

An improved empirical model for predicting postfire debris-flow volume in the western United States

Alexander N. Gorr¹ ORCID: 0000-0002-3239-7773, Francis K. Rengers¹ ORCID: 0000-0002-1825-0943, Katherine R. Barnhart¹ ORCID: 0000-0001-5682-455X, Matthew A. Thomas¹
5 ORCID: 0000-0002-9828-5539, Jason W. Kean¹ ORCID: 0000-0003-3089-0369

¹U.S. Geological Survey, Geologic Hazards Science Center, Golden, CO, USA

Correspondence to: Alexander N. Gorr (agorr@usgs.gov)

10

15

20

25

Abstract. Reliable estimates of debris-flow volume can be used to help predict the magnitude of debris-flow hazards following wildfire in the western United States. In this study, we compiled and used a database of 227 postfire debris-flow volumes that were collected across the western United States to develop a multiple linear regression model for predicting postfire debris-flow volume. We explored 36 predictor variables related to rainfall, terrain, and fire characteristics, and selected the model with the combination of variables that yielded the most accurate predictions of debris-flow volume. We evaluated model performance against the entire volume database, as well as against four subsets of volume data from southern California, the Intermountain West, the Southwest, and regions with limited volume data, such as northern California and Washington. We also compared model performance against three existing postfire debris-flow volume models that were developed for use in southern California, the Intermountain West, and the Southwest. We demonstrate that the new volume model performs as well as the regional models in the regions for which they were developed and outperforms existing models when applied to volumes from data-limited regions in the western United States. These results indicate that the debris-flow volume model introduced in this study can be used to improve postfire hazard assessments across the western United States, especially outside of southern California.

1 Introduction

Debris flows are a common hazard in mountainous areas around the world (e.g., Rickenmann and Zimmermann, 1993; Wang et al., 2003; Cannon and Gartner, 2005; Sepúlveda et al., 2006; Guthrie et al., 2012; Gartner et al., 2024) but are particularly prevalent in steep landscapes that have been recently burned by wildfire. Wildfire reduces vegetation cover (e.g., McGuire et al., 2024a) and alters soil-hydraulic properties (e.g., Ebel et al., 2012; Hoch et al., 2021), which promotes the initiation of runoff-generated debris flows in burned watersheds (e.g., Cannon et al., 2001; Parise and Cannon, 2012; Wall et al., 2020). As a result, burned watersheds are more likely to produce debris flows than comparable unburned watersheds given similar rainfall conditions (McGuire et al., 2021). Burned watersheds also tend to produce larger debris flows than unburned watersheds (Santi and Morandi, 2013), resulting in elevated downstream effects, including the loss of human life (Dowling and Santi, 2014; Kean et al., 2019; Daurio, 2025), damage to infrastructure (e.g., Lancaster et al., 2021), and degradation of water quality (e.g., Smith et al., 2011; Langhans et al., 2016), for communities in fire-prone regions of the western United States (U.S.).

Recent increases in postfire debris-flow activity in the western United States, driven by changes in wildfire activity (Westerling, 2016) and growth in the wildland-urban interface (Radeloff et al., 2018), have motivated the development of a postfire hazard assessment framework that is used by the U.S. Geological Survey (USGS) to mitigate the impact of potential postfire debris flows. The USGS framework uses postfire debris-flow likelihood (Staley et al., 2017) and volume (Gartner et al., 2014) models to generate a combined hazard map for all watersheds within a fire perimeter and identifies the most hazardous watersheds as those that have a high likelihood of debris-flow occurrence and are likely to produce a debris flow that mobilizes a large volume of sediment (Cannon et al., 2010; Landslide Hazards Program, 2018). Methods for predicting debris-flow likelihood can be used to identify which upstream watersheds are likely to produce postfire debris flows. Methods for predicting debris-flow volume, on the other hand, provide insight into the potential magnitude of downstream effects of postfire debris flows, as multiple studies have found that the

area inundated by a debris flow scales with volume (Iverson et al., 1998; Berti and Simoni, 2007; Griswold and Iverson, 2008). Accurate predictions of volume are also used to inform runout models that can evaluate the potential downstream effects of postfire debris flows (Barnhart et al., 2021; Gorr et al., 2022).

65 Although numerous methods for predicting postfire debris-flow volume have been developed in recent years (e.g., Gartner et al., 2008; Pak and Lee, 2008; Cannon et al., 2010; Santi and Morandi, 2013; Gartner et al., 2014; Pelletier and Orem, 2014; Donovan and Santi, 2017; Wall et al., 2023; Gorr et al., 2024a), none are ideally suited for use in postfire hazard assessment frameworks that are applied across the entire western United States. Multiple volume models have been developed for broad use across the western United States (e.g., Gartner et al., 2008; Cannon et al., 70 2010; Santi and Morandi, 2013; Pelletier and Orem, 2014) but have deficiencies that limit their use in hazard assessment scenarios. Specifically, existing broadly applicable volume models do not include rainfall variables, despite the fact the volume of postfire debris flows in the western United States is known to scale with short-duration (≤ 1 h) rainfall intensity (Gartner et al., 2008; Pak and Lee, 2008; Cannon et al., 2010; Gartner et al., 2014; Gorr et al., 2024a). The lack of rainfall variables limits the accuracy of these models, particularly when compared to volume 75 models that do include rainfall variables (Gorr et al., 2024a). Volume models that do not consider rainfall are also unable to predict postfire debris-flow volume based on rainfall forecasts, which is a practical benefit offered by volume models that do (Prescott et al., 2024).

There are several postfire debris-flow volume models that include rainfall variables (e.g., Gatwood et al., 2000; Pak and Lee, 2008; Gartner et al., 2008, 2014; Gorr et al., 2024a), but they also have shortcomings that limit their 80 applicability in widespread postfire hazard assessments across the western United States. First, most existing volume models that include rainfall variables are regionally focused to predict volume within a specific area, such as southern California (e.g., Gartner et al., 2014), the Intermountain West (e.g., Wall et al., 2023), or the Southwest (Arizona and New Mexico) (Gorr et al., 2024a). Previous studies have found that these regional models perform well in the areas for which they were developed (e.g., Kean et al., 2019; Wall et al., 2023; Gorr et al., 2024a) but are considerably less 85 accurate when applied to areas outside of their training datasets (e.g., Gorr et al., 2023, 2024a; Rengers et al., 2023, 2024). For example, previous studies have found that a volume model developed for use in southern California (Gartner et al., 2014) overpredicted volumes in other regions of the western United States by up to several orders of magnitude (Gorr et al., 2024a). The decreased performance of regional models in these scenarios may be partially attributed to the fact they use rainfall intensity, even though the intensity of debris-flow-producing rainfall varies 90 widely across the western United States (Staley et al., 2017). As a result, models developed for use in areas where the rainfall intensity required to generate postfire debris flows is low (e.g., Gartner et al., 2014), such as the Transverse Ranges of southern California, where the 15-minute rainfall intensity-duration threshold for debris-flow occurrence is less than 20 mm/h (Staley et al., 2013), tend to overpredict volumes when applied in areas where the intensity required to generate postfire debris flows is much higher, such as northern Arizona, where the 15-minute rainfall intensity-duration threshold for debris-flow occurrence is more than 60 mm/h (Youberg, 2014). Conversely, models developed 95 for use in areas with intense debris-flow-generating rainfall (Gorr et al., 2024a) tend to underpredict volumes in areas with less intense debris-flow-generating rainfall. Furthermore, parts of the western United States lack the volume data

needed to develop regional models. For instance, although the Pacific Northwest (Oregon and Washington) and the northern Rockies (Idaho, Montana, and Wyoming) are susceptible to postfire debris flows (e.g., Meyer and Wells, 1997; Gabet and Bookter, 2008; Wall et al., 2020; Selander et al., 2025), insufficient data has prohibited the development of volume models in these regions. The shortcomings of existing postfire debris-flow volume models indicate that a model that includes a rainfall variable and can be applied broadly across the western United States would be beneficial for improving postfire hazard assessments, particularly in regions with limited volume data.

In this study, we developed a new method for predicting postfire debris-flow volume in the western United States for the purpose of improving postfire hazard assessments. Specifically, we compiled and used the largest known database of postfire debris-flow volumes with associated rainfall data (Gorr et al., 2025) to develop a multiple linear regression model that predicts postfire debris-flow sediment volume. We explored 36 potential predictor variables related to rainfall, terrain, and fire characteristics, and selected the combination of three variables that yielded the most accurate predictions of debris-flow volume. We assessed model performance against the entire volume database, which includes 227 postfire debris-flow volumes across six states, as well as against three subsets of data from regions in the western United States that have published regional volume models: southern California (Gartner et al., 2014), the Intermountain West (Wall et al., 2023), and the Southwest (Gorr et al., 2024a). We then compared the performance of the new model with three existing regional models. Finally, we evaluated the performance of all four models when applied to volumes from data-limited regions, which we define as areas that do not have enough volume data to develop regional volume models. Results from this study can improve our ability to accurately predict postfire debris-flow volume across the western United States, particularly in data-limited regions.

2 Data

2.1 Debris-flow volumes

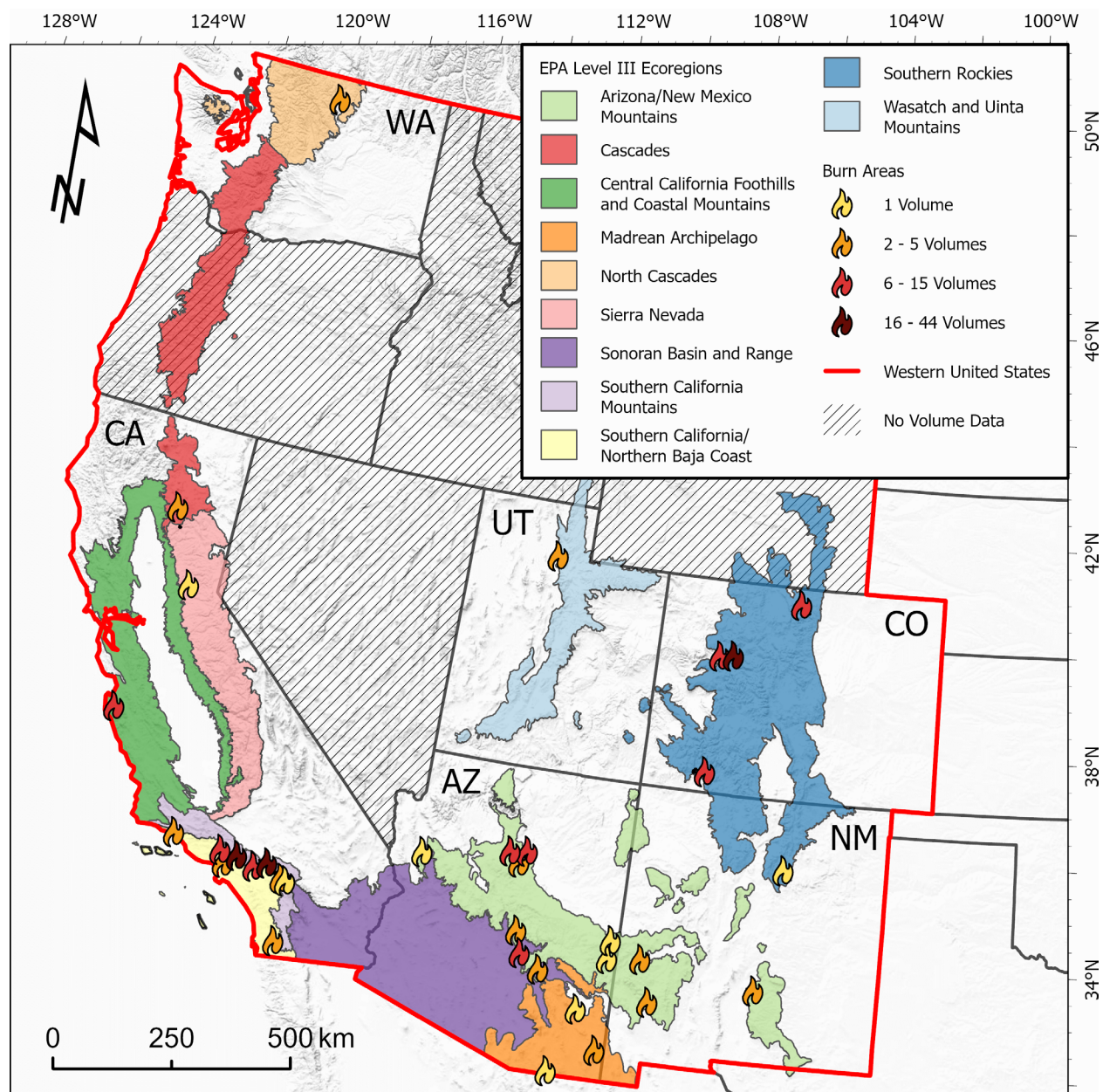
We compiled a database of 227 postfire debris-flow volumes from across the western United States (Figure 1) to develop the new volume model introduced in this study. Roughly 85% of the database (192 of 227 volumes) consists of previously published postfire debris-flow volumes from Arizona (Gorr et al., 2024a); California (Gartner et al., 2008; Gartner et al., 2014; Kean et al., 2019; Smith et al., 2021; Swanson et al., 2024), Colorado (Gartner et al., 2008; Rengers et al., 2023), New Mexico (Gorr et al., 2024a), and Utah (Gartner et al., 2008). We collected the remaining 15% of volumes (35) from sites in Arizona, northern California, Colorado, New Mexico, and Washington as part of this study (Table 1). All volumes represent the volume of sediment deposited downstream from the watershed outlet. We did not consider the volume of water mobilized by flows, nor any sediment that may have been mobilized and deposited upstream from the watershed outlet. This is consistent with the data used to develop previous postfire debris-flow volume models (e.g., Gartner et al., 2014; Gorr et al., 2024a).

130

Table 1: Fire and associated debris-flow information. Fire information includes the fire name, year of occurrence, and location, including Arizona (AZ), California (CA), Colorado (CO), New Mexico (NM), Utah (UT), and Washington (WA). Debris-flow information includes the number of volume measurements for each fire, the range of associated sediment volumes, and the original sources of the volume measurements.

Fire Name	Year	State	No. Measurements	Range of Debris- Flow Volumes (m³)	Source
Apple	2020	CA	1	103,400	Swanson et al. (2024)
Bush	2020	AZ	3	140 – 680	Gorr et al. (2024a)
Buzzard	2018	NM	5	140 – 740	Gorr et al. (2024a)
Cameron Peak	2020	CO	14	10 – 87,000	Gorr et al. (2025)
Carmel	2020	CA	11	24 – 990	Smith et al. (2021)
Cedar	2003	CA	2	383 – 444	Gartner et al. (2008)
Coal Seam	2002	CO	6	237 – 1,387	Gartner et al. (2008)
Cub Creek 2	2021	WA	2	4,932 – 10,276	Gorr et al. (2025)
Dixie	2021	CA	2	2,044 – 13,356	Thomas et al. (2023)
El Dorado	2020	CA	2	17,196 – 75,000	Swanson et al. (2024)
Farmington	2003	UT	3	511 – 1,515	Gartner et al. (2008)
Flag	2021	AZ	1	1,130	Gorr et al. (2023)
Frye	2017	AZ	1	10,000	Gorr et al. (2024a)
Grand Prix	2003	CA	7	570 – 801,770	Gartner et al. (2008)
Grizzly Creek	2020	CO	19	478 – 57,228	Rengers et al. (2024)
Harvard	2005	CA	4	22,116 – 65,284	Gartner et al. (2014)
Hermits Peak	2022	NM	1	4,000	Gorr et al. (2025)
Horseshoe 2	2011	AZ	4	650 – 3,000	Gorr et al. (2024a)
Horton	2021	AZ	1	1,170	Gorr et al. (2024a)
Missionary Ridge	2002	CO	8	752 – 23,571	Gartner et al. (2008)
Monument	2011	AZ	1	10,000	Gorr et al. (2024a)
Mosquito	2022	CA	1	16,561	Gorr et al. (2025)
Museum	2019	AZ	4	20 – 2,980	Gorr et al. (2024a)
Old	2003	CA	17	1,622 – 488,463	Gartner et al. (2008)
Pipeline	2022	AZ	14	372 – 35,737	Gorr et al. (2025)
Sayre	2008	CA	10	1,239 – 172,205	Gartner et al. (2014)
Schultz	2010	AZ	11	1,000 – 14,000	Gorr et al. (2024a)
Station	2009	CA	45	75 – 91,800	Gartner et al. (2014)
Tadpole	2020	NM	4	25 – 150	Gorr et al. (2024a)
Telegraph	2021	AZ	4	540 – 1,700	Gorr et al. (2024a)

Thomas	2017	CA	5	10,000 – 297,000	Kean et al. (2019)
Three Rivers	2021	NM	2	200 – 1,400	Gorr et al. (2024a)
Wallow	2011	AZ	1	7,500	Gorr et al. (2024a)
Woodbury	2019	AZ	11	30 – 1,200	Gorr et al. (2024a)



135 **Figure 1: Map of the locations of the 34 burn areas included in this study. The burn areas span six states across the western United States (US), including Arizona (AZ), California (CA), Colorado (CO), New Mexico (NM),**

Utah (UT), and Washington (WA), and 11 Environmental Protection Agency (EPA) Level III Ecoregions. The names of the ecoregions shown in this figure are derived directly from the EPA (U.S. Environmental Protection Agency, 2013). Basemap credits: United States Geological Survey The National Map: 3D Elevation Program, United States Geological Survey Earth Resources Observation & Science Center: GMTED2010.

