



1 **Dissimilar Roles of Aerosols, Nitrogen Deposition and Ozone on the Terres-**
2 **trial Carbon Sink in China during 2010-2020**

3 Nanhong Xie¹, Tijian Wang^{1*}, Shu Li¹, Bingliang Zhuang¹, Mengmeng Li¹, Min Xie²,
4 Qian Zhang¹, Danyang Ma², Jane Liu³, Jing M. Chen³, Zhaozhong Feng⁴, Dimitrios Me-
5 las⁵, Kostas Karatzas⁶

6 ¹School of Atmospheric Sciences, Nanjing University, Nanjing, China

7 ²School of Environment, Nanjing Normal University, Nanjing, China

8 ³Department of Geography and Planning, University of Toronto, Toronto, Ontario, Canada

9 ⁴Key Laboratory of Ecosystem Carbon Source and Sink, China Meteorological Administra-
10 tion (ECSS-CMA), School of Ecology and Applied Meteorology, Nanjing University of In-
11 formation Science and Technology, Nanjing, China

12 ⁵Laboratory of Atmospheric Physics, School of Physics, Aristotle University of Thessaloniki,
13 Thessaloniki, Greece

14 ⁶Environmental Informatics Research Group, School of Mechanical Engineering, Aristotle
15 University of Thessaloniki, Thessaloniki, Greece

16

17 *Corresponding to:* Tijian Wang (tjwang@nju.edu.cn)

18

19 **Abstract**

20 China's Clean Air Action (CAA) plan implemented since 2013 has significantly altered atmos-
21 pheric composition, and yet its impact on the terrestrial carbon sink remains unclear. This study
22 employed the Regional Earth System Model (RegESM), an online-coupled climate–chemistry–
23 ecosystem modeling framework, to quantify the impacts of aerosols, surface ozone (O₃), and
24 nitrogen deposition on China's net ecosystem productivity (NEP) from 2010 to 2020. The re-
25 sults show that aerosols enhanced China's NEP by 17.93 TgC yr⁻¹ (4.49% of the total NEP),
26 primarily by increasing diffuse radiation, with the most pronounced effects in Southern and
27 Eastern China. Nitrogen deposition further increased NEP by 37.98 TgC yr⁻¹ (9.52%),



concentrated in Central and Southern regions. In contrast, O_3 pollution reduced NEP by 51.33 $TgC\ yr^{-1}$ (12.9%), particularly in the forest-dominated Southeast. The positive impacts of aerosols and nitrogen deposition on the carbon sink weakened over time, whereas the negative influence of O_3 was increasing. The combined effects indicate that CAA-induced atmospheric chemistry changes reversed the dominant atmospheric drivers of China's terrestrial carbon sink, from enhancement by aerosols and nitrogen deposition to suppression by ozone. Our findings highlight the need for stronger O_3 pollution control to achieve co-benefits between air-quality improvement and carbon neutrality.

1 Introduction

Terrestrial ecosystems act as major carbon sinks, sequestering atmospheric carbon dioxide (CO_2) through plant photosynthesis, and constitute a fundamental natural process for mitigating global climate change (Friedlingstein et al., 2023; Piao et al., 2013; Yuan et al., 2025). Under ongoing global warming, the dynamics of carbon sinks are regulated not only by climatic factors such as temperature and precipitation (Cao et al., 2023; Post et al., 2018; Ren et al., 2020), but also by variations in atmospheric composition (Zhou et al., 2021). Among these, aerosols, O_3 , and atmospheric nitrogen deposition have been identified as key atmospheric pollutants affecting terrestrial carbon sequestration (Liu et al., 2022; Zhou et al., 2024). As a crucial component of the global carbon cycle, terrestrial ecosystems in China sequester approximately 0.20–0.25 $PgC\ yr^{-1}$, playing an essential role in supporting the achievement of the national carbon neutrality target (Piao et al., 2022; Xia et al., 2025; Yue et al., 2021). Therefore, assessing the responses of carbon sinks to multiple atmospheric composition changes is of great scientific significance for understanding both the global carbon cycle and climate feedback mechanisms.

Aerosols influence vegetation photosynthesis and carbon sequestration primarily through radiative forcing (Shu et al., 2022; Zhou et al., 2022). Aerosol scattering and absorption reduce surface solar radiation and can suppress vegetation photosynthesis (Doughty et al., 2010; Kuniyal and Guleria, 2019). In the meantime, enhanced diffuse radiation increases light use efficiency of plants, leading to the diffuse fertilization effect (Gu et al., 2003; Mercado et al., 2009). Aerosols also influence cloud microphysics by modifying droplet formation and lifetime,



56 which further affects regional precipitation and water availability for vegetation (Li et al., 2020;
 57 Unger et al., 2017). Consequently, the net effect of aerosols on photosynthesis exhibits marked
 58 spatial heterogeneity, with both enhancement and suppression reported in highly polluted re-
 59 gions such as eastern China (Strada and Unger, 2016; Wang et al., 2018; Xie et al., 2020).

60 In addition, near-surface O₃ impairs plant carbon uptake through direct physiological
 61 damage (Lei et al., 2022; Unger et al., 2020). O₃ enters leaves through stomata and induces
 62 reactive oxygen species at the cellular level, leading to degradation of photosynthetic pigments,
 63 suppressed Rubisco activity, premature leaf senescence, and defoliation, all of which inhibit
 64 photosynthetic carbon assimilation (Wittig et al., 2007). Evidence from O₃-FACE (free-air O₃
 65 concentration enrichment) experiments shows that a 10 ppb increase in O₃ concentration can
 66 lower crop productivity by 5–15% (Feng et al., 2015). In China, summertime O₃ peaks often
 67 coincide with the peak growing season of vegetation, particularly in the North China Plain and
 68 the Yangtze River Delta, posing a notable threat to regional carbon sequestration (Lei et al.,
 69 2022; Li et al., 2024; Yue et al., 2017).

70 Furthermore, atmospheric nitrogen deposition is a major external nitrogen source for ter-
 71 restrial ecosystems and exerts both positive and negative effects on carbon sinks (Chen et al.,
 72 2015; Lu et al., 2021). In nitrogen-limited systems, such as temperate forests and grasslands,
 73 moderate deposition can enhance photosynthesis and biomass accumulation, thereby increas-
 74 ing carbon sequestration (Cen et al., 2025; Lu et al., 2016; Peng et al., 2025). When inputs
 75 become excessive, however, they can induce soil acidification, biodiversity loss, and broader
 76 ecosystem degradation, a condition known as nitrogen saturation (Chen et al., 2015; Yue et al.,
 77 2016). China receives some of the highest nitrogen deposition levels globally, with annual av-
 78 erages of 15–20 kg N ha⁻¹ yr⁻¹ and hotspots surpassing 30 kg N ha⁻¹ yr⁻¹, raising increasing
 79 concerns about long-term ecological impacts (Liu et al., 2022; Liu et al., 2013; Yu et al., 2019).

80 China has long faced the dual pressures of severe air pollution and growing greenhouse
 81 gas emissions (Tu et al., 2019; Wang et al., 2024). Rapid economic expansion in the early 2010s
 82 was accompanied by persistent increases in fine particulate matter (PM_{2.5}) concentrations (Hao
 83 et al., 2020). Since 2013, successive Clean Air Action Plans have led to a substantial decline in
 84 PM_{2.5} levels (Xue et al., 2019; Yue et al., 2020; Zheng et al., 2018). At the same time, near-



85 surface summertime O_3 has risen sharply (Liu et al., 2018; Zhou et al., 2024), while nitrogen
86 deposition has slowed in growth but remains at a high level (Liu et al., 2024). These changes
87 not only reflect the outcomes of emission control policies but also reshape the regional atmos-
88 pheric chemical environment, potentially exerting complex and combined effects on carbon
89 sinks (Liu et al., 2022; Zhou et al., 2024). The rapid transition in atmospheric composition
90 during 2010–2020 provides an unprecedented large-scale natural experiment for disentangling
91 the relative roles of aerosols, O_3 , and nitrogen deposition in altering China’s carbon sink. How-
92 ever, most existing studies have examined these drivers in isolation, relied on offline or statis-
93 tical frameworks that cannot capture dynamic climate–chemistry–ecosystem feedbacks, and
94 rarely compared responses across ecological regions (Unger et al., 2020; Yue et al., 2017; Zhou
95 et al., 2024).

96 Here, we employ an improved regional climate–chemistry–ecosystem online-coupling
97 model, RegESM (Xie et al., 2024; Zhang et al., 2025), to quantify the impacts of aerosols, O_3 ,
98 and nitrogen deposition on China’s terrestrial carbon sinks during 2010–2020. RegESM incor-
99 porates two-way interactions among climate, atmospheric chemistry, and biogeochemical pro-
100 cesses and has been extensively evaluated over East Asia (Ma et al., 2023; Xie et al., 2024; Xie
101 et al., 2020; Zhang et al., 2025). Our objective is to isolate the contributions of individual at-
102 mospheric components to changes in China’s carbon sinks based on RegESM after its assess-
103 ment using multiple observational datasets. These results offer new insight into the ecological
104 consequences of rapid atmospheric composition changes and provide a scientific foundation
105 for coordinated multi-pollutant control and ecosystem management under China’s carbon-neu-
106 trality goals.

107 **2 Data and Methods**

108 **2.1 The RegESM model**

109 In this study, we employed the RegESM, an improved extension of the RegCM-Chem-
110 YIBs regional climate–chemistry–ecosystem modeling framework (Xie et al., 2024; Xie et al.,
111 2019; Zhang et al., 2025). The original RegCM-Chem-YIBs couples the RegCM4 regional



112 climate model (Giorgi et al., 2012), the radiative interactive gas-phase chemistry module Chem
 113 (Shalaby et al., 2012), and the YIBs terrestrial ecosystem model (Yue and Unger, 2015) to
 114 represent interactive processes among atmospheric dynamics, chemistry, and terrestrial carbon
 115 cycles (Xie et al., 2024). Building upon this foundation, RegESM strengthens two-way feed-
 116 back among the atmosphere, atmospheric chemistry, and land surface processes, enabling a
 117 more realistic simulation of biogeochemical cycles (Zhang et al., 2025). The enhanced coupling
 118 allows land surface changes, such as vegetation dynamics and soil moisture variations, to more
 119 directly influence atmospheric composition, radiation, and meteorology, while atmospheric
 120 and chemical variations simultaneously affect ecosystem processes (Xie et al., 2024; Zhang et
 121 al., 2025). This bidirectional integration improves the model’s capability to capture transient
 122 and spatially heterogeneous climate–ecosystem–chemistry interactions, which are crucial for
 123 regional climate change and carbon budget assessments (Zhang et al., 2025).

124 The RegESM framework used in this study integrates RegCM4 as the dynamical core for
 125 simulating regional climate processes at a high resolution, the Chem module for interactive
 126 gas-phase and aerosol chemistry coupled with radiation and meteorology, and the YIBs land
 127 surface model for calculating biophysical processes such as photosynthesis, transpiration, and
 128 energy balance, along with biogeochemical cycles of carbon and nitrogen (Giorgi et al., 2012;
 129 Shalaby et al., 2012; Xie et al., 2024; Yue and Unger, 2015). These components are linked
 130 through an improved coupling mechanism that ensures the consistent exchange of meteorolog-
 131 ical, chemical, and biogeophysical variables at each model timestep, enabling fully interactive
 132 simulations in which land, atmosphere, and chemistry evolve in a physically coherent manner
 133 (Xie et al., 2024; Zhang et al., 2025). This model has been widely applied in East Asia (Xie et
 134 al., 2025; Xie et al., 2019; Zhang et al., 2025; Zhang et al., 2024).

135 We used net ecosystem productivity (NEP) as an indicator for characterizing carbon
 136 sources and sinks ($NEP > 0$ suggests a carbon sink). NEP was calculated as the difference
 137 between gross primary production (GPP) and the sum of autotrophic respiration (R_a) and het-
 138 erotrophic respiration (R_h) (Xie et al., 2025; Yue et al., 2021). It is noteworthy that the NEP
 139 estimated in this study does not account for lateral carbon transfers.

140 **2.2 Ozone Damage Scheme**



141 Once surface O_3 enters plants through the stomata, it directly damages plant cellular struc-
 142 tures and suppresses the photosynthetic rate, thereby reducing vegetation productivity. In the
 143 YIBs vegetation module of the RegESM model, a semi-mechanistic parameterization scheme
 144 is employed to represent the impacts of O_3 on plants (Sitch et al., 2007; Yue and Unger, 2015):

$$145 \quad B = B_{tot} \cdot K, \quad (1)$$

146 where B denotes the photosynthetic rate under O_3 exposure, B_{tot} represents the total leaf pho-
 147 tosynthetic rate, and K is the remaining proportion of photosynthetic capacity after O_3 stress.
 148 This proportion is determined by the stomatal O_3 flux that exceeds a specified threshold:

$$149 \quad K = 1 - b \cdot \max[(K_{ozn} - K_{ozncrit}), 0], \quad (2)$$

150 where b denotes the vegetation sensitivity parameter to O_3 derived from observational data.
 151 $K_{ozncrit}$ represents the threshold of O_3 -induced damage to vegetation, and K_{ozn} denotes the
 152 O_3 flux entering the leaf through stomata:

$$153 \quad K_{ozn} = \frac{[O_3]}{r_b + \frac{\kappa_{O_3}}{r_s}}, \quad (3)$$

154 where $[O_3]$ denotes the O_3 concentration at the canopy top, r_b is the boundary layer resistance,
 155 κ_{O_3} is the ratio of O_3 leaf resistance to water vapor blade resistance, and r_s is the stomatal
 156 resistance accounting for the effects of O_3 :

$$157 \quad r_s = g_s \cdot K. \quad (4)$$

158 g_s denotes the leaf conductance unaffected by O_3 exposure. By simultaneously solving Equa-
 159 tions (2), (3), and (4), a quadratic term with respect to K is obtained, which can be solved
 160 analytically.

161 2.3 Experimental design and input data

162 The simulation domain covers most of East Asia (Fig. S1), centered at 36° N and 107° E.
 163 The horizontal resolution is 30 km, with 18 vertical layers. To quantify the independent con-
 164 tributions of aerosol, O_3 damage, and atmospheric nitrogen deposition to China's terrestrial
 165 carbon sink during 2010–2020, four sensitivity experiments were conducted (Table 1): a base-
 166 line simulation without these effects (Base), and three single-factor cases that enabled only
 167 aerosol (Ctrl_AOD), O_3 -induced vegetation damage (Ctrl_ O_3), and nitrogen deposition



168 impacts (Ctrl_Ndep). The difference between each sensitivity case and the Base run represents
 169 the corresponding individual effect. All simulations were preceded by a one-year spin-up to
 170 reduce the influence of initial conditions. To further assess regional responses, China was di-
 171 vided into six representative subregions (Fig. S2), and statistical analyses were performed for
 172 each.

