



Assessing the contribution of extratropical cyclones to river floods that caused property damage in Quebec, Canada

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Abstract: In the past few decades, the province of Quebec in eastern Canada has experienced widespread and costly flood events, yet the contribution of extratropical cyclones to these floods remains unquantified. This paper presents a methodology for analysing the contribution of extratropical cyclones (ETCs) to 498 fluvial floods that occurred in watersheds in southern Quebec, Canada, between 1991 and 2020, and highlights key characteristics shared by the 200 flood ETCs (storms that contributed to flooding) identified in the study. This analysis combined reconstituted river outputs, government financial aid claims following flooding, ETC tracks, and precipitation data to identify flood events and their relevant ETCs. ETC contribution was defined as the percentage of rainfall during a search window surrounding each flood event that was associated with the relevant ETCs. Most of these floods occurred in the spring (74.7 %). The majority (72.7 %) of flood events had a high (50–75 %) to very high (> 75 %) ETC contribution, and only 2.6 % of events (mostly occurring in summer) had a negligible (< 5 %) ETC contribution. Flood ETCs had larger percentages of tracks originating from the Central US & Gulf of Mexico (+30.3 %) and lower percentages of tracks originating from Quebec & Maritimes (–16.7 %) and Western Canada & Pacific Ocean (–12.8 %) compared to non-flood ETCs (storms that did not contribute to flooding but spent at least one hour inside the *Baseline Southern Quebec domain*, $N = 6,027$). Flood ETCs also spent on average 18.2 h more in the domain and generally tracked over southern Quebec as opposed to over the Atlantic Ocean along the east coast of North America or over northern Quebec like non-flood ETCs. The percentages of financial aid claims allotted to flood ETCs were highly variable. Just five of the 200 flood ETCs were associated with 51.0% of all financial aid claims filed during the 30-year period, indicating that a high to very high ETC contribution rarely resulted in widespread flooding. The method developed in this study could be applied to other regions or types of storms analysed (e.g. tropical cyclones and mesoscale convective systems) to further our understanding of the atmospheric causes of flooding and ultimately enhance flood preparedness.



1 Introduction

30 Globally, between 2003 and 2022, floods affected 74.6 million people, had an associated cost of 41.1 billion US dollars, and led to 5518 deaths, on average, per year (Centre for Research on the Epidemiology of Disasters et al., 2024). According to the Canadian Disaster Database, flooding was the most frequent natural disaster in Canada between 1900 and 2005 (public Safety Canada, 2025b; Sandink et al., 2010). Canada's largest and second most populous province, Quebec, received 51 % of Canada's flood disaster assistance between 1988 and 2001 (Roy et al., 2003), despite accounting for only 23.8 % of the
 35 country's population in 2001 (Statistics Canada, 2025b). In the past 30 years, Quebec has experienced deadly, widespread and costly flood events. The Saguenay flood of July 1996 claimed ten lives (Tremblay and Guillaud, 2019), while the spring of 2017 and 2019 floods affected over 250 municipalities each time (Bourgault et al., 2022; Ministère de l'Environnement, de la Lutte contre les changements climatiques, de la Faune et des Parcs, 2025). More recently, in the summer of 2024, flooding caused by the remnants of Hurricane Debby resulted in the costliest natural disaster in the history of the province (Insurance
 40 Bureau of Canada, 2024).

Spring freshet (fluvial) flooding, which is caused by a combination of snowmelt and rainfall, is the most common type of flooding in Quebec (Javelle et al., 2003; Sandink et al., 2010), with the floods of spring 2011, 2017, and 2019 being prime examples (Benoit et al., 2022; Riboust and Brissette, 2016; Saad et al., 2014; Teufel et al., 2019). In the snow-free season, heavy or prolonged rainfall can also lead to fluvial flooding, as was seen during the Saguenay flood of 1996 (Tremblay and
 45 Guillaud, 2019). Flash flooding can also occur in the summer due to intense, short-lived rainfall. This is not surprising, given that Quebec experiences some of the highest short duration (less than 1 hour) maximum rainfall intensities in Canada (Buttle et al., 2016). Other mechanisms that can cause flooding in the province include rain on snow, ice jams, and storm surges (Buttle et al., 2016; Danard et al., 2003). Ultimately, fluvial flooding is most often caused by excessive precipitation (Bourgault et al., 2022), which is set to increase by 10–40 % in Quebec by the end of the century (Martel et al., 2021) in combination with
 50 snowmelt. Since extratropical cyclones (ETCs) are responsible for most of the precipitation in the middle latitudes, they are at the forefront of flood-triggering mechanisms in regions dominated by fluvial floods (Hawcroft et al., 2012).

Over Quebec, the precipitation associated with ETCs accounts for up to 85 % of the total winter (DJF) precipitation and at least 70 % of the total summer (JJA) precipitation (Hawcroft et al., 2012). Additionally, the proportion of extreme hourly precipitation events associated with ETCs ranges from 50–95 % throughout the year, being closer to 80–95 % in DJF and
 55 reaching the lowest values in JJA (Chen and Di Luca, 2025; Pfahl and Wernli, 2012). For instance, in the summer of 1996, an ETC travelled from southern Manitoba, traveled over the Great Lakes, and then moved over the Saguenay region, where it brought more than 250 mm of rainfall to some areas in less than 48 h. The cyclone remained stationary for 24 h due to blocking by two anticyclones located to the east, which prolonged precipitation over the Saguenay region (Milbrandt and Yau, 2001; Tremblay and Guillaud, 2019). In spring of 2017, two consecutive ETCs were associated with the flooding of the Ontario
 60 River basin. The first one originated from the central US, while the other travelled up the east coast of North America before reaching Quebec (Teufel et al., 2019). Similarly, the fall 1996 flood, which resulted in the evacuation of 1,000 residents in the

Montreal and Mauricie regions (Public Safety Canada, 2025a), was linked to an ETC moving up the east coast of North America and becoming blocked by a strong anticyclone (Perrier and Slivitzky, 1999). Thus, a thorough analysis of the causes of flooding in Quebec must include the quantification of the contribution of ETCs, given that they are associated with most of the precipitation, including extreme precipitation events.

The province of Quebec is surrounded by major bodies of water such as the Hudson Bay, the Great Lakes, and the eastern seaboard, which can increase heat fluxes and intensify cyclones (Chen et al., 2022a; Klock et al., 2002). Thus, storms that cause flooding can develop in many regions, typically to the south or west of the province, before moving east or northeast. Mackenzie lows, which form over the Mackenzie River Basin, and Alberta lows, also known as Alberta clippers, originate over the Northern Rockies (Klock et al., 2002). Colorado lows form to the southwest of Quebec, in the Midwestern US. As they both originate to the west of Quebec, the tracks of the Colorado and Alberta lows sometimes converge over the Great Lakes region (Chen et al., 2022a). Great Lakes lows and Hudson Bay lows form over large bodies of water to the southwest and northwest of Quebec, respectively. Hatteras lows form to the south of Quebec, near Cape Hatteras, over the US eastern seaboard (Klock et al., 2002; Poan et al., 2018) and can cause hurricane-level damage to the east coast of North America during the cold season (October–April) (Hirsch et al., 2001). Finally, Gulf lows form over the Gulf of Mexico, to the south of Quebec (Klock et al., 2002).

Although some studies have linked the presence of cyclones to specific floods in Quebec (Lin et al., 2019; Milbrandt and Yau, 2001; Rickard et al., 2025; Teufel et al., 2019), these have only considered a few, high-impact, well-known cases. Similarly, case studies examining the role of ETCs or storms in the eastern US (near our study region) have also focused on a small number of larger floods (Barros and Kuligowski, 1998; Miller, 1990; Smith et al., 2010; Su et al., 2023). River discharge (flood peak) and rainfall data obtained from rain gauges or radar estimates are often used to characterize floods. Therefore, historically, floods occurring in ungauged watersheds could not be analyzed and had to be excluded from case studies, even when they had caused significant damage. Furthermore, using rain gauge data to characterize floods has its limitations as it does not indicate where the consequences of flooding (primarily material loss) were felt in the watershed. Data on material losses following flooding is rarely made available, as it contains sensitive information on claimants, which is why there have been limited large-scale studies on the contribution of ETCs to flooding.

This paper aims to quantify the contribution of ETCs to fluvial flooding and its seasonality in southern Quebec between 1991 and 2020. More specifically, we investigate whether ETCs that trigger flood events have any common characteristics (e.g. cyclogenesis location, trajectories and lifetime). Combining data on governmental financial assistance with reconstituted river outputs enables us to identify floods objectively and systematically during the period of interest. Government financial assistance data provides evidence of where and when floods caused damage, while reconstituted river output data allows us to cover ungauged watersheds. As financial aid data depends on the presence of human populations, it should be noted that the study only focuses on river floods that caused property damage. Furthermore, as the focus is put on the atmospheric processes leading to flooding, soil moisture and snowpack aspects are not considered in this analysis.



95 The article is structured as follows. First, Sect. 2 details the various types of data included in this project. Section 3 describes the methodology used to identify flood events, as well as to find and classify relevant ETCs. Then, Sect. 4 presents findings on the distribution of flood events, the contribution of ETCs to flooding, the classification and differences of ETCs as well as the sensitivity analysis of key variables. Finally, Sect. 5 discusses the seasonality of floods and ETC contribution and contextualises the results on flood ETCs within the existing literature.

100 2 **Data**

Figure 1 provides an overview of the article's methodological approach, indicating the baseline datasets (in green), the methods applied (in white, which are described in Sect. 3) and the resulting data of interest (in blue).



