



1 **Complexation strength between organic carbon and transition metal ions dominates the**
2 **photochemical conversion of SO₂ to sulfates**

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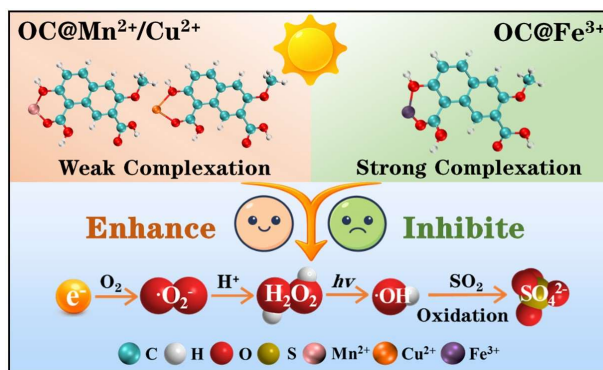
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8 **Abstract**

9 The photooxidation of SO₂ on organic carbon (OC) is a critical pathway for sulfate formation,
10 yet the role of interactions between OC and transition metal ions (TMIs) in this process remains
11 scarcely understood. We systematically investigated potential influences of TMIs (Fe³⁺, Cu²⁺
12 and Mn²⁺) on the conversion of SO₂ to sulfates on OC from coal combustion under irradiation.
13 Fe³⁺ exhibited a significantly antagonistic effect, whereas Cu²⁺ and Mn²⁺ enhanced SO₂ uptake
14 and sulfate generation on OC. Spectroscopic evidences, including absorbance decreasing and
15 fluorescence quenching, confirmed the complexes formation of TMI with chromophores in OC.
16 Fe³⁺ owned the strongest binding affinity with chromophores, followed by Cu²⁺ and Mn²⁺. This
17 variation in the coordination strength dominated the generation of reactive species, such as free
18 electrons (e⁻), superoxide radicals (•O₂⁻), H₂O₂ and hydroxyl radicals (•OH), and •OH acted as
19 a pivotal trigger to drive the sulfate production. Extended investigations containing more metal
20 ions demonstrated that the regulatory effects on the sulfate production were generally governed
21 through their ability to suppress or facilitate the •OH generation.



22 **Graphical abstract:**



23

24 **1 Introduction**

25 Fine particulate matter (PM_{2.5}) is a complex mixture of inorganic and organic species, and
 26 plays a central role in the persistent haze events (Fan et al., 2016; Huang et al., 2014; Peng et
 27 al., 2021). The rapid accumulation of sulfates in PM_{2.5} significantly influences atmospheric
 28 radiative forcing, air quality and human health (Zhang et al., 2020a; Salana et al., 2024).
 29 Sulfates primarily stem from SO₂ oxidation, including reactions with hydroxyl radicals (•OH)
 30 in the gas phase and aqueous oxidation in cloud and fog droplets involving various oxidizing
 31 species (Tsona and Du, 2019; Liu et al., 2020a; Jiang et al., 2009; Hoyle et al., 2016; Liu and
 32 Abbatt, 2021). However, only considering these pathways usually leads to a pronounced gap
 33 between the modeled and observed sulfate concentrations during the haze events (Li et al.,
 34 2017; Eckhardt et al., 2015; Zheng et al., 2015).

35 To bridge this gap, several heterogeneous pathways have been proposed, involving reactions
 36 of SO₂ on PM_{2.5}, mineral dust (e.g., Arizona test dust and metal oxides), inorganic salt (e.g.,
 37 nitrates and NaCl), soot and organic aerosols (Zhang et al., 2020c; Zhang et al., 2019; Dupart
 38 et al., 2012; Yang et al., 2024a; Cao et al., 2024b; Zhang et al., 2022). Although these
 39 heterogeneous reactions have narrowed the discrepancies between model prediction and field
 40 observation concentrations of sulfates, substantial uncertainties and controversies still remain
 41 (Wang et al., 2014). This may be attributed to the fact that most studies mainly focus on
 42 simplified single-component systems and overlook the extreme complexity of multicomponent
 43 aerosols under realistic atmospheric conditions. Recent evidence highlighted that the potential



44 interactions among distinct components in aerosols can significantly influence the
45 transformation pathways of atmospheric pollutants (Zhang and Chan, 2023; Zhang et al., 2023).
46 For instance, organic carbon (OC) in soot can donate electrons to elemental carbon (EC) and
47 enhance the electron-hole separation under irradiation, thereby generating more •OH and
48 promoting the conversion of SO₂ to sulfates (Zhu et al., 2022). OC-derived photosensitizers
49 activated both O₂ and Cl⁻ in NaCl-OC under ultraviolet light, producing •OH and Cl• that
50 synergistically oxidized SO₂ to sulfates (Tang et al., 2023).

51 OC from coal combustion can serve as a photosensitive electron donor under irradiation,
52 initiating the formation of reactive oxygen species (ROS), which subsequently oxidized SO₂ to
53 sulfates (Yang et al., 2025). In atmospheric aerosols, OC often coexisted with various TMIs,
54 especially in some regions influenced by the combustion, vehicular emissions, and industrial
55 processes (Yang et al., 2024b; Deng et al., 2022; Zhang et al., 2020b; Shiraiwa et al., 2017;
56 Hua et al., 2025). This coexistence was not a simple superposition of their specific roles. OC
57 usually contained diverse functional groups such as carboxyl, hydroxyl, and carbonyl moieties,
58 which can interact with TMIs via the complexation, metal-proton ion exchange and
59 electrostatic adsorption, forming organometallic complexes in aerosols (Li et al., 2022; Wang
60 et al., 2021c; Wang et al., 2021b). Furthermore, the binding strength between OC and TMIs
61 varied with the specific type of TMIs, which was related to the differences in ionic radius,
62 oxidation state, and electronic configuration that influenced coordination chemistry and redox
63 reactivity (Wang et al., 2021a; Pan et al., 2020). These interactions were expected to exert
64 antagonistic or synergistic effects on the ROS formation (Li et al., 2022; Pan et al., 2020; Wan
65 et al., 2019; Lu et al., 2019). The potential roles of OC and TMIs interactions in the
66 heterogeneous photochemical oxidation of SO₂ haven't been systematically investigated, which
67 may restrict the exact incorporation of OC photochemistry into the sulfate formation prediction
68 models.

