

1 **The 2025 Mediterranean marine heatwaves throughout sea surface temperature**
2 **analogues**

a supprimé: heatwave in the context of climate change through ...

3 Davide Faranda (1,2,3), Thibault Guinaldo (4), Francisco Pastor (5), Tommaso Alberti (6), Samira Khodayar (5)

- 4 (1) Laboratoire des Sciences du Climat et de l'Environnement, UMR 8212 CEA-CNRS-UVSQ, Université Paris-
5 Saclay & IPSL, CEA Saclay l'Orme des Merisiers, 91191 Gif-sur-Yvette, France
6 (2) London Mathematical Laboratory, 8 Margrave Gardens, London W6 8RH, UK
7 (3) LMD/IPSL, ENS, Université PSL, École Polytechnique, Institut Polytechnique de Paris, Sorbonne Université,
8 CNRS, Paris France
9 (4) Météo-France, CNRS, Univ. Toulouse, CNRM, Toulouse, France.
10 (5) Mediterranean Center for Environmental Studies (CEAM), Meteorology and Climatology Area,
11 Valencia, Spain
12 (6) Istituto Nazionale di Geofisica e Vulcanologia, via di Vigna Murata 605, 00143, Rome, Italy

13 **Abstract**

14 The Mediterranean Sea experienced an exceptional marine heatwave in early summer 2025, which developed through
15 June, reached its peak between late June and early July, and persisted until mid-July. Here we analyse this event using
16 an analogue framework that differs from the common atmospheric circulation-based analogue attribution approach
17 Instead of identifying atmospheric circulation analogues first, we search directly for historical analogues of the observed
18 sea surface temperature (SST) anomaly pattern of the 20 to 30 June 2025 build-up and peak phase, and then compare
19 the associated atmospheric and air-sea surface conditions in a past (1950 to 1986) and a present (1987 to 2024) period
20 using ERA5 daily fields over the Mediterranean between June and September. Our results suggest that the present
21 period exhibits higher SST anomalies than the past one, with the strongest positive differences over the western
22 Mediterranean and the Adriatic. Warmer 2 m air temperatures over surrounding land areas and higher incoming
23 shortwave radiation over broad parts of the basin are also observed, while sea level pressure, wind speed, latent heat
24 flux, sensible heat flux, and longwave radiation display smaller and more spatially heterogeneous changes. Thus, the
25 2025 event occurred in a warmer thermodynamic background than analogous SST patterns in earlier decades. Since
26 the analogue framework cannot fully separate externally forced trends from internal variability and it does not explicitly
27 resolve ocean dynamical processes, these results do not constitute a complete formal attribution on their own.
28 Nevertheless, the analogue analysis supports the interpretation that long term warming contributed substantially to the
29 intensity of the 2025 Mediterranean marine heatwave and shows how SST analogue comparisons can complement
30 existing attribution approaches for marine extremes.

a supprimé: 2025 lasting from 20 to 30 June.
a supprimé: more
a supprimé: circulation based
a supprimé: period
a supprimé: one
a supprimé: 2024), based on
a supprimé: during
a supprimé: to
a supprimé: They do not constitute a complete formal attribution on their own, since
a supprimé: processes such as advection, mixing, and stratification changes. ...

31 **1. Introduction**

32 Anthropogenic climate change has unequivocally warmed the oceans and the atmosphere, increasing the
33 frequency and intensity of different extreme events (Forster et al., 2025; Seneviratne et al., 2021). Among
34 them, marine heatwaves (MHWs) are prolonged periods of unusually warm ocean temperatures and represent
35 one of the clearest manifestations of the interaction between climate internal variability and long-term changes
36 in the ocean (Hobday et al., 2016; Holbrook et al., 2020). Their frequency, duration, and intensity have
37 increased over the last century at the global scale (Frölicher et al., 2018; Oliver et al., 2018; Laufkötter et al.,
38 2020) as evidenced by the recent surge in both global and regional extreme heat events (Gyuleva et al., 2025,
39 England et al., 2025, Guinaldo et al., 2025). These events can produce severe ecological and socioeconomic
40 impacts, including mass mortality of marine organisms, coral bleaching, seagrass decline, and fisheries
41 disruption (Smale et al., 2019; Smith et al., 2021; Garrabou et al., 2022; Wernberg et al., 2025).

56 The Mediterranean Sea is a particularly sensitive region in this context. It is widely recognized as a climate
 57 change hotspot (Giorgi 2006), with regional warming trend exceeding the global average (Cramer et al., 2018;
 58 Tuel & Eltahir, 2020; Pastor et al., 2020; Seneviratne et al., 2021). Consistent with this background trend,
 59 Mediterranean MHWs have become more frequent, longer, and more intense in recent decades (Darmaraki
 60 et al., 2024) with an annual mean number of days increasing from 1 day/year over the period 1982-1999 to
 61 5.5 days/years over the period 2000-2021 (Pastor & Khodayar, 2023). Several remarkable events have been
 62 documented during the last two decades (Guinaldo et al., 2023; Martínez et al., 2023; Pastor et al., 2024) w ^{a supprimé: (Simon et al., 2023, Guinaldo}
 63 major impacts on local marine ecosystems and coastal environments (Coll et al. 2010, Garrabou et al., 2022;
 64 Artana et al., 2026). Recent work has also highlighted that Mediterranean MHWs can be intensified when th ^{a supprimé: 2026} .
 65 co-occur with atmospheric heat extremes and anomalous summer circulation patterns that favour enhanced
 66 radiative input and weak surface winds (Sen Gupta et al., 2020; Pastor et al., 2024; Bonino et al., 2025;
 67 Paredes-Fortuny et al., 2025), conditions whose frequency is increasing in recent years (d'Andrea et al. 202 ^{a supprimé: 2025}
 68 The contribution of anthropogenic climate change to the increasing likelihood and intensity of MHWs is no ^{a supprimé: itself}
 69 unequivocally established, with the vast majority of observed events attributable to human influence (Frölicher ^{a supprimé: 2024}
 70 et al. 2018). The question addressed here is more specific. We examine the June 2025 Mediterranean event
 71 using an analogue framework designed to compare physically similar SST anomaly patterns across different
 72 climate backgrounds. The aim is not to replace formal probabilistic attribution studies based on climate model
 73 ensembles, but to provide an event-specific observational line of evidence and a process-oriented diagnostic
 74 of how the background thermodynamic environment differs between earlier and more recent analogues.

75 Analogue-based attribution has become an established complementary approach in the extreme event
 76 attribution field (Vautard et al., 2016; Faranda et al., 2022; Faranda et al., 2024). In its more classical form,
 77 the method first identifies atmospheric circulation analogues and then evaluates how associated impacts differ
 78 between climate states. Here we adopt a different perspective. We first identify analogues of the observed
 79 SST anomaly pattern itself, and only then examine the atmospheric and surface fields associated with those
 80 SST analogue dates. This choice is motivated by the nature of the event under study, namely a MHW, for
 81 which the SST pattern is the central feature to be matched.

82 The study has three objectives. First, we document the exceptional character of the June 2025 event (Section
 83 3). Second, we assess how the associated SST, atmospheric, and air-sea flux anomaly composites differ
 84 between past and present SST analogues (Section 4). Third, we discuss the extent to which these differences
 85 are consistent with thermodynamic amplification in a warmer climate, while explicitly acknowledging the limits
 86 of the analogue framework, in particular its inability to isolate forced and internal components fully and its
 87 omission of explicit ocean dynamical processes (Section 5).

88 2. Data and Methods

89 2.1 Reanalysis and observational datasets

90 The main dataset used in this study is the ECMWF ERA5 reanalysis (Hersbach et al., 2020). ERA5 provides
 91 globally gridded atmospheric and surface fields at 0.25° resolution from 1940 onward. In this work, ERA5 is
 92 used as a temporally consistent, observation constrained dataset for both the characterization of the 2025
 93 event and the identification of historical analogues. All analyses are restricted to the June to September
 94 season (JJAS).

95 For SST, the ERA5 product relies on observationally constrained analyses rather than on a fully coupled
 96 ocean reanalysis. In practice, the SST field combines HadISST for the earlier part (before 2007) of the record

and OSTIA in the more recent period (after 2007), both regridded within the ERA5 framework. We therefore refer to ERA5 SST as a consistent gridded SST product used within ERA5, rather than as an ocean reanalysis field in the strict dynamical sense.