3. Methods

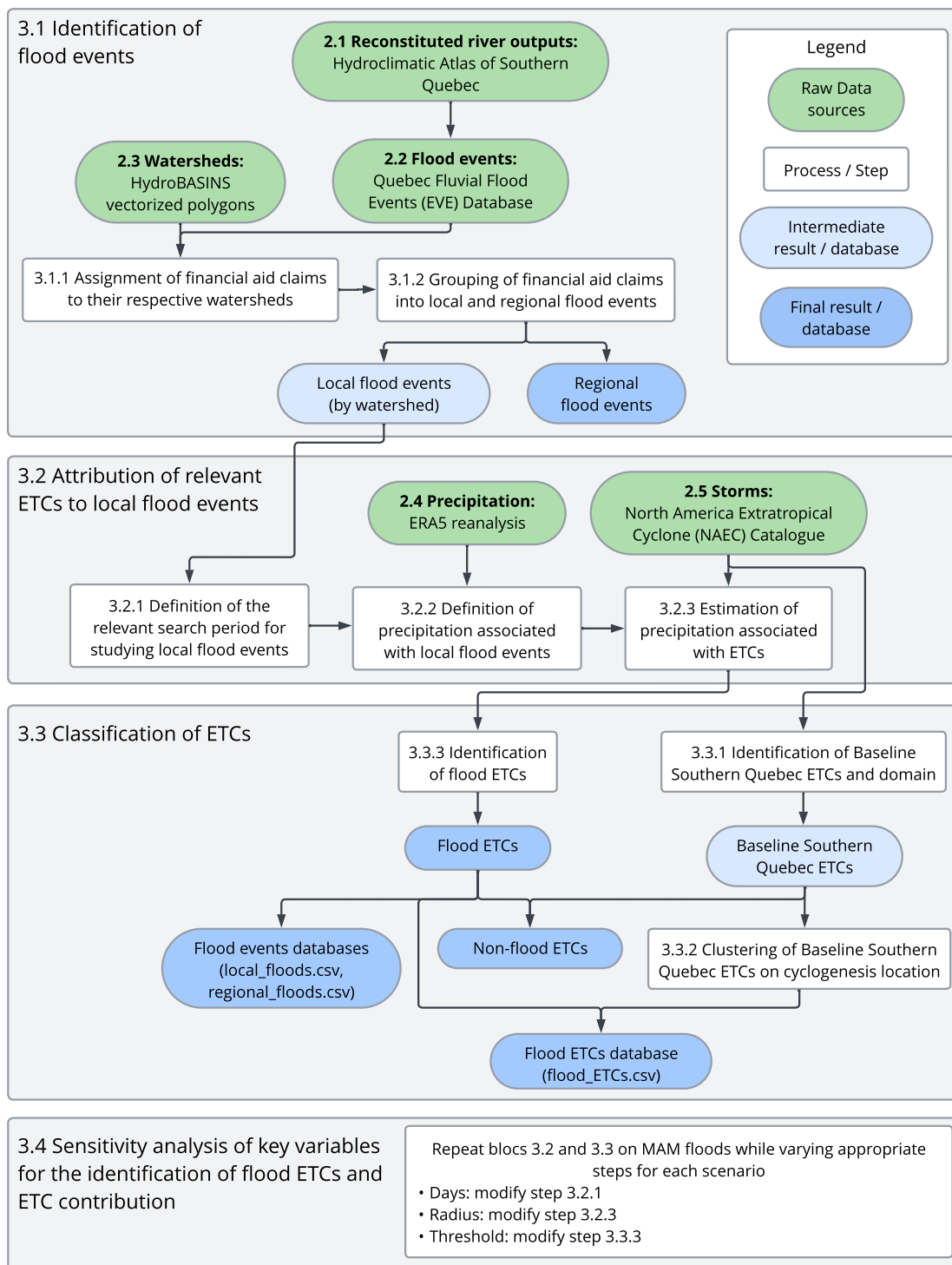


Figure 1. Overview of the methodological steps (described in Sect. 3) and sources of data (described in Sect. 2) used in the study.



105 2.1 Reconstituted river outputs: Hydroclimatic Atlas of Southern Quebec

The *Hydroclimatic Atlas of Southern Quebec* contains time series of river discharge for over 10,000 mostly ungauged river sections across meridional Quebec. It is the study's primary source of hydrological data (Centre d'Expertise Hydrique du Québec, 2025; Lachance-Cloutier et al., 2017). These time series of retrospective daily streamflow estimates are available from 1970 to 2024. They focus on a region of southern Quebec spanning approximately 640,000 km² and located between 45–
 110 53°N and 57–80°W (see Figure 1 in Lachance-Cloutier et al., 2017). For ungauged river sections, discharge time series were reconstructed using statistical interpolation based on observations from 75 gauged sites and outputs from a hydrological model used as the background field. The model field was adjusted to correct spatial discrepancies between observations and initial model estimates. The hydrological model used for this purpose is Hydrotel: a spatially distributed, mostly physically-based model. Hydrotel divides watersheds into relatively homogeneous hydrological units (RHHUs) and river reaches. Rainfall–
 115 runoff processes are simulated using RHHUs, and streamflows are simulated using the sub-sections of river reaches (Lucas-Picher et al., 2020). Hydrotel was developed to enable remote-sensed and GIS data (e.g. digital elevation models, soil type and land use maps) to be used as input. Physical or empirical relations are used in the model to simulate the hydrological processes occurring in watersheds, such as snowmelt, surface runoff, river routing and evapotranspiration (Fortin et al., 2001).

The 2-year recurrence flood values for the river sections included in the study were calculated using the average annual
 120 maximum discharge values available from the Hydroclimatic Atlas of Southern Quebec between 1991 and 2020. These values were then inputted in the Quebec Fluvial Flood Events (EVE) database to provide additional context surrounding the flood events.

2.2 Flood events: Quebec Fluvial Flood Events (EVE) Database

Flood events in southern Quebec were identified based on 17,587 financial aid claims available in the EVE database of
 125 the Quebec Ministry of Public Security (Roy et al., 2024). This database contains claims filed by individuals or businesses in Quebec following flood events between 1991 and 2020. Where a claim was filed, the government provided financial aid to either repair damaged buildings and homes or help people relocate and settle elsewhere. As the database relies on the presence of flood-damaged communities, it only contains data pertaining to populated areas. Consequently, in Quebec, the available data are concentrated in the southern part of the province, along the Saint Lawrence River and Lac Saint-Jean, where most of
 130 the population lives. This area is located between latitudes 44.99°N and 50.47°N, and longitudes 59.61°W and 79.42°W. The rest of Quebec is sparsely populated (Statistics Canada, 2025a).

The database contains the location of the damaged building for each claim, as well as important information relating to the nearest river section associated with the flood event. It also lists the dates on which the rising and falling limbs of the hydrographs cross the 2-year recurrence flood value (calculated using data from the Hydroclimatic Atlas of southern Quebec)
 135 for each river section. These dates serve as a reference point for the timing and duration of flooding in relation to each claim.



2.3 Watersheds: HydroBASINS vectorized polygons

The watersheds delineations used to define the research domain were extracted from HydroBASINS, a secondary data product from the hydroSHEDS project (Lehner and Grill, 2025). This dataset contains sub-basin boundaries at various scales covering the entire globe and is available in polygon format (Lehner et al., 2008; Lehner and Grill, 2013). HydroBASINS features 12 nested, hierarchical sub-basins levels, ranging from largest (level 01) to the smallest (level 12). The average size of level 12 sub-basins globally is 130.6 km² (Lehner and Grill, 2025).

2.4 Precipitation: ERA5 reanalysis

Precipitation data were taken from the ECMWF ERA5 reanalysis (Hersbach et al., 2020, 2025). This dataset combines a weather model with billions of observations around the globe from in-situ, upper air, and satellite data to reproduce actual and past meteorological conditions. This study uses hourly total precipitation and snowfall data with a spatial resolution of 0.25° × 0.25°, covering the period from 1991 to 2020.

2.5 Storms: North America Extratropical Cyclone (NAEC) Catalogue

The extratropical cyclones (storms) used in this study were extracted from the NAEC Catalogue between 1991 and 2020 (Chen et al., 2022b). This catalogue contains hourly tracks of extratropical cyclones over North America (20–80°N, 0–180°W) as well as various variables of interest, such as total precipitation, wind speed, and mean sea level pressure (MSLP), averaged over different radii around cyclone centers. Extratropical cyclones are tracked by monitoring the position of their centers over time. The algorithm used for storm detection is applied to the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA5 (Hersbach et al., 2025) reanalysis that has a horizontal grid spacing of 0.25° × 0.25°. Storms in the catalogue are identified through a MSLP local minimum, and must have a maximum vorticity (a measure of the spin of air particles, indicative of a region of convergence when positive) greater than 10⁻⁵ s⁻¹ within a 200 km radius of the local minimum (Chen et al., 2022a; 2022b).

3 Methods

Figure 2 shows how the methodology described in this section (and illustrated in Fig. 1) was applied in practice, using a sample representative local flood event (contained to one watershed) that occurred in November 1996 and resulted in 268 claims being filed.

