

1 **Aerosol Scavenging in DC3 and SEAC<sup>4</sup>RS Deep Convective Storms**

2 Mary C. Barth et al.

3 **Supplementary Text**

4 S1. Description of storm cases included in the analysis

5 **18 May 2012 DC3 case**

6 The 18 May 2012 case, sampled near Ogallala, Nebraska in the vicinity of a cold front, had a SWEAT Index of 283  
7 categorizing it as a moderate, single cell storm. The storm was quite strong with respect to the 20 dBZ cloud top  
8 heights of 14 km MSL, reports of hail of one inch diameter (Herndon, 2012a), and periods of high lightning flash  
9 rates. The storm had relatively higher anthropogenic and agriculture signatures and lower biogenic signatures based  
10 on VOC measurements. The environment had a dry aerosol extinction of  $31 \text{ Mm}^{-1}$ , which is slightly higher than typical  
11 rural conditions (Barth et al., 2015). The BL  $f(OA)$  was 36%, underscoring the weaker influence of biogenic VOCs.

12 **29 May 2012 DC3 case**

13 The 29 May 2012 storm, sampled north of Oklahoma City, Oklahoma as a line of supercell storms, had a SWEAT  
14 Index of 422 producing large hail (up to 3 inch diameter) and a weak tornado. With its high convective available  
15 potential energy and 0-6 km wind shear, 20 dBZ cloud top heights exceeded 17 km MSL altitude. Vertical velocities  
16 exceeded  $50 \text{ m s}^{-1}$  (DiGangi et al., 2016). The central Oklahoma region has characteristics of both anthropogenic and  
17 biogenic VOCs with nearby oil and gas fields in central Oklahoma and north Texas, urban influences (Oklahoma  
18 City), and vegetation in eastern Oklahoma. The BL  $f(OA)$  was 53% and the dry aerosol extinction was  $36 \text{ Mm}^{-1}$ ,  
19 slightly higher than typical rural conditions. This case is the same as that studied by Yang et al. (2015).

20 **02 June 2012 DC3 case**

21 On 2 June 2012, isolated convection was sampled between Greeley and Fort Morgan northeast of Denver, Colorado.  
22 The storm can be characterized as moderate convection (SWEAT Index of 256). While storm reports do not include  
23 observations of hail, there were reports of surface winds  $> 33 \text{ m s}^{-1}$  (Herndon, 2012b). The 20 dBZ cloud top heights  
24 reached 15 km MSL. Like the 18 May case, the chemical environment showed signatures of anthropogenic influences  
25 from the Front Range of Colorado and smaller biogenic VOC mixing ratios. The dry aerosol extinction was low ( $12 \text{ Mm}^{-1}$ )  
26 and the BL  $f(OA)$  was 58%.

27 **06 June 2012 DC3 case**

28 Convection on 6 June 2012 was associated with the “Denver cyclone,” where low-level flow is southeasterly on the  
29 plains east of Denver and is northwesterly to the west of Denver, Colorado. The strong convection occurred just  
30 northeast of Denver and had a SWEAT Index of 296, 20 dBZ cloud top heights of 15 km MSL, and storm reports of  
31 one-inch diameter hail (Herndon, 2012b). Both the biogenic and anthropogenic VOC mixing ratios in the BL were  
32 low relative to other Colorado cases. The dry aerosol extinction was typical of rural background conditions ( $24 \text{ Mm}^{-1}$ )  
33 and the BL  $f(OA)$  was 50%.

34 **16 June 2012 DC3 case**

35 Multicell convection occurred on 16 June 2012 in central Oklahoma. Its SWEAT Index of 360 categorized it as a  
36 severe storm. The 20 dBZ cloud top heights reached 15 km MSL. In contrast to the 29 May storm in Oklahoma, both  
37 the anthropogenic and biogenic signatures of VOCs were low in the vicinity of this storm. The dry aerosol extinction  
38 was low ( $20 \text{ Mm}^{-1}$ ), typical of rural conditions. The  $\text{BL } f(OA)$  was 45%.

