



Diffusion kinetics of ^3He in pyroxene and plagioclase and applications to cosmogenic exposure dating and paleothermometry in mafic rocks.

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10 **Abstract.** In this study, we investigate the diffusivity of cosmogenic ^3He in a variety of plagioclase and pyroxene compositions, and its application to paleothermometry and exposure dating in these minerals, through stepwise degassing experiments. While cosmogenic ^3He has been utilized for exposure dating in pyroxene for decades due to its retentivity, plagioclase, often found along with pyroxene in mafic rocks, is generally less retentive of cosmogenic noble gas. However, the diffusivity of ^3He in either plagioclase or pyroxene has not yet been measured quantitatively. A challenge in measuring
15 diffusion kinetics by step-degassing experiments in poorly retentive minerals is the fact that significant amounts of He can be lost prior to the experiment. To address this issue, we apply a forward ‘multiple diffusion domain’ (MDD) inversion model that includes model predictions of initial gas loss during irradiation and storage of the samples to account for this observation and add constraints to the diffusion parameters. We find that ^3He diffusivity in plagioclase appears to be highly variable. This variability can be explained by the MDD inversion models’ inability to constrain the diffusion parameters when
20 significant gas has been lost during irradiation and/or prolonged storage prior to experiment analysis, resulting in an overestimation of ^3He retentivity. Plagioclase samples that were kept frozen after irradiation to limit the initial gas loss yielded the most reliable estimate of diffusion kinetics. We find that ^3He in plagioclase is diffusively lost at Earth’s surface temperatures on a timescale of hundred years, and therefore, unsuitable for surface temperature paleothermometry. Contrary, we find cosmogenic ^3He in pyroxene to be retentive at Earth’s surface temperatures on a million-year timescale.

25 1 Introduction

We measured ^3He diffusivity in various plagioclase and pyroxene grains to determine their diffusion kinetic parameters and relevance to the application of surface temperature thermochronology and exposure dating. Cosmogenic ^3He has been used for cosmogenic nuclide dating for decades and is produced in mineral matter near Earth’s surface by spallation and muon production by cosmic rays. The utilization of ^3He in pyroxene for exposure dating is well established (e.g., Craig and Poreda,
30 1986; Bruno et al., 1997; Ackert et al., 1999; Schäfer et al., 1999; Balter-Kennedy et al., 2020; Tremblay et al., 2014b), on



the assumption that ^3He is well retained at Earth's surface temperatures. While these studies do not quantify the diffusion kinetics of ^3He , they do indeed confirm that ^3He is well retained, as measured ^3He concentrations in pyroxene are consistent with exposure ages determined from other geological evidence. Plagioclase is often found along with pyroxene in mafic rocks, but is less retentive of cosmogenic noble gases (Megrue, 1966; Cerling, 1990; Bruno et al., 1997; Gournbet et al., 2012; 35 Cassata and Renne, 2013; Tremblay et al., 2017). This limits the application of plagioclase for cosmogenic nuclide exposure dating but provides potential usefulness for surface temperature thermochronometry, which relies on temperature-dependent diffusion of noble gases in mineral matter (Tremblay et al., 2017). However, the diffusion kinetics of cosmogenic ^3He in both pyroxene and plagioclase have not yet been determined.

40 Using predictive models of cosmogenic noble gas production and diffusive loss through time and temperature in mineral matter, a quantitative constraint on the thermal history of an exposed rock surface can be inferred from a surface exposure age derived from, e.g., cosmogenic nuclide dating (Tremblay et al., 2014b, a; Tremblay et al., 2017; Gribenski et al., 2022). In other words, even if the loss of a diffusively mobile nuclide precludes using it to determine the exposure age, it might allow one to determine the surface temperature. Currently, paleothermometry relies on the open-system behavior of 45 cosmogenic ^3He in quartz at Earth surface temperatures (Brook et al., 1993; Shuster and Farley, 2005). Since He is also produced by the decay of uranium and thorium, its diffusivity in mineral matter has been studied extensively as chronometers in terrestrial and extraterrestrial materials. In a review, Baxter (2010) summarizes the current literature on diffusion kinetics of noble gasses. While the diffusion kinetics of ^3He and ^4He have been determined in various minerals, we are not aware of any measurements of He diffusion in plagioclase. The only existing constraint on He diffusion in pyroxene 50 was reported by Lippolt and Weigel (1988), who determined ^4He in pyroxene to be generally retentive at low temperatures.

The diffusivity of noble gases can be measured from a stepwise degassing experiment of a single mineral grain in which the grain is repeatedly heated at steadily increasing temperatures until the included gas is exhausted. In simple cases, diffusion coefficients derived from the fractional release of gas in each step form a linear array in an Arrhenius plot, in which the slope 55 of the array is proportional to the material activation energy (e.g., Gournbet et al., 2012). When some data diverge from this linear array, diffusion kinetics are often estimated simply by fitting a straight line to the linear portion of the data set and disregarding any divergent data (e.g., Shuster and Farley, 2005; Gournbet et al., 2012; Cassata and Renne, 2013). Because thermochronological studies rely on applying extrapolation of laboratory-determined diffusion properties to geological settings, the magnitude of the diffusion parameters, such as activation energy (E_a) and pre-exponential factor (D_0), are 60 particularly significant for determining temperature-dependent gas release. However, since a subset of the step-heating experimental data is used for determining these parameters, the excluded data may hold important information on diffusion kinetics that is not accounted for.



Previous diffusion experiments of ^{21}Ne , ^{37}Ar , and ^{39}Ar in plagioclase and pyroxene are known to show complex diffusion
65 behavior, expressed as diffusion coefficients that form multiple linear or curvilinear arrays on Arrhenius plots (Gourbet et al., 2012; Cassata and Renne, 2013; Tremblay et al., 2017), and has commonly been explained by so-called ‘multiple diffusion domain’ (MDD) models to describe the overall diffusion kinetics (Lovera et al., 1989; Lovera et al., 1997). In the MDD model, the nonlinearity is explained by the systematic exhaustion of sub-grain domains that are usually envisioned to consist of the same material with the same activation energy, but having different effective sizes. Since its development, the
70 MDD model has been widely used in thermochronology (e.g., Reiners et al., 2005), including surface temperature thermochronology (Gribenski et al., 2022).

Originally, the MDD model was developed to determine the diffusion kinetics for minerals with high closure temperatures compared to Earth’s surface temperatures. Thus, zero degassing would occur during storage, and gas would only be released
75 during experimental heating steps. However, applying any method for determining diffusion kinetics (MDD or single Arrhenius domain) to minerals that have closure temperatures near room temperature requires accounting for the fact that the sample loses gas at ambient conditions. Any loss of gas before a step-heating experiment results in an initially heterogeneous distribution of the gas concentration, which in turn yields an anomalously elevated slope at low temperatures on Arrhenius arrays calculated assuming an initially homogeneous distribution (Shuster and Farley, 2005; Tremblay et al., 2014b;
80 Tremblay et al., 2017).

Previous studies of diffusion kinetics by step-degassing in non-retentive minerals (e.g., ^3He in quartz) have found that the first few initial heating steps $<100^\circ\text{C}$ often lie below the linear array in the Arrhenius plot and exhibit increasing apparent diffusivity (Tremblay et al., 2014b). This has been explained by the diffusive loss of gas either during or after proton
85 irradiation (Shuster and Farley, 2005). As a result, the first few heating steps are often neglected from the regression line used to determine the activation energy (e.g., Gourbet et al., 2012; Cassata and Renne, 2013), or two different diffusion kinetics are modeled separately for a higher and lower temperature domain (Tremblay et al., 2014b).

To evaluate the use of cosmogenic ^3He in paleothermometry and exposure dating, we therefore, measured the ^3He diffusivity
90 through a stepwise degassing experiment in plagioclase and pyroxene fragments from the Ferrar Dolerite in Antarctica and grains from a variety of feldspar and pyroxene end-member compositions. We then apply a similar optimization approach as Gorin et al. (2024) for determining the diffusion kinetics, by predicting the gas releases using a MDD forward inversion model and comparing the modeled results with the measurements. In addition, we include model predictions of gas loss during and after irradiation. We find that accounting for such diffusive loss during and after irradiation has a significant
95 effect on both (i) the diffusion kinetics inferred from model inversions, and (ii) inferred geologic conditions that attend cosmogenic production.



2 Method

100 We conducted stepwise diffusion experiments on single feldspar and pyroxene grains from various rock samples. For a diffusive mineral, natural cosmic-ray-produced ^3He in a sample grain will have a nonuniform distribution and be unsuitable for step-heating experiments used for determining diffusion kinetics. Therefore, all sample grains were either proton- or neutron-irradiated. Proton irradiation of the grains results in the production and homogeneous distribution of ^3He through nuclear reactions similar to those produced from naturally induced cosmic rays, but with an order of magnitude higher
105 concentration than naturally produced cosmogenic ^3He . Neutron irradiation is intended for the production of ^{39}Ar from K for purposes of $^{40}\text{Ar}/^{39}\text{Ar}$ dating, and the neutron-irradiated samples described here were irradiated for that purpose. However, ^3He is also produced as a byproduct, so these samples can also be used for ^3He diffusion experiments. In both cases, the production of ^3He at levels well above naturally occurring concentrations allows for more precise and detailed stepwise degassing experiments on single grains (Shuster et al., 2004; Shuster and Farley, 2005; Tremblay et al., 2014b).

110 2.1 Samples and irradiation

In this study, we analyzed feldspar and pyroxene grains from two distinctly different sample groups (Table 1). One group of samples consists of plagioclase and pyroxene grains from the Ferrar Dolerite, an intrusive gabbroic rock found in the Transantarctic Mountains (TAM), Antarctica. The Ferrar consists predominantly of subequal parts of plagioclase and pyroxene (Harvey, 2001; Elliot and Fleming, 2021), and its high resistance to weathering makes it one of the most dominant
115 rocks exposed at the surface in Antarctica. Therefore, it is commonly used for exposure-dating applications. For this study, we obtained grains from two separate Ferrar Dolerite samples. The LABCO sample is a core sample from the Labyrinth in the McMurdo Dry Valleys, collected in 2009 as part of the CRONUS-Earth project and described in Balter-Kennedy et al. (2023). The ROB samples are from glacial moraine boulders collected at Roberts Massif in the central TAM for ^3He exposure dating (Balter-Kennedy et al., 2020).

120 Each rock sample was crushed, and individual crystal fragments of plagioclase and pyroxene devoid of any obvious microfractures under a microscope were handpicked. Numerous grains from a given sample were packed in aluminium foil, put into small (9x3 mm) HDPE capsules, stacked into quartz tubes, and proton-irradiated for ~8 hours at temperatures of <45°C at the Harvard Cyclotron Facility following an established methodology (Shuster et al., 2004). The LABCO
125 plagioclase and pyroxene samples were proton irradiated in December 2014 and have since been stored at room temperature (~21°C). The ROB samples were proton-irradiated in March 2024, cooled to -25°C after irradiation, and kept frozen at -20°C until analysis.



Another group of samples consists of various feldspar and pyroxene grains having known composition (Table 1). These samples have been the subject of extensive mineralogical research for decades and used in previous diffusion studies for determining the diffusion kinetics of Ar (Cassata et al., 2011; Cassata and Renne, 2013) and Ne (Gourbet et al., 2012; Tremblay et al., 2017). The feldspar samples were neutron-irradiated by Cassata and Renne (2013, see reference for more details) and pyroxene samples by Cassata et al. (2011, see reference for more details). The temperature during neutron irradiation is unknown, but has been estimated to reach at least 150°C and perhaps as high as 200°C for a duration of either 50 hr and 100 hr (Personal communication with the Cadmium-Lined In-Core Irradiation Tube (CLICIT) facility at Oregon State University where samples were irradiated). Since irradiation, the samples have been stored at room temperature (~21°C) at the Berkeley Geochronology Center (BGC) for variable durations (Table 1).

