

Response to the Referees #1 and #2 for the “Insights on Ozone Formation Sensitivity in Southeast and East Asian Megacities during ASIA-AQ”

Cho et al.

We would like to thank the reviewers for their careful reading of the paper and for providing useful comments which improved the manuscript. We reply here to each reviewer comment (comment shown in italics). Blue text represents the corresponding correction in the manuscript.

Anonymous Referee #1

Comment #1: The choice of a 750 m altitude threshold to represent near-surface conditions is reasonable and supported by vertical profiles (Fig. 2). However, the authors should test the sensitivity of this altitude to selecting another altitude and how it would impact the 4 cities considered in this study, particularly given potential vertical gradients in NO_x, VOCs, and photolysis rates. A brief sensitivity test or justification based on boundary-layer height variability would improve the methodology.

Response #1: We thank the reviewer for this helpful suggestion. To evaluate the sensitivity of the analysis to the selected altitude threshold, we performed additional simulations using precursor constraints derived from observations below alternative altitude thresholds (500, 750, 1000, and 1500 m). This sensitivity analysis was conducted using the flight steady-state (FSS) configuration of the box model (Sect. 2.2 and SI Sect. S2). The results show that the resulting PO_x estimates are broadly consistent across the four cities for thresholds between 500 and 1000 m. Larger differences appear only when extending the threshold to 1500 m, where observations include air masses increasingly influenced by conditions above the boundary layer. These results support the use of 750 m as a reasonable threshold for representing near-surface urban air masses. The results of this sensitivity test are now provided in the Supplement (Table S3 and text), and a brief reference has been added in the manuscript. In Lines 147-148, we have revised the text to be “Inspection of precursor vertical profiles (Fig. 2) and sensitivity tests with different altitude cuts (see [Supplementary Information \(SI\) Sect. S2](#)) supports 750 m as a sensible and conservative choice that keeps the analysis centered on boundary-layer chemistry.”

Table S3. Mean and standard deviations of precursor concentrations and simulated PO_x to the selected altitude threshold for the four megacities.

Location	Altitude threshold (m)	NO _x (ppbv)	O ₃ (ppbv)	CO (ppbv)	VOCR (s ⁻¹)	PO _x (ppbv hr ⁻¹)
MM, PHL	500	4.26 ± 6.11	46.90 ± 13.51	210.44 ± 61.65	1.15 ± 1.32	10.58 ± 7.46
	750	3.72 ± 5.30	47.72 ± 15.19	205.53 ± 62.46	1.16 ± 1.39	10.44 ± 6.91
	1000	3.65 ± 5.24	47.28 ± 15.02	203.40 ± 62.95	1.15 ± 1.38	10.28 ± 6.80
	1500	3.49 ± 5.15	46.70 ± 14.80	199.66 ± 63.41	1.11 ± 1.36	9.77 ± 6.82
BMR, THAI	500	10.32 ± 7.92	51.91 ± 19.98	324.01 ± 87.44	1.60 ± 1.23	14.68 ± 7.33
	750	8.07 ± 7.24	58.06 ± 22.33	310.55 ± 81.66	1.45 ± 1.13	15.62 ± 6.38
	1000	6.58 ± 6.61	59.73 ± 20.62	308.52 ± 75.20	1.35 ± 1.05	14.35 ± 7.19
	1500	5.73 ± 6.40	60.22 ± 19.64	305.82 ± 75.46	1.25 ± 1.00	13.81 ± 7.47
TKMA, TWN	500	3.24 ± 2.73	64.28 ± 9.19	277.00 ± 61.20	0.98 ± 0.79	13.38 ± 4.67
	750	2.53 ± 3.03	62.58 ± 14.29	275.27 ± 102.85	0.78 ± 0.73	11.02 ± 5.63
	1000	2.09 ± 2.83	61.05 ± 14.20	260.39 ± 93.18	0.66 ± 0.65	8.86 ± 5.38
	1500	1.73 ± 2.70	56.73 ± 15.20	236.59 ± 93.46	0.56 ± 0.60	8.09 ± 5.29
SMA, KOR	500	12.01 ± 11.59	49.08 ± 8.69	256.32 ± 71.46	1.02 ± 1.23	1.36 ± 1.05
	750	10.66 ± 10.59	50.38 ± 8.75	248.60 ± 66.81	0.94 ± 1.16	1.30 ± 1.07
	1000	8.66 ± 9.57	51.84 ± 8.30	234.47 ± 63.81	0.79 ± 1.03	1.27 ± 1.06
	1500	5.33 ± 7.84	53.00 ± 7.53	204.76 ± 62.25	0.52 ± 0.81	0.73 ± 0.81