173

174 **Table 1.** Numerical model experiments.

| Simulations | Periods | Aerosol radiative effect | O ₃ damage | atmospheric nitrogen deposition |
|---------------------|-----------|-----------------------------|-----------------------|---------------------------------------|
| Base | 2010-2020 | off | off | off |
| Ctrl_AOD | 2010-2020 | open | off | off |
| Ctrl_O ₃ | 2010-2020 | off | open | off |
| Ctrl_Ndep | 2010-2020 | off | off | open |

175

176 The initial and boundary meteorological fields were taken from the ECMWF (European
 177 Centre for Medium-Range Weather Forecasts) ERA-Interim reanalysis with a temporal reso-
 178 lution of 6 h and a horizontal resolution of $1.5^\circ \times 1.5^\circ$ (Hersbach et al., 2020). Aerosol initial
 179 and boundary conditions were provided by the global chemical transport model (MOZART)
 180 (Emmons et al., 2010; Horowitz et al., 2003). Background CO₂ fields were constrained by
 181 three-dimensional concentrations from NOAA CarbonTracker (CT) reanalysis (Peters et al.,
 182 2007). The initial parameters for the YIBs model were derived from soil carbon stocks based
 183 on equilibrium tree height and a 30-year harvest cycle (Yue and Unger, 2015). Vegetation
 184 cover was prescribed from MODIS and AVHRR (Advanced Very High Resolution Radiometer)
 185 datasets (Lawrence and Chase, 2007). Anthropogenic emissions in China were taken from the
 186 Multi-resolution Emission Inventory for China (MEIC) (Geng et al., 2024; Li et al., 2017;
 187 Zheng et al., 2018).

188 2.4 Validation data



189 We employed monthly mean aerosol optical depth (AOD) data from the MODIS sensor
190 onboard NASA's Terra satellite (MOD08_M3.061). The data have a spatial resolution of $1^\circ \times$
191 1° and are retrieved using three algorithms: the Dark Target, Deep Blue, and combined ap-
192 proaches (Levy et al., 2013). Ground-level O_3 observations were obtained from 366 monitoring
193 stations operated by the China National Environmental Monitoring Center (CNEMC). To eval-
194 uate the model's capability in simulating atmospheric nitrogen deposition, we employed pub-
195 licly available datasets (Liu et al., 2024; Zhu et al., 2025). These datasets integrate observations
196 with model outputs to provide nitrogen deposition estimates at both global and regional scales
197 over China. To assess the reliability of simulated CO_2 , we used observations from the World
198 Data Centre for Greenhouse Gases (WDCGG). This dataset provides measured surface atmos-
199 pheric CO_2 concentrations and was used to evaluate the model's ability to reproduce observed
200 CO_2 levels. For the spatial distribution of CO_2 , we additionally used CO_2 concentration fields
201 from CT (Peters et al., 2007). For GPP and net primary production (NPP) validation, we used
202 the global MODIS products MOD17A2H and MOD17A3H (Collection 6). The GPP data, at
203 8-day resolution, were derived using the radiation use efficiency algorithm, while NPP
204 ($NPP = GPP - R_a$) data were produced by annually accumulating GPP values, with a spatial res-
205 olution of 500 m (He et al., 2018; Madani et al., 2014).

206 2.5 Analytical Approach

207 Aerosol-induced meteorological changes are highly interdependent, making it challenging
208 to isolate their individual effects on terrestrial carbon cycling. To quantify the relative contri-
209 butions of these meteorological responses to vegetation carbon fluxes, we applied a multiple
210 linear regression framework. Standardized regression coefficients were used to assess the rel-
211 ative influence of each climate variable. This approach has been widely demonstrated as effec-
212 tive for disentangling the impacts of multiple environmental drivers on ecosystem processes
213 (Jung et al., 2017; Xie et al., 2025; Zhang et al., 2024).
214 The regression model is expressed as follows:

$$\begin{aligned} 215 \quad \Delta Y = & A_1 \times \Delta X_1^{RadD} + A_2 \times \Delta X_2^{RadF} + A_3 \times \Delta X_3^{Temp} + A_4 \times \Delta X_4^{Precip} + A_5 \times \Delta X_5^{VPD} \\ 216 \quad & + \varepsilon \quad (5) \end{aligned}$$



where ΔY denotes the difference in terrestrial carbon flux between the simulations Base and Ctrl_AOD, respectively. ΔX_1^{RadD} , ΔX_2^{RadF} , ΔX_3^{Temp} , ΔX_4^{Precip} and ΔX_5^{VPD} denote the differences in direct radiation, diffuse radiation, temperature, precipitation, and vapor pressure deficit (VPD) between the simulations Base and Ctrl_AOD, respectively. A_i represents the partial regression coefficient for different meteorological factors, indicating the sensitivity of carbon flux to variations in these factors. ε is the residual term of the regression model. We use the following equation to calculate the standardized regression coefficient B_i for comparing the relative impacts of different meteorological factors:

$$B_i = A_i \times SD(\Delta X_i) \div SD(\Delta Y) \quad (6)$$

where $SD(\Delta X_i)$ and $SD(\Delta Y)$ represent the standard deviations of the changes in each meteorological factor and carbon flux, respectively. B_i quantifies the relative contribution of different meteorological factors to variations in carbon flux. This approach enables a quantitative assessment of the individual impacts of changes in each meteorological factor induced by aerosol radiative effects on terrestrial carbon flux.

3 Results

3.1 Model validations

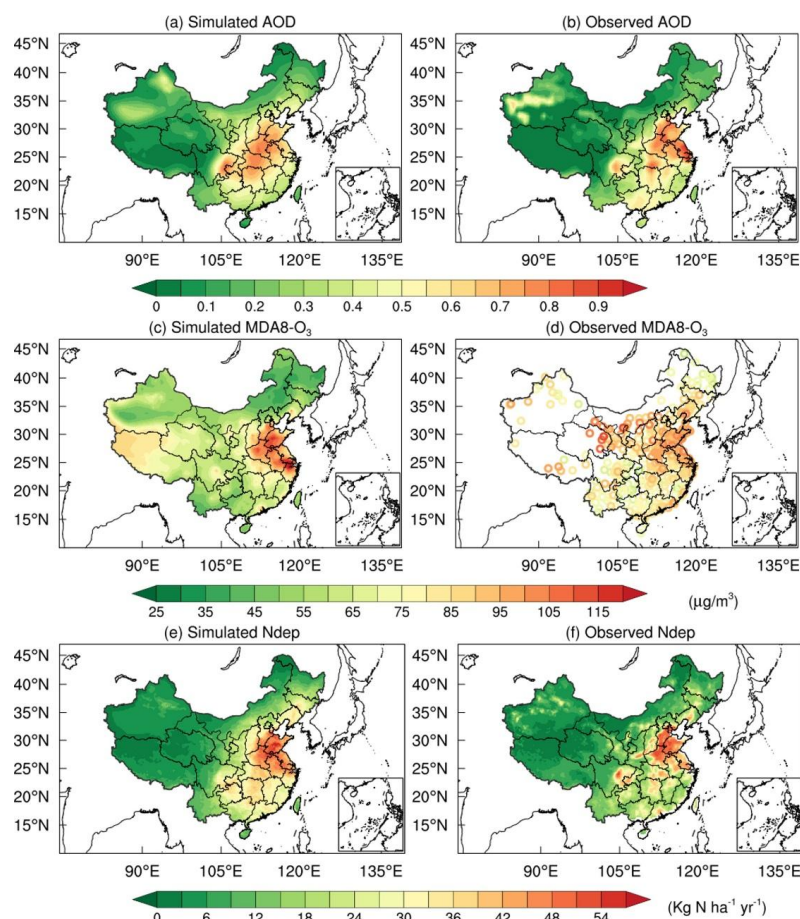
3.1.1 Aerosols, surface ozone, and atmospheric nitrogen deposition

We assessed the RegESM performance by comparing the 2010–2020 simulations with multi-source observations. Simulated AOD showed good agreement with MODIS products in both spatial distribution and magnitude (Fig. 1a, b). High AOD values are located over the North China Plain and the Sichuan Basin, consistent with dense anthropogenic emissions in these regions (Luo et al., 2014). The observations indicate that the national mean AOD decreased from 0.36 in 2010 to 0.28 in 2020, driven by air quality improvement policies. We calculated statistical metrics, including the correlation coefficient (R), mean bias (MB), and root mean square error (RMSE), to evaluate the model performance (Fig. S3). The RegESM captures this trend with a correlation coefficient (R) of 0.71. However, compared with monthly MODIS AOD, the model shows a minor underestimation (MB = −0.02), which can be



244 attributed primarily to uncertainties in the anthropogenic emission inventories (Xie et al., 2020).
245 Surface O₃ simulations reproduce both spatial patterns (Fig. 1c, d). The correlation with site
246 observations reaches 0.72 (Fig. S4). High concentrations in the North China Plain, the Yangtze
247 River Delta, and the Sichuan Basin are captured well, highlighting the model's skill in simu-
248 lating O₃ fields. The simulated annual mean atmospheric nitrogen deposition flux ranges from
249 20 to 40 kg N ha⁻¹ yr⁻¹ in eastern agricultural and urban areas, consistent with reported values
250 of 25–35 kg N ha⁻¹ yr⁻¹ (Fig. 1e, f). The simulated national mean of 15.09 kg N ha⁻¹ yr⁻¹ is close
251 to the dataset range of 13.45–15.39 kg N ha⁻¹ yr⁻¹. The model also reproduces the observed
252 decline after the implementation of air pollution control policies in 2013, with a gradual de-
253 crease after 2015 (Fig. S5). These evaluations indicate that RegESM reliably simulates AOD,
254 O₃, and nitrogen deposition fields across China.

255



256

257 **Figure 1.** Annual mean AOD (a, b), maximum daily 8 h average (MDA8) O₃ (c, d), and At-
 258 mospheric nitrogen deposition (e, f) from model simulation (a, c, e) and observations (b, d, f).

259

260 3.1.2 Atmospheric CO₂ concentrations, GPP, and NPP

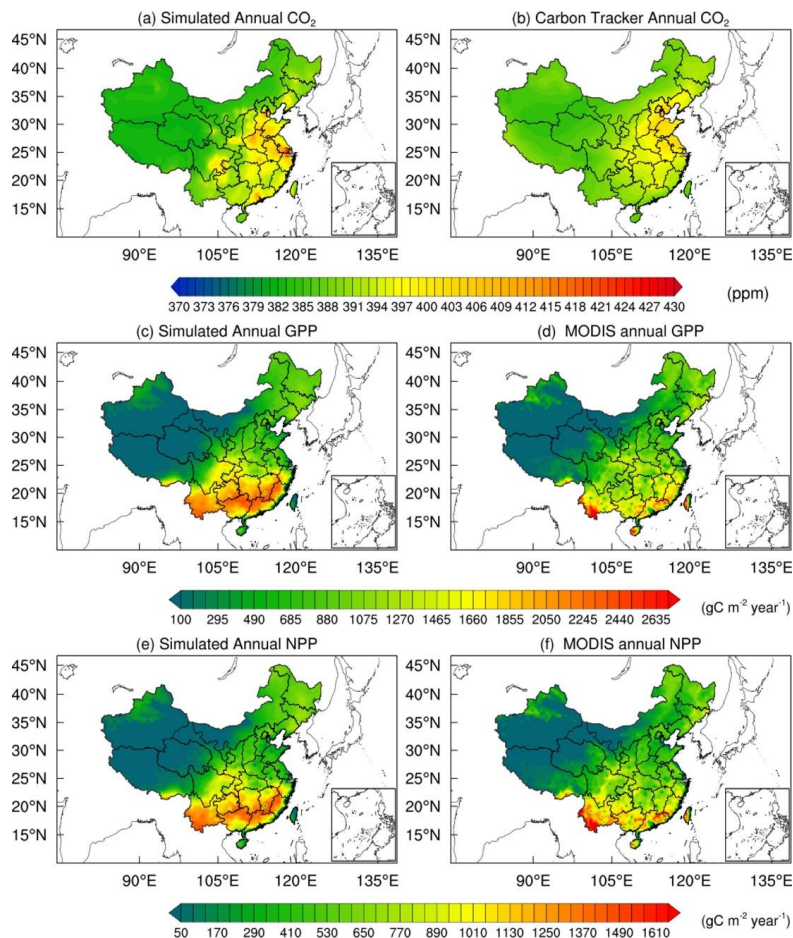
261 Simulated CO₂ concentrations were compared with six stations from the WDCGG. The
 262 correlation coefficients range from 0.83 to 0.96 (Table S1). The YON site shows the best agree-
 263 ment ($R = 0.96$, $MB = -1.1$ ppm), likely due to minimal influence from terrestrial emissions.
 264 In contrast, HKG and HKO show larger biases, with overestimates of 3.1 ppm ($R = 0.83$) and
 265 3.3 ppm ($R = 0.84$), probably linked to unaccounted variability in urban sources in monthly



266 inventories. Nevertheless, the seasonal cycle is reproduced well at all sites (Fig. S6).

267 We further compared simulated CO₂ with the CarbonTracker CT2022 assimilation dataset
268 (Peters et al., 2007). The spatial correlation coefficient reaches 0.72 (Fig. 2a, b). High CO₂
269 concentrations appear over the Beijing–Tianjin–Hebei region, the Yangtze River Delta, the
270 Pearl River Delta, and the Sichuan Basin, consistent with intense industrial emissions. The
271 model slightly overestimates values in the Pearl River Delta, likely due to underrepresented
272 local sources and complex topography. Overall, RegESM effectively captures the spatial dis-
273 tribution of CO₂ concentrations.

274 Simulated GPP agrees well with MODIS in spatial distribution (Fig. 2c, d), with a spatial
275 correlation of 0.89. However, GPP from this study is larger than MODIS GPP by 7.4%, with
276 largest differences in Central (11.6%) and Southeast China (5.7%). Other studies also found
277 that MODIS GPP was underestimated at high values (Xie et al., 2019; Zhang et al., 2012). The
278 southeast-to-northwest decreasing gradient is reproduced, with high values over regions dom-
279 inated by forest ecosystems. The seasonal cycle of GPP is also captured (Fig. S7). The simu-
280 lated NPP exhibits a spatial distribution consistent with the MODIS (Fig. 2e, f), with a spatial
281 correlation coefficient of 0.86. Similar to GPP, the model overestimates NPP by 8.4%, mainly
282 due to the overestimation in Central (14.3%) and Northeastern (6.2%) China. These results
283 confirm the model’s ability to represent terrestrial carbon fluxes.



