

# Insights on Ozone Formation Sensitivity in Southeast and East Asian Megacities during ASIA-AQ

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## 35 **Abstract.**

Controlling ozone (O<sub>3</sub>) in rapidly urbanizing megacities in Southeast and East Asia remains a challenge. O<sub>3</sub> is a secondary pollutant formed through nonlinear photochemical reactions with its precursors: nitrogen oxides (NO<sub>x</sub>) and volatile organic compounds (VOCs). Observation-based quantification of precursor sensitivity remains scarce, limiting actionable O<sub>3</sub> control. To address this, we leverage airborne observations from the NASA DC-8 during the ASIA-AQ campaign conducted in February and March 2024 across four Asian megacities: Metro Manila, the Seoul Metropolitan Area, the Tainan-Kaohsiung

Metropolitan Area, and the Bangkok Metropolitan Region. These extensive measurements of various trace gases were used to constrain a zero-dimensional box model and estimate the net production rates of  $O_x$  ( $PO_x$ ,  $O_x = O_3 + NO_2$ ). Precursor sensitivity regimes were characterized for each megacity by generating isopleths of  $PO_x$  across varying levels of  $NO_x$  and VOCs. The analysis revealed that Manila and Tainan-Kaohsiung exhibited predominantly  $NO_x$ -sensitive conditions, favoring  $NO_x$  reduction as an effective  $O_3$  mitigation strategy, while Bangkok showed a more mixed sensitivity, suggesting combined  $NO_x$  and VOC reductions. In contrast, Seoul, ~~under colder and low solar irradiance conditions~~, exhibited a primarily VOC-sensitive regime associated with its higher  $NO_x$  conditions relative to the other cities, underscoring the importance of VOC-focused strategies. In addition, to quantitatively assess sensitivity transitions, we computed orthogonal distances from the isopleth transition boundaries for all four study areas. Diurnal analyses of these distances revealed a shift from more VOC-sensitive conditions in the morning toward more  $NO_x$ -sensitive regimes in the afternoon. These findings provide critical insights for formulating effective, city-specific  $O_3$  control policies in urban environments.

## 1 Introduction

Megacities in Southeast and East Asia, such as those in the Philippines, Thailand, Taiwan, and South Korea, have undergone rapid urbanization and industrialization in recent decades, leading to severe air quality issues. These Asian megacities are characterized by high population densities, diverse emission profiles, and unique meteorological conditions associated with the timing of monsoon dynamics and transboundary transport of pollutants, making their air pollution issues complex and multifaceted (Chin et al., 2017). High levels of primary and secondary pollutants, particularly ozone ( $O_3$ ), in these densely populated areas pose significant risks to public health and agricultural production (Dörter et al., 2022; Gaudel et al., 2018; Gu et al., 2024; Wang et al., 2022b).

Surface  $O_3$  is produced through the photochemical reactions of nitrogen oxides ( $NO_x$ ) and volatile organic compounds (VOCs) involving radical propagation reactions (Finlayson-Pitts and Pitts Jr., 2000). The net production of  $O_3$  is driven by the reactions of nitric oxide (NO) with peroxy radicals formed from the oxidation of VOCs by hydroxyl radicals (OH), leading to the formation of nitrogen dioxide ( $NO_2$ ), which subsequently photolyzes to produce  $O_3$ .

The balance between the mixing ratios of the precursors, VOCs and  $NO_x$ , alongside meteorological conditions such as wind, temperature, humidity, and sunlight intensity, influences  $O_3$  mixing ratios. Particularly, the nonlinear dependence of  $O_3$  formation on its precursors causes  $O_3$  to respond differently to changes in VOC and  $NO_x$  emissions depending on local atmospheric conditions, leading to either  $NO_x$ -sensitive or VOC-sensitive regimes (Cohan et al., 2005; Liu and Shi, 2021; Shen et al., 2021). In  $NO_x$ -sensitive areas, reducing  $NO_x$  emissions leads to a marked decrease in  $O_3$  production. In VOC-sensitive regimes (also referred to as  $NO_x$ -saturated), reductions in VOC emissions are effective at controlling  $O_3$  levels in the short term. If VOC mixing ratios are reduced in a  $NO_x$ -sensitive regime (or vice versa)  $O_3$  production can remain constant or even increase rather than decrease, causing a counterproductive effect. It is therefore critical to identify the prevailing sensitivity regime, while recognizing that sustainable long-term  $O_3$  mitigation ultimately requires reductions in both  $NO_x$  and VOCs

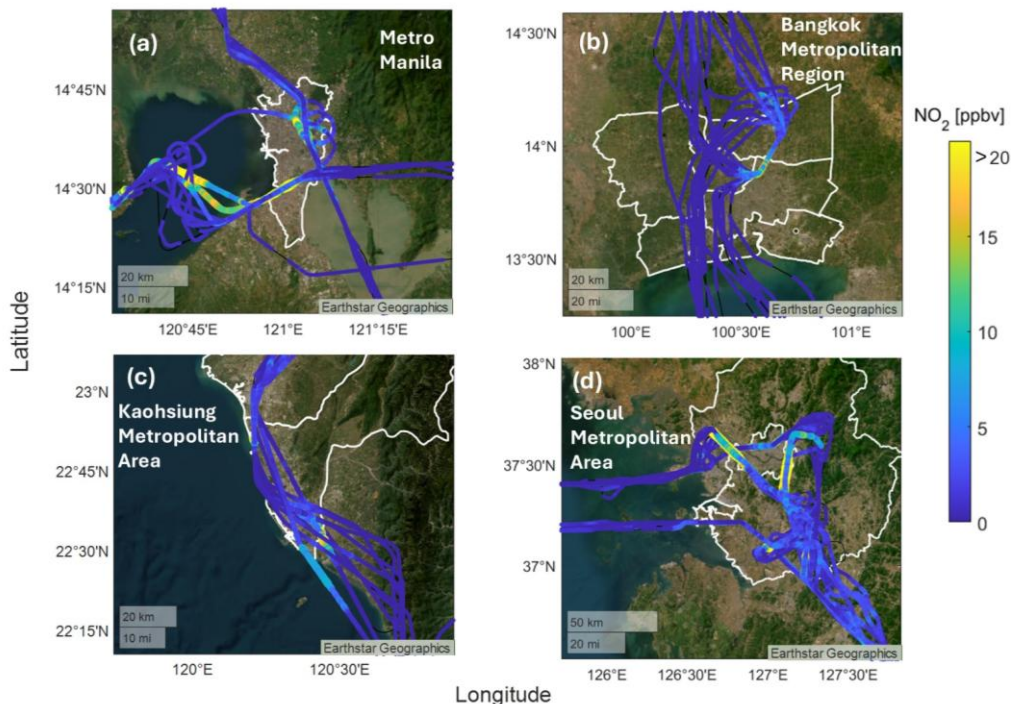
emissions. Between these two regimes, transitional conditions may occur when precursor levels lie near the ridge line separating NO<sub>x</sub>- and VOC-sensitive regimes. Under such conditions, ozone formation sensitivity can shift as precursor levels and photochemical conditions evolve, even on local and diurnal scales (Mazzuca et al., 2016; Guo et al., 2021; Tan et al., 2018; Stockwell et al., 2025). This nonlinearity makes it more difficult to navigate and inform the direction of emission reduction strategies (Ojha et al., 2022; Yarwood et al., 2013). Therefore, a thorough understanding of the chemical sensitivity of O<sub>3</sub> formation to these precursor levels is essential for policy makers to develop effective control strategies.

The complexity of O<sub>3</sub> formation, due to nonlinear relationships with precursor levels, has been the subject of extensive research. Three-dimensional (3-D) chemical transport models (CTMs) have been widely used to simulate O<sub>3</sub> formation by combining emission inventories, meteorological conditions, and chemical processes (Wang and Jacob, 1998; Lee et al., 2014; Li et al., 2020; Dang et al., 2023). Although CTMs can simulate large spatial and temporal scales, allowing for testing of emission reduction scenarios and their effects on air quality over extensive regions, the accuracy of model evaluations is often limited by large uncertainties in boundary layer dynamics, emission inventories, and chemical mechanisms (Choi and Souri, 2015; Liu and Shi, 2021). Satellites have become an increasingly important tool for studying O<sub>3</sub> precursors across wide geographic areas. Formaldehyde (HCHO) and NO<sub>2</sub> are often used as proxies for VOCs and NO<sub>x</sub>, respectively (Tonnesen and Dennis, 2000), and the ratio of HCHO/NO<sub>2</sub> is widely used to diagnose whether O<sub>3</sub> production is in a NO<sub>x</sub>- or VOC-sensitive regime (Martin et al., 2004; Duncan et al., 2010; Jin and Holloway, 2015; Souri et al., 2017; Jin et al., 2017; Ren et al., 2022). However, using HCHO as a ~~substitution~~ for total VOC reactivity carries intrinsic uncertainty, since direct HCHO emissions and secondary production may not fully capture the diversity of reactive VOC species. This limitation is further compounded when the ratio is derived from satellite retrievals, where both retrieval errors and the reliance on models to relate total column abundances to near-surface conditions can introduce discrepancies (Schroeder et al., 2017; Souri et al., 2020, 2023). In addition, the HCHO/NO<sub>2</sub> ratio can vary substantially between environments, so regime classifications based on a single threshold value should be treated with caution (Schroeder et al., 2017; Roozitalab et al., 2022). Therefore, to minimize uncertainties in O<sub>3</sub>-formation sensitivities, various ground-based observations (Sillman and He, 2002; Chen and Brune, 2012; Chen et al., 2022; Wang et al., 2022a) and airborne campaigns (Mazzuca et al., 2016; Schroeder et al., 2017, 2020; Womack et al., 2019; Nussbaumer et al., 2024) have been performed with comprehensive sets of measurements of O<sub>3</sub> and its precursors. Ground-based measurements provide continuous monitoring of pollutant levels over time, while airborne observations offer detailed vertical profiles and capture spatial variability, bridging the gap between total column satellite measurements and near-surface conditions.

In this study, we investigate the precursor sensitivities of the chemical production rates of ozone (PO<sub>x</sub>; O<sub>x</sub> = NO<sub>2</sub> + O<sub>3</sub>) in four Asian megacities during the Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ) campaign from February to March 2024. Using comprehensive airborne measurements from the NASA DC-8 aircraft, we constrain a zero-dimensional chemical box model to estimate the PO<sub>x</sub>. Isopleths of PO<sub>x</sub> across a varying range of NO<sub>x</sub> and VOC levels are used to show the sensitivity regimes of ozone formations in each megacity to their precursors. These results offer a scientific basis for informing region-specific precursor control strategies across different urban atmospheres in Asia.

## 2 Methods

### 2.1 The ASIA-AQ Campaign



**Figure 1.** Flight tracks (Thin black lines) during the ASIA-AQ campaign over four major Asian megacities: (a) Metro Manila, Philippines; (b) Bangkok Metropolitan Region, Thailand; (c) Tainan-Kaohsiung Metropolitan Area, Taiwan; and (d) Seoul Metropolitan Area, South Korea. The flight tracks are colored by the observed NO<sub>2</sub> mixing ratios from the DC-8. White boundaries indicate the administrative or urban extents of each megacity. Basemap source: Esri, Maxar, Earthstar Geographics, and the GIS User Community (Esri World Imagery).

The ASIA-AQ campaign, conducted in February and March 2024, targeted specific regions across major urban  
110 centers in Southeast and East Asia: the Philippines (6–15 February), South Korea (17 February–11 March), Thailand (16–25  
March), and Taiwan (15 February, 13 and 27 March). The timing of the campaign was determined by the broader scientific  
objectives of ASIA-AQ as described in the ASIA-AQ white paper ([https://espo.nasa.gov/asia-aq/document/ASIA-AQ\\_White\\_Paper](https://espo.nasa.gov/asia-aq/document/ASIA-AQ_White_Paper)).

