



Emission characteristics of greenhouse gases and air pollutants in a Qinghai-Tibetan Plateau city using a portable Fourier transform spectrometer and TROPOMI observations

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Abstract

Despite the critical need to understand greenhouse gas and air pollutant concentrations and their emissions characteristics in urban and industrial areas, limited assessments have been conducted in the Qinghai-Tibetan Plateau (QTP) cities. Herein, for the first time, we present CO₂, CH₄ and CO column abundances using a portable Fourier-transform infrared spectrometer (EM27/SUN) in Ganhe Industrial Park (36.546°N, 101.518°E, 2603 m a.s.l.), located in the suburbs of Xining, Qinghai Province, during May – June 2024. Ground-based measurements found to be higher than spaceborne measurements (TROPOMI and IASI) and model forecast (CAMS) across all investigated species, indicating higher local emissions. Notably, significant discrepancies in CO levels are observed, particularly under easterly wind conditions, which transport polluted airmasses from Xining city. To further quantify emissions, we applied a simple dispersion model to the EM27/SUN data and TROPOMI products, estimating an average CO emission rate of 12.3 ± 9.6 kg/s and 8.9 ± 7.5 kg/s, respectively. A wind-assigned anomaly method further applied to the TROPOMI dataset yielded a CO emission rate of 8.5 kg/s. Additionally, the ground-based observations of $\Delta\text{XCO}/\Delta\text{XCO}_2$ ratio exhibits a strong correlation under easterly winds, which suggests an average CO₂ emission rate of 550 kg/s from Xining city. These findings underscore the utility of portable FTIR spectrometers to enhance our understanding of urban emissions at QTP.

1. Introduction

Carbon dioxide (CO₂) and methane (CH₄) are two primary greenhouse gases (GHGs) whose atmospheric concentrations have surged to unprecedented levels since 1750, largely driven by anthropogenic activities. According to the WMO's annual Greenhouse Gas Bulletin (<https://library.wmo.int/idurl/4/69057>, last access: January, 2024), the globally averaged surface concentration of CO₂ reached 420 parts per million (ppm), and CH₄ reached 1934 parts per billion (ppb) in 2023. Carbon monoxide (CO) is one of the most important atmospheric pollutants, which is primarily produced by inefficient combustion, such as biomass burning (Griffin et al., 2024), traffic and industrial activity (Dils et al., 2011). CO plays a crucial role in atmospheric chemistry, especially in the troposphere, where it reacts with the



hydroxyl radical (OH) in the reaction $\text{CO} + \cdot\text{OH} \rightarrow \text{CO}_2 + \cdot\text{H}$ (Spivakovsky et al., 2000; Thompson, 1992). This oxidation process serves as the dominant sink for atmospheric CO, accounting for 90-95% of its total removal. Consequently, CO indirectly affects the lifetime of CH₄ by reducing the availability of OH that would otherwise oxidize CH₄. This, in turn, indirectly contributes to global warming (IPCC, 2007).

To achieve highly accurate and precise measurements of total column abundances of CO₂, CH₄ and CO, the Total Carbon Column Observing Network (TCCON) and the Collaborative Carbon Column Observing Network (COCCON) have been established. Both networks utilize Fourier-transform infrared (FTIR) spectrometers. TCCON employs high-resolution FTIR instruments for exceptionally precise CO₂, CH₄ and CO measurements amongst other gases (Wunch et al., 2011). In contrast, COCCON employs portable, low-resolution FTIR spectrometers to measure CO₂, CH₄ and CO, providing a valuable complementary extension to the TCCON network (Alberti et al., 2022; Frey et al., 2015; Herkommer et al., 2024).

The Qinghai-Tibetan Plateau (QTP), with its unique topography, plays a crucial role in Earth's climate system and has become a key region for monitoring climate trends and global air quality (Kang et al., 2021; Zhang et al., 2022). A recent study highlights that surface pollutants in Asia can be transported into the stratosphere from the Tibetan Plateau region, potentially exerting significant impacts on global climate (Bian et al., 2020). There are only two stations on the Tibetan Plateau—Waliguan, the only global station in Eurasia, and Shangri-La—that measure surface concentrations of CO₂, CH₄, and CO, making them important for understanding atmospheric composition in this critical region (Guo et al., 2020; Xiong et al., 2022). These stations conduct in-situ measurements and provide highly accurate surface observations; however, they are influenced by surface exchanges and vertical transport, which can limit their ability to estimate sources and sinks. Additionally, surface observations are less effective for satellite validation compared to column-based measurements, such as those obtained from TCCON and COCCON. However, there are currently no TCCON or COCCON stations on the QTP. Zhou et al. (2023) conducted the first FTIR column observations in a small city on QTP, but measurements in larger cities, such as Xining—the capital of Qinghai Province, remain sparse. This is notable since most anthropogenic emissions of GHGs and CO are concentrated in urban centers (Crippa et al., 2021).

This study presents the EM27/SUN FTIR remote sensing observations conducted in Ganhe Industrial Park (36.546°N, 101.518°E, 2603 m a.s.l.), located in the southwest suburbs of Xining, Qinghai Province during May-June 2024. It includes a comparison to the co-located TROPOMI and IASI observations and CAMS forecasts. Additionally, a simple dispersion model is introduced and applied to estimate the CO and CO₂ emissions.

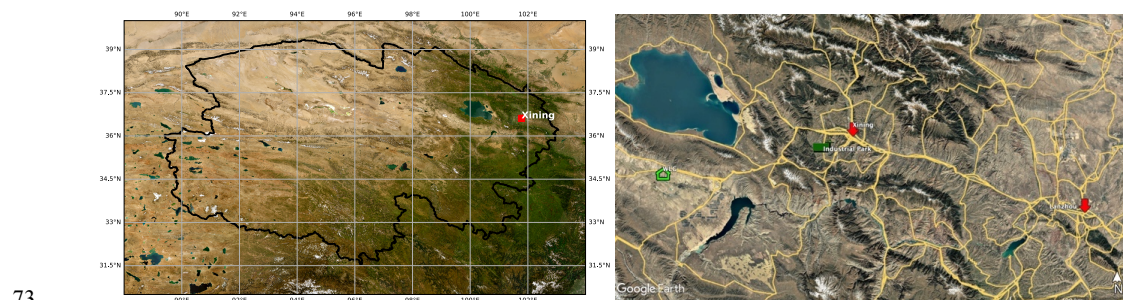
2. Methods and materials

2.1 Site description

A three-week field campaign was conducted in Ganhe Industrial Park (36.546°N, 101.518°E, 2603 m a.s.l.) from May 23 to June 14, 2024. The industrial park is located southwest of Xining, the capital of Qinghai Province, China (Figure 1) and was established in 2002 in the Huangzhong District. Xining, situated on the eastern edge of the Tibetan Plateau



70 and upstream of the Huangshui River, experiences a plateau mountain climate. In 2020, the city has a resident
71 population of approximately 2.47 million, with a density of 324 persons/km² (Xining Statistical Bureau, 2021), making
72 it the most densely populated area in Qinghai Province with nearly 80% of its population living in the urban center.