39 **22 June 2012 DC3 case**

40 The 22 June 2012 case had two severe convective storms occur along the Colorado-Nebraska border (at  $41^\circ\text{N}$ ). The  
41 north storm was unique in that it ingested a wildfire smoke plume at 7 km MSL elevation, while the south storm did  
42 not ingest this smoke plume (Barth et al., 2015; Apel et al., 2015). The scavenging efficiency is estimated for only the  
43 south storm. The SWEAT Index for these storms was very high (442), cloud top heights reached 18 km MSL and  
44 there were reports of 1–2-inch hail for both storms (Herndon, 2012b). The smoke plume came from the High Park  
45 Fire west of Fort Collins, Colorado which had been burning for  $\sim 2$  weeks. Biomass burning was occurring throughout  
46 the Rocky Mountain region resulting in its higher contribution to the atmospheric composition in the troposphere. The  
47 dry aerosol extinction was high ( $41 \text{ Mm}^{-1}$ ) for the region and  $\text{BL } f(OA)$  was 59%. Like other northeast Colorado  
48 storms, the biogenic VOCs were low relative to the anthropogenic VOCs (Barth et al., 2015).

49 **02 September 2013 SEAC<sup>4</sup>RS case**

50 On 2 September 2013 a frontal system extending from north of the Great Lakes to southwest Texas was moving  
51 eastward across the US (Cuchiara et al., 2020). Associated with this front was the development of pre-frontal  
52 convection in Mississippi mostly of the form of airmass and multicell storms. The SWEAT Index calculated from the  
53 1200 UTC Jackson, Mississippi NWS sounding was 214, while the SWEAT Index calculated from aircraft  
54 observations near the two storms was 225 for the airmass storm and 209 for the multicell storm (Table 1). In general,  
55 these storms were weaker than the storms sampled in Colorado and Oklahoma during DC3. The airmass and multicell  
56 storm cloud top heights were 8 km MSL and 13 km MSL, respectively, and the maximum radar reflectivity was lower  
57 in the airmass storm (45 dBZ) compared to the multicell storm (55 dBZ). The pre-frontal convection occurred mid-  
58 day (1800–1900 UTC; 1300–1400 LT) in a region rich in biogenic VOCs and low in anthropogenic VOCs (average  
59 BL mixing ratios were 1200 pptv isoprene and 39 pptv toluene). The dry aerosol extinction was high ( $47 \text{ Mm}^{-1}$ ) for  
60 the region with  $\text{BL } f(OA)$  of 54%.

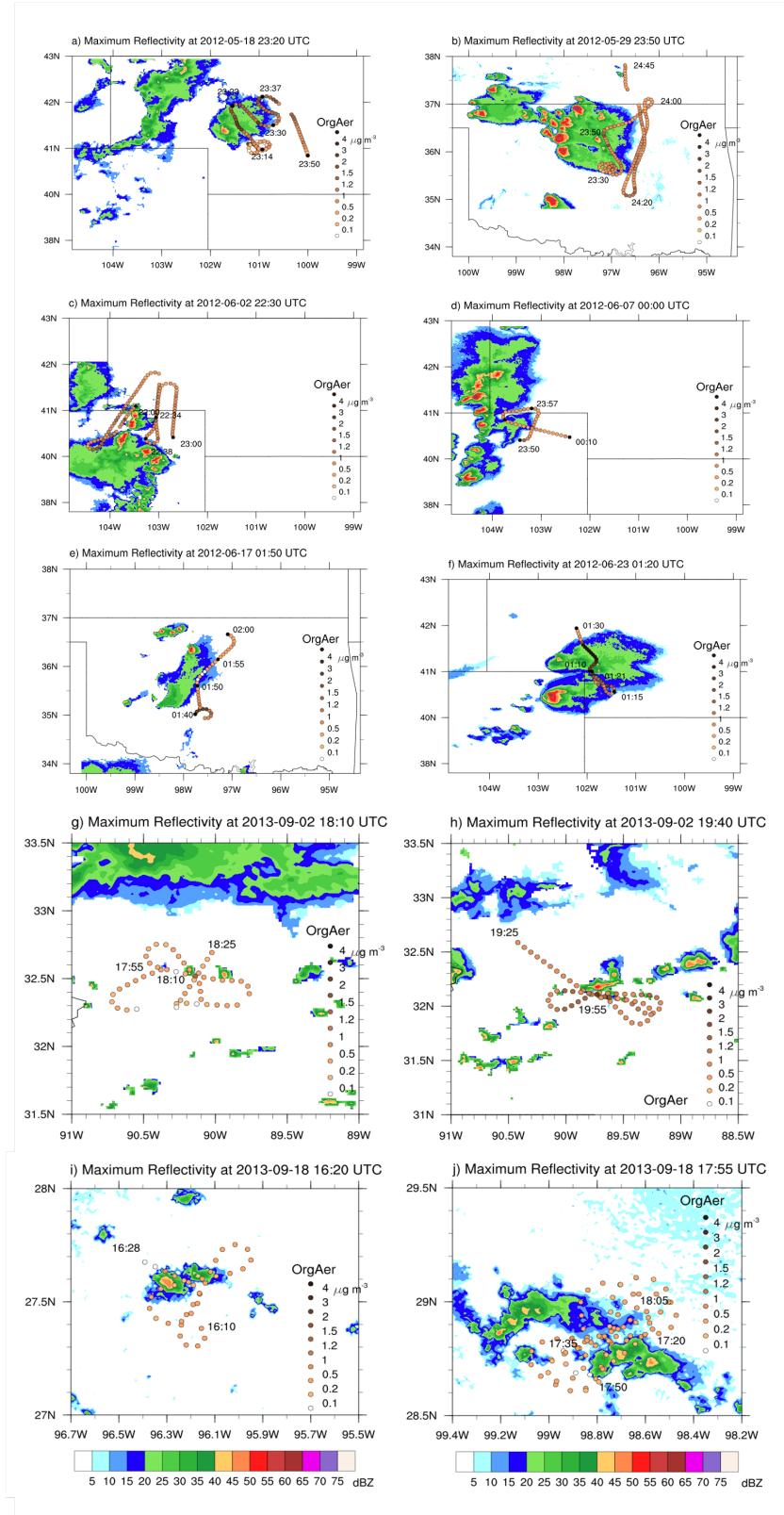
61 **18 September 2013 SEAC<sup>4</sup>RS case**