The variables analysed in the analogue-based attribution are SST, 2 m air temperature (T2m), mean sea level pressure (MSL), surface net shortwave radiation (SSR), surface net longwave radiation (STR), surface latent heat flux (SLHF), surface sensible heat flux (SSHF), and 10 m wind speed (WSPD, derived from U10 and V10). These variables were selected because they are relevant to the development and maintenance of summer MHWs in the Mediterranean Sea through radiative forcing, air-sea exchanges, and near surface circulation (Capotondi et al., 2024; Darmaraki et al., 2024, Bonino et al. 2025).

For descriptive purposes in the event characterization section, we also use the Mediterranean Sea High Resolution L4 Sea Surface Temperature Reprocessed product (MED REP L4; Copernicus Marine Service Information, <https://doi.org/10.48670/moi-00173>). This observational product is not used in the analogue protocol and its role is only to document the timing, spatial structure, and magnitude of the 2025 event ~~from~~ an independent observational perspective.

2.2 Definition of anomalies and target event

For ERA5, daily anomalies are computed at each grid point relative to the 1950 to 2024 JJAS daily climatology. Thus, for each calendar day, the corresponding climatological mean for that day is removed before further analysis. All variables are then smoothed with an 11 day centred moving average, following the Climameter framework and ensuring that the target event and candidate analogue fields are compared at the same temporal scale (Faranda et al., 2024). The target event is defined as the temporal mean SST anomaly over 20 to 30 June 2025. Daily SST anomalies are first computed, then averaged over this 11 day period. This period was selected because it corresponds to the most rapid build up and peak phase of the exceptional western Mediterranean warming documented by the observational datasets and by ERA5. The target is therefore a temporally averaged SST anomaly field, not a single day snapshot.

2.3 Analogue selection

We search for analogues of the 20 to 30 June 2025 SST anomaly pattern over the Mediterranean basin. The matching domain is the entire Mediterranean basin and immediate coastal grid points represented in ERA5. For each JJAS day in the historical record, we compute the Euclidean distance between the candidate daily SST anomaly field and the target event pattern. The distance is calculated over all grid points in the domain, with cosine of latitude weighting to account for the different area represented by each grid cell. Candidate days are then ranked from the smallest to the largest distance. The analogue days are those with the smallest distances (i.e., the first 1% of dates), that is, the SST anomaly fields most similar to the target 2025 pattern. The search is performed independently for two subperiods, a past period, 1950 to 1986, and a present period, 1987 to 2024. The year 1987 is used as the split because it lies close to a documented acceleration in Mediterranean warming and also divides the record into two comparably sized samples (Lionello & Scarascia, 2018). We stress that the past period is not a pre industrial baseline. It is simply a cooler segment of the observational era, that could be cooler due to the decrease in aerosol emissions over western Europe or a different phase of the multidecadal variability (Pisano et al. 2020).

In each period we retain the top 42 analogue days. This corresponds to about 1% of all available JJAS days in each subperiod and follows the standard Climameter choice adopted in previous sensitivity analyses (Faranda et al., 2024). Fixing the number of analogues in both periods allows direct comparison of composite means between the two climate backgrounds.

For each period, we compute composites by averaging the selected analogue days for SST, T2m, MSL, SSR, STR, SLHF, SSHF, and WSPD. We then compare the past and present composites as well as the present minus past difference. This difference is interpreted as a diagnostic of how the environmental conditions associated with similar SST anomaly patterns differ between the earlier and later parts of the record. A positive difference indicates that the present climate analogues are associated with stronger positive anomalies—consistent with a warming signal. For instance, a positive SST difference implies that recent analogues occur under warmer SST conditions than similar events in the past. Conversely, near-zero or negative differences suggest either no significant change or a reduction in anomaly amplitude under current climate conditions, though this is less common for variables associated with heat extremes. We remark that this framework does not justify a claim that identical atmospheric or oceanic states are being compared. It compares atmospheric and surface conditions associated with similar SST anomaly patterns. The method is therefore best understood as conditional and diagnostic. It provides information on how the thermodynamic and near surface atmospheric context of SST analogue events has shifted between the two periods, but it does not in itself isolate external forcing from internal variability.

2.4 Statistical significance

To assess whether present minus past composite differences are robust to sampling uncertainty, we use a non parametric bootstrap test following Faranda et al. (2022, 2024). The 84 analogue dates, 42 from the past and 42 from the present, are pooled together. We then generate 10,000 random resamples, each time dividing the 84 dates into two groups of 42 without preserving their original period labels. For each resample we compute the difference between the two composite means at each grid point. The resulting empirical distribution provides the expected range of differences under the null hypothesis that the observed present minus past difference arises from random sampling within the pooled analogue set. Grid point differences are considered significant at the 95% level when they fall outside the 2.5th to 97.5th percentile interval of the bootstrap distribution. A false discovery rate correction is applied to account for multiple testing across the spatial domain (Wilks, 2016).

2.5 Climate variability indices and analogue quality

To characterize the background large scale variability associated with the analogue dates, we analyse six standard climate indices, El Niño Southern Oscillation (ENSO), North Atlantic Oscillation (NAO), Pacific Decadal Oscillation (PDO), Atlantic Multidecadal Oscillation (AMO), East Atlantic pattern (EA), and Scandinavian pattern (SCAND). The indices are taken from the following sources: ENSO 3.4 from ERSSTv5 (Extended Reconstructed Sea Surface Temperature version 5); PDO from SWFSC (Southwest Fisheries Science Center); NAO from ICPC (International Climate Prediction Center); AMO from ERSST TS (Extended Reconstructed Sea Surface Temperature time series); EA and SCAND from CPC (Climate Prediction Center, NOAA). For each analogue date we extract the value of the corresponding index time series and compare the past and present distributions using violin plots and a two tailed Cramér von Mises test. These comparisons are purely descriptive and inform on the large-scale context without demonstrating a causal role of any given mode in the 2025 event. We also evaluate analogue quality using the mean Euclidean distance between the target 2025 SST pattern and its selected analogues, following Faranda et al. (2022). This quality metric indicates how closely the selected analogue set matches the target pattern relative to typical analogue distances within each period.

3. 3. Event characterisation: the June 2025 Mediterranean marine heatwave

a supprimé: The

The June to July 2025 MHW was exceptional in both duration and spatial extent. Figure 1a shows the daily mean western Mediterranean SST together with the 1991 to 2020 mean climatology, the 90th percentile

threshold used to define MHW conditions and the MHW category (Hobday et al. 2016, Hobday et al. 2018). The event developed progressively during June 2025, with SST remaining above the climatological threshold for the whole month and then intensifying sharply in the second half of June and the start of July. According to this, western Mediterranean presented moderate MHW conditions persisting from early June to late July, with two periods of stronger intensity. A short strong phase occurred in mid June, followed by a more sustained severe phase around the turn of June to July. The daily western Mediterranean SST reached its maximum close to 28 °C in early July. MHW intensity rapidly decayed until its end in mid-July. This time evolution of the MHW event classifies it as a long-duration event, significantly exceeding the average of 16 days for Mediterranean MHWs in the period 1982–2021 (Pastor et al., 2023).

