



## Exploring the Mechanisms of Dust Emission and Transport based on Observations and GEOS-Chem Simulations

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### 10 Abstract:

Dust aerosols play a significant role in climate and air quality, yet understanding of their emission and long-range transport mechanisms remains incomplete. By looking into a severe dust event occurred in northern China on April 2025, and conducting the comparative analysis against a 30-year climatic average and the historical dust events using multi-source observations and the GEOS-Chem model simulations, we systematically investigate its meteorological conditions, emission 15 mechanisms, and transport processes. Results show that the dust event in April was originated in the western Inner Mongolia (WIM) source region, accompanied by wind speeds exceeding 8 m/s and hourly PM<sub>10</sub> concentrations above 1900 µg/m<sup>3</sup>, and affected the southern China including Yangtze River Basin and Hainan Province. Under the influence of the Siberian high-pressure system and the Mongolian cyclone, the WIM experienced persistent dry-cold advection (relative humidity around 20%, wind speeds exceeding 10 m/s). Three months preceding the dust event, the WIM exhibited significantly high 20 temperatures (~2 °C), reduced precipitation (~25 mm) and low volumetric soil water (~0.02 m<sup>3</sup>/m<sup>3</sup>). Comparison with two other severe historical dust events in year 2021 and 2023, demonstrating that long-range transport in 2025 was primarily due to strong northerly winds that effectively guided southward transport of dust aerosols. Furthermore, the dust in 2025 consistently moved southward but generally behind the rainband, which imply relatively low wet scavenging and thereby 25 enabling stable long-range transport. The study confirms that persistent drought and strong winds triggered intense dust emission, and airflow transport under specific synoptic conditions dominated the long-range dust transport.

### 1 Introduction

Dust aerosols, which commonly originate from arid and semi-arid regions, significantly reduce visibility (Seinfeld et al., 2004; Tang et al., 2018), directly threat socioeconomic activities and public health (Griffin et al., 2004; Miri et al., 2022), alter regional climate through radiative effects (Twomey et al., 1977; Seinfeld et al., 2004), and by acting as ice nuclei modulate 30 cloud microphysical processes and precipitation (Huang et al., 2006; Wang et al., 2010; Huang et al., 2014; Wang et al., 2015). Furthermore, dust-carried nutrients and pollutants can affect marine ecosystems, soil fertility, and vegetation growth, triggering complex ecological responses (Griffin et al., 2004; Wang et al., 2006; Gao et al., 2009; Gassó et al., 2010; Field et al., 2010). Observational studies indicate that dust aerosols emitted from East Asia can undergo long-range transport to Eastern 35 China (Tan et al., 2012), Japan (Iwasaka et al., 1983), South Korea (Kim et al., 2008), and even North America (Guo et al., 2017), exerting multifaceted impacts on the climate, environment, and economy of these regions.

The Inner Mongolia Autonomous Region of China is one of the most important dust source regions in East Asia (Zhang et al., 2003; Tan et al., 2017). The underlying surface in its western part is primarily desert, while the central part is mainly grassland, providing favorable underlying surface conditions for the occurrence of dust events. Previous studies have indicated that on longer time scales, the frequency of dust events in the Inner Mongolia region is primarily correlated with soil moisture 40 and vegetation changes (Lee et al., 2011; Munkhsetseg et al., 2016; An et al., 2018; Bao et al., 2021). The occurrence of dust



weather normally relates to unstable atmosphere (Idso et al., 1976; Knippertz et al., 2006; Shao et al., 2020) and soil condition in dust source regions (Sun et al., 2001; Tegen et al., 2002). Most dust storms in the Inner Mongolia Autonomous Region occur mainly in spring, due to low vegetation coverage, scarce spring precipitation, and the influence of Mongolian cyclones (Gao et al., 2009; Liu et al., 2015; Borjigin et al., 2024), since strong winds associated with intense Mongolian cyclones create  
45 favorable dynamic conditions for the initiation and development of dust events (Takemi et al., 2005; An et al., 2018; Liu et al., 2024), and low vegetation coverage combined with minimal rainfall leads to dry and loose surface soil, providing ample material for dust emission.

The horizontal and vertical of dust transport depend on the meteorological conditions over the source region and the synoptic features in the downwind areas (Chen et al., 1987; McKendry et al., 2001), as well as the associated dry and wet  
50 deposition processes (Tsai et al., 2008; Liu et al., 2009; Fu et al., 2014; Chen et al., 2017). In dust source regions, the vertical transport of dust aerosols is primarily associated with turbulence and convection (Wang et al., 1990; Park et al., 2015; Richter et al., 2018). Brief bursts of turbulence facilitate the initial entrainment of dust into the atmosphere to become dust aerosols, while sustained turbulence promotes their vertical transport (Zhang et al., 2022). In the Asian Pacific Regional Aerosol Characterization Experiment (ACE-Asia), dust aerosols are typically lifted to approximately 3 km above the source region  
55 (Tsai et al., 2008). Under certain weather conditions, strong upward currents further elevate dust aerosols, thereby facilitating their long-range transport (Tsai et al., 2008). Long-range dust transport from Inner Mongolia relies on the dynamic forcing of northwesterly airflows. Previous studies indicated that under the influence of the Mongolian cyclone, northwesterly flows in the rear of upper-level troughs guide the dust aerosols from Mongolia and Inner Mongolia toward lower latitudes (Liu et al., 2009). Park et al. (2010) found dry deposition was approximately ten times greater than wet deposition near the source area,  
60 whereas wet deposition accounted for over 50% of the total in most downstream marine areas. Xiong et al. (2020) conducted a 20-year dust simulation and confirmed a significant seasonal variability in dust aerosol deposition over East Asia, with wet deposition consistently exceeding dry deposition in all months except December and January. Liang et al. (2022) analyzed an observed dust event in the Inner Mongolia Gobi Desert in March 2021 and found that wet deposition in the downwind North China Plain was approximately twice as effective as dry deposition.

65 In addition, the pathway and strengthens of dust long-range transport remains uncertain. Previous studies suggested that dust aerosols originating from Inner Mongolia primarily affect the North China Plain (Wang et al., 2004). For instance, during several typical dust events in March 2021, dust aerosols from northern China and southeastern Mongolia impacted regions including North China, southern Northeast China, and the northern Huang-Huai area (Yu et al., 2023). Dust aerosols from major Asian source regions (northern China and Mongolia) can be transported to southeastern Asia and the Pacific Ocean  
70 under the influence of westerly and southwesterly winds (Husar et al., 2001; Kai et al., 2007). Moreover, under specific meteorological conditions such as intense frontal cyclones, strong westerly jets, and limited precipitation-dominated wet deposition, dust aerosols from Mongolia and Inner Mongolia can undergo long-range transport to South Korea (Park et al., 2015), and Japan (Tsedendamba et al., 2019), and even North America (Zhao et al., 2008). The dust storm event occurred in western Inner Mongolia on April 11, 2025 is one of most severe dust events in recent years, which unusually penetrated  
75 southward across the Yangtze River Basin, eventually swept southwestern China and the South China Sea, so provides a comprehensive observations and opportunity for researchers to study dust-related processes. This study integrates multi-source observational data with model simulations based on an improved GEOS-Chem model (Tian et al., 2020) to systematically analyze the dust emission characteristics, formation mechanisms, transport pathways, and the primary causes enabling the ultra-long-range transport of this dust event.



