compressed air from a cylinder both dry and over a range of typical atmospheric water vapour levels (see Stavert et al., 2019, for details) alongside the use Nafion membrane dryers to reduce the water vapour levels within the sampled air.

165 3.1.2 Off-axis integrated cavity output spectroscopy

At TAC, N₂O and CO are measured using off-axis integrated cavity output spectroscopy (OA-ICOS; model EP30/913 0015-810; Los Gatos Research (LGR), USA, Lebeque et al., 2016). This technique is similar to CRDS in that it involves passing a laser through an optical cavity containing the sample and measuring absorption as a function of wavelength while sweeping the laser across the relevant spectral lines. Lebeque et al., 2016 provide details on the differences between CRDS and OA-
170 ICOS techniques. The same sensitivity to water vapour interference exists for OA-ICOS as for CRDS. At TAC an Earth Networks sample selection system, including a Nafion dryer, was used to dry the sample (see Welp et al., 2013 for details). From this point on we will abbreviate the OA-ICOS instrument name to “LGR”.

3.1.3 Gas chromatography

Within the UK and Ireland atmospheric monitoring network, GC systems are used to measure several species, using an electron
175 capture detector (ECD) for N₂O and SF₆, a flame ionisation detector (FID) for CH₄, and a pulsed discharge detector (PDD) or reduction gas analyser (RGA) for hydrogen (H₂), where the latter can also measure CO. These instruments incorporate automated sample and standard selection, as well as Nafion drying systems. The focus here is on the GC-ECD system used to measure N₂O.



In a GC-ECD system, the sample is transported by an inert carrier gas (a 5% CH₄ in Argon (Ar) mix is used in the UK and
180 Ireland atmospheric monitoring network) through chromatographic columns to separate target compounds. The eluate then
enters a detector cell containing a weak radioactive β -emitter (usually Ni-63), which produces free electrons collected as a
steady current. Analytes with a high electron affinity (e.g., halogenated or nitro compounds) capture the electrons, decreasing
the current in proportion to their abundance. Because the response in an ECD is based on electron capture kinetics and
recombination processes, certain nonidealities need to be corrected or accounted for such as nonlinearity, detector saturation
185 effects and carrier gas purity (Stanley et al., 2018).

3.1.4 Fourier transform infra-red spectroscopy

The Fourier Transform Infrared (FTIR) spectrometer (Ecotech Spectronus) at WAO is a high precision instrument capable of
simultaneously measuring mole fractions of CO₂, CH₄, N₂O, and CO. The FTIR operates by passing a broadband infrared light
source through an air sample at low pressure in the spectrometer cell, where individual gas species absorb characteristic
190 wavelengths of infrared radiation. The transmitted light is recorded as an interferogram, which is subsequently transformed
into a high-resolution spectrum using Fourier Transform techniques. Mole fractions of each gas are then derived from the
specific absorption features corresponding to their molecular signatures. A detailed description of the Spectronus FTIR can be
found in Griffith et al., 2012.

3.1.5 Calibration

195 The full calibration procedure for the UK and Ireland atmospheric monitoring network instruments, excluding WAO, can be
found in Stanley et al., 2018, however it is described briefly here. The CRDS/LGR measurements taken during the round robin
intercomparison were subjected to a multi-point calibration using a suite of three or four calibration gases that ranged
atmospheric mole fractions for CO₂, CH₄ and N₂O. A working standard is also used of approximately ambient mole fractions
to perform regular drift corrections (a drift-corrected sensitivity is then calculated, which is a function of the measured dry
200 mole fraction divided by its assigned value. This is then divided by the measured standard dry mole fraction divided by the
assigned/adjusted standard value (see Stanely et al. 2018 for details)). Connections of all working standards, and calibration
gases are made using 1/16 in. O.D., 0.03 in. I.D. 304 stainless steel tubing (Supelco, Sigma-Aldrich, UK) to minimise dead
volumes and wasted gas. Whole air is used along with high pressure aluminium cylinders to ensure the stability of CO₂ within
the gas mixtures (Luxfer Gas Cylinders, UK). These cylinders are filled either at the Mace Head atmospheric research station,
205 the University of Bristol or the University of East Anglia (UEA) Cylinder Filling Facility (CFF). Since 2021 RGL has gradually
transitioned to receiving working standards/calibration gases filled at the ICOS Flask and Calibration Laboratory (FCL). The
FCL is an entity of the Central Analytical Laboratory (CAL), located at the Max Plank Institute (MPI), Jena, and provides
whole air mixtures for stations within the ICOS network (Yver Kwok et al., 2015; Yver-Kwok et al., 2021). The FCL also
assigns mole fractions on the relevant WMO-NOAA scales for the RGL working standards/calibration gases. At non-ICOS
210 sites the WMO-NOAA traceable mole fractions for the CRDS/LGR working standards/calibration gases are assigned at the



Swiss Federal Laboratories for Materials Science and Technology (Empa; see Sect. 3.2.2 for more details). Note that RGL CRDS data is currently processed using both the UK and Ireland atmospheric monitoring network and ICOS procedures, but only the results from the UK and Ireland network processing are included in this study.

215 A full calibration procedure of GC-ECD N₂O measurements is also described by Stanley et al., 2018 but is described here briefly. Mole fractions for the ambient air samples are derived from either the peak height or peak area on the chromatogram, relative to the height/area of the peak measured in a working standard whose mole fraction is known. The working standards are prepared in 34 L electropolished stainless steel cylinders (Essex Cryogenics, Missouri, USA). They are generally filled at MHD, except for the cylinders used on the MHD GC-ECD itself which are filled at the SIO in California. The working standards used at MHD are assigned values on the SIO-16 scale in the SIO central calibration laboratory. The working
220 standards for other sites are then measured on the MHD ECD against these SIO-assigned standards (and in addition each working standard is measured against the previous working standard when it arrives on site). Hence all GC-ECD measurements of N₂O presented here are fundamentally traceable to the SIO-16 scale. For consistency with the other analysers in this study, these GC-ECD data have been converted to the WMO-N₂O-X2006A scale by subtracting a constant offset of 0.41 nmol mol⁻¹. This offset is calculated as the mean difference between measurements of the SIO-filled working standards at SIO and at
225 Empa. It is important to note that this factor is appropriate for this network and this time period and should not be considered as a universal conversion factor between the two scales (Joe Pitt, private communication, 2026). Nonlinearity corrections for the GC-ECD analysers are applied based either on measurements of scale-traceable RMs spanning the typical atmospheric range or dynamic dilution experiments (see Stanley et al., 2018, for details).

