



1 Antarctic Sea Ice Variations and Their Linkages with 2 Global Extreme Weather Events

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8 Abstract

9 Antarctic sea ice is a critical indicator of global climate dynamics, yet its post-2015 accelerated retreat
10 and links to extreme weather/key climate modes remain insufficiently characterized. To address these
11 gaps, we analyse 2010–2024 multi-source data (NSIDC sea ice extent (SIE), CMA extreme weather
12 records, NOAA ENSO indices) via spatiotemporal decomposition, non-linear modelling, and
13 mechanistic dissection. Antarctic SIE showed a "stable-then-decline" trend: minimal variability 2010–
14 2014 (peak: 20.16×10^6 km², Sep 2014), followed by unprecedented post-2015 retreat to a Feb 2023
15 historic minimum (1.85×10^6 km², ~20% below 2010–2014 summer mean). The Amundsen–
16 Bellingshausen Sea and Antarctic Peninsula were main retreat zones, with 2020–2024 declines (10–
17 15%/20–25% winter/summer) exceeding East Antarctica (5–10%) and Ross Sea (8–12%). We
18 identified tiered negative covariance between SIE and extreme weather, presumably co-modulated by
19 the El Niño–Southern Oscillation (ENSO) and Southern Annular Mode (SAM): strong associations
20 ($R^2 \geq 0.6$, $P < 0.001$) for extreme heat (400% frequency increase) and cold waves (700% increase),
21 potentially via albedo–circulation feedback; moderate associations ($0.5 \leq R^2 < 0.6$, $0.001 \leq P < 0.01$)
22 for floods/rainstorm/typhoons, likely from ACC heat transport changes and convective propagation;
23 weak associations ($R^2 < 0.4$, $P \approx 0.05$) for blizzards, possibly due to spatially constrained El Niño–
24 SAM effects. La Niña-positive SAM may have amplified these linkages (e.g., 400% extreme heat
25 increase in 2021–2023 triple La Niña), while El Niño-negative SAM suppressed them. These findings
26 advance polar-low latitude coupling understanding, aiding extreme weather prediction and IPCC AR6-
27 aligned adaptation.

28 **Keywords:** Antarctic Sea Ice Extent; Global Extreme Weather; ENSO; Southern Annular Mode;
29 Climate Correlation

30 1 Introduction

31 Antarctic sea ice, as a key regulator of the Earth's climate system, profoundly influences global energy
32 balance and ocean circulation patterns by governing heat, momentum, freshwater, and gas exchange
33 between the atmosphere and ocean (Morioka et al., 2024; Zhu et al., 2023). Climate coupled models
34 predict that sea ice melt will cause localized surface warming, which propagates to the global
35 troposphere via the 'polar amplification effect' (Tewari et al., 2023). Satellite observations from 1979
36 to 2024 reveal a phased abrupt change in Antarctic sea ice: a slight increase from 1979 to 2012
37 (Parkinson, 2019), followed by unprecedented rapid retreat after 2015, with successive record lows



38 occurring in February 2022 and the winter of 2023. Notably, the SIE in February 2023 reached a low
39 of 1.85 million km², representing a reduction of approximately 20% compared to the 2010–2014
40 summer mean, with a retreat rate significantly exceeding that of Arctic Sea ice (Purich & Doddridge,
41 2023; Diamond et al., 2024). This ‘initial expansion followed by contraction’ pattern defies traditional
42 climate model predictions, highlighting the Antarctic sea ice’s high sensitivity to global climate change.

43 Spatially, Antarctic sea ice retreat is not uniform: during 2020–2024, the ‘core retreat zones’
44 encompassed the western Antarctic (Amundsen–Bellingshausen Sea) and the Antarctic Peninsula
45 periphery. Winter sea ice concentration in these areas decreased by 10–15% compared to 2010–2019,
46 and summer decline by 20–25%, significantly exceeding the retreat rates observed in the East Antarctic
47 (5–10%) and Ross Sea (8–12%) (Josey et al., 2024; Cordero et al., 2023). This regional differentiation
48 reflects both the synergistic interaction between local ocean circulation (such as the Antarctic
49 Circumpolar Current (ACC)) and atmospheric processes, and suggests that sea ice changes may exert
50 spatially differentiated impacts on climate systems at different latitudes (Wang et al., 2021)

51 Amidst the dramatic changes in Antarctic Sea ice, extreme weather events worldwide are exhibiting a
52 trend of increasing frequency and intensity (Intergovernmental Panel on Climate Change [IPCC],
53 2022). Observational data indicate that the frequency of global extreme heat events in 2024 increased
54 by 37% compared to the 2010–2020 average, with 2024 surpassing 2023 to become the hottest year
55 on record (Jha et al., 2025). Regional extreme events have also intensified. For instance, during the
56 winter of 2023, the number of storm days in the northern Weddell Sea and northwestern Ross Sea
57 increased by 7 days per month compared to the 1990–2015 average, with the northern Weddell Sea
58 rising from 9.1 days to 11.6 days (Josey et al., 2024). By event type, the intensity and frequency of
59 extreme heat, rainstorm, and typhoons/storms have significantly increased (Chen et al., 2023).
60 Meanwhile, extreme cold events remain more frequent and difficult to predict under global warming
61 (He et al., 2019). Between 2010 and 2024, the annual frequency of these extreme events significantly
62 exceeded the 1981–2010 baseline period, with further increases observed from 2021 to 2024,
63 underscoring the severity of global extreme climate conditions.

64 Existing research confirms a link between Antarctic sea ice changes and global extreme weather events:
65 sea ice retreat alters polar albedo and the phase of the SAM, triggering anomalies in mid-to-low-
66 latitude atmospheric circulation (such as increased north-south oscillation of the westerly jet stream).
67 This, in turn, modulates the environmental conditions conducive to extreme event generation (Bader
68 et al., 2012; Eabry et al., 2024; Cordero et al., 2023). For instance, Zhu et al. (2023) numerically
69 simulated that a 10% reduction in Antarctic sea ice correlates with an 8–12% increase in mid-to-low-
70 latitude extreme precipitation events; Josey et al. (2024) observed that during the record-low SIE in
71 2023, when tropical western Pacific sea surface temperature anomalies reached +1.5°C, typhoon
72 generation frequency increased by 40% compared to the long-term average. However, current research
73 still faces three core gaps: first, insufficient quantitative characterization of the relationship between
74 multi-scale sea ice evolution (particularly decadal variability) and extreme events; existing studies
75 predominantly focus on qualitative trends at the global scale, lacking hierarchical analyses of the
76 strength of associations between different types of extreme events (such as heatwaves, cold spells, and
77 typhoons) and sea ice indicators (SIE). Furthermore, the identification of key influence periods (such
78 as the spring SAM positive phase) remains unclear (Cordero et al., 2023). Second, the climate mode
79 modulation mechanism remains unresolved (Morioka et al., 2024; Dou et al., 2022).

