

1 **Technical note: 12-km resolution capability for the global GEOS-Chem model of atmospheric**  
2 **composition**

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10

11 **Abstract**

12

13 We present a new 12-km nested resolution capability in the GEOS-Chem global model of  
14 atmospheric composition. This capability can be applied to simulations for any user-selected  
15 domain worldwide from March 2021 onward by accessing a new hourly cubed-sphere C720  
16 (approximately 12 km × 12 km resolution) global wind archive from the NASA GEOS-FP  
17 meteorological data assimilation system. We regridded the archive to support rectilinear GEOS-  
18 Chem Classic nested grid simulations worldwide at 0.125° × 0.15625° resolution and denote this  
19 as the 12-km GEOS-Chem capability. We evaluate this 12-km configuration of GEOS-Chem by  
20 comparison with the standard 25-km (0.25° × 0.3125° resolution) nested configuration in  
21 simulations of transport tracers, oxidant-aerosol chemistry, and inversions of satellite data using  
22 the Integrated Methane Inversion (IMI). The 12-km simulation features stronger vertical  
23 transport (up to 20% lower surface <sup>222</sup>Rn concentrations) because it better captures horizontal  
24 convergence both spatially and temporally. Aerosol lifetimes against deposition are shorter by a  
25 few percent. The 12-km oxidant-aerosol chemistry can better simulate urban observations of  
26 NO<sub>2</sub>, and shows stronger ozone urban titration together with slightly higher surface ozone  
27 background due to enhanced vertical transport. 12-km and 25-km inversions using the IMI show  
28 highly consistent results on the regional scale, but the 12-km inversion provides greater  
29 information and improved spatial detail to resolve emissions from different sectors.

30

31 **1. Introduction**

32 GEOS-Chem (<http://geos-chem.org>) is an open-source global 3-D model of atmospheric  
33 chemistry originally described by Bey et al. (2001) and used by hundreds of research groups  
34 around the world for a wide range of applications. The standard offline configuration of GEOS-  
35 Chem is driven by archived Goddard Earth Observing System (GEOS) meteorological data from  
36 the NASA Global Modeling and Assimilation Office (GMAO), with no feedback of chemistry on  
37 meteorology. GEOS-Chem can also operate online in dynamical coupling with Earth system  
38 models (Fritz et al., 2022; Hu et al., 2018; Lu et al., 2020) or the WRF regional model (Lin et al.,

39 2020). Here we introduce the capability to conduct offline GEOS-Chem simulations at  $0.125^\circ \times$   
40  $0.15625^\circ$  ( $\approx 12 \text{ km} \times 12 \text{ km}$ ) resolution by exploiting a new GEOS advection data archive (grid-  
41 scale winds) available from March 2021 onward. This enables low-cost, reproducible, high-  
42 resolution simulations of atmospheric chemistry and air quality anywhere in the world through  
43 the nested capability of GEOS-Chem. In what follows we refer to it as the 12-km capability in  
44 GEOS-Chem.

45 The open-access GEOS global meteorological datasets used to drive GEOS-Chem are produced  
46 at GMAO by data assimilation on the cubed-sphere grid of the underlying GEOS Earth system  
47 model (ESM) and then archived on a rectilinear latitude-longitude grid for public dissemination.  
48 The GEOS-Chem Support Team extracts from these archives the data needed to run GEOS-  
49 Chem and distributes them to users as an open dataset library through the Amazon Web Services  
50 (AWS) cloud (Zhuang et al., 2019; Martin et al., 2022). GEOS-Chem users may choose from  
51 three main datasets: the Modern-Era Retrospective Analysis for Research and Applications,  
52 version 2 (MERRA-2; 1980-present) at  $0.5^\circ \times 0.625^\circ$  resolution, the GEOS Forward Processing  
53 (GEOS-FP; 2014-present) at  $0.25^\circ \times 0.3125^\circ$  resolution, and the GEOS for Instrument Teams  
54 (GEOS-IT; 1998-present) at  $0.5^\circ \times 0.625^\circ$  or native cubed-sphere C180 resolution (180 grid cells  
55 per cubed-sphere side, approximately  $50 \text{ km} \times 50 \text{ km}$  resolution). MERRA-2 provides a long  
56 stable record with fixed physics and data assimilation algorithms, GEOS-IT is the next-  
57 generation stable record with updated physics and data assimilation, and GEOS-FP is the  
58 operational product generated in near real time using the latest validated GEOS system. GEOS-  
59 FP operates at native C720 resolution (approximately  $12 \text{ km} \times 12 \text{ km}$ ) but the data archive was  
60 previously made available only at  $0.25^\circ \times 0.3125^\circ$  resolution. Since March 2021, GMAO has been  
61 producing a native-resolution hourly C720 GEOS-FP advection archive specifically to serve  
62 GEOS-Chem needs, and this is what we use to enable the 12-km capability in GEOS-Chem.

63 GEOS-Chem operates in two modes, Classic (GC-Classic) and High-Performance (GCHP)  
64 (Martin et al., 2022). GC-Classic is designed for easy use on the rectilinear longitude-latitude  
65 grid and operates on single-node mode with shared-memory parallelization. GCHP enables high-  
66 resolution simulations using distributed-memory parallelization (MPI) on the cubed-sphere grid  
67 with efficient multi-node scalability extending to thousands of cores (Eastham et al., 2018). Both  
68 GC-Classic and GCHP can use the original rectilinear GEOS archive available at resolutions  
69 down to  $0.25^\circ \times 0.3125^\circ$ , with GCHP converting these inputs to its cubed-sphere grid at runtime.

70 They can also operate at coarser resolutions for computational economy using regridded GEOS  
71 data archived on AWS as part of the GEOS-Chem input dataset library. Global GC-Classic  
72 simulations thus typically use  $2^\circ \times 2.5^\circ$  or  $4^\circ \times 5^\circ$  GEOS data. GC-Classic includes a one-way  
73 nested capability to conduct regional simulations over limited domains with archived dynamic  
74 boundary conditions from a coarse-resolution simulation (Wang et al., 2004). GCHP includes a  
75 stretched-grid capability to enable higher-resolution two-way nesting over target regions (Bindle  
76 et al., 2021). Emission and surface type information are generally available at  $0.1^\circ \times 0.1^\circ$

77 resolution from the GEOS-Chem input dataset library, from which they are regridded on the fly  
78 within GEOS-Chem at the desired resolution using the HEMCO software tool (Lin et al., 2021).  
79 The one-way nested GC-Classic capability at GEOS-FP  $0.25^\circ \times 0.3125^\circ$  (referred to as 25-km)  
80 resolution is widely used for air quality applications (Kim et al., 2015; Zhang et al., 2015) and  
81 for inversions of greenhouse gas data to infer surface fluxes (Varon et al., 2022).

82 Here, we implement the GEOS-FP native C720 advection archive for use in GC-Classic nested-  
83 grid simulations. This 12-km capability can be applied over any user-selected domain  
84 worldwide, and for any period from March 2021 onward. It has been released in GEOS-Chem  
85 14.6.0 (<https://doi.org/10.5281/zenodo.15243271>). We describe the 12-km capability in Section 2  
86 and compare its transport to the standard 25-km nested GC-Classic configuration in Section 3.  
87 We then demonstrate its application to a full-chemistry simulation over the North China Plain  
88 (Section 4) and to the inversion of satellite observations using the Integrated Methane Inversion  
89 (IMI) framework (Section 5).

