

# Vertical Structure and Driving Mechanism of PM<sub>2.5</sub> and PM<sub>10</sub> Aerosols in Hefei Based on LiDAR Observations (2021–2023)

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## **Abstract:**

Aerosol pollution remains a significant environmental concern in China. However, the vertical structure and evolution of particulate matter are poorly understood due to the lack of long-term, high-resolution observations. In Hefei, the aerosols during the study period were dominated by a mixture of fine particulate matter (PM<sub>2.5</sub>) and coarse particulate matter (PM<sub>10</sub>), mainly originating from urban traffic emissions, industrial activities, and regional transport, with significant contributions from secondary inorganic aerosols and occasional dust events. To address the knowledge gap in aerosol vertical distribution during different pollution episodes, this study employed an aerosol LiDAR system with 532 nm band to investigate the vertical profile characteristics of aerosols, with a focus on comparing the stratification differences of optical properties between PM<sub>2.5</sub> and PM<sub>10</sub> pollution events over Hefei. The seasonal and diurnal variations of aerosol profiles were investigated on polluted and clean days. The relationship between near-surface particulate matter concentrations and aerosol vertical properties was analyzed at different heights, alongside the dynamic evolution of aerosol layers during typical pollution events. Our results demonstrated that the extinction coefficient (532 nm) of PM<sub>2.5</sub>-polluted days below 0.6 km was approximately three times that of PM<sub>10</sub>-polluted days. In contrast, the depolarization ratio of PM<sub>10</sub>-polluted episodes remains consistently higher than

28 that of PM<sub>2.5</sub>-polluted cases throughout the entire observed altitude range. The differences in extinction between  
29 polluted and clean days for PM<sub>2.5</sub> were most pronounced below 0.9 km and subsequently decreased as altitude  
30 increased, whereas the differences in PM<sub>10</sub> remained significant below 1.2 km. For PM<sub>2.5</sub>, the strongest enhancement  
31 appeared between 7:00 and 14:00 (Beijing time, BJT). A subtle lifting with height was observed around midday.  
32 PM<sub>10</sub>-polluted days were characterized by a greater vertical extension of high aerosol extinction (up to ~1.2–1.4 km)  
33 but a shorter duration of strong extinction. In contrast, PM<sub>2.5</sub>-polluted days exhibited a persistent but vertically  
34 confined aerosol layer. The vertical wind shear (VWS) was weaker on PM<sub>2.5</sub>-polluted days compared to clean days.  
35 On PM<sub>10</sub>-polluted days, the VWS in the near-surface layer (1000–900 hPa) was significantly stronger than that on  
36 clean days, especially during the early morning and evening periods. The PM<sub>2.5</sub> pollution in Hefei was mostly  
37 contributed by temperature inversion and high relative humidity, while PM<sub>10</sub> pollution was driven by long-range  
38 transport of aerosol particles under the cold front system and dry conditions. These findings reveal complex  
39 interactions between aerosol optical properties, boundary-layer dynamics, and synoptic conditions, providing new  
40 insights into the vertical air quality processes in eastern China.

41 **Keywords:** Aerosol LiDAR, Vertical profile, Heavy pollution episodes, Vertical wind shear, Fine and coarse  
42 particulate matter.

## 43 1. Introduction

44 Air pollution remains one of the most pressing environmental challenges globally, with fine particulate matter  
45 and inhalable coarse particles posing serious risks to public health and atmospheric visibility ([Chen et al., 2023](#); [Deng  
46 et al., 2023](#)). Air quality is closely related to the concentration of pollutants suspended in the atmosphere ([Wang et  
47 al., 2024](#)). Solid and liquid phase pollutants, known as aerosols, not only impact air quality and visibility through  
48 multiple mechanisms but also affect the climate by altering Earth's radiation budget and water cycle processes ([Chen  
49 et al., 2016](#); [Miao et al., 2018](#)). In China, substantial improvements in air quality have been achieved in recent years  
50 through stringent emission control policies. However, severe particle pollution events still occur frequently, especially  
51 in winter haze episodes and spring dust storms across the Yangtze River Delta (YRD) region ([Han and Cao 2022](#);  
52 [Wang and Wang 2021](#)). To date, most studies have focused on surface-level air quality, with particular attention to  
53 individual PM<sub>2.5</sub> and PM<sub>10</sub> event in eastern China and urban agglomerations from ground monitoring networks ([He  
54 et al., 2021](#)). These data provide critical insights into near-surface concentrations. However, they lack information on  
55 the vertical distribution and formation mechanisms of pollutants, which is essential for understanding aerosol  
56 processes in the troposphere ([Mehta et al., 2021](#); [Mishra and Shibata 2012](#); [Wang et al., 2018](#)).

57 While extensive ground-based monitoring has provided insights into surface-level pollution, the understanding

58 of the vertical distribution and evolution of aerosols remains limited due to a lack of long-term and high-resolution  
59 vertical observations ([He et al., 2022](#); [Ou et al., 2021](#); [Shen et al., 2022](#)). The atmospheric boundary layer (ABL)  
60 plays a crucial role in regulating aerosol dynamics through physical mechanisms such as turbulence, convection, and  
61 mixing ([Gao et al., 2011](#); [Garratt 1994](#); [Tombrou et al., 2007](#)). As the interface between the surface and the free  
62 troposphere, the ABL governs the vertical exchange of energy, moisture, and pollutants. Numerous studies have  
63 demonstrated that lower ABL heights could trap pollutants near the surface, leading to elevated PM<sub>2.5</sub> levels. These  
64 lower heights are often associated with thermal inversions and stagnant synoptic conditions ([Liu et al., 2020](#); [Sun et  
65 al., 2024a](#)). Under such conditions, limited mixing suppresses the vertical dispersion of aerosols, allowing surface  
66 emissions to accumulate rapidly, particularly in urban areas with high anthropogenic activity. In contrast, strong  
67 vertical mixing and higher ABL heights enhance dispersion and dilution of pollutants, which frequently results in  
68 improved surface air quality ([Jin et al., 2021](#)). Moreover, interactions between the lower ABL and the overlying free  
69 troposphere, including vertical wind shear, subsidence, and entrainment processes, also significantly influence  
70 aerosol layering and transboundary transport ([Deng et al., 2023](#); [Li et al., 2022](#)). These interactions are significant  
71 during transition periods such as the morning boundary layer growth phase or evening collapse, which strongly affect  
72 aerosol vertical distribution ([Li et al., 2018](#)). Despite these known mechanisms, detailed observations of the vertical  
73 structure of aerosols and their relation to meteorological dynamics remain limited under varying pollution conditions  
74 and across different seasons ([Yang et al., 2025](#)). Comprehensive long-term observations are still needed to  
75 characterize how meteorology influences aerosol stratification and transformation over time.

76 Active remote sensing techniques, such as aerosol Light Detection and Ranging (LiDAR), have emerged as  
77 powerful tools to fill this observational gap ([Ansmann et al., 2013](#); [Chen et al., 2024a](#)). Unlike passive satellite-based  
78 instruments, LiDAR systems provide high-resolution vertical profiles of aerosol optical properties at fine temporal  
79 scales and under both day and night conditions, capable of capturing sub-kilometer vertical gradients and diurnal  
80 variability that are critical for understanding boundary-layer dynamics ([Wang et al., 2024](#); [Zhang et al., 2020b](#)). It  
81 enables continuous monitoring of aerosol structure, boundary layer formation, and pollutant layering in response to  
82 atmospheric dynamics ([Fan et al., 2024](#); [Fang et al., 2024](#); [Li et al., 2024](#)). Specifically, polarization-sensitive aerosol  
83 LiDAR can simultaneously retrieve extinction coefficients and depolarization ratios, which serve as indications of  
84 aerosol concentration and morphology ([Cairo et al., 2024](#); [Chen et al., 2024a](#); [Kumar et al., 2024](#)). The extinction  
85 coefficient quantifies the total attenuation of light caused by scattering and absorption by particles, and it is directly  
86 related to aerosol optical depth and visibility ([Chen et al., 2024b](#); [Sun et al., 2024b](#)). The depolarization ratio  
87 distinguishes the spherical and non-spherical particles, providing insights into aerosol types such as secondary

88 inorganic aerosols, biomass burning smoke, and desert dust. These parameters are crucial for understanding not only  
89 the optical but also the physical properties of aerosols ([Chen et al., 2024b](#); [Gebauer et al., 2024](#)). However, most  
90 previous studies using such measurements have been limited to short-term campaigns or case studies of individual  
91 pollution episodes ([Chen et al., 2022](#); [Zhong et al., 2018](#)). These attempts often lack the long-term temporal continuity  
92 and seasonal representativeness necessary for climatological or process-level understanding, particularly in regions  
93 subject to strong seasonal and meteorological variability.

94 Long-term monitoring of air pollutants is essential for the comprehensive understanding of their temporal  
95 characteristics ([Fan et al., 2021](#); [Xiang et al., 2021](#)). Unlike ground-based air quality monitoring networks, vertically  
96 resolved data are required to evaluate the entire process of pollution events, vertical mixing, and potential long-range  
97 transport ([Wang et al., 2024](#)). Passive satellite remote sensing techniques have proven useful in bridging spatial gaps  
98 in air quality data, but are limited in their ability to resolve vertical structures with high temporal accuracy due to  
99 their coarse resolution and fixed overpass times ([Chen et al., 2023](#)). Chemical transport models (CTMs) have been  
100 widely used to simulate the spatiotemporal behavior of aerosols and their interactions with meteorology ([Wang et al.,](#)  
101 [2025](#); [Xiong et al., 2023](#)). However, their accuracy is constrained by uncertainties in emissions inventories, boundary  
102 conditions, meteorological input fields, and the representation of aerosol microphysics. In addition, model  
103 performance varies significantly by region and pollution type ([Zhan et al., 2024](#)). LiDAR technology provides high  
104 temporal resolution and fine vertical resolution measurements of aerosol and meteorological variables, including  
105 aerosol extinction, backscattering, and wind profiles ([Chen et al., 2023](#); [Chouza et al., 2015](#); [Zhang et al., 2020b](#)).  
106 This capability makes LiDAR particularly suitable for investigating aerosol evolution and transport within the ABL  
107 and lower troposphere. Nonetheless, long-term time series analyses based on continuous LiDAR measurements  
108 remain limited in the literature.

109 To address these limitations, we conducted multi-year observations using an aerosol LiDAR system in Hefei.  
110 The deployed LiDAR system provides real-time vertical profiles of aerosol extinction and depolarization ratios,  
111 allowing for a detailed assessment of aerosol structure under different pollution levels. This study fills the  
112 observational gap in the long-term vertical characteristics of both fine and coarse particulate pollution over this region.  
113 We characterized the vertical and temporal evolution of aerosol properties and linked them with pollution episodes  
114 and boundary-layer processes by integrating LiDAR data with surface PM<sub>2.5</sub> and PM<sub>10</sub> concentrations,  
115 meteorological observations, and reanalysis products. Therefore, the objectives of this study were (1) to analyze the  
116 seasonal and diurnal variations of vertical profiles with aerosol extinction and depolarization data; (2) to compare the  
117 vertical structures on polluted and clean days to clarify the differences in aerosol optical properties between these

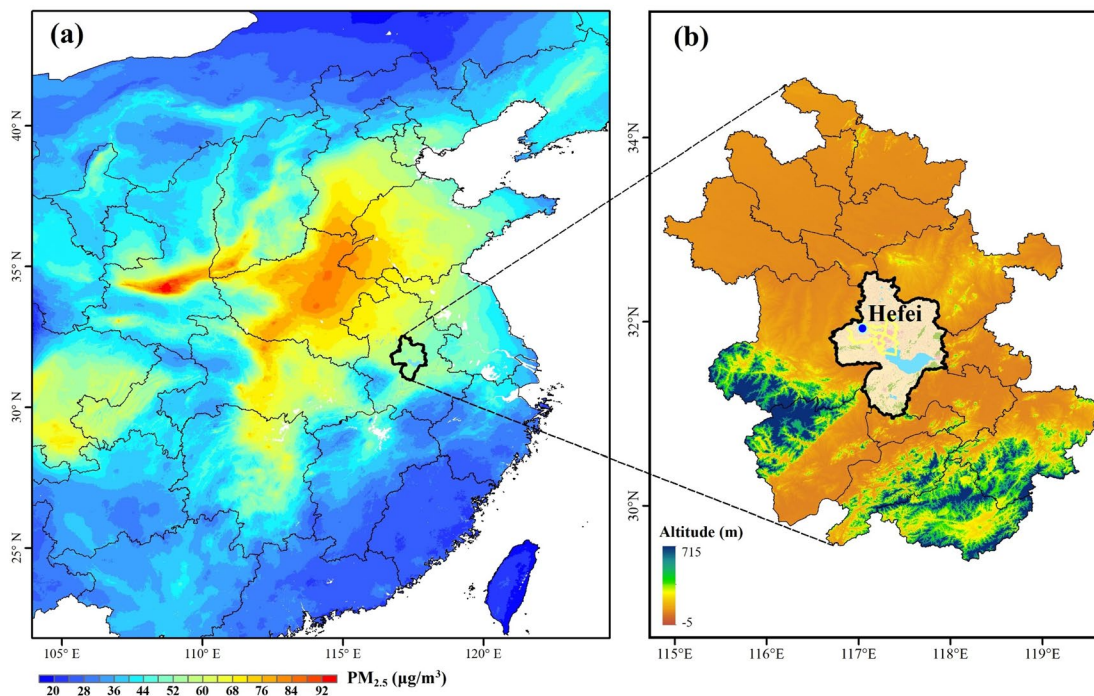
118 two conditions; (3) to assess the relationship between aerosol layering and near-surface pollutant concentrations at  
119 different heights; and (4) to investigate the spatiotemporal evolution of aerosols during specific pollution episodes  
120 and explore the mechanisms of their vertical structure, meteorological conditions, and accumulation process. The  
121 findings are expected to enhance our understanding of aerosol vertical structure in eastern China and provide  
122 scientific support for air quality management and modeling efforts.