69 Our previous work mainly focused on the intrinsic photo-reactivity of OC and its dependence
70 on the physicochemical properties (Yang et al., 2025). In this study, laboratory experiments and
71 theoretical calculations were conducted to elucidate the influences of TMIs (Fe³⁺, Mn²⁺ and
72 Cu²⁺) on the photooxidation of SO₂ to sulfates on OC from coal combustion. Differences in the
73 photochemical activity were quantified by measuring SO₂ uptake coefficients and sulfate yields.



74 Changes in the optical properties were examined to evaluate the potential interactions between
75 OC and TMI. Mechanism insights into the reactive species generation, including e^- , $\bullet O_2^-$,
76 H_2O_2 and $\bullet OH$, clarified how these interactions regulated the SO_2 photooxidation. Finally,
77 further investigation involving multiple metal ions commonly within atmospheric aerosols
78 highlighted the broader significance of OC and TMI interactions in the sulfate formation.

79 **2 Materials and methods**

80 **2.1 Preparation of OC and OC@TMI samples**

81 As shown in Fig. S1, OC was obtained with a custom combustion and sampling apparatus.
82 The coal in the combustion experiment originated from Shanxi Province, China. The details of
83 OC sampling and preparation are provided in Text S1. 1 mL of the extracted OC solution was
84 dropped onto 6.0×1.5 cm prebaked quartz-fiber filters, and the solvents were dried under a N_2
85 stream (100 mL min^{-1}) in the dark at 298 K. Then, the organic carbon content was measured
86 to be $818.4 \pm 53.0 \text{ } \mu\text{gC mL}^{-1}$ using a total organic carbon analyzer (TOC-5000 RD, Metash).

87 The sources and purities of salt chlorides used here are provided in Text S2. Stock solutions
88 of metal salts (1.5 mM, in ultrapure water) were mixed with the OC solution at a 1:2 volume
89 ratio, yielding mixtures containing 0.5 mM metal ions and $545.6 \pm 35.3 \text{ } \mu\text{gC mL}^{-1}$ (Fig. S1).
90 This metal-ion concentration was close to the levels reported in heavily polluted atmospheric
91 conditions. (Li et al., 2022) Each mixture was ultrasonicated for 30 min at 25 °C in the dark.
92 The resulting products were designated as OC@TMI, where OC@TMI referred to a series of
93 samples prepared with different TMIs (e.g., OC@ Fe^{3+} , OC@ Cu^{2+} and OC@ Mn^{2+}). 1.5 mL of
94 the mixed solution was deposited on the inner surface of a quartz flow tube (20 cm length, 1.0
95 cm inner diameter). The flow tube was rotated slowly to ensure the uniform wetting of the
96 surface. The coating was obtained by drying the samples under a N_2 stream (100 mL min^{-1}) in
97 the dark environment at 298 K to prevent photochemical interference.

98 **2.2 Flow tube experiments and SO_2 uptake coefficients**

99 The SO_2 uptake experiments were performed at ambient pressure in a horizontal cylindrical
100 quartz flow tube reactor (34 cm length, 1.6 cm i.d.), as shown in Fig.S1. The details of the flow
101 reactor and reaction procedure are given in Text S3. The changes in the SO_2 concentration were
102 measured with a SO_2 analyzer (Thermo 43i). The calculation of the SO_2 uptake coefficient is



103 described in Text S4.

104 **2.3 Analysis methods**

105 **2.3.1 Measurements of sulfate products**

106 The functional group changes during the photochemical reaction of SO₂ on the samples were
107 characterized using in situ attenuated total reflection infrared (ATR-IR) spectroscopy (Nicolet
108 iS50, Thermo Scientific), which equipped with a mercury-cadmium-telluride (MCT) detector.
109 Descriptions of experimental process are given in Text S5.

110 The sulfate ions were quantified using an ion chromatography (IC) system (CIC-D120+,
111 Shenghan) with an analytical column (AS11-HC, Thermo) and a conductivity detector. The
112 testing procedures are described in Text S6.

113 **2.3.2 Optical property measurements**

114 The optical properties of OC, TMIs, and OC@TMIs were analyzed using the UV-visible
115 absorption spectroscopy and the excitation-emission matrix (EEM) fluorescence spectroscopy.
116 The UV-visible spectra were recorded by the UV-visible spectrophotometer (UV-2550,
117 Shimadzu) with a 1.0 cm pathlength cell in the wavelength range of 200–600 nm. A mixed
118 solution of methanol and water (2:1, v/v) was used as the blank background.

119 The EEM fluorescence spectroscopy were measured via a fluorescence spectrophotometer
120 (FP8550, Jasco) at 700 V voltage and 3-D mode in the wavelength ranges of 200–400 nm for
121 excitation (Ex.) and 300–520 nm for emission (Em.). The fluorescence spectra were recorded
122 at 5 nm intervals and a scan speed of 2400 nm min⁻¹ in a 1 cm path-length quartz cuvette.
123 Samples were diluted to one-fifteenth of the original concentrations to minimize the inner-filter
124 effect (IFE) and reabsorption artifacts, which can distort the excitation–emission spectra at
125 high sample concentrations. Instrument calibration, correction for inner-filter effects and
126 normalization of fluorescence intensity to the Raman units (RU) of solvent were conducted
127 before EEM spectral analysis. The EEM data were modeled with the parallel factor analysis
128 (PARAFAC) through the DOMFluo toolbox (version 0.2.0) for MATLAB. Based on the
129 evaluation of 2–4 components using EEM profiles and residual error patterns, a 3-component
130 model was ultimately selected (Fig. S4).

131 **2.3.3 Electrochemical measurements**

132 The electron donating capacities (EDC) of the samples were determined using an



133 electrochemical workstation (CHI760F, Shanghai Chenhua Co., Ltd.) equipped with a three-
134 electrode system. The testing procedures are described in Text S7. The EDC was calculated as
135 the following Eq. (1) (Ma et al., 2024),

136
$$\text{EDC} = \frac{\int \frac{I_{\text{ox}}}{F} dt}{m_{\text{organic carbon}}} \quad (1)$$

137 where I_{ox} means the OC oxidation curves; F is the constant of Faraday, which is equal to 96485
138 $\text{s A mol}^{-1} \text{e}^{-}$; $m_{\text{organic carbon}}$ represents the amount of organic carbon of samples.