## 90 **2. GEOS-Chem simulation at 12-km resolution**

91 GC-Classic allows users to conduct nested simulations over any domain of interest using the  
92 GEOS meteorological archives as input and with rectilinear coordinates specified at runtime.  
93 This framework was first introduced by Wang et al. (2004) and updated to  $0.25^\circ \times 0.3125^\circ$   
94 resolution by Zhang et al. (2015) and Kim et al. (2015). The nested simulations are conducted as  
95 a regional model with boundary conditions provided by a separate global simulation at  $2^\circ \times 2.5^\circ$   
96 or  $4^\circ \times 5^\circ$  resolution that provides dynamic chemical fields updated every three hours.

97 Here we use the new global C720 GEOS-FP hourly advection archive, regridded to  $0.125^\circ \times$   
98  $0.15625^\circ$ , to produce the 12-km advection archive that enables GC-Classic nested-grid  
99 simulations at  $0.125^\circ \times 0.15625^\circ$  resolution. The 12-km advection archive includes horizontal  
100 wind vectors, surface pressure, and specific humidity on 72 vertical levels extending from the  
101 surface to 0.01 hPa (model top). Horizontal wind vectors are converted from cubed-sphere C720  
102 mass flux data for use in the GC-Classic transport scheme (Lin and Rood, 1996). Surface  
103 pressure is needed to infer the vertical air mass fluxes from the horizontal air mass fluxes by  
104 mass conservation. Specific humidity is needed to convert wet air fluxes and pressure in the  
105 GEOS data to dry air fluxes and pressures used in GEOS-Chem. The regridding from cubed-  
106 sphere C720 to the rectilinear  $0.125^\circ \times 0.15625^\circ$  grid yields an approximately 12 km  $\times$  12 km  
107 horizontal scale over midlatitudes, slightly coarser toward the equator and finer toward the poles.

108 Table 1 summarizes the GEOS-FP input data for the 12-km configuration as compared to the 25-  
109 km configuration. The idea behind the advection archive is that higher resolution is most needed  
110 for the winds to better represent eddy flows and to leverage the  $0.1^\circ \times 0.1^\circ$  resolution of the  
111 emission data. Limiting the number of variables in the advection archive enables global data  
112 storage at C720. The 12-km advection archive has hourly temporal resolution, compared to 3-

113 hourly resolution for the 3-D variables in the 25-km archive. Other non-advection meteorological  
 114 inputs including convective mass fluxes and vertical mixing depths are from the standard GEOS-  
 115 FP archive at  $0.25^\circ \times 0.3125^\circ$  resolution and are dynamically regridded to  $0.125^\circ \times 0.15625^\circ$  at  
 116 runtime. Most emission datasets are available on a  $0.1^\circ \times 0.1^\circ$  grid and are regridded at runtime to  
 117 the simulation resolution using HEMCO (Lin et al., 2021). The 12-km archive is available  
 118 globally. We also maintain separate 12-km advection archives for five continental regions  
 119 (Africa, Asia, Europe, North America, and South America) to reduce data size and thereby speed  
 120 up data access and processing over the user-selected nested domains  
 121 (<https://registry.opendata.aws/geoschem-nested-input-data>; latest access: 21 October 2025).

122 All simulations presented here were carried out on the Harvard Cannon v2.0 supercomputing  
 123 cluster using compute nodes equipped with dual-socket Intel Xeon Platinum 8480CL CPUs (56  
 124 cores each, base frequency  $\sim 2.9$  GHz). Each simulation used 48 physical CPU cores and ran on a  
 125 single node. We find that the 12-km simulation wall-time is about 7 times that of a 25-km  
 126 simulation over the same domain. This is driven by the 4 times increase in grid cells and the 2  
 127 times reduction in timestep. The 1-month full-chemistry simulation presented in Section 4 took  
 128 26 wall-time hours to complete at 12-km resolution.

129

130 **Table 1.** GEOS-FP meteorological data archives available for driving GEOS-Chem at  $0.125^\circ \times$   
 131  $0.15625^\circ$  (12-km) and  $0.25^\circ \times 0.3125^\circ$  (25-km) resolutions<sup>a</sup>.

<b>GEOS-Chem simulation</b>	<b>12-km</b>	<b>25-km</b>
Archive period	March 2021-present	January 2014-present
Horizontal resolution (advection)	$0.125^\circ \times 0.15625^\circ$	$0.25^\circ \times 0.3125^\circ$
Temporal resolution (advection)	Hourly	3-hourly
Horizontal resolution (other)	$0.25^\circ \times 0.3125^\circ$ regridded to $0.125^\circ \times 0.15625^\circ$	$0.25^\circ \times 0.3125^\circ$
Temporal resolution (other)	3-hourly or hourly <sup>b</sup>	3-hourly or hourly
Timestep <sup>c</sup>	Transport 150 s Non-transport 300 s	Transport 300 s Non-transport 600 s

132 a. For GC-Classic simulations at  $0.125^\circ \times 0.15625^\circ$  and  $0.25^\circ \times 0.3125^\circ$  resolutions with 72  
 133 hybrid sigma-pressure vertical levels extending up to 0.01 hPa. GEOS-FP meteorological  
 134 data used as GEOS-Chem input include advection variables (horizontal wind vectors,  
 135 surface pressure, specific humidity) and other variables driving convective transport,  
 136 planetary boundary layer (PBL) mixing, emissions, radiation, chemistry, and deposition.  
 137 The full list of variables is at: <https://wiki.seas.harvard.edu/geos->

138 [chem/index.php/List\\_of\\_GEOS-FP\\_met\\_fields](chem/index.php/List_of_GEOS-FP_met_fields). The advection variables are grouped in a  
139 separate advection archive for 12-km applications. The 12-km and 25-km archives are  
140 available globally, and also for individual continents to speed up data extraction.  
141 b. Hourly for two-dimensional fields such as surface properties and PBL depth; 3-hourly for  
142 three-dimensional fields such as temperature and convective mass flux.  
143 c. Recommended timesteps for operator splitting (Philip et al., 2016), can be adjusted by  
144 user. Non-transport operators include emissions, chemistry, and deposition.

145

### 146 3. Transport tracer simulations

147 We use the TransportTracers simulation of GEOS-Chem (Zhang et al., 2021) to evaluate model  
148 transport and scavenging processes. This simulation includes 21 generic species (tracers) to test  
149 the different components of GEOS-Chem transport. We focus here on the radionuclide tracers  
150 radon-222 ( $^{222}\text{Rn}$ ), lead-210 ( $^{210}\text{Pb}$ ), and beryllium-7 ( $^7\text{Be}$ ), which are routinely used to  
151 benchmark transport and wet deposition in GEOS-Chem (Liu et al., 2001; Yu et al., 2018). These  
152 tracers are particularly useful for evaluating model vertical transport, whereas grid-scale vertical  
153 winds are calculated from horizontal mass flux convergence and is not directly available in the  
154 model. The nested-grid simulations are conducted over eastern China (100–125°E, 17–45°N,  
155 domain shown in Figure 1) at both  $0.125^\circ \times 0.15625^\circ$  and  $0.25^\circ \times 0.3125^\circ$  resolutions for February  
156 and June 2022. Initial conditions are generated from spin-up simulations at the same resolutions,  
157 starting from October 2021 (corresponding to spin-up periods of 4 months for February and 8  
158 months for June). The boundary conditions are updated every three hours from a global  
159 simulation at  $2^\circ \times 2.5^\circ$  resolution.