Table 1: Sample description

Sample	Location	Phase	Irradiation Type	Storage time since irradiation* (yr)
<i>Plagioclase</i>				
LABCO (Da, Db)	Gabbro, Labyrinth Valley, Antarctica	Plagioclase	Proton	7.70, 8.13
ROB (Da, Db)	Gabbro, Roberts Massif, Antarctica	Plagioclase	Proton	0.16, 0.17 (-20°C)
FCs	Rhyolitic ignimbrite, Fish Canyon Tuff, CO USA	Sanidine	Neutron	12.64
ML	Rhyolitic Ash, Mono Lake, CA USA	Oligoclase	Neutron	13.72
SURTp	Basaltic lava, Surtsey, Iceland	Labradorite	Neutron	13.85
SW-1	Anorthosite, Stillwater Complex, MT, USA	Bytownite	Neutron	12.82
<i>Pyroxene</i>				
LABCO (Db, Dc)	Gabbro, Labyrinth Valley, Antarctica	Pyroxene	Proton	7.76, 7.78
ROB	Gabbro, Roberts Massif, Antarctica	Pyroxene	Proton	0.18 (-20°C)
GEM CPx	Unknown	Clinopyroxene	Neutron	12.70
GEM OPx	Unknown	Orthopyroxene	Neutron	12.73

*Assumed to be at room temperature (~21°C), except for ROB (Da, Db)

2.2 Step-heating experiment

Single grains used for step-heating experiments were selected under a microscope after irradiation. To eliminate any complications arising from variation in grain size, shape, and other properties, the grains were selected by the following criteria: i) rounded shape, ii) single crystal, and iii) void of any visually apparent fractures and impurities. Photos from each selected grain were obtained and its dimensions were recorded before and after the experiment (Fig. S1-2).



At BGC, irradiated single sample grains were loaded into small (~5mm) Pt-Ir alloy or Ta packages connected to a type-K thermocouple and attached to the “Ohio” noble gas mass spectrometer extraction line under high vacuum ($< 10^{-8}$ torr). The package containing the sample is heated with a 30W diode laser focused on the package to obtain uniform heating. The laser and thermocouple are connected in a control loop using a Watlow PID controller that achieves a typical accuracy of $< 3^{\circ}\text{C}$.
 150 The exact duration and temperature of the heating are determined by the assigned heating schedule (Table S1-2) and adjusted for the type of mineral being analyzed. None of the diffusion experiments resulted in melting or visual changes for either plagioclase or pyroxene grains.

All diffusion experiments included an initial heating step of 25°C for 30-60 min depending on mineral composition, followed
 155 by an increasing heating schedule that consists of a few repeating temperature steps where heating duration is increased, and at least one retrograde at higher temperatures. While all steps are sensitive to the distribution of domain sizes, retrograde temperature cycling emphasizes their presence and is useful for constraining the activation energy (Lovera et al., 1989; McDougall and Harrison, 1999; Shuster and Farley, 2005).

The ^3He measurement method follows that described in Tremblay et al. (2014b). Throughout the heating experiment, we measured 5-8 sample chamber blanks as a 30-60 min ‘heating’ step at 25°C . The initial heating step at this temperature commonly yielded significant amounts of ^3He and was not considered a blank. A linear fitting of the sample chamber blanks was then subtracted from the raw signal in other heating steps. The magnitude of the blank varies between each experiment but averages 10^4 atoms of ^3He . For heating steps where the gas release is insignificant from the blank level, the measured
 165 value is set to zero and included in the data set. The measured fractional gas release and the duration during each heating step are then used to calculate diffusion coefficients using the equations for spherical geometry derived by Ginster and Reiners (2018).

2.3 Forward MDD model

We fit a forward multiple diffusion domain (MDD) model to the step-degassing data to determine the diffusion kinetics for
 170 the various mineral grains. One of the main assumptions for the forward MDD model is that the same activation energy applies to all sub-grain domains and that each domain can exhibit varying diffusive length scales (e.g., different radii). The diffusion domain, n , in a grain having multiple diffusion domains, can be described by (i) the activation energy, E_a (kJ mol^{-1}), which applies to all domains, (ii) the pre-exponential factor, D_0/a_n^2 (s^{-1}), and (iii) the proportion of the total gas contained in the domain. The fractional loss is calculated using the equations in Fechtig and Kalbitzer (1966, Equation 4a-c) for each
 175 domain, and scaled by the proportion of gas contained in the domain, such that the total cumulative fractional release is the sum of all releases within each domain.

Given a step-heating experiment, the forward MDD model then predicts the fractional gas release and calculates the apparent



diffusion constant, D/a^2 (s^{-1}) (Fechtig and Kalbitzer, 1966, Equation 5a-c), during each heating step for n number of domains.

180 This allows for a direct comparison of the measured and predicted fractional release between all data points for a given diffusion experiment. When a single domain ($n = 1$) is assigned to the forward MDD model, this converts to a simple Arrhenius linearity with 3 free parameters. For all other ($n > 1$) domains, this will convert to a non-linear Arrhenius behavior, with $2+2n$ free parameters in the MDD model.

2.3.1 Diffusion loss during irradiation and storage

185 Since these samples have experienced a long period of storage time between irradiation and analysis (Table 1), it is expected that non-retentive samples have lost gas in the outer grain boundary due to years of storage, even at room temperature ($\sim 21^\circ\text{C}$). Therefore, in the forward MDD model, we account for this initial gas loss prior to the step heating experiment (referred to as initial loss). Within the forward MDD model, the initial irradiation and storage loss are temporally added as two separate heating steps to be included in the calculation of cumulative fractional release, such that the predicted fractional
190 release from the first experimental heating steps reflects such initial loss. Once the predicted cumulative fractional release is determined, the two additional heating steps are removed, and the fractional release from each of the experimentally heating steps is re-scaled to sum 1 for direct comparison to the measured fractional gas release data points. When included, the two steps increase the degree of freedom in the forward MDD optimization model.

195 Gas loss during storage (we refer to this as storage loss) occurs between the irradiation and time of the diffusion experiment and is simply dependent on the diffusivity of gas within the mineral, duration, and temperature of the storage. Gas loss during irradiation (we refer to this as irradiation loss) is more complex and is assumed to occur for diffusive samples as the temperature reaches up to 45°C for 8 hr during proton irradiation (Shuster et al., 2004). However, the exact temperature throughout the irradiation is not well known. For the samples that have been proton-irradiated, we include a modeled initial
200 heating of an estimated 40°C for 8 hr in the forward MDD model. While we acknowledge that there is simultaneous production and diffusion of ^3He during the irradiation, we only account for diffusion loss (see discussion for more details), resulting in a maximum gas loss value. The true build-up and amount of gas present after irradiation lies between the two end-members: including diffusion loss only and not including any production or gas loss at all. The overall effect on the total amount of ^3He is insignificant unless the sample is diffusive at room temperature. However, in such a case, any storage (e.g.,
205 1 month) of the sample at room temperature prior to step heating analysis would suppress any measurable signal. The group of samples with known feldspar and pyroxene compositions were all neutron-irradiated for either 50 hr or 100 hr (Cassata et al., 2011; Cassata and Renne, 2013) in which temperatures could reach up to 200°C . All samples contain measurable He, indicating that for a diffusive mineral, simultaneous production and gas loss do result in a measurable amount of ^3He . Since the exact temperature of neutron-irradiation is unknown, we do not account for any irradiation loss when optimizing for the
210 diffusion kinetics but instead only account for storage at room temperature (See discussion section 4.2.1 for more details). We acknowledge that for neutron irradiation (i) significantly higher temperatures are reached during irradiation than proton



irradiation, (ii) the production reaction of ^3He is not well understood, and (iii) zoning in plagioclase may cause initial heterogeneous ^3He distribution. Consequently, for a diffusive mineral, this results in an initial non-homogeneous distribution of ^3He in the grain during irradiation. While this is not ideal and it is clear that the high temperatures reached during irradiation make neutron irradiation a poor choice for measuring diffusion kinetics of poorly retentive minerals, we use the approach here because it allows us to take advantage of samples that were already irradiated for other purposes and to explore the diffusion kinetics in a variety of feldspar and pyroxene compositions.

In Fig. 1 we show the effect of a non-homogeneous gas distribution from proton-irradiation followed by storage loss modeled for a diffusive mineral by applying arbitrary MDD diffusion kinetics for three domains to a simple increasing heating schedule with a high-temperature retrograde. Any gas loss prior to the diffusion experiment results in an apparent diffusivity that is lower than expected and increases for repeated temperature heating steps $< 250^\circ\text{C}$. Accounting for irradiation loss of 40°C for 8 hr results in a gas loss of 0.51% prior to the experimental analysis (Fig. 1a). Such the same sample would lose 6.26% and 18.64% of its gas during storage only at room temperature (21°C) for 1 yr and 10 yr, respectively (Fig. 1b). An interesting aspect of this effect is the observed change in apparent activation energy for low-temperature heating steps when determined from a linear regression. The magnitude of bias in apparent activation energy then depends on the amount of initial data points excluded from the regression. As it is not clear how to choose these points in an unbiased way, this raises a significant disadvantage in the conventional linear regression approach. Our approach, instead, aims to fit all data, thus eliminating the need for potentially arbitrary data selection.

As we show in the following sections, our forward MDD model inversion that includes free parameters for storage loss showed that the majority of gas produced in the neutron-irradiated feldspar samples was lost during storage. Because this was unexpected, we conducted an additional experiment in which we measured total ^3He and ^4He concentrations in a set of Fish Canyon sanidine (FCs) samples with different storage durations. For many decades, the FCs grains (Cebula, 1986) have routinely been used as a reference mineral for argon studies (Phillips et al., 2022, see references therein). Therefore, at BGC, neutron-irradiated FCs grains are available from various irradiation batches going back 15 years, which all have experienced comparable irradiation but varying storage duration. To evaluate major gas loss during prolonged storage time for diffusive minerals, we measured the total ^3He and ^4He concentrations in 7 sample sets of grains having experienced between 1.86 and 14.69 years of storage time at $\sim 21^\circ\text{C}$.