139 2.3.4 EPR measurements

140 Reactive oxygen species (ROS) under irradiation were analyzed using an electron
141 paramagnetic resonance spectrometer (A300, Bruker). For the radical detection, 20 μL of 0.05
142 M DMPO (in methanol) and BMPO (in water) were used as spin traps for superoxide radicals
143 ($\bullet\text{O}_2^-$) and $\bullet\text{OH}$, respectively. The instrument parameters were set as follows: magnetic field
144 range of 3450–3550 G, central field range of 3500 G, scan time of 30 s and microwave power
145 of 20 mW.

146 2.3.5 H_2O_2 detection

147 The H_2O_2 generation was analyzed by the TiOSO_4 colorimetric method (Liu et al., 2023; Li
148 et al., 2023). TiOSO_4 reacted with H_2O_2 to form a yellow titanium-peroxide complex
149 (Ti(IV)O_2^{2+}), which exhibited a characteristic absorbance peak at 405 nm. The H_2O_2
150 concentration was determined using the H_2O_2 standard curve (Fig. S5), and the testing
151 procedure is given in Text S8.

152 2.3.6 X-ray photoelectron spectroscopy

153 The X-ray photoelectron spectroscopy (XPS, ESCALAB 250Xi, Thermo) was employed to
154 investigate the valence state variations of metal ions under irradiation. Fresh samples (3 mL,
155 prepared by 2 mL OC extraction with 1 mL ionic solution) or the ones photoaged by SO_2 were
156 combined with SiO_2 (100 mg) and dried by rotary evaporation under ultra-high vacuum
157 conditions (10^{-9} mbar). For XPS analysis, aluminum $K\alpha$ radiation ($h\nu = 1486.68$ eV, spot size
158 $500 \mu\text{m}$) was used as the excitation source. The instrument was operated at 15 kV and 10 mA.
159 Signals were accumulated over 5–10 scans, and the spectra were recorded with a pass energy
160 of 30 eV and a step size of 0.05 eV. Peak deconvolution was conducted to resolve the
161 contributions of different valence states.



162 2.3.7 Density functional theory calculations

163 Density functional theory (DFT) calculations were carried out through the quantum
164 chemistry software Gaussian 16. Geometry optimizations and energy calculations were
165 conducted at the B3LYP/LANL2DZ level and B3LYP-D3/def2-TZVP level, respectively. O₂
166 adsorption on OC and the corresponding IGMH diagrams were analyzed using Multiwfn (Lu
167 and Chen, 2011; Lu, 2024). The calculation about the complexation energy of OC with TMIs
168 and O₂ adsorption energy are described in Text S9.

169 3 Results and discussion

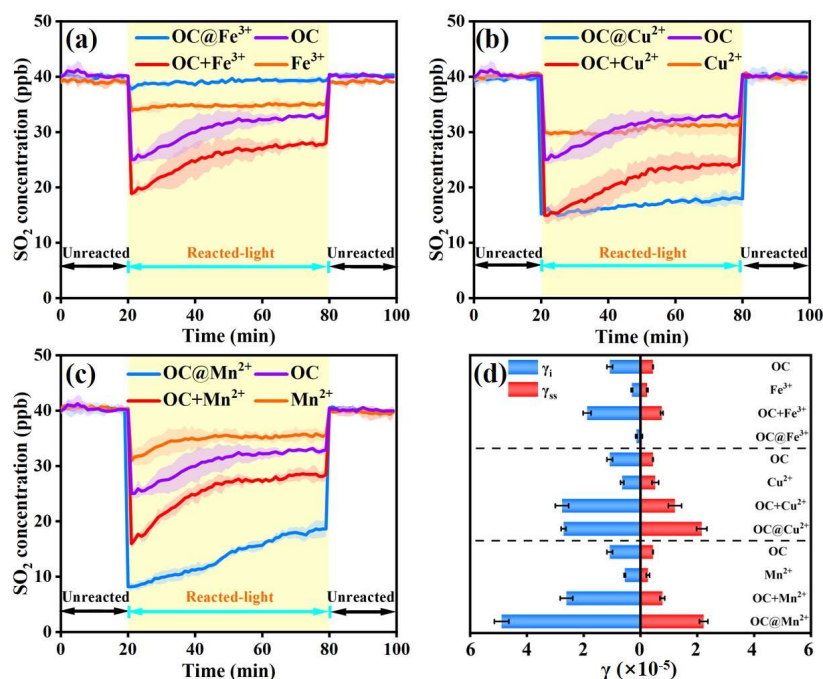
170 3.1 Dependence of SO₂ uptake on metal ions

171 Fig. 1a–c illustrates the temporal evolution of SO₂ concentration on OC, TMIs and
172 OC@TMIs under irradiation. Here, the SO₂ signal for OC+TMIs represented the arithmetic
173 sum of the SO₂ uptake measured for OC and TMIs in the separate experiments. All errors
174 denoted the standard deviations of three independent replicates. Because chloride salts were
175 used here, OC@NaCl control experiments were conducted. The decreasing trend of SO₂
176 concentration on OC@NaCl was the same as that on OC, suggesting no measurable effect of
177 Cl⁻ on the SO₂ uptake (Fig. S6). The SO₂ loss on FeCl₃ or MnCl₂ was less than that on OC,
178 while it was similar for CuCl₂ and OC. When SO₂ was exposed to OC@Fe³⁺, the SO₂
179 concentration decreased by 1–2 ppb, which was obviously lower than that for OC+Fe³⁺. This
180 indicates that the coexistence of Fe³⁺ with OC markedly suppressed the photochemical uptake
181 of SO₂. The SO₂ loss on OC@Cu²⁺ and OC@Mn²⁺ was markedly greater than that on OC+Cu²⁺
182 and OC+Mn²⁺, respectively. This suggests that both Cu²⁺ and Mn²⁺ enhanced the
183 photochemical reaction of SO₂ with OC.

184 The initial (γ_i) and steady-state (γ_{ss}) uptake coefficients of SO₂ are summarized in Fig. 1d.
185 To better quantify TMIs effects, the uptake coefficient for OC+TMIs was calculated using the
186 SO₂ uptake amount for OC+TMIs as the reference. The γ_{ss} was consistently lower than the γ_i
187 due to the gradual aging and the depletion of photoactive components during the reaction. The
188 γ_{ss} on OC@Fe³⁺ only accounted for 5% of that on OC+Fe³⁺, confirming that Fe³⁺ significantly
189 restrained the photochemical SO₂ uptake on OC. In contrast, γ_{ss} for OC@Cu²⁺ and OC@Mn²⁺
190 was 1.83 and 2.93 times of the corresponding one for OC+Cu²⁺ and OC+Mn²⁺, respectively.



