



1 Global Ionospheric Response to the 17 March 2015 Geomagnetic Storm: A Multi-Station and  
2 Multi-Parameter Study

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7 Abstract

8 The global ionospheric response to the St. Patrick's Day Storm during 15–20 March 2015 is  
9 investigated using a comprehensive set of ground-based and satellite observations. Ionospheric  
10 parameters (foF2, hmF2, and h'F) from ionosonde stations distributed across different longitude  
11 sectors are analyzed in conjunction with geomagnetic field variations (H-component),  
12 thermospheric composition (O/N<sub>2</sub>), and geomagnetic indices (Dst and Kp) to examine the  
13 coupled magnetosphere–ionosphere–thermosphere system.

14 During the main phase of the storm (17 March), a pronounced uplift of the F-region is observed  
15 at equatorial and low-latitude stations in both western and eastern longitude sectors, indicating  
16 the presence of enhanced eastward prompt penetration electric fields (PPEF). The associated  
17 increase in hmF2 and h'F, followed by enhancement in foF2, confirms the occurrence of a  
18 positive ionospheric storm driven by storm-time electrodynamics. The H-component exhibits a  
19 significant decrease across all stations, reflecting the intensification of the ring current, while  
20 enhanced fluctuations at higher latitudes indicate the contribution of auroral electrojets and field-  
21 aligned currents.

22 The thermospheric response, characterized by variations in the O/N<sub>2</sub> ratio, shows depletion  
23 during the main phase due to storm-time upwelling, followed by a marked increase during the  
24 recovery phase (18–19 March), indicating a transition toward an oxygen-rich thermosphere. This  
25 compositional change contributes to reduced recombination rates and sustained enhancement in  
26 electron density. The evolution of the equatorial electric field further reveals a transition from  
27 PPEF-dominated dynamics during the main phase to disturbance dynamo electric field (DDEF)  
28 effects during the recovery phase.



29 Overall, the results demonstrate that the ionospheric response to the storm is governed by a  
30 combination of prompt electrodynamic forcing and delayed thermospheric processes, with global  
31 coherence across longitude sectors and modulation by latitude and local time.

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34 Keywords

35 Geomagnetic storm, Ionosphere, PPEF, DDEF, O/N<sub>2</sub> ratio, Equatorial electrodynamics, St.

36 Patrick's Day Storm 2015.

37 1. Introduction

38 Geomagnetic storms represent one of the most significant drivers of variability in the near-Earth  
39 space environment, resulting from enhanced solar wind–magnetosphere coupling during  
40 disturbed solar conditions. These storms are typically initiated by the arrival of interplanetary  
41 coronal mass ejections (CMEs) or high-speed solar wind streams, which lead to intensified  
42 magnetospheric convection and energy deposition into the high-latitude ionosphere (Gonzalez et  
43 al., 1994; Tsurutani et al., 2006). The consequent disturbances propagate through the coupled  
44 magnetosphere–ionosphere–thermosphere system, producing complex and highly variable  
45 responses in ionospheric density, composition, and electrodynamics.

46 At low and equatorial latitudes, the ionospheric response to geomagnetic storms is primarily  
47 governed by storm-time electric fields and thermospheric dynamics. Two key electrodynamic  
48 processes are known to dominate: prompt penetration electric fields (PPEF) and disturbance  
49 dynamo electric fields (DDEF). PPEFs arise from rapid changes in magnetospheric convection  
50 driven by southward interplanetary magnetic field (IMF B<sub>z</sub>) and can penetrate to low latitudes  
51 with minimal shielding, producing immediate and global ionospheric responses (Kelley et al.,  
52 2003; Fejer et al., 2007). In contrast, DDEFs are generated by storm-time thermospheric winds  
53 driven by high-latitude Joule heating and particle precipitation, leading to delayed and longer-  
54 lasting modifications in the ionospheric electric field (Blanc and Richmond, 1980; Fuller-Rowell  
55 et al., 1994).

