

Evaluation of a Coupled Regional Reanalysis for the Mediterranean Region Covering the period 1993-2024

Andrea Storto^{1,2}, Vincenzo de Toma^{1,2}, Chunxue Yang^{1,2}

¹Institute of Marine Science (ISMAR), National Research Council (CNR), Rome, I-00133, Rome

²National Research Center for High Performance Computing, Big Data and Quantum Computing (ICSC), Italy

Correspondence to: Andrea Storto (andrea.storto@cnr.it), Chunxue Yang (chunxue.yang@cnr.it)

Abstract. Regional reanalyses provide physically consistent reconstructions of past climate, offering a powerful tool for understanding the processes that shape regional responses to climate change. Here, we introduce MESMAR-R (Mediterranean Earth System Model at ISMAR - Reanalysis), a newly developed coupled atmosphere-ocean-hydrology reanalysis for the Mediterranean region. The system couples state-of-the-art numerical models with data assimilation in both the ocean and the atmosphere, producing a dynamically coherent reconstruction of the 1993-2024 period. MESMAR-R shows strong performance in the ocean, particularly for sea level, sea-surface salinity, and upper-layer temperature, confirming its skill in the most relevant layers for air-sea interactions. The atmospheric component performs reasonably well given its simplified assimilation based on three-dimensional scale-selective nudging, yielding basin-scale fields comparable to state-of-the-art reanalyses and maintaining internally consistent coupled air-sea interactions. Sensitivity experiments show that data assimilation in both the ocean and atmosphere is essential for optimal performance: coupling alone does not necessarily improve reanalysis skill, but its benefits are relevant for salinity and surface fluxes and amplify when complemented by data assimilation. The ocean and atmospheric trends reconstructed by MESMAR-R are physically coherent and consistent with observations, reproducing basin-wide salinification, thermo-halo compensation in steric sea level, mass-driven sea-level rise, and pronounced near-surface atmospheric warming and drying. MESMAR-R therefore provides a unique, balanced reconstruction of past Mediterranean climate, suitable for investigating coupled air-sea processes, compound events, marine extremes, and other phenomena where dynamically consistent atmosphere-ocean interactions are essential.

Short Summary. The Mediterranean region is one of the fastest-warming areas on Earth, and understanding its changes is thus crucial. We produced a new reconstruction of the Mediterranean climate for the 1993-2024 period, using a modelling system that combines weather, ocean, and hydrology information with observations. This new reanalysis offers a clearer picture of how the region has been warming, drying, and changing over the past 30 years, supporting better studies of extreme events and climate impacts.

30 1. Introduction

31 Reanalysis datasets provide temporally and spatially continuous representations of the Earth system, or any of its components,
32 by combining numerical model simulations with historical observations through data assimilation techniques (e.g., Storto et
33 al., 2019a). These datasets are essential for climate monitoring, model validation, and understanding past climate variability.
34 While global reanalyses, such as ERA5 (Hersbach et al., 2020) and MERRA-2 (Gelaro et al., 2017) in the atmosphere or GREP
35 (Storto et al., 2019b) in the ocean, offer broad coverage, their limited horizontal resolution may hinder their ability to accurately
36 represent mesoscale processes, especially in regions with complex geography such as the Mediterranean, which is the focus
37 of the present study. The Mediterranean basin is widely recognized as a climate change hotspot, experiencing rates of warming
38 exceeding the global average and undergoing rapid changes in hydro-climatic extremes (Cramer et al., 2018; Lionello and
39 Scarascia, 2018). Projected increases in temperature, decreases in precipitation, and enhanced variability in oceanographic and
40 atmospheric conditions pose serious challenges for ecosystems, water resources, and human activities across the region.

41 In contrast to global products, high-resolution regional climate models (RCMs) have been developed to dynamically downscale
42 global reanalyses and provide a refined depiction of climate processes at local to regional scales (e.g., Ruti et al., 2016). These
43 regional reanalyses are particularly important in areas with significant land-sea contrasts, steep topography, and densely
44 populated coastal zones, where the impact of climate extremes can be severe (e.g., Senatore et al., 2020; Fox-Hughes et al.,
45 2022).

46 The importance of regional reanalyses is testified by their increasing development and adoption. The Copernicus Marine
47 Service provides a Mediterranean ocean reanalysis (e.g., Escudier et al., 2021), while the Copernicus Climate Service has
48 developed European-scale regional atmospheric reanalyses (e.g., CERRA; Ridal et al., 2024). However, to date, most available
49 products are either uncoupled or limited to single components: ocean-only or atmosphere-only. Furthermore, some downscaled
50 reanalyses lack any form of data assimilation, limiting their utility for monitoring and decision-making. Regional coupled
51 reanalyses over the Mediterranean Sea are expected to become critically important for climate services and downstream
52 applications because they provide a physically consistent, high-resolution reconstruction of the complex air-sea-land
53 interactions, enabling improved assessment of marine heatwaves (e.g., Camus et al., 2021), storm surge and coastal flooding
54 risks (e.g., Juza et al., 2022), which directly support sectors like fisheries management, renewable energy planning, and coastal
55 adaptation strategies (e.g., Soukissian et al., 2021). Furthermore, regional coupled reanalyses will be increasingly vital for
56 training deep-learning emulators (Doury et al., 2023) because they offer high-resolution, physically consistent input-output
57 datasets that capture air-sea-land feedback, enabling efficient and accurate surrogate modeling of regional climate processes
58 for applications such as bias correction, super-resolution downscaling, and rapid scenario exploration (e.g., van der Meer et
59 al., 2023).

60 Coupled reanalysis systems, which combine ocean and atmosphere models with concurrent assimilation, offer a more
61 physically consistent reconstruction of the Earth system by representing feedback across components. While global coupled
62 reanalyses such as CERA-20C (Laloyaux et al., 2018) have demonstrated the value of this approach, no such product has been
63 available at the regional scale for the Mediterranean. Recent studies highlight the potential of coupled regional systems to
64 improve the representation of events like medicanes (Storto et al., 2023).

65 To fill this gap, we present a new high-resolution regional coupled reanalysis for the Mediterranean Sea over the period 1993-
66 2024, with an approximately 15 km atmospheric grid and an approximately 7 km ocean grid with 72 vertical levels. The system
67 is based on the MESMAR (Mediterranean Earth System Model at ISMAR) coupled model configuration (Storto et al., 2023),
68 which integrates the WRF atmospheric model and the NEMO ocean model, with interactive river runoff and land surface

modelling components. Ocean data assimilation is implemented via a 3D variational (3DVAR) scheme, while the atmosphere is constrained at the large scale, through spectral nudging, toward ERA5. The configuration thus ensures that large-scale consistency with global reanalyses is preserved while allowing the model to resolve regional-scale features (Storto et al., 2023). The system assimilates in-situ and satellite observations in the ocean, providing a dynamically consistent estimate of the Mediterranean climate state across the atmosphere and ocean components.

This paper introduces the dataset (section 2) and presents a first evaluation against many observational datasets (in section 3). Additionally, comparisons with existing regional and global reanalyses provided by Copernicus Marine and Climate Services are included for selected metrics, together with an analysis of sensitivity experiments (section 4) and key climatic trends (section 5). The coupled nature of this reanalysis and the inclusion of data assimilation in both ocean and atmosphere, although the latter as large-scale constraint, represent a significant advancement in regional Earth system reanalyses for the Mediterranean region.

2. Data and Methods

2.1 Earth System Model configuration

The MESMAR modeling system couples three components: an atmosphere (WRF), ocean (NEMO), and hydrology (HD) models, via the OASIS3-MCT coupler, and it is described in detail in Storto et al. (2023). The atmospheric core is WRF (v4.3.3, Skamarock et al., 2021), implemented over the Euro-Mediterranean domain (~15 km horizontal grid, 41 hybrid vertical levels). WRF is a non-hydrostatic model with a 60-second timestep. Physics options include Thompson microphysics, RRTMG radiation, the Noah-MP land surface model (4 soil layers), the Mellor-Yamada (Nakanishi and Niino 2006), PBL closure, and the Grell-Freitas cumulus scheme for subgrid convection. Lateral atmospheric forcing is provided via a 10-point sponge layer relaxing toward three-hourly ERA5 reanalysis fields (Hersbach et al., 2020).

The ocean component is NEMO (v4.0.7, Madec et al., 2017) covering the Mediterranean (Atlantic box to Dardanelles) on a regular ~7 km grid with 72 vertical levels (with partial steps at bottom). The NEMO baroclinic timestep is 450s, while the barotropic timestep is 6s, with the split-explicit free-surface formulation. Shortwave radiation follows a three-band scheme (Morel and Maritorena, 2001) using monthly satellite chlorophyll data (Brewin et al., 2015). Horizontal mixing uses a Laplacian diffusion (viscosity via bi-Laplacian) with coefficients ~80 m²/s (4.5×10^9 m⁴/s³), increased by 20% near Gibraltar and the Aegean. Vertical mixing is handled by a General Length Scale (GLS) scheme (Mellor-Yamada, 1982; closure with Canuto et al., 2001, stability). Ocean boundaries are forced by the ECMWF ORAS5 reanalysis (Zuo et al., 2019) with Flather barotropic conditions and prescribed baroclinic velocities, and a 10-point relaxation of temperature and salinity toward ORAS5 fields. River discharge from the HD model is injected at coastlines; the Black Sea (Dardanelles) outflow uses climatology (Kourafalou & Barbopoulos, 2003). No additional basin-wide volume-restoration term is imposed, thus Mediterranean volume changes arise from the split-explicit free-surface equation, the net surface freshwater flux, river input, and the net transport through the open boundaries. The astronomical tidal forcing is not included in the present NEMO configuration.

Interactive rivers are provided by the hydrological discharge model (HD v5.1, Hagemann et al., 2020). HD is based on the MPI hydrology scheme (MPI-HM; Hagemann & Dümenil Gates, 2001) with specialized developments for coupling. It runs at 0.5° resolution over Europe, with a 30-minute timestep. To improve stability, HD implements a conservative smoothing of river outflow: each discharge is spread to a 5×5 grid around the target ocean point. HD collects surface and subsurface runoff from WRF, and routes it to NEMO coastal points via the coupler.

The coupling between the models is handled by OASIS3-MCT (v5.0, Craig et al., 2017). All exchanged fields are mapped with first-order conservative interpolation, at 30-minute intervals. Key coupling variables include surface fluxes of momentum, heat, and freshwater (computed in WRF using the Janjić surface scheme), sea surface temperature (SST) and salinity (SSS) from NEMO, surface pressure from WRF (for inverse-barometer correction), and river discharge from HD to NEMO. Thus, air-sea fluxes are generated in the atmosphere and passed to the ocean, while SST and currents and atmospheric pressure influence the atmosphere and ocean, respectively, and runoff links hydrology and ocean. Air-sea momentum, heat, and freshwater fluxes are indeed computed in WRF to ensure consistency with its surface-layer formulation, PBL closure, and implicit vertical turbulent-diffusion scheme. The shorter atmospheric timestep also better resolves rapid boundary-layer and diurnal adjustments; fluxes are then accumulated over the coupling interval and conservatively remapped to NEMO. Computing them on the finer ocean grid would instead require interpolation of the atmospheric state and a separate bulk formulation, without adding atmospheric information. Temporally synchronized atmospheric fluxes also reduce coupling-splitting and solar-radiation phase errors (Marti et al., 2021; Schüller et al., 2025).

2.2 Data assimilation

MESMAR-R employs a weakly coupled data assimilation system; details of the formulation are provided in Appendix A, while here we summarize the main features. The ocean uses a 3DVAR scheme (Storto et al., 2018) updated every 3 days. Background-error covariances are derived from multi-year monthly mean anomalies with respect to the long-term climatology (1998-2021), from a previous iteration of the coupled reanalysis. Observations include subsurface profiles (XBT, CTD, moored instruments, Argo/glider floats) from the UK Met Office EN4 dataset (Good et al., 2013), with variational quality control enabled (Storto, 2016), and along-track altimetry data from CMEMS (Pujol et al., 2023), after the standard geophysical and tidal corrections, assimilated with a dynamic height observation operator (Storto et al., 2011) and the latest Mediterranean Sea Mean Dynamic Topography (MDT) from the Copernicus Marine Service (Jousset et al., 2025). A weak surface relaxation of SST and SSS is applied toward reference analyses (CNR-ISMAR SST and EN4 SSS), with e-folding times of ~ 15 days (SST) and 300 days (SSS). Additionally, a large-scale bias correction scheme, as in Storto and Yang (2024), is applied to minimize drifts in the deep ocean layers.

Atmospheric assimilation is achieved by spectral nudging in the WRF model. Using spectral nudging is appropriate for a 32-year coupled downscaling because it provides a computationally affordable and temporally homogeneous large-scale constraint from ERA5, while allowing atmospheric small scales to evolve according to the regional model dynamics, physics, and air-sea feedbacks. Nudging also avoids the need for a full observation collection and re-processing, which is extremely demanding for a long reanalysis exercise. Large-scale (synoptic) components of wind, temperature, and humidity are continuously relaxed toward three-hourly ECMWF ERA5 fields. Scales longer than ~ 850 km (wavenumbers ≤ 6 zonally and ≤ 5 meridionally) are constrained, with nudging timescales of order 1 h for winds and temperature, and 1 day for humidity. This approach maintains large-scale consistency with the ERA5 reanalysis (Hersbach et al., 2020), while mesoscale features can develop freely (in both the atmosphere and the ocean, like medicanes, see Storto et al., 2023). Tests with the MESMAR configuration showed that spectral nudging better retained the effective regional resolution than full-field nudging and provided improved coupled ocean and medicane behavior (Storto et al., 2023).

2.3 Changes with respect to the pilot version

The MESMAR-R reanalysis is an improved version compared to a first, unpublished, iteration (covering the period 1998-2022), with a few upgrades that mostly concern the ocean component. Besides the temporal extension, the present version now includes i) the assimilation of altimetry data in the ocean data assimilation system; ii) a different set of background-error

covariances, calculated from the long-term mean anomalies of the previous reanalysis iteration, thus better representing the error structures of the coupled reanalysis and leading to better verification skill scores; iii) a deep-ocean large-scale bias correction scheme. Additionally, the NEMO model now uses XIOS 2.5 (Yepes-Arbós et al., 2022) for input/output operations, permitting a more flexible choice of ocean output fields.

2.4 Experiments

In addition to the nominal reanalysis setup (MESMAR-R), covering the 32 years from 1993 to 2024, we have also performed a few other sensitivity experiments, mostly devoted to understanding the relative impact of data assimilation and coupling for the enhancement of the reanalysis skill scores and key processes, with emphasis on the ocean component of the coupled reanalysis. Experiments are summarized below, and are all run for the shorter period from 2000 to 2020.

We performed a set of experiments switching on/off the data assimilation in both components and the coupling to the atmosphere. Such a set allows us to disentangle the effects of ocean data assimilation or atmosphere relaxation and the coupling, on key oceanic skill scores and metrics that will be introduced in section 4. Note that, in UOS and UOA, the atmospheric forcing is not interactive and provided by the ERA5 reanalysis, meaning that it differs from COS and COA experiments not only by the coupling on/off, but also by the underlying atmospheric forcing source, and the bulk flux formulation.

<i>NAME</i>	<i>Atm</i>	<i>Oce</i>	<i>Purpose</i>	<i>Period</i>
MESMAR-R	Data Assimilation	Data Assimilation	Coupled Reanalysis	1993-2024
COA	Simulation	Data Assimilation	Impact of atmospheric data assimilation	2000-2020
COS	Simulation	Simulation	Coupled Simulation	2000-2020
UOS	Not present	Simulation	Forced ocean Simulation	2000-2020
UOA	Not present	Data Assimilation	Forced ocean Reanalysis	2000-2020

Table 1. List of reanalysis experiments performed. *Simulation* means that no data assimilation is employed.

2.5 Verification data and methods

Several datasets were used for verifying the reanalysis against. We used in particular:

i) gridded sea surface temperature, salinity, and sea level (absolute dynamic topography, ADT) data from the Copernicus Marine Service (L4 ADT, Ballarotta et al., 2013, from which also EKE is calculated; L4 SST, Pisano et al., 2020; L4 SSS, Sammartino et al., 2022);

ii) in-situ profiles extracted from the UK MetOffice EN4 dataset (Good et al., 2013);

iii) e-OBS gridded near-surface atmospheric variables from the Copernicus Climate Change Service (Haylock et al., 2008). Here, we used monthly means from version 30.0e at 0.1° of spatial resolution, and all model data were compared to e-OBS after interpolation onto the e-OBS native grid;

iv) Mixed layer depth (MLD) data from the SeaDataCloud (SDC_MED_DP1);

v) ocean heat content from the Copernicus Marine Service Ocean Monitor Indicator (von Schuckmann et al., 2018);

vi) precipitation over ocean from the Global Precipitation Climatology Project (GPCP, Adler et al., 2003); note that, while e-OBS provide near-surface atmospheric variables over land only, verification of precipitation is seamless over land and ocean because of the joint use of e-OBS and GPCP.

Some of these datasets are not independent, strictly speaking, as they are assimilated into the system. However, we use lower-frequency based metrics for SST, SSH, and SLA validation (monthly), while for in-situ profiles, statistics are computed from innovation data before the assimilation of the observations. RMSEs in section 3 are calculated from the daily mean fields; for altimetry, given the fact that each product may use different MDTs, the long-term mean SSH from each experiment is subtracted, meaning in facts that RMSE corresponds to the standard deviation of the innovations.

As reference datasets to compare skill score metrics, we used the Mediterranean Sea ocean reanalysis MEDREA24 (Escudier et al., 2021) and the Mercator Ocean International global ocean reanalysis GLORYS12 (Lellouche et al., 2021) for the ocean component, and the ECMWF ERA5 reanalysis (Hersbach et al., 2020) and CERRA (Ridal et al., 2024) for the atmosphere. All these reanalyses represent state-of-the-science systems, either with global or regional focus.

Note that MEDREA24 and the UOA experiment (Table 1) are both uncoupled, NEMO-based ocean reanalyses, but they are not equivalent configurations. MEDREA24 uses NEMO v3.6 at approximately $1/24^\circ$ horizontal resolution with 141 vertical levels and the Mediterranean OceanVar implementation described by Escudier et al. (2021). There are several differences with our UOA configuration: beside model version and spatial resolution, bulk formulas, lateral boundary conditions and frequency of input atmospheric forcing, setting and estimates of background-error covariances, river runoff specification, vertical mixing and advection schemes, use of sea surface relaxation, all contributing significantly to the differences between the two datasets, whose detailed investigation is beyond the scope of the present manuscript. For strict temporal consistency, all intercomparisons that include MEDREA24 are evaluated over the common period 2000–2019; the 2000–2020 period is used only where all compared datasets are available.

Climate diagnostics presented later mostly consist of linear trends over the full period. Total trends are calculated from the monthly mean data. Additionally, for ocean diagnostics, we subdivide the Mediterranean Sea into the western Mediterranean Sea (WMED) and the eastern Mediterranean Sea (EMED), divided by a straight line across the Strait of Sicily (at 37°N from Tunisia to Sicily).