191 This definitely suggests that the coexistence of Cu^{2+} and Mn^{2+} with OC has synergistic effects
 192 on the photochemical uptake of SO_2 , with Mn^{2+} exhibiting a stronger enhancement than Cu^{2+} .
 193 Overall, these contrasting behaviors reveal that the addition of TMIs to OC can either restrict
 194 (Fe^{3+}) or facilitate (Cu^{2+} and Mn^{2+}) the photochemical uptake of SO_2 , highlighting the distinct
 195 roles of TMIs in modulating the OC reactivity.



196

197 Figure 1. (a–c) Temporal variations of the SO_2 concentration during the heterogeneous
 198 reaction of SO_2 with OC, TMIs, OC+TMIs and OC@TMIs under irradiation. (OC+TMIs
 199 denoted the arithmetic sum of the SO_2 uptake measured for OC and TMIs in the separate
 200 experiments). (d) Initial uptake coefficients (γ_i) and steady-state uptake coefficients (γ_{ss}) of
 201 SO_2 under irradiation. Reaction conditions: irradiance of 1.05×10^{16} photons $\text{cm}^{-2} \text{s}^{-1}$, 40 ppb
 202 SO_2 , 298 K and 60% RH.

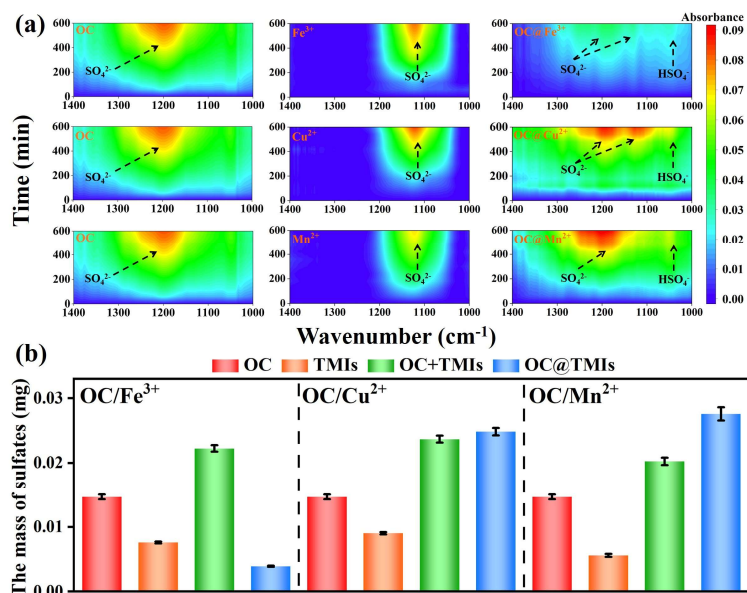
203 3.2 Dependence of sulfate formation on metal ions

204 The contour maps of in-situ ATR-IR spectra of OC, TMIs and OC@TMIs exposed to SO_2
 205 under irradiation are displayed in Fig. 2a. Several characteristic sulfate (SO_4^{2-} or HSO_4^-) bands
 206 were identified for all samples in the $1300\text{--}1000 \text{ cm}^{-1}$ region (Zhang et al., 2023; Zhang et al.,



207 2019). It was noted that these bands intensified progressively with the time, reflecting the
 208 continuous accumulation of sulfates.

209 The IC results provided quantitative evidence of sulfate formation under irradiation (Fig. 2b).
 210 For OC@Fe³⁺, sulfate production was much lower than that on OC or FeCl₃. Specifically, the
 211 sulfate mass on OC@Fe³⁺ was merely 18% of that for OC+Fe³⁺, confirming that Fe³⁺ restrained
 212 the formation of sulfates on OC. In contrast, the sulfate mass on OC@Cu²⁺ and OC@Mn²⁺ was
 213 1.08 and 1.36 times larger than that for OC+Cu²⁺ and OC+Mn²⁺, respectively. These confirm
 214 that both Cu²⁺ and Mn²⁺ promote the photochemical oxidation of SO₂ to sulfates on OC. Taken
 215 together with the SO₂ uptake data in Fig.1, TMIs exerted markedly different influences on SO₂
 216 conversion to sulfates on OC.



217
 218 Figure 2. (a) Contour map of in-situ ATR-IR spectra of OC, FeCl₃, CuCl₂, MnCl₂, OC@Fe³⁺,
 219 OC@Cu²⁺ and OC@Mn²⁺ exposed to SO₂ under irradiation. Reaction conditions: irradiance
 220 of 7.40×10^{15} photons cm⁻² s⁻¹, 2 ppm SO₂, 298 K and 60% RH. (b) Mass of sulfates
 221 produced on OC, TMIs, OC+TMIs and OC@TMIs after the 10 h reaction with SO₂ under
 222 irradiation (OC+TMIs denoted the arithmetic sum of the mass of sulfates measured for OC
 223 and TMIs in the separate experiments). Reaction conditions: irradiance of 1.05×10^{16}
 224 photons cm⁻² s⁻¹, 200 ppb SO₂, 298 K and 60% RH.



225 3.3 Spectroscopic evidence for complexation between OC and TMIs

226 The complexation between OC and TMIs may occur and alter the optical properties of OC
227 (Wang et al., 2024). Changes in optical absorption and steady-state fluorescence signals were
228 used to probe these interactions. Fig. S7a–c display the UV-vis spectra of OC, TMIs, OC+TMIs
229 and OC@TMIs. The absorbance for OC+TMIs was defined as the arithmetic sum of the
230 absorbance measured separately for OC and TMIs. OC and OC@TMIs exhibited a sharp
231 increase in the absorbance at the shorter wavelengths. FeCl_3 , CuCl_2 and MnCl_2 owned very
232 weak absorbance. OC@Fe^{3+} , OC@Cu^{2+} and OC@Mn^{2+} presented lower absorbance than
233 OC+Fe^{3+} , OC+Cu^{2+} and OC+Mn^{2+} , respectively, which can be attributed to the complexation
234 between OC and TMIs (Li et al., 2022). The FT-IR spectra of OC are displayed in Fig. S8, and
235 the peak assignments are summarized in Table S1. OC contained carboxyl, hydroxyl and
236 carbonyl groups. These oxygen-containing groups with π bonds electrons and lone pairs can
237 served as the complexation sites by donating electrons to the vacant orbitals of TMIs (Wang et
238 al., 2021c). Notably, as shown in Fig. S7d, there was the largest absorbance difference (ΔAbs)
239 between OC@Fe^{3+} and OC+Fe^{3+} in the entire spectral range, indicating the strongest
240 complexation of OC with Fe^{3+} among three TMIs.