73
74 **Figure 1: A terrain map showing Qinghai Province and the location of its capital city of Xining (Terrain information**
75 **originates from World Imagery). A map illustrating the locations of the EM27/SUN instrument within the Ganhe Industrial**
76 **Park, Xining city (Qinghai province), Lanzhou city (Gansu province), and Waliguan station. The base map is sourced from**
77 **© Google Earth, Image © 2024 Maxar Technologies; Image © 2024 CNES / Airbus.**

78 2.2 COCCON GHG products

79 In this study, a portable ground-based Fourier-transform infrared (FTIR) spectrometer, the EM27/SUN, was used.
80 Developed by the Karlsruhe Institute of Technology (KIT) in collaboration with Bruker Optics GmbH, Ettlingen,
81 Germany, the EM27/SUN is designed to measure solar absorption spectra in the near-infrared spectral range (covering
82 5500–11,000 cm⁻¹ and 4000–5500 cm⁻¹) with a spectral resolution of 0.5 cm⁻¹ (Gisi et al., 2012). The spectrometer
83 records double-sided, DC-coupled interferograms using two indium gallium arsenide (InGaAs) detectors at room
84 temperature (Hase et al., 2016). These interferograms are processed using a preprocessing tool and the PROFFAST
85 nonlinear least squares fitting algorithm, which was developed by KIT in the framework of the COCCON-
86 PROCEEDS project funded by the European Space Agency (ESA) (Alberti et al., 2022; Herkommer et al., 2024).
87 This approach enables the retrieval of atmospheric concentrations of trace gases such as O₂, CO₂, CH₄, CO and H₂O.
88 The gas column is then converted to column-averaged dry-air mole fractions of the gas (X_{gas}) by dividing the gas
89 column by the co-observed O₂ column, using the well-known mole fraction of O₂ in dry air.

90 With meanwhile about 200 EM27/SUN instruments deployed worldwide, these spectrometers are widely used for
91 monitoring GHG concentrations and air pollutant, for estimating emissions from various sources, and for validating
92 satellite measurements (Chen et al., 2020; Frey et al., 2015; Hase et al., 2015; Herkommer et al., 2024; Luther et al.,
93 2019; Tu et al., 2021, 2020). Due to its excellent level of robustness and reliability, the EM27/SUN instrument can be
94 applied in both field campaigns and long-term deployment at fixed sites. These observations complement TCCON at
95 various locations enhances global GHG monitoring efforts. In this context, the COCCON was established to further
96 advance and standardize these observational efforts.



97 **2.3 TROPOMI CH₄ and CO products**

98 The TROPospheric Monitoring Instrument (TROPOMI), single payload of the Sentinel-5 Precursor (S5P) satellite,
99 has been in orbit since October 2017. It operates in a low Earth polar orbit with a planned operational lifespan of 7
100 years. S5P is the first Copernicus mission and is designed to deliver daily global information on concentrations of
101 traces gases (e.g., CH₄ and CO) and aerosols, aiming to monitor air quality, climate forcing and ozone abundances
102 with high spatial and temporal resolution (Veefkind et al., 2012).

103 TROPOMI, currently the most advanced nadir-viewing and multispectral imaging spectrometer, was developed jointly
104 by the European Space Agency (ESA) and the Netherlands Space Office. It measures across several spectral bands:
105 ultraviolet (UV) and visible (VIS) (270–500 nm), near-infrared (NIR, 675–775 nm) and shortwave infrared (SWIR,
106 2305–2385 nm). With a wide swath of 2600 km across track, TROPOMI provides operational level 2 (L2) CH₄ and
107 CO products with a very high spatial resolution of approximately 5.5 km × 7 km since August 2019.

108 The retrieval of TROPOMI total column abundances of CH₄ is conducted using the RemoTec-S5P algorithm (Butz et
109 al., 2009; Hasekamp and Butz, 2008). For CO, the total column retrieval is performed by a modified SWIR CO
110 retrieval (SICOR) algorithm, which is based on the CO absorption band between 2305 nm and 2385 nm, with
111 interfering trace gases and effective cloud parameters (Landgraf et al., 2016).

112 In this study, the TROPOMI L2 CH₄ and CO data with quality assurance values (qa_value) greater than 0.5, as
113 recommend in the S5P product readme files, are utilized. Detailed documentation can be accessed for the CH₄
114 (<https://sentinel.esa.int/documents/247904/3541451/Sentinel-5P-Methane-ProductReadme-File>, last accessed: June
115 9, 2025) and CO data products ([https://sentinel.esa.int/documents/247904/3541451/Sentinel-5P-CarbonMonoxide-](https://sentinel.esa.int/documents/247904/3541451/Sentinel-5P-CarbonMonoxide-Level-2-Product-Readme-File)
116 Level-2-Product-Readme-File, last accessed: June 9, 2025).

117 **2.4 IASI CO products**

118 The Infrared Atmospheric Sounding Interferometer (IASI) is the primary payload carried on the EUMETSAT's
119 MetOp series of polar-orbiting satellites and delivers meteorological parameters (e.g., water vapour and atmospheric
120 temperature), and trace species with an unprecedented spatial and temporal coverage (Clerbaux et al., 2009). IASI
121 sensors are nadir looking thermal infrared sensors and there are currently three instruments in operation, which were
122 launched in 2006, 2012 and 2018.

123 The IASI CO dataset is processed using the Fast Optimal Retrievals on Layers for IASI (FORLI) software (George et
124 al., 2009; Hurtmans et al., 2012). The retrievals are performed within the 2143–2181.25 cm⁻¹ spectral range based on
125 the optimal estimation theory and tabulated absorption cross sections at various pressures and temperatures to enhance
126 the efficiency of the radiative transfer calculation. In this study, total column abundances of CO from IASI onboard
127 Metop-C Level 2 (version 6.7) are used.