Comment #2: *Averaging observations over 1–2 h flight tracks may obscure rapid chemical transitions and plume heterogeneity, especially in high-NO_x urban environments. The authors should discuss the implications of this temporal smoothing for nonlinear ozone chemistry and whether shorter averaging windows were tested.*

Response #3: We thank the reviewer for this insightful comment. The potential impact of temporal averaging on the isopleth analysis was evaluated using the FSS box model described in Sect. 2.2 and Supplementary Sect. S2. Unlike the isopleth framework, which uses precursor concentrations averaged over each 1–2 h flight track, the FSS model treats each observation independently at 1-minute temporal resolution and diagnoses instantaneous PO_x along the flight track. As described in Lines 276–282, the PO_x values derived from the FSS model show generally good agreement with those inferred from the isopleth framework across regions (Slope: 0.83-1.2; Fig. S5). This comparison indicates that the main features of the PO_x sensitivity diagnosed by the isopleth

analysis are robust to the temporal averaging applied in the track-based approach. We have clarified this point in Lines 276-282 as following:

“In addition, as noted in Sect. 2.2, the representativeness of this isopleth-based PO_x estimation was further evaluated by comparison with PO_x simulated from a flight steady-state (FSS) box model, which treats each observation independently to estimate instantaneous PO_x along the flight track (Sect. 2.2 and SI Sect. S2). The two approaches showed generally good agreement across regions, indicating that the key features of the inferred PO_x sensitivity are not strongly affected by the 1–2 h temporal averaging used in the isopleth framework supporting the reliability of the isopleth framework for diagnosing PO_x sensitivity (Fig. S5). Using the FSS box model results, the contributions of each O_x production or destruction term in Eq. (1) were computed and are shown in Fig. S6. City-specific PO_x budget features are discussed in Sect. 3.2.”

Comment #3: *MCM is a widely accepted chemical mechanism. However, it can be biased due to the neglect of heterogeneous chemistry, for example, HO_2 uptake on aerosols, and the impact of halogens on secondary chemistry and radical budget. The authors can include a brief discussion on the direction and impact of these uncertainties.*

Response #3: We thank the reviewer for this helpful comment. The box model simulations in this study employ MCMv3.3.1, which provides a near-explicit representation of gas-phase VOC oxidation chemistry and is widely used in atmospheric box-model studies of ozone formation. Processes such as heterogeneous uptake of radicals on aerosols or halogen-mediated chemistry are not explicitly represented in this mechanism and could influence radical budgets under certain conditions. However, evaluating these processes quantitatively would require additional observational constraints (OH, HO_2 , halogen species, etc.) and model development beyond the scope of the present study. To acknowledge this limitation, we added a brief statement in Lines 151-154 as following: “..., was applied to represent VOC oxidation chemistry and its role in ozone formation. Processes not represented in the current gas-phase mechanism (e.g., heterogeneous radical uptake or updated halogen chemistry) may influence radical budgets under certain conditions and therefore represent a potential source of uncertainty.”

Comment #4: *"A gridded parameter space was created by scaling the observed mean mixing ratios of NO_x and total VOCs from 1 to 300%." Please clarify this scaling approach to improve reproducibility.*

Response #4: Corrected. To improve clarity and reproducibility, we have added “Total VOCs were scaled linearly in 10% increments, with all individual VOC species adjusted proportionally to preserve VOC speciation. NO_x was scaled using a step size of 0.2 in log space, while maintaining a fixed NO/NO_2 ratio.” in Lines 171-173.

Comment #5: *Section 2.3 is conceptually clear but not sufficiently detailed for full reproducibility. For example, grid resolution, ridge identification algorithm, and transition-line fitting are not sufficiently specified to allow full reproducibility of the orthogonal distance metric.*

Response #5: The grid resolution has been clarified as described in Response #4. In addition, we have revised Lines 191–195 to more explicitly describe the ridge identification procedure and transition-line fitting approach. The revised text now reads: “The ridge of the PO_x surface was identified by locating the PO_x maxima along both NO_x and VOC dimensions of the grid, and the resulting ridge points were fit using a least-squares linear regression to define the transition line. Observed data points were then overlaid onto this space to quantify their position relative to the transition line and associated PO_x sensitivity regimes, and the transition line between NO_x and VOC-sensitive regimes was defined as the ridge that follows the direction of steepest descent of PO_x values.”.