62 On 18 September 2013, the position of the “Bermuda High” pressure system over the mid-Atlantic coast and two low  
63 pressure systems over the Yucatan Peninsula and west coast of Mexico favored southeasterly flow from the Gulf of  
64 Mexico toward southern Texas giving moderate instability off the Texas coast and inland (Cuchiara et al., 2023). As  
65 a result, convection occurred over the Gulf of Mexico near Corpus Christi, Texas as well as to the south of San  
66 Antonio, Texas. The convection was sampled in the late morning to midday (16:00 – 19:00 UTC, 11:00 – 14:00 LT).  
67 To calculate the SWEAT Index, we used aircraft vertical profiles near the convection since the NWS Corpus Christi  
68 radiosonde occurred at least four hours before the convection was sampled. The SWEAT Index indicated weak to  
69 moderate convection for both the marine (SWEAT = 242) and land (SWEAT = 257) convection. The 20 dBZ cloud

70 top heights for both marine and land convection reached 8 km MSL and the maximum radar reflectivities were 45-50  
71 dBZ. Being in the same southeasterly flow, the chemical environment was clean for both marine and land convection.  
72 Average BL mixing ratios for isoprene and toluene were <60 pptv and the aerosol dry extinction was <12  $Mm^{-1}$  over  
73 both the marine and land regions. This contrasts with the findings by Cuchiara et al. (2023) who reported differences  
74 in cloud condensation nuclei (CCN) concentrations (282  $cm^{-3}$  and 484  $cm^{-3}$  in the marine and land convective inflow  
75 regions, respectively) and cloud characteristics of LWC, cloud base height, and updraft velocity. The  $BL_f(OA)$  was  
76 22% and 34% for the marine and land regions, respectively, which is smaller than the other cases analyzed in this  
77 study.