284
285 **Figure 2.** Annual mean CO₂ (a, b), GPP (c, d), and NPP (e, f) from model simulation (a, c, e)
286 and observations (b, d, f).
287

288 3.2 Impacts of Aerosols on Meteorology and Carbon Sinks

289 3.2.1 Impacts of Aerosols on Meteorological Factors

290 During 2010–2020, the aerosol exerted a substantial influence on China’s surface radia-
291 tion and near-surface climate (Fig. 3). Nationally, aerosols reduced downward direct solar ra-
292 diation by 8.81 W m⁻², while increasing diffuse radiation by 3.04 W m⁻², leading to an overall
293 reduction of 5.77 W m⁻² in total shortwave radiation reaching the surface. Spatially, these



294 radiative changes were most pronounced over major urban agglomerations such as the North
 295 China Plain, the Yangtze River Delta, and the Sichuan Basin, coinciding with regions of high
 296 AOD associated with intensive anthropogenic emissions. These results are consistent with pre-
 297 vious modeling and satellite-based analyses (Wang et al., 2017; Xie et al., 2020), confirming
 298 the robustness of the simulated radiative forcing patterns.

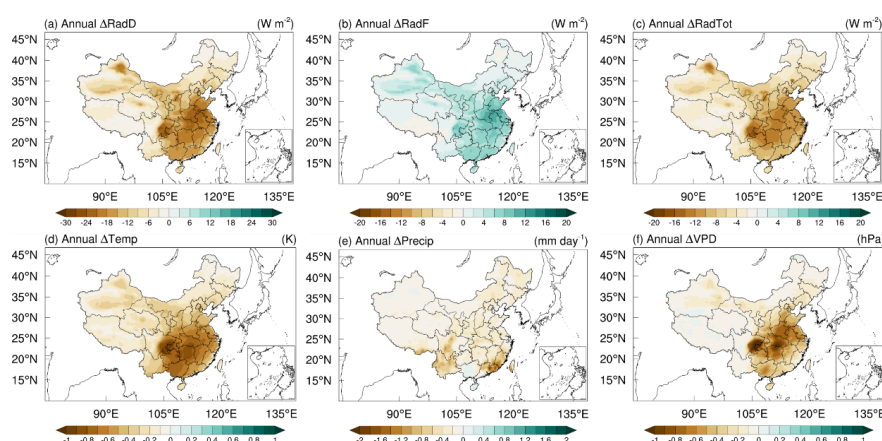
299 The reduction in surface solar radiation directly perturbed the regional energy balance and
 300 atmospheric thermodynamics, resulting in a cooling effect over most of eastern and central
 301 China. As shown in Fig. 3d, surface air temperature decreased significantly in the Sichuan Ba-
 302 sin and coastal regions, with local maxima reaching -1.1°C . In contrast, western and north-
 303 eastern China experienced weaker changes, consistent with lower AOD levels. The simulated
 304 national mean temperature decline of 0.32°C agrees well with previous RegCM-based studies
 305 (Wang et al., 2015; Xie et al., 2020). This widespread cooling is primarily attributed to aerosol-
 306 induced dimming, which suppresses surface shortwave absorption and weakens boundary-
 307 layer turbulence, thereby inhibiting vertical heat exchange and reducing near-surface tempera-
 308 tures.

309 Aerosol also exerted a marked influence on regional hydrological processes. Precipitation
 310 decreased across much of southern and southwestern China, with notable reductions in Guang-
 311 dong, Fujian, Yunnan, and Sichuan provinces, where daily rainfall decreased by up to 2 mm day^{-1}
 312 (Fig. 3e). On average, national precipitation declined by 0.23 mm day^{-1} . The reduction in
 313 rainfall reflects the combined effects of radiative cooling and weakened convective activity.
 314 Specifically, aerosol-induced surface dimming stabilizes the lower atmosphere and suppresses
 315 the upward transport of moisture, while reduced latent heating further limits convective cloud
 316 formation. These mechanisms together explain the widespread drying observed in the simula-
 317 tions.

318 The VPD, a key indicator of plant water stress, also responded sensitively to aerosol forc-
 319 ing. As shown in Fig. 3f, aerosols significantly reduced VPD over central and southeastern
 320 China, with decreases of -0.3 to -0.6 hPa , and locally up to -1.2 hPa in Sichuan, Hebei, and
 321 Jiangsu. The national mean reduction was -0.11 hPa . Lower VPD values imply a moister near-
 322 surface environment and weaker atmospheric demand for evapotranspiration. Ecologically, this



323 alleviation of plant water stress can enhance stomatal conductance and facilitate photosynthetic
 324 carbon uptake, thereby partially compensating for the productivity loss caused by reduced solar
 325 radiation. Thus, the aerosol-induced decline in VPD represents an important indirect pathway
 326 through which aerosols modulate the terrestrial carbon cycle, linking atmospheric radiative
 327 forcing to ecosystem function.
 328



329 **Figure 3.** Annual mean changes in meteorological variables due to aerosol direct radiative
 330 effect during 2010–2020. (a) RadD, direct radiation; (b) RadF, diffuse radiation; (c) RadTot,
 331 total radiation; (d) Temp, air temperature; (e) Precip, precipitation; (f) VPD, vapor pressure
 332 deficit.
 333
 334

335 3.2.2 Effects of Aerosols on the Terrestrial Carbon Sink

336 During 2010–2020, the aerosol overall enhanced the productivity of China’s terrestrial
 337 ecosystems, increasing GPP and NEP by $293.28 \text{ TgC yr}^{-1}$ and $17.93 \text{ TgC yr}^{-1}$, accounting for
 338 3.98% and 4.49% of the national totals, respectively. Spatially, the responses of GPP and NEP
 339 to the aerosol radiative effect displayed significant heterogeneity, with pronounced enhance-
 340 ments in southern and eastern China (Fig. 4a, b). The most significant enhancements occurred
 341 in South-Central and East China, where GPP increased by $0.32 \text{ gC m}^{-2} \text{ day}^{-1}$ and 0.31 gC m^{-2}
 342 day^{-1} , respectively. These regions are characterized by dense forests and cropland ecosystems
 343 with high leaf area index, enabling them to fully exploit the additional diffuse radiation induced



344 by aerosols. Meanwhile, the high aerosol loading in these regions ensured sufficient radiative
 345 perturbation, amplifying the improvement in canopy light-use efficiency. In the Southwest, the
 346 response was more complex. Although the mean GPP increased by $0.20 \text{ gC m}^{-2} \text{ day}^{-1}$, parts of
 347 Yunnan showed a negative effect. This reduction likely results from excessive attenuation of
 348 solar radiation under the region's unique topographic and climatic conditions, which con-
 349 strained photosynthetic activity. Nevertheless, NEP in this region remained positive (approxi-
 350 mately $0.01 \text{ gC m}^{-2} \text{ day}^{-1}$), suggesting that the cooling effect of aerosols substantially sup-
 351 pressed ecosystem respiration, thereby compensating for the reduced photosynthesis. In con-
 352 trast, the North and Northwest exhibited weak positive responses ($<0.07 \text{ gC m}^{-2} \text{ day}^{-1}$), while
 353 the Northeast showed slight inhibition ($-0.04 \text{ gC m}^{-2} \text{ day}^{-1}$), probably due to aerosol-induced
 354 cooling delaying the onset of the growing season. Overall, the spatial patterns of GPP and NEP
 355 responses to the aerosol radiative effect show a clear latitudinal gradient: the humid, high-bio-
 356 mass ecosystems in southern and eastern China are most sensitive to diffuse radiation enhance-
 357 ment, whereas the high-latitude and arid regions experience limited or even negative responses
 358 due to temperature and radiation constraints.

359 From 2010 to 2020, the influence of aerosols on carbon fluxes exhibited distinct interan-
 360 nual variability (Fig. 4c, d). Both GPP and NEP showed an upward trend before 2016, with
 361 GPP increasing from $214.66 \text{ TgC yr}^{-1}$ in 2010 to 384 TgC yr^{-1} in 2016, and NEP rising from
 362 $13.54 \text{ TgC yr}^{-1}$ to $21.31 \text{ TgC yr}^{-1}$. The synchronous growth of GPP and NEP indicates that the
 363 aerosol radiative effect enhanced terrestrial carbon uptake mainly through photosynthetic ac-
 364 tivity. The strong enhancement during 2015–2017 coincided with years of high aerosol loading
 365 and a greater proportion of diffuse radiation, which improved canopy light-use efficiency under
 366 humid and cloudy conditions. After 2018, the positive effect weakened slightly and stabilized
 367 at a lower level. This reduction likely reflects the combined influence of cleaner atmospheric
 368 conditions and changing meteorological patterns, including increased direct radiation and a
 369 reduced diffuse fraction. Year-to-year variations were further modulated by hydroclimatic con-
 370 ditions: higher humidity and cloud cover enhanced aerosol scattering efficiency, while drier or
 371 cleaner years favored direct radiation and weakened the diffuse light advantage. Moreover, the
 372 smaller NEP fluctuations compared to GPP imply a delayed response of ecosystem respiration,



373 as aerosol-induced cooling moderates' respiration more gradually than photosynthesis. Overall,
 374 the interannual variability of GPP and NEP responses to the aerosol radiative effect highlights
 375 the coupled influences of aerosol loading, radiation balance, and regional climate variability
 376 on China's terrestrial carbon sink dynamics.

377 The effects of aerosols on GPP and NEP show pronounced seasonal variation (Fig. S8),
 378 driven by the dynamic coupling between vegetation phenology and environmental factors. In
 379 spring (March–May), aerosols increase GPP by 42.35 TgC (14.4%) and NEP by 2.82 TgC
 380 (15.7%), making a notable contribution at the start of the growing season as rising temperatures
 381 and rapid canopy expansion enhance diffuse radiation benefits, improving light-use efficiency;
 382 meanwhile, moderate cooling suppresses respiration without causing thermal stress, further
 383 boosting NEP. In summer (June–August), positive effects peak, with GPP rising by 173.62 TgC
 384 (59.2%) and NEP by 10.15 TgC (56.6%); under high solar radiation and full canopy closure,
 385 diffuse light penetration reaches its maximum, while cooling alleviates heat stress and reduces
 386 respiration, driving NEP to its annual maximum. In autumn (September–November), aerosols
 387 add 88.38 TgC to GPP (30.1%) and 3.95 TgC to NEP (22.0%), effectively extending the pho-
 388 tosynthetic period as shorter days and reduced total radiation increase the proportion of diffuse
 389 light, sustaining carbon storage. In winter (December–February), GPP declines slightly (-
 390 11.07 TgC, -3.8%), but NEP shows a small positive gain (1.01 TgC, 5.6%) because cooling
 391 strongly suppresses respiration, offsetting reduced photosynthesis. Overall, aerosol radiative
 392 effects regulate seasonal carbon cycling by modifying radiation and thermal conditions. The
 393 net impact depends on the trade-off between the fertilization effect of diffuse radiation and the
 394 opposing effects of reduced total radiation and cooling. Summer emerges as the primary driver
 395 of the annual net positive effect. Accurately quantifying this seasonal dynamic is crucial for
 396 assessing the ecological and climatic consequences of anthropogenic aerosols.

397

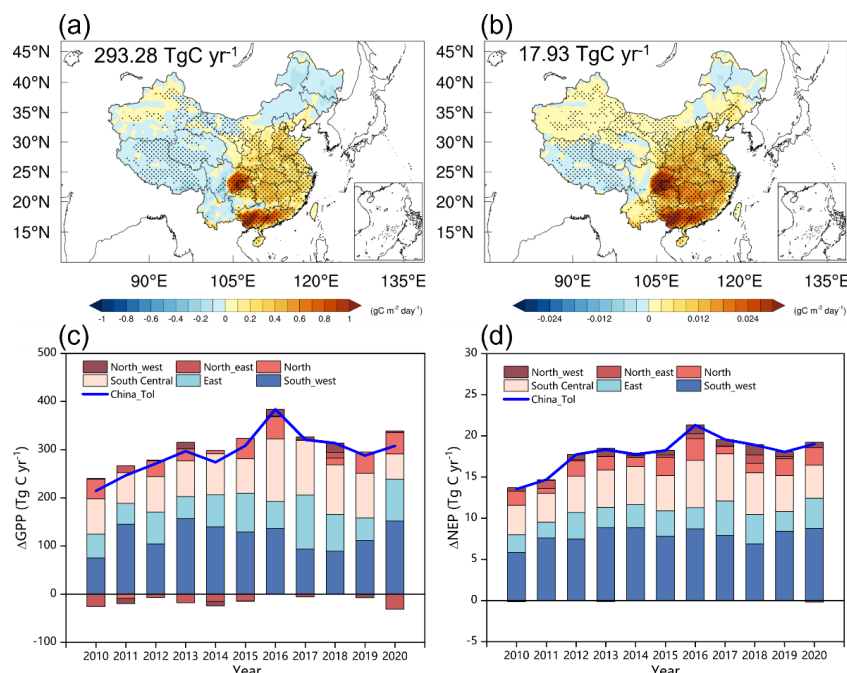


Figure 4. Spatiotemporal variations in carbon flux changes caused by the aerosol radiative effect during 2010–2020. (a–b) Multi-year mean spatial patterns of GPP and NEP changes caused by the aerosol radiative effect. National totals are shown in each panel. Black dots denote significant changes ($p < 0.01$). (c–d) Interannual variations of GPP and NEP changes caused by the aerosol radiative effect.

3.2.3 Contributions of Meteorological Factors to Carbon Sink Changes

We quantified the independent contributions of aerosol-induced meteorological changes to carbon fluxes using the multiple linear regression analysis described in Section 2.5 (Fig. 5a, b). Overall, aerosol substantially influenced China’s terrestrial carbon uptake by altering radiation composition and meteorological conditions. At the national scale, the increase in diffuse radiation emerged as the dominant positive driver, contributing to GPP (325.07 TgC yr⁻¹) and NEP (11.46 TgC yr⁻¹). This highlights the crucial role of the diffuse radiation fertilization effect, particularly in regions with high aerosol loading across eastern and southwestern China, where enhanced diffuse light improves canopy light distribution and photosynthetic efficiency. In



414 contrast, the reduction in direct radiation suppressed GPP ($94.78 \text{ TgC yr}^{-1}$) and NEP (2.59 TgC
 415 yr^{-1}) due to insufficient illumination, though the impact on NEP was weaker, reflecting partial
 416 offset by the reduction in ecosystem respiration under aerosol-induced cooling. Cooling alone
 417 reduced GPP by $59.62 \text{ TgC yr}^{-1}$ and NEP by 4.73 TgC yr^{-1} . This decline in NEP occurred
 418 because the decrease in GPP (driven by reduced transpiration and stomatal conductance) out-
 419 weighed the concurrent reduction in ecosystem respiration. Meanwhile, lower VPD enhanced
 420 GPP by $114.44 \text{ TgC yr}^{-1}$ and NEP by 8.25 TgC yr^{-1} by alleviating water stress, reinforcing
 421 photosynthetic carbon uptake. Changes in precipitation played only a minor role, slightly re-
 422 ducing GPP (8.17 TgC yr^{-1}) and NEP (0.62 TgC yr^{-1}), with limited influence even in the mon-
 423 soon regions of southern China. These findings indicate that variations in radiation components,
 424 rather than hydrometeorological perturbations, serve as the primary pathway through which
 425 aerosols modulate terrestrial carbon sinks.