The Spectronous FTIR at WAO is calibrated using whole air standards prepared and calibrated by the ICOS FCL. These
230 calibration standards contain CO₂, CH₄, N₂O and CO traceable to WMO-NOAA calibration scales. In addition, the δ¹³C and δ¹⁸O isotopic signatures of each gas mixture are known to enable the accurate determination of CO₂ mole fraction using its three most abundant isotopologues (Griffith et al., 2012). Every 14 days a response factor is determined for the analyser for each species, including the CO₂ isotopologues, by introducing the calibration gases to the analyser. The raw data is then sent to the ICOS Atmospheric Thematic Centre (ATC) for processing. Analyser performance is routinely assessed in between
235 calibrations by introducing a target tank; a whole air cylinder with mole fractions of CO₂, CH₄, N₂O and CO₂ assigned by the ICOS FCL.

3.2 Preparation of whole air reference materials

3.2.1 Production of whole air reference materials

The whole air RMs used in this study were filled at the UEA CFF, a specialised facility for filling high-pressure gas cylinders
240 with whole air. The facility is primarily used to routinely produce RMs containing CO₂, CH₄, N₂O, CO and O₂ and supports a wide range of research projects and monitoring programmes; however, the facility can make whole air mixtures of many trace gas species, if the species is stable (e.g., ethane (C₂H₆), H₂, SF₆).



The first step in the procedure is to condition the gas cylinder prior to filling. This is achieved by evacuating the cylinders to 0.5 mbar, which removes ambient air and water vapour. The conditioned cylinder can then be filled to a maximum of 300 bar with a heavy-duty, 3-stage, oil-free ‘RIX’ compressor. During this phase, the mole fractions in the gas can be ‘tailored’ to add or remove species of interest. For whole air mixtures requiring mole fractions at ambient levels or below, the gas can be passed through various chemical scrubbers. For example, a portion of the gas can be passed through a Molsieve 13X trap to remove CO₂, or through a Zero Air Generator to remove CH₄. For cylinders requiring higher than ambient mole fractions, a precisely calibrated volume of the ‘target’ species is introduced to the conditioned gas cylinder prior to filling. During filling, the gas is passed through a coalescent filter, a 3A Molsieve trap and a magnesium perchlorate trap to dry the air to reduce the water content to less than 1 μmol mol⁻¹. This is essential to ensure the long-term stability of the mole fractions of the gases within the ambient air sample in the cylinder. The dew point of the cylinders is routinely measured with an Accupoint moisture meter. In addition to filling, the CFF is equipped with state-of-the-art equipment to analyse the mole fractions of CO₂, CH₄, N₂O and CO in gas mixtures, traceable to WMO-NOAA calibration scales.

3.2.2 Value assignment of whole air reference materials

Three 20 L Luxfer cylinders were filled, scrubbed/spiked and their water contents were checked at the UEA CFF following the methods in Sect. 3.2.1. The cylinders were then measured at Empa, who maintain strong links to the WMO-NOAA scales through their role as a WMO World Calibration Centre (WCC) for CO₂ and CH₄ (in addition to surface ozone and CO). World Calibration Centres are tasked with ensuring data quality and the traceability of network observations across GAW stations through regular intercomparison exercises (Buchmann et al., 2009; Zellweger et al., 2019, 2016). Filled cylinders are analysed at Empa over a 15-minute period where only the last three minutes of measurements are used for CO₂ and CH₄ and the last ten minutes for N₂O to ensure measurement stability. The value assignments on the WMO-NOAA scales are derived from multiple measurement sequences which include NOAA reference and working standards. The uncertainties are derived from the square root of the variance of all known sources of uncertainty excluding the WMO-NOAA scale uncertainties themselves, however typically the distribution of the variance is not well known so standard uncertainty is multiplied by a coverage factor. This coverage factor provides expanded uncertainties that encompass confidence limits which likely cover the ‘true’ value. For a normal distribution, a coverage factor $k = 2$ yields confidence limits of about 95%. The combined uncertainties take into account uncertainties arising the propagation of the WMO-NOAA scale onto the calibrated RMs and the uncertainties of the analytical instrumentation. Table 1 lists the assigned values of the whole air RMs.

270



275 **Table 1. The mole fraction of CH₄, CO₂, and N₂O assigned to each of three whole air RMs (given to 1 decimal place), traceable to the WMO-NOAA scales (WMO-CO₂-X2019 for CO₂, WMO-CH₄-X2004A for CH₄, and WMO-N₂O-X2006A for N₂O). The reported expanded uncertainty of the measurement is stated as the standard uncertainty of the measurement multiplied by the coverage factor $k = 2$, which for a normal distribution corresponds to a coverage probability of approximately 95%.**

Cylinder ID	CO ₂ ($\mu\text{mol mol}^{-1}$)	CH ₄ (nmol mol^{-1})	N ₂ O (nmol mol^{-1})
D664636	398.10 \pm 0.08	1838.08 \pm 0.86	324.44 \pm 0.18
D664637	439.62 \pm 0.08	2210.07 \pm 0.86	336.18 \pm 0.18
D664638	502.73 \pm 0.08	2425.12 \pm 0.86	345.22 \pm 0.18

280 3.3 Preparation of synthetic air reference materials

The term "synthetic air" refers to a mixture of gases that is artificially created to mimic the composition of natural, ambient air. The matrix gas in this study contains N₂, O₂ and Ar.

All synthetic air RMs were prepared at the National Physical Laboratory (NPL) by gravimetry, in accordance with ISO 6142-1:2015 (ISO, 6142-1) in 10 L Luxfer aluminium cylinders. This was done by addition of CO₂, CH₄, and N₂O via intermediate mole fraction mixtures in a N₂ matrix. The cylinders were treated internally with Spectraseal (BOC) proprietary passivation processes to inhibit adsorption of target components to the internal cylinder surfaces. Cylinders were evacuated overnight using an oil-free pump (Scrollvac SC15D, Leybold Vacuum) and turbo molecular pump with magnetic bearing (Turbovac 340M, Leybold Vacuum) to a pressure of $< 3 \times 10^{-7}$ mbar.