160 Figure 1 shows surface  $^{222}\text{Rn}$  concentrations from the 12-km simulation and relative differences  
161 with the 25-km simulation.  $^{222}\text{Rn}$  in the simulation has a spatially uniform soil source and is  
162 removed by radioactive decay with a half-life of 3.8 days, making it a sensitive tracer for vertical  
163 transport in the troposphere (Liu et al., 2001; Yu et al., 2018). Surface  $^{222}\text{Rn}$  concentrations are  
164 lower by 0–20% at 12-km than at 25-km resolution. Differences are most pronounced over  
165 complex terrain such as Sichuan and Taiwan. The total  $^{222}\text{Rn}$  burden in the two simulations is the  
166 same, and the difference is in the vertical distribution (Figures 2 and 3). The 12-km simulation  
167 shows lower  $^{222}\text{Rn}$  concentrations in the lower troposphere and higher concentrations in the  
168 middle-to-upper troposphere, indicating stronger vertical transport. The lower  $^{222}\text{Rn}$   
169 concentrations in the tropical stratosphere in June, combined with the lower  $^7\text{Be}$  concentrations  
170 below the tropopause, may reflect greater restriction of transport across the tropopause at higher  
171 resolution (Stanevich et al., 2020) but would require further investigation because the absolute  
172 changes are very small.

173 Transport processes in GEOS-Chem include grid-resolved advection (winds), sub-grid  
174 parameterized convection (convective mass fluxes), and PBL-mixing (Lin and McElroy, 2010).  
175 Differences between the 12-km and 25-km simulations could reflect stronger vertical winds at  
176 12-km due to better resolved horizontal convergence (Yu et al., 2018) and better resolution of  
177 concentration gradients leading to stronger tracer convergence for the same winds. To separate  
178 these two effects, we conducted a 12-km sensitivity simulation driven by the 3-hourly 25-km  
179 advection archive regrided to 12-km resolution during runtime (right columns of Figures 1-3).  
180 We find that both effects contribute, with 12-km winds accounting for about 77% of the surface  
181  $^{222}\text{Rn}$  difference in February and 32% in June, while better resolved concentration gradients  
182 account for the remainder.

183  $^{210}\text{Pb}$  (half-life of 22.3 years) is produced by the decay of  $^{222}\text{Rn}$ , and  $^7\text{Be}$  (half-life of 53.3 days)  
184 is generated by cosmic-ray interactions with atmospheric oxygen and nitrogen at high altitudes.  
185 Both radionuclides rapidly attach to aerosol particles and are subsequently transported and  
186 removed by wet and dry deposition. This makes  $^{210}\text{Pb}$  useful to evaluate aerosol transport and  
187 removal processes, with  $^7\text{Be}$  providing complementary information on stratosphere–troposphere  
188 exchange and tropospheric subsidence (Liu et al., 2001). Comparisons of 12-km and 25-km  
189 simulations for  $^{210}\text{Pb}$  show similar differences as for  $^{222}\text{Rn}$  (Figures 2 and 3) but weaker in  
190 magnitude because the  $^{210}\text{Pb}$  source is more diffuse. Lifetimes against deposition differ by less  
191 than 2% between the 12-km and 25-km simulations. Comparisons for  $^7\text{Be}$  also show weaker  
192 differences than for  $^{222}\text{Rn}$  because vertical transport of  $^7\text{Be}$  in the troposphere is mainly by large-  
193 scale subsidence, which is less sensitive to eddy motions. The  $^7\text{Be}$  lifetime against deposition is  
194 3% higher at 25-km resolution, explaining the slightly higher overall tropospheric  
195 concentrations.

196

#### 197 **4. Full-chemistry simulations**

198 We perform 1-month full-chemistry simulations for February and June of 2022 over the North  
199 China Plain (NCP; domain shown in Figure 4) at 12-km and 25-km horizontal resolutions for  
200 comparison. The full-chemistry configuration of GEOS-Chem includes detailed ozone–NO<sub>x</sub>–  
201 VOCs–aerosol–halogen tropospheric and stratospheric chemistry (Wang et al., 2021). Boundary  
202 conditions are provided by a global  $2^\circ \times 2.5^\circ$  GEOS-Chem simulation and updated every three  
203 hours, and spin-up simulations are conducted from October 2021 to generate initial conditions.  
204 Monthly anthropogenic emissions are from the MIXv2 Asian emission inventory at  $0.1^\circ \times 0.1^\circ$   
205 resolution for 2017, scaled to 2022 using province-level emission data from the MEIC v1.4  
206 inventory (Zheng et al., 2018) and mapped to the 12-km and 25-km grids using HEMCO. We  
207 compare the model simulations of surface NO<sub>2</sub>, ozone, and fine particulate matter (PM<sub>2.5</sub>)  
208 concentrations to hourly observations from 361 sites operated by the China National  
209 Environmental Monitoring Centre (CNEMC; <http://www.cnemc.cn>, last assess: 23 May 2023).

210 The sites are mainly urban. We remove anomalous observations at each site following the quality  
211 control protocols described in Lu et al. (2018). We sample model outputs at the observation sites  
212 for comparisons.

213 Figure 4 shows the afternoon (13-18 local time) surface  $\text{NO}_2$  concentrations over the NCP in  
214 June 2022 and the effect of model resolution. The finer structure at 12-km resolution is evident  
215 and largely reflects the ability to exploit the higher resolution of emissions. Comparison to  
216 CNEMC observations (Figure 5a) shows a low bias in the model, likely reflecting the near-  
217 source locations of the sites, but the bias is reduced at 12-km resolution. The 12-km simulation  
218 does not improve the correlation with observations for individual CNEMC sites, which could  
219 reflect errors in model transport or in the spatial distribution of emissions.

220 Figure 6 compares simulated maximum daily 8-h average (MDA8) surface ozone concentrations  
221 at 12-km and 25-km resolutions in June 2022. The differences between the two resolutions are  
222 generally smaller than 5 ppb, and there is no significant difference in the fit to observations  
223 (Figure 5b). Background ozone concentrations increase by up to 3 ppb over the northern NCP  
224 region, likely driven by increased vertical transport of ozone from aloft as seen in the simulated  
225  $^7\text{Be}$  concentrations in Figure 1. Increasing the model resolution to 12-km decreases surface  
226 ozone concentrations by about 3 ppb in Beijing and 6 ppb in Tianjin city core areas, as expected  
227 from higher  $\text{NO}_x$  concentrations driving stronger ozone titration, whereas suburban areas exhibit  
228 ozone increases of no more than 2 ppb. These effects are too small to be effectively evaluated in  
229 the comparison to observations (Figure 5b).