3. Validation

In this section, we revise basic validation metrics from the MESMAR-R reanalysis and compare them with reference reanalysis datasets, introduced in section 2.5, for comparison. Table 2 shows the RMSE for ocean parameters (temperature and salinity for different depth ranges, and sea surface temperature, salinity, and height), over the 2000-2020 period.

The MESMAR-R reanalysis shows particularly strong skill at the sea surface. Compared with the regional MEDREA24 and the global GLORYS12, it delivers the lowest RMSE for sea surface height, with values slightly better than GLORYS12 and markedly better than MEDREA24. Similarly, MESMAR-R clearly outperforms both competitors in sea surface salinity, where the error is reduced by more than half compared with GLORYS12. For sea surface temperature, MESMAR-R is nearly indistinguishable from MEDREA24, both substantially outperforming GLORYS12. These results highlight the robustness of the coupled reanalysis in capturing surface variability, which is critical for mesoscale dynamics, air-sea fluxes, and the freshwater balance of the Mediterranean.

In the upper ocean, MESMAR-R maintains a strong performance. Temperature in the 20–60 m layer shows the lowest RMSE among all products, while at 0–20 m the skill is essentially comparable to MEDREA24 and better than GLORYS12. Salinity

in the upper 210 m is well represented, too, with MESMAR-R either leading or effectively tied for the best performance. Below 210 m, MEDREA24 generally yields slightly lower RMSE than MESMAR-R, particularly for temperature between 210 and 600 m. At deeper layers (400-1000 m), differences among the three systems are small, and MESMAR-R typically places second behind MEDREA24 but ahead of GLORYS12.

Looking at profiles of bias (Figure 1), all three reanalyses show a small fresh bias in the upper layers, most pronounced in the top 50 m, but MESMAR-R (blue) exhibits the smallest amplitude at the sea surface. Below about 100 m, all products converge toward near-zero bias. For temperature, the three systems exhibit a weak warm bias near the surface that decreases with depth. MESMAR-R shows a bias magnitude comparable to or slightly smaller than MEDREA24 above 100 m, and both regional reanalyses outperform GLORYS12, which displays a more persistent warm bias extending deeper into the thermocline. Below 200 m, MESMAR-R's bias remains very close to zero, suggesting good thermal consistency with in-situ observations and minimal drift at depth.

<i>Metric</i>	<i>MESMAR-R</i>	<i>MEDREA24</i>	<i>GLORYS12</i>
S0-20 T0-20 (psu °C)	0.222 1.189	0.254 1.175	0.367 1.227
S20-60 T20-60 (psu °C)	0.209 1.087	0.228 1.131	0.229 1.208
S60-210 T60-210 (psu °C)	0.113 0.417	0.114 0.358	0.123 0.402
S210-400 T210-400 (psu °C)	0.054 0.261	0.048 0.210	0.055 0.240
S400-600 T400-600 (psu °C)	0.037 0.173	0.033 0.145	0.041 0.186
S600-1000 T600-1000 (psu °C)	0.027 0.104	0.025 0.091	0.028 0.105
SST (°C)	0.342	0.331	0.490
SSS (psu)	0.314	0.331	0.784
SSH (cm)	3.4	5.3	3.5

Table 2. Root mean square error for the ocean, in comparison with other state-of-the-art reanalyses. The parameters include temperature (T, in °C) and salinity (S, in psu), sea surface temperature (SST, in °C), sea surface salinity (SSS, in psu), and sea surface height (SSH, in m). The calculation covers the period 2000-2020.

The spatial distribution of RMSE (Figure 2) confirms the overall strengths of MESMAR-R, while also highlighting regional differences. For sea surface temperature (SST), the absolute RMSE map shows widespread errors below 0.5 °C across most of the basin, with slightly elevated values along the northern Adriatic and Aegean Seas, and locally in the western basin. The difference plots (middle and right panels of Figure 2) indicate that MESMAR-R generally performs as good as MEDREA24, with small areas of improvement in the Levantine Basin and central Mediterranean, but slightly higher errors along parts of the western basin and, locally, elsewhere. In contrast, MESMAR-R systematically outperforms GLORYS12 almost everywhere, with particularly marked improvements in the Adriatic and Aegean Sea, the Levantine Basin, and the Alboran Sea.

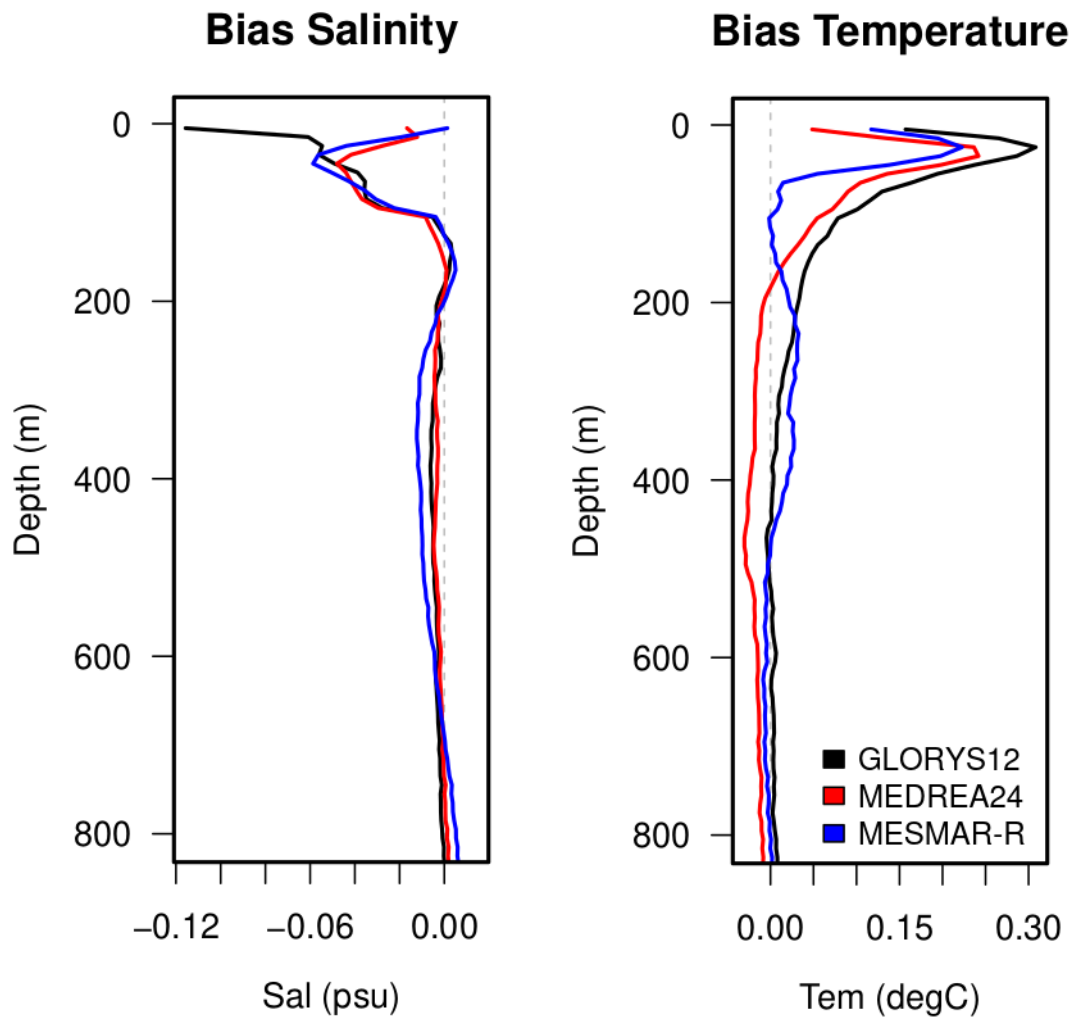
For sea surface salinity (SSS), MESMAR-R achieves basin-wide errors below 0.5 psu except for localized peaks in the northern Aegean. Compared with MEDREA24, there are uniform and widespread improvements, with peaks near the Po river mouth and local deterioration in proximity to the Dardanelles, likely amplified by differences in the boundary conditions therein. The localized difference near the Po River mouth is most plausibly linked to the different river-discharge specifications and coastal

240 mapping used by the systems, rather than any possible indirect effect of the coupled data assimilation, as suggested also by
 241 pure simulation runs (see Storto et al., 2023). Against GLORYS12, MESMAR-R demonstrates a clear gain almost everywhere,
 242 with the largest improvements in the northern Adriatic and Aegean seas, likely driven by the better river discharge specification
 243 than in the global reanalysis.

244 The sea surface height (SSH) comparison is especially promising, as MESMAR-R exhibits low RMSE values across the entire
 245 basin, typically under 0.04 m. The difference maps confirm that MESMAR-R consistently outperforms MEDREA24
 246 throughout the Mediterranean, with the most pronounced improvements in the western basin and Levantine Sea. Compared
 247 with GLORYS12, MESMAR-R shows a systematic advantage in the eastern Mediterranean, while locally GLORYS12 tends
 248 to outperform MESMAR-R.

249

250

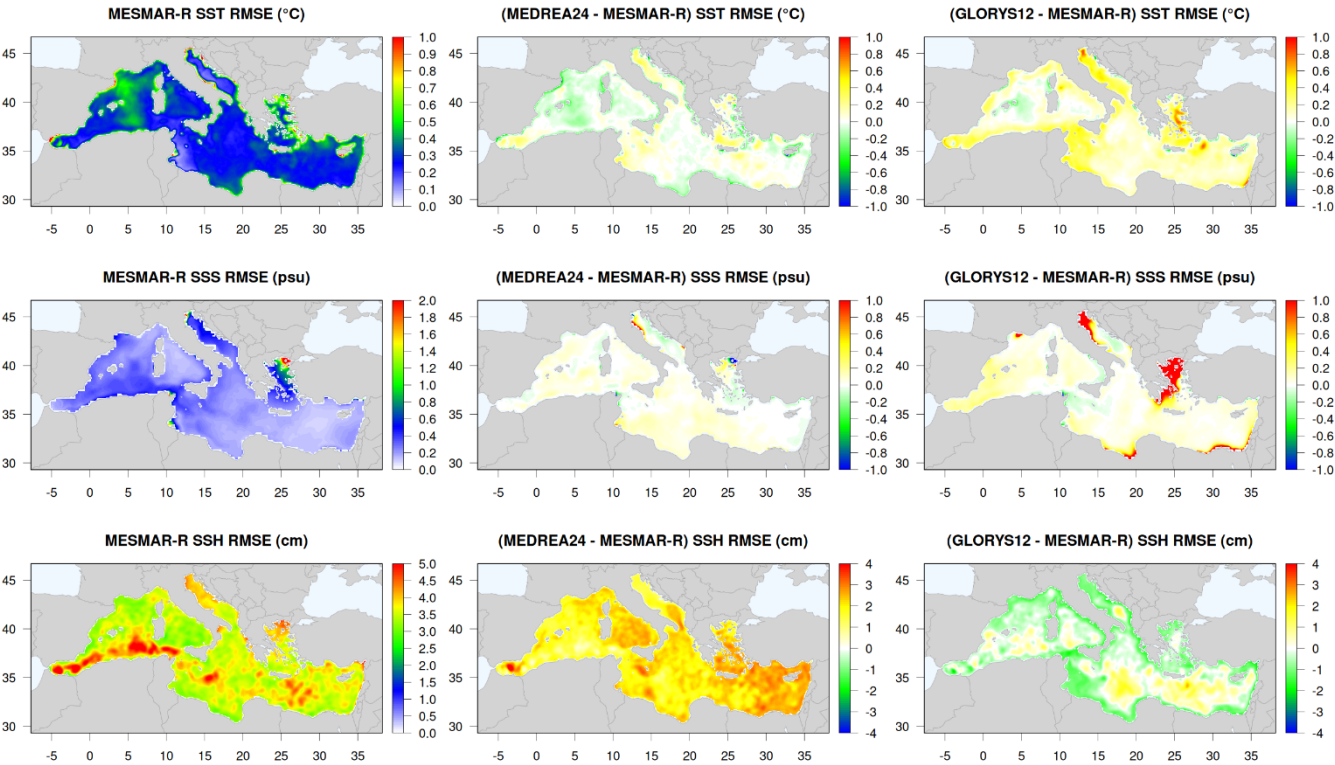


251

252 **Figure 1. Profiles of salinity (left panel) and temperature (right panel) bias, calculated against all available in-situ**
 253 **profiles, for MESMAR-R in comparison with other state-of-the-art reanalyses. The calculation covers the period 2000-**
 254 **2020, and the bias is defined as model minus observations.**

255

256
257



258
259
260
261
262
263
264
265

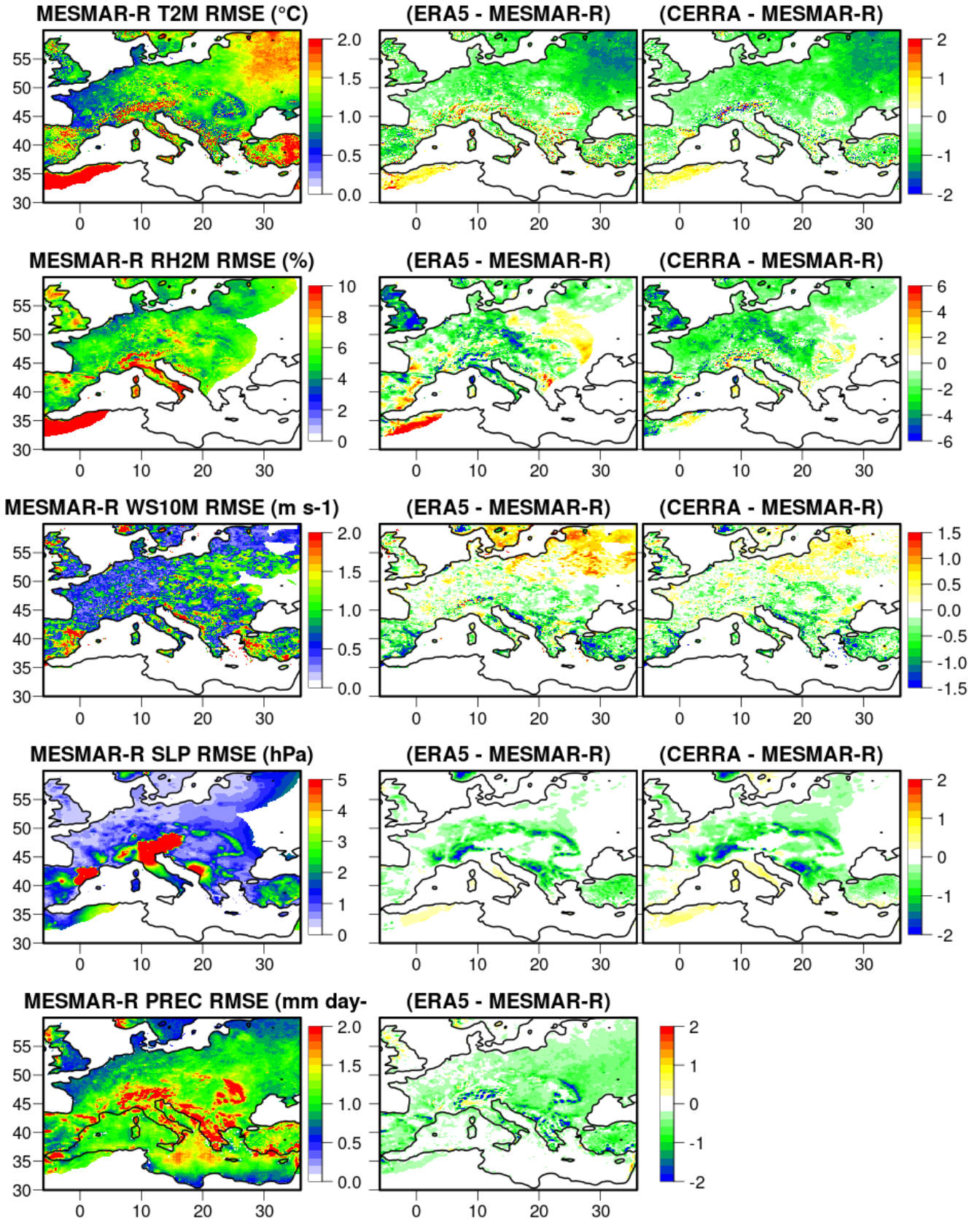
Figure 2. Maps of RMSE against sea surface observations (SST in °C, SSS in psu, and SSH in cm), in comparison with other state-of-the-art reanalyses. Left panels show the MESMAR-R RMSE, while middle and right panels show the RMSE difference between either MEDREA24 or GLORYS12 and MESMAR-R (where positive, MESMAR-R provides the smallest RMSE). The calculation covers the period 2000-2020.

<i>Metric</i>	<i>MESMAR-R</i>	<i>ERA5</i>	<i>CERRA</i>
T2M (°C)	1.38	1.24	1.14
RH2M (%)	7.63	7.26	6.61
MSLP (hPa)	2.61	2.35	2.38
WS10M (m s⁻¹)	0.91	0.64	0.67
PRECIP-LAND (mm/day)	1.37	0.89	-
PRECIP-OCE (mm/day)	1.17	1.06	-

Table 3. Root mean square error for selected atmospheric parameters, in comparison with other state-of-the-art reanalyses. T2M: 2m air temperature; RH2M: 2m relative humidity; MSLP: mean sea level pressure; WS10M: wind speed at 10m; PRECIP: total precipitation (either over land or ocean).

Moving to atmospheric parameter validation, area-averaged scores show MESMAR-R generally has slightly higher RMSE than ERA5 or CERRA for near-surface variables (Table 3, see the caption therein for the abbreviation of parameters): T2M (1.38 °C vs 1.24/1.14), RH2M (7.63% vs 7.26/6.61), WS10M (0.91 m s⁻¹ vs 0.64/0.67), and MSLP (2.61 hPa vs 2.35/2.38). For precipitation, the gap is larger over land (1.37 vs 0.89 mm day⁻¹ to ERA5) and modest over ocean (1.17 vs 1.06 mm day⁻¹), where evaluation uses GPCP. MESMAR-R is worse overall but still within the same order of magnitude as the state-of-the-art reanalyses. Note that the comparison may also be affected by representativeness inaccuracies, due to the smoothing procedure in e-OBS and the different spatial resolutions of data and models; yet the comparison provides a qualitative indication of the model-data agreement.

The maps and RMSE-difference panels, in Figure 3, confirm the picture. For all variables, most of the domain shows ERA5/CERRA with equal or lower RMSE than MESMAR-R (negative differences), but MESMAR-R achieves localized improvements (positive differences). The fine-scale RMSE-difference patterns contain both positive and negative values; they are concentrated near complex coastlines and orography, where representativeness differences between a 15 km regional model and gridded observations are large and where the WRF surface-layer, land-surface, boundary-layer, and precipitation schemes may generate gradients that are absent or smoothed in the reference datasets. Over the sea, coupled SST and surface-current feedback can additionally modify near-surface temperature, humidity, and wind, while spectral nudging constrains only long scales and therefore permits mesoscale departures from ERA5. A large cold bias near Siberia affects the skill scores for MESMAR-R in the eastern part of the domain, due to suboptimal land model performances therein (Storto et al., 2023); for WS10M, the northern part of the domain and scattered marine areas exhibit marginal gains; MSLP shows very narrow bands in the southernmost part of the MESMAR domain, where MESMAR-R reduces error relative to ERA5/CERRA; and for precipitation, ERA5 provides the best skill scores almost everywhere. Taken together, the maps indicate MESMAR-R worsens most of the time, but can outperform locally, typically in small, dynamically or observationally challenging sub-regions.