80 As key Southern Hemisphere climate modes, El Niño–Southern Oscillation (ENSO) and SAM (Wang



81 et al., 2023) exert regulatory influence on sea ice-extreme event linkages, though their precise
82 pathways remain contentious — for instance, El Niño years may locally increase Ross Sea sea ice via
83 Rossby wave teleconnections, temporarily weakening the negative correlation between sea ice and
84 extreme heat (Zhang et al., 2021), while La Niña years may amplify sea ice retreat trends (Wang et al.,
85 2025). However, the specific impacts of both on regional-scale extreme events (such as the West
86 Antarctic rainstorm) remain unquantified; Thirdly, observational and modelling constraints limit
87 research depth: at the observational level, Antarctic sea ice data suffers from ‘marginal zone
88 uncertainty’—satellite sensors (such as SMMR and SSMIS) exhibit insufficient identification
89 accuracy for areas with sea ice concentration below 15%, leading to biases in the National Snow and
90 Ice Data Center (NSIDC) data for marginal seas like the Antarctic Peninsula (Crosta et al., 2021); At
91 the modelling level, while CMIP6 climate models (e.g., CanESM5, CESM2) can simulate the overall
92 negative correlation between sea ice and extreme climate events, they inadequately capture the abrupt
93 retreat events post-2016. Furthermore, their relatively low resolution (typically $1^\circ \times 1^\circ$) hinders
94 accurate characterisation of interactions between deep convection in the Southern Ocean and sea ice
95 (Heuzé et al., 2021; Morioka et al., 2024). Furthermore, existing research has paid scant attention to
96 regional extreme climate responses in areas adjacent to Antarctica (such as the West Antarctic
97 Peninsula and Ross Sea), hindering the formulation of regional-scale climate adaptation strategies.

98 This study integrates multi-source data from 2010 to 2024 (NSIDC sea ice data, China National
99 Climate Centre extreme event data, NOAA ENSO indices) within an observational-simulation-
100 mechanism three-dimensional analytical framework. It aims to achieve three core objectives: (1) To
101 characterize the spatiotemporal evolution of Antarctic SIE from 2010 to 2024, **laying the foundation**
102 **for linkage analysis**; (2) to quantify the strength of associations between different types of extreme
103 climate events and sea ice indicators, identifying key sensitive regions and time periods; (3) to
104 elucidate the **potential** modulation mechanisms of ENSO and SAM on the sea ice–extreme event
105 linkage. The findings offer new insights into the coupling mechanisms between polar and low-latitude
106 climates, the interannual variability of sea ice, and extreme weather events. They provide scientific
107 support for global extreme event forecasting and early warning systems, as well as for regional climate
108 adaptation strategies in line with the IPCC Sixth Assessment Report (AR6).

109 **2 Data and Methods**

110 **Data Sources:**

111 SIE and selected sea ice concentration (SIC) data for 2010–2024 are sourced from NSIDC (Fetterer et
112 al., 2017).

113 Data on major global weather and climate disaster events are based on monthly releases from the
114 China National Climate Centre (2010–2024), covering occurrence times, locations, and intensity
115 information for global extreme heatwaves, extreme rainfall events, floods, typhoons/storms, cold
116 waves, heavy snowfall, and tornadoes.

117 Monthly Niño Eastern Pacific Index (NEPI) and Niño Central Pacific Index (NCPI) for 2010–2024
118 are sourced from the Global Ocean Data Assimilation System (GODAS; Behringer & Xue, 2004),
119 provided by the National Centers for Environmental Prediction (NCEP) of the U.S. National Oceanic
120 and Atmospheric Administration (NOAA). The data cover latitudes 75°S – 65°N with a horizontal



121 resolution of $1^{\circ} \times 1^{\circ}$.

122 **Data processing:**

123 Using the median SIE values for 1981–2010 as the baseline, daily SIE data from 2010 to 2024 were
124 employed to calculate annual SIE daily anomalies for this period according to Equation (1).

$$125 \quad SIE_{DN} = \frac{SIE_X - SIE_A}{SIE_A} \times 100\% \quad (1)$$

126 Where: SIE_{DN} — Daily Anomaly of SIE (%); SIE_X — Known daily SIE (Unit: Millions km^2); SIE_A —
127 Daily average SIE from 1981 to 2010 (Unit: Millions km^2)

128 After obtaining the SIE_{DN} values, the annual average SIE anomaly (SIE_{YN}) from 2010 to 2024 was
129 calculated using Equation (2).

$$130 \quad \overline{SIE_{YN}} = \frac{1}{n} \sum_{i=1}^n SIE_{DNI} \quad (2)$$

131 Where: SIE_{YN} — Annual average SIE anomaly

132 The annual average SIE (SIE_Y) was obtained using Equation (3)

$$133 \quad \overline{SIE_Y} = \frac{1}{n} \sum_{i=1}^n SIE_{Di} \quad (3)$$

134 Where: SIE_Y — Annual average SIE; SIE_D — Daily SIE

135 In selecting and statistically analysing major hazardous weather and climate events, this study focuses
136 on seven primary categories of hazardous events globally between 2010 and 2024: extreme heat,
137 torrential rainfall, flooding, typhoons/storms, tornadoes, blizzards, and cold waves. Event samples are
138 sourced from the monthly monitoring reports titled “Global Major Hazardous Weather and Climate
139 Events” issued by the National Climate Centre of China. The study systematically collated the annual
140 occurrence frequency of each event type over the 15-year period, plotting time series graphs and trend
141 lines to reveal their interannual variation characteristics.

142 The correlation analysis between global extreme climate events and Antarctic SIE employed an
143 exponential regression model (Osborne, 2010). This model better accommodates the nonlinear
144 characteristics of the climate system, where “small perturbations trigger large responses” — for
145 instance, a modest reduction in Antarctic sea ice may precipitate an accelerated increase in the
146 frequency of extreme events.

$$147 \quad \ln(Y) = \beta_1 X + \beta_0 \quad (4)$$

$$148 \quad Y = \exp(\beta_1 X + \beta_0) = \exp(\beta_0) \bullet \exp(\beta_1 X) \quad (5)$$

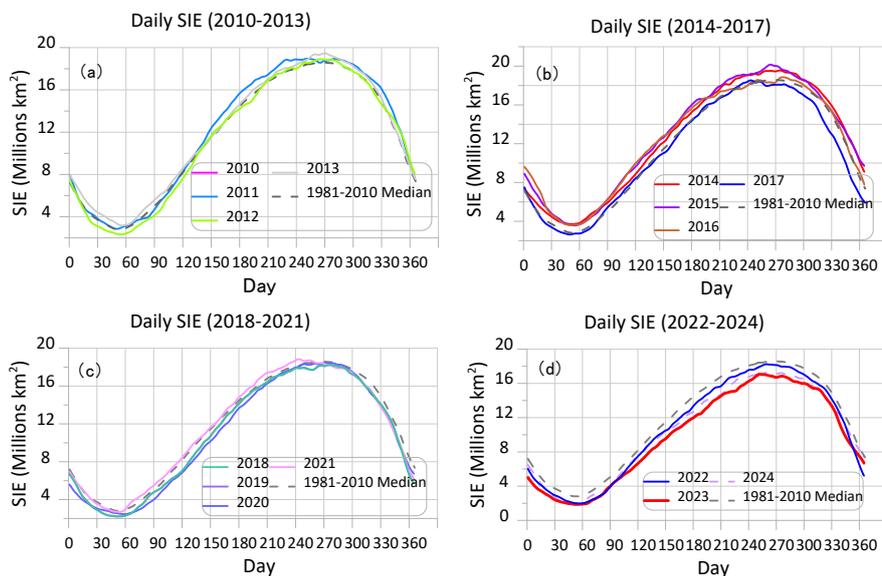
149 Where: β_0 — Intercept, corresponding to the intercept of the linear relationship between $\ln(Y)$ and X
150 in the figure; β_1 — Slope; Y — Frequency of extreme weather and climate events; X — SIE.