56 The F-region ionosphere exhibits significant structural changes during geomagnetic storms, often  
57 characterized by vertical plasma redistribution and variations in peak electron density. The uplift



58 of the F-layer, reflected in increases in  $hmF2$  and  $h'F$ , is primarily driven by enhanced eastward  
59 electric fields through upward  $E \times B$  drift, which can result in a positive ionospheric storm due to  
60 reduced recombination at higher altitudes (Prölss, 1995; Buonsanto, 1999). Conversely, storm-  
61 time thermospheric composition changes, particularly variations in the atomic-to-molecular ratio  
62 ( $O/N_2$ ), play a crucial role in controlling ionospheric density by modulating recombination rates.  
63 A decrease in  $O/N_2$  due to thermospheric upwelling leads to enhanced recombination and  
64 negative storm effects, whereas an increase during the recovery phase supports positive storm  
65 conditions (Rishbeth and Müller-Wodarg, 2006).

66 The geomagnetic field response during storms, commonly represented by the Dst index and  
67 ground-based H-component measurements, reflects the development of the ring current and  
68 associated current systems. A significant decrease in the H-component is indicative of ring  
69 current intensification, while superimposed short-term fluctuations are associated with auroral  
70 electrojets and field-aligned currents (Daglis et al., 1999; Kamide et al., 1998). These current  
71 systems play a critical role in coupling high-latitude energy inputs to low-latitude ionospheric  
72 dynamics.

73 The St. Patrick's Day Storm, which occurred on 17 March 2015, is one of the most intense  
74 geomagnetic storms of solar cycle 24, with a minimum Dst of approximately  $-223$  nT. This event  
75 has been widely studied due to its pronounced global ionospheric and thermospheric impacts,  
76 including significant modifications in electron density, electric fields, and neutral composition  
77 (Astafyeva et al., 2015; Nayak et al., 2016). The storm provides an excellent opportunity to  
78 investigate the interplay between electrodynamic and thermospheric processes across different  
79 latitude and longitude sectors.

80 Although several studies have examined the ionospheric response to this event, most analyses  
81 are limited to specific regions or individual parameters. A comprehensive, multi-parameter  
82 investigation combining ionospheric, geomagnetic, and thermospheric observations across  
83 multiple longitude sectors remains limited. In particular, the role of longitudinal variability and  
84 local time modulation in shaping the global ionospheric response requires further investigation.

85 In this context, the present study aims to examine the global ionospheric response to the March  
86 2015 geomagnetic storm using observations from multiple stations distributed across western  
87 and eastern longitude sectors. By integrating ionosonde measurements, geomagnetic field data,  
88 thermospheric composition, and satellite-derived electron density profiles, this study seeks to (1)  
89 characterize the temporal evolution of ionospheric parameters during different storm phases, (2)



90 identify the dominant electrodynamic processes controlling the response, and (3) assess the role  
91 of longitudinal and latitudinal variability in modulating storm-time ionospheric dynamics.

92 Unlike previous studies, which primarily focus on regional or single-parameter analyses, the  
93 present work provides a comprehensive multi-parameter investigation of the storm-time  
94 ionospheric response by integrating ionospheric, geomagnetic, and thermospheric observations  
95 across both western and eastern longitude sectors. The study highlights the global coherence of  
96 prompt penetration electric fields and quantifies the role of longitudinal and local time variability  
97 in modulating the ionospheric response.

## 98 2. Data and Methodology

99 The present study focuses on the intense geomagnetic disturbance associated with the St.  
100 Patrick's Day Storm during 15–20 March 2015. The storm was initiated by the arrival of an  
101 interplanetary coronal mass ejection (CME) at ~14:00 UT on 15 March 2015, followed by a  
102 prolonged southward turning of the interplanetary magnetic field (IMF  $B_z$ ), reaching values of  
103 approximately  $-30$  nT. This facilitated enhanced dayside magnetic reconnection and intensified  
104 magnetospheric convection.

105 The storm reached its peak intensity on 17 March, with the Dst index attaining a minimum value  
106 of  $\sim -223$  nT, indicating a severe geomagnetic storm (G4 level). The corresponding decrease in  
107 the H-component reflects the development of a strong ring current, while short-term fluctuations  
108 are associated with ionospheric current systems.

109 For analysis, the storm interval is categorized into three phases: initial phase (15 March 2015,  
110 ~14:00 UT), main phase (17 March 2015), and recovery phase (18–19 March 2015).

111 A combination of ground-based and satellite observations is utilized to examine the coupled  
112 magnetosphere–ionosphere–thermosphere system. The ionospheric parameters  $foF2$ ,  $hmF2$ , and  
113  $h'F$  are obtained from the Global Ionospheric Radio Observatory (GIRO). Here,  $foF2$  represents



114 the critical frequency proportional to peak electron density ( $NmF2$ ),  $hmF2$  denotes the height of  
115 maximum electron density (km), and  $h'F$  indicates the virtual height associated with rapid  
116 vertical plasma motion.