294

295

296

297

298

Figure 3. Maps of RMSE against atmospheric near-surface gridded observations (T2M in $^{\circ}\text{C}$, RH2M in %, WS10M in m s^{-1} , SLP in hPa, and PREC in mm day^{-1}), in comparison with other state-of-the-art reanalyses. Left panels show the MESMAR-R RMSE, while middle and right panels show the RMSE difference between either ERA5 or CERRA and MESMAR-R (where positive, MESMAR-R provides the smallest RMSE). The calculation covers the period 2000-2020.

299

300

301

4. Sensitivity to coupling and data assimilation

In this section, we compare the MESMAR-R reanalysis with the sensitivity experiments described in section 2.4, to assess the relative merits of data assimilation and interactive coupling in the MESMAR-R performances. This section includes both a comparison of skill scores and key oceanic processes.

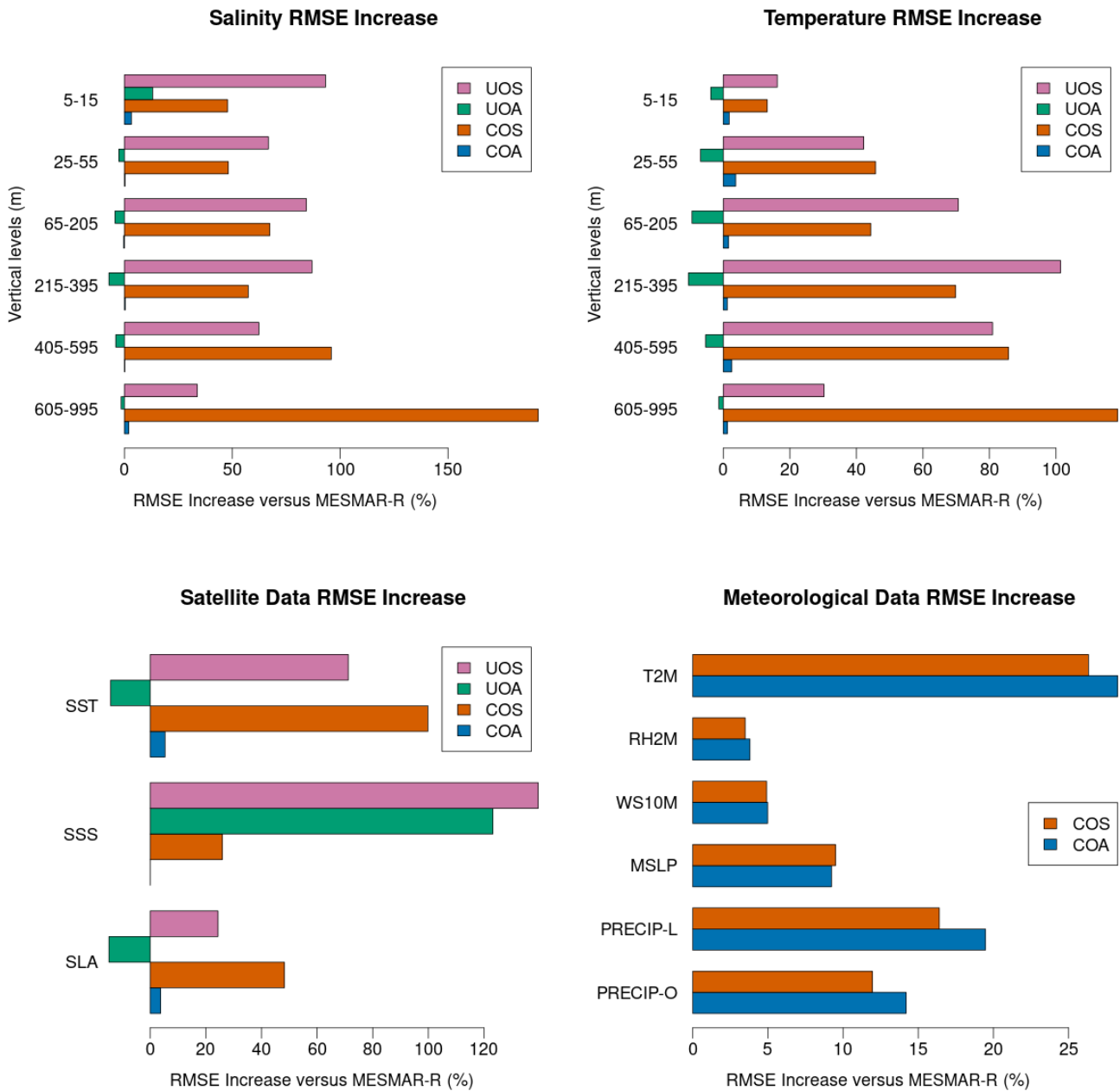


Figure 4. RMSE increase (in %) compared to MESMAR-R for the experiments described in Table 1, for several oceanic and atmospheric parameters.

The sensitivity experiments confirm the expected hierarchy among configurations in the comparison of RMSE (see the RMSE increase bar plots in Figure 4). The fully coupled free simulation (COS) consistently shows the largest RMSE increases,

particularly at depth, indicating that the absence of any data assimilation in coupled mode leads to a rapid degradation of the subsurface structure. In contrast, the uncoupled ocean simulation (UOS) performs better than COS but is still significantly worse than MESMAR-R, reflecting the stabilizing influence of the ERA5 atmospheric forcing, which effectively acts as a weak atmospheric constraint. Interestingly, for some parameters (most notably sea surface salinity (SSS) and upper-ocean salinity), the coupling provides a measurable improvement, suggesting that the explicit exchange of freshwater fluxes helps preserve a more realistic surface salinity balance. This may also be due to the different bulk formulas employed in the coupled configuration compared to the NEMO uncoupled setup, although it is not technically possible to disentangle the two changes.

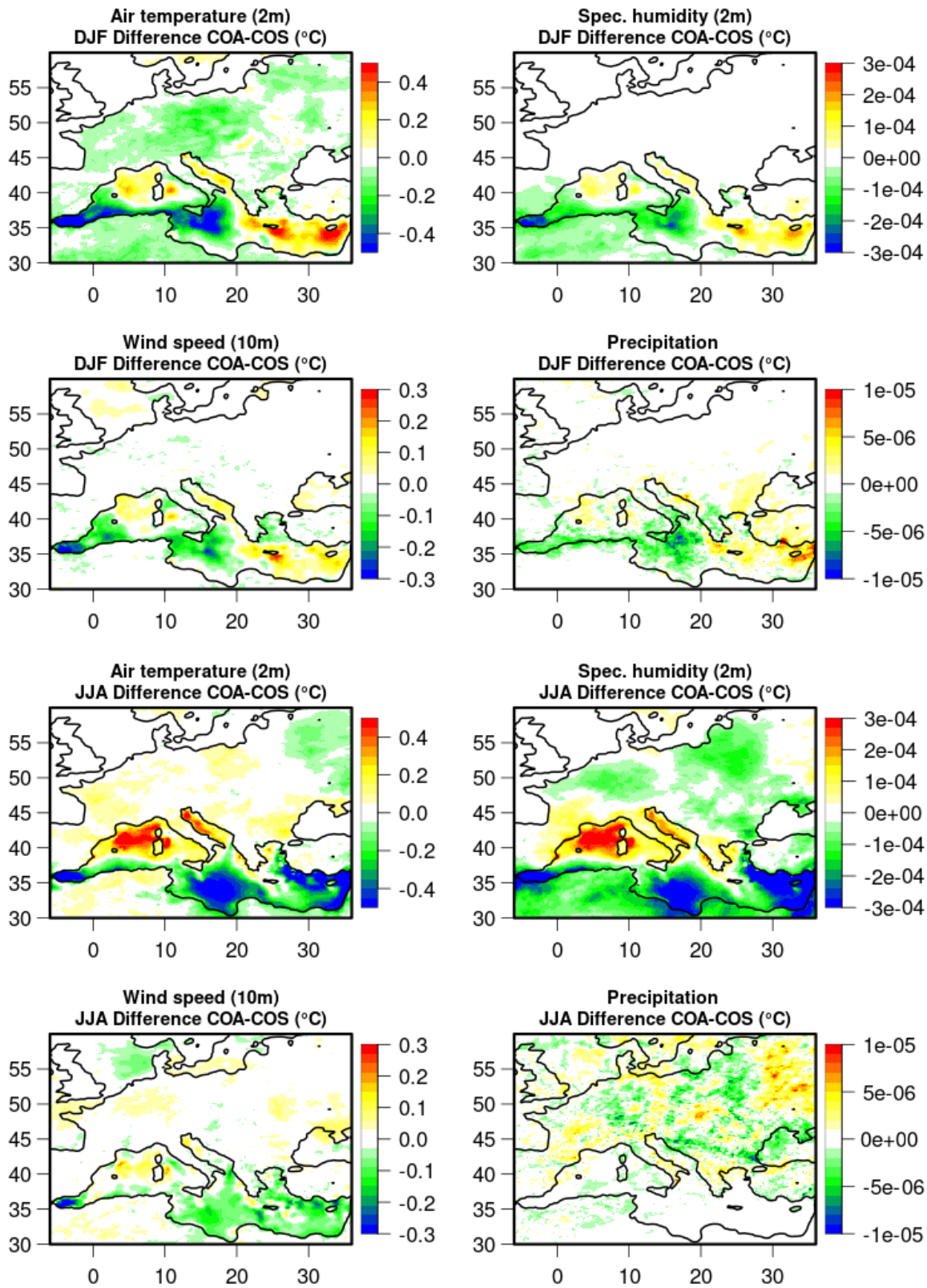
The uncoupled ocean reanalysis (UOA) exhibits slight overall improvements compared with MESMAR-R, except for SSS and salinity in-situ verification in the top 15m, where the coupled configuration performs marginally better, as explained earlier. This outcome indicates that the combination of ocean data assimilation and prescribed atmospheric reanalysis forcing tends to outperform the coupled system when coupling errors outweigh flux benefits, except for metrics directly linked to freshwater closure. Conversely, the coupled ocean analysis (COA), which retains ocean data assimilation but excludes atmospheric assimilation, shows small but statistically significant degradations in a few key parameters, namely temperature near the base of the mixed layer, sea level anomaly (SLA), and sea surface temperature (SST). These changes imply that the atmospheric large-scale constraint employed in MESMAR-R exerts a positive influence on the air–sea interface, improving consistency in surface heat and momentum fluxes.

The verification against meteorological observations (Figure 4, bottom right panel) reveals moderate RMSE increases for both coupled configurations (COS and COA) compared with the full reanalysis. The largest degradation occurs for near-surface air temperature (T2M), where RMSE increases exceed 20%, followed by precipitation, particularly over land (PRECIP-L), where errors rise by about 15–20%. In contrast, relative humidity (RH2M) and 10 m wind speed (WS10M) show only minor increases of a few percent, while mean sea level pressure (MSLP) exhibits intermediate sensitivity (around 10%). Notably, COS generally performs slightly better than COA for several parameters, especially T2M and precipitation; the COA-COS differences, summarized in Figure 5, are concentrated mainly over the Mediterranean Sea and broadly follow the SST adjustments induced by ocean data assimilation, with the clearest responses in 2 m temperature and specific humidity and generally weaker changes in 10 m wind speed. SST anomalies modify the air–sea temperature and moisture contrasts and lower-atmospheric stability, thereby affecting turbulent momentum mixing and producing locally coherent wind-speed changes. The signal occasionally extends inland, particularly for winter air temperature, summer humidity, and precipitation in both seasons, indicating a limited downstream propagation of the oceanic constraint into the atmosphere. Noting that verification of atmospheric parameters against e-OBS T2M is over land, this indicates that such inland propagation of ocean observations may require further constraints in the system.

The surface flux differences (Figure 6) between the coupled (COS) and uncoupled (UOS) simulations highlight the physical mechanisms through which coupling alters SST and SSS, and they are broadly consistent with the performance differences documented in the sensitivity experiments. In the Aegean Sea, and to a lesser extent in the area south of Sicily, the signal is particularly clear: the net heat flux difference is negative (reduced downward heat flux), while the net freshwater flux (upward) difference is positive, indicating enhanced evaporation. Together, these patterns point to stronger latent heat loss in COS, which cools the surface and simultaneously increases salinity through evaporative concentration. This mechanism aligns well with the SSS RMSE improvement observed for COS in this region and explains the localized SST differences. In other regions, the coupling influence is more heterogeneous. Along the Alboran and Levantine basins, patches of increased downward heat flux coincide with small reductions in evaporation, suggesting coupling-induced adjustments in wind speed and humidity that alter the surface energy balance without producing a uniform SST or SSS response. The runoff anomalies shown in the bottom

352 panel are spatially limited and generally small but mostly positive, implying that river-input differences may not be the
 353 dominant driver of basin-wide changes but can affect local variations in salinity.

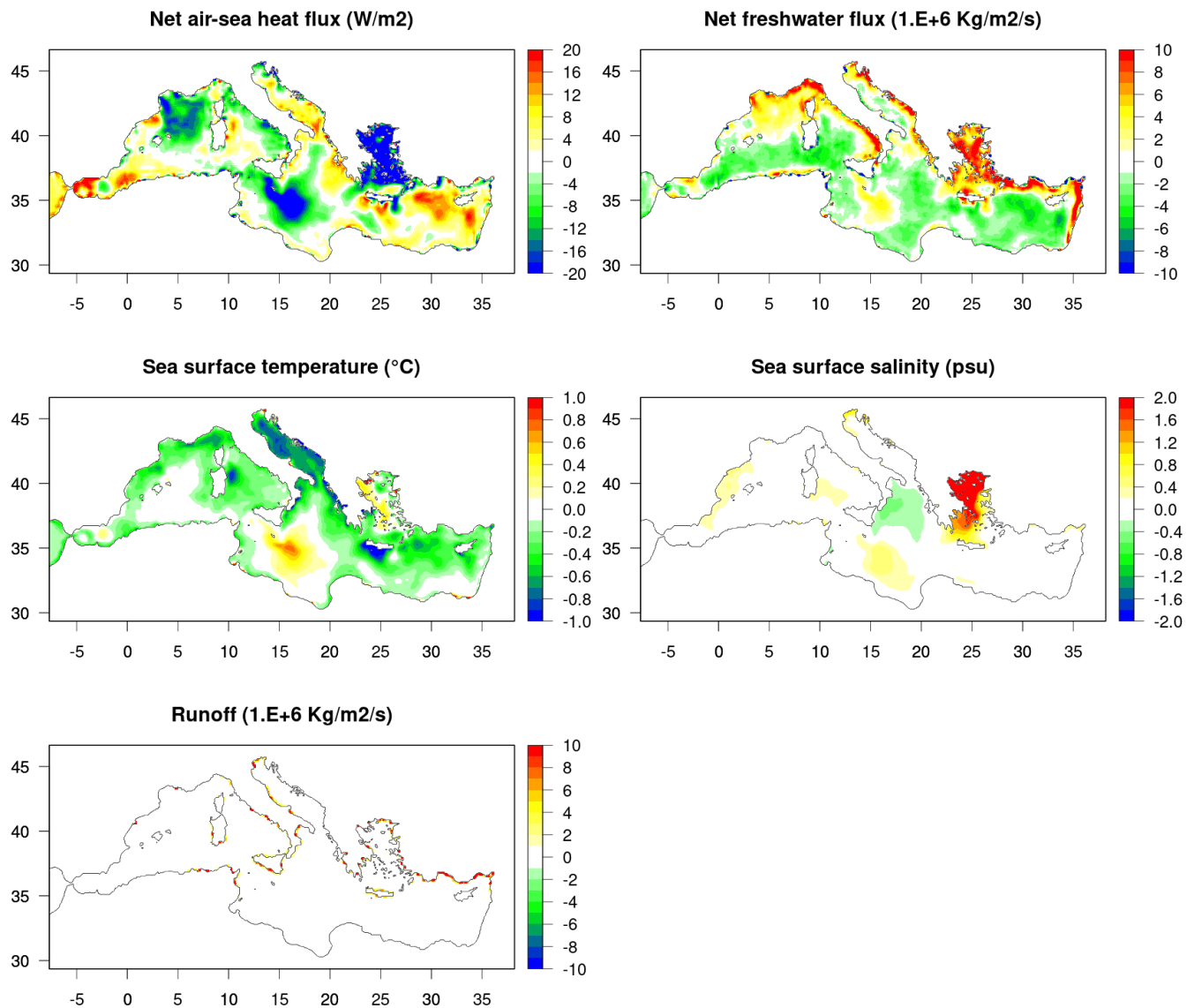
354



355

356 **Figure 5. Long-term averaged (2000-2020) near-surface atmospheric parameter differences between COA and COS**
 357 **sensitivity experiments (namely, showcasing the impact of the ocean data assimilation on atmospheric variables).**

358
359



360
361
362
363
364
365
366
367

Figure 6. Long-term averaged (2000-2020) air-sea flux and surface parameter differences between the COS and UOS sensitivity experiments (namely, showcasing the impact of the interactive coupling with WRF). Positive net heat flux is directed downward into the ocean and therefore represents ocean heat gain. Positive net freshwater flux is directed upward out of the ocean and therefore represents net freshwater loss through evaporation minus precipitation and runoff.

