



1 **Assessment and prediction of dust emissions, deposition** 2 **and radiation forcing in Central Asia**

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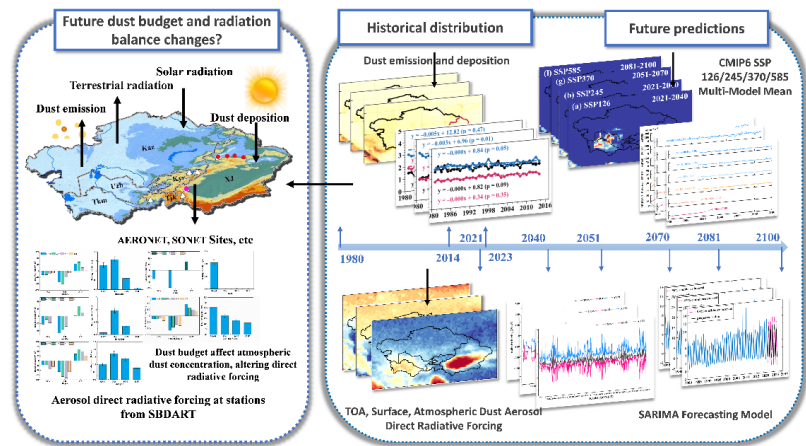
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14 **Abstract.** Dust aerosols significantly influence climate by modulating radiative balance and cloud
15 processes. This study integrates MERRA-2 reanalysis data and the CMIP6 multi-model ensemble to
16 assess the spatiotemporal evolution of dust emissions, deposition, and associated radiative effects in
17 Central Asia from 1980 to 2100. Four SSP scenarios project that dust emissions in Central Asia exhibit
18 a high-emission, high-deposition pattern with primary sources exceeding $15 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$. The deposition
19 area significantly exceeds the source area (maximum $>8 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$). Cross-scenario analysis
20 demonstrates that dust emissions are highly sensitive to climate policy, with end-of-century emissions in
21 the SSP5-8.5 high-emission scenario increasing by 94.9% relative to the baseline period. In contrast,
22 emissions under the SSP1-2.6 low-carbon pathway vary by only 4.5%. Simulations using the SBDART
23 model show that aerosol direct radiative forcing (ADRF) from dust in Central Asia under clear-sky
24 conditions exhibits a vertical gradient, with cooling at the top of the atmosphere (TOA) and heating near
25 the surface, yielding a net negative forcing at the TOA, with a minimum of $<-10 \text{ W/m}^2$ near the Caspian
26 Sea. Peak positive forcing within the atmosphere, observed in spring, reaches 10.0 W/m^2 . Increased dust
27 emissions reduce shortwave radiation at the surface by up to -20 W/m^2 . Ground-based observations
28 indicate seasonal variations in the dust-induced heating rate, with peak near-surface heating in spring at
29 Kashgar (93.0 W/m^2) and a maximum heating rate of 2.6 K/day . In contrast, the near-surface heating rate
30 at Issyk-Kul Lake in autumn (0.34 K/day) is approximately four times higher than in spring (0.08 K/day).



31 **Graphical Abstract**



32 **Keywords:** Dust cycle; CMIP6 multi-model ensemble (MME) ; Direct radiation forcing of dust;
33 SBDART model

34 **1.introduction**

35 Dust aerosols are a significant component of the mass load of tropospheric aerosols, accounting for up
36 to approximately 50%, and have a profound impact on the functioning of the Earth system (Mahowald
37 et al. 2010, Ramanathan et al. 2001). Their trans-circulation process (lithosphere-atmosphere-cryosphere)
38 and interaction with the climate system have become cutting-edge research areas in Earth system science.
39 The release, transport, and deposition of dust aerosols not only involve multiple geospheres but also have
40 a significant impact on weather and climate, air quality, and even human health after entering the
41 atmosphere (Tegen et al. 2004, Penner et al. 2006, Pozzer et al. 2012).
42 Global annual dust emissions are enormous, ranging from approximately 1000 to 2150 Tg, with 30% to
43 40% originating from arid regions of Asia (Tanaka and Chiba 2006). Dust is transported across continents
44 by the westerly wind circulation, which significantly impacts the atmospheric radiation balance in East
45 Asia, North America, and even the Arctic region (Wallace and Hobbs 2006). Although studies have
46 confirmed that dust regulates the land-atmosphere energy budget through direct radiative forcing
47 (including scattering and absorption of short-wave and long-wave radiation) and indirect effects (such as
48 changing the efficiency of precipitation as cloud condensation nuclei), significant uncertainty remains
49 regarding the vertical distribution of dust, the amplification mechanism of anthropogenic emissions, and



50 its regional climate feedback (IPCC AR6).

51 Due to the challenges associated with dust observation, our understanding of the behavior of dust

52 throughout its life cycle remains insufficient, hindering a complete understanding and accurate modeling

53 of its complex mechanism of action (Kok et al. 2023). Numerous studies have used a variety of methods,

54 including in situ observations, satellite remote sensing, and model simulations, to thoroughly examine

55 the spatiotemporal changes, optical properties, and radiative forcing of dust aerosols (Wang et al. 2018,

56 Song et al. 2021, Chen, Zhao and Fan 2022). For example, global dust are primarily confined to the "dust

57 belt," with approximately one-third originating from the Asian region (Kok et al. 2023). Dai et al. utilized

58 a variety of remote sensing and ground-based data to study the sources, microphysical characteristics,

59 and optical properties (Dai et al. 2022, Salvador et al. 2022). Zhao et al. investigated the simulation of

60 global and regional dust by 16 CMIP6 models in the Atmospheric Model Intercomparison Project (AMIP)

61 experiment and compared the results with observational and reanalysis data (Zhao et al. 2023, Liu et al.

62 2024).

63 Mode simulation provides information on the temporal and spatial changes of dust aerosols worldwide

64 and helps predict future trends (Li et al. 2021). The results of climate models such as CMIP5 and CMIP6

65 have enabled us to understand the main characteristics of dust aerosols better. These models have

66 increasing resolutions and increasingly complex physical processes and parameterizations,

67 demonstrating their ability to simulate dust events and processes on meso- to global scales (Zhao, Ryder

68 and Wilcox 2022). In particular, the CMIP6 experiment has provided a valuable opportunity to

69 understand the impact of dust emissions on climate and the role of dust in the latest generation of climate

70 models (Braconnot et al. 2021, Zhao, Wilcox and Ryder 2024).

71 The flux of dust emissions (Shen et al. 2016). However, most current research focuses on the

72 spatiotemporal distribution and transport processes of dust (Li et al. 2022, Tao et al. 2022). A systematic

73 understanding of key aspects of the local dust life cycle in this region remains lacking, including the

74 long-term evolution of the dust emission-deposition budget, the high dependence of the vertical dust

75 profile on direct radiative forcing, and the modal differences in the climatic feedback of dust under

76 different carbon emission scenarios. These deficiencies in understanding seriously limit the reliability of

77 climate models in Central Asia, with the uncertainty in radiation-forcing estimates primarily stemming

78 from insufficient ground verification due to a lack of sites (Brown et al. 2021, Wu and Boor 2021).



79 To overcome the aforementioned bottleneck, this study establishes a multi-source data fusion framework
80 that seamlessly integrates the following observational and modeling resources: (1) MERRA-2 reanalysis
81 data and the CMIP6 multi-model ensemble are employed to construct high-resolution data on the dust
82 budget in Central Asia through dynamic downscaling and forecast the evolutionary trend of the climatic
83 effect of dust under different scenarios; (2) The SBDART radiative transfer model is integrated with the
84 valuable measured data from the SONET Asian Dust Observation Network and the Jinghe CE318
85 ground-based remote sensing station to quantify the long-term trend of the shortwave radiation forcing
86 (ADRF) of dust under clear-sky conditions; (3) The SARIMA statistical model is used to predict the
87 short-term evolutionary trend of the radiative effect of dust. This method is the first to achieve a closed-
88 loop analysis of the entire "emission-deposition-radiation" chain of dust in the source region using
89 multiple datasets, providing key constraints for elucidating the physical mechanisms of the dust life cycle.
90 The structure of this paper is as follows. Section 2 presents the data sources, the downscaling method for
91 the CMIP6 dust budget, and the calculation method for clear-sky aerosol radiative forcing. Section 3
92 examines the detailed characteristics of the dust budget, projections of future changes, and the radiative
93 forcing of dust aerosols. Finally, the main conclusions and a discussion are presented in Section 4.

94 **2. Data and Methods**

95 **2.1 Data sources**

96 The study area is situated between 35°-57°N and 48°-96°E, encompassing the five Central Asian
97 countries (Kazakhstan, Uzbekistan, Tajikistan, Turkmenistan, and Kyrgyzstan) and the Xinjiang region
98 of China (comprising both its northern and southern parts). This region is positioned in the hinterland of
99 the Eurasian continent and is characterized by a temperate continental climate with extreme aridity. The
100 region features a highly heterogeneous surface, with the Taklamakan Desert and the surrounding Gobi
101 (comprising over 40% of the study area) interspersed with mountain ranges, such as the Tianshan and
102 Pamir, forming a unique landform (Shen et al. 2016, Hetzel et al. 2002). As the world's second-largest
103 source of dust, strong thermal and dynamic coupling drives intense dust activities (Zhang et al. 2020),
104 with emission hotspots concentrated in the Tarim Basin, the dried bed of the Aral Sea, and the Kazakh
105 steppe belt. This study focuses on the regional dust budget and radiative effects, utilizing MERRA-2



reanalysis data, the CMIP6 multimodel ensemble, AERONET, SONET, and hand-held photometer data.