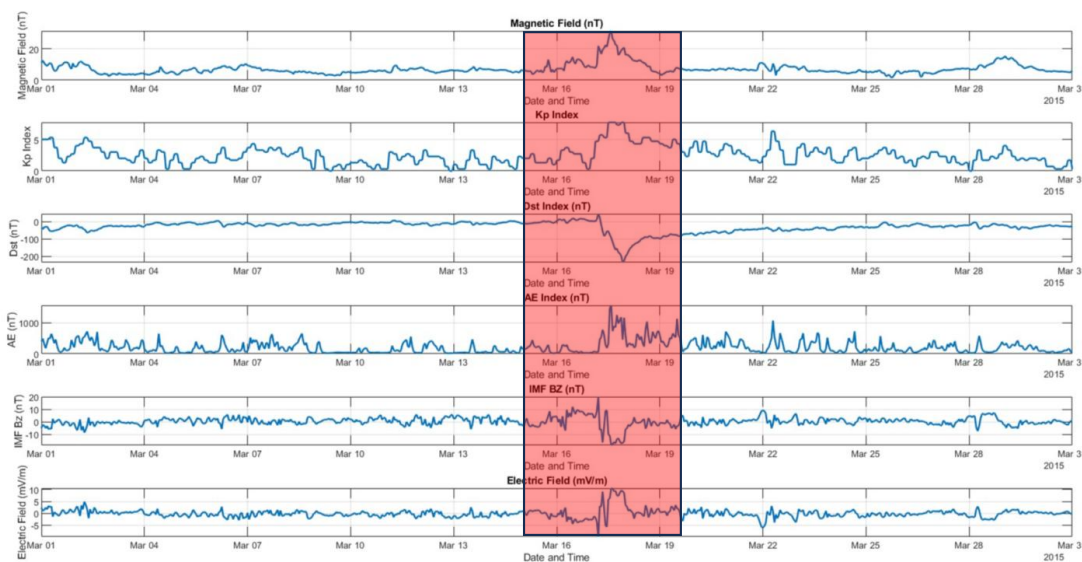
117 The thermospheric composition is characterized using the  $O/N_2$  ratio derived from the Global  
118 Ultraviolet Imager (GUVI) onboard the TIMED satellite. The  $O/N_2$  ratio serves as an indicator  
119 of neutral composition changes that influence ionospheric recombination processes. The  
120 horizontal component (H) of the geomagnetic field is obtained from ground-based magnetometer  
121 stations provided by the INTERMAGNET. These data are used to examine storm-time variations  
122 associated with ring current development and ionospheric current systems. Global geomagnetic  
123 activity is characterized using the Dst and Kp indices obtained from the World Data Center for  
124 Geomagnetism Kyoto. Where Dst represents the intensity of ring current and Kp index indicates  
125 the planetary scale geomagnetic disturbances. To examine the global and longitudinal  
126 characteristics of storm-time ionospheric response, data from seven stations are selected and  
127 grouped into two longitude sectors: western longitudinal sector (American sector) and eastern  
128 longitude sector (Asian- African sector), This classification enables the investigation of  
129 longitudinal differences in ionospheric response, particularly those arising from local time  
130 dependence of electrodynamic processes.

131 The analysis is carried out using a multi-parameter and multi-station approach. Temporal  
132 variations of ionospheric parameters ( $foF2$ ,  $hmF2$ ,  $h'F$ ), geomagnetic field (H-component),  $O/N_2$   
133 ratio, and geomagnetic indices are examined over the storm period in Universal Time (UT).  
134 Storm-time observations are compared with corresponding quiet-day conditions ( $Kp \leq 2$ ) to  
135 isolate storm-induced disturbances.

136 Electrodynamic signatures are identified using variations in F-region parameters: increases in  
137  $hmF2$  and  $h'F$  indicate eastward electric fields (upward  $E \times B$  drift), while decreases indicate  
138 westward electric fields (downward drift). The observed variations are interpreted in conjunction



139 with IMF Bz, Dst, and Kp indices to identify the role of prompt penetration electric fields (PPEF)  
140 and disturbance dynamo electric fields (DDEF). Latitudinal and longitudinal analyses are  
141 performed to distinguish between global electrodynamic effects and localized thermospheric  
142 processes.



