

Response to reviewers' comments

Response to Reviewer #1:

[Comment 1] In this paper, the authors hypothesize that the mismatch between modelled and measured global silicate-weathering rates are due to uncertainties in the underlying erosion models. The authors therefore present a global erosion-rate model trained against cosmogenic-nuclide derived erosion rates using a random forest approach. They show that this model can reproduce erosion rates much better than a basic stream-power-erosion approach. The authors then incorporate their predicted erosion patterns into an existing steady-state silicate-weathering model and demonstrate that the new product can predict both the total global weathering flux and the spatial distribution of weathering rates better than previous models.

I found the study to be very well motivated, well written, and convincingly argued. The analysis is sound and provides a significant contribution for modelling global silicate-weathering fluxes. The figures and text are of a high quality. In particular, I appreciate the explanations of complex terms and approaches (e.g. the SHAP work) for the unfamiliar reader. I only have a few relatively minor comments that may be interesting to consider before publication.

Response: Thank you very much for your thoughtful and constructive review of our manuscript. We are gratified by your positive assessment of the study's motivation and the clarity of our explanations. Your comments have helped us identify areas where our terminology was ambiguous or where the data presentation required more rigorous quantification.

[Comment 2] L478 – 496: You write that Fig. 9 “clearly illustrates the trade-off”, that the value in Fig. 9a is “far to the right of the observed range” and that its value within the square-band is “low”. To my eyes, the peak is just to the right of the optimum and there are still reasonable values within the grey band. Next, you write that in Fig. 9b, “peaks of all three metric curves are now located squarely within the grey observational band”. For the red and blue curves that is true, but the peak of the yellow curve is clearly not in the square band and actually looks quite close to the location in Fig 9b. Can you revise this text so that the text matches the data more clearly?

Response: We sincerely thank the reviewer for this rigorous and insightful evaluation of Figure 9. We appreciate the reviewer pointing out that our descriptions of the curves' positions relative to the observational grey band were not sufficiently precise.

We agree that the term "far to the right" in L478-496 is subjective. In Figure 9a (SPIM), the R_{log}^2 peak (representing the optimal fit to the spatial pattern of observations evaluated with the original metric Park et al. used in the paper) occurs at a global flux of $\sim 4.2 \times 10^{12} \text{ mol yr}^{-1}$ which is approximately 1.7 times of the observation-based central estimate of $2.5 \times 10^{12} \text{ mol yr}^{-1}$. While some R_{log}^2 values within the grey band are technically "reasonable" (0.41

37 compared to its optimal of 0.53), the model's best spatial performance (e.g., the peak) cannot be
38 achieved without overestimating the global total flux.

39 The reviewer is correct that the peak of the yellow curve (R_{log}^2) in Figure 9b sits at
40 $\sim 3.5 \times 10^{12} \text{ mol yr}^{-1}$, which is indeed outside the $2.0 - 3.0 \times 10^{12} \text{ mol yr}^{-1}$ grey band.
41 Our previous statement that "all three metric curves" were "squarely within" the band was
42 inaccurate.

43 We have revised the text to clarify that the significant improvement in MErSiM (Fig. 9b) is the
44 convergence of the metrics. Specifically, the peak of the joint metric ($R^2 + R_{log}^2$, blue curve) and
45 the standard R^2 (red curve) now fall within the observational range. The final parameter set we
46 used for simulation is those who optimize $R^2 + R_{log}^2$.

47

48 [Comment 3] Perhaps, you can also explicitly quantify (1) the global weathering fluxes predicted
49 by the two different models (MErSiM versus Park20) including an estimate of uncertainty for
50 both, and (2) the difference between each one of these predictions and the measured total
51 weathering fluxes.

52 Response: In the original manuscript, we have explicitly quantified the predicted flux in Section
53 3.5.1 and in Table 3 where the sensitivity test is presented, but the reviewer is right that this
54 quantification could be mentioned in this section. Following the reviewer's suggestion, we have
55 added a text summary explicitly comparing the model predictions to observations. To account for
56 uncertainty in parameter selection, we consider all parameter sets whose metric value is within
57 0.05 of the maximum. The prediction uncertainty is then quantified by the variability of model
58 outputs across these acceptable parameter sets.

59 Measured target: $2.5 \pm 0.5 \times 10^{12} \text{ mol yr}^{-1}$

60 Park20 (GM09-SPIM): $4.54 \pm 0.4 \times 10^{12} \text{ mol yr}^{-1}$ (Error: $2.04 \times 10^{12} \text{ mol yr}^{-1}$, 82%
61 bias).