426 Regionally, among these factors, diffuse radiation exerted the strongest positive influence
 427 on GPP across all regions, particularly in the southwest ($115.92 \text{ TgC yr}^{-1}$), east ($67.04 \text{ TgC yr}^{-1}$),
 428 and south-central ($93.08 \text{ TgC yr}^{-1}$) China (Fig. 5a, b). Enhanced diffuse light under elevated
 429 aerosol loading improved the vertical distribution of photosynthetically active radiation within
 430 the canopy and increased photosynthetic efficiency. In contrast, direct radiation consistently
 431 exhibited negative effects, most evident in the southwest ($-42.32 \text{ TgC yr}^{-1}$) and east (-22.92
 432 TgC yr^{-1}), indicating that aerosol-induced solar dimming partly offset the diffuse radiation fer-
 433 tilization benefit. Temperature changes associated with aerosol cooling suppressed GPP na-
 434 tionwide, especially in the southwest ($-29.07 \text{ TgC yr}^{-1}$) and south-central regions (-16.19 TgC
 435 yr^{-1}), by lowering canopy temperature and reducing evapotranspiration. The contributions of
 436 precipitation were minor (-1 to -3 TgC yr^{-1}), while VPD exerted a positive effect, particularly
 437 in humid southern regions ($55.76 \text{ TgC yr}^{-1}$ in the southwest), suggesting that aerosol-induced
 438 cooling and moistening alleviated water stress and indirectly promoted carbon uptake. For NEP,
 439 diffuse radiation remained the dominant positive driver, with the largest increases in the south-
 440 west (4.06 TgC yr^{-1}) and south-central (3.35 TgC yr^{-1}) China, while direct radiation continued
 441 to exert negative effects. The temperature effect was moderate but consistent with GPP,



implying that aerosol cooling simultaneously suppressed photosynthesis and respiration, with a net positive outcome for NEP. Taken together, these spatial contrasts highlight the combined effects of aerosol composition, vegetation structure, and regional hydroclimate, emphasizing that radiative forcing dominates in humid, high-biomass ecosystems, whereas climatic constraints prevail in arid zones.

To further clarify the dominant controls of these spatial differences, we identified the primary meteorological drivers of GPP and NEP based on the standardized regression coefficients (Fig. 5c, d, and Table 2). The results indicate that for GPP, diffuse radiation accounts for the largest proportion (77.83%), followed by vapor pressure deficit (9.27%) and direct radiation (8.4%), while the influence of temperature (4.45%) and precipitation (0.05%) is relatively small. For NEP, diffuse radiation remains the dominant driver (72.2%), followed by direct radiation (15.92%) and temperature (5.64%). These results highlight that aerosols modify the radiation composition, particularly by enhancing diffuse radiation, which substantially increases photosynthetic efficiency and strengthens the regional carbon sink. In contrast, the effects of temperature and VPD are weaker overall but more pronounced in northern arid and semi-arid regions, where water limitation constrains carbon uptake. Collectively, these findings confirm that radiation composition primarily controls the spatiotemporal dynamics of China's terrestrial carbon sink, while temperature and moisture factors exert region-dependent modulations.

Table 2. Proportion of dominant meteorological factors for GPP and NEP across China (Units: %).

| Factors | RadD (%) | RadF (%) | Temp (%) | Precip (%) | VPD(%) |
|---------|----------|----------|----------|------------|--------|
| GPP | 8.4 | 77.83 | 4.45 | 0.05 | 9.27 |
| NEP | 15.92 | 72.2 | 5.64 | 0.14 | 6.1 |

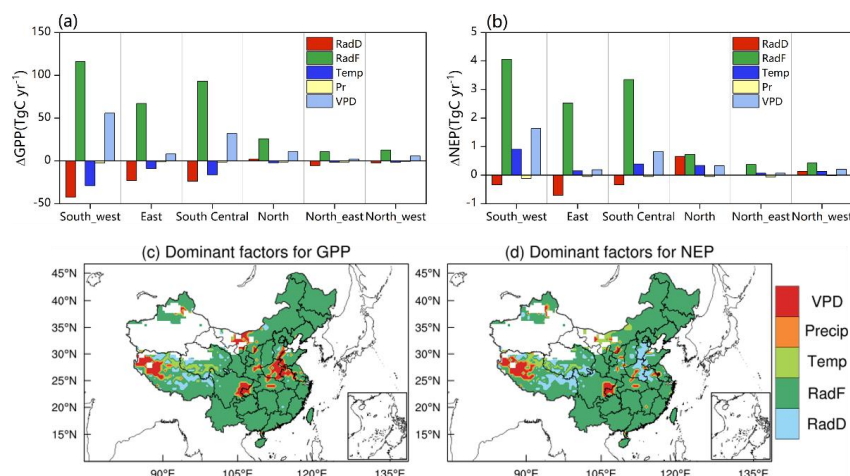


Figure 5. Effect of aerosol-induced changes in meteorological factors on GPP and NEP, and spatial patterns of dominant factors. (a) Regional contributions of individual meteorological factors to GPP; (b) Regional contributions of individual meteorological factors to NEP; (c) Spatial distribution of the dominant meteorological factor for GPP; (d) Spatial distribution of the dominant meteorological factor for NEP.

3.3 Effects of surface ozone on carbon sinks

During 2010–2020, surface O_3 in China increased and imposed a persistent suppression on terrestrial carbon sinks. Simulations show a strong reduction of GPP by $0.4\text{--}0.6\text{ gC m}^{-2}\text{ day}^{-1}$ in most regions, with more than $0.8\text{ gC m}^{-2}\text{ day}^{-1}$ in Southeast and Southwest China (Fig. 6a). NEP shows a similar spatial pattern (Fig. 6b). The largest decline occurs in the southeast (Yangtze River basin and South China coast), with NEP reduced by $0.06\text{--}0.08\text{ gC m}^{-2}\text{ day}^{-1}$ and locally above $0.1\text{ gC m}^{-2}\text{ day}^{-1}$, consistent with high O_3 and evergreen broadleaf forests (Yue et al., 2017). In the southwest (Sichuan Basin and Yunnan–Guizhou Plateau), NEP decreases by $0.03\text{--}0.06\text{ gC m}^{-2}\text{ day}^{-1}$, related to complex terrain and dense forests. Impacts are weaker in Northeast and Northwest China, mostly below $0.02\text{ gC m}^{-2}\text{ day}^{-1}$. In Shandong, Henan, and northern Jiangsu, the simulated losses are small, reflecting cropland-dominated land cover.



489 However, earlier studies reported strong O_3 effects on crops (Ren et al., 2012), suggesting pos-
 490 sible underestimation. This bias may stem from the simplified crop representation in the model
 491 (Fig. S2). Nationwide, O_3 reduces GPP and NEP by $749.44 \text{ TgC yr}^{-1}$ and $51.33 \text{ TgC yr}^{-1}$, ac-
 492 counting for 10.17 % and 12.9 % of the totals. The suppression is attributed to reduced photo-
 493 synthesis, altered stomatal conductance, and shifts in carbon allocation, which together weaken
 494 ecosystem sinks.

495 The annual effect intensifies until 2018 and then weakens (Fig. 6c). In 2010, O_3 reduces
 496 NEP by $42.93 \text{ TgC yr}^{-1}$, reaching $55.71 \text{ TgC yr}^{-1}$ in 2018. It then decreases to $51.98 \text{ TgC yr}^{-1}$
 497 in 2019 and $51.77 \text{ TgC yr}^{-1}$ in 2020. These variations reflect air pollution control policies. Be-
 498 tween 2013 and 2017, the first Clean Air Action reduced $PM_{2.5}$ and NO_x but left volatile or-
 499 ganic compounds (VOCs) largely uncontrolled, thereby enhancing O_3 formation, especially in
 500 VOCs-limited regions (Lu et al., 2020). Both model and observations show higher O_3 during
 501 this stage (Fig. S4). After 2018, the second Clean Air Action introduced coordinated control of
 502 NO_x and VOCs in the Yangtze River Delta and Pearl River Delta, reducing O_3 during summer
 503 and easing sink suppression in 2019–2020. In contrast, O_3 continued to rise in North China,
 504 indicating uneven policy outcomes across regions.

505 Seasonal effects are distinct (Fig. 6d and Fig. S9). Summer shows the strongest suppres-
 506 sion, with NEP reduced by 29.1 TgC (56.69 % of the annual effect). This results from the
 507 overlap of peak O_3 and peak photosynthesis, when high temperature and humidity keep stomata
 508 open and allow O_3 uptake. Spring is second, with NEP reduced by 11.67 TgC (22.74 %). The
 509 effect is linked to leaf expansion, rapid growth, and frequent transport events. Autumn and
 510 winter show weaker impacts due to lower photosynthesis, unfavorable O_3 chemistry, and re-
 511 duced stomatal conductance. Regional differences are evident: in the south, suppression ex-
 512 tends from spring to late autumn, while in the north it is confined to summer. This highlights
 513 the role of climate and phenology in modulating the impact of O_3 on carbon sinks.

514

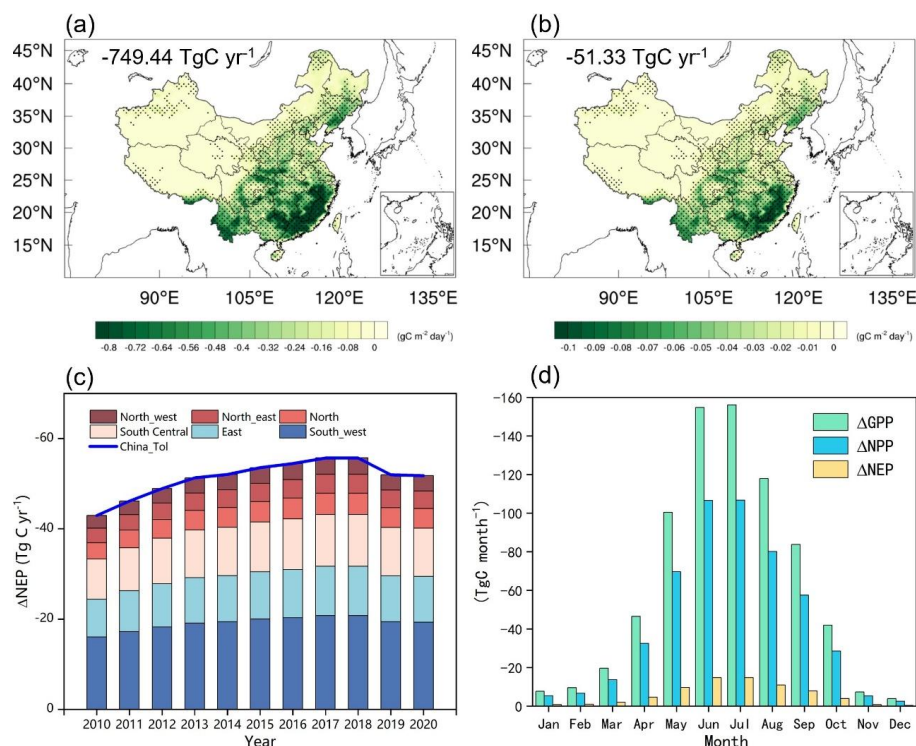


Figure 6. Spatiotemporal variations in O_3 -induced changes in carbon fluxes during 2010–2020. (a–b) Multi-year mean spatial patterns of O_3 -induced changes in GPP and NEP. National totals are shown in each panel. Black dots denote significant changes ($p < 0.01$). (c) Interannual variation of O_3 -induced NEP. (d) O_3 -induced monthly variations in GPP, NPP, and NEP.

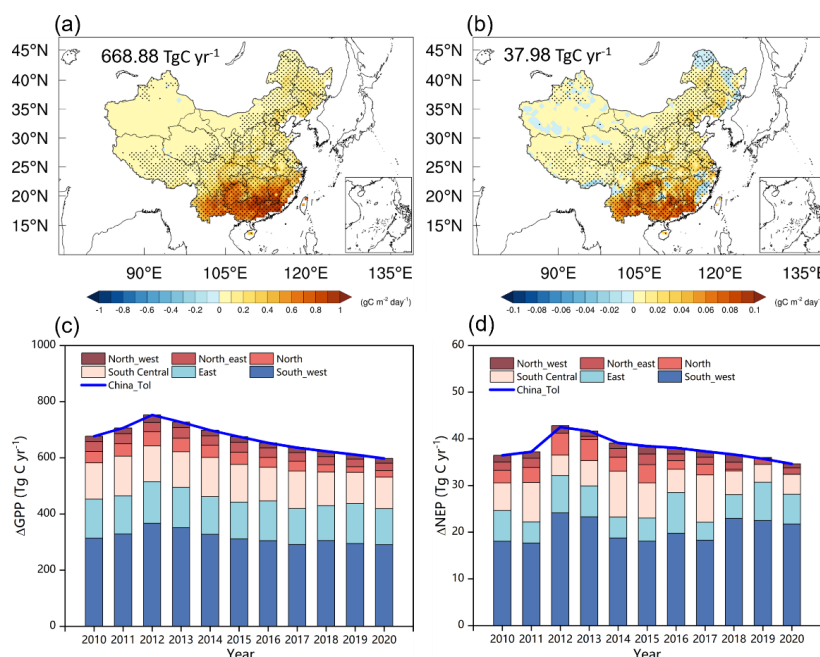
3.4 Effects of atmospheric nitrogen deposition on carbon sinks

The response of China's terrestrial ecosystems to atmospheric nitrogen deposition during 2010–2020 shows pronounced spatial heterogeneity (Fig. 7a, b). At the national scale, nitrogen deposition increased GPP and NEP by $668.88 \text{ TgC yr}^{-1}$ and $37.98 \text{ TgC yr}^{-1}$, respectively. These increases account for 9.08% of total GPP and 9.52% of total NEP. The net gains were mainly concentrated in the southeastern, southwestern, and central regions. In these areas, NEP increased by $0.04\text{--}0.08 \text{ g C m}^{-2} \text{ day}^{-1}$, forming the dominant contribution to the nitrogen-induced carbon sink. Although atmospheric nitrogen deposition is highest in eastern China (Fig. 2e, f), the regional variations in GPP and NEP induced by nitrogen deposition are more pronounced



530 in southern China than in the east. The strong spatial gradient highlights that the ecological
 531 effects of nitrogen deposition are not uniform, but tightly linked to anthropogenic nitrogen
 532 emissions and ecosystem sensitivity (Shang et al., 2024). High responses were observed in
 533 regions with intensive agriculture and industry, where deposition exceeded $15 \text{ kg N ha}^{-1} \text{ yr}^{-1}$.
 534 Vegetation dominated by subtropical evergreen broadleaf forests, mixed forests, and croplands
 535 is generally nitrogen-limited. Additional nitrogen input alleviated nutrient constraints, en-
 536 hanced photosynthesis and biomass accumulation, and shifted soil microbial processes. When
 537 stimulation of GPP and NPP outweighed the increase in ER, NEP rose. Warm and humid cli-
 538 mates, together with long growing seasons, further amplified these effects.