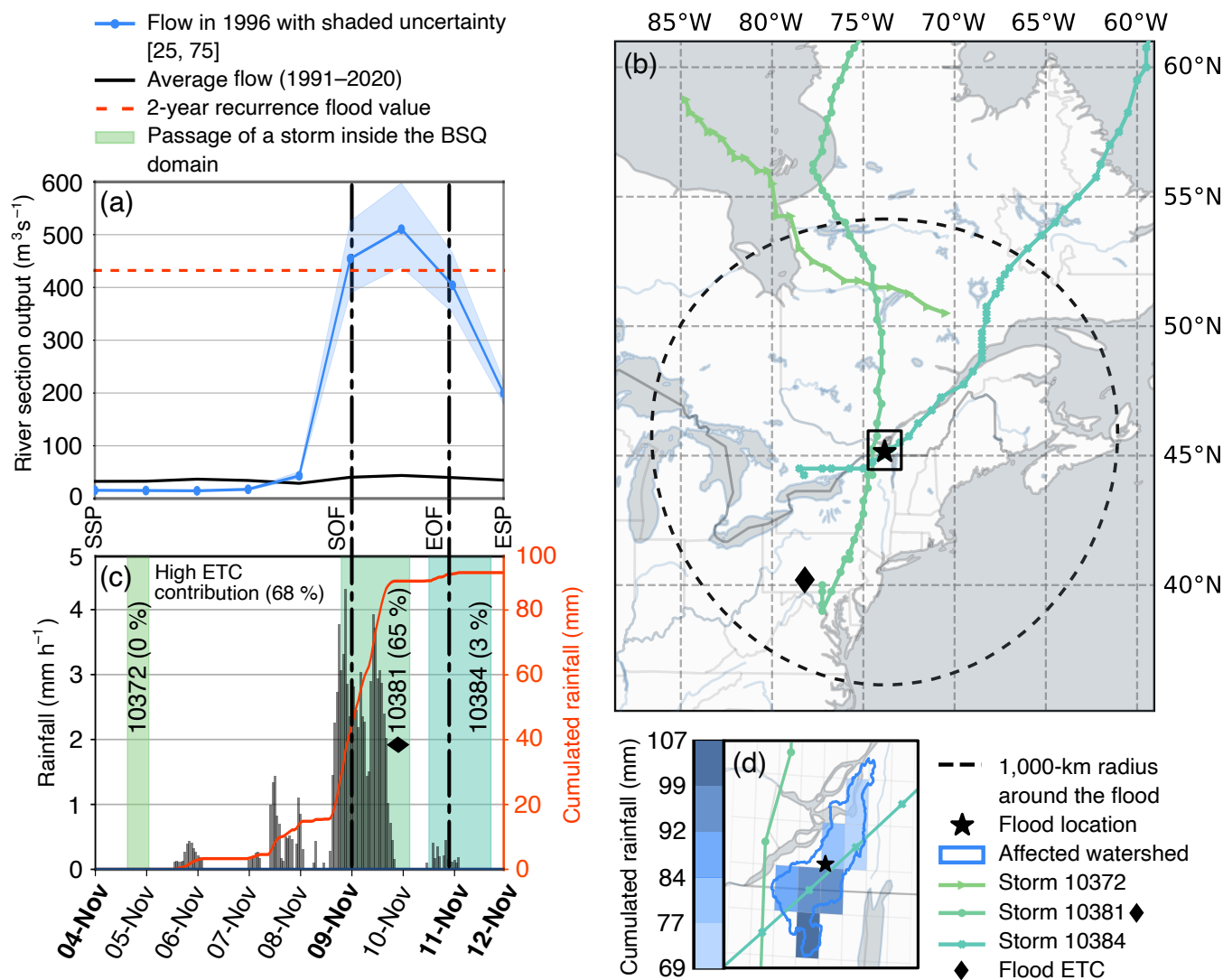


Figure 2. Application of the methodology to the flood event of November 1996 in the Montérégie region. (a) River section discharge (blue) compared with the 2-year recurrence flood value (dashed red) for the flood event. (b) Storm tracks of ETCs entering the search domain (dashed black) between the SSP (start of the search period) and ESP (end of the search period) dates for the flood event and location of the flood (star). Tracks are shown for their entire lifetime (even components before and after they leave the domain). (c) Rainfall in the affected watershed for the local flood event with the contribution of the different ETCs and overall ETC contribution. The SSP, SOF (start of flood), EOF (end of flood), and ESP dates are shown in (a) and (c) and highlighted in bold in (c). The SOF and EOF dates are shown by the dash-dot lines in (a) and (c). The identified flood ETC is highlighted with a diamond (storm #10381) in (b) and (c). (d) Close up view of the Montérégie watershed affected by flooding, flood location, and cumulated rainfall during the search period.

3.1 Identification of flood events

Flood events were identified by grouping financial aid claims from the EVE database according to two variables: location of the damaged buildings (sensitive data) and the Start of Flood (SOF) date. For each financial aid claim entry, the SOF was defined as the date on which the discharge of the river section associated with the claim exceeded the 2-year recurrence flood



value for the first time. This is an crude estimate of the true SOF, considering that each construction is subject to flooding at
175 its respective return period, which is usually greater than two years. The river section associated with the claim was the one
closest to the damaged buildings. The End of Flood (EOF) corresponded to the date of the last day on which the river section's
discharge was above the 2-year recurrence flood value – also a crude estimate. The SOF was always set at the beginning of
the day, and the EOF at the end of the day, as this still placed the two time markers 23 h apart if they occurred on the same
day (as daily river output values are used). Grouping financial aid claims in time and space was necessary to identify flood
180 events and subsequently quantify the contribution of ETCs to those events.

3.1.1 Assignment of financial aid claims to their respective watersheds

Financial aid claims were assigned to watersheds based on the level-7 hydroBASINS polygons (Lehner and Grill, 2025).
The level 7 polygons watersheds were selected for this purpose because they provided relatively small watersheds (average
size of 2,920 km² for those included in the study), yet were large enough to be accurately described using the ERA5 reanalysis
185 precipitation data. Of all North American hydroBASINS polygons, only the 98 watersheds containing at least one Quebec
financial aid claim were retained.

3.1.2 Grouping of financial aid claims into local and regional flood events

After grouping the financial aid claims by their watershed, they were organised by their SOF dates. If two (or more)
claims had SOF dates that fell within seven days of each other within the same watershed, they were grouped together as one
190 *local flood*. Not all claims aggregated into one local flood event had the same SOF and EOF dates, as the claims sometimes
were associated with more than one river section (with its unique SOF and EOF dates). Thus, the local flood's SOF and EOF
dates were defined as the earliest and latest dates, respectively, amongst their aggregated claims. This process is pictured for
a fictitious flood in Fig. 3, where seven financial aid claims, close to two flooded river sections are grouped into one local
flood on the affected watershed.

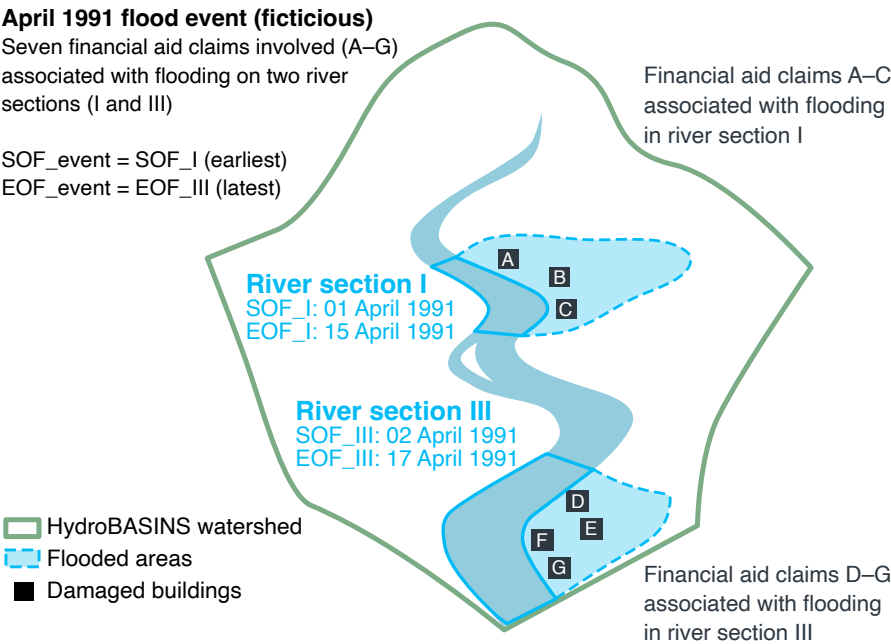


Figure 3. Visual representation of grouping financial aid claims into a local fictitious flood and selecting the correct SOF and EOF dates among the flooded areas’ associated river sections’ dates.

Finally, the locations of local flood events were defined by the centers of their assigned watersheds. This formed the basis of the search domain for ETCs related to the local flood. Local floods occurring simultaneously in different watersheds were grouped together as a *regional flood* if they had a difference of up to seven days in their SOF dates.

3.2 Attribution of relevant extratropical cyclones to local flood events

This section outlines the methods used to determine the time interval and physical domain of analysis for each local flood event. It also explains how we estimated the precipitation associated with the local floods, and consequently the different ETCs. These steps ensure that we correctly attribute the most appropriate ETC to each local flood event, accurately representing its contribution to flooding.

3.2.1 Definition of the relevant search period for studying local flood events

In order to quantify the contribution of ETCs to flooding, we first needed to define the time period during which we analysed the flood. This so-called *search period* is bound by the Start of the Search Period (SSP) and End of the Search Period (ESP). For the summer (JJA), autumn (SON), and winter (DJF) seasons, the SSP date for the data search was set to five days before the SOF date. For spring, the SSP date for the data search was set to 15 days before the SOF date (Bourgault et al., 2022). These different SSP dates were chosen based on known seasonal flooding patterns in the region and a careful inspection of the hydrographs of flooded river sections by season. Please note that we performed a sensitivity test with a shorter time interval between the SSP and SOF dates (10 days) on spring flood events, which is described in Sect. 3.4. For all seasons, the



ESP date for the data search was set to one day after the EOF date. These time markers can be visualized in Fig. 2a and Fig. 2c where the search period lasted eight days, from 4 November (SSP) to 12 November (ESP) and the flood occurred between 9 November (SOF) and 11 November (EOF).

3.2.2 Definition of precipitation associated with local flood events

Once the search period for studying floods had been defined, ERA5 precipitation data was gathered for the affected watershed of each local flood event. At each time step (hourly), the total precipitation and snowfall on the grid points within the watershed were averaged. This was done throughout the search period (between the SSP and ESP dates). The total cumulative values for snowfall, total precipitation, and rainfall were recorded for each analysed event. Rainfall was calculated by subtracting snowfall from total precipitation.

3.2.3 Estimation of precipitation associated with extratropical cyclones

Flood-related precipitation was attributed to ETCs using a *search radius* of 1,000 km around the affected watershed. For each event, ETCs from the NAEC catalogue that entered the search radius during the search period were considered relevant, and all watershed precipitation was assigned to them for as long as they remained within the domain. If there was more than one ETC within the search radius at any moment, the precipitation was associated with the ETC closest to the centre of the studied watershed.

The choice of the search radius was made following several publications on ETCs over North America (Chen et al., 2024; Chen and Di Luca, 2025; Chen et al., 2022a; Hawcroft et al., 2012) and it was decided to keep it constant independently of the seasons, to simplify the method. The choice of a 1,000 km radius is discussed further and compared to a smaller one (750 km) in Sect. 3.4.

The following variables were recorded for each storm occurring during a flood event: time interval in the search radius, quantity and percentage of total precipitation, snowfall, and rainfall over the search period. For each analysed flood event, the *ETC contribution* was defined as the percentage of the total rainfall associated with all ETCs in the search radius over the search period. The ETC contribution was then categorised according to five levels: negligible (0–5 %), mild (5–25 %), moderate (25–50 %), high (50–75 %), and very high (75–100 %). For the sample flood event shown in Fig. 2, the three ETCs present in the search domain accounted for 68.1 % of the total rainfall for the event, indicating a high ETC contribution.