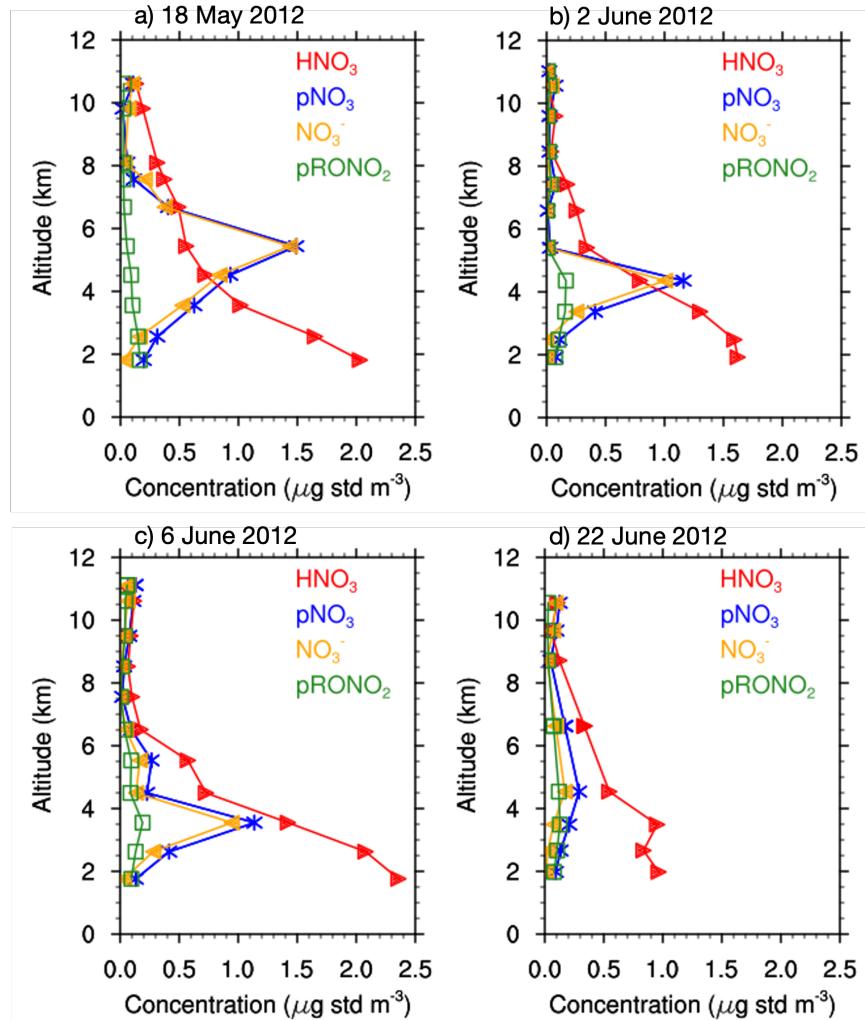
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79 Supplementary Figures



81 Figure S1. Maximum radar reflectivity in each vertical column for the ten cases analyzed, a) 18 May 2012, b) 29  
82 May 2012, c), 2 June 2012, d) 6 June 2012, e) 16 June 2012, f) 22 June 2012, g) 2 September 2013 air mass storm,  
83 h) 2 September 2013 multicell storm, i) 18 September 2013 Gulf of Mexico storm, and j) 18 September 2013 South  
84 Texas storm. Overlaid is the DC-8 flight track in the storm outflow colored by the organic aerosol mass  
85 concentration ( $\mu\text{g std m}^{-3}$ ).

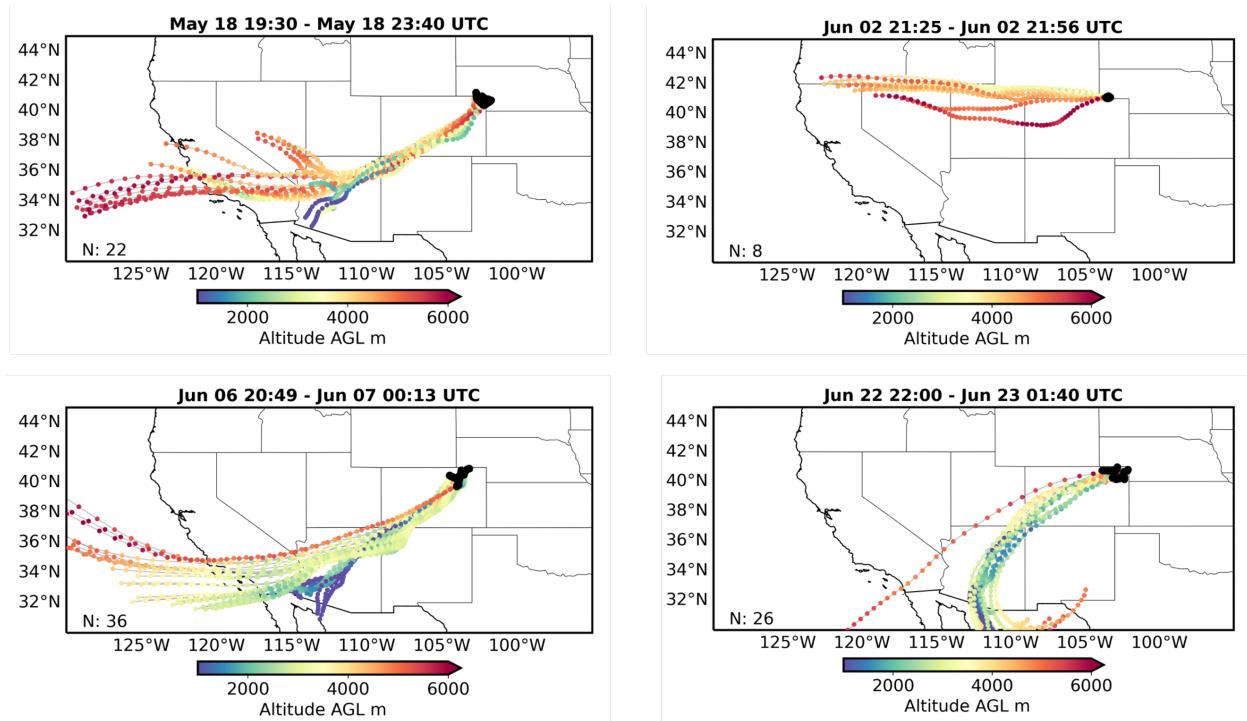
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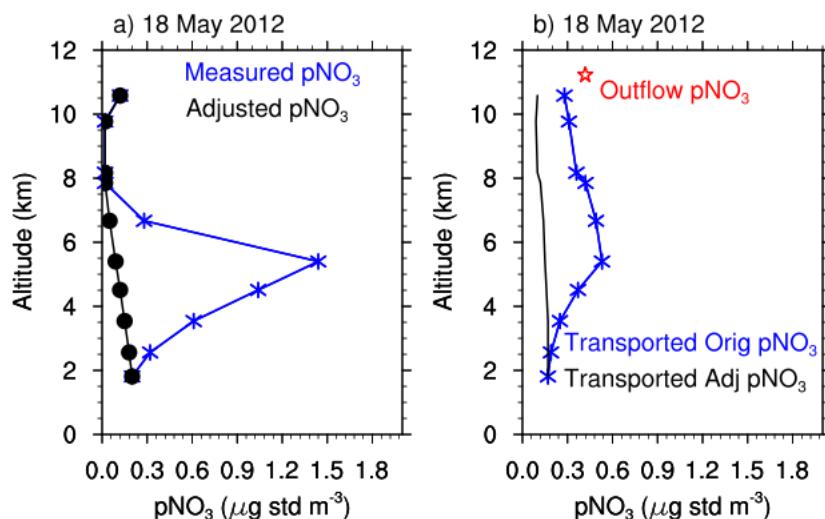
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90 Figure S2. Clear air vertical profiles of  $\text{HNO}_3$  (red), particulate nitrate (blue), inorganic particulate nitrate (gold), and  
91 organic particulate nitrate (green) for a) 18 May, b) 2 June, c) 6 June, and d) 22 June DC3 storms.