Briefly, for tropical and subtropical the Southeast Asian regions (Philippines, and Thailand, and Taiwan), this period  
115 corresponds to the pre-monsoon season, characterized by a relatively dry conditions and warm season with enhanced solar  
radiation, which are favorable for active ozone formation photochemistry (Toh et al., 2013; Marvin et al., 2021). In Taiwan,  
the campaign period represents a transitional phase prior to the peak ozone season, during which photochemical activity is  
already significant (Lin et al., 2014). In contrast, South Korea was in its late winter season, outside the typical peak ozone  
season, with generally reduced photochemical activity with peak aerosol pollution. Lower temperatures and weaker solar  
120 radiation limited photochemical activity compared to the tropical other regions (Lee and Park, 2022; Kim et al., 2023). Despite

these seasonal differences, the campaign sampled a broad range of NO<sub>x</sub> and VOC conditions across the studied regions, enabling characterization of ozone production sensitivity across diverse chemical environments.

We ~~performed~~ conducted a total of 16 research flights with the NASA DC-8 aircraft, including 4 flights over the Philippines, 5 over South Korea, 4 over Thailand, and 3 over Taiwan (Table S1). The DC-8 was equipped with 26 state-of-the-art instruments to measure over 100 trace gases, aerosol properties, and meteorological parameters (Table S2). The full ASIA-AQ dataset has been archived and is publicly available from the NASA LaRC Airborne Science Data for Atmospheric Composition (Project DOI: 10.5067/SUBORBITAL/ASIA-AQ/DATA001). ~~In addition, detailed descriptions of the campaign can be found in the ASIA-AQ white paper (<https://espo.nasa.gov/asia-aq/document/ASIA-AQ-White-Paper>).~~

This study focuses on the data collected during the flights over four Asian megacities: Metro Manila (MM), the Seoul Metropolitan Area (SMA), the Tainan-Kaohsiung Metropolitan Area (TKMA), and the Bangkok Metropolitan Region (BMR), shown in Fig. 1. It is important to note that measurements in TKMA were conducted during transit flights between other country deployments, therefore early morning sampling was not performed, and flight dates over TKMA were sparser compared to other cities (see Sect. 3.1). The DC-8 regularly executed missed-approaches, in which the aircraft flew over several regional airport runways without landing to obtain vertical profiles of trace gases from a few tens of meters to several kilometers. The flight paths were centered on each megacity, ensuring that the aircraft flew over the city center at least twice per flight pattern. This pattern was repeated two to three times during each research flight at different times of day (typically morning, midday, and/or afternoon) to identify key patterns in chemical evolution of pollutants as a function of both altitude and time of day.

## 2.2 Box Model Setup

We used the Framework for 0-D Atmospheric Modeling (F0AM v4.3; Wolfe et al., 2016) to simulate PO<sub>x</sub> and examine the sensitivity of ozone formation to the changes in precursor mixing ratios (e.g., NO<sub>x</sub> and VOCs) constrained by the airborne observations from the DC-8 (Table S2).

The model was constrained using near-surface observations below 750 m (pressure altitude above sea level) over each megacity ~~averaged within each flight track (typically spanning ~1–2 hours per track)~~, including trace gases (e.g., O<sub>3</sub>, CO, NO<sub>x</sub>, H<sub>2</sub>O, and VOCs), photolysis frequencies, and meteorological conditions such as temperature and pressure (Table S2). The altitude threshold was chosen to focus on air masses most representative of near-surface conditions typically influenced by local emissions. Inspection of precursor vertical profiles (Fig. 2) and sensitivity tests with different altitude cuts (see Supplementary Information (SI) Sect. S2 and Table S3) supports 750 m as a sensible and conservative choice that keeps the analysis centered on boundary-layer chemistry.

For gas-phase chemical reactions, the Master Chemical Mechanism (MCMv3.3.1; Jenkin et al., 2015), a near-explicit chemical mechanism widely used for atmospheric box modeling, was applied to represent VOC oxidation chemistry and its role in ozone formation. Processes not represented in the current gas-phase mechanism (e.g., heterogeneous radical uptake or

updated halogen chemistry) may influence radical budgets under certain conditions and therefore represent a potential source of uncertainty.

155 To achieve a steady-state representation of radical species and other unmeasured compounds, the concentrations of these species were allowed to evolve over a three-day period, using the diurnal steady-state mode. In this configuration, the measured photolysis frequencies were adjusted according to the evolving solar zenith angle every 20-minute intervals. This approach yielded results within  $\pm 5\%$  of those obtained using the convergence mode, where simulations are continued until each species stabilizes within 1% (Schroeder et al., 2020; Brune et al., 2022; Nault et al., 2024; Lee et al., 2026). In addition, 160 the first-order physical loss lifetime for dilution is set to 12 hours, similar to studies from the Korea-U.S. air quality (KORUS-AQ) field campaign using the DC-8 over South Korea (Brune et al., 2022; Nault et al., 2024).

For the calculation of  $\text{PO}_x$ , the model-estimated turnover rates of reactions contributing to the production or destruction of  $\text{O}_x$  are summed up from the model output following Eq. (1):

$$165 \text{PO}_x = k_{\text{NO}+\text{HO}_2}[\text{NO}][\text{HO}_2] + \sum(1 - \alpha_i)k_{\text{NO}+\text{RO}_2i}[\text{NO}][\text{RO}_2]_i - k_{\text{NO}_2+\text{OH}}[\text{NO}_2][\text{OH}] - \phi_{\text{O}_3}[\text{O}_3] - k_{\text{HO}_2+\text{O}_3}[\text{HO}_2][\text{O}_3] - k_{\text{OH}+\text{O}_3}[\text{OH}][\text{O}_3] - \sum k_i[\text{alkene}]_i[\text{O}_3] - \text{netP}_{\text{PAN}}, \quad (1)$$

where  $k$  and  $\alpha$  indicate the reaction rate coefficient and organic nitrate yield from MCMv3.3.1, respectively.  $\phi$  is the fraction of  $\text{O}(^1\text{D})$  reacting with  $\text{H}_2\text{O}$ . Effective  $\text{NO}_2$  loss due to PAN production subtracting the thermal dissociation of PAN ( $\text{netP}_{\text{PAN}}$ ) is also included in the calculation of  $\text{PO}_x$ .

170 For the construction of  $\text{PO}_x$  isopleths, the model was constrained over each megacity using precursor concentrations averaged within each flight track (typically spanning ~1–2 hours per track). A gridded parameter space was created by scaling the observed mean mixing ratios of  $\text{NO}_x$  and total VOCs from 1 to 300%. Total VOCs were scaled linearly in 10% increments, with all individual VOC species adjusted proportionally to preserve VOC speciation.  $\text{NO}_x$  was scaled using a step size of 0.2 in log space, while maintaining a fixed  $\text{NO}/\text{NO}_2$  ratio. Each  $\text{NO}_x$ –VOC grid point was used as input to the box model to compute  $\text{PO}_x$ . The  $\text{NO}_x$ –VOC grid enabled simulation of a wide range of conditions, reflecting both low and high precursor levels, while holding their speciation constant.

180 In addition to the isopleth simulations, a flight steady-state (FSS) configuration was implemented using the same chemical mechanism and constraint strategy described above (see SI Sect. S2 for details). Unlike the isopleth approach, which applies track-averaged precursor constraints to construct a gridded  $\text{NO}_x$ –VOC scaling space, the FSS simulations were performed using time-resolved in-situ observations along the aircraft track (1 minute average). For each data point, the model was run to achieve steady state using the corresponding observed constraints to diagnose instantaneous  $\text{PO}_x$  along the flight track. This configuration enables evaluation of  $\text{PO}_x$  without applying precursor scaling or track-averaging, thereby providing a complementary perspective on how the 1–2 hour averaging and scaling assumptions used in the isopleth framework influence the inferred  $\text{PO}_x$  values.

## 2.3 Isopleth ~~construction~~ projection and orthogonal distance calculation

185 To diagnose the sensitivity of  $\text{PO}_x$  to its precursors, we constructed  $\text{PO}_x$  isopleths as a function of  $\text{NO}_x$  and total VOC mixing ratios using the box model setup described in Sect. 2.2. ~~A gridded parameter space was created by scaling the observed mean mixing ratios of  $\text{NO}_x$  and total VOCs from 1 to 300%. This grid enabled simulation of a wide range of conditions, reflecting both low and high precursor levels, while holding their speciations.~~ Each grid point in this two-dimensional  $\text{NO}_x$ -VOC space was individually used as an input to the box model to calculate  $\text{PO}_x$  (Eq. (1)). The resulting  $\text{PO}_x$  matrix was compiled into the two-dimensional surface, and contour plots of this surface were generated to produce isopleths (Sect. 3.2). The ridge of the  $\text{PO}_x$  surface was identified by locating the  $\text{PO}_x$  maxima along both  $\text{NO}_x$  and VOC dimensions of the grid, and the resulting ridge points were fit using a least-squares linear regression to define the transition line. Observed data points were then overlaid onto this space ~~to quantify their position relative to the transition line and associated  $\text{PO}_x$  sensitivity regimes, and the transition line between  $\text{NO}_x$ - and VOC-sensitive regimes was defined as the ridge that follows the direction of steepest descent of  $\text{PO}_x$  values.~~

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For each observed data point at  $(x_0, y_0)$  in the isopleth, the shortest orthogonal distance to the transition line was calculated as follows:

$$d = |ax_0 + by_0 + c|/\sqrt{a^2 + b^2}, \quad (2)$$

where  $ax + by + c$  is the linear form of the transition line. Distances were defined as negative on the  $\text{NO}_x$ -sensitive side and positive on the VOC-sensitive side. The absolute, unitless magnitude of  $|d|$  reflects the strength of the precursor sensitivity, with larger values indicating a more robust regime and smaller values indicating proximity to the transition.

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## 3 Results

### 3.1 Overview of vertical distributions and diurnal patterns

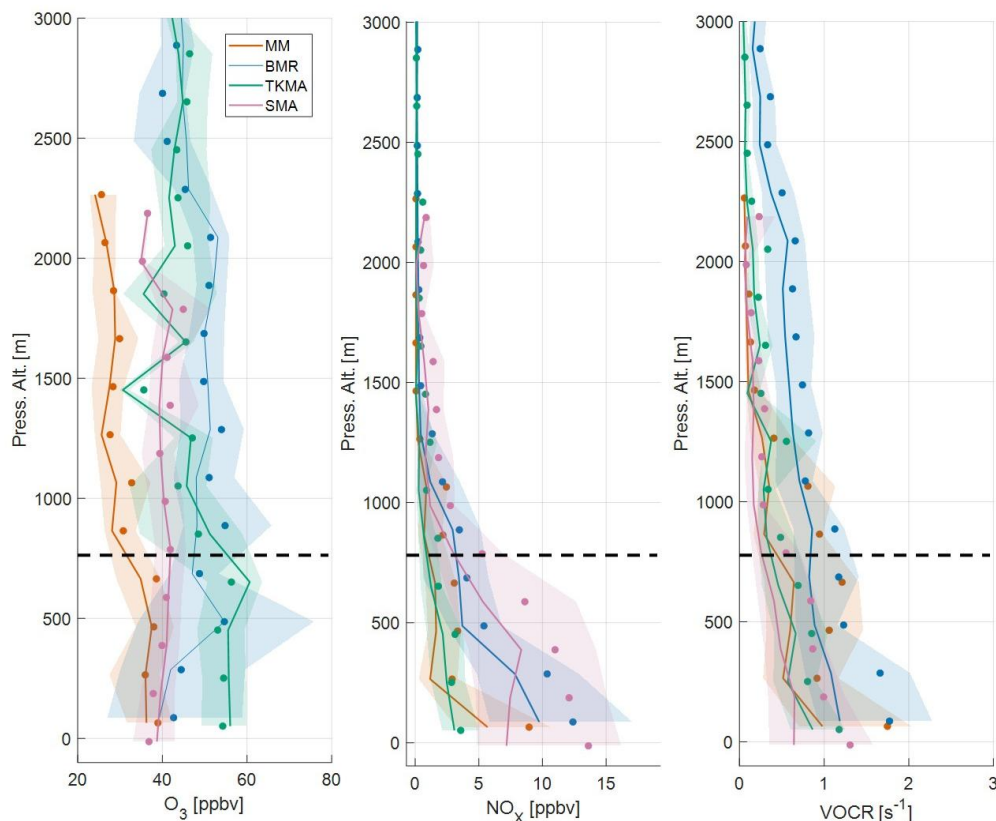
The comprehensive measurements collected during the ASIA-AQ campaign, including multiple missed-approaches over the studied megacities, provide a detailed view of the chemical conditions as a function of both altitude (Fig. 2) and time of day (Fig. 3).