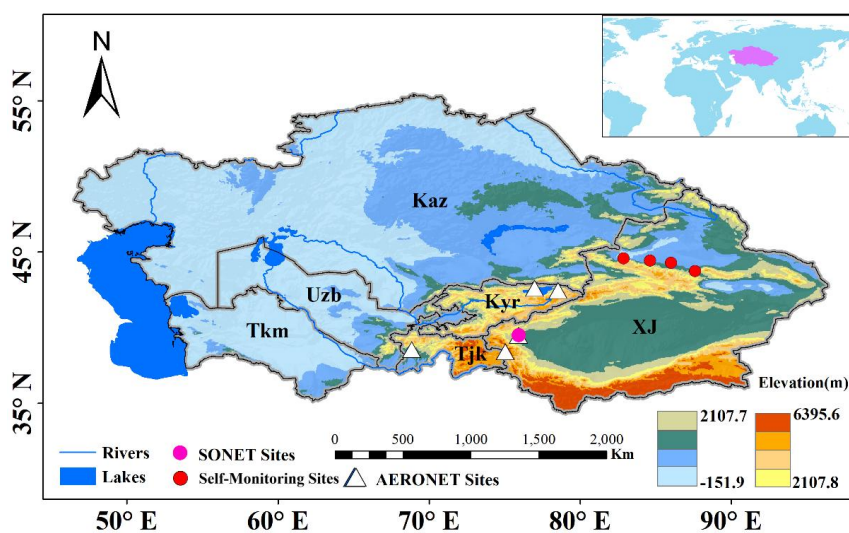


Figure 1 Location of the Study area.

2.1.1 Data from ground-based heliometers

AERONET (Aerosol RObotic NETwork) employs a CE-318 solar photometer to measure aerosol optical depth (AOD) across 8 bands in the range of 340–1640 nm and to derive microphysical parameters, including single scattering albedo (SSA), refractive index (m), and particle size spectrum (Holben et al. 1998, Holben et al. 2001). The Level 2 data exhibit an uncertainty of less than 5%. As an internationally recognized standard for ground-based aerosol observations, its long-term stability and algorithmic consistency offer a reliable input for radiative forcing calculations (García et al. 2012).

The Chinese Academy of Sciences-led SONET (Sun-sky radiometer Observation NETWORK) employs the CE318-DP instrument to provide information on the chemical composition and vertical profile of aerosols while adhering to the stringent quality control procedures of AERONET. The establishment of SONET sites has effectively addressed the gaps in AERONET's spatial coverage in this source region (Li et al. 2018). Cross-validation demonstrated that the correlation coefficient between SONET and AERONET AOD was 0.98 ($RMSE < 0.02$), confirming a seamless connection between the two data sets (She et al. 2024).

To supplement the insufficient temporal and spatial coverage of fixed stations, this study employs CE-



124 318 and Microtops II handheld photometer to obtain transient AOD observations in the 550–870 nm
125 band (accuracy ± 0.01) for verifying the local applicability of satellite inversion products. By integrating
126 the aforementioned multi-scale observational data, this paper employs AERONET and SONET Level 2
127 data to provide vertical profiles of the optical-physical properties of aerosols, calculate the direct radiative
128 forcing of aerosols, and validate the satellite data on AOD and radiation flux in Central Asia
129 (Supplementary Figure 1).

130 **2.1.2 MERRA-2 reanalysis data**

131 The MERRA-2 reanalysis data used in this study was developed by the NASA Goddard Space Flight
132 Center. Its core is based on the GEOS-5 atmospheric circulation model and the ADAS-5.12.4
133 assimilation system. A global multi-element dataset with 72 vertical layers (surface to 80 km) and a
134 horizontal resolution of $0.625^\circ \times 0.5^\circ$ has been constructed from 1980 to the present by fusing satellite
135 remote sensing (MODIS/AVHRR aerosol optical thickness), ground-based observations (soundings,
136 aircraft observations), and the GOCART aerosol chemical transport model output (Gelaro et al. 2017).
137 In addition to covering variables related to cloud, radiation, and hydrological cycles, the coupled
138 GOCART model distinguishes the interaction mechanisms of five types of aerosols in this dataset: dust
139 (DU), sea salt (SS), sulfate (SO_4), black carbon (BC), and organic carbon (OC). For the first time, the
140 entire life cycle of dust aerosols has been analyzed, providing key parameters such as monthly average
141 dust emission flux, dry/wet deposition rate, particle size-classified loads, and single scattering albedo at
142 483.5 nm, ensuring physical consistency for quantifying the radiative forcing of dust (Buchard et al.
143 2017). Based on the advantages of this data, this study extracts radiation flux and dust cycle parameters
144 under clear sky conditions in Central Asia and systematically constructs a collaborative analysis
145 framework for dust emissions, deposition, and radiative forcing.

146 **2.1.3 CMIP6 model simulations**

147 The Sixth Coupled Model Intercomparison Project (CMIP6) integrates 112 climate models from 33
148 institutions worldwide, with its multi-scenario simulation significantly surpassing previous studies in
149 both breadth and depth. To analyze the interdecadal variations in dust emissions and wet and dry
150 deposition in Central Asia, 10 models were selected from CMIP6 based on the principle of data



completeness (Eyring et al. 2016). The selection criteria include key variables of the dust cycle: monthly mean dust emission fields and dry/wet deposition fluxes from 1980 to 2014 for the historical period, and from 2015 to 2100 for four shared socioeconomic pathways (SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5).

To ensure spatial consistency in the comparison of multi-source data, all model output data were subjected to statistical downscaling and aligned with MERRA-2 reanalysis data (spatial resolution $0.625^\circ \times 0.5^\circ$). This multi-model ensemble effectively characterizes the uncertainty in climate responses while controlling computational costs, providing reliable data support for analyzing the long-term evolution of the dust cycle in the arid region of Central Asia.

Table.1 Overview of the models and simulations used in this study.

Model	Nation	Resolution	Hist	SSP126	SSP245	SSP370	SSP585	Dust emission scheme	Model references
CESM2-WACCM	USA	$1.25^\circ \times 0.94^\circ$	3	1	5	3	5	Zender et al. (2003)	Danabasoglu et al. (2020)
CESM2	USA	$1.25^\circ \times 0.95^\circ$	11	3	3	3	3	Zender et al. (2003)	Wu et al. (2016)
CNRM-ESM2-1	France	$1.25^\circ \times 0.96^\circ$	3	5	10	5	5	Marticorena et al. (1997)	Séférian et al. (2019)
GFDL-ESM4	USA	$1.25^\circ \times 0.97^\circ$	1	1	1	1	1	Evans et al. (2016)	Dunne et al. (2020)
GISS-E2-1-G	USA	$1.25^\circ \times 0.98^\circ$	19	10	25	17	10	Ginoux et al. (2004)	Bauer et al. (2020)
GISS-E2-1-H	USA	$1.25^\circ \times 0.99^\circ$	10	5	5	1	5	Bauer and Koch. (2005)	Kelley et al. (2020)
GISS-E2-2-G	USA	$1.25^\circ \times 1.00^\circ$	5	5	5	5	5	Cakmur et al. (2006)	Rind et al. (2020)
MRI-ESM2-0	Japan	$1.25^\circ \times 1.01^\circ$	12	5	10	5	6	Tanaka and Chiba. (2005)	Yukimoto et al. (2019)
HadGEM3-GC31-LL	UK	$1.875^\circ \times 1.25^\circ$	5	3	4	2	3	Marticorena. (1995)	Williams et al. (2020)
UKESM1-0-LL	UK	$1.25^\circ \times 0.103^\circ$	3	5	5	3	4	Marticorena. (1995)	Senior et al. (2020)



161 2.2 Methodology

162 2.2.1 Delta statistical downscaling

163 Due to the limited original spatial resolution of the CMIP6 model (with a typical horizontal grid of
164 approximately $1.25^\circ \times 1^\circ$), direct application to the analysis of regional-scale dust cycles may introduce
165 systematic biases. Therefore, the Delta Change Factor method is employed in this study for statistical
166 downscaling. The core of this method is to separate the historical bias of the climate model from the
167 future change signal, followed by the reconstruction of high-resolution climate elements (Maraun et al.
168 2010, Gutmann et al. 2014).

169 First, the baseline period deviation is calculated, and the monthly mean dust emission $P_{m,his}$ of the
170 historical simulation data (1980-2014) of each CMIP6 model is extracted. The grid is then matched with
171 the MERRA-2 reanalysis observations P_{obs} for the same period to calculate the model's systematic
172 deviation ratio:

$$173 \quad B_m = \frac{P_{m,his}}{P_{obs}} \quad (1)$$

174 where $\overline{P_{obs}}$ is the monthly average of the observation period, and B_m represents the spatial deviation of
175 model m in the reference period.

176 Second, the relative change factor for future scenarios is extracted, and the ratio of dust emissions for
177 each model during the future scenario period (2015-2100) relative to its own historical simulation is
178 calculated.

$$179 \quad R_{m,fut} = \frac{P_{m,fut}}{P_{m,his}} \quad (2)$$

180 Among them, $P_{m,fut}$ is the average monthly emission of mode m in the future, and $P_{m,his}$ is the average
181 monthly emission of mode m over multiple years in the historical period.

182 This method decouples the historical deviation from the climate change signal, retaining the physical
183 response characteristics of CMIP6 to future climate forcing while improving the simulation accuracy at
184 the regional scale through the use of high-resolution observational data. Compared with dynamic
185 downscaling, it significantly reduces computational costs and is well-suited for multi-model uncertainty
186 quantification research.