230 Figure 7 shows the daily-averaged total  $\text{PM}_{2.5}$  mass concentrations in February 2022, where  
231  $\text{PM}_{2.5}$  is computed in the model as the sum of fine aerosol components (Zhai et al., 2021). We  
232 focus on February here because  $\text{PM}_{2.5}$  concentrations are higher in winter than in summer (Zhai  
233 et al., 2019).  $\text{PM}_{2.5}$  concentrations are in general slightly lower at 12-km resolution because of  
234 the enhanced vertical transport. However, higher concentrations are found in a few urban  
235 hotspots due to primary organic aerosol emissions from combustion that are better resolved  
236 spatially. Again, differences are too small to be arbitrated by the observations (Figure 5c).

237

## 238 **5. Application to the Integrated Methane Inversion (IMI)**

239 The IMI applies the nested GEOS-Chem as forward model in regional analytical inversions of  
240 TROPOMI satellite observations of methane columns to optimize methane emissions (Estrada et  
241 al., 2025; Varon et al., 2022). It uses Bayesian inference to minimize the mismatch between the  
242 TROPOMI methane column observations and the corresponding concentrations simulated by  
243 GEOS-Chem, regularized by prior information on emissions. GEOS-Chem is applied to  
244 construct the Jacobian matrix describing the sensitivity of TROPOMI observations to emissions  
245 as simulated by the model transport. This is then used together with the TROPOMI observations,  
246 prior emission estimates, and prior and observational error statistics to derive optimized

247 (posterior) emissions. The TROPOMI satellite observations are at  $5.5 \times 7 \text{ km}^2$  pixel resolution so  
248 there is potential benefit for conducting the inversion at 12-km resolution using our new GEOS-  
249 Chem capability. The 12-km IMI configuration was previously applied in Wang et al. (2026) to  
250 quantify methane emissions across 12 U.S. urban areas. Here we examine how inversion results  
251 vary between 12-km and 25-km resolution, focusing on the Houston urban area in eastern Texas  
252 as an example.

253 The IMI inversion procedure is described by Estrada et al. (2025) and Hancock et al. (2025),  
254 including the design of state vector, error estimates, and optimization strategy. Here we optimize  
255 annual methane emissions in 2022 over a  $3^\circ \times 4^\circ$  (latitude  $\times$  longitude) domain encompassing  
256 Houston and its surrounding area as shown in Figure 8. The state vector to be optimized consists  
257 of emissions in each land-containing grid cell of the  $3^\circ \times 4^\circ$  domain (496 elements at 12-km  
258 resolution and 164 elements at 25-km resolution) and boundary conditions on each lateral edge  
259 (4 elements). Our prior estimates of anthropogenic emissions are from the U.S. Environmental  
260 Protection Agency Greenhouse Gas Inventory (GHGI) at  $0.1^\circ \times 0.1^\circ$  resolution for 2020  
261 (Maasakkers et al., 2023). Natural emissions follow the default configuration as described in  
262 Estrada et al. (2025). We assume a lognormal prior error probability density function (PDF) for  
263 emissions with a geometric error standard deviation of 2.0 (Bruno et al., 2025), and a normal  
264 error PDF for boundary conditions with a 10 ppb error standard deviation. The IMI analytical  
265 inversion returns posterior emissions and a posterior error covariance matrix from which the  
266 averaging kernel matrix can be derived as a measure of information content from the  
267 observations (Brasseur and Jacob, 2017).

268 The prior estimate of total emissions over the  $3^\circ \times 4^\circ$  inversion domain is  $790 \text{ Gg a}^{-1}$  (Figure 8).  
269 Total posterior emissions are 50% higher than the prior estimate and agree closely between the  
270 12-km inversion ( $1260 \text{ Gg a}^{-1}$ ) and the 25-km inversion ( $1170 \text{ Gg a}^{-1}$ ). The 12-km inversion  
271 results, when averaged over the 25-km grid, show a high degree of consistency with the 25-km  
272 inversion results over the inversion domain (spatial correlation coefficient = 0.92). The trace of  
273 the averaging kernel matrix defines the Degrees of Information for Signal (DOFS) indicating the  
274 number of pieces of information that can be obtained from the observations through the inversion  
275 independently from the prior estimate. The DOFS for the  $3^\circ \times 4^\circ$  domain is 24.4 for the 12-km  
276 inversion, higher than the 16.8 for the 25-km inversion. Conducting the inversion at higher  
277 resolution allows for more information on emissions to be extracted from the observations.

278 Another advantage of the higher-resolution inversion is better separation of the sectors  
279 contributing to methane emissions. Sectoral information in the inversion is obtained by  
280 attributing the posterior/prior emission ratios for each grid cell to the different sectors  
281 contributing emissions to that grid cell in the prior estimate (Wecht et al., 2014). Higher spatial  
282 resolution in the inversion reduces spatial overlap between sectors.

## 283 6. Conclusion

284 We have implemented a 12-km resolution nested capability in the GEOS-Chem global model of  
285 atmospheric chemistry by taking advantage of a new hourly GEOS advection archive available  
286 globally from March 2021 onward and freely distributed to GEOS-Chem users through the AWS  
287 cloud. We compared the 12-km simulation to the standard 25-km nested simulation in GEOS-  
288 Chem for transport tracers ( $^{222}\text{Rn}$ ,  $^{210}\text{Pb}$ ,  $^7\text{Be}$ ), ozone-aerosol chemistry, and inversion of methane  
289 satellite data where GEOS-Chem provides the forward model. The 12-km simulation has  
290 stronger vertical transport (up to 20% decrease in  $^{222}\text{Rn}$  surface concentrations) because of  
291 improved representation of horizontal convergence. It shows finer spatial structure in surface  
292 pollutants ( $\text{NO}_2$ ), with improved capability to reproduce urban observations. It better represents  
293 surface ozone titration in urban air and slightly increases the surface ozone background by  
294 increasing vertical transport. Application to the Integrated Methane Inversion (IMI) shows  
295 regional-scale results consistent with a 25-km inversion but higher information content and  
296 greater spatial detail. Overall, this 12-km capability further extends the application of GEOS-  
297 Chem in regional atmospheric chemistry and in the use of high-resolution satellite observations  
298 to quantify emissions of air pollutants and greenhouse gases at fine spatial scales. Higher model  
299 resolution improves the representation of finer concentration gradients and inverse model  
300 sensitivity through improved transport and chemistry.

301

302 **Code and data availability.**

303 The source code of GEOS-Chem 14.6.0 is publicly available at  
304 <https://doi.org/10.5281/zenodo.15243271>. The GEOS-FP meteorological fields used to drive  
305 GEOS-Chem are available on the Amazon Web Services (AWS) cloud  
306 (<https://s3.amazonaws.com/gcgrid/index.html>; last access: 21 October 2025).

307

308 **Author contributions.**

309 XW and DJJ conceptualized the research. XW performed the analyses and data visualization.  
310 XW and MPS developed the model code. MPS, YZ and RVM contributed to the data collection.  
311 XW and DJJ wrote the manuscript with input from all authors.