## 80 2 Data and measurements

### 2.1 Observational data

The hourly  $PM_{2.5}$  and  $PM_{10}$  concentration provided by the China National Environmental Monitoring Center (<http://www.cnemc.cn/>) over 2,016 monitoring stations including 54 stations in western Inner Mongolia ( $37^{\circ}\text{N}$ - $43^{\circ}\text{N}$ ,  $98^{\circ}\text{E}$ - $112^{\circ}\text{E}$ ) (Figure 1a), and the daily Aerosol Optical Depth (AOD) at 550 nm from MODIS (Moderate-resolution Imaging Spectroradiometer)/Aqua Level 3 Dark Target Deep Blue Combined product (MOD08\_M3, Collection 6.1) with  $1^{\circ}\times 1^{\circ}$  spatial resolution, are employed to look into the dust events during the study period.

Meteorological variables such as near-surface wind speed, temperature, pressure, humidity, and precipitation, along with soil data, are obtained from the ECMWF ERA5 hourly reanalysis dataset from 1940 (<https://cds.climate.copernicus.eu/>), with a spatial resolution of  $0.25^{\circ}\times 0.25^{\circ}$  and temporal resolution of one hour. These data are utilized to analyze the transport characteristics of dust aerosols, the synoptic conditions during dust events, and to conduct comparative analyses against both the 30-year climatological baseline and historical dust cases. Additionally, near-surface temperature and relative humidity from 2,167 meteorological stations (station distribution shown in Figure 1b) operated by the China Meteorological Administration (CMA, <http://data.cma.cn/>), with a temporal resolution of 3 hours, are used to evaluate the model performance. It is noteworthy that, due to the lack of precipitation observations from meteorological stations, ERA5 reanalysis data are used instead to assess the model's simulation of precipitation.

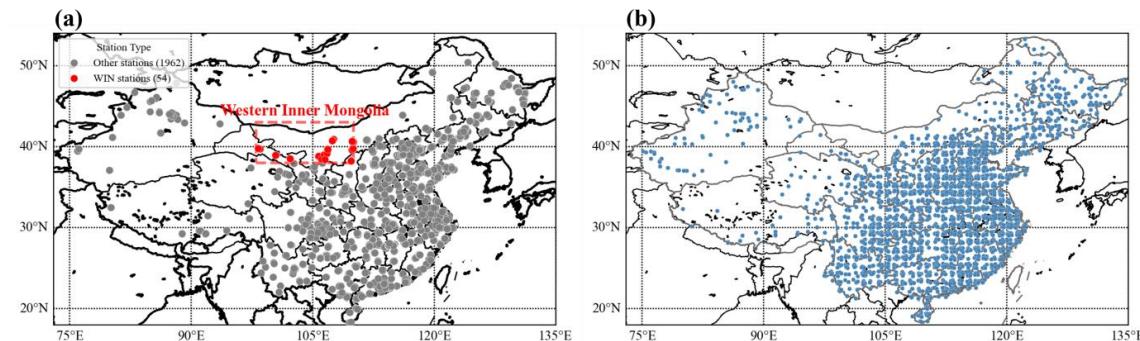


Figure 1. (a) Distribution of air quality monitoring stations in China. The area within the red rectangle ( $38^{\circ}\text{N}$ - $43^{\circ}\text{N}$ ,  $98^{\circ}\text{E}$ - $110^{\circ}\text{E}$ ) illustrates the distribution of stations in WIM. Stations marked in red are located in western Inner Mongolia, while those in gray represent stations in other parts of China. (b) Distribution of Meteorological Stations in China.

### 100 2.2 Model description

The numerical simulations in this study were based on the GEOS-Chem model (<http://acmg.seas.harvard.edu/geos/>). As a global three-dimensional atmospheric chemical transport model, GEOS-Chem possesses the capability to simulate atmospheric components from local to global scales and has been widely applied in various atmospheric composition studies.

This study used GEOS-Chem version 12.6.0, incorporating the dust emission scheme revised by Tian et al. (2020), which introduced spatially varying surface roughness lengths, updated soil property data, and physically based parameterizations of dust emission processes. These modifications enable the model to simulate threshold friction velocity and dust emission fluxes more accurately, reduce biases in simulated  $PM_{10}$  concentrations and aerosol optical depth (AOD), and consequently better capture the spatiotemporal distribution characteristics of dust storms.



110 The meteorological field data used in the model were obtained from the GEOS-FP reanalysis product provided by  
NASA's Global Modeling and Assimilation Office. The simulation period spanned from 00:30 UTC on April 8 to 11:30 UTC  
on April 15, 2025. The model configuration featured a horizontal resolution of  $2^\circ \times 2.5^\circ$ , 72 vertical layers, a global simulation  
domain, and an output time interval of one hour.

### 3 A severe dust storm during 11~14 April, 2025 in Northern China

115 During April 2025 a severe dust storm swept across the most regions of China, from the Northern China to the Southern  
China. As one of the most severe dust storms in recent years, this dust event provides an excellent opportunity to study the  
related dynamical and physical on the dust processes.

120 The PM<sub>10</sub> concentration monitored at surface and AOD from satellite retrievals are used here to represent the dust mass  
concentrations. The spatiotemporal variations of PM<sub>10</sub> and AOD, as well as wind fields during the dust periods are shown in  
Figure 2. It is evident that on April 10, both PM<sub>10</sub> and AOD in China remain quite low, with PM<sub>10</sub> concentrations below 200  
125  $\mu\text{g}/\text{m}^3$  and AOD below 0.4 in most of China. The regionally-averaged PM<sub>10</sub>, PM<sub>2.5</sub>, and the ratio of PM<sub>2.5</sub> to PM<sub>10</sub> in the WIM  
region (Figure 3) remains around 100  $\mu\text{g}/\text{m}^3$ , 20  $\mu\text{g}/\text{m}^3$  and 0.3, respectively before 17:00 BJT (Beijing Time) on April 11th.  
At 17:00 BJT on April 11th, a dust storm started to occur in WIM, accompanied by strong north winds over 8 m/s. It is noticed  
that the hourly PM<sub>10</sub> in WIM is greater than 1900  $\mu\text{g}/\text{m}^3$ , while the ratio of PM<sub>2.5</sub> to PM<sub>10</sub> dropped to approximately 0.2. The  
130 dust was rapidly transported southward, driven by strong northerly winds exceeding 10 m/s. During the subsequent 72 hours,  
the dust aerosols showed distinct spatiotemporal evolution, i.e. On April 12, the dust reached the Yangtze River Basin, where  
the PM<sub>10</sub> concentration exceeded 1000  $\mu\text{g}/\text{m}^3$  and the ratio of PM<sub>2.5</sub> to PM<sub>10</sub> dropped below 0.2. Meanwhile, the AOD in WIM  
exceeded 2 (Figure 2 and 4). By April 13, the dust arrived at Hainan Island, with the PM<sub>10</sub> concentration exceeding 300  $\mu\text{g}/\text{m}^3$   
(Figure 2) and the ratio of PM<sub>2.5</sub> to PM<sub>10</sub> being below 0.3 (Figure 4), marking the completion of long-range transport spanning  
20 degrees of latitude. By April 14, the intensity of the dust had significantly weakened in the affected areas of Southeastern  
135 China, with PM<sub>10</sub> concentration mostly dropping below 300  $\mu\text{g}/\text{m}^3$ , indicating the end of the dust event (Figure 2). In summary,  
this dust event, which originated in the WIM region on April 11 and concluded by April 14, significantly impacted a vast  
expanse of China through southward transport, affecting areas from the north and central to the southern parts of the country.