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Figure 2 illustrates the average vertical profiles of  $\text{O}_3$ ,  $\text{NO}_x$ , and non-methane VOC reactivity (VOCR), the OH loss rate due to VOCs (see Sect. S1), observed over the four megacities, MM, BMR, TKMA, and SMA, during the entire campaign. In MM,  $\text{O}_3$  exhibits the lowest mixing ratios of the four locations at all altitudes, increasing from about 25 ppbv near 2 km to 39 ppbv close to the surface. BMR and TKMA show generally higher  $\text{O}_3$  in the range of 40–60 ppbv with a 1-minute-averaged maximum of 91 and 72 ppbv, respectively. ~~In the At~~ lower altitudes ( $< 750$  m),  $\text{O}_3$  mixing ratios at TKMA remain nearly constant at around 55 ppbv, whereas in BMR they decrease from about 55 to 42 ppbv due to  $\text{O}_3$  titration by NO. SMA shows comparatively uniform  $\text{O}_3$  mixing ratios of about 40 ppbv throughout the lower 3 km. The vertical distributions of  $\text{NO}_x$  display a steep decline with altitude in all cities, indicative of dominant near-surface-level sources and rapid chemical processes. SMA

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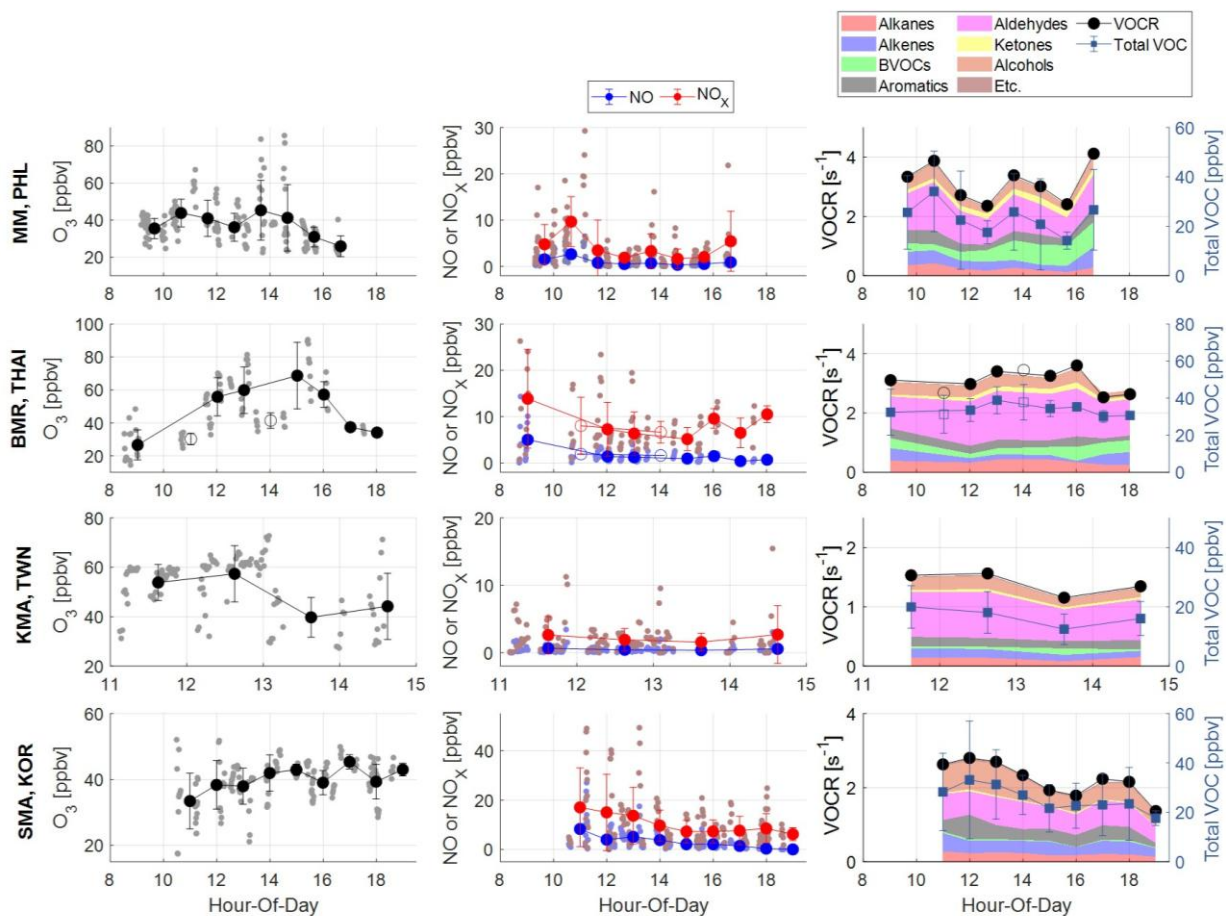
215 and BMR exhibited high  $\text{NO}_x$  mixing ratios at low altitudes ( $< 750$  m), with averages of 9.2 and 7.5 ppbv, respectively, whereas MM and TKMA showed lower levels on average, at 4 and 2.6 ppbv, respectively. VOCR exhibits distinct characteristics specific to each city. Below 750 m, MM and BMR exhibited the highest VOCR values of the four cities, on average 1.3 and 1.4  $\text{s}^{-1}$ , respectively, followed by TKMA (1.0  $\text{s}^{-1}$ ). SMA showed the lowest VOCR values, on average 0.8  $\text{s}^{-1}$  below 750 m, likely attributed to reduced VOC emissions during the colder and low solar irradiance conditions, when lower biogenic VOC  
220 (BVOCs) emissions (Fig. S4) and limited VOC evaporation are expected.



**Figure 2.** Average vertical profiles of the mixing ratios of  $\text{O}_3$ ,  $\text{NO}_x$ , and non-methane VOC reactivity (VOCR) observed over Asian megacities (red: Metro Manila, magenta: Bangkok Metropolitan Region, green: Tainan-Kaohsiung Metropolitan Area, black: Seoul Metropolitan Area) during ASIA-AQ. Dots indicate the mean values. Lines and shaded areas show the median and interquartile (25<sup>th</sup> and 75<sup>th</sup>), respectively. The black dashed line denotes the altitude cut-off; only data below this line were selected for the analysis.

Beyond the differences in VOCR levels, their compositions also varied among cities. Across the study domain, oxygenated VOCs (OVOCs), especially aldehydes and alcohols, were consistently dominant contributors to overall reactivity (Figs. S1–S4), due to both direct urban emissions (e.g., fuel use, solvents) and secondary formation from VOC oxidation, highlighting their importance in urban photochemistry. Isoprene made a notable contribution to VOCR, particularly in MM  
225 (15%; Fig. S1). BMR and TKMA showed stronger contributions from OVOCs (62% and 66%, respectively; Figs. S2 and S3),

mainly driven by HCHO and acetaldehyde. In SMA (Fig. S4), VOCR is distinguished by enhanced contributions from alcohols (26%) and elevated alkenes (15%) and aromatics (16%), while aldehydes contribute less than in the other cities.



**Figure 3.** Mean diurnal patterns of the mixing ratios of  $O_3$ , NO,  $NO_x$ , VOCs, and non-methane VOC reactivity (VOCR) near the surface (< 750 m) in Asian megacities. Pale-colored dots represent 1-minute averaged observations. Colored dots and error bars denote hourly averages and standard deviations. Stacked colored areas indicate the contributions of different VOC classes to VOCR. The open circles in the BMR panel indicate data that were sampled at relatively high altitudes. All times are local time (LT).

Figure 3 shows the diurnal variations of mixing ratios of  $O_3$ , NO,  $NO_x$ , and total VOC along with VOCR below 750 m. In the BMR panel, the two open symbols indicate data collected during low altitude legs (450–750 m) rather than near-surface missed approaches (50–750 m) as for all the other points. While these measurements are included for completeness, the higher altitude sampling may not fully represent conditions near the surface and ~~are~~ is thus marked differently to distinguish them from the other data points. As noted in Sect. 2.2, measurements in TKMA were obtained with a limited local-time window of approximately 11:00–14:30 local time (LT).

The amplitude of the diurnal  $NO_x$  and  $O_3$  cycle varies across cities. A distinct afternoon enhancement of  $O_3$  is most pronounced in the BMR (13:00–16:00 LT), where it is likely facilitated by efficient  $O_3$  formation from the oxidation of elevated

VOC and NO<sub>x</sub> levels. In contrast, the other cities show weaker diurnal variations. In MM, BMR, and SMA, NO<sub>x</sub> mixing ratios were higher during the morning rush hours, decreased through midday, and increased again toward the afternoon rush hour, while NO exhibited a clearer morning maximum. In contrast, in TKMA, NO<sub>x</sub> levels remained consistently low (< 3 ppbv on average) throughout the limited sampling window. This limitation and its implications for PO<sub>x</sub> sensitivity are further discussed in Sect. 4.1.1.

Within each city, the diurnal pattern of total VOC mixing ratios ~~are~~is similar to that of VOCR. This noteworthy relationship supports using total VOC as a practical proxy for reactivity in subsequent analyses (Sect. 3.2). As mentioned above, in all cities, overall VOCR is dominated by low carbon (< C3) aldehydes and alcohols, throughout the day (Figs S1–S4). In MM, both total VOC mixing ratios and VOCR exhibit three distinct peaks: during the morning rush hour (10:00–11:00 LT), early afternoon (13:00–15:00 LT), and late-afternoon rush hour (after 16:00 LT), where these enhancements were associated with different source contributions. The morning peak is primarily driven by anthropogenic VOCs (AVOCs) such as alkanes, aromatics, and alkenes, the early-afternoon peak is associated with enhanced biogenic contributions, mainly isoprene (Fig. S1), and the late-afternoon peak reflects contributions from both anthropogenic and biogenic VOCs. In BMR, total VOC mixing ratios and VOCR are relatively steady but show a modest afternoon enhancement driven by aldehydes. In contrast, TKMA exhibits overall lower total VOC mixing ratios and VOCR than other tropical Asian cities during the observed time window. Lastly, in SMA, VOCR showed distinct peaks around noon (12:00 LT) and near sunset (after 17:00 LT) coincided with ~~the~~ enhancements of alkenes and aromatic levelss.