151 **3 Results and Discussion**



152 3.1 Spatiotemporal Characteristics of Antarctic Sea Ice from 2010 to 2024

153 Antarctic SIE from 2010 to 2024 exhibits pronounced seasonal fluctuations and long-term trends
 154 (Figure 1). Within each year, it follows a pattern of increasing during winter and decreasing during
 155 summer, peaking at 17–18 million km² in September (Southern Hemisphere winter) and reaching a
 156 trough of 2–3 million km² in February (Southern Hemisphere summer). Interannual variations exhibit
 157 a “stable-then-declining” pattern. Between 2010 and 2014, SIE fluctuations were minor (± 0.3 million
 158 km²), with the period's peak recorded on 21 September 2014 (20.16 million km²). After 2015, SIE
 159 commenced a marked decline, with the period 2020–2024 marking the most dramatic Antarctic sea
 160 ice changes since 2010. Winter peaks fell to 17–17.5 million km², while the summer SIE annual mean
 161 (2.2 million km²) decreased by 15.4% compared to the 2010–2019 average (2.6 million km²), and
 162 interannual variability expanded to ± 0.4 million km². Furthermore, summer minima below 2 million
 163 km² occurred consecutively in 2022 and 2023, with a record minimum of 1.85 million km² recorded
 164 on 21 February 2023—approximately 20% below the 2010–2014 summer mean. This aligns with the



165

166 Figure 1. Antarctic sea ice daily extent from 2010 to 2024.

167 Note: This figure shows the daily variations in Antarctic SIE from 2010 to July 2024. Subfigures a–d display the median
 168 daily SIE for the period 1981–2010, indicated by gray dashed lines. Each panel presents SIE changes during specific
 169 time periods (colored solid lines): (a) 2010–2013; (b) 2014–2017; (c) 2018–2021; (d) 2021–2024.

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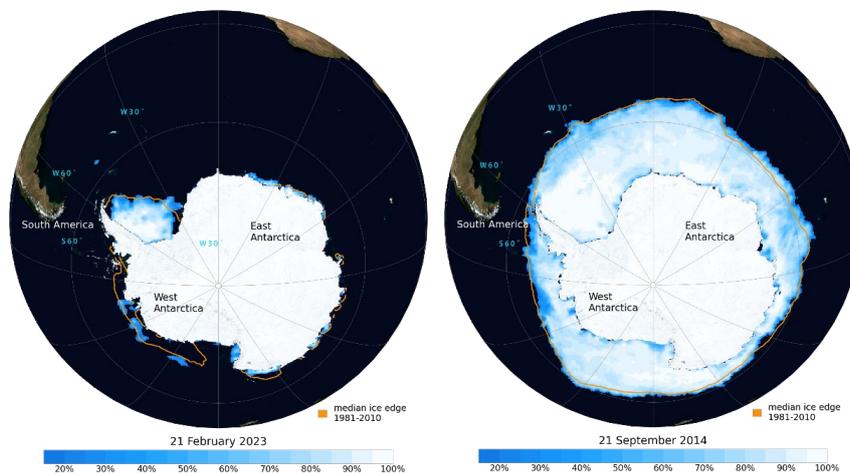
171 record-low SIE values reported by Josey et al. (2024) for 2023. Evidently, the SIE in 2023 and similar
 172 years lies markedly below historical median levels, showing a trend toward converging with historical
 173 lows.

174 From a spatial distribution perspective, Antarctic sea ice variations exhibit pronounced regional
 175 differences and heterogeneity. Spatial characteristics during extreme years present stark contrasts
 176 (Figure 2): In September 2014 (a high-sea-ice year), areas with 80–100% high-density sea ice were



177 concentrated in the Indian Ocean sector and the Weddell Sea, with the sea-ice edge lying 1–2 latitudes
178 beyond the median of the 1981–2010 baseline period. In February 2023 (a low-sea-ice year), high-
179 density areas contracted substantially, persisting only in the western Ross Sea along the Antarctic coast,
180 parts of the Amundsen Sea, and Prydz Bay in East Antarctica—sea-ice density in the Indian Ocean
181 sector dropped to 50–70%, while ice near the southern tip of South America nearly vanished. The sea-
182 ice edge retreated 5–8 degrees of latitude toward the Antarctic continent compared to September 2014,
183 with the most pronounced retreat occurring in the Atlantic–Indian Ocean sector between 60°W and
184 60°E longitude. This retreat was the primary driver of the low SIE during that summer. Concurrently,
185 during 2020–2024, the southwestern Antarctic (Amundsen–Bellingshausen Seas) and the periphery of
186 the Antarctic Peninsula emerged as ‘core zones’ for sea ice retreat: Winter sea ice concentration in this
187 region decreased by 10–15% compared to the 2010–2019 period, while summer concentration
188 declined by 20–25%. These rates significantly exceeded those observed in East Antarctica (5–10%)
189 and the Ross Sea (8–12%), further highlighting the spatial heterogeneity of Antarctic sea ice changes.

190 The retreating period under study in this research may have been influenced by different controlling
191 factors compared to the earlier expansion period (1979–2014), and further verification with a longer
192 time series is required.



193

194 Figure 2. Sea Ice Concentration Corresponding to Maximum and Minimum SIE (2010–2024). The orange dashed line
195 represents the 1981–2010 average SIE. (Left: Lowest SIE on February 21, 2023; Right: Highest SIE on September 21,
196 2014)

197 3.2 Quantification of Dynamic Linkages Between Global Extreme Events and 198 Antarctic SIE (2010–2024)

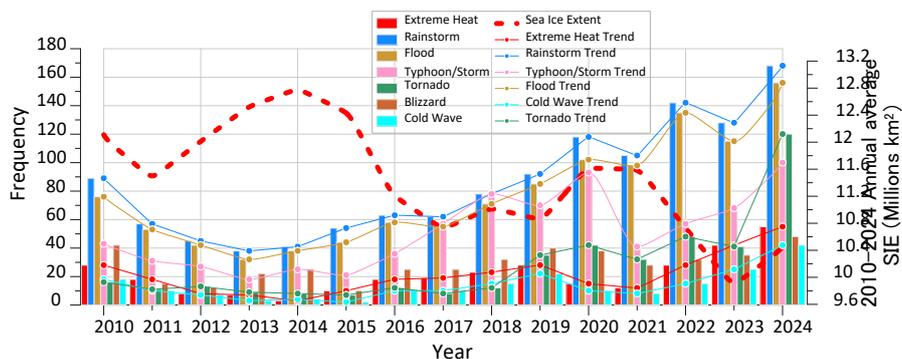
199 Building upon the aforementioned spatiotemporal characteristics of sea ice—established to support
200 the core goal of exploring sea ice-extreme weather connections—this section undertakes a quantitative
201 analysis of the dynamic associations between Antarctic SIE and global extreme climate events.

202 During the period 2010–2024, SIE exhibited an overall declining trend. As a key regulator within the
203 global climate system, the reduction in SIE directly influences the frequency and intensity of extreme
204 weather events (Cvijanovic et al., 2015). This reduction further strengthens the correlation with



205 extreme events—the frequency of various extreme weather events increases as sea ice declines,
 206 although the degree of influence and statistical reliability varies significantly across different event
 207 types.