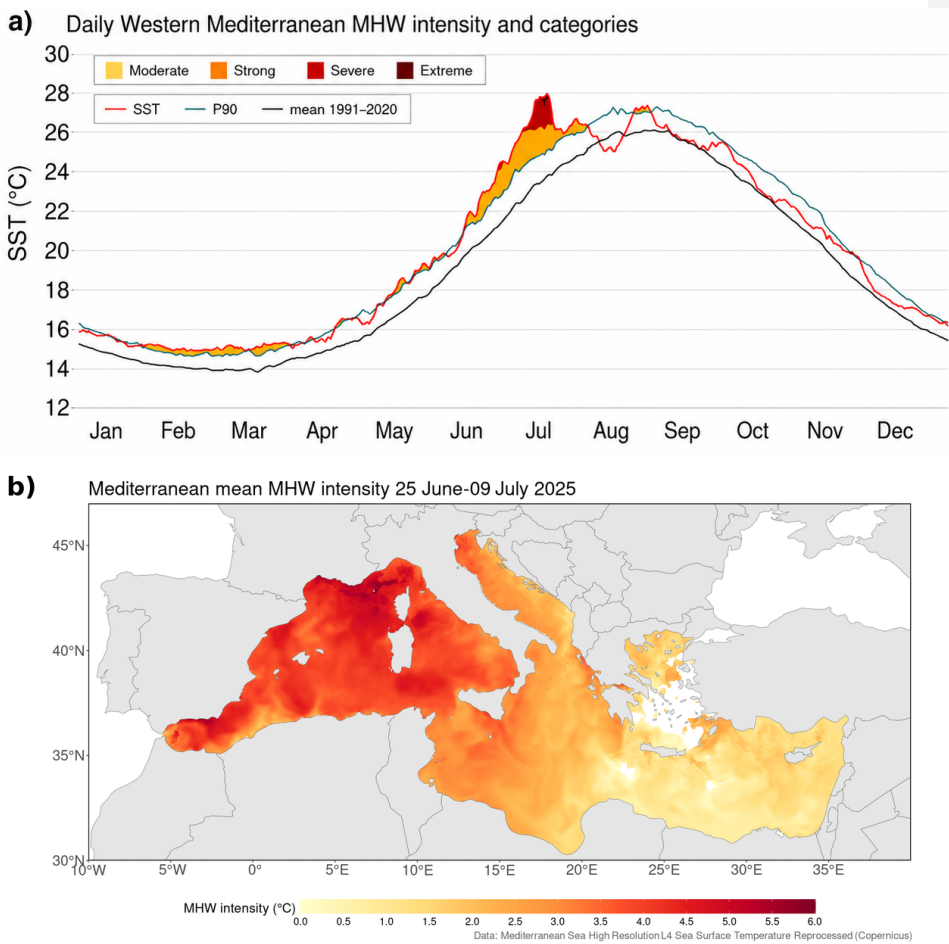


Figure 1. a) Daily SST averaged over the Western Mediterranean Sea for 2025 (red), together with the 1991–2020 daily climatological mean (black line) and the 90th percentile MHW threshold (P90, teal line). Shading indicates the MHW category (moderate, strong, severe, extreme) following Hobday et al. (2016, 2018). **b) Mean MHW intensity (temperature anomaly above the 1991–2020 climatology, in °C).**

206 over the Mediterranean Sea for the 25 June–09 July 2025 period, i.e. the strong-category pha a supprimé: June-09
 207 highlighted in panel a. White areas over the sea are grid points where no MHW was detected. SST da a supprimé: period (strong category
 208 obtained from the Mediterranean Sea-High Resolution L4 Sea Surface Temperature Reprocess a supprimé: a).
 209 (MED-REP-L4) dataset from Copernicus.

210 Figure 1b shows the mean MHW intensity over the Mediterranean Sea for the period 25 June to 09 July 2025,
 211 corresponding to the peak phase identified in Figure 1a. The strongest intensities were concentrated in the
 212 western Mediterranean, particularly between the Gulf of Lions, the Balearic Sea, Corsica, Sardinia, and the
 213 Tyrrhenian sub-basin, with local values above +4.5 °C and in some areas approaching +5.5 to 6 °C. Positive
 214 intensities extended across most of the basin, but with a clear west to east decrease. The central
 215 Mediterranean showed widespread moderate to strong conditions, whereas the eastern Mediterranean
 216 exhibited weaker intensities overall. The spatial distribution of MHW intensity therefore qualifies the event a supprimé: distribution
 217 a western Mediterranean-centred one (Fig. 1b): although it affected much of the basin, its most exception a supprimé: it
 218 expression was clearly located in the western Mediterranean, which motivates focusing the following sectio a supprimé: event (Fig 1b), which motivates the following
 219 on this sub-basin. The temporal window 20 to 30 June 2025 used for the analogue target therefore sections to be focused on this basin. The event
 220 corresponds to the onset of the strongest basin-wide phase and captures the spatial structure that la a supprimé: Mediterranean, but
 221 culminated during the 25 June to 09 July interval shown in Figure 1b. a supprimé: basin.

222 ERA5 shows that the event coincided with positive MSL anomalies over parts of southwestern and central
 223 Europe (Fig. 2i) and with widespread positive SSR anomalies over the Mediterranean region (Fig. 2m). T2m
 224 anomalies were strongly positive over surrounding land areas, especially over the Iberian Peninsula, southern
 225 France, Italy, and North Africa (Fig. 2e). WSPD anomalies were weak to moderate and spatially
 226 heterogeneous (Fig. 2ac). These features are consistent with a hot and radiatively favourable large scale
 227 environment during the peak phase of the event, but the present paper does not attempt a formal classification
 228 of the associated atmospheric circulation pattern or of concurrent atmospheric heatwaves.

229 4. Results

230 4.1 SST analogue composites

231 Figure 2a–d compares the target SST anomaly field for 20 to 30 June 2025 with the past and present SST
 232 analogue composites and with their difference. The past and present composites both reproduce the broad
 233 spatial structure of the target event, with stronger anomalies in the western Mediterranean than in the eastern
 234 basin. This indicates that the 2025 pattern is not unprecedented in terms of its large-scale spatial organization.
 235 The amplitudes, however, differ between the two analogue periods. The past composite shows weaker
 236 positive SST anomalies than the present composite over much of the basin. The present minus past SST
 237 difference is positive over large areas, with the strongest values in the western Mediterranean and parts of
 238 the Adriatic. Many of these positive differences are significant at the 95% level after bootstrap testing. Thus,
 239 within the analogue framework, similar SST anomaly patterns are associated with systematically higher SST
 240 anomalies in the later part of the record.

241 4.2 Near surface air temperature and pressure

242 Figure 2e–h shows the T2m fields. The target event exhibits strong positive air temperature anomalies over
 243 land surrounding the Mediterranean. Both past and present analogue composites retain positive T2m
 244 anomalies over southern Europe and North Africa, but the present composite is generally warmer. The present
 245 minus past T2m difference is positive over substantial land areas, with significant values over parts of the
 246 Iberian Peninsula, southern France, Italy, the Balkans, and North Africa. Over the sea, T2m differences are
 247 smaller and less spatially coherent. The target event is associated with positive pressure anomalies over the

257 western and central Mediterranean sector and negative anomalies farther east. The past and present
 258 analogue composites show broadly similar large scale pressure structures. In contrast to SST and T2m, the
 259 present minus past MSL differences are weaker, spatially heterogeneous, and often not significant over the
 260 basin. This suggests that the pressure patterns associated with the SST analogues do not differ as strongly
 261 between the two periods as the thermodynamic fields do.

262

263 4.3 Radiative fluxes

264 Figure 2m–p shows SSR. During the target event, positive shortwave anomalies cover much of the
 265 Mediterranean and adjacent land areas. Both analogue composites also display positive SSR anomalies, but
 266 the present composite is generally stronger over broad parts of the basin and neighbouring regions. The
 267 present minus past SSR differences are positive over much of the western and central Mediterranean and are
 268 significant at many grid points. Figure 2q–t shows STR. The target event presents negative STR anomalies
 269 over wide marine areas, while the analogue composites are weaker and more spatially mixed. The present
 270 minus past STR difference is not uniformly positive over the basin. Instead, it shows a heterogeneous
 271 structure, with positive values in some areas and near zero or negative values elsewhere. The signal is
 272 therefore less spatially robust than for SST, T2m, or SSR, and its interpretation requires caution.

273 4.4 Turbulent heat fluxes and wind speed

274 Figure 2u–x shows SLHF. The event field and the analogue composites display mixed spatial structures. The
 275 present minus past difference is generally weak and heterogeneous, with localized positive and negative
 276 regions and limited significance. Within this framework, SLHF does not show a basin wide shift comparable
 277 to that seen in SST or SSR. Figure 2y–ab shows SSHF. Here again the present minus past differences are
 278 modest and spatially variable. Over parts of the Mediterranean basin the differences are close to zero, while
 279 surrounding land and coastal regions show a mix of positive and negative values. The SSHF response is
 280 therefore not spatially coherent at the basin scale. Figure 2ac–af shows WSPD. The target event is associated
 281 with generally weak wind anomalies over much of the basin, but with pronounced regional contrasts. The past
 282 and present composites also show weak and heterogeneous wind anomaly structures. Present minus past
 283 wind speed differences are small in most regions and only locally significant.