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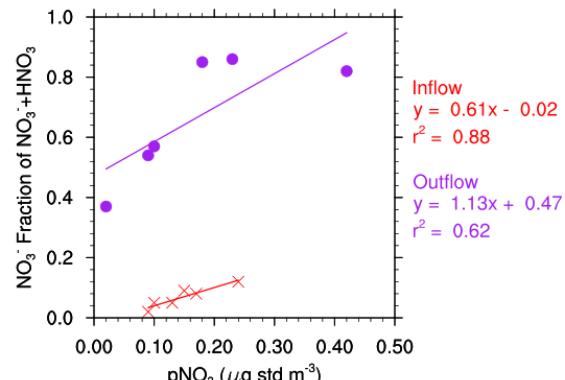


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95 Figure S3. HYSPLIT 48-hour back trajectories calculated for four DC3 storms.



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100 Figure S4. Vertical profiles of particulate nitrate (left) that were measured (blue) and adjusted to remove the mid-  
101 troposphere layer (black) and of transported particulate nitrate calculated by the entrainment model (right) with  
102 calculated values using the measured clear sky vertical profile (blue) and those using the adjusted clear sky vertical  
103 profile (black) for the 18 May 2012 case. The red circle marks the measured concentration of the outflow particulate  
104 nitrate concentration.

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Figure S5. Scatter plot of pNO<sub>3</sub> and the NO<sub>3</sub><sup>-</sup> fraction of NO<sub>3</sub><sup>-</sup> + HNO<sub>3</sub> for the six DC3 cases. Purple solid circles are for the convective outflow regions and red crosses are for the inflow regions.

