

This manuscript presents a robust and methodologically significant paleolimnological study that refines diatom-based transfer functions for reconstructing Southern Hemisphere westerly wind (SHW) variability on sub-Antarctic Macquarie Island. By integrating contemporary diatom data with multi-year hydrogeochemical and isotopic datasets, the work effectively updates earlier conductivity-inference models and capitalizes on a unique post-eradication ecological context to establish a more representative diatom-environment baseline. The rigorous assessment of hydrochemical stability and evaporation solidifies electrical conductivity (EC) as a reliable proxy for wind-driven sea-spray aerosol deposition. The resulting model provides an ecologically grounded and statistically robust tool for reconstructing long-term SHW dynamics, advancing paleoclimatic methodology in the data-sparse sub-Antarctic region. However, few aspects require clarification and expansion to fully support the conclusions and maximize the manuscript's impact. I consider this paper to be highly important and valuable, and I strongly recommend it for publication with major revisions.

Detailed assessment and suggestions for improvement

Major issue:

Comment 1: Clarification on the concept of a "recovered baseline" and its validation (lines 61–75).

The manuscript highlights that earlier diatom-environment models were developed during a period of significant ecosystem disturbance from invasive rabbits and argues that the post-eradication dataset provides a better representation of a "natural" or "pre-invasion baseline." This is a critical premise for claiming an improved and more ecologically relevant transfer function. However, the current argument is largely inferential. To substantiate this central claim, more direct evidence or a more rigorous conceptual framework is required. Specifically, the text should address the following questions:

What defines the "recovered" state? Is it the mere absence of rabbits, or are there specific, measurable limnological parameters (e.g., nutrient levels, sediment composition, vegetation cover) that have demonstrably returned to a defined range? Please reference specific post-eradication recovery studies to define the criteria for "recovery" as applied to lacustrine systems.

How do we know this state approximates a "pre-invasion baseline"? The strongest evidence would be a direct comparison. If available, please discuss whether diatom assemblages from your 2022 surface sediments show greater similarity to subfossil diatom assemblages from sediment core intervals dated to pre-1900 (i.e., pre-rabbit introduction) than to assemblages from core intervals representing the peak disturbance period. If such core data is not available, this limitation should be explicitly

acknowledged, and the argument should be reframed more cautiously. Instead of claiming to represent a "pre-invasion baseline," it may be more accurate to state that the model reflects "post-distribution recovery conditions" which are assumed to be moving toward a pre-disturbance state, thereby reducing the confounding noise of extreme eutrophication and erosion in the calibration dataset.

CS: We thank the reviewer for this detailed and constructive comment and agree that the concept of a "recovered baseline" requires clearer definition and more cautious framing.

In the revised manuscript, we endeavour to clarify that by "recovered" we do not imply a fully demonstrable return to pre-invasion conditions, but rather a post-eradication state in which the dominant anthropogenic disturbance (introduced rabbits) has been removed, and ecosystem processes are no longer characterised by the extreme erosion and associated changes documented during the peak disturbance period.

Direct limnological recovery studies for individual lakes following rabbit eradication are limited. However, strong independent evidence for catchment-scale recovery across the island is provided by the documented, rapid and widespread recovery of vegetation following rabbit eradication (e.g. *ref*). Given the previously severe overgrazing pressure, this vegetation recovery is expected to have substantially reduced soil erosion, landsliding, and sediment and nutrient delivery to lakes, even though quantitative runoff or nutrient time series are not available to directly test this assumption. We will explicitly acknowledge this limitation in the updated text.

Importantly, we are able to provide site-specific evidence from one sedimentary diatom record. At Emerald Lake (LK6), downcore diatom assemblages show a clear ecological shift coincident with the onset and persistence of rabbits, with *Fragilaria capucina* and *Psammothidium abundans* dominating downcore intervals across this time (Saunders *et al.* 2013). Diatom assemblages in recent (2022) surface sediments from this site exhibit higher diversity (48 species) and greater similarity to pre-rabbit sediment intervals than assemblages from surface sediments collected in 2006 (15 species), which was dominated by *F. capucina* (48% relative abundance), consistent with elevated nutrient conditions and disturbance. Contrastingly *F. capucina* was not found to occur at LK6 in the 2022 surface sample. This comparison supports the interpretation that recent assemblages no longer reflect peak disturbance conditions.