128 **2.5 CAMS high-resolution GHG forecasts and CAMS-GLOB-ANT inventory**

129 The Copernicus Atmosphere Monitoring Service (CAMS), which is implemented by the European Centre for
130 Medium-Range Weather Forecasts (ECMWF), produces daily global forecasts for the two main long-lived GHGs. As
131 part of the CAMS GHG services, the ECMWF Integrated Forecasting System (IFS) delivers 5-day high-resolution
132 forecasts of CO₂, CH₄ (Agustí-Panareda et al., 2014, 2017, 2019), as well as CO and meteorological parameters
133 essential for GHG forecasting (Flemming et al., 2015). The forecast is generated a few hours behind real time, with
134 initial conditions derived from a 4-day forecast of the analysis experiment. It is run at a horizontal resolution TCo1279,
135 corresponding to a cubic octahedral reduced Gaussian grid with an approximately spatial resolution of 9 km (Holm et
136 al., 2016). The model includes 137 vertical levels extending from the surface to 0.01 hPa (Agustí-Panareda et al.,
137 2019).

138 In this study, the forecasting suite is based on the IFS model cycle CY48R1, which was upgraded in June 2023. This
139 update introduced several system enhancements, particularly in composition modelling, emissions and assimilation
140 (Eskes et al., 2024). Notably, CY48R1 includes the assimilation of TROPOMI CO and the performance has generally
141 improved compared to all observations, such as surface observations, vertical profiles from IAGOS aircraft and
142 NDACC FTIR measurements, and satellite total column retrievals (Eskes et al., 2024).

143 CAMS also delivers global anthropogenic emissions, referred to as CAMS-GLOB-ANT. This includes emissions of
144 both air pollutants (e.g., CO) and greenhouse gases (e.g., CO₂ and CH₄) for real-time forecasts (Soulie et al., 2024).
145 Emissions are provided for 17 sectors and 35 species as monthly averages, with a spatial resolution of 0.1° × 0.1°,
146 covering the period from 2000 to 2025. This study uses the latest version v6.2, which is based on the Emissions
147 Database for Global Atmospheric Research (EDGARv6) and employs the same methodology as v5.3.

148 **3. Results and discussion**

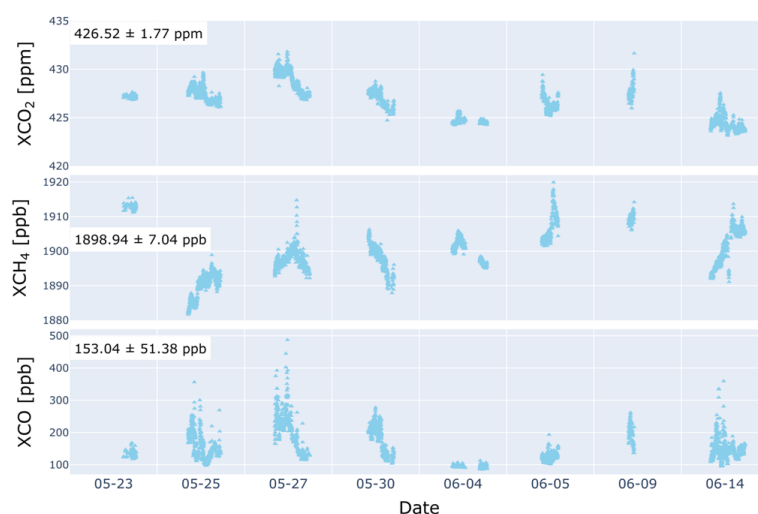
149 **3.1 Ground-based observations**

150 Figure 2 illustrates the time-series of XCO₂, XCH₄ and XCO observed by the EM27/SUN over 8 intermittent days.
151 Among these gases, XCO exhibits the most significant intraday variability, particularly on May 25, 27 and June 14.
152 During these days, prevailing easterly winds transported airmasses from Xining City and the nearby highway (Xiong
153 et al., 2022), leading to the elevated XCO levels. In contrast, on June 4, when airmasses primarily originated from the
154 west and northwest—regions with less urbanized areas—XCO levels were notably stable and low, averaging $93.93 \pm$
155 4.23 ppb (see back trajectories on Figure A 2). Additionally, higher wind speed on June 4 facilitated the dispersion of
156 CO, contributing to the lower CO level.

157 Similarly, XCO₂ on June 4 showed relatively less intraday variability and a lower average concentration of $424.64 \pm$
158 0.25 ppm, except for a slight enhancement (~ 1 ppm) in the morning, which was also observed in XCH₄ (~ 6 ppb). This
159 suggests different CO sources in the western areas relative to the EM27/SUN location.



160 The observed average XCH₄ concentration during the measurement period was 1899.947 ± 7.04 ppb. This relatively
161 low variability suggests an absence of significant local methane sources, such as coal mining activities (Tu et al.,
162 2024a), which are known to emit substantial amounts of methane.



163

164 **Figure 2: time-series of EM27/SUN observations from May 23 to June 14 2024.**

165 3.2 Comparison of COCCON data with TROPOMI and IASI products

166 Figure 3a-b illustrates the correlations between TROPOMI and COCCON measurements on different days. The spatial
167 collocation criterion requires the TROPOMI observations to fall within a radius of 200 km for XCH₄ and 100 km for
168 XCO. The temporal collocation criterion is set to ± 2 h for COCCON measurements.

169 The results indicate that TROPOMI tends to slightly overestimate XCH₄ by an average bias of -2.84 ± 12.74 ppb,
170 while it significantly underestimates of XCO, showing a bias of 34.31 ± 24.45 ppb. The correlation between
171 TROPOMI and COCCON XCH₄ measurements improves as the spatial distance between their respective locations
172 decreases. The largest observed difference in XCH₄, approximately -20.66 ppb, occurs on June 14, when the minimal
173 distance between TROPOMI and COCCON locations is around 150 km (Figure A 3a).

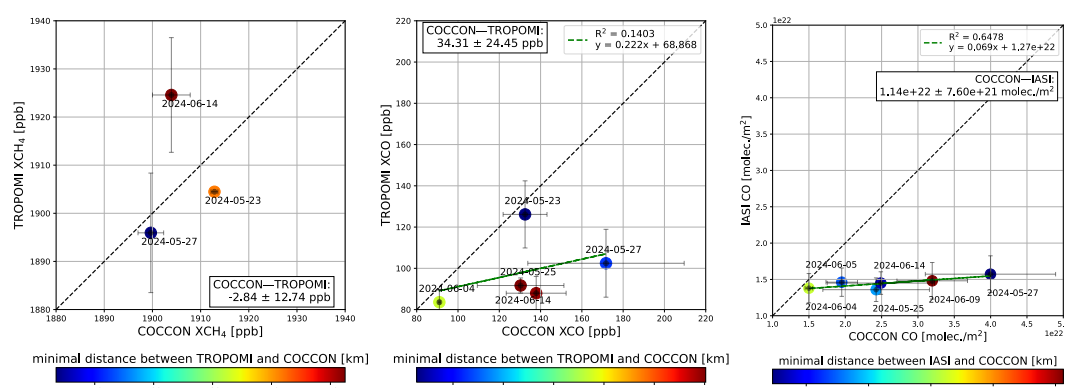
174 A similar trend is observed for XCO, where larger biases are generally associated with greater distances. However,
175 on May 27, the largest bias (69.23 ppb) occurs despite a relatively small minimum distance of 13 km. This significant
176 discrepancy is likely influenced by the spatial distribution of TROPOMI observations relative to the EM27/SUN
177 location. Notably, on May 27, TROPOMI XCO levels exhibit a negative correlation with distance for observations
178 within approximately 100 km (Figure A 3b). The closest observation to the EM7/SUN location records a TROPOMI
179 XCO value of 157.20 ppb, resulting in a significantly reduced bias of 16.72 ppb.