Eddy Kinetic Energy (from SSH)

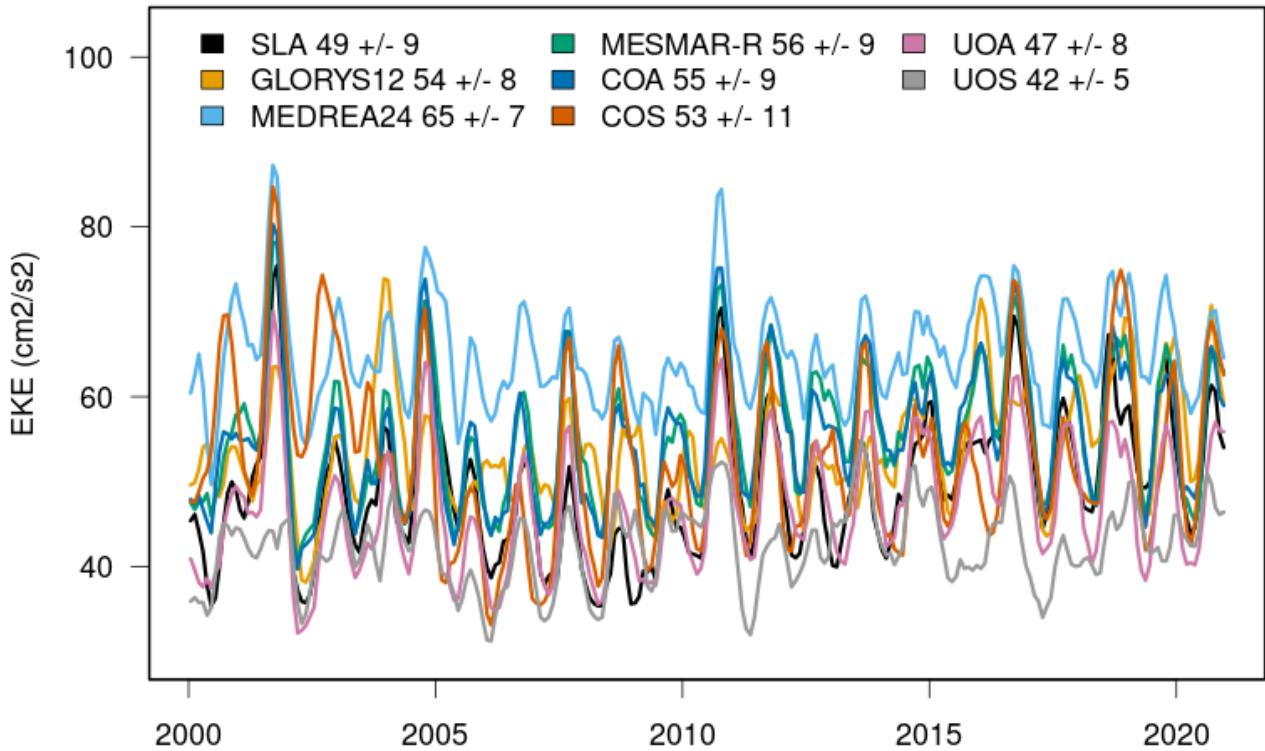


Figure 7. Eddy kinetic energy (EKE) timeseries, calculated from the MESMAR-R reanalysis, gridded altimetry data, the MEDREA24 and GLORYS12, and the sensitivity experiments, over the period 2000-2020. For all datasets, EKE is calculated from sea level, adopting the usual geostrophic relationship to calculate geostrophic currents.

Next, we review the impact of the different configurations on mesoscale eddy activity, ocean heat content, and deep-water formation. Eddy Kinetic Energy (EKE), computed from SSH under the geostrophic assumption for all datasets for the sake of consistency, reveals distinct dynamical signatures across the experiments (see Figure 7 for EKE timeseries in the Mediterranean Sea, and Figure 8 for EKE maps and differences). MESMAR-R exhibits consistently elevated EKE, both in the spatial maps and the basin-averaged time series, indicating that both data assimilation and coupling actively energize the mesoscale. Its values fall reasonably close to the observational benchmark derived from L4 altimetric sea-level anomalies (SLA), with the main discrepancy appearing near the Strait of Gibraltar, where steep topography, and intense barotropic flows may amplify representation errors in either observations or the model. In contrast, MEDREA24 shows EKE values that are substantially higher than all other systems, consistently with its much finer horizontal resolution ($1/24^\circ$), which can resolve smaller eddies and sharper gradients that naturally increase EKE. The spatial patterns suggest that MESMAR-R differs most strongly from SLA and GLORYS12 in parts of the southern Mediterranean.

The sensitivity experiments highlight the contributions of individual modelling components. COA and COS exhibit higher EKE than uncoupled systems, underscoring the tendency of coupling alone to energize the mesoscale; nevertheless, without

the constraint of ocean DA (COS), this may lead to less realistic variability. Conversely, UOA and UOS retain lower EKE levels, with UOS showing the weakest mesoscale activity overall, reflecting the damping imposed by prescribed atmospheric forcing and the absence of DA-induced eddy enhancement. The time series confirms this hierarchy: MESMAR-R $\approx$ COA > COS > UOA > UOS, mirroring the combined influence of coupling and data assimilation.

Together, these results highlight that MESMAR-R achieves a dynamically active mesoscale field that remains close to altimetric estimates, while coupling and data assimilation jointly act to enhance EKE relative to uncoupled or unassimilated configurations.

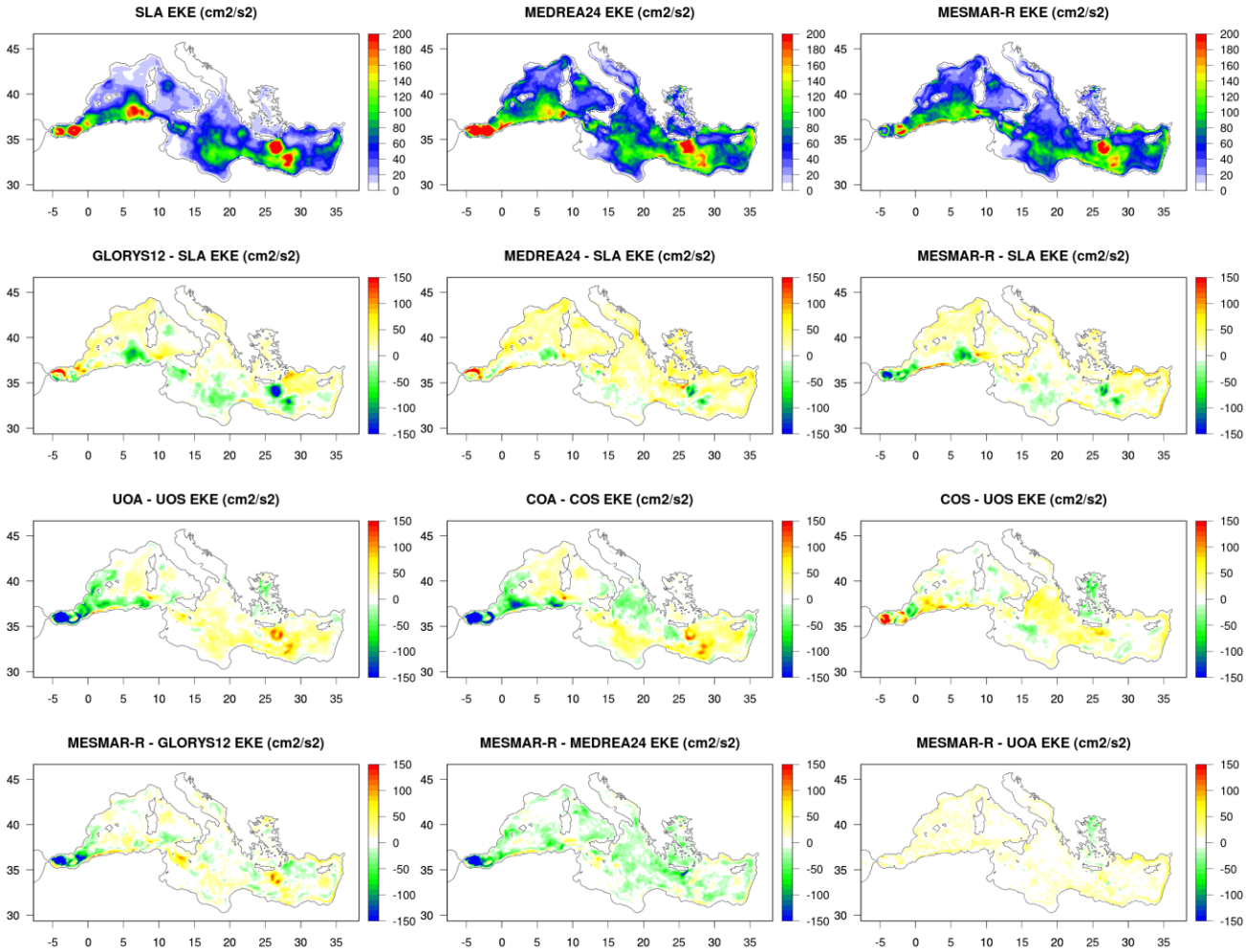


Figure 8. Eddy kinetic energy (EKE) maps, calculated from the MESMAR-R reanalysis, gridded altimetry data, the MEDREA24 and GLORYS12, and the sensitivity experiments, over the period 2000–2020. For all datasets, EKE is calculated from sea level, adopting the usual geostrophic relationship to calculate geostrophic currents. The top panels show the absolute EKE values, while the panels below show differences of EKE between pairs of experiments as indicated in each panel title.

Yearly Ocean Heat Content Anomaly 0-700m

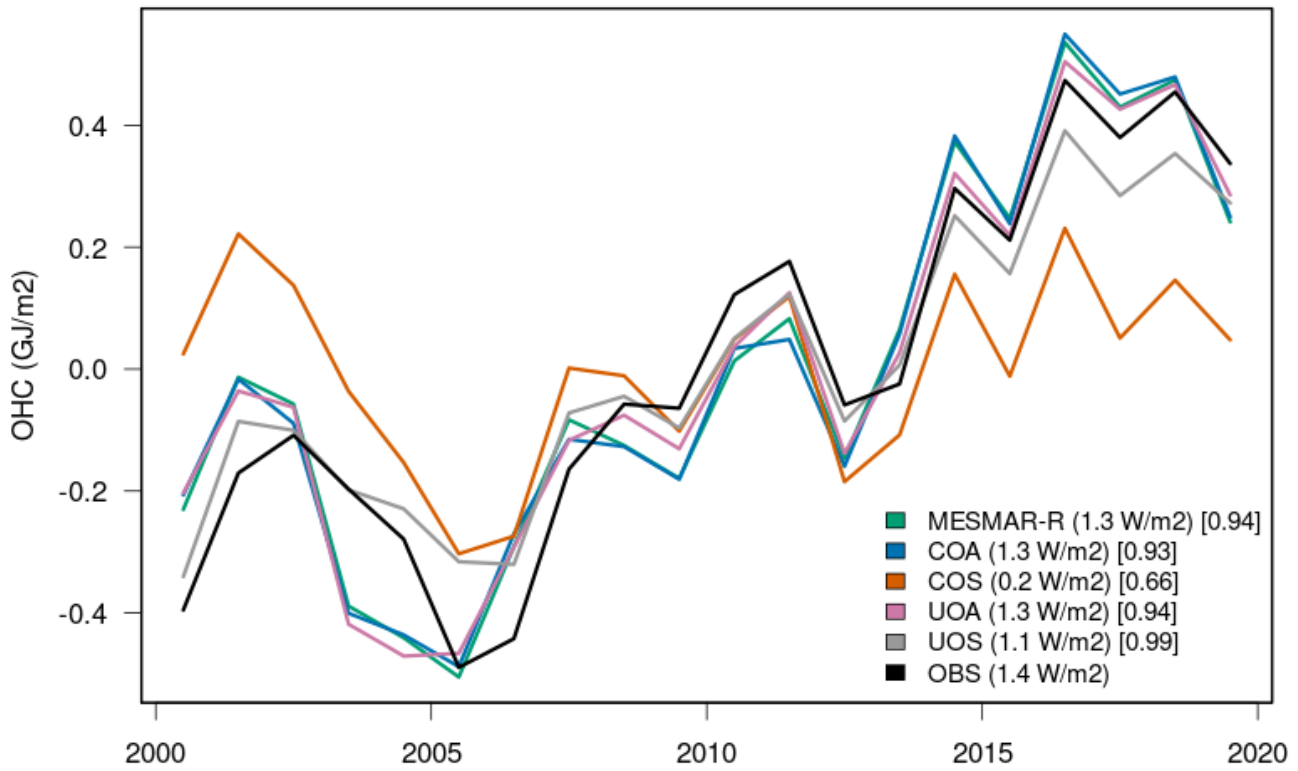


Figure 9. Yearly ocean heat content anomalies for the period 2000-2020 from MESMAR-R, the sensitivity experiments, and observation-based values. The legend also reports the trend in W m^{-2} , and the correlation with the observation-based time series.

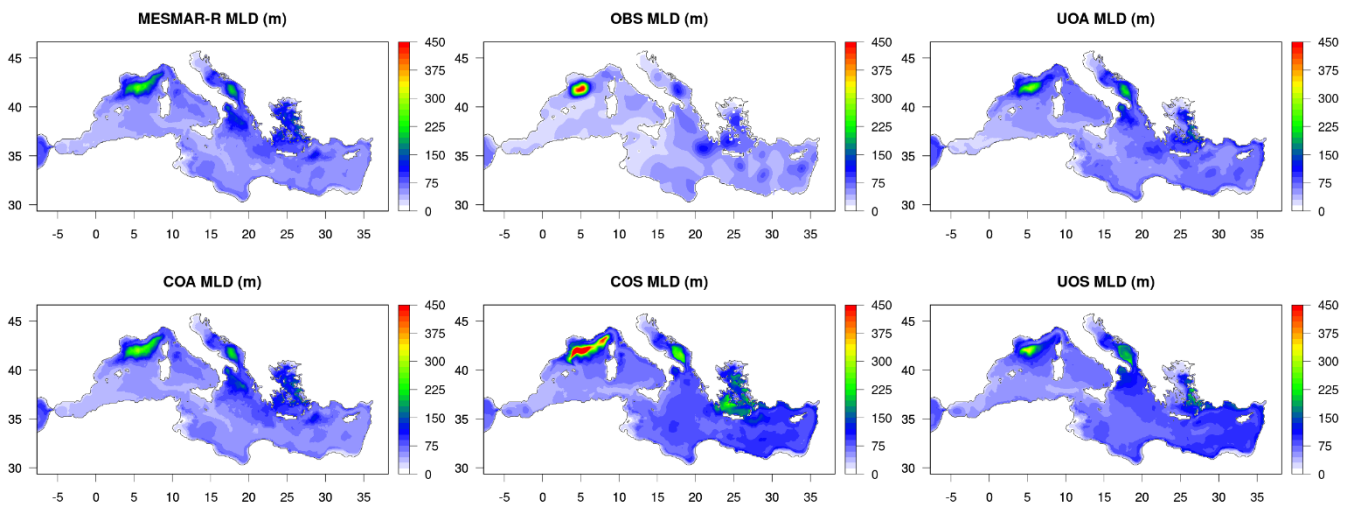


Figure 10. Long-term mean winter mixed layer depth (JFM) over 2000-2020 for the MESMAR-R reanalysis and the sensitivity experiments, together with values derived from in-situ observations.

The evolution of ocean heat content (OHC, Figure 9 as yearly anomalies) in the upper 700 m is reproduced remarkably well by all configurations except COS, with very high correlations (0.93–0.99) relative to observed OHC estimates. The MESMAR-R, COA, and UOA simulations each display a correlation of 0.93–0.94, closely matching the amplitude and timing of observed decadal variability, while UOS performs similarly well (0.99), albeit with slightly damped interannual variability and a smaller trend. In contrast, COS shows a markedly weaker correlation (0.66), reflecting the impact of the unconstrained coupled atmosphere–ocean system, which introduces spurious variability and phase errors, whereas COS ingests indirectly the atmospheric observations as it is forced by the ERA5 atmospheric reanalysis. The fact that MESMAR-R, COA, and UOA all produce nearly identical heat-content trends highlights that ocean data assimilation exerts the primary control on the basin-scale OHC signal, whereas coupling without DA (COS) deviates substantially. The long-term warming rate inferred from MESMAR-R, COA, and UOA ($\sim 1.3 \text{ W m}^{-2}$) also closely matches the observational estimate ($\sim 1.4 \text{ W m}^{-2}$), suggesting that the system’s heat uptake is robustly constrained by the assimilation of temperature and sea-level observations.

The mixed layer depth (MLD) fields, shown for winter (JFM), reveal a clearer impact of coupling on vertical structure (Figure 10). The comparison between COS and UOS highlights that coupling deepens the mixed layer in large areas of the western Mediterranean and Levantine Basin, consistent with the enhanced wind stress and modified heat-flux patterns seen earlier. However, the COS-UOS contrast does not isolate coupling alone. COS contains an interactive WRF–NEMO–HD system, whereas UOS is an ocean-only simulation forced by ERA5-derived atmospheric fields and uses a different surface-forcing pathway. Their difference therefore combines interactive feedbacks, atmospheric-forcing differences, and bulk-formulation differences. Consequently, different ventilation and ocean stratification cannot unambiguously be attributed to the coupling itself, but to the combination of the above factors.

The addition of ocean data assimilation in COA considerably reduces these differences, bringing the MLD patterns closer to those of the uncoupled configurations. Similarly, MESMAR-R exhibits a more moderate MLD distribution, showing that data assimilation effectively damps the excessive deepening (or shoaling) introduced by the free coupled dynamics. This behavior mirrors the EKE results, in which coupling tends to energize the system while data assimilation restores balance: together, these findings indicate that coupling intensifies upper-ocean overturning and stratification adjustments, but assimilation acts as a strong physical constraint that controls spurious variability.

Taken together, the OHC and MLD diagnostics show that ocean DA provides the dominant control on basin-scale heat-content evolution, ensuring close agreement with observations, while coupling primarily modifies vertical mixing and mixed-layer dynamics, with effects that become exaggerated in the absence of data constraints. The combined MESMAR-R system therefore achieves a balanced compromise: it preserves the physically consistent surface-flux feedback of a coupled model while maintaining a realistic representation of heat storage and vertical structure through assimilation.

The last comparison involves ocean data assimilation increments, which are in turn a measure of systematic model errors when averaged over time (e.g., Bhargava et al., 2018). The comparison of analysis increments among the three ocean–DA-enabled configurations (MESMAR-R, COA, and UOA) provides direct insight into how strongly the data assimilation system must correct the model state in each experiment. The comparison is summarized in Figures 11 (time-averaged increments) and 12 (Hovmöller diagram of spatial standard deviation over the Mediterranean Sea). Because increments represent the adjustments

required to reconcile the model with observations, their time-averaged values reveal systematic model biases, whereas their spatial standard deviation quantifies how strongly and how frequently the assimilation acts. By comparing the increments across experiments, we can diagnose how coupling and atmospheric forcing influence the ocean state: if increments are smaller, the model is closer to observations, meaning that the combination of physics, coupling, and forcing used in that configuration is more self-consistent.

The time-averaged increments at the surface highlight a robust pattern present in all experiments: positive temperature increments in the western Mediterranean and negative increments in the eastern basin. This structure indicates a cold bias in the west and a warm bias in the east before analysis, consistent with the SST RMSE patterns discussed earlier. Notably, UOA shows smaller surface temperature increments than MESMAR-R or COA, especially in the central and eastern Mediterranean, confirming that the prescribed ERA5 atmospheric forcing provides a more accurate near-surface heat flux constraint than the interactive coupling. UOA increments are, however, consistent in sign with those from MESMAR-R, showcasing persistent biases likely originating in the ocean model component. This result aligns with the surface flux analysis, where coupling produced regional discrepancies in latent and sensible heat fluxes, particularly in the Levantine and Aegean Seas.

For salinity at 0 m, the increments again show broad similarities across experiments, but UOA displays a distinct freshening increment in the Aegean Sea that is not present in the coupled configurations. This feature likely reflects the fresh bias introduced by ERA5 freshwater forcing, which lacks the dynamically consistent evaporation–precipitation adjustments of the coupled system. It also echoes earlier findings that coupling improves SSS in the Aegean by enhancing realistic evaporative fluxes and, to a lesser extent, river discharge. These consistencies strengthen the interpretation that salinity is especially sensitive to air–sea freshwater coupling, whereas temperature is more sensitive to heat-flux accuracy.