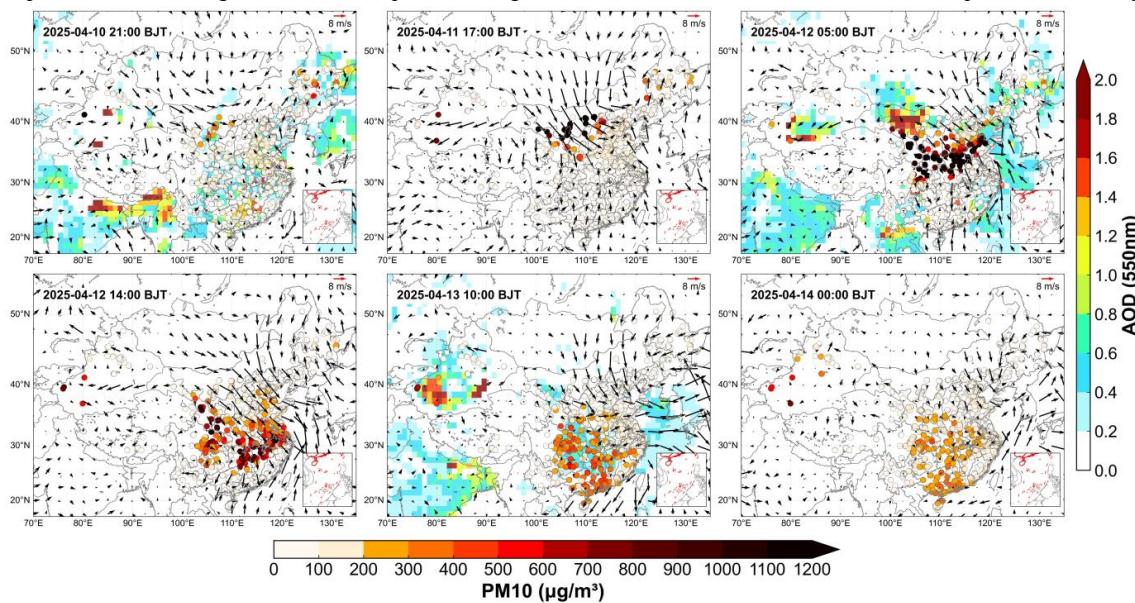




Figure 2. Spatiotemporal variations of  $\text{PM}_{10}$  concentration ( $\mu\text{g}/\text{m}^3$ ) in China, AOD (MODIS) and wind fields (ERA5) over China and its surrounding areas from April 10 to 14, 2025 (BJT). The scatter plot represents  $\text{PM}_{10}$  concentration, and the shaded plot represents AOD.

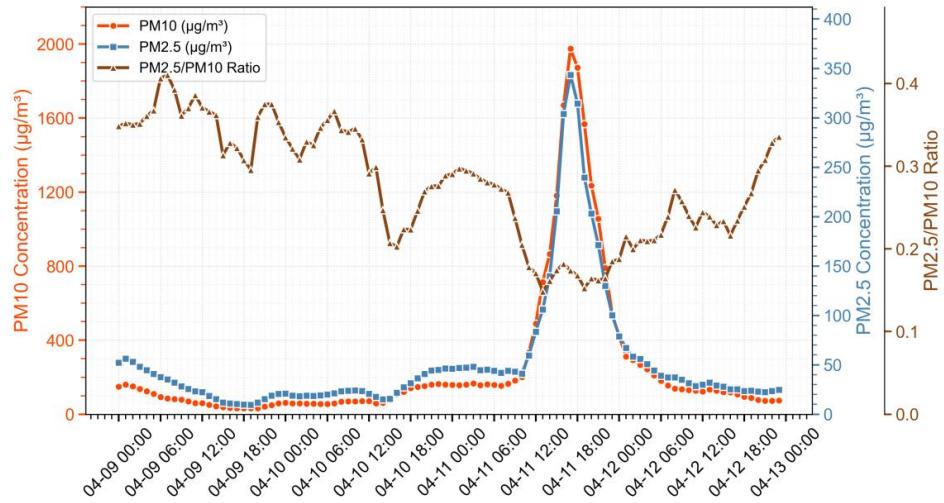
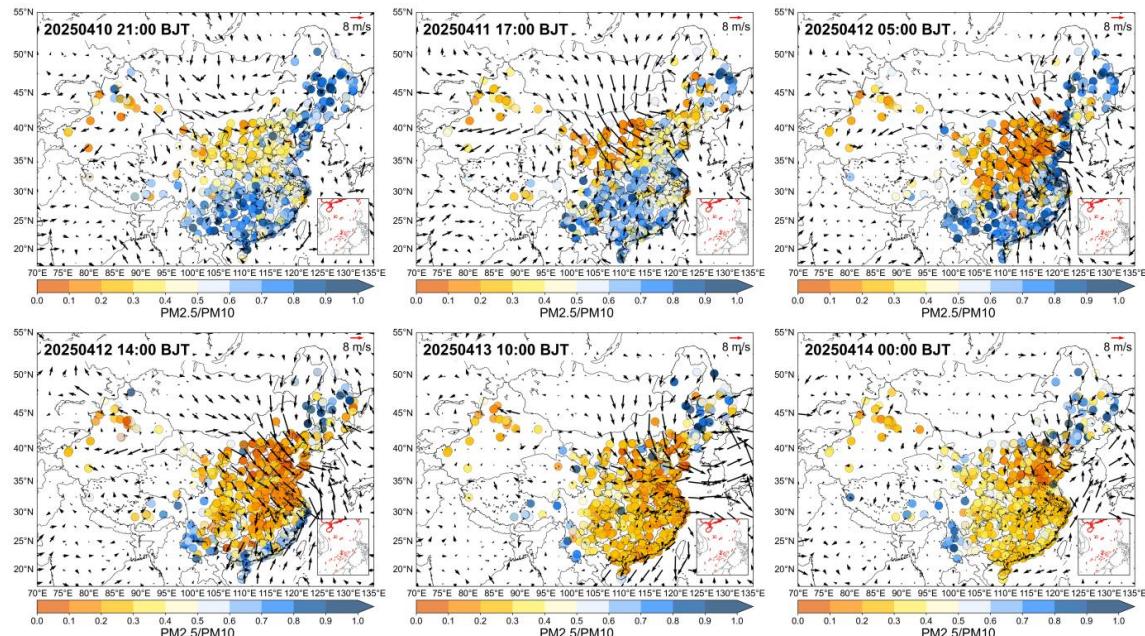


Figure 3. The hourly variations of  $\text{PM}_{10}$ ,  $\text{PM}_{2.5}$  concentrations ( $\mu\text{g}/\text{m}^3$ ) and the  $\text{PM}_{2.5}$ -to- $\text{PM}_{10}$  ratio within the WIM region ( $38^\circ\text{N}$ - $43^\circ\text{N}$ ,  $98^\circ\text{E}$ - $110^\circ\text{E}$ ) during the dust events from April 9 to 12, 2025 (BJT).



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Figure 4. During April 10-14, 2025 (BJT), the spatiotemporal variations of  $\text{PM}_{2.5}$ -to- $\text{PM}_{10}$  ratio and wind fields (ERA5) over China.