143  
144 Figure 1. Temporal variations of geomagnetic indices (Dst and Kp), IMF Bz, and H-component  
145 during 15–20 March 2015. The shaded region highlights the main phase of the storm. The sharp  
146 southward turning of IMF Bz and corresponding decrease in Dst indicate the onset of intense  
147 geomagnetic activity.

148 A key feature observed during the main phase is the simultaneous disturbance at ~06:50 UT  
149 across all stations, including Yakutsk, Beijing, Guam, Jicamarca, Fortaleza, Hermanus, and Port  
150 Stanley. This near-simultaneity across a wide longitudinal sector indicates the penetration of  
151 magnetospheric electric fields to low latitudes, characteristic of prompt penetration electric fields  
152 (PPEF).



153 The immediate ionospheric response is manifested as abrupt variations in hmF2 and h'F,  
154 indicating rapid vertical plasma redistribution.

### 155 3. Results and Discussions

156 At equatorial and low-latitude stations within the western longitude sector, namely Jicamarca and  
157 Fortaleza, along with the southern high-latitude station Port Stanley, a pronounced storm-time  
158 disturbance is observed during the main phase.

159 At the equatorial and low-latitude stations, a clear uplift of the F-region is evident, manifested as  
160 a near-simultaneous increase in hmF2 and h'F around ~06:50 UT on 17 March 2015. This rapid  
161 uplift indicates the presence of an enhanced eastward electric field associated with prompt  
162 penetration electric fields (PPEF), which drives upward  $E \times B$  drift.

163 The vertical transport of plasma to higher altitudes reduces recombination rates, resulting in a  
164 subsequent enhancement of foF2, characteristic of a positive ionospheric storm. The coherent  
165 response between Jicamarca and Fortaleza confirms the strong influence of storm-time  
166 electrodynamics in the equatorial region.

167 In contrast, the response at Port Stanley exhibits comparatively larger variability, reflecting the  
168 combined influence of electrodynamic processes and high-latitude energy inputs such as particle  
169 precipitation and Joule heating. Despite these differences, the simultaneous disturbance across  
170 the western sector indicates that the initial ionospheric response is driven by a globally coherent  
171 penetration electric field, with latitude-dependent modulation.

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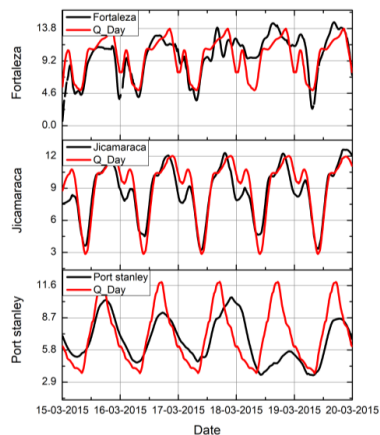
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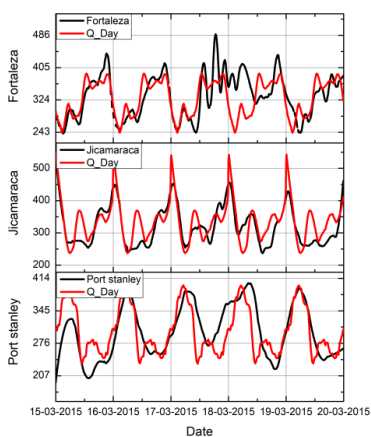


176 2 (a)

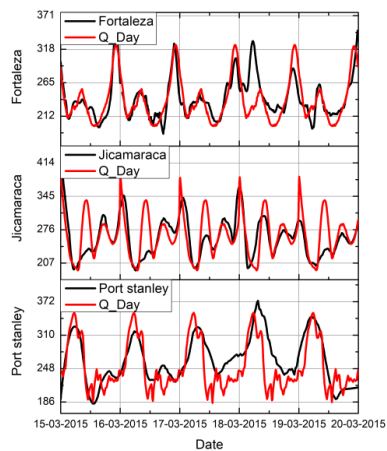


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2(b)



178 2(c)



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180 Figure 2. Temporal variations of ionospheric parameters (a) foF2, (b) h'F and (c) hmF2 during