The volumes that we compiled were collected using a range of field and remote-sensing techniques. Most volumes were measured using some variation of the field survey methods outlined in Gorr et al. (2024a). In short, measurements of deposit area and average thickness were made in the field and then multiplied to determine debris-flow volume (e.g., Gorr et al., 2024a; Swanson et al., 2024). Other methods used to measure postfire debris-flow volume in the field included surveys of closely spaced channel cross-sections (e.g., Gartner et al., 2008) and counting the number of trucks filled with sediment when emptying a debris-retention basin (truck counts) (Gartner et al., 2014). The remaining volumes were measured primarily using remote-sensing techniques. Most commonly, these volumes were calculated using digital elevation models (DEM) of difference (DoD) that were generated by differencing pre-event and post-event light detection and ranging (lidar) (Smith et al., 2021; Rengers et al., 2024; Swanson et al., 2024). In other cases, high-resolution aerial imagery was used to help constrain the area of larger debris flows, and volume was calculated by multiplying the area by depth measurements made in the field (e.g., Gorr et al., 2024a).

Variations in the size of the debris-flow volumes included in this database, and the techniques used to measure them, mean that the uncertainty associated with each volume varies widely. Santi (2014) determined that the uncertainty associated with deposit boundary and thickness measurements, the most used method to measure volumes in this database, was -25% to +35% for small debris flows (~1,500 m³), -28% to +30% for medium debris flows (~15,000 m³), and -9% to +17% for large debris flows (~150,000 m³). Other field measurement techniques, such as truck counts and channel cross-section surveys, have a lower degree of uncertainty, but still vary between -25% and +20%, depending on debris-flow size (Santi, 2014). The uncertainty associated with volumes measured by remote-sensing techniques are less constrained, but we estimate that the volumes calculated by lidar differencing have an uncertainty of -14% to +14%, based on a ±10 cm level of detection (LoD) (Rengers et al., 2024). Overall, given the wide range of debris-flow sizes and measurement techniques included in the volume database, we conservatively estimate the uncertainty associated with these volume measurements to be ±25%.

The volume database includes data from 195 watersheds from 34 burn areas across six states in the western United States, a region we define as the states of Arizona, California, Colorado, Idaho, Montana, Nevada, New Mexico, Oregon, Utah, Washington, and Wyoming (Figure 1; Table S1). Specifically, the volume database includes volumes from Arizona, California, Colorado, New Mexico, Utah, and Washington (Figure 1). The burn areas included in this study range in size from 4.2 km² to 3,965 km² and span a wide range of climatological settings. The mean annual precipitation at the burn areas ranges from 396 mm to 1,343 mm, and the mean annual temperature ranges from 3.7 °C to 17.8 °C (PRISM Climate Group, 2025) (Table S1).

The burn areas are also geographically and ecologically diverse, as they span 11 Level III ecoregions, or areas where ecosystems and ecosystem components, including geology, vegetation, climate, and hydrology, are generally similar (U.S. Environmental Protection Agency, 2013) (Figure 1). The names of the ecoregions presented here are derived

directly from U.S. Environmental Protection Agency (2013). The Arizona/New Mexico Mountains ecoregion, which is characterized by steep foothills, mountains, and dissected plateaus (Wilken et al., 2011), contains 12 burn areas. Grassland, chaparral, and pinyon-juniper and oak woodlands grow at lower elevations in this ecoregion, whereas ponderosa pine and mixed-conifer forests are common at higher elevations (Wilken et al., 2011). The Southern California Mountains ecoregion contains seven burn areas (Figure 1). This ecoregion contains the high-elevation Transverse Ranges, which serve as a buffer between a coastal Mediterranean climate to the west and a dry, desert climate to the east. Chaparral and oak woodlands are the predominant vegetation communities in this region, although coniferous forests are found at higher elevations (Griffith et al., 2016). The Southern Rockies ecoregion, which includes most of western Colorado, as well as parts of southern Wyoming and northern New Mexico, contains an additional five burn areas (Figure 1). It consists primarily of steep, high-elevation mountain ranges, with some intermontane valleys, and the dominant vegetation communities vary based on a steep elevation gradient. Grasslands and shrublands are common at lower elevations, ponderosa pine, aspen, juniper, and oak forests at middle elevations, mixed-conifer forests at higher elevations, and alpine vegetation at the highest elevations (Wilken et al., 2011; Drummond, 2012). The remaining 12 burn areas are spread across an additional eight ecoregions that contain between one to three burn areas each (Figure 1). These ecoregions range from the high, rugged mountains and dense coniferous forests of the North Cascades ecoregion to the low, broad basins and microphyllous scrubland of the Sonoran Basin and Range ecoregion (Wilken et al., 2011).

2.2 Rainfall, topography, and fire severity data

In addition to debris-flow volume data, we also collected data related to rainfall, terrain, and fire characteristics to calculate 36 potential predictor variables for use in the development of the new volume model, as described in more detail in Section 3.1. We collected rainfall data for every debris-flow-producing storm using a series of rain gages located near watersheds with volume measurements. The rain gages we used were installed and maintained by local, state, and federal government agencies including, but not limited to, the Los Angeles County Department of Public Works (Gartner et al., 2014), Arizona Department of Water Resources (Gorr et al., 2024a), U.S. Forest Service (Gorr et al., 2024a), and the USGS (Gartner et al., 2014), as well as universities (Smith et al., 2021), and private consulting firms (Gorr et al., 2024a). To ensure that the recorded rainfall was representative of the debris-flow-producing storms, we used rain gages located within 4 km of watersheds with debris-flow volume measurements, as suggested by Staley et al. (2017). Most rain gages, however, were located within 2 km of the debris-flow-producing watersheds. We could attribute most debris-flow volumes to a single storm, but when there were multiple storms prior to a volume measurement, we followed the methods of Gartner et al. (2014) and attributed the volume to the most intense storm that occurred between the assumed debris-flow initiation date and the volume measurement. We defined individual storms as events that were separated by at least eight hours without rainfall (Staley et al., 2020).

We used national datasets to calculate metrics related to terrain and fire characteristics for each debris-flow-producing watershed included in our database. Specifically, we resampled the 1/3 arc-second seamless DEM dataset from the USGS 3D Elevation Program (3DEP) to create a series of 10-m resolution DEMs that we used to delineate watershed boundaries and calculate terrain metrics for each watershed. We manually defined the outlet of each watershed as the

point immediately upstream from the debris-flow deposit used to calculate volume, ensuring that all terrain metrics
 210 only considered the watershed area that contributed to debris-flow volume. This also ensured consistency among
 metrics related to fire characteristics, which we calculated using data from the Monitoring Trends in Burn Severity
 (MTBS) program (Monitoring Trends in Burn Severity, 2025). MTBS provides information for all fires 1,000 acres
 and larger in the western United States that burned from 1984 to present, including ignition date, fire severity, and
 215 differenced Normalized Burn Ratio (dNBR) data (Monitoring Trends in Burn Severity, 2025). The differenced
 Normalized Burn Ratio is a remote-sensing index that measures fire-induced changes in vegetation by comparing pre-
 and post-fire satellite imagery and is commonly used to classify burn severity (Parsons et al., 2010).

3 Methods

3.1 Calculation of predictor variables

We calculated 36 potential predictor variables for use in model development: six rainfall variables, 13 terrain variables,
 220 and 17 variables related to fire characteristics. We analyzed six variables related to peak rainfall intensity and rainfall
 ratios (Table 2) because previous work indicates that hourly or sub-hourly rainfall intensity data can be used to more
 accurately constrain postfire debris-flow volume (e.g., Pak and Lee, 2008; Gartner et al., 2014; Gorr et al., 2024a).
 We define a rainfall ratio as a recorded rainfall metric normalized by that same metric associated with a 1-year
 recurrence interval storm at a given location (Cavagnaro et al., 2025a). For example, we define the rainfall ratio of the
 225 peak rainfall intensity measured over a 15-minute duration (i_{15}) as the recorded i_{15} normalized by the i_{15} associated
 with a 1-year recurrence interval storm at a given watershed. Similarly, we define the rainfall ratios of the peak rainfall
 intensity measured over 30-minute (i_{30}) and 60-minute (i_{60}) durations as the recorded i_{30} or i_{60} normalized by the
 i_{30} or i_{60} associated with a 1-year recurrence interval storm at a given watershed.

Table 2: Summary statistics for rainfall predictor variables, as well as the transformation of each variable (e.g., no
 transformation (None) or natural log (Ln)) that yielded the most linear relationship with debris-flow volume, which
 we determined using the Pearson product-moment correlation coefficient (ρ).

Predictor Variable	Min.	Max.	Mean	Median	Transform	ρ
Peak 15-minute rainfall intensity (i_{15}) (mm/h)	5	124	49	37	None	-0.11
Peak 30-minute rainfall intensity (i_{30}) (mm/h)	4	82	34	26	None	-0.13
Peak 60-minute rainfall intensity (i_{60}) (mm/h)	2	51	22	21	None	-0.13
i_{15} Rainfall ratio	0.16	2.96	1.36	1.21	Ln	0.04
i_{30} Rainfall ratio	0.14	3.19	1.39	1.32	None	-0.03
i_{60} Rainfall ratio	0.13	3.00	1.37	1.27	None	-0.09

230 We selected a 1-year recurrence interval to calculate rainfall ratio, as postfire debris flows in the western United States
 are generated by storms with a 1-year recurrence interval, on average (Staley et al., 2020). When available, we used
 National Oceanic and Atmospheric Administration (NOAA) Atlas 14 precipitation frequency data (Bonnin et al., 2006;
 Perica et al., 2013, 2014) to determine the i_{15} , i_{30} , and i_{60} of 1-year recurrence interval storms at each debris-flow-

235 producing watershed. For one burn area where Atlas 14 data were not available (Cub Creek 2), we used NOAA Atlas
 2 data (Miller et al., 1973) to estimate the i_{15} , i_{30} , and i_{60} of a 1-year recurrence interval storm. When known, we
 determined the 1-year recurrence interval values at the location of the rain gage where rainfall was recorded. If we did
 not know the location of the rain gage (i.e., the coordinates of the gage were not provided by a previous study), we
 used the 1-year recurrence interval rainfall associated with the centroid of the associated watershed. We considered
 240 rainfall ratio metrics because previous studies have identified a relationship between postfire debris-flow likelihood
 and rainfall ratio (referred to as “rainfall anomaly” in Cavagnaro et al., 2025a), suggesting that there may also be a
 relationship with postfire debris-flow volume.

Table 3: Summary statistics for watershed terrain predictor variables, as well as the transformation of each
 variable (e.g., natural log (Ln) or square root ($\sqrt{}$)) that yielded the most linear relationship with debris-flow
 volume, which we determined using the Pearson product-moment correlation coefficient (ρ).

Predictor Variable	Min.	Max.	Mean	Median	Transform	ρ
Watershed area (km ²)	0.01	28.0	1.55	0.41	Ln	0.79
Relief (m)	88	2,031	538	481	$\sqrt{}$	0.72
Mean elevation (m) ^a	253	3,000	1,626	1,377	Ln	0.05
Mean slope (°) ^a	11.1	50.6	30.1	29.7	Ln	0.05
Area with slopes $\geq 23^\circ$ (km ²)	0.00	24.7	1.02	0.29	$\sqrt{}$	0.74
Mean slope (%) ^a	19.7	155.5	60.5	60.4	Ln	0.03
Area with slopes $\geq 30\%$ (km ²)	0.00	26.1	1.21	0.34	$\sqrt{}$	0.74
Area with slopes $\geq 50\%$ (km ²)	0.00	23.4	0.89	0.26	$\sqrt{}$	0.74
Maximum flow path (m)	217	10,669	1,945	1,456	Ln	0.74
Total channel length (m)	22	189,451	10,160	2,400	$\sqrt{}$	0.69
Drainage density (km ⁻¹) ^a	1.37	12.6	6.32	6.20	$\sqrt{}$	0.10
Relief ratio	0.11	1.76	0.36	0.32	Ln	-0.50
Ruggedness	0.20	2.86	0.80	0.66	Ln	-0.58

^aVariable removed because it was not linearly related to volume

245 We also calculated 13 terrain variables that previous studies found were correlated with postfire debris-flow volume
 (Gartner et al., 2014; Wall et al., 2023; Gorr et al., 2024a) for all 195 debris-flow-producing watersheds using ArcGIS
 Pro 3.3.0 (Table 3). Here we define relief (Table 3) as the difference between the maximum elevation and minimum
 elevation within a watershed, maximum flow path as the longest flow path within a watershed, as measured from the
 watershed outlet to the top of the drainage divide, and total channel length as the combined length of all channels
 within a watershed. Drainage density is defined as the total channel length divided by the watershed area, relief ratio
 as the length of the maximum flow path divided by watershed relief, and ruggedness, also known as the Melton ratio,
 250 as watershed relief divided by the square root of watershed area. We used 10-m DEMs to calculate these variables
 because this was the highest resolution data available for every watershed. Ensuring consistency across sites was

necessary, as several of the terrain variables that we calculated (e.g., slope) were dependent on DEM resolution (Smith et al., 2019).

Table 4: Summary statistics for fire predictor variables, as well as the transformation of each variable (e.g., natural log (Ln) or square root ($\sqrt{\quad}$)) that yielded the most linear relationship with debris-flow volume, which we determined using the Pearson product-moment correlation coefficient (ρ).