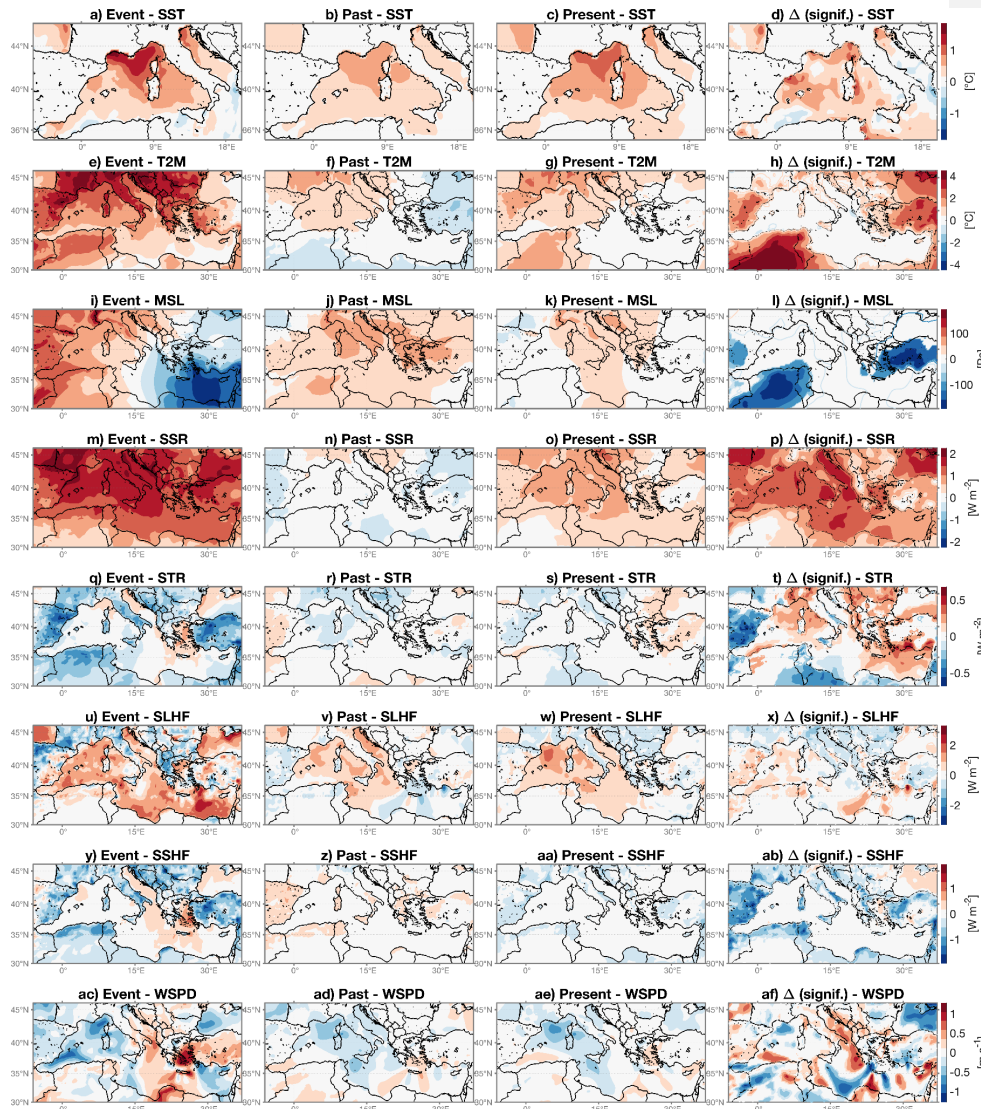
284 4.5 Climate variability indices and analogue quality

285 The distributions of ENSO (El Niño–Southern Oscillation), NAO (North Atlantic Oscillation), PDO (Pacific
 286 Decadal Oscillation), AMO (Atlantic Multidecadal Oscillation), EA (East Atlantic pattern), and SCAND
 287 (Scandinavian pattern) differ between past and present analogue dates according to the two-tailed Cramér–
 288 von Mises test, with $p < 0.001$ for all six indices (Figure 3). Because the sample is limited to 42 dates per
 289 period, these p-values quantify a clear distributional shift but should not be over-interpreted as a precise
 290 probability statement. The Figure 3 shows consistent shifts between past and present distributions for
 291 modes. ENSO moves from a distribution centred on negative values in the past to values closer to neutral
 292 slightly positive in the present, with a reduced spread. NAO shifts from a predominantly positive distribution
 293 the past to a distribution centred on negative values in the present, with a broader range. PDO changes from
 294 a distribution centred on positive values to one centred on negative values. AMO shifts from weakly negative
 295 values toward values closer to zero or slightly positive. EA changes from a past distribution with two preferred
 296 states, one negative and one positive, to a present distribution concentrated near neutral values. SCAND
 297 shifts from positive values in the past to negative values in the present. These differences indicate that the
 298 analogue dates in the two periods sample different large-scale states of oceanic and atmospheric variability.

a supprimé: The indices are taken from the following sources: ENSO 3.4 from ERSSTv5 (Extended Reconstructed Sea Surface Temperature version 5); PDO from SWFSC (Southwest Fisheries Science Center); NAO from ICPC (International Climate Prediction Center); AMO from ERSST_TS (Extended Reconstructed Sea Surface Temperature time series); EA and SCAND from CPC (Climate Prediction Center, NOAA).

a supprimé: present. These

308 The result is statistical and does not imply that any of these modes individually explains or causes the 2025
 309 event but they might have contributed to the occurrence of the event. Finally, the analogue quality analysis
 310 (Figure 4) shows that the selected analogue sets reproduce the target circulation pattern with comparable skill
 311 in both periods. The 2025 event lies within the range of analogue quality values but toward the upper part of
 312 the distributions. This indicates that the event is not outside the analogue space but is not among the most
 313 typical configurations. The comparison between periods does not show a systematic improvement of analogue
 314 quality in the present climate, despite the temporal proximity of the event.



316

317

318

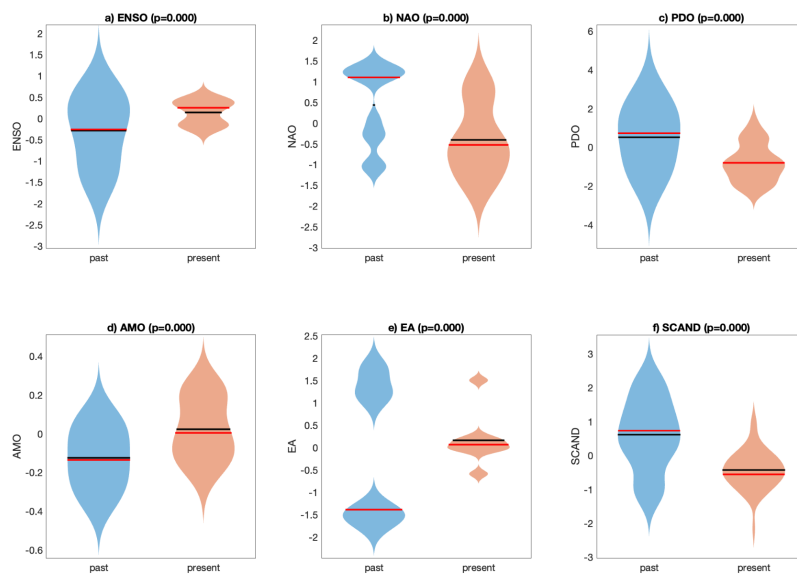
319

320

321

322

Figure 2. Multi-variable analogue analysis for the period 20-30 June 2025. Rows: sea-surface temperature, SST (a–d); 2-m air temperature, T2m (e–h); mean sea-level pressure, MSL (i–l); surface net shortwave radiation, SSR (m–p); surface net longwave radiation, STR (q–t); surface latent heat flux, SLHF (u–x); surface sensible heat flux, SSHF (y–ab); 10-m wind speed, WSPD (ac–af). Columns: event anomaly; Past analogue composite (1950–1986); Present analogue composite (1987–2024); Present–Past difference. Color-filled areas indicate 95% bootstrap significance. Temperatures are in °C, fluxes in $W m^{-2}$, and wind speed in $m s^{-1}$.