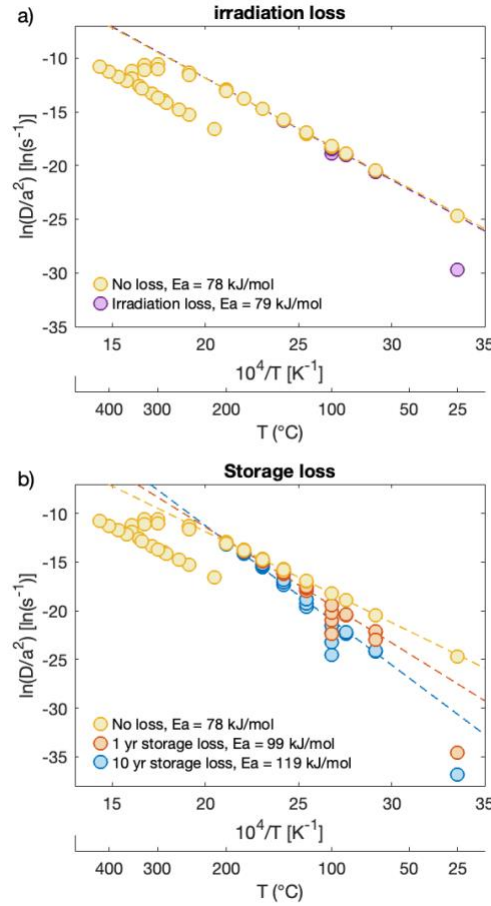


Figure 1: Calculated apparent diffusivities for (a) proton-irradiation and (b) storage at room temperature (21°C) for a diffusive mineral. All scenarios use an activation energy of 80 kJ mol⁻¹, and three diffusive domains with an ln(D/a²) of 8, 6, and 3 s⁻¹, and a fractional contribution of 0.70, 0.27, and 0.03 to the total gas release, respectively. The incorrect, but apparent activation energy obtained from a linear regression of the low-temperature heating steps is reported in the legend and fitted using the same data point for all scenarios. The difference in apparent diffusivity and activation energy is solely due to multiple diffusion domains, gas loss during irradiation, and storage time before the step-heating experiment.

2.3.2 Fitting statistics

We compare the observed fractional gas release from the diffusion experiment (f_{obs}) with that of the predicted (f_{pred}) MDD forward model using evenly weighted fractional misfit statistics based on the gas released from each heating step, i , and is described by the misfit,

$$M = \sum_{i=1}^N \frac{f_{i,obs} - f_{i,pred}}{f_{i,obs}}, \quad (1)$$

The diffusion kinetics is determined by the best fit found using a constrained nonlinear multivariable optimization function (fmincon) within a global search multi-start algorithm, MultiStart (Ugray et al., 2007) in MATLAB, while optimizing for the



free parameters: (i) Total modeled gas release in atoms, (ii) activation energy, (iii) the diffusion constant for each diffusive
 255 domains present and (iv) the proportional gas releases from each of the domains. When diffusion loss during irradiation and
 storage loss is applied to the forward MDD model, the fractional release from these initial steps also becomes free
 parameters.

The model fitting is constrained by the following: (i) The domains are kept in order, such that $\ln D_0/a_1^2 > \ln D_0/a_2^2 > \dots >$
 260 $\ln D_0/a_n^2$, and where n is the number of domains, (ii) the sum of the fractional release from each domain must equal 1, in
 which the fractional release from each domain must be $> 10^{-9}$ as we consider any less in these domains to be insignificant to
 the overall diffusion kinetics of the mineral gas release and, (iii) the total predicted gas release results in a complete
 degassing.

265 The optimal number of domains is then determined by first optimizing the diffusion kinetics for a varying number of
 domains between 2 and 10. For each domain size, the optimization was repeated 20 times for which only the minimum misfit
 results were considered. Then we calculate a reduced misfit (similar to a reduced χ^2) by dividing the misfit value by the
 degree of freedom, which is defined by the number of modeled parameters subtracted from the number of data points used in
 the fitting. The number of modeled domains that result in the lowest reduced misfit value is found to be the optimal number
 270 of domains needed to explain the data set without overfitting.

3 Results

3.1 ^3He diffusion kinetics in plagioclase

Results for ^3He step-heating experiments for the two plagioclase and feldspars sample groups are shown in Table S1 and Fig.
 2. All experiments display nonlinear behavior in the Arrhenius plot. For the samples stored at room temperature prior to
 275 analysis, there is a consistent increase in apparent diffusivity in isothermal heating steps below 250°C. This is consistent with
 gas loss occurring during irradiation and/or storage. The forward-modeled MDD diffusion kinetics are reported in Table 2
 and the predicted apparent diffusivity is shown in Fig. 3 and Fig. 4.

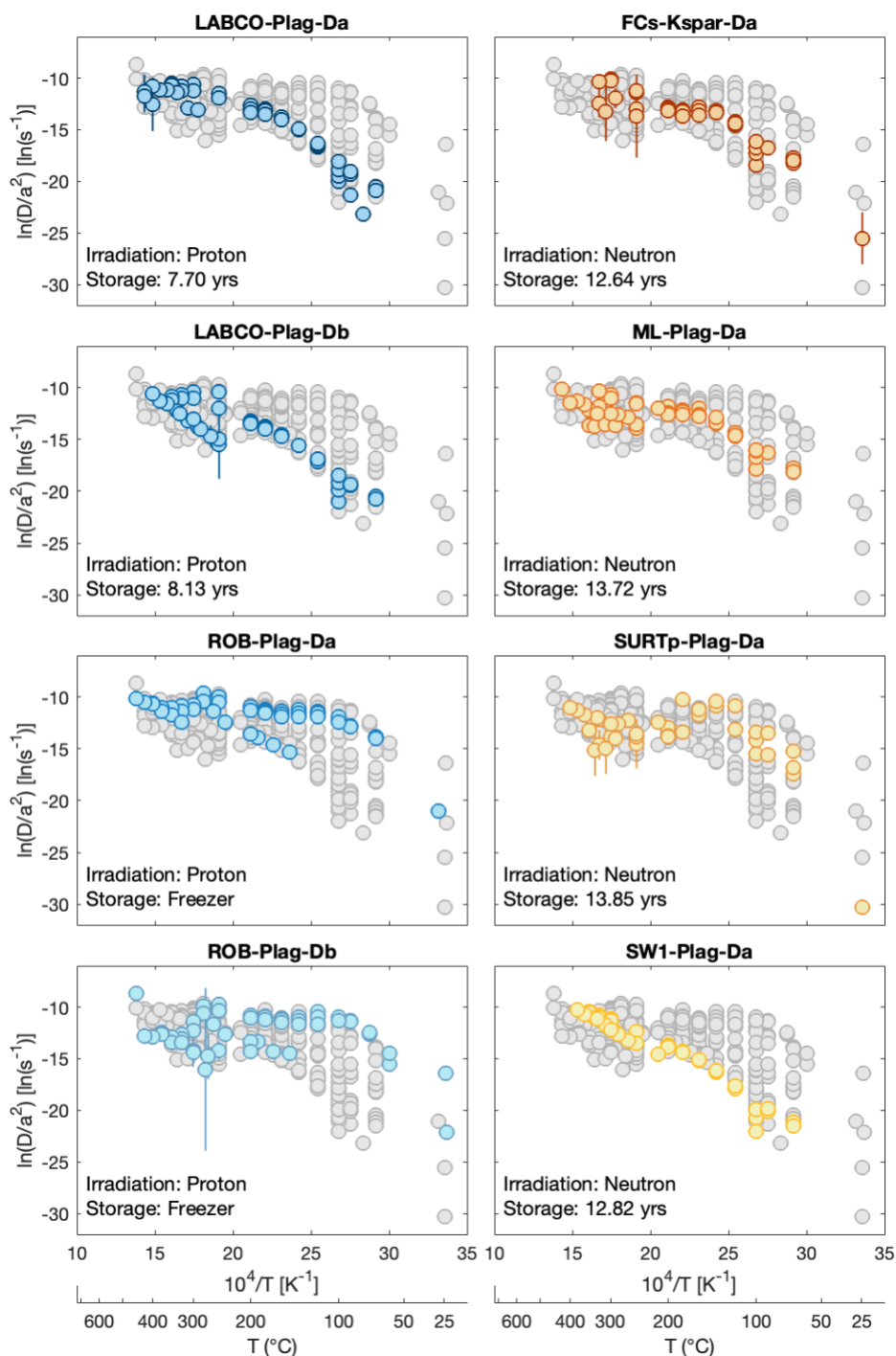


Figure 2: Arrhenius plot for apparent ^3He diffusivity in plagioclase grains. Grey data points represent all plagioclase diffusion experiments for comparison, and colored data points represent a unique sample. Vertical lines show the uncertainty in $\ln(D/a^2)$ and are estimated using error propagation derived by Ginster and Reiners (2018), although the uncertainty is usually smaller than the symbol.

280



The Ferrar Dolerite plagioclase samples, LABCO and ROB, were proton-irradiated, and we, therefore, account for loss during the irradiation stage for both samples and only include storage loss for the LABCO sample, as ROB was kept frozen (ca. -20°C) immediately after irradiation. The total modeled irradiation and storage gas loss for LABCO is ~47%, in which ~2.5% was released during irradiation (Table 2). On the contrary, the ROB samples show 25-36% gas release during irradiation. This variation in the predicted gas loss during irradiation can be attributed to the difference in modeled diffusion kinetics between the two samples, with activation energies of ~90 kJ mol⁻¹ for LABCO and ~67 kJ mol⁻¹ for ROB. If the two Ferrar plagioclase grains are comparable, this discrepancy suggests that the model inversion may be compensating for storage loss by adjusting the activation energy, reflecting a potential trade-off between the two free parameters. This is discussed in more detail below.

The sample group consisting of various feldspar grains has been neutron irradiated. As mentioned in the method section, we do not account for loss during neutron irradiation but only consider the gas lost during storage at 21°C. The Arrhenius behavior and modeled diffusion kinetics for these samples show significant variation depending on the composition. However, the total modeled gas loss during storage exceeds 99% across all feldspar compositions (Table 2). As a result, the step-degassing experiment for these samples only involved ~1% of the total amount of gas that was present at the end of the neutron irradiation.