180 For May 23, four TROPOMI XCO observations, with distances ranging from 2.8 km to 10.4 km, result in a mean
181 value of 132.43 ppb. These closer observations show an overestimation in XCO, with a difference of 8 ppb compared



182 to the EM27/SUN measurements. However, this discrepancy could be influenced by the fact that EM27/SUN
183 observations were available only one hour after the TROPOMI overpass, potentially introducing temporal variations
184 into the comparison.

185 The correlation between IASI and COCCON is relatively higher ($R^2 = 0.6478$) compared to the correlation between
186 TROPOMI and COCCON XCO. Similar to TROPOMI, IASI shows the highest bias on May 27 with a value of
187 $2.43E22$ molec./m². Enhanced CO levels are observed within 20 km radius, with the peak CO concentration reaching
188 $2.66E22$ molec./m² at 15.4 km (Figure A 3c). As a result, the bias decreases by nearly half to $1.32E22$ molec./m².



189

190 **Figure 3: Correlation plot between TROPOMI and COCCON for XCH₄ (a) and XCO (b) and between IASI and COCCON**
191 **for CO column (c). The color bar represents the minimum distance between the TROPOMI and COCCON locations.**

192 3.3 Comparison of COCCON data with CAMS products

193 The CAMS GHG forecast provides a high spatial resolution of approximately 9 km. For comparison with the
194 EM27/SUN site, the following coincidence criteria were used: CAMS data within a 20 km radius of the EM27/SUN
195 location and EM27/SUN observations within ± 2 h around noon.

196 The CAMS forecast XCH₄ demonstrates a strong correlation ($R^2 = 0.7930$) with COCCON observations and an
197 average bias of 4.86 ± 3.81 ppb. The mean bias increases to 8.87 ppb when a larger coincidence radius of 50 km is
198 used for the CAMS dataset. The correlation slightly decreases, with the R^2 value dropping to 0.7877.

199 CAMS generally underestimates XCO₂ and the CO column, exhibiting lower correlations. The mean biases are 1.66
200 ± 1.50 ppm for XCO₂ and $1.01E22 \pm 6.31E21$ molec./m² for CO. The largest biases in XCO₂ and CO are recorded on
201 May 27. This discrepancy is also observed in TROPOMI and IASI CO observations.

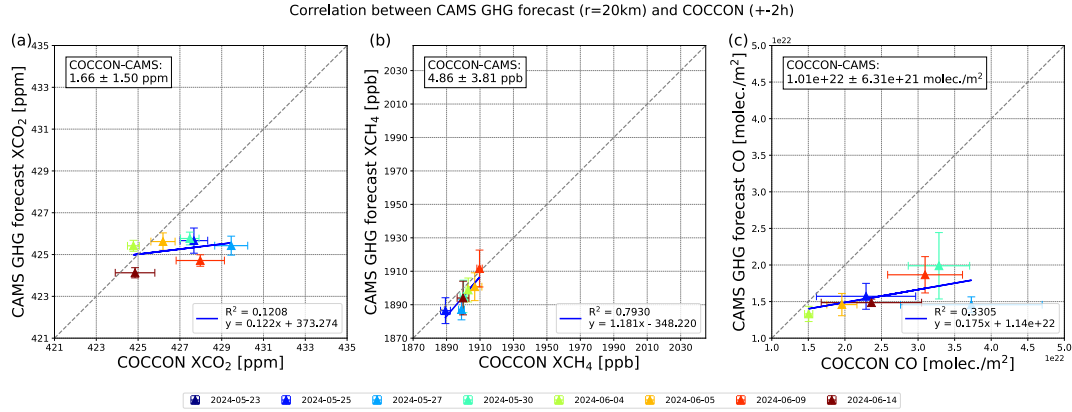


Figure 4: Correlation between CAMS forecast and COCCON for XCO₂ (a), XCH₄ (b) and CO column (c).

3.4 CO emission estimates based on ground-based observations

Satellite observations, such as TROPOMI and IASI, and CAMS forecasts, consistently underestimate CO levels during the entire field campaign. Thus, ground-based measurements are essential for accurately estimating CO emissions. Wind directions predominantly originated from the east (Figure A 1) on five days, indicating that emissions from the Xining city were being transported to the downwind site where the EM27/SUN located.

Figure 5 highlights the correlation between ΔXCO multiplied by wind speed as a function of wind direction. Wind directions within the range of $80^\circ - 120^\circ$, associated with relatively higher multiplication between ΔXCO and wind speed, were selected to represent the predominant winds transporting emissions from Xining. In this range, the average wind speed is approximately 2.3 m/s, with a standard deviation of 0.7 m/s. To account for variability, an uncertainty of $\pm 20^\circ$ is applied. The wind spreading angle ∂ is defined as:

$$\partial = 120^\circ - 80^\circ = 40^\circ (\pm 20^\circ) = 0.7 (\pm 0.35) \text{ rad}$$

Based on this wind spreading angle, the CO emission is calculated using the dispersion model as shown in the following equation (Babenhauserheide et al., 2020; Tu et al., 2022a):

$$\varepsilon = \Delta CO \times d \times v \times \partial \quad \text{Eq. 1}$$

where ΔCO represents the enhanced CO column observed at the EM27/SUN site, d is the average distance from Xining city to the EM27/SUN location (approximately 25 km) and v is the wind speed.