### 3.2 Ozone production sensitivities in megacities

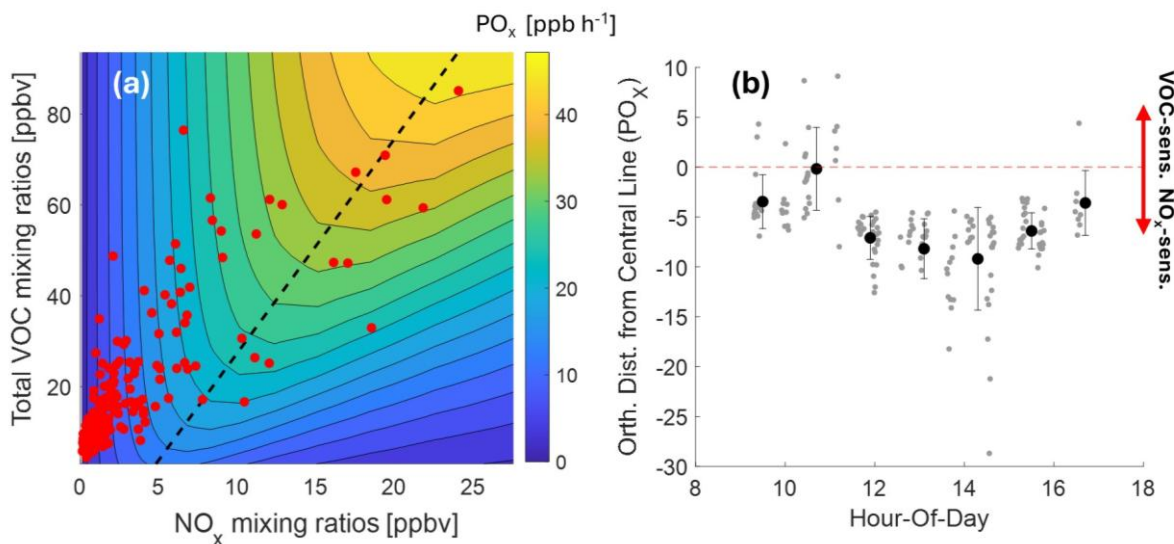
To investigate the sensitivity of O<sub>3</sub> formation to precursor levels, we generated isopleths of the PO<sub>x</sub> as a function of NO<sub>x</sub> and total VOC mixing ratios as described in Sect. 2.3. Figures 4a, 5a, 6a, and 7a show the model-generated PO<sub>x</sub> isopleths (background) with overlaid observations (red dots). The dashed black line marks the transition line between two regimes. In the NO<sub>x</sub>-sensitive regime (left side of the transition line), increasing NO<sub>x</sub> leads to a sharp increase in O<sub>x</sub> production, while in the VOC-sensitive regime (right side of the transition line), O<sub>x</sub> production becomes limited by the availability of radicals, which are produced from the oxidation of reactive VOCs.

In the following sections (Sects. 3.2.1–3.2.4), for the general isopleth representation of PO<sub>x</sub> sensitivity to the VOCs and NO<sub>x</sub> levels in each city (Figs. 4a, 5a, 6-a, and 7a), we used observations below 750 m across flights, regardless of flight date and time of day, and constrained the model using the values of chemical and meteorological parameters, averaged over all the flights at a given location. While the averaged isopleths cannot capture the instantaneous PO<sub>x</sub> sensitivity to precursors, they provide a more generalized view of chemical regimes that is useful for broader interpretation and policy discussions. However, this averaging process may introduce biases in precise PO<sub>x</sub> estimations and therefore their sensitivities to precursors, as it smooths out variations caused by diurnal changes and differences in meteorological conditions between research flights and time of day. To assess the potential bias from this large-scale averaging, we conducted additional model runs using data averaged within each flight track for each research flight (typically spanning ~1–2 hours per track; Supplementary

Information (SI) Sect. S3). Because each flight track was repeated two to three times per day, we categorized the model simulations into distinct time periods: morning and afternoon for BMR and TKMA, and morning, noon, and afternoon for MM and SMA. Given the large number of resulting isopleths, the individual figures are provided in the SI (SI Sect. S3). Although we primarily utilize the representative isopleths shown in Figs. 4a, 5a, 6a, and 7a for general visualization and interpretation in the following sections, it is important to note that all chemical sensitivity identifications (Table 1) and orthogonal distance-based calculations using Eq. (2) (Figs. 4b, 5b, 6b, and 7b) are based on the time- and flight-specific model outputs shown in Figs. S7–S10.

In addition, as noted in Sect. 2.2, the representativeness of this isopleth-based  $\text{PO}_x$  estimation was further evaluated by comparison with  $\text{PO}_x$  simulated from the a flight steady state (FSS) box model, which treats each observation independently to estimate instantaneous  $\text{PO}_x$  along the flight track (Sect. 2.2 and SI Sect. S2). The two approaches showed generally good agreement across regions, indicating that the key features of the inferred  $\text{PO}_x$  sensitivity are not strongly affected by the 1–2 h temporal averaging used in the isopleth framework supporting the reliability of the isopleth framework for diagnosing  $\text{PO}_x$  sensitivity (Fig. S5). Using the FSS box model results, the contributions of each  $\text{O}_x$  production or destruction term in Eq. (1) were computed and are shown in Fig. S6. City-specific  $\text{PO}_x$  budget features are discussed in Sect. 3.2.

### 3.2.1 Ozone production sensitivity in Metro Manila



**Figure 4.** (a) Colored isopleth of ozone production rate ( $\text{PO}_x$ ) as functions of  $\text{NO}_x$  and VOC mixing ratios in Metro Manila (MM), Philippines. The dashed lines denote the transition between  $\text{NO}_x$ -sensitive (left area of the transition) and VOC-sensitive (right area of the transition) regimes. Red dots indicate the observed data points (1-min average) sampled below 750 m. (b) Diurnal variation of the orthogonal distance of observation from the  $\text{PO}_x$  transition line (negative values:  $\text{NO}_x$ -sensitive; positive values: VOC-sensitive). The red dashed lines denote the transition line. Grey dots indicate individual data points (1-min average), and black dots and error bars represent the hourly mean values and standard deviations, respectively. All times are local time (LT).

Figure 4a presents the  $\text{PO}_x$  isopleth, generated using averaged observations below 750 m in MM. The simulated  $\text{PO}_x$  ranges up to 45  $\text{ppbv h}^{-1}$ , with regions of the highest  $\text{PO}_x$  high production rates (e.g., > 40  $\text{ppbv h}^{-1}$ ) occurring under moderate

high NO<sub>x</sub> (> 15 ppbv) and high VOC conditions (> 70 ppbv). Individual isopleths in Fig. S7 illustrate how the shape and gradient of the isopleths vary across different research flights and flight tracks at different times of day. The absolute levels of precursors and PO<sub>x</sub> vary between flights, and the isopleth shapes and transition lines likewise exhibit clear diurnal changes among morning, noon, and afternoon flight tracks (Figs. S7, ~~Table S3 and S11~~). Morning isopleths (Track 1 in Fig. S7) generally show broader VOC-sensitive regions and steeper PO<sub>x</sub> transition line ~~slopes~~ (~~Table S3 Fig. S11~~), driven by elevated NO<sub>x</sub> levels and fresh anthropogenic VOC emissions (Fig. 3). In noon and afternoon tracks (Tracks 2 and 3 in Fig. S7), the transition-line slopes become progressively smaller (~~Table S Fig. S1143~~) under

lower NO<sub>x</sub> but sustained VOCR with an increasing biogenic (isoprene) influence (Figs. 3 and S1).

As summarized in Table 1, 93% of the observed data points fall within the NO<sub>x</sub>-sensitive regime, contributing 79% of total PO<sub>x</sub>, where the PO<sub>x</sub> contribution is defined as the sum of PO<sub>x</sub> values over all NO<sub>x</sub>-sensitive points divided by the sum over all observations. Even among the highest PO<sub>x</sub>-producing conditions (upper 50% in the PO<sub>x</sub> distribution), the NO<sub>x</sub>-sensitive regime remains dominant, with 85% of the data points contributing 74% of total PO<sub>x</sub>, but these proportions are lower than those for the full dataset. This shift suggests that although data points in the VOC-sensitive condition are relatively few (7%), they tend to be associated with higher PO<sub>x</sub>. These VOC-sensitive points are clustered in the morning (~11:00 LT; Figs. 4b and S7, Track 1), when PO<sub>x</sub> levels are highest (Fig. S6a). It is important to note that all numbers in Table 1 are derived from individual isopleths generated for different flights and tracks (Fig. S7), rather than from the composite isopleths shown in Fig. 4a to ensure a more precise representation of varying atmospheric conditions throughout different flights and times of day. As a result, there are some discrepancies between Table 1 and Fig. 4a. For example, Fig. 4a shows fewer VOC-sensitive data points (5% of all data points), whereas Table 1 indicates that slightly more observed data points (7%) fall within the VOC-

Location	Regime	% of Points (All data)	%PO <sub>x</sub> Contribution (All data)	% of Points (High PO <sub>x</sub> #)	%PO <sub>x</sub> Contribution (High PO <sub>x</sub> #)
<del>MM, PHH</del>	NO <sub>x</sub> -sens.	93%	79%	85%	74%
	VOC-sens.	7%	21%	15%	26%
<del>BMR, THAI</del>	NO <sub>x</sub> -sens.	69%	77%	77%	79%
	VOC-sens.	31%	23%	23%	21%
<del>TKMA, TWN</del>	NO <sub>x</sub> -sens.	87%	83%	80%	79%
	VOC-sens.	13%	17%	20%	21%
<del>SMA, KOR</del>	NO <sub>x</sub> -sens.	4%	4%	4%	4%
	VOC-sens.	96%	96%	96%	96%

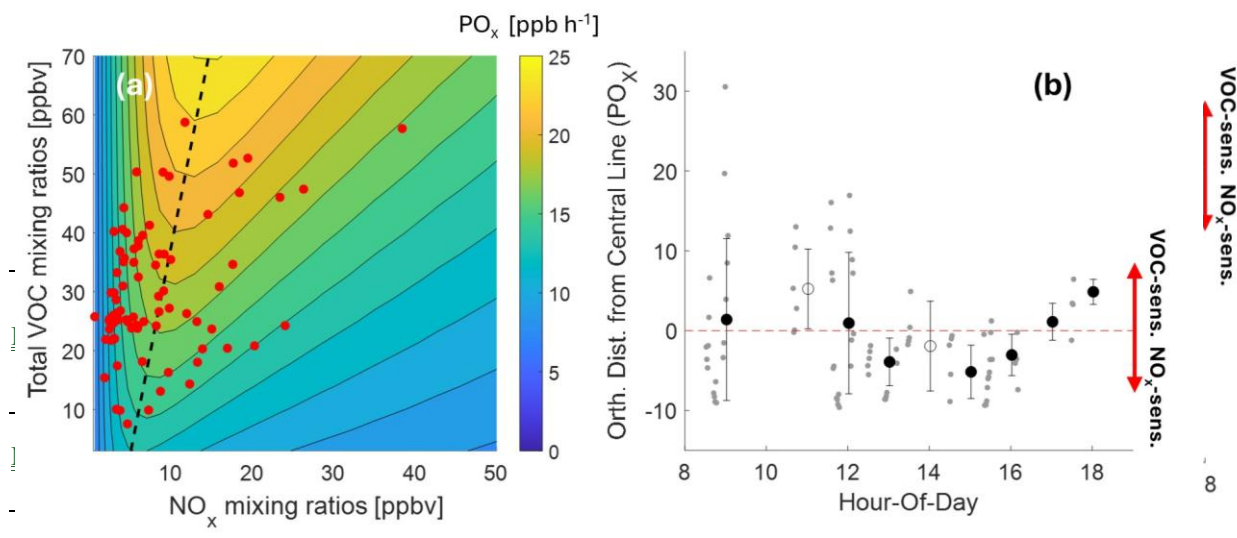
sensitive regime.