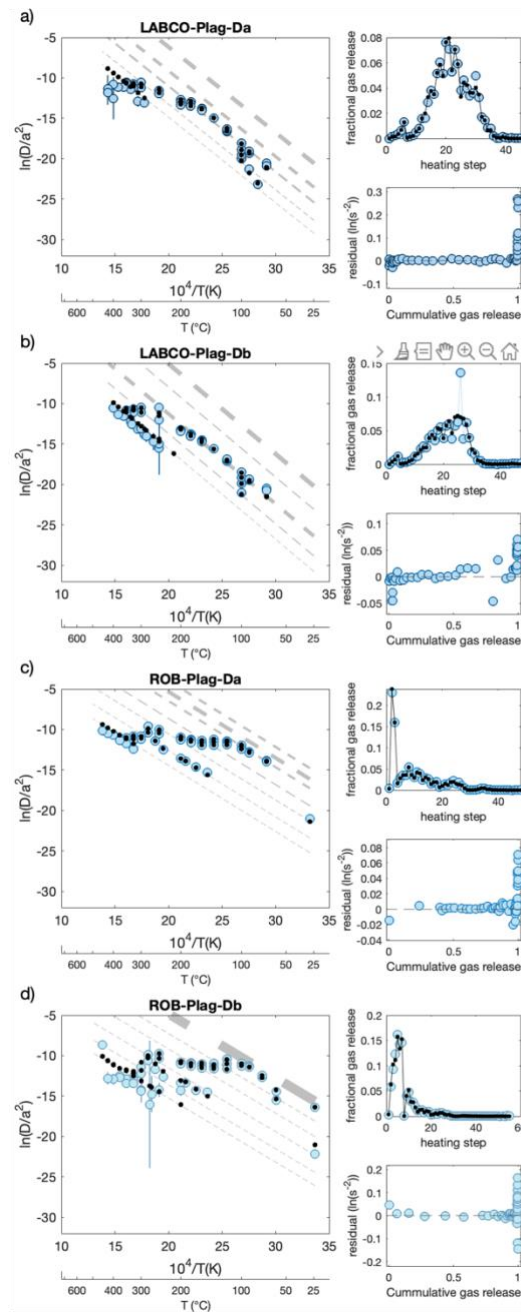
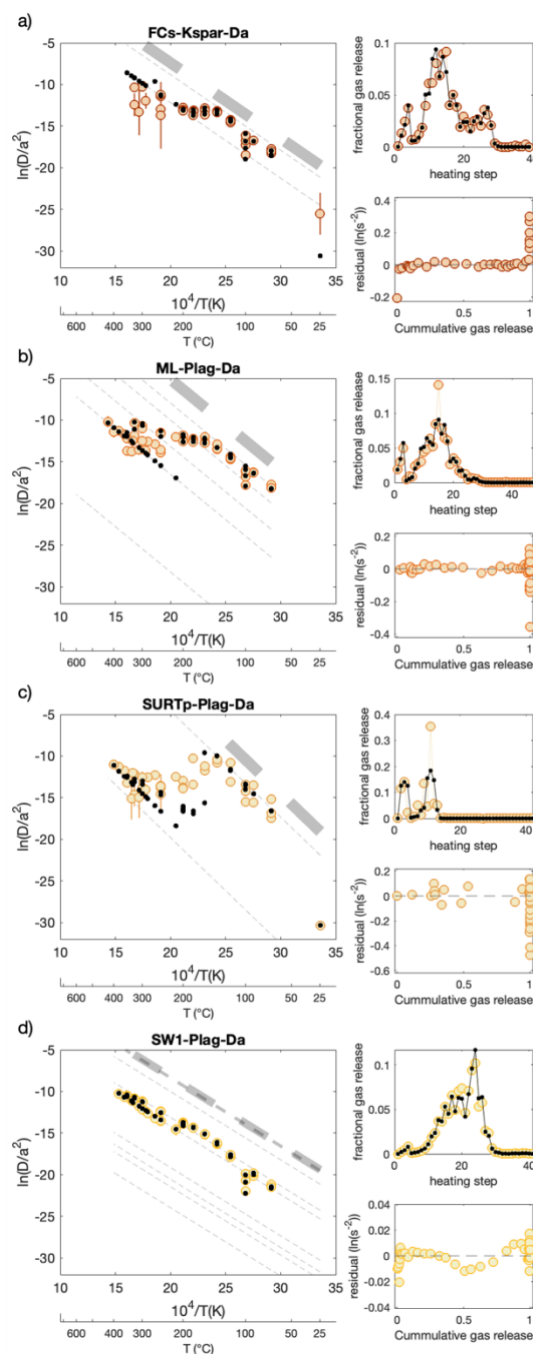


Figure 3: Forward-modeled MDD results for Ferrar Dolerite plagioclase. Each sample result (a-d) is displayed in the Arrhenius, fractional release, and residual plots. Colored circles show the observed apparent ^3He diffusivity from the step-heating experiment. Vertical lines show the uncertainty in $\ln(D/a^2)$ (see description in Fig. 2 for more details). Black dots show the modeled predicted apparent diffusivity from the best reduced misfit, where the grey lines represent the diffusion kinetics of each individual domain modeled, with the line thickness being proportional to the fractional size of the domain. The fractional release plot compares the measured and predicted fractional gas release from each heating step during the diffusion experiment. The residuals are defined as the difference between the calculated $\ln(D/a^2)$ from a given heating step and the expected $\ln(D/a^2)$ from the first MDD model domain at that same temperature and plotted against the cumulative gas release.



310 Figure 4: Forward-modeled MDD results for a variety of feldspar compositions. Each sample result (a-d) is displayed in the
 Arrhenius, fractional release, and residual plots. Colored circles show the observed apparent ^3He diffusivity from the step-heating
 experiment. Vertical lines show the uncertainty in $\ln(D/a^2)$ (see description in Fig. 2 for more details). Black dots show the modeled
 predicted apparent diffusivity from the best reduced misfit, where the grey lines represent the diffusion kinetics of each individual
 domain modeled, with the line thickness being proportional to the fractional size of the domain. See Fig. 3 for more details on
 315 plotting.



Table 2: Summary of modeled MDD diffusion kinetics results for all feldspar grains.

Sample name	Radii (cm)	Misfit	Ea (KJ mol ⁻¹)	ln(D0/a ²) ln(s ⁻¹)	Fractional release	Irradiation loss	Storage loss
LABCO-Plag-Da (protons; 7.67 yr, room temp)	0.0173	8.3	88.1	15.0	0.39	0.0225	0.4456
				12.1	0.29		
				10.2	0.19		
				7.9	0.082		
				6.3	0.040		
LABCO-Plag-Db (protons; 7.70 yr, room temp)	0.0269	8.8	92.7	17.2	0.42	0.0270	0.4405
				14.3	0.11		
				11.0	0.34		
				8.6	0.11		
				6.7	0.018		
ROB-Plag-Da (protons; 0.16 yr, -20°C)	0.0337	7.0	67.9	12.4	0.22	0.2512	-
				11.0	0.36		
				9.9	0.20		
				7.7	0.11		
				5.8	0.075		
				3.7	0.030		
				1.9	6.82 x 10 ⁻³		
ROB-Plag-Db (protons; 0.17 yr, -20°C)	0.0217	12.9	65.9	11.0	0.87	0.3625	-
				8.6	0.069		
				6.7	0.041		
				4.4	0.015		
				2.2	3.56 x 10 ⁻⁴		
				0.6	1.54 x 10 ⁻³		
FCs-Kspar-Da (neutrons; 12.64 yr, room temp)	0.0461	9.1	75.5	10.8	0.997	-	0.9990
				9.4	2.95 x 10 ⁻³		
				6.0	3.93 x 10 ⁻⁴		
ML-Plag-Da (neutrons; 13.72 yr, room temp)	0.0380	14.4	90.4	16.6	0.99	-	0.9957
				13.8	4.45 x 10 ⁻³		
				11.7	1.39 x 10 ⁻³		
				8.6	1.40 x 10 ⁻⁴		
				5.2	8.24 x 10 ⁻⁶		
SURTp-Plag-Da (neutrons; 13.85 yr, room temp)	0.0483	33.0	106.0	-6.5	1.00 x 10 ⁻⁹	-	0.9999
				23.7	0.99999		
				20.9	1.13 x 10 ⁻⁵		
				5.7	1.00 x 10 ⁻⁹		
SW-1-Plag-Da (neutrons; 12.82 yr, room temp)	0.0594	3.5	68.3	8.2	0.72	-	0.9997
				8.1	0.28		
				6.4	5.72 x 10 ⁻⁴		
				3.3	2.16 x 10 ⁻⁴		
				2.3	4.09 x 10 ⁻⁵		
				-2.7	2.0 x 10 ⁻⁹		
				-4.1	1.0 x 10 ⁻⁹		
				-5.0	1.0 x 10 ⁻⁹		
				-7.5	1.0 x 10 ⁻⁹		



3.2 ^3He diffusion kinetics in pyroxene

320 Results from the ^3He step-heating experiment applied to two pyroxene sample groups are shown in Table S2 and Fig. 5. All
pyroxene samples display non-linear Arrhenius behavior. While the effect of gas loss during irradiation and storage may be
negligible or non-existent for the retentive pyroxene grains, we apply similar modeled irradiation and storage conditions to
these samples as for the feldspar samples, as outlined in Table 1. The forward-modeled MDD diffusion kinetics for pyroxene
are shown in Fig. 6 and Fig. 7.

325

Both gem-quality orthopyroxene and clinopyroxene grains show similar diffusion behavior, with the majority of the gas
being released in just a few heating steps, resulting in linearity in the Arrhenius plot (Fig. 7). Although the MDD model
inversion assumes that all gas at the beginning and end of the diffusion experiment belongs to the same material, which is
likely an inaccurate simplification due to natural crystal fragments deviating from perfect spheres (Huber et al., 2011), the
330 model still assigns more than 95% of the gas to a single domain without directed by arbitrary data selection. This suggests
that both gem-quality pyroxene grains likely exhibit a single domain. In effect, the remaining small domains (<3.5%) are
used as nuisance parameters so that model fitting is not overly restricted by arbitrary data selection.

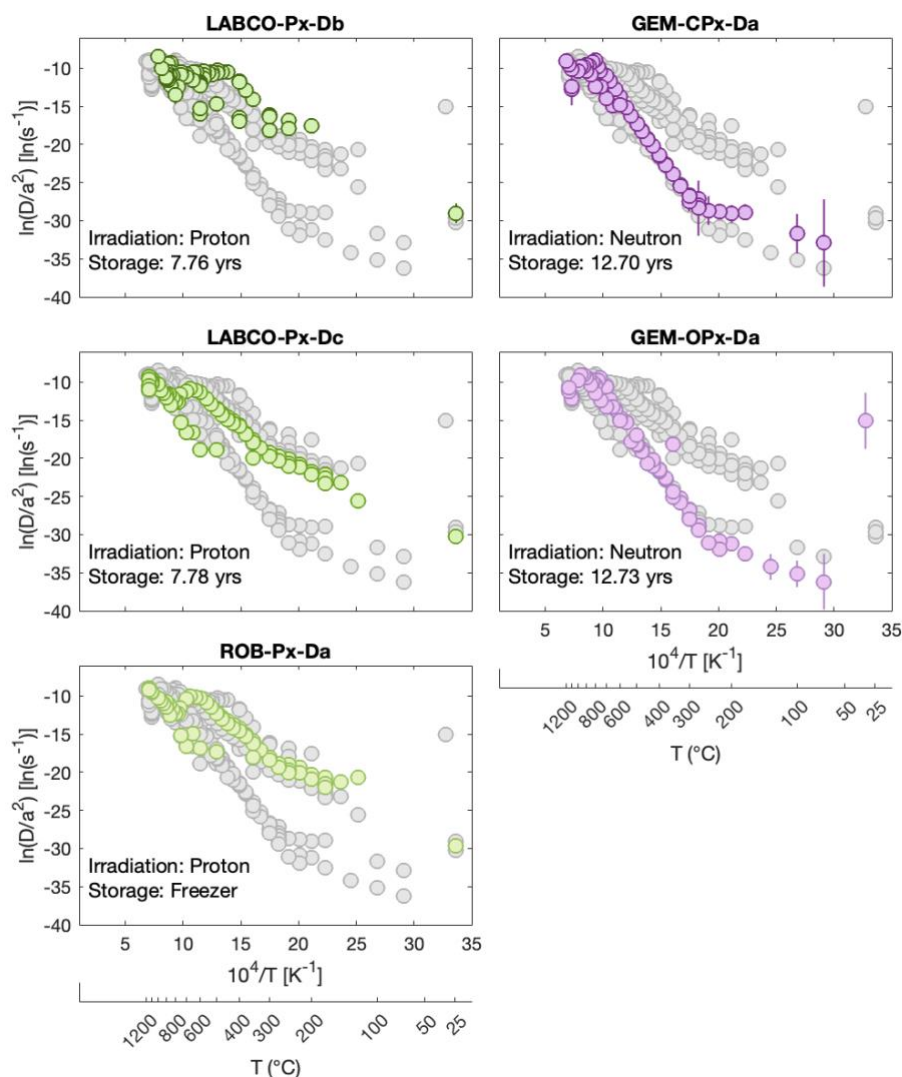


Figure 5: Arrhenius plot for ^3He diffusivity in pyroxene grains. Grey data points represent all pyroxene diffusion experiments for comparison, and colored data points represent a unique sample. Vertical lines show the uncertainty in $\ln(D/a^2)$ and are estimated using error propagation derived by Ginster and Reiners (2018), although the uncertainty is usually smaller than the symbol.