123

## 124 **2. Data and methods**

### 125 **2.1 Study region**

126 Hefei occupies a strategic location, bridging the Beijing-Tianjin-Hebei metropolitan cluster and the Yangtze  
127 River Delta economic hub. The study area lies in a transitional climatic zone, where temperature inversions occur  
128 frequently in winter. Such conditions often result in stagnant air that limits pollutant dispersion. In addition, the  
129 geography of Hefei plays a critical role in exacerbating air quality challenges. Hefei is surrounded by the Dabie  
130 Mountains in the west and the Huangshan ranges in the south. The open plains dominate its northern and eastern  
131 frontiers (**Fig. 1b**). These natural barriers hinder the southward dispersion of airborne pollutants while trapping those  
132 transported from industrialized northern regions within the basin. The unique combination of geographic constraints  
133 and atmospheric conditions has made Hefei a focal point for studying complex pollution mechanisms. Researchers  
134 here focus particularly on the synergistic effects of regional transportation, local emissions, and meteorological  
135 drivers on particulate matter formation ([Fang et al., 2024](#); [Huang et al., 2016](#); [Liu et al., 2024](#)).



136

137 **Figure 1.** The spatial distribution of (a) the averaged  $PM_{2.5}$  concentration during winter from 2021 to 2023 (Ministry  
 138 of Environmental Protection in China, <https://air.cnemc.cn:18007/>) and (b) the location of the study area (Hefei,  
 139 China). The digital elevation model (DEM) data are derived from the NASA Shuttle Radar Topography Mission  
 140 (SRTM) 30 m product (<https://glovis.usgs.gov/>), and the land use map of the study area is obtained from the 30 m  
 141 resolution annual China Land Cover Dataset (CLCD) dataset. The monitoring station is marked by the blue dot.

## 142 2.2 Aerosol LiDAR data

143 The LGJ-05 aerosol LiDAR combines traditional radar technology and modern laser technology, with 355 nm,  
 144 532 nm, and 1064 nm as detection light sources. Measurements at 532 nm were utilized exclusively for this study.  
 145 The LiDAR emits a laser pulse into the atmosphere to capture backscattered signals, thereby obtaining key aerosol  
 146 optical characteristics. As a polarization-sensitive system, it can retrieve extinction coefficients and depolarization  
 147 ratios to characterize aerosols in the atmosphere. During the transmission process, the laser pulse is scattered and  
 148 extinguished by atmospheric aerosol particles. The intensity of backscattered light at different altitudes correlates  
 149 with the scattering and extinction properties of aerosol and cloud particles at those altitudes. The backscattered light  
 150 from spherical particles can maintain its polarization characteristics of the emitted pulse since the emitted laser pulse  
 151 is linearly polarized. In contrast, non-spherical particles (such as dust particles and ice crystals in cirrus clouds)  
 152 depolarize the backscattered light due to their irregular shapes and asymmetric interaction with linearly polarized  
 153 laser pulses. The LiDAR could detect the echo signals of the parallel and vertical components in the backscattered  
 154 light, enabling acquisition of the vertical profile of the depolarization ratio of atmospheric aerosol particles. It could

155 enable three-dimensional monitoring of the atmosphere in real time through active remote sensing, with the capacity  
156 to invert spatiotemporal distribution information such as aerosol extinction, depolarization, and water vapor mixing  
157 ratio in the atmosphere. The aerosol extinction coefficient reflects the attenuation capacity of aerosol particles to  
158 incident light, while the depolarization ratio indicates the non-spherical nature of aerosol and cloud particles. A higher  
159 extinction coefficient typically indicates a larger concentration of aerosol particles or stronger light-  
160 scattering/absorbing properties. However, greater depolarization values denote a higher abundance of non-spherical  
161 aerosol particles (e.g., dust, rugged industrial aerosols) or mixed-phase cloud constituents. By contrast, lower  
162 depolarization ratios indicate a dominance of spherical particles or simpler particle structures.

163 Since the instrument is capable of receiving data below 30 km, the effective detection range of the aerosol  
164 LiDAR can cover 0.2-6 km and 0.2-15 km for daytime and nighttime, respectively. It offers a vertical spatial  
165 resolution of 7.5 meters and a temporal resolution of 10 minutes. The observations were conducted at the National  
166 Meteorological Observation Station in the northwest of Hefei (117.06°N, 31.96°E) from March 2021 to May 2023.

167

### 168 **2.3 Ground-based observational data**

169 Hourly concentrations of PM<sub>2.5</sub> and PM<sub>10</sub> were obtained from ground-based monitoring stations in Hefei, which  
170 are operated and maintained by the China National Environmental Monitoring Centre (CNEMC) ([Liu et al., 2017](#)).  
171 These pollutant data were obtained from professional monitoring instruments, with the LGH-01E aerosol mass  
172 concentration monitor used for PM<sub>10</sub> and the LGH-01B for PM<sub>2.5</sub>, both of which apply the beta attenuation method.  
173 All instruments are calibrated regularly according to national standards, and the data undergo strict quality control  
174 procedures, including hourly, daily, and annual audits, as described in the China Environmental Monitoring Quality  
175 Assurance and Quality Control Manual. It should be noted that we focused on PM<sub>2.5</sub> and PM<sub>10</sub> concentrations during  
176 winter and spring, respectively. All comparisons between PM<sub>2.5</sub> and PM<sub>10</sub> pollution episodes were conducted across  
177 both seasons. The hourly resolution of the dataset allows for capturing diurnal variations in pollutant levels, while  
178 the multi-pollutant data facilitates analysis of the differences in meteorological responses of pollutants in Hefei.

179 The observed meteorological variables were obtained from the China Meteorological Administration (CMA)  
180 (<http://data.cma.cn/en>), measured by the Vaisala PTB210 Digital Barometer for atmospheric pressure, the Vaisala  
181 HMP155A Temperature and Relative Humidity Probe for air temperature (T) and relative humidity (RH), the EL15-  
182 2C Wind Direction Sensor for wind direction (WD), and the EL15-1C Wind Speed Sensor for wind speed (WS). All  
183 meteorological instruments are routinely maintained and calibrated to ensure the accuracy and reliability of  
184 observational data.

## 2.4 ECMWF reanalysis data

This study aims to investigate the impact of the synoptic system on PM<sub>2.5</sub> and PM<sub>10</sub> pollution. The ERA5 reanalysis dataset, freely accessible via the Copernicus Climate Change Services platform (<https://cds.climate.copernicus.eu/datasets>), serves as the data source. Vertical atmospheric data within the ERA5 dataset are interpolated to 37 pressure levels. This interpolation yields comprehensive data ranging from the Earth's surface to the upper atmosphere. Given its high spatial and temporal resolution, the ERA5 dataset has been widely employed in extreme weather events, climate prediction, and air pollution studies (Fan et al., 2021; Zhang et al., 2020b).

This study collected hourly geopotential height, vertical wind velocity, temperature, and humidity data. The data correspond to multi-pressure levels (500 hPa, 700 hPa, 725 hPa, 750 hPa, 775 hPa, 800 hPa, 825 hPa, 850 hPa, 875 hPa, 900 hPa, 925 hPa, 950 hPa, 975 hPa, and 1000 hPa). The spatial resolution of our dataset was 0.25°×0.25° (zonal and meridional) and supports a detailed and accurate exploration of the synoptic-particulate matter relationship. The data were extracted for winter and spring during 2021–2023, consistent with the periods of the surface pollutant and aerosol LiDAR observations.

## 2.5 Quality control

To explore the potential of aerosol LiDAR observations and improve data reliability, this study extracted the extinction coefficient and depolarization ratio data at 532 nm from the LiDAR system for the period from March 2021 to May 2023. In the quality control process, specific criteria were established to ensure the reliability of the retrieved parameters. Specifically, it covers three steps for outlier detection, removal of spurious points, and temporal consistency analysis. Outlier detection was conducted to eliminate all records outside the normal value range according to the parameter intervals provided by the instrument, with specific thresholds applied based on the signal-to-noise ratios (SNR) of the parallel (P) and perpendicular (S) channels. The extinction coefficient at 532 nm was calculated using data from the P channel, which was considered reliable only when the corresponding SNR exceeded three. The depolarization ratio was derived from the P and S channels, and data were accepted only when both channels had SNRs greater than three. The P channel detects linearly polarized light in the parallel direction, while the S channel detects linearly polarized light in the perpendicular direction. Based on the time series of SNR records, the individual extinction and depolarization data were cross-referenced. Temporal consistency analysis was conducted by identifying records that deviated by more than three standard deviations from the mean as outliers. The original LiDAR data have a 10-minute temporal resolution. After quality control, these data were averaged to hourly

215 resolution to be consistent with the hourly ground-based pollutant observations. To further improve data reliability,  
 216 this study calculated the credibility ratio of the extinction coefficient and depolarization ratio at different altitudes.  
 217 Based on the criterion of a 60% valid data availability rate, the effective detection ranges for the 532 nm of the  
 218 extinction coefficient and depolarization ratio were 0.2~1.8 km.

## 220 2.6 Methodology

### 221 2.6.1 Definition of polluted and clean

222 In this study, the data from 2021-2023 are gathered and categorized into four seasons: spring (Mar.-May),  
 223 summer (Jun.-Aug.), autumn (Sep.-Nov.), and winter (Dec.-Feb.). Given that PM<sub>2.5</sub> pollution is most severe in winter  
 224 and sand-dust events are frequent in spring, the definitions of polluted and clean days in this study primarily focus  
 225 on these two seasons. Days in winter when the daily average concentration of PM<sub>2.5</sub> exceeds 75 µg/m<sup>3</sup> are defined as  
 226 PM<sub>2.5</sub>-polluted days. Conversely, days in winter with a daily average PM<sub>2.5</sub> concentration below 50 µg/m<sup>3</sup> are  
 227 classified as PM<sub>2.5</sub>-clean days. For PM<sub>10</sub>, spring days with daily average concentrations above 150 µg/m<sup>3</sup> are defined  
 228 as polluted days, while those below 100 µg/m<sup>3</sup> are considered clean days (Ministry of Environmental Protection of  
 229 the People's Republic of China, Technical Regulation on Ambient Air Quality Index, 2012). To eliminate the wet  
 230 deposition effect of precipitation, all observations in this study excluded precipitation daily data.

### 231 2.6.2 Calculation of Vertical Wind Shear

232 The vertical wind shear (VWS) is a crucial factor in characterizing the change in wind velocity with height. It  
 233 is defined as the variation of wind vector across atmospheric layers, quantified by the magnitude of the difference  
 234 between upper and lower levels. The VWS can be calculated using the following formula:

$$235 \text{ VWS} = \frac{\sqrt{(u_t - u_l)^2 + (v_t - v_l)^2}}{(z_t - z_l)} \times 1000 \quad (1)$$

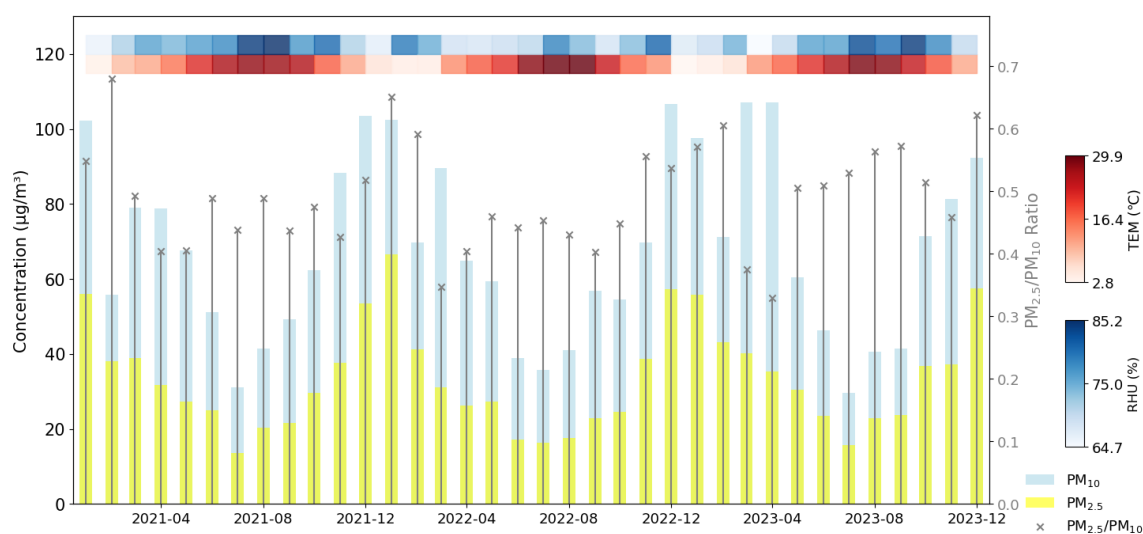
236 Where VWS is measured in units of m/s·km<sup>-1</sup>.  $u_t$  and  $u_l$  denote the zonal wind components at the upper  $z_t$  and lower  
 237  $z_l$  levels, respectively, while  $v_t$  and  $v_l$  represent the corresponding meridional wind components. This calculation aids  
 238 in quantifying the vertical dynamic forces that may impact the dispersion and transportation of pollutants.