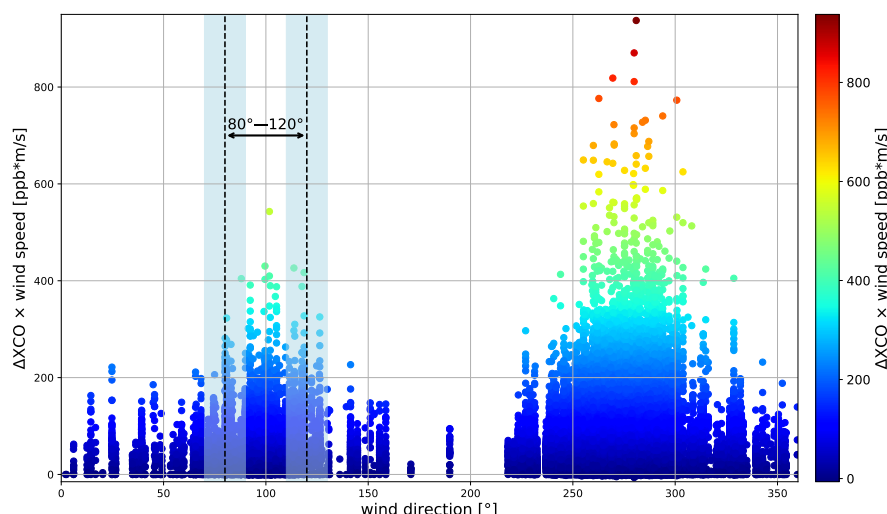
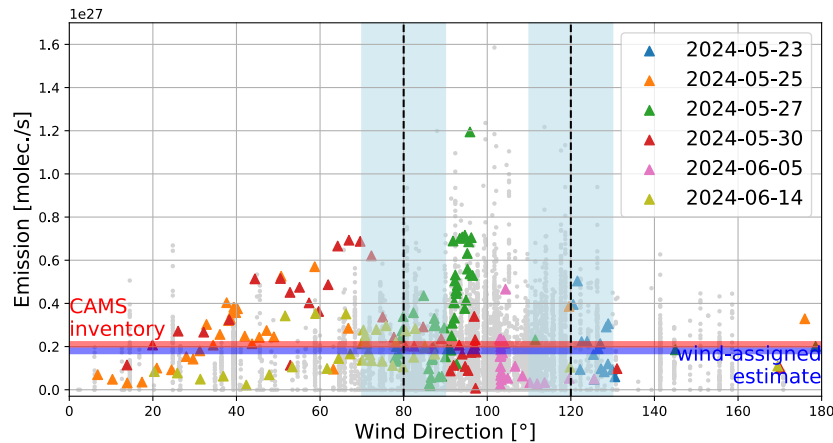


Figure 5: TROPOMI ΔXCO multiplied by wind speed against the wind direction for wind speed greater than 1.5 m/s, covering the period from May 2018 to June 2024. Wind data are derived from ERA5 at a height of 100m above ground level at the EM27/SUN station at 14:00 local time. The ΔXCO dataset represents XCO residuals after background subtraction. The wind direction range of 80°–120° (delineated by dashed lines) captures the predominant wind influence from Xining city. The shaded areas represent the uncertainty of 10°.

The daily background concentration of CO is defined as the lowest 10th percentile of measurements. To minimize accidental bias, a 10-minuten averaged EM27/SUN CO dataset was used. Significant enhancements are observed in ΔXCO and ΔXCO_2 in the 80°–120° wind direction range (Figure A 4). Using Equation (2), the estimated CO emissions for five days are presented in Figure 6. Notably, relatively high emissions were observed in the 80°–120° wind direction range, corresponding to areas of higher wind speed. A similar enhancement of CO associated with easterly winds was reported by Xiong et al. (2022), where in situ observations at the Waliguan station (36.28°N, 100.09°E, 3816 m a.s.l.), located approximately 155 km southwest of Xining city. This highlights the influence of regional air transport on the observed CO concentrations.

The CO emissions in the wind direction interval of 80°–120° are likely attributed to Xining city, with a mean estimated emission rate of $2.7E26 \pm 2.1E26$ molec./s (i.e., 12.3 ± 9.6 kg/s). The maximum emission rate reached $1.2E27$ molec./s (i.e., 55.6 kg/s) on May 27, a day when both satellite observations and forecasts underestimated the CO levels. The peak emission on May 27 occurred around noon, coinciding with a significant enhancement in XCO, which reached an approximate 10-minute averaged signal of 300 ppb.

Another notable high emission event was observed on May 25 and 30, under wind directions around 60°. These elevated emissions are likely to attributed to emission from the north-northeast region of the EM27/SUN location. On May 25, an enhanced XCO signal (~80 ppb) was observed in the late afternoon, coinciding with a relatively higher wind speed of ~2.8 m/s, which contributed to the increased emissions. On March 30, the XCO signal peaked at ~120 ppb around noon, leading to even higher emission rates.

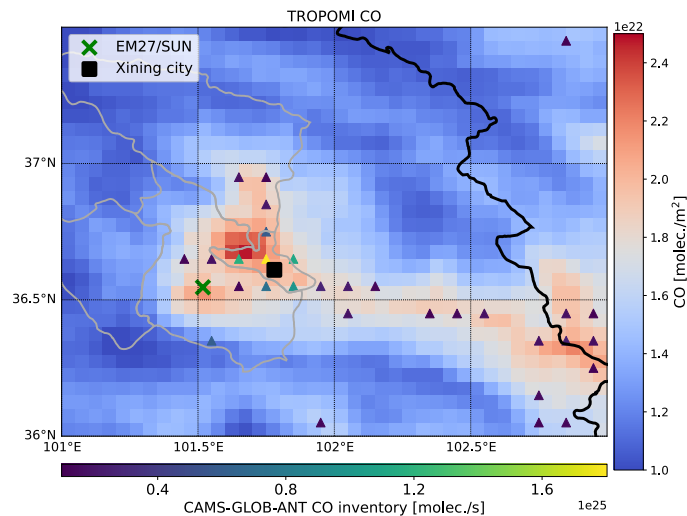


243

244 **Figure 6: Estimated emission relative to wind direction on six days of ground-based measurements (triangle symbols) when**
245 **winds were predominantly easterly. Grey dots represent emissions derived from TROPOMI XCO. Two horizontal lines**
246 **indicate CO emission rates derived from the wind-assigned anomaly method and the CAMS-GLOB-ANT inventory,**
247 **respectively.**

248 **3.5 CO emission estimates based on satellite-based observations**

249 TROPOMI observes relative high CO levels near the central region of Xining. A distinct streak of elevated CO
250 concentrations extends eastward from Xining toward the Qinghai province border and continues toward Lanzhou in
251 Gansu province (not shown in the figure). This pattern aligns with the location of densely populated residential regions
252 and coincides with a gradual decrease in altitude from west to east (Figure A 5).