**Table 1.** Fractions of NO<sub>x</sub>-sensitive and VOC-sensitive regimes for PO<sub>x</sub> in each megacity.

<sup>a</sup> High  $PO_x$  denotes upper 50% in the  $PO_x$  distribution (above the median). Above-median  $PO_x$  refers to  $PO_x$  values greater than the median of the observed data points (shown as dots in the isopleth figures) for each region.

Figure 4b presents the diurnal variation of the orthogonal distance calculated using Eq. (2). While most observations remain in the  $NO_x$ -sensitive regime (negative distances), some exhibit smaller negative distances or even fall into the VOC-sensitive regime (positive distances) in the morning (~10:00–12:00 LT) and late afternoon (after 14:30 LT). This is most likely associated with high  $NO_x$  levels (Fig. 3) from traffic and urban sources during rush hour, temporarily shifting the chemical environment toward the transition boundary. As the day progresses and photochemical activity increases, ozone production shifts further into the  $NO_x$ -sensitive regime, with distances becoming more negative toward the afternoon. Near sunset, the distances shift slightly toward the transition, likely due to the evening rush hour emissions introducing additional  $NO_x$ .

### 3.2.2 Ozone production sensitivity in Bangkok Metropolitan Region



**Figure 5.** (a) Colored isopleth of ozone production rate ( $PO_x$ ) as functions of  $NO_x$  and VOC mixing ratios in the Bangkok Metropolitan Region (BMR), Thailand. The dashed lines denote the transition between  $NO_x$ -sensitive (left area of the transition) and VOC-sensitive (right area of the transition) regimes. Red dots indicate the observed data points (1-min average) sampled below 750 m. (b) Diurnal variation of the orthogonal distance of observation from the  $PO_x$  transition line (negative values:  $NO_x$ -sensitive; positive values: VOC-sensitive). The red dashed lines denote the transition line. Grey dots indicate individual data points (1-min average), and black dots and error bars represent the hourly mean values and standard deviations, respectively. The open circles at 11:00 and 14:00 LT indicates the data were sampled at relatively high altitude. All times are local time (LT).

VOC-SENS. 70/0 70/0 70/0 70/0

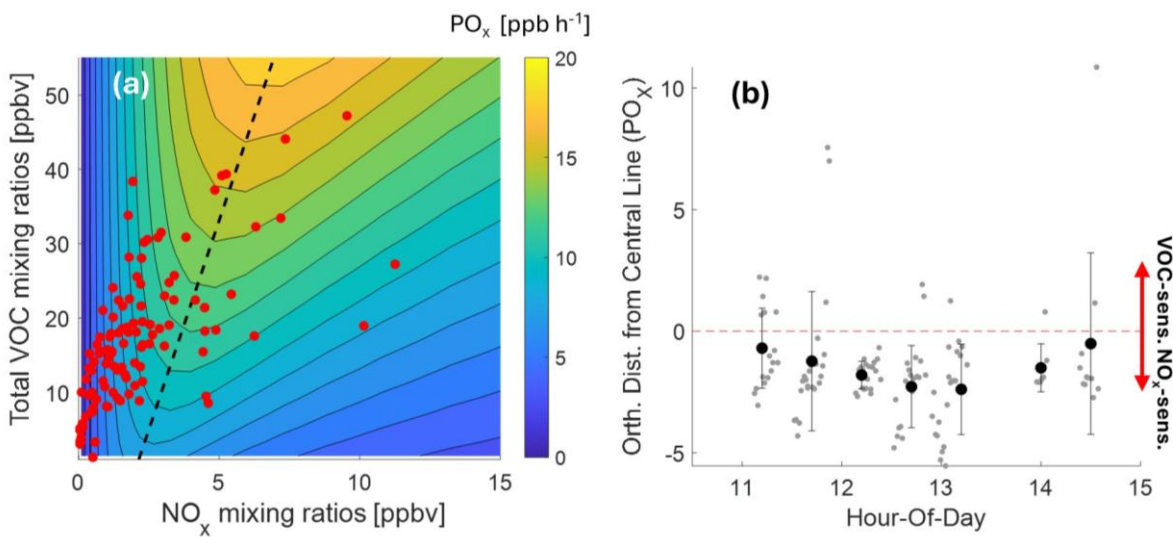
Figure 5a presents the generalized  $PO_x$  isopleth for BMR where simulated  $PO_x$  ranges up to 25  $ppbv\ h^{-1}$ . Regions of the highest  $PO_x$ . The high production rates (e.g., > 20  $ppbv\ h^{-1}$ ) occur under relatively low/moderate-to-high  $NO_x$  (7–18 ppbv) and moderate-to-high VOC mixing ratios (> 40 ppbv), consistent with a steeper transition line slope than in MM (Table S3). As summarized in Table 1, BMR exhibits a more evenly distributed chemical regime compared to MM, with a notable fraction

of data points falling in both NO<sub>x</sub>-sensitive (69%) and VOC-sensitive regimes (31%). NO<sub>x</sub>-sensitive and VOC-sensitive data points contribute to 77% and 23% of total PO<sub>x</sub>, respectively. However, among the highest PO<sub>x</sub>-producing conditions (upper 50% in the PO<sub>x</sub> distribution), NO<sub>x</sub>-sensitive data points are even more dominant (77% of the data points), accounting for 79% of total PO<sub>x</sub>. Consistent with this, the diurnal pattern of PO<sub>x</sub> (Fig. S6b) shows an increasing trend from morning to late afternoon and peaking around 16:00 LT, when NO<sub>x</sub>-sensitive conditions dominate (Fig. 5b). This suggests that while VOC-sensitive conditions are not negligible, ozone production in BMR is primarily limited by NO<sub>x</sub> availability, particularly during peak photochemical periods.

Compared to MM, BMR shows relatively less variability in PO<sub>x</sub> levels (Fig. S8) and transition-line slopes (Fig. S11) across both flight days and tracks (Fig. S8 and Table S3). In Track 1 (09:00–13:00 LT, Fig. S8), all flights exhibit similarly shaped isopleths, with the NO<sub>x</sub> levels at which PO<sub>x</sub> is maximized and ridge lines appearing in comparable positions, though RF13 and RF15 (defined in Table S1) show higher PO<sub>x</sub> values under comparable NO<sub>x</sub> and VOC levels. In Track 2 (13:00–18:00 LT, Fig. S8), the isopleths maintain consistent curvature with slightly reduced PO<sub>x</sub> levels.

To examine diurnal variations in PO<sub>x</sub> sensitivity, Fig. 5b illustrates the orthogonal distance of observed data points from the transition line within the PO<sub>x</sub> isopleths. The diurnal pattern shows a clear transition from VOC sensitive conditions in the morning to NO<sub>x</sub> sensitive conditions later in the day, likely driven by high NO<sub>x</sub> levels from rush hour emissions (Fig. 3). Decreasing NO<sub>x</sub> and increasing VOC levels in the afternoon results in the shift into the NO<sub>x</sub>-sensitive regime. As mentioned in Sect. 3.1, a brief increase in the distance is observed around 11:00 and 14:00 LT (open circles). These are based on indicating observations sampled at relatively higher altitudes with and are likely influenced by slightly different chemical conditions aloft that differ slight from than near-surface observations.

### 3.2.3 Ozone production sensitivity in Tainan-Kaohsiung Metropolitan Area



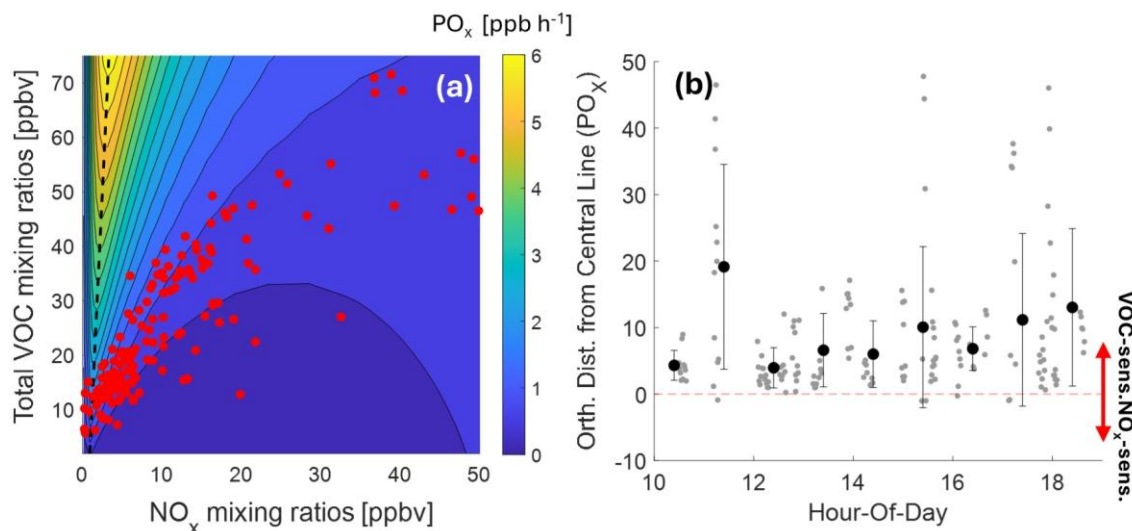
**Figure 6.** (a) Colored isopleth of ozone production rate ( $PO_x$ ) as functions of  $NO_x$  and VOC mixing ratios in the Tainan-Kaohsiung Metropolitan Area (KMA), Taiwan. The dashed lines denote the transition between  $NO_x$ -sensitive (left area of the transition) and VOC-sensitive (right area of the transition) regimes. Red dots indicate the observed data points (1-min average) sampled below 750 m. (b) Diurnal variation of the orthogonal distance of observation from the  $PO_x$  transition line (negative values:  $NO_x$ -sensitive; positive values: VOC-sensitive). The red dashed lines denote the transition line. Grey dots indicate individual data points (1-min average), and black dots and error bars represent the hourly mean values and standard deviations, respectively. All times are local time (LT).

Figure 6a shows the averaged  $PO_x$  isopleth for TKMA, showing the simulated  $PO_x$  ranges up to  $20 \text{ ppbv h}^{-1}$ . High Regions of the highest  $PO_x$  values ( $> 15 \text{ ppbv h}^{-1}$ ) occur at low-to-moderate  $NO_x$  (4–10 ppbv) and moderate-to-high VOC mixing ratios ( $> 40 \text{ ppbv}$ ). TKMA exhibits a predominantly  $NO_x$ -sensitive regime, with 87% of total data points falling within this condition, contributing 83% of total  $PO_x$  (Table 1). However, similar to MM (Table 1 and Sect. 3.2.1), on the upper 50%  $PO_x$  values,  $NO_x$ -sensitive points are less dominant (80% of data points, contributing 79% of total  $PO_x$ ), and points with-in the VOC-sensitive regime becomes significant, contributing 21% of total  $PO_x$ . As noted in Sect. 3.1, Table 1 does not include any early morning (~11:00 LT) and late afternoon (17:00–18:00 LT) data due to the constraints of the transit flights. Given that early morning and late afternoon points tend to be more VOC-sensitive due to high  $NO_x$  emissions from rush hour traffic, as observed in MM (Fig. 4b) and BMR (Fig. 5b), the actual VOC-sensitive contribution to  $PO_x$  in TKMA may be underestimated relative to by the values reported in Table 1.