539 The impacts of nitrogen deposition on GPP and NEP varied strongly over time (Fig. 7c,
 540 d). In 2010, deposition enhanced NEP by $36.45 \text{ TgC yr}^{-1}$. The effect increased to a peak of 42.5
 541 TgC yr^{-1} in 2012, but then declined, reaching $34.65 \text{ TgC yr}^{-1}$ by 2020. This trajectory reflects
 542 the influence of China's air pollution control policies on ecosystem carbon dynamics. The tem-
 543 poral trend corresponds to changes in nitrogen deposition fluxes. Between 2010 and 2012, rapid
 544 industrialization and agriculture raised deposition from 15.85 to $17.91 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (+13%).
 545 After 2013, emission reduction policies reduced nitrogen deposition, which fell to 13.25 kg N
 546 $\text{ha}^{-1} \text{ yr}^{-1}$ in 2020 (−26.02%). Notably, the effect of nitrogen deposition on NEP leveled off after
 547 2015, which can be attributed to the slower decline rate of atmospheric nitrogen deposition
 548 since 2015 (Fig. S5). The reduction in NEP (−18.47%) was smaller than that in nitrogen input.
 549 This lagged response suggests that soil nitrogen pools accumulated from long-term deposition
 550 continued to supply nitrogen to vegetation, buffering the decline.



551

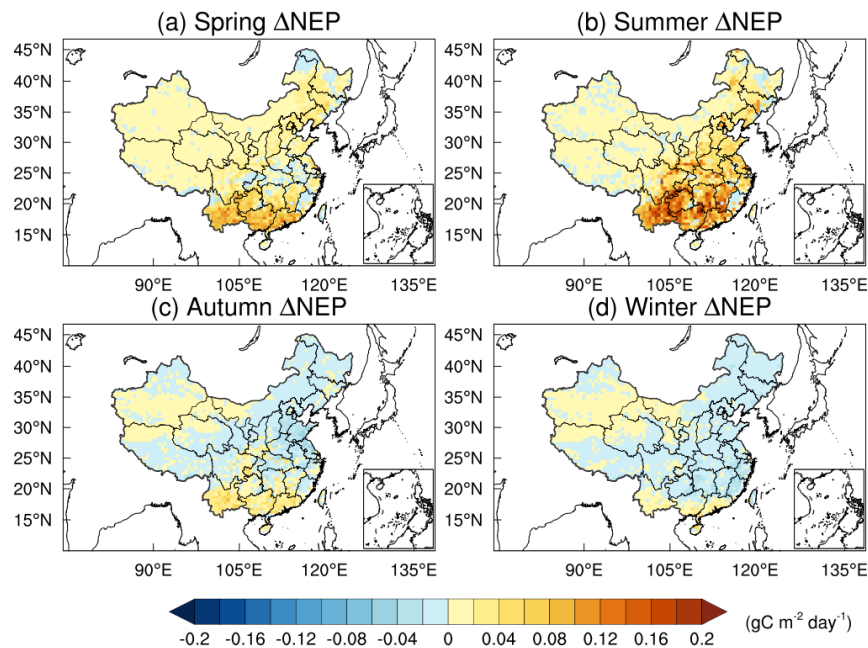
552 **Figure 7.** Spatiotemporal variations in Ndep-induced changes in carbon fluxes during 2010–
 553 2020. (a–b) Multi-year mean spatial patterns of Ndep-induced changes in GPP and NEP. Na-
 554 tional totals are shown in each panel. Black dots denote significant changes ($p < 0.01$). (c–d)
 555 Interannual variation of Ndep-induced in GPP and NEP.

556

557 The influence of nitrogen deposition on NEP displayed clear seasonality (Fig. 8). Strong
 558 positive effects occurred in summer and spring, while autumn and winter showed suppression.
 559 Summer accounted for the largest gain, with an NEP increase of 27.16 TgC. Spring followed
 560 with 14.12 TgC. In contrast, autumn and winter reduced NEP by 0.2 and 3.1 TgC, respectively.
 561 These seasonal differences result from the combined influence of multiple factors. During sum-
 562 mer, optimal temperature, light, and water supported vigorous canopy photosynthesis. Plants
 563 assimilated nitrogen efficiently, leading to higher GPP and biomass accumulation. Spring
 564 growth stages were also nitrogen-sensitive, producing strong positive responses. In autumn and
 565 winter, however, plant activity slowed. Nitrogen inputs mainly stimulated heterotrophic respi-
 566 ration, while GPP and NPP remained low. As a result, NEP decreased, and carbon sink strength
 567 weakened outside the growing season.



568



569

570 **Figure 8.** Spatial distribution of Ndep-induced seasonal variations in NEP during 2010–2020
571 (units: $\text{gC m}^{-2} \text{ day}^{-1}$). (a) Spring, including March, April, and May. (b) Summer, including June,
572 July, and August. (c) Autumn, including September, October, and November. (d) Winter, in-
573 cluding January, February, and December.

574

575 3.5 Integrated Impact Analysis

576 To assess the combined influence of aerosols, O_3 , and atmospheric nitrogen deposition on
577 China's terrestrial carbon sink, the three independent effects were algebraically summed. Dur-
578 ing 2010–2020, the co-evolution of these atmospheric factors jointly drove substantial interan-
579 nual variability and stage-dependent changes in carbon uptake, closely linked to the implemen-
580 tation of the CAA plan. The interannual trend (Fig. 9) shows that although aerosols and nitrogen
581 deposition generally enhanced carbon sequestration, the strong carbon loss caused by O_3
582 largely offset these positive effects. The mean net effect was 4.58 TgC yr^{-1} , exhibiting pro-
583 nounced fluctuations and a declining trend. Net enhancement was strong in the early years of
584 the decade but weakened steadily and approached neutral levels by 2018–2020, when a slight



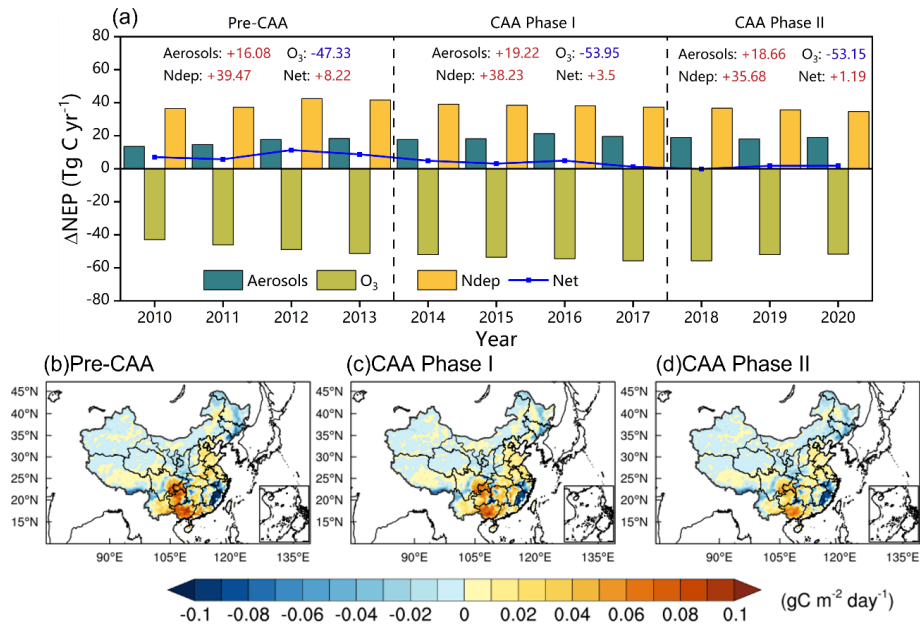
585 negative value ($-0.14 \text{ TgC yr}^{-1}$) first appeared. These changes indicate a gradual transition from
 586 an enhancement-dominated to an inhibition-dominated regime.

587 To further interpret this transition, the study period was divided into three phases accord-
 588 ing to key CAA milestones, and the dominant factors were identified (Fig. 9a). In pre-CAA
 589 (2010–2013), the mean annual net effect reached 8.22 TgC yr^{-1} , characterized by nitrogen-
 590 deposition-dominated enhancement. Nitrogen deposition provided the largest positive contri-
 591 bution ($+39.47 \text{ TgC yr}^{-1}$), while the diffuse-radiation fertilization effect of aerosols offered a
 592 secondary gain ($+16.08 \text{ TgC yr}^{-1}$). The negative impact of O_3 ($-47.33 \text{ TgC yr}^{-1}$) was largely
 593 compensated by the two positive drivers, resulting in a pronounced increase in carbon sink
 594 strength. During CAA Phase I (2014–2017), the mean net effect decreased sharply to 3.50 TgC
 595 yr^{-1} , marking a transitional stage with competing influences. The positive effect of aerosols
 596 peaked ($+19.22 \text{ TgC yr}^{-1}$), likely due to enhanced scattering as absorbing components were
 597 preferentially reduced. However, this gain was largely offset by intensified O_3 -induced inhibi-
 598 tion ($-53.80 \text{ TgC yr}^{-1}$). In CAA Phase II (2018–2020), the mean net effect further declined to
 599 1.19 TgC yr^{-1} , forming an O_3 -dominated pattern. With continued emission control, the aerosol-
 600 induced enhancement decreased from its peak ($+18.66 \text{ TgC yr}^{-1}$), and the nitrogen-deposition
 601 gain weakened ($+35.68 \text{ TgC yr}^{-1}$). Although O_3 suppression slightly eased ($-53.15 \text{ TgC yr}^{-1}$), it
 602 still nearly balanced the combined positive contributions, indicating a fundamental shift in at-
 603 mospheric drivers controlling China’s terrestrial carbon sink.

604 The spatial overlay further supports these findings (Fig. 9b, c, d). In forested and indus-
 605 trialized regions of eastern and southern China, the cancellation between positive and negative
 606 effects was most pronounced. These areas, benefiting from nitrogen and aerosol fertilization
 607 but suffering from intense O_3 pollution, became hotspots of weakened or even reversed net
 608 effects. Overall, the CAA plan not only improved air quality but also altered atmospheric com-
 609 position in ways that substantially affected China’s terrestrial carbon sinks. Policy-driven emis-
 610 sion changes transformed the system from a nitrogen–aerosol-enhanced regime to an O_3 -dom-
 611 inated offset pattern. These results suggest that achieving synergistic benefits between air-qual-
 612 ity improvement and carbon neutrality requires elevating O_3 mitigation to a higher strategic
 613 priority.



614



615

616 **Figure 9.** The overall impacts of aerosols, O₃, and atmospheric nitrogen deposition on the ter-
617 restrial carbon sink in China during 2010–2020. (a) Interannual variations of the combined
618 effects. Pre-CAA represents the period before the implementation of the Clean Air Action
619 (2010–2013); CAA Phase I and CAA Phase II represent the first (2014–2017) and second
620 (2018–2020) stages of the CAA, respectively. (b–d) Spatial distributions of the annual means
621 during the Pre-CAA, CAA Phase I, and CAA Phase II periods.

622

623 3.6 Uncertainties

624 Although the RegESM framework captures the overall spatiotemporal variations of
625 China’s terrestrial carbon sink in response to atmospheric composition changes, several uncer-
626 tainties remain that may influence the quantitative assessment of the individual and combined
627 effects of aerosols, O₃, and nitrogen deposition.

628 First, this study only considered the direct radiative effects of aerosols, while aerosol–
629 cloud interactions were excluded. The first and second indirect effects of aerosols on cloud
630 formation and albedo involve large uncertainties (Haywood and Boucher, 2000) and were



631 therefore not represented in our simulations. However, observations have shown that terrestrial
 632 carbon fluxes are highly sensitive to sky conditions and diffuse radiation changes (Oliphant et
 633 al., 2011; Yue and Unger, 2017). The omission of aerosol–cloud interactions may affect the
 634 magnitude and spatial pattern of aerosol impacts on radiation and photosynthesis, as cloud-
 635 mediated diffuse radiation responses remain uncertain. Future work should explicitly include
 636 aerosol–cloud–radiation feedbacks to better quantify their effects on ecosystem carbon ex-
 637 change.

638 Second, uncertainties remain in evaluating vegetation responses to O₃ exposure. Field-
 639 based O₃ fumigation experiments across China are still limited, making it difficult to compre-
 640 hensively assess ecosystem-level damage. In this study, the YIBs model applied different O₃
 641 damage coefficients for plant functional types, and the parameterization has shown reasonable
 642 regional performance in simulating GPP–O₃ responses (Yue et al., 2017). Nevertheless, a wider
 643 range of site-level observations is required to constrain the O₃ damage functions across various
 644 vegetation types and climate zones at the national scale.

645 Third, nonlinear coupling among aerosols, O₃, and nitrogen deposition introduces sys-
 646 temic uncertainty in estimating their combined effects. Aerosol reduction alters photolysis rates
 647 and thereby affects O₃ formation (Tang et al., 2017; Yan et al., 2023; Yang et al., 2022), while
 648 O₃ and nitrogen jointly regulate stomatal conductance, photosynthetic efficiency, and water-
 649 use dynamics (Sitch et al., 2007; Zhang et al., 2018). These interactions may amplify or offset
 650 each other under changing climatic conditions, emphasizing the need for high-resolution, fully
 651 coupled chemistry–ecosystem modeling frameworks to capture the co-evolution of multiple
 652 atmospheric processes.

653 In summary, despite these uncertainties, this study provides robust quantitative evidence
 654 that aerosols, O₃, and nitrogen deposition jointly modified the magnitude and spatial distribu-
 655 tion of China’s terrestrial carbon sink during 2010–2020. Future efforts should focus on incor-
 656 porating aerosol–cloud interactions, expanding field-based O₃ response networks, and improv-
 657 ing representation of multi-process coupling to further constrain atmospheric–biosphere feed-
 658 backs under China’s evolving air quality and carbon neutrality goals.