2.2.2 SBDART Radiative Transfer Model Calculation of Direct Radiative Forcing of Aerosols

The SBDART radiation transfer model (Ricchiazzi et al. 1998) was employed in this study to quantitatively assess the direct radiative effect of aerosols. SBDART solves the atmospheric radiation transfer equation using the four-stream approximation method. Its core architecture comprises three modules: First, the discrete ordinates radiation transfer (DISORT) module calculates the radiative fluxes of the 45-layer atmosphere (with a vertical resolution of 0.3 km); second, the spectral parameterization module integrates the LOWTRAN-7 atmospheric absorption spectrum and Mie scattering theory to cover the shortwave band from 0.25 to 4.0 μm ; and finally, the surface-atmosphere coupling module analyzes the radiative interaction between surface albedo and atmospheric constituents such as water vapor and ozone.

Our research is based on a comprehensive data collection, with key input parameters including the optical properties (e.g., optical depth τ , single scattering albedo SSA, asymmetry factor ASY), and the vertical profiles of aerosols. These are obtained from the solar photometer observation network at the Central Asia site, which offers a significant advantage in temporal and spatial resolution over satellite retrieval products (Dubovik and King 2000). To quantify the radiative forcing of dust aerosols, all simulations were performed under clear sky conditions, and the solar zenith angle was constrained according to the seasonal average value of the study area to ensure the comparability of regional radiative effects (Halthore et al. 2005). The aerosol direct radiative forcing (ADRF) is calculated using the standard scientific approach to determine the difference in net radiative flux with and without aerosols under cloud-free conditions. Specifically, the ADRF at a given altitude z , at the top of the atmosphere (TOA), at the surface (SFC), and in the atmosphere (ATM), can be defined as follows:

$$NF_z = F_{z,down} - F_{z,up} \quad (3)$$

$$ADRF_z = NF_z^{aer} - NF_z^{noaer} \quad (4)$$

$$ADRF_{TOA} = NF_{TOA}^{aer} - NF_{TOA}^{noaer} \quad (5)$$

$$ADRF_{SFC} = NF_{SFC}^{aer} - NF_{SFC}^{noaer} \quad (6)$$

$$ADRF_{ATM} = ADRF_{TOA} - ADRF_{SFC} \quad (7)$$

$$ADRF_{dust} = ADRF \times \left(\frac{DAOD}{AOD} \right) \quad (8)$$

Among them, $F_{z,down}$ and $F_{z,up}$ are the downward and upward radiative fluxes, NF_z^{aer} and NF_z^{noaer} are the net radiative fluxes with and without aerosols, and ADRF is the aerosol direct radiative forcing.



216 2.2.3 SARIMA prediction model

217 Given the non-stationarity and interannual cycle characteristics of the radiation forcing time series of
218 Central Asian dust, this study employs the seasonal autoregressive integrated moving average model
219 (SARIMA) for modeling and analysis. First, the augmented Dickey-Fuller test (ADF, $p < 0.05$) was used
220 to identify the non-stationarity of the series. A compound differencing strategy (first-order conventional
221 difference $d=1$, first-order seasonal difference $D=1$, period $s=12$) was employed to eliminate the trend
222 and interannual fluctuations, resulting in a stationary residual series (KPSS test $p > 0.1$).
223 The non-seasonal order ($p=2$, $q=1$) was determined based on the autocorrelation function (ACF) and
224 partial autocorrelation function (PACF), while the seasonal order ($P=1$, $Q=1$) was optimized using grid
225 search, resulting in the final SARIMA(2,1,1)(1,1,1)₁₂ model (AIC=112.3, BIC=125.7). Model validation
226 demonstrated that the goodness of fit was $R^2=0.87$ for annual cycle dynamics, and the prediction error
227 for extreme event peaks was less than 15%, confirming its effectiveness in the analysis of non-stationary
228 sequences (Sirisha, Belavagi and Attigeri 2022).

229 3. Results and analysis

230 3.1 Spatial pattern and multimode prediction of dust emissions in Central Asia

231 Figure 2 compares MERRA-2 observations with CMIP6 multi-model ensemble (MME) dust emissions
232 from 1980 to 2014. The historical spatial distribution of the ten models is shown in Supplementary Figure
233 2. The reanalysis data agrees with the MME simulations, with the Taylor skill score (SS) close to 1. Dust
234 emissions in the study area exhibit significant temporal and spatial variation. In terms of spatial
235 distribution (Figure 2a), both datasets consistently identify the three primary core emission sources in
236 the Tarim Basin, the dried-up Aral Sea area, and the Gobi Desert, with maximum emission fluxes
237 exceeding $15 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$. Regarding the trend of dust loads (Figure 2b), dust emissions in the Aral Sea
238 region have increased significantly (>0.5) over the past 34 years, while those in the Tarim Basin have
239 slightly decreased (≈ -0.3).

240 The Aral Sea region has experienced a 68% reduction in lake area since 1960, resulting in 54,000 km² of
241 exposed lakebed (Wang et al. 2020). Under arid climatic conditions with an annual average precipitation
242 of less than 100 mm and potential evaporation of more than 2000 mm, the dust emission flux has
243 increased significantly at a rate of approximately $0.5 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}\cdot\text{yr}^{-1}$ over 34 years. In contrast, the Tarim



Basin has benefited from ecological restoration projects and increased precipitation during the growing season (Fu et al. 2021), leading to a decrease in emission flux at a rate of $\approx 0.3 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}\cdot\text{yr}^{-1}$. Time series analysis (Figure 2c) shows that overall dust emissions fluctuate gently without significant annual changes. Dust emissions in the southern Tarim Basin of Xinjiang increase and decrease annually, consistent with the spatial distribution map of the trend. Dust emissions in northern Xinjiang are similar to those in Central Asia, with northern Xinjiang slightly higher than other regions in Central Asia. This may be attributed to local differences in surface roughness and land use, reflecting regional disparities in emission characteristics.

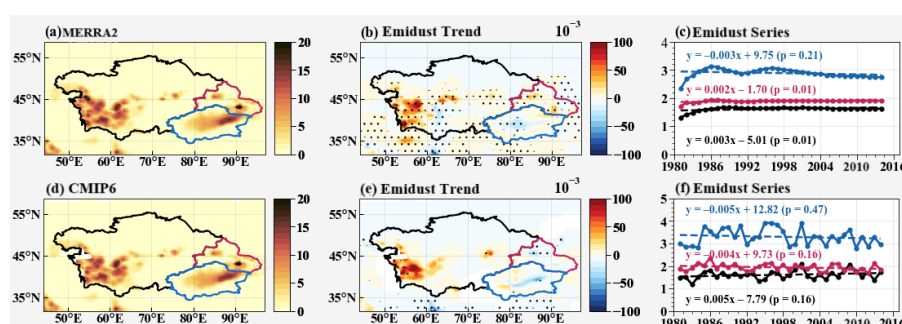


Figure 2 Spatial distribution, linear trend, and time series of dust emissions from MERRA-2 and CMIP6 MME in Central Asia from 1980 to 2014. Red highlights the northern part of Xinjiang, blue indicates the southern part of Xinjiang, and black indicates the five Central Asian countries. The black dots in (b) and (e) indicate the 90% confidence level regions.

Figure 3 illustrates the relative changes in Central Asian dust emissions in the near term (2021–2040), medium-term (2051–2070), and long-term (2081–2100), compared to the reference period (1980–2014), with significant temporal and spatial differences observed. In all scenarios, areas with high dust values ($>50 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) are consistently distributed across the Aral Sea hinterland, Turkmenistan, and the eastern edge of the Tarim Basin. Dust emission intensity is positively correlated with the radiative forcing scenario, with intensity continuing to increase over time within a given scenario (long term > medium term > short term). Specifically, short-term emissions in the Aral Sea region range from 17.8 (SSP370) to 26.0 (SSP245) $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, with relatively minor differences. However, in the long term, under the high-radiation scenario (SSP585), emissions surge to 387.1 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, representing a 94.9% increase compared to the reference period. This sharp increase is directly attributed to the exposure of saline sediments on the lake bed, soil loosening due to rising surface temperatures, and increased wind erosion (Lioubimtseva and Cole 2006). In contrast, the Tarim Basin experienced a long-term decrease in



emissions ranging from 18.7% (SSP245) to 29.3% (SSP370) due to ecological restoration (with a 10-year NDVI increase of 0.12) and an increase in growing season precipitation (Xu et al. 2019). The SSP585 scenario shows a decrease from 27.2 (short-term) to 20.1 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, reflecting a 26.1% reduction. Regional comparisons reveal significant differences in the sensitivity of climate responses. Emissions from the Aral Sea increase exponentially with the intensity of radiative forcing ($R^2 = 0.93$). Meanwhile, the southern Xinjiang region shows a gradual decreasing trend, confirming the potential of human intervention in regulating the dust process.