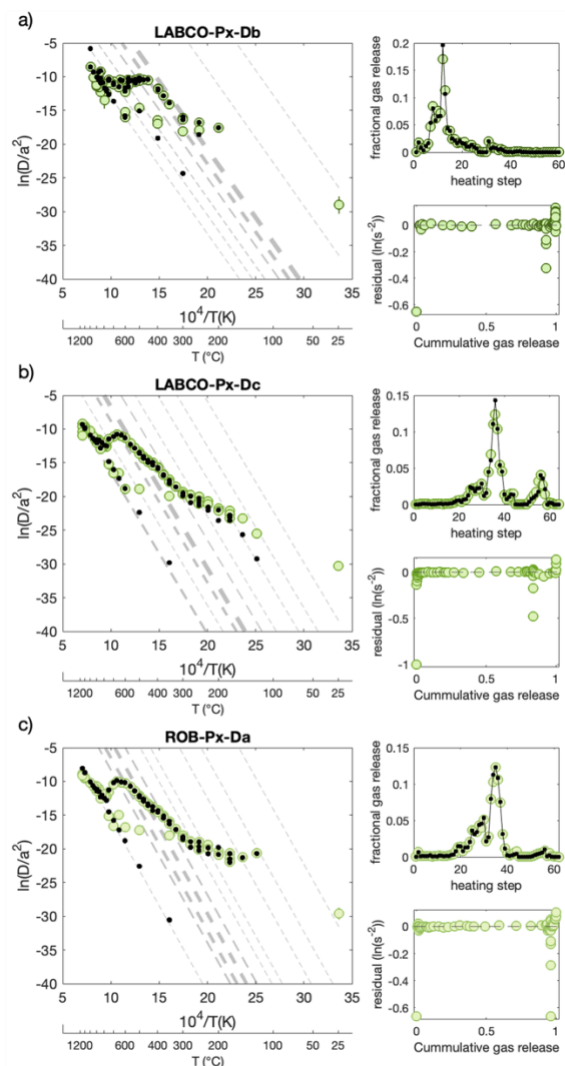


Figure 6: Forward-modeled MDD results for the Ferrar Dolerite pyroxene. Each sample result (a-c) is displayed in the Arrhenius, fractional release, and residual plots. Colored circles show the observed apparent ^3He diffusivity from the step-heating experiment. Vertical lines show the uncertainty in $\ln(D/a^2)$ (see description in Fig. 5 for more details). Black dots show the modeled predicted apparent diffusivity from the best reduced misfit, where the grey lines represent the diffusion kinetics of each individual domain modeled, with the line thickness being proportional to the fractional size of the domain. See Fig. 3 for more details on plotting.

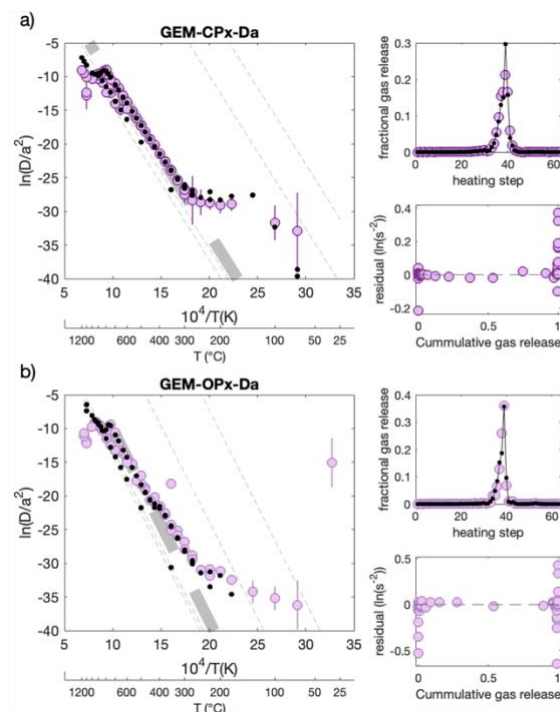


Fig. 7 Forward modeled MDD results for the gem-quality pyroxene. Each sample result (a-b) is displayed in the Arrhenius, fractional release, and residual plots. Colored circles show the observed apparent ^3He diffusivity from the step-heating experiment. Vertical lines show the uncertainty in $\ln(D/a^2)$ (see description in Fig. 5 for more details). Black dots show the modeled predicted apparent diffusivity from the best reduced misfit, where the grey lines represent the diffusion kinetics of each individual domain modeled, with the line thickness being proportional to the fractional size of the domain. See Fig. 3 for more details on plotting.



350 **Table 3: Summary of the modeled MDD diffusion kinetics results for all pyroxene grains**

Sample name	Radii (cm)	Misfit	Ea (KJ mol ⁻¹)	ln(D0/a ²) ln(s ⁻¹)	Fractional release	Irradiation loss	Storage loss
LABCO-Px-Db (protons; 7.76 yr, room temp)	0.0270	14.63	167.84	41.43	0.02	8.80 x 10 ⁻⁵	9.28 x 10 ⁻⁴
				31.40	0.01		
				19.60	0.39		
				17.47	0.29		
				15.37	0.13		
				12.39	0.09		
				9.84	0.07		
				7.10	0.02		
LABCO-Px-Dc (protons; 7.78 yr, room temp)	0.0261	15.27	198.92	41.96	0.0038	7.51 x 10 ⁻⁸	5.18 x 10 ⁻⁷
				35.27	0.0046		
				30.50	0.0045		
				24.91	0.0487		
				21.39	0.1013		
				17.19	0.3729		
				15.09	0.2210		
				11.71	0.0452		
ROB-Px-Da (protons; 0.18 yr, -20°C)	0.0267	17.08	212.81	54.517	0.007	4.99 x 10 ⁻⁶	-
				44.483	0.006		
				38.724	0.005		
				31.061	0.025		
				28.326	0.053		
				24.362	0.147		
				21.578	0.293		
				19.062	0.263		
GEM-CPx-Da (neutrons; 12.70 yr, room temp)	0.0426	13.96	189.15	45.23	0.0003		1.97 x 10 ⁻⁶
				35.68	0.0001		
				12.19	0.9819		
				9.99	0.0163		
				8.27	0.0014		
GEM-Opx-Da (neutrons; 12.73 yr, room temp)	0.1983	27.85	236.96	49.74	0.0000		1.29 x 10 ⁻¹⁰
				33.13	0.0022		
				18.45	0.9566		
				15.32	0.0346		
				14.32	0.0066		



4 Discussion

4.1 Fitting of the forward MDD model to the data set

The forward MDD model predicts the fractional gas release and apparent diffusion constant for each heating step across a given number of domains for a degassing experiment. Overall, the predicted apparent diffusivity shows good agreement with that observed during the step-heating experiments. However, for the high-temperature heating steps, diffusivity does not align well with the observed data for both plagioclase and pyroxene grains. This discrepancy is primarily due to the fact that the fractional loss from these high-temperature steps accounts for less than 1% of the total gas release. For example, heating steps 27 to 55 for ROB-Plag-Db (Fig. 3d and Table S1D) contribute only 1% of the total gas release, making the model fitting statistically insignificant for these steps.

For samples that have experienced prolonged storage periods prior to analysis, the optimized diffusion kinetic parameters become less constrained. We show this in Fig. 8, where a best fit is achieved across a range of activation energies. Previous misfit statistics for MDD inversion models have relied on either fractional misfit, as used in the study, or Chi-square misfit (Tremblay et al., 2014b; Tremblay et al., 2017; Gorin et al., 2024). While both methods effectively constrain the activation energy for samples experiencing minimal initial gas loss (e.g. ROB-Plag), they provide a weaker constraint on the diffusion kinetics for samples having lost significant gas (>25%, see LABCO-Plag in Table 2). Thus, as expected from the general principle that diffusion is an information-destroying process, gas loss prior to the experiment degrades the ability of a model inversion to correctly estimate diffusion kinetics.

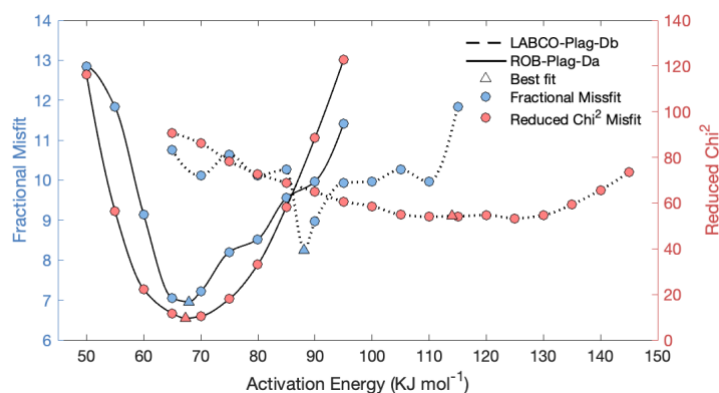


Figure 8: Forward-modeled MDD inversion misfit values for a range of activation energies. To generate these results, the activation energy was fixed at a given value and the model inversion was allowed to optimize all other parameters. Blue dots show the fractional misfit, as used in this study. Red dots show the reduced chi-square misfit value as implemented in previous MDD studies (Gorin et al. 2024). Triangles show the best misfit obtained and displayed in Table 2 for LABCO-Plag-Db (dashed line) and ROB-Plag-Da (solid line).



375 4.2 Diffusive loss prior to step-heating experiment

4.2.1 Short-term loss during irradiation

It is evident that gas loss from irradiation only affects the apparent diffusivity for the low-temperature heating steps (Fig. 1a). This behavior is observed for the ROB plagioclase samples (Fig. 3c-d), where the first 25°C heating step has a measurable amount of gas, but the apparent diffusivity is lower than expected if no loss occurred. These samples were kept frozen after
 380 irradiation to prevent any diffusive loss of ^3He during their short storage time (~2 months) prior to the step heating analysis. Consequently, any gas loss prior to analysis is assumed to occur only during irradiation. In the modeled irradiation step, we find that, for example, ROB-Plag-Da has lost 25% of its gas. Since we do not account for the simultaneous production of ^3He gas during irradiation, this is likely an overestimate. If we apply the resulting diffusion kinetics and account for simultaneous production and diffusion of ^3He during irradiation (using the model for production-diffusion forward simulation: Gribenski
 385 et al., 2022), this loss is reduced to 17%.

One could argue, given the high diffusivity of ROB plagioclase, gas loss occurs in three stages prior to analysis: (1) irradiation loss, (2) loss during freezer storage, and (3) loss during a 24-hour period at room temperature for when the sample is transferred to the mass spectrometer and placed under high vacuum. Therefore, the true gas loss prior to analysis likely
 390 falls between the modeled gas loss that includes all three stages and that modeled by excluding the irradiation step. If we exclude irradiation loss and consider only gas loss during the 2-month freezer storage and 24-hour of sample transfer to the mass spectrometer at room temperature (two heating steps; -20°C for 2 months and 21°C for 24 hr), the forward modeled MDD kinetics predicts a gas loss of 13% during freezer storage and 7% during sample transfer, with a total loss of 20% and activation energy of 69.6 kJ mol⁻¹. When we include diffusion loss during irradiation (not including production), the modeled
 395 gas losses for all three stages are 30%, 5%, and 5%, respectively, with a total loss of 39% and reduced activation energy of 66.9 kJ mol⁻¹.