It is known that the meteorological conditions including temperature, relative humidity, and wind fields significantly influence dust emissions and transport (Chepil et al., 1956; Zhang et al., 2002; Ravi et al., 2004; Hussein et al., 2006; Zhu et



145 al., 2024; Wang et al., 2021). Temperature and humidity affect dust emissions by modulating soil moisture, surface friability, and vegetation coverage in source regions, while strong winds provide the necessary dynamic forcing for aerosol entrainment into the atmosphere and subsequent long-range transport (Zou et al., 2004; Liu et al., 2004; Xu et al., 2006; Guan et al., 2017). During transport, aerosols are primarily subject to dry and wet deposition (Bergametti et al., 2014). Dry deposition removes the aerosol particles from the atmosphere and toward the surface due to gravitational settling and turbulence or diffusion 150 (Pryor et al., 2004; Bergametti et al., 2014; Ma and von salzen, 2006; Farmer et al., 2021). Wet deposition involves the incorporation of aerosols into hydrometeors (such as raindrops or cloud droplets), followed by their removal through precipitation (Zannetti et al., 1999; Pryor et al., 2004; Bergametti et al., 2014). Since turbulence intensity is often driven by thermal and mechanical energy (Roth et al., 2000; Hu et al., 2007), and droplet formation and fallout depend largely on precipitation, it is essential to consider meteorological factors when discussing aerosol transport.

155 **4 Meteorological factors and dust processes**

**4.1 Synoptic processes**

160 Previous studies found that large-scale circulation systems such as the polar vortex and the westerly jet influence the occurrence and transport of dust by affecting the development of the Mongolian cyclone and the southward movement of cold air masses, thereby regulating the upper-level jet stream and surface wind speeds (Zhao et al., 2004; Yang et al., 2008). In East Asia, Mongolian cyclones are the primary synoptic system responsible for most spring dust events (Li et al., 2022). More than 50% of the dust events occurring in Mongolia and northern China are solely triggered by Mongolian cyclones, while the remainder result from the combined influence of Mongolian cyclones and cold high-pressure systems (Borjigin et al., 2024).

165 For this dust event originating from the WIM region, particular attention should be paid to the role of the Mongolian cyclone and whether a cold Siberian high-pressure system is interacting with it. Figure 5 illustrates the evolution of synoptic conditions during the dust event from April 10 to 13, 2025. On April 10, eastern Outer Mongolia was influenced by a low-pressure system (central pressure was approximately 996 hPa), while the Mongolian cyclone had not yet fully developed. A Siberian high-pressure system (central pressure was exceeding 1028 hPa) was located along the northwestern border of Mongolia. Inner Mongolia was dominated by dry air mass (relative humidity around 30%) and influenced by cold northerly flow from Siberia, with westerly winds in the western region exceeding 10 m/s. By April 11, as the Siberian high (central 170 pressure was exceeding 1036 hPa) moved southward to the northern border of Mongolia, the Mongolian cyclone formed over eastern Inner Mongolia with a central pressure below 996 hPa. The combined influence of these systems generated strong northerly winds exceeding 10 m/s, transporting dry, cold air southward. Relative humidity in Inner Mongolia dropped below 30%, with temperatures approaching 0°C. On April 12, the Mongolian cyclone migrated into northeastern China and intensified (central pressure near 996 hPa), while the Siberian high continued moving southward and weakened, with its 175 central pressure dropping below 1030 hPa. Strong northerly winds persisted over Inner Mongolia and extended to the Yangtze River Basin, facilitating the southward transport of dry, cold air. The low-humidity center shifted to central China (relative humidity around 20%), and temperature in Inner Mongolia fell below 0°C. By April 13, the Mongolian cyclone further intensified, exhibiting a central pressure below 988 hPa. This development resulted in strong northeasterly winds dominating the southeastern coastal regions, which efficiently transported the dry, cold air mass from central China toward the South 180 China Sea, causing relative humidity in southern China to decrease to 40%.

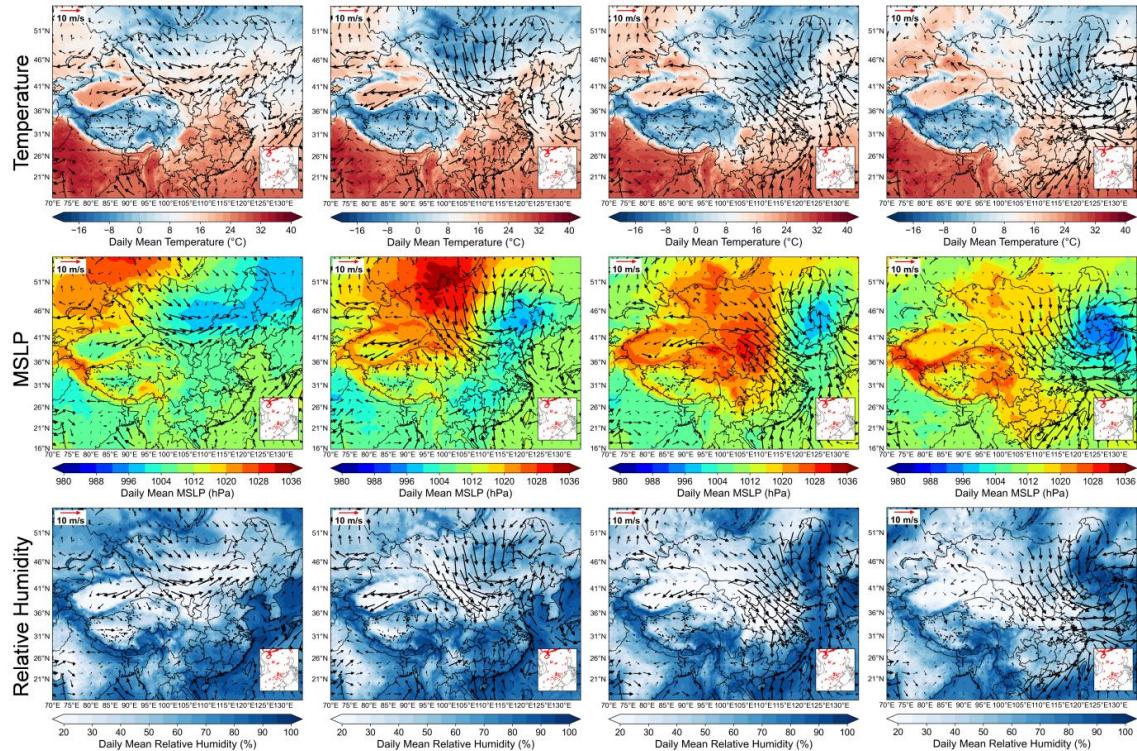


Figure 5. Spatiotemporal variations of daily mean temperature ( $^{\circ}\text{C}$ ), mean sea level pressure (MSLP, unite: hPa), relative humidity (%), and wind fields over China from ERA5 during 10-13 April 2025 (BJT).

Throughout the event, WIM region remained under the influence of dry, cold air advection on the cyclone's western flank.

185 The persistent occurrence of weather conditions characterized by low relative humidity and high wind speeds collectively created an optimal environment for dust transport. Meanwhile, the strong northerly winds generated by the synergistic effect of the Siberian high and Mongolian cyclone effectively transported dust toward the North China Plain (Figure 2). Additionally, the southward extension of dry, cold air masses reaching the Yangtze River Delta provided favorable conditions for the unusual long-range transport of dust to Hainan. In summary, the synergistic interaction between the Siberian high and the Mongolian cyclone served as the primary driving mechanism for this dust event, promoting the southward propagation of dry, cold air masses and enabling the long-range transport of dust aerosols.