323

324

325

326

327

328

329

330

331

332

Figure 3. Violin plots comparing the distributions of natural variability indices during analogue dates in the past (1950–1986, blue) and present (1987–2024, orange) periods. The indices include: a) El Niño–Southern Oscillation (ENSO), b) North Atlantic Oscillation (NAO), c) Pacific Decadal Oscillation (PDO), d) Atlantic Multidecadal Oscillation (AMO), e) East Atlantic pattern (EA), and f) Scandinavian pattern (SCAND). Each violin shows the full probability density of the distribution, with horizontal red and black lines indicating the mean and median, respectively. The width of the violin reflects the relative frequency of values. All comparisons yield statistically significant differences between past and present periods (two-tailed Cramér–von Mises test, $p < 0.001$), suggesting that the underlying state of large-scale climate modes has shifted over time.

333

5. Discussion

334

335

336

337

338

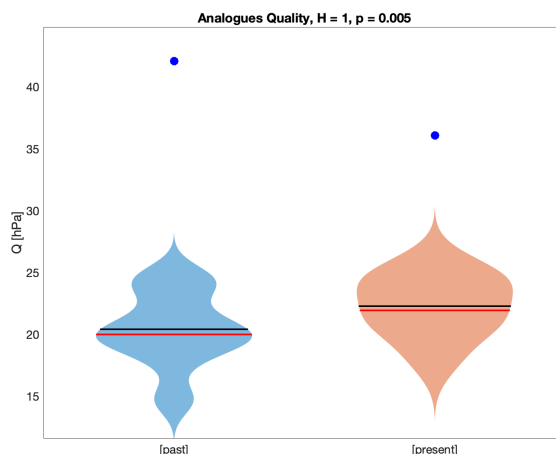
The analogue comparison reveals a physically coherent picture: the clearest differences between past and present analogues of the June 2025 MHW lie in the thermodynamic fields, SST itself, near-surface temperature over surrounding land, and incoming shortwave radiation, rather than in the dynamic fields. S level pressure, wind speed, latent heat flux and sensible heat flux show weaker, spatially heterogeneous present-minus-past differences with limited basin-wide coherence. This is consistent with the expectation that

a supprimé: fields —

a supprimé: radiation —

341 long-term warming primarily elevates the background thermodynamic state, whereas forced changes in t
 342 large-scale atmospheric circulation are smaller and harder to detect against internal variability over t
 343 observational record (Deser et al., 2012; Shepherd, 2014; Faranda et al., 2023). We stress, however, th
 344 internal variability is only one of several plausible readings of the patchy present-minus-past differences
 345 sea level pressure, wind speed, and latent and sensible heat fluxes. These heterogeneous patterns may
 346 equally reflect a spatially complex atmospheric response, forced circulation changes whose amplitude is
 347 simply weaker than the internal variability sampled by 42 analogue dates per period, or the influence of
 348 regional processes such as land-sea thermal contrasts, orographic channelling of the regional wind systems,
 349 and coastal upwelling, which are not well captured by basin-scale composites. Discriminating between these
 350 explanations would require large initial-condition model ensembles rather than the observational analogue
 351 framework used here. Our results complement recent work showing that concurrent atmospheric heatwaves
 352 and weakened surface winds under subtropical ridges can intensify MHW through modified air-sea heat flux
 353 exchanges (Bonino et al., 2025; Paredes-Fortuny et al., 2025; Pastor et al., 2024). The relatively weak and
 354 spatially incoherent wind speed and turbulent flux differences in the present analogue composites suggest
 355 that the 2025 event's exceptional character, such as other recent MHW (Gyuleva et al. 2025, Dong et al.
 356 2025), was not primarily driven by a shift of the large-scale dynamical regime, but rather by the superpositi
 357 of a recurrent SST anomaly pattern onto a warmer thermodynamic background without any change of variance
 358 (Xu et al. 2022): a warmer ocean, a warmer atmosphere, and more available solar energy reaching a stratifi
 359 upper-ocean.

a supprimé: while
 a supprimé: remains more constrained by
 a supprimé: that has not yet shifted significantly
 a supprimé: record. It complements
 a supprimé: 2025),
 a supprimé: 2022) :
 a supprimé: upper-ocean →



360

361 **Figure 4.** Analogue quality scores (Q) for the set of analogues selected dates in the past (1950–1986, blue)
 362 and present (1987–2024, orange) periods. Violin plots show the distribution of mean Euclidean distances
 363 between 42 analogue SST patterns and randomly selected JJAS days from the same period. The horizontal
 364 red and black lines indicate the mean and median of each distribution, respectively. The filled blue dots mark
 365 the actual Q values computed between the target 2025 SST event and its selected analogues. A lower Q
 366 indicates a better match. The statistical difference between past and present distributions is confirmed by a
 367 two-tailed Cramér–von Mises test ($p = 0.005$).

375 The positive SSR differences between present and past analogues are the most spatially coherent and robust
 376 signals over the western and central Mediterranean basin. This is consistent with enhanced clear-sky
 377 conditions and reduced cloudiness in summer as documented both regionally (Guinaldo et al. 2023, Bonino
 378 et al. 2025, Behr et al. 2025) and at a global-scale (Sen Gupta et al. 2020, Goessling et al. 2025) but this point
 379 cannot be quantified directly with the present methodology. By contrast, STR differences are spatially
 380 heterogeneous and cannot be interpreted as a simple basin wide increase in longwave forcing. Several
 381 mechanisms may contribute to the observed radiation patterns, including changes in the thermodynamic state
 382 of the lower atmosphere, changes in humidity, and changes in cloudiness. We did not explicitly analyse cloud
 383 cover or specific humidity, and multiple mechanisms can contribute to the observed radiation changes,
 384 including thermodynamic modifications of the lower atmosphere and long-term drying trends. The radiation
 385 results are best interpreted as part of a consistent surface energy budget signal rather than as an isolated
 386 causal diagnosis. This interpretation should, however, be stated with care. The method does not compare
 387 identical atmospheric or coupled ocean atmosphere states. It compares atmospheric and surface conditions
 388 associated with similar SST anomaly patterns. The results therefore support a conditional statement. In the
 389 later part of the historical record, SST patterns resembling the June 2025 event are associated with stronger
 390 marine warming and, in several regions, with stronger atmospheric warming and stronger shortwave input
 391 than in the earlier part of the record. This is consistent with thermodynamic amplification in a warmer climate
 392 background, but it is not a complete attribution proof on its own.

393 A first-order quantification of the conditional shift can be obtained directly from the analogue composites in
 394 Figure 2. Over the grid points where the present-minus-past difference is significant at the 95% level after
 395 bootstrap and FDR correction, the SST shift typically reaches +1 to +1.5 °C across broad sectors of the
 396 western Mediterranean and parts of the Adriatic, while T2m shifts over surrounding land are of comparable to
 397 larger magnitude, locally in the +1.5 to +2.5 °C range over the Iberian Peninsula, southern France, Italy and
 398 the Balkans. The corresponding present-minus-past SSR differences are of ~~the order of~~ a few W m^{-2} over
 399 much of the western and central basin. These figures should be interpreted as conditional metrics: they
 400 describe the typical shift of the thermodynamic environment in which SST patterns analogous to the June
 401 2025 event occur, and not as a probabilistic attribution estimate of the human contribution to the 2025 event
 402 itself. Internal variability associated with the different distributions of low-frequency modes sampled by the two
 403 analogue populations (Figure 3) can also contribute to these magnitudes, which is why we present them as
 404 ranges rather than as point estimates.