The lower apparent diffusivity observed in the first few heating steps, caused by a diffusively rounded profile of ^3He in the grain, can be explained by a 20-39% gas loss prior to step-heating analysis. While the specific contribution of each degassing
 400 stage remains unclear, accounting for initial loss and non-homogeneous distribution prior to the diffusion experiment explains the observed gas loss in the initial 25°C heating step, with negligible effects at temperatures > 70°C. This approach ensures that all data points are used for modeling the diffusion kinetics.

We acknowledge that assigning diffusion loss to either irradiation and/or storage loss may appear arbitrary, as the initial
 405 conditions affect the total inferred gas loss. This becomes particularly significant as we lack precision on the temperatures during irradiation and extended storage time. However, the details of the initial gas loss conditions in the model inversion have less of an effect on the activation energy and overall diffusion kinetics. In summary, for samples exhibiting loss



reflected as apparent diffusivity in the Arrhenius plot the inclusion of initial gas loss is important for constraining the activation energy, however, the precise allocation of gas for these initial steps within the inversion model is dependent on the gas loss conditions but have less effect on the estimation of the activation energy

4.2.2 Neutron-irradiation complications

For a diffusive mineral, neutron irradiation is expected to result in a non-homogeneous distribution of ^3He concentration. Although the exact temperature during neutron irradiation has not been quantified, and therefore not included in the inversion model to predict irradiation loss, it is estimated to reach at least 150°C during a 50-hour irradiation period. Consequently, the simultaneous production and diffusion of ^3He and ^4He likely result in a heterogeneous distribution within the various feldspar grains. When combined with prolonged storage, exceeding three years, this leads to $>90\%$ gas loss prior to the diffusion experiment analysis (Fig. 9).

As previously mentioned, the model inference of more than 99% gas loss for the neutron-irradiated samples is unexpected. Thus, to validate this, we measured total ^3He and ^4He concentrations in a series of Fish Canyon sanidine (FCs) grains having experienced varying storage durations (Table S3).

In Fig. 9, we show He retention in the FCs grains over 15 years of storage. Because we cannot independently estimate the initial He concentration at the end of irradiation, we assume that the samples with the shortest storage time (1.8 yr) experienced an average of 79% gas loss, as predicted by the forward MDD model using the diffusion kinetic for FCs-Plag-Da (Table 2), and back-calculate the initial concentration for all grains on this basis. All samples were neutron-irradiated for a duration of 50 hr in the same reactor position. However, the grain size and amount in each sample varies for the 7 samples and therefore contain various subgrain domains that do not scale with He diffusion. Note that the J-factors obtained from $^{39}\text{Ar}/^{40}\text{Ar}$ analysis of simultaneously irradiated grains are within 3% of each other. With this normalization, it is evident that the observed decrease in He concentration with storage time conforms to the model predicted, which supports the finding from model inversion of nearly complete gas loss after 12-14 years of storage (Table 2).

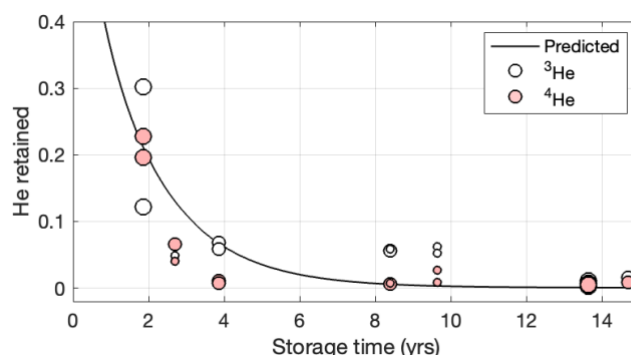


Figure 9: He retention in Fish Canyon sanidine (FCs) grains. The black line shows the ^3He retention predicted from the diffusion kinetics modeled for FCs (Table 2). The circles show the observed ^3He (white) and ^4He (red) retention from measured concentrations. The size of the circles is related to the average radius of the grain sizes of the FCs sample analyzed (small: $<200\ \mu\text{m}$, medium: $200\text{--}400\ \mu\text{m}$, large: $>400\ \mu\text{m}$). The measured He concentrations and estimated fractional retention have been normalized to a storage of 1.8 years with a 79% loss.

The MDD inversion model can compensate storage time for activation energy, thereby reducing the constraints on diffusion kinetic parameters and potentially overestimating the activation energy due to the non-unique inversion solution, as shown in Fig. 8. Further, the MDD inversion model accounts for initial gas loss by assigning a fractional loss to two initially predefined steps within the inversion process. This approach allows the model to account for any heterogeneous distribution reflected by the increased apparent diffusivity observed in repeated low-temperature heating steps during the diffusion experiment. However, the model does not differentiate whether this heterogeneous distribution arises from irradiation, storage, or a combination of both but instead adjusts the assigned fractional loss to match the observed. Therefore, lack of constraint on the temperature during neutron irradiation, hence, the temperature assigned to the initial predefined steps, complicates quantifying the fractional loss from neutron irradiation. Consequently, for diffusive minerals and cases where spallogenic-produced nuclides are of interest, proton irradiation is preferred to achieve a more homogenous initial He distribution.

4.2.3 Long-term storage loss and implications

Although the model inversion obtains different activation energies for LABCO and ROB plagioclase (Table 2), we will argue that the modeled activation energy and domain parameters for LABCO are likely inaccurate because any signal from low-retentivity domains has been erased by significant gas loss during prolonged storage. To demonstrate this, we take the best-fit diffusion kinetics results for the ROB-Plag-Da sample (which was frozen) and simulate 7.7 years of extended storage experienced by the LABCO-Plag-Db sample using its step-heating experiment. This forward MDD model simulation predicts the experimental step-degassing results that would have occurred if the ROB sample had the same extended storage history as LABCO. Figure 10 shows that the same diffusion kinetics (from ROB) can produce an experimental result similar to that of LABCO if we include extended storage. If we take this prediction result and invert it using the MDD inversion model, we fail to recover the initial MDD diffusion kinetics parameters for ROB.

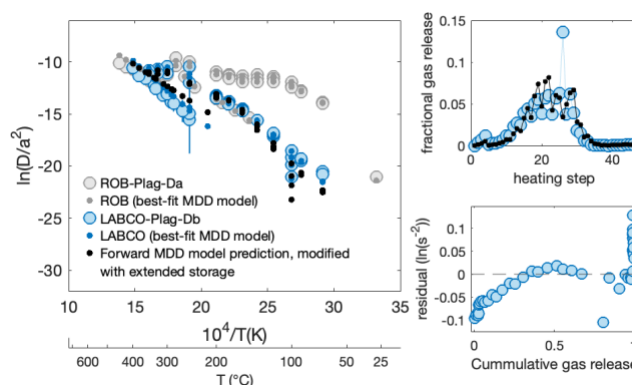


Figure 10: Comparison between the forward-modeled MDD results for the Ferrar Dolerite plagioclase, when including storage loss displayed in the Arrhenius, fractional release, and residual plot (See Fig. 3 for more details on plotting). Circles show the measured apparent diffusivity, and dots represent the forward-modeled predicted apparent diffusivity for LABCO-Plag-Db (blue) and ROB-Plag-Da (grey). Black dots show the predicted apparent diffusivity for LABCO-Plag-Db when applying the diffusion kinetics of ROB-Plag-Da and including initial gas loss during proton irradiation and 7.7 years of storage time at ~21°C.

During prolonged storage time, signals from the low-temperature domains can be erased by complete degassing of low-retentive diffusion domains. While the MDD inversion model can account for initial gas and heterogeneous distribution of He during irradiation and storage, it is not able to recover the diffusion kinetics from domains that have been completely degassed and information has been lost prior to the diffusion experiment. This is evident in the weak constraint on the activation energy, even when accounting for loss during and after irradiation (Fig. 8), leading to a higher apparent activation energy and, overall, an apparently more retentive mineral.

This raises the question of whether the diffusion kinetics from the various feldspar grains (Fig. 4 and Table 2) can be reliably interpreted as being unique to the compositions or if we are observing the effect of completely degassed low-retentive diffusion domains during storage, and therefore overestimating the activation energy and retentivity of ^3He in feldspars during inversion.

In summary, this suggests that (i) not including storage loss in the inversion for the diffusion kinetics produces incorrect kinetics, (ii) significant storage loss makes it impossible to get accurate kinetics even with a more complete MDD inversion model, or (iii) both.

4.3 Application of ^3He diffusion in plagioclase in paleothermometry

Measured ^3He diffusivity on a range of plagioclase compositions shows variable and complex diffusion behaviors. This is reflected in the retentivity of cosmogenic ^3He in plagioclase at Earth's surface temperatures, which varies greatly on a timescale $< 10^6$ years (Fig. 11). Here, we use the inverted MDD diffusion kinetics to model the long-term retentivity of ^3He



in plagioclase as a function of exposure time and constant effective diffusion temperature (EDT: Tremblay et al., 2017; Christodoulides et al., 1971) for a cold region climate (1°C), such as polar or high-altitude environments.

The retentivity between the two Ferrar Dolerite samples (LABCO and ROB) displays the largest discrepancy in retentivity. This can, in part, be explained by the signals from the low-temperature domains in LABCO being erased by complete degassing of diffusion domains during storage of the sample at room temperature, as discussed above. As a result, the MDD inversion will overestimate the diffusion kinetics, predicting a higher retentivity. Other factors such as effective diffusion radius, crystal structure, crystal chemistry, defective density, and state of order, can cause variable diffusivity between the grains of similar mineral groups and compositions (Gourbet et al., 2012). The ROB plagioclase samples are the least retentive and display significant diffusion loss within 100 years. As the ROB plagioclase samples have experienced the least amount of diffusive loss prior to diffusion analysis, the diffusion kinetics are presumed to be accurately represented for those samples. The LABCO plagioclase samples appear to be partially retentive beyond 10^3 years at 1°C and potentially useful as a paleothermometry for cold climate regions. However, the apparent diffusion kinetics for LABCO is potentially an overestimation and, therefore, not a true representation.

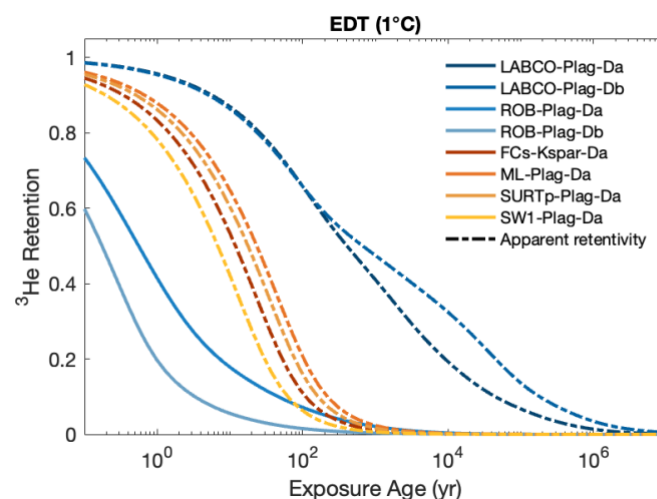


Figure 11: Cosmogenic ^3He retention in plagioclase with exposure age for an effective diffusion temperature (EDT) of 1°C. The ^3He retentivity is calculated based on the accumulation-diffusion equation in Wolf et al. (1998) and modified to accept multiple diffusion domains. Solid lines represent diffusion kinetics that are presumed to be accurate representation for the samples. Dashed lines indicate samples where the apparent diffusivity is likely an overestimation due to prolonged storage loss and therefore not a true representation.