Precipitation not associated with ETCs (such as 31.9 % of the rainfall that occurred during the November 1996 flood) could be attributed to different phenomena, i.e. convective systems, ETCs with a radius of influence greater than 1,000 km, or weak systems not reaching the necessary MSLP and vorticity thresholds to be featured in the NAEC Catalogue.

3.3 Classification of extratropical cyclones

This section outlines the process of classifying ETCs. This involves distinguishing between storms that contribute to flooding (*flood ETCs*) and those that do not (*non-flood ETCs*), and grouping them according to where they originated.



3.3.1 Identification of Baseline Southern Quebec extratropical cyclones and domain

To determine whether the storms identified as flood ETCs exhibited unique characteristics or resembled typical storms in the area, a general climatology of ETCs was conducted for the study area. To achieve this, a broader search domain with a radius of 1,000 km was established around the edges of all the watersheds involved in the study. This domain is referred to as the *Baseline Southern Quebec* domain. Any ETC that entered the Baseline Southern Quebec domain for at least one hour between January 1991 and December 2020 was thus tagged as a Baseline Southern Quebec ETC. Figure 4 shows this domain along with all watersheds included in the study (i.e., where financial aid claims were filed).

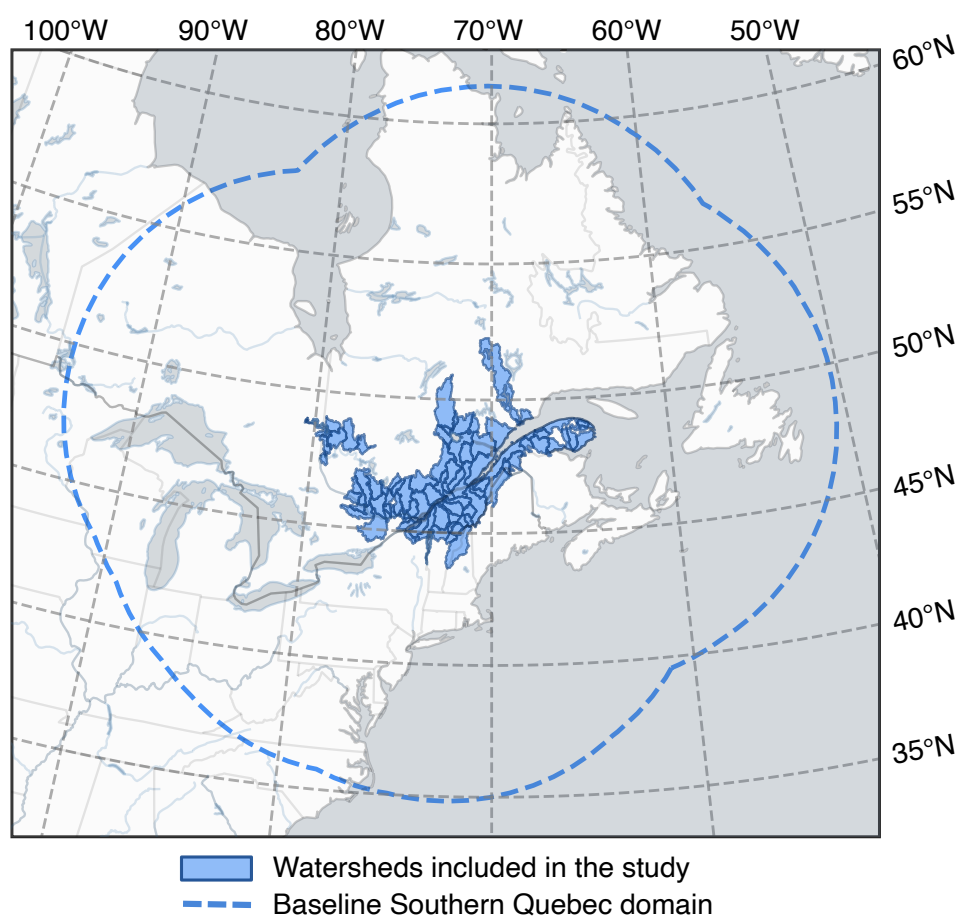


Figure 4. Watersheds included in the study (filled blue), and the Baseline Southern Quebec domain (dashed blue).

3.3.2 Clustering of Baseline Southern Quebec extratropical cyclones on cyclogenesis location

The Baseline Southern Quebec ETCs were clustered according to their genesis locations (i.e. their first recorded position in the NAEC Catalogue) using *k*-means clustering. This method limits the distance between the cluster centers and data points,



creating tightly packed clusters. The optimal number of clusters was determined by examining the silhouette score and elbow method plot, and by taking into account the typical regions of cyclogenesis in North America.

3.3.3 Identification of flood extratropical cyclones

260 The percentage of rainfall associated with individual storms during local flood events was used to distinguish the flood ETCs from the Baseline Southern Quebec ETCs. Hence, Baseline Southern Quebec ETCs were classified as flood ETCs if they accounted for more than 20 % (10 % in MAM) of the total rainfall for any local flood event. Otherwise, they were labelled as non-flood ETCs. The choice of the exact threshold (10 % vs 20 %) was based on observations of the maximal percentage of associated rainfall per storm by season. The longer search periods in MAM reduced the percentage of rainfall associated
 265 with individual storms, as the total rainfall accumulated over a longer period was usually greater. This motivated the need for a different threshold in spring. This threshold is compared to a higher one (15 %) in Sect. 3.4. An example of this distinction between storms is also shown in Fig. 2b–c where only one of the three ETCs present in the search domain qualified as a flood ETC (storm 10381, highlighted by a diamond).

To illustrate the magnitude of the consequences associated with the flood ETCs, their impact was assessed based on the
 270 number of related financial aid claims. For each local flood event, claims were assigned to the flood ETC that contributed the most to the cumulative rainfall of the event. This also meant that a specific storm could be allocated claims relating to more than one local flood event, since local flood events often occurred concurrently in different watersheds. It should also be noted that not all financial aid claims could be associated with a flood ETC, as some flood events did not feature one.

3.4 Sensitivity analysis of key variables for the identification of flood ETCs and ETC contribution

275 A sensitivity analysis was carried out on spring flood events to assess the impact of varying some key variables on identified flood ETCs and overall ETC contribution. Three different scenarios were studied, each with one key parameter varying from the default value, defined herein as the *control* case. The spring season was chosen for this analysis due to its higher number of flood events, allowing better comparison between the scenarios. The first scenario involved reducing the number of days between the SSP and SOF dates from 15 to 10 days, referred to as the *days* case. The second scenario involved
 280 reducing the search radius from 1,000 km to 750 km, referred to as the *radius* case. The third scenario involved increasing the threshold for storms to be considered flood ETCs from 10 % to 15 % of cumulated rainfall during the search period for a flood event, referred to as the *threshold* case. The set of variables associated with the three scenarios studied and those of the control case are shown in Table S1 in the Supplement.



4 Results

285 4.1 Annual and seasonal distribution of flood events

Table 1 shows the seasonal distribution of the 498 local floods identified at the watershed level in southern Quebec between January 1991 and December 2020. The majority of these floods (74.7 %) occurred during the spring freshet season (372 events). The remaining flood events were distributed as follows: 48 (9.6 %) in summer, 58 (11.6 %) in fall, and 20 (4.0 %) in winter. The average number of financial aid claims per local flood event were similar for spring, fall, and winter (between 26.1–32.3 claims per flood), but was higher for summer (71.7 claims per flood). This larger number in summer was due to a single summer flood event cumulating 1,668 claims (Saguenay flood of 1996), skewing the mean towards higher values. Otherwise, the median number of claims per local flood event were similar between the seasons (4–7 claims), with winter having the highest value. On average, 16.6 local flood events occurred each year (12.4 in spring, 1.9 in fall, 1.6 in summer, and 0.7 in winter). The worst year was 2017, when 75 local flood events occurred throughout the study area.

295 **Table 1.** Statistics on the seasonal distribution of flood events, their ETC contribution in percentage and qualitative level (in parenthesis), and number of financial aid claims associated

Season	ETC contribution		Number of financial aid claims				Number of flood events
	Mean	Median	Min	Median	Mean	Max	
MAM	61.8 % (high)	66.1 % (high)	1	4	32.3	1668	372
JJA	65.7 % (high)	83.5 % (very high)	1	5	71.7	1375	48
SON	69.5 % (high)	72.9 % (high)	1	4	27.5	446	58
DJF	71.2 % (high)	74.9 % (high)	1	7	26.1	197	20

Figure 5 shows the annual and monthly distribution of the 85 regional flood events (grouped local floods happening simultaneously in different watersheds) identified between January 1991 and December 2020. The figure indicates the number of financial aid claims (colour) and the number of watersheds (circle size) associated with each regional flood event, with both variables increasing non-linearly. More widespread regional events were typically associated with more claims, although not always. Regional events affecting over 35 watersheds and cumulating more than 1,000 financial aid claims all occurred in the spring season and were rare. Spring regional flood events were more widespread, involving 7.6 watersheds per event on average. The number of watersheds involved by regional flood event were similar in fall and summer, with 4.5 and 3.4 watersheds on average, respectively. Winter regional flood events were more contained, involving an average of 2.2 watersheds per event. The most widespread event involved 71 watersheds in spring 2017, spread over two months.

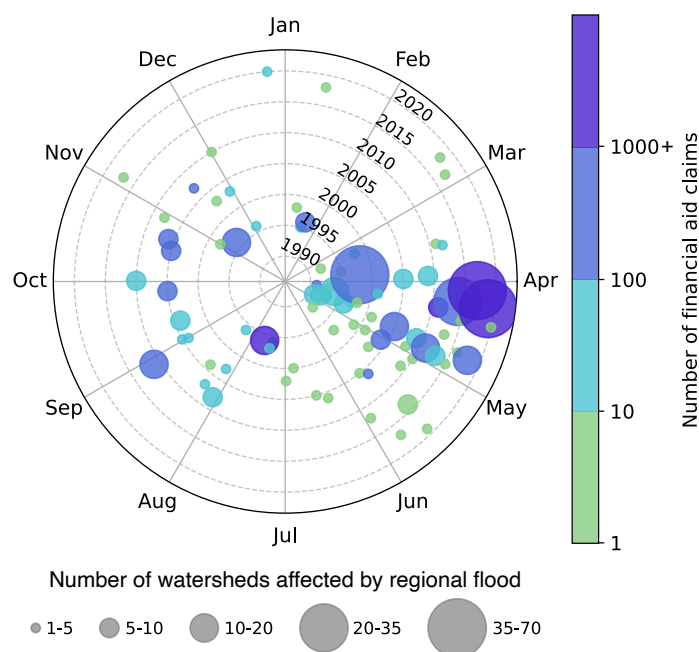


Figure 5. Yearly and monthly distribution of regional flood events (N = 85) between January 1991 and December 2020. The size of the circle indicates the number of watersheds (or local floods) involved in the event while the colour denotes the number of financial aid claims associated to the regional flood events.