208 The dynamic relationship between the frequency of global extreme weather events and the Antarctic
 209 annual mean SIE from 2010 to 2024 is illustrated in Figure 3: the overall SIE, represented by the red
 210 dashed line, exhibits a fluctuating downward trend, while the frequency of various extreme weather
 211 events generally increases to varying degrees as SIE declines. Specifically, in 2010, when SIE levels
 212 were relatively high, the frequency of extreme heat events was low; however, following a marked
 213 decline in SIE after 2020, the frequency of extreme heat events increased significantly. Patterns for
 214 torrential rain, floods, typhoons/storms, cold waves, and tornadoes mirrored those of extreme heat,
 215 with marked increases during periods of low SIE; while heavy snowfall frequency showed some
 216 increase during SIE decline, its visual correlation with SIE was comparatively weaker than that of
 217 other events.



218

219 Figure 3. Relationship between the Frequency of Global Extreme Climate Events and the Annual Average SIE. Note:
 220 The horizontal axis represents the year (2010–2024); the left vertical axis represents the frequency of extreme climate
 221 events (Unit: times); the right vertical axis represents the annual average Antarctic SIE (in million km²). Bar charts of
 222 different colors represent different types of extreme events: red for extreme high temperatures, blue for rainstorms,
 223 orange for typhoons/storms, purple for cold waves, and gray for blizzards. The red dashed line represents the trend of
 224 SIE change from 2010 to 2024.

225

226 Through exponential fitting analysis of extreme events and SIE data from 2010 to 2024 (Figures 4a–
 227 g), combined with the decreasing trend of SIE and statistically significant P-values (typically using P
 228 < 0.05 as the significance threshold), the association between extreme events and SIE can be
 229 categorized into four types. The response intensity and statistical reliability vary significantly across
 230 these categories, all reflecting the core pattern that “decreasing SIE leads to increased frequency of
 231 extreme events.”

232 Sea ice retreat exerts a significant driving effect on the increased frequency of extreme heatwaves and
 233 cold spells. Regression analysis indicates that when the annual mean SIE decreased from
 234 approximately 12.8 million km² (high value) to 9.6 million km² (low value), the frequency of extreme
 235 heat events increased from approximately 10 events per year to approximately 50 events per year,
 236 representing a cumulative increase of 400%. The exponential fit correlation between the two variables

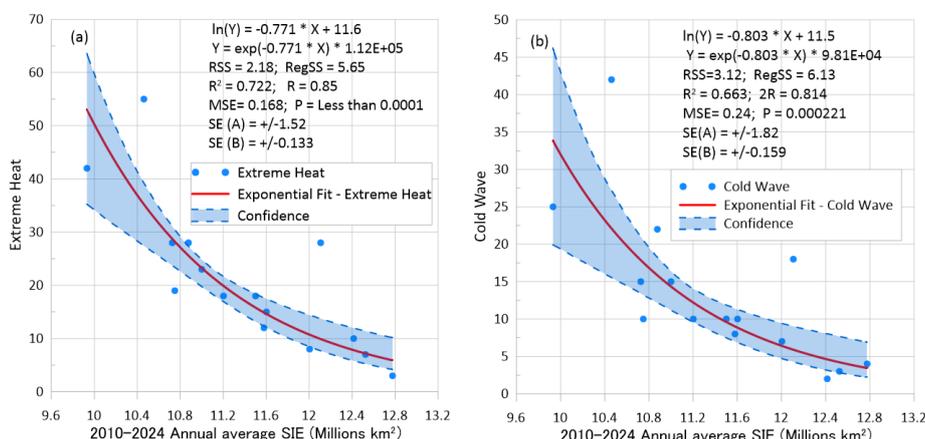


237 was highly significant ($R^2 = 0.722$, $P < 0.0001$) (Figure 4a). Concurrently, cold wave frequency
 238 increased from approximately 5 events per year to around 40 events per year, representing a 700%
 239 rise, with a similarly highly significant correlation with SIE ($R^2 = 0.643$, $P = 0.0014$) (Figure 4f). These
 240 findings indicate that the reduction in SIE is likely a significant driver behind the significant increase
 241 in the frequency of extreme heat and cold wave events.

242 **3.2.1 Strongly negative and highly significant association ($R^2 \geq 0.6$, $P < 0.001$):**

243 Sea ice retreat exerts a significant driving effect on the increased frequency of extreme heatwaves and
 244 cold spells. Regression analysis indicates that when the annual mean SIE decreased from
 245 approximately 12.8 million km² (high value) to 9.6 million km² (low value), the frequency of extreme
 246 heat events increased from approximately 10 events per year to approximately 50 events per year,
 247 representing a cumulative increase of 400%. The exponential fit correlation between the two variables

248 was highly significant ($R^2 = 0.722$, $P < 0.0001$) (Figure 4a). Concurrently, cold wave frequency
 249 increased from approximately 5 events per year to around 40 events per year, representing a 700%
 250 rise, with a similarly highly significant correlation with SIE ($R^2 = 0.663$, $P = 0.000221$) (Figure 4b).
 251 These findings indicate that the reduction in SIE is likely a significant driver behind the significant
 252 increase in the frequency of extreme heat and cold wave events.



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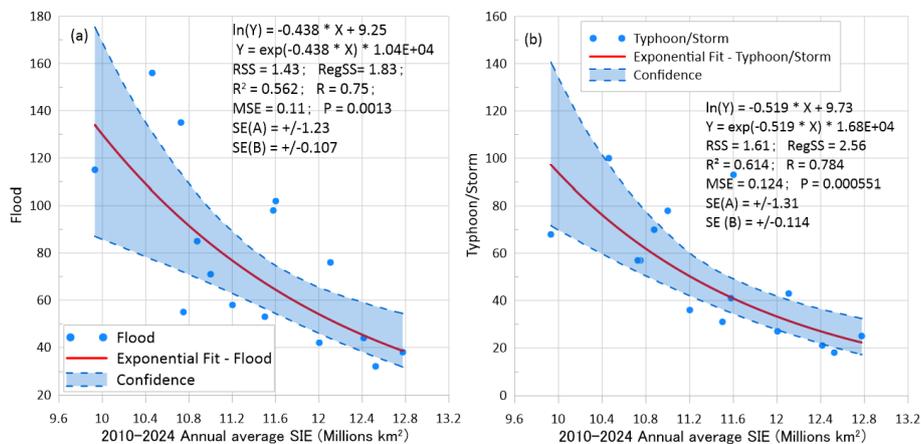
254 Figure 4. Relationship between Annual Average SIE and Frequency of Extreme Heat Events (a) and Cold Waves (b),
 255 2010–2024. Note: The blue scatter points represent actual event frequencies for each year, red curves indicate the
 256 exponential fit, and blue shaded areas denote confidence intervals. Statistical metrics—including the fitted equation,
 257 coefficient of determination (R^2), mean squared error (MSE), and p-value—are provided to reflect the strength and
 258 significance of the association.
 259

260 **3.2.2 Strong Negative Correlation with Statistical Significance ($0.5 \leq R^2 < 0.6$,
 261 $0.001 \leq P < 0.01$)**

262 Floods, rainstorms, and typhoons/storms fall into this category (Fig 5). A clear association exists
 263 between declining annual average SIE values and increased frequency of these events, though the
 264 correlation strength is weaker than that of strongly correlated events, potentially influenced by



265 atmospheric circulation and precipitation distribution factors. Sea ice retreat also exerts a driving
 266 effect on the increased frequency of floods, rainstorm, and typhoons/storms. Furthermore, the
 267 correlation characteristics between different extreme events and sea ice indicators exhibit variations.