Comment #6: *Section 3.2 discusses the individual PO_x isopleths. However, the authors may consider adding a brief comparative discussion of regime differences across cities and times of day, as this could further strengthen the interpretation*

Response #6: We thank the reviewer for this helpful suggestion. Following this comment, we have added a new subsection (Sect. 3.2.5, “Comparative overview of PO_x isopleths across cities and times of day”) to provide a concise cross-city and diurnal comparison of the PO_x isopleths.

“3.2.5 Comparative overview of PO_x isopleths across cities and times of day

Across the four megacities, the PO_x isopleths differ systematically, as reflected in differences in transition-line slopes (Table S4), the background isopleth patterns (Figs. 4a–7a and S7–S10), and the distributions of observed data points (Figs. 4b-7b).

The slopes of the transition line calculated from the PO_x isopleths range from approximately 3 to 40 across the four cities. MM shows the shallowest transition line (slope = 4.71, Table S3), followed by progressively steeper slopes in BMR (7.01) and TKMA (11.24). SMA exhibits a substantially steeper transition line (30.27), indicating a markedly different background isopleth geometry under colder conditions. In addition, a clear diurnal variation in the individual isopleths is evident primarily in MM, with larger transition slopes during morning tracks and decrease toward afternoon tracks. In contrast, BMR and TKMA show relatively minor diurnal changes within individual flights (except for RF12 Track 2), with transition-line slopes remaining broadly consistent across tracks. SMA shows no systematic diurnal pattern in the transition-line slopes.

The distributions of the observed data points further highlight differences in how each city samples on its isopleths. In MM, the morning tracks exhibit a relatively broad spread of observed

points across the NO_x–VOC space, whereas observations during later tracks become more concentrated and are largely confined to a narrow NO_x range, predominantly below 5 ppbv, on the NO_x-sensitive side of the transition line. In contrast, BMR and TKMA show a more dispersed distribution of observed points across the NO_x–VOC space, spanning a wider range of precursor conditions relative to the transition line. In SMA, the observed points also appear widely distributed in NO_x–VOC space. However, because of the substantially steeper transition line, nearly all observations fall on the VOC-sensitive side of the isopleths despite this apparent spread.”

Comment #7: *The FSS model and its setup should be included in the methodology section.*

Response #7: We have revised Sect. 2.2 to explicitly describe the flight steady-state (FSS) model configuration within the methodology section as follows:

In Line 144, we have removed “~~averaged within each flight track (typically spanning ~1–2 hours per track)~~”.

In Lines 169-183, we have added “For the construction of PO_x isopleths, the model was constrained over each megacity using precursor concentrations averaged within each flight track (typically spanning ~1–2 hours per track). A gridded parameter space was created by scaling the observed mean mixing ratios of NO_x and total VOCs from 1 to 300%. Total VOCs were scaled linearly in 10% increments, with all individual VOC species adjusted proportionally to preserve VOC speciation. NO_x was scaled using a step size of 0.2 in log space, while maintaining a fixed NO/NO₂ ratio. Each NO_x–VOC grid point was used as input to the box model to compute PO_x. The NO_x–VOC grid enabled simulation of a wide range of conditions, reflecting both low and high precursor levels, while holding their speciation constant.

In addition to the isopleth simulations, a flight steady-state (FSS) configuration was implemented using the same chemical mechanism and constraint strategy described above (see Supplementary Sect. S2 for details). Unlike the isopleth approach, which applies track-averaged precursor constraints to construct a gridded NO_x–VOC scaling space, the FSS simulations were performed using time-resolved in-situ observations along the aircraft track (1 minute average). For each data point, the model was run to achieve steady state using the corresponding observed constraints to diagnose instantaneous PO_x along the flight track. This configuration enables evaluation of PO_x without applying precursor scaling or track-averaging, thereby providing a complementary perspective on how the 1–2 hour averaging and scaling assumptions used in the isopleth framework influence the inferred PO_x values.”