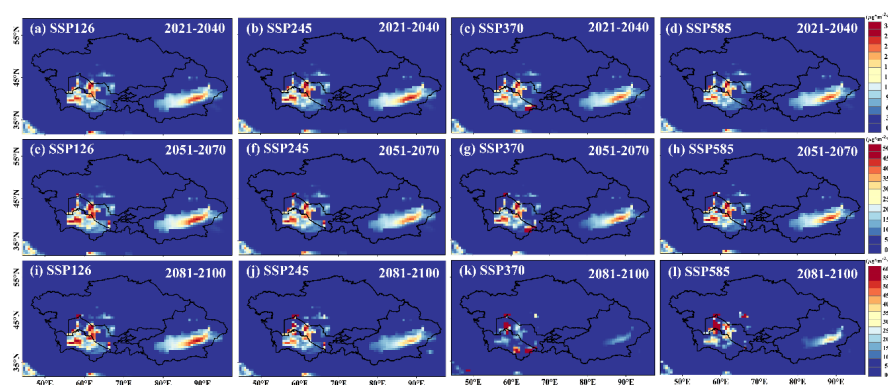


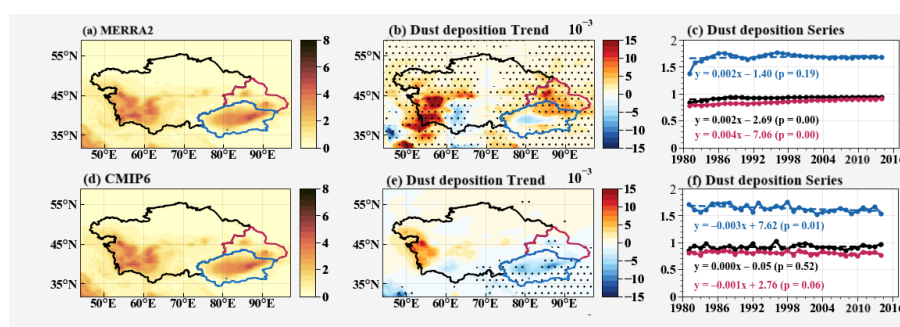
Figure.3 Future changes in dust emissions over time. Spatial variability of dust emissions in Central Asia under the four SSP scenarios of CMIP 6 MME for (a-d) the near term (2021-2040), (e-h) the medium term (2051 - 2070), and (i-l) the long term (2081 - 2100) relative to the historical period (2000-2014).

3.2 Spatial pattern and multi-model prediction of dust deposition in Central Asia

Dust emissions and deposition together constitute the complete dust budget, with deposition serving as the final outcome of emissions. Dust particles entering the atmosphere are redistributed at the surface-atmosphere interface through gravity-dominated dry deposition and precipitation-driven wet deposition (see Supplementary Figures 3-4 for the historical distribution of dry and wet deposition across 10 models) (Marticorena and Bergametti 1995, Shao et al. 2011). Figure 4 shows that the multimodel ensemble (MME) and the observed data are in good agreement in simulating total dust deposition in Central Asia, though the deposition trend observed by MERRA-2 is significantly stronger than that of the model ensemble. In terms of spatial distribution, the high-value sedimentation areas ($>5 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) heavily overlap with emission hotspots, concentrated in the western part of Central Asia and the Tarim Basin in southern Xinjiang, confirming the local coupling mechanism of dust "generation-deposition." The trend analysis (Figure 4b) shows that the eastern edge of the Aral Sea and the East Caspian Sea exhibit the



292 most substantial positive trend ($\Delta S = +0.15 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$). Meanwhile, southern Xinjiang is characterized
293 by a negative trend ($\Delta S = -0.10 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$). Time series analysis (Figure 4c) indicates that the observed
294 data for the period 1980-2014 exhibit a slowly increasing trend in sedimentation flux in Central Asia,
295 with an annual rate of change of 0.002. In contrast, the MME simulation for the Xinjiang region shows
296 a slight decrease of -0.003 yr^{-1} . This discrepancy between observations and simulations may arise from
297 the model's uncertainty regarding the boundary layer's dynamic processes and the parameterization of
298 precipitation microphysics in Central Asia's arid region. Specifically, the quantification of wet deposition
299 efficiency for dust requires further improvement.

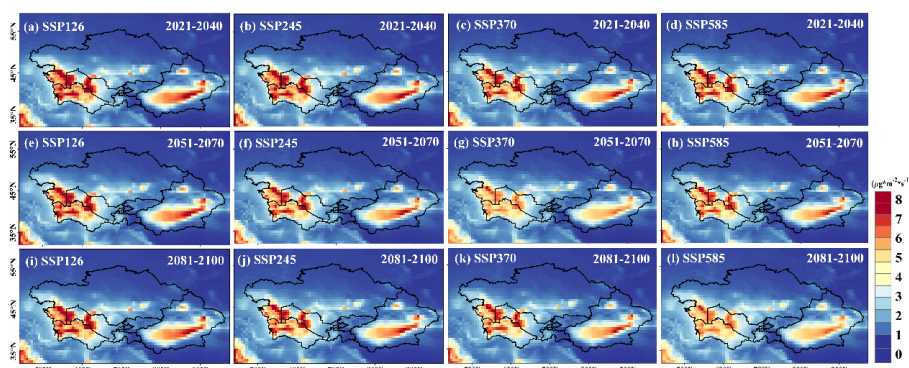


300 **Figure. 4** Spatial distribution, linear trend, and time series of dust deposition (dry deposition + wet
301 deposition) for MERRA-2 and CMIP6 MME in Central Asia from 1980 to 2014. Red highlights the northern
302 part of Xinjiang, blue indicates the southern part of Xinjiang, and black indicates the five Central Asian
303 countries. The black dots in (b) and (e) indicate the regions with a 90% confidence level.

304 Figure 5 illustrates the relative changes in dust deposition projections over time for the four scenarios
305 (see Supplementary Figures 5-6 for future changes in dry and wet deposition). In contrast to the
306 distribution of the dust emission source area, the influence of deposition extends to the surrounding
307 regions, primarily covering the southwestern part of Central Asia, the southeastern edge of the Tarim
308 Basin, and the Junggar Basin, with the maximum deposition flux exceeding $8 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, presenting a
309 spatial pattern of "deposition domain > emission source." In terms of temporal evolution, the mean values
310 ranged from 9.3 (SSP585) to 10.4 (SSP245) $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in the near term (2021-2040) and from 9.6
311 (SSP370) to 10.0 (SSP126) $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in the long term (2081-2100), with an overall variation of less
312 than 12%. This phenomenon may stem from the compensatory effect of wet and dry deposition processes,
313 with dry deposition fluxes decreasing by approximately 0.2% per year under medium- to high-level
314 radiative forcing in the southern Xinjiang region due to precipitation patterns. In contrast, dry deposition
315 increases in western Central Asia due to the enhancement of near-surface winds resulting from reduced



316 surface roughness. However, the spatial and temporal stability of the wet deposition process mitigates
317 the overall deposition fluctuations.



318 **Figure. 5** Future changes in dust deposition over different periods. Spatial variability of dust deposition in
319 Central Asia under the four SSP scenarios of CMIP 6 MME for (a-d) the near term (2021-2040), (e-h) the
320 mid-term (2051 - 2070) and (i-l) the long term (2081 - 2100) relative to the historical period (2000-2014).

321 In order to more accurately assess the trend simulation performance of the dust cycle, we constructed a
322 time series of dust emission and wet and dry deposition between 1980 and 2100 based on the MERRA-
323 2 and CMIP6 multimodel ensembles (MME) (see Fig. 6). Overall, the simulation results show that dust
324 emissions in Xinjiang remain relatively stable over the next 120 years. In contrast, in the five Central
325 Asian countries, especially under scenarios with higher radiative forcing (e.g., SSP370, SSP585), dust
326 emissions increase significantly between 2081 and 2100, accompanied by a fluctuating gradual increase.
327 In contrast, dust deposition (both dry and wet deposition) shows a smoother trend with less volatility.
328 In the specific analyses, dust emissions from MERRA-2 exhibited a smooth trend, with an average of 30
329 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in the Tarim Basin and 15 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in other regions. In contrast, simulated data from the
330 multi-model ensemble exhibited slight fluctuations, with peaks exceeding 45 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ at certain times.
331 Some deviation is observed in the temporal variability between the two. The volatility of dry deposition
332 of dust is relatively small, with a slope of less than 0.1, indicating that the dry deposition process is
333 smooth. Additionally, both dry and wet deposition do not exhibit significant volatility in the long-term
334 trend. Wet deposition exhibited slight deviation in northern Xinjiang and remained relatively smooth in
335 other regions. The average wet deposition flux was approximately 1.5 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, with an overall slope
336 of less than 0.2, indicating relatively small variation. Notably, the wet deposition data from MERRA-2
337 exhibited a significant increase in the northern border region around 2000. This change may be related
338 to the assimilation of MODIS satellite and other observations by MERRA-2. Therefore, MERRA-2 data



from 2000 to 2014 were selected to calibrate the model and ensure the accuracy of the simulation results. In summary, although future dust emissions vary significantly under different climate scenarios, the overall dust deposition process remains relatively stable. The simulation results from MERRA-2 and the multi-model ensemble exhibit spatial and temporal differences in different regions.