253

254 **Figure 7: Spatial distribution of the average TROPOMI CO column at a latitude-longitude resolution of 0.05°, covering the**
255 **period from May 2018 to June 2024. The triangle symbols indicate locations where CO emissions exceed 1E24 molec./s**
256 **based on the CAMS-GLOB-ANT inventory. The green cross marks the location of the EM27/SUN and the black square**

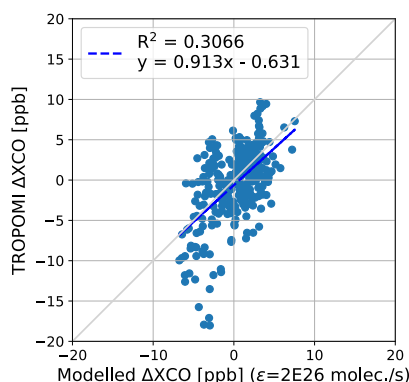


257 represents the center of Xining city. The thick black line outlines the Qinghai province boundary and the thin grey lines
258 delineate subregions within Xining.

259 Emissions derived from TROPOMI CO products, based on the Eq1. are presented in Figure 6 (grey dots), showing
260 relatively higher values for wind directions between 80° and 120°. The average emission rate is approximately $8.9 \pm$
261 7.5 kg/s, which closely matches the emissions derived from the COCCON measurements.

262 To further estimate regional CO emissions, a wind-assigned method in combination with the dispersion model
263 described in Eq.1 is applied to the TROPOMI observations. The wind-assigned method, previously used for estimating
264 regional CH₄ and NO₂ emissions (Tu et al., 2024b, 2023, 2022a, 2022b), by fitting the wind-assigned anomalies
265 between the satellite observations and the modeled anomaly based on the dispersion model. The wind-assigned
266 anomaly is defined as the difference in the enhancements under two opposite wind fields, e.g., E (>0° and <180°) and
267 W (180°-360°) fields. This approach eliminates the uncertainty associated with calculating the background for long-
268 lifetime gases (e.g., for CH₄), thus improving the reliability of emission estimates.

269 The estimated CO emission based on the wind-assigned anomaly method is approximately $1.8\text{E}26$ molec./s (8.5 kg/s),
270 which is close to the CAMS-GLOB-ANT inventory ($2.1\text{E}26$ molec./s, i.e., 9.8 kg/s) and the average emission derived
271 from the six-day EM27/SUN measurements ($2.7\text{E}26$ molec./s, i.e., 12.3 kg/s). However, scatters are observed in the
272 correlation between the TROPOMI and modeled wind-assigned anomalies, leading to a lower R^2 value (Figure 8).
273 This discrepancy could originate from uncertainties in the CAMS inventory, which may not account for some some-
274 high emission sources. Additionally, the simple dispersion model introduces its own uncertainties, which limits the
275 ability to accurately model CO enhancements.



276
277 **Figure 8: Correlation in the wind-assigned anomalies between the TROPOMI observations and modeled anomalies.**

278 **3.6 CO₂ estimates based on combustion efficiency**

279 In urban environments, CO is primarily produced through the incomplete combustion of fossil fuels, such as those
280 from traffic emissions and industrial activities, often occurring alongside the formation of CO₂. The CO to CO₂ ratio
281 serves as a valuable indicator for atmospheric fossil-fuel CO₂ emissions and combustion efficiency, providing insights
282 into the sources and effectiveness of fuel use (Che et al., 2022; Lee et al., 2024; Sim et al., 2020).



Figure 9 presents the relationship between ΔXCO and ΔXCO_2 under predominant wind of wind direction of $80^\circ - 120^\circ$ and wind speed larger than 1.5 m/s for different days. The daily background concentrations of XCO_2 and XCO were determined as the lowest 10th percentile of the respective observations. ΔXCO and ΔXCO_2 was calculated by subtracting the daily background concentrations from the daily observed values. Under easterly wind conditions, ΔXCO exhibits a stronger correlation with ΔXCO_2 , with a slope of 35.14 (ppb/ppm) and an R^2 value of 0.7552, indicating distinct source contributions and atmospheric transport processes from eastern regions, such as Xining city. The observed $\Delta\text{XCO}/\Delta\text{XCO}_2$ ratio is higher than those reported in other studies (Che et al., 2022; Lee et al., 2024). This discrepancy is attributed to significantly elevated CO levels under easterly wind conditions, suggesting lower combustion efficiency in this region.

Using CO as a proxy, the fossil fuel-derived CO_2 emissions can be computed as follows:

$$E_{\text{CO}_2} = \left(\frac{1}{ER_{\text{CO}/\text{CO}_2}} \times \frac{M_{\text{CO}_2}}{M_{\text{CO}}} \right) E_{\text{CO}} \quad \text{Eq. 2}$$

Where E_{CO_2} is the emissions of CO_2 in kg/s, $ER_{\text{CO}/\text{CO}_2}$ is the ratio of $\Delta\text{XCO}:\Delta\text{XCO}_2$ in ppb/ppm, E_{CO} is the CO emission rate derived based on Sec. 3.4, and M_{CO_2} and M_{CO} are the molar mass of CO_2 and CO , respectively.

We estimate an average CO_2 emission rate of approximately 550 kg/s. The maximum $\Delta\text{XCO}:\Delta\text{XCO}_2$ ratio was observed on May 27, with a slope of 40.08 ($R^2 = 0.8544$), corresponding to a maximum CO emission rate of 55.6 kg/s and resulting in a maximum CO_2 emission rate of 2180 kg/s. This estimate aligns reasonably with the Carbon Emission Accounts and Datasets (CEADs) inventory, with reports a CO_2 emission rate of 726 kg/s for Xining in 2015 (“Methodology and applications of city level CO_2 emission accounts in China,” 2017; Shan et al., 2018). The observed discrepancies may be attributed to differences in temporal coverage, methodological approaches, and potential changes in emission patterns over time. A longer period of ground-based observations and running several spectrometers upwind and downwind may improve our results. Our findings so far demonstrate the potential of the EM27/SUN spectrometer as a promising tool for comprehensively evaluating greenhouse gas (GHG) and air pollutant emissions in urban areas (Che et al., 2022; Lee et al., 2024).

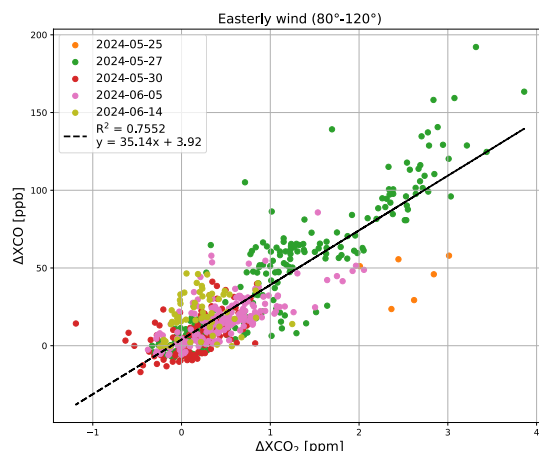


Figure 9: $\Delta\text{XCO}/\Delta\text{XCO}_2$ ratios under easterly wind conditions covering direction between 80° and 120° . Different colors indicate observations from individual days.