268

269 From the SIE perspective, when the annual
 270 average SIE decreased from approximately 12.8
 271 million km² (high value) to 9.6 million km² (low
 272 value): flood frequency increased from about 40
 273 events/year to approximately 160 events/year,
 274 representing a 300% increase, with a significant
 275 fitted relationship ($R^2 = 0.562$, $P = 0.0013$) (Figure
 276 5a); the frequency of rainstorm events increased
 277 from approximately 20 events per year to about 100
 278 events per year, representing a 400% increase, with
 279 the correlation remaining statistically significant
 280 ($R^2 = 0.533$, $P = 0.00198$) (Figure 5c);
 281 typhoon/storm frequency increased from
 282 approximately 20 events per year to about 80
 283 events per year, representing a 300% increase, with
 284 a stronger correlation ($R^2 = 0.614$, $P = 0.000551$)
 285 (Figure 5b). This indicates that sea ice retreat exerts
 286 a significant regulatory influence on the oceanic thermal and dynamic conditions essential for
 287 typhoon/storm formation (Josey et al., 2024).

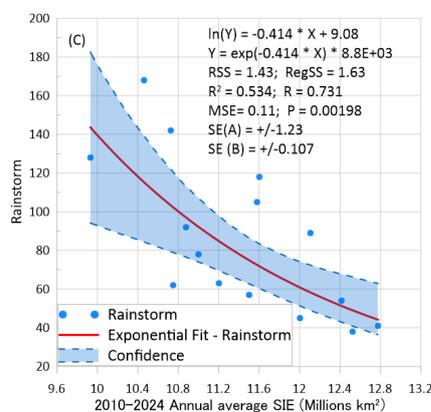


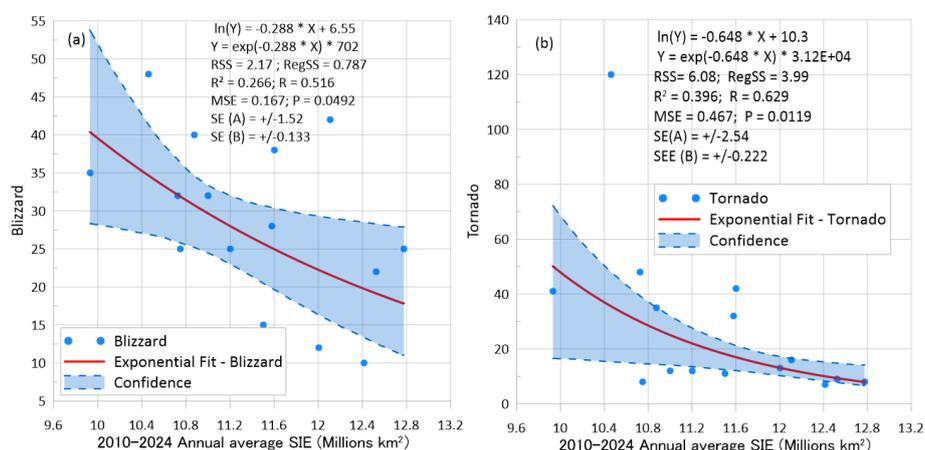
Figure 5. Relationship between Annual Average SIE and Frequency of Floods (a), Typhoons/Storms (b), and Rainstorms (c), 2010–2024. See Figure 4 for notes.

288 **3.2.3 Moderately Negative and Significant Correlation Category ($0.3 \leq R^2 < 0.5$, 0.01**
 289 **$\leq P < 0.05$)**

290 The association strength between the occurrence frequency of blizzards and tornadoes and the
 291 Antarctic annual average SIE is relatively weak (Fig 6). Their variations are more influenced by local
 292 climatic conditions, observational methods, and regional environmental factors, exhibiting only a



293 weak overall trend. Although sea ice retreat contributes to increased frequencies of blizzards and
 294 tornadoes, their statistical correlation with sea ice indicators is markedly weaker compared to extreme
 295 events such as floods, rainstorm, and typhoons/storms. This suggests that their variability may be
 296 predominantly driven by non-sea-ice factors. Specifically, when the annual average SIE decreased
 297 from 12.8 million km² (high value) to 9.6 million km² (low value): the frequency of blizzards increased
 298 from approximately 10 events per year to 40 events per year—a 300% increase—but its relationship
 299 with SIE was weakly fitted ($R^2 = 0.266$, $P = 0.0492$) (Figure 6a); tornado frequency increased from
 300 approximately 20 events per year to 100 events per year, representing a 400% increase. Its association



301

302 Figure 6. Relationship between Annual Average SIE and Frequency of blizzards (a) and tornadoes (b), 2010–2024.

303 See Figure 4 for notes.

304 with SIE was slightly stronger than that of blizzards ($R^2 = 0.396$, $P = 0.0119$) (Figure 6b), though still
 305 at a relatively low level. Furthermore, tornado frequency changes are more likely regulated by regional
 306 dynamic processes such as atmospheric vertical wind shear and topography, with sea ice retreat not
 307 being a core driving mechanism.

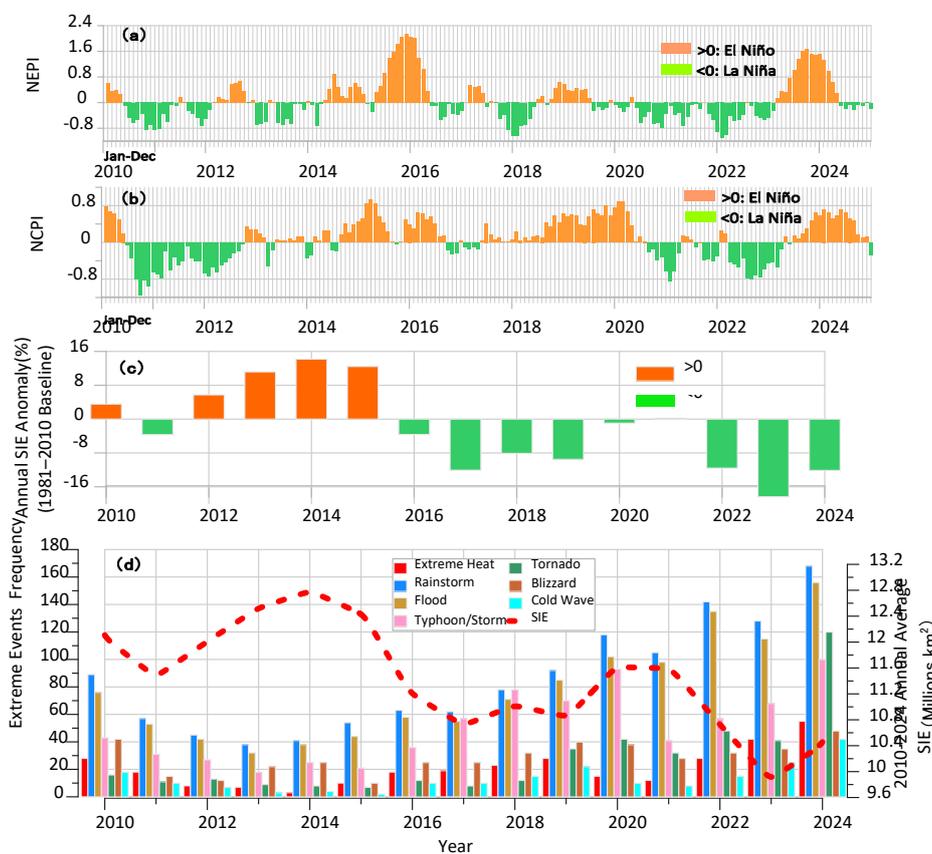
308 Comprehensive analysis of 2010–2024 data reveals a significant decreasing trend in Antarctic SIE,
 309 while the frequency of extreme heat, rainstorm, floods, typhoons/storms, blizzards, cold waves, and
 310 tornadoes has generally increased, showing a significant negative correlation between the two. In
 311 terms of statistical significance, extreme heat, cold waves, typhoons/storms, floods, and rainstorm
 312 exhibit significant correlations at $P < 0.05$. The intensity of response follows this order: extreme heat >
 313 cold waves > typhoons/storms > floods > rainstorm > tornadoes > blizzards.