## 239 3. Results and discussion

### 240 3.1 Variations of inter-annual PM<sub>2.5</sub> and PM<sub>10</sub>

241 **Figure 2** shows the variations in the concentrations of PM<sub>2.5</sub> and PM<sub>10</sub> pollutants from 2021 to 2023. Both PM<sub>2.5</sub>  
 242 and PM<sub>10</sub> pollutants exhibited a clear seasonal pattern, with concentrations higher in winter and lower in summer

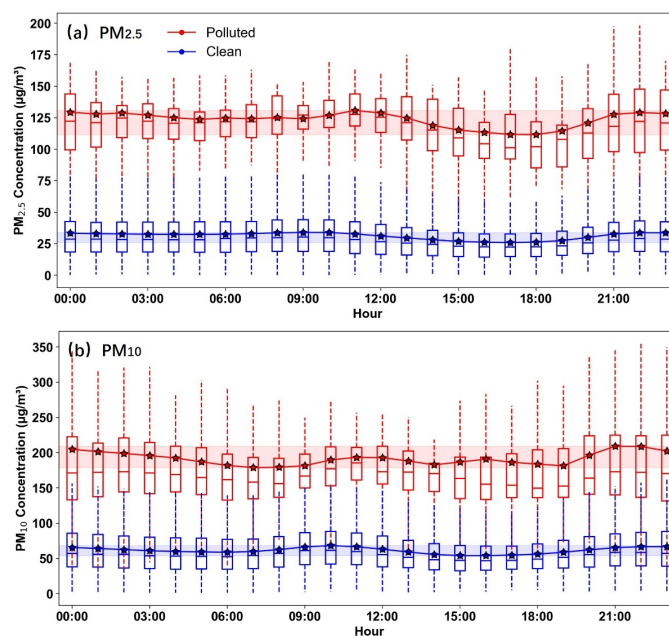
243 (Liu et al., 2021; Wang et al., 2024). In addition, PM<sub>10</sub> concentrations were relatively high during the spring months  
 244 of 2022 and 2023, likely driven by increased dust events such as sand-dust storms that commonly occurred during  
 245 this season. High PM<sub>2.5</sub>/PM<sub>10</sub> ratios ( $\geq 0.5$ ) during winter suggest a dominant contribution from fine particulate  
 246 matter, whereas lower ratios ( $\leq 0.5$ ) in spring indicate an increased presence of coarse particles (Liu et al., 2015).  
 247 The episodic PM<sub>10</sub> peaks observed in spring 2022–2023 coincided with a pronounced decrease in the PM<sub>2.5</sub>/PM<sub>10</sub>  
 248 ratio in Hefei. This phenomenon reflected intensified pollution of coarse-mode aerosols. Although winter months  
 249 also suffered elevated PM<sub>10</sub> pollution, the concurrent high PM<sub>2.5</sub> concentrations complicated the separation of dust-  
 250 related contributions. In contrast, the lower PM<sub>2.5</sub>/PM<sub>10</sub> ratios in spring facilitated the distinct observation of dust  
 251 events due to the reduced interference from fine particulate matter.



252  
 253 **Figure 2.** Monthly variation of PM<sub>2.5</sub> (marked as the yellow bars) and PM<sub>10</sub> (marked as the blue bars)  
 254 concentrations, PM<sub>2.5</sub>/PM<sub>10</sub> ratio (the grey lines with markers on top), temperature (TEM), and relative humidity  
 255 (RH) during 2021-2023 in Hefei (Beijing time, BJT = UTC + 8 h).

256  
 257 The hourly PM<sub>2.5</sub> and PM<sub>10</sub> concentrations exhibited clear diurnal variation on both polluted and clean days (**Fig.**  
 258 **3**). Polluted days showed significantly higher PM<sub>2.5</sub> levels, with hourly averages ranging from 100 to 125 µg/m<sup>3</sup>,  
 259 compared to the stable 35–55 µg/m<sup>3</sup> range observed on clean days. The PM<sub>2.5</sub> concentrations on polluted days  
 260 displayed a bimodal pattern. The concentrations remained consistently high throughout the morning period from  
 261 04:00 to 11:00 (Beijing time, BJT), with the peak at 10:00–11:00 (BJT) (Dai et al., 2020). Valleys in PM<sub>2.5</sub>  
 262 concentrations occurred at 16:00 (BJT), possibly associated with enhanced atmospheric dispersion from midday  
 263 heating that facilitates the dilution of particulate matter. Concentrations initiated an upward trend from 17:00 (BJT),  
 264 with a gradual increase culminating in a peak between 21:00 and 22:00 (BJT). In contrast, clean days exhibited a

265 subtle morning and overnight peaks. For  $PM_{10}$ , polluted days presented pronounced peaks at 10:00–12:00 and 21:00–  
 266 22:00 (BJT), coinciding with increased dust resuspension or construction activities (Yu et al., 2020). However, valleys  
 267 occurred at 07:00 and 19:00 (BJT). These may be due to reduced surface disturbances during dawn and enhanced  
 268 vertical mixing in the afternoon. Clean days maintained relatively low hourly mean  $PM_{10}$  concentrations (50–70  
 269  $\mu\text{g}/\text{m}^3$ ) with minor fluctuations. The time series curves of pollutants had different peak times and amplitudes,  
 270 indicating distinct origins.  $PM_{2.5}$  was more influenced by continuous anthropogenic emissions, while  $PM_{10}$  was  
 271 sensitive to episodic coarse-particle events (Deng et al., 2023; Wang et al., 2024). These hourly dynamics highlighted  
 272 the role of diurnal meteorological cycles in modulating particulate matter distribution, with polluted days amplifying  
 273 both primary emissions and secondary aerosol formation (Liu et al., 2021).



274  
 275 **Figure 3.** Hourly mean (a)  $PM_{2.5}$  and (b)  $PM_{10}$  concentrations on polluted and clean days (Beijing time, BJT = UTC  
 276 + 8 h). The blue and red shaded areas represent the ranges of mean concentration values for polluted and clean days,  
 277 respectively. The whiskers and boxes represent the 95th, 75th, 50th, 25th, and 5th percentiles, respectively. The stars  
 278 represent the mean values of the pollutant concentrations. The solid lines connect the hourly mean values to show  
 279 the diurnal variation trends.

280

### 281 3.2 Temporal variation of vertical distribution structure

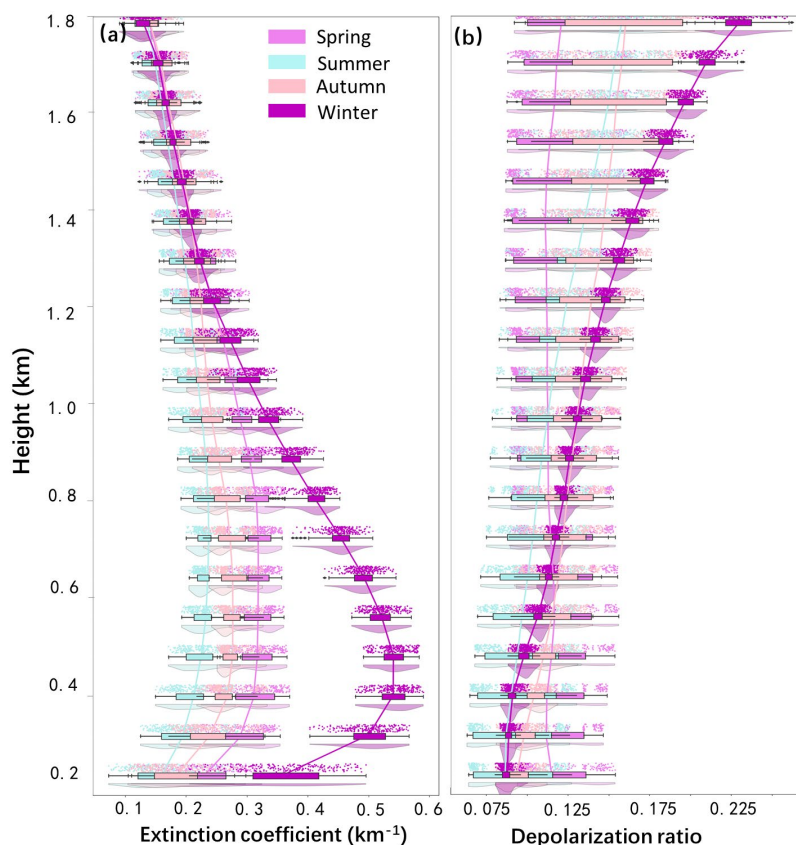
282 In addition to the impacts exerted by surface-level pollution, the investigation of vertical distribution is also of  
 283 paramount significance for a comprehensive understanding of air quality dynamics. The extinction coefficient at 532  
 284 nm shows seasonal and vertical variation. Vertically, extinction coefficients were highest near the surface and

285 gradually decreased with altitude in all seasons. This vertical decline arises from rapid reductions in aerosol loading  
286 due to vertical mixing, gravitational sedimentation, and dilution with increasing height, as well as the dominance of  
287 weaker molecular scattering above the planetary boundary layer ([Wang et al., 2021](#)). Winter presented the highest  
288 extinction values among the four seasons (**Fig. 4a**) ([Chen et al., 2023](#); [Wang et al., 2024](#)). This was likely related to  
289 enhanced emissions from domestic heating and stagnant meteorological conditions that restrict vertical mixing  
290 ([Zhong et al., 2018](#)). In contrast, summer exhibited the lowest extinction coefficients at 532 nm across the profile,  
291 suggesting the effective dispersion of aerosols due to strong convection and a high planetary boundary layer ([Li et](#)  
292 [al., 2015](#)). Spring and autumn showed intermediate levels. The most significant differences between seasons appeared  
293 below 1 km, especially in the 0.2–0.6 km range ([Zhong et al., 2018](#)). Above 1.0 km, the extinction values in all  
294 seasons converged toward high altitude, indicating a reduced aerosol presence in the upper layer ([Liu et al., 2024](#)).

295 The depolarization ratio also exhibited seasonal differences (**Fig. 4b**). With increasing altitude, the contribution  
296 of non-spherical particles increases, leading to a steady rise in depolarization ratio. This is attributed to the fact that  
297 non-spherical particles exhibit stronger anisotropic light scattering than nearly spherical fine-mode aerosols that  
298 dominate near the surface ([Wang et al., 2020b](#)). In the lower layer below 0.6 km, spring showed the highest  
299 depolarization ratios, which were attributed to a greater proportion of non-spherical particles such as mineral dust or  
300 internally mixed soot ([Biuki et al., 2022](#)). This may result from springtime dust resuspension and aged industrial  
301 emissions. In contrast, summer had the lowest depolarization ratios, implying that aerosols were dominated by more  
302 spherical particles, such as those formed through secondary processes (e.g., sulfate or organic aerosols) ([Sun et al.,](#)  
303 [2013](#)). In autumn and winter, values were moderate and stayed between the two extremes. However, the seasonal  
304 pattern changed above 0.6 km. Winter showed the highest depolarization ratio, while spring became the lowest. The  
305 vertical profile revealed that the change in particle shape with height strongly depends on seasons ([Wang et al., 2004](#)).  
306 In winter, the low ambient temperatures in the upper atmosphere promote the formation of non-spherical particles  
307 (e.g., ice crystals and irregular aerosols), which substantially enhance the depolarization ratio due to their anisotropic  
308 light-scattering properties ([Haarig et al., 2017](#); [Wang et al., 2021](#); [Yu et al., 2021](#)). In addition, elevated aerosol layers  
309 were dominated by long-range transported dust with pronounced nonspherical properties. However, aerosols in the  
310 upper layer were more spherical and less influenced by dust in spring ([Wang et al., 2020b](#)).

311

312



313

314 **Figure 4.** Seasonal variation of vertical aerosol (a) extinction coefficient and (b) depolarization profiles at 532 nm  
 315 from LiDAR measurements. The whiskers represent the 90th and 10th percentiles, respectively. The shaded areas  
 316 represent the 25th-75th percentiles. The connected lines depict the trend of mean values across different heights.