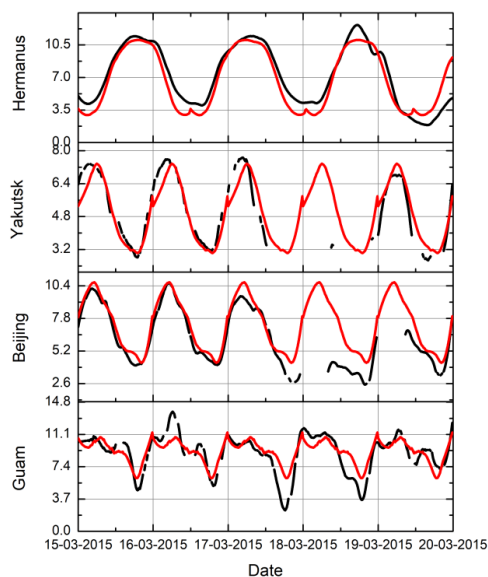
181 storm and quiet conditions at Western Longitude Sector (American sector) (Fortaleza, Jicamarca,

182 and Port Stanley) for 15–20 March 2015.



183 During the later phase (~15 UT on 17 March), a reversal to westward electric fields is inferred  
184 from the downward movement of the F-region. This leads to suppression of the pre-reversal  
185 enhancement (PRE), a key driver of equatorial plasma instabilities. The inhibition of PRE  
186 indicates unfavorable conditions for the development of equatorial spread-F during this interval.  
187 Overall, the equatorial and low-latitude ionospheric response during the storm is predominantly  
188 controlled by storm-time electrodynamics, with PPEF governing the initial uplift and subsequent  
189 electric field reversal modulating the evening ionospheric dynamics.

190 3(a)



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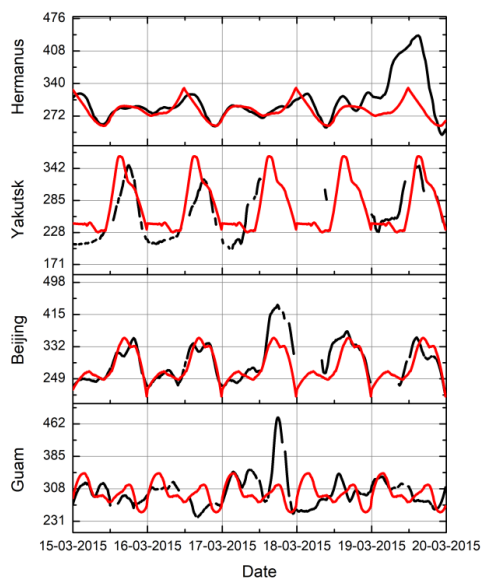
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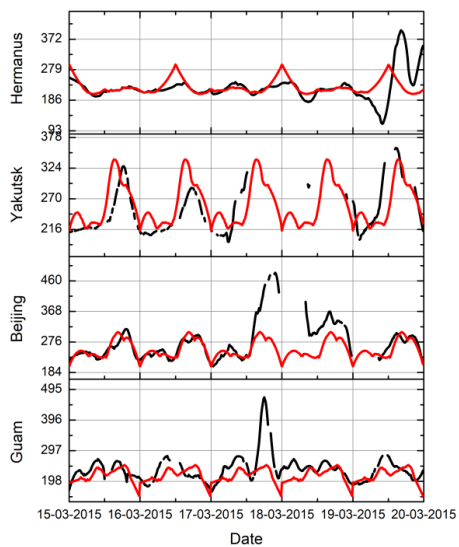


196 3(b)



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198 3(c)



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200 Figure 3. Temporal variations of ionospheric parameters (a) foF2, (b) h'F and (c) hmF2 during  
201 storm-time and quiet-day conditions at the Eastern Longitude Sector (Asian–African sector)  
202 (Yakutsk, Beijing, Guam, and Hermanus) for 15–20 March 2015.

203 The variations in the horizontal component (H) of the geomagnetic field clearly capture the  
204 evolution of storm-time magnetospheric and ionospheric current systems during the St. Patrick's  
205 Day Storm. The multi-station observations reveal both global signatures of ring current  
206 development and localized effects associated with ionospheric and auroral processes.