To prepare the intermediate mole fraction mixtures, each component was added via either a direct addition to the cylinder, through purged 1/16 in. tubing (Swagelok, electropolished stainless steel) or through using a transfer vessel and then diluting with synthetic air by the direct addition of Ar (BIP, Air Products), O₂ (N6.0, BOC) and N₂ (BIP+, Air Products) to blend to nominally atmospheric ratios. Argon was added to the cylinder from a 30 cmol mol^{-1} Ar-in-N₂ intermediate mixture for more precise targeting. The transfer vessel was weighed against a tare vessel matched for size and shape before and after addition into the evacuated cylinder (Mettler-Toledo XP2004S, ± 0.3 mg weighing uncertainty). The mass added via direct addition was determined by weighing the cylinder before and after addition against a tare cylinder on an automatic weighing facility, developed by the Korean Research Institute for Standards and Science (KRISS) (Mettler-Toledo XP26003L, ± 20 mg weighing uncertainty) (Lim et al., 2017). Table 2 lists the certified values of the gases within the synthetic air RMs.

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Table 2. The gravimetrically assigned mole fractions of CH₄, CO₂ and N₂O for each of the four synthetic air RMs produced at NPL (given to 1 decimal place), traceable to the mole. The reported expanded uncertainties are based on standard uncertainties multiplied by a coverage factor $k = 2$, providing a coverage probability of approximately 95%.

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Cylinder ID	CO ₂ ($\mu\text{mol mol}^{-1}$)	CH ₄ (nmol mol^{-1})	N ₂ O (nmol mol^{-1})
NPL-2892	384.0 \pm 0.8	1849.8 \pm 3.7	322.5 \pm 3.2
NPL-3097	409.5 \pm 0.8	1987.8 \pm 4.0	336.6 \pm 3.4
NPL-3064	445.4 \pm 0.9	2256.8 \pm 5.0	341.0 \pm 3.4
NPL-2997	480.6 \pm 1.0	2509.1 \pm 5.0	348.2 \pm 3.5

3.4 Measurement procedure

The CRDS instruments used in the UK and Ireland atmospheric monitoring network measure mole fractions of CO₂, CH₄ and N₂O at approximately 3 second intervals, while the LGR measurements at TAC are made at 1 Hz. Both the CRDS and LGR data were averaged into 1 min periods for analysis. The GC-ECDs used within the network collected data at 10 min intervals, except for MHD where a sample was measured every 20 min. For these GC-ECD measurements a whole air working standard was sampled at least every 60 mins during the intercomparison to correct for potential drift of the instrument (every 40 min at MHD). The FTIR at WAO produces a data point every minute.

Both the whole air and synthetic air RMs were measured at BSD, RGL, TAC, WAO, HFD and MHD using a procedure designed to ensure that the instrument response had stabilised. All measurements were conducted for sufficient durations to allow the sample stream to reach equilibrium within the analysers and only data collected after equilibration were retained for analysis and comparison purposes. During the periods the intercomparison cylinders were measured, for both whole air and synthetic air, the variability is taken to be the standard deviation of the measurements over the period they were measured.

4 Results and discussion

4.1 Site-to-site measurement offsets observed using whole air reference materials

This section examines the differences between measurements of the whole air RMs made at each of the atmospheric monitoring sites, where each site may use a different type of instrument. These results provide an assessment of internal network consistency and allow identification of offsets associated with individual sites or instruments when using whole air.

Figure 2 shows the mole fraction of three whole air RMs measured by the various instruments at each of the atmospheric monitoring sites that took part in the intercomparison exercise. The results are shown in ascending mole fraction order of CO₂ (left), CH₄ (middle) and N₂O (right).

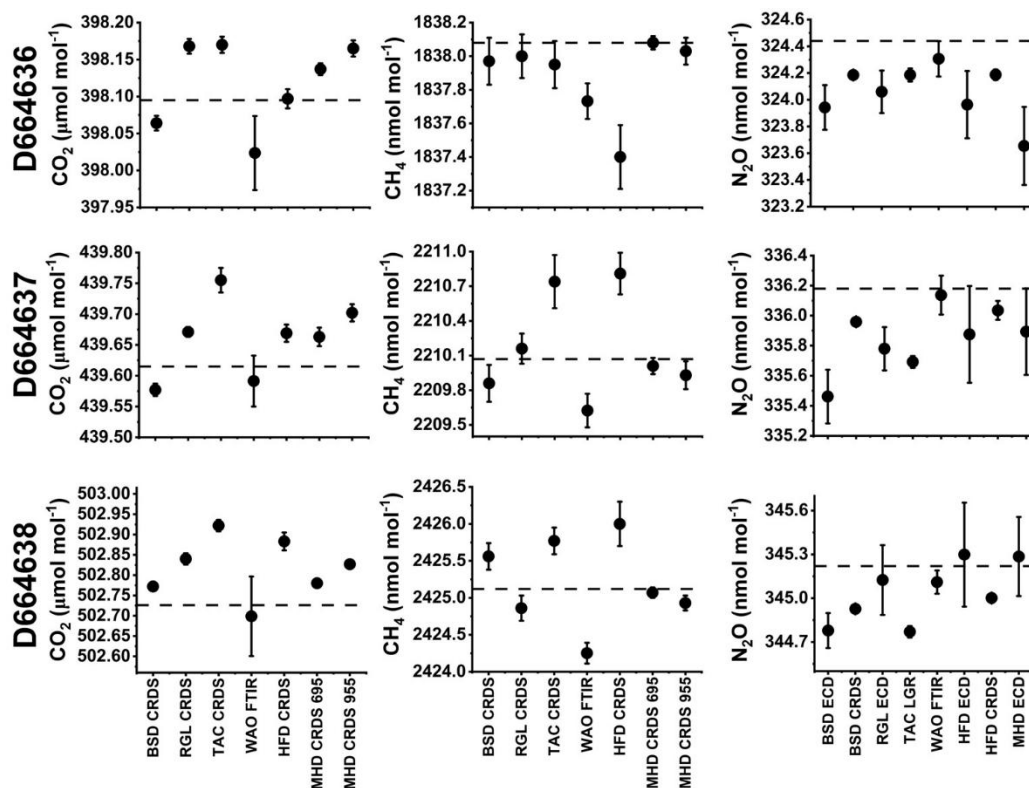


Figure 2. Measurement results of the three whole air RMs at each site and instrument that took part in the intercomparison exercise (Bilsdale (BSD), Ridgehill (RGL), Tacolneston (TAC), Weybourne (WAO), Heathfield (HFD) and Mace Head (MHD)). The MHD site had two CRDS instruments of the same make and model (Picarro Inc. G2401) for CO₂ and CH₄ measurements owned by the University of Galway and Laboratory for Climate and Environmental Sciences (LSCE) in France. They were given identification numbers (695 and 955 respectively) to distinguish between them. The results are the mean of the measurements collected from each instrument at each site. They are obtained using optical instrumentation (CRDS, LGR and FTIR) and GC-ECDs (denoted “ECD” on the x-axis) and are shown in ascending mole fraction order of CO₂ (left), CH₄ (middle) and N₂O (right). The top panel shows results for D664636 (assigned values of 398.10 μmol mol⁻¹, 1838.08 nmol mol⁻¹ and 324.44 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively), middle panel showing results for D664637 (assigned values of 439.62 μmol mol⁻¹, 2210.07 nmol mol⁻¹ and 336.18 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively) and the bottom panel showing results for D664638 (assigned values of 502.73 μmol mol⁻¹, 2425.12 nmol mol⁻¹ and 345.22 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively). Dashed lines denote the assigned mole fraction of CO₂, CH₄ and N₂O, traceable to the respective WMO-NOAA scales. The error bar represents the standard deviation of the measurements over the time period they were measured. Where the error bars cannot be seen, they are smaller than the plotted points.