317

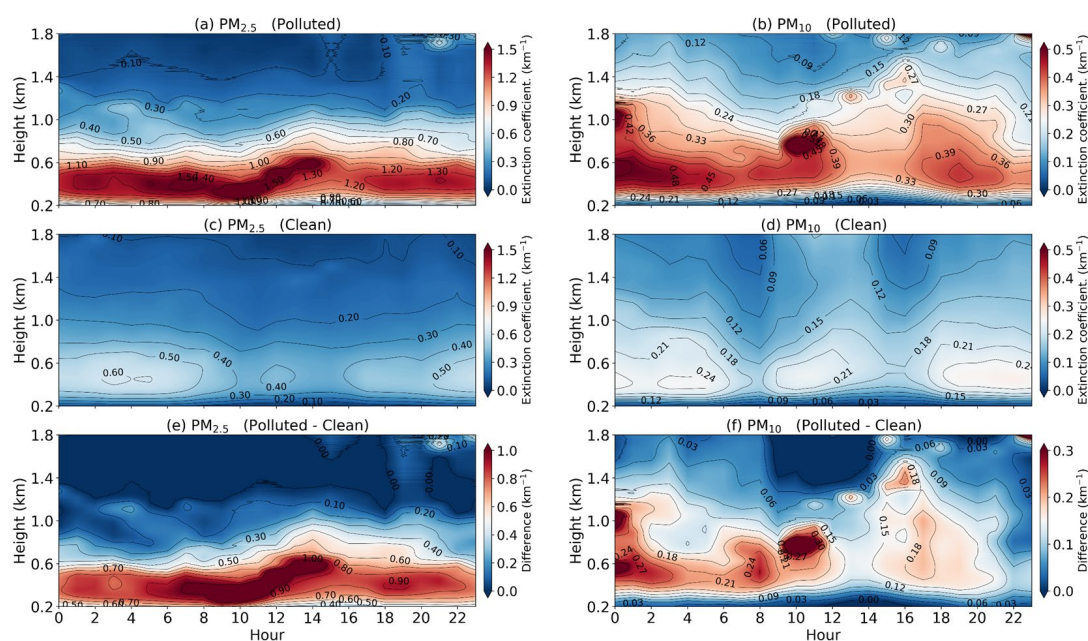
318 The extinction coefficient is significantly higher on PM<sub>2.5</sub>-polluted days than on PM<sub>10</sub> polluted days. This is  
 319 attributed to the optimal size match between fine-mode particles ( $\leq 2.5 \mu\text{m}$ ) and visible light (0.4–0.7  $\mu\text{m}$ ), as well  
 320 as their chemical components, which are more efficient at scattering and absorbing light (Wang et al., 2021; Chen et  
 321 al., 2023). The extinction coefficients at 532 nm exhibited clear differences between polluted and clean days for both  
 322 PM<sub>2.5</sub> and PM<sub>10</sub>, with strong vertical and diurnal variation (Fig. 5). For PM<sub>2.5</sub>, polluted days were characterized by  
 323 significantly higher extinction values below 0.9 km, especially during the nighttime and the period from early  
 324 morning to near noon (Wang et al., 2024). This enhancement was mainly attributed to the stable boundary layer and  
 325 limited vertical mixing during the night, which favored the accumulation of aerosols near the surface. Strong solar  
 326 radiation enhances photochemical reactions, leading to substantial secondary aerosol formation at noon and a lift in  
 327 altitude (Fig. 5a). However, extinction values near the surface (below 0.4 km) were relatively low during 12:00-15:00  
 328 (BJT), with high values concentrated mainly around 0.6 km. This midday vertical uplift of the aerosol layer  
 329 corresponds to the decrease in near-surface PM<sub>2.5</sub> shown in Fig. 3, as enhanced convective mixing lifts fine particles

330 upward and weakens near-surface extinction. Hence, strong daytime turbulence promotes the vertical transport of  
331 aerosols, resulting in an elevated and vertically expanded aerosol layer (**Fig. 5a**) ([Wang et al., 2024](#)). During this  
332 period (12:00–15:00 BJT), the enhanced extinction was concentrated around 0.6 km, extending vertically from the  
333 surface to this altitude, rather than peaking near the ground. The upward expansion of aerosols was driven by  
334 boundary layer development and persistent emission sources ([Dai et al., 2020](#)). In contrast, clean days showed low  
335 extinction values throughout the day, with little vertical variation and no significant peak. For PM<sub>10</sub>, extinction  
336 coefficients at 532 nm were also increased on polluted days (**Fig. 5b**). However, the enhancement of extinction values  
337 extended to a higher altitude compared to that of PM<sub>2.5</sub>, reaching up to 1.2 km. The primary increase occurred during  
338 the night and early morning, with a minor secondary rise observed around midday. ([Li et al., 2020](#)). Clean-day  
339 extinction remained uniformly low with minimal temporal variation (**Fig. 5d**).

340 For the depolarization ratio, distinct patterns also emerged between polluted and clean days for both PM<sub>2.5</sub> and  
341 PM<sub>10</sub> (**Fig. 6**). For PM<sub>2.5</sub>, depolarization ratios near the surface (below ~0.6 km) were relatively low on polluted days,  
342 particularly during nighttime and early morning (**Fig. 6a**). This can be attributed to the accumulation of fine, spherical  
343 particles in the stable boundary layer, which tend to have lower depolarization ratios. As altitude increased,  
344 depolarization ratios rose. This indicates the presence of more irregularly shaped or coarser particles aloft, possibly  
345 from transported dust or aged aerosols. In contrast, clean days showed uniformly low depolarization ratios across  
346 most heights, consistent with the dominance of fine, spherical aerosols under favorable dispersion conditions (**Fig.**  
347 **6c**). For PM<sub>10</sub>, depolarization ratios at 532 nm on polluted days exhibited a different behavior (**Fig. 6b**). Starting  
348 around 8:00 (BJT), elevated depolarization values were observed across a broad vertical layer, extending up to ~1.2  
349 km throughout the day. This suggests that coarser particles, which typically have higher depolarization ratios, were  
350 more effectively mixed vertically and persisted at higher altitudes compared to PM<sub>2.5</sub>. The broader vertical  
351 distribution of PM<sub>10</sub> depolarization signals reflects the larger size and potentially different sources (e.g., resuspended  
352 dust, industrial emissions) of these particles, which are less confined by boundary layer dynamics than finer PM<sub>2.5</sub>.  
353 The increased depolarization ratio at noon during PM<sub>10</sub>-polluted days is attributed to midday boundary layer growth  
354 and frontal dust transport, elevating non-spherical coarse particle concentrations.

355 The difference between polluted and clean days highlighted a strong increase in near-surface extinction during  
356 early morning on polluted days, indicating suppressed vertical mixing and increased emission accumulation (**Fig.**  
357 **5e&f**). For PM<sub>2.5</sub>, the differences were concentrated near the ground, especially below 0.9 km. The strongest  
358 enhancement appeared between 7:00 and 14:00 (BJT). A subtle lifting with height was observed around midday.  
359 Noon solar radiation maximizes surface heating, which causes severe thermal convection. The lower atmosphere

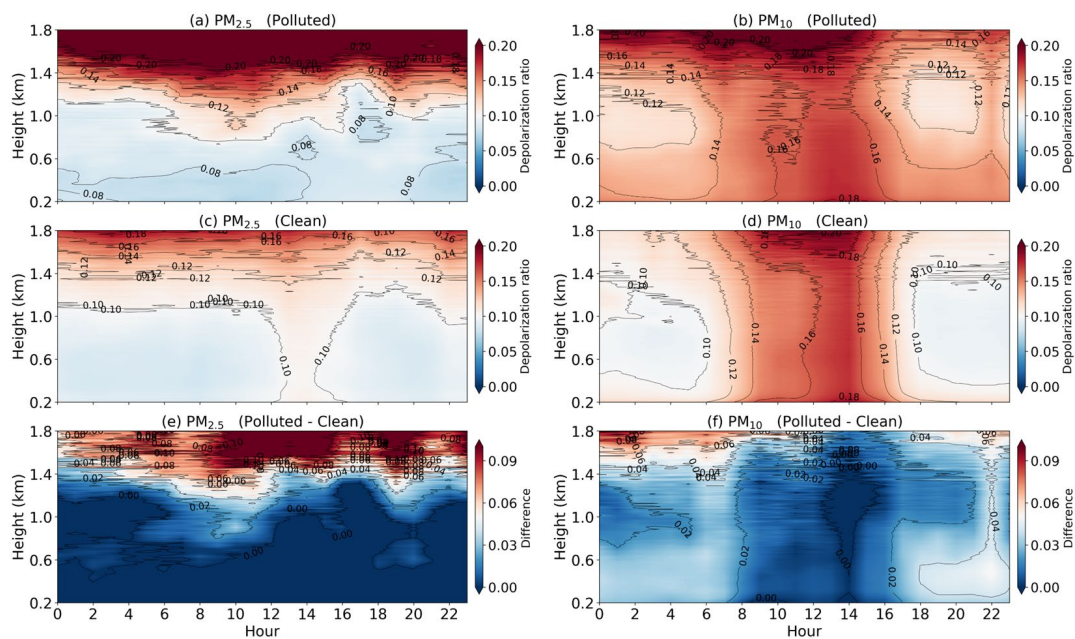
360 becomes unstable as a result of the hot Earth transferring energy upward. It could break the stable stratification and  
 361 drive upward air motion. Pollutants trapped near the surface are lifted, reducing extinction in the lower layer (<0.4  
 362 km) while increasing it aloft as vertical mixing intensifies (Wang et al., 2024). For PM<sub>10</sub>, the enhancement tended to  
 363 be localized and less vertically extended. The coarse particles may be lifted in upstream source regions, but their  
 364 vertical transport weakens significantly along the trajectory. As a result, the transport altitude of PM<sub>10</sub> reaching Hefei  
 365 was largely confined below 1.2 km (Shim et al., 2022). Compared to PM<sub>2.5</sub>, the high-value region for PM<sub>10</sub> on polluted  
 366 days has a higher vertical extension, and the period of strong extinction is shorter. The particles of PM<sub>10</sub> are relatively  
 367 large and heavy, which leads to rapid sedimentation and thus a short duration of heavy pollution. Pollutants tend to  
 368 accumulate in the middle and low altitudes (0.4-1.2 km) on polluted days.



369

370 **Figure 5.** Diurnal variation in the vertical distribution of aerosol extinction coefficient (Winter days for PM<sub>2.5</sub>: First  
 371 column; Spring days for PM<sub>10</sub>: Second column) under days with (a, b) polluted, (c, d) clean, and (e, f) the difference  
 372 between the two. All data are presented in Beijing time (BJT = UTC + 8 h) for the 532 nm channel.

373



374

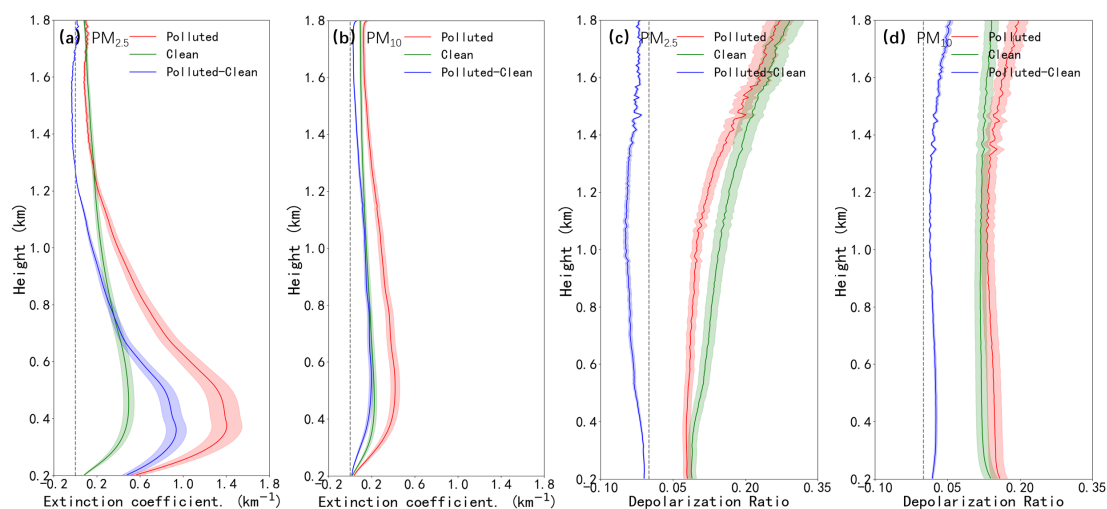
375 **Figure 6.** Same as Fig. 5, but for depolarization ratio.