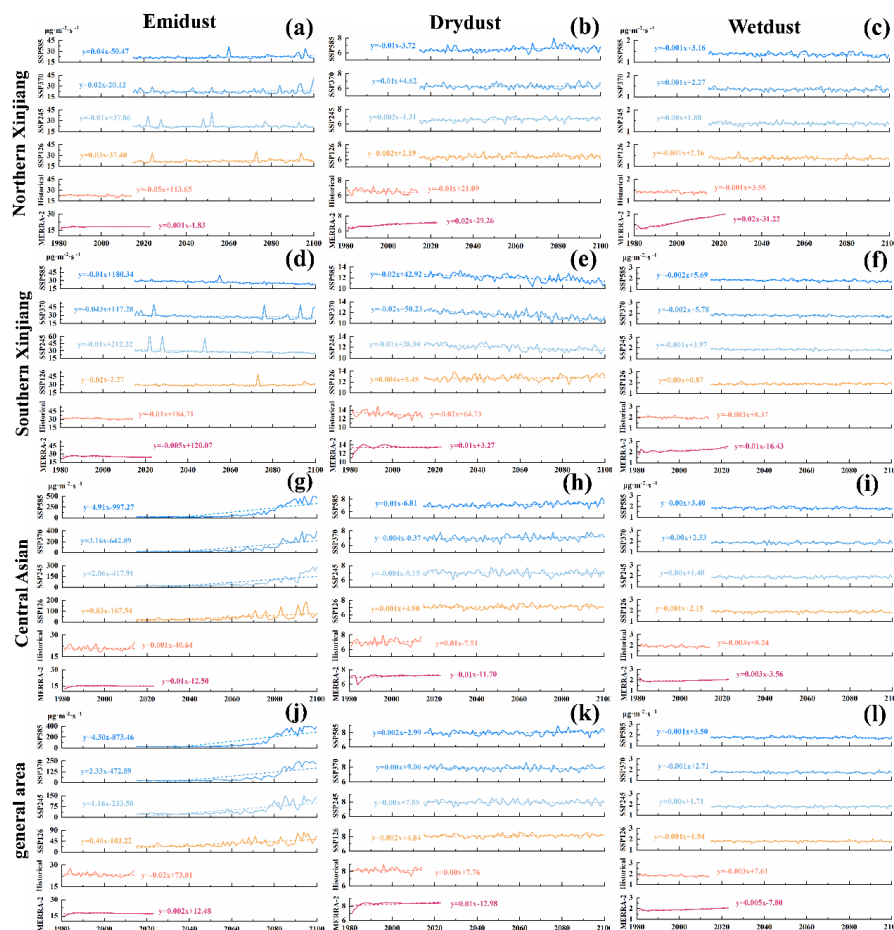


Figure 6 Time evolution of dust receipts and payments. Dust emissions, and wet and dry deposition ($\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) in the northern (a-c), southern (d-f) Central Asia (g-i), and the whole region (j-l). From CMIP 6 MME 1980-2100, MERRA-2 results 1980-2023.

3.3 Analysis of changes in direct radiative forcing of dust aerosols

3.3.1 Monthly average changes in direct radiative forcing by dust aerosols

Based on the above quantitative characterization of dust emission sources and deposition processes, further investigation is required to elucidate the perturbation mechanism of dust aerosols on the energy



balance of the surface-atmosphere system. This study quantifies the radiative balance impacts of Central Asian dust aerosols at different spatial and temporal scales through the short-wave direct radiative forcing (ADRF) using MERRA-2 observations of dust aerosols under clear-sky conditions from 1980 to 2023. As shown in Fig. 7a-d, the top-of-atmosphere (TOA) radiative forcing exhibits considerable spatial heterogeneity. The overall negative forcing is characterized by the lowest value ($< -10 \text{ W/m}^2$) in the Caspian Sea region, followed by the Tarim Basin and the Aral Sea region ($< -8 \text{ W/m}^2$), confirming that dust aerosols exert a significant cooling effect by enhancing shortwave reflection. The seasonal analysis showed that the negative TOA forcing intensity followed a decreasing pattern: spring (-3.32 W/m^2) > summer (-3.21 W/m^2) > fall (-3.07 W/m^2) > winter (-1.94 W/m^2), which was closely associated with the seasonal characteristics of dust activity.

The spatial pattern of surface (SFC) radiative forcing (Fig. 7e-h) shows a more pronounced negative distribution, with two cooling centers forming in the Tarim Basin and southwestern Central Asia, peaking at shortwave radiation losses of -20 W/m^2 . This phenomenon results from the synergistic effect of scattering and absorption within the atmosphere, a dual attenuation mechanism of incident solar radiation by dust particles (Li et al., 2022a), which significantly diminishes the surface energy balance. Notably, the atmospheric radiative forcing (ADRF) is consistent with TOA and SFC in spatial distribution. However, its positive characteristics (10.02 W/m^2 in spring and 9.89 W/m^2 in summer) reveal the energy redistribution role of dust aerosols, which attenuate surface radiation while trapping energy in the atmospheric system through absorption processes.

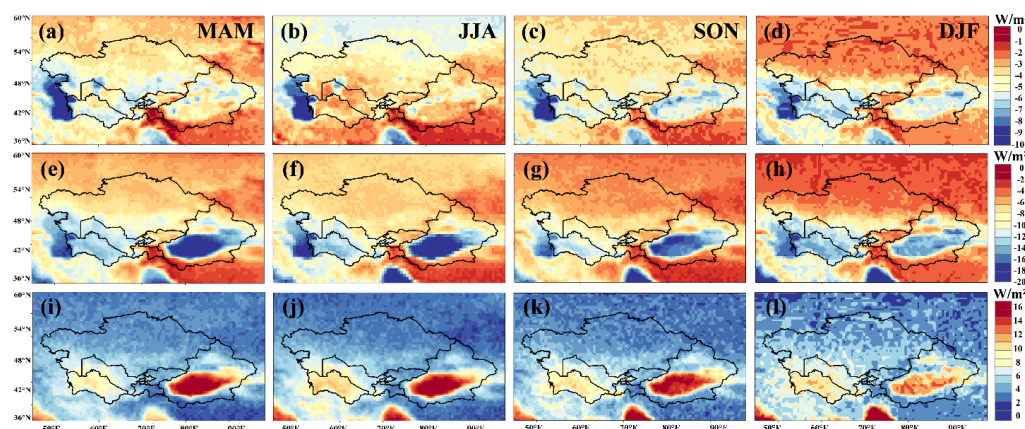


Figure. 7 Seasonal spatial distribution of direct radiative forcing of dust aerosols in Central Asia, 1980-2023, taking into account clear-sky shortwave aerosol direct radiative forcing at the top of the atmosphere (a-d), at the surface (e-h) and in the atmosphere (i-l).



3.3.2 Refinement of aerosol direct radiative forcing in dusty weather

Following an in-depth exploration of the spatial distribution characteristics of dust aerosol direct radiative forcing in the atmosphere, as revealed by MERRA2 reanalysis data, this study was further refined to investigate the temporal divergence characteristics of the direct radiative effect of dust aerosols and its physical mechanisms at typical sites in Central Asia, in conjunction with the SBDART atmospheric transport model (Fig. 8). The dynamic response of site-scale aerosol direct radiative forcing (ADRF) to the atmospheric heating rate is quantitatively analyzed.

Observations indicate that seasonal changes characterize ADRF. The Dushanbe, Issyk-Kul, and Jinghe sites exhibit peak radiative forcing in summer (56.72, 34.22, and 61.17 W/m²) and subsequently decrease to the annual minimum in winter (approximately 2.33 W/m² in Dushanbe and 27.36 W/m² in Jinghe, respectively). This pattern coincides with the period of frequent dust occurrence during summer in western Central Asia, influenced by the westerly wind circulation (Li et al. 2022). Notably, the Kashgar site exhibits a unique spring-dominant pattern (92.99 W/m²), which may be attributed to the synergistic effect of surface exposure following spring snowmelt and vigorous Mongolian cyclone activity, along with the unique sand initiation mechanism in the Tarim Basin.

Changes in the atmospheric heating rate maintain a significant positive correlation with ADRF, confirming the central role of radiation absorption by dust aerosols. The peaks in heating rates at all stations occur during the active dust period: Dushanbe (1.29 K/day in summer) and Jinghe (1.72 K/day in summer) align with the westerly transport paths, while the anomalously high value in spring at Kashgar (2.61 K/day) corresponds to the significant sand uplift event in the Taklamakan Desert. It is noteworthy that the heating rate at Issyk-Kul in spring (0.08 K/day) is significantly lower than that in fall (0.34 K/day), which may be related to the site being shielded by mountainous terrain, limiting vertical transport of dust in spring. This may also impact the accuracy of the results, considering the relative scarcity of observational data at the Issyk-Kul site.



In this study, it was found that the spatial and temporal divergence of regional radiative effects is primarily controlled by two major factors: (1) seasonal modulation of emission intensity in dust source regions, such as the enhanced transport of dust from the westerly rapids to the Aral Sea basin in summer, and (2) modulation of localized atmospheric boundary layer processes, typically manifested as the difference in thermal response between a mountainous site (Issyk-Kul) and a basin site (Kashgar). These findings provide essential observational constraints for improving the dust-radiation parameterization scheme in regional climate models.

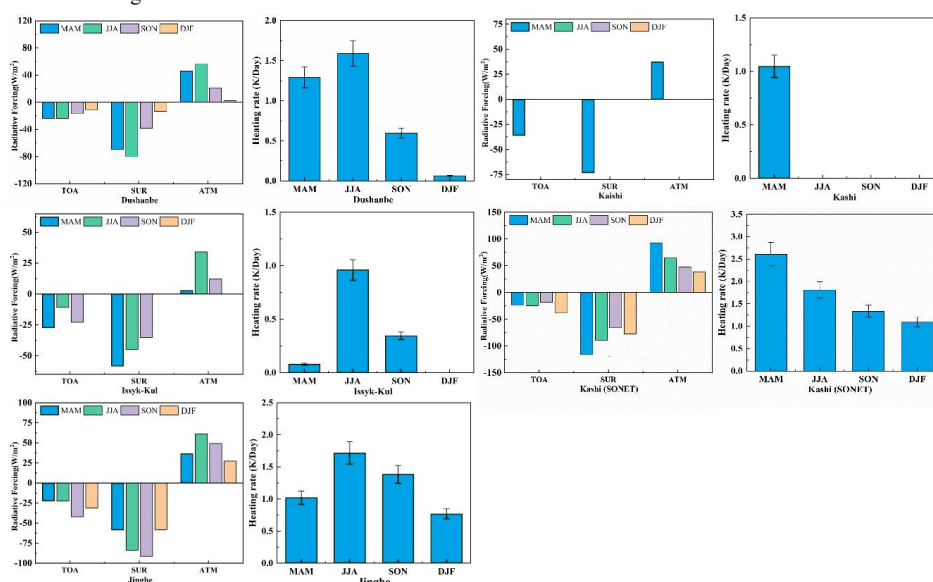


Figure 8 Seasonally averaged shortwave radiative forcing and atmospheric heating rate (including direct radiative forcing at the top of the atmosphere (TOA), the surface (SUR), and the atmosphere (ATM)) for dust aerosols at stations in Central Asia.