4.2 Contribution of extratropical cyclones to flooding

Figure 6 shows the level of ETC contribution to flooding by season for all local flood events. Considering all seasons combined, 40.2 % of flood events had a very high ETC contribution. Then, the proportion of events decreased with ETC contribution: 32.5 % of events had a high contribution, 18.5 % had a moderate one, 6.2 % had a mild one, and only 2.6 % of events had a negligible ETC contribution. Summer recorded the highest proportion of local flood events with a negligible ETC contribution (14.6 % of all summer floods). The summer season also exhibited an ETC contribution pattern distinct from those of the spring and fall, with post hoc tests confirming a statistically significant difference with an adjusted α of 0.0083. In contrast, the summer ETC contribution did not differ significantly from that of winter. Table 1 also shows that, overall, the mean ETC contribution for all seasons was in the high (50–75 %) range. However, summer was the only season with a median ETC contribution in the very high range (83.5 %).

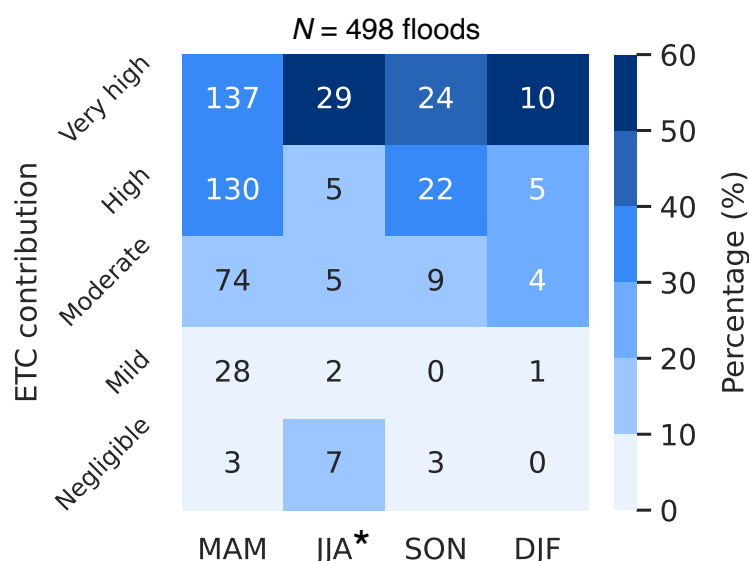


Figure 6. Level of ETC contribution to local floods, shown as percentages of the total number of local flood events by season (in shades of blue) and absolute number of events in each category (numbers in the cells). Significant statistical differences in the distribution by seasons is identified with an asterisk (*).

325 4.3 Comparison of flood and non-flood ETCs

4.3.1 Seasonality, origin, and lifetime of flood ETCs

Figure 7a shows the number of storms in each season and indicates whether or not they contributed to a flood event. Of the 6,277 Baseline Southern Quebec storms, 602 occurred during a flood event. Only 200 of these contributed significant rainfall, qualifying as flood ETCs (i.e. contributing at least 10 % of cumulative rainfall during the search period for a flood event in MAM, and at least 20 % in the other seasons). The majority of these flood ETCs were identified in spring, primarily due to the increased number of flood events during this season due to the contribution of snowmelt, as well as the longer search periods (averaging 24.0 days in spring, compared to 8.1, 6.8, and 6.6 days in summer, fall, and winter, respectively).

Of the 17,587 financial aid claims included in the study, 15,491 (88.1 %) could be associated with at least one flood ETC (meaning that 2,096 claims (11.9 %), distributed across 30 (6.0 %) flood events could not). Most of these flood events without a flood ETC occurred in spring (17 events, cumulating 1973 claims), some occurred in summer (8 events, 105 claims), and very few occurred in winter (2 events, 10 claims) and fall (3 events, 8 claims).

The typical lifetime of a storm in the Baseline Southern Quebec search domain depended on its type (flood vs non-flood ETC). Figure 7b shows the lifetime distribution per type, along with their respective medians (white line). The lifetime distributions per type were found to be statistically significantly different from one another (shown with an asterisk), with flood ETCs spending an average of 18.2 h longer in the Baseline Southern Quebec search domain than their non-flood counterparts. This was determined using a Kolmogorov-Smirnov (K-S) test with an α value of 0.05.

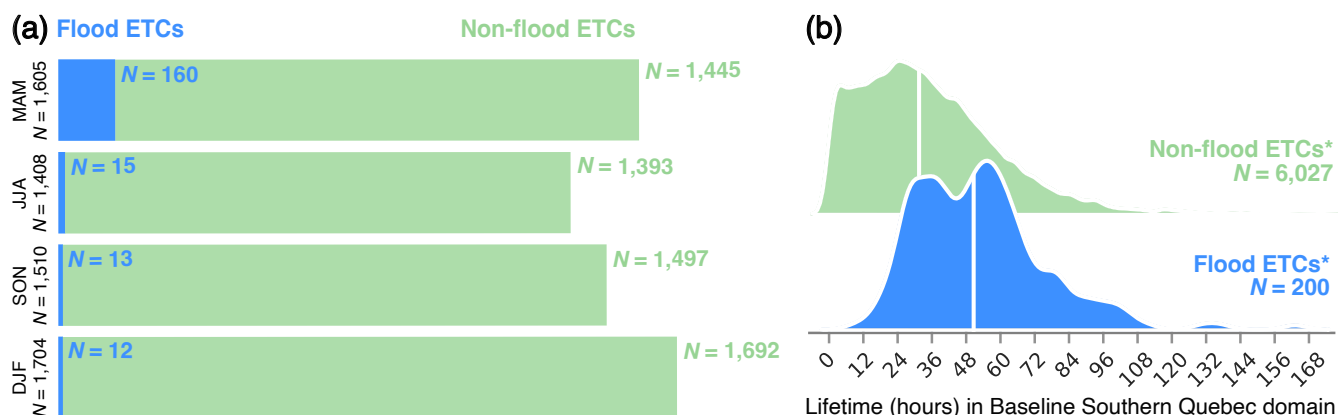


Figure 7. (a) ETCs included in the study with their season and type (flood ETC and non-flood ETC). (b) Lifetime of all ETCs in the Baseline Southern Quebec search domain, shown for flood and non-flood ETCs. The medians are shown by the white lines. The asterisk (*) denotes a statistically significant difference in the distribution according to the K-S test.

Figure 8 shows the density of cyclogenesis locations for both flood (Fig. 8a) and non-flood (Fig. 8c) ETCs. The darkest areas represent the regions with the highest density, which differed between the two groups of storms. For flood ETCs, the regions with the highest density of cyclogenesis locations were the central US (over eastern Colorado) and, to a lesser extent, the east coast of the US (over northern Virginia). For non-flood ETCs, these regions were found in western Canada (over Alberta), as well as off the east coast of the US and over the Atlantic Ocean. There are clear differences in the density of cyclogenesis locations between the two groups, with maximum densities only reaching up to 6 % for the non-flood ETCs group.