Since the loss of ^3He during prolonged storage time prior to step-heating analysis can explain the discrepancy between the observed variation in retentivity, it is questionable if cosmogenic ^3He is retentive enough in plagioclase to be used as a paleothermometer in polar and high-altitude environments. However, the retention shown in Fig. 11 is specific to the analyzed grain size and has not been normalized. Larger grains corresponding to larger diffusion domains would, therefore, display larger retention, which could be advantageous for ^3He diffusivity in plagioclase for paleothermometry.



510 4.4 Implication of irradiation and storage loss on paleothermometry

In this study, cosmogenic ^3He in plagioclase grains are found to be diffusive at Earth's surface temperature and experience significant diffusive loss on timescales of a hundred years, making it unsuitable as a paleoclimate indicator, even in polar and high-altitude regions with persistent subzero surface temperature. Conversely, cosmogenic ^3He in quartz grains has been found to be partly diffusive at Earth's surface temperature on timescales $>10^3$ yr at subzero (Tremblay et al., 2014b) and, therefore, applicable to paleothermometry. However, the step-heating data sets from which this is determined all display an increased apparent diffusivity in the Arrhenius trend for the low-temperature heating steps, similar to that observed in plagioclase grains in this study (Fig. 3). The behavior in these first few heating steps was attributed to loss during or after irradiation and excluded from the regression fitting and the retentivity was established from a two-domain MDD model, including two separate activation energies determined from the high- and low-temperature Arrhenius linearity.

520 As the fitting scheme used by Tremblay et al. (2014b) has some properties (e.g., the use of multiple activation energies) that have the effect of compensating for spuriously low diffusivity in early steps affected by gas loss, we cannot directly compare the result of our MDD model that includes gas loss to the results of Tremblay et al. (2014b). However, we can use our forward MDD model to evaluate the effective magnitude of not including initial gas loss on the application for paleothermometry. Thus, we inverted the step-degassing data of the quartz sample HU-08-03 obtained by Tremblay et al. (2014b) with our (i) forward MDD model inversion lacking irradiation and storage loss, and (ii) forward MDD model inversion including both irradiation and storage loss (See Table S4 and Fig. S3 for details and diffusion kinetic results). This quartz sample was proton-irradiated for 5 hr in which the temperature did not exceed 45°C and kept in storage at room temperature for 1.3 years between irradiation and step-heating analysis. Further, radiocarbon ages from the moraine on which this sample was collected provide an exposure age of $12,350 \pm 200/-20$ yr (Kelly et al., 2012; Kelly et al., 2013), with cosmogenic ^3He measurements indicating a retentivity of 0.22 ± 0.03 .

535 In Fig. 12, we show that the modeled MDD inversion results, including initial gas loss, predict ^3He retention closely matching the observed Holocene value. Contrary, a model inversion lacking initial gas loss greatly overpredicts ^3He retention. This effect is most significant for any diffusive mineral that displays increased apparent diffusivity in the Arrhenius trend and resides within a partial retention zone at surface temperatures. While this does not suggest that the diffusion kinetics estimated by Tremblay et al. (2014b) are incorrect (since their fitting approach indirectly compensates for initial gas loss), it does highlight that not accounting for initial gas loss when determining the diffusion kinetics will have a significant effect on the paleothermometry estimates.

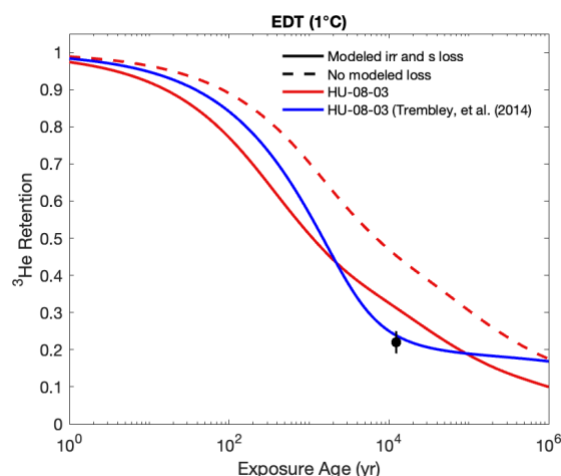


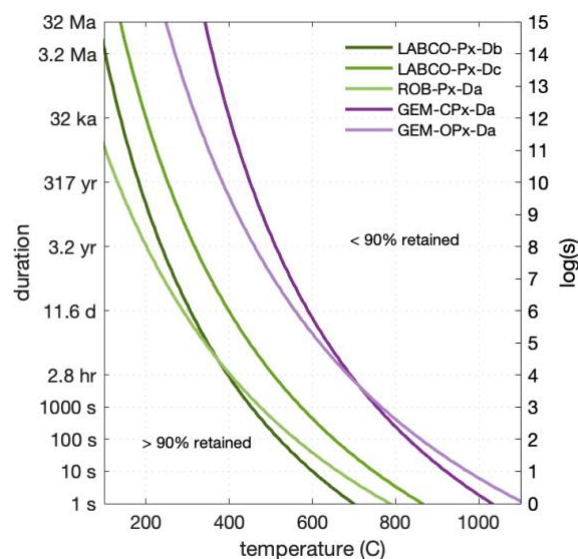
Figure 12: Cosmogenic ^3He retention in quartz sample, HU-08-03 (Tremblay et al., 2014b), with exposure age for an effective diffusion temperature (EDT) of 1°C . The ^3He retentivity is calculated based on the accumulation-diffusion equation in Wolf et al. (1998) and modified to accept multiple diffusion domains. The black data point shows the observed ^3He retentivity in HU-08-03 and its radiocarbon exposure age of $12,350 \pm 200/-20$ yr (Kelly et al., 2012; Kelly et al., 2013).

4.5 Retentivity of ^3He in pyroxene and its implication to exposure-dating

We find the activation energy for cosmogenic ^3He diffusion in pyroxene to be $168\text{--}237$ kJ mol^{-1} . The variability observed in the kinetics among pyroxene samples is less important from the perspective of cosmogenic ^3He retention at Earth surface temperatures, as the least retentive pyroxene will retain 95% of cosmogenic ^3He at 50°C on a timescale of millions of years, and other pyroxenes will retain $>>99\%$.

Further, all pyroxene samples are found to be more retentive in cosmogenic ^3He , compared to the only known estimates for ^4He diffusion in pyroxene, which has published activation energies of 116 kJ mol^{-1} and 124 kJ mol^{-1} (Lippolt and Weigel, 1988)(Lippolt and Weigle, 1988). While Lippolt and Weigel (1988) found pyroxene to be retentive on a > 100 Ma timescale, these lower activation energies were derived from linear regression of low-temperature heating steps. If a similar methodology were applied to the ROB and LABCO pyroxene, those would likely show comparable low activation energies (Fig. 6b-c). The activation energy and retentivity for ^3He diffusion are higher in orthopyroxene than in clinopyroxene. While less retentive, this observation is consistent with that observed for cosmogenic ^{21}Ne diffusivity in orthopyroxene and clinopyroxene (Gourbet et al., 2012), attributing this difference to the atomic structure of the pyroxene phases.

In Fig. 11, we show the duration and temperature curves that will result in a 10% diffusive loss of accumulated cosmogenic ^3He in the various pyroxene samples. While retentive at Earth's surface temperature, a brief preheating (<1000 s), such as an impact event (Shuster et al., 2010), volcanic activity, or wildfires (Balco et al., 2023) will significantly disturb the cosmogenic ^3He concentration for temperatures as low as 400°C .



565 **Figure 13: Cosmogenic ^3He retention in pyroxene for duration and temperature pairs that predict 10% diffusive loss of cosmogenic ^3He in pyroxene. The retentivity is calculated based on the accumulation-diffusion equation in Wolf et al. (1998) and modified to accept multiple diffusion domains.**

Given that few places on Earth have surface temperatures that consistently exceed 50°C , cosmogenic ^3He will generally be quantitatively retained in pyroxene over timescales and temperatures relevant for exposure dating. This is consistent with numerous measurements of cosmogenic ^3He exposure ages in pyroxene that show (i) high temperatures are required to extract He from pyroxene and (ii) measured ^3He concentrations in pyroxene are consistent with exposure ages determined from other geological evidence (e.g., Kurz, 1986; Schäfer et al., 1999; Balter-Kennedy et al., 2020).

The distinct difference in the retentivity between cosmogenic ^3He in pyroxene and plagioclase could be beneficial for exposure dating of ^3He in pyroxene. The Antarctic Ferrar dolerite samples consist predominantly of plagioclase and pyroxene, with minor traces of oxides (Harvey, 2001; Elliot and Fleming, 2021). The crystal texture is mostly fine-grained, making the separation and isolation procedure tedious to obtain monomineralic pyroxene grains of $100 - 500 \mu\text{m}$ fragments, as often used in cosmogenic nuclide exposure dating. However, as cosmogenic-produced ^3He in plagioclase is diffusive at Earth's surface temperature, a separation process may not be necessary as an insignificant amount of ^3He is naturally retained in the plagioclase grains. Thus, measuring the concentration of ^3He in a piece of whole rock of Ferrar dolerite could be directly converted into a ^3He concentration within the pyroxene fragment, eliminating the laboratory procedure for separating and isolating pyroxene grains.



5 Conclusion

In this study, we measured ^3He diffusivity in various plagioclase and pyroxene grains through a series of stepwise degassing experiments and applied a ‘multiple-diffusion-domain’ (MDD) inversion model to determine the diffusion kinetics. We find the apparent diffusion kinetics of ^3He in plagioclase to be greatly variable. However, this variability can be explained by the MDD inversion models’ inability to constrain the diffusion parameters when significant gas has been lost during either irradiation and/or prolonged storage prior to the diffusion experiment. In such cases, the forward MDD model overestimates the activation energy by compensating for storage loss. Further, for diffusive minerals (e.g., $<70 \text{ KJ mol}^{-1}$), prolonged storage time at room temperature can result in complete degassing of low-retentive domains, limiting the MDD inversion model’s ability to constrain the diffusion kinetics. As a result, samples that were kept frozen for a short period of time immediately after irradiation until experimental analysis (ROB-Plag) yielded the most reliable estimate of the activation energy and diffusion parameters. This highlights the importance of (i) limiting storage time between irradiation and analysis, (ii) accounting for any diffusive loss occurring before the diffusion experiment when determining the diffusion kinetics of less retentive minerals, and (ii) freeze samples to limit gas loss during storage time.

Overall, our findings suggest that ^3He in plagioclase is diffusive at Earth’s surface temperatures on a timescale of hundred years, making it unsuitable for paleothermometry. On the contrary, pyroxene retains cosmogenic ^3He at Earth’s surface temperatures over millions of years, though brief preheating ($<1000\text{s}$) could significantly disturb its concentration. The distinct difference in cosmogenic ^3He retentivity between pyroxene and plagioclase could, however, be beneficial for exposure dating of ^3He in mafic rock consisting predominantly of pyroxene and plagioclase.

Code and data availability

All data described in the paper as well as the MATLAB code for forward modeling, model fitting, and figure generation, are included in the Supplement.

605 Author contribution

MB carried out the He diffusion experiments and measurements, and GB and ALG assisted in data reduction. MB developed the forward MDD model with contributions from GB and ALG on the theory. MB led the data interpretation and prepared the manuscript with contributions from all authors.