#### 4.2 Dust emission and transport

In order to facilitate a more comprehensive discussion of the characteristics of this dust event, model simulations were utilized to the analysis of the dust emission process. The GEOS-Chem model employing a revised dust emission version by 195 Tian et al. (2020), in which the geographical variation of aerodynamic roughness length, smooth roughness length and soil texture, the Owen effect, and the formulation of the sandblasting efficiency  $\alpha$  by Lu and Shao (1999) are incorporated to improve dust emission over China. We performed a systematic model evaluation before utilizing model outputs for analysis of the dust event. The simulated meteorological conditions including temperature, relative humidity, precipitation, and wind fields during this dust event are evaluated against the observations on the weather stations and ERA5 reanalysis data, while 200 the simulated dust mass concentrations are evaluated with observed  $\text{PM}_{10}$  concentration, respectively (Figure S1, S2). Our



evaluation indicated that the model effectively captured the key characteristics of meteorological fields and the spatiotemporal variations of dust aerosols during the April 2025 dust event. Specifically, from April 10 to 11, strong northerly winds (exceeding 8 m/s) dominated northern and central China. A dry-cold air mass intruded into the Inner Mongolia region on April 11, reducing relative humidity to below 30% and temperatures to near 0°C. High dust concentrations were located in western Inner Mongolia, with observed  $PM_{10}$  levels surpassing 950  $\mu g/m^3$  and simulated dust aerosol concentrations exceeding 1000  $\mu g/m^3$ , while precipitation was primarily concentrated in central China. Between April 12 and 13, the model continued to accurately capture the features of temperature, humidity, wind fields, and precipitation, as well as the general southward transport of dust aerosols. In a word, the model reproduced the southward movement of the dry-cold air mass, the transport of dust aerosols driven by northerly winds, and the concurrent southward shift of the precipitation.

Figure 6 shows the temporal variations of dust emission fluxes based on GEOS-Chem simulation. Relatively low dust emissions (less than 150  $\mu g/m^2/s$ ) were present in WIM region at 08:00 BJT on April 11, and dust emission fluxes exceeded 400  $\mu g/m^2/s$  by 16:00 BJT, after which the emissions weakened. It is known that both magnitude and direction of dust transport could be impacted by dust emission fluxes and its vertical distribution. CALIPSO satellite retrievals are widely used to examine the vertical profiles of aerosols including dust (Liu et al., 2008; Uno et al., 2008; Ma et al., 2013; Zhao et al., 2020; Chaibou et al., 2020). Unfortunately, such CALIPSO data are not available since August 1, 2023. In order to look into the vertical distribution of dust and its temporal variations, the simulated vertical distribution of dust mass concentrations along 110°E at selected times from April 11 to 13 are presented (Figure 7). At 16:30 BJT on April 11, dust aerosols were concentrated between 35°N and 40°N, with vertical transport heights exceeding 3.6 km. Subsequently, the dust aerosols transported toward lower latitudes, reaching 30°N by 08:30 BJT on April 12, with transport heights decreasing to approximately 1.6 km. By 04:30 BJT on April 13, the dust aerosols had transported to regions below 25°N, with transport heights further declining to around 0.8 km. Combined with Figure 3, it can be observed that during this dust event, significant dust aerosol emissions occurred in WIM region at 16:30 BJT on April 11, indicating the onset of the dust emission process. The dust aerosols entered the atmosphere and were lifted to altitudes above 3.6 km under the influence of strong winds and other factors. By 17:00 BJT on April 11, the  $PM_{10}$  concentration in WIM region sharply increased, reaching its peak value. Intense dust emissions provided abundant source materials for this dust event, while the lifting of dust aerosols over 3 km in the source region facilitated long-range transport (Idso et al., 1976; Tsai et al., 2008).

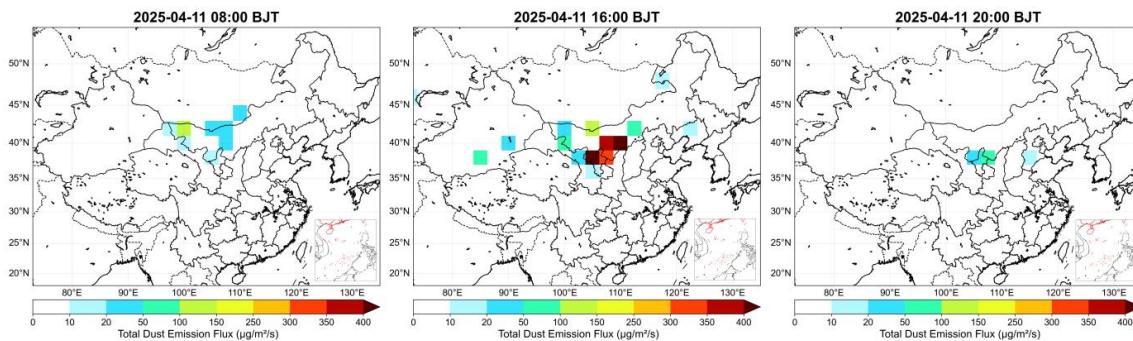
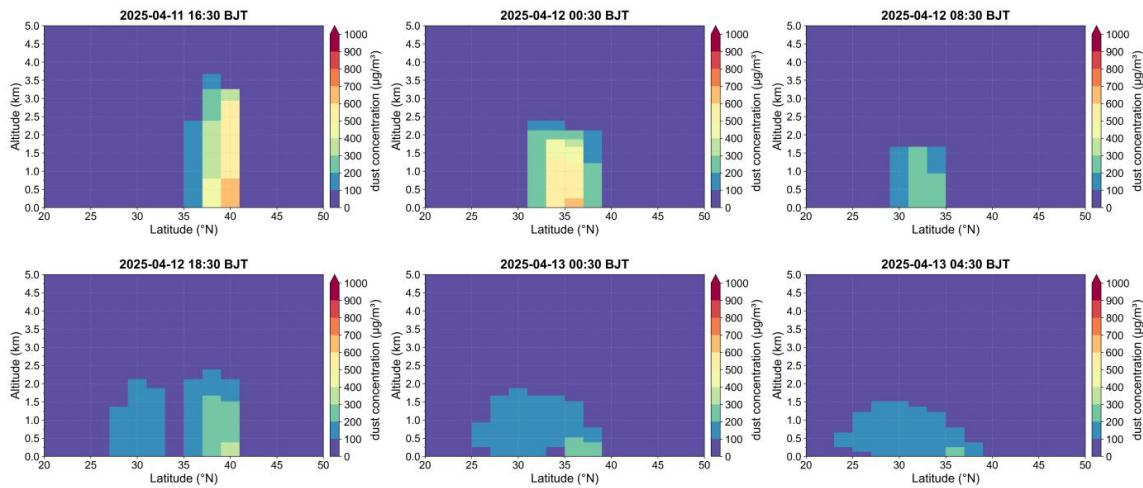


Figure 6. Spatial distribution of model-simulated dust emission fluxes ( $\mu g/m^2/s$ ) on April 11, 2025 (BJT).



230 Figure 7. Vertical cross-section of dust concentrations ( $\mu\text{g}/\text{m}^3$ ) along 110°E between 20°N and 50°N during 11-13 April 2025 (BJT).