Conclusion

A three-week field campaign using a portal FTIR spectrometer (EM27/SUN) was conducted at Ganhe Industrial Park, located in the southwestern suburbs of Xining city, from May 23 to June 14, 2024. The mean and standard deviation values for XCO_2 , XCH_4 and XCO were 426.52 ± 1.77 ppm, 1898.94 ± 7.04 ppb and 153.04 ± 51.38 ppb, respectively. Among these gases, XCO exhibited significant intraday variability, particularly on days dominated by easterly winds, which transported air mass from Xining city.

Ground-based observations were compared with co-located datasets from TROPOMI, IASI and CAMS. Results indicate that TROPOMI slightly overestimates XCH_4 , with an average bias of -2.84 ± 12.74 ppb, but significantly underestimates XCO , showing a bias of 34.31 ± 24.45 ppb. IASI generally underestimates the CO column relative to COCCON observations, with an average bias of $1.14\text{E}22 \pm 7.60\text{E}21$ molec./m². CAMS forecasts also underestimated XCO_2 (1.66 ± 1.50 ppm), XCH_4 (4.86 ± 3.81 ppb) and CO column ($1.01\text{E}22 \pm 6.31\text{E}21$ molec./m²). A high correlation ($R^2=0.7930$) was found for XCH_4 between CAMS and EM27/SUN datasets. The largest discrepancies in CO products occurred on May 27, when significant enhancements were observed at the co-located time by EM27/SUN. These discrepancies are partly attributed to satellite observations, as their concentrations decreased with distance from the EM27/SUN location.

Wind directions predominantly originated from the east on five measurement days, indicating that emissions from the Xining city were being transported to the downwind site where the EM27/SUN spectrometer was located. A simple dispersion model, incorporating wind information and enhanced CO column data from EM27/SUN instrument, was applied to estimate CO emissions. The CO emissions from Xining city are estimated at 12.3 ± 9.6 kg/s with a peak emission rate of 55.6 kg/s on May 27, when both satellite observations and forecasts underestimated CO levels. When this model was applied to TROPOMI CO data, the resulting average emission rate was 8.9 ± 7.5 kg/s. A wind-assigned



329 anomaly method was also applied to the TROPOMI CO dataset, resulting in an estimate emission rate of 8.5 kg/s.
330 Both EM27/SUN-based and TROPOMI-based estimates are comparable to the CAMS inventory value of 9.8 kg/s.
331 CO can serve as a proxy for fossil fuel-derived CO₂ emissions. Ground-based observations of ΔXCO and ΔXCO_2 exhibit
332 a stronger correlation under wind direction of 80°–120°, with a slope of 35.14 ppb/ppm and an R^2 value of 0.7552,
333 compared to conditions under westerly winds. Using these correlations, we estimated the CO₂ emission rate to have
334 an average value of 550 kg/s with a maximum value of 2180 kg/s. These estimates align reasonably with the CEASs
335 inventory, which reports a CO₂ emission rate of 726 kg/s for Xining in 2015. Observed discrepancies may be attributed
336 to differences in temporal coverage, methodological approaches, and potential changes in emission patterns over time.
337 Note that long-term ground-based measurements of trace gases may improve the precision of estimated emissions.
338 Our findings highlight the potential of the EM27/SUN spectrometer as a valuable tool for detecting local emission
339 and supporting satellite validation, particularly in high-altitude regions such as Tibetan Plateau. It also demonstrates
340 the ability to efficiently measure both GHG and air pollutant emissions, offering a cost-efficient approach.

341



342 **Data availability.**

343 The TROPOMI CH₄ and CO dataset is a Copernicus operational product and is available at
344 <https://doi.org/10.5270/S5P-3lcdqiv> (Copernicus Sentinel-5P, 2021). The access and use of any Copernicus Sentinel
345 data available through the Copernicus Open Access Hub are governed by the legal notice on the use of Copernicus
346 Sentinel Data and Service Information, which is given at
347 https://sentinels.copernicus.eu/documents/247904/690755/Sentinel_Data_Legal_Notice (last access: January 2025).
348 IASI CO dataset is available at <https://iasi.aeris-data.fr/co/> (last access: January 2025). CAMS forecasts are
349 available at 10.24381/93910310 (last access: January 2025) and CAMS-GLOB-ANT inventory data are available at
350 <https://doi.org/10.24380/eets-qd81> (last access: January 2025). CEADS inventory is available at:
351 <https://www.ceads.net/data/city?#1280>. last access: January 2024.

352 **Author contributions.**

353 QT and FH conceived the research question. QT wrote the manuscript and conducted the data analysis with input from
354 FH. XL and ZY carried out the measurements, while YJ and JF was responsible for the data processing. All authors
355 contributed to the interpretation of the results and the improvement of the paper.

356 **Declaration of Competing Interest.**

357 The authors declare that they have no known competing financial interests or personal relationships that could have
358 appeared to influence the work reported in this paper.

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365



366 **Appendix**

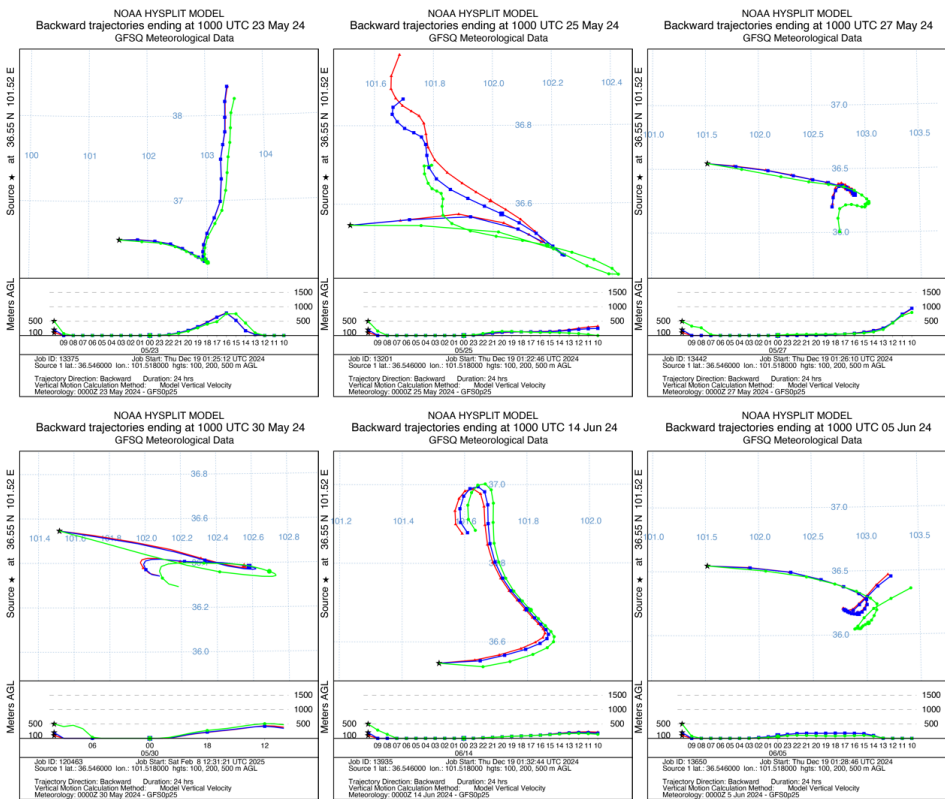
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Figure A 1 The back trajectories originating from eastern areas and arriving at EM27/SUN station at 10 UTC (18 Local time) on different days. The trajectories, calculated using the HYSPLIT model (Hybrid Single-Particle Lagrangian Integrated Trajectory), and shown for height of 100, 200 and 500 m above ground level.