207 In the Western longitude sector (American sector) Jicamarca, Fortaleza, and Port Stanley a  
208 pronounced decrease in the H-component is observed during the main phase (17 March),  
209 consistent with the intensification of the storm-time ring current. The depletion is relatively  
210 smooth at the low- and equatorial-latitude stations (Jicamarca and Fortaleza), reflecting the  
211 dominant influence of the symmetric ring current.

212 In contrast, the response at Port Stanley exhibits enhanced fluctuations superimposed on the main  
213 depression. These rapid variations are indicative of high-latitude processes, including auroral  
214 electrojet activity, particle precipitation, and intensified Joule heating. This highlights the  
215 transition from globally driven magnetospheric currents at low latitudes to locally driven  
216 electrodynamic processes at higher latitudes.

217 In the Eastern longitude sector (Asian–African sector)—Yakutsk, Beijing, Guam, and  
218 Hermanus—a comparable large-scale decrease in the H-component is observed, confirming the  
219 global development of the ring current system. However, the temporal evolution and amplitude  
220 of the disturbance show noticeable variability across stations.

221 At higher latitudes, particularly at Yakutsk and Hermanus, the H-component exhibits strong  
222 irregular fluctuations, reflecting the influence of auroral current systems and field-aligned  
223 currents. At mid- and low-latitude stations such as Beijing and Guam, the response is



224 comparatively smoother, dominated by large-scale magnetospheric currents with limited local  
225 perturbations.

226 The simultaneous decrease in the H-component across both longitude sectors during the main  
227 phase indicates the global enhancement of the ring current, driven by sustained southward IMF  
228 Bz and enhanced magnetospheric convection. Superimposed short-term perturbations,  
229 particularly at higher latitudes, highlight the role of auroral electrojets, ionospheric currents, and  
230 field-aligned current systems.

231 Overall, the observations demonstrate that while the ring current produces a globally coherent  
232 depression in the H-component, the superimposed variability is strongly latitude-dependent and  
233 modulated by local electrodynamic processes, with minimal longitudinal dependence in the  
234 large-scale structure.

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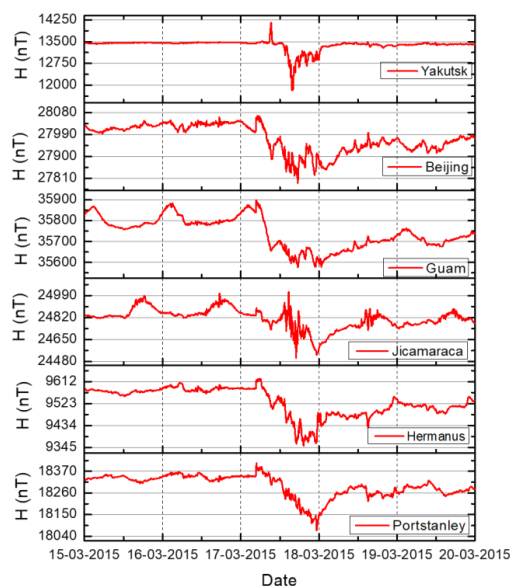
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242 Figure:4 Temporal variations of the horizontal component (H) of the geomagnetic field at  
243 multiple stations during 15–20 March 2015.

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245 The thermospheric composition, represented by the  $O/N_2$  ratio, exhibits significant variability  
246 during the storm period and plays a crucial role in modulating ionospheric density through  
247 changes in recombination rates.

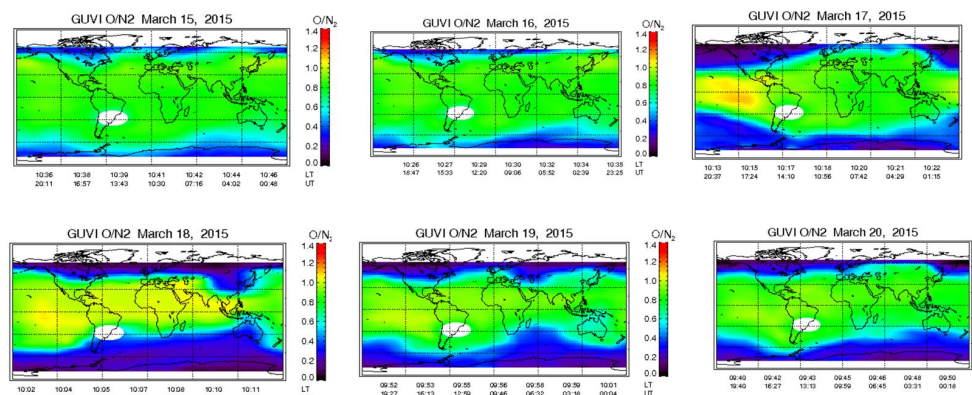
248 During the main phase (17 March 2015), a decrease in  $O/N_2$  is inferred, particularly at mid- and  
249 high-latitudes, indicating an enhancement of molecular species ( $N_2$  and  $O_2$ ) driven by storm-time  
250 thermospheric upwelling. This compositional disturbance leads to increased recombination rates  
251 and contributes to the observed negative ionospheric storm effects, especially at stations in both  
252 longitude sectors.