Figure 9 provides further refinement of the direct radiative forcing of aerosols at the stations, and the daily variation of the ADRF demonstrates that the radiative forcing at the top of the atmosphere (TOA), at the surface (SFC), and in the entire atmosphere exhibits a clear temporal divergence pattern. The ADRF time series at each site demonstrates a differentiated response: Dushanbe (2011-2023) displays characteristics typical of inland Central Asia, with the radiative forcing at the TOA and SFC oscillating within the interval of $\pm 200 \text{ W/m}^2$ and the atmospheric heating rate peaking at 8 K/day . The short-term variations are primarily controlled by intermittent dust transport induced by disturbances in the westerly jet. The Jinghe site demonstrates a generally stable trend, with transient, strong negative forcing ($\text{SFC} < -400 \text{ W/m}^2$) occurring during extreme dust events. Kashgar, on the other hand, demonstrates significant



temporal variability (TOA/SFC forcing ranging from ± 400 W/m² and heating rates from 0-8 K/day during 2016-2022), particularly high-frequency oscillations during the spring and summer afternoons, which are directly linked to the "afternoon mixed layer development-dust vertical uplift" mechanism unique to the Tarim Basin (Nakamae and Takemi 2022).
Notably, the irregular fluctuations in the enhancement of ADRF observed in recent years (2020-2023) can be attributed to the synergistic effect of changes in surface cover and the frequency of extreme weather events in the arid zones of Central Asia. Specifically, the significant day-to-day variability ($\Delta \text{ADRF} > 50$ W/m²) at the Kashgar site demonstrates the sensitive feedback of aerosol loading on boundary layer thermal processes in the source region of the Taklamakan Desert. These refined observations provide critical process evidence for elucidating the transient effects of dust radiative effects on the regional energy balance.

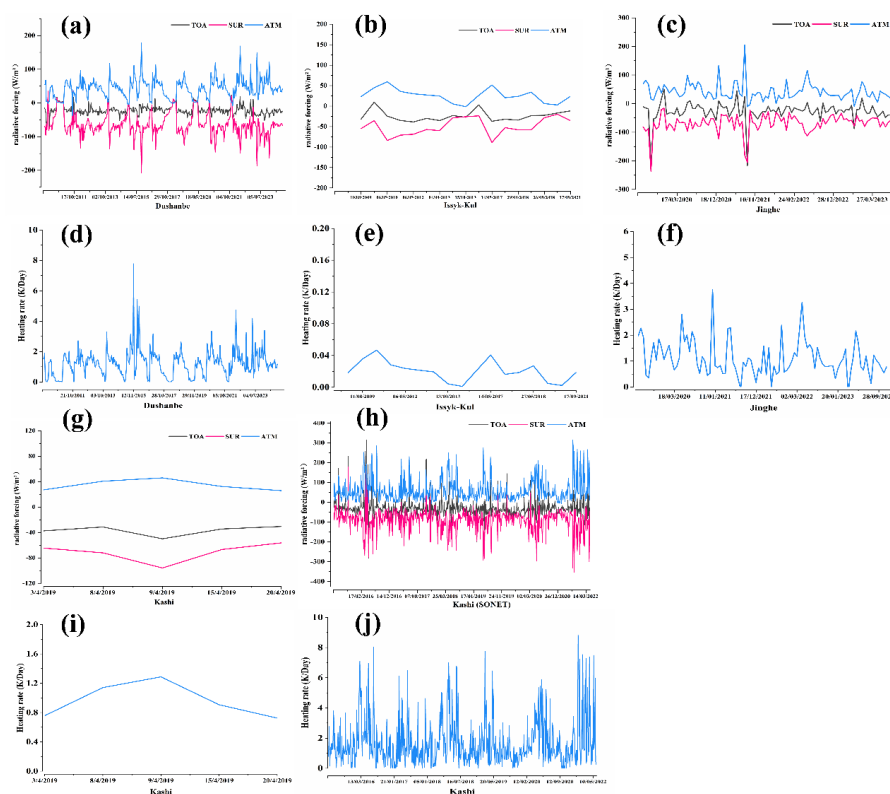


Figure. 9 Short-wave radiative forcing and atmospheric heating rates (including direct radiative forcing and heating rates at the top of the atmosphere (TOA), surface (SUR), and atmosphere (ATM) for daily dust aerosols at stations in Central Asia.



430 4. Conclusion and discussion

431 4.1 Conclusion

432 Dust aerosols are a key factor in the climate system and are characterized by both complexity and regional
433 variations. This paper compares the spatial distribution and temporal trends, forecasting the future trends
434 of the two data sets in Central Asia based on MERRA-2 observational data and the dust balance data
435 generated by the Multi-model Ensemble of CMIP6 (MME). The comparative analysis from 1980 to 2014
436 indicates that the reanalyzed data are consistent with the results of the MME simulations. The primary
437 areas of dust emissions include the Tarim Basin, the Aral Sea dry zone, and the Gobi Desert, with
438 maximum emission fluxes greater than $15 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$. Dust emissions in the Aral Sea region increased
439 substantially (>0.5) over the 34 years. In contrast, the emission fluxes in the Tarim Basin increased at an
440 approximate rate of $-0.3 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}\cdot\text{yr}^{-1}$, with a declining trend.

441 Regarding short-, medium-, and long-term projections, the area of high dust values in Central Asia
442 remains stable in the hinterland of the Aral Sea, Turkmenistan, and along the eastern edge of the Tarim
443 Basin. Short-term emissions in the Aral Sea region range from 17.8 to $26.0 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, showing minimal
444 variation; however, long-term dust emissions in Central Asia under high radiative forcing scenarios (e.g.,
445 SSP585) increase to $387.1 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, representing an increase of up to 94.9% compared to the reference
446 period. Long-term emissions in the Tarim Basin, on the other hand, demonstrate a declining trend,
447 ranging from 18.7% (SSP245 scenario) to 29.3% (SSP370 scenario), particularly in the SSP585 scenario,
448 with short-term emissions at $27.2 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$, which then decrease to $20.1 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ in the long term,
449 representing a decrease of 26.1%.

450 The area of high dust deposition values ($>5 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) overlaps significantly with the emission hotspots.
451 The trend analysis shows that the Aral Sea and the eastern edge of the Caspian Sea exhibit the strongest
452 positive trend ($\Delta S = +0.15 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$), while the southern border experiences a negative trend ($\Delta S = -$
453 $0.10 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$). Under the four future scenarios, the influence of dust deposition spreads across
454 southwestern Central Asia, the southeastern edge of the Tarim Basin, and the Junggar Basin, with the
455 maximum deposition flux exceeding $8 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$. The mean values in the near term (2021-2040) range
456 from $9.3 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ (SSP585) to $10.4 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ (SSP245), and from $9.6 \mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ (SSP370) to 10.0
457 $\mu\text{g}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$ (SSP126) in the long term (2081-2100), with an overall variation of less than 12%.

458 The ADRF of dust aerosols in the clear skies of Central Asia exhibits notable spatial patterns. The TOA



radiative forcing is negative overall, with the lowest value observed in the region occurring in the Caspian Sea ($<-10 \text{ W/m}^2$), followed by the Tarim Basin and around the Aral Sea ($<-8 \text{ W/m}^2$), with the seasonal minimum in spring (-3.32 W/m^2) $>$ summer (-3.21 W/m^2) $>$ fall (-3.07 W/m^2) $>$ winter (-1.94 W/m^2). The SFC reaches its peak at -20 W/m^2 in the Tarim Basin as well as in southwestern Central Asia. The peak of SFC shortwave radiative losses in the Tarim Basin and southwestern Central Asia reaches a maximum of -20 W/m^2 . The atmospheric shortwave radiative forcing aligns with TOA and SFC in spatial distribution, with a maximum value of 10.02 W/m^2 in spring, which closely correlates with the seasonal characteristics of dust activities.

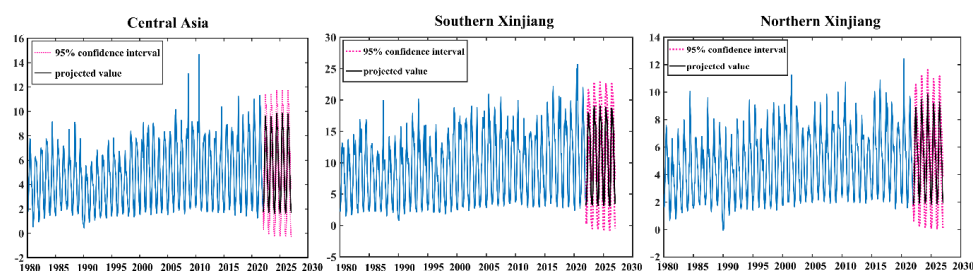
The direct aerosol radiative forcing at the sites simulated by the SBDART model was calculated, peaking at the Dushanbe, Issyk-Kul, and Jinghe sites in summer (56.72 , 34.22 , and 61.17 W/m^2), and then dropping to an annual minimum in winter (about 2.33 W/m^2 in Dushanbe and 27.36 W/m^2 in Jinghe, respectively), with a peak observed at the Kashgar site in spring (92.99 W/m^2). Atmospheric heating rate variations exhibited a strong positive correlation with ADRF. The peaks of heating rates at all sites were observed during the active dust season: Dushanbe (1.29 K/day in summer), Jinghe (1.72 K/day in summer), and Kashgar exhibited a peak in spring (2.61 K/day). Notably, the heating rate at Issyk-Kul was significantly lower in spring (0.08 K/day) than in fall (0.34 K/day), indicating seasonal modulation of dust emission intensity and the influence of local boundary layer processes.