The individual  $PO_x$  isopleths for TKMA (Fig. S9) exhibit quite consistent contour shapes and transition-line slopes within each individual flight (Table S3Fig. S11), while the absolute  $PO_x$  levels vary significantly between flights, with RF16 showing notably higher production rates than RF05 and RF11. This pronounced variability in  $PO_x$  magnitude is likely driven by large day-to-day differences in meteorological and photochemical conditions associated with the sparse flight dates in TKMA (Sect. 3.1).

Figure 6b illustrates the diurnal variations (11:00–15:00 LT) of the  $PO_x$  based on orthogonal distances of observed data points from the transition line. As noted above, TKMA displays-illustrates less pronounced diurnal variations in  $PO_x$  sensitivity compared to MM and BMR due to the limited flight sampling times without early morning data. Although most data points fall on the  $NO_x$ -sensitive side, the distances tend to decrease from morning to afternoon, consistent with the patterns observed in MM (Fig. 4b) and BMR (Fig. 5b).

### 3.2.4 Ozone production sensitivity in Seoul Metropolitan Area



**Figure 7.** (a) Colored isopleth of ozone production rate ( $PO_x$ ) as functions of  $NO_x$  and VOC mixing ratios in Seoul Metropolitan Area (SMA), South Korea. The dashed lines denote the transition between  $NO_x$ -sensitive (left area of the transition) and VOC-sensitive (right area of the transition) regimes. Red dots indicate the observed data points (1-min average) sampled below 750 m. (b) Diurnal variation of the orthogonal distance of observation from the  $PO_x$  transition line (negative values:  $NO_x$ -sensitive; positive values: VOC-sensitive). The red dashed lines denote the transition line. Grey dots indicate individual data points (1-min average), and black dots and error bars represent the hourly mean values and standard deviations, respectively. All times are local time (LT).

As mentioned above, the Seoul Metropolitan Area (SMA) deployment during the ASIA-AQ campaign was under low-temperature and limited solar radiation, in contrast to the tropical and subtropical Asian cities (MM, BMR, and TKMA), resulting in reduced reaction rates (Fig. S6d).

Figure 7a presents the generalized isopleths of  $PO_x$  for SMA. The simulated  $PO_x$  ranges from near zero to 6  $ppbv\ h^{-1}$ , which is notably lower than those found in other cities. A distinct feature of the SMA isopleths is that the transition line is highly inclined toward the VOC axis, implying a generally stronger VOC dependence compared to  $NO_x$ . The model simulation background isopleth shows that regions of highest  $PO_x$  values occur at low  $NO_x$  levels below 5  $ppbv$  and high VOC conditions ( $> 65\ ppbv$ ), with ozone production rapidly declining as  $NO_x$  increases. The observed data shows that overall  $PO_x$  levels were below 2  $ppbv\ h^{-1}$  and most of them are in VOC-sensitive regime, with 96% of all data points falling in the VOC-sensitive zone, contributing 96% of total  $PO_x$  (Table 1).

375 Consistent with the general isopleth, the individual PO<sub>x</sub> isopleths for SMA (Fig. S10) exhibit highly inclined  
ridgelines and transition lines across all flights and tracks in SMA, with slopes more than twice those in other cities (Fig.  
S11 Table S3). Notably, the absolute PO<sub>x</sub> production rates are substantially lower than those observed in the other megacities,  
with most observations remaining under the contours below 2 ppbv h<sup>-1</sup> except for RF06 and further decreasing as the day  
progresses. In Track 3 (15:00–18:00 LT, Fig. S10), PO<sub>x</sub> levels decline even more markedly, likely due to diminished  
photochemical activity and precursor availability during the late winter afternoon.

380 Correspondingly, the orthogonal distances in Fig. 7b show consistently large positive values, indicating that most  
observations remained well within the VOC-sensitive regime regardless of time. While there is some hour-to-hour variability,  
no clear diurnal trend or regime transition is observed, unlike in the more dynamic tropical cities.

### **3.2.5 Comparative overview of PO<sub>x</sub> isopleths across cities and times of day**

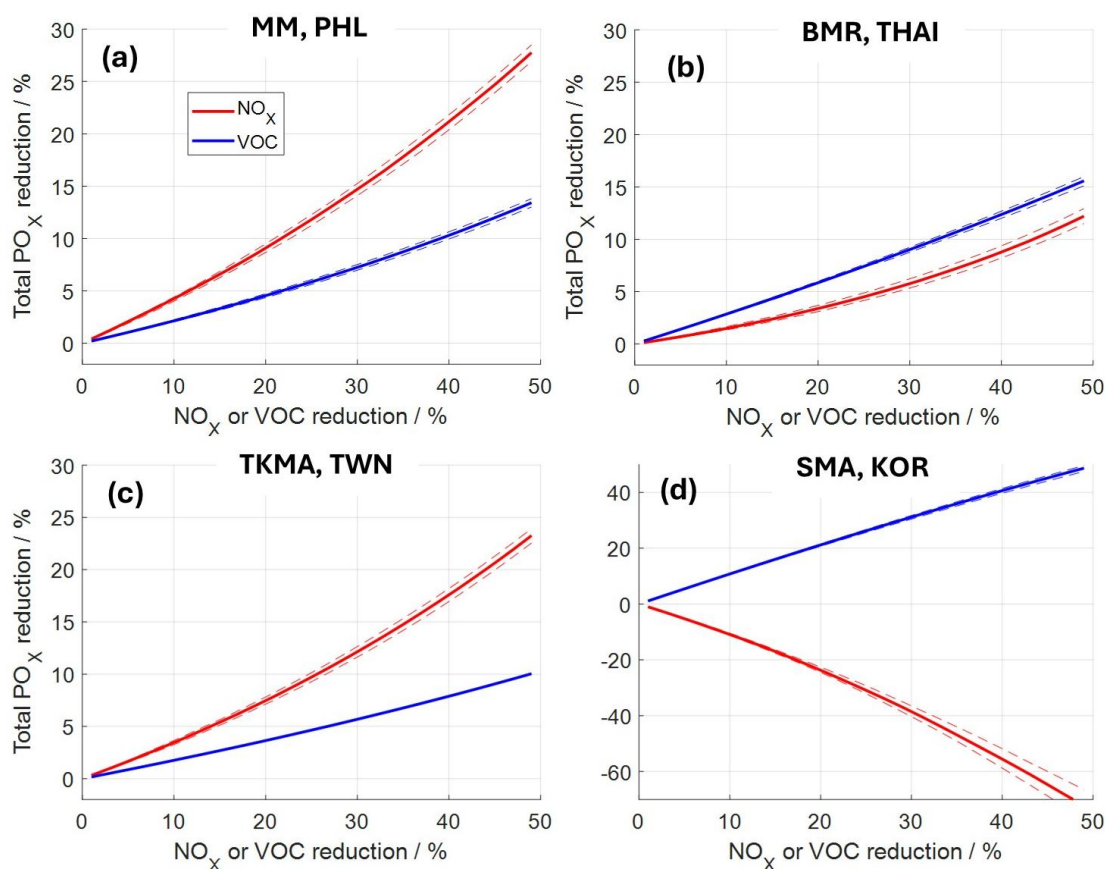
385 Across the four megacities, the PO<sub>x</sub> isopleths differ systematically, as reflected in differences in transition-line slopes  
(Fig. S11), the background isopleth patterns (Figs. 4a–7a and S7–S10), and the distributions of observed data points (Figs. 4b–  
7b).

390 The slopes of the transition line calculated from the PO<sub>x</sub> isopleths range from approximately 3 to 40 across the four  
cities. MM shows the shallowest transition line (slope = 4.71, Fig. S11), followed by progressively steeper slopes in BMR  
(7.01) and TKMA (11.24). SMA exhibits a substantially steeper transition line (30.27), indicating a markedly different  
background isopleth geometry under colder conditions. In addition, a clear diurnal variation in the individual isopleths is  
evident primarily in MM, with larger transition slopes during morning tracks and decrease toward afternoon tracks. In contrast,  
BMR and TKMA show relatively minor diurnal changes within individual flights (except for RF12 Track 2), with transition-  
line slopes remaining broadly consistent across tracks. SMA shows no systematic diurnal pattern in the transition-line slopes.

395 The distributions of the observed data points further highlight differences in how each city samples on its isopleths.  
In MM, the morning tracks exhibit a relatively broad spread of observed points across the NO<sub>x</sub>–VOC space, whereas  
observations during later tracks become more concentrated and are largely confined to a narrow NO<sub>x</sub> range, predominantly  
below 5 ppbv, on the NO<sub>x</sub>-sensitive side of the transition line. In contrast, BMR and TKMA show a more dispersed distribution  
of observed points across the NO<sub>x</sub>–VOC space, spanning a wider range of precursor conditions relative to the transition line.  
In SMA, the observed points also appear widely distributed in NO<sub>x</sub>–VOC space. However, because of the substantially steeper  
400 transition line, nearly all observations fall on the VOC-sensitive side of the isopleths despite this apparent spread.

## 4 Discussion

### 4.1 Ozone reduction strategies



**Figure 8.** Sensitivity of total PO<sub>x</sub> reduction (%) to NO<sub>x</sub> and VOC reductions (%) in four megacities: (a) Metro Manila (MM), (b) the Bangkok Metropolitan Region (BMR), (c) the Tainan-Kaohsiung Metropolitan Area (TKMA), and (d) the Seoul Metropolitan Area (SMA). Red and blue lines and markers represent the response of PO<sub>x</sub> to NO<sub>x</sub> and VOC reductions, respectively. The dashed lines indicate the range of PO<sub>x</sub> reduction responses obtained by propagating measurement uncertainties in NO<sub>x</sub> and VOC observations.

As described in Sect. 3.2, the analysis of ozone production rates reveals distinct precursor sensitivity regimes across the four studied megacities, emphasizing the need for location-specific precursor control strategies. In this section, we discuss the effectiveness of NO<sub>x</sub> and VOC reductions in mitigating ozone formation depending on NO<sub>x</sub> or VOC reduction scenario across the four megacities (Fig. 8). It should be noted that the analysis diagnoses local photochemical ozone production across the NO<sub>x</sub>-VOC precursor space. Transport processes such as regional advection or vertical mixing may influence the observed ozone mixing ratios but are not explicitly represented in this study.

410 To evaluate the sensitivity of total PO<sub>x</sub> reduction to NO<sub>x</sub> or VOC reductions, we used the individual PO<sub>x</sub> isopleths from each city (SI Sect. S3). For each isopleth, we first computed the baseline total PO<sub>x</sub> by summing PO<sub>x</sub> values at all observed data point locations in the isopleth. We then applied hypothetical reductions of NO<sub>x</sub> or VOC levels from 0% to 50%, adjusting the location of each data point accordingly along the relevant axis (NO<sub>x</sub> or VOC). At each reduction level, the total PO<sub>x</sub> was recalculated for each reduction scenario by summing the PO<sub>x</sub> value at the shifted positions. The resulting curves (Fig. 8) show  
415 the percentage change in total PO<sub>x</sub> as a function of precursor reduction. Here, positive percentage changes indicate that PO<sub>x</sub> decreases relative to the baseline, whereas negative percentage changes indicate that PO<sub>x</sub> increases, reflecting a counterproductive response to the applied precursor reduction.