At 30 m and 100 m (not shown), the differences among experiments are substantially reduced. The increments remain spatially coherent but become weaker and less distinct across configurations, indicating that the influence of atmospheric forcing and coupling diminishes rapidly below the surface mixed layer. Nevertheless, a persistent feature across all three experiments is the salinity correction in the southwestern Mediterranean, which likely reflects biases in the water exchanged through the Strait of Gibraltar, consistently with the EKE and SSS anomalies noted near the entrance of the basin. Beyond this, the experiments converge, confirming that the main differences between configurations are confined to the upper ~50 m, where coupling and surface fluxes exert the strongest control.

472

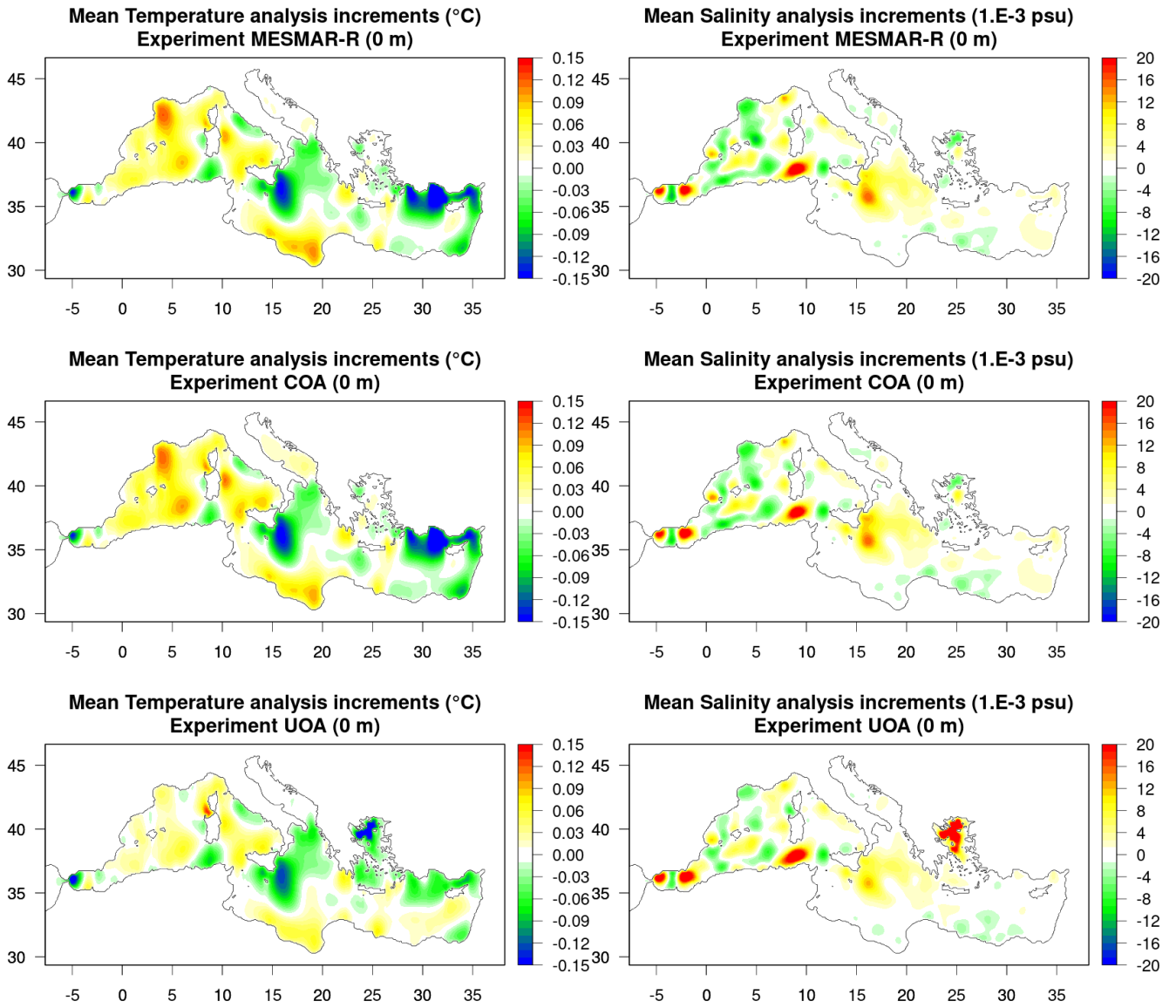


Figure 11. Time-averaged analysis increments at the sea surface for MESMAR-R and the other two experiments with ocean data assimilation enabled. The period is 2000-2020.

The Hovmöller diagrams (Figure 12) further clarify the behavior of the assimilation system. In MESMAR-R, temperature increments peak in summer, when strong stratification makes the model more sensitive to surface forcing errors and limits vertical redistribution of heat. Over time, a modest decrease in increment magnitude is evident, suggesting that the reanalysis improves internally throughout the period, likely via better-constrained large-scale circulation and gradual spin-up of the coupled and data assimilation systems.

For salinity, MESMAR-R exhibits larger increment variability, particularly in the upper 100 m. Part of this arises from the greater density of salinity observations (ARGO, gliders, ship-based), but it also reflects the intrinsic sensitivity of Mediterranean salinity to evaporation, runoff, and mixing, all of which contribute to persistent local biases.

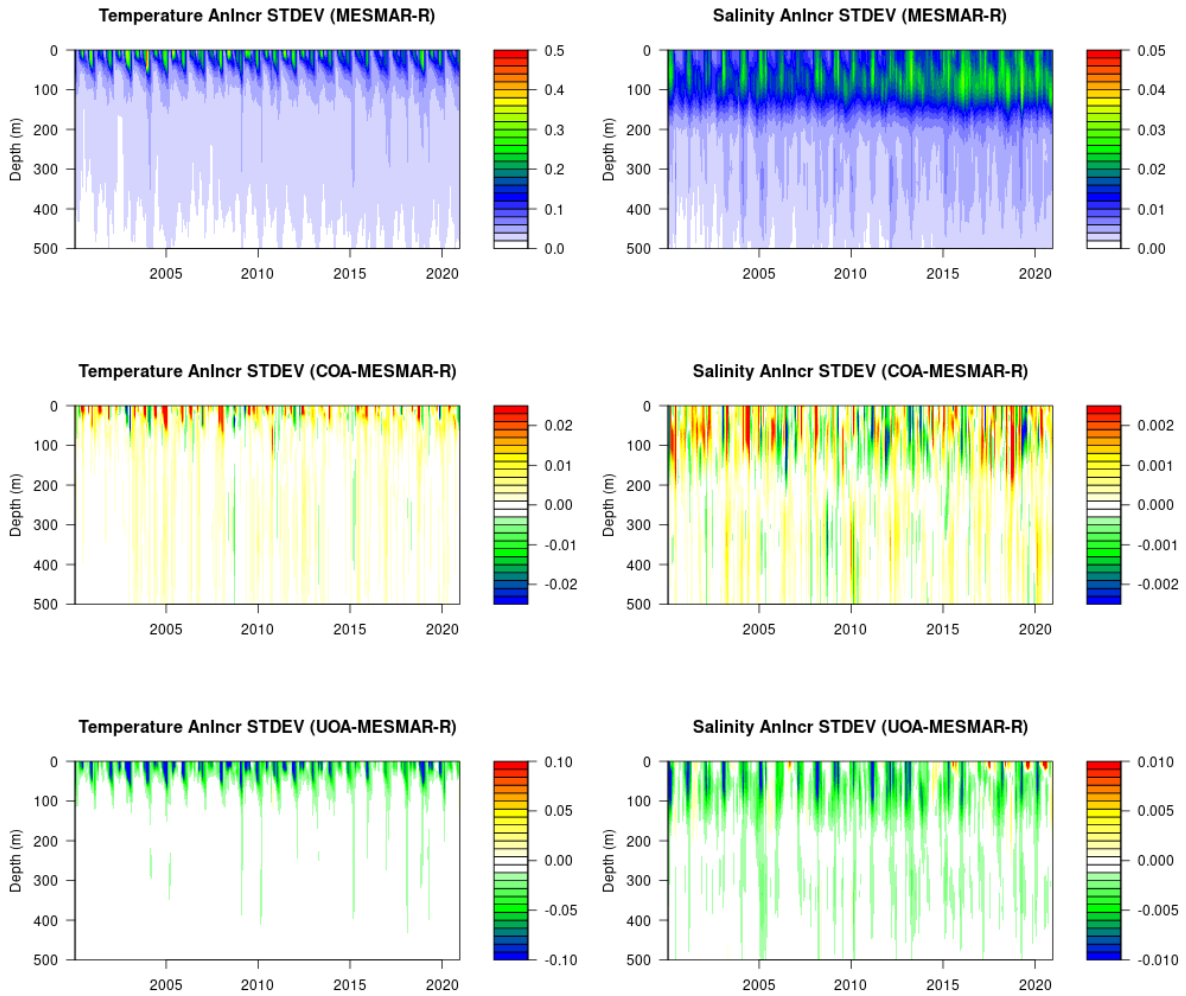


Figure 12. Hovmöller diagram (time-depth) of the analysis increments spatial standard deviation for MESMAR-R and the other two experiments with ocean data assimilation enabled. The period is 2000-2020.

Differences relative to MESMAR-R provide further insight. The COA minus MESMAR-R differences are slightly positive almost everywhere, confirming that the atmospheric DA included in MESMAR-R marginally reduces the need for ocean corrections. This reinforces earlier conclusions that atmospheric assimilation improves the air-sea flux balance, particularly for heat and freshwater, thereby generally reducing ocean misfits.

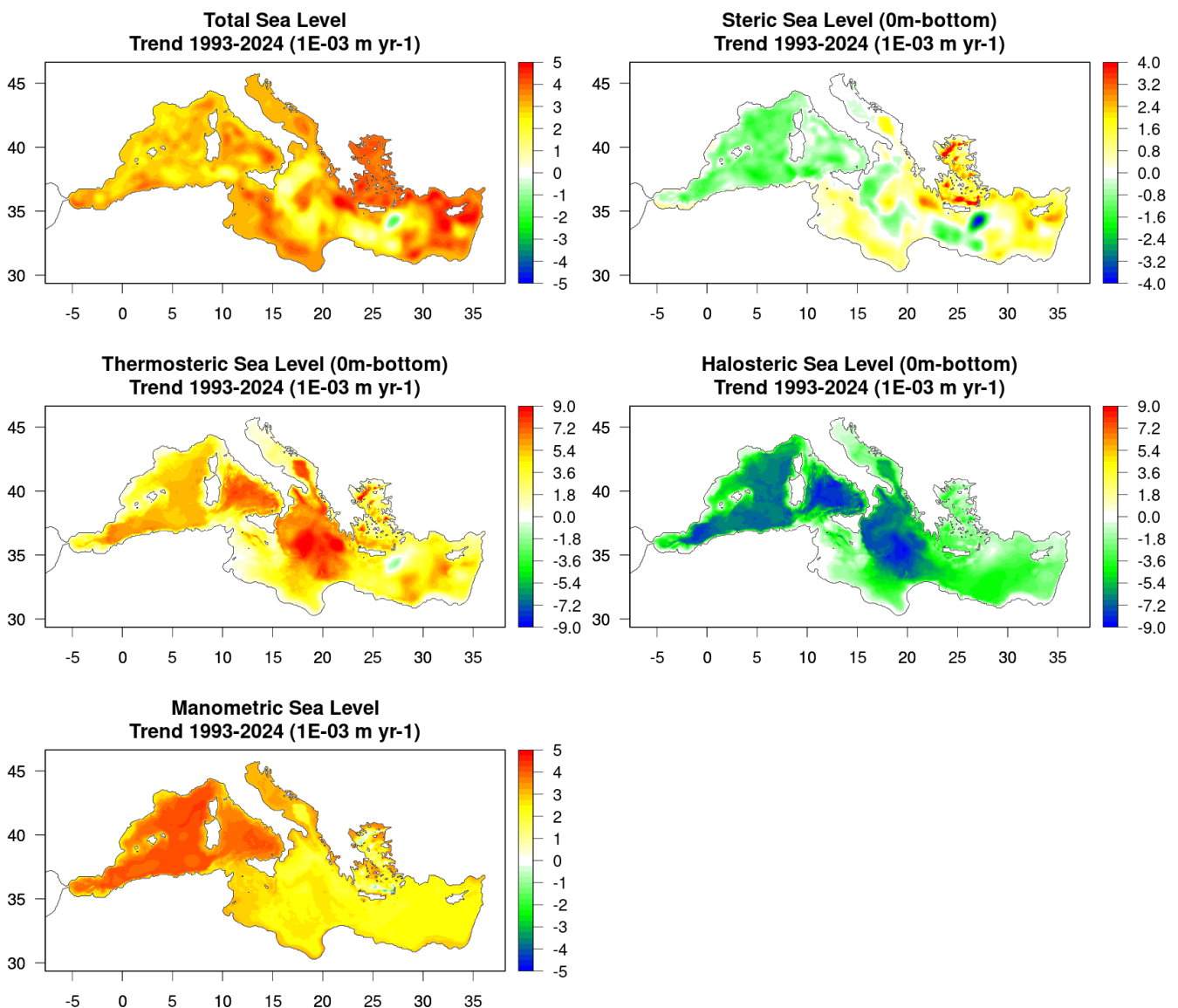
Conversely, UOA minus MESMAR-R differences show generally negative values in the upper 50 m, consistent with the higher fidelity of ERA5 forcing for near-surface fields. These differences peak in winter, when atmospheric forcing dominates mixed-layer dynamics, and coupling errors can accumulate most easily. Below ~50 m, the differences vanish, reinforcing that ERA5 and coupling primarily diverge near the surface.

500 **5. Climate indices**

501 **5.1 Oceanic long-term changes from the MESMAR-R Reanalysis**

502 The MESMAR-R coupled reanalysis provides a dynamically consistent multidecadal reconstruction of the Mediterranean
503 circulation and its climate evolution over 1993–2024. Beyond the validation metrics discussed previously, the long-term trends
504 in sea level, temperature, salinity, and heat/salt content offer a representation of the climate response of the Mediterranean
505 basin. The Mediterranean is a semi-enclosed concentration basin whose thermohaline characteristics evolve rapidly under
506 climate forcing, and where salinity-driven density changes can dominate over thermal effects. The climate indices derived
507 from MESMAR-R are therefore particularly informative in assessing regional climate changes.

508



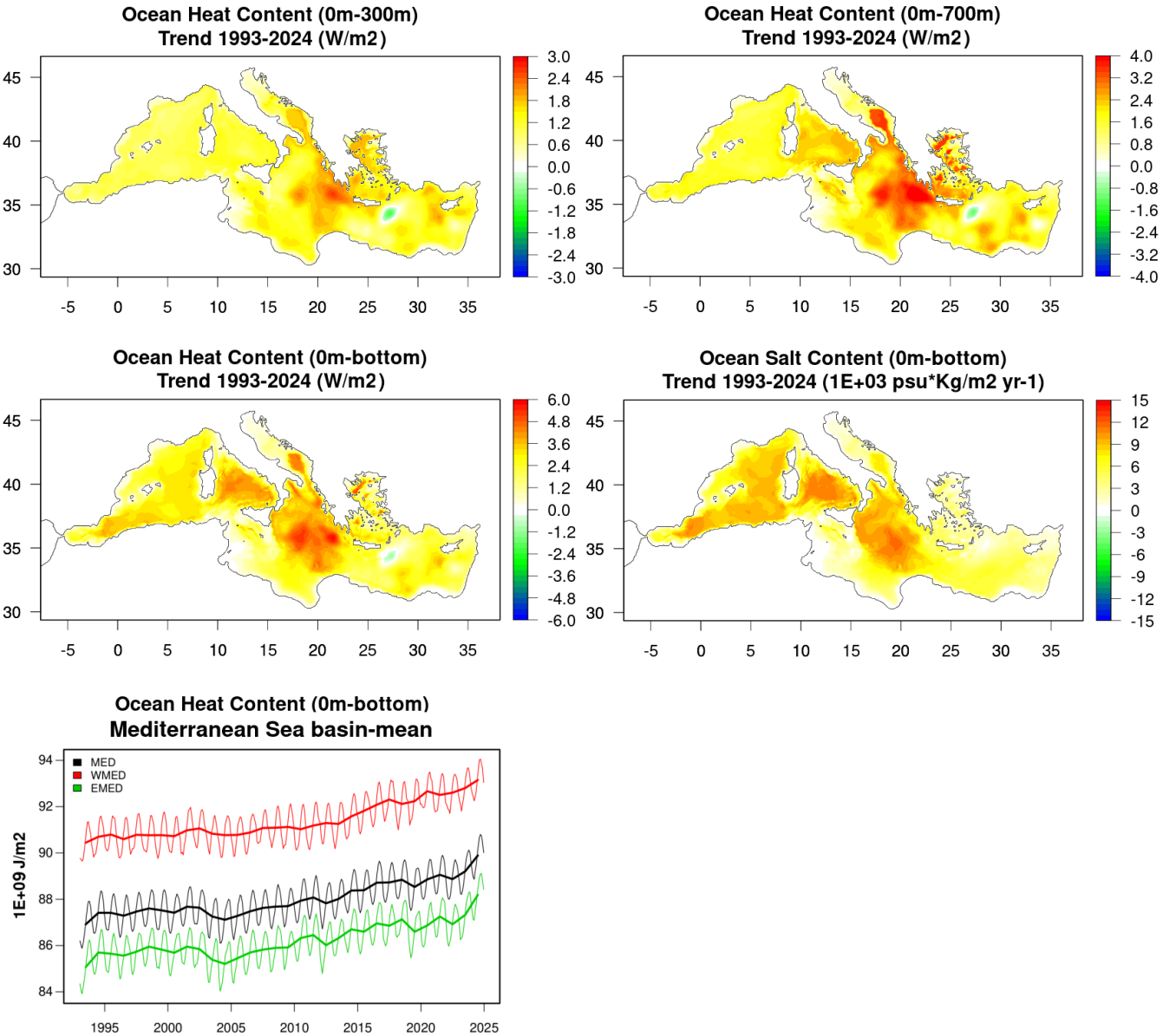
509

510 **Figure 13. Maps of linear trends (1993-2024) for sea level and its components, from MESMAR-R.**

511

512 The halosteric component of sea level exhibits a strong, basin-wide negative trend in MESMAR-R, with a basin-mean decrease
 513 of approximately -3.87 mm yr^{-1} over 1993–2024 and even stronger halosteric decline in the western basin (-5.02 mm yr^{-1})
 514 relative to the eastern basin (-3.27 mm yr^{-1}) (see Figure 13 and Table 4). These values are fully consistent with the intensifying
 515 salinification of the Mediterranean reported in observational and modelling studies (Liu et al., 2025; Skliris et al., 2018;
 516 Aydogdu et al., 2023). The trend maps show particularly strong halosteric lowering in the Alboran Sea and western
 517 Mediterranean gyre, a region identified in previous work as a hotspot of salinity-driven density increase, likely associated with
 518 enhanced evaporation and changes in Atlantic inflow properties (Skliris et al., 2018). A pronounced east–west contrast persists
 519 across depth, indicating that salinity anomalies penetrate well into intermediate waters, consistent with evidence of progressive
 520 warming and salinification of Levantine Intermediate Water (Pinardi and Masetti, 2000; Margirier et al., 2020). This halosteric
 521 signal strongly compensates for the thermosteric rise, leading to the well-known thermo–halo compensation characteristic of
 522 the Mediterranean (Liu et al., 2025). Local features such as weaker halosteric trends in the Adriatic, influenced by river
 523 discharge and dense-water formation, further corroborate the realism of MESMAR-R.