4.2 Discussion

In this study, a coupled "emission-deposition-radiation" pathway model was developed for the dust cycle in Central Asia by integrating the MERRA-2 reanalysis data, CMIP6 Multi-Model Ensemble (MME), and ground-based solar photometer observations, with the aim of revealing the radiative forcing mechanism of dust aerosols on the geosphere and the atmosphere. Given the spatial and temporal heterogeneity of aerosol radiative forcing and the scarcity of observational data, a seasonal autoregressive integrated sliding average model (SARIMA) was introduced to forecast and model regional dust radiative forcing using monthly resolution data from 1980 to 2023.



484 The prediction results (Fig. 10) show that the overall radiative forcing of dust in the Central Asian dry
485 zone from 2024 to 2029 remains quasi-steady, with interannual fluctuations of ADRF in the range of 1.6-
486 9.8 W/m², peaking in 2026, and with no extreme event signals detected. The regional differentiation is
487 characterized by significant features: the southern border, as a substantial radiological response area, is
488 predicted to reach 2.8-18.9 W/m², while the northern border shows a non-stationary trend (1.6-10.0
489 W/m²), increasing and then decreasing, possibly due to the bidirectional modulation of dust emissions
490 by changes in snow cover in the region. Model validation (Supplementary Figure 7) indicated that the
491 residual series of SARIMA(1,1,0)X(1,0,2)₁₂ satisfied the white noise assumption (Ljung-Box Q-test $p >$
492 0.05), the probability distributions conformed to N(0,1) normality (K-S test $D = 0.12$), and the
493 autocorrelation coefficients of the ACFs lay within the 95% confidence intervals, confirming the model's
494 predictive reliability.



495 **Figure.10 Dust and sand aerosol direct radiative forcing SARIMA model predictions.**

496 As a key framework for describing the entire life cycle of the dust cycle, the dust balance encompasses
497 critical aspects such as emission, dispersion, deposition, mass loading, lifetime, and optical depth (DOD).
498 Despite progress in this field, a deeper understanding of the complex interactions between dust with land,
499 vegetation, and climate still faces significant challenges. In particular, the diversity in assumptions
500 regarding dust particle size in CMIP6 models significantly affects the consistency of simulation results,
501 thereby increasing the uncertainty in simulating dust cycling processes.
502 Specifically, the influence of the spatial and temporal variability of the dust particle size spectrum on
503 radiative forcing has often been neglected, particularly the rapid settling of coarse particles ($>10 \mu\text{m}$),
504 which may result in an underestimation of the longwave radiative effect of dust transported over long
505 distances. Therefore, an in-depth investigation of the diversity in the particle size spectrum and its
506 mechanisms on the radiative effect of dust is crucial for understanding regional differences in the dust
507 cycle and their radiative impacts, particularly in the context of extreme climatic events, such as



508 sandstorms and dust storms.

509 In addition, the current uncertainty in the simulation of dust aerosols in CMIP6 models not only persists
510 but is also increasing (Wang et al. 2021), which makes it more difficult to understand the full life cycle
511 of dust and its interactions with other components of the climate system (e.g., radiation budget, cloud
512 processes, precipitation, and atmospheric circulation). This reflects the lack of practical constraints on
513 the links between dust processes, highlighting the urgent need to develop more accurate models of dust
514 aerosol processes to enhance the precision of climate predictions.

515 It is worth noting that existing radiation transfer models, such as SBDART, while valuable in their
516 applications, fail to fully incorporate aerosol-cloud interactions, which are particularly important in
517 regions with high dust concentrations. Aerosol-cloud interactions significantly affect radiative forcing,
518 and neglecting this factor may lead to discrepancies in radiative forcing estimates. Therefore, future
519 research should integrate high-resolution models that account for aerosol-cloud interactions and refine
520 the vertical distribution and particle size spectrum characteristics of dust to improve the accuracy of
521 radiative forcing calculations and enhance scientific understanding of the climatic effects of dust.

522 **Author contributions**

523 All authors contributed to the manuscript and approved the final version. YG designed the study,
524 performed the data analysis, and wrote the original draft. WC, JD, and YR assisted with data collection
525 and software processing. YR also contributed to the validation and interpretation of results. ZZ
526 supervised the research and contributed to manuscript revision and funding acquisition.

527 **Competing interests.**

528 The contact author has declared that none of the authors has any competing interests.

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537 **References:**

538 Braconnot, P., S. Albani, Y. Balkanski, A. Cozic, M. Kageyama, A. Sima, O. Marti & J.-Y. Peterschmitt
539 (2021) Impact of dust in PMIP-CMIP6 mid-Holocene simulations with the IPSL model. *Climate of the*
540 *Past*, 17, 1091-1117.

541 Brown, H., X. Liu, R. Pokhrel, S. Murphy, Z. Lu, R. Saleh, T. Mielonen, H. Kokkola, T. Bergman & G.
542 Myhre (2021) Biomass burning aerosols in most climate models are too absorbing. *Nature*
543 *communications*, 12, 277.

544 Buchard, V., C. Randles, A. Da Silva, A. Darmenov, P. Colarco, R. Govindaraju, R. Ferrare, J. Hair, A.
545 Beyersdorf & L. Ziemba (2017) The MERRA-2 aerosol reanalysis, 1980 onward. Part II: Evaluation and
546 case studies. *Journal of climate*, 30, 6851-6872.

547 Chen, A., C. Zhao & T. Fan (2022) Spatio-temporal distribution of aerosol direct radiative forcing over
548 mid-latitude regions in north hemisphere estimated from satellite observations. *Atmospheric Research*,
549 266, 105938.

550 Dai, Y., P. Hitchcock, N. M. Mahowald, D. I. Domeisen, D. S. Hamilton, L. Li, B. Marticorena, M.
551 Kanakidou, N. Mihalopoulos & A. Aboagye-Okyere (2022) Stratospheric impacts on dust transport and
552 air pollution in West Africa and the Eastern Mediterranean. *Nature Communications*, 13, 7744.

553 Dubovik, O. & M. D. King (2000) A flexible inversion algorithm for retrieval of aerosol optical
554 properties from Sun and sky radiance measurements. *Journal of Geophysical Research: Atmospheres*,
555 105, 20673-20696.

556 Eyring, V., S. Bony, G. A. Meehl, C. A. Senior, B. Stevens, R. J. Stouffer & K. E. Taylor (2016)
557 Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental design and
558 organization. *Geoscientific Model Development*, 9, 1937-1958.



559 Fu, A., W. Li, Y. Chen, Y. Wang, H. Hao, Y. Li, F. Sun, H. Zhou, C. Zhu & X. Hao (2021) The effects
560 of ecological rehabilitation projects on the resilience of an extremely drought-prone desert riparian forest
561 ecosystem in the Tarim River Basin, Xinjiang, China. *Scientific Reports*, 11, 18485.

562 García, O., J. Díaz, F. Expósito, A. Díaz, O. Dubovik, Y. Derimian, P. Dubuisson & J.-C. Roger (2012)
563 Shortwave radiative forcing and efficiency of key aerosol types using AERONET data. *Atmospheric*
564 *Chemistry and Physics*, 12, 5129-5145.

565 Gelaro, R., W. McCarty, M. J. Suárez, R. Todling, A. Molod, L. Takacs, C. A. Randles, A. Darmenov,
566 M. G. Bosilovich & R. Reichle (2017) The modern-era retrospective analysis for research and
567 applications, version 2 (MERRA-2). *Journal of climate*, 30, 5419-5454.

568 Gutmann, E., T. Pruitt, M. P. Clark, L. Brekke, J. R. Arnold, D. A. Raff & R. M. Rasmussen (2014) An
569 intercomparison of statistical downscaling methods used for water resource assessments in the U nited S
570 tates. *Water Resources Research*, 50, 7167-7186.

571 Halthore, R. N., D. Crisp, S. E. Schwartz, G. Anderson, A. Berk, B. Bonnel, O. Boucher, F. L. Chang,
572 M. D. Chou & E. E. Clothiaux (2005) Intercomparison of shortwave radiative transfer codes and
573 measurements. *Journal of Geophysical Research: Atmospheres*, 110.

574 Hetzel, R., S. Niedermann, M. Tao, P. W. Kubik, S. Ivy-Ochs, B. Gao & M. R. Strecker (2002) Low slip
575 rates and long-term preservation of geomorphic features in Central Asia. *Nature*, 417, 428-432.

576 Holben, B. N., T. F. Eck, I. a. Slutsker, D. Tanre, J. Buis, A. Setzer, E. Vermote, J. A. Reagan, Y.
577 Kaufman & T. Nakajima (1998) AERONET—A federated instrument network and data archive for
578 aerosol characterization. *Remote sensing of environment*, 66, 1-16.