Across the three whole air RMs, site-to-site agreement varied by species, showing instrumental differences, the inherent measurement precision of each species and potential scale offsets. To quantify the spread of measurements between sites, the coefficient of variation (standard deviation of measurement divided by mean measurement result) for each RM for each gas is shown in Table 3 and the standard deviation of each RM across all sites and instruments in Table 4. For CO₂, the average



345 standard deviation across all sites is $0.07 \mu\text{mol mol}^{-1}$ (i.e. the mean of the standard deviations for the three cylinders listed in Table 4), with a maximum offset from the assigned cylinder values of $0.20 \mu\text{mol mol}^{-1}$ (TAC CRDS, D664638). For CH_4 , the average standard deviation is $0.66 \text{ nmol mol}^{-1}$ with a maximum offset of $0.88 \text{ nmol mol}^{-1}$ (HFD CRDS, D664638) and for N_2O the average standard deviation is $0.21 \text{ nmol mol}^{-1}$ with a maximum offset of $-0.79 \text{ nmol mol}^{-1}$ (MHD ECD, D664636). These results show that for CH_4 the WMO compatibility goal of 2 nmol mol^{-1} is met in all cases and the spread across the network is also within the compatibility goal, while for CO_2 some results lie outside the (northern hemisphere) compatibility goal of $0.1 \mu\text{mol mol}^{-1}$, but all are within the extended compatibility goal of $0.2 \mu\text{mol mol}^{-1}$, with the spread of results across the network being within the northern hemispheric compatibility goal. For N_2O the spread across the network is significantly larger than the compatibility goal of $0.1 \text{ nmol mol}^{-1}$ but lies within the extended goal of $0.3 \text{ nmol mol}^{-1}$ and shows that, some results lie outside both the compatibility and extended compatibility goal. Additionally, compared to the CO_2 and CH_4 measurements the coefficient of variation for N_2O is an order of magnitude larger. This reflects the challenges posed in reaching these goals for N_2O given the limitations imposed by the performance of currently available instrumentation, which impacts both the uncertainty associated with site measurements and the uncertainty associated with scale propagation.

360 **Table 3. Coefficient of variation for each RM that took part in the intercomparison exercise. The values are calculated using the standard deviation of the measurements over the measurement period (taken to be the uncertainty of the result) divided by the average measurement result over the measurement period.**

Cylinder ID	CO_2	CH_4	N_2O
D664636	6.25×10^{-5}	6.92×10^{-5}	5.62×10^{-4}
D664637	5.26×10^{-5}	8.10×10^{-5}	5.78×10^{-4}
D664638	8.15×10^{-5}	7.09×10^{-5}	5.08×10^{-4}
NPL-2892	5.97×10^{-5}	6.45×10^{-5}	4.13×10^{-4}
NPL-3097	5.71×10^{-5}	8.44×10^{-5}	4.54×10^{-4}
NPL-3064	3.02×10^{-5}	6.00×10^{-5}	3.59×10^{-4}
NPL-2997	3.30×10^{-5}	6.86×10^{-5}	5.36×10^{-4}

Table 4. The standard deviation of each RM across all sites and instruments that took part in the intercomparison exercise.

Cylinder ID	$\text{CO}_2 (\mu\text{mol mol}^{-1})$	$\text{CH}_4 (\text{nmol mol}^{-1})$	$\text{N}_2\text{O} (\text{nmol mol}^{-1})$
D664636	0.06	0.38	0.21
D664637	0.06	0.84	0.21
D664638	0.07	0.74	0.21
NPL-2892	0.12	0.67	0.28
NPL-3097	0.11	1.14	0.21



NPL-3064	0.15	0.40	0.13
NPL-2997	0.10	0.69	0.25

4.2 Assessing results from the whole air reference materials and implications for traceability

365 In this section, the mean of the measurement results for the whole air RMs from all instruments and sites presented in Sect. 4.1 are compared with the assigned values of each whole air RM (D664636, D664637 and D664638). This comparison is used to quantify the average network offset from the assigned (scale-traceable) values. The assigned values of each RM are traceable to WMO-CO₂-X2019 for CO₂, WMO-CH₄-X2004A for CH₄, and WMO-N₂O-X2006A for N₂O and each assigned value has an associated uncertainty. Below shows step-by-step calculations on how the measurement values from each site were brought together, along with their uncertainties and compared to their scale-traceable assigned values and associated uncertainties. Weighted mean mole fractions were calculated for each RM using the measurements at each site and from each instrument (Montgomery and Runger, 2014):

$$370 \text{ Weighted mean} = \frac{\sum(x_i w_i)}{\sum w_i} \quad (1)$$

Where x_i represents the measured value from an instrument and w_i represents its weight. The weight is expressed as (Montgomery and Runger, 2014):

$$375 w_i = \frac{1}{(u_{instrument})^2} \quad (2)$$

Where $u_{instrument}$ is the measurement variability from a particular instrument from a particular site, taken to be the standard deviation of the measurements over the measurement period. The standard error of the weighted mean is expressed as (Montgomery and Runger, 2014):

$$380 u(\text{weighted mean}) = \sqrt{\frac{1}{\sum w_i}} \quad (3)$$

Using the weighted mean ensures that measurements with smaller variabilities contribute more strongly to the RM average, providing a better estimate of the estimate of each site's performance (as the poor estimates are down-weighted). The residual is calculated by subtracting the assigned value from this weighted mean:

$$\text{residual} = \text{weighted mean} - \text{assigned value} \quad (4)$$

385 The assigned value uncertainty is given including the coverage factor of $k=2$ to ensure the "true value" is encompassed at a confidence level of 95%. Therefore, the $k=2$ value of the uncertainty of the weighted mean is also used to calculate the combined uncertainty. Then, the combined uncertainty on the residual is:

$$U(x) = \sqrt{2 \times (u(\text{weighted mean}))^2 + u(\text{RM assigned value})^2} \quad (5)$$

390 Where $u(\text{RM assigned value})$ is the uncertainty on the assigned value of the RM (which for the whole air RMs does not include the scale uncertainty itself, it is composed of the propagation uncertainty and the uncertainty of the analytical instrumentation used to assign the values, see Tables 1 and 2). The $u(\text{weighted mean})$ values are given in Table 5.