Modify title of Section 2.3 “~~Isopleth construction~~ **projection** and orthogonal distance calculation”.

In Lines 186-188, we have removed “~~A gridded parameter space was created by scaling the observed mean mixing ratios of NO_x and total VOCs from 1 to 300%. This grid enabled simulation~~”

of a wide range of conditions, reflecting both low and high precursor levels, while holding their speciation”

Comment #8: *The authors state that the box model is constrained using meteorological variables and photolysis rates. Please clarify the source of these parameters (e.g., in situ aircraft measurements, radiative transfer calculations, or model-derived values), and briefly describe how they were implemented in the box model and potentially the bias associated with this.*

Response #8: Photolysis frequencies and meteorological parameters used to constrain the box model are obtained from in-situ DC-8 aircraft measurements, as noted in Line 145 and Table S2. To clarify how photolysis was implemented in the model, we added the following sentence in Lines 156-157: “In this configuration, the measured photolysis frequencies were adjusted according to the evolving solar zenith angle at 20-minute intervals.”

Comment #9: *The manuscript would also benefit from a brief discussion of how uncertainties in observed NO_x, VOCs, and photolysis rates propagate into PO_x estimates and regime classification, particularly near the transition line.*

Response #9: Measurement uncertainties of the instruments used to constrain the box model have now been explicitly summarized in Table S2 in the Supplement. To evaluate how these uncertainties propagate into the analysis, we estimated the combined uncertainty of the observed NO_x and total VOC mixing ratios and added the corresponding propagated error bars to the observational data points in the individual isopleth plots (Supplementary Figs. S7–S10).

Furthermore, to illustrate how these observational uncertainties influence the inferred PO_x reduction responses, the resulting uncertainty range has been propagated into the PO_x response curves shown in Fig. 8, where dashed lines indicate the range of possible responses associated with the measurement uncertainties in NO_x and VOC observations.

These show that the propagated uncertainty ranges do not substantially shift the observational constraints across the transition line between NO_x- and VOC-sensitive regimes. Consequently, the inferred regime classification and the relative effectiveness of NO_x versus VOC reductions remain robust within the observational uncertainty range considered in this study.

We revised the captions in Figs. S7–S10 by adding the following sentence: “Horizontal and vertical error bars represent the measurement uncertainties of NO_x and total VOC mixing ratios, respectively.” In addition, “The dashed lines indicate the range of PO_x reduction responses obtained by propagating measurement uncertainties in NO_x and VOC observations.” was added to the caption of Fig. 8.

Anonymous Referee #2

Specific comments

Comment #1: *In Introduction, Line 65-75, after discussing NO_x- and VOC-sensitive regimes, please consider adding a few sentences about the “transitional regime” because ozone formation is not either or, and more studies in recent years reveal that the formation regime at local scale could shift within a day, modulating peak ozone formation rates (Mazzuca et al., 2016; Guo et al., 2021; Tan et al., 2024; Stockwell et al., 2025). Such context will facilitate your point that “...nonlinearity makes it more difficult to navigate and inform the direction of emission reduction strategies” and better pave the road for the subsequent discussions in Section 4.*

-Mazzuca et al., 2016 (<https://doi.org/10.5194/acp-16-14463-2016>)

-Guo et al., 2021 (<https://doi.org/10.1016/j.atmosenv.2021.118624>)

-Tan et al., 2024 (<https://doi.org/10.1016/j.scib.2018.07.001>)

-Stockwell et al., 2025 (<https://doi.org/10.5194/acp-25-1121-2025>)

Response #1: We agree with the reviewer that ozone formation sensitivity is not strictly limited to two discrete regimes and that transitional conditions can occur. To highlight this point, we added a brief description of transitional regimes in the Introduction (Lines 73–75): “[Between these two regimes, transitional conditions may occur when precursor levels lie near the ridge line separating NO_x- and VOC-sensitive regimes. Under such conditions, ozone formation sensitivity can shift as precursor levels and photochemical conditions evolve, even on local and diurnal scales \(Mazzuca et al., 2016; Guo et al., 2021; Tan et al., 2018; Stockwell et al., 2025\).](#)”