Predictor Variable	Min.	Max.	Mean	Median	Transform	ρ
Time since fire (yrs) ^a	0.04	3.17	0.48	0.25	Ln	-0.18
Area burned (km ²)	0.00	23.3	1.36	0.38	$\sqrt{\quad}$	0.72
Area burned at low severity (km ²)	0.00	4.82	0.33	0.08	$\sqrt{\quad}$	0.61
Area burned at moderate severity (km ²)	0.00	17.7	0.82	0.19	$\sqrt{\quad}$	0.72
Area burned at high severity (km ²)	0.00	2.48	0.21	0.02	$\sqrt{\quad}$	0.46
Area burned at mod/high severity (km ²)	0.00	18.5	1.03	0.26	$\sqrt{\quad}$	0.72
Area burned with slopes $\geq 23^\circ$ (km ²)	0.00	20.3	0.92	0.28	$\sqrt{\quad}$	0.73
Area burned mod/high with slopes $\geq 23^\circ$ (km ²)	0.00	15.8	0.71	0.17	$\sqrt{\quad}$	0.74
Area burned with slopes $\geq 30\%$ (km ²)	0.00	21.6	1.09	0.31	$\sqrt{\quad}$	0.74
Area burned mod/high with slopes $\geq 30\%$ (km ²)	0.00	17.0	0.84	0.21	$\sqrt{\quad}$	0.73
Area burned with slopes $\geq 50\%$ (km ²)	0.00	19.1	0.79	0.24	$\sqrt{\quad}$	0.73
Area burned mod/high with slopes $\geq 50\%$ (km ²)	0.00	14.8	0.61	0.14	$\sqrt{\quad}$	0.74
2 x area burned mod/high + 1 x area low (km ²)	0.00	41.8	2.39	0.68	$\sqrt{\quad}$	0.72
4 x area burned mod/high + 1 x area low (km ²)	0.00	78.9	4.44	1.18	$\sqrt{\quad}$	0.72
Fraction of watershed burned ^a	0.03	1.00	0.90	0.98	Ln	0.11
Fraction of watershed burned mod/high ^a	0.00	1.00	0.68	0.74	$\sqrt{\quad}$	0.21
Mean differenced Normalized Burn Ratio ^a	6	842	353	340	Ln	0.17

^aVariable removed because it was not linearly related to volume

255 We calculated another 17 variables for each watershed related to fire characteristics (Table 4) that previous studies
 have found to be correlated with postfire debris-flow volume using data from MTBS (Gartner et al., 2014; Wall et al.,
 2023; Gorr et al., 2024a). We define time since fire (Table 4) as the time between the date of fire ignition and the date
 of debris-flow initiation. Mean dNBR is the only fire variable we considered that has not been explored by previous
 260 volume studies (e.g., Gartner et al., 2014; Wall et al., 2023; Gorr et al., 2024a). We included it in this analysis because
 it has been identified as an important control on postfire debris-flow likelihood (Staley et al., 2017), and because it
 provides an objective measure of how severely a watershed has been affected by fire.

3.2 Model development

3.2.1 Initial screening of predictor variables

265 After identifying 36 potential predictor variables related to rainfall (Table 2), watershed terrain (Table 3), and fire characteristics (Table 4), we used a multiple linear regression analysis to develop a model for predicting postfire debris-flow volume in the western United States. Multiple linear regression is a statistical technique that uses multiple predictor variables to estimate the value of a response variable (debris-flow volume) following the general form:

$$y = \beta_0 + \beta_1x_1 + \beta_2x_2 + \dots + \beta_kx_k + \varepsilon \quad (1)$$

270 where y is the response variable, β_0 is the intercept, x_i is the i th predictor variable, β_i is the slope coefficient for the i th predictor variable, x_k and β_k are the k th predictor variable and the slope coefficient for the k th predictor variable, respectively, and ε is the error term.

We started the model development process by ensuring that each potential predictor variable was linearly related to debris-flow volume, as a linear relationship between predictor variable and response variable is a requirement of multiple linear regression (Helsel et al., 2020). First, we used the Pearson product-moment correlation coefficient (ρ) to quantify the relationship between the response variable and each predictor variable. We then took the square root and natural log of the response and predictor variables to assess whether transforming one, or both, variables resulted in a more linear relationship between the two. This process resulted in nine correlation coefficients, representing the relations between the response variable and the predictor variable after applying each of three transformations (no transform, square root, and natural log) to both variables. Using this information, we selected the transformations that yielded the highest value of ρ , and thus the most linear relationship between the variables (Tables 2-4). Additionally, because ρ can be heavily influenced by outliers or a curved relationship between response and predictor variables (Helsel et al., 2020), we used scatter plots to visually confirm that debris-flow volume and each predictor variable exhibited a linear relationship. Using these plots, we determined that predictor variables that had a ρ value between -0.3 and 0.3 did not exhibit a convincing linear relationship with debris-flow volume. As a result, we removed these variables from our analysis (Tables 3 and 4).

285 We made an exception to the requirement that each predictor variable be linearly related to debris-flow volume for variables related to rainfall. Although none of the rainfall variables explored here had a correlation coefficient stronger than ± 0.3 (Table 2), we did not remove them from our analysis, as previous studies have found that including a rainfall variable can result in more accurate estimates of postfire debris-flow volume (e.g., Pak and Lee, 2008; Gartner et al., 2014; Gorr et al., 2024a). For instance, Gorr et al. (2024a) found that volume models that contained a rainfall variable considerably outperformed those that did not, despite a weak relationship between the rainfall variables they considered and debris-flow volume. We attribute the weak relationship between rainfall variables and debris-flow volume in this study to uncertainty in the rainfall data. Though we only used data from rain gages within 4 km of debris-flow-producing watersheds, spatial variations in rainfall may have still resulted in substantial differences between what was measured by a rain gage and actual rainfall conditions in the watershed (Figure S1). This situation was likely more common in states like Arizona, Colorado, and New Mexico, where most debris flows initiate as the result of highly localized convective storms (e.g., Cannon et al., 2008; Gorr et al., 2023; McGuire et al., 2024b). However, this uncertainty and lack of linearity is in line with that of rainfall data used in previous volume studies (e.g.,

Gartner et al., 2014; Gorr et al., 2024a), so we did not remove any rainfall variables from our analysis based on the linear relationship requirement. As a result, we were left with 28 potential predictor variables for model development.

3.2.2 Predictor variable selection and model calibration

We selected the predictor variables for the new volume model using a multi-step procedure designed to maximize model performance, minimize the number of predictor variables used, and ensure that the final model met all requirements for multiple linear regression. Because we considered 28 predictor variables, there were 2^{28} potential variable combinations that we could have evaluated. Instead of considering all 2^{28} potential variable combinations, we grouped the variables into three bins (rainfall, terrain, and fire) and only considered models that contained one variable from each bin ($n=702$). Following the methods outlined below, we fit each of the 702 models, selected those that met the requirements of multiple linear regression (i.e., had residuals that were normally distributed and had a constant variance) (Helsel et al., 2020), identified a subset of similarly performing top models ($n=29$), and made final variable selection based on additional multiple linear regression requirements and the relative frequency of occurrence of predictor variables within the top model subset.

We started the variable selection process by separating each of the remaining 28 predictor variables into three bins (six rainfall variables, nine terrain variables, and 13 fire variables) and only considered models that selected one variable from each bin to prevent multicollinearity. Multicollinearity occurs when one predictor variable is closely related to another, and it can result in unrealistically large slope coefficients and illogical relationships between predictor and response variables (Eq. 1), negatively impacting model performance (Alin, 2010; Helsel et al., 2020). Separating the predictor variables into bins reduced the likelihood of selecting two variables that exhibited multicollinearity (e.g., watershed area with slopes $\geq 23^\circ$ and watershed area with slopes $\geq 30\%$). This process yielded 702 unique combinations of rainfall, terrain, and fire predictor variables. We then fit a multiple linear regression model to each combination using the Statistics and Machine Learning Toolbox in MATLAB R2024b, resulting in 702 unique, three-variable models.

After independently fitting all 702 models, we evaluated each to ensure they met the following requirements of multiple linear regression: that the residuals were normally distributed and that the residuals had a constant variance (Helsel et al., 2020). These requirements ensure valid hypothesis tests and reliable confidence and prediction intervals for the model (Helsel et al., 2020). We used the Anderson-Darling (AD) test (Anderson and Darling, 1954) to assess the normality of model residuals, and the Brown-Forsythe (BF) test (Brown and Forsythe, 1974) to assess the variance of the residuals. The null hypothesis for the AD test is that the residuals follow a normal distribution. Therefore, an AD p-value >0.05 indicates that the null hypothesis cannot be rejected and that the residuals are normally distributed. The null hypothesis for the BF test is that the residuals have a constant variance, so a BF p-value >0.05 means that the null hypothesis cannot be rejected and that there is a constant variance in the residuals. To ensure our final model met these requirements of multiple linear regression, we removed 570 models that did not pass the AD and/or BF tests from consideration, leaving 132 models for further analysis.

After removing the models that did not fit our statistical requirements, we evaluated the performance of the remaining 132 models against the entire volume database using metrics including R^2 and root mean square error (RMSE). Higher R^2 values and lower RMSE values reflected better model performance. We also calculated the percentage of volumes predicted within an order of magnitude by each model, as having a first order estimate of debris-flow magnitude is useful for rapid hazard assessment scenarios. We used these metrics to further reduce the number of models we considered during our final model selection process by removing all models where the R^2 and RMSE values were not within 10% of those of the best-performing model. This resulted in 29 models to consider for final evaluation.

From the remaining 29 models, we selected one final model using several factors in addition to the metrics outlined above. First, we determined how often each rainfall, terrain, and fire variable appeared in the 29 best-performing models, and prioritized models that used more commonly selected variables. Given the similar quantitative performance of the remaining 29 models, we interpreted variables that appeared more frequently as those that were more important for constraining postfire debris-flow volume using our dataset. We also ensured that there was no multicollinearity between the selected predictor variables for each model using the variance inflation factor (VIF) (Marquardt, 1970). We interpreted VIF values over 10 as indicative of a strong relationship between predictor variables (Helsel et al., 2020). We used a p-value of 0.1 to assess whether the predictor variables included in each model were statistically significant and removed any models that contained one or more predictor variables with a p-value > 0.1 from consideration. Finally, we assessed whether the predictor variables included in each model fit our conceptual understanding of postfire debris-flow growth. For example, it is well-established that more intense rainfall tends to produce larger debris-flow volumes (e.g., Gartner et al., 2014; Gorr et al., 2024a), so we did not consider models that exhibited a negative relationship between rainfall intensity and volume. Using these considerations, in addition to the quantitative performance metrics, we selected a final model for predicting postfire debris-flow volume in the western United States, which we refer to hereafter as the western United States (WEST) model.

3.3 Model validation

We ensured the WEST model was not overfit using iterated fivefold cross validation (Kohavi, 1995), a method that has been used to validate previous postfire debris-flow volume models (e.g., Gorr et al., 2024a). We started this process by randomly separating the volume database into five similarly sized groups, four of which we classified as the training dataset and one as the testing dataset. We then fit the model on the training dataset and evaluated its performance against the testing dataset using R^2 and RMSE. We repeated this process four more times so that each group of volumes was used as part of the training dataset four times and as the testing dataset once, resulting in five R^2 and RMSE values that we averaged to determine a mean R^2 and RMSE for that iteration of the fivefold cross validation. Then, we started the entire process over again by randomly splitting the volume database into five new groups. In total, we completed 20 iterations of fivefold cross validation to more robustly evaluate model performance when applied to different subsets of data. This process yielded 100 distinct groups of volume data that were used for both training and testing, as well as 20 averaged R^2 and RMSE values. We once again averaged the mean R^2 and RMSE values to determine a single cross-validated (CV) R^2 and RMSE, which we used to evaluate how well the model performed against unseen data. We also assessed the distribution of the R^2 and RMSE values associated with all 100 folds to determine how

370 generalizable the model was to different subsets of volume data. We interpreted CV R^2 and RMSE values similar to the R^2 and RMSE of the WEST model trained on the entire dataset as an indication that the model was not overfit, and a narrow range of R^2 and RMSE values as an indication that the model was not overly sensitive to volumes from specific geographic regions.

3.4 Comparison with existing models

375 We compared the performance of the WEST model against the performance of three existing postfire debris-flow volume models: the Emergency Assessment volume (EAV) model (Gartner et al., 2014), the Intermountain West (IMW) volume model (Wall et al., 2023), and the V1 volume model (Gorr et al., 2024a). Although other methods for predicting postfire debris-flow volume exist (Santi and Morandi, 2013; Pelletier and Orem, 2014; Donovan and Santi, 2017), we selected the EAV, IMW, and V1 models, in particular, for comparison because they were developed for the purpose of postfire hazard assessment using at least in part, subsets of volume data from the larger volume database used in this study (Gartner et al., 2014; Wall et al., 2023; Gorr et al., 2024a). However, unlike the WEST model, which was developed using data from across the western United States (Figure 1), these models were developed using data from more specific geographic regions, including southern California (Gartner et al., 2014), the Intermountain West, defined as the states of Arizona, Colorado, Idaho, Montana, Nevada, New Mexico, Utah, and Wyoming (Wall et al., 2023), and the Southwest, defined as the states of Arizona and New Mexico (Gorr et al., 2024a). Note that the Southwest is a smaller region within the larger Intermountain West, and both regions include the states of Arizona and New Mexico. The IMW model was developed for a broad region that includes Arizona and New Mexico (Wall et al., 2023), whereas the V1 model was developed for use in Arizona and New Mexico, specifically (Gorr et al., 2024a). We compared the WEST model to these three regional models to further evaluate model performance against existing methods for constraining postfire debris-flow volume.

390 We evaluated and compared the performance of each of the models when applied to the entire western United States database. Additionally, to more fairly compare the existing models to the WEST model, we also evaluated the performance of each model when applied to subsets of data from the regions for which the existing models were developed: southern California (Gartner et al., 2014), the Intermountain West (Wall et al., 2023), and the Southwest (Gorr et al., 2024a). The southern California dataset consisted of 93 debris-flow volumes from the Transverse Ranges (Table 1), the Intermountain West dataset of 118 volumes from the states of Arizona, Colorado, New Mexico, and Utah, and the Southwest dataset of 68 volumes from Arizona and New Mexico (Table 1). By assessing the performance of each model against these subsets of volume data, we were able to evaluate how the WEST model performed against regional models when applied to the regions in which those models were developed. We also evaluated how each of the models performed against volumes from data-limited regions, which we define as regions where there is currently not enough volume data to develop a regional volume model. The data-limited dataset included 19 total volumes: 14 from northern California, 3 from Utah, and 2 from Washington (Table S2).

We used multiple metrics to evaluate the performance of each model across the five subsets of volume data. We visually assessed the goodness of fit of each model by plotting the probability density function of model residuals and quantified it by calculating the mean (μ) and standard deviation (σ) of the residuals. Residual mean values closer to

405 zero and smaller σ values indicate better model performance. We also calculated the median absolute error (MAE) and the percentage of volumes predicted within an order of magnitude to further assess model performance. Because all four volume models were developed in natural logarithmic space (Gartner et al., 2014; Wall et al., 2023; Gorr et al., 2024a), we present μ and σ values in their natural log transformed form. However, we present the MAE and percentage of volumes predicted within an order of magnitude in dimensional space for better interpretability.