Table 5. The standard error of the weighted mean from all instruments at all sites for each RM that took part in the intercomparison exercise. This is calculated using Eq. (3).

Cylinder ID	CO ₂ (μmol mol ⁻¹)	CH ₄ (nmol mol ⁻¹)	N ₂ O (nmol mol ⁻¹)
D664636	0.004	0.031	0.017
D664637	0.005	0.046	0.017
D664638	0.005	0.047	0.017
NPL-2892	0.005	0.023	0.014
NPL-3097	0.005	0.044	0.016
NPL-3064	0.006	0.063	0.018
NPL-2997	0.005	0.044	0.020

395 In Fig. 3, residuals for the whole air RMs (traceable to WMO-NOAA scales) are plotted on the left-hand side (LHS), titled
 “scale traceability”. The complete uncertainty on the assigned mole fraction should include a contribution from the calibration
 scale itself, which is the largest contributor to the total uncertainty. However, as this scale uncertainty is common to all whole
 air RMs (depending on the species in question this is the uncertainty in either the WMO-CO₂-X2019, WMO-CH₄-X2004A
 or WMO-N₂O-X2006A scale) it represents a systematic component and is correlated across all the whole air RMs. Due to
 400 this, the calibration scale uncertainty was removed from the error bars on the LHS of Fig. 3, therefore the error bars represent
 the uncertainty from the measurements, the uncertainties of the analytical instrumentation used to assign values and the scale
 propagation uncertainty. The solid black line represents the WMO network compatibility goals (0.1 μmol mol⁻¹ for CO₂, 2
 nmol mol⁻¹ for CH₄ and 0.1 nmol mol⁻¹ for N₂O) and the dashed lines represent the WMO extended network compatibility
 goal for each gas (0.2 μmol mol⁻¹ for CO₂, 5 nmol mol⁻¹ for CH₄ and 0.3 nmol mol⁻¹ for N₂O). The dotted pink line denotes
 405 zero residual. The WMO/GAW network compatibility goal is intended for studies that include well-mixed background air,
 where the smallest possible bias between datasets is needed, and the WMO/GAW extended network compatibility goal was
 established for studies where larger biases are acceptable (e.g. in more heavily polluted air masses). This includes
 measurements in urban areas, where anthropogenic emissions are more frequent.

The values assigned to the whole air RMs arise from using the scale method (see Sect. 3.2.2). As mentioned in Sect. 1, the
 410 scale is a set of primary standards with agreed stable reference values. This method allows for unrivalled precision as the
 measurements are compared to a single point of reference so that the scale accuracy has a limited impact on the compatibility
 of the measurements (Brewer et al., 2018). The scale approach is adopted within the GHG monitoring community as it can
 achieve precisions (defined in this context as how measurements change relative to a single set of primaries) that are more
 suited to atmospheric monitoring requirements than the uncertainties associated with SI traceable, gravimetrically prepared
 415 mixtures.



The LHS of Fig. 3 evaluates the agreement between the network-average measurements and the assigned values of the whole air RMs, providing an indication of any systematic network bias relative to the calibration scale (i.e. the WMO-NOAA scales). This assessment is distinct from the assessment of network consistency presented in Fig. 2, which is determined solely from the spread of measurements between instruments and sites. Figure 3 (LHS) examines whether the network, as a whole, is offset
420 from the scale to which the RMs are traceable. For CO₂ and CH₄, the weighted mean residuals are close to zero, indicating no significant network-wide bias relative to the scale. For CO₂, the error bars are slightly larger than the compatibility goal but lie well within the extended compatibility goal, whilst for CH₄, the error bars lie well within the compatibility goal. The N₂O results show residuals that are close to the assigned values but are noticeably larger compared with the CO₂ and CH₄ residuals and are accompanied by smaller relative uncertainties to the N₂O compatibility goals. Achieving the WMO/GAW extended
425 network compatibility goal for N₂O is considerably more challenging than CO₂ or CH₄. Trends in atmospheric N₂O vary only weakly in space and time, so even sub-ppb calibration or instrumental biases can obscure atmospheric signals.

4.3 Site-to-site measurement offsets observed using synthetic air reference materials

This section examines the differences between measurements of the synthetic air RMs made at each of the atmospheric monitoring sites. These results provide an assessment of internal network consistency and allow identification of offsets
430 associated with individual sites or instruments when using synthetic air.

Figure 4 shows the results from the measurements of four synthetic air RMs at each of the atmospheric monitoring sites for the instruments that took part in the intercomparison exercise. The top panel (NPL-2892) shows the results for the lowest mole fraction RM measured for CO₂ (left), CH₄ (middle) and N₂O (right). The second and third panels (NPL-3097 and NPL-3064) show results for the two intermediate mole fraction RMs used and the bottom panel (NPL-2997) for the highest mole fraction
435 RM used.