659



660 4 Conclusions

661 This study employed the RegESM to quantify the effects of aerosol, surface O₃, and ni-
 662 trogen deposition on China's terrestrial carbon sink during 2010–2020. The model effectively
 663 reproduced the spatial and temporal variations of aerosol optical depth, O₃, nitrogen deposition,
 664 and carbon fluxes, providing a solid basis for process-level attribution analysis.

665 Aerosols exerted a substantial positive influence on China's terrestrial carbon sink. On
 666 average, aerosols enhanced GPP and NEP by 293.28 TgC yr⁻¹ (3.98%) and 17.93 TgC yr⁻¹
 667 (4.49%), respectively, primarily through the diffuse radiation fertilization effect. The strongest
 668 enhancement appeared in southern and eastern China, where high aerosol loading and dense
 669 vegetation synergistically improved canopy light-use efficiency. Aerosol-induced surface cool-
 670 ing and reduced VPD further alleviated water stress and stimulated carbon uptake. The en-
 671 hancement peaked during 2015–2017, coinciding with elevated diffuse radiation fractions, and
 672 weakened slightly under cleaner atmospheric conditions after 2018.

673 In contrast, surface O₃ persistently suppressed ecosystem carbon uptake, reducing GPP
 674 and NEP by 749.44 TgC yr⁻¹ (10.17%) and 51.33 TgC yr⁻¹ (12.9%), respectively. The strongest
 675 suppression occurred in southeastern and southwestern China, where dense forest ecosystems
 676 coincided with high O₃ concentrations. O₃-induced damage peaked in 2018, consistent with the
 677 exceptionally high O₃ levels. Subsequent coordinated NO_x–VOCs management under the sec-
 678 ond Clean Air Action partially mitigated O₃ levels and NEP suppression. O₃ exerted a strongly
 679 seasonal negative impact on NEP, with the strongest suppression occurring in summer.

680 Atmospheric nitrogen deposition enhanced the terrestrial carbon sink by 9.08% for GPP
 681 and 9.52% for NEP, with effects concentrated in southern and central China. The enhancement
 682 peaked around 2012, declined gradually after 2013 following reduced anthropogenic emissions,
 683 and leveled off after 2015, corresponding to a slower decline in deposition and a lagged eco-
 684 system response due to soil nitrogen accumulation. Seasonal variations showed stronger stim-
 685 ulation in summer and spring, while autumn and winter exhibited minor reductions linked to
 686 enhanced respiration.

687 During 2010–2020, the combined effects of aerosols, surface O₃, and atmospheric nitro-
 688 gen deposition on China's terrestrial carbon sink exhibited marked interannual variability and



689 a distinct transition under the Clean Air Action (CAA). The net atmospheric contribution de-
 690 clined from $+8.22 \text{ Tg C yr}^{-1}$ during the Pre-CAA period (2010–2013) to $+1.19 \text{ Tg C yr}^{-1}$ in
 691 Phase II (2018–2020), as the increasing suppression from O_3 ($-53.15 \text{ Tg C yr}^{-1}$) gradually offset
 692 the positive impacts of aerosols and nitrogen deposition. These results indicate that China’s air-
 693 pollution control not only improved air quality but also altered atmospheric chemical compo-
 694 sition in ways that significantly affected ecosystem carbon uptake, with O_3 becoming the dom-
 695 inant limiting factor in the later period.

696 Overall, aerosols, O_3 , and nitrogen deposition exerted interconnected yet contrasting in-
 697 fluences on China’s terrestrial carbon sink. Aerosols and nitrogen deposition enhanced carbon
 698 uptake through diffuse radiation and nutrient input, whereas O_3 caused physiological damage
 699 that suppressed it. The evolving interplay among these factors illustrates how emission reduc-
 700 tions, atmospheric chemistry, and ecosystem feedbacks jointly impact carbon sink dynamics
 701 under China’s clean-air policies. Strengthening integrated O_3 control is therefore essential to
 702 secure co-benefits for air quality improvement and carbon neutrality goals.

703 **Data Availability Statement**

704 MODIS data are provided at <https://doi.org/10.5067/MODIS/MOD17A2H.006> (last access: 11
 705 May 2024; Running et al., 2015). AOD data come from the MODIS Level-3 monthly product
 706 (MOD08_M3, Collection 6.1) available at https://doi.org/10.5067/MODIS/MOD08_M3.061
 707 (last access: 5 November 2024; Platnick et al., 2015). WDCGG data are available at
 708 <https://doi.org/10.15138/wkgj-f215> (last access: 10 June 2024; Lan, 2023). ERA-Interim data
 709 are available at <https://doi.org/10.24381/cds.f2f5241d> (last access: 8 June 2024; Dee et al.,
 710 2011). The Carbon Tracker data can be accessed at <https://gml.noaa.gov/ccgg/carbontracker/>
 711 (last access: 21 June 2024; Jacobson et al., 2023). CNEMC data are provided at
 712 <https://www.cnemc.cn/en> (Kong et al., 2021).

713 **Author contributions**

714 TW and NX designed the study. NX performed the analysis, conducted the evaluation, and
 715 drafted the manuscript. TW supervised the research and acquired funding. Both authors con-
 716 tributed to the interpretation of results and to reviewing and editing the manuscript.



717 **Competing interests**

718 The corresponding author has stated that all the authors have no conflicts of interest.

719 **Disclaimer**

720 Publisher's note: Copernicus Publications remains neutral with regard to jurisdictional claims
 721 made in the text, published maps, institutional affiliations, or any other geographical represen-
 722 tation in this paper. The authors bear the ultimate responsibility for providing appropriate place
 723 names. Views expressed in the text are those of the authors and do not necessarily reflect the
 724 views of the publisher.

725 **Acknowledgments**

726 We would like to acknowledge the anthropogenic emission inventory support from Tsinghua
 727 University and the observed data from the China National Environmental Monitoring Center.
 728 We also gratefully acknowledge a wide range of other institutional partners.

729 **Financial support**

730 This work was supported by the National Key Basic Research Development Program of China
 731 (2024YFC3711905) and the National Natural Science Foundation of China (42477103).

732 **References**

- 733 Cao, S., He, Y., Zhang, L., Sun, Q., Zhang, Y., Li, H., Wei, X., and Liu, Y.: Spatiotemporal
 734 dynamics of vegetation net ecosystem productivity and its response to drought in
 735 Northwest China, *GIScience & Remote Sensing*, 60, 2194597,
 736 <https://doi.org/10.1080/15481603.2023.2194597>, 2023.
- 737 Cen, X., He, N., Van Sundert, K., Yu, K., Li, M., Xu, L., He, L., and Butterbach-Bahl, K.:
 738 Global patterns of nitrogen saturation in forests, *One Earth*, 8,
 739 <https://doi.org/10.1016/j.oneear.2024.10.007>, 2025.
- 740 Chen, H., Li, D., Gurmesa, G. A., Yu, G., Li, L., Zhang, W., Fang, H., and Mo, J.: Effects of
 741 nitrogen deposition on carbon cycle in terrestrial ecosystems of China: A meta-analysis,



- 742 Environ. Pollut., 206, 352-360, <https://doi.org/10.1016/j.envpol.2015.07.033>, 2015.
- 743 Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U.,
 744 Balmaseda, M. A., Balsamo, G., Bauer, P., Bechtold, P., Beljaars, A. C. M., van de Berg,
 745 L., Bidlot, J., Bormann, N., Delsol, C., Dragani, R., Fuentes, M., Geer, A. J., Haimberger,
 746 L., Healy, S. B., Hersbach, H., Hólm, E. V., Isaksen, L., Kållberg, P., Köhler, M.,
 747 Matricardi, M., McNally, A. P., Monge-Sanz, B. M., Morcrette, J. J., Park, B. K., Peubey,
 748 C., de Rosnay, P., Tavolato, C., Thépaut, J. N., and Vitart, F.: The ERA-Interim reanalysis:
 749 configuration and performance of the data assimilation system, *QJRM*, 137, 553-597,
 750 <https://doi.org/10.1002/qj.828>, 2011.
- 751 Doughty, C. E., Flanner, M. G., and Goulden, M. L.: Effect of smoke on subcanopy shaded
 752 light, canopy temperature, and carbon dioxide uptake in an Amazon rainforest, *Global*
 753 *Biogeochemical Cycles*, 24, <https://doi.org/10.1029/2009gb003670>, 2010.
- 754 Emmons, L. K., Walters, S., Hess, P. G., Lamarque, J. F., Pfister, G. G., Fillmore, D., Granier,
 755 C., Guenther, A., Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall, G., Wiedinmyer,
 756 C., Baughcum, S. L., and Kloster, S.: Description and evaluation of the Model for Ozone
 757 and Related chemical Tracers, version 4 (MOZART-4), *Geosci. Model Dev.*, 3, 43-67,
 758 <https://doi.org/10.5194/gmd-3-43-2010>, 2010.
- 759 Feng, Z., Hu, E., Wang, X., Jiang, L., and Liu, X.: Ground-level O₃ pollution and its impacts
 760 on food crops in China: A review, *Environ. Pollut.*, 199, 42-48,
 761 <https://doi.org/10.1016/j.envpol.2015.01.016>, 2015.
- 762 Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Bakker, D. C. E., Hauck, J.,
 763 Landschützer, P., Le Quéré, C., Luijkx, I. T., Peters, G. P., Peters, W., Pongratz, J.,
 764 Schwingshackl, C., Sitch, S., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., Anthoni,
 765 P., Barbero, L., Bates, N. R., Becker, M., Bellouin, N., Decharme, B., Bopp, L., Brasika,
 766 I. B. M., Cadule, P., Chamberlain, M. A., Chandra, N., Chau, T.-T.-T., Chevallier, F., Chini,
 767 L. P., Cronin, M., Dou, X., Enyo, K., Evans, W., Falk, S., Feely, R. A., Feng, L., Ford, D.
 768 J., Gasser, T., Ghattas, J., Gkritzalis, T., Grassi, G., Gregor, L., Gruber, N., Gürses, Ö.,
 769 Harris, I., Hefner, M., Heinke, J., Houghton, R. A., Hurtt, G. C., Iida, Y., Ilyina, T.,
 770 Jacobson, A. R., Jain, A., Jarníková, T., Jersild, A., Jiang, F., Jin, Z., Joos, F., Kato, E.,



- 771 Keeling, R. F., Kennedy, D., Klein Goldewijk, K., Knauer, J., Korsbakken, J. I., Körtzinger,
 772 A., Lan, X., Lefèvre, N., Li, H., Liu, J., Liu, Z., Ma, L., Marland, G., Mayot, N., McGuire,
 773 P. C., McKinley, G. A., Meyer, G., Morgan, E. J., Munro, D. R., Nakaoka, S.-I., Niwa, Y.,
 774 O'Brien, K. M., Olsen, A., Omar, A. M., Ono, T., Paulsen, M., Pierrot, D., Pocock, K.,
 775 Poulter, B., Powis, C. M., Rehder, G., Resplandy, L., Robertson, E., Rödenbeck, C., Rosan,
 776 T. M., Schwinger, J., Séférian, R., Smallman, T. L., Smith, S. M., Sospedra-Alfonso, R.,
 777 Sun, Q., Sutton, A. J., Sweeney, C., Takao, S., Tans, P. P., Tian, H., Tilbrook, B., Tsujino,
 778 H., Tubiello, F., van der Werf, G. R., van Ooijen, E., Wanninkhof, R., Watanabe, M.,
 779 Wimart-Rousseau, C., Yang, D., Yang, X., Yuan, W., Yue, X., Zaehle, S., Zeng, J., and
 780 Zheng, B.: Global Carbon Budget 2023, *Earth System Science Data*, 15, 5301-5369,
 781 <https://doi.org/10.5194/essd-15-5301-2023>, 2023.
- 782 Geng, G. N., Liu, Y. X., Liu, Y., Liu, S. G., Cheng, J., Yan, L., Wu, N. N., Hu, H. W., Tong, D.,
 783 Zheng, B., Yin, Z. C., He, K. B., and Zhang, Q.: Efficacy of China's clean air actions to
 784 tackle PM_{2.5} pollution between 2013 and 2020, *Nature Geoscience*, 17,
 785 <https://doi.org/10.1038/s41561-024-01540-z>, 2024.
- 786 Giorgi, F., Coppola, E., Solmon, F., Mariotti, L., Sylla, M. B., Bi, X., Elguindi, N., Diro, G. T.,
 787 Nair, V., Giuliani, G., Turuncoglu, U. U., Cozzini, S., Guttler, I., O'Brien, T. A., Tawfik,
 788 A. B., Shalaby, A., Zakey, A. S., Steiner, A. L., Stordal, F., Sloan, L. C., and Brankovic,
 789 C.: RegCM4: model description and preliminary tests over multiple CORDEX domains,
 790 *Clim. Res.*, 52, 7-29, <https://doi.org/10.3354/cr01018>, 2012.
- 791 Gu, L. H., Baldocchi, D. D., Wofsy, S. C., Munger, J. W., Michalsky, J. J., Urbanski, S. P., and
 792 Boden, T. A.: Response of a deciduous forest to the Mount Pinatubo eruption: Enhanced
 793 photosynthesis, *Science*, 299, 2035-2038, <https://doi.org/10.1126/science.1078366>, 2003.
- 794 Hao, Y., Meng, X., Yu, X., Lei, M., Li, W., Yang, W., Shi, F., and Xie, S.: Quantification of
 795 primary and secondary sources to PM_{2.5} using an improved source regional apportionment
 796 method in an industrial city, China, *Sci. Total Environ.*, 706,
 797 <https://doi.org/10.1016/j.scitotenv.2019.135715>, 2020.
- 798 Haywood, J. and Boucher, O.: Estimates of the direct and indirect radiative forcing due to
 799 tropospheric aerosols: A review, *Reviews of Geophysics*, 38, 513-543,