#### 4.1.1 Ozone reduction strategies in MM and TKMA

MM and TKMA exhibit a predominantly NO<sub>x</sub>-sensitive ozone formation regime with 93% and 87% of the observed  
420 data, respectively (Table 1). This NO<sub>x</sub>-sensitive dominance in MM and TKMA is further supported by Fig. 8a and c, where PO<sub>x</sub> reduction shows a stronger response to NO<sub>x</sub> reductions than to VOC reductions with slightly steeper slopes of NO<sub>x</sub>- and VOC-response curves in MM than in TKMA. A 20% reduction in NO<sub>x</sub> levels would result in 9% and 7% decreases in PO<sub>x</sub> in MM and TKMA, respectively. However, the impact of VOC reductions would not be negligible. As shown in Table 1, under high-O<sub>3</sub>-producing conditions, the contribution of VOC-sensitive points increases to 26% in MM and 21% in TKMA.  
425 Particularly in MM, the morning PO<sub>x</sub> maximum aligns with relatively frequent VOC-sensitive conditions (Figs. 4, S6a, and S7) and enhanced contributions from anthropogenic alkanes, alkenes, and aromatics (Fig. 3). This indicates that VOC controls targeting morning anthropogenic VOC emissions could be especially effective in reducing peak ozone production. The reduction of VOCs would result in approximately half the PO<sub>x</sub> decrease compared to NO<sub>x</sub> reduction in both MM and TKMA (Fig. 8a and c). Given these findings, a combined NO<sub>x</sub> and VOC control strategy would be the most effective approach for O<sub>3</sub>  
430 mitigation in MM and TKMA, with priority given to NO<sub>x</sub> reductions due to their dominant influence on PO<sub>x</sub> formation.

It is important to recognize that the sampling window in TKMA was restricted to 11:00–15:00 LT, meaning that early morning and late afternoon data were not included (Figs. 6b and S9). This limitation could have a significant impact on the characterization of O<sub>3</sub> sensitivity shown in Fig. 8c. As shown in tropical Asian megacities (MM and BMR), early morning and late afternoon periods tend to exhibit higher NO<sub>x</sub> mixing ratios (Fig. 3), leading to an increased occurrence of VOC-sensitive  
435 conditions or points closer to the transition line (Figs. 4a and 5a). The absence of these time periods in the TKMA dataset suggests that VOC-sensitive conditions in Table 1 and the effect of VOC reduction to PO<sub>x</sub> in Fig. 8c may be underestimated, potentially leading to an overemphasis on NO<sub>x</sub> reduction in TKMA.

~~In addition, sampling inhomogeneity in terms of time of day and altitude can introduce a bias in this analysis,~~  
particularly when investigating the statistical distribution of PO<sub>x</sub> sensitivities in Table 1 and evaluating the potential impact of  
440 precursor reductions in Fig. 8. ~~The detailed description of t~~his potential bias is discussed in detail in the SI (Sects. S4 and S5).

#### 4.1.2 Ozone reduction strategies in BMR

As described in Sect. 3.2.2, BMR exhibits a mixed sensitivity regime, with  $\text{NO}_x$ -sensitive conditions prevailing overall (69% of the observed data classified as  $\text{NO}_x$ -sensitive and 31% as VOC-sensitive). However, under higher  $\text{PO}_x$  conditions,  $\text{NO}_x$ -sensitive regime becomes more predominant (Table 1), suggesting that ozone formation remains predominantly  $\text{NO}_x$ -driven during peak  $\text{O}_3$  build-up periods. Based on this, it appears that  $\text{NO}_x$  reductions would be the most effective strategy for  $\text{O}_3$  mitigation in BMR, similarly to MM and TKMA.

Figure 8b presents a contrasting perspective to the above, however, showing that  $\text{PO}_x$  reduction in BMR responds more strongly to VOC reductions than to  $\text{NO}_x$  reductions, particularly at higher precursor reduction levels. A 20% reduction in VOC levels results in a 6% decrease in  $\text{PO}_x$ , while the same reduction in  $\text{NO}_x$  leads to only a 3% decrease. Compared to the results in MM and TKMA, the greater effectiveness of VOC reductions in BMR is attributed to the higher proportion of data points classified as VOC-sensitive. Additionally, the presence of multiple VOC-sensitive points implies that reducing  $\text{NO}_x$  levels can, in some cases, lead to an increase in  $\text{PO}_x$  (Figs. 5a and S8), which partially offsets the  $\text{PO}_x$  reductions achieved by the  $\text{NO}_x$ -sensitive points and thus moderates the net  $\text{PO}_x$  decrease. This counteracting effect is due to the nature of  $\text{PO}_x$  isopleth geometry in VOC-sensitive regime in BMR. A slight upward curvature of the contour lines in the VOC-sensitive area can shift data points toward regions of higher  $\text{PO}_x$ , when  $\text{NO}_x$  decreases. This can be explained by an increased OH level driven by reduced OH loss via the  $\text{OH} + \text{NO}_2$  reaction, which facilitates VOC oxidation and enhances ozone production. This counterbalancing effect explains why  $\text{NO}_x$  reductions are less effective in BMR than in other (sub)tropical Asian cities (MM and TKMA).

Moreover, considering the sampling inhomogeneity (Sect. S5), the effect of VOC reduction on  $\text{PO}_x$  decrease may be underestimated. The sampling distribution shown in Figs. 5b and S143 indicates that data collection was relatively sparse in 16:00–18:00 LT with several observations near the transition line, which may correspond to transitional or VOC-sensitive conditions, further shifting its classification toward a more balanced  $\text{NO}_x$ -VOC sensitivity regime. This suggests that while the statistics in Table 1 alone make BMR appear predominantly  $\text{NO}_x$ -sensitive, a more comprehensive analysis incorporating Fig. 8b and the sampling biases indicates that VOC reductions may be more effective than initially perceived.

#### 4.1.3 Ozone reduction strategies in SMA

As shown in Sect. 3.2.4, the SMA exhibits a distinctly VOC-sensitive regime for ozone formation (see Table 1 and Fig. 7). This behavior is clearly illustrated in Fig. 8d, where a VOC reduction results in a decrease in  $\text{PO}_x$ , while a  $\text{NO}_x$  reduction leads to an increase in  $\text{PO}_x$ . This highlights that  $\text{NO}_x$  reductions could be counterproductive in VOC-sensitive environments like wintertime SMA if not accompanied by VOC controls.

However, since winter photochemistry in the SMA is less active, and  $\text{O}_3$  pollution is typically less severe, it is important to assess whether this VOC sensitivity persists during periods of stronger  $\text{O}_3$  formation. To address this, we applied the same isopleth analysis used for ASIA-AQ to the observations from the KORUS-AQ campaign conducted in April–June

2016, a time of enhanced photochemical activity (Crawford et al., 2021). The isopleth diagram in Fig. S176 shows that, although the isopleths and central transition line are slightly less severe than during ASIA-AQ, the majority of the observed data points are located in the VOC-sensitive regime. According to Table S44, SMA remains VOC-sensitive even during KORUS-AQ, with 59% of observations classified as VOC-sensitive and contributing 63% of total PO<sub>x</sub>. The importance of VOC-sensitive cases is amplified under high-O<sub>3</sub> producing conditions, where 68% of the upper 50% PO<sub>x</sub> cases (above median) fall in the VOC-sensitive regime, contributing 68% of total PO<sub>x</sub>. Further quantitative support comes from Fig. S187, which shows the response of total PO<sub>x</sub> to precursor reductions. A VOC reduction effectively reduces the ozone production, while a NO<sub>x</sub> reduction leads to a counteractive increase in PO<sub>x</sub>.

Such precursor control strategies for ozone pollution are consistent with major findings from the KORUS-AQ campaign. For example, the KORUS-AQ overview study (Crawford et al., 2021) highlighted ~~that~~ the complexity of ozone formation in the SMA, emphasizing the importance of local VOC emissions and regional transport as drivers of elevated O<sub>3</sub> pollution. More specifically, Schroeder et al. (2020) investigated regional O<sub>3</sub> formation by comparing observations from SMA with those from a downwind remote site (Taehwa Forest), focusing on chemical evolution driven by precursor oxidation. The comparison indicated that O<sub>3</sub> production in SMA was in the VOC-sensitive regime, particularly influenced by oxidation of C<sub>7+</sub> aromatics, isoprene, and anthropogenic alkenes. Consistent with this study, the counteracting effects of NO<sub>x</sub> reduction in O<sub>3</sub> production ~~were~~ found in Seoul, but areas outside of Seoul were found to be in the transition regime. In addition, Kim et al. (2020) reported high O<sub>3</sub> episodes and their controlling factors at a ground site in Seoul during KORUS-AQ study. They showed that the O<sub>3</sub> formation for high surface-level O<sub>3</sub> cases was sensitive to VOC, consistent with our findings. Furthermore, Cho et al. (2021) and Fried et al. (2020) quantified VOC emissions from a large petrochemical complex located ~~in~~ on the coast of the yellow-West Ssea and traced their downwind effects on formaldehyde and O<sub>3</sub> production, demonstrating how local VOC-rich plumes can significantly influence the oxidation capacity and O<sub>3</sub> levels in urban regions of South Korea (Lee et al., 2022). Together, these studies build a compelling case that ozone production in SMA is consistently VOC-sensitive, regardless of the method or platform used.

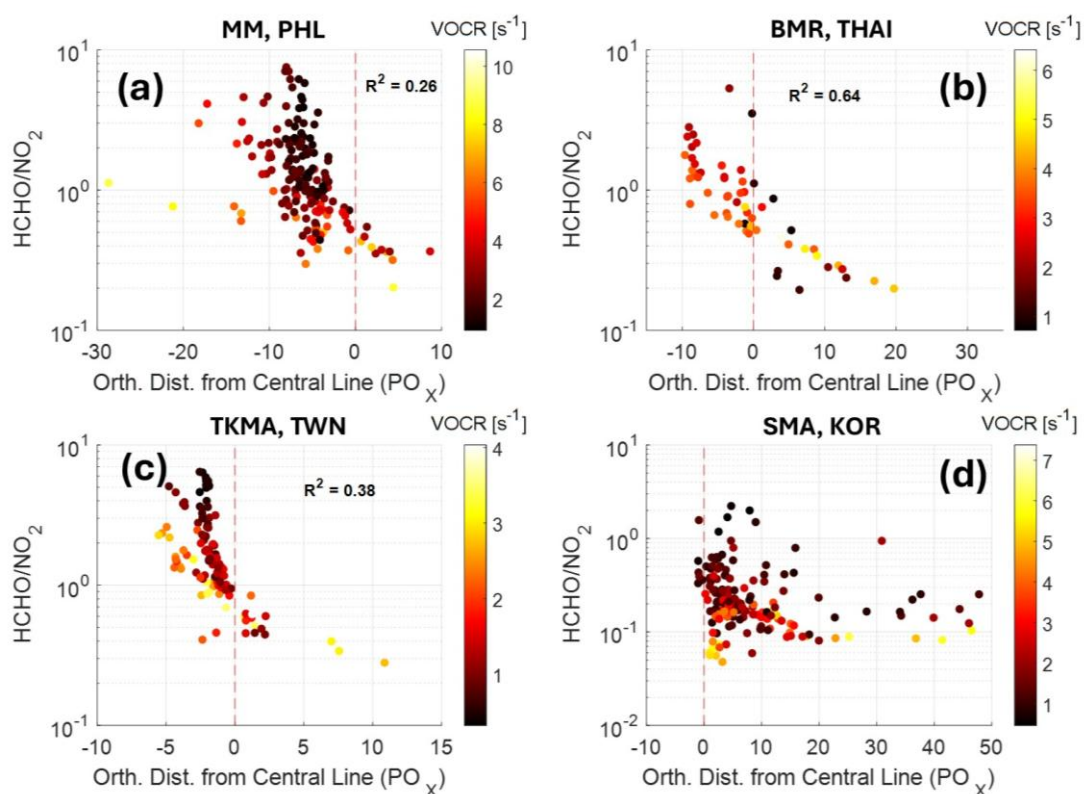
Taken together, the results from ASIA-AQ (late winter) and KORUS-AQ (spring ~~to~~ through early summer) demonstrate a seasonally consistent VOC-sensitive PO<sub>x</sub> regime in SMA. Despite variations in photochemical activity and emission dynamics between seasons and years, both campaigns consistently identify VOC reductions as the most effective pathway for mitigating ozone formation in SMA. However, this VOC-sensitive classification reflects the marginal response of ozone under current conditions and should not be interpreted as an absence of NO<sub>x</sub> pollution, as NO<sub>x</sub> mixing ratios in Seoul remain elevated. VOC-focused controls are therefore needed to avoid short-term ozone disbenefits when NO<sub>x</sub> emissions are reduced, while sustained NO<sub>x</sub> reductions remain essential as part of a long-term air-quality strategy.