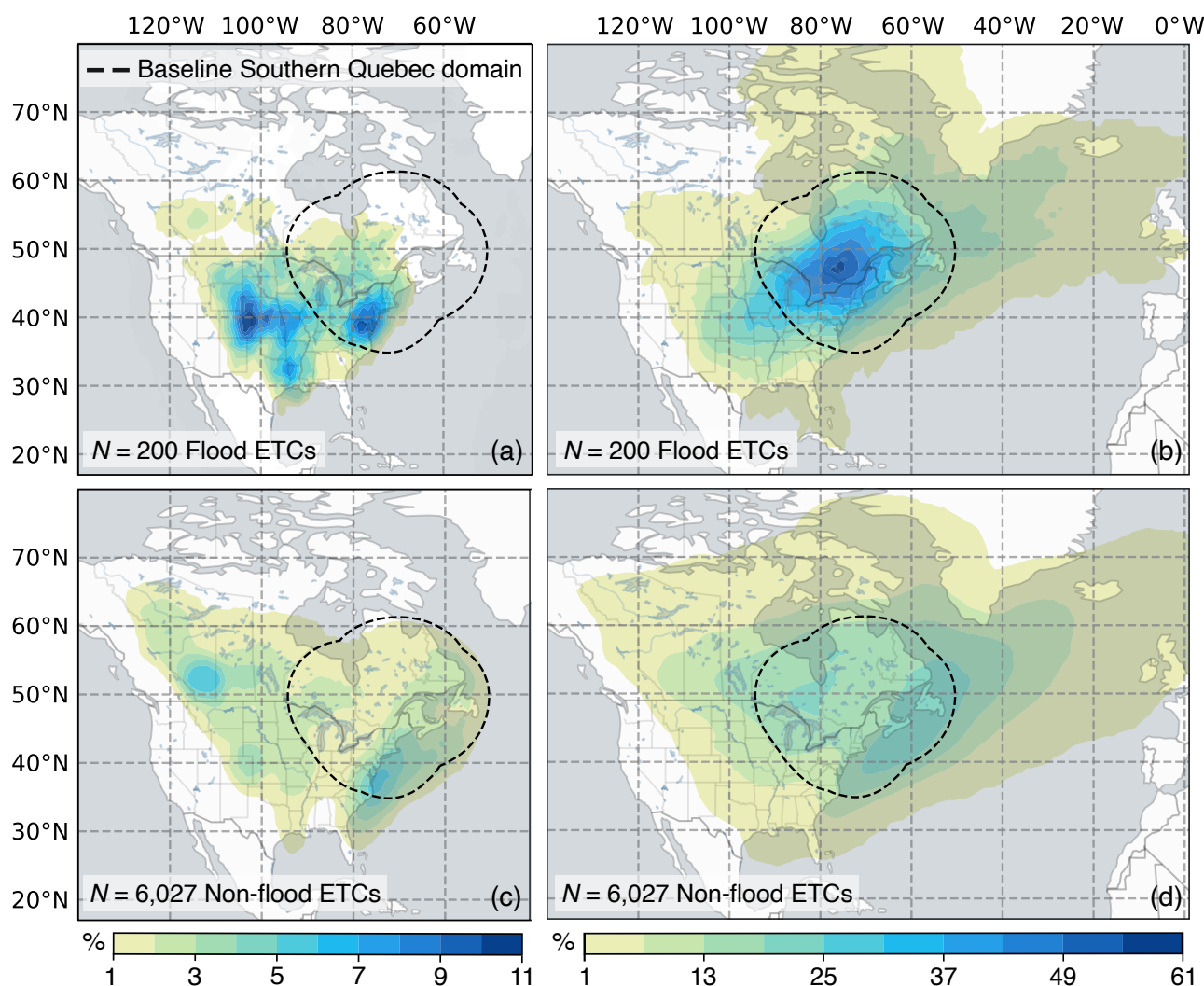


Figure 8. (a) Cyclogenesis locations density map for flood ETCs. (b) Track density map for flood ETCs. (c) Cyclogenesis locations density map for non-flood ETCs. (d) Track density map for non-flood ETCs. The Baseline Southern Quebec search domain is shown in dashed black, colors indicate the density (as a percentage of total storms for the sub-plot tracking/generating within 350 km of grid points) from highest (dark) to lowest (light).

Figure 8 also shows the track density for flood (Fig. 8b) and non-flood (Fig. 8d) ETCs. The flood ETCs tracks, which originated mostly in the central and eastern US, had streamlined trajectories that went roughly northeast, passing through the Great Lakes region and fanning out as they entered Quebec from the SW. Many, though not all, flood ETCs track over southern Quebec (> 49 % of the tracks). For non-flood ETCs, the area of greatest tracks density (up to 25 % of all tracks) ran along the eastern seaboard of North America (from the US coast of Massachusetts to the Canadian coast of Newfoundland and Labrador) over the Atlantic Ocean. A secondary, smaller area of high track density for non-flood ETCs ran above the Great Lakes, near the border between Ontario and Quebec. Areas of high track density for non-flood ETCs can also be visualized in Fig. S3 in



365 the Supplement, where the colorbar is unique to each plot, allowing for better visualisation of the tracks and genesis location densities for non-flood ETCs.

Another noticeable difference between the tracks of the two groups is the area of highest track density compared to the Baseline Southern Quebec domain. For non-flood ETCs, the highest track density was located to the southeastern edge of the domain, whereas for flood ETCs, it was located in its center, over southern Quebec. Thus, flood ETCs generally tracked
 370 closer to the study area.

4.3.2 Clustering of cyclogenesis location among flood ETCs

Figure 9a shows the cyclogenesis locations and the clusters (colored dots) of all identified Baseline Southern Quebec storms (storms that crossed the Baseline Southern Quebec search domain during the 1991–2020 period) along with the search domain (dashed black). Dots circled in black show flood ETCs. Clustering the genesis locations of the storms yielded five
 375 different regions of origin: Western Canada & Pacific Ocean, Ontario & Hudson Bay, Quebec & Maritimes, Central US & Gulf of Mexico, and US east coast & the Caribbean. The number of storms in each cluster ranged from 984 (Ontario & Hudson Bay) to 1,425 (US east coast & the Caribbean).

Figure 9b shows the distribution of genesis locations for the two groups of storms (non-flood ETC vs flood ETC). The non-flood ETCs that were identified were mostly uniform in terms of their genesis location distribution, with percentages
 380 ranging from 15.9 % (Ontario & Hudson Bay) to 22.8 % (US east coast & the Caribbean) per cluster. In contrast, the distribution of genesis locations for flood ETCs was largely non-uniform. The increased proportion of flood ETCs originated from the Central US & Gulf of Mexico cluster (50.5 %), whereas the decreased proportions came from the Western Canada & Pacific Ocean (7.0 %), and Quebec & Maritimes (4.5 %) clusters. These differences in the distribution of genesis points compared with non-flood ETCs were statistically significant for these clusters (as determined by a chi-squared test with
 385 $\alpha = 0.05$ and standardized residuals above/below 2). However, the differences in the proportion of non-flood and flood ETCs originating from the US east coast & the Caribbean, as well as for the Ontario & Hudson Bay clusters were not significant.

While the differences between the two groups of storms were significant when all seasons were considered together, the results were somewhat different when the seasons were analysed independently. Figure S4 in the Supplement shows the distribution of storms genesis points per season. Differences between non-flood ETCs and flood ETCs for spring were
 390 significant, and the results reflected those for all seasons combined, with an even higher proportion of flood ETCs originating from the Central US & Gulf of Mexico (57.5 %). Differences between the two groups in summer and fall were also significant, but with increased proportions of storms originating from the US east coast & the Caribbean (40.0 % of summer flood ETCs and 76.9 % of fall flood ETCs). Finally, the differences were not significant in winter, which exhibited a more uniform distribution of origin among flood ETCs, ranging from 8.3 % (Quebec & Maritimes) to 33.3 % (Central US & Gulf of Mexico).

395 The typical lifetime of a storm in the Baseline Southern Quebec search domain also depended on its region of origin (cluster). Figure 9c shows the distribution of lifetimes per cluster along with their respective medians (white line). Using a K-S test with an adjusted α value of 0.005, it was determined that some clusters had unique lifetime distributions. Both Quebec

& Maritimes and Western Canada & Pacific Ocean clusters had distinctive distributions that differed from those of the other clusters. These two clusters had lower lifetime medians: 17 h for Quebec & Maritimes storms, and 28 h for Western Canada & Pacific Ocean storms. However, the test determined that the last three clusters in Fig. 9c (Ontario & Hudson Bay, Central US & Gulf of Mexico, US east coast & the Caribbean) came from the same distribution, with medians of 38, 39, and 39 h.

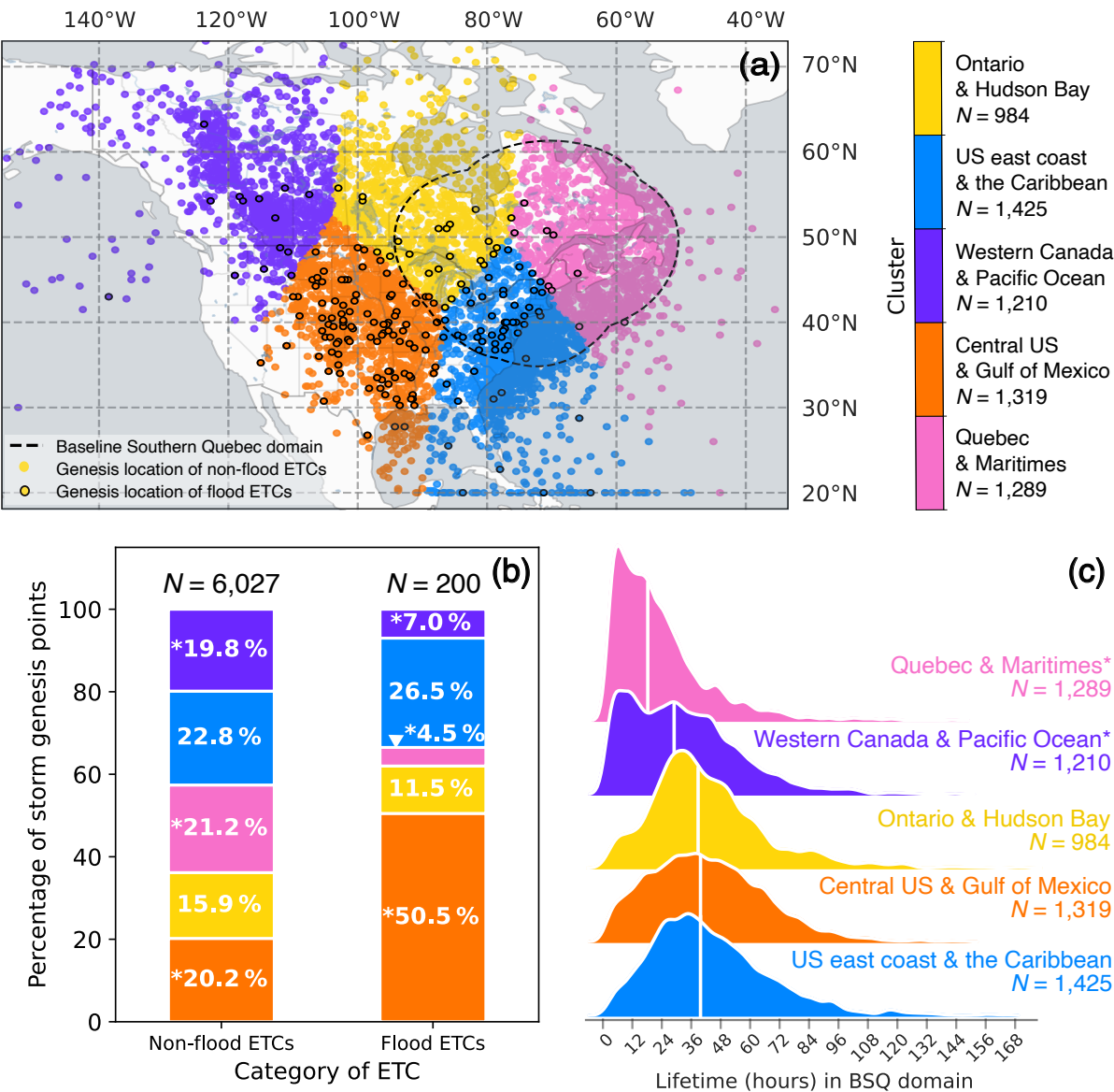


Figure 9. (a) Clusters on cyclogenesis of the 6,227 Baseline Southern Quebec ETCs included in the study spending at least one hour in the Baseline Southern Quebec search domain (dashed black) between January 1991 and December 2020. Dots with a black edge represent the genesis locations of flood ETCs. (b) Comparison of the distribution of genesis location percentages for the two categories of ETCs. Statistically significant differences are indicated by an asterisk (*) preceding the percentages shown. (c) Lifetime of storms in the Baseline Southern Quebec search domain by cluster of genesis location. The medians are shown by the white lines. The asterisk (*) denotes a statistically significant difference in the distribution.