114 **Supplementary Tables**

115 Table S1. List of data and instruments used in the analysis.

Species/Parameter	Instrument <sup>a</sup>	Uncertainty	Reference
SO <sub>4</sub> <sup>2-</sup> , NH <sub>4</sub> <sup>+</sup> , NO <sub>3</sub> <sup>-</sup> , OA, <i>m/z</i> 44 for submicron aerosol mass	CU aircraft AMS	Inorganics 34%, Organics 38% (2 sigma)	DeCarlo et al. (2006), Guo et al. (2021)
CO	DACOM	5% or 1 ppmv	Sachse et al. (1987)
CO <sub>2</sub>	AVOCET	0.25 ppmv	Vay et al. (2011)
n-butane, i-butane, n-pentane, i-pentane, isoprene, toluene	WAS	5% or 3 pptv	Simpson et al. (2011)
isoprene, toluene	PTR-MS	5%	de Gouw and Warneke (2007)
Aerosol dry extinction at 532 nm	NASA/LaRC optical aerosol measurements	5%	Wagner et al. (2015)
BC	HD-SP2	30%	Schwarz et al. (2013)
HCN	CIT-CIMS	50% + 50 pptv	Crounse et al. (2006)
CH <sub>3</sub> CN	PTR-MS	25 pptv	de Gouw and Warneke (2007)
MPN, ANs, PNs	TD-LIF	MPN, ANs: 15% PNs: 10%	Nault et al. (2015)
HNO <sub>3</sub>	CIT-CIMS	30% + 50 pptv	Crounse et al. (2006)
O <sub>3</sub>	CSD CL	0.040 ppbv + 3%	Ryerson et al. (2000); Pollack et al. (2011)
OH	ATHOS	32%	Faloona et al. (2004)
Aerosol number	SMPS	10%	Wang and Flagan (1993)
Pressure, Temperature, 3D Winds	MMS	Pressure: 0.5% Temperature: 0.2% Winds: 3%	Chan et al. (1998)
Ice water content	2D-S	Not available	Lawson et al. (2011)

<sup>a</sup>CU-HR-TOF-AMS is the University of Colorado High Resolution, Time of Flight, Aerosol Mass Spectrometer; DACOM is the Differential Absorption CO Measurement; AVOCET is the Atmospheric Vertical Observation of CO<sub>2</sub> in the Earth's Troposphere; WAS is the Whole Air Sampler that uses gas chromatography; PTR-MS is the Institut fuer Ionenphysik und Angewandte Physik Proton Transfer Mass Spectrometer; HD-SP2 is the Humidified Dual Single Particle Soot Photometer; CIT-CIMS is California Institute of Technology chemical ionization mass spectrometry; TD-LIF is thermal dissociation – laser induced fluorescence; CSD CL is NOAA Chemical Science Division chemiluminescence; ATHOS is the Airborne Tropospheric Hydrogen Oxides Sensor; SMPS is the Scanning Mobility Particle Sizer; MMS is Meteorological Measurement System; 2D-S is two-dimensional stereo probe.

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126 Table S2. NASA DC-8 aircraft inflow and outflow times, their associated altitudes, and entrainment rates for  
 127 each storm analyzed.

Campaign and Date (mm/dd/yy)	Inflow Altitude (km)	Inflow Time (UTC)	Outflow Altitude (km)	Outflow Time (UTC)	Entrain. Rate (% km <sup>-1</sup> )
DC3 05/18/12	1.7	22:48:29-22:51:10	11.3	23:17:50-23:22:00	14.7
DC3 05/29/12	1.3	23:10:21-23:15:53	11.0	23:48:30-23:58:13	7.8
DC3 06/02/12	1.9	21:16:18-21:27:38	11.1	22:34:14-22:46:10	11.5
DC3 06/06/12	1.7	22:13:40-22:25:12	12.4	23:56:00-24:10:00	4.1
DC3 06/16/12	0.95	24:15:00-24:20:00	11.9	25:50:00-25:55:00	15.4
DC3 06/22/12	2.0	22:31:27-22:45:54	11.2	25:13:52-25:14:38 25:16:51-25:19:24	4.1
SEAC <sup>4</sup> RS 09/02/13 Airmass	0.8	16:53:11-16:53:551	8.0 8.0 8.0	18:09:45-18:10:15 18:16:03-18:16:32 18:23:16-18:23:33	33.3 11.3 18.2
SEAC <sup>4</sup> RS 09/02/13 Multicell	0.4	22:12:01-22:18:14	12.0 12.0	19:31:10-19:31:55 19:53:34-19:54:11	10.2 12.5
SEAC <sup>4</sup> RS 09/18/13 Marine	0.6	15:38:35-15:39:18	8.1 8.1 8.1	16:15:13-16:15:16 16:20:34-16:20:36 16:21:01-16:21:10	9.6
SEAC <sup>4</sup> RS 09/18/13 Land	0.9	18:30:58-18:35:38	10.1 10.2 10.7 11.3 12.0 12.0 11.1	17:16:40-17:16:47 17:17:01-17:17:19 17:22:23-17:22:30 17:27:48-17:27:58 17:35:31-17:35:40 17:40:28-17:40:41 17:46:39-17:46:48	7.8

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133 Table S3. Inflow, outflow, and calculated aerosol concentrations ( $\mu\text{g std m}^{-3}$ ) for each case and the estimated  
 134 scavenging efficiency (%) given as averages  $\pm$  standard deviation.