#### 4.1.4 Policy-relevant framing of VOC controls

The PO<sub>x</sub>-reduction curves in Fig. 8 apply uniform reductions to all VOC classes and should therefore be interpreted as an upper bound of achievable VOC control benefits. In practice, regulations affect AVOCs, whereas BVOCs are not directly

controllable. Guided by the VOC speciation (Figs. S1–S4), the sum of primary VOCs (alkanes, alkenes, and aromatics) comprises approximately 30% (MM), 29% (BMR), 28% (TKMA), and 29% (SMA) of total VOCR. Accordingly, real-world VOC benefits will generally be smaller than those indicated by the all-VOC curves. Importantly, OVOCs dominate VOCR, and they are only partially controllable. A portion is primary anthropogenic (e.g., solvent/industry sources) and a substantial portion is secondarily produced, which will decline indirectly as AVOCs are reduced. Consequently, the attainable VOC impact should be viewed as less than the upper bound represented by the all-VOC curves (Fig. 8). As a forward step, combining this sensitivity framework with VOC source apportionment (e.g., PMF analysis; He et al., 2019; Chen et al., 2022) could more clearly identify the specific controllable source factors dominating VOCR in each city, enabling source-targeted VOC controls that translate sensitivity into actionable policy.

#### 4.2 Comparison with *in-situ* HCHO/NO<sub>2</sub> ratio



**Figure 9.** Comparisons between *in situ* HCHO/NO<sub>2</sub> ratios and orthogonal distances (negative: NO<sub>x</sub>-sensitive; positive: VOC-sensitive) from the PO<sub>x</sub> transition line (red dashed line) for (a) MM, (b) BMR, (c) TKMA, and (d) SMA. Colors denote VOCR values. All data points are based on 1-minute average.

The HCHO/NO<sub>2</sub> ratio has widely-been widely used as an empirical indicator to diagnose ozone production sensitivity. Many previous studies have suggested a transition range of approximately 1–2 for HCHO/NO<sub>2</sub>, although this value can vary depending on chemical environment (e.g., VOC speciation, direct HCHO emission) (Martin et al., 2004; Duncan et al., 2010;

520 Chang et al., 2016; Jin et al., 2017; Schroeder et al., 2017; Souri et al., 2020, 2023). In contrast, the orthogonal distance  
obtained from the isopleth approach inherently accounts for VOC speciation and radical chemistry based on box model  
simulations. This offers a more mechanistic and potentially robust diagnostic, conceptually similar to LNO<sub>x</sub>/LRO<sub>x</sub> (total NO<sub>x</sub>  
loss/total RO<sub>x</sub> radical loss) approach, which is often used to verify HCHO/NO<sub>2</sub> thresholds (Kleinman, 2005; Tonnesen and  
Dennis, 2000; Duncan et al., 2010; Schroeder et al., 2017; Souri et al., 2020, 2023). While the primary focus of this study is  
525 to diagnose PO<sub>x</sub> sensitivity in different Asian megacities, we further examine how our derived sensitivity regimes compare  
with the widely used empirical indicator, the HCHO/NO<sub>2</sub> ratio.

Figure 9 shows *in situ* HCHO/NO<sub>2</sub> ratios from DC-8 observations as a function of the orthogonal distances from the  
PO<sub>x</sub> transition line obtained from the isopleth analysis. Across tropical and subtropical Asian megacities (MM, BMR, and  
TKMA), we observed a generally negative correlation, indicating that HCHO/NO<sub>2</sub> ratios decrease as the orthogonal distance  
530 shifts toward VOC sensitivity. This relationship is strongest in BMR ( $r^2 = 0.64$ ) and moderate in MM and TKMA ( $r^2 = 0.26$   
and 0.38, respectively). In contrast, SMA shows a low correlation, which we attribute to low photochemical activity during  
the wintertime. Interestingly, the data with very low VOCCR ( $< 2 \text{ s}^{-1}$ ) show both a broader vertical spread in orthogonal distance  
and a steeper slope compared to higher-VOCCR conditions. This suggests that when overall VOC-reactivity is extremely low,  
HCHO/NO<sub>2</sub> ratios become less reliable in delineating sensitivity regimes. This finding aligns with the results of Souri et al.  
535 (2020), who showed that HCHO does not fully capture VOCCR in pristine or heavily polluted environments.

In our analysis, the transition point (orthogonal distance = 0) occurs at an HCHO/NO<sub>2</sub> ratio of about 0.3–1.0, which  
is notably lower than the conventional range reported in previous studies. This discrepancy likely arises from methodological  
differences: we use *in situ* HCHO and NO<sub>2</sub> mixing ratios from the DC-8 observations within the boundary layer, whereas most  
previous studies employed satellite column densities.

## 540 5 Conclusions

Understanding the nonlinear responses of ozone formation to its precursors such as NO<sub>x</sub> and VOCs is essential for  
designing effective air quality management strategies. In rapidly urbanizing regions of Asia, particularly in urban centers or  
major megacities, where emissions from transportation, industry, and biomass burning are intense and diverse, the nonlinearity  
poses a major challenge for targeted policy implementation.

545 Previous studies have examined ozone formation sensitivity in urban environments using a variety of observational  
and modeling approaches. Building on these studies, we use airborne observations and an isopleth approach to diagnose ozone  
formation sensitivity and its diurnal variation. Specifically, this study utilizes measurements from the NASA DC-8 aircraft  
during the Airborne and Satellite Investigation of Asian-Air Quality (ASIA-AQ) for a comprehensive assessment of the  
sensitivity of ozone production (PO<sub>x</sub>) to NO<sub>x</sub> and VOCs across four Asian megacities in Metro Manila (MM), Bangkok  
550 Metropolitan Region (BMR), Tainan-Kaohsiung Metropolitan Area (TKMA), and Seoul Metropolitan Area (SMA).

Across all four cities, extensive observations were collected within the lower boundary layer below 750 m through multiple missed approaches during aircraft operations over megacities, providing high-resolution vertical profiles of NO<sub>x</sub>, VOCs, and O<sub>3</sub> down to near-surface levels with daytime diurnal information. The three (sub)tropical Asian cities (MM, BMR, and TKMA) were sampled during the dry hot season, characterized by active photochemistry. MM exhibited relatively low NO<sub>x</sub> and high VOC reactivity (VOCR), BMR showed high NO and high VOCR, and TKMA displayed low NO<sub>x</sub> and moderate VOCR. In contrast, SMA was sampled under conditions with high NO<sub>x</sub> and low VOCR, reflecting suppressed photochemistry typical of cold-season urban environments.

Using a chemical box model constrained by the averages from in situ observations, we constructed isopleths of PO<sub>x</sub> as a function of NO<sub>x</sub> and VOC mixing ratios. Observed data points were then overlaid onto the model-generated isopleths to diagnose the chemical sensitivity regimes for each city. The isopleth analysis revealed clear regional contrasts in ozone production sensitivity, and based on this analysis, ozone reduction strategies were derived for each city.

- In MM and TKMA, most data points fell into the NO<sub>x</sub>-sensitive regimes. Therefore, NO<sub>x</sub> reduction should be prioritized to mitigate ozone pollution. However, under high PO<sub>x</sub> conditions, in MM, the fraction of VOC-sensitive points increased significantly, indicating that VOC controls may still yield considerable benefits.
- BMR displayed a more mixed sensitivity regime than in MM and TKMA. Despite a larger proportion of NO<sub>x</sub>-sensitive data points, ozone formation shows greater responsiveness to VOC reductions due to the counteractive effects of NO<sub>x</sub> reductions under high-PO<sub>x</sub> and VOC-sensitive conditions. Therefore, a simultaneous reduction of both NO<sub>x</sub> and VOCs would be the most effective strategy for addressing ozone pollution in BMR.
- SMA was consistently VOC-sensitive under both wintertime (ASIA-AQ) and springtime (KORUS-AQ) conditions. VOC reductions yielded clear decreases in PO<sub>x</sub>, while NO<sub>x</sub> reductions alone led to increased ozone production.

We calculated the orthogonal distance, defined as the shortest distance of each observation from the PO<sub>x</sub> transition line on the isopleth. This metric indicates both the regime and the degree of sensitivity relative to the transition line. Using the distances, we analyzed the diurnal variations of PO<sub>x</sub> sensitivity and found a general pattern across tropical and subtropical Asian megacities: distances were closer to the transition line or slightly VOC-sensitive in the morning but shifted toward stronger NO<sub>x</sub> sensitivity in the afternoon. In addition, we applied the orthogonal distance to compare with the widely used HCHO/NO<sub>2</sub> ratio derived from *in situ* observations for ~~the~~ identifying the PO<sub>x</sub> sensitivity regime. ~~Both~~ The relationship between the two indicators showed ~~a~~ negative correlation with transition occurring at an HCHO/NO<sub>2</sub> range of approximately 0.3–1.0, notably lower than the previously reported HCHO/NO<sub>2</sub> thresholds (1–2).

While this study provides important insights into the chemical sensitivity of ozone formation, several limitations should be noted. The analysis is based on changes in precursor concentrations rather than emission-based controls, and it focuses on chemical production rates rather than ambient mixing ratios that also influenced by transport, mixing, and meteorology. As a result, the inferred control responses should be interpreted as chemical tendencies rather than direct predictions of surface ozone changes. Future work using 3-D chemical transport models will be necessary to fully capture the impacts on real-world air quality by perturbing precursor emission controls. In addition, given that only a subset of VOCs is

585 realistically controllable, rigorous VOC source apportionment is required to identify and prioritize the controllable, high-  
reactivity sources and to bound the attainable benefits. Nevertheless, this study highlights ozone formation sensitivity varies  
substantially across Asian megacities, emphasizing that coordinated NO<sub>x</sub> and/or VOC reductions are essential to effectively  
mitigate urban O<sub>3</sub> pollution.

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## Data and availability

The airborne observational data from the ASIA-AQ campaign used in this study are publicly available at  
595 <https://doi.org/10.5067/SUBORBITAL/ASIA-AQ/DATA001>.

## Competing interests

The authors declare that they have no conflict of interest.

## Author Contributions

600 CC and AF designed the analysis, performed the PO<sub>x</sub> isopleth and sensitivity calculations, and led the interpretation of the results. CC wrote the first draft of the manuscript with substantial inputs from AF. KEM and FF provided additional scientific guidance on the PO<sub>x</sub> isopleth analysis and contributed to the discussion and revision of the manuscript. JD and JC, as ASIA-AQ principal investigators, led the overall planning and execution of the campaign. KL, CO, SH, KU, ECA, AJH, RSH, BR, DJ, GSD, YC, JPD, JM, GMW, TFH, JMSC, ERD, AS, RAH, POW, KB, YRL, LGH, DJT, LA, DRB, NJB, BB, SM, KEM, HK, WN, AW, FP, and WW contributed to the execution of the campaign and to the provision and quality control of the  
605 observational datasets used in this study. All authors discussed the results, commented on the manuscript, and contributed to its final version.

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