524



525

526 **Figure 14. Maps of linear trends (1993-2024) for ocean heat and salt content, and OHC timeseries from MESMAR-R**

527 The decomposition of sea-level rise into steric and manometric components (by difference) shows that basin-mean

528 Mediterranean sea-level rise in MESMAR-R is dominated by mass changes rather than steric contributions, i.e., the basin's

529 response to the Atlantic sea-level rise via exchanges through the Strait of Gibraltar. The basin-average total sea-level trend of

530 approximately 3.06 mm yr^{-1} is in excellent agreement with the Copernicus Marine Service Mediterranean Sea Level Indicator

531 and with external literature estimates ranging from 2.5 to 4.0 mm yr^{-1} (Calafat et al., 2012; Marcos & Tsimplis, 2008). By

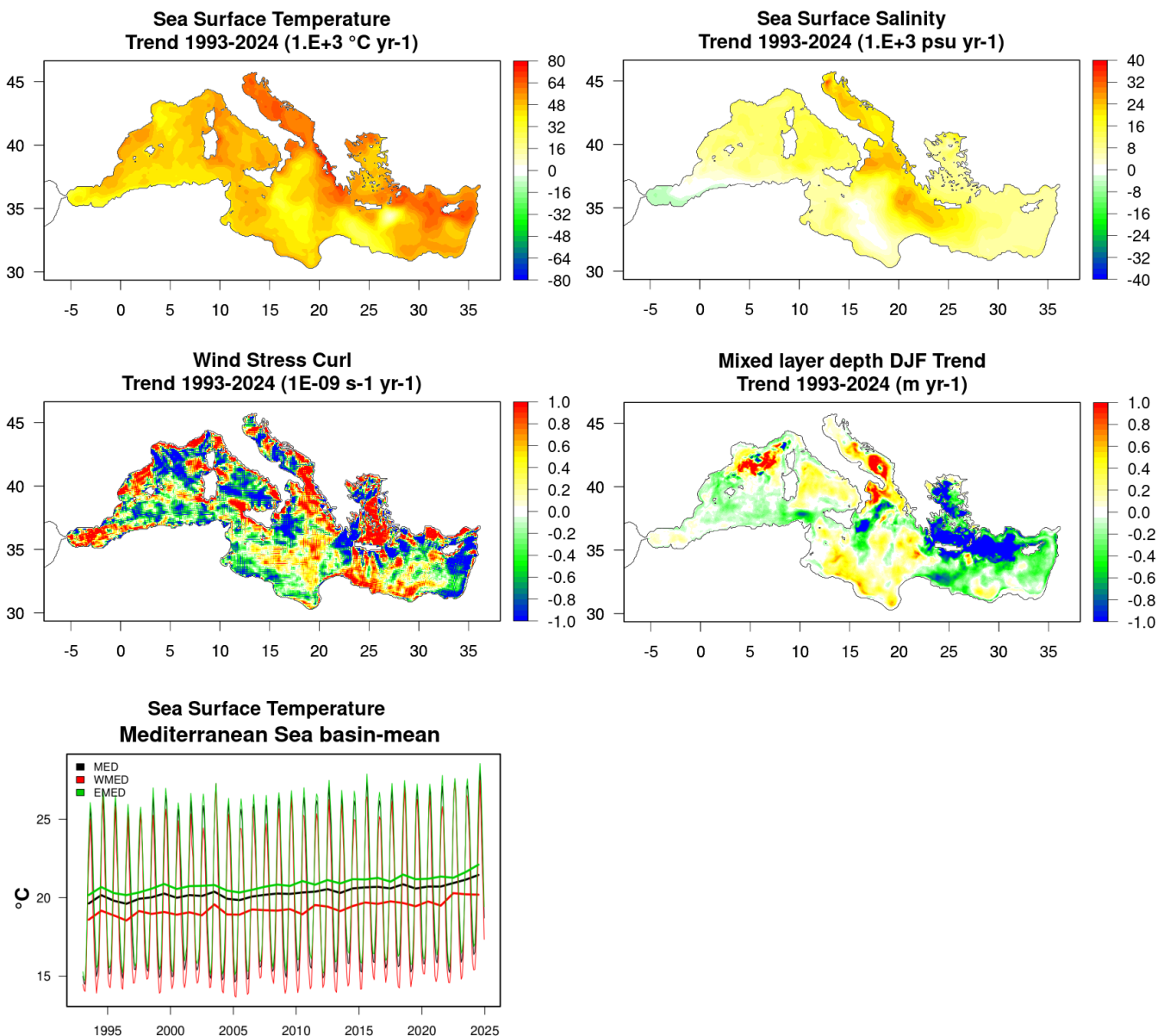
532 contrast, the manometric sea level rise is equal to 2.88 mm yr^{-1} , accounting for 94% of the sea level rise over 1993-2024 and

533 in agreement with independent estimates based on sea level budget analysis (Storto et al., 2024). Wind-stress-curl trends in

534 MESMAR-R do not appear to exert first-order control on the basin-mean sea-level rise; however, positive curl anomalies over

535 parts of the eastern Mediterranean (Figure 15) may contribute locally to the intermediate-layer warming signal.

536



537

538 **Figure 15. Maps of linear trends (1993-2024) for selected variables, and time series of SST, from MESMAR-R.**

539 The basin-average ocean heat content (OHC in the top 700m, Figure 14 and Table 4) trend from MESMAR-R is 1.10 (1.34)
 540 W m^{-2} for the period 1993-2024 (1999-2024), in line with recent observational and reanalysis-based assessments (Pisano et
 541 al., 2020; von Schuckmann et al., 2016). MESMAR-R also captures the characteristic east–west contrast in heat-content
 542 evolution: the eastern basin exhibits the stronger long-term warming trend, whereas the western-basin time series shows the
 543 larger seasonal-to-interannual amplitude. Trend maps show spatially coherent warming extending from the surface to
 544 intermediate depths ($\sim 300\text{--}700$ m), associated with the warming and deep penetration of LIW. The implied vertical structure
 545 is consistent with the findings of Margirier et al. (2020), who documented abrupt warming and salinification of intermediate
 546 waters beginning in the early 2000s. The warming pattern in MESMAR-R aligns with the basin’s sensitivity to changes in
 547 surface heat fluxes and stratification, and the weak wind-stress-curl trends suggest that mechanical forcing plays a secondary
 548 role compared with thermohaline adjustments.

549

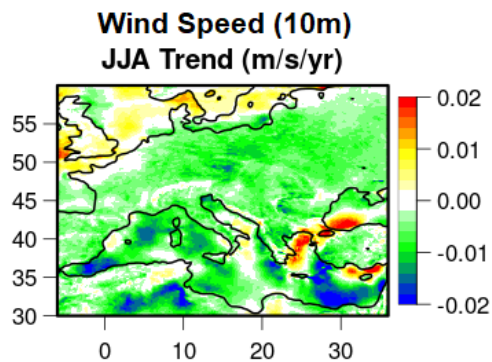
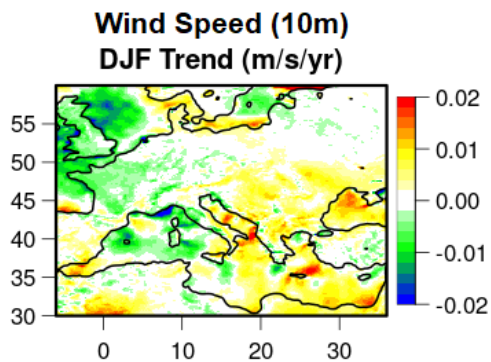
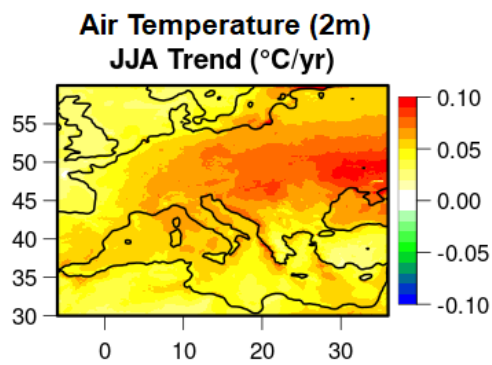
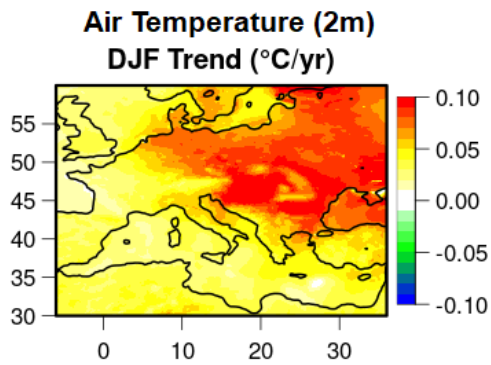
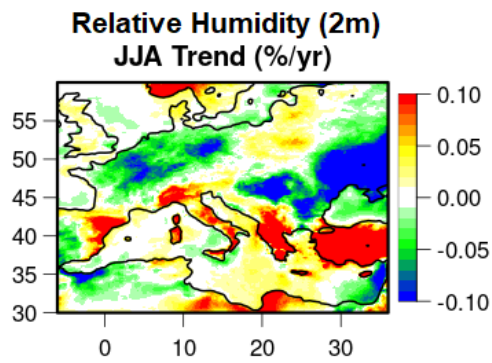
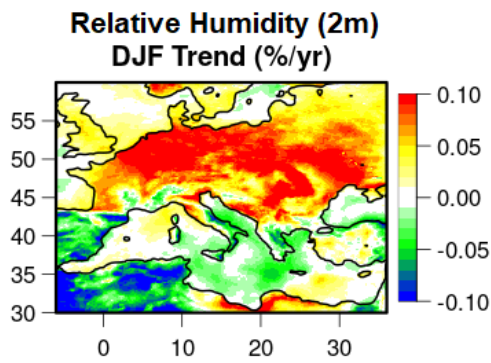
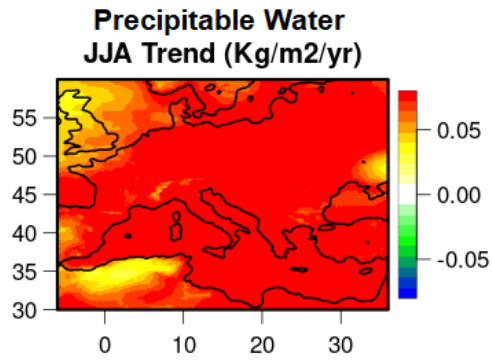
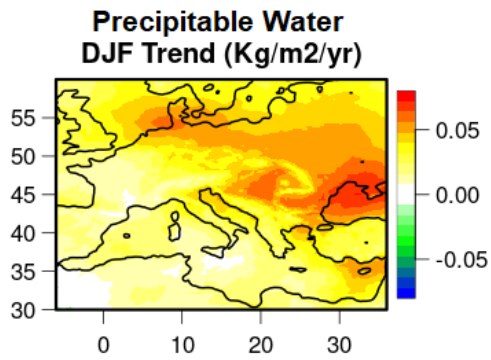
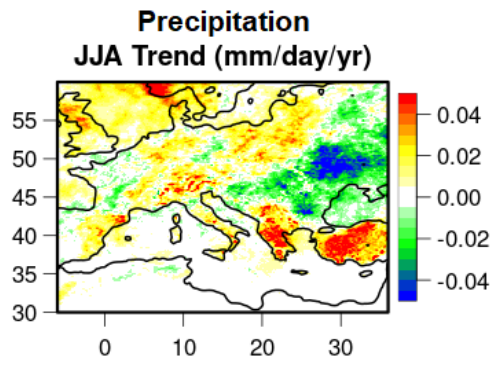
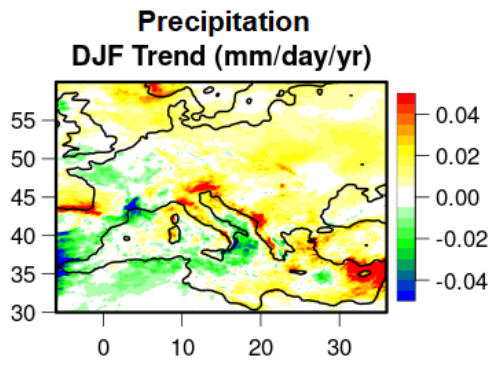
Variable	Basin-mean trend	Western Mediterranean	Eastern Mediterranean
Total Sea Level trend (mm yr ⁻¹)	3.61 (3.06)	3.73 (2.94)	3.54 (3.13)
Thermosteric Sea Level trend (mm yr ⁻¹)	5.04 (4.03)	5.29 (4.29)	4.91 (3.89)
Halosteric Sea Level trend (mm yr ⁻¹)	-4.21 (-3.87)	-5.23 (-5.02)	-3.68 (-3.27)
(Total) Steric Sea Level trend (mm yr ⁻¹)	0.82 (0.18)	0.11 (-0.67)	1.19 (0.62)
Manometric Sea Level trend (mm yr ⁻¹)	2.79 (2.88)	3.62 (3.61)	2.35 (2.51)
Ocean Heat Content trend (0–700 m) (W m^{-2})	1.34 (1.10)	1.28 (0.94)	1.37 (1.18)
Ocean Heat Content trend (0–bottom) (W m^{-2})	2.05 (1.68)	2.05 (1.55)	2.04 (1.75)
Sea Surface Salinity trend (psu yr ⁻¹)	0.008 (0.011)	0.007 (0.009)	0.009 (0.012)
Sea Surface Temperature trend ($^{\circ}\text{C yr}^{-1}$)	0.055 (0.048)	0.056 (0.047)	0.055 (0.049)
Ocean Salt Content trend (psu·kg m ⁻² yr ⁻¹)	5.68×10^3 (5.22×10^3)	7.02×10^3 (6.75×10^3)	4.98×10^3 (4.42×10^3)
DJF Mixed Layer Depth (m / yr ⁻¹)	-0.305 (0.019)	-0.173 (-0.234)	-0.218 (-0.147)

550 **Table 4. Mediterranean Sea averaged long-term trends over the periods 1999-2024 (1993-2024 in parentheses),**
 551 **separated for the entire basin, the western and eastern basins as defined in section 2.5.**

552

5.2 Atmospheric long-term changes from the MESMAR-R Reanalysis

The trends in near-surface atmospheric variables over 1993–2024 (Figure 16 and Table 5) depict a Mediterranean–southern Europe climate signal that is fully consistent with the “hot-spot” character of the region described in previous work (Giorgi & Lionello, 2008; MedECC, 2020). In winter (DJF), 2-m air temperature shows a robust warming over almost the entire domain, with maxima over central and eastern Europe and weaker positive trends over the Mediterranean Sea itself; this corresponds well to the MED-average DJF warming of $+0.040\text{ }^{\circ}\text{C yr}^{-1}$. Precipitation trends are more heterogeneous, with modest drying over much of the western and central Mediterranean and parts of North Africa, and localized increases along the northern mountain ranges (e.g., Alps-Dinaric arc), in line with earlier observational and modelling studies that report winter precipitation decreases over the Mediterranean land areas but more mixed signals at higher latitudes and elevations (Hoerling et al., 2012; Hertig, 2017; Beranová et al., 2025). Localized coastal trends are also in agreement with global reanalyses such as ERA5 (not shown). Column-integrated precipitable water increases almost everywhere in DJF ($+0.023\text{ }\%\text{ yr}^{-1}$ increase on average), while near-surface relative humidity shows a dipole structure, with slight moistening over northern Europe and drying over the southern flanks of the basin. These changes point to a more moisture-laden but thermally driven winter atmosphere in which enhanced water vapour does not systematically translate into increased mean precipitation (with an average DJF precipitation trend of $+2.2\times 10^{-3}\text{ mm day}^{-1}\text{ yr}^{-1}$), consistent with recent moisture-budget analyses based on ERA5 that highlight the dominant role of transient processes in transporting oceanic moisture onto surrounding land (Tootoonchi et al., 2025).



571 **Figure 16. Winter (DJF) and summer (JJA) trends of selected atmospheric variables from MESMAR-R, computed**
572 **over the period 1993-2024.**

573

574 Summer (JJA) trends display an even clearer expression of Mediterranean aridification. Near-surface temperature increases
575 are larger than in winter in southern Europe and the Mediterranean region and exhibit a strong meridional gradient (average
576 JJA warming of $+0.047\text{ }^{\circ}\text{C yr}^{-1}$, in line with projections and recent assessments of compound hot–dry summers over the region,
577 i.e., Giorgi & Lionello, 2008; Lhotka et al., 2023; MedECC, 2020). Mean precipitation shows heterogeneous behavior, with
578 neutral tendency in the southern Mediterranean and mixed above; at the same time, precipitable water exhibits strong positive
579 trends over essentially the whole domain (average JJA trend of $+0.100\text{ }\%\text{ yr}^{-1}$) while near-surface relative humidity decreases
580 over large parts of Europe while the southern Mediterranean shows a small increase. This combination of higher column water
581 vapour, lower relative humidity, and reduced mean precipitation is typical of a warming, drying regime where the atmosphere
582 can hold more moisture (Hoerling et al., 2012; González-Hidalgo et al., 2025; Figueiredo et al., 2025) but does not necessarily
583 precipitate more often, instead favouring longer dry spells punctuated by intense events (Zittis et al., 2021; Senatore et al.,
584 2025).