405 Compared to formal probabilistic attribution based on CMIP-class model ensembles, the SST-analogue
 406 framework offers complementary, rather than redundant, information for an event such as the June 2025
 407 Mediterranean MHW. Model-ensemble attribution typically targets a scalar return-period or risk-ratio change
 408 in a regional SST or MHW index. Here, by holding the SST anomaly pattern approximately constant by
 409 construction, we obtain spatially resolved, multi-variable composite differences for SST, T2m, MSL, SSR,
 410 STR, SLHF, SSHF, and WSPD, all from observation-constrained reanalysis fields. This disaggregation makes
 411 it possible to identify which surface energy budget terms shift most coherently between past and present
 412 analogue populations — in this case clearly SST, T2m and SSR — and which terms remain weak and spatially
 413 heterogeneous — here MSL, STR, SLHF, SSHF and WSPD. ~~Model ensembles can of course also deli~~
 414 variable-by-variable, grid-point-resolved ~~diagnostics. The distinctive feature of the present framework is rath~~
 415 ~~that these diagnostics are conditioned on observed analogues of one specific event and are therefore deriv~~
 416 ~~directly from the observational record, instead of being obtained from model simulations or attributi~~
 417 ~~ensembles.~~ The two approaches are therefore best used jointly, as already suggested by hybrid framework
 418 combining observational analogues with large ensembles.

a supprimé: CMIP-based attribution

a supprimé: quantify the risk-ratio, but does not provide this kind

a supprimé: diagnostic

a supprimé: analogue environment for

a supprimé: event.

425 The shifts in ENSO, PDO, AMO, EA, and SCAND distributions indicate an important and complex interaction
 426 in the observational analogue method: the analogue populations are embedded in different low frequency
 427 variability states. The observed present minus past differences reflect a combination of externally forced long-
 428 term warming and changes in internal variability. It is a key aspect of every attribution study in order to separate
 429 such shifts from the externally forced signal. Numerous methods have been proposed (Trenberth et al. 2015,
 430 Naveau et al. 2020, Otto 2023) to overcome these limitations and sample in the best way possible the
 431 decennial to multi-decennial variability including the use of Large Ensemble (Deser et al. 2020). The
 432 observed differences provide a physically interpretable signal, but they cannot be partitioned into forced a
 433 unforced components. As such, hybrid approaches that combine observational analogues with large
 434 ensemble model diagnostics could help address this in future work.

a supprimé: 2020) or the UNSEEN methods (Atkins et al. 2025). →

435 A second important limitation concerns ocean dynamics. Our analysis does not explicitly include ocean
 436 advection, mixed layer depth, vertical mixing, or subsurface heat content. These processes can be important
 437 in the Mediterranean, and in some seasons or subregions they may substantially modulate the onset, intensity,
 438 or persistence of marine heatwaves. The omission is less severe in the summer Mediterranean than it would
 439 be in other contexts because stratification is typically strong and mixed layers are shallow and less variable in
 440 summer (d'Ortenzio et al., 2005; Juza et al., 2022). Even so, ocean dynamical contributions cannot be ruled
 441 out and the present study should not be interpreted as a full process closure of the 2025 event.

442 A key assumption of the analogue method is that variable structures remain physically sufficiently comparable
 443 across climate states, meaning that a pattern of SST anomalies in 1980 represents a dynamically simi
 444 configuration to the same pattern in 2020, differing primarily in the background state. In the Mediterrane
 445 Sea, the basin geometry and recurrent large scale forcing constrain the broad spatial organization of
 446 anomalies, making this assumption plausible. Nevertheless, long-term warming can modify the magnitude
 447 SST anomalies and their west east gradients, which may reduce the representativeness of historical
 448 analogues, especially when recent extremes are compared with earlier decades (Marullo et al., 2011; Pastor
 449 et al., 2020; Pisano et al., 2020). The analogue quality analysis confirms that the 2025 event lies toward the
 450 upper end of analogue distances in both periods, indicating that the event is only partially captured even by
 451 its best historical analogues. This also illustrates the limits of this approach alone as attribution evidence.

a supprimé: states

a supprimé: changes.

a supprimé: anomalies make

452 Within these limits, the analogue analysis provides an event specific observational comparison of the
 453 environmental context associated with similar SST anomaly patterns in two parts of the historical record. It
 454 therefore complements the broader literature showing that Mediterranean MHWs are intensifying under
 455 climate change by showing, for this particular event, where and in which variables the past and present
 456 analogue contexts differ most strongly.

457 6. Conclusions

458 The Mediterranean Sea experienced one of its most intense early-summer MHWs on record in June 2025.
 459 Using an SST analogue framework, in which the target of the analogue search is the observed SST anomaly
 460 pattern itself, applied to ERA5 reanalysis data covering 1950–2024, we have compared the atmospheric and
 461 surface energy budget conditions associated with historical SST patterns appearing like the June 2025 event
 462 across two climate periods: a past period (1950–1986) and a present period (1987–2024). This distinguishes
 463 the study from the more commonly used circulation-based analogue approach and is better aligned with t
 464 marine nature of the event.

a supprimé: usual circulation first

465 Three main findings emerge from the analysis. First, the June 2025 event was exceptional in the observational
 466 record, with the strongest SST anomalies concentrated in the western Mediterranean, locally exceeding +4.5
 467 to 5.5°C, and with elevated temperatures extending over much of the basin.

Second, the analogue comparison reveals a clear thermodynamic fingerprint of long-term warming. Present-period analogues of the 2025 SST pattern are systematically associated with stronger SST anomalies, warmer near-surface air temperatures over surrounding land areas, and greater incoming shortwave radiation than past-period analogues. These differences are spatially coherent and statistically significant across broad portions of the basin. By contrast, differences in sea level pressure, wind speed, latent and sensible heat fluxes, and longwave radiation are weaker and spatially heterogeneous, indicating that the dominant shift between earlier and more recent climate backgrounds lies in the thermodynamic environment rather than in a change of the associated large-scale circulation.

Third, the analogue framework demonstrates both its utility and its limitations for marine heat extreme event attribution. It provides a physically transparent, observation-based, and event-specific line of evidence consistent with thermodynamic amplification of the 2025 MHW in the context of long-term Mediterranean warming. It cannot, however, fully disentangle contributions of the externally forced warming from internal variability, and it does not resolve ocean dynamical processes that may have contributed to the event's onset and persistence.

More generally, the study shows that SST analogue comparisons offer a valuable complement to model-based attribution for marine extremes. Their observational grounding, and event-specific objectives make them well-suited for rapid scientific assessments following major events. Future developments of this approach should explore its combination with large climate model ensembles to explicitly disentangle forced and internal contributions, and its extension to include subsurface ocean variables to better capture the role of ocean dynamics in MHW intensification.

Corresponding author

Correspondence to Davide Faranda: davide.faranda@cea.fr

Data availability

ERA5 reanalysis data used in this study are publicly available from the Copernicus Climate Data Store. The Mediterranean Sea High Resolution L4 Sea Surface Temperature Reprocessed product is available from the Copernicus Marine Service Information (<https://doi.org/10.48670/moi-00173>). All derived data supporting the findings of this study are available from the corresponding author upon reasonable request.

Code availability

The analogue analysis workflow builds on publicly available functions implemented in the ClimaMeter framework (Faranda et al., 2024). Custom scripts used in this study are available from the corresponding author upon reasonable request.

Author contributions

D.F. designed the study and coordinated the analysis. T.G. and F.P. provided data and analysis for observational datasets. T.G., F.P., and S.K. provided expertise on Mediterranean marine heatwaves and contributed to interpretation. T.A. contributed to methodological development. All authors discussed the results and contributed to writing and revising the manuscript.

Competing interests

The authors declare no competing interests.

Acknowledgments

This article is based on collaborative work within the COST Action CA22162 FutureMed, supported by COST, European Cooperation in Science and Technology. It was conducted as part of the PowDev project (Strategic development of the power grids of the future, ANR reference: 22-PETA-0016) and the ANR project Templex (ANR-23-CE56-0002), the ANR project TRACCS-PC4-EXTENDING (ANR-22-EXTR-0005). The work of Samira Khodayar and Francisco Pastor was funded by the research project EVER_PROMETEO (CIPROM/2022/37) funded by the Generalitat Valenciana.