312

313 **Competing interests.**

314 The authors declare no competing interests.

315

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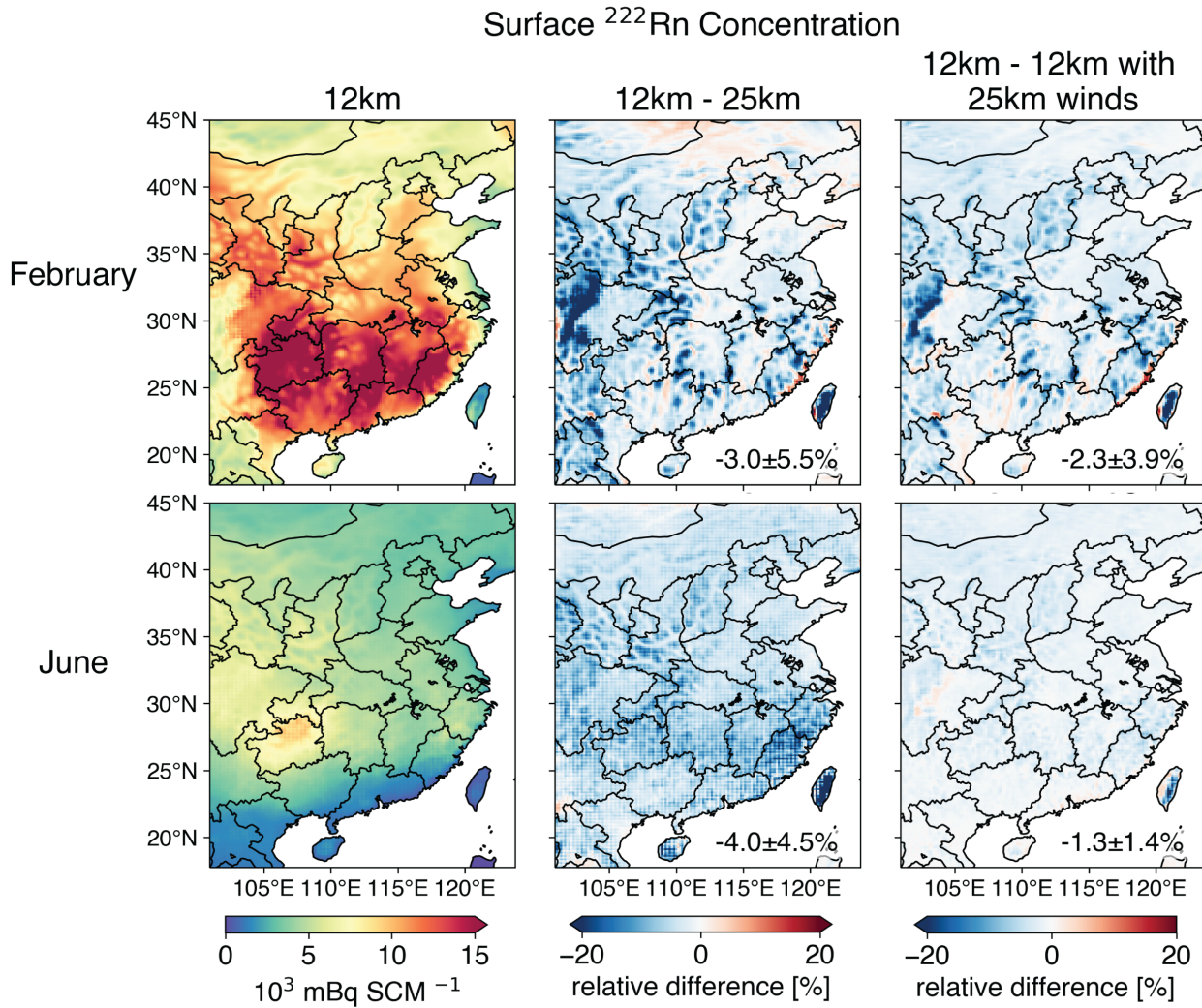
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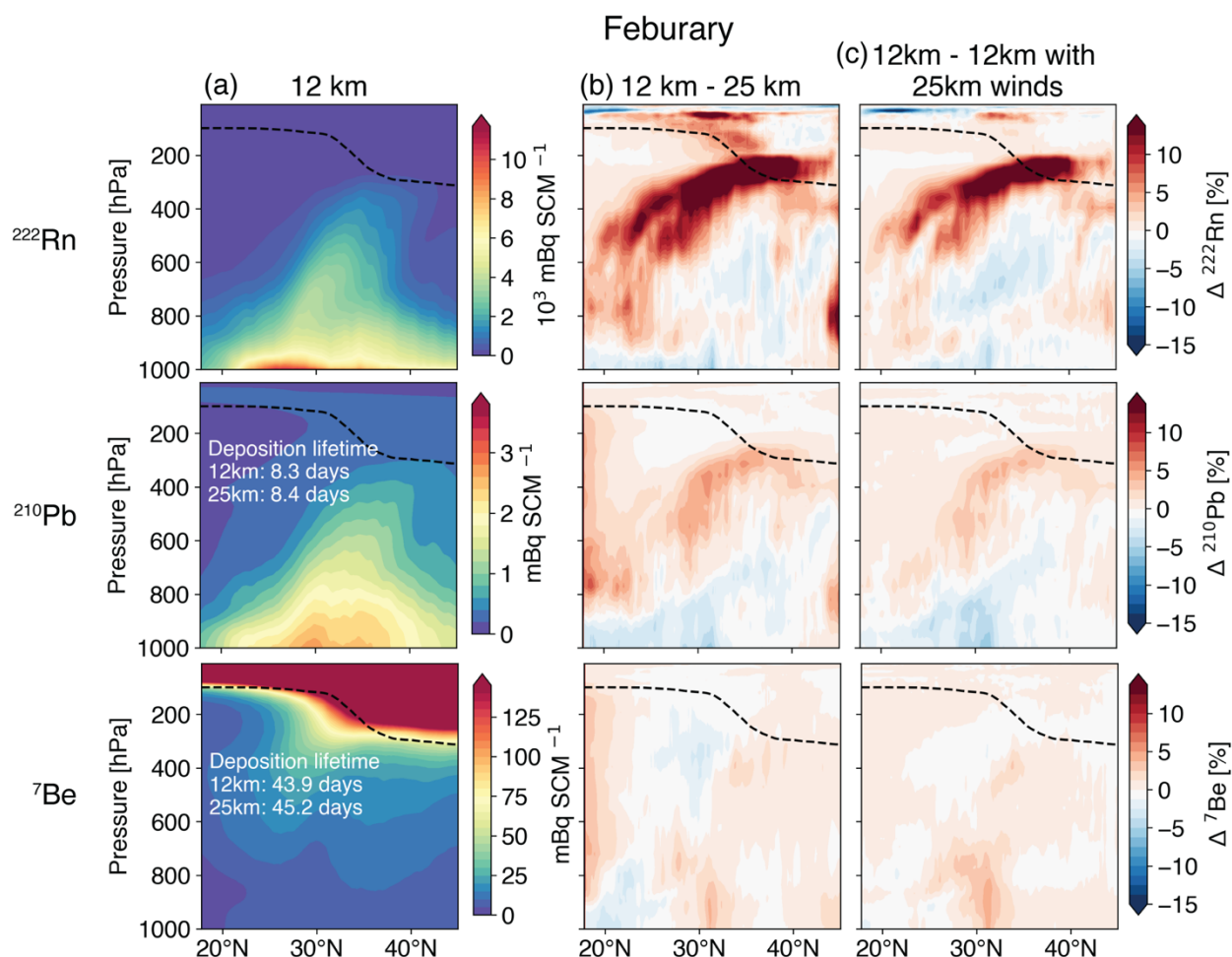