376

377 To further analyze the vertical aerosol properties under different pollution levels, the mean extinction coefficient  
 378 and depolarization ratio at 532 nm were calculated across three representative conditions (**Fig. 7**). Unlike the temporal  
 379 variation that emphasized vertical profile distributions (**Fig. 5&6**), these curves highlight altitude-related trends (e.g.,  
 380 gradual changes, peak positions, and inflection points) for the comparison between polluted, clean, and difference  
 381 conditions. For  $PM_{2.5}$ , the extinction coefficient decreased exponentially with height and intensified sharply below  
 382 0.4 km on polluted days (**Fig. 7a**) ([Wang et al., 2020a](#)). The near-surface value reached up to  $1.4 \text{ km}^{-1}$  at 0.4 km,  
 383 which was nearly three times higher than that on clean days. The difference was most significant within the boundary  
 384 layer, especially between 0.3-0.6 km. For  $PM_{10}$  (**Fig. 7b**), although the overall extinction values were lower than  
 385 those of  $PM_{2.5}$ , the difference between polluted and clean days remained obvious. The polluted-clean difference for  
 386  $PM_{10}$  was more confined below 1.2 km. This discrepancy may be attributed to that the coarse-mode particles tend to  
 387 concentrate closer to the surface due to their limited vertical transport.

388 The vertical distributions of the depolarization ratio under different pollution levels are presented in **Figs. 7**  
 389 **(c&d)** for  $PM_{2.5}$  and  $PM_{10}$ , respectively. For  $PM_{2.5}$ , the depolarization ratio on clean days was higher than that on  
 390 polluted days across the entire vertical column up to 1.8 km (**Fig. 7c**). The depolarization ratio of  $PM_{2.5}$  increases  
 391 with height. The most significant difference in depolarization ratios between polluted and clean days was observed  
 392 around 1 km. Above this height, the difference narrowed but clean days still showed slightly higher values. Unlike  
 393 near-surface  $PM_{2.5}$ , which is dominated by local pollution sources,  $PM_{2.5}$  at higher altitudes is mainly associated with

394 long-range transport in the free troposphere, leading to higher depolarization ratios. Higher altitudes have more PM<sub>2.5</sub>  
 395 particles from long-range transport in the free troposphere, which possess higher depolarization ratios (Vakkari et al.,  
 396 2021; Wang et al., 2021). Even at higher altitudes, the residual influence of cleaner, non-spherical particle sources in  
 397 the background air kept the clean-day depolarization ratio higher (Li et al., 2020). For PM<sub>10</sub>, the depolarization ratio  
 398 of polluted days was higher than that of clean days across the entire vertical range (Fig. 7d), reflecting a greater  
 399 abundance of coarse-mode and irregularly shaped particles on polluted days. Below 1.2 km, the depolarization ratio  
 400 of PM<sub>10</sub> was distinctly higher than that of PM<sub>2.5</sub> on polluted days. This difference emphasized the dominant influence  
 401 of coarse, non-spherical particles (e.g., dust and mechanically suspended material) in the PM<sub>10</sub> fraction, particularly  
 402 near the surface (Shim et al., 2022). On PM<sub>10</sub>-polluted days in autumn, the depolarization ratio increased with height,  
 403 which is mainly also affected by the long-distance transport of dust in the free troposphere.  
 404

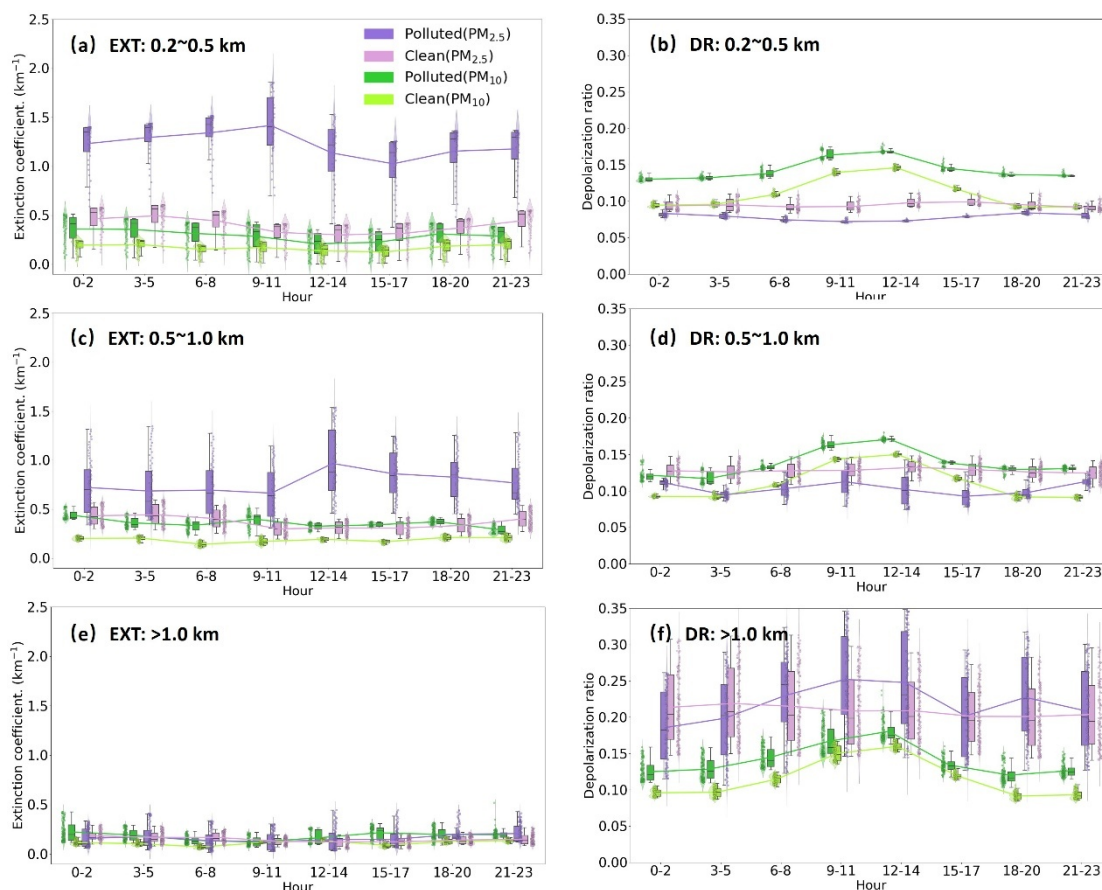


405  
 406 **Figure 7.** Vertical profiles of aerosol (a, b) extinction coefficient and (c, d) depolarization ratio at 532 nm for (a, c)  
 407 PM<sub>2.5</sub> in winter and (b, d) PM<sub>10</sub> in spring on polluted, clean days, and the difference between the two. The gray  
 408 dashed line represents the zero line. Note: The extinction coefficient and depolarization ratio of individual polluted  
 409 and clean days are positive physical quantities. Negative values in the "Polluted-Clean" panels indicate that the  
 410 aerosol optical parameters (extinction coefficient or depolarization ratio) are higher on clean days than on polluted  
 411 days. All data are presented in Beijing time (BJT = UTC + 8 h) for the 532 nm channel.

412  
 413 **Figure 8** presents the 3-hourly variation of extinction coefficient and depolarization ratio at 532 nm on polluted  
 414 and clean days across three altitude layers (0-0.5 km, 0.5-1 km, and 1-1.8 km). Below 0.5 km, the extinction  
 415 coefficients were markedly higher on polluted days for PM<sub>2.5</sub> than that on clean days, with peak values occurring at  
 416 03:00-12:00 (BJT) (Fig. 8a), which was consistent with the variation in surface pollutant concentrations (Fig. 3). The

417 depolarization ratio for  $PM_{2.5}$  was lower on polluted days in contrast with clean days. This is due to the dominance  
418 of spherical fine particles (e.g., sulfates, organic aerosols) from anthropogenic and secondary sources (**Fig. 8a&c**).  
419 These particles scatter light less directionally compared to the irregular coarse particles (e.g., dust) prevalent on clean  
420 days. In contrast, both the extinction coefficient and depolarization ratio for  $PM_{10}$  were high on polluted days ([Zhang  
421 et al., 2020a](#)). At the layer between 0.5 and 1 km, the extinction coefficient for  $PM_{2.5}$  showed a significant peak  
422 around 12:00 to 15:00 (BJT) on polluted days, which was consistent with the upward transport of pollutants to around  
423 0.6 km during the same period as shown in the diurnal vertical structure (**Fig. 5**). The extinction values decreased  
424 significantly and remained at low levels above 1 km (**Fig. 8e**). However, the depolarization ratio remained relatively  
425 high during polluted periods. The depolarization ratio in winter was higher than that of spring, mainly because  $PM_{2.5}$   
426 pollution contains a higher proportion of irregular particles from coal-fired fly ash and industrial emissions (**Fig. 8f**).  
427 The low temperatures inhibit the diffusion and settlement of coarse particles. In contrast, days in spring for  $PM_{10}$   
428 pollution were more affected by coarse particles such as windblown dust. Although these particles are irregular in  
429 shape, their large particle size makes them prone to settling. Moreover, the strong atmospheric fluidity in spring led  
430 to a slightly lower depolarization ratio compared to winter. Overall, the greatest contrast between polluted and clean  
431 days was observed in the lowest layer, where both particle concentration and shape varied most distinctly with  
432 pollution level and time of day ([Zhong et al., 2018](#)).

433



434

435 **Figure 8.** Diurnal variations of (EXT: a, c, e) extinction coefficient and (DR: b, d, f) depolarization ratio at 532 nm  
 436 on polluted and clean days for PM<sub>2.5</sub> and PM<sub>10</sub> at (a, b) 0.2-0.5 km, (c, d) 0.5-1.0km, and (e, f) above 1.0 km,  
 437 respectively. The whiskers represent the 90th and 10th percentiles, respectively. The shaded areas represent the 25th-  
 438 75th percentiles. The connected lines depict the trend of mean values across different heights. All data are presented  
 439 in Beijing time (BJT = UTC + 8 h) for the 532 nm channel.

440

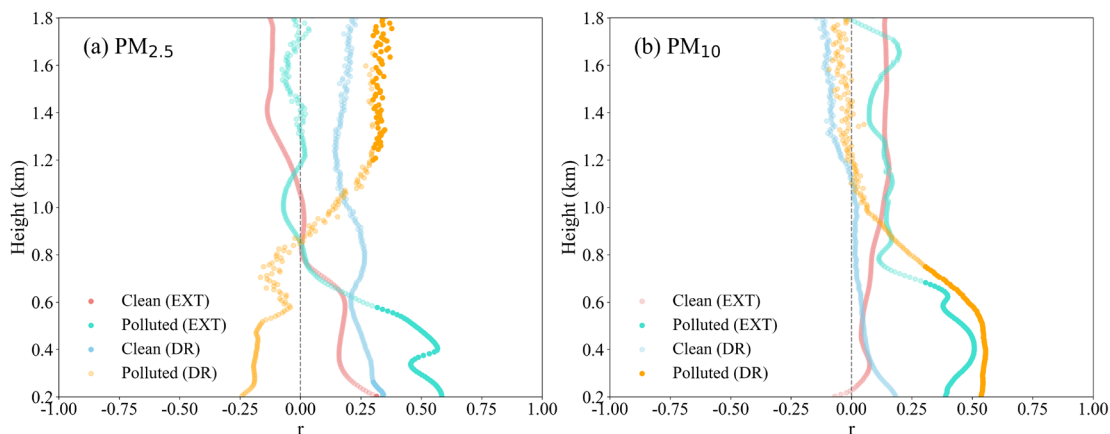
### 441 3.3 Role of aerosol extinction coefficient and depolarization ratio at different altitudes on 442 ground-based PM<sub>2.5</sub> and PM<sub>10</sub> concentrations

443 **Figure 9** shows the vertical profiles of the relationships between surface pollutant concentrations and aerosol  
 444 optical properties (extinction coefficient and depolarization ratio at 532 nm) on polluted and clean days. For PM<sub>2.5</sub>,  
 445 the extinction coefficients showed a clear positive correlation with surface PM<sub>2.5</sub> concentration throughout the profile  
 446 below 0.9 km on polluted days, marking a key boundary for vertical variation in particle composition (**Fig. 9a**). The  
 447 values on polluted days increase sharply near the surface, suggesting a local accumulation. A negative correlation  
 448 was observed between the depolarization ratio and PM<sub>2.5</sub> concentration below 0.9 km on polluted days, as elevated  
 449 PM<sub>2.5</sub> levels corresponded to a higher fraction of fine, spherical particles with low depolarization ([Zhang et al., 2020a](#)).

450 In contrast, clean days showed a consistently positive correlation between depolarization ratio and surface pollutants  
 451 throughout the vertical profile.

452 The increased depolarization levels are linked to greater proportions of non-spherical particles, such as dust or  
 453 mechanically suspended matter (Vakkari et al., 2021). The depolarization ratio showed consistently positive  
 454 correlations with surface PM<sub>10</sub> concentration on polluted days, especially below 1.2 km (Fig. 9b). On clean days, the  
 455 correlation between the depolarization ratio and PM<sub>10</sub> concentration remained near zero, reinforcing that the  
 456 depolarization ratio is a distinguishing feature of coarse-mode aerosols. These results revealed the contrasting aerosol  
 457 properties between PM<sub>2.5</sub> and PM<sub>10</sub> pollution. PM<sub>2.5</sub> events were dominated by fine, spherical particles that reduce  
 458 depolarization, while PM<sub>10</sub> events were closely associated with the presence of coarse, non-spherical particles that  
 459 enhanced depolarization signals (Biuki et al., 2022).