585

Variable	DJF	MAM	JJA	SON
10m Wind Speed ($\text{cm s}^{-1}\text{ yr}^{-1}$)	0.20	0.30	-0.45	-0.37
2m Temperature ($^{\circ}\text{C yr}^{-1}$)	0.040	0.037	0.047	0.053
2m Relative Humidity ($\%\text{ yr}^{-1}$)	-0.017	0.008	0.027	-0.006
Precipitable Water ($\%\text{ yr}^{-1}$)	0.023	0.043	0.100	0.070
Total Precipitation (mm $\text{day}^{-1}\text{ yr}^{-1}$)	2.2 E-3	4.3 E-3	4.2 E-3	-0.36 E-3

586 **Table 5. Mediterranean region-averaged trends for the four seasons and selected variables, from MESMAR-R over the**
587 **period 1993-2024. The region is defined as 30–45°N, 10°W–40°E (land and sea combined), close to the domain used by**
588 **Giorgi and Lionello (2008) but without the Alpine region.**

589 The trends in 10-m wind speed are comparatively modest, but they exhibit coherent regional structures that can lead to changes
590 in winter oceanic mixed-layer depth (MLD, Figure 15). In DJF, MESMAR-R shows localized increases in surface wind speed
591 over the Gulf of Lion and parts of the northern Ionian/Aegean, consistent with the average DJF trend of $+0.20\text{ cm s}^{-1}\text{ yr}^{-1}$
592 together with small negative temperature trends relative to surrounding regions. These areas coincide with weakly positive
593 DJF MLD trends for the full 1993-2024 period, indicating a deepening of the winter mixed layer that is consistent with
594 enhanced mechanical mixing and stronger buoyancy loss. Conversely, in the Levantine and southeastern Mediterranean, winter
595 MLD trends are negative (shoaling) and co-occur with marked warming and slight wind weakening, suggesting reduced
596 buoyancy loss and a decline in deep or intermediate-water formation. The large-scale wind-stress-curl trends derived earlier

597 show only modest changes, but positive curl anomalies in parts of the eastern basin support a dynamical picture in which
598 circulation adjustments and stratification changes modulate the mixed-layer response on top of local surface-flux trends.
599

600 6. Discussion and Conclusions

601 Achieving a dynamically consistent reconstruction of the recent atmosphere-ocean-hydrology evolution over the
602 Mediterranean region, one of the world's most prominent climate-change hotspots, is essential for disentangling the processes
603 that govern its regional response to global warming. Here, we have presented one of the first regional coupled reanalysis
604 produced to date, to our knowledge. The regional reanalysis system, called MESMAR-R, features state-of-the-art ocean,
605 weather, and river discharge numerical models, complemented with data assimilation systems in both the atmosphere and the
606 ocean, although these are weakly coupled in the sense that observations in one medium do not directly affect analysis
607 corrections in the other medium. Atmospheric data assimilation is somewhat simple (based on large-scale three-dimensional
608 nudging) and not observation-space-based; yet the system provides an observation-aware atmosphere-ocean-consistent
609 reconstruction of the latest 32 years (1993-2024).

610 The ocean component of MESMAR-R embeds an advanced variational data assimilation system and demonstrates clear
611 strengths at the surface and in the upper ocean, particularly for sea level, sea surface salinity, and upper-layer temperature.
612 These results confirm the quality of the new reanalysis in the layer most relevant to air-sea interaction and marine variability.
613 Remaining differences in mid-depth temperature suggest areas for further improvement, possibly linked to intermediate-water
614 representation. The spatial patterns of skill scores reinforce the basin-wide good performances of MESMAR-R at the sea
615 surface, with the clearest benefits relative to external products, especially for salinity and sea level. Localized degradations
616 (e.g., in the northern Adriatic and Aegean Seas) suggest areas where representation of strong riverine inputs and complex
617 circulation could be further optimized.

618 Overall, the atmospheric evaluation confirms that MESMAR-R performs reasonably well given its simplified atmospheric data
619 assimilation approach based on large-scale nudging. While its near-surface atmospheric variables generally exhibit higher
620 RMSE than those from ERA5 and CERRA, the differences remain moderate, and the fields are of comparable quality at the
621 basin's scale. Localized improvements indicate that the coupled system can maintain internally consistent air-sea interactions
622 despite the limited atmospheric assimilation. These results demonstrate that, although MESMAR-R does not aim to compete
623 with dedicated atmospheric reanalyses, its atmospheric component provides a physically coherent and sufficiently accurate
624 representation of the Mediterranean climate to support further process-oriented investigations.

625 Several sensitivity tests were conducted to disentangle the impact of coupling and data assimilation in MESMAR-R. These
626 demonstrate that data assimilation in both components is essential for optimal coupled performance. MESMAR-R achieves
627 the best compromise between realism and stability, with coupling improving certain surface metrics (notably salinity) and
628 atmospheric relaxation further enhancing the representation of SST and SLA. The results highlight that while coupling alone
629 is not sufficient to match the reanalysis skill, it becomes most effective when supported by data assimilation in both the ocean
630 and atmosphere. The sensitivity experiments indicate that coupling most strongly affects regions where the air-sea humidity
631 and wind-speed fields are highly sensitive to ocean-atmosphere feedback, with evaporation acting as the primary mediator of
632 salinity changes. The results are consistent with the RMSE analysis, where coupling intensifies latent heat losses (e.g., Aegean
633 and parts of the eastern Mediterranean), and coupled configurations outperform uncoupled ones. This was also strengthened
634 by the assessment of the analysis increments whose average signal primarily responds to systematic error corrections.

635 The analysis of oceanic trends demonstrates that MESMAR-R has provided a coherent and physically consistent reconstruction
636 of Mediterranean climate evolution over the past three decades. The reanalysis captures the basin-wide salinification, the
637 thermo-halo compensation of steric sea-level change, the mass-driven nature of Mediterranean sea-level rise, and a heat-
638 content increase fully aligned with observational estimates. Long-term atmospheric trends from MESMAR-R (1993–2024)

639 reveal a climate signal fully consistent with the well-documented Mediterranean hot-spot behavior. Winter and summer both
640 exhibit pronounced near-surface warming, widespread drying, and strong moistening of the atmospheric column, confirming
641 the thermodynamic intensification of the regional hydrological cycle.

642 An important caveat is that the observational information is not homogeneous over 1993–2024. Changes in the availability,
643 type, and spatial distribution of ocean observations can affect both reanalysis accuracy and the temporal comparability of
644 observation-based verification statistics, particularly during earlier or sparsely sampled periods (Storto et al., 2019a). A similar
645 limitation applies to the atmospheric component: because MESMAR-R is spectrally nudged towards ERA5, its large-scale
646 atmospheric constraint partly inherits the time-varying information content of the observing system assimilated in ERA5
647 (Hersbach et al., 2020). Interdecadal variability and trends should therefore be interpreted while accounting for this temporally
648 evolving observational data.

649 MESMAR-R, while presenting room for further improvement in the future (observation space-based atmospheric data
650 assimilation and air-sea balanced corrections), represents a unique reconstruction of the past climate in the Mediterranean
651 region with more dynamically consistent atmospheric and oceanic fields than in uncoupled systems. Its ocean-focused design
652 mirrors state-of-the-art performance for the ocean component, despite its moderate spatial resolution. The primary advantage
653 of the MESMAR-R reanalysis is, however, the balanced response across the atmosphere and the ocean in the Mediterranean
654 Sea, which makes it a valuable product to assess compound events in the region, and any other strongly coupled air-sea process,
655 including, e.g., medicanes and heavy precipitation events; the present paper, however, provides a first climatological and mean-
656 state evaluation and does not constitute a dedicated assessment of strongly coupled events, which is left for future analyses.
657 The six-hourly and daily MESMAR-R output is suitable for these applications, and future work will quantify event timing,
658 amplitude, mixed-layer response, and coupled feedback against independent observations. Nonetheless, the ability of the
659 underlying MESMAR system to represent coupled extreme events has already been examined for Mediterranean tropical-like
660 cyclones by Storto et al. (2023).

661

662 Appendix A. Data-assimilation formulation in MESMAR-R

663 MESMAR-R combines an incremental three-dimensional variational analysis in the ocean with scale-selective spectral
 664 nudging of the atmospheric prognostic variables toward ERA5. The two procedures are applied independently in their
 665 respective components. Consequently, the system is weakly coupled: ocean observations do not generate an instantaneous
 666 atmospheric analysis increment, and the atmospheric large-scale constraint does not generate an instantaneous ocean analysis
 667 increment. Cross-component information is transferred dynamically during the subsequent coupled model integration through
 668 the exchanged surface fluxes, sea-surface state variables, atmospheric pressure, and river discharge.

669 Let $\mathbf{x}^b$ be the ocean background state at an analysis time, $\delta\mathbf{x} = \mathbf{x} - \mathbf{x}^b$ the model-space increment, $\mathbf{y}$ the observation vector,
 670 $\mathcal{H}$ the nonlinear observation operator, and $\mathbf{H}$ its tangent-linear approximation about $\mathbf{x}^b$. The incremental 3DVAR analysis is
 671 obtained by minimizing

$$J(\delta\mathbf{x}) = \frac{1}{2} \delta\mathbf{x}^T \mathbf{B}^{-1} \delta\mathbf{x} + \frac{1}{2} (\mathbf{d} - \mathbf{H}\delta\mathbf{x})^T \mathbf{R}^{-1} (\mathbf{d} - \mathbf{H}\delta\mathbf{x}) \quad (\text{A1})$$

672 where the innovation vector is

$$\mathbf{d} = \mathbf{y} - \mathcal{H}(\mathbf{x}^b) \quad (\text{A2})$$

673 and $\mathbf{B}$ and $\mathbf{R}$ are the background- and observation-error covariance matrices, respectively. Equation (A1) is three-dimensional
 674 because no model trajectory is included in the cost function. However, the scheme is usually referred to as FGAT (First Guess
 675 at Appropriate Time) as the innovations are formed by comparing the observations entering a given analysis with the
 676 background state closer to observation time (at a 6-hour frequency within the 3-day assimilation window).

677 For numerical efficiency, the background-error covariance is represented through a control-variable transform,

$$\mathbf{B} = \mathbf{V}\mathbf{V}^T, \quad \delta\mathbf{x} = \mathbf{V}\mathbf{v} \quad (\text{A3})$$

678 where $\mathbf{v}$ has approximately unit covariance. In MESMAR-R, $\mathbf{V}$ combines the multivariate vertical covariance structures
 679 represented by spatially varying empirical orthogonal functions with horizontal recursive-filter operators and the balance
 680 relationships used by the ocean analysis system. The covariance statistics are derived from monthly anomalies of the previous
 681 coupled-reanalysis iteration relative to the 1998–2021 climatology.

682 Substitution of Eq. (A3) into Eq. (A1) gives the preconditioned cost function

$$J(\mathbf{v}) = \frac{1}{2} \mathbf{v}^T \mathbf{v} + \frac{1}{2} (\mathbf{d} - \mathbf{H}\mathbf{V}\mathbf{v})^T \mathbf{R}^{-1} (\mathbf{d} - \mathbf{H}\mathbf{V}\mathbf{v}) \quad (\text{A4})$$

683 with gradient

$$\nabla_{\mathbf{v}} J = \mathbf{v} - \mathbf{V}^T \mathbf{H}^T \mathbf{R}^{-1} (\mathbf{d} - \mathbf{H}\mathbf{V}\mathbf{v}) \quad (\text{A5})$$

684 The minimizing control vector $\mathbf{v}^a$ provides the analyzed ocean state

$$\mathbf{x}^a = \mathbf{x}^b + \mathbf{V}\mathbf{v}^a \quad (\text{A6})$$

685 The observation vector includes temperature and salinity profiles and along-track altimetric observations. The altimetric
 686 component uses the dynamic-height observation operator and the Mediterranean mean dynamic topography described in the

687 main text. Variational quality control reduces the influence of observations whose departures are inconsistent with the assumed
 688 background and observation errors, without the need for rejecting them.

689 In addition to the discrete 3DVAR analyses, weak surface constraints are applied to sea-surface temperature and salinity as
 690 Newtonian relaxation tendencies,

$$\left. \frac{\partial \phi}{\partial t} \right|_{\text{relax}} = \frac{\phi^{\text{ref}} - \phi}{\tau_{\phi}}, \quad \phi \in \{\text{SST}, \text{SSS}\} \quad (\text{A7})$$

691 with e-folding times of approximately 15 d for SST and 300 d for SSS. A separate large-scale deep-ocean bias-correction term
 692 is applied as described by Storto and Yang (2024).

693 The atmospheric component is not analyzed through an observation-space variational method. Instead, the large spatial scales
 694 of selected WRF prognostic fields are continuously constrained toward three-hourly ERA5 fields. For a generic horizontal
 695 atmospheric field $\psi(\lambda, \varphi, z, t)$, its two-dimensional spectral representation at each model level may be written as

$$\psi(\lambda, \varphi, z, t) = \sum_k \sum_l \hat{\psi}_{kl}(z, t) \exp[i(k\lambda + l\varphi)] \quad (\text{A8})$$

696 The large-scale projection retained by the nudging is

$$\mathcal{P}_L \psi = \sum_{|k| \leq K_x} \sum_{|l| \leq K_y} \hat{\psi}_{kl} \exp[i(k\lambda + l\varphi)] \quad (\text{A9})$$

697 where $K_x = 6$ and $K_y = 5$ in the MESMAR-R configuration, corresponding approximately to wavelengths longer than 850
 698 km over the model domain. The prognostic equation is augmented by the scale-selective relaxation tendency

$$\frac{\partial \psi}{\partial t} = \mathcal{M}_{\psi}(\mathbf{x}) + G_{\psi} \mathcal{P}_L(\psi^{\text{ERA5}} - \psi) \quad (\text{A10})$$

699 where $\mathcal{M}_{\psi}$ contains the unnudged WRF dynamical and physical tendencies and G_{ψ} is the nudging coefficient. It is convenient
 700 to express the latter as

$$G_{\psi} = \tau_{\psi}^{-1} \quad (\text{A11})$$

701 with $\tau_{\psi} \simeq 1$ h for horizontal wind and temperature and $\tau_{\psi} \simeq 1$ d for humidity. ERA5 fields are temporally interpolated
 702 between their three-hourly input times before the nudging tendency is evaluated. Spatial scales shorter than the spectral cutoff
 703 are not directly relaxed, allowing regional circulation, boundary-layer structures, orographically forced features, and coupled
 704 mesoscale variability to develop according to the regional model dynamics and physics.

705 The term *weakly coupled* refers to the absence of cross-component terms in the instantaneous analysis operators. In compact
 706 form, the analysis step can be represented as

$$\mathbf{x}_o^a = \mathcal{A}_o(\mathbf{x}_o^b, \mathbf{y}_o), \quad \mathbf{x}_a^c = \mathcal{N}_a(\mathbf{x}_a, \mathbf{x}_a^{\text{ERA5}}) \quad (\text{A12})$$

707 where $\mathcal{A}_o$ denotes the ocean 3DVAR analysis and $\mathcal{N}_a$ the continuous atmospheric spectral-nudging operator. There is no
 708 coupled covariance term that directly maps ocean innovations into an atmospheric increment, or vice versa. Nevertheless, the
 709 ensuing model integration transfers the analyzed information between components through dynamically consistent air–sea and
 710 land–ocean exchanges (formally, $\mathbf{x}_o^a = \mathcal{A}_o(\mathbf{x}_o^b, \mathbf{y}_o) = \mathcal{A}_o(\mathcal{M}(\mathbf{x}_o^{a-}, \mathbf{x}_a^{c-}), \mathbf{y}_o)$, namely the background itself depends on the

711 previous cycle's analyses of both Earth system components ($\mathbf{x}_o^{a-}$ and $\mathbf{x}_a^{c-}$), so that the cross-medium observation propagation
712 is still present, but with delay.

Data availability: The MESMAR-R reanalysis data files used in this study are available for download from Storto et al. (2025) [<https://doi.org/10.5281/zenodo.17829847>]. Such data records contain monthly mean fields, and data assimilation files (observation innovation and statistics, analysis increments), together with post-processed data as used in the manuscripts (linear trends). Higher-frequency fields (6-hourly), not used in this study but produced, are available on request from the authors and are planned to be published in institutional data hubs soon.

Code availability: The NEMO Ocean model code (v4.0.7) is available at <https://forge.ipsl.jussieu.fr/nemo/wiki> (NEMO, 2023), the WRF atmospheric model code (v4.3.3) is available at <https://github.com/wrf-model/WRF> (wrf-model, 2023), and the HD hydrological discharge model (v5.1) is available at <https://doi.org/10.5281/zenodo.5707587> (Hagemann and Hagemann, 2021). The frozen version of the MESMAR v1 code used in this paper is available at <https://doi.org/10.5281/zenodo.7898938> (Storto et al., 2023). The ocean variational data assimilation is available at: https://baltig.cnr.it/nemo_ismar-rm/3dvar. Specific configuration files of the models and coupler used within this study are contained in the dataset record (Storto et al., 2025).

Author contribution: AS and VdT developed the regional coupled model; AS and CY designed the study; AS performed the reanalysis and the sensitivity experiments; AS prepared the manuscript with contributions from all co-authors.

Competing interests: At least one of the (co-)authors serves as editor for the special issue to which this paper belongs.

Acknowledgements

Financial support

This research is supported by programme CN00000013, the “National Centre for HPC, Big Data and Quantum Computing” Directorial Decree (grant no. 1031 of 17 June 2022), from the resources of the PNRR MUR – M4C2 – Investment 1.4 – “National Centers” Directorial Decree (grant no. 3138 of 16 December 2021).

734 **References**

- 735 Adler, R.F., Huffman, G.J., Chang, A., Ferraro, R., Xie, P., Janowiak, J., Rudolf, B., Schneider, U., Curtis, S., Bolvin, D.,
736 Gruber, A., Susskind, J., and Arkin, P.: The Version-2 Global Precipitation Climatology Project (GPCP) Monthly Precipitation
737 Analysis (1979–Present). *J. Hydrometeor.*, 4, 1147–1167, [https://doi.org/10.1175/1525-7541\(2003\)004<1147:TVGPCP>2.0.CO;2](https://doi.org/10.1175/1525-7541(2003)004<1147:TVGPCP>2.0.CO;2), 2003.
- 739 Akhtar, N., Brauch, J., Dobler, A., Béranger, K., and Ahrens, B.: Medicanes in an ocean–atmosphere coupled regional climate
740 model, *Nat. Hazards Earth Syst. Sci.*, 14, 2189–2201, <https://doi.org/10.5194/nhess-14-2189-2014>, 2014.
- 741 Aydogdu, A., Miraglio, P., Escudier, R., Clementi, E., and Masina, S.: The dynamical role of upper layer salinity in the
742 Mediterranean Sea, in: 7th edition of the Copernicus Ocean State Report (OSR7), von Schuckmann, K. et al. (Eds.), Copernicus
743 Publications, State Planet, 1-osr7, 6, <https://doi.org/10.5194/sp-1-osr7-6-2023>, 2023.
- 744 Ballarotta, M., Ubelmann, C., Veillard, P., Prandi, P., Etienne, H., Mulet, S., Faugère, Y., Dibarboure, G., Morrow, R., and
745 Picot, N.: Improved global sea surface height and current maps from remote sensing and in situ observations, *Earth Syst. Sci.*
746 *Data*, 15, 295–315, <https://doi.org/10.5194/essd-15-295-2023>, 2023.
- 747 Beranová, R., Huth, R., and Vít, V.: A multidataset analysis of precipitation trends in Europe, *J. Hydrometeorol.*, 26, 1003–
748 1015, <https://doi.org/10.1175/JHM-D-24-0114.1>, 2025.
- 749 Bhargava, K., Kalnay, E., Carton, J. A., and Yang, F.: Estimation of systematic errors in the GFS using analysis increments.
750 *J. Geophys. Res.-Atmos.*, 123, 1626–1637, <https://doi.org/10.1002/2017JD027423>, 2018.
- 751 Brewin, R. J. W., Sathyendranath, S., Jackson, T., Barlow, R., Brotas, V., Airs, R., and Lamont, T.: Influence of light in the
752 mixed-layer on the parameters of a three-component model of phytoplankton size class, *Remote Sens. Environ.*, 168, 437–
753 450, <https://doi.org/10.1016/j.rse.2015.07.004>, 2015.
- 754 Calafat, F. M., Chambers, D. P., and Tsimplis, M. N.: Mechanisms of decadal sea level variability in the eastern North Atlantic
755 and the Mediterranean Sea, *J. Geophys. Res.-Oceans*, 117, C09022, <https://doi.org/10.1029/2012JC008285>, 2012.
- 756 Camus, P., Haigh, I. D., Nasr, A. A., Wahl, T., Darby, S. E., and Nicholls, R. J.: Regional analysis of multivariate compound
757 coastal flooding potential around Europe and environs: sensitivity analysis and spatial patterns, *Nat. Hazards Earth Syst. Sci.*,
758 21, 2021–2040, <https://doi.org/10.5194/nhess-21-2021-2021>, 2021.
- 759 Canuto, V. M., Howard, A., Cheng, Y., and Dubovikov, M. S.: Ocean turbulence. Part I: One-point closure model—momentum
760 and heat vertical diffusivities, *J. Phys. Oceanogr.*, 31, 1413–1426, 2001.
- 761 Choi, Y.-S., and Lee, D.-K.: Spectral nudging to the global reanalysis data in the WRF regional model, *Asia-Pac. J. Atmos.*
762 *Sci.*, 52, 135–150, <https://doi.org/10.1007/s13143-016-0004-9>, 2016.
- 763 Craig, A. P., Valcke, S., and Coquart, L.: Development and performance of a new version of the OASIS coupler, OASIS3-
764 MCT_3.0, *Geosci. Model Dev.*, 10, 3297–3308, <https://doi.org/10.5194/gmd-10-3297-2017>, 2017.
- 765 Cramer, W., Guiot, J., Fader, M., et al.: Climate change and interconnected risks to sustainable development in the
766 Mediterranean, *Nat. Clim. Change*, 8, 972–980, <https://doi.org/10.1038/s41558-018-0299-2>, 2018.
- 767 de Figueiredo, D. R., Bento, S., and Condesso de Melo, M. T.: Freshwater quality challenges in southern Europe under an
768 increasingly warmer and drier climate scenario, *Water*, 17, 1873, <https://doi.org/10.3390/w17131873>, 2025.