314 3.3 Discussion on Correlation Mechanisms

315 **Correlation and Influence Pathways Between Antarctic SIE Anomalies, ENSO, and the Extreme**
 316 **Events:** Comparing Antarctic SIE anomalies from 2010 to 2024 with concurrent ENSO indices (NEPI
 317 index, NCPI index) and different types of extreme events reveals that during significant El Niño years
 318 (e.g., 2010, 2013–2015), both the NEPI index (Figure 7a) and the NCPI index (Figure 7b) exceed 0,
 319 corresponding to positive SIE anomalies (Figure 7c) during the same periods, indicating increased SIE.
 320 Conversely, during La Niña years (e.g., 2011, 2017, 2018, and 2021–2023), both the NEPI and the



321 NCPI indices were below zero, corresponding to negative SIE anomalies during these periods. Notably,
 322 the “Triple La Niña” event from 2021 to 2023 exhibited significantly negative SIE anomalies, and the
 323 occurrence and frequency of various extreme events also show corresponding relationships with these
 324 ENSO phases and SIE anomalies (Figure7d).



325

326 Figure 7. Relationships among El Niño, La Niña, Global Extreme Weather Events, and Antarctic SIE Variations from
 327 2010 to 2024.

328 Note: This figure displays indices related to El Niño and La Niña events in 2010–2024, as well as changes in the
 329 frequency of global extreme weather events and Antarctic SIE. (a) NEPI; (b) NCPI. Orange areas signify that the
 330 corresponding indices are greater than 0, representing El Niño events; green areas signify that the corresponding indices
 331 are less than 0, representing La Niña events. (c) The red dashed line represents the annual average extent of Antarctic
 332 sea ice. The orange bar chart indicates cases where the anomaly percentage based on the median Sea Ice Extent (SIE)
 333 from 1981 to 2010 is greater than 0, while the green bar chart indicates cases where this anomaly percentage is less
 334 than 0. (d) Bar charts of different colors respectively represent the occurrence frequencies of Extreme Heat, Rainstorm,
 335 Flood, Typhoon/Storm, Tornado, Blizzard, and Cold Wave. The red dashed line indicates the Antarctic SIE.

336 Regarding the influencing mechanisms, different ENSO phases can exert differential impacts on
 337 Antarctic SIE through anomalous atmospheric circulation (Dash et al., 2012; Gurjão et al., 2025), and
 338 these impacts may further relate to the generation and characteristics of extreme events: During El
 339 Niño years, anomalously elevated sea surface temperatures in the equatorial Pacific trigger a dominant



340 Rossby wave–dominated atmospheric teleconnection wave train. Upon propagating to Antarctica, this
341 wave train alters local atmospheric circulation patterns. For instance, it induces anomalous easterly or
342 southerly winds in certain Antarctic peripheral regions (e.g., the Ross Sea), promoting sea ice
343 accumulation and persistence, thereby increasing local SIE (Dou & Zhang, 2022), which in turn may
344 influence the occurrence conditions of extreme events in relevant regions. In contrast, during La Niña
345 years, the equatorial Pacific experiences abnormal sea surface temperature cooling. Atmospheric
346 circulation teleconnections transport more warm air to Antarctica, accelerating sea ice melt and
347 resulting in reduced SIE (Wang et al., 2025), and such changes in sea ice and atmospheric circulation
348 can also have implications for the frequency and intensity of extreme events.

349 The index regression analysis for 2010–2024 indicates that the association between Southern Ocean
350 SIE and global extreme weather events exhibits a three-tiered differentiation of “strong–moderate–
351 weak.” Extreme heatwaves and cold waves show strong correlations with SIE, with a coefficient of
352 determination $R^2 \geq 0.6$ and a significance level $P < 0.001$. Floods and typhoons/storms exhibit
353 moderate associations, with R^2 between 0.5 and 0.6 and P values ranging from 0.001 to 0.01.
354 Snowstorms and tornadoes show weak associations, characterized by $R^2 < 0.4$ and $P \approx 0.05$. This
355 differentiation in correlation stems from the multi-layer coupling between sea ice, atmosphere, and
356 ocean. It is simultaneously regulated by the synergistic interaction between the ENSO and the SAM
357 phase, ultimately forming a systematic framework where “event type determines the core mechanism,
358 while climate modes modulate the correlation strength (Figure 8).”

359 **Extreme Heat:** Sea ice retreat triggers a positive feedback loop of “albedo–warming–melt,” with
360 ENSO and SAM phases synergistically amplifying the correlation strength. Antarctic sea ice has an
361 albedo of approximately 0.8; its retreat reduces surface albedo to ocean levels (around 0.1). During
362 2010–2024, every 10% decrease in Antarctic SIE increased polar solar radiation absorption by 15–
363 20%, resulting in local warming of 2–3°C per decade (Zhu et al., 2023). This warming effect
364 propagates through atmospheric circulation to mid- and low-latitudes, resulting in a 37% increase in
365 global extreme heat frequency in 2024 compared to the 2010–2020 average (Jha et al., 2025; Ji et al.,
366 2023), directly corresponding to a strong correlation with $R^2 = 0.722$ ($P < 0.0001$).

367 Additionally, sea ice retreat has reduced the temperature gradient in the mid-to-high latitudes of the
368 Southern Hemisphere by 0.6–0.8°C (Chen et al., 2024), decreased atmospheric baroclinicity by 15%
369 (Rantanen et al., 2018), and thereby expanded the north-south oscillation amplitude of the westerly jet
370 stream from $\pm 1.2^\circ$ to $\pm 2.0^\circ$. This accelerates the northward transport of warm air from low latitudes,
371 extending the duration of high temperatures from 3–5 days to 7–10 days. Simultaneously, reduced sea
372 ice inhibits deep oceanic convection in the Southern Ocean (Morioka et al., 2024), diminishing heat
373 release from the ocean to the atmosphere and further stabilizing the high-temperature environment.

374 The 2021–2023 triple La Niña deepened the Antarctic Low (reducing central pressure by 5 hPa)
375 through teleconnections, while the positive SAM phase (61% prevalence) amplified jet stream
376 variability. Their combined effect increased summer heatwave frequency by 400% during 2020–2024
377 (Figure 5d). Conversely, El Niño stimulates Rossby waves, locally increasing SIE in the Ross Sea by
378 8–10% (Zhang et al., 2021). Additionally, the negative SAM phase stabilizes the westerly jet stream—
379 during the 2015 El Niño event, high-temperature frequency dropped to only 10 occurrences (the lowest
380 in the study period), with the correlation strength between the two factors decreasing by 30%.