410 4 Results

4.1 WEST model

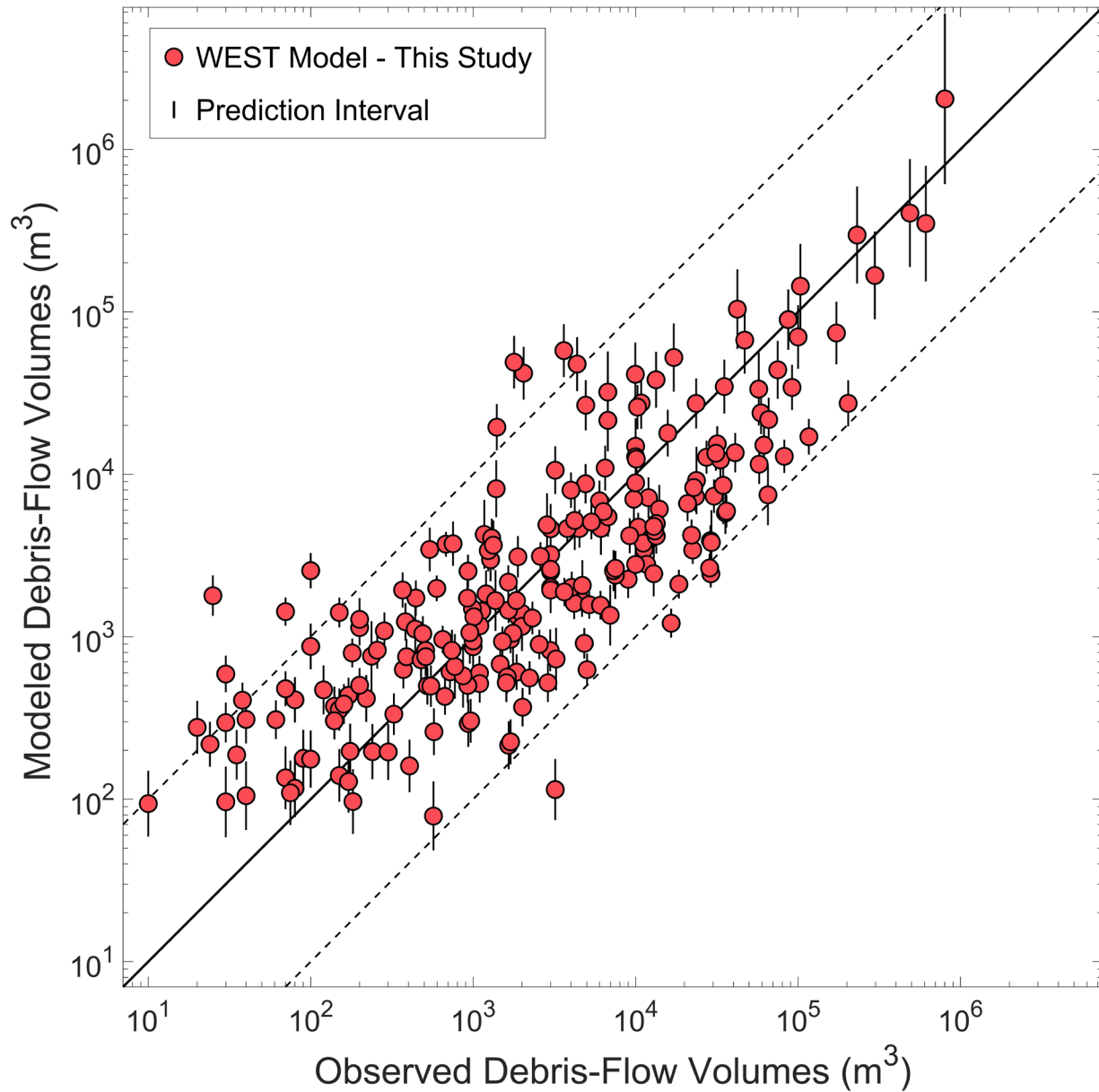
Using the methods outlined in Section 3.2, we selected one model for predicting postfire debris-flow volume in the western United States. The WEST model predicts the volume of sediment deposited by postfire debris flows using the equation:

$$415 \quad \ln V = 7.56 + 0.20i_{30rr} + 0.75 \ln a + 1.11\sqrt{mh_{50}} \quad (2)$$

where V is debris-flow volume (m^3), i_{30rr} is the $i30$ rainfall ratio, a is watershed area (km^2), and mh_{50} is watershed area burned at moderate or high severity with slopes $\geq 50\%$ ($\sim 27^\circ$) (km^2). The WEST model had an $R^2 = 0.66$ and a $\text{RMSE} = 1.31$, both of which were the best among the 29 final models discussed in Section 3.2.2 (Tables S3 and S4). The variance inflation factor (VIF) for each predictor variable in the WEST model was less than three, indicating that there was no multicollinearity, and the p-value for all three predictor variables was < 0.1 , indicating that each was statistically significant. Note that although we selected the WEST model as the final model for this study using the criteria outlined above, many of the final 29 models offer similar performance metrics (Table S3) and may be viable alternatives to the WEST model in some scenarios.

420 The WEST model overpredicted 48% of volumes in the database and underpredicted the remaining 52% (Figure 2). It predicted 41% of volumes within $1,000 \text{ m}^3$, 75% within $10,000 \text{ m}^3$, and 98% within $100,000 \text{ m}^3$ of what was observed. It also predicted 93% of volumes within an order of magnitude (Figure 2). Additionally, the relations among debris-flow volume and each of the predictor variables selected for inclusion in the WEST model agreed with our conceptual understanding of postfire debris-flow growth, as more intense rainfall, larger watersheds, steeper slopes, and higher burn severity yielded greater sediment volumes (Eq. 2).

430 Results of the cross-validation (CV) evaluation indicated that the WEST model was not overfit and was not overly sensitive to volumes from any particular geographic region (Figure S2). The CV R^2 and RMSE values were 0.63 and 1.32, respectively, closely matching the R^2 (0.66) and RMSE (1.31) values of the WEST model trained on the entire volume database. This demonstrated that the model generalized well to unseen data. Additionally, the distributions of the 100 fold-level R^2 and RMSE were relatively narrow (Figure S2), with standard deviations of 0.08 and 0.13, respectively, indicating that, although there was some fold-to-fold variability, most splits produced broadly similar performance, regardless of the geographic distributions of the volumes. Furthermore, the 20 mean R^2 and RMSE values (one associated with each iteration of fivefold cross validation) varied only slightly (Figure S2), providing additional evidence that model performance was stable across random splits of the volume database.



440 **Figure 2: A comparison between the observed volume of all 227 postfire debris flows and the corresponding volume predicted by the Western United States (WEST) model. Vertical lines represent the 95% prediction interval associated with each point. The thick, black line is a 1:1 line, and the thin, dashed lines represent an order of magnitude envelope.**

4.2 Comparison with existing models

445 The WEST model outperformed the EAV (Gartner et al., 2014), IMW (Wall et al., 2023), and V1 (Gorr et al., 2024a) models when applied to the entire western United States volume database. Probability density functions of model residuals revealed that the WEST model provided the best fit between observed and modeled postfire debris-flow volumes in the western United States (Figure 3). The WEST model had a residual mean (μ) nearly equal to zero (Table

5), indicating that it did not systemically overpredict or underpredict debris-flow volumes in the western United States (Figure 3a). In contrast, the EAV model (Figure 3b) had a residual mean greater than zero, and the IMW (Figure 3c) and V1 (Figure 3d) models had residual means less than zero (Table 5), revealing that they tended to overestimate and underestimate postfire debris-flow volumes in this dataset, respectively. The WEST model also had the lowest standard deviation (σ) of all four models (Table 6), indicating the variability of the residuals was lower compared to the other three models. Finally, the WEST model had the lowest MAE (Table 7) and predicted the greatest percentage of volumes within an order of magnitude (Table 8), further indicating that it provided the best fit between modeled and observed volumes in the western United States.

Table 5: Residual means for the western United States (WEST), Emergency Assessment volume (EAV), Intermountain West (IMW), and V1 models (subset by region)

Model	Residual Mean (μ)				
	Western United States	Southern California	Intermountain West	Southwest	Data-Limited Regions
WEST	-0.003	-0.70	0.42	0.41	0.85
EAV	1.32	-0.01	2.31	2.52	2.25
IMW	-2.91	-3.83	-2.61	-2.96	-0.05
V1	-2.02	-3.26	-1.12	-0.50	-1.41

The WEST model also performed well relative to existing models, when evaluated against subsets of data from regions where the other volume models were developed, including southern California (Figure S3), the Intermountain West (Figure S4), and the Southwest (Figure S5). In southern California, the WEST model was the second best-performing model, just behind the EAV model, which was developed for use in this region. Although the EAV model had a lower MAE (Table 7) and predicted a higher percentage of volumes within an order of magnitude (Table 8), the difference in performance between the EAV and WEST models was marginal, especially when compared to the IMW and V1 models (Figure S3). The IMW and V1 models both had MAE values nearly double that of the EAV and WEST models (Table 7) and predicted less than 20% of southern California volumes within an order of magnitude (Table 8). Both models also tended to substantially underpredict debris-flow volumes in southern California, whereas the WEST model only slightly underpredicted volumes in this region, on average (Table 5). The EAV model neither systemically overpredicted nor underpredicted volumes in southern California, as evidenced by a residual mean value of -0.01 (Table 5).

470

Western United States

Predicted / Observed Volume

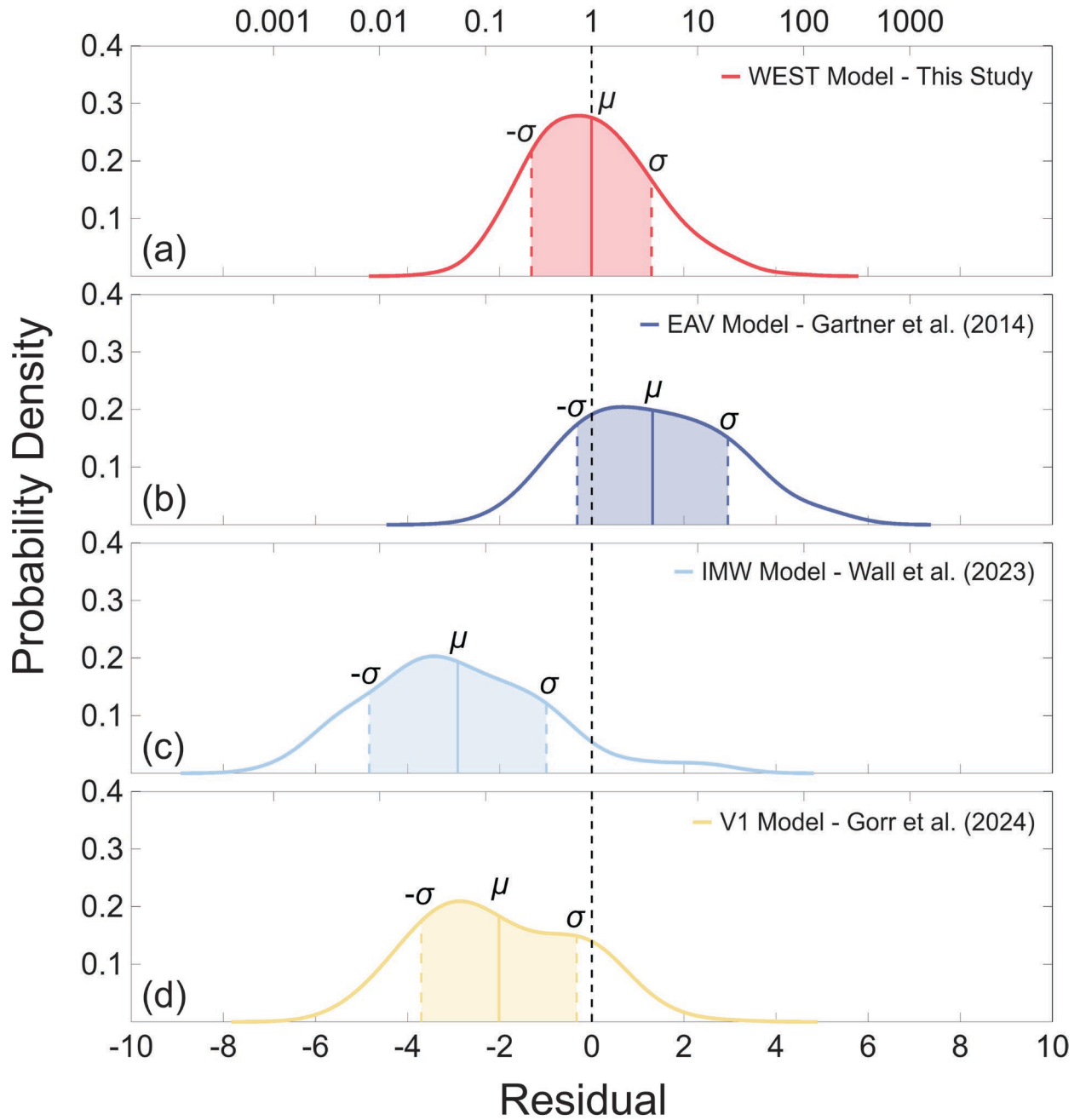


Figure 3: Probability density functions for the residuals of the (a) western United States (WEST), (b) Emergency Assessment volume (EAV), (c) Intermountain West (IMW), and (d) V1 models when applied to the entire western United States dataset.

Table 6: Standard deviation of the residuals for the western United States (WEST), Emergency Assessment volume (EAV), Intermountain West (IMW), and V1 models (subset by region)

Model	Standard Deviation (σ)				
	Western United States	Southern California	Intermountain West	Southwest	Data-Limited Regions
WEST	1.30	1.07	1.20	1.22	1.32
EAV	1.63	0.95	1.31	1.24	1.21
IMW	1.92	1.52	1.62	1.63	2.04
V1	1.69	0.96	1.54	1.25	1.44

475

Table 7: Median absolute errors for the western United States (WEST), Emergency Assessment volume (EAV), Intermountain West (IMW), and V1 models (subset by region)

Model	Median Absolute Error (m^3)				
	Western United States	Southern California	Intermountain West	Southwest	Data-Limited Regions
WEST	1,581	5,037	794	753	269
EAV	7,088	4,407	10,070	10,717	2,495
IMW	2,256	9,627	1,165	1,097	411
V1	1,738	9,160	664	594	245

Table 8: Percentage of volumes predicted within an order of magnitude by the western United States (WEST), Emergency Assessment volume (EAV), Intermountain West (IMW), and V1 models (subset by region)

Model	Percentage of Volumes Predicted within an Order of Magnitude				
	Western United States	Southern California	Intermountain West	Southwest	Data-Limited Regions
WEST	93%	96%	93%	93%	84%
EAV	67%	98%	44%	34%	63%
IMW	35%	18%	45%	38%	58%
V1	52%	17%	76%	90%	74%

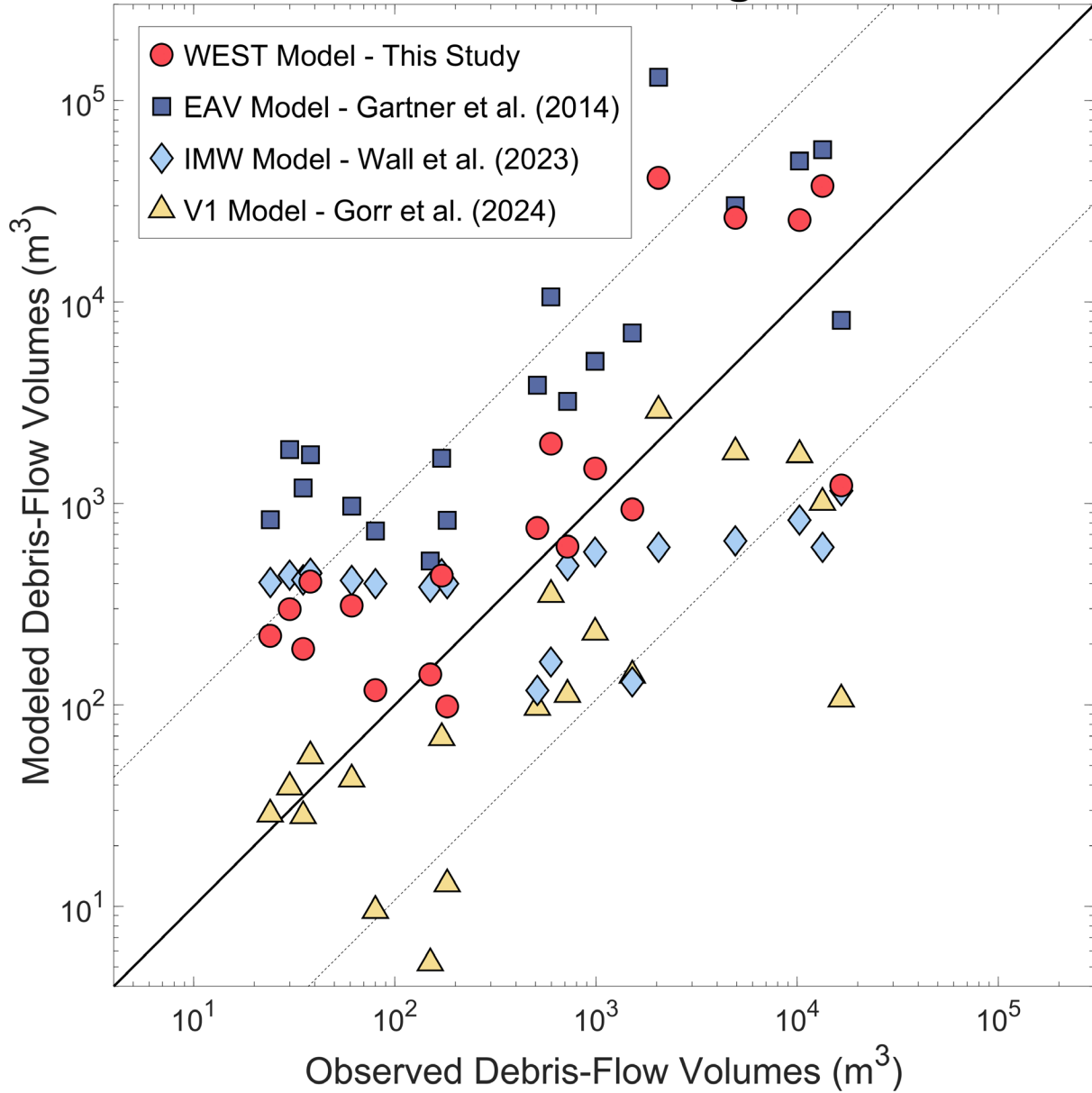
In some scenarios, the WEST model even outperformed existing models in the regions for which they were developed, including the Intermountain West (Figure S4) and the Southwest (Figure S5). In the Intermountain West, the WEST model outperformed the IMW model (Figure S4). It had a lower MAE (Table 7) than the IMW model and predicted a

480

greater percentage of volumes within an order of magnitude (Table 8) in this region. Furthermore, the residual mean of the WEST model was closer to zero (Table 5) and its standard deviation was smaller than that of the IMW model (Table 6), which systemically underpredicted volumes in the Intermountain West (Figure S4). The WEST model also outperformed both the EAV and V1 models in the Intermountain West, as the EAV model greatly overpredicted debris-flow volume, on average, and the V1 model underpredicted debris-flow volume, on average (Figure S4). In the Southwest, the WEST model outperformed the V1 model, according to most metrics (Figure S5). Although the V1 model had a slightly lower MAE (Table 7), the WEST model predicted a greater percentage of volumes in the Southwest within an order of magnitude (Table 8), had a smaller standard deviation (Table 6), and had a residual mean closer to zero (Table 5). The WEST model tended to slightly overpredict debris-flow volumes in the Southwest (Figure S5a), whereas the V1 model tended to slightly underpredict volumes in this region (Figure S5d). The EAV and IMW models, on the other hand, more severely overpredicted and underpredicted volumes in the Southwest, respectively (Figure S5). Differences between the volumes predicted by these models and observed volumes in the Southwest routinely exceeded an order of magnitude (Table 8).