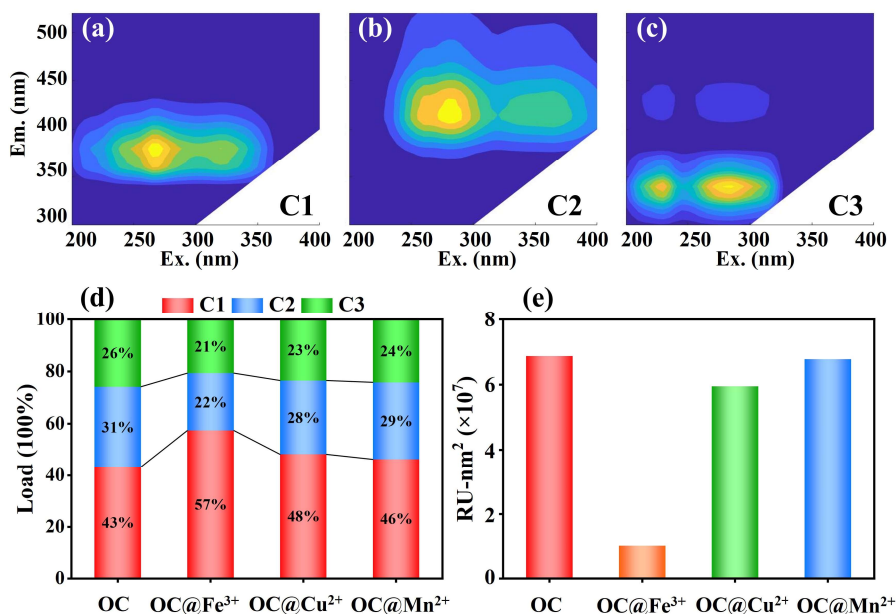
241 EEM fluorescence spectroscopy combining with the PARAFAC analysis were employed to
242 further investigate the complexation interactions between TMIs on OC. Three chromophores
243 are identified in Fig. 3a–c and Fig. S9, and their peak positions and assignments are given in
244 Table S2. C1 (Ex/Em = 265(320)/380 nm) and C2 (Ex/Em = 285(370)/420 nm) were assigned
245 to chromophores with low and high oxidation states, respectively. C3 (Ex/Em = 200(260)/340
246 nm) belonged to protein-like chromophores (Li et al., 2022; Wang et al., 2021b). Fluorophores
247 at longer emission wavelengths, such as C2, typically owned abundant polar groups, including
248 hydroxyl and carboxyl, which acted as primary binding sites for TMIs (Pan et al., 2020; Lu et
249 al., 2019). The relative contributions of different components to OC shifted upon the addition
250 of TMIs, reflecting the varied complexation strengths of TMIs with OC (Fig. 3d). For
251 OC@Fe^{3+} in comparison with OC, the ratio of C1 obviously increased by 14%, while it
252 decreased by 9% and 5% for C2 and C3, respectively. By contrast, OC@Cu^{2+} and OC@Mn^{2+}
253 showed smaller changes in the C1–C3 proportions. These mean that Fe^{3+} forms the most stable
254 complexes with OC due to its high charge density and trivalent oxidation state, which would



255 enhance the ligand attraction and enable higher coordination numbers (Liu et al., 2022). Cu^{2+}
256 and Mn^{2+} , with lower charge densities and less favorable coordination geometries, exerted
257 weaker complexation effects with OC.

258 As a comprehensive measure of fluorescence intensity, the total fluorescence volume (TFV)
259 was used to evaluate the overall impact of metal complexations (Fig. 3e). The addition of TMIs
260 led to a reduction in TFV, confirming the fluorescence quenching occurrence. Fe^{3+} caused the
261 most pronounced reduction of TFV (85.2%), followed by Cu^{2+} (13.4%) and Mn^{2+} (1.4%). This
262 strong quenching effect was attributed to the high binding affinity of Fe^{3+} with highly oxidative
263 and protein-like fluorophores. The complexation between OC and TMIs can disrupt the $\pi-\pi^*$
264 conjugation and deactivate fluorescence centers, which contributed to a significant reduction
265 in TFV (Kuramochi et al., 2018). In addition, the complexation of Fe^{3+} with OC could promote
266 the intersystem crossing (ISC) from the singlet state ($^1\text{OC}^*$) to the triplet state ($^3\text{OC}^*$) (Ruzi et
267 al., 2017; Treacy and Rovis, 2024), enhancing nonradiative decay and further suppressing the
268 fluorescence emission.

269 To directly demonstrate the complexation strength between TMIs with species in OC, DFT
270 calculations were conducted using phenol (PH), benzoic acid (BA) and a model substance (MS)
271 (Fig. S10). PH and BA were chosen according to FT-IR analysis (Fig. S8), as they represented
272 species with hydroxyl and carboxyl groups commonly presenting in OC. MS simultaneously
273 contained these functional groups with a π -conjugated aromatic backbone to better reflect the
274 structural features of ambient OC. Text S10 and Fig. S11 illustrate the complexation structure
275 of three molecular models with Fe^{3+} , Cu^{2+} and Mn^{2+} . As summarized in Table S3, all molecular
276 models exhibited the highest complexation energy with Fe^{3+} , followed by Cu^{2+} and Mn^{2+} . This
277 was well consistent with the results observed in the UV-Vis and fluorescence spectra.

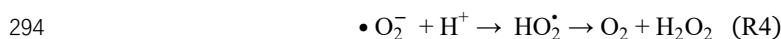


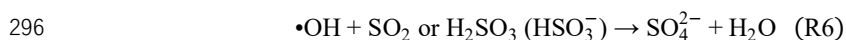
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279 Figure 3. (a–c) Three components extracted by the PARAFAC analysis. (d) Relative
280 fluorescence contributions of C1–C3 in OC, OC@Fe³⁺, OC@Cu²⁺ and OC@Mn²⁺. (e) Total
281 fluorescence volume of OC, OC@Fe³⁺, OC@Cu²⁺ and OC@Mn²⁺.