Competing interests

610 Greg Balco is a member of the editorial board of Geochronology.



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References

- Ackert, R. P., Jr., Barclay, D. J., Borns, H. W., Jr., Calkin, P. E., Kurz, M. D., Fastook, J. L., and Steig, E. J.: Measurements of Past Ice Sheet Elevations in Interior West Antarctica, *Science*, 286, 276-280, 10.1126/science.286.5438.276, 1999.
- 620 Balco, G., Hidy, A., Struble, W. T., and Roering, J. J.: Short communication: Cosmogenic noble gas depletion in soils by wildfire heating, *geochronology Discussions*, 10.5194/gchron-2023-23, 2023.
- Balter-Kennedy, A., Bromley, G., Balco, G., Thomas, H., and Jackson, M. S.: A 14.5-million-year record of East Antarctic Ice Sheet fluctuations from the central Transantarctic Mountains, constrained with cosmogenic ^3He , ^{10}Be , ^{21}Ne , and ^{26}Al , *The Cryosphere*, 14, 2647-2672, 10.5194/tc-14-2647-2020, 2020.
- 625 Balter-Kennedy, A., Schaefer, J. M., Schwartz, R., Lamp, J. L., Penrose, L., Middleton, J., Hanley, J., Tibari, B., Blard, P.-H., Winckler, G., Hidy, A. J., and Balco, G.: Cosmogenic ^{10}Be in pyroxene: laboratory progress, production rate systematics, and application of the ^{10}Be - ^3He nuclide pair in the Antarctic Dry Valleys, *Geochronology*, 5, 301-321, 10.5194/gchron-5-301-2023, 2023.
- Baxter, E. F.: Diffusion of Noble Gases in Minerals, *Reviews in Mineralogy and Geochemistry*, 72, 509-557, 10.2138/rmg.2010.72.11, 2010.
- 630 Brook, E. J., Kurz, M. D., Ackert, R. P., Denton, G. H., Brown, E. T., Raisbeck, G. M., and Yiou, F.: Chronology of Taylor Glacier Advances in Arena Valley, Antarctica, Using in Situ Cosmogenic ^3He and ^{10}Be , *Quaternary Research*, 39, 11-23, 10.1006/qres.1993.1002, 1993.
- Bruno, L. A., Baur, H., Graf, T., Schlu, C., Signer, P., and Wieler, R.: Dating of Sirius Group tillites in the Antarctic Dry Valleys with cosmogenic ^3He and ^{21}Ne , *Earth and Planetary Science Letters*, 147, 37-54, 1997.
- 635 Cassata, W. S. and Renne, P. R.: Systematic variations of argon diffusion in feldspars and implications for thermochronometry, *Geochimica et Cosmochimica Acta*, 112, 251-287, 10.1016/j.gca.2013.02.030, 2013.
- Cassata, W. S., Renne, P. R., and Shuster, D. L.: Argon diffusion in pyroxenes: Implications for thermochronometry and mantle degassing, *Earth and Planetary Science Letters*, 304, 407-416, 10.1016/j.epsl.2011.02.019, 2011.
- 640 Cebula, G. T.: The Fish Canyon Tuff, a potential standard for the ^{40}Ar - ^{39}Ar and fission-track dating methods, 6th Int. Conf. Geochr. Cosmochr. Isot. Geol. Abstr. 139, 1986.
- Cerling, T. E.: Dating Geomorphologic Surfaces Using Cosmogenic ^3He , *Quaternary Research*, 33, 148-156, 10.1016/0033-5894(90)90015-d, 1990.
- Christodoulides, C., Ettinger, K. V., and Fremlin, J. H.: The use of TL glow peaks at equilibrium in the examination of the thermal and radiation history of materials, *Modern Geology*, 2, 275-280, 1971.
- 645 Craig, H. and Poreda, R. J.: Cosmogenic ^3He in terrestrial rocks: The summit lavas of Maui, *Proc Natl Acad Sci U S A*, 83, 1970-1974, 10.1073/pnas.83.7.1970, 1986.
- Elliot, D. H. and Fleming, T. H.: Chapter 2.1b Ferrar Large Igneous Province: petrology, Geological Society, London, *Memoirs*, 55, 93-119, 10.1144/m55-2018-39, 2021.
- 650 Fechtig, H. and Kalbitzer, S.: The diffusion of argon in potassium-bearing solids, in: *Potassium-Argon Dating*, edited by: Schaeffer, O. A., and Zahringer, J., Heidelberg, Springer, 68-106, 1966.



- Ginster, U. and Reiners, P. W.: Error Propagation in the Derivation of Noble Gas Diffusion Parameters for Minerals From Step Heating Experiments, *Geochemistry, Geophysics, Geosystems*, 19, 3706-3720, 10.1029/2018gc007531, 2018.
- Gorin, A. L., Gorin, J. M., Bergelin, M., and Shuster, D. L.: An optimization tool for identifying multiple-diffusion domain model parameters, *Geochronology*, 6, 521-540, 10.5194/gchron-6-521-2024, 2024.
- 655 Gourbet, L., Shuster, D. L., Balco, G., Cassata, W. S., Renne, P. R., and Rood, D.: Neon diffusion kinetics in olivine, pyroxene and feldspar: Retentivity of cosmogenic and nucleogenic neon, *Geochimica et Cosmochimica Acta*, 86, 21-36, 10.1016/j.gca.2012.03.002, 2012.
- Gribenski, N., Tremblay, M. M., Valla, P. G., Balco, G., Guralnik, B., and Shuster, D. L.: Cosmogenic ^3He paleothermometry on post-LGM glacial bedrock within the central European Alps, *Geochronology Discussions*, 10.5194/gchron-2022-1, 2022.
- 660 Harvey, R. P.: The Ferrar Dolerite: An Antarctic analog for martian basaltic lithologies and weathering processes, *Field Trip and Workshop on the Martian Highlands and Mojave Desert Analogs*, 25, 2001.
- Huber, C., Cassata, W. S., and Renne, P. R.: A lattice Boltzmann model for noble gas diffusion in solids: The importance of domain shape and diffusive anisotropy and implications for thermochronometry, *Geochimica et Cosmochimica Acta*, 75, 2170-2186, 10.1016/j.gca.2011.01.039, 2011.
- 665 Kelly, M. A., Lowell, T. V., Applegate, P. J., Smith, C. A., Phillips, F. M., and Hudson, A. M.: Late glacial fluctuations of Quelccaya Ice Cap, southeastern Peru, *Geology*, 40, 991-994, 10.1130/g33430.1, 2012.
- Kelly, M. A., Lowell, T. V., Applegate, P. J., Phillips, F. M., Schaefer, J. M., Smith, C. A., Kim, H., Leonard, K. C., and Hudson, A. M.: A locally calibrated, late glacial ^{10}Be production rate from a low-latitude, high-altitude site in the Peruvian Andes, *Quat Geochronol*, 26, 70-85, 10.1016/j.quageo.2013.10.007, 2013.
- 670 Kurz, M. D.: In situ production of terrestrial cosmogenic helium and some applications to geochronology, *Geochimica et Cosmochimica Acta*, 50, 2855-2862, 10.1016/0016-7037(86)90232-2, 1986.
- Lippolt, H. J. and Weigel, E.: ^4He diffusion in ^{40}Ar -retentive minerals, *Geochimica et Cosmochimica Acta*, 52, 1449-1458, 10.1016/0016-7037(88)90215-3, 1988.
- 675 Lovera, O. M., Richter, F. M., and Harrison, T. M.: The $^{40}\text{Ar}/^{39}\text{Ar}$ thermochronometry for slowly cooled samples having a distribution of diffusion domain sizes, *Journal of Geophysical Research: Solid Earth*, 94, 17917-17935, 10.1029/JB094iB12p17917, 1989.
- Lovera, O. M., Grove, M., Mark Harrison, T., and Mahon, K. I.: Systematic analysis of K-feldspar step heating results: I. Significance of activation energy determinations, *Geochimica et Cosmochimica Acta*, 61, 3171-3192, 10.1016/s0016-7037(97)00147-6, 1997.
- 680 McDougall, I. and Harrison, T. M.: *Geochronology and Thermochronology by the $^{40}\text{Ar}/^{39}\text{Ar}$ Method*, Oxford University Press 1999.
- Megrue, G. H.: Rare-gas chronology of calcium-rich achondrites, *Journal of Geophysical Research*, 71, 4021-4027, 10.1029/JZ071i016p04021, 1966.
- 685 Phillips, D., Matchan, E. L., Dalton, H., and Kuiper, K. F.: Revised astronomically calibrated $^{40}\text{Ar}/^{39}\text{Ar}$ ages for the Fish Canyon Tuff sanidine – Closing the interlaboratory gap, *Chemical Geology*, 597, 10.1016/j.chemgeo.2022.120815, 2022.
- Reiners, P. W., Ehlers, T. A., and Zeitler, P. K.: Past, Present, and Future of Thermochronology, *Reviews in Mineralogy and Geochemistry*, 58, 1-18, 10.2138/rmg.2005.58.1, 2005.
- 690 Schäfer, J. M., Ivy-Ochs, S., Wieler, R., Leya, I., Baur, H., Denton, G. H., and Schlüchter, C.: Cosmogenic noble gas studies in the oldest landscape on earth: surface exposure ages of the Dry Valleys, Antarctica, *Earth and Planetary Science Letters*, 167, 215-226, 10.1016/s0012-821x(99)00029-1, 1999.
- Shuster, D. L. and Farley, K. A.: Diffusion kinetics of proton-induced ^{21}Ne , ^3He , and ^4He in quartz, *Geochimica et Cosmochimica Acta*, 69, 2349-2359, 10.1016/j.gca.2004.11.002, 2005.
- 695 Shuster, D. L., Farley, K. A., Sisterson, J. M., and Burnett, D. S.: Quantifying the diffusion kinetics and spatial distributions of radiogenic ^4He in minerals containing proton-induced ^3He , *Earth and Planetary Science Letters*, 217, 19-32, 10.1016/s0012-821x(03)00594-6, 2004.
- Shuster, D. L., Balco, G., Cassata, W. S., Fernandes, V. A., Garrick-Bethell, I., and Weiss, B. P.: A record of impacts preserved in the lunar regolith, *Earth and Planetary Science Letters*, 290, 155-165, 10.1016/j.epsl.2009.12.016, 2010.
- 700 Tremblay, M. M., Shuster, D. L., and Balco, G.: Cosmogenic noble gas paleothermometry, *Earth and Planetary Science Letters*, 400, 195-205, 10.1016/j.epsl.2014.05.040, 2014a.



- Tremblay, M. M., Shuster, D. L., and Balco, G.: Diffusion kinetics of ^3He and ^{21}Ne in quartz and implications for cosmogenic noble gas paleothermometry, *Geochimica et Cosmochimica Acta*, 142, 186-204, 10.1016/j.gca.2014.08.010, 2014b.
- 705 Tremblay, M. M., Shuster, D. L., Balco, G., and Cassata, W. S.: Neon diffusion kinetics and implications for cosmogenic neon paleothermometry in feldspars, *Geochimica et Cosmochimica Acta*, 205, 14-30, 10.1016/j.gca.2017.02.013, 2017.
- Ugray, Z., Lasdon, L., Plummer, J., Glover, F., Kelly, J., and Martí, R.: Scatter Search and Local NLP Solvers: A Multistart Framework for Global Optimization, *INFORMS Journal on Computing*, 19, 328-340, 10.1287/ijoc.1060.0175, 2007.
- 710 Wolf, R. A., Farley, K. A., and Kass, D. M.: Modeling of the temperature sensitivity of the apatite (U–Th)/He thermochronometer, *Chemical Geology*, 148, 105-114, 10.1016/s0009-2541(98)00024-2, 1998.