- 800 <https://doi.org/10.1029/1999rg000078>, 2000.
- 801 He, M. Z., Kimball, J. S., Maneta, M. P., Maxwell, B. D., Moreno, A., Begueria, S., and Wu,
 802 X. C.: Regional Crop Gross Primary Productivity and Yield Estimation Using Fused
 803 Landsat-MODIS Data, *Remote Sens.*, 10, 372, <https://doi.org/10.3390/rs10030372>, 2018.
- 804 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J.,
 805 Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X.,
 806 Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren,
 807 P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer,
 808 A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux,
 809 P., Lopez, P., Lupu, C., Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S.,
 810 and Thépaut, J. N.: The ERA5 global reanalysis, *QJRMS*, 146, 1999-2049,
 811 <https://doi.org/10.1002/qj.3803>, 2020.
- 812 Horowitz, L. W., Walters, S., Mauzerall, D. L., Emmons, L. K., Rasch, P. J., Granier, C., Tie,
 813 X. X., Lamarque, J. F., Schultz, M. G., Tyndall, G. S., Orlando, J. J., and Brasseur, G. P.:
 814 A global simulation of tropospheric ozone and related tracers: Description and evaluation
 815 of MOZART, version 2, *J. Geophys. Res.: Atmos.*, 108, 4784,
 816 <https://doi.org/10.1029/2002jd002853>, 2003.
- 817 Jacobson, A. R., Schuldt, K. N., Tans, P., Arlyn Andrews, Miller, J. B., Oda, T., Mund, J., Weir,
 818 B., Ott, L., Aalto, T., Abshire, J. B., Aikin, K., Aoki, S., Apadula, F., Arnold, S., Baier, B.,
 819 Bartyzel, J., Beyersdorf, A., Biermann, T., ... Mirosław Zimnoch.: CarbonTracker
 820 CT2022, NOAA Global Monitoring Laboratory, [data set],
 821 <https://doi.org/10.25925/Z1GJ-3254>, 2023.
- 822 Jung, M., Reichstein, M., Schwalm, C. R., Huntingford, C., Sitch, S., Ahlstrom, A., Arneth, A.,
 823 Camps-Valls, G., Ciais, P., Friedlingstein, P., Gans, F., Ichii, K., Ain, A. K. J., Kato, E.,
 824 Papale, D., Poulter, B., Raduly, B., Rodenbeck, C., Tramontana, G., Viovy, N., Wang, Y.-
 825 P., Weber, U., Zaehle, S., and Zeng, N.: Compensatory water effects link yearly global
 826 land CO₂ sink changes to temperature, *Nature*, 541, 516-520,
 827 <https://doi.org/10.1038/nature20780>, 2017.
- 828 Kong, L., Tang, X., Zhu, J., Wang, Z. F., Li, J. J., Wu, H. J., Wu, Q. Z., Chen, H. S., Zhu, L. L.,



- 829 Wang, W., Liu, B., Wang, Q., Chen, D. H., Pan, Y. P., Song, T., Li, F., Zheng, H. T., Jia,
 830 G. L., Lu, M. M., Wu, L., and Carmichael, G. R.: A 6-year-long (2013-2018) high-
 831 resolution air quality reanalysis dataset in China based on the assimilation of surface
 832 observations from CNEMC, *Earth System Science Data*, 13, 529-570,
 833 <https://doi.org/10.5194/essd-13-529-2021>, 2021.
- 834 Kuniyal, J. C. and Guleria, R. P.: The current state of aerosol-radiation interactions: A mini
 835 review, *Journal of Aerosol Science*, 130, 45-54,
 836 <https://doi.org/10.1016/j.jaerosci.2018.12.010>, 2019.
- 837 Lan, X.: Atmospheric Carbon Dioxide Dry Air Mole Fractions from the NOAA GML Carbon
 838 Cycle Cooperative Global Air Sampling Network, 1968-2022, Version: 2023-08-28, [data
 839 set], <https://doi.org/10.15138/wkgj-f215>, 2023.
- 840 Lawrence, P. J. and Chase, T. N.: Representing a new MODIS consistent land surface in the
 841 Community Land Model (CLM 3.0), *J. Geophys. Res.: Biogeosci.*, 112, G01023,
 842 <https://doi.org/10.1029/2006jg000168>, 2007.
- 843 Lei, Y., Yue, X., Wang, Z., Liao, H., Zhang, L., Tian, C., Zhou, H., Zhong, J., Guo, L., Che, H.,
 844 and Zhang, X.: Mitigating ozone damage to ecosystem productivity through sectoral and
 845 regional emission controls: a case study in the Yangtze River Delta, China, *Environmental*
 846 *Research Letters*, 17, <https://doi.org/10.1088/1748-9326/ac6ff7>, 2022.
- 847 Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F., and Hsu, N.
 848 C.: The Collection 6 MODIS aerosol products over land and ocean, *Atmospheric*
 849 *Measurement Techniques*, 6, 2989-3034, <https://doi.org/10.5194/amt-6-2989-2013>, 2013.
- 850 Li, M., Huang, X., Yan, D., Lai, S., Zhang, Z., Zhu, L., Lu, Y., Jiang, X., Wang, N., Wang, T.,
 851 Song, Y., and Ding, A.: Coping with the concurrent heatwaves and ozone extremes in
 852 China under a warming climate, *Science Bulletin*, 69, 2938-2947,
 853 <https://doi.org/10.1016/j.scib.2024.05.034>, 2024.
- 854 Li, M., Liu, H., Geng, G. N., Hong, C. P., Liu, F., Song, Y., Tong, D., Zheng, B., Cui, H. Y.,
 855 Man, H. Y., Zhang, Q., and He, K. B.: Anthropogenic emission inventories in China: a
 856 review, *National Science Review*, 4, 834-866, <https://doi.org/10.1093/nsr/nwx150>, 2017.
- 857 Li, X., Liang, H., and Cheng, W.: Spatio-Temporal Variation in AOD and Correlation Analysis



- 858 with PAR and NPP in China from 2001 to 2017, *Remote Sens.*, 12,
 859 <https://doi.org/10.3390/rs12060976>, 2020.
- 860 Liu, H., Liu, S., Xue, B., Lv, Z., Meng, Z., Yang, X., Xue, T., Yu, Q., and He, K.: Ground-level
 861 ozone pollution and its health impacts in China, *Atmos. Environ.*, 173, 223-230,
 862 <https://doi.org/10.1016/j.atmosenv.2017.11.014>, 2018.
- 863 Liu, L., Wen, Z., Liu, S., Zhang, X., and Liu, X.: Decline in atmospheric nitrogen deposition
 864 in China between 2010 and 2020, *Nature Geoscience*, 17, [https://doi.org/10.1038/s41561-](https://doi.org/10.1038/s41561-024-01484-4)
 865 024-01484-4, 2024.
- 866 Liu, M., Shang, F., Lu, X., Huang, X., Song, Y., Liu, B., Zhang, Q., Liu, X., Cao, J., Xu, T.,
 867 Wang, T., Xu, Z., Xu, W., Liao, W., Kang, L., Cai, X., Zhang, H., Dai, Y., and Zhu, T.:
 868 Unexpected response of nitrogen deposition to nitrogen oxide controls and implications
 869 for land carbon sink, *Nature Communications*, 13, [https://doi.org/10.1038/s41467-022-](https://doi.org/10.1038/s41467-022-30854-y)
 870 30854-y, 2022.
- 871 Liu, X., Zhang, Y., Han, W., Tang, A., Shen, J., Cui, Z., Vitousek, P., Erisman, J. W., Goulding,
 872 K., Christie, P., Fangmeier, A., and Zhang, F.: Enhanced nitrogen deposition over China,
 873 *Nature*, 494, 459-462, <https://doi.org/10.1038/nature11917>, 2013.
- 874 Lu, X., Hou, E., Guo, J., Gilliam, F. S., Li, J., Tang, S., and Kuang, Y.: Nitrogen addition
 875 stimulates soil aggregation and enhances carbon storage in terrestrial ecosystems of China:
 876 A meta-analysis, *Global Change Biology*, 27, 2780-2792,
 877 <https://doi.org/10.1111/gcb.15604>, 2021.
- 878 Lu, X., Jiang, H., Liu, J., Zhang, X., Jin, J., Zhu, Q., Zhang, Z., and Peng, C.: Simulated effects
 879 of nitrogen saturation on the global carbon budget using the IBIS model, *Sci. Rep.*, 6,
 880 <https://doi.org/10.1038/srep39173>, 2016.
- 881 Lu, X., Zhang, L., Wang, X. L., Gao, M., Li, K., Zhang, Y. Z., Yue, X., and Zhang, Y. H.: Rapid
 882 Increases in Warm-Season Surface Ozone and Resulting Health Impact in China Since
 883 2013, *Environmental Science & Technology Letters*, 7, 240-247,
 884 <https://doi.org/10.1021/acs.estlett.0c00171>, 2020.
- 885 Luo, Y. X., Zheng, X. B., Zhao, T. L., and Chen, J.: A climatology of aerosol optical depth over
 886 China from recent 10 years of MODIS remote sensing data, *IJCLI*, 34, 863-870,



- 887 <https://doi.org/10.1002/joc.3728>, 2014.
- 888 Ma, D. Y., Wang, T. J., Wu, H., Qu, Y. W., Liu, J., Liu, J. E., Li, S., Zhuang, B. L., Li, M. M.,
 889 and Xie, M.: The effect of anthropogenic emission, meteorological factors, and
 890 carbondioxide on the surface ozone increase in China from 2008 to 2018 during theEast
 891 Asia summer monsoon season, *Atmos. Chem. Phys.*, 23, 6525-6544,
 892 <https://doi.org/10.5194/acp-23-6525-2023>, 2023.
- 893 Madani, N., Kimball, J. S., Affleck, D. L. R., Kattge, J., Graham, J., van Bodegom, P. M., Reich,
 894 P. B., and Running, S. W.: Improving ecosystem productivity modeling through spatially
 895 explicit estimation of optimal light use efficiency, *J. Geophys. Res.: Biogeosci.*, 119,
 896 1755-1769, <https://doi.org/10.1002/2014jg002709>, 2014.
- 897 Mercado, L. M., Bellouin, N., Sitch, S., Boucher, O., Huntingford, C., Wild, M., and Cox, P.
 898 M.: Impact of changes in diffuse radiation on the global land carbon sink, *Nature*, 458,
 899 1014-U1087, <https://doi.org/10.1038/nature07949>, 2009.
- 900 Oliphant, A. J., Dragoni, D., Deng, B., Grimmond, C. S. B., Schmid, H. P., and Scott, S. L.:
 901 The role of sky conditions on gross primary production in a mixed deciduous forest, *Agric.*
 902 *For. Meteorol.*, 151, 781-791, <https://doi.org/10.1016/j.agrformet.2011.01.005>, 2011.
- 903 Peng, X., Wei, W., Niu, S., Huang, Y., and Chen, L.: Divergent impact of long-term
 904 anthropogenic nitrogen inputs on global particulate and mineral-associated organic carbon,
 905 *Ecological Processes*, 14, <https://doi.org/10.1186/s13717-025-00624-x>, 2025.
- 906 Peters, W., Jacobson, A. R., Sweeney, C., Andrews, A. E., Conway, T. J., Masarie, K., Miller,
 907 J. B., Bruhwiler, L. M. P., Petron, G., Hirsch, A. I., Worthy, D. E. J., van der Werf, G. R.,
 908 Randerson, J. T., Wennberg, P. O., Krol, M. C., and Tans, P. P.: An atmospheric perspective
 909 on North American carbon dioxide exchange: CarbonTracker, *Proc. Natl. Acad. Sci.*
 910 *U.S.A.*, 104, 18925-18930, <https://doi.org/10.1073/pnas.0708986104>, 2007.
- 911 Piao, S., He, Y., Wang, X., and Chen, F.: Estimation of China's terrestrial ecosystem carbon
 912 sink: Methods, progress and prospects, *Science China Earth Sciences*, 65, 641-651,
 913 <https://doi.org/10.1007/s11430-021-9892-6>, 2022.
- 914 Piao, S., Sitch, S., Ciais, P., Friedlingstein, P., Peylin, P., Wang, X., Ahlstrom, A., Anav, A.,
 915 Canadell, J. G., Cong, N., Huntingford, C., Jung, M., Levis, S., Levy, P. E., Li, J., Lin, X.,



- 916 Lomas, M. R., Lu, M., Luo, Y., Ma, Y., Myneni, R. B., Poulter, B., Sun, Z., Wang, T.,
 917 Viovy, N., Zaehle, S., and Zeng, N.: Evaluation of terrestrial carbon cycle models for their
 918 response to climate variability and to CO₂ trends, *Glob. Chang. Biol.*, 19, 2117-2132,
 919 <https://doi.org/10.1111/gcb.12187>, 2013.
- 920 Platnick, S., King, M., and Hubanks, P.: MOD08_M3 - MODIS/Terra Aerosol Cloud Water
 921 Vapor Ozone Monthly L3 Global 1Deg CMG, NASA MODIS Adaptive Processing
 922 System, Goddard Space Flight Center [data set], USA,
 923 https://doi.org/10.5067/MODIS/MOD08_M3.061, 2015.
- 924 Post, E., Steinman, B. A., and Mann, M. E.: Acceleration of phenological advance and warming
 925 with latitude over the past century, *Sci. Rep.*, 8, 3927, [https://doi.org/10.1038/s41598-018-](https://doi.org/10.1038/s41598-018-22258-0)
 926 [22258-0](https://doi.org/10.1038/s41598-018-22258-0), 2018.
- 927 Ren, W., Tian, H. Q., Tao, B., Huang, Y., and Pan, S. F.: China's crop productivity and soil
 928 carbon storage as influenced by multifactor global change, *Global Change Biology*, 18,
 929 2945-2957, <https://doi.org/10.1111/j.1365-2486.2012.02741.x>, 2012.
- 930 Ren, W., Banger, K., Tao, B., Yang, J., Huang, Y., and Tian, H.: Global pattern and change of
 931 cropland soil organic carbon during 1901-2010: Roles of climate, atmospheric chemistry,
 932 land use and management, *Geography and Sustainability*, 1, 59-69,
 933 <https://doi.org/10.1016/j.geosus.2020.03.001>, 2020.
- 934 Running, S., Mu, Q., Zhao, M.: MOD17A2H MODIS/Terra Gross Primary Productivity 8-Day
 935 L4 Global 500m SIN Grid V006, NASA EOSDIS Land Processes Distributed Active
 936 Archive Center, [data set], <https://doi.org/10.5067/MODIS/MOD17A2H.006>, 2015.
- 937 Shalaby, A., Zakey, A. S., Tawfik, A. B., Solmon, F., Giorgi, F., Stordal, F., Sillman, S., Zaveri,
 938 R. A., and Steiner, A. L.: Implementation and evaluation of online gas-phase chemistry
 939 within a regional climate model (RegCM-CHEM4), *Geosci. Model Dev.*, 5, 741-760,
 940 <https://doi.org/10.5194/gmd-5-741-2012>, 2012.
- 941 Shang, F., Liu, M. X., Song, Y., Lu, X. J., Zhang, Q., Matsui, H., Liu, L. L., Ding, A. J., Huang,
 942 X., Liu, X. J., Cao, J. J., Wang, Z. F., Dai, Y. J., Kang, L., Cai, X. H., Zhang, H. S., and
 943 Zhu, T.: Substantial nitrogen abatement accompanying decarbonization suppresses
 944 terrestrial carbon sinks in China, *Nature Communications*, 15,