We agree with the reviewer that the term "pre-invasion baseline" may overstate the certainty of this interpretation. Accordingly, we suggest revising the manuscript to more cautiously describe the updated calibration dataset as representing **post-eradication recovery conditions**, which are assumed to be moving toward a pre-disturbance state and, critically, are less influenced by the extreme disturbance characterised by the

earlier dataset. We believe this distinction strengthens the conceptual foundation of the transfer function and more accurately reflects the available evidence.

Comment 2: Lines 88-89 & throughout Discussion: Clarifying the mechanistic link between diatom-inferred EC and SHW strength.

The manuscript's core hypothesis is that diatom-inferred Electrical Conductivity (EC) serves as a proxy for Southern Hemisphere Westerly Wind (SHW) strength via wind-driven Sea Spray Aerosol (SSA) deposition. While the study excellently establishes the diatom-EC relationship and demonstrates that EC in western lakes reflects SSA inputs, it presents a circular argument when applying this to paleoclimate. The logic is: 1) SHW drives SSA, 2) SSA increases lake EC, 3) Diatoms record EC. Therefore, fossil diatoms can reconstruct past EC, which *implies* past SSA, which implies past SHW.

However, a critical intermediate step is missing: a quantitative or semi-quantitative demonstration that the observed spatial and interannual EC gradient is directly forced by measurable wind parameters. The discussion relies on citations (e.g., Saunders et al., 2009, 2018) for this link but does not independently validate it with the new, richer 2018-2022 dataset. To resolve this, the following major addition is required:

Perform and present a correlation analysis between measured EC values (or marine ion concentrations like Na⁺, Cl⁻) from your key SSA-influenced lakes and instrumental wind data (e.g., mean seasonal wind speed, frequency of high-wind events) from the Macquarie Island station (BOM, 2025) for the corresponding periods (2018, 2022-23). This analysis would test the fundamental assumption that higher EC in a given year or at a given site correlates with stronger westerly wind metrics.

Discussion of the Wind-EC Transfer Function: If a correlation exists, discuss its strength and the potential variance explained by wind speed alone. If the correlation is weak or non-existent for the short instrumental record, discuss other modulating factors (e.g., rainfall dilution history, wave state affecting SSA generation) and what this means for the certainty of reconstructing *wind speed* (a dynamic variable) from *SSA deposition* (an integrated flux variable). This moves the discussion from a simple cause-and-effect to a more nuanced, process-based understanding of the proxy.

CS: We thank the reviewer for this comment and agree that the mechanistic link between wind variability and electrical conductivity (EC) requires clear framing and appropriate consideration of scale.

In this study, EC is interpreted as an integrated limnological variable reflecting the balance between solute inputs, soil-water-air processes operating over multi-annual to decadal timescales. The inferred influence of wind variability on EC is therefore not intended to reflect short-term or event-scale wind forcing, but rather persistent, large-

scale changes in atmospheric circulation. Diatom assemblages respond to these longer-term EC states rather than to transient meteorological variability.

The type of direct correlation suggested (e.g. between individual wind events and EC or diatom assemblages) is not feasible at the study location. Macquarie Island is extremely remote, and continuous limnological monitoring data are not available or feasible. Moreover, even if such data existed, event-scale correlations would be inconsistent with the temporal integration inherent in sedimentary diatom assemblages and EC reconstructions and would therefore not provide meaningful validation of the proxy–environment relationship.

The conceptual link between large-scale wind regimes and limnological and ecological responses in Southern Hemisphere lake systems is well established in previous studies (e.g. Saunders et al. 2009; 2016; 2018; Perren et al. 2020; 2025; Van Nieuwenhuyze 2020; Humphries et al. 2021; Meredith et al. 2022) and our interpretation follows this existing framework rather than proposing a new or untested mechanism. We proposed to revise the manuscript introduction to clarify the timescale over which wind variability is relevant and to explicitly state that the inferred wind influence reflects longer-term circulation change, capable of influencing EC through sustained hydrogeochemical changes rather than event-scale forcing.