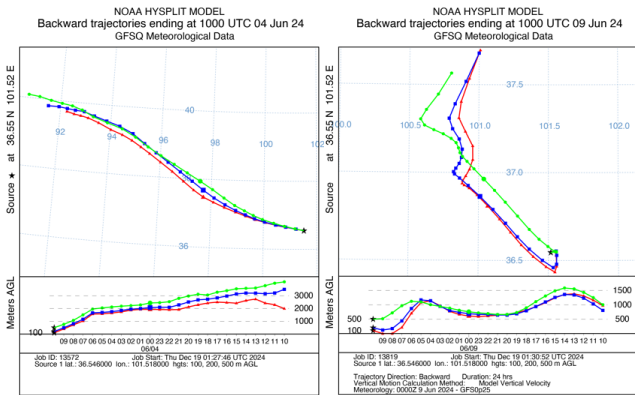


Figure A 2 Similar to Figure A 1, but for trajectories generally from western areas on June 4 and 9, 2024.

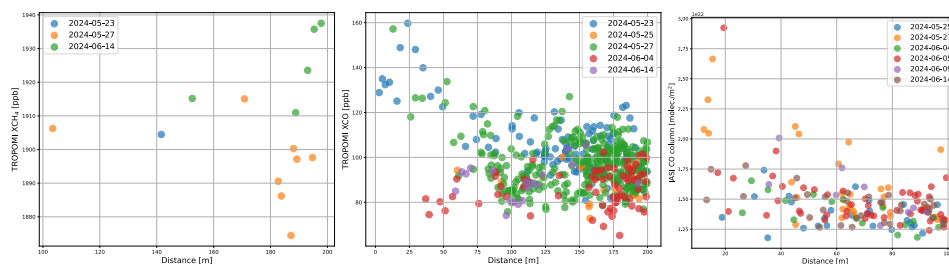


Figure A 3: Correlation between TROPOMI XCH₄ (a) and XCO (b) with the distance from each TROPOMI observation to the EM27/SUN location. Panel (c) shows the correlation between IASI CO and the distance from each IASI observation to the EM27/SUN location.

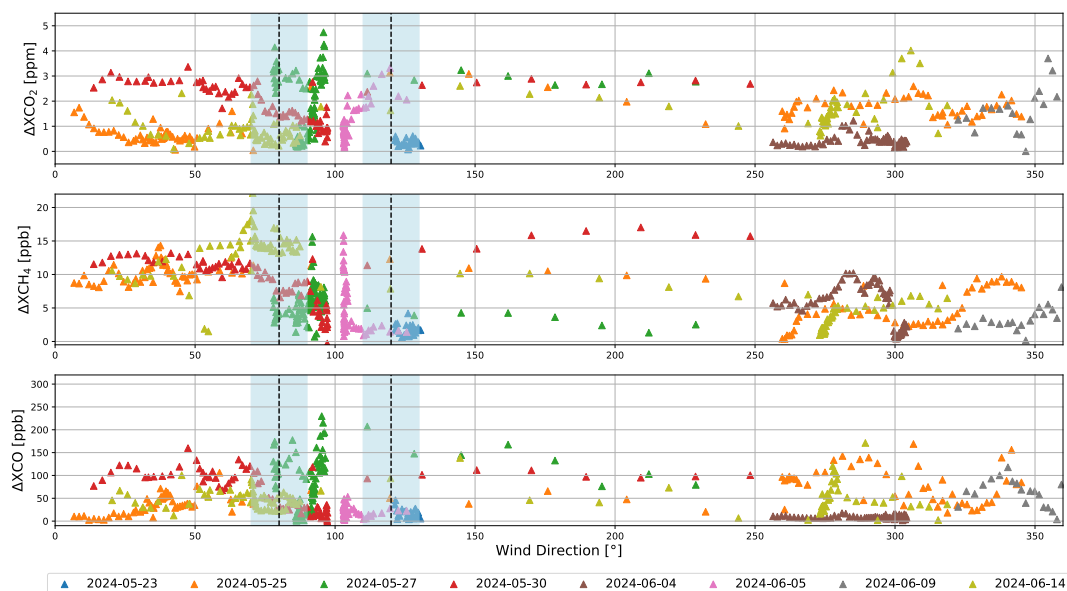
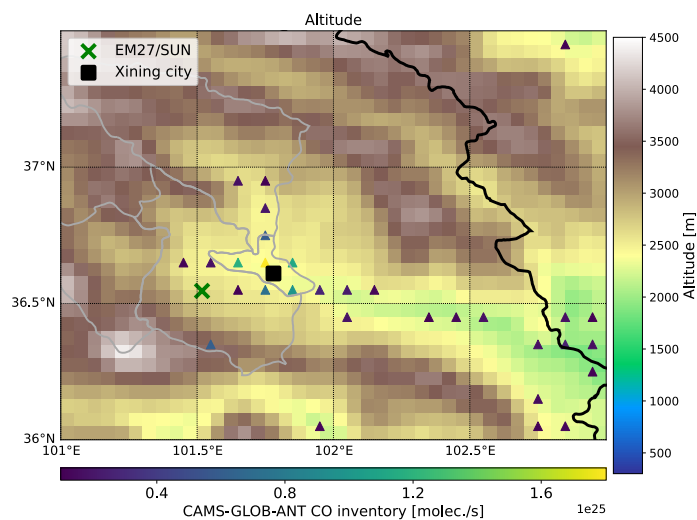


Figure A 4: correlations between ΔXCO_2 , ΔXCH_4 and ΔXCO with wind direction. The wind direction range of 80°-120° is delineated by dashed lines.



381

382 **Figure A 5: similar to Figure 7 but for terrain distribution.**

383



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