#### 4.3 Why is this dust storm so strong?

235 Previous studies found that the frequency of dust events in northern China has decreased in recent years, but the occurrence of severe dust storms has increased (Wang et al., 2018; Bao et al., 2021; Gavrouzou et al., 2021; Zhu et al., 2024).  
 Earlier studies indicated that declining wind speeds, reduced precipitation, and increased vegetation coverage might contribute to the decrease in dust events across China (Guan et al., 2017; Wang et al., 2021). Prolonged drought, high temperatures, and short-term strong winds are key drivers of severe dust storms (Littmann et al., 1991; Attiya et al., 2020; Yong et al., 2021; Wang et al., 2022). The long-range transport of dust aerosols is primarily influenced by meteorological conditions in the source region and removal mechanisms during transport (Chen et al., 1987; Liu et al., 2009; Fu et al., 2014; Chen et al., 2017).  
 240 In East Asia, precipitation-dominated wet deposition might play a more critical role than dry deposition in the long-range transport of dust aerosols (Zhao et al., 2003; Park et al., 2010; Liang et al., 2022). Therefore, it is essential to consider the long-term influences of humidity, precipitation and temperature, the short-term dynamic effects of wind speed, and the impact of precipitation-dominated wet deposition when analyzing the driving mechanisms and causes of long-range transport in this dust event.

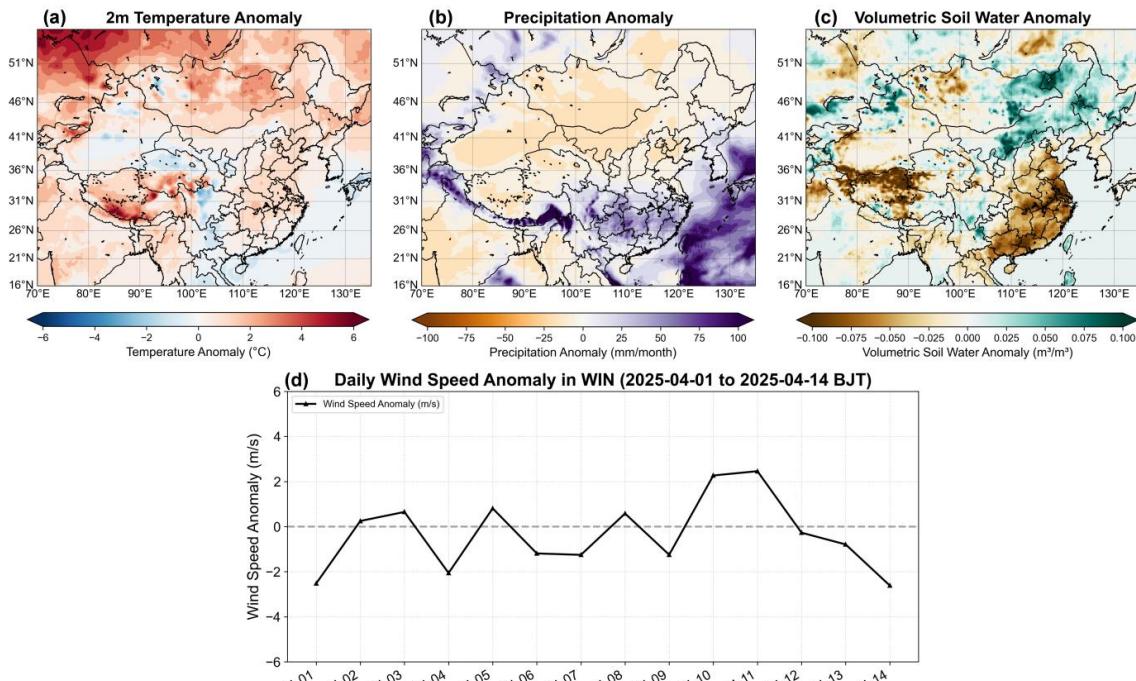
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##### 4.3.1 Comparison of meteorological conditions with 30-year climatological mean

To understand the driving role of meteorological factors in this dust event, we analyzed the anomalies of temperature, volumetric soil water, and precipitation during the four months preceding the dust event. These meteorological anomalies are relative to the 30-year climatological baseline. The spatial distribution of temperature, volumetric soil water, and total precipitation anomalies from January 1, 2025, to April 14, 2025 (Figure 8a-c) indicates that relative to the 30-year climatological baseline, the WIM region (40-45°N, 95-105°E) exhibited higher-than-normal temperatures (Approximately +2 °C), lower-than-normal precipitation (Approximately -25 mm), and lower-than-normal volumetric soil water (Approximately -0.02  $\text{m}^3/\text{m}^3$ ) both during and for the three months preceding the dust event. Moreover, it is essential to focus on wind speed anomalies on shorter time scales immediately preceding the dust event. Figure 8d shows the daily mean wind speed anomalies in the WIM region from April 5 to 14, 2025. Approximately +2 m/s wind speed anomalies occurred before the dust outbreak (April 10), and the daily mean wind speed anomaly peaked on the day of the dust event (April 11). To



summarize, compared to the climatological baseline, the WIM region experienced generally warm and dry conditions from January to April 2025, accompanied by pronounced short-term strong winds preceding the dust event. These factors served as critical drivers of this intense dust episode, providing the necessary material and dynamic conditions for dust emission.



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Figure 8. Spatial distribution of mean anomalies for temperature ( $^{\circ}\text{C}$ ), total precipitation (mm) and top-layer volumetric soil water ( $\text{m}^3/\text{m}^3$ ) over China from January 1, 2025 to April 13, 2025 BJT relative to the 30-year baseline of 1995-2024 (a-c). Daily anomaly time series of wind speed (m/s) from April 1 to 14, 2025 BJT (Relative to 30-year baseline: 1995-2024) in WIM (d).

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#### 4.3.2 Comparison with historical severe dust storms

In order to further explore the dynamical and physical processes during this dust event, the historical dust events in northern China since 2000 are collected and compiled, and the two most severe dust events (one on March 14, 2021, and the other on April 9, 2023) are selected for comparison. Similar to the 2025 event, these historical events originated from source areas including WIM region and were primarily driven by Mongolian cyclones (Mikalai et al., 2022; Mikalai et al., 2024). However, neither historical event reached Hainan Province (Figure S3). By looking into the meteorological condition and soil moisture, to compare and understand the differences in the dust emission, transport, and scavenging processes for the selected dust events.

To compare the intensity of the three dust events occurring in March 2021, April 2023, and April 2025, the regional averages of both daily maximum  $\text{PM}_{10}$  concentrations in the WIM region during these events were analysed. As shown in Figure 9, the peak regional average daily maximum  $\text{PM}_{10}$  concentrations for the 2021, 2023, and 2025 dust events were  $749 \mu\text{g}/\text{m}^3$ ,  $708 \mu\text{g}/\text{m}^3$ , and  $841 \mu\text{g}/\text{m}^3$ , respectively. This indicates that all three events were relatively intense, with no significant difference in their outbreak intensity. The meteorological conditions including precipitation, wind speed and surface temperature, as well as soil moisture in the dust source regions strongly influence dust emissions (Ishizuka et al., 2008; Kim



280 et al., 2015; Yang et al., 2019). To further investigate the causes of the three dust events, Figure 10 shows the time series of  
 precipitation, volumetric soil water, surface temperature for the four months preceding the three dust events, as well as the  
 daily mean wind speed for the five days before and after the dust events in the WIM region. The data indicate that during the  
 four months prior to the dust events, there was no significant difference in surface temperature between 2025 and the years  
 285 2021 or 2023. However, in the month when the dust event occurred, the surface temperature in 2025 was higher than during  
 both historical events (approximately 8 K higher than in 2021 and 4 K higher than in 2023). Precipitation levels during the  
 four months preceding the three events were relatively similar. However, in the month when the dust events occurred, the  
 precipitation in 2025 (approximately 10 mm) was significantly lower than that in the two historical events (30 mm in 2023  
 and 20 mm in 2021). Correspondingly, the soil moisture in the month of the 2025 dust event (approximately  $0.1 \text{ m}^3/\text{m}^3$ ) was  
 significantly lower than that during the months of the two historical events (approximately  $0.15 \text{ m}^3/\text{m}^3$ ). Furthermore, on the  
 290 day of the dust outbreak, the daily mean wind speed of the 2025 dust event (approximately 3.5 m/s) was significantly higher  
 than that of the two historical events (approximately 2 m/s).