253 In the Western longitude sector (American sector)—Jicamarca, Fortaleza, and Port Stanley—the  
254 recovery phase (18 March) is characterized by a marked increase in  $O/N_2$ , particularly at low and



255 equatorial latitudes. This enhancement indicates a transition toward an oxygen-rich  
256 thermosphere, reducing recombination rates and facilitating an increase in electron density. The  
257 effect is most evident at Jicamarca and Fortaleza, where the positive ionospheric storm response  
258 is more pronounced.

259 In the Eastern longitude sector (Asian–African sector)—Yakutsk, Beijing, Guam, and  
260 Hermanus—a similar recovery-phase enhancement in  $O/N_2$  is observed, although with slight  
261 temporal differences across stations. These variations can be attributed to longitudinal differences  
262 in thermospheric circulation, driven by high-latitude energy input and subsequent equatorward  
263 propagation of disturbed neutral winds.



264

265 Figure:5 Temporal variations of the thermospheric  $O/N_2$  ratio during 15–20 March 2015 over  
266 multiple locations, including Jicamarca, Fortaleza, Port Stanley, Yakutsk, Beijing, Guam, and  
267 Hermanus.

268 Overall, the  $O/N_2$  variations indicate that thermospheric composition changes dominate during  
269 the recovery phase, contributing significantly to the positive ionospheric storm effect, while their  
270 influence during the main phase is secondary to electrodynamic forcing. The recovery-phase  
271 enhancement in  $O/N_2$  is temporally consistent with the observed increase in foF2, confirming the  
272 dominant role of thermospheric composition in driving positive ionospheric storm effects

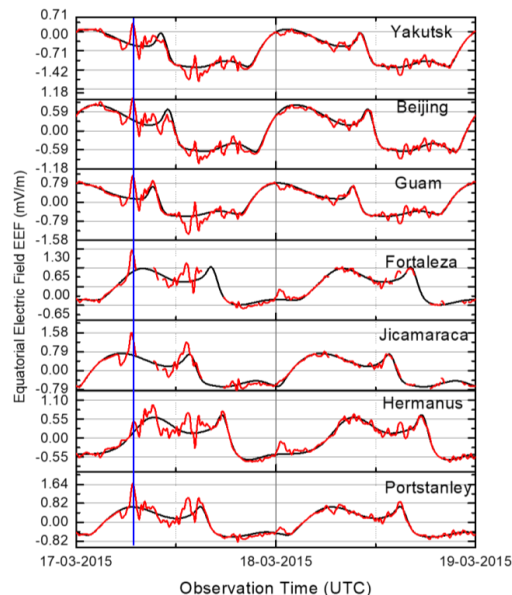


273 The equatorial electric field (EEF), inferred from the temporal variations in hmF2 and h'F,  
274 exhibits distinct signatures corresponding to storm-time electrodynamic processes and plays a  
275 dominant role in controlling the low-latitude ionospheric response.

276 During the main phase (~06:50 UT on 17 March 2015), a clear enhancement of the eastward  
277 electric field is observed simultaneously across both longitude sectors. This is evidenced by the  
278 rapid uplift of the F-region at equatorial stations in the Western sector—Jicamarca and  
279 Fortaleza—and in the Eastern sector at Guam. The simultaneity of this response confirms the  
280 action of prompt penetration electric fields (PPEF), which are global in extent and rapidly  
281 transmitted from high to low latitudes.

282 The enhanced eastward EEF drives upward  $E \times B$  drift, leading to a rapid increase in hmF2 and  
283 h'F, and subsequently contributing to the enhancement of foF2. This process is responsible for  
284 the observed positive ionospheric storm effect at low and equatorial latitudes during the initial  
285 stage of the main phase.