When applied to 19 postfire debris-flow volumes from data-limited regions, including northern California, Utah, and Washington, (Table S2), the WEST model again outperformed the EAV, IMW, and V1 models. The WEST model predicted the greatest percentage of volumes within an order of magnitude (Figure 4; Table 8) and had one of the lowest MAE values (Table 7). Although the IMW model had the residual mean closest to zero (Table 5), it also had the largest standard deviation (Table 6), indicating high variability in the residuals compared to other models (Figure 5). The WEST model slightly overpredicted volumes from data-limited regions but had the second lowest standard deviation of the four models (Table 6). The EAV model had a slightly lower standard deviation than the WEST model but overpredicted volumes from data-limited regions more substantially (Figure 5b). Unlike the WEST and EAV models, the V1 model underpredicted volumes from data-limited regions, on average (Figure 5d).

Data-Limited Regions



505 **Figure 4: A comparison between the observed volume of 19 postfire debris flows from data-limited regions and the corresponding volume predicted by the western United States (WEST), Emergency Assessment volume (EAV), Intermountain West (IMW), and V1 models. The thick, black line is a 1:1 line, and the thin, dashed lines represent an order of magnitude envelope.**

510

Data-Limited Regions

Predicted / Observed Volume

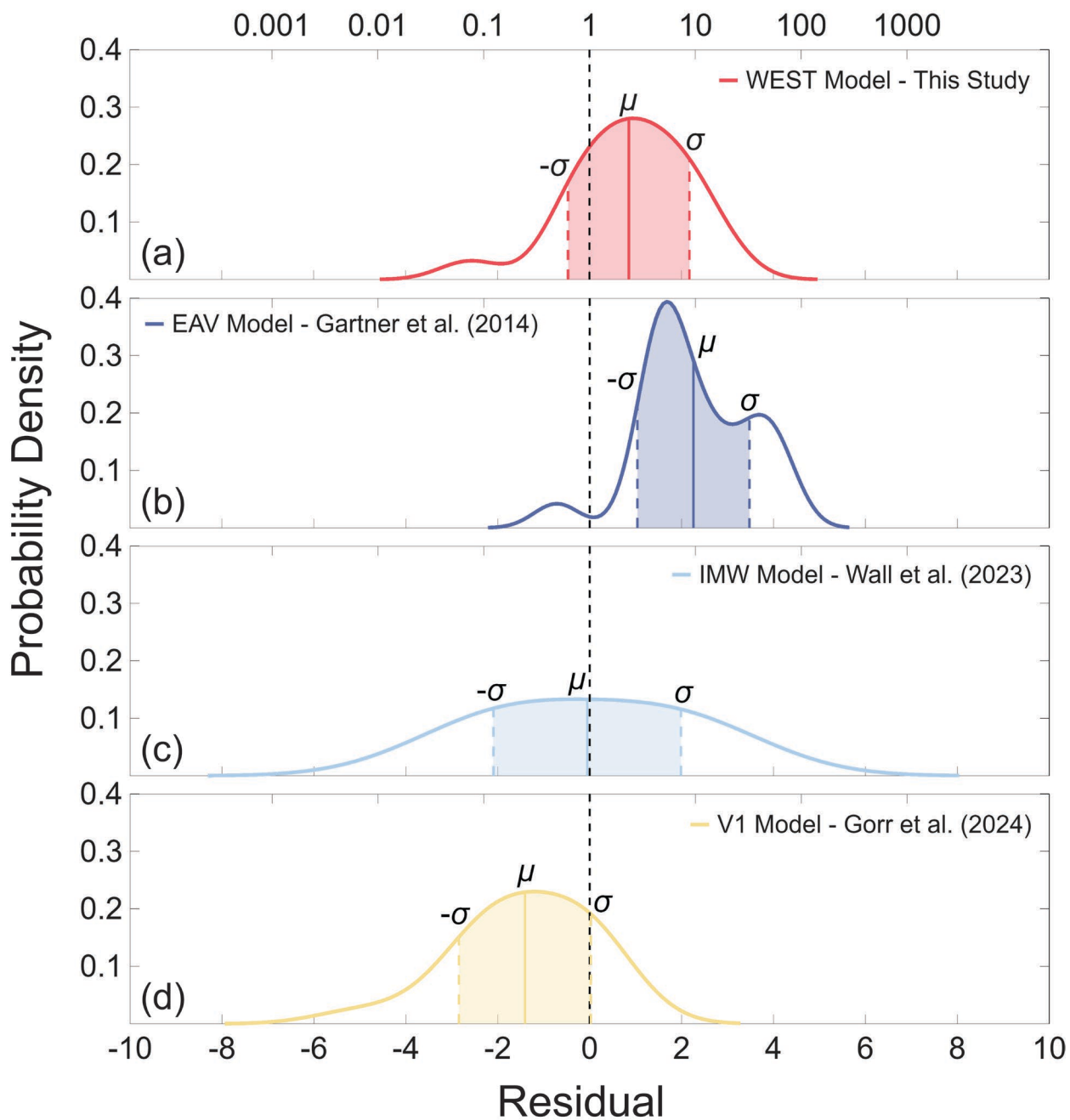


Figure 5: Probability density functions for the residuals of the (a) western United States (WEST), (b) Emergency Assessment volume (EAV), (c) Intermountain West (IMW), and (d) V1 models when applied to volumes from data-limited regions.

We introduced a new empirical model for predicting postfire debris-flow volume in the western United States. This model, referred to as the WEST model, predicts the volume of sediment deposited by postfire debris flows as a function of i_{30} rainfall ratio, watershed area, and watershed area burned at moderate or high severity with slopes greater than or equal to 50% (Eq. 2). It offers an improvement over existing volume models because it accounts for regional differences in rainfall characteristics with a rainfall ratio metric and because it was trained on a larger dataset of debris-flow volumes from across the western United States. Specifically, the WEST model outperforms three existing debris-flow volume models (Gartner et al., 2014; Wall et al., 2023; Gorr et al., 2024a) when applied to the entire western United States database (Figure 3), as well as to subsets of data from the Intermountain West (Figure S4), the Southwest (Figure S5), and data-limited regions (Figures 4 and 5). It also maintains a similar level of performance to that of a model developed for use in southern California when applied in that region (Figure S3). These results indicate that the WEST model is more broadly applicable than existing volume models, particularly in data-limited regions, and that it may be a promising tool for postfire hazard assessment in the western United States.

5.1 Improvements over existing models

5.1.1 Rainfall ratio

The WEST model offers an improvement over existing volume models because it accounts for regional differences in rainfall intensity and because it was trained on the largest dataset of debris-flow volumes. The WEST model uses a rainfall ratio metric that normalizes for regional variations in rainfall and is consequently able to achieve equivalent performance to several models developed for different regions of the western United States using a single regression equation. Prior regional volume models, on the other hand, use rainfall intensity (e.g., Gartner et al., 2008, 2014; Gorr et al., 2024a), and are thus limited when applied to areas outside of their training datasets due to regional differences in the intensity of debris-flow-generating rainfall (e.g., Gorr et al., 2023, 2024a; Rengers et al., 2023, 2024). Previous studies have found that, although hourly (Pak and Lee, 2008) or sub-hourly (e.g., Gartner et al., 2014; Gorr et al., 2024a) rainfall intensity is an important control on postfire debris-flow volume across the western United States, the rainfall intensity needed to generate postfire debris flows varies between regions (Cavagnaro et al., 2025b). For example, the 15-minute rainfall intensity (i_{15}) needed to generate a postfire debris flow is less than 20 mm/h in the Transverse Ranges of southern California (Staley et al., 2013), roughly 30 mm/h in the Front Range of Colorado (Staley et al., 2015), and more than 60 mm/h in northern Arizona (Youberg, 2014).

Regional volume models tend to be biased when applied outside of their training regions such that they overpredict volumes in areas with higher average rainfall intensities than their training region and underpredict volumes in areas with lower average rainfall intensities. For instance, southern California requires some of the least intense rainfall to generate postfire debris flows (Staley et al., 2017), so the EAV model (Gartner et al., 2014), which was developed using data from southern California (Gartner et al., 2014), tends to overpredict debris-flow volume in other regions of the western United States and Canada. Gorr et al. (2024a) found that the EAV model overpredicted postfire debris-flow volumes in the Southwest by roughly 3,500%, on average, and Rengers et al. (2024) found that the model overpredicted observed volumes in Colorado by more than a factor of four. Additionally, Gartner et al., (2024) found that the EAV model overpredicted postfire debris-flow volumes in British Columbia, Canada by a factor of 2 to 4. We

observed similar model behavior in this study, as the EAV model consistently overpredicted postfire debris-flow volume in all regions other than southern California (Figures 3, S4, and S5). Conversely, the Southwest requires some of the most intense rainfall to generate postfire debris flows (Staley et al., 2017), so the V1 model consistently underpredicts postfire debris-flow volumes in other parts of the western United States (Figures 3, S3, and S4). In this study, the V1 model underestimated debris-flow volumes on all four subsets of data, as well as when applied to the entire western United States database (Figure 3). It performed particularly poorly when applied to southern California (Figure S3), the region with the least similar rainfall characteristics to the Southwest. The IMW model also consistently underpredicted postfire debris-flow volume in this study, including in the Intermountain West, the region for which it was developed (Figure S4). However, this model does not include a rainfall variable (Wall et al., 2023), indicating that other regional differences or model limitations are responsible for reduced performance against this volume dataset.

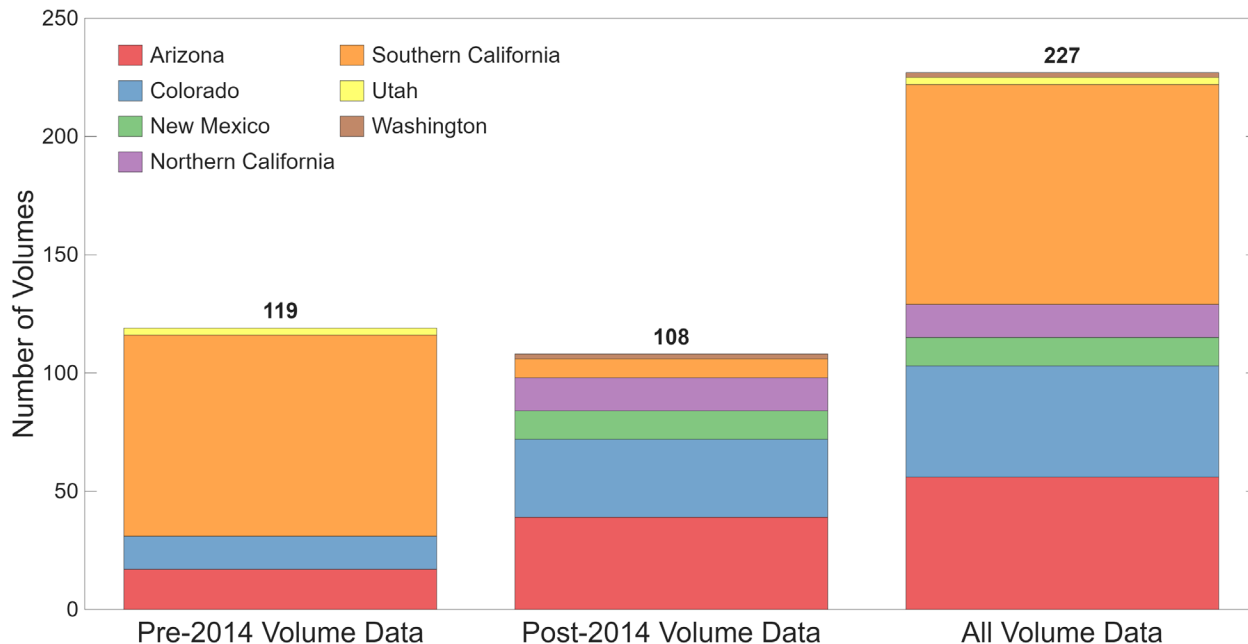
The WEST model, however, does not exhibit large variations in model performance between geographic regions (Figures S3-S5) because it uses *i*30 rainfall ratio instead of rainfall intensity (Eq. 2). This allows the WEST model to incorporate regional differences in rainfall intensity without the limitations associated with regional volume models. Because the *i*30 rainfall ratio normalizes the peak 30-minute rainfall intensity of a debris-flow-producing storm by the peak 30-minute rainfall intensity associated with a 1-year recurrence interval storm at the location of a debris-flow-producing watershed, it is consistent across regions that have different rainfall characteristics. As a result, the WEST model performs similarly when applied to different geographic regions, including southern California (Figure S3), the Intermountain West (Figure S4), and the Southwest (Figure S5).

570 5.1.2 Training dataset

The WEST model also offers an improvement over existing postfire debris-flow volume models, as it was developed using a more comprehensive training dataset. The WEST model was developed using a dataset of 227 postfire debris-flow volumes from 34 burn areas across six states (Figure 1). The three regional models evaluated in this study, on the other hand, were developed using smaller, more geographically limited datasets. Specifically, the EAV model was developed using 92 volumes from southern California (Gartner et al., 2014), the IMW model using 47 volumes from four states in the Intermountain West (Arizona, Colorado, Utah, and Wyoming), although 39 volumes were from Utah alone (Wall et al., 2023), and the V1 model using 54 volumes from Arizona and New Mexico (Gorr et al., 2024a). The geographic diversity of its training dataset contributes to the broader applicability of the WEST model relative to existing regional volume models.

The broad applicability of the WEST model indicates that it can be a substantial improvement over volume models that are currently used for postfire hazard assessments, including the EAV model. The EAV model is currently the most-commonly used method for predicting postfire debris-flow volume in the western United States, as it is used as part of the USGS operational postfire hazard assessment framework (Landslide Hazards Program, 2018). Although the accuracy of the EAV model is limited outside of southern California, as discussed above, a lack of postfire debris-flow volume data, and associated rainfall data, in many parts of the western United States has historically prevented the development of a viable alternative. When the EAV model was published in 2014, nearly all measured postfire debris-flow volumes with associated rainfall data were from southern California (Gartner et al., 2014), with minor

590 exceptions from Arizona (Youberg, 2014), Colorado (Gartner et al., 2008), and Utah (Gartner et al., 2008). Taking the database used in this study (Gorr et al., 2025) as an example, more than 70% of postfire debris-flow volumes measured prior to 2014 came from southern California (Figure 6). The limited geographic scope of volume data available at the time therefore prevented the development of a more broadly applicable postfire debris-flow volume model.