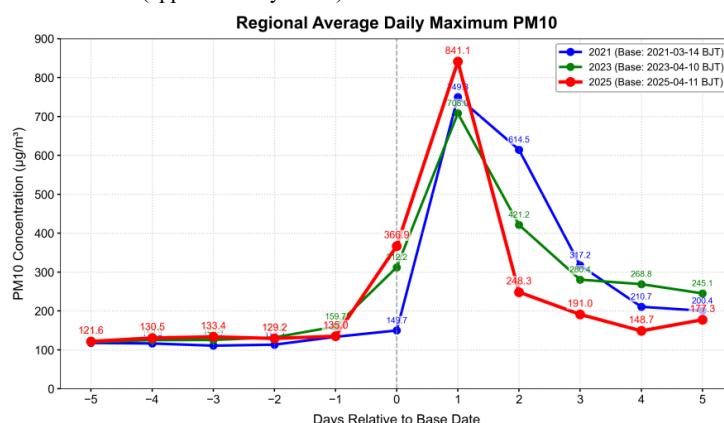


Figure 9. The temporal evolution of regionally averaged  $\text{PM}_{10}$  concentrations (daily maximum, unite:  $\mu\text{g}/\text{m}^3$ ) from monitoring stations during dust events in March 2021, April 2023, and April 2025. Data are plotted relative to the dust event onset date  
 295 (day 0) for each year, with  $\pm 5$ -day analysis windows. Gray dashed line marks the dust onset day (day 0).

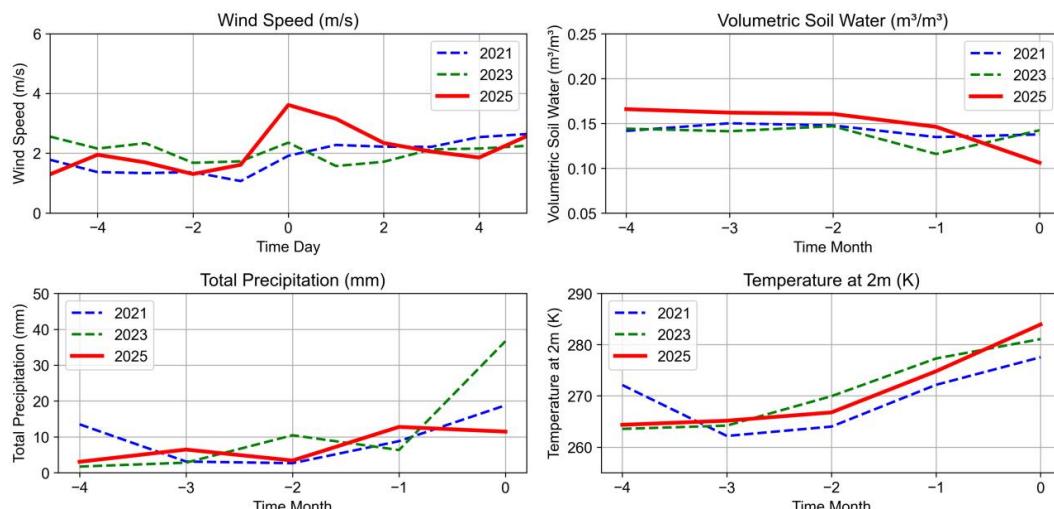
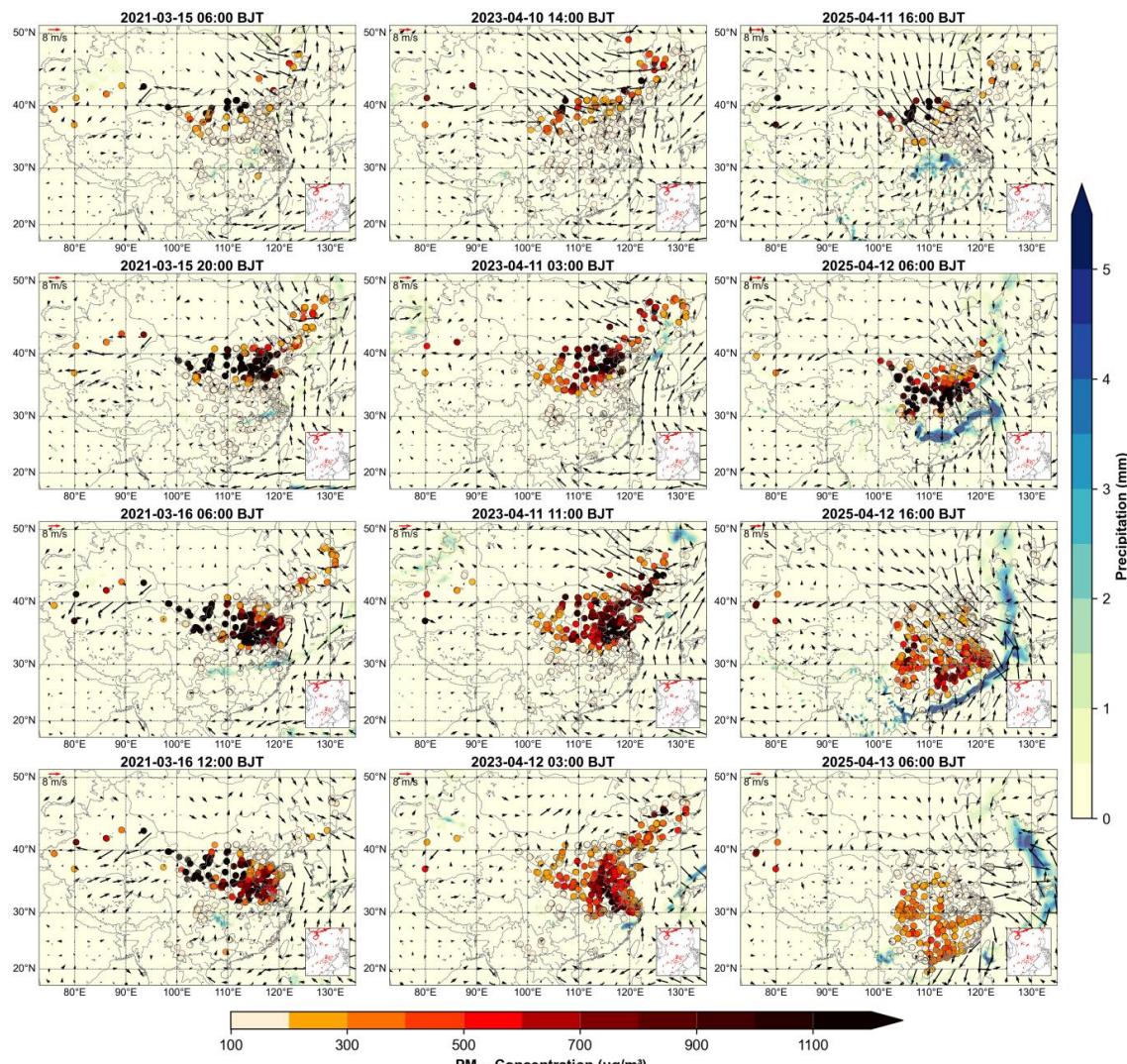




Figure 10. Daily time series of wind speed (m/s) in western Inner Mongolia (5 days pre- and post-dust event onset). Monthly time series of key parameters in western Inner Mongolia (4 months pre-dust event onset), including top-layer volumetric soil water ( $\text{m}^3/\text{m}^3$ ), total precipitation (mm), surface temperature (K).