460



461

462 **Figure 9.** Vertical distributions of correlation ( $r$ ) between aerosol optical properties (532 nm) and ground-based  
 463 pollutants (a: PM<sub>2.5</sub> in winter, b: PM<sub>10</sub> in spring) on clean and polluted days. EXT = extinction coefficient; DR =  
 464 depolarization ratio. The gray dashed line represents the zero line.

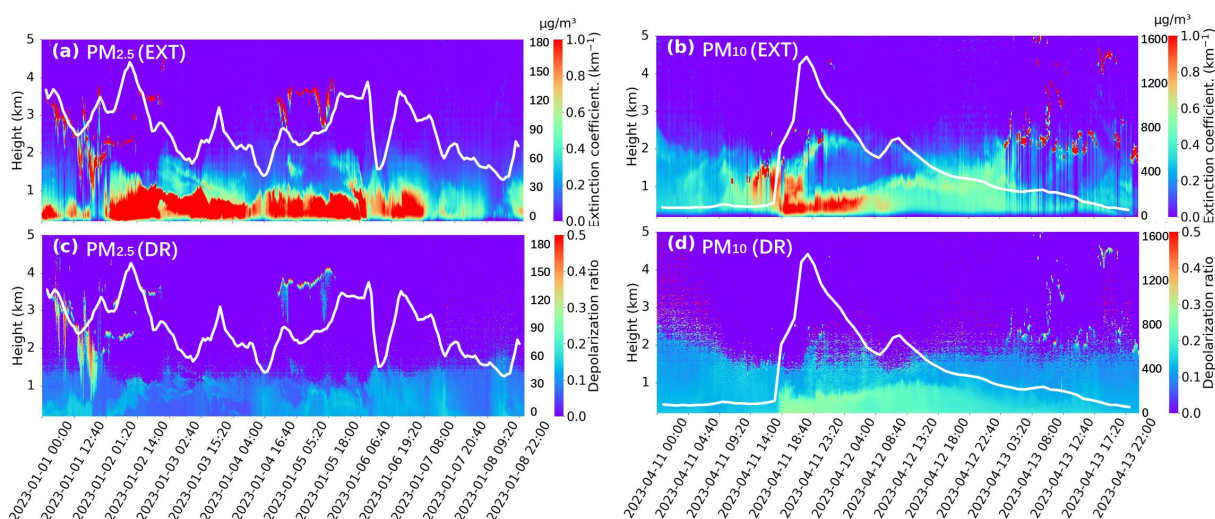
465

### 466 3.4 Analysis of LiDAR observations during heavy pollution episodes

467 Temporal variations of the extinction coefficient and depolarization ratio at 532 nm exhibited a clear association  
 468 with PM<sub>2.5</sub> and PM<sub>10</sub> concentrations, effectively highlighting severe pollution periods (Fig. 10). Severe PM<sub>2.5</sub>  
 469 pollution was evident at 10:00 on January 2, 2023 (BJT), with the maximum hourly concentration reaching 159.6  
 470  $\mu\text{g}/\text{m}^3$ . The peak of the extinction coefficient matched the actual maximum concentration of PM<sub>2.5</sub>, with fine  
 471 particulate matter exhibiting strong light-scattering properties at this time (Fig. 10a). However, the depolarization  
 472 ratio of PM<sub>2.5</sub>-polluted days was markedly decreased during this peak period (Fig. 10c). Physically, the extinction  
 473 coefficient scales with aerosol concentration and size, while the depolarization ratio is governed by particle

474 morphology. Spherical fine particles such as hygroscopic sulfates and nitrates dominating this  $PM_{2.5}$  event, leading  
 475 to increased extinction and decreased depolarization. A dramatic peak in  $PM_{10}$  concentration was observed, with the  
 476 highest hourly value of  $1440 \mu\text{g}/\text{m}^3$  recorded at 22:00 p.m. (BJT) on April 11, 2023. The severe  $PM_{10}$  event was  
 477 reflected by a pronounced increase in the extinction coefficient (**Fig. 10b**) ([Sun et al., 2013](#)). In addition, the  
 478 depolarization ratio exhibited high values during this period, which aligned with the elevated extinction coefficient  
 479 and was consistent with the irregular morphology of coarse particles that contributed to both stronger light scattering  
 480 and greater depolarization (**Fig. 10d**). Overall, the profiles of the extinction coefficient and depolarization ratio  
 481 correspond directly to the temporal patterns of  $PM_{2.5}$  (January 1–8, 2023) and  $PM_{10}$  (April 11–13, 2023)  
 482 concentrations (**Fig. 10**). Peaks in pollutant concentrations were reflected in enhanced extinction coefficients and  
 483 characteristic depolarization ratio patterns ([Zhong et al., 2018](#)). This confirms that these optical metrics effectively  
 484 identify severe pollution episodes.

485



486

487 **Figure 10.** Pollution Episodes: Temporal dynamics of (a, b) extinction and (c, d) depolarization parameters for the  
 488 532 nm channel during severe (a, c)  $PM_{2.5}$  (January 1–8, 2023) and (b, d)  $PM_{10}$  (April 11–13, 2023) periods in Hefei.  
 489 The white lines represent the hourly (a)  $PM_{2.5}$  and (b)  $PM_{10}$  concentrations, with their scale on the secondary y-axis.  
 490 All data are presented in Beijing time (BJT = UTC + 8 h). EXT = extinction coefficient; DR = depolarization ratio.

491

### 492 3.5 Analysis of meteorological conditions and their driving role in air pollution

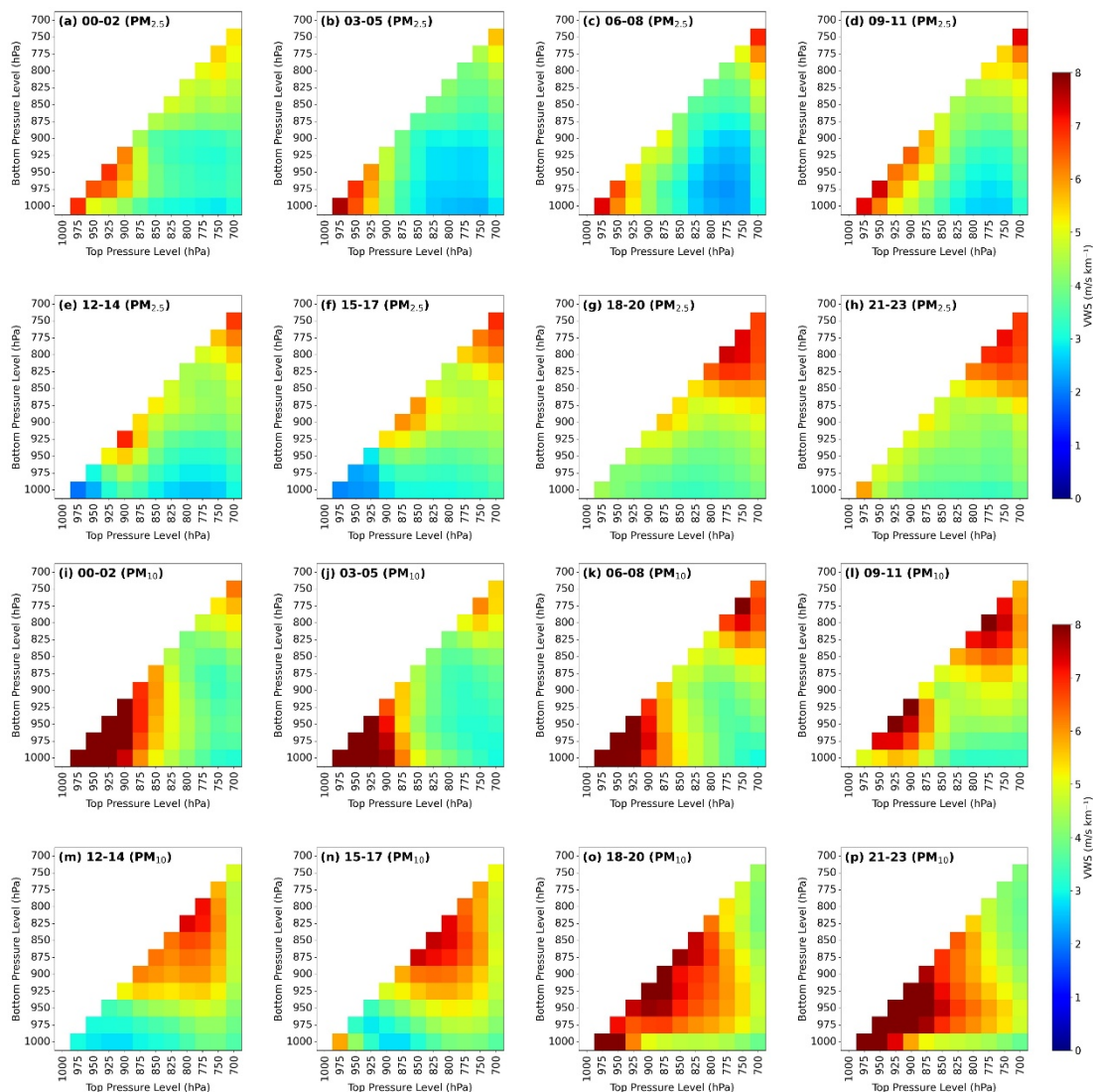
493 **Figure S1** shows boxplots of the concentrations of  $PM_{2.5}$  and  $PM_{10}$  in temperature (T), relative humidity (RH),  
 494 surface pressure (PRS), and wind speed (WS) bins, respectively. Overall, the relationships of temperature with  $PM_{2.5}$ ,  
 495  $PM_{10}$  were not significant (**Fig. S1a&e**). The RH exerted distinct influences on  $PM_{10}$  and  $PM_{2.5}$  concentrations (**Fig.**  
 496 **S1b&f**). Higher  $PM_{2.5}$  concentrations were associated with increased RH, primarily due to elevated humidity

497 promoting the condensation of gaseous precursors (e.g., sulfur dioxide, nitrogen oxides) onto pre-existing PM<sub>2.5</sub>  
498 particles (**Fig. S1b**) ([Wang et al., 2004](#)). Additionally, it accelerated secondary aerosol formation via aqueous-phase  
499 chemical reactions in the atmosphere, thereby elevating PM<sub>2.5</sub> levels ([Yang et al., 2022](#)). In contrast to PM<sub>2.5</sub>, the  
500 relationship between PM<sub>10</sub> concentrations and RH was inverse (**Fig. S1f**). Higher RH promotes moisture uptake  
501 and growth of coarse PM<sub>10</sub> particles, accelerating their sedimentation and wet deposition, thereby reducing ambient  
502 concentrations ([Gao et al., 2020](#); [Ma et al., 2023](#)). PRS tended to increase with rising concentrations of PM<sub>2.5</sub> and  
503 PM<sub>10</sub> (**Fig. S1c&g**). The average pressure values were lower on clean days (PM<sub>2.5</sub> ≤ 75 μg/m<sup>3</sup>, PM<sub>10</sub> ≤ 150 μg/m<sup>3</sup>)  
504 and became higher on polluted days. The high-pressure systems promote atmospheric stability and suppress vertical  
505 mixing, thereby facilitating the accumulation of particulate matter near the surface ([Li et al., 2015](#)). Wind speed  
506 exhibited a decreasing trend with increasing PM<sub>2.5</sub> concentrations (**Fig. S1d&h**). Lower wind speeds during heavy  
507 pollution periods indicate weak horizontal dispersion, which contributes to the persistence and accumulation of fine  
508 particles in the boundary layer. In addition, the polar plot of PM<sub>2.5</sub> concentrations shows a directional pattern, with  
509 the highest values predominantly associated with winds from the northwest (**Fig. S2a**). The cities of northern Anhui  
510 Province, Henan Province, and even further north were characterized by dense industrial activities and frequent  
511 wintertime heating emissions, which contribute to elevated PM<sub>2.5</sub> levels ([Qian et al., 2024](#); [Shi et al., 2018](#)). For PM<sub>10</sub>,  
512 the concentration distribution was more spatially confined but still shows a dominant contribution from the northwest  
513 area ([Huang et al., 2016](#)). The most intense PM<sub>10</sub> concentrations were clustered in the northwest, appearing localized  
514 and patchy (**Fig. S2b**).

515 The vertical wind shear (VWS) is a key dynamic factor affecting pollutant dispersion, as it influences mechanical  
516 turbulence and vertical mixing within the boundary layer ([Deng et al., 2023](#)). **Figures 11** and **12** illustrate the diurnal  
517 variation of VWS on polluted and clean days, respectively. For PM<sub>2.5</sub>, weak shear persisted throughout the boundary  
518 layer on polluted days when temperature inversions and stable stratification typically dominate the lower atmosphere  
519 (**Fig. 11a-h**). Clean days showed stronger shear during most periods, enhancing upward transport and dilution of fine  
520 particles (**Fig. 12**). Notably, the difference in VWS between polluted and clean days was more pronounced in the  
521 lower boundary layer (1000–900 hPa) than in the upper boundary layer (875–700 hPa) ([Zhang et al., 2020b](#)). Hence,  
522 suppressed vertical mixing near the surface contributed significantly to the accumulation of PM<sub>2.5</sub>. Surface air quality  
523 tends to deteriorate significantly due to pollutant accumulation near the ground, especially in cases where shear is  
524 insufficient in the lower layer ([Wang et al., 2024](#)), which was consistent with the trend that the extinction coefficient  
525 at 532 nm near the surface is higher than that in the upper layer (**Figs. 7 & 8**).