References

- Adloff, F., Somot, S., Sevault, F., Jordà, G., Aznar, R., Déqué, M., ... & Gomis, D. (2015). Mediterranean Sea response to climate change in an ensemble of twenty first century scenarios. *Climate Dynamics*, 45(9), 2775-2802.
- Alberti, T., Faranda, D., Rapella, L., Coppola, E., Lepreti, F., Dubrulle, B., & Carbone, V. (2024). Impacts of changing atmospheric circulation patterns on aviation turbulence over Europe. *Geophysical Research Letters*, 51, e2024GL111618. <https://doi.org/10.1029/2024GL111618>
- Alexander, M. A., Bladé, I., Newman, M., Lanzante, J. R., Lau, N. C., & Scott, J. D. (2002). The atmospheric bridge: The influence of ENSO teleconnections on air sea interaction over the global oceans. *Journal of Climate*, 15(16), 2205-2231.
- Artana, C., Kaplan, A., Ramírez, F., Ortega, M., Steenbeek, J. G., & Coll, M. (2026). Marine heatwaves are transforming Western mediterranean marine ecosystems. *Scientific Reports*.
- Banzon, V., Smith, T. M., Steele, M., Huang, B., & Zhang, H. M. (2016). Improved estimation of proxy sea surface temperature in the Arctic. *Journal of Atmospheric and Oceanic Technology*, 33(11), 2493-2499.
- Behr, L., Xoplaki, E., Luther, N., Josey, S. A., & Luterbacher, J. (2025). Atmospheric patterns drive marine heatwaves in the North Atlantic and Mediterranean Sea during summer 2023. *Environmental Research Letters*, 20(10), 104020.
- Bonino, G., McAdam, R., Athanasiadis, P., Cavicchia, L., Rodrigues, R. R., Scoccimarro, E., ... & Masina, S. (2025). Mediterranean summer marine heatwaves triggered by weaker winds under subtropical ridges. *Nature Geoscience*, 18, 1-6.
- Capotondi, A., Rodrigues, R. R., Sen Gupta, A., Benthuisen, J. A., Deser, C., Frölicher, T. L., ... & Wang, C. (2024). A global overview of marine heatwaves in a changing climate. *Communications Earth & Environment*, 5(1), 701.
- Chen, X., & Tung, K. K. (2014). Varying planetary heat sink led to global warming slowdown and acceleration. *Science*, 345(6199), 897-903. <https://doi.org/10.1126/science.1254937>
- Coll, M., Piroddi, C., Steenbeek, J., Kaschner, K., Ben Rais Lasram, F., Aguzzi, J., ... & Voultsiadou, E. (2010). The biodiversity of the Mediterranean Sea: estimates, patterns, and threats. *PloS one*, 5(8), e11842.

- Cramer, W., Guiot, J., Fader, M., Garrabou, J., Gattuso, J. P., Iglesias, A., ... & Xoplaki, E. (2018). Climate change and interconnected risks to sustainable development in the Mediterranean. *Nature Climate Change*, 8(11), 972-980.
- Darmaraki, S., Somot, S., Sevault, F., Nabat, P., Cabos Narvaez, W. D., Cavicchia, L., ... & Sein, D. V. (2019). Future evolution of marine heatwaves in the Mediterranean Sea. *Climate Dynamics*, 53(3), 1371-1392.
- Darmaraki, S., Denaxa, D., Theodorou, I., Livanou, E., Rigatou, D., Raitsos, D., ... & Parasyris, A. (2024). Marine heatwaves in the Mediterranean Sea: A literature review. *Mediterranean Marine Science*, 25(3), 586-620.
- D'Andrea, F., Duvel, J. P., Rivière, G., Vautard, R., Cassou, C., Cattiaux, J., ... & Yiou, P. (2024). Summer deep depressions increase over the Eastern North Atlantic. *Geophysical Research Letters*, 51(5), e2023GL104435.
- Deser, C., Lehner, F., Rodgers, K. B., Ault, T., Delworth, T. L., DiNezio, P. N., ... & Ting, M. (2020). Insights from Earth system model initial-condition large ensembles and future prospects. *Nature climate change*, 10(4), 277-286.
- Deser, C., Phillips, A., Bourdette, V., & Teng, H. (2012). Uncertainty in climate change projections: the role of internal variability. *Climate Dynamics*, 38(3), 527-546. <https://doi.org/10.1007/s00382-010-0977-x>
- d'Ortenzio, F., Iudicone, D., de Boyer Montégut, C., Testor, P., Antoine, D., Marullo, S., ... & Madec, G. (2005). Seasonal variability of the mixed layer depth in the Mediterranean Sea as derived from in situ profiles. *Geophysical Research Letters*, 32(12), L12605.
- Dong, T., Zeng, Z., Pan, M., Wang, D., Chen, Y., Liang, L., ... & Zhang, D. (2025). Record-breaking 2023 marine heatwaves. *Science*, 389(6758), 369-374.
- Donlon, C. J., Martin, M., Stark, J., Roberts-Jones, J., Fiedler, E., & Wimmer, W. (2012). The operational sea surface temperature and sea ice analysis (OSTIA) system. *Remote Sensing of Environment*, 116, 140-158.
- England, M. H., Li, Z., Huguenin, M. F., Kiss, A. E., Sen Gupta, A., Holmes, R. M., & Rahmstorf, S. (2025). Drivers of the extreme North Atlantic marine heatwave during 2023. *Nature*, 639, 1-8.
- Estaque, T., Richaume, J., Bianchimani, O., Schull, Q., Méricot, B., Bensoussan, N., ... & Garrabou, J. (2023). Marine heatwaves on the rise: One of the strongest ever observed mass mortality event in temperate gorgonians.
- Faranda, D., Messori, G., & Yiou, P. (2020). Diagnosing concurrent drivers of weather extremes: Application to warm and cold days in North America. *Climate Dynamics*, 54(3-4), 2187-2201. <https://doi.org/10.1007/s00382-019-05106-3>
- Faranda, D., Bourdin, S., Ginesta, M., Krouma, M., Noyelle, R., Pons, F., Yiou, P., & Messori, G. (2022). A climate change attribution retrospective of some impactful weather extremes of 2021. *Weather and Climate Dynamics*, 3(4), 1311-1340. <https://doi.org/10.5194/wcd-3-1311-2022>

- Faranda, D., Messori, G., Jézéquel, A., Vrac, M., & Yiou, P. (2023). Atmospheric circulation compounds anthropogenic warming and impacts of climate extremes in Europe. *Proceedings of the National Academy of Sciences*, 120(3), e2214525120. <https://doi.org/10.1073/pnas.2214525120>
- Faranda, D., Coppola, E., Alberti, T., Vrac, M., Pons, F., Yiou, P., et al. (2024). ClimaMeter: contextualizing extreme weather in a changing climate. *Weather and Climate Dynamics*, 5(3), 959-983. <https://doi.org/10.5194/wcd-5-959-2024>
- Forster, P. M., Smith, C., Walsh, T., Lamb, W. F., Lamboll, R., Cassou, C., ... & Zhai, P. (2025). Indicators of global climate change 2024: Annual update of key indicators of the state of the climate system and human influence. *Earth System Science Data Discussions*, 2025, 1-72.
- Frölicher, T. L., Fischer, E. M., & Gruber, N. (2018). Marine heatwaves under global warming. *Nature*, 560(7718), 360-364.
- Garrabou, J., Gómez-Gras, D., Medrano, A., Cerrano, C., Ponti, M., Schlegel, R., ... & Harmelin, J. G. (2022). Marine heatwaves drive recurrent mass mortalities in the Mediterranean Sea. *Global Change Biology*, 28(19), 5708-5725.
- Giorgi, F. (2006), Climate change hot-spots, *Geophys. Res. Lett.*, 33, L08707, doi:10.1029/2006GL025734.
- Goessling, H. F., Rackow, T., & Jung, T. (2025). Recent global temperature surge intensified by record-low planetary albedo. *Science*, 387(6729), 68-73.
- Guinaldo, T., Voldoire, A., Waldman, R., Saux Picart, S., & Roquet, H. (2023). Response of the sea surface temperature to heatwaves during the France 2022 meteorological summer. *Ocean Science*, 19(3), 629-647.
- Guinaldo, T., Cassou, C., Sallée, J. B., & Liné, A. (2025). Internal variability effect doped by climate change drove the 2023 marine heat extreme in the North Atlantic. *Communications Earth & Environment*, 6(1), 291.
- Gyuleva, G., Knutti, R., & Sippel, S. (2025). Combination of internal variability and forced response reconciles observed 2023-2024 warming. *Geophysical Research Letters*, 52(14), e2025GL115270.
- Hersbach, H., Bell, B., Berrisford, P., et al. (2020). The ERA5 global reanalysis. *Quarterly Journal of the Royal Meteorological Society*, 146(730), 1999-2049. <https://doi.org/10.1002/qj.3803>
- Hobday, A. J., Alexander, L. V., Perkins, S. E., et al. (2016). A hierarchical approach to defining marine heatwaves. *Progress in Oceanography*, 141, 227-238. <https://doi.org/10.1016/j.pocean.2015.12.014>
- Hobday, A. J., Oliver, E. C. J., Sen Gupta, A., Benthuyssen, J. A., Burrows, M. T., Donat, M. G., ... & Smale, D. A. (2018). Categorizing and naming marine heatwaves. *Oceanography*, 31(2), 162-173. <https://doi.org/10.5670/oceanog.2018.205>
- Holbrook, N. J., Scannell, H. A., Sen Gupta, A., et al. (2019). A global assessment of marine heatwaves and their drivers. *Nature Communications*, 10(1), 2624. <https://doi.org/10.1038/s41467-019-10206-z>
- Holbrook, N. J., Sen Gupta, A., Oliver, E. C. J., Hobday, A. J., Benthuyssen, J. A., Scannell, H. A., ... & Wernberg, T. (2020). Keeping pace with marine heatwaves. *Nature Reviews Earth & Environment*, 1(9), 482-493.