62 MErSiM (Revised): $3.11 \pm 0.3 \times 10^{12} \text{ mol yr}^{-1}$ (Error: $0.61 \times 10^{12} \text{ mol yr}^{-1}$, 24% bias).

63

64 [Comment 4] Finally, in the section L478 – 496 you claim that these results speak to the trade-
65 off between matching total weathering fluxes versus the pattern of weathering fluxes. I did not
66 understand how Figure 9 relates to (or informs) the spatial pattern of silicate weathering fluxes.

67 Response: We appreciate the reviewer's request for clarification on the relationship between the
68 metrics in Figure 9 and the "spatial pattern" of weathering. In our evaluation framework, the
69 performance metrics (R^2 and R_{log}^2) shown in Figure 9 are not calculated for a single location, but
70 represent the agreement between modeled and observed weathering fluxes across 81 major river

71 basins distributed globally. Because these 81 basins encompass a vast range of tectonic settings
72 (from active orogens like the Andes to stable cratons like the Amazon) and climatic zones (from
73 the tropics to high latitudes), a high R^2 value indicates that the model successfully reproduces the
74 observed spatial heterogeneity and relative differences in weathering rates among these diverse
75 regions. Therefore, the "model performance" plotted on the y-axis of Figure 9 is a direct proxy
76 for how well the model captures the global spatial pattern of silicate weathering, while the value
77 on the x-axis indicates the performance in matching the total weathering flux. A tradeoff is thus
78 obtained if the maximum y value is obtained for an x value within the grey band. We have
79 clarified this definition in the revised manuscript to avoid ambiguity.

80

81 [Comment 5] L497 – 504: Similar to above, the text here seems a bit more pushed than
82 warranted based on Fig 10. I agree that the errors are clearly smaller in the RF model, but there
83 are still quite large errors, in particular in the tropical regions. When you write things like "large"
84 errors and is "much less" severe, can you explain quantitatively what you understand by a large
85 and small error. How much were errors reduced on average etc.?

86 Response: We thank the reviewer for this constructive critique. We agree that using qualitative
87 descriptors like "large" or "much less severe" without clear benchmarks can be subjective. In the
88 revised manuscript, we have explicitly defined "large errors" as those exceeding
89 $5 \times 10^{10} \text{ mol yr}^{-1}$.

90 To provide a clearer comparison, we have calculated the following metrics across the 81 major
91 basins. First, MErSiM reduced the Mean Absolute Error (MAE) across all 81 basins by
92 approximately 35% compared to the original GM09-SPIM ($1.77 \times 10^{10} \text{ mol yr}^{-1}$ compared to
93 $2.72 \times 10^{10} \text{ mol yr}^{-1}$). Second, the number of basins exhibiting "large" positive errors
94 ($>5 \times 10^{10} \text{ mol yr}^{-1}$) dropped from 10 in GM09-SPIM to 4 in MErSiM. Third, specifically in
95 the tropics (basins located between 23S-23N), where GM09-SPIM errors often exceeded
96 $7 \times 10^{10} \text{ mol yr}^{-1}$, MErSiM reduced the average basin-scale bias by 90% ($0.41 \times$
97 $10^{10} \text{ mol yr}^{-1}$ compared to $4.12 \times 10^{10} \text{ mol yr}^{-1}$). We have added these quantitative
98 description in the explanation paragraph of Fig. 10.

99 We appreciate the reviewer for pointing out that significant errors remain in the tropics. We have
100 added a paragraph noting that these persistent biases likely stem from uncertainties in the cation
101 concentrations of sedimentary and metamorphic rocks, which remain as adjustable bulk
102 parameters in the current framework.

103

104 [Comment 6] I was confused by the discussion about the effect of CO2 increases. You explain
105 your models small response to the CO2 increase by widespread chemostasis and a resulting
106 runoff insensitivity of the weathering flux. My understanding is exactly the opposite: To me,
107 chemostasis means that concentrations remain constant with changes in runoff – contrary to what
108 is suggested in L568 (e.g. Godsey et al., 2019; Godsey et al., 2009). Under these conditions
109 (common in active mountains (Godsey et al., 2019)) the weathering fluxes should strongly

110 increase with runoff. Hence, the system should have a high sensitivity to increased CO₂ and
111 runoff. Something is off here, and maybe it is just about the term chemostasis? I guess your
112 model also has a negative temperature response, does that matter here?