282 3.4 Mechanistic insights into TMIs-regulated oxidation of SO₂ to sulfates

283 As described through R1–6, previous studies have proposed the photochemical conversion
284 pathways of SO₂ to sulfates on OC (Yang et al., 2025; Yang et al., 2024a). Upon accepting
285 photons, OC is excited to ³OC*, which subsequently generates free electrons (R1–2) (Yang et
286 al., 2024a; Wang et al., 2020). These electrons would reduce adsorbed O₂ to •O₂⁻ (R3) (Zhang
287 et al., 2022). •O₂⁻ can combine with protons to form hydroperoxyl radicals (HO₂•), which are
288 converted into O₂ and hydrogen peroxides (H₂O₂) (R4) (Zhang et al., 2020c). H₂O₂ undergoes
289 the photolysis to produce •OH (R5), and •OH drives the oxidation of adsorbed SO₂ or H₂SO₃
290 or HSO₃⁻ to sulfates (R6) (Liu et al., 2020b).





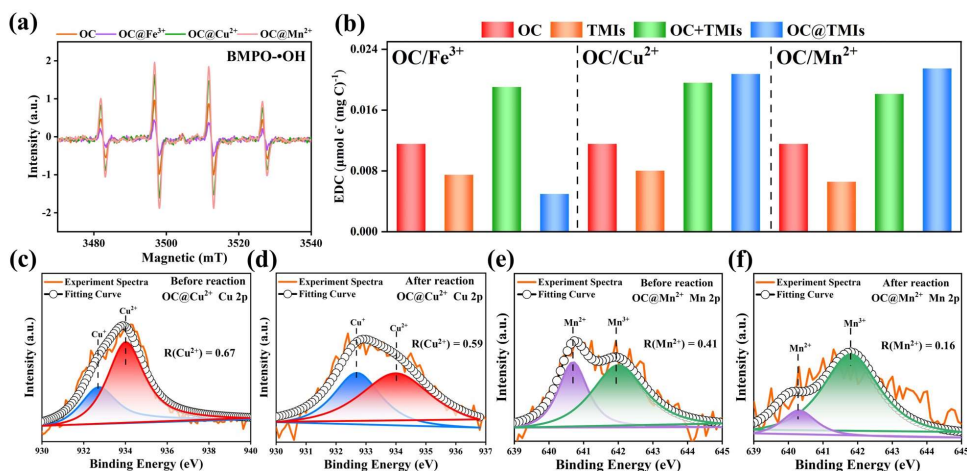
297 To verify whether the above mechanism remained operative after the complexation between
298 TMIs and OC, electrochemical and EPR analyses were conducted. Mediated electrochemical
299 oxidation (MEO) data of OC without or with TMIs showed persistent occurrence of the photo-
300 generated electron transfer (Fig. S12 and S13). BMPO was used as the spin-trapping agent. As
301 displayed in Fig. 4a, the characteristic EPR signal of BMPO-•OH adducts was detected under
302 irradiation, confirming the generation of •OH. Moreover, EPR signals of BMPO-•OH adducts
303 followed the order: OC@Mn²⁺ > OC@Cu²⁺ > OC > OC@Fe³⁺, which well aligned with γ_{ss} and
304 sulfate production. The critical role of •OH in the conversion of SO₂ to sulfates was further
305 determined using NaHCO₃ to scavenge •OH (Yang et al., 2024a), where the sulfate formation
306 was markedly restricted (Fig. S14). These demonstrate that •OH generation capacity is the
307 primary determinant of sulfate formation on OC@TMIs.

308 According to the reactions R1–2, the photogeneration of electrons was an initial step that
309 can affect the reaction activity of OC. Fig. 4b summarizes the distinct regulatory roles of Fe³⁺,
310 Cu²⁺ and Mn²⁺ in the EDC of OC under irradiation. The EDC of OC@Fe³⁺ (0.00496 $\mu\text{mol e}^-$
311 (mg C)⁻¹) was substantially less than that of OC+Fe³⁺ (0.01904 $\mu\text{mol e}^-$ (mg C)⁻¹), suggesting
312 that Fe³⁺ significantly suppressed the photogeneration of electrons. This mainly stemmed from
313 the strong complexation of Fe³⁺ and OC, which masked the redox-active moieties and
314 dynamically quenched ³OC*, restraining free electron formation (Li et al., 2022; Liu et al.,
315 2024).

316 Fig. 4b shows that EDC of OC@Cu²⁺ (0.02074 $\mu\text{mol e}^-$ (mg C)⁻¹) and OC@Mn²⁺ (0.02147
317 $\mu\text{mol e}^-$ (mg C)⁻¹) was higher than that of OC+Cu²⁺ (0.01954 $\mu\text{mol e}^-$ (mg C)⁻¹) and OC+Mn²⁺
318 (0.01811 $\mu\text{mol e}^-$ (mg C)⁻¹), respectively. This means that Cu²⁺ and Mn²⁺ enhance the electron-
319 donating ability of OC. These larger EDC may be related to reversible redox cycling of Cu²⁺/Cu⁺
320 and Mn²⁺/Mn³⁺ in the photochemical process. In contrast, Fe³⁺ formed more stable complexes
321 with OC that trapped the photogenerated electrons, thereby suppressing efficient redox cycling
322 of Fe³⁺/Fe²⁺. Under irradiation, Cu²⁺ was reduced to Cu⁺ by accepting photogenerated electrons,
323 which was subsequently re-oxidized to Cu²⁺ by O₂ or electron-deficient functional groups (e.g.,



324 quinones) (Pan et al., 2020). As displayed in XPS spectra of Cu 2p for OC@Cu²⁺ in Fig. 5a
 325 and b, the peaks at 932.5 and 934.3 eV corresponded to Cu⁺ and Cu²⁺, respectively (Biesinger
 326 et al., 2011). The proportion of Cu²⁺, denoted as R(Cu²⁺), decreased from 0.67 to 0.59 after
 327 photochemical SO₂ aging. This confirms the existence of reversible redox cycling of Cu²⁺/Cu⁺.
 328 Similarly, Mn²⁺ can be oxidized to Mn³⁺ by photogenerated holes (OC⁺) or reactive oxygen-
 329 containing intermediates on OC, and then Mn³⁺ was reduced by phenolic or aromatic groups
 330 on OC (Hansard et al., 2011). The Mn 2p_{3/2} spectra of OC@Mn²⁺ exhibited two peaks at 640.4
 331 and 641.9 eV (Fig. 5c and d), which were ascribed to Mn²⁺ and Mn³⁺, respectively (Xing et al.,
 332 2025). The proportion of Mn²⁺ (R(Mn²⁺)) decreased from 0.41 to 0.16 after the reaction,
 333 indicating Mn²⁺/Mn³⁺ redox interconversion. These reversible redox cycles facilitated dynamic
 334 interfacial electron transfer by accepting and releasing photogenerated electrons. This
 335 bidirectional electron flow can amplify the outward electron flux from OC, thereby increasing
 336 the electron-donating capacity (Fulda et al., 2013; Li et al., 2021).