- Juza, M., Fernández-Mora, A., & Tintoré, J. (2022). Sub regional marine heat waves in the Mediterranean Sea from observations: Long term surface changes, sub surface and coastal responses. *Frontiers in Marine Science*, 9, 785771.
- Laufkötter, C., Zscheischler, J., & Frölicher, T. L. (2020). High impact marine heatwaves attributable to human induced global warming. *Science*, 369(6511), 1621-1625. <https://doi.org/10.1126/science.aba0690>
- Lionello, P., & Scarascia, L. (2018). The relation between climate change in the Mediterranean region and global warming. *Regional Environmental Change*, 18(5), 1481-1493. <https://doi.org/10.1007/s10113-018-1290-1>
- Marullo, S., Artale, V., & Santoleri, R. (2011). The SST multidecadal variability in the Atlantic Mediterranean region and its relation to AMO. *Journal of Climate*, 24(16), 4385-4401.
- Martínez, J., Leonelli, F. E., García-Ladona, E., Garrabou, J., Kersting, D. K., Bensoussan, N., & Pisano, A. (2023). Evolution of marine heatwaves in warming seas: The Mediterranean Sea case study. *Frontiers in Marine Science*, 10, 1193164.
- Masson-Delmotte, V., Zhai, P., Pirani, A., Connors, S. L., Péan, C., Berger, S., ... & Zhou, B. (2021). Climate change 2021: The physical science basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change.
- Naveau, P., Hannart, A., & Ribes, A. (2020). Statistical methods for extreme event attribution in climate science. *Annual Review of Statistics and Its Application*, 7(1), 89-110.
- Oliver, E. C. J., Donat, M. G., Burrows, M. T., et al. (2018). Longer and more frequent marine heatwaves over the past century. *Nature Communications*, 9(1), 1324. <https://doi.org/10.1038/s41467-018-03732-9>
- Otto, F. E. L. (2016). The art of attribution. *Nature Climate Change*, 6(4), 342-343. <https://doi.org/10.1038/nclimate2971>
- Otto, F. E. (2023). Attribution of extreme events to climate change. *Annual Review of Environment and Resources*, 48(1), 813-828.
- Paredes-Fortuny, L., Pastor, F., & Khodayar, S. (2025). Concurrent atmospheric heatwaves intensify marine heatwaves through air-sea heat flux change in the Mediterranean Sea. *Communications Earth & Environment*, 6, 638. <https://doi.org/10.1038/s43247-025-02633-2>
- Pastor, F., Valiente, J. A., & Khodayar, S. (2020). A warming Mediterranean: 38 years of increasing sea surface temperature. *Remote Sensing*, 12, 2687.
- Pastor, F., & Khodayar, S. (2023). Marine heat waves: Characterizing a major climate impact in the Mediterranean. *Science of the Total Environment*, 861, 160621.
- Pastor, F., Paredes-Fortuny, L., & Khodayar, S. (2024). Mediterranean marine heatwaves intensify in the presence of concurrent atmospheric heatwaves. *Communications Earth & Environment*, 5(1), 797. <https://doi.org/10.1038/s43247-024-01982-8>
- Pisano, A., Marullo, S., Artale, V., Falcini, F., Yang, C., Leonelli, F. E., ... & Buongiorno Nardelli, B. (2020). New evidence of Mediterranean climate change and variability from sea surface temperature observations. *Remote Sensing*, 12(1), 132.

- Sen Gupta, A., Thomsen, M., Benthuisen, J. A., Hobday, A. J., Oliver, E. C. J., Alexander, L. V., ... & Wernberg, T. (2020). Drivers and impacts of the most extreme marine heatwave events. *Scientific Reports*, 10, 19359. <https://doi.org/10.1038/s41598-020-75445-3>
- Seneviratne, S. I., Zhang, X., Adnan, M., et al. (2021). Weather and climate extreme events in a changing climate. In *Climate Change 2021: The physical science basis (IPCC AR6 WGI, Chapter 11)*. Cambridge University Press. <https://doi.org/10.1017/9781009157896.013>
- Shepherd, T. G. (2014). Atmospheric circulation as a source of uncertainty in climate change projections. *Nature Geoscience*, 7(10), 703-708. <https://doi.org/10.1038/ngeo2253>
- Shepherd, T. G. (2016). A common framework for approaches to extreme event attribution. *Current Climate Change Reports*, 2(1), 28-38. <https://doi.org/10.1007/s40641-016-0033-y>
- Smale, D. A., Wernberg, T., Oliver, E. C. J., et al. (2019). Marine heatwaves threaten global biodiversity and the provision of ecosystem services. *Nature Climate Change*, 9(4), 306-312. <https://doi.org/10.1038/s41558-019-0412-1>
- Smith, K. E., Burrows, M. T., Hobday, A. J., Sen Gupta, A., Moore, P. J., Thomsen, M., ... & Smale, D. A. (2021). Socioeconomic impacts of marine heatwaves: Global issues and opportunities. *Science*, 374(6566), eabj3593.
- Stott, P. A., Christidis, N., Otto, F. E. L., et al. (2016). Attribution of extreme weather and climate related events. *Wiley Interdisciplinary Reviews: Climate Change*, 7(1), 23-41. <https://doi.org/10.1002/wcc.380>
- Sutton, R. T., & Hodson, D. L. R. (2005). Atlantic Ocean forcing of North American and European summer climate. *Science*, 309(5731), 115-118.
- Trenberth, K. E., Fasullo, J. T., & Shepherd, T. G. (2015). Attribution of climate extreme events. *Nature climate change*, 5(8), 725-730.
- Tuel, A., & Eltahir, E. A. B. (2020). Why is the Mediterranean a climate change hot spot? *Journal of Climate*, 33, 5829-5843.
- van Garderen, L., Feser, F., & Shepherd, T. G. (2021). A methodology for attributing the role of climate change in extreme events: A global spectrally nudged storyline. *Natural Hazards and Earth System Sciences*, 21(1), 171-186. <https://doi.org/10.5194/nhess-21-171-2021>
- Vautard, R., Yiou, P., Otto, F., Stott, P., Christidis, N., van Oldenborgh, G. J., & Schaller, N. (2016). Attribution of human induced dynamical and thermodynamical contributions in extreme weather events. *Environmental Research Letters*, 11(11), 114009.
- Wernberg, T., Thomsen, M. S., Burrows, M. T., Filbee-Dexter, K., Hobday, A. J., Holbrook, N. J., ... & Smith, K. (2025). Marine heatwaves as hot spots of climate change and impacts on biodiversity and ecosystem services. *Nature Reviews Biodiversity*, 1-19.
- Wilks, D. S. (2016). The stippling shows statistically significant grid points: How research results are routinely overstated and overinterpreted, and what to do about it. *Bulletin of the American Meteorological Society*, 97(12), 2263-2273.
- Xu, T., Newman, M., Capotondi, A., Stevenson, S., Di Lorenzo, E., & Alexander, M. A. (2022). An increase in marine heatwaves without significant changes in surface ocean temperature variability. *Nature Communications*, 13(1), 7396.

- Yiou, P. (2014). AnaWEGE: A weather generator based on analogues of atmospheric circulation. Geoscientific Model Development, 7(2), 531-543. <https://doi.org/10.5194/gmd-7-531-2014>