595 **Figure 6: Geographic distributions of the volume data used to develop the western United States (WEST) model, separated by date of volume measurement. Specifically, the distributions of volume data collected prior to 2014, volume data collected after 2014, and all volume data included in the database used to develop the WEST model.**

600 However, efforts to measure postfire debris-flow volumes across the western United States have expanded since the EAV model was published. Of the 227 postfire debris-flow volumes used to develop the WEST model, 108 (48%) have been measured since 2014. Many of the volumes collected during this time are from regions where volume data have been historically sparse, and only 7% of post-2014 volumes used in this study are from southern California (Figure 6). All volumes from New Mexico, northern California, and Washington have been collected since 2014, and data from Arizona and Colorado have expanded by 229% and 236% relative to the pre-2014 dataset, respectively (Figure 6). The improved performance of the WEST model across the western United States compared to existing models can therefore be partially attributed to the fact that it was trained on the largest and most geographically diverse postfire debris-flow volume dataset with associated rainfall data.

605

5.2 Implications for hazard assessment

Because it accounts for regional differences in rainfall characteristics and was trained on a geographically diverse volume dataset, the WEST model can be used for widespread postfire hazard assessments in the western United States. One of the most encouraging signs for improved postfire hazard assessments is the model's performance in data-

610 limited regions. Results show that the WEST model performed similarly in data-limited regions as it did elsewhere in the western United States (Tables 5-8), and that it outperformed the EAV, IMW, and V1 models in these areas (Figures 4 and 5). The WEST model performed particularly well, relative to existing models, against two debris-flow volumes from the Dixie Fire, which burned in the Sierra Nevada of northern California, and two volumes from the Cub Creek 2 Fire, which burned in the Pacific Northwest (Washington) (Figure 1; Table 1).

615 The WEST model overpredicted the four volumes in the Sierra Nevada and Pacific Northwest by an average of 25,007 m³, which represents a substantial improvement over the EAV model, which overpredicted the same four volumes by an average of 59,284 m³. The IMW and V1 models, on the other hand, underpredicted these volumes by an average of 6,981 m³ and 5,792 m³, respectively. Though the IMW and V1 models provided a smaller absolute difference between observed and modeled volumes in these regions, they severely underpredicted observed volumes in the Sierra
620 Nevada and Pacific Northwest. Specifically, the IMW model underpredicted these four volumes by between 70.5% and 95.5%. The V1 model slightly overpredicted one of the four volumes by 41% but underpredicted the remaining three by between 63.5% and 92.4%. Because previous studies have found that debris-flow volume scales with runoff distance and area inundated (Iverson et al., 1998; Rickenmann, 1999; Griswold and Iverson, 2008), underpredicting debris-flow volume limits the effectiveness of postfire hazard assessments by underestimating the downstream effects
625 of postfire debris flows. Based on the results of this study, the WEST model is less likely to underestimate the extent of potential downstream effects from postfire debris flows in data-limited regions, as it offers more conservative predictions of postfire debris-flow volume relative to the IMW and V1 models. At the same time, it provides more accurate volume estimates than the EAV model in these areas, making it less likely to predict unrealistically severe downstream effects.

630 Having a method that outperforms existing models in data-limited regions, such as the Sierra Nevada and Pacific Northwest, can improve postfire hazard assessments in these areas, particularly as fire activity increases. Wildfire activity across the western United States has increased considerably in recent decades, but this change has been particularly pronounced in the Pacific Northwest (Westerling, 2016). According to Westerling (2016), the area burned by wildfire in the Pacific Northwest between 2003 and 2012 increased by nearly 5,000% relative to the area burned
635 in this region between 1973 and 1982. Furthermore, since this study was published in 2016, the Pacific Northwest has experienced multiple historic fire seasons, including the 2020 season, which burned nearly as much forest in the western Cascade Range in two weeks as in the previous five decades combined (Reilly et al., 2022). Fire activity is also increasing in California's Sierra Nevada (Westerling, 2016). Not only are fires in the Sierra Nevada increasing in size and frequency (Westerling, 2016), but the severity of fires is also changing. Miller and Safford (2012) found that
640 proportion of annual wildfire burned at high severity increased considerably in parts of the Sierra Nevada between 1984 and 2010. This has important implications for postfire debris-flow hazards, as the effects of fire on infiltration and erosion, which affects debris-flow initiation and growth, are most pronounced in areas burned at high severity (e.g., Vieira et al., 2015; McGuire and Youberg, 2019). As fire activity in these regions increases, so does the potential for postfire debris-flow hazards (e.g., DeGraff et al., 2011; Neptune et al., 2021; Wall et al., 2020; Selander et al.,

645 2025), underscoring the use of a volume model, like the WEST model, that can provide accurate estimates of debris-flow volume in these areas.

Additional volume data from the Sierra Nevada and Pacific Northwest will further improve predictions of postfire debris-flow volumes in these regions. However this study represents a first step, as it is the first to include data from the Sierra Nevada and Pacific Northwest in the development and evaluation of volume models. In particular, volume
650 data from larger-magnitude debris flows in data-limited regions are needed to more fully evaluate the performance of the WEST model in these settings, as the median volume of the 19 debris flows from data-limited regions included in this study was 511 m³ (Table S2), nearly five times lower than the median volume of 2,550 m³ associated with the entire volume database. Additionally, volume data from the northern Rockies (Idaho, Montana, and Wyoming) are needed to evaluate the WEST model in this region. Although the northern Rockies are susceptible to postfire debris
655 flows (e.g., Meyer and Wells, 1997; Gabet and Bookter, 2008), we are unaware of any postfire debris-flow volumes with associated rainfall data from this region that can be used to improve model performance.

Similarly, more data are needed to evaluate the performance of the WEST model when applied outside of the western United States. Although postfire debris flows are common hazards in fire-prone regions around the world, including Australia (e.g., Nyman et al., 2011), Canada (Hancock and Wlodarczyk, 2025), Italy (e.g., Esposito et al., 2023), and
660 Spain (García-Ruiz et al., 2013), debris-flow volumes and associated rainfall data remain scarce outside of the western United States. This lack of data has limited the development of postfire debris-flow volume models for these regions and has prevented the evaluation of most volume models developed for use in the western United States, including the WEST model, when applied elsewhere. However, previous studies suggest that the primary controls on postfire debris-flow volume may be consistent across geographic regions. For example, Nyman et al. (2015) found that
665 watershed area was the most important control on the volume of 10 postfire debris flows that initiated in southeast Australia and that volumes could be accurately predicted using an empirical volume model developed for use in the western United States (Gartner et al., 2008). Similar to the WEST model, this model predicts postfire debris-flow volume as a function of rainfall characteristics, watershed area and slope, and burn severity (Gartner et al., 2008). These similarities indicate that the primary controls on postfire debris-flow volume are generally transferrable across
670 geographic settings, but additional volume data from fire-prone regions around the world are needed to evaluate the performance of the WEST model beyond the western United States.

Finally, although we selected the WEST model using the criteria outlined in Section 3.2.2, several alternative models that we developed as part of this study offer similar performance metrics (Table S3) and may be preferred to the WEST model in some situations. For instance, we determined that rainfall ratio calculated over a 30-minute duration yielded
675 the best results for our dataset, but there may be scenarios where models that incorporate rainfall ratio calculated over 15 or 60-minute durations are more suitable. It may be more practical, for example, to use a model that includes *i*60 rainfall ratio, instead of *i*30 rainfall ratio, to estimate volumes for a mitigation project if only *i*60 design storms are available. Alternatively, it may be easier to implement a model that uses *i*15 rainfall ratio within a hazard assessment framework that also predicts postfire debris-flow likelihood using rainfall characteristics measured over a 15-minute
680 duration (Landslide Hazards Program, 2018). Given their potential applicability in these scenarios, we present

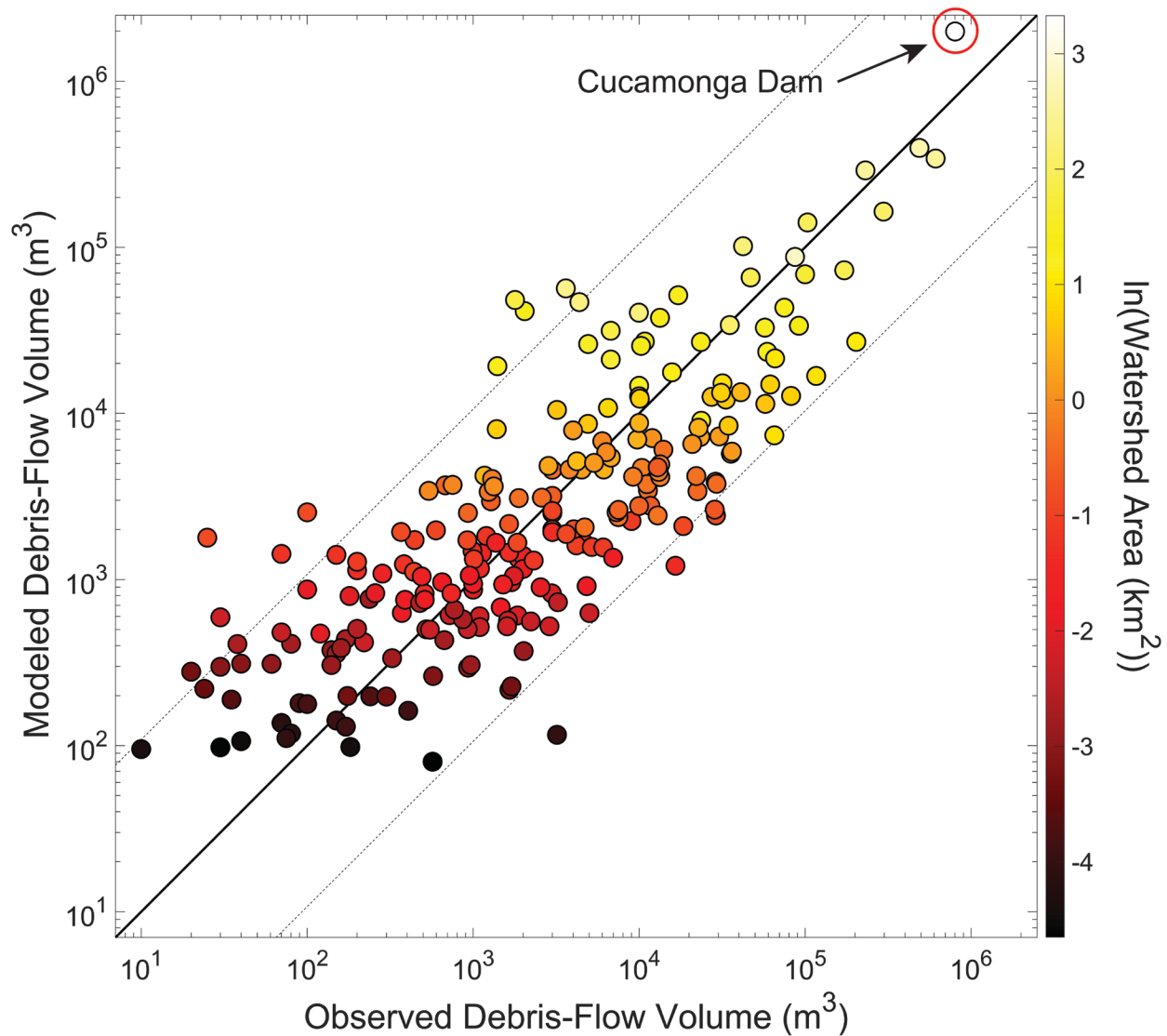
alternative models that use $i15$ and $i60$ rainfall ratio, along with the same terrain and fire variables used by the WEST model (Equation 2), to predict postfire debris-flow volume in Table 9. Note that, although it marginally outperforms the WEST model, the $i15$ model does not pass the Anderson-Darling test (p -value < 0.05), so was not considered for final model selection. The $i60$ model, on the other hand, performs similarly to the WEST model and meets all requirements of multiple linear regression (Table 9).

Table 9: Equations for volume models that use rainfall ratios calculated at different durations, including 15 minutes ($i15_{rr}$), 30 minutes ($i30_{rr}$), and 60 minutes ($i60_{rr}$). All models predict debris-flow volume (V) as a function of rainfall ratio, watershed area (a), and watershed area burned at moderate or high severity with slopes $\geq 50\%$ (mh_{50}). Performance metrics include R^2 and root mean square error (RMSE), as well as the p -values for the Anderson-Darling (AD) and Brown-Forsythe (BF) tests.

ID	Model	R^2	RMSE	AD Test p-value	BF Test p-value
$i15$	$\ln V = 7.82 + 0.35 \ln i15_{rr} + 0.76 \ln a + 1.09\sqrt{mh_{50}}$	0.67	1.30	0.03	0.17
WEST	$\ln V = 7.56 + 0.20i30_{rr} + 0.75 \ln a + 1.11\sqrt{mh_{50}}$	0.66	1.31	0.07	0.17
$i60$	$\ln V = 7.61 + 0.18i60_{rr} + 0.76 \ln a + 1.10\sqrt{mh_{50}}$	0.66	1.31	0.07	0.15

5.3 Model limitations and future work

Although the WEST model improves our ability to predict postfire debris-flow volume in many regions of the western United States, there are some scenarios where the effectiveness of the model may be limited. For example, because the WEST model predicts postfire debris-flow volume as a function of both watershed area and watershed area burned at moderate or high severity with slopes greater than or equal to 50%, it tends to substantially overpredict volumes from large watersheds ($>20 \text{ km}^2$) (Figure 7). We observed this model behavior when applied to the largest watershed in the dataset used in this study: Cucamonga Dam (Gorr et al., 2025). The Cucamonga Dam watershed, which burned in the 2003 Grand Prix Fire in southern California (Figure 1), had an area of 28.0 km^2 and an area burned at moderate or high severity with slopes greater than or equal to 50% of 14.8 km^2 , both the largest in the volume dataset (Tables 3 and 4). It also produced the largest postfire debris flow in the dataset, with a volume of $801,770 \text{ m}^3$ (Gorr et al., 2025). When applied to this watershed, the WEST model predicted a volume of $1,993,100 \text{ m}^3$, an overprediction of more than $1,000,000 \text{ m}^3$ (Figure 7). However, we only expect this behavior in the watersheds larger than 20 km^2 , like the Cucamonga Dam watershed, as the model slightly underpredicted the volumes for the next two largest watersheds in the dataset (Figure 7), which had areas of 17.3 km^2 and 14.4 km^2 . This limitation is unlikely to affect postfire hazard assessments in the western United States, as the USGS operational postfire hazard assessment framework is typically applied to watersheds with areas less than 8.0 km^2 (Landslide Hazards Program, 2018).



705 **Figure 7: Comparison between observed debris-flow volumes and debris-flow volumes predicted by the Western United States (WEST) model when applied to the entire western United States dataset. Points are colored by the natural log of the area of the debris-flow-producing watershed. The thick, black line is a 1:1 line, and the thin, dashed lines represent an order of magnitude envelope.**

Another potential limitation of the WEST model is that, aside from differences in rainfall characteristics, it does not account for other regional factors that may influence postfire debris-flow volume, including sediment availability. In southern California, dry ravel (i.e., the transport of sediment by gravity without rainfall) is a common process in burned landscapes (e.g., Lamb et al., 2011). Following fire, but before rainfall, dry ravel can load channels with up to three meters of unconsolidated sediment that serves as an important sediment source for postfire debris flows (e.g., Wells, 1987; DiBiase and Lamb, 2020; Palucis et al., 2021). Although prevalent in southern California, dry ravel is less common in other regions of the western United States (Perkins et al., 2022). Previous studies have found that

715 rilling on hillslopes and/or channel incision, not dry ravel, are the primary sources of sediment for postfire debris flows
in many areas, including elsewhere in California (e.g., DeGraff et al., 2011), Colorado (Rengers et al., 2024), the
Pacific Northwest (Wall et al., 2020), and the Southwest (e.g., Tillery and Rengers, 2020; Gorr et al., 2024b), among
others. The absence of dry ravel in these regions may limit sediment availability and result in smaller postfire debris-
flow volumes relative to southern California. Regional differences in sediment availability may account for differences
720 in model behavior that we observed in this study. In areas with dry ravel (i.e., southern California), the WEST model
is biased low (Figure S3), and in areas where dry ravel is less prevalent (i.e., the Intermountain West and Southwest),
the WEST model is biased high (Figures S4 and S5). This indicates that future postfire debris-flow volume models
may benefit from the inclusion of a sediment availability variable.