Comment #2: *In Methods, Line 106-110, please consider adding more contexts about the early ozone season of East/Southeast Asian for the months of February and March. What are the knowns and unknowns from the existing literature? On the other hand, despite earlier understanding that the summer monsoon of Southeast Asia could alleviate local air pollution, recent studies identified summer ozone episodes in the region as well, especially across South China and transported to Southeast Asia (Zhou et al., 2025). Please briefly justify the 1) scientific logic behind the timing of your campaign, 2) key questions it targeted in the scope of ozone chemistry, and 3) its limitations and caveats (e.g. how representative is the winter ozone formation regime from an annual perspective?), so that the broader audience who are not familiar with the region could have a clearer picture. If deemed necessary, please consider moving a portion of this description to the Introduction.*

-Zhou et al., 2025 (<https://doi.org/10.1021/acs.est.5c10258>)

Response #2: We thank the reviewer for this insightful comment regarding the scientific rationale and seasonal context of the ASIA-AQ campaign timing. The timing of the campaign was

determined by the broader scientific objectives of ASIA-AQ, as outlined in the mission white paper, which aimed to characterize regional air quality, including both gaseous and particulate pollutants, as well as their precursors, sources, and transport processes across diverse meteorological and chemical regimes in Southeast and East Asia, rather than focusing on a single pollutant or season.

To address the reviewer's points, we have revised Section 2.1 to clarify the regional seasonal context of the observations. Specifically, the measurement period corresponds to the pre-monsoon season in the Philippines and Thailand, which is favorable for active ozone formation. In Taiwan, the campaign represents a transitional period prior to the peak ozone season, during which photochemical activity is already significant. Lastly, South Korea was in a late winter period, outside the typical peak ozone season.

We acknowledge that the campaign does not uniformly represent peak ozone season conditions across all regions, particularly for South Korea. However, this study does not aim to provide a comprehensive seasonal characterization. Instead, it focuses on the chemical regimes sampled during the campaign period. Ozone production sensitivity is primarily governed by precursor conditions, and the ASIA-AQ observations captured a broad range of these conditions across the studied regions. This allows for robust characterization of ozone production sensitivity across diverse chemical environments.

To further support the interpretation for South Korea, we additionally compare our results with observations from the KORUS-AQ campaign (April–June 2016), which represents peak ozone season conditions in the region (Sect. 4.1.3).

Lines 109-122: “The ASIA-AQ campaign, conducted in February and March 2024, targeted specific regions across major urban centers in Southeast and East Asia: the Philippines (6–15 February), South Korea (17 February–11 March), Thailand (16–25 March), and Taiwan (15 February, 13 and 27 March). [The timing of the campaign was determined by the broader scientific objectives of ASIA-AQ as described in the ASIA-AQ white paper \(https://espo.nasa.gov/asia-aq/document/ASIA-AQ_White_Paper\).](https://espo.nasa.gov/asia-aq/document/ASIA-AQ_White_Paper)

~~Briefly, For tropical and subtropical the Southeast Asian regions (Philippines, and Thailand, and Taiwan), this period corresponds to the pre-monsoon season, characterized by a relatively dry conditions and warm season with enhanced solar radiation, which are is favorable for active ozone formation photochemistry (Toh et al., 2013; Marvin et al., 2021). In Taiwan, the campaign period represents a transitional phase prior to the peak ozone season, during which photochemical activity is already significant (Lin et al., 2014). In contrast, South Korea was in its late winter season, outside the typical peak ozone season, with generally reduced photochemical activity with peak aerosol pollution. Lower temperatures and weaker solar radiation limited photochemical activity compared to the tropical other regions (Lee and Park, 2022; Kim et al., 2023). Despite these seasonal differences, the campaign sampled a broad range of NOx and VOC conditions across the~~

studied regions, enabling characterization of ozone production sensitivity across diverse chemical environments.”

Comment #3 & #4: *In Results, Section 3.2, isopleths of PO_x across a varying range of NO_x and VOC levels are used to show the sensitivity regimes of ozone formations in each area. Yet your criteria of “high production rates” is not consistent across areas (>40 ppbv/hr for MM in Line 249, >20 ppbv/hr for BMR in Line 280, >15 ppbv/hr for TKMA in Line 303, and not explicitly written out for SMA). It is even more confusing when you summarize the fractions into Table 1. The note “a” of the table specifies that “High PO_x denotes upper 50% in the PO_x distribution (above the median)” which seems again inconsistent with your text descriptions. From your Figure 4 for example, the color scale is around 0-50 ppbv/hr. Clearly your median PO_x in MM cannot be 40 ppbv/hr. Please revise and provide a less arbitrary criteria for area-by-area interpretation and intercomparison.*

Same issue for NO_x and VOCs, for example, your Line 303 describes TKMA at “moderate NO_x (4-10 ppbv) and “moderate VOCs (> 40 ppbv)” but Line 280 calls BMR under “relatively low NO_x (7-18 ppbv)” and “moderate-to-high VOC (> 40 ppbv)”. Please revise and make them consistent for interpretation.