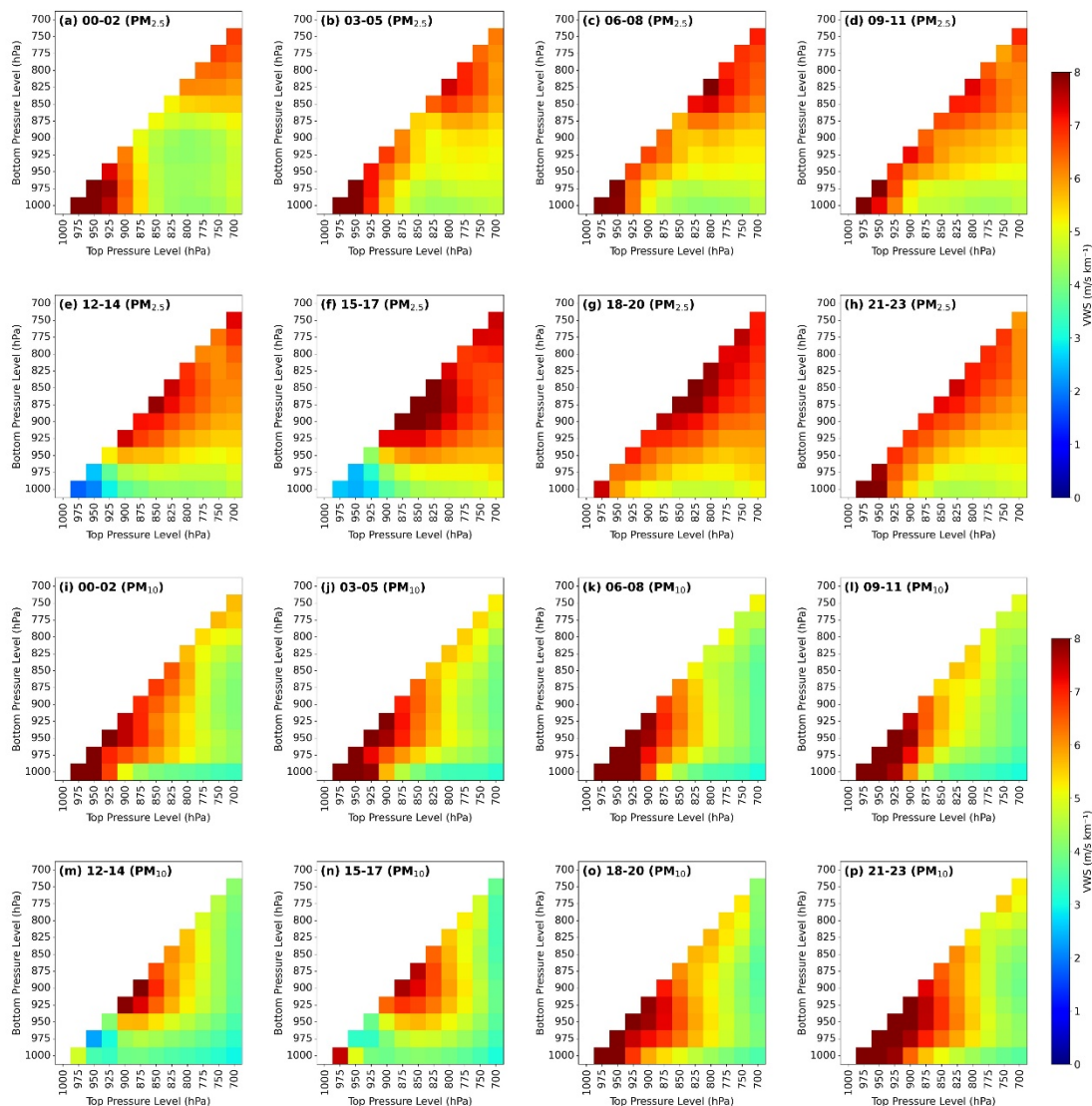
526 Contrary to the expectation that polluted days generally exhibited weak VWS, the VWS associated with PM<sub>10</sub>

527 on polluted days was not consistently lower than that on clean days (**Fig. 11 i-p, Fig. 12 i-p**). In particular, dust storms  
528 with high  $PM_{10}$  levels are typically accompanied by intensified VWS, which is conducive to dust uplift and transport  
529 under dynamic atmospheric conditions ([Yang et al., 2019](#)). The large VWS associated with  $PM_{10}$  would clear the air  
530 in the upper air at the beginning but later contributed to raising surface  $PM_{10}$  levels through regional transport at the  
531 near-surface and through downward transport of aerosol particles from the upper air. During the onset phase of dust  
532 intrusion, intensified shear facilitates both horizontal and vertical transport of dust particles ([Biuki et al., 2022](#)). From  
533 00:00 to 11:00 (BJT), the VWS in the upper layer (850 hPa-700 hPa) gradually increases on  $PM_{2.5}$ -polluted days. At  
534 night, the near-surface air cools, forming a stable temperature inversion layer that inhibits vertical mixing. Hence,  
535 the stable structure would lead to weak VWS in the upper layer. As the sun rises in the mornings, the inversion layer  
536 gradually dissipates, and vertical mixing shifts from being suppressed to active (**Fig. 11 i-l**). The VWS in the upper  
537 layer gradually increases due to the mixing effect. In addition,  $PM_{10}$  pollution often occurs during cold front passage.  
538 The vertical gradient of wind direction and speed changes suddenly near the frontal surface when cold and warm air  
539 masses converge. Once the front is located in a certain pressure range, the wind shear may be significantly higher  
540 than that in the surrounding layers. The weaker VWS between adjacent pressure layers (e.g., near 825–850 hPa) may  
541 correspond to the fault distribution (**Fig. 11 i-l**). The sinking of cold air also leads to an increase in VWS at the lower  
542 level.



543

544 **Figure 11.** Three-hourly variation of vertical wind shear for (a-h)  $PM_{2.5}$  and (i-p)  $PM_{10}$  on polluted days in Hefei at  
 545 (a&i) 00:00–02:00 BJT, (b&j) 03:00–05:00 BJT, (c&k) 06:00–08:00 BJT, (d&l) 09:00–11:00 BJT, (e&m) 12:00–  
 546 14:00 BJT, (f&n) 15:00–17:00 BJT, (g&o) 18:00–20:00 BJT, and (h&p) 21:00–23:00 BJT. Beijing Time (BJT) is  
 547 UTC+8, the local standard time used in this study. The axes indicate the vertical bounds (upper and lower pressure  
 548 levels) used to compute VWS, where the main diagonal (from top-right to bottom-left) corresponds to self-referential  
 549 values for single-layer calculations. The color gradient represents the magnitude of VWS between adjacent  
 550 atmospheric layers.



551

552 **Figure 12.** Same as Fig. 11, but under clean days.

553

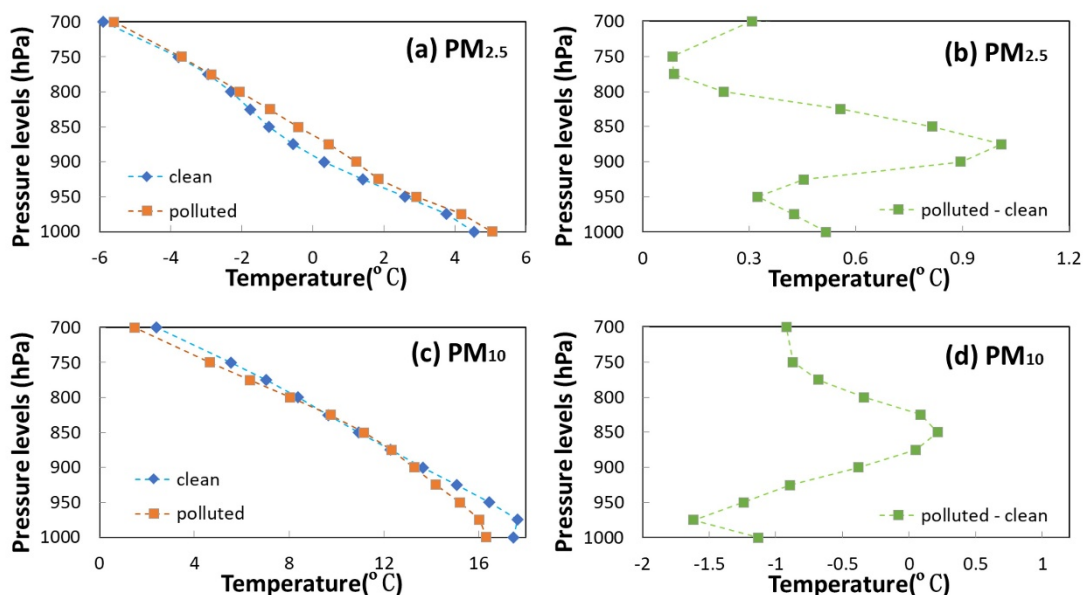
554 To clarify the synoptic influences on  $PM_{2.5}$  and  $PM_{10}$  pollution, we analyzed the distributions of air temperature  
 555 across different pressure levels from March 2021 to May 2023. During  $PM_{2.5}$  pollution days, temperatures at all levels  
 556 were consistently higher than those on clean days. The warming showed a notable pattern, with the most pronounced  
 557 trend initially observed at 875 hPa, (**Fig. 13b**). The enhanced upper-level warming relative to the lower-level led to  
 558 a thermally stable stratification, which inhibited vertical mixing and promoted the accumulation of  $PM_{2.5}$  near the  
 559 surface. In contrast,  $PM_{10}$ -polluted days exhibited negative temperature differences below 875 hPa (**Fig. 13d**),  
 560 indicating the intrusion of cold air masses. These cold air masses, often accompanied by long-range dust transport,  
 561 not only bring in  $PM_{10}$  particles from upwind regions but also suppress vertical mixing in the lower atmosphere due  
 562 to their higher density. Notably, the altitude of the cooling (875 hPa) coincided with the effective aerosol layer height  
 563 identified in our LiDAR-based analyses (**Figs. 8, 11, 12**), which supported the conclusion that dust particles were

564 primarily transported within this level.

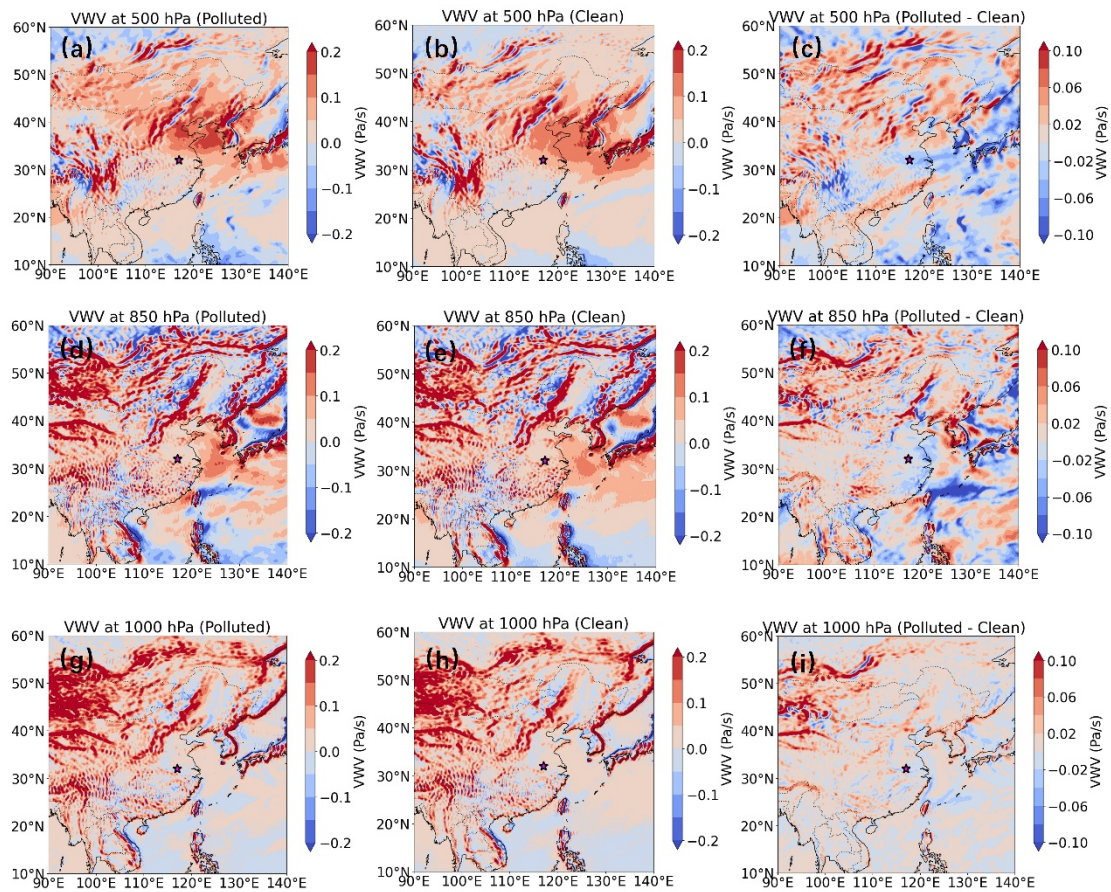
565 To further investigate the vertical thermal and humidity differences between clean and polluted days, we focus  
566 on three key pressure levels (1000, 850, and 500 hPa). During PM<sub>2.5</sub>-polluted periods, the warming at 850 hPa was  
567 greater than that near the surface at 1000 hPa in Hefei (**Fig. S3f&i**). This vertical temperature structure created a  
568 pronounced inversion, which inhibited the vertical mixing and enhanced the atmospheric stability. The lower  
569 troposphere also exhibited elevated relative humidity, particularly in Hefei and nearby regions. Elevated relative  
570 humidity suppresses turbulent exchange and reduces dilution capacity, and hence forms a shallow, moist, and stable  
571 boundary layer that favors pollutant persistence and contributes to poor air quality. In contrast, the clean days showed  
572 a lower humidity in the boundary layer and a weaker vertical temperature gradient, allowing enhanced upward motion  
573 and effective removal of surface aerosols ([Deng et al., 2023](#)). PM<sub>10</sub> pollution episodes occurred under a markedly  
574 different meteorological background (**Fig. S4**). Specifically, the relative humidity at both 1000 hPa and 850 hPa was  
575 significantly lower on PM<sub>10</sub>-polluted days than on clean days (**Fig. S4f&i**). Aloft cold air intrusions occurred on PM<sub>10</sub>  
576 pollution days, creating favorable conditions for the large-scale transport of dust into Hefei. Under such dry and calm  
577 conditions, the accumulation of PM<sub>10</sub> is likely driven by mechanical resuspension or regional dust transport rather  
578 than secondary aerosol formation processes enhanced by moisture ([Li et al., 2020](#)). Overall, these findings reveal that  
579 PM<sub>2.5</sub> episodes over Hefei are typically governed by moist and thermodynamically stable boundary layers under  
580 subsiding air masses, while PM<sub>10</sub> events are more influenced by dust-laden airflow and dry boundary-layer dynamics.  
581 The distinct thermal and humidity structures observed across the vertical profile emphasize the importance of  
582 differentiating pollution types when diagnosing meteorological drivers.