6 Conclusions

725 Debris-flow volume is a critical component of postfire hazard assessments in the western United States. However,
existing methods for predicting postfire debris-flow volume have various shortcomings that limit their applicability
for the purpose of postfire hazard assessment. In this study, we introduced a new model for predicting postfire debris-
flow volume in the western United States. Using a database of 227 postfire debris-flow volumes collected across six
states and 34 burn areas, we developed a multiple linear regression model that predicts postfire debris-flow volume as
730 a function of rainfall, watershed terrain, and fire characteristics. This model offers an improvement over existing
volume models, as it accounts for regional differences in debris-flow-generating rainfall and was trained on a larger,
more geographically diverse volume dataset. Results from this study demonstrate that the new model outperforms
three existing regional volume models when applied to the entire western United States database, and either
outperforms or performs similarly to existing regional volume models when applied to subsets of volume data from
735 the regions for which they were developed. The new model also outperforms existing models when applied to volumes
from data-limited regions where there is not enough data to develop region-specific models. The broad applicability
of the model introduced in this study indicates that it can be used to improve widespread postfire hazard assessments
across the western United States.

Data availability

740 The data used to develop the WEST model are publicly available in Gorr et al. (2025).

Author contributions

ANG, FKR, KRB, MAT, and JWK conceptualized the study. ANG and FKR collected data with assistance from co-
authors. ANG performed data analysis with assistance from all co-authors. ANG prepared the original draft of the
manuscript, and FKR, KRB, MAT, and JWK reviewed and edited the draft.

745 Competing Interests

The authors declare that they have no conflict of interest.

Acknowledgements

We thank Joseph Gartner (BGC Engineering) and Jacob Woodard (U.S. Geological Survey) for their work reviewing this paper. We also thank those who contributed new volume data, beyond what has already been reported in previous studies, including Corey Crowder and Luke McGuire (University of Arizona), Andrew Graber, Olivia Hoch, and Jaime Kostelnik (U.S. Geological Survey), and Don Lindsay and Paul Richardson (California Geological Survey). Any use of trade, firm, or product names is for descriptive purposes only and does not imply endorsement by the U.S. Government.

References

- 755 Alin, A.: Multicollinearity, *WIREs Computational Statistics*, 2, 370–374, <https://doi.org/10.1002/wics.84>, 2010.
- Anderson, T. W. and Darling, D. A.: A Test of Goodness of Fit, *Journal of the American Statistical Association*, 49, 765–769, <https://doi.org/10.1080/01621459.1954.10501232>, 1954.
- Barnhart, K. R., Jones, R. P., George, D. L., McArdeell, B. W., Rengers, F. K., Staley, D. M., and Kean, J. W.: Multi-Model Comparison of Computed Debris Flow Runout for the 9 January 2018 Montecito, California Post-Wildfire Event, *Journal of Geophysical Research: Earth Surface*, 126, e2021JF006245, <https://doi.org/10.1029/2021JF006245>, 2021.
- 760 Berti, M. and Simoni, A.: Prediction of debris flow inundation areas using empirical mobility relationships, *Geomorphology*, 90, 144–161, <https://doi.org/10.1016/j.geomorph.2007.01.014>, 2007.
- Bonnin, G. M., Martin, D., Lin, B., Parzybok, T., Yekta, M., and Riley, D.: Precipitation-Frequency Atlas of the United States. Volume 1, Version 5.0. Semiarid Southwest (Arizona, Southeast California, Nevada, New Mexico, Utah), 2006.
- 765 Brown, M. B. and Forsythe, A. B.: Robust Tests for the Equality of Variances, *Journal of the American Statistical Association*, 69, 364–367, <https://doi.org/10.1080/01621459.1974.10482955>, 1974.
- Cannon, S., Gartner, J., Rupert, M., Michael, J., Rea, A., and Parrett, C.: Predicting the probability and volume of postwildfire debris flows in the intermountain western United States, *Geological Society of America Bulletin*, 122, 127–144, <https://doi.org/10.1130/B26459.1>, 2010.
- 770 Cannon, S. H., Kirkham, R. M., and Parise, M.: Wildfire-related debris-flow initiation processes, Storm King Mountain, Colorado, *Geomorphology*, 39, 171–188, [https://doi.org/10.1016/S0169-555X\(00\)00108-2](https://doi.org/10.1016/S0169-555X(00)00108-2), 2001.
- Cannon, S. H., Gartner, J. E., Wilson, R. C., Bowers, J. C., and Laber, J. L.: Storm rainfall conditions for floods and debris flows from recently burned areas in southwestern Colorado and southern California, *Geomorphology*, 96, 250–269, <https://doi.org/10.1016/j.geomorph.2007.03.019>, 2008.
- 775 Cannon, S. H. and Gartner, J. E.: Wildfire-related debris flow from a hazards perspective, in: *Debris-flow Hazards and Related Phenomena*, edited by: Jakob, M. and Hungr, O., Springer, Berlin, Heidelberg, 363–385, https://doi.org/10.1007/3-540-27129-5_15, 2005.

- 780 Cavagnaro, D. B., McCoy, S. W., Thomas, M. A., Kostelnik, J., and Lindsay, D. N.: Improved Prediction of Postfire Debris Flows Through Rainfall Anomaly Maps, *Geophysical Research Letters*, 52, e2025GL114791, <https://doi.org/10.1029/2025GL114791>, 2025a.
- Cavagnaro, D. B., McCoy, S. W., Lindsay, D. N., McGuire, L. A., Kean, J. W., and Trugman, D. T.: Rainfall Thresholds for Postfire Debris-Flow Initiation Vary With Short-Duration Rainfall Climatology, *Journal of Geophysical Research: Earth Surface*, 130, e2024JF007781, <https://doi.org/10.1029/2024JF007781>, 2025b.
- 785 Daurio, M.: Cascading hazards, cascading consequences: Linking social-ecological systems in post-fire recovery, *Human Organization*, 84, 339–350, <https://doi.org/10.1080/00187259.2025.2499680>, 2025.
- DeGraff, J. V., Wagner, D. L., Gallegos, A. J., DeRose, M., Shannon, C., and Ellsworth, T.: The remarkable occurrence of large rainfall-induced debris flows at two different locations on July 12, 2008, Southern Sierra Nevada, CA, USA, *Landslides*, 8, 343–353, <https://doi.org/10.1007/s10346-010-0245-5>, 2011.
- 790 DiBiase, R. A. and Lamb, M. P.: Dry sediment loading of headwater channels fuels post-wildfire debris flows in bedrock landscapes, *Geology*, 48, 189–193, <https://doi.org/10.1130/G46847.1>, 2020.
- Donovan, I. P. and Santi, P. M.: A probabilistic approach to post-wildfire debris-flow volume modeling, *Landslides*, 14, 1345–1360, <https://doi.org/10.1007/s10346-016-0786-3>, 2017.
- Dowling, C. A. and Santi, P. M.: Debris flows and their toll on human life: a global analysis of debris-flow fatalities from 1950 to 2011, *Nat Hazards*, 71, 203–227, <https://doi.org/10.1007/s11069-013-0907-4>, 2014.
- 795 Drummond, M. A.: Southern Rockies Ecoregion, in *Status and Trends of Land Change in the Western United States – 1973 to 2000*, edited by: Sleeter B. M., Wilson, T. S., and Acevedo, W., U.S. Geological Survey, 95-103, <https://pubs.usgs.gov/pp/1794/a/>, 2012.
- Ebel, B. A., Moody, J. A., and Martin, D. A.: Hydrologic conditions controlling runoff generation immediately after wildfire, *Water Resources Research*, 48, <https://doi.org/10.1029/2011WR011470>, 2012.
- 800 Esposito, G., Gariano, S. L., Masi, R., Alfano, S., and Giannatiempo, G.: Rainfall conditions leading to runoff-initiated post-fire debris flows in Campania, Southern Italy, *Geomorphology*, 423, 108557, <https://doi.org/10.1016/j.geomorph.2022.108557>, 2023.
- Gabet, E. J. and Bookter, A.: A morphometric analysis of gullies scoured by post-fire progressively bulked debris flows in southwest Montana, USA, *Geomorphology*, 96, 298–309, <https://doi.org/10.1016/j.geomorph.2007.03.016>, 2008.
- 805 García-Ruiz, J. M., Arnáez, J., Gómez-Villar, A., Ortigosa, L., and Lana-Renault, N.: Fire-related debris flows in the Iberian Range, Spain, *Geomorphology*, 196, 221–230, <https://doi.org/10.1016/j.geomorph.2012.03.032>, 2013.

- 810 Gartner, J. E., Cannon, S. H., Santi, P. M., and Dewolfe, V. G.: Empirical models to predict the volumes of debris flows generated by recently burned basins in the western U.S., *Geomorphology*, 96, 339–354, <https://doi.org/10.1016/j.geomorph.2007.02.033>, 2008.
- Gartner, J. E., Cannon, S. H., and Santi, P. M.: Empirical models for predicting volumes of sediment deposited by debris flows and sediment-laden floods in the transverse ranges of southern California, *Engineering Geology*, 176, 45–56, <https://doi.org/10.1016/j.enggeo.2014.04.008>, 2014.
- 815 Gartner, J. E., Kean, J. W., Rengers, F. K., McCoy, S. W., Oakley, N., and Sheridan, G.: Post-Wildfire Debris Flows, in: *Advances in Debris-flow Science and Practice*, edited by: Jakob, M., McDougall, S., and Santi, P., Springer International Publishing, Cham, 309–345, https://doi.org/10.1007/978-3-031-48691-3_11, 2024.
- Gatwood, E., Pederson, J., and Casey, K.: Los Angeles district method for prediction of debris yield. US Army Corps of Engineers, Los Angeles District, 145, 2000. Gorr, A.N., Rengers, F.K., Barnhart, K.R., Thomas, M.A., Kean, J.W., and Crowder, C.A.: Inventory of 227 postfire debris-flow volumes for 34 fires in the western United States, United States Geological Survey Data Release, <https://doi.org/10.5066/P13EZSWW>, 2025.
- 820 Gorr, A. N., McGuire, L. A., and Youberg, A. M.: Empirical Models for Postfire Debris-Flow Volume in the Southwest United States, *Journal of Geophysical Research: Earth Surface*, 129, e2024JF007825, <https://doi.org/10.1029/2024JF007825>, 2024a.
- 825 Gorr, A. N., McGuire, L. A., Youberg, A. M., Beers, R., and Liu, T.: Inundation and flow properties of a runoff-generated debris flow following successive high-severity wildfires in northern Arizona, USA, *Earth Surface Processes and Landforms*, 49, 622–641, <https://doi.org/10.1002/esp.5724>, 2024b.
- Gorr, A. N., McGuire, L. A., Youberg, A. M., and Rengers, F. K.: A progressive flow-routing model for rapid assessment of debris-flow inundation, *Landslides*, 19, 2055–2073, <https://doi.org/10.1007/s10346-022-01890-y>, 2022.
- 830 Gorr, A. N., McGuire, L. A., Beers, R., and Hoch, O. J.: Triggering conditions, runout, and downstream impacts of debris flows following the 2021 Flag Fire, Arizona, USA, *Nat Hazards*, 117, 2473–2504, <https://doi.org/10.1007/s11069-023-05952-9>, 2023.
- Griffith, G. E., Omernik, J. M., Smith, D. W., Cook, T. D., Tallyn, E., Moseley, K., and Johnson, C. B.: Ecoregions of California, U.S. Geological Survey, Open-File Report 2016-1021, <https://doi.org/10.3133/ofr20161021>, 2016.
- 835 Griswold, J. P. and Iverson, R. M.: Mobility statistics and automated hazard mapping for debris flows and rock avalanches, Scientific Investigations Report, U.S. Geological Survey, <https://doi.org/10.3133/sir20075276>, 2008.
- Guthrie, R. H., Friele, P., Allstadt, K., Roberts, N., Evans, S. G., Delaney, K. B., Roche, D., Clague, J. J., and Jakob, M.: The 6 August 2010 Mount Meager rock slide-debris flow, Coast Mountains, British Columbia: characteristics, dynamics, and implications for hazard and risk assessment, *Natural Hazards and Earth System Sciences*, 12, 1277–
840 1294, <https://doi.org/10.5194/nhess-12-1277-2012>, 2012.

- Hancock, C.-A. and Wlodarczyk, K.: The role of wildfires and forest harvesting on geohazards and channel instability during the November 2021 atmospheric river in southwestern British Columbia, Canada, *Earth Surface Processes and Landforms*, 50, e6065, <https://doi.org/10.1002/esp.6065>, 2025.
- 845 Helsel, D. R., Hirsch, R. M., Ryberg, K. R., Archfield, A., and Gilroy, E. J.: *Statistical Methods in Water Resources*, in: U.S. Geological Survey Techniques and Methods, <https://doi.org/10.3133/tm4A3>, 2020.
- Hoch, O. J., McGuire, L. A., Youberg, A. M., and Rengers, F. K.: Hydrogeomorphic Recovery and Temporal Changes in Rainfall Thresholds for Debris Flows Following Wildfire, *Journal of Geophysical Research: Earth Surface*, 126, e2021JF006374, <https://doi.org/10.1029/2021JF006374>, 2021.
- 850 Iverson, R. M., Schilling, S. P., and Vallance, J. W.: Objective delineation of lahar-inundation hazard zones, *GSA Bulletin*, 110, 972–984, [https://doi.org/10.1130/0016-7606\(1998\)110<0972:ODOLIH>2.3.CO;2](https://doi.org/10.1130/0016-7606(1998)110<0972:ODOLIH>2.3.CO;2), 1998.
- Kean, J. W., Staley, D. M., Lancaster, J. T., Rengers, F. K., Swanson, B. J., Coe, J. A., Hernandez, J. L., Sigman, A. J., Allstadt, K. E., and Lindsay, D. N.: Inundation, flow dynamics, and damage in the 9 January 2018 Montecito debris-flow event, California, USA: Opportunities and challenges for post-wildfire risk assessment, *Geosphere*, 15, 1140–1163, <https://doi.org/10.1130/GES02048.1>, 2019.
- 855 Kohavi, R.: A study of cross-validation and bootstrap for accuracy estimation and model selection, in: *Proceedings of the 14th international joint conference on Artificial intelligence - Volume 2*, San Francisco, CA, USA, 1137–1143, 1995.
- Lamb, M. P., Scheingross, J. S., Amidon, W. H., Swanson, E., and Limaye, A.: A model for fire-induced sediment yield by dry ravel in steep landscapes, *Journal of Geophysical Research: Earth Surface*, 116, <https://doi.org/10.1029/2010JF001878>, 2011.
- 860 Lancaster, J. T., Swanson, B. J., Lukashov, S. G., Oakley, N. S., Lee, J. B., Spangler, E. R., Hernandez, J. L., Olson, B. P. E., DeFrisco, M. J., Lindsay, D. N., Schwartz, Y. J., McCrea, S. E., Roffers, P. D., and Tran, C. M.: Observations and Analyses of the 9 January 2018 Debris-Flow Disaster, Santa Barbara County, California, *Environmental & Engineering Geoscience*, 27, 3–27, <https://doi.org/10.2113/EEG-D-20-00015>, 2021.
- 865 Landslide Hazards Program: Scientific Background, <https://www.usgs.gov/programs/landslide-hazards/science/scientific-background> (last access 19 May 2025), 2018.
- Langhans, C., Smith, H. G., Chong, D. M. O., Nyman, P., Lane, P. N. J., and Sheridan, G. J.: A model for assessing water quality risk in catchments prone to wildfire, *Journal of Hydrology*, 534, 407–426, <https://doi.org/10.1016/j.jhydrol.2015.12.048>, 2016.
- 870 Marquardt, D. W.: Generalized Inverses, Ridge Regression, Biased Linear Estimation, and Nonlinear Estimation, *Technometrics*, 12, 591–612, <https://doi.org/10.1080/00401706.1970.10488699>, 1970.