113 Response: We are very grateful to the reviewer for pointing out the potential confusion regarding
114 the term "chemostasis." We agree that in the hydrological community (e.g., Godsey et al., 2009,
115 2019), chemostasis typically refers to a state where solute concentrations remain nearly constant
116 despite changes in discharge, which implies that weathering fluxes should scale linearly with
117 runoff.

118 In the original manuscript, we used "chemostatic" in the context of the Damköhler number
119 framework (Maher and Chamberlain, 2014) to describe a state of weathering flux saturation. We
120 acknowledge that this usage is ambiguous and contradicts the definition favored by the reviewer.

121 In our revised model, under greenhouse conditions (), the system enters a regime where the fluid
122 travel time (τ_f) is much shorter than the reaction equilibration time (τ_{eq}). This leads to a
123 "sluggish" response, causing the total weathering flux to plateau despite in higher temperature
124 scenario.

125 The reviewer correctly intuited the role of the negative temperature response. In our data-driven
126 erosion module, warming in high-latitude or high-altitude regions can move the environment out
127 of the optimal "frost-cracking window (Anderson and Anderson, 2010; Zhao et al., 2026),"
128 leading to a localized decrease in erosion rates. This reduction in mineral supply effectively
129 decouples the weathering flux from the kinetic acceleration expected from rising temperatures.

130 We have revised the text to replace "chemostatic" with "kinetically limited" and added a
131 discussion on the coupling between the negative temperature-erosion response and the attenuated
132 weathering feedback.

133

134 [Comment 7] Your paper demonstrates that an erosion model based on 7 input parameters
135 performs equally well to a model based on 14 parameters. That makes me wonder: would a
136 model with even fewer parameters also work? For example, if I understand Fig. 6 correctly, the
137 lithology may play a minor role. Would it be feasible to progressively eliminate parameters from
138 your model (starting with the least important) and see when the model starts breaking down?

139 Response: We thank the reviewer for this insightful suggestion regarding model robustness.
140 Achieving a balance between model simplicity and predictive accuracy is particularly important
141 for deep-time applications where boundary conditions are less certain.

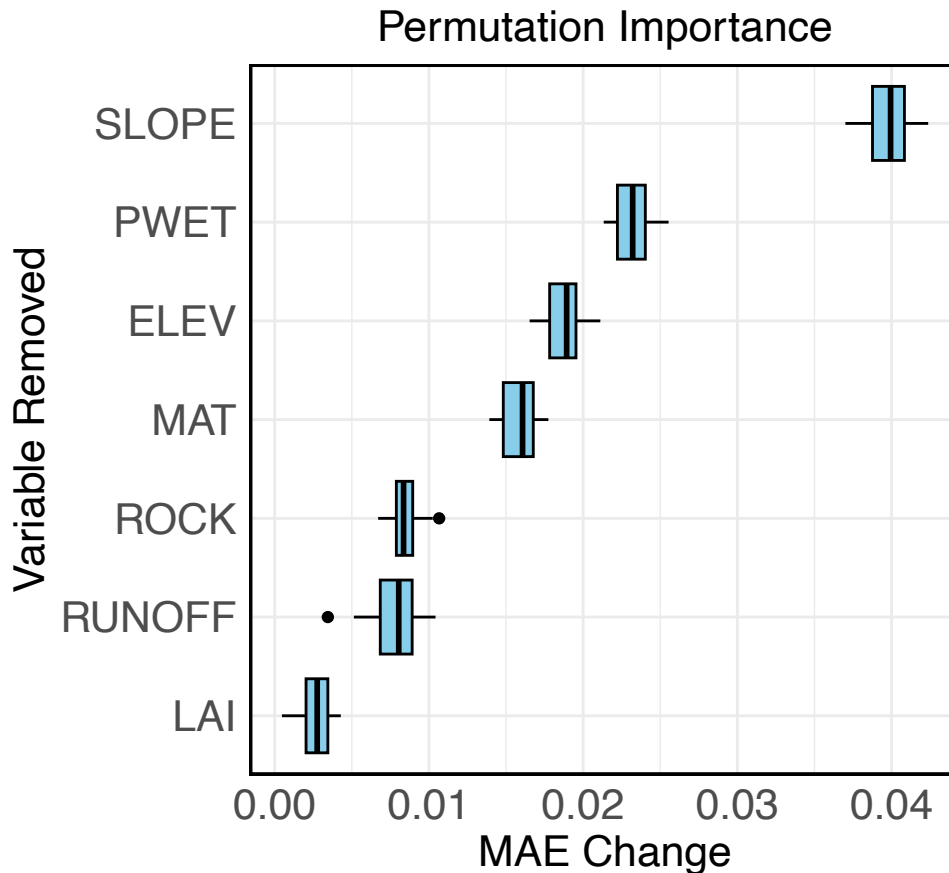
142 While our SHAP analysis (Fig. 6) suggests that lithology (ROCK) has a lower GLOBAL
143 importance compared to other predictors, it remains a fundamental factor in geomorphology as it
144 dictates rock strength and initial cation concentrations (Bluth and Kump, 1994; Bufe et al.,
145 2022). When modeling LOCAL erosion rates, lithology sometimes serves as the only other