Comment 3: Interannual variability and its implications for the calibration dataset

The manuscript correctly observes that the primary clustering in the PCA (Fig. 4) is by lake type (SSA, catchment, rainfall), which indicates the dominant spatial hydrogeochemical processes are consistent between 2018 and 2022. This is a crucial point supporting site selection. However, the same figure also reveals a secondary pattern: for a given lake type, samples often separate by year along PC1 (e.g., SSA lakes from 2018 have more positive scores than those from 2022). Table 2 quantifies this, showing significantly higher mean concentrations of marine-derived ions (Cl, SO₄, Br, Mg) in 2018. This interannual variability in the intensity of the sea-spray signal is an important finding, but its implications for the transfer function need further exploration to fully support the conclusion of robust "hydrogeochemical stability." Please address the following

Impact on model calibration: the diatom-EC transfer function is calibrated on a composite dataset spanning years with demonstrably different SSA intensities (2006, 2018, 2022). This is standard practice, but the observed variability raises a critical methodological question: Does this interannual variation in the modern calibration gradient introduce uncertainty or bias into the species' inferred environmental optima, particularly at the high-EC end critical for reconstructing strong wind periods?

To conclusively demonstrate the model's robustness and the validity of the "stable conditions", a **quantitative sensitivity analysis** must be added. This analysis should

test how the model's core parameters change when calibrated on subsets representing different SSA conditions. **Specifically, we request you perform subset calibration:** Recalculate the WA transfer function using two distinct modern calibration sets

Set A (Higher SSA): Combine the 2006 coastal data with the 2018 plateau data.

Set B (Lower SSA): Use the 2022 plateau data only.

For key high-EC indicator taxa (e.g., *Planothidium lanceolatum*, *Fragilaria capucina*), compare their inferred conductivity optima and tolerances between Set A and Set B. Present this in a supplementary table.

Compare the overall model performance metrics (RMSEP) and the observed vs. predicted scatter for the two subset models. If species optima and model performance are consistent between subsets, it provides quantitative proof that the model is robust to interannual SSA variability, strongly reinforcing the conclusion of a stable, transferable relationship. If significant shifts in optima occur, it quantifies a potential source of error. This would necessitate a discussion on the implications for reconstructing absolute wind strength and would strengthen the manuscript by rigorously defining the model's calibration uncertainty.

CS: We thank the reviewer for highlighting the interannual variability in SSA intensity and its potential implications for the calibration dataset. Fig. 4 and Table 2 show that, within lake types, samples from 2018 generally have higher concentrations of marine-derived ions (Cl, SO₄, Br, Mg) compared with 2022. This reflects real interannual differences in sea-spray input, likely driven by both the timing of sampling and natural year-to-year variability. Specifically, the 2018 sample represents a single late-summer sampling, whereas the 2022–23 dataset averages multiple summer samples. Stable isotope data indicate enhanced evaporation for the 2018 point, but even when comparing this to the late-summer 2022 samples (E3), 2018 shows higher ion concentrations than 2022, suggesting stronger SSA input during that year.

Importantly, the diatom–EC transfer function is calibrated only on the 2006 coastal dataset and mean 2022 plateau dataset; the 2018 data were not included. Therefore, interannual variability between 2018 and 2022 does not directly influence the calibration of the transfer function. Within 2022, ion concentrations are consistent across the multiple sampling events, confirming that the modern calibration dataset represents stable hydrogeochemical conditions.

Taken together, these observations support the conclusion that the transfer function is robust, and the inferred species optima are not affected by interannual SSA variability outside the calibration dataset. We will clearly state the data used in the quantitative

models to clarify this issue. However if another sensitivity test is still recommended we can include this.

Comment 4: Disentangling EC and nutrient signals (Line 567-574 & Figure 8).

While the manuscript appropriately attributes low plateau-lake nutrient variability to recovered conditions, the combined calibration dataset explicitly includes a high-nutrient gradient via the 2006 coastal lakes. The variance partitioning (Fig. 8) shows major shared variance between EC and nutrients (TON, Phosphate), yet the partial CCA (Table 5) indicates negligible unique nutrient effects. This raises a critical, unresolved question: does the dominant EC signal ecologically overwhelm nutrient influences in this system, or does it statistically mask a co-varying nutrient effect, particularly in coastal lakes where both marine ions and organic nutrients are elevated? Please expand this discussion to explicitly interpret Fig. 8, clarifying whether the weak unique nutrient signal is an ecological reality of the recovered state or a statistical artefact of collinearity, thereby justifying the exclusive focus on EC for paleoclimatic inference.

CS: We agree that the co-variation between EC and nutrients requires clearer interpretation. In this study, EC represents the dominant environmental gradient structuring diatom assemblages across Macquarie Island lakes.