583 The vertical wind velocity was investigated at three pressure levels to reveal the dynamic mechanisms on PM<sub>2.5</sub>  
584 and PM<sub>10</sub> pollution days (**Figs. 14 & 15**). To better elucidate the meteorological characteristics between polluted and  
585 clean conditions, we calculated the differences in vertical velocity between these two scenarios, as their broadly  
586 similar large-scale background circulation patterns obscure the subtle but critical dynamic anomalies that drive  
587 pollution formation. On PM<sub>2.5</sub>-polluted days, the atmosphere around Hefei exhibits a clear subsidence trend at 850  
588 and 1000 hPa (**Fig. 14**). The subsidence reflects the existence of high-pressure systems and stagnant synoptic  
589 conditions, which contribute to the accumulation of fine particles. This inhibition of the upward motion of vertical  
590 wind corresponds with a stable thermal structure, which can effectively restrict vertical exchange and compress the  
591 boundary layer. During PM<sub>10</sub> pollution events, subsidence at 500 hPa is significantly stronger than that on clean days  
592 (**Fig. 15**). However, Hefei is located in a unique region defined by upward motion at 850 hPa, with a comparable  
593 weak upward movement also observed at 1000 hPa. The subsidence at 500 hPa stabilizes the atmosphere and

594 suppresses the upward dispersion of particles. Meanwhile, upward motion at 850 hPa and near the surface (1000 hPa)  
 595 may be caused by local convergence or orographic lifting, allowing the uplift and recirculation of dust within the  
 596 lower troposphere. This dynamic difference could reflect more complex interactions between local surface sources  
 597 and large-scale synoptic. Upward motion at 850 hPa supports the interpretation that  $PM_{10}$  events are influenced by  
 598 weak convective movements or vertical recycling of particles within a confined layer rather than stagnation alone.  
 599 These findings reveal that  $PM_{2.5}$  and  $PM_{10}$  pollution in Hefei are governed by contrasting vertical dynamic regimes.  
 600  $PM_{2.5}$  events are closely tied to a uniformly subsiding atmosphere, which favors the trapping and accumulation of  
 601 fine particles. In contrast,  $PM_{10}$  events are characterized by a layered structure, where upper-level subsidence and  
 602 lower-level ascent work in tandem to recirculate coarse particles, rather than simply trapping them. This distinction  
 603 underscores the need for targeted, pollutant-specific strategies in air quality forecasting and management.



604  
 605 **Figure 13.** Vertical distributions of temperature for polluted, clean, and their difference in (a, b)  $PM_{2.5}$  and (c, d)  
 606  $PM_{10}$ .



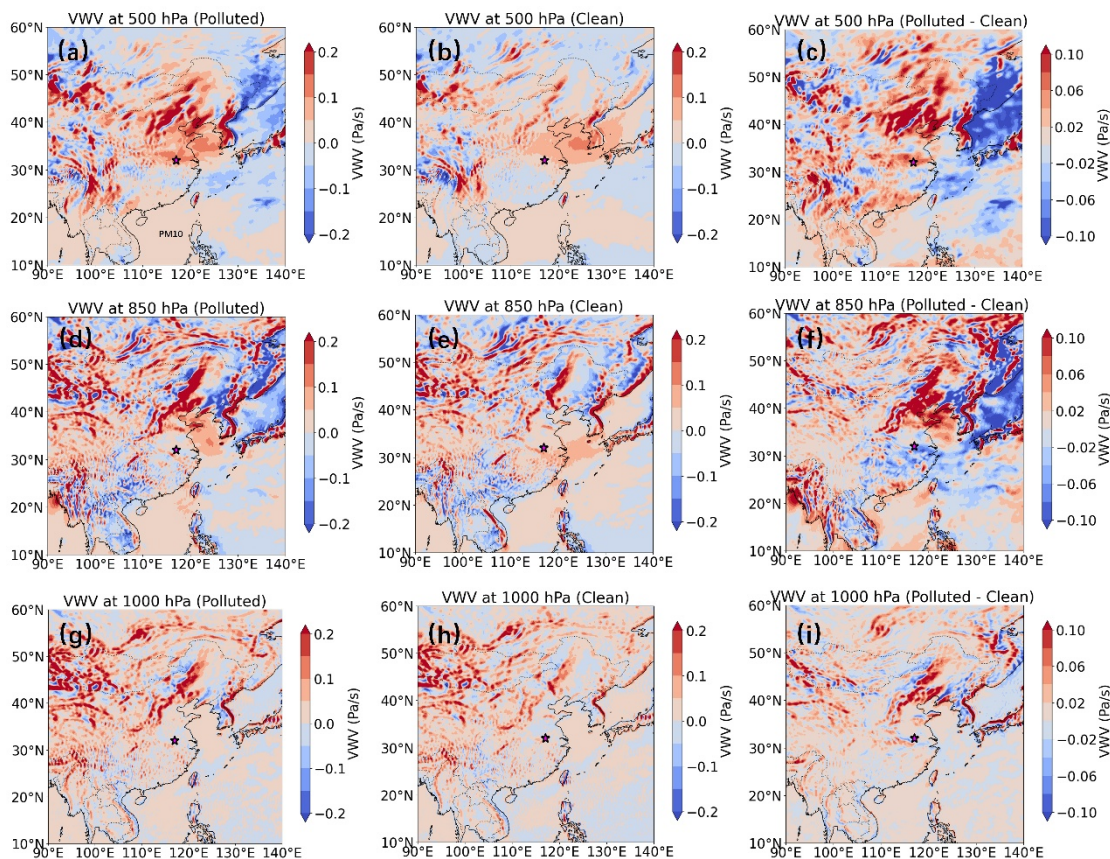
607

608 **Figure 14.** Spatial distribution of the vertical wind velocity (VWV) at three pressure levels (500 hPa: a–c;609 d–f; 1000 hPa: g–i) on PM<sub>2.5</sub>-polluted (left column), clean (middle column) days and the difference between them

610 (right column), respectively. The areas highlighted with a red star represents the location of Hefei.

611

612



613

614 **Figure 15.** Same as Fig. 14, but on PM<sub>10</sub>-polluted (left column), clean (middle column) days and the difference  
 615 between them (right column), respectively.

616

#### 617 4. Conclusions

618 This study comprehensively assesses the vertical structures of aerosols in Hefei by comparing fine (PM<sub>2.5</sub>) and  
 619 coarse (PM<sub>10</sub>) particulate matter on clean and polluted days. We utilized long-term aerosol LiDAR measurements  
 620 and reanalysis meteorological data to compare the vertical distribution characteristics of different particulate pollution  
 621 types as well as the underlying thermodynamic and dynamic mechanisms that drive their vertical development.  
 622 Aerosol extinction coefficients at 532 nm for PM<sub>2.5</sub> and PM<sub>10</sub> were consistently higher on polluted days than on clean  
 623 days. Since fine-mode particles predominate the aerosol population on PM<sub>2.5</sub> pollution days, the depolarization ratios  
 624 fell within a range lower than those observed on clean days. In contrast, PM<sub>10</sub>-polluted days had larger depolarization  
 625 ratios than clean days, indicating significant contributions from non-spherical coarse particles.

626 Our findings reveal that the accumulation of PM<sub>2.5</sub> is closely linked to stable boundary layer structures  
 627 characterized by high humidity, weak vertical shear, and a mid-level inversion temperature. These conditions  
 628 collectively inhibit vertical exchange and confine fine particles near the surface. The vertical extinction and  
 629 depolarization profiles support the dominance of aerosols during PM<sub>2.5</sub> pollution, particularly in the lower 0.9 km. In

630 contrast, PM<sub>10</sub> pollution is typically associated with low humidity and mechanically driven transport of dust from  
631 upwind sources. The vertical profiles show enhanced depolarization ratios, indicating the presence of non-spherical,  
632 coarse particles such as dust or resuspended material. Distinct vertical motion patterns identify the two types of  
633 pollution. PM<sub>2.5</sub> episodes showed upward mobility in the mid-troposphere and near-surface subsidence, which  
634 reinforced stratification and pollution trapping. However, PM<sub>10</sub> events were distinguished by upper-level subsidence  
635 coupled with low-level ascent, which might facilitate vertical transport of coarse particles during regional dust  
636 episodes. These dynamic differences highlight that PM<sub>2.5</sub> and PM<sub>10</sub> are generated by fundamentally separate  
637 atmospheric processes, and a unified treatment of particulate pollution risks oversimplifies their behavior.

638 Expanding such analyses across multiple sites would also aid in identifying geographical heterogeneity and  
639 testing the generalizability of the discovered trends. Future research may be directed toward integrating long-term,  
640 high-resolution vertical observations with advanced modeling approaches to better understand the multiscale  
641 dynamics behind various forms of particle pollution and create tailored mitigation solutions.

642

#### 643 **Data availability**

644 Hourly concentrations of major air pollutants were obtained from the China National Environmental Monitoring  
645 Centre (CNEMC) (<https://www.cnemc.cn/>). Surface meteorological variables were obtained from the China  
646 Meteorological Administration (CMA) (<http://data.cma.cn/en>). The ERA5 reanalysis dataset, used to investigate the  
647 impact of synoptic systems on PM<sub>2.5</sub> and PM<sub>10</sub> pollution, was freely accessible via the Copernicus Climate Change  
648 Services platform (<https://cds.climate.copernicus.eu/datasets>).

649

#### 650 **Author contributions**

651 YY acquired and analyzed the data and drafted the manuscript; XD conceived the study; RD and QX performed  
652 data processing and investigation; QH, YL and CW provided theoretical support; JX and YS discussed the results  
653 and revised the manuscript; YL acquired funding; all co-authors reviewed and approved the final version of the  
654 manuscript.

655

#### 656 **Competing interests**

657 The contact author has declared that none of the authors has any competing interests.

658

## 659 Acknowledgments

660 The authors thank the editor and reviewers for their constructive comments and valuable suggestions, which  
 661 have greatly improved the quality of this manuscript. We also acknowledge the China National Environmental  
 662 Monitoring Centre (CNEMC), the China Meteorological Administration (CMA), and the Copernicus Climate Change  
 663 Service for providing observation and reanalysis datasets used in this study.

## 665 Financial support

666 This work was supported by the Innovation and Development Special Project of Anhui Meteorological Bureau  
 667 (CXB202404), the 2025 Independent Innovation Scientific Research Project of Anhui Public Meteorological Service  
 668 Center (GFCX202504), the Open Fund Project for Heavy Rain (BYKJ2025D03), the Science and Technology  
 669 Projects of Xizang Autonomous Region, China (XZ202502JD0041), and the Special Project for Forecasters of Anhui  
 670 Meteorological Bureau (KY202003).

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