A comparison of tracks of flood vs non-flood ETCs for clusters Central US & Gulf of Mexico and US east coast & the Caribbean was also done but not shown, available in the Supplement in Figs. S5 and S6 in the Supplement. This comparison was only done for these two clusters as the number of flood ETCs from the other three clusters were not sufficient ($N = 23$, 14, and 9).

4.3.3 Uneven contribution of flood ETCs to financial aid claims totals

Figure 10a shows the impact (number of financial aid claims associated) and cluster of all flood ETCs associated with at least 100 financial aid claims. The impact of flood ETCs varied greatly. The five most impactful flood ETCs were associated with 51.0 % of all financial aid claims over the 30-year study. Furthermore, the top 18 storms were associated with 75.7 % of all financial aid claims (see bars left of the 75 % *claims* in Fig. 10a), while the remaining 182 storms were associated with only 12.4 % of claims (as 11.9 % of claims were not associated with a flood ETC). In order of importance, the most impactful ETCs originated predominantly from the Central US & Gulf of Mexico, US east coast & the Caribbean, and Ontario & Hudson Bay clusters.

Figure 10b shows all flood ETCs (represented by colored dots corresponding to their cluster) with their associated number of financial aid claims and watersheds affected (number of local floods they are linked to). Similarly to the associated with financial aid claims, the association of flood ETCs with local flood events was also highly uneven, with only a few flood ETCs linked to widespread flooding affecting a large number of watersheds. Specifically, one flood ETC was associated with 48 local flood events, which was more than double the value of any other flood ETC (17 watersheds). The right half of Fig. 10b shows that the storms that lead to the most widespread flooding also originated from the same two clusters (Central US & Gulf of Mexico, and US east coast & the Caribbean). All storms associated with flooding in more than five watersheds and leading to more than 345 financial aid claims originated from these two clusters.

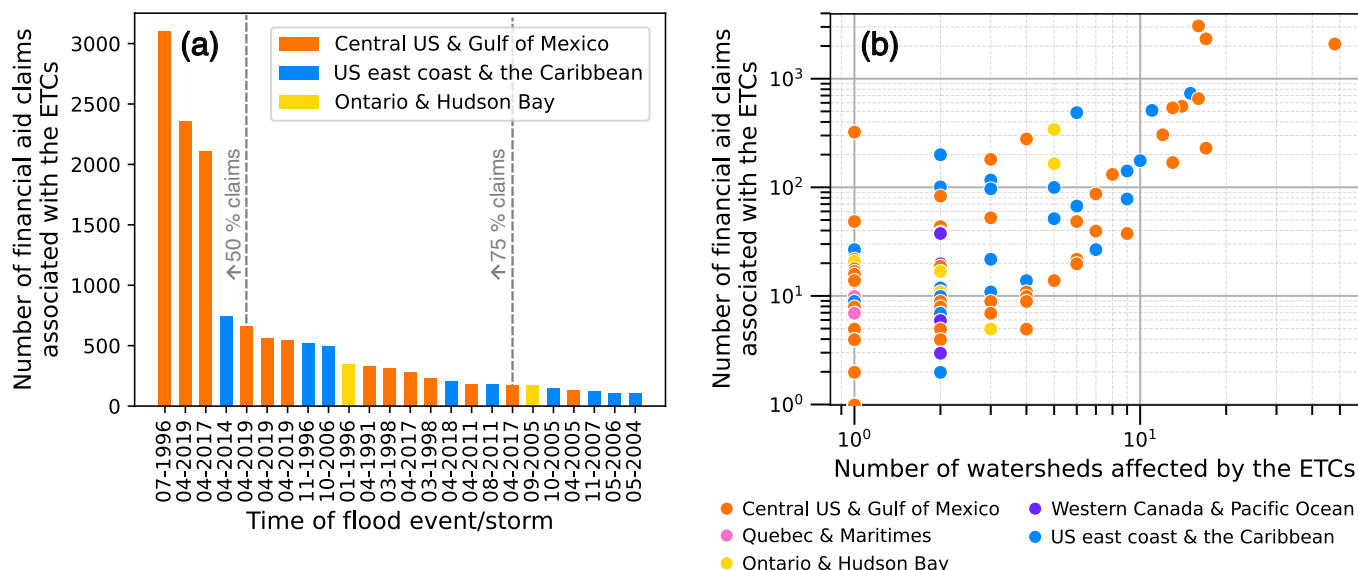


Figure 10. (a) Impact (number of financial aid claims associated) of flood ETCs, by importance, showing all flood ETCs with at least 100 financial aid claims associated ($N = 24$), their respective cluster of genesis location and the time of the flood event/storm. (b) Number of financial aid claims and watersheds affected (number of local floods) of flood ETCs and their cluster.

4.4 Impact of varying the days, radius, and threshold on identified flood ETCs

Figure 11 shows the percentage distribution of genesis location of flood ETCs, as identified through the three alternative scenarios and the control case, resulting from the application of the method to the 372 spring flood events. All three alternative scenarios showed very little differences in the distribution of genesis location among identified flood ETCs: the majority of flood ETCs still came from the Central US & Gulf of Mexico cluster, as can be seen in Fig. 11. The ‘days’ scenario only identified three fewer flood ETC. This suggests that the chosen search period was conservative, since few flood ETCs were identified during the first five days of the search period for spring flood events.

For the two other scenarios, the number of identified flood ETCs were lower than for the control case, with 29 (18.1 %) fewer flood ETCs identified in the radius scenario, and 26 (16.3 %) fewer in the threshold scenario. Differences between the control and radius scenarios signify that a small proportion (18.1 %) of flood ETCs identified in the control case track relatively far – between 750 and 1,000 km – from the center of the flooded watersheds. This could suggest that events of this kind (where flood ETCs track far from the affected watershed) are associated with fronts, given that these can be located far from an ETC’s center (Ahrens, 2009). Differences between the control and threshold scenarios indicate that the majority (83.7 %) of flood ETCs identified in the control case had a cumulated rainfall percentage already higher than 15 % in their respective search periods.

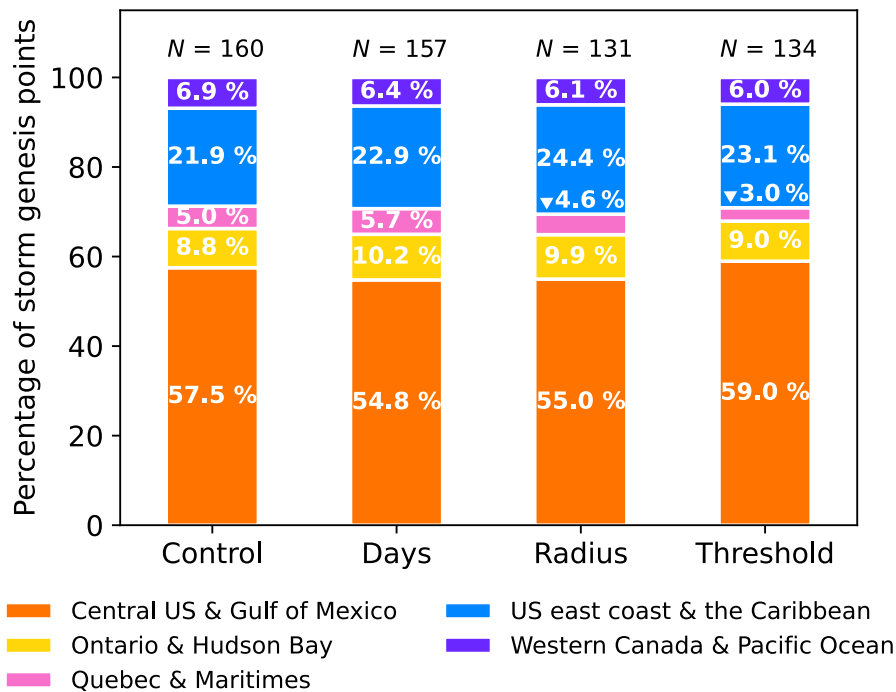


Figure 11. Distribution of genesis location of flood ETCs identified in MAM flood events for the control case compared to the three other scenarios

Differences in ETC contribution from the control to the ‘days’ scenarios were also minimal, with the median ETC contribution falling from 66.1 % to 63.9 % (see Fig. S7 in the Supplement). In contrast, but expectedly, reducing the search radius impacted ETC contribution, decreasing it from a median of 66.1 % in the control case to 44.6 % in the ‘radius’ case (see Fig. S7). The higher value for the control scenario is closer to the expected precipitation associated with ETCs (70–85 % depending on the season) for the province (Hawcroft et al., 2012). This suggests that the search radius was not greatly overestimated or underestimated by using a fixed value of 1,000 km.

5 Discussion

5.1 Seasonality of floods in Quebec and ETC contribution

Due to the large number of spring flood events, spring storms were more commonly identified as flood ETCs. This is why the seasonal distribution of storm genesis locations was examined (see Fig. S4 in the Supplement). Although a formal statistical attribution was not performed, the findings nonetheless reveal a marked increase in the proportion of U.S. east coast & the Caribbean storms among flood ETCs in summer and fall (with a stronger signal in fall) relative to non-flood ETCs. We suspect that with more observations, these differences would be statistically significant, as these storms were found among the



465 most impactful (with high numbers of financial aid claims associated) and linked to widespread (with high numbers of watersheds affected) flooding (see Fig. 10). The rarity of winter flood events makes findings related to this season difficult to interpret. Perhaps as a result, no differences in the provenance of ETCs have been found in winter when comparing flood and non-flood ETCs.

The number of local flood events with negligible ETC contribution was highest in summer, as can be seen in Fig. 6. In summer, high-intensity rainfall is often associated with smaller systems, such as mesoscale-convective systems and isolated thunderstorms, which can lead to flash flooding and could explain the occurrence of floods with negligible and mild ETC contribution (Alpizar et al., 2025; Bourgault et al., 2022; Prein et al., 2023). This is unique to the summer season, which is why its ETC contribution distribution was found to be the only one that was statistically different from those of the other seasons. Aside from the occurrence of sporadic extreme precipitation, summer is generally a season when flooding is less likely to occur, as river levels are low and can accommodate more rainfall. For flooding to occur in summer, significant amounts of rainfall are required, which explains the very median high ETC contribution (83.5 %) of summer floods.