Response #3 & #4:

We thank the reviewer and agree that the original descriptions could be confusing.

The “high production rates” reported in the text were intended to qualitatively describe regions of high PO_x within each background isopleth (contours) in each city, rather than to define a consistent quantitative threshold across regions. To clarify this, we have revised the text from “high production rates” to “regions of the highest PO_x” together with indicative values (e.g., using “e.g.”). In contrast, Table 1 is based on the statistical distribution of observed PO_x values. To remove ambiguity, we have revised the terminology in Table 1 to “Above-median PO_x” and updated the footnote to indicate that this refers to values above the median of the observed data points (shown as dots in the isopleths).

We also acknowledge that the use of qualitative descriptors such as “low”, “moderate”, and “high” for NO_x and VOC could imply inconsistent classification across regions. To address this, we have revised Section 3.2 to consistently report numerical ranges of NO_x and VOC across all regions as follows:

Lines 285-286: “...under ~~moderate~~ **high** NO_x (> 15 ppbv) ...”

Line 319: “...under ~~relatively low~~ **moderate-to-high** NO_x (7–18 ppbv) ...”

Lines 343-344: “...occur at **low-to-moderate** NO_x (4–10 ppbv) and **moderate-to-high** VOC mixing ratios (> 40 ppbv).”

Lines 375-376: "...occur at low NO_x levels below 5 ppbv and high VOC conditions (> 65 ppbv) ..."

Technical corrections:

Comment #5: *In Abstract, Line 43-47, "In contrast, Seoul, under colder and low solar irradiance conditions, exhibited a primarily VOC-sensitive regime..." is confusing. Even though Seoul is in a higher latitude compared to the other three areas, it is not necessarily the determining factor of its ozone formation regime (or at least not shown in the data analysis). The current description is prone to be mistaken as drawing connections between lower temperature & solar irradiance and the VOC-sensitive regime, which is not the case. SMA's VOC-sensitive regime could be simply due to its highest ground-level NO_x as shown in your Figure 2. Please rephrase.*

Response #5: We thank the reviewer for pointing this out. We agree that the previous wording could be interpreted as implying a causal relationship between colder temperatures or lower solar irradiance and the VOC-sensitive regime observed in Seoul. Our intention was only to describe the contrasting environmental conditions among the cities rather than to attribute the regime classification to these factors. We have therefore revised the sentence as follows:

"In contrast, Seoul, ~~under colder and low solar irradiance conditions~~, exhibited a primarily VOC-sensitive regime **associated with its higher NO_x conditions relative to the other cities**, underscoring the importance of VOC-focused strategies."

Comment #6: *Somewhere in the Methods/Results, please make it clear that your VOCR represents non-methane VOCs (NMVOC) because your Table S2 contains CH₄.*

Response #6: Corrected: Line 207: "..., and **non-methane** VOC reactivity (VOCR),..."

Comment #7: *Somewhere in the manuscript, please make it clear that your 0-D box model and the subsequent result analysis assumed all measured O₃ are secondary and formed from your co-measured precursors. In reality, transboundary vertical transport and regional transport could modulate your observed O₃ greatly. Detailed wind analysis coupled with night-time surface background O₃ analysis shall be preferred, but it might be well out of the scope of this manuscript.*

Response #7: We agree with the reviewer that the scope of the analysis should be stated more clearly. The purpose of this study is to diagnose local photochemical ozone production sensitivity, rather than to attribute the observed ozone mixing ratios themselves to local versus transported sources. We have therefore added the following statement in Lines 407-409:

“It should be noted that the analysis diagnoses local photochemical ozone production across the NO_x –VOC precursor space. Transport processes such as regional advection or vertical mixing may influence the observed ozone mixing ratios but are not explicitly represented in this study.”