Variance partitioning (Fig. 8) shows that EC shares a substantial proportion of explained variance with TON and phosphate, particularly in coastal lakes where marine ions and organic inputs co-occur. However, partial CCA results (Table 5) demonstrate that nutrients explain negligible unique variance once EC is accounted for, whereas EC retains a substantial independent contribution. Statistically, this pattern reflects collinearity along a shared coastal–marine exposure gradient, rather than an independent nutrient effect on diatom composition.

Ecologically, this collinearity is consistent with the recovered state of the system. In plateau lakes, nutrient concentrations have little variation and are persistently low and near detection limits and are therefore unlikely to exert a primary control on diatom assemblages. Varying and elevated nutrient concentrations only occur in a subset of coastal lakes influenced by marine inputs and animal activity, where nutrients covary with EC as part of a broader marine signal. Thus, the weak unique nutrient signal is not merely a statistical artefact but reflects an ecological reality in which EC exerts a first-order control on diatom communities.

Importantly, this collinearity does not compromise the transfer function, as nutrients are not included as predictor variables and EC retains a strong, independent relationship with diatom assemblage composition, and captures changes across the dataset as a whole, which other variables are not capable of. We will expand the Discussion to explicitly interpret Fig. 8 and Table 5 and to clarify that the apparent nutrient influence arises from statistical collinearity rooted in system ecology, thereby

justifying our exclusive focus on EC for palaeoclimatic inference. We also suggest including some simple linear regression biplots between EC and nutrient variables (Si, TON, Phosphate) from coastal sites in supplementary material to further illustrate the low correlation between them (rough figures included in attached document). While these variables both increase across coastal sites there is little relationship between them.

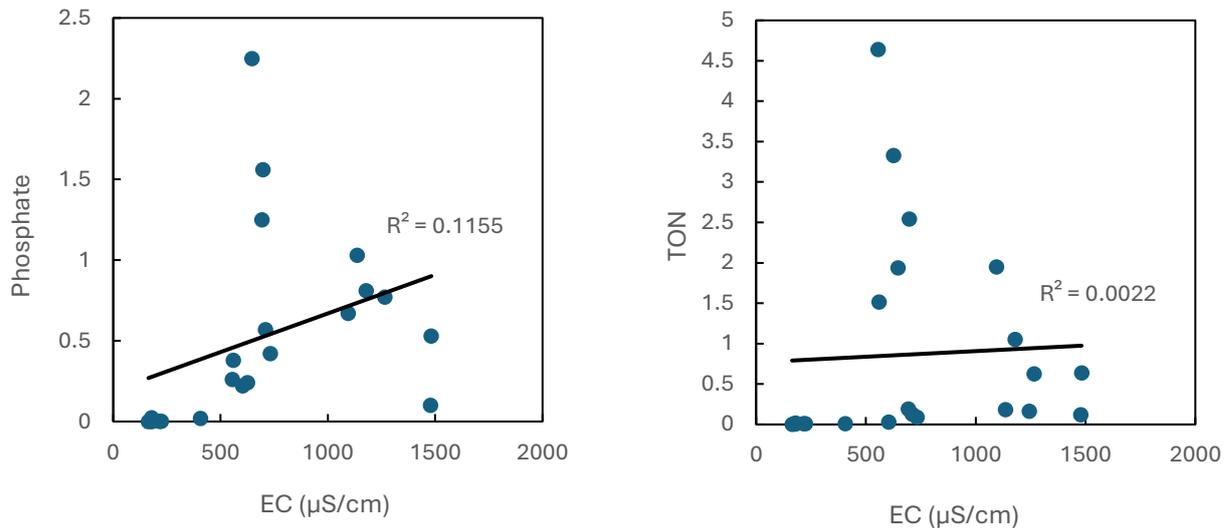


Fig 1: linear correlations between electrical conductivity and nutrient variable from coastal sites.

Comment 5: From suggestion to synthesis for multiproxy research (Line 625–632)

The conclusion on multi-proxy applications is currently vague. Please expand this into a concise synthesis paragraph proposing a specific framework. For example, explain how $\delta^2\text{H}/\delta^{18}\text{O}$ could constrain evaporative effects on EC, or how mercury isotopes could independently validate wind-driven aerosol deposition. Outline one or two testable hypotheses that future multi-proxy studies on Macquarie Island could address to resolve the interplay between SHW strength, precipitation, and evaporation.