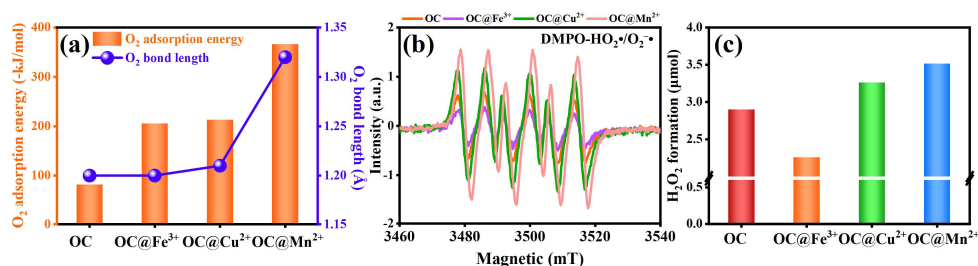
337
 338 Figure 4. (a) BMPO-•OH adduct EPR spectra of OC, OC@Fe³⁺, OC@Cu²⁺ and OC@Mn²⁺
 339 under irradiation. (b) Electron donating capacity (EDC) of OC, TMIs, OC+TMIs and
 340 OC@TMIs (OC+TMIs denoted the arithmetic sum of the EDC measured for OC and TMIs in
 341 separate experiments). High-resolution XPS spectra of Cu 2p for OC@Cu²⁺ (c) before and
 342 (d) after the reaction, and Mn 2p for OC@Mn²⁺ (e) before and (f) after the reaction (R(Cu²⁺)
 343 and R(Mn²⁺) denoted the proportion of Cu²⁺ and Mn²⁺ on the surface, respectively).

344 As shown in the reaction R3, •O₂⁻ formation relied not only on the photogenerated electrons



345 but also on the O₂ activation. To elucidate the effect of TMIs complexation on the O₂ activation,
 346 DFT calculations were carried out in the interaction of O₂ with MS (Fig. S15). Fig. 5a displays
 347 that O₂ adsorption energy on MS was $-82.03 \text{ kJ mol}^{-1}$, with O–O bond length of 1.20 Å. For
 348 MS@Fe³⁺, O₂ adsorption energy became more negative ($-205.59 \text{ kJ mol}^{-1}$), reflecting stronger
 349 electrostatic binding due to high positive charge of metal center (Zhang et al., 2024). However,
 350 O–O bond length remained unchanged, suggesting negligible electron transfer into the π^*
 351 antibonding orbital of O₂. Thus, O₂ is strongly bound yet not appreciably activated.
 352 Complexation between OC with Cu²⁺ or Mn²⁺ obviously enhanced the adsorption of O₂
 353 ($-212.93 \text{ kJ mol}^{-1}$ for Cu²⁺ and $-366.32 \text{ kJ mol}^{-1}$ for Mn²⁺), but also elongated the O–O bond
 354 length to 1.21 Å and 1.32 Å, respectively. These changes suggest electron transfer from the
 355 metal centers to the π^* antibonding orbital of O₂, boosting effective activation of O₂ and
 356 facilitating the $\bullet\text{O}_2^-$ generation. Compared to Cu²⁺, Mn²⁺ exhibited stronger O₂ activation,
 357 which was attributed to its half-filled 3d⁵ configuration and moderate redox potential, both of
 358 which favored efficient π^* electron donation.

359 Accordingly, Fe³⁺ inhibited electron generation and failed to activate O₂, whereas Cu²⁺ and
 360 Mn²⁺ promoted these two processes, with Mn²⁺ showing stronger enhancement effect. This led
 361 to the $\bullet\text{O}_2^-$ signal intensity of OC@Mn²⁺ > OC@Cu²⁺ > OC > OC@Fe³⁺ (Fig. 5b). $\bullet\text{O}_2^-$ directly
 362 contributed to H₂O₂ formation (R4), which was quantified with a TiOSO₄ chromogenic method
 363 (Fig. S16). Fig. 5c shows that H₂O₂ production trend had the consistency with $\bullet\text{OH}$ generation
 364 ability among these samples (Fig. 4a), which ultimately governed the photoconversion of SO₂
 365 to sulfates.



366 Figure 5. (a) O₂ adsorption energy and the bond length of O₂ for MS, MS@Fe³⁺, MS@Cu²⁺
 367 and MS@Mn²⁺. (b) EPR spectra of DMPO-HO₂•/•O₂⁻ adduct for OC, OC@Fe³⁺, OC@Cu²⁺
 368 and OC@Mn²⁺ under irradiation. (c) H₂O₂ formation from the photochemical reactions on
 369



370 OC, OC@Fe³⁺, OC@Cu²⁺ and OC@Mn²⁺.

371 **4 Conclusions and atmospheric implications**

372 It was well confirmed that the heterogeneous photooxidation of SO₂ to sulfates was strongly
373 modulated by the interactions of OC with TMIs. As evidenced by SO₂ uptake coefficient and
374 sulfate formation mass, Fe³⁺ exhibited a significantly inhibitory effect, whereas Cu²⁺ and Mn²⁺
375 promoted the conversion of SO₂ to sulfates, with Mn²⁺ playing a more obvious enhancement
376 role. Spectroscopic analyses and DFT calculations consistently demonstrated that Fe³⁺ had the
377 strongest binding affinity to OC chromophores, followed by Cu²⁺ and Mn²⁺. This difference in
378 the complexation strength controlled the electron generation, O₂ activation, and the sequential
379 formation of •O₂⁻, H₂O₂, and •OH. The regulation of sulfate formation on OC by TMIs was
380 largely governed by their capacity to suppress or enhance •OH generation.