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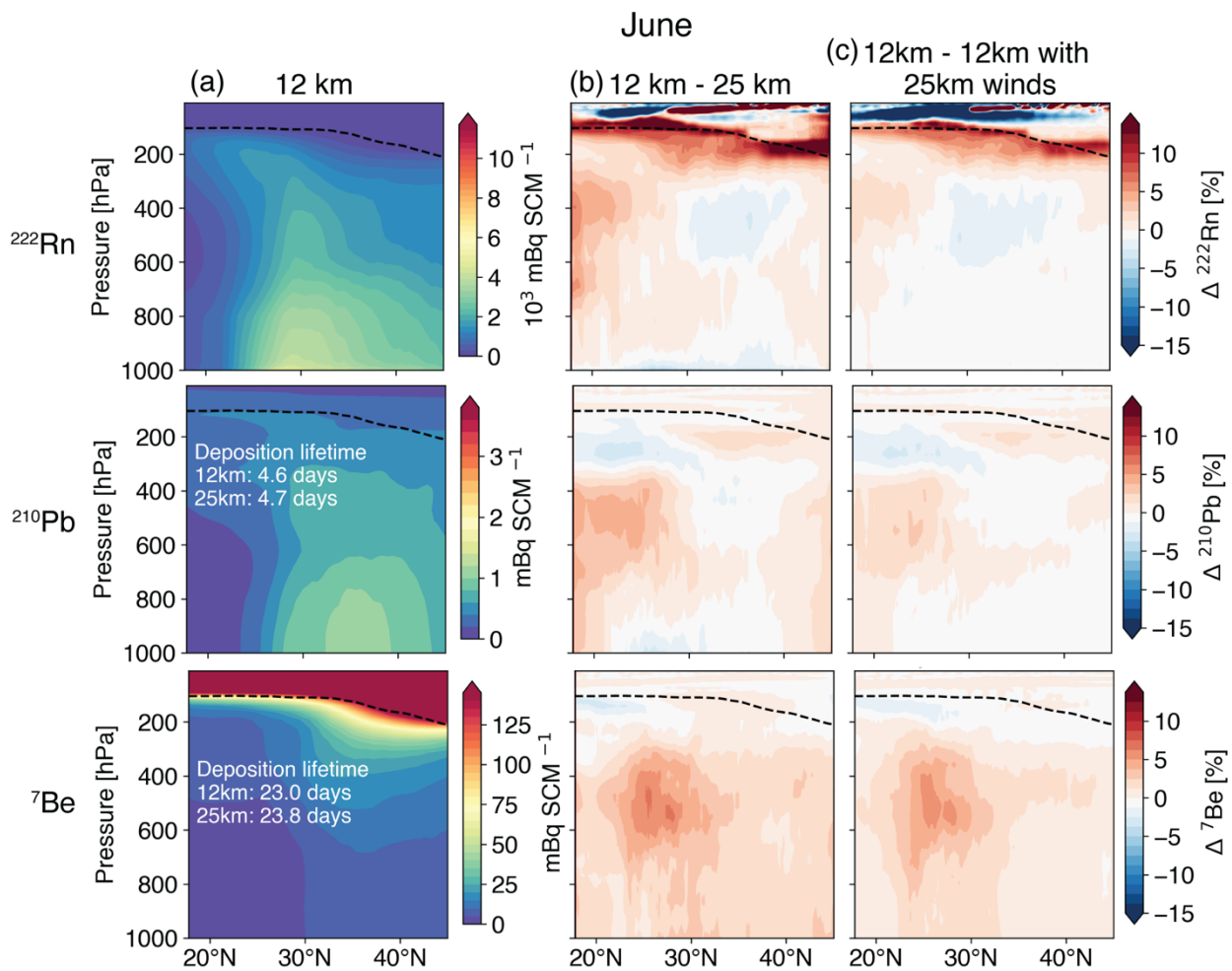
474 **Figure 1.** Monthly mean surface  $^{222}\text{Rn}$  mixing ratios ( $\text{mBq SCM}^{-1}$ ) simulated over eastern China  
 475 in February and June 2022. Columns from left to right show the concentrations simulated by  
 476 GEOS-Chem at 12-km resolution, the relative differences with a simulation at 25-km resolution,  
 477 and the relative differences between two 12-km simulations driven by the 12-km and 25-km  
 478 GEOS-FP advection archives, respectively. SCM is a standard cubic meter of air at 0 °C (273.15  
 479 K) and 1 atm (1013.25 hPa). The mean relative difference and standard deviation are shown  
 480 inset.

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484 **Figure 2.** Zonally averaged latitude-pressure monthly mean mixing ratios of  $^{222}\text{Rn}$ ,  $^{210}\text{Pb}$ , and  
 485  $^{7}\text{Be}$  simulated by GEOS-Chem for February 2022 over the eastern China domain of Figure 1 at  
 486 12-km resolution (left column), the relative differences with a 25-km simulation (middle column)  
 487 and relative differences with a 12-km simulation driven by the 25-km advection archive (right  
 488 column) Lifetimes of tropospheric  $^{210}\text{Pb}$  and  $^{7}\text{Be}$  against deposition are inset. The dashed black  
 489 lines indicate the tropopause.



491 **Figure 3.** Same as Figure 2 but for June 2022.