CS: We thank the reviewer for this constructive suggestion. We agree that the discussion of multi-proxy applications would benefit from a more explicit and testable framework. In the revised manuscript, we will expand this section into a concise synthesis focused primarily on mercury (Hg) isotopes as an independent proxy for atmospheric transport and deposition, which is particularly well suited to the remote setting of Macquarie Island, and such work is currently underway (Schnieder et al. 2024; Li et al. 2025). While stable water isotope ($\delta^2\text{H}/\delta^{18}\text{O}$) data are not currently available for Macquarie Island, we acknowledge their potential for constraining precipitation and evaporation and their influence on EC. We identify $\delta^2\text{H}/\delta^{18}\text{O}$ as an important avenue for future research rather than a component of the present work.

Comment 6: Some important environmental variables are missing, for example, the effect of water depth on diatom assemblages (Farqan Muhammad et al., 2025, Quaternary International).

CS: We thank the reviewer for this suggestion and for highlighting the potential importance of water depth as an environmental control on diatom assemblages. We agree that water depth can be an important variable in many lacustrine systems, however it is often a composite variable that can reflect (Birks, 1998) complex, underlying gradients of habitat (sediment type, macrophyte cover), light, water chemistry (salinity, nutrients, oxygen), and taphonomy, all of which influence diatom composition. Depth often acts as a surrogate variable for what are complex environmental gradients, and the precise causal relationships underlying the observed correlation between depth and diatom distribution are often largely unknown and unquantified (Birks 1998; Juggins et al. 2013).

Furthermore, water depth was not included here as an explanatory variable because all surface sediment samples were collected within a very narrow depth range (0–1.5 m). Within this restricted range, variation in water depth is minimal and unlikely to generate meaningful gradients in light availability or other depth-related ecological controls. As a result, water depth would act as a largely invariant parameter in the calibration dataset and would not contribute explanatory power to the transfer function.

While the narrow sampling depth range was reported in the Methods, we recognise that the rationale for excluding water depth as an environmental variable was not made sufficiently explicit. We will revise the manuscript to clearly state that the limited and shallow depth range of the samples renders water depth ecologically negligible in this dataset, and that associated variables for which depth often acts as a surrogate (e.g. light availability) are therefore assumed to be broadly comparable across sites. This was avoided as including non-causal variables which act as surrogates for unknown or unquantified ecological factors can lead to spurious and misleading results (Juggins et al. 2013).

Comment 7: Lines 340-341: How many diatom taxa are classified as ‘unknown’?

CS: 21 unknown species remained in the > 1% relative abundance dataset used to develop the quantitative models. We will include such information in the results.

Comment 8: In the Discussion section, figure numbers should be inserted where appropriate.

CS: We thank the reviewer for drawing our attention to comment 8 and the comments listed below (comment 9 and minor issues). We will revise the manuscript to address them as suggested.

Comment 9: In the Introduction and Discussion sections, more studies on diatom–EC transfer functions and the indicative significance of diatoms for salinity should be discussed, for example, Farqan Muhammad et al. (2025, Ecological Indicators).

Minor issue:

Line 21: “Sea-spray inputs” – Specify whether this refers to dry deposition, wet deposition, or both, as this affects the wind-salinity relationship.

Line 63: “from the late 1900s to early 2000's” change to “late 1900s to early 2000s” (no apostrophe).

Line 70: “were conducted during a period of disturbance related to introduced invasive vertebrates rather than when the island was in a natural state” → awkward; rephrase to “were conducted during a period of vertebrate-induced disturbance rather than under near-natural conditions”.

Lines 79–82: Mention whether lake type (SSA, catchment, rainfall) was validated with independent data or is based solely on the Meredith et al. (2022) classification.

Line 121: “drier windier summers” add comma “drier, windier summers”.

Line 126: Figure 1 caption “modern core SHW belt (50–55S°)” correct to 50–55°S.

Lines 203–207: Justify the use of both parametric and non-parametric statistical tests in the same dataset. Consider providing a table of test results for clarity.

Lines 347–355: Mention whether any taxa (Species distributions) are exclusive to coastal vs. plateau lakes, and the implications for salinity inferences.

Lines 437–440: Figure 11 – Label panels (a–d) clearly in the caption and refer to them in the text when discussing model performance.

General – units and formatting: EC is sometimes $\mu\text{S cm}^{-1}$, sometimes not specified. Standardize to $\mu\text{S cm}^{-1}$ everywhere. Also check consistency of $\delta^{18}\text{O}$ / $\delta^2\text{H}$ notation (superscript vs plain text).

Line 345: In the taxon name *Psammothidium confusum var. atomoides*, ‘var.’ should be in regular font rather than italics.