- McGuire, L. A. and Youberg, A. M.: Impacts of successive wildfire on soil hydraulic properties: Implications for debris flow hazards and system resilience, *Earth Surface Processes and Landforms*, 44, 2236–2250, <https://doi.org/10.1002/esp.4632>, 2019.
- 875 McGuire, L. A., Youberg, A. M., Rengers, F. K., Abramson, N. S., Ganesh, I., Gorr, A. N., Hoch, O., Johnson, J. C., Lamom, P., Prescott, A. B., Zanetell, J., and Fenerty, B.: Extreme Precipitation Across Adjacent Burned and Unburned Watersheds Reveals Impacts of Low Severity Wildfire on Debris-Flow Processes, *Journal of Geophysical Research: Earth Surface*, 126, e2020JF005997, <https://doi.org/10.1029/2020JF005997>, 2021.
- McGuire, L. A., Rengers, F. K., Youberg, A. M., Gorr, A. N., Hoch, O. J., Beers, R., and Porter, R.: Characteristics of debris-flow-prone watersheds and debris-flow-triggering rainstorms following the Tadpole Fire, New Mexico, USA, *Natural Hazards and Earth System Sciences*, 24, 1357–1379, <https://doi.org/10.5194/nhess-24-1357-2024>, 2024a.
- 880 McGuire, L. A., Ebel, B. A., Rengers, F. K., Vieira, D. C. S., and Nyman, P.: Fire effects on geomorphic processes, *Nature Reviews Earth & Environment*, 5, 486–503, <https://doi.org/10.1038/s43017-024-00557-7>, 2024b.
- Meyer, G. A. and Wells, S. G.: Fire-related sedimentation events on alluvial fans, Yellowstone National Park, U.S.A., *Journal of Sedimentary Research*, 67, 776–791, <https://doi.org/10.1306/D426863A-2B26-11D7-8648000102C1865D>, 1997.
- 885 Miller, J. D. and Safford, H.: Trends in Wildfire Severity: 1984 to 2010 in the Sierra Nevada, Modoc Plateau, and Southern Cascades, California, USA, *Fire Ecology*, 8, 41–57, <https://doi.org/10.4996/fireecology.0803041>, 2012.
- Miller, J. F., Frederick, R. H., and Tracy, R. J.: *Precipitation-Frequency Atlas of the Western United States. Volume IX-Washington*, 1973.
- 890 Monitoring Trends in Burn Severity: Direct download [data set], <https://www.mtbs.gov/direct-download>, 2025.
- Neptune, C. K., Degraff, J. V., Pluhar, C. J., Lancaster, J. T., and Staley, D. M.: Rainfall Thresholds for Post-Fire Debris-Flow Generation, Western Sierra Nevada, CA, *Environmental & Engineering Geoscience*, 27, 439–453, <https://doi.org/10.2113/EEG-D-21-00039>, 2021.
- 895 Nyman, P., Sheridan, G. J., Smith, H. G., and Lane, P. N. J.: Evidence of debris flow occurrence after wildfire in upland catchments of south-east Australia, *Geomorphology*, 125, 383–401, <https://doi.org/10.1016/j.geomorph.2010.10.016>, 2011.
- Pak, J. H. and Lee, J. J.: A Statistical Sediment Yield Prediction Model Incorporating the Effect of Fires and Subsequent Storm Events, *Journal of the American Water Resources Association*, 44, 689–699, <https://doi.org/10.1111/j.1752-1688.2008.00199.x>, 2008.
- 900 Palucis, M. C., Ulizio, T. P., and Lamb, M. P.: Debris flow initiation from ravel-filled channel bed failure following wildfire in a bedrock landscape with limited sediment supply, *GSA Bulletin*, 133, 2079–2096, <https://doi.org/10.1130/B35822.1>, 2021.

- 905 Parise, M. and Cannon, S. H.: Wildfire impacts on the processes that generate debris flows in burned watersheds, *Natural Hazards*, 61, 217–227, <https://doi.org/10.1007/s11069-011-9769-9>, 2012.
- Parsons, A., Robichaud, P. R., Lewis, S. A., Napper, C., and Clark, J. T.: Field guide for mapping post-fire soil burn severity, General Technical Report RMRS-GTR-243. Fort Collins, CO: U.S. Department of Agriculture, Forest Service, Rocky Mountain Research Station. 49 p., 2010.
- 910 Pelletier, J. D. and Orem, C. A.: How do sediment yields from post-wildfire debris-laden flows depend on terrain slope, soil burn severity class, and drainage basin area? Insights from airborne-LiDAR change detection, *Earth Surface Processes and Landforms*, 39, 1822–1832, <https://doi.org/10.1002/esp.3570>, 2014.
- Perica, S., Dietz, S., Heim, S., Hiner, L., Maitaria, K., Martin, D., Pavlovic, S., Roy, I., Trypaluk, C., Unruh, D., Yan, F., Yekta, M., Zhao, T., Bonnin, G., Brewer, D., Chen, L.-C., Parzybok, T., and Yarchoan, J.: Precipitation-Frequency Atlas of the United States. Volume 6 Version 2.3. California, 2014.
- 915 Perica, S., Martin, D., Pavlovic, S., Roy, I., St. Laurent, M., Trypaluk, C., Unruh, D., Yekta, M., and Bonnin, G.: Precipitation-Frequency Atlas of the United States. Volume 8 Version 2.0. Midwestern States (Colorado, Iowa, Kansas, Michigan, Minnesota, Missouri, Nebraska, North Dakota, Oklahoma, South Dakota, Wisconsin), 2013.
- 920 Perkins, J. P., Diaz, C., Corbett, S. C., Cerovski-Darriau, C., Stock, J. D., Prancevic, J. P., Micheli, E., and Jasperse, J.: Multi-Stage Soil-Hydraulic Recovery and Limited Ravel Accumulations Following the 2017 Nuns and Tubbs Wildfires in Northern California, *Journal of Geophysical Research: Earth Surface*, 127, e2022JF006591, <https://doi.org/10.1029/2022JF006591>, 2022.
- Prescott, A. B., McGuire, L. A., Jun, K.-S., Barnhart, K. R., and Oakley, N. S.: Probabilistic assessment of postfire debris-flow inundation in response to forecast rainfall, *Natural Hazards and Earth System Sciences*, 24, 2359–2374, <https://doi.org/10.5194/nhess-24-2359-2024>, 2024.
- 925 PRISM Climate Group, Oregon State University [data set], <https://prism.oregonstate.edu>, 2025.
- Radeloff, V. C., Helmers, D. P., Kramer, H. A., Mockrin, M. H., Alexandre, P. M., Bar-Massada, A., Butsic, V., Hawbaker, T. J., Martinuzzi, S., Syphard, A. D., and Stewart, S. I.: Rapid growth of the US wildland-urban interface raises wildfire risk, *Proceedings of the National Academy of Sciences*, 115, 3314–3319, <https://doi.org/10.1073/pnas.1718850115>, 2018.
- 930 Reilly, M. J., Zuspan, A., Halofsky, J. S., Raymond, C., McEvoy, A., Dye, A. W., Donato, D. C., Kim, J. B., Potter, B. E., Walker, N., Davis, R. J., Dunn, C. J., Bell, D. M., Gregory, M. J., Johnston, J. D., Harvey, B. J., Halofsky, J. E., and Kerns, B. K.: Cascadia Burning: The historic, but not historically unprecedented, 2020 wildfires in the Pacific Northwest, USA, *Ecosphere*, 13, e4070, <https://doi.org/10.1002/ecs2.4070>, 2022.
- 935 Rengers, F. K., McGuire, L. A., Barnhart, K. R., Youberg, A. M., Cadol, D., Gorr, A. N., Hoch, O. J., Beers, R., and Kean, J. W.: The influence of large woody debris on post-wildfire debris flow sediment storage, *Natural Hazards and Earth System Sciences*, 23, 2075–2088, <https://doi.org/10.5194/nhess-23-2075-2023>, 2023.

- Rengers, F. K., Bower, S., Knapp, A., Kean, J. W., von Lembke, D. W., Thomas, M. A., Kostelnik, J., Barnhart, K. R., Bethel, M., Gartner, J. E., Hille, M., Staley, D. M., Anderson, J. K., Roberts, E. K., DeLong, S. B., Lane, B., Ridgway, P., and Murphy, B. P.: Evaluating post-wildfire debris-flow rainfall thresholds and volume models at the 2020 Grizzly Creek Fire in Glenwood Canyon, Colorado, USA, *Natural Hazards and Earth System Sciences*, 24, 2093–2114, <https://doi.org/10.5194/nhess-24-2093-2024>, 2024.
- Rickenmann, D. and Zimmermann, M.: The 1987 debris flows in Switzerland: documentation and analysis, *Geomorphology*, 8, 175–189, [https://doi.org/10.1016/0169-555X\(93\)90036-2](https://doi.org/10.1016/0169-555X(93)90036-2), 1993.
- Rickenmann, D.: Empirical Relationships for Debris Flows, *Natural Hazards*, 19, 47–77, <https://doi.org/10.1023/A:1008064220727>, 1999.
- Santi, P. M. and Morandi, L.: Comparison of debris-flow volumes from burned and unburned areas, *Landslides*, 10, 757–769, <https://doi.org/10.1007/s10346-012-0354-4>, 2013.
- Santi, P. M.: Precision and Accuracy in Debris-Flow Volume Measurement, *Environmental & Engineering Geoscience*, 20, 349–359, <https://doi.org/10.2113/gseegeosci.20.4.349>, 2014.
- 940 Selander, B. D., Calhoun, N., Burns, W. J., Kean, J. W., and Rengers, F. K.: Assessment of western Oregon debris-flow hazards in burned and unburned environments, *Earth Surface Processes and Landforms*, 50, e70045, <https://doi.org/10.1002/esp.70045>, 2025.
- Sepúlveda, S. A., Rebolledo, S., and Vargas, G.: Recent catastrophic debris flows in Chile: Geological hazard, climatic relationships and human response, *Quaternary International*, 158, 83–95, <https://doi.org/10.1016/j.quaint.2006.05.031>, 2006.
- 955 Smith, D. P., Schnieders, J., Marshall, L., Melchor, K., Wolfe, S., Campbell, D., French, A., Randolph, J., Whitaker, M., Klein, J., Steinmetz, C., and Kwan, R.: Influence of a Post-dam Sediment Pulse and Post-fire Debris Flows on Steelhead Spawning Gravel in the Carmel River, California, *Front. Earth Sci.*, 9, <https://doi.org/10.3389/feart.2021.802825>, 2021.
- 960 Smith, H. G., Sheridan, G. J., Lane, P. N. J., Nyman, P., and Haydon, S.: Wildfire effects on water quality in forest catchments: A review with implications for water supply, *Journal of Hydrology*, 396, 170–192, <https://doi.org/10.1016/j.jhydrol.2010.10.043>, 2011.
- Smith, T., Rheinwalt, A., and Bookhagen, B.: Determining the optimal grid resolution for topographic analysis on an airborne lidar dataset, *Earth Surface Dynamics*, 7, 475–489, <https://doi.org/10.5194/esurf-7-475-2019>, 2019.
- 965 Staley, D. M., Kean, J. W., Cannon, S. H., Schmidt, K. M., and Laber, J. L.: Objective definition of rainfall intensity–duration thresholds for the initiation of post-fire debris flows in southern California, *Landslides*, 10, 547–562, <https://doi.org/10.1007/s10346-012-0341-9>, 2013.

- Staley, D. M., Gartner, J. E., and Kean, J. W.: Objective Definition of Rainfall Intensity-Duration Thresholds for Post-fire Flash Floods and Debris Flows in the Area Burned by the Waldo Canyon Fire, Colorado, USA, in: Engineering Geology for Society and Territory - Volume 2, Springer, Cham, 621–624, https://doi.org/10.1007/978-3-319-09057-3_103, 2015.
- 970 Staley, D. M., Negri, J. A., Kean, J. W., Laber, J. L., Tillery, A. C., and Youberg, A. M.: Prediction of spatially explicit rainfall intensity–duration thresholds for post-fire debris-flow generation in the western United States, *Geomorphology*, 278, 149–162, <https://doi.org/10.1016/j.geomorph.2016.10.019>, 2017.
- 975 Staley, D. M., Kean, J. W., and Rengers, F. K.: The recurrence interval of post-fire debris-flow generating rainfall in the southwestern United States, *Geomorphology*, 370, 107392, <https://doi.org/10.1016/j.geomorph.2020.107392>, 2020.
- Swanson, B. J., Lindsay, D. N., Cato, K., DiBiase, R. A., and Neely, A. B.: Debris flows and sediment transport at Yucaipa Ridge and impacts to Oak Glen and Forest Falls area, southern California, following the 2020 El Dorado and Apple Fires, in: *From Coastal Geomorphology to Magmatism: Guides to GSA Connects 2024 Field Trips in Southern California and Beyond*, vol. 70, edited by: Van Buer, N. J. and Schwartz, J. J., Geological Society of America, [https://doi.org/10.1130/2024.0070\(03\)](https://doi.org/10.1130/2024.0070(03)), 2024.
- 980 Thomas, M. A., Lindsay, D. N., Cavagnaro, D. B., Kean, J. W., McCoy, S. W., and Graber, A. P.: The Rainfall Intensity-Duration Control of Debris Flows After Wildfire, *Geophysical Research Letters*, 50, e2023GL103645, <https://doi.org/10.1029/2023GL103645>, 2023.
- 985 Tillery, A. C. and Rengers, F. K.: Controls on debris-flow initiation on burned and unburned hillslopes during an exceptional rainstorm in southern New Mexico, USA, *Earth Surface Processes and Landforms*, 45, 1051–1066, <https://doi.org/10.1002/esp.4761>, 2020.
- U.S. Environmental Protection Agency: Level III ecoregions of the continental United States, U.S. EPA – National Health and Environmental Effects Research Laboratory, Corvallis, Oregon, <https://www.epa.gov/eco-research/level-iii-and-iv-ecoregions-continental-united-states>, 2013.
- 990 Vieira, D. C. S., Fernández, C., Vega, J. A., and Keizer, J. J.: Does soil burn severity affect the post-fire runoff and interrill erosion response? A review based on meta-analysis of field rainfall simulation data, *Journal of Hydrology*, 523, 452–464, <https://doi.org/10.1016/j.jhydrol.2015.01.071>, 2015.
- 995 Wall, S., Murphy, B. P., Belmont, P., and Yocom, L.: Predicting post-fire debris flow grain sizes and depositional volumes in the Intermountain West, United States, *Earth Surface Processes and Landforms*, 48, 179–197, <https://doi.org/10.1002/esp.5480>, 2023.
- Wall, S. A., Roering, J. J., and Rengers, F. K.: Runoff-initiated post-fire debris flow Western Cascades, Oregon, *Landslides*, 17, 1649–1661, <https://doi.org/10.1007/s10346-020-01376-9>, 2020.

- 1000 Wang, G., Sassa, K., and Fukuoka, H.: Downslope volume enlargement of a debris slide–debris flow in the 1999 Hiroshima, Japan, rainstorm, *Engineering Geology*, 69, 309–330, [https://doi.org/10.1016/S0013-7952\(02\)00289-2](https://doi.org/10.1016/S0013-7952(02)00289-2), 2003.
- Wells, W. G., II: The effects of fire on the generation of debris flows in southern California, in: *Debris Flows/Avalanches: Process, Recognition, and Mitigation*, vol. 7, edited by: Costa, J. E. and Wieczorek, G. F., Geological Society of America, <https://doi.org/10.1130/REG7-p105>, 1987.
- 1005 Westerling, A. L.: Increasing western US forest wildfire activity: sensitivity to changes in the timing of spring, *Philosophical Transactions of the Royal Society B: Biological Sciences*, 371, 20150178, <https://doi.org/10.1098/rstb.2015.0178>, 2016.
- Wilken, E., Nava, F. N., and Griffith, G.: *North America Terrestrial Ecoregions – Level III*. Commission for Environmental Cooperation, 2011.
- 1010 Youberg, A. M.: *Prehistoric and modern debris flows in semi-arid watersheds: Implications for hazard assessments in a changing climate*, 2014.