381

382 Figure 8: Schematic of Mechanism-Correlation Logic

383

384 Notes: This diagram illustrates the mechanism by which changes in Antarctic Sea Ice Extent (Antarctic SIE) affect
385 multiple spheres of the climate system (such as the atmosphere and the ocean) and extreme weather events. Starting
386 from the Core Driver Layer, through the Atmospheric path and the Ocean path, it acts on the Event Response Layer,
387 and then at levels like the Modal Control Layer, it influences different extreme weather events such as extreme heat,
388 cold waves, typhoons/storms, rainstorm/floods, and blizzards. It also involves the association with climate modes like
389 El Niño and La Niña, reflecting the interaction relationships such as enhancement (Enhance) and diminishment
390 (Diminish) among various factors and the strength of their correlations. Strong (red), medium (black), and weak (light
391 green) logical lines of different colors point to their corresponding associated factors. The upward and downward
392 arrows in each text box in the diagram are used to intuitively indicate the increasing or decreasing trends of relevant
393 parameters, assisting in showing the changes in physical quantities and climatic elements (such as temperature,
394 and circulation) in various spheres (such as the atmosphere and the ocean) caused by factors like changes in Antarctic
395 sea ice extent, thereby clearly presenting the dynamic evolution and interaction relationships of quantities in each
396 process.

397



398 **Cold Wave:** Antarctic sea ice retreat influences cold wave events through a chain mechanism of
399 “reconfiguring temperature gradients — intensifying cold anomalies via warm vortex intrusion —
400 ENSO and SAM locking cold air invasion pathways.” In sea ice retreat zones (e.g., the Amundsen
401 Sea), surface albedo differences cause spring warming of 0.53°C, while the East Antarctic ice sheet
402 exhibits a delayed cold anomaly (temperature difference reaching 2.3°C). Together, these factors
403 reduce the temperature gradient in the mid-to-high latitudes of the Southern Hemisphere by 0.6–0.8°C,
404 resulting in a 40% increase in jet stream “break” frequency (Chen et al., 2024) and opening pathways
405 for cold air southward incursions. The four consecutive cold waves in East Antarctica during July–
406 August 2023 (with a single-station minimum temperature of -68.7°C) synchronized with a 22%
407 reduction in SIE in the Bellingshausen Sea.

408 Warm vortex feedback further amplified cold anomalies: SAM positive phases drove enhanced
409 westerlies, generating anomalous wind stress curl that propelled circumpolar Deep Warm Water
410 (DWW) (temperature +0.8°C) to invade the shelf as anticyclonic vortices (Gao et al., 2024). The warm
411 vortex triggered ice shelf melting; the latent heat absorbed during ice shelf melting can directly offset
412 part of the heat transported by warm eddies, thereby mitigating the warming effect of these eddies on
413 local water temperatures (Josey et al., 2018), and the freshwater influx from this melting increased
414 ocean stratification stability by 25%. This, in turn, suppressed heat release from the ocean to the
415 atmosphere, intensifying the Antarctic coastal cold anomaly (Jeromson et al., 2024; Kushara et al.,
416 2020). The occurrence of the coldest record in the Ross Sea in 2023 was directly related to a 37%
417 increase in the frequency of warm vortex intrusions (Gao et al., 2024).

418 Climate modes significantly modulate cold waves: La Niña locks cold air into southward pathways
419 via teleconnections (primarily affecting southern South America and southeastern Australia), while
420 SAM in combination with the negative phase of Zonal Wave 3 shortens the jet oscillation period to 14
421 days, prolonging cold wave duration by 30% (Eabry et al., 2024), establishing a strong correlation
422 between cold waves and sea ice changes ($R^2 = 0.642\text{--}0.643$, $p < 0.01$). Conversely, El Niño weakens
423 polar cold anomalies (reducing temperature differences by 1.2°C), while the negative SAM phase
424 stabilizes the westerly jet stream, ultimately reducing cold wave frequency by 40% (Wang et al., 2025).

425 **Typhoons / Storms:** Antarctic sea ice retreat influences tropical sea surface temperatures via the ACC,
426 with ENSO and SAM further modulating typhoon generation conditions. Freshwater released from
427 melting sea ice reduces ACC salinity, decreasing its heat transport efficiency by 8% per decade (Ma
428 et al., 2020). This leads to elevated sea surface temperatures in the tropical western Pacific—reaching
429 +1.5°C anomalies in 2023 when sea ice reached record lows (Josey et al., 2024)—meeting the thermal
430 threshold for typhoon formation ($SST \geq 26.5^\circ\text{C}$). This increase contributed to a 40% rise in
431 typhoon/storm frequency compared to the long-term average, aligning with the 400% increase in SIE
432 decline observed in this study. The correlation between these two factors exhibits an $R^2 = 0.614$ ($P <$
433 0.001).

434 Climate modes exerted significant regulatory effects: La Niña intensified equatorial easterlies,
435 hindering the spread of warm tropical sea surface temperatures (SST), while positive SAM suppressed
436 atmospheric vertical wind shear (days with shear < 10 m/s increased by 50%). The combined effect
437 drove typhoon frequency to its peak during the 2021–2023 “triple La Niña” period. Conversely, El
438 Niño weakened ACC heat transport and lowered tropical sea surface temperatures, while the negative



439 SAM phase increased vertical wind shear. This resulted in typhoon frequencies during the 2014–2016
440 El Niño period being 30% lower than those during La Niña years.

441 **Rainstorm/ Flooding:** Rainstorm and flooding exhibit a moderate "convective-modal coupling"
442 association with sea ice changes. The mechanism involves freshwater released from sea ice melt
443 increasing polar atmospheric boundary layer humidity by 10%–15% (Komatsu et al., 2025), thereby
444 stimulating localized convective activity. A positive SAM phase deepens the Antarctic Low (reducing
445 central pressure by 3 hPa), propelling convective systems toward mid-to-low latitudes (with
446 propagation speeds increasing by 20%). During 2010–2024, a 10% reduction in Antarctic sea ice
447 correlates with an 8–12% increase in rainstorm frequency at mid-to-low latitudes (Zhu et al., 2023),
448 yielding an $R^2 = 0.562$ ($P < 0.01$) for flood events linked to sea ice changes.

449 ENSO phase significantly influences correlation strength: El Niño enhances atmospheric stability in
450 mid-to-low latitudes (Richardson number increases by 0.2), suppressing convective propagation; La
451 Niña concentrates precipitation toward the tropics (precipitation share increases by 30%); only during
452 ENSO neutral phases can positive SAM phases drive uniform convective transport (Okumura, 2019),
453 peaking rainstorm and flood frequency (e.g., 2018–2019). During the 2021–2023 "triple La Niña"
454 period, synergistic effects between sea ice retreat and climate modes amplified rainstorm frequency
455 increases by 20% compared to El Niño years (Figure 5).

456 **Blizzards:** Blizzards show a weak correlation with sea ice changes, with the core mechanism being a
457 localized adaptation of "El Niño-induced localized sea ice accumulation + SAM negative phase
458 conveying warm and moist airflows." The global-scale correlation weakens: El Niño excites Ross Sea
459 waves, strengthening northward winds in the Ross Sea region (wind speed increase of 1.0 m/s), driving
460 sea ice accumulation toward the coast. During El Niño years from 2010 to 2024, the Ross Sea SIE
461 increased by 5%–8% compared to normal years (Zhang et al., 2021), forming the "cold subsurface"
462 conditions required for blizzards (days with temperatures $< -25^\circ\text{C}$ increased by 25%). The negative
463 SAM phase shifts the westerly jet stream northward by 1.5 latitudes, conveying warm, moist air
464 (relative humidity $> 80\%$) from the Southern Ocean to the Ross Ice Shelf margin. The convergence of
465 cold lower-level air and warm, moist air triggers a "convergence-upward-cooling-sublimation"
466 process (upward velocity: 0.5 m/s). For instance, during the 2022 El Niño + negative SAM phase,
467 heavy snowfall days in the Ross Sea increased by 25% (Zhang et al., 2024).