- 945 <https://doi.org/10.1038/s41467-024-52152-5>, 2024.
- 946 Shu, Y., Liu, S., Wang, Z., Xiao, J., Shi, Y., Peng, X., Gao, H., Wang, Y., Yuan, W., Yan, W.,
947 Ning, Y., and Li, Q.: Effects of Aerosols on Gross Primary Production from Ecosystems
948 to the Globe, *Remote Sens.*, 14, <https://doi.org/10.3390/rs14122759>, 2022.
- 949 Sitch, S., Cox, P. M., Collins, W. J., and Huntingford, C.: Indirect radiative forcing of climate
950 change through ozone effects on the land-carbon sink, *Nature*, 448, 791-U794,
951 <https://doi.org/10.1038/nature06059>, 2007.
- 952 Strada, S. and Unger, N.: Potential sensitivity of photosynthesis and isoprene emission to direct
953 radiative effects of atmospheric aerosol pollution, *Atmos. Chem. Phys.*, 16, 4213-4234,
954 <https://doi.org/10.5194/acp-16-4213-2016>, 2016.
- 955 Tang, M. J., Huang, X., Lu, K. D., Ge, M. F., Li, Y. J., Cheng, P., Zhu, T., Ding, A. J., Zhang,
956 Y. H., Gligorovski, S., Song, W., Ding, X., Bi, X. H., and Wang, X. M.: Heterogeneous
957 reactions of mineral dust aerosol: implications for tropospheric oxidation capacity, *Atmos.*
958 *Chem. Phys.*, 17, 11727-11777, <https://doi.org/10.5194/acp-17-11727-2017>, 2017.
- 959 Tu, M., Liu, Z., He, C., Fang, Z., and Lu, W.: The relationships between urban landscape
960 patterns and fine particulate pollution in China: A multiscale investigation using a
961 geographically weighted regression model, *Journal of Cleaner Production*, 237,
962 <https://doi.org/10.1016/j.jclepro.2019.117744>, 2019.
- 963 Unger, N., Yue, X., and Harper, K. L.: Aerosol climate change effects on land ecosystem
964 services, *Faraday Discuss.*, 200, 121-142, <https://doi.org/10.1039/c7fd00033b>, 2017.
- 965 Unger, N., Zheng, Y., Yue, X., and Harper, K. L.: Mitigation of ozone damage to the world's
966 land ecosystems by source sector, *Nature Climate Change*, 10, 134-+,
967 <https://doi.org/10.1038/s41558-019-0678-3>, 2020.
- 968 Wang, J., Dong, J., Yi, Y., Lu, G., Oyler, J., Smith, W. K., Zhao, M., Liu, J., and Running, S.:
969 Decreasing net primary production due to drought and slight decreases in solar radiation
970 in China from 2000 to 2012, *J. Geophys. Res.: Biogeosci.*, 122, 261-278,
971 <https://doi.org/10.1002/2016jg003417>, 2017.
- 972 Wang, K., Zhang, Y., Yahya, K., Wu, S. Y., and Grell, G.: Implementation and initial application
973 of new chemistry-aerosol options in WRF/Chem for simulating secondary organic



- 974 aerosols and aerosol indirect effects for regional air quality, *Atmos. Environ.*, 115, 716-
975 732, <https://doi.org/10.1016/j.atmosenv.2014.12.007>, 2015.
- 976 Wang, X., Wu, J., Chen, M., Xu, X., Wang, Z., Wang, B., Wang, C., Piao, S., Lin, W., Miao,
977 G., Deng, M., Qiao, C., Wang, J., Xu, S., and Liu, L.: Field evidences for the positive
978 effects of aerosols on tree growth, *Global Change Biology*, 24, 4983-4992,
979 <https://doi.org/10.1111/gcb.14339>, 2018.
- 980 Wang, Y., Ni, J., Xu, K., Zhang, H., Gong, X., and He, C.: Intricate synergistic effects between
981 air pollution and carbon emission: An emerging evidence from China, *Environ. Pollut.*,
982 349, <https://doi.org/10.1016/j.envpol.2024.123851>, 2024.
- 983 Wittig, V. E., Ainsworth, E. A., and Long, S. P.: To what extent do current and projected
984 increases in surface ozone affect photosynthesis and stomatal conductance of trees? A
985 meta-analytic review of the last 3 decades of experiments, *Plant Cell and Environment*,
986 30, 1150-1162, <https://doi.org/10.1111/j.1365-3040.2007.01717.x>, 2007.
- 987 Xia, J., Xia, X., Wang, X., Ju, W., Lin, Z., Qin, Z., Sang, Y., Yan, Y., Yuan, W., Yue, X., Zhang,
988 H., Zhou, H., and Zhu, Q.: China Land Carbon Budget (CLCB1.0): a comprehensive
989 estimate of the land carbon budget in China, *National Science Review*, 12,
990 <https://doi.org/10.1093/nsr/nwaf052>, 2025.
- 991 Xie, N., Wang, T., Xie, X., Yue, X., Giorgi, F., Zhang, Q., Ma, D., Song, R., Xu, B., Li, S.,
992 Zhuang, B., Li, M., Xie, M., Kilifarska, N. A., Gadzhev, G., and Dimitrova, R.: The
993 regional climate-chemistry-ecology coupling model RegCM-Chem (v4.6)-YIBs (v1.0):
994 development and application, *Geosci. Model Dev.*, 17, 3259-3277,
995 <https://doi.org/10.5194/gmd-17-3259-2024>, 2024.
- 996 Xie, N. H., Wang, T. J., Xie, M., Ma, D. Y., Zhang, Q., Li, M. M., Li, S., Zhuang, B. L., Kalsoom,
997 U., Kilifarska, N. A., Gadzhev, G., Dimitrova, R., Melas, D., and Karatzas, K.: Carbon
998 Sink of Terrestrial Ecosystems in China During 2010-2020: Spatiotemporal Variability
999 and Climate Impact, *J. Geophys. Res.: Atmos.*, 130,
1000 <https://doi.org/10.1029/2025jd043405>, 2025.
- 1001 Xie, X., Wang, T., Yue, X., Li, S., Zhuang, B., and Wang, M.: Effects of atmospheric aerosols
1002 on terrestrial carbon fluxes and CO₂ concentrations in China, *Atmos. Res.*,



- 1003 237, <https://doi.org/10.1016/j.atmosres.2020.104859>, 2020.
- 1004 Xie, X. D., Wang, T. J., Yue, X., Li, S., Zhuang, B. L., Wang, M. H., and Yang, X. Q.: Numerical
1005 modeling of ozone damage to plants and its effects on atmospheric CO₂ in China, *Atmos.*
1006 *Environ.*, 217, 116970, <https://doi.org/10.1016/j.atmosenv.2019.116970>, 2019.
- 1007 Xue, T., Liu, J., Zhang, Q., Geng, G., Zheng, Y., Tong, D., Liu, Z., Guan, D., Bo, Y., Zhu, T.,
1008 He, K., and Hao, J.: Rapid improvement of PM_{2.5} pollution and associated
1009 health benefits in China during 2013–2017, *Science China-Earth Sciences*, 62, 1847–1856,
1010 <https://doi.org/10.1007/s11430-018-9348-2>, 2019.
- 1011 Yan, S. Q., Zhu, B., Shi, S. S., Lu, W., Gao, J. H., Kang, H. Q., and Liu, D. Y.: Impact of aerosol
1012 optics on vertical distribution of ozone in autumn over Yangtze River Delta, *Atmos. Chem.*
1013 *Phys.*, 23, 5177–5190, <https://doi.org/10.5194/acp-23-5177-2023>, 2023.
- 1014 Yang, H., Chen, L., Liao, H., Zhu, J., Wang, W. J., and Li, X.: Impacts of aerosol-photolysis
1015 interaction and aerosol-radiation feedback on surface-layer ozone in North China during
1016 multi-pollutant air pollution episodes, *Atmos. Chem. Phys.*, 22, 4101–4116,
1017 <https://doi.org/10.5194/acp-22-4101-2022>, 2022.
- 1018 Yu, G., Jia, Y., He, N., Zhu, J., Chen, Z., Wang, Q., Piao, S., Liu, X., He, H., Guo, X., Wen, Z.,
1019 Li, P., Ding, G., and Goulding, K.: Stabilization of atmospheric nitrogen deposition in
1020 China over the past decade, *Nature Geoscience*, 12, 424–+,
1021 <https://doi.org/10.1038/s41561-019-0352-4>, 2019.
- 1022 Yuan, X., Chen, X., Ochege, F. U., Hamdi, R., Tabari, H., Li, B., He, B., Zhang, C., De Maeyer,
1023 P., and Luo, G.: Weakening of global terrestrial carbon sequestration capacity under
1024 increasing intensity of warm extremes, *Nature Ecology & Evolution*, 9,
1025 <https://doi.org/10.1038/s41559-024-02576-5>, 2025.
- 1026 Yue, H., He, C., Huang, Q., Yin, D., and Bryan, B. A.: Stronger policy required to substantially
1027 reduce deaths from PM_{2.5} pollution in China, *Nature Communications*, 11,
1028 <https://doi.org/10.1038/s41467-020-15319-4>, 2020.
- 1029 Yue, K., Peng, Y., Peng, C., Yang, W., Peng, X., and Wu, F.: Stimulation of terrestrial ecosystem
1030 carbon storage by nitrogen addition: a meta-analysis, *Sci. Rep.*, 6,
1031 <https://doi.org/10.1038/srep19895>, 2016.



- 1032 Yue, X. and Unger, N.: The Yale Interactive terrestrial Biosphere model version 1.0: description,
1033 evaluation and implementation into NASA GISS ModelE2, *Geosci. Model Dev.*, 8, 2399-
1034 2417, <https://doi.org/10.5194/gmd-8-2399-2015>, 2015.
- 1035 Yue, X. and Unger, N.: Aerosol optical depth thresholds as a tool to assess diffuse radiation
1036 fertilization of the land carbon uptake in China, *Atmos. Chem. Phys.*, 17, 1329-1342,
1037 <https://doi.org/10.5194/acp-17-1329-2017>, 2017.
- 1038 Yue, X., Zhang, T., and Shao, C.: Afforestation increases ecosystem productivity and carbon
1039 storage in China during the 2000s, *Agric. For. Meteorol.*, 296, 108227,
1040 <https://doi.org/10.1016/j.agrformet.2020.108227>, 2021.
- 1041 Yue, X., Unger, N., Harper, K., Xia, X., Liao, H., Zhu, T., Xiao, J., Feng, Z., and Li, J.: Ozone
1042 and haze pollution weakens net primary productivity in China, *Atmos. Chem. Phys.*, 17,
1043 6073-6089, <https://doi.org/10.5194/acp-17-6073-2017>, 2017.
- 1044 Zhang, F. M., Chen, J. M., Chen, J. Q., Gough, C. M., Martin, T. A., and Dragoni, D.:
1045 Evaluating spatial and temporal patterns of MODIS GPP over the conterminous US
1046 against flux measurements and a process model, *Remote Sensing of Environment*, 124,
1047 717-729, <https://doi.org/10.1016/j.rse.2012.06.023>, 2012.
- 1048 Zhang, Q., Wang, T. J., Wu, H., Qu, Y. W., Xie, M., Li, S., Zhuang, B. L., Li, M. M., and
1049 Kilifarska, N. A.: Radiative and Chemical Effects of Non-Homogeneous Methane on
1050 Terrestrial Carbon Fluxes in Asia, *J. Geophys. Res.: Atmos.*, 129,
1051 <https://doi.org/10.1029/2023jd040204>, 2024.
- 1052 Zhang, Q., Wang, T., Zhang, Z., Xu, X., Xie, N., Zhuang, B., Li, S., Gao, L., Li, M., and Xie,
1053 M.: Methane Emissions in Asian Wetlands During 2010–2020: Insights From an Online-
1054 Coupled Microbial Functional-Group-Based Model, *Earth's Future*, 13,
1055 <https://doi.org/10.1029/2025ef005991>, 2025.
- 1056 Zhang, W. W., Wang, M., Wang, A. Y., Yin, X. H., Feng, Z. Z., and Hao, G. Y.: Elevated ozone
1057 concentration decreases whole-plant hydraulic conductance and disturbs water use
1058 regulation in soybean plants, *Physiologia Plantarum*, 163, 183-195,
1059 <https://doi.org/10.1111/ppl.12673>, 2018.
- 1060 Zheng, B., Tong, D., Li, M., Liu, F., Hong, C. P., Geng, G. N., Li, H. Y., Li, X., Peng, L. Q., Qi,



1061 J., Yan, L., Zhang, Y. X., Zhao, H. Y., Zheng, Y. X., He, K. B., and Zhang, Q.: Trends in
 1062 China's anthropogenic emissions since 2010 as the consequence of clean air actions,
 1063 Atmos. Chem. Phys., 18, 14095-14111, <https://doi.org/10.5194/acp-18-14095-2018>, 2018.
 1064 Zhou, H., Yue, X., Lei, Y., Tian, C., Ma, Y., and Cao, Y.: Aerosol radiative and climatic effects
 1065 on ecosystem productivity and evapotranspiration, Current Opinion in Environmental
 1066 Science & Health, 19, <https://doi.org/10.1016/j.coesh.2020.10.006>, 2021.
 1067 Zhou, H., Yue, X., Lei, Y., Tian, C., Zhu, J., Ma, Y., Cao, Y., Yin, X., and Zhang, Z.:
 1068 Distinguishing the impacts of natural and anthropogenic aerosols on global gross primary
 1069 productivity through diffuse fertilization effect, Atmos. Chem. Phys., 22, 693-709,
 1070 <https://doi.org/10.5194/acp-22-693-2022>, 2022.
 1071 Zhou, H., Yue, X., Dai, H., Geng, G., Yuan, W., Chen, J., Shen, G., Zhang, T., Zhu, J., and Liao,
 1072 H.: Recovery of ecosystem productivity in China due to the Clean Air Action plan, Nature
 1073 Geoscience, 17, <https://doi.org/10.1038/s41561-024-01586-z>, 2024.
 1074 Zhu, J. X., Jia, Y. L., Yu, G. R., Wang, Q. F., He, N. P., Chen, Z., He, H. L., Zhu, X. J., Li, P.,
 1075 Zhang, F. S., Liu, X. J., Goulding, K., Fowler, D., and Vitousek, P.: Changing patterns of
 1076 global nitrogen deposition driven by socio-economic development, Nature
 1077 Communications, 16, <https://doi.org/10.1038/s41467-024-55606-y>, 2025.
 1078
 1079