300 Figure 11 presents the spatial-temporal variations of the  $\text{PM}_{10}$  mass concentration, wind field, and precipitation during three dust event periods. It is clearly shown that although strong dust emissions occurred in WIM for three dust events (first row), the magnitude and pathway of dust transport exhibits quite different behaviour (second-fourth rows). For the case in March 2021, strong northerly winds pushed dust particles to central-eastern China on Mar 15, and then remain stagnant and had no further southern-ward transport until March 16 due to quite weak wind speed. For the case in April 2023, the wind 305 speed was apparently stronger than that in 2021 case, so dust particles could be transported to eastern China and even northeast China. The lack of sustained strong northerly winds over the North China Plain on April 11-12 limited the southward spread of dust to southern China. Different from two above cases, dust aerosols in April 2025 reached the Yangtze River Basin, and sustained strong northerly winds exceeding 8 m/s provided sufficient momentum and make dust particles crossed the Yangtze River and further transport southward to the South China Sea. Wet scavenging of dust particles are negligible due to very 310 limited precipitation for the 2021 and 2023 cases. A rainbelt over eastern China was observed during the 2025 case, but dust southward transport was not affected wet scavenging due to precipitation since dust aerosols consistently remained behind the rainbelt. Therefore, the reason why the dust storm during April 2025 can be transported to the most southern China is mostly attributed to favourable meteorological conditions, i.e. strong and persistent northerly winds, for dust emission and long-range transport, and combined with weak wet scavenging by precipitation.



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Figure 11. Spatial distribution of hourly accumulated precipitation (mm), wind fields, and  $\text{PM}_{10}$  concentrations ( $\mu\text{g}/\text{m}^3$ ) over China during dust events: March 2021, April 2023, and April 2025. Precipitation is shown as color-filled maps, with  $\text{PM}_{10}$  displayed as scatter plots.

## 5 Summary and Conclusion

320 We analyze the basic characteristics of the outbreak and transport processes of a severe dust event that occurred in northern China on April 11, 2025, especially focus on its emission and vertical transport, and then investigate the dust emission mechanisms and transport pathways involved by comparing meteorological conditions against the 30-year climatological average and contrasting the event with two typical historical dust storms.

325 Starting from 17:00 Beijing Time on April 11, 2025, a dust storm occurred in WIM accompanied by strong northerly winds exceeding 8 m/s, and the hourly  $\text{PM}_{10}$  concentration surged to over  $1900 \mu\text{g}/\text{m}^3$ . The GEOS-Chem model simulations



indicate that the emission flux exceeded  $400 \mu\text{g}/\text{m}^2/\text{s}$ , and vertical transport reached 3.6 km. Under the synergistic effect of the Siberian cold high-pressure system and the Mongolian cyclone, WIM remained within the dry-cold advection zone west of the cyclone. The combination of low relative humidity and wind speeds over 10 m/s provided favorable dynamic conditions for dust emission. Furthermore, strong northerly winds in the WIM region guided the dust toward the North China Plain. The 330 dust reached the Yangtze River Basin on April 12 and arrived at Hainan Island by April 13, where observed  $\text{PM}_{10}$  exceeded  $200 \mu\text{g}/\text{m}^3$  and simulated dust concentration surpassed  $20 \mu\text{g}/\text{m}^3$ . During this event, dust aerosols originating from WIM were transported an exceptionally long distance to Hainan Province.

The comparisons of meteorological conditions during this dust event against the 30-year climatological mean from ERA5 reanalysis data indicate that the dust source region (WIM) exhibited persistent high-temperature and drought conditions during 335 the four months preceding the event, i.e. air temperature was approximately  $+2^\circ\text{C}$  above the baseline, precipitation was about 25 mm below normal, and volumetric soil water content was approximately  $0.02 \text{ m}^3/\text{m}^3$  lower than the baseline.. Short-term wind field analysis before dust emission showed a positive wind speed anomaly of approximately  $+2 \text{ m/s}$  on the day before the outbreak (April 10), which peaked on the dust event day (April 11). Prolonged high temperature and drought led to desiccated and loosened surface soil, providing ample source material for dust emission, while the strong winds preceding the 340 event supplied crucial dynamic conditions. The synergistic interaction of these factors constituted the core triggering mechanism for this dust event.

To further discuss the reasons causing the long-range transport during this dust event, we selected two typical historical dust events in northern China (March 14, 2021, and April 9, 2023) for comparative analysis. Although all three events originated from the same source regions including WIM and were primarily driven by Mongolian cyclones, the 2025 event 345 achieved ultra-long-range transport to Hainan, while the historical events did not affect southern China. Comparison of the daily maximum  $\text{PM}_{10}$  concentrations in the WIM region revealed values of  $749 \mu\text{g}/\text{m}^3$  (2021) and  $708 \mu\text{g}/\text{m}^3$  (2023), whereas the 2025 event reached  $841 \mu\text{g}/\text{m}^3$ , indicating slightly higher intensity. Analysis of dust emission conditions for the three events based on ERA5 reanalysis data showed that the volumetric soil water content during the month of the 2025 event (approximately  $0.1 \text{ m}^3/\text{m}^3$ ) was significantly lower than during the months of the historical events. This is primarily attributed to the much lower precipitation in April 2025 (near 10 mm) compared to April 2023 (over 30 mm) and March 2021 (about 20 350 mm), while temperature, humidity, and precipitation anomalies over the preceding four months showed no significant differences among the three events. Meanwhile, the regional average wind speed on the day of the 2025 event (close to 4 m/s) was significantly stronger than during the historical events.

To look into long-range transport conditions, the  $\text{PM}_{10}$  concentrations, and wind fields and precipitation patterns during 355 the emission and transport periods of the three dust events were compared. Our study found that both the 2021 and 2023 events were limited by the lack of persistent strong northerly winds along their transport paths (over North and Central China), ultimately preventing long-range transport to southern China. In contrast, the 2025 event exhibited two key features: (1) dust aerosols consistently remained behind the rainband and moved southward synchronously with it, effectively avoiding wet scavenging; and (2) sustained strong northerly winds exceeding 8 m/s continuously pushed the dust over the Yangtze River 360 Basin, enabling ultra-long-range transport to Hainan.

In summary, this study demonstrates that the southward movement of the dry-cold air mass driven by persistent strong northerly winds, combined with the coordinated dust-rainband configuration, collectively facilitated the stable long-range transport of dust in the 2025 event. These results highlight the critical leading role of airflow transport under specific synoptic conditions.

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#### Author contributions

PZ was responsible for data analysis and drafted the manuscript. XM participated in the conceptualization of the study,  
data analysis, and article writing. RT contributed to the model simulations and, together with JZ, TY, and YK, assisted in the  
375 manuscript preparation.

#### Competing interests

The authors declare that they have no conflict of interest.

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