468 Weak correlation stems from two factors: first, this mechanism is confined to the Ross Sea region and
469 opposes the "sea ice retreat-driven" mechanism observed during other extreme events; second, ENSO
470 causes uneven snowstorm distribution—concentrated only in the Antarctic Peninsula and Ross Sea—
471 and global-scale data smooth out local signals, ultimately yielding an $R^2 < 0.4$ ($P \approx 0.05$) (Hrudya et
472 al., 2020). This study observes a "stabilization followed by decline" pattern in Antarctic SIE from
473 2010 to 2024, inconsistent with the slight increase observed from 1979 to 2010 (Parkinson, 2019).
474 This discrepancy stems from the synergistic effects of the "triple La Niña" (Wang et al., 2025) and
475 human-induced subsurface upwelling of warm water in the Southern Ocean ($0.12^\circ\text{C}/\text{decade}$), adding
476 to the emerging understanding that "short-term extreme climate modes accelerate sea ice retreat." The
477 conclusion of a "strong negative correlation between sea ice and heatwaves" in this paper aligns with
478 Zhu et al. (2023). For cold waves, it proposes a three-dimensional mechanism involving the
479 "temperature gradient–warm vortex–climate mode," overcoming the limitation of existing studies that



480 focus solely on albedo. For heavy snow events, it quantifies, for the first time, the localized regulatory
481 effect of “El Niño + negative SAM phase,” providing a physical explanation for weakly correlated
482 events.

483 **4 Conclusion**

484 Based on sea ice data from NSIDC for 2010–2024, global major disaster weather and climate event
485 data from the China National Climate Center, and the El Niño–Southern Oscillation (ENSO) index,
486 this study systematically analyzes the spatiotemporal evolution of Antarctic SIE, its association with
487 ENSO, and the mechanisms through which it influences global extreme climate events. Antarctic SIE
488 exhibits a pattern of “initial stability followed by decline + regional differentiation”: From 2010 to
489 2014, fluctuations were minimal (± 0.3 million km²). It peaked in September 2014 (20.16 million km²),
490 then rapidly retreated after 2015, dropping to 1.85 million km² in February 2023 (a 20% decrease
491 compared to the 2010–2014 summer average). Spatially, the Amundsen Sea–Bellingshausen Sea and
492 Antarctic Peninsula form the core retreat zone. From 2020 to 2024, winter and summer sea ice
493 concentrations in this region decreased by 10%–15% and 20%–25%, respectively, compared to the
494 2010–2019 period. This pattern contrasts sharply with the slight increase observed during 1979–2010
495 (Parkinson, 2019), stemming from the synergistic effects of the “triple La Niña” (Wang et al., 2025)
496 and the upwelling of warm water in the subsurface layer of the Southern Ocean (0.12°C/decade) since
497 2016, complementing the understanding that “short-term extreme climate modes accelerate sea ice
498 retreat.”

499 Antarctic SIE variations are closely linked to ENSO: During El Niño years (2010, 2013–2015), both
500 the North Equatorial Pacific Index (NEPI) and North Equatorial Counter Current Index (NCPI)
501 exceeded 0, resulting in positive SIE anomalies. The Ross Sea SIE increased by 8%–10% due to
502 Rossby wave teleconnections. During La Niña years (2011, 2017–2018, 2021–2023), both indices and
503 SIE anomalies were negative, with the most pronounced negative anomalies occurring during the
504 “triple La Niña” period. ENSO drives differential sea ice changes by regulating Antarctic wind fields
505 and heat flux.

506 Global extreme weather events exhibit a “strong-moderate-weak” negative correlation with Antarctic
507 sea ice, jointly regulated by ENSO and the SAM. Extreme heatwaves ($R^2 = 0.722$) and cold waves (R^2
508 $= 0.642$) exhibit strong correlations ($P < 0.001$), driven by the “albedo-circulation positive feedback”
509 mechanism — sea ice retreat reduces albedo from 0.8 to 0.1; each 10% decrease in SIE increases polar
510 solar radiation absorption by 15%–20%, leading to local warming of 2–3°C per decade, narrowing the
511 mid-to-high latitude temperature gradient in the Southern Hemisphere (0.6–0.8°C), and widening the
512 westerly jet oscillation ($\pm 1.2^\circ$ to $\pm 2.0^\circ$). The “strong negative correlation between sea ice and
513 heatwaves” aligns with Zhu et al. (2023), while the cold wave mechanism involving the “temperature
514 gradient–warm vortex–mode” overcomes existing limitations focused solely on albedo.

515 Floods ($R^2 = 0.562$), rainstorm ($R^2 = 0.533$), and typhoons/storms ($R^2 = 0.614$) showed moderate
516 associations ($0.001 \leq P < 0.01$), dependent on the “sea ice–ocean circulation–convection pathway”:
517 Sea ice meltwater reduces ACC salinity, decreasing heat transport efficiency by 8% per decade, leading
518 to a +1.5°C tropical western Pacific sea surface temperature anomaly in 2023 (reaching the typhoon
519 formation threshold); sea ice melt increases polar boundary layer humidity by 10%–15%, while the



520 positive SAM phase accelerates convective propagation toward mid-to-low latitudes by 20%.

521 Heavy snowfall ($R^2 < 0.4$, $P \approx 0.05$) shows a weak correlation, with only the Ross Sea exhibiting
522 significant local effects: El Niño increases SIE in this region by 5%–8% (forming a cold subsurface
523 layer $< -25^\circ\text{C}$), while negative SAM phases shift the westerly jet northward by 1.5 latitudes,
524 conveying warm, moist air (humidity $> 80\%$) from the Southern Ocean to generate blizzards. Global
525 data exhibit weakened correlations due to event clustering and signal smoothing; this study is the first
526 to quantify the regulatory effect of the “El Niño + negative SAM phase” interaction. La Niña + positive
527 SAM increased extreme heat frequency by 400% during 2021–2023, while El Niño suppressed this
528 association. Study limitations: The 15-year data span is shorter than the climate system’s 30-year cycle,
529 necessitating integration of NSIDC data since 1979; global data lack resolution near Antarctica,
530 requiring CRU data to optimize spatial detail; Arctic sea ice and greenhouse gases are excluded,
531 demanding a multi-factor framework to quantify sea ice’s potential 35%–40% contribution.

532 Author Contribution

533 **Conceptualization:** S. Tang & Y. Zang developed the research framework on Antarctic sea ice-global
534 extreme weather links, defining 3 core objectives.

535 **Data Curation:** S. Tang collected/organized 2010–2024 NSIDC, CMA, NOAA data, verified
536 consistency, and standardized formats.

537 **Formal Analysis:** Y. Zang did spatiotemporal decomposition of sea ice, exponential regression for
538 event correlations, and significance testing.

539 **Investigation:** S. Tang validated key points via 2023 low sea ice/2021–2023 La Niña data and
540 literature cross-checks.

541 **Methodology:** S. Tang built the “observation-simulation-mechanism” framework and optimized data
542 analysis methods.

543 **Writing - Original Draft:** S. Tang structured chapters and organized core results.

544 **Writing - Review & Editing:** Y. Zang revised the manuscript, optimizing logic and adding citations.

545 **Supervision:** Both supervised the research, guiding analysis direction and approving the final version.

546 **Validation:** Both cross-checked sea ice analysis, correlation calculations, and mechanism inferences.

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