146 predictor other than slope (e.g., Zondervan et al., 2023). Removing this variable could lead to
147 systematic spatial biases, especially in regions with distinct lithological controls.

148 To address this, we have previously conducted experiments by progressively removing variables.
149 When ROCK was eliminated (leaving SLOPE, ELEVATION, MAT, LAI, RUNOFF, and
150 PWET), the erosion model's validation R^2 decreased to 0.78 (from 0.81 in the baseline), and the
151 weathering model's maximum $R^2 + R_{log}^2$ dropped to 0.72 (from 0.86 in the 7-parameter version).
152 Although the performance decline is not "catastrophic," the 7-parameter model consistently
153 provides a superior fit to both erosion and weathering datasets. So are the results of experiments
154 eliminating other 6 predictors.

155 To further evaluate the necessity of these 7 variables, we have performed a Permutation
156 Importance Test. Unlike SHAP, which measures local contributions, this test assesses global
157 importance by randomly shuffling a single feature's values and measuring the resulting change in
158 the model's Mean Absolute Error (MAE).

159 In this test, all seven variables yielded a positive MAE change, meaning the exclusion of any
160 single variable increases the model's prediction error (Figure R1). In terms of its impact on
161 global error (MAE change), ROCK was found to be more significant than both RUNOFF and
162 LAI. This confirms that while lithology may not be the primary driver of local variation in all
163 cases, it is essential for the model's overall stability and global accuracy. Therefore, we believe
164 the 7-parameter configuration represents the "optimal" set of predictors for MERSiM v1.0.



165

166 Figure R1 Mean Absolute Error (MAE) change of randomly shuffling a single feature's value

167

168 [Comment 8] L10: Which “widely used weathering model”- can you be specific? There are
 169 several out there.

170 Response: We referred to the GM09 framework. The rephrased L10 is “The widely-used
 171 weathering model framework of Gabet and Mudd (2009), when driven by stream power erosion
 172 laws...”

173 [Comment 9] L51: became

174 Response: Fixed to "became." Thanks.

175 [Comment 10] L60: How can a model “bias [be] supported by [...] data”? Wasn’t your point that
 176 the bias is a bias because it doesn’t match global fluxes?

177 Response: Thanks. We have rephrased. The "bias" is supported by data in the sense that
 178 independent proxies (like Osmium isotopes) also suggest the model's spatial distribution of
 179 weathering is incorrect. The rephrased L60 is “It stems from a systematic overestimation of

180 weathering fluxes specifically in tropical regions, a spatial mismatch corroborated by osmium
181 isotope data (Rugenstein et al., 2021).”

182 [Comment 11] L102 – 136: Please systematically define all variables (even if they are obvious);
183 e.g. missing t in equation 1 (definition comes only for equation 2) and many of the parameters in
184 the Arrhenius relationship (equation 6).

185 Response: We have defined all parameters in the manuscript. The rephrased explanation is “The
186 core of the GM09 model describes the change as a function of time (t , in years) in regolith
187 thickness (h , in meters) as a balance between...” for equation (1) and “Where E_a is the activation
188 energy (J/mol), R is the ideal gas constant (8.314 J/(mol·K)), T_0 is the reference temperature (K),
189 and k_{rp} is a proportionality constant” for equation (6).

190 [Comment 12] L140: I guess you could cite West (2012) here who explores that steady-state
191 expression that you show

192 Response: Thanks. We have added the citation here.

193 [Comment 13] L149: You have the variable name R already in the Arrhenius relationship. It
194 would be useful to choose a different name for runoff – for example q_w .