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288 Figure:6 Temporal variations of the equatorial electric field (EEF) during 15–20 March 2015 at  
289 low- and equatorial-latitude stations, including Jicamarca, Fortaleza, and Guam.

290 During the later phase (~15 UT on 17 March), a reversal to westward electric fields is inferred  
291 from the observed downward movement of the F-region. This westward EEF suppresses the pre-  
292 reversal enhancement (PRE), thereby inhibiting the development of equatorial plasma  
293 instabilities. The consistency of this reversal across both longitude sectors indicates that the  
294 penetration of westward electric fields is also a globally coherent process.

295 In the recovery phase (18–19 March), the electric field is primarily governed by disturbance  
296 dynamo electric fields (DDEF), generated by storm-time thermospheric winds. The DDEF,  
297 typically eastward during nighttime, contributes to sustained uplift of the F-region and enhances  
298 electron density, particularly at low latitudes.

299 Although the primary features of EEF variation are consistent across both longitude sectors,  
300 minor differences in timing and magnitude can be attributed to local time dependence, which



301 modulates the efficiency of electrodynamic coupling and plasma transport. This transition from  
302 PPEF-dominated to DDEF-dominated electrodynamics highlights the temporal evolution of  
303 storm-time ionospheric forcing mechanisms.

#### 304 4. Conclusion

305 The present study investigates the global ionospheric response to the St. Patrick's Day  
306 Storm during 15–20 March 2015 using a combination of ground-based and satellite observations  
307 across multiple longitude sectors. The analysis reveals a complex interplay between  
308 magnetospheric forcing, ionospheric electrodynamics, and thermospheric composition.

309 During the main phase of the storm, a near-simultaneous uplift of the F-region is observed at  
310 equatorial and low-latitude stations in both longitude sectors, indicating the action of prompt  
311 penetration electric fields (PPEF). The associated enhancement in  $hmF_2$  and  $h'F$ , followed by an  
312 increase in  $foF_2$ , confirms that storm-time electrodynamics plays a dominant role in driving the  
313 initial positive ionospheric response. The global coherence of this response highlights the rapid  
314 transmission of high-latitude electric fields to low latitudes.

315 The geomagnetic field variations show a pronounced decrease in the H-component across all  
316 stations, reflecting the development of a strong ring current during the main phase. Superimposed  
317 fluctuations at higher latitudes emphasize the contribution of auroral electrojets and field-aligned  
318 currents, demonstrating the combined influence of global and localized current systems.

319 Thermospheric composition, represented by the  $O/N_2$  ratio, exhibits a depletion during the main  
320 phase due to storm-time upwelling, followed by a significant enhancement during the recovery  
321 phase. The increase in  $O/N_2$ , particularly at low and equatorial latitudes, reduces recombination  
322 rates and contributes to the sustained positive ionospheric storm effect.

323 The evolution of the equatorial electric field further illustrates the transition of dominant physical  
324 processes. While PPEF governs the initial phase through enhanced eastward electric fields and



325 upward plasma drift, a subsequent reversal to westward electric fields leads to suppression of the  
326 pre-reversal enhancement. During the recovery phase, disturbance dynamo electric fields  
327 (DDEF) become dominant, contributing to continued ionospheric uplift and delayed recovery.

328 Overall, the results demonstrate that the ionospheric response to the storm is governed by a  
329 combination of prompt electrodynamic forcing during the main phase and thermospheric  
330 composition changes coupled with disturbance dynamo effects during the recovery phase. The  
331 observed consistency across longitude sectors confirms the global nature of storm-time  
332 electrodynamics, while variations in magnitude and timing highlight the role of latitude and local  
333 time in modulating the response.

#### 334 5. Data availability

335 The data used in this study are publicly available. TEC data were obtained from GIRO,  
336 Dst index data from the World Data Center for Geomagnetism, Kyoto, and Kp index data  
337 from the GFZ German Research Centre for Geosciences. FORMOSAT/COSMIC data  
338 were accessed from the CDAAC repository. All datasets are available from the respective  
339 data centers upon request or via their official websites.

#### 340 6. Author contribution

341 The author confirms sole responsibility for the study conception and design, data collection,  
342 analysis and interpretation of results, and manuscript preparation.

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