The lowest average and median ETC contributions were recorded in spring, which is not surprising considering that other factors, such as snowmelt triggered by sudden temperature increases, can influence spring flooding. This effectively reduces the importance of ETC-associated precipitation needed to trigger flooding in that season. One of the study's main limitations is that it does not consider land surface processes that could increase the likelihood of flooding at certain times of year or in particular locations, such as soil moisture content and snowpack characteristics.

5.2 Characteristics of flood ETCs

The increased presence of storms from the Central US & Gulf of Mexico among flood ETCs cannot be entirely attributed to their longer lifetimes. A study of extratropical cyclone climatology across eastern Canada by Plante et al. (2015) showed that, although storms originating in the Gulf of Mexico and southern US (Central Rockies) did not frequently reach eastern Canada, they had high relative vorticity when they did, suggesting a high potential for damage. A climatology of cold season (October–March) ETCs leading to socially disruptive extreme weather events (EWEs) over central and eastern North America between 1979–2016 found that ETCs leading to EWEs typically formed in three areas: lee of the Rocky Mountains, over the south central US, and along the east coast of North America (Bentley et al., 2019). One reason these storms were found among ETCs leading to EWEs is their proximity to the Gulf of Mexico, which provides additional moisture and enhances the cyclones' diabatic processes.

The regions of enhanced genesis density for ETCs leading to EWEs are analogous to our clusters in the Central US & Gulf of Mexico, as well as in the US east coast & the Caribbean (see Figure 6 in Bentley et al., 2019). Although their study did not specifically examine ETCs leading to flooding, we can hypothesize that the characteristics of our flood ETCs would resemble those of storms leading to EWEs. Furthermore, the authors found that their ETCs leading to EWEs most frequently formed in November and March. March coincides with the onset of the spring freshet and the majority of floods in Quebec. Therefore, the increased formation of southern storms leading to EWEs during a time when Quebec is vulnerable to flooding



could explain why Central US & Gulf of Mexico storms were more prevalent among our flood ETCs, in addition to having long lifetimes and high intensities.

500 The aforementioned study (Bentley et al., 2019) also found that ETCs leading to EWEs rarely originated in the southern Canadian Rocky Mountains, as they remained weak when tracking over central and eastern Canada. This was also true for a climatology of extratropical cyclones over the 1979–2009 period in eastern Canada by Plante et al., (2015), which showed that cyclones that formed in high-latitude regions (Northern Rockies and Hudson Bay) were typically weak. Considering this and the shorter lifetimes of these storms in the Baseline Southern Quebec domain, a lower presence of Western Canada & Pacific
 505 Ocean storms among flood ETCs is to be expected, and coincides with our findings.

Forming within the Baseline Southern Quebec domain and typically tracking eastward, Quebec & Maritimes storms have shorter lifetimes (22 h less on average than Central US & Gulf of Mexico storms, see Fig. 9c). This makes them poor flood ETCs candidates, which explains their limited presence among flood ETCs.

Furthermore, Bentley et al. (2019) have identified that ETCs leading to EWEs forming along the east coast of North
 510 America occur more frequently in November and December. This is consistent with our findings, given that 76.9 % of our flood ETCs identified in the fall originated from the US east coast & the Caribbean cluster (see Fig. S4c). Cold season ETCs have also frequently been associated with flooding in eastern US (close to our study site) (Smith et al., 2010), and are typically strong and frequent in winter (Plante et al., 2015), which further explains their contribution to Quebec floods.

As can be seen in Fig. 9a, some storms form along 20°N latitude (the southernmost limit of the NAEC Catalogue), which
 515 indicates that the Catalogue includes tropical cyclones that have transitioned into the extratropics. The timing and tracks of the nine flood ETCs that generated (first time they were identified in the NAEC Catalogue) between latitudes of 20°N and 30°N (circled in black in Fig. 9a) were compared to NOAA historical hurricane tracks to identify potential matches (NOAA, 2025). Out of these nine flood ETCs, which all belonged to the US east coast & the Caribbean cluster, four had a match among hurricanes or tropical storms. Table 2 provides information on these extratropical transition storms, with some showing above
 520 seasonal average number of financial aid claims associated and/or watersheds affected (shown in bold). Figure S8 in the Supplement shows the tracks of these four extratropical transition flood ETCs (from the NAEC Catalogue) and their tropical cyclone matches (from the NOAA historical hurricanes).

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Table 2. Extratropical transition ETCs identified among flood ETCs with their associated number of financial aid claims and watersheds affected. Numbers in the *Storm id* column refer to their id in the NAEC Catalogue. Values in bold indicate an above average value for the given season. The levels indicated in parenthesis in the *Hurricane (H) / tropical storm (TS)* column refers to the maximum category reached by the storm during its lifetime.

Storm id	Timing	Season	Hurricane (H) / tropical storm (TS)	Number financial aid claims	Number watersheds affected
15013	August 2004	JJA	Frances (H4)	27	7
15595	August 2005	JJA	Katrina (H5)	27	1
18622	September 2010	SON	Nicole (TS)	68	6
19159	August 2011	JJA	Irene (H3)	178	10

Conclusions

This study quantified the contribution of extratropical cyclones to 498 fluvial flood events that occurred in 98 watersheds in southern Quebec between January 1991 and December 2020 and investigated whether flood ETCs (extratropical cyclones that contributed to flooding) had unique genesis locations, lifetimes, or tracks using multiple sources of data (reconstituted river outputs, ERA5 precipitation data, storm tracks, and government financial aid data).

Results on the quantification of ETC contribution to flooding in Quebec showed that the majority of flood events (72.7 %) had a high or very high ETC contribution (> 50 % of the rainfall during the search period was associated with ETCs), while only 2.6 % of events, mostly occurring in the summer, had a negligible ETC contribution (< 5 % of the rainfall during the search period was associated with ETCs). Although summer exhibited more extreme values, the lowest average ETC contribution (61.8 %) occurred in spring, and is hypothesised to be due to the additional contribution of snowmelt to flooding during this season.

This study revealed that flood ETCs differed from non-flood ETCs (ETCs that did not contribute to flooding, but entered the domain during the 30-year study period) in some respects. Flood ETCs had longer lifetimes, spending an average of 18.2 h more in the Baseline Southern Quebec domain than their non-flood counterparts. Lifetime in the domain and cluster of origin were closely related: both Quebec & Maritimes and Western Canada & Pacific Ocean storms spent less time in the domain (an average of 22 and 11 h less, respectively) than US east coast & the Caribbean and Central US & Gulf of Mexico storms, and rarely appeared among flood ETCs (constituting only 11.5 % of flood ETCs combined). Given that storms originating from Western Canada & Pacific Ocean are known to weaken as they reach eastern Canada, and that Quebec & Maritimes storms naturally spent less time in the domain because they originated inside it, this is not surprising.

The proportion of storms originating from the Central US & Gulf of Mexico increased from 20.2 % among non-flood ETCs to 50.5 % among flood ETCs. This is a clear indication that the region of origin (and thus the trajectory) was a key factor in determining whether storms were linked to flooding in the province. It is suspected that the longer lifetime, known intensity, and increased cyclogenesis of these storms in spring – when most flooding occurs – made them ideal flood ETC candidates.



555 Not all flood ETCs led to the same consequences: Central US & Gulf of Mexico and US east coast & the Caribbean
storms both typically resulted in more damage. Storms originating from these two regions were the only ones associated with
flood events involving more than 345 financial aid claims or affecting more than five watersheds. Furthermore, the percentage
of financial aid claims allocated to flood ETCs was highly variable: the 18 (9 %) most impactful storms were associated with
75.7 % of the financial aid claims. This indicates that, even though floods occurred almost annually during the 30-year study
560 period, a high ETC contribution (which occurred during most floods) did not necessarily lead to widespread flooding in the
province. Among US east coast & the Caribbean flood ETCs, 7.5 % turned out to be extratropical transition cyclones, raising
the question of the extent to which these types of storms were involved in flooding compared to extratropical cyclones.

Finally, the trajectories of flood ETCs also differed from those of non-flood ETCs. They were more concentrated over
the study area, typically entering Quebec from the south-west over land rather than tracking over the Atlantic Ocean along the
565 east coast of North America or over northern Quebec.

This study proposes a methodology for systematically identifying extratropical cyclones that contribute to flooding (flood
ETCs), using data characterizing damages (government financial aid claims) and reconstituted river outputs to determine the
timing and locations of floods. Studying these flood ETCs in detail can provide valuable insight into the atmospheric causes
of flooding and help understand the processes that led to past catastrophes. This could enable forecasters to monitor the types
570 of storms likely to lead to flooding more closely, such as those originating from the Central US & Gulf of Mexico in spring.
Ultimately, this could enhance flood preparedness.

As these datasets become more widely available, we hope this work will encourage further large-scale case studies of
flood events and motivate the adaptation of this methodology – for example, by considering soil moisture and snowmelt – to
improve our understanding of the impacts associated with different types of storms, including tropical cyclones, ETCs or
575 mesoscale convective systems.

Data availability

All data resulting from this analysis are available at: <https://doi.org/10.5683/SP3/98HDVT> (Gagnon et al., 2025). The raw data
on financial aid claims (Roy et al., 2024) is not available as it contains sensitive information.

Supplement link

580 The link to the supplement will be included by Copernicus.



Author contribution

Clarence Gagnon: Conceptualization, Methodology, Data curation, Formal analysis, Visualization, Writing – original draft.
Daniel F. Nadeau: Conceptualization, Methodology, Writing – review and editing, Supervision. Alejandro Di Luca:
585 Conceptualization, Methodology, Writing – review and editing, Supervision. François Anctil: Funding acquisition,
Conceptualization, Supervision. Marc-André Bourgault: Supervision. Benoit Brault: Data curation, Visualisation. Romane
Hamon: Data curation. Nicolas L. Roy: Data curation.

Competing interests

The authors declare that they have no conflict of interest.

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600 physical aspects (*réalisation de retours d'expérience sur les inondations ayant causé des sinistres sur le territoire québécois
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