579 Holben, B. N., D. Tanré, A. Smirnov, T. Eck, I. Slutsker, N. Abuhassan, W. Newcomb, J. Schafer, B.
580 Chatenet & F. Lavenu (2001) An emerging ground-based aerosol climatology: Aerosol optical depth
581 from AERONET. *Journal of Geophysical Research: Atmospheres*, 106, 12067-12097.

582 Kok, J. F., T. Storelvmo, V. A. Karydis, A. A. Adebisi, N. M. Mahowald, A. T. Evan, C. He & D. M.
583 Leung (2023) Mineral dust aerosol impacts on global climate and climate change. *Nature Reviews Earth*
584 *& Environment*, 4, 71-86.

585 Li, X., Y. Liu, M. Wang, Y. Jiang & X. Dong (2021) Assessment of the Coupled Model Intercomparison
586 Project phase 6 (CMIP6) Model performance in simulating the spatial-temporal variation of aerosol
587 optical depth over Eastern Central China. *Atmospheric Research*, 261, 105747.



588 Li, Y., Y. Song, D. G. Kaskaoutis, X. Zhang, X. Chen, N. Shukurov & R. Orozbaev (2022) Atmospheric
589 dust dynamics over Central Asia: A perspective view from loess deposits. *Gondwana Research*, 109,
590 150-165.

591 Li, Z., H. Xu, K. Li, D. Li, Y. Xie, L. Li, Y. Zhang, X. Gu, W. Zhao & Q. Tian (2018) Comprehensive
592 study of optical, physical, chemical, and radiative properties of total columnar atmospheric aerosols over
593 China: An overview of Sun–Sky Radiometer Observation Network (SONET) measurements. *Bulletin of*
594 *the American Meteorological Society*, 99, 739-755.

595 Lioubimtseva, E. & R. Cole (2006) Uncertainties of climate change in arid environments of Central Asia.
596 *Reviews in Fisheries Science*, 14, 29-49.

597 Liu, J., X. Wang, D. Wu, H. Wei, Y. Li & M. Ji (2024) Historical footprints and future projections of
598 global dust burden from bias-corrected CMIP6 models. *npj Climate and Atmospheric Science*, 7, 1.

599 Mahowald, N. M., S. Kloster, S. Engelstaedter, J. K. Moore, S. Mukhopadhyay, J. R. McConnell, S.
600 Albani, S. C. Doney, A. Bhattacharya, M. A. J. Curran, M. G. Flanner, F. M. Hoffman, D. M. Lawrence,
601 K. Lindsay, P. A. Mayewski, J. Neff, D. Rothenberg, E. Thomas, P. E. Thornton & C. S. Zender (2010)
602 Observed 20th century desert dust variability: impact on climate and biogeochemistry. *Atmos. Chem.*
603 *Phys.*, 10, 10875-10893.

604 Maraun, D., F. Wetterhall, A. Ireson, R. Chandler, E. Kendon, M. Widmann, S. Brien, H. Rust, T.
605 Sauter & M. Themeßl (2010) Precipitation downscaling under climate change: Recent developments to
606 bridge the gap between dynamical models and the end user. *Reviews of geophysics*, 48.

607 Marticorena, B. & G. Bergametti (1995) Modeling the atmospheric dust cycle: 1. Design of a soil-derived
608 dust emission scheme. *Journal of geophysical research: atmospheres*, 100, 16415-16430.

609 Nakamae, K. & T. Takemi (2022) Relationship between the development of a convective mixed layer
610 and dust weather in arid and semi-arid regions of East Asia. *International Journal of Climatology*, 42,
611 3076-3093.

612 Penner, J. E., J. Quaas, T. Storelvmo, T. Takemura, O. Boucher, H. Guo, A. Kirkevåg, J. E. Kristjánsson
613 & Ø. Seland (2006) Model intercomparison of indirect aerosol effects. *Atmospheric Chemistry and*
614 *Physics*, 6, 3391-3405.



615 Pozzer, A., A. De Meij, K. Pringle, H. Tost, U. Doering, J. Van Aardenne & J. Lelieveld (2012)
616 Distributions and regional budgets of aerosols and their precursors simulated with the EMAC chemistry-
617 climate model. *Atmospheric Chemistry and Physics*, 12, 961-987.
618 Ramanathan, V., P. J. Crutzen, J. Kiehl & D. Rosenfeld (2001) Aerosols, climate, and the hydrological
619 cycle. *science*, 294, 2119-2124.
620 Ricchiazzi, P., S. Yang, C. Gautier & D. Sowle (1998) SBDART: A research and teaching software tool
621 for plane-parallel radiative transfer in the Earth's atmosphere. *Bulletin of the American Meteorological*
622 *Society*, 79, 2101-2114.
623 Salvador, P., J. Pey, N. Pérez, X. Querol & B. Artñano (2022) Increasing atmospheric dust transport
624 towards the western Mediterranean over 1948–2020. *NPJ Climate and Atmospheric Science*, 5, 34.
625 Shao, Y., K.-H. Wyrwoll, A. Chappell, J. Huang, Z. Lin, G. H. McTainsh, M. Mikami, T. Y. Tanaka, X.
626 Wang & S. Yoon (2011) Dust cycle: An emerging core theme in Earth system science. *Aeolian Research*,
627 2, 181-204.
628 She, L., Z. Li, G. de Leeuw, W. Wang, Y. Wang, L. Yang, Z. Feng, C. Yang & Y. Shi (2024) Time series
629 retrieval of Multi-wavelength Aerosol optical depth by adapting Transformer (TMAT) using Himawari-
630 8 AHI data. *Remote Sensing of Environment*, 305, 114115.
631 Shen, H., J. Abuduwaili, A. Samat & L. Ma (2016) A review on the research of modern aeolian dust in
632 Central Asia. *Arabian Journal of Geosciences*, 9, 1-16.
633 Sirisha, U. M., M. C. Belavagi & G. Attigeri (2022) Profit prediction using ARIMA, SARIMA and
634 LSTM models in time series forecasting: A comparison. *IEEE Access*, 10, 124715-124727.
635 Song, Q., Z. Zhang, H. Yu, P. Ginoux & J. Shen (2021) Global dust optical depth climatology derived
636 from CALIOP and MODIS aerosol retrievals on decadal timescales: regional and interannual variability.
637 *Atmospheric Chemistry and Physics*, 21, 13369-13395.
638 Tanaka, T. Y. & M. Chiba (2006) A numerical study of the contributions of dust source regions to the
639 global dust budget. *Global and Planetary Change*, 52, 88-104.
640 Tao, M., L. Chen, J. Wang, L. Wang, W. Wang, C. Lin, L. Gui, L. Wang, C. Yu & Y. Wang (2022)
641 Characterization of dust activation and their prevailing transport over East Asia based on multi-satellite
642 observations. *Atmospheric research*, 265, 105886.



643 Tegen, I., M. Werner, S. Harrison & K. E. Kohfeld (2004) Relative importance of climate and land use
644 in determining present and future global soil dust emission. *Geophysical research letters*, 31.
645 Wallace, J. M. & P. V. Hobbs. 2006. *Atmospheric science: an introductory survey*. Elsevier.
646 Wang, X., Y. Chen, Z. Li, G. Fang, F. Wang & H. Liu (2020) The impact of climate change and human
647 activities on the Aral Sea Basin over the past 50 years. *Atmospheric Research*, 245, 105125.
648 Wang, X., J. Liu, H. Che, F. Ji & J. Liu (2018) Spatial and temporal evolution of natural and
649 anthropogenic dust events over northern China. *Scientific Reports*, 8, 2141.
650 Wang, Z., L. Lin, Y. Xu, H. Che, X. Zhang, H. Zhang, W. Dong, C. Wang, K. Gui & B. Xie (2021)
651 Incorrect Asian aerosols affecting the attribution and projection of regional climate change in CMIP6
652 models. *npj Climate and Atmospheric Science*, 4, 2.
653 Wu, T. & B. E. Boor (2021) Urban aerosol size distributions: a global perspective. *Atmospheric*
654 *Chemistry and Physics*, 21, 8883-8914.
655 Xu, M., X. Wang, T. Sun, H. Wu, X. Li & S. Kang (2019) Water balance change and its implications to
656 vegetation in the Tarim River Basin, Central Asia. *Quaternary International*, 523, 25-36.
657 Zhang, X.-X., C. Claiborn, J.-Q. Lei, J. Vaughan, S.-X. Wu, S.-Y. Li, L.-Y. Liu, Z.-F. Wang, Y.-D. Wang
658 & S.-Y. Huang (2020) Aeolian dust in Central Asia: Spatial distribution and temporal variability.
659 *Atmospheric Environment*, 238, 117734.
660 Zhao, A., C. L. Ryder & L. J. Wilcox (2022) How well do the CMIP6 models simulate dust aerosols?
661 *Atmospheric Chemistry and Physics*, 22, 2095-2119.
662 Zhao, A., L. J. Wilcox & C. L. Ryder (2024) The key role of atmospheric absorption in the Asian summer
663 monsoon response to dust emissions in CMIP6 models. *Atmospheric Chemistry and Physics*, 24, 13385-
664 13402.
665 Zhao, Y., X. Yue, Y. Cao, J. Zhu, C. Tian, H. Zhou, Y. Chen, Y. Hu, W. Fu & X. Zhao (2023) Multi-
666 model ensemble projection of the global dust cycle by the end of 21st century using the Coupled Model
667 Intercomparison Project version 6 data. *Atmos. Chem. Phys.*, 23, 7823-7838.
668