195 Response: We have renamed the runoff variable from R to q to avoid confusion with the
196 universal gas constant.

197 [Comment 14] L263: By “this compilation” you refer to the Gaillardet et al. (1999) compilation,
198 right? Maybe specify to clarify sentence

199 Response: We will specify this compilation as “the compilation of Gaillardet et al. (1999) and
200 HYBAM”

201 [Comment 15] Figure 3: Please define the abbreviations in the figure caption

202 Response: We have added definitions for all abbreviations (e.g., PWET, LAI) and units (e.g.,
203 mm/kyr, °C) to the caption. Thanks for bringing this to our attention.

204 [Comment 16] Figure 8: Can you indicate the units of these variables?

205 Response: We have added the units in the caption.

206 Reference

207 Anderson, R.S., Anderson, S.P., 2010. Geomorphology : The Mechanics and Chemistry of
208 Landscapes.

209 Bluth, G., Kump, L., 1994. Lithologic and Climatologic Controls of River Chemistry.
210 Geochimica Et Cosmochimica Acta 58, 2341–2359. [https://doi.org/10.1016/0016-](https://doi.org/10.1016/0016-7037(94)90015-9)
211 [7037\(94\)90015-9](https://doi.org/10.1016/0016-7037(94)90015-9)

212 Bufe, A., Cook, K.L., Galy, A., Wittmann, H., Hovius, N., 2022. The effect of lithology on the
213 relationship between denudation rate and chemical weathering pathways – evidence from
214 the eastern Tibetan Plateau. *Earth Surface Dynamics* 10, 513–530.
215 <https://doi.org/10.5194/esurf-10-513-2022>

216 Coogan, L.A., Caves Rügenstein, J.K., 2025. Regulation of the carbon cycle on geological
217 timescales, in: *Treatise on Geochemistry*. Elsevier, pp. 419–465.
218 <https://doi.org/10.1016/B978-0-323-99762-1.00060-7>

219 Dessert, C., Dupré, B., Gaillardet, J., François, L.M., Allègre, C.J., 2003. Basalt weathering laws
220 and the impact of basalt weathering on the global carbon cycle. *Chemical Geology, Controls on Chemical Weathering* 202, 257–273.
221 <https://doi.org/10.1016/j.chemgeo.2002.10.001>

222 Gabet, E.J., Mudd, S.M., 2009. A theoretical model coupling chemical weathering rates with
223 denudation rates. *Geology* 37, 151–154. <https://doi.org/10.1130/G25270A.1>

224 Gaillardet, J., Dupré, B., Louvat, P., Allègre, C.J., 1999. Global silicate weathering and CO₂
225 consumption rates deduced from the chemistry of large rivers. *Chemical Geology* 159, 3–
226 30. [https://doi.org/10.1016/S0009-2541\(99\)00031-5](https://doi.org/10.1016/S0009-2541(99)00031-5)

227
228 Maher, K., Chamberlain, C.P., 2014. Hydrologic Regulation of Chemical Weathering and the
229 Geologic Carbon Cycle. *Science* 343, 1502–1504.
230 <https://doi.org/10.1126/science.1250770>

231 Moon, S., Chamberlain, C.P., Hilley, G.E., 2014. New estimates of silicate weathering rates and
232 their uncertainties in global rivers. *Geochimica et Cosmochimica Acta* 134, 257–274.
233 <https://doi.org/10.1016/j.gca.2014.02.033>

234 Rügenstein, J.K.C., Ibarra, D.E., Zhang, S., Planavsky, N.J., Von Blanckenburg, F., 2021.
235 Isotope mass-balance constraints preclude that mafic weathering drove Neogene cooling.
236 *Proc. Natl. Acad. Sci. U.S.A.* 118, e2026345118.
237 <https://doi.org/10.1073/pnas.2026345118>

238 Zhao, J., Liu, Y., Li, G., Zuo, H., 2026. Thresholds in the controls of denudation rates: A global
239 analysis of tectonic, climatic and biological factors based on machine learning. *Earth and
240 Planetary Science Letters* 674, 119750. <https://doi.org/10.1016/j.epsl.2025.119750>

241 Zondervan, J.R., Hilton, R.G., Dellinger, M., Clubb, F.J., Roylands, T., Ogrič, M., 2023. Rock
242 organic carbon oxidation CO₂ release offsets silicate weathering sink. *Nature* 1–5.
243 <https://doi.org/10.1038/s41586-023-06581-9>

244