381 The photochemical lifetime of SO₂ (τ_{SO₂}) on OC@TMIs was calculated with Eq. (2) (Yang
382 et al., 2021),

$$383 \tau_{\text{SO}_2} = \frac{4}{\gamma v A} \quad (2)$$

384 where γ represents the uptake coefficient of SO₂; v denotes the mean SO₂ molecular velocity;
385 A corresponds to the aerosol surface area concentration (6.25 × 10⁻⁵ cm² cm⁻³) (Liu et al.,
386 2020b; Yang et al., 2021). Under typical atmospheric conditions (40 ppb SO₂, 60% RH), γ_{ss} for
387 SO₂ on OC@TMIs were determined to be (0.39–22.6) × 10⁻⁶. If organic matters contributed
388 to 20% of aerosol mass, τ_{SO₂} was estimated to be 5.2–330 days in the coexistence case of OC
389 and TMIs. By contrast, τ_{SO₂} on OC from coal combustion was much narrower, which ranged
390 from 17 to 53 days under the same conditions (Yang et al., 2025).

391 To assess the atmospheric production of sulfates during the photooxidation of SO₂ on
392 OC@TMIs, the sulfate formation rate (R) was estimated using Eq. (3),

$$393 R = \frac{d[\text{SO}_4^{2-}]}{dt} = \left[\frac{R_p}{D_g} + \frac{4}{\gamma v} \right]^{-1} A [\text{SO}_2] \quad (3)$$

394 where D_g refers to the diffusion coefficient of SO₂ (1.12 × 10⁻⁵ m² s⁻¹); R_p denotes the aerosol
395 particle radius and is estimated via Eq. (4) (Li et al., 2020),

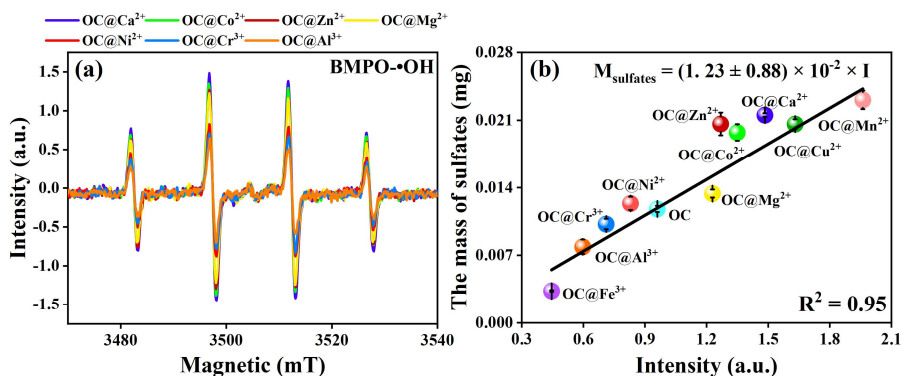
$$396 R_p = \left(0.254 \times \frac{[\text{PM}_{2.5}]}{(\mu\text{g m}^{-3})} + 10.259 \right) \times 10^{-9} \quad (4)$$

397 where [PM_{2.5}] denotes the mean mass concentration of PM_{2.5}, and a representative



398 concentration of $300 \mu\text{g m}^{-3}$ for the polluted episodes in typical Chinese cities was adopted
399 here (Li et al., 2020). Under these conditions, R was determined to be $0.08\text{--}4.56 \mu\text{g m}^{-3} \text{h}^{-1}$,
400 which was a substantially broader range than $0.43\text{--}1.33 \mu\text{g m}^{-3} \text{h}^{-1}$ for OC (Yang et al., 2025).
401 Broader ranges of τ_{SO_2} and R on OC@TMIs indicate that the coexistence of OC and TMIs
402 significantly modulates the photochemical conversion of SO_2 to sulfates. These findings
403 suggest that current model simulations may underestimate or overestimate sulfate formation
404 on OC by overlooking such antagonistic or synergistic processes.

405 It has been well demonstrated that $\bullet\text{OH}$ was the dominant active species responsible for the
406 oxidation of SO_2 to sulfates on OC@TMIs (Fig. 4 and S14). Other metal ions (Ca^{2+} , Co^{2+} , Zn^{2+} ,
407 Mg^{2+} , Ni^{2+} , Cr^{3+} and Al^{3+}) commonly presenting in atmospheric aerosols were expanded to
408 examine whether this mechanism exhibited broader generality (Hua et al., 2024). As shown in
409 Fig. 6a, the peaks of BMPO- $\bullet\text{OH}$ adducts were observed under irradiation, confirming the $\bullet\text{OH}$
410 generation on all OC@TMIs. Relative to OC, the coexistence of Ca^{2+} , Co^{2+} , Zn^{2+} and Mg^{2+}
411 obviously promoted the $\bullet\text{OH}$ generation, while Ni^{2+} , Cr^{3+} and Al^{3+} exerted an inhibitory effect.
412 Fig. S17 summarized the mass of sulfates formed on OC@TMIs, which was consistent with
413 the $\bullet\text{OH}$ generation trend. Importantly, a linearly positive relationship ($R^2 = 0.95$) between $\bullet\text{OH}$
414 intensity and sulfate mass (M_{sulfates}) was determined in Fig. 6b. M_{sulfates} can be parametrized as
415 a function of $\bullet\text{OH}$ signal intensity (I) with an Eq.: $M_{\text{sulfates}} = (1.23 \pm 0.88) \times 10^{-2} \times I$. This
416 quantitative relationship highlighted the potential dual roles of metal ion types in the sulfate
417 formation, challenging the traditional view that metal ions often promote the SO_2 oxidation
418 (Cao et al., 2024a; Wang et al., 2022; Wang et al., 2021a). It also emphasized the critical role
419 of $\bullet\text{OH}$ in the sulfate formation on OC@TMIs. Accordingly, incorporating the $\bullet\text{OH}$ generation
420 ability of diverse OC@TMIs complexes into the atmospheric models is essential to accurately
421 constrain sulfate source strength, predict haze evolution and assess associated health impacts
422 in polluted environments.



423

424

Figure 6. (a) BMPO-•OH adduct EPR spectra. (b) The correlation of M_{sulfates} with the

425

generation of •OH.

426

Data availability. The data used in this study are available from the corresponding author upon

427

request (hanch@smm.neu.edu.cn).

428

Supplement. Figures, texts and tables about experimental section, background experiment of

429

SO₂ uptake, complexation evidences of OC with metal ions, the mechanism verification and

430

atmospheric implication.

431

Author contributions. CH and SY designed the experiment; SY and SL conducted the

432

experiments; SY, SL, JZ, HN, FL and WY performed the data interpretation; CH and SY wrote

433

the paper. CH supervised the project. All authors contributed to the paper with useful scientific.

434

Competing interests. The contact author has declared that none of the authors has any

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competing interests.

436

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