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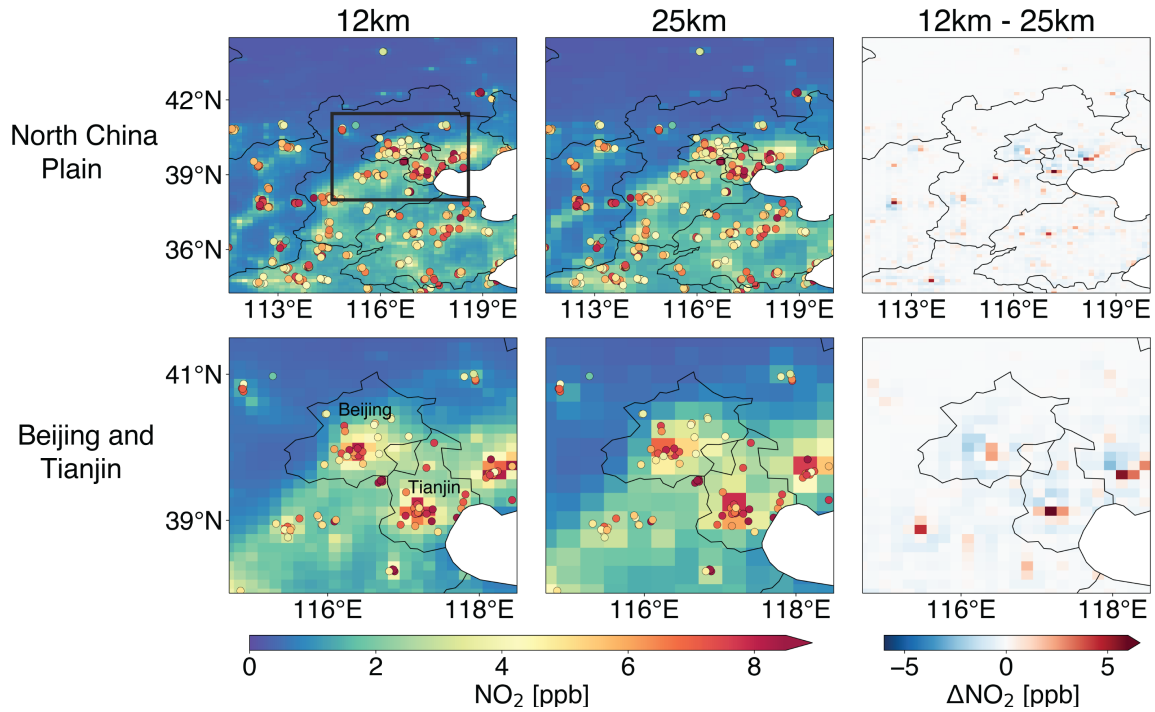
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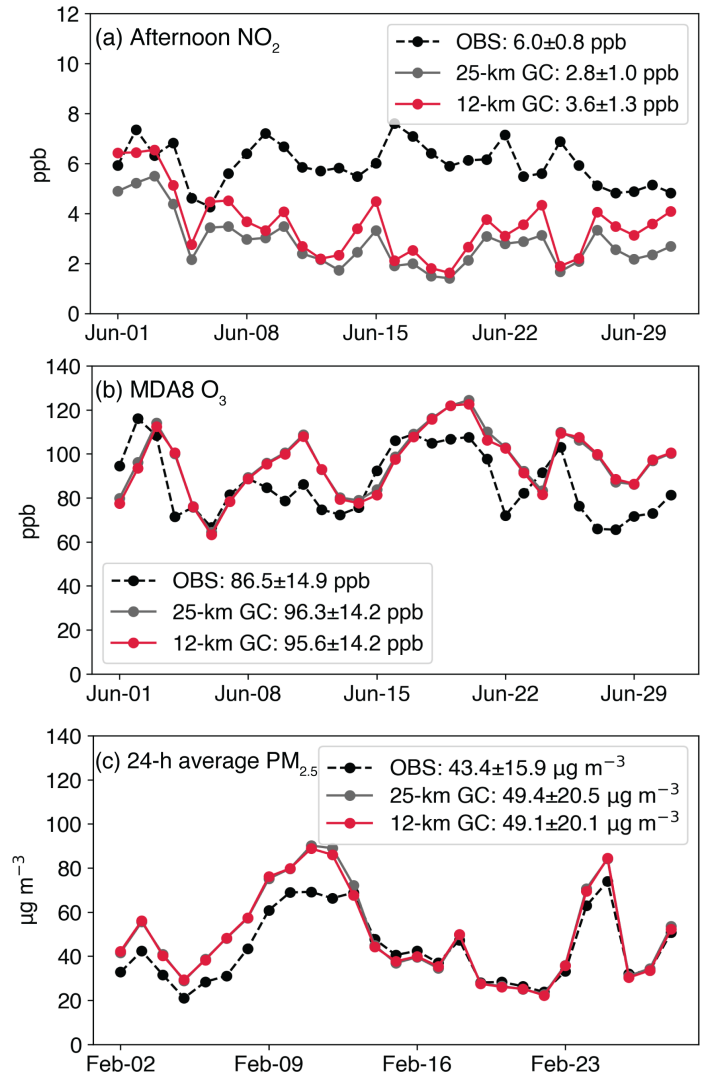
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498 **Figure 4.** Afternoon (13–18 local time) monthly mean surface NO<sub>2</sub> concentrations in June 2022  
499 simulated by GEOS-Chem at 12-km and 25-km resolution, for the North China Plain (NCP) and  
500 for Beijing and Tianjin (box in top left panel). Circles show CNEMC network observations.

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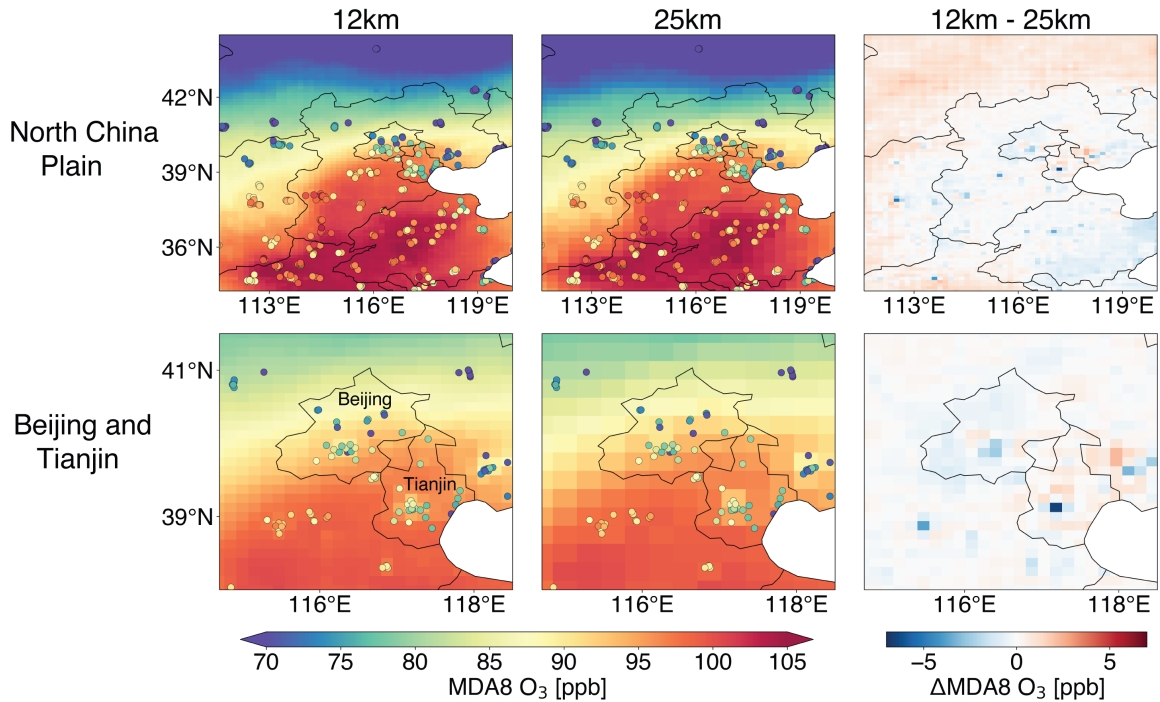
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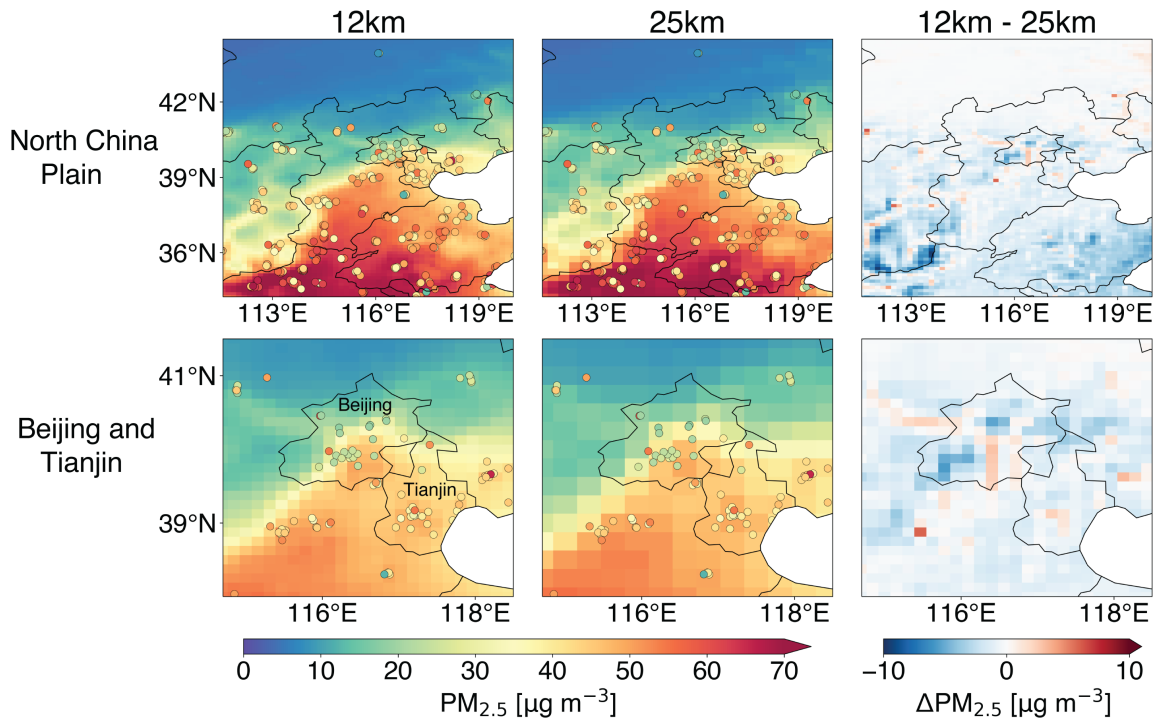
504 **Figure 5.** Daily time series of surface concentrations in the North China Plain (NCP): (a)  
 505 afternoon (13–18 local time)  $\text{NO}_2$  in June 2022, (b) maximum daily 8-h average (MDA8) ozone  
 506 in June 2022, and (c) 24-h average  $\text{PM}_{2.5}$  in February 2022. Observations at China National  
 507 Environmental Monitoring Center (CNEMC) sites, shown as circles in Figure 4 and averaged  
 508 over the NCP domain, are compared to GEOS-Chem simulations at 12- and 25-km resolution for  
 509 the same sites. Monthly mean values and standard deviations across all domain sites are shown  
 510 inset.