769 Doury, A., Somot, S., Gadat, S., et al.: Regional climate model emulator based on deep learning: concept and first evaluation
770 of a novel hybrid downscaling approach, *Clim. Dyn.*, 60, 1751–1779, <https://doi.org/10.1007/s00382-022-06343-9>, 2023.

771 Escudier, R., Clementi, E., Cipollone, A., Pistoia, J., Drudi, M., Grandi, A., Lyubartsev, V., Lecci, R., Aydogdu, A., Delrosso,
772 D., Omar, M., Masina, S., Coppini, G., and Pinardi, N.: A high-resolution reanalysis for the Mediterranean Sea, *Front. Earth*
773 *Sci.*, 9, 702285, <https://doi.org/10.3389/feart.2021.702285>, 2021.

774 Fox-Hughes, P., Su, C.-H., Eizenberg, N., et al.: A review of early severe weather applications of high-resolution regional
775 reanalysis in Australia, *Meteorol. Appl.*, 29, e2087, <https://doi.org/10.1002/met.2087>, 2022.

776 Giorgi, F., and Lionello, P.: Climate change projections for the Mediterranean region, *Global Planet. Change*, 63, 90–104,
777 2008.

778 González-Hidalgo, J. C., and Vicente-Serrano, S. M.: Is there a precipitation decline in the Mediterranean region? An
779 assessment based on the scientific literature, *Int. J. Climatol.*, 45, e8918, <https://doi.org/10.1002/joc.8918>, 2025.

780 Good, S. A., Martin, M. J., and Rayner, N. A.: EN4: quality controlled ocean temperature and salinity profiles and monthly
781 objective analyses with uncertainty estimates, *J. Geophys. Res.-Oceans*, 118, 6704–6716,
782 <https://doi.org/10.1002/2013JC009067>, 2013.

783 Hagemann, S., and Dümenil Gates, L.: Validation of the hydrological cycle of ECMWF and NCEP reanalyses using the MPI
784 hydrological discharge model, *J. Geophys. Res.-Atmos.*, 106, 17901–17917, <https://doi.org/10.1029/2000JD900568>, 2001.

785 Hagemann, S., Ho-Hagemann, H. T. M., and Rechid, D.: The hydrological discharge model (HD model, version 5.1) –
786 description and technical documentation, Zenodo, <https://doi.org/10.5281/zenodo.5707587>, 2020.

787 Haylock, M. R., Hofstra, N., Klein Tank, A. M. G., Klok, E. J., Jones, P. D., and New, M.: A European daily high-resolution
788 gridded data set of surface temperature and precipitation for 1950–2006, *J. Geophys. Res.-Atmos.*, 113, D20119, 2008.

789 Hersbach, H., Bell, B., Berrisford, P., et al.: The ERA5 global reanalysis, *Q. J. Roy. Meteor. Soc.*, 146, 1999–2049,
790 <https://doi.org/10.1002/qj.3803>, 2020.

791 Hertig, E., and Trambly, Y.: Regional downscaling of Mediterranean droughts under past and future climatic conditions,
792 *Global Planet. Change*, 151, 36–48, <https://doi.org/10.1016/j.gloplacha.2016.10.015>, 2017.

793 Hoerling, M., Eischeid, J., Perlwitz, J., Quan, X., Zhang, T., and Pegion, P.: On the increased frequency of Mediterranean
794 drought, *J. Climate*, 25, 2146–2161, <https://doi.org/10.1175/JCLI-D-11-00296.1>, 2012.

795 Jousset, S., Mulet, S., Greiner, E., et al.: New global mean dynamic topography CNES-CLS-22 combining drifters,
796 hydrography profiles and high frequency radar data, ESS Open Archive,
797 <https://doi.org/10.22541/essoar.170158328.85804859/v2>, 2025.

798 Juza, M., Fernández-Mora, À., and Tintoré, J.: Sub-regional marine heat waves in the Mediterranean Sea from observations:
799 long-term surface changes, sub-surface and coastal responses, *Front. Mar. Sci.*, 9, 785771,
800 <https://doi.org/10.3389/fmars.2022.785771>, 2022.

801 Kourafalou, V. H., and Barbopoulos, K.: High resolution simulations on the North Aegean Sea seasonal circulation, *Ann.*
802 *Geophys.*, 21, 251–265, <https://doi.org/10.5194/angeo-21-251-2003>, 2003.

803 Laloyaux, P., de Boisseson, E., Balmaseda, M., Bidlot, J.-R., Broennimann, S., Buizza, R., et al.: CERA-20C: A coupled
804 reanalysis of the twentieth century. *Journal of Advances in Modeling Earth Systems*, 10, 1172–1195,
805 <https://doi.org/10.1029/2018MS001273>, 2018.

806 Lellouche, J.-M., Greiner, E., Romain, B., et al.: The Copernicus global 1/12° oceanic and sea ice GLORYS12 reanalysis,
807 *Front. Earth Sci.*, 9, 698876, <https://doi.org/10.3389/feart.2021.698876>, 2021.

808 Lhotka, O., Bešťáková, Z., and Kyselý, J.: Prolongation of compound dry–hot seasons over Europe under climate change
809 scenarios, *Earth’s Future*, 11, e2023EF003557, <https://doi.org/10.1029/2023EF003557>, 2023.

810 Lionello, P., and Scarascia, L.: The relation between climate change in the Mediterranean region and global warming, *Reg.*
811 *Environ. Change*, 18, 1481–1493, <https://doi.org/10.1007/s10113-018-1290-1>, 2018.

812 Liu, C., Liang, X., and Yu, L.: Salinity trends and mass balances in the Mediterranean Sea: revisit the role of air–sea freshwater
813 fluxes and oceanic exchange, *Ocean Sci.*, 21, 2069–2083, <https://doi.org/10.5194/os-21-2069-2025>, 2025.

814 Madec, G., and the NEMO Team: NEMO ocean engine – version 4.0, Institut Pierre-Simon Laplace (IPSL), France, 2017.

815 Marcos, M., and Tsimplis, M. N.: Coastal sea level trends in southern Europe, *Geophys. J. Int.*, 175, 70–82,
816 <https://doi.org/10.1111/j.1365-246X.2008.03892.x>, 2008.

817 Margirier, F., Testor, P., Heslop, E., et al.: Abrupt warming and salinification of intermediate waters interplays with decline
818 of deep convection in the northwestern Mediterranean Sea, *Sci. Rep.*, 10, 20923, [https://doi.org/10.1038/s41598-020-77859-](https://doi.org/10.1038/s41598-020-77859-5)
819 5, 2020.

820 Marti, O., Nguyen, S., Braconnot, P., Valcke, S., Lemarié, F., and Blayo, E.: A Schwarz iterative method to evaluate ocean–
821 atmosphere coupling schemes: implementation and diagnostics in IPSL-CM6-SW-VLR, *Geosci. Model Dev.*, 14, 2959–2975,
822 <https://doi.org/10.5194/gmd-14-2959-2021>, 2021.

823 MedECC: Climate and Environmental Change in the Mediterranean Basin – Current Situation and Risks for the Future, First
824 Mediterranean Assessment Report, Union for the Mediterranean, Plan Bleu, UNEP/MAP, Marseille, France, 632 pp.,
825 <https://doi.org/10.5281/zenodo.4768833>, 2020.

826 Morel, A., and Maritorena, S.: Bio-optical properties of oceanic waters: a reappraisal, *J. Geophys. Res.-Oceans*, 106, 7163–
827 7180, <https://doi.org/10.1029/2000JC000319>, 2001.

828 Nakanishi, M., and Niino, H.: An improved Mellor–Yamada level-3 model: its numerical stability and application to a regional
829 prediction of advection fog, *Bound.-Lay. Meteorol.*, 119, 397–407, <https://doi.org/10.1007/s10546-005-9030-8>, 2006.

830 Pinardi, N., and Masetti, E.: Variability of the large scale general circulation of the Mediterranean Sea from observations and
831 modelling: A review. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 158, 153–173, [https://doi.org/10.1016/S0031-0182\(00\)00048-](https://doi.org/10.1016/S0031-0182(00)00048-1)
832 1, 2000.

833 Pisano, A., et al.: New high-resolution Mediterranean SST reanalysis products from the Copernicus Marine Environment
834 Monitoring Service, *Remote Sens. Environ.*, 243, 111781, <https://doi.org/10.1016/j.rse.2020.111781>, 2020.

835 Pisano, A., Marullo, S., Artale, V., et al.: New evidence of Mediterranean climate change and variability from sea surface
836 temperature observations, *Remote Sens.*, 12, 132, <https://doi.org/10.3390/rs12010132>, 2020.

837 Pujol, M.-I., Dupuy, S., Vergara, O., et al.: Refining the resolution of DUACS along-track Level-3 sea level altimetry products,
838 *Remote Sens.*, 15, 793, <https://doi.org/10.3390/rs15030793>, 2023.

839 Ridal, M., Bazile, E., Le Moigne, P., et al.: CERRA, the Copernicus European Regional Reanalysis system, *Q. J. Roy. Meteor.*
840 *Soc.*, 150, 3385–3411, <https://doi.org/10.1002/qj.4764>, 2024.

841 Ruti, P. M., et al.: Med-CORDEX Initiative for Mediterranean climate studies, *Bull. Amer. Meteor. Soc.*, 97, 1187–1208,
842 <https://doi.org/10.1175/BAMS-D-14-00176.1>, 2016.

843 Sammartino, M., Aronica, S., Santoleri, R., and Buongiorno Nardelli, B.: Retrieving Mediterranean Sea Surface Salinity
844 Distribution and Interannual Trends from Multi-Sensor Satellite and In Situ Data. *Remote Sensing*, 14(10), 2502,
845 <https://doi.org/10.3390/rs14102502>, 2022.

846 Schüller, V., et al.: Quantifying coupling errors in atmosphere–ocean–sea-ice models: a study of iterative and non-iterative
847 approaches in the EC-Earth AOSCM, *Geosci. Model Dev.*, 18, 9167–9187, <https://doi.org/10.5194/gmd-18-9167-2025>, 2025.

848 Senatore, A., Davolio, S., Furnari, L., and Mendicino, G.: Reconstructing flood events in Mediterranean coastal areas using
849 different reanalyses and high-resolution meteorological models, *J. Hydrometeorol.*, 21, 1865–1887,
850 <https://doi.org/10.1175/JHM-D-19-0270.1>, 2020.

851 Skamarock, W. C., Klemp, J. B., Duda, M. G., et al.: A description of the Advanced Research WRF model version 4.3.3,
852 NCAR Tech. Note, <https://doi.org/10.5065/1dfh-6p97>, 2021.

853 Skliris, N., Zika, J. D., Herold, L., et al.: Mediterranean Sea water budget long-term trend inferred from salinity observations,
854 *Clim. Dyn.*, 51, 2857–2876, <https://doi.org/10.1007/s00382-017-4053-7>, 2018.

855 Soukissian, T. H., Karathanasi, F., and Axaopoulos, P.: Offshore renewable energy in the Mediterranean Sea: a review of the
856 resource, existing projects and prospects, *Renew. Energy*, 179, 1225–1241, <https://doi.org/10.1016/j.renene.2021.07.109>,
857 2021.

858 Storto, A., Dobricic, S., Masina, S., and Di Pietro, P.: Assimilating along-track altimetric observations through local
859 hydrostatic adjustment in a global ocean variational assimilation system, *Mon. Weather Rev.*, 139, 738–754,
860 <https://doi.org/10.1175/2010MWR3350.1>, 2011.

861 Storto, A.: Variational quality control of hydrographic profile data with non-Gaussian errors for global ocean variational data
862 assimilation systems, *Ocean Model.*, 104, 226–241, <https://doi.org/10.1016/j.ocemod.2016.06.011>, 2016.

863 Storto, A., Oddo, P., Cipollone, A., Mirouze, I., and Lemieux-Dudon, B.: Extending an oceanographic variational scheme to
864 allow for affordable hybrid and four-dimensional data assimilation, *Ocean Model.*, 128, 67–86,
865 <https://doi.org/10.1016/j.ocemod.2018.06.005>, 2018.

866 Storto, A., Alvera-Azcárate, A., Balmaseda, M. A., et al.: Ocean reanalyses: recent advances and unsolved challenges, *Front.*
867 *Mar. Sci.*, 6, 418, <https://doi.org/10.3389/fmars.2019.00418>, 2019a.

868 Storto, A., Masina, S., Simoncelli, S., et al.: The added value of the multi-system spread information for ocean heat content
869 and steric sea level investigations in the CMEMS GREP ensemble reanalysis product, *Clim. Dyn.*, 53, 287–312,
870 <https://doi.org/10.1007/s00382-018-4585-5>, 2019b.

871 Storto, A., Hesham Essa, Y., de Toma, V., Anav, A., Sannino, G., Santoleri, R., and Yang, C.: MESMAR v1: a new regional
872 coupled climate model for downscaling, predictability, and data assimilation studies in the Mediterranean region, *Geosci.*
873 *Model Dev.*, 16, 4811–4833, <https://doi.org/10.5194/gmd-16-4811-2023>, 2023.

874 Storto, A., and Yang, C.: Acceleration of the ocean warming from 1961 to 2022 unveiled by large-ensemble reanalyses, *Nat.*
875 *Commun.*, 15, 545, <https://doi.org/10.1038/s41467-024-44749-7>, 2024.

876 Storto, A., Chierici, G., Pfeffer, J., Barnoud, A., Bourdalle-Badie, R., Blazquez, A., Cavaliere, D., Lalau, N., Couprie, B.,
877 Drevillon, M., Fourest, S., Larnicol, G., and Yang, C.: Variability in manometric sea level from reanalyses and observation-
878 based products over the Arctic and North Atlantic oceans and the Mediterranean Sea, in: 8th edition of the Copernicus Ocean
879 State Report (OSR8), edited by: von Schuckmann, K., Moreira, L., Grégoire, M., Marcos, M., Staneva, J., Brasseur, P., Garric,
880 G., Lionello, P., Karstensen, J., and Neukermans, G., Copernicus Publications, State Planet, 4-osr8, 12,
881 <https://doi.org/10.5194/sp-4-osr8-12-2024>, 2024.

882 Storto, A., de Toma, V., and Yang, C.: MESMAR-R: A Coupled Regional Reanalysis for the Mediterranean Region Covering
883 the period 1993-2024 [Data set]. Zenodo. <https://doi.org/10.5281/zenodo.17829847>, 2025.

884 Tootoonchi, R., Bordoni, S., and D’Agostino, R.: Revisiting the moisture budget of the Mediterranean region in the ERA5
885 reanalysis, *Weather Clim. Dynam.*, 6, 245–263, <https://doi.org/10.5194/wcd-6-245-2025>, 2025.

886 van der Meer, M., de Roda Husman, S., and Lhermitte, S.: Deep learning regional climate model emulators: a comparison of
887 two downscaling training frameworks, *J. Adv. Model. Earth Syst.*, 15, e2022MS003593,
888 <https://doi.org/10.1029/2022MS003593>, 2023.

889 von Schuckmann, K., Storto, A., Simoncelli, S., Raj, R.P., Samuelson, A., de Pascual Collar, A., Garcia Sotillo, M., Szerkely,
890 T., Mayer, M., Peterson, K.A., et al.: Ocean heat content, in CMEMS ocean state report 2, *J. Oper. Oceanogr.* 11 (sup1), S1–
891 S142, <https://doi.org/10.1080/1755876X.2018.1489208>, 2018.

892 von Schuckmann, K., Palmer, M., Trenberth, K., et al.: An imperative to monitor Earth’s energy imbalance, *Nat. Clim. Change*,
893 6, 138–144, <https://doi.org/10.1038/nclimate2876>, 2016.

894 Yepes-Arbós, X., van den Oord, G., Acosta, M. C., and Carver, G. D.: Evaluation and optimisation of the I/O scalability for
895 the next generation of Earth system models: IFS CY43R3 and XIOS 2.0 integration as a case study, *Geoscientific Model*
896 *Development*, 15, 379–394, [doi:10.5194/gmd-15-379-2022](https://doi.org/10.5194/gmd-15-379-2022), 2022.

897 Zittis, G., Bruggeman, A., and Lelieveld, J.: Revisiting future extreme precipitation trends in the Mediterranean, *Weather Clim.*
898 *Extrem.*, 34, 100380, <https://doi.org/10.1016/j.wace.2021.100380>, 2021.

899 Zuo, H., Balmaseda, M. A., Tietsche, S., Mogensen, K., and Mayer, M.: The ECMWF operational ensemble reanalysis–
900 analysis system for ocean and sea ice: a description of the system and assessment, *Ocean Sci.*, 15, 779–808,
901 <https://doi.org/10.5194/os-15-779-2019>, 2019.