Date	Measured Inflow Concentration <sup>a</sup>	Measured Outflow Concentration <sup>a</sup>	Calculated Cloud Top Concentration <sup>b</sup>	Scavenging Efficiency <sup>b,c,d</sup>
<i>Sulfate</i>				
18 May 2012	$2.93 \pm 0.16$	$0.44 \pm 0.14$	1.19	63.5
29 May 2012	$2.11 \pm 0.24$	$0.16 \pm 0.07$	1.19	86.8
02 June 2012	$0.80 \pm 0.09$	$0.18 \pm 0.08$	0.42	56.7
06 June 2012	$2.00 \pm 0.26$	$0.15 \pm 0.07$	1.49	89.7
16 June 2012	$1.58 \pm 0.16$	$0.06 \pm 0.05$	0.67	90.7
22 June 2012	$1.33 \pm 0.16$	$0.26 \pm 0.02$	1.13	$77.3 \pm 1.5$
02 Sept 2013 <sup>e</sup>	$3.45 \pm 0.27$	$0.10 \pm 0.04$	0.93	88.9
02 Sept 2013	$3.22 \pm 0.46$	$0.12 \pm 0.04$	$1.20 \pm 0.17$	$90.3 \pm 1.6$
18 Sept 2013 <sup>e</sup>	$1.71 \pm 0.37$	$0.11 \pm 0.06$	0.89	100
18 Sept 2013	$1.48 \pm 0.45$	$0.08 \pm 0.02$	$0.75 \pm 0.03$	$89.8 \pm 3.1$
<i>Ammonium</i>				
18 May 2012	$1.15 \pm 0.07$	$0.31 \pm 0.09$	0.51	39.2
29 May 2012	$0.87 \pm 0.09$	$0.09 \pm 0.03$	0.48	81.2
02 June 2012	$0.33 \pm 0.04$	$0.12 \pm 0.05$	0.18	37.0
06 June 2012	$0.72 \pm 0.09$	$0.08 \pm 0.02$	0.55	86.0
16 June 2012	$0.58 \pm 0.05$	$0.03 \pm 0.02$	0.24	87.6
22 June 2012	$0.50 \pm 0.05$	$0.15 \pm 0.01$	0.46	$66.9 \pm 1.8$
02 Sept 2013 <sup>e</sup>	$0.95 \pm 0.07$	$0.02 \pm 0.02$	0.22	100
02 Sept 2013	$0.83 \pm 0.14$	$0.07 \pm 0.04$	$0.30 \pm 0.04$	$75.0 \pm 15.8$
18 Sept 2013 <sup>e</sup>	$0.10 \pm 0.03$	$0.001 \pm 0.001$	0.05	100
18 Sept 2013	$0.40 \pm 0.10$	$0.001 \pm 0.001$	$0.18 \pm 0.01$	100
<i>Nitrate</i>				
18 May 2012	$0.17 \pm 0.05$	$0.42 \pm 0.12$	0.28	–
29 May 2012	$0.24 \pm 0.04$	$0.09 \pm 0.05$	0.15	39.6
02 June 2012	$0.10 \pm 0.02$	$0.18 \pm 0.08$	0.13	–
06 June 2012	$0.13 \pm 0.03$	$0.10 \pm 0.05$	0.16	40.8
16 June 2012	$0.15 \pm 0.03$	$0.02 \pm 0.04$	0.13	82.8
22 June 2012	$0.09 \pm 0.03$	$0.23 \pm 0.02$	0.18	–
02 Sept 2013 <sup>e</sup>	$0.18 \pm 0.04$	$0.08 \pm 0.14$	0.05	–
02 Sept 2013	$0.07 \pm 0.03$	$0.33 \pm 0.03$	$0.09 \pm 0.00$	–
18 Sept 2013 <sup>e</sup>	$0.03 \pm 0.01$	$0.01 \pm 0.01$	0.02	NaN
18 Sept 2013	$0.08 \pm 0.07$	$0.03 \pm 0.01$	$0.04 \pm 0.00$	100
<i>Organic Aerosol</i>				
18 May 2012	$6.05 \pm 0.42$	$1.92 \pm 1.06$	2.53	24.3
29 May 2012	$8.76 \pm 0.71$	$0.74 \pm 0.51$	4.55	83.8
02 June 2012	$3.48 \pm 0.40$	$0.90 \pm 0.59$	1.68	46.4
06 June 2012	$4.89 \pm 0.38$	$0.51 \pm 0.31$	3.70	86.3
16 June 2012	$4.12 \pm 0.54$	$0.19 \pm 0.39$	1.80	89.5
22 June 2012	$5.80 \pm 0.39$	$1.64 \pm 0.52$	5.19	$68.4 \pm 10.0$
02 Sept 2013 <sup>e</sup>	$5.92 \pm 0.46$	$0.86 \pm 1.97$	1.68	49.0
02 Sept 2013	$6.46 \pm 1.51$	$1.49 \pm 0.88$	$2.50 \pm 0.32$	$42.3 \pm 27.9$
18 Sept 2013 <sup>e</sup>	$0.66 \pm 0.18$	$0.31 \pm 0.16$	0.42	100
18 Sept 2013	$1.05 \pm 0.33$	$0.28 \pm 0.13$	$0.54 \pm 0.01$	100