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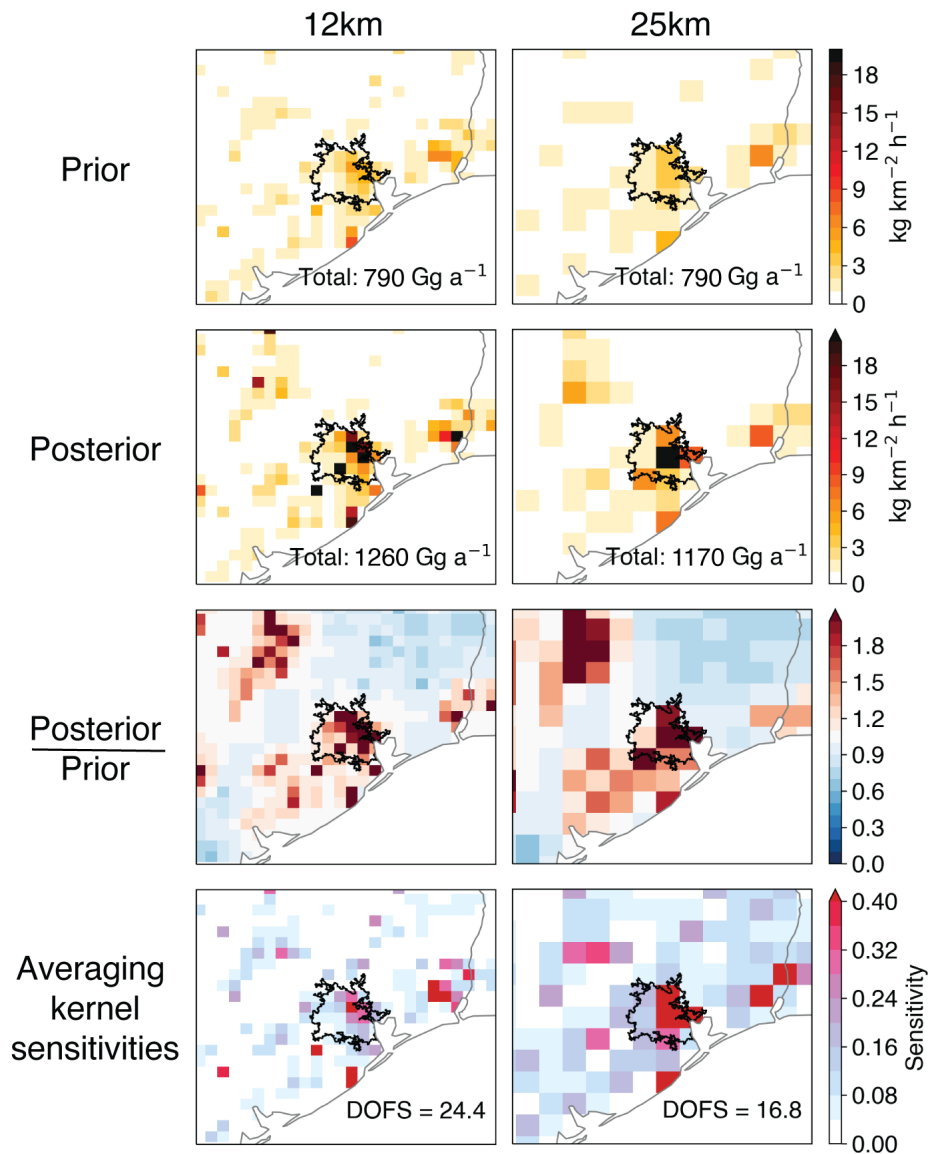
513 **Figure 6.** Same as Figure 4 but for surface MDA8 ozone concentrations in June 2022.



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515 **Figure 7.** Same as Figure 4 but for surface 24-h average PM<sub>2.5</sub> concentrations in February 2022.

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519 **Figure 8.** Methane emissions in the Houston area of eastern Texas ( $3^{\circ}\times 4^{\circ}$  domain) inferred from  
 520 TROPOMI satellite observations using the Integrated Methane Inversion (IMI) at 12- and 25-km  
 521 resolution. Panels from top to bottom are prior emissions from bottom-up inventories; posterior  
 522 emissions from the inversion; ratio of posterior to prior emissions; and averaging kernel  
 523 sensitivities that quantify the sensitivity of the posterior estimates to the true state. The Houston  
 524 urban boundary (US Census Bureau, 2017) is delineated in black. Emission totals over the whole  
 525  $3^{\circ}\times 4^{\circ}$  domain are shown inset in the top two rows. The sum of area-weighted averaging kernel  
 526 sensitivities for the whole domain defines the Degrees of Freedom for Signal (DOFS) inset in the  
 527 bottom panels.