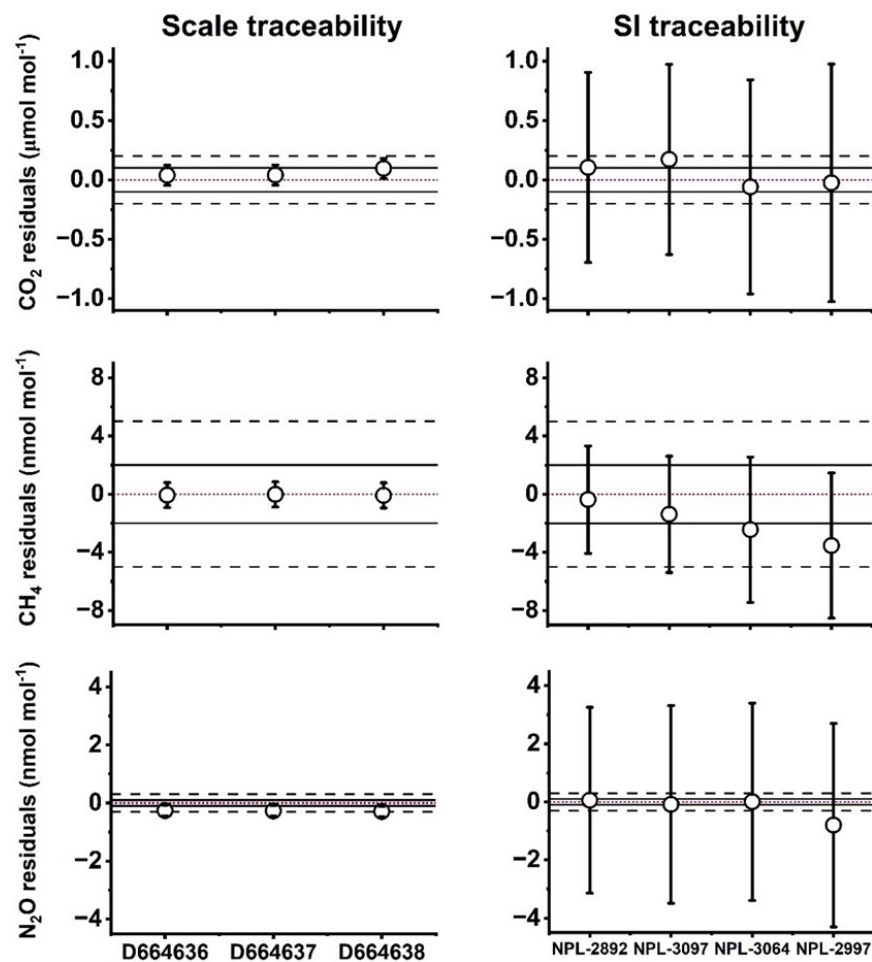


Figure 3. The weighted mean measurement of each RM from each site and instrument subtracted from the assigned mole fraction (residual). Top panels show the results for CO₂; the middle panel shows results for CH₄, and bottom panels show results for N₂O. Left: results from the three whole air RMs. The y-axis shows the residuals, the solid black lines show the range of the WMO compatibility goals (0.1 μmol mol⁻¹ for CO₂, 2 nmol mol⁻¹ for CH₄ and 0.1 nmol mol⁻¹ for N₂O), the dashed lines show the range of the extended WMO compatibility goals (0.2 μmol mol⁻¹ for CO₂, 5 nmol mol⁻¹ for CH₄ and 0.3 nmol mol⁻¹ for N₂O) and the x-axis shows the RM IDs. As this scale uncertainty is common to all cylinders (depending on the species in question this is the uncertainty in either the WMO-CO₂-X2019, WMO-CH₄-X2004A or WMO-N₂O-X2006A scale) it represents a systematic component and is correlated across all the whole air RMs. Due to this, the calibration scale uncertainty is not included in the error bars, therefore the error bars represent the uncertainty from the measurements, the uncertainties of the analytical instrumentation used to assign the values and scale propagation uncertainty. Right: results from the four synthetic air RMs, where the solid and dashed lines are as in the LHS of the plot. The error bars represent the combined measurement and certified SI-traceable uncertainty for each RM. The dotted pink line denotes zero residuals. For the N₂O residuals on the LHS (scale traceability), the error bars are smaller than the plotted points.

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Across the four synthetic air RMs, the degree of site-to-site agreement varies by species, typically showing slightly larger differences from the assigned value than observed for whole air. To quantify the variability of measurements across sites, the coefficient of variation (standard deviation of measurement divided by mean measurement result) for each RM for each gas is shown in Table 3. For CO₂, we observe a large offset for the WAO FTIR. The potential contribution of matrix effects to this offset is discussed in Sect. 4.4.1. Excluding WAO FTIR CO₂, the average standard deviation across sites is 0.07 μmol mol⁻¹ for CO₂, 0.73 nmol mol⁻¹ for CH₄ and 0.22 nmol mol⁻¹ for N₂O. The maximum offset from the RM certified values are 0.21 μmol mol⁻¹ for CO₂ (RGL CRDS, NPL-3097), -3.51 nmol mol⁻¹ for CH₄ (MHD CRDS 955, NPL-2997) and -1.09 nmol mol⁻¹ for N₂O (HFD ECD). The average standard deviation (spread) of values are slightly larger compared to the equivalent values calculated using the whole air RMs, indicating that, with the notable exception of the WAO FTIR CO₂ measurement, the synthetic air matrix is not significantly impacting the intra-network compatibility for other species and analysers. The maximum offset (excluding the WAO FTIR) for CO₂ is only very slightly larger than that for the whole air, however for CH₄ there appears to be a systematic dependence of the offsets on mole fraction, this is discussed further in Sect. 4.4. For N₂O, the maximum offset is significantly larger than that for whole air although, all measurement results are within the uncertainty of the assigned values of the cylinders. Note, that for NPL-3064, it was not measured on either of the MHD CRDS instruments due to issues with regulator stem.

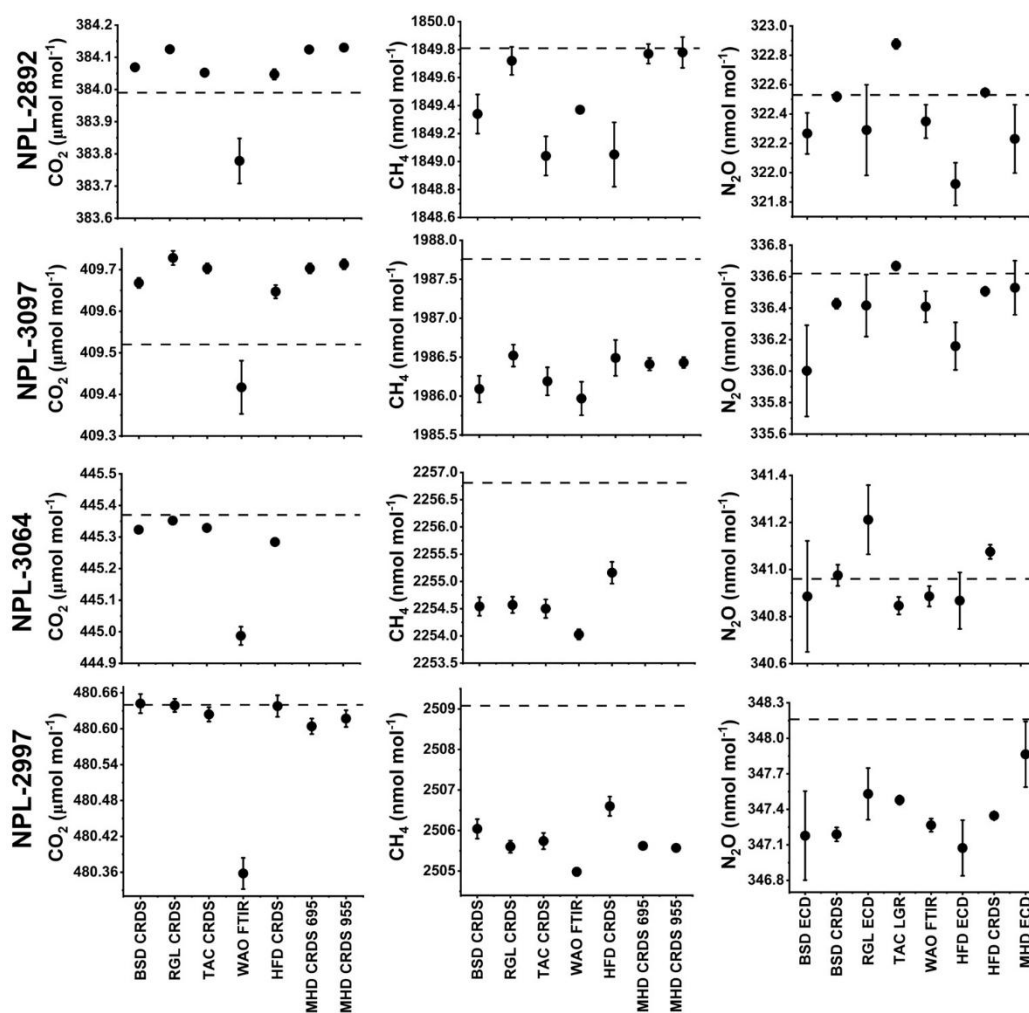
4.4 Assessment of results from the synthetic air reference materials and implications for traceability

The synthetic air RMs were prepared gravimetrically, where a thorough understanding of the method can be achieved i.e. it takes into account the uncertainty arising from the preparation, loss of material to the cylinder walls, and validation (Brewer et al., 2018). The values and associated uncertainty assigned to RMs prepared this way allows for traceability of mole fractions to the SI with an uncertainty that is validated through key comparison exercises (Flores et al., 2019, 2015; Viallon et al., 2023). The measurements of the synthetic air RMs were made following the standard calibration procedures at each site, i.e. instruments are calibrated using whole air standards, which are themselves traceable to WMO-NOAA scales. As a result, the measured values obtained for the synthetic air RMs are traceable to the scale-defined values rather than an independent, SI-traceable realisation. In other words, the calibration chain effectively links the synthetic air measurements back to the WMO-NOAA scale through the whole air standards used for instrument calibration. Consequently, while the measurements generally demonstrate good internal agreement amongst the RMs, they do not provide a direct assessment of the absolute accuracy of the synthetic air standards themselves, but rather of their alignment to the existing whole air calibration scale.

Figure 4 shows an offset between the SI-assigned value of the CH₄ RMs and the measured value. This offset lies within the uncertainty of the SI-assigned values (see Table 2), suggesting that the two traceability chains, the SI realisation and the WMO-NOAA scale, for these RMs are statistically consistent within their stated uncertainties. However, the monotonic increase in the magnitude of the offset with increasing mole fraction (see CH₄ on the right-hand side (RHS) of Fig. 3) could indicate the presence of a systematic difference between the two approaches. Comparison against a larger suite of SI-traceable RMs would



485 be required to begin to confirm the existence of this effect. For CO₂ and N₂O, Fig. 3 (RHS) shows the uncertainties on the SI-
 traceable RMs are significantly larger compared to the compatibility goals, demonstrating that these SI-traceable RMs would
 not be suitable for use as calibration gases or working standards within an atmospheric monitoring network. It is interesting to
 note that the actual agreement between the network-average weighted means and the SI-traceable assigned values is much
 better than would be expected based on these uncertainties and the mean coefficient of variability for the synthetic RMs is
 smaller compared to the whole air RMs. Again, further measurement of a larger suite of SI-traceable RMs is needed to begin
 490 to establish whether this is reflective of overestimation of the SI-traceable uncertainties.



495 **Figure 4.** Measurement results of the four synthetic air RMs at each site and instrument that took part in the intercomparison exercise (Bilsdale (BSD), Ridgehill (RGL), Tacolnестon (TAC), Weybourne (WAO), Heathfield (HFD) and Mace Head (MHD)). The MHD site had two CRDS instruments of the same make and model (Picarro Inc. G2401) for CO₂ and CH₄ measurements owned by the University of Galway and Laboratory for Climate and Environmental Sciences (LSCE) in France. They were given identification numbers (695 and 955 respectively) to distinguish between them. The results are the mean of the measurements collected from each



instrument at each site. The results are obtained using optical instrumentation (CRDS, LGR and FTIR) and GC-ECDs (denoted “ECD” on the x-axis) and are shown in ascending mole fraction order of CO₂ (left), CH₄ (middle) and N₂O (right) in each whole air RM. The top panel showing results for NPL-2892 (assigned values of 383.99 μmol mol⁻¹, 1849.81 nmol mol⁻¹ and 322.53 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively), second from top panel showing results for NPL-3097 (assigned values of 409.52 μmol mol⁻¹, 1987.76 nmol mol⁻¹ and 336.62 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively), second from bottom panel showing results for NPL-3064 (assigned values of 445.37 μmol mol⁻¹, 2256.81 nmol mol⁻¹ and 340.96 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively) and bottom panel showing results for NPL-2997 (assigned values of 480.64 μmol mol⁻¹, 2509.08 nmol mol⁻¹ and 348.16 nmol mol⁻¹ for CO₂, CH₄ and N₂O respectively). Dashed lines denote the assigned values of CO₂, CH₄ and N₂O. The error bars represent the standard deviation of the measurements over the time period they were measured. Where the error bars cannot be seen, they are smaller than the plotted points.

4.4.1 Unquantified matrix effects

In this study, the compositions of the major components of the synthetic air RMs were chosen to closely match atmospheric levels, thereby minimising the potential for matrix-induced effects. There is a large potential for offsets between assigned and measured values due to matrix effects that arise from differences in the proportions of the major constituents that make up the synthetic air for synthetic RMs. The absence of a particular inert constituent (e.g., Ar) can lead to pressure broadening for spectroscopic analysers as the analyte of interest interacts with surrounding molecules, thus broadening the spectral line (Griffith, 1996, 1982; Griffith et al., 1982). The strength of this effect will depend on the measurement technique and how the spectrum is being analysed. For example, an analyser that uses peak height (e.g., CRDS) will not be affected in the same way as an analyser that fits the whole peak shape (e.g., FTIR). For this reason, it has long been the practice and recommendation of atmospheric measurement networks to prepare calibration material using whole air to best match the makeup composition between the calibration gas/working standard and the sample air being measured (WMO, 1982). It seems evident from Fig. 2 and 3 that assessing the network using synthetic standards can lead to increased variability between sites for certain species and that it is likely that matrix effects contribute to this. The matrix effects appear to be most apparent for CO₂, where the CRDS instruments are in close agreement with each other, but are significantly different from the FTIR. Therefore, it is essential that any matrix effects on the network are fully understood and that the potential for these is minimised during the production of the synthetic air RMs; this could include adding other trace species (e.g., helium or neon) to better match the composition of whole air.

4.4.2 Key comparisons

Key comparisons, conducted under the framework of the CIPM Mutual Recognition Arrangement (MRA), are designed to assess the equivalence of measurement standards maintained by NMIs. Their purpose is to demonstrate global comparability of RMs and calibration capabilities, ensuring that measurements made across different countries remain consistent and traceable to the SI within stated uncertainties. The results establish the best-case performance achievable by the RMs, independent of field-specific influences such as instrument drift, site-to-site and instrument-to-instrument variability, or



530 calibration-chain differences. In contrast, the results from this intercomparison exercise reflect the applied performance of those RMs when deployed to atmospheric monitoring sites. Evaluating how applied performance of the RMs compare to the best-case performance allows the assessment of how uncertainties arising from the RMs translate into uncertainties observed under operational conditions.

The three comparisons, CCQM-K120 for CO₂ (Flores et al., 2019), CCQM-K82 for CH₄ (Flores et al., 2015), and CCQM-K68
535 for N₂O (Viallon et al., 2023), all show good consistency between NMIs, but the degree of agreement varies depending on the gas. For CO₂ (K120), the internal consistency between standards submitted by each participant was relatively tight but only marginally met the extended WMO compatibility goal. This suggests that for CO₂, where atmospheric measurement networks require very high precision, maintaining a common calibration scale such as the WMO-NOAA scale remains essential to ensure inter-comparability across laboratories and long-term stability. In contrast, the results for CH₄ (K82) showed somewhat
540 larger variability between participants compared to K120, yet all measurements still fell within both WMO network compatibility goals. This indicates that, while a scale approach provides improved precision, the current level of agreement between independent realisations is sufficient for CH₄ monitoring needs as they meet the WMO/GAW goals. For N₂O (K68), the degree of internal consistency between standards was notably high, with differences well below the spread observed across all participants. This strong internal agreement supports the effectiveness of a scale-based approach. Such an approach is
545 particularly advantageous for N₂O, where the required WMO/GAW compatibility goals are the most stringent and the atmospheric variability small.

Figure 3 (RHS) shows that for CO₂ in synthetic air the WMO extended network compatibility goals are not met due to the associated uncertainties being too large (where the main contributing uncertainty is from the SI-traceable assigned value, as is the case with CH₄ and N₂O, see Tables 2 and 4 for measurement value and SI-traceable value uncertainties). For CH₄, Fig. 3
550 (RHS) shows that the uncertainties are on the same order of magnitude as the extended WMO/GAW compatibility goal, which aligns with K82 in that it is the gas with SI-traceable uncertainties closest to meeting WMO/GAW compatibility goals. These results are again consistent with the key comparison results, in that we are unable to achieve WMO/GAW extended network compatibility goals, and a scale-based approach would be able to provide higher levels of precision for monitoring requirements. While the key comparison results show agreement between NMIs on the same order of magnitude as the
555 WMO/GAW network compatibility goals, it is important to recognise that these results represent a best-case performance. To ensure negligible offsets and maintain comparability at the level achieved by a scale-based approach, measurement uncertainties in key comparisons would need to be reduced by a further factor of three to four to account for the additional sources of variability encountered in field conditions.

5 Conclusions

560 This intercomparison, for the first time, used both scale-traceable whole air and SI-traceable synthetic air RMs to provide insight into the performance and internal consistency of GHG measurements within the UK and Ireland atmospheric



monitoring network. By evaluating measurements of CO₂, CH₄ and N₂O across multiple sites and instruments, the study offers an assessment of site-to-site and instrument-to-instrument variability under field conditions.

565 Whole air RMs, which are traceable to the WMO-NOAA calibration scales, showed minimal network-average biases for CO₂ and CH₄ that are within or very close to the WMO/GAW compatibility goals. This reflects the stability and linearity of the measurement techniques for these gases, the stability of whole air mixtures, and the robustness of scale-traceable calibration procedures. However, it is important to note that instrument-specific effects, including drift, non-linearity, and differences in calibration implementation, also contribute to differences between sites and should be considered when interpreting network compatibility. In contrast, N₂O shows network-average biases that were outside the compatibility goals and also exhibited
570 larger spread across sites relative to these goals. This is consistent with known challenges associated with the relatively poorer measurement precision for this species.

For the synthetic RMs, CH₄ was shown to have uncertainties on the SI-traceable assigned values that are comparable to the extended compatibility goals, but these uncertainties for CO₂ and N₂O remain considerably larger compared to the compatibility goals (and extended compatibility goals). Despite this, the spread of the measurements across the sites (the
575 standard deviation) of the synthetic RMs are in closer agreement to that of the whole RMs than what would be expected from the SI-traceable uncertainties. Taken together, the results show that whole air RMs remain essential for high-precision calibration and achieving WMO/GAW compatibility goals. Continued reductions in synthetic RM uncertainties, along with improved characterisation of matrix effects, mole fraction dependant instrument responses, and long-term stability, are critical in order for synthetic, SI-traceable air to become suitable for high-precision atmospheric monitoring.

580 **Code, data, or code and data availability**

Data used in this paper are available in the supplementary information.

Author contributions

ES led the research, performed all data visualizations, and developed the concepts and results. Round robin measurements at HFD were conducted by CR, ES and DK, at WAO by GF, at BSD, RGL and TAC by JP, SOD, DY, KC and KS and at MHD
585 by GS. AM and NG ran the UEA CFF and filled the whole air reference materials respectively. DM, ML and MR provided the instrumentation used at the MHD site. ES wrote the original draft of the paper, with JP, GF, CR, DK and KS contributing to the editing and review process.



Competing interests

The authors also have no other competing interests to declare.

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