135 <sup>a</sup>Average concentrations below the detection limit are shown in italics.

136 <sup>b</sup>For cases where the outflow leg is below the detection limit, the scavenging efficiency is shown as 100% since the  
 137 measured outflow concentration is not statistically different than zero.

138 <sup>c</sup>For cases with one outflow leg (Table S2), there is no scavenging efficiency standard deviation calculated.

139 <sup>d</sup>For cases marked with a dash, the scavenging efficiency is not given because the outflow concentration is greater  
 140 than the inflow concentration. The case marked NaN is not a number because both the inflow and outflow measured  
 141 concentrations are below the detection limit.

142 <sup>e</sup>Measurements from the outflow intercepts were combined to have a larger sample size for the scavenging  
 143 efficiency calculation.

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Table S4. Inorganic nitrate partitioning ratios:  $NO_3^-/(NO_3^- + HNO_3)$ .

Date	Inflow	Outflow
18 May 2012	0.08	0.82
29 May 2012	0.12	0.54
02 June 2012	0.05	0.85
06 June 2012	0.05	0.57
16 June 2012	0.09	0.37
22 June 2012	0.02	0.86

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151 Table S5. Inflow and outflow particulate inorganic nitrate ( $NO_3^-$ ), and particulate organic nitrate (p $RONO_2$ ) average  
152  $\pm$  standard deviation concentrations ( $\mu\text{g std m}^{-3}$ ) and their estimated scavenging efficiency (%) for the DC3 cases.

Date	Measured Inflow Concentration <sup>a</sup>	Measured Outflow Concentration <sup>a</sup>	Scavenging Efficiency <sup>b</sup>
$NO_3^-$			
18 May 2012	$0.02 \pm 0.02$	$0.31 \pm 0.14$	—
29 May 2012	$0.07 \pm 0.01$	$0.03 \pm 0.01$	42.6
02 June 2012	$0.00 \pm 0.00$	$0.11 \pm 0.01$	—
06 June 2012	$0.03 \pm 0.01$	$0.04 \pm 0.02$	39.8
16 June 2012	$0.00 \pm 0.00$	—	—
22 June 2012	$0.01 \pm 0.01$	$0.15 \pm 0.01$	—
$pRONO_2$			
18 May 2012	$0.16 \pm 0.02$	$0.13 \pm 0.08$	—
29 May 2012	$0.16 \pm 0.03$	$0.05 \pm 0.03$	56.9
02 June 2012	$0.09 \pm 0.01$	$0.07 \pm 0.03$	10.5
06 June 2012	$0.09 \pm 0.01$	$0.05 \pm 0.02$	35.7
16 June 2012	$0.15 \pm 0.02$	—	—
22 June 2012	$0.08 \pm 0.01$	$0.06 \pm 0.01$	17.2 <sup>c</sup>

153 <sup>a</sup>Concentrations where total particulate nitrate (p $NO_3$ ) was greater than the detection limit but apportioned  $NO_3^-$  or  
154 p $RONO_2$  was below the detection limit are shown in italics. An apportionment is not possible when p $NO_3$  is below  
155 the detection limit (Day et al., 2022); these cases are marked as missing values.156 <sup>b</sup>Unreported scavenging efficiencies are due to outflow concentrations being greater than inflow concentrations or  
157 the calculated cloud top concentration is less than the outflow concentration (18 May 2012 case for p $RONO_2$ ).158 <sup>c</sup>Uncertain result as the clear-sky concentrations contained missing data in the mid-troposphere that was filled with  
159